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25

Abstract

26 The Tibetan Plateau (TP), referred to as the “roof of the world” is also known as the
27 “world water tower”, because it contains a large amount of water resources and
28 ceaselessly transports these waters to its surrounding areas. However, it is not
29 clear how these waters are being supplied and replenished. In particular, how
30 plausible hydrological cycles can be realized between tropical oceans and the TP. In
31 order to explore the mechanism sustaining the atmospheric “water tower” over the
32 TP, the relationship of a “heat source column” over the plateau and moist flows in
33 the Asian summer monsoon circulation is investigated, here we show that the
34 plateau’s thermal structure leads to dynamic processes with an integration of two
35 couples of lower convergences and upper divergences, respectively, over the
36 plateau’s southern slopes and main platform, which relay moist air in two ladders
37 up to the plateau. Similarly to the CISK (Conditional Instability of the Second Kind)
38 mechanism of tropical cyclones, the elevated warm-moist air, in turn, forces
39 convective weather systems, hence building a water cycle over the plateau. An
40 integration of mechanical and thermal TP-forcing is revealed in relation to the Asian
41 summer monsoon circulation knitting a close tie of vapor transport from tropical
42 oceans to the atmospheric “water tower” over the TP.

43

1. Introduction

44 It has long been known that the Tibetan Plateau (TP) as the third pole and “the
45 world water tower” (Xu et al., 2008; Qiu, 2008) plays an important and special role in
46

47 global climate and energy/water cycle. In particular, due to its elevated land surface
48 and thus enhanced sensible heating, the TP becomes a unique heat source,
49 nonexistent in any other part of the world (Flohn, 1957;Yeh et al., 1957;Yanai et al.,
50 1992;Webster et al., 1998;Wu and Zhang, 1998;An et al., 2001;Sugimoto and Ueno,
51 2010). From its topographic structure, we know that the TP possesses steep slopes
52 with dramatic rising of land surfaces on its south and east rims. Over the plateau,
53 however, the TP extends into north and west extensively in a relatively flat fashion,
54 thus being presented as an oversized “mesa”, although there are large mountains
55 over the TP triggering convective cloud formations. In the boreal summer, this
56 massive “mesa” is strongly heated by solar radiation. One of the consequences of
57 this thermal structure is its virtual functionality serving as an “air pump”, which
58 attracts warm and moist air from low-latitude oceans up to the north into the Asian
59 continent(Wu et al., 1997;Wu et al., 2012). During boreal winter, this flow pattern
60 reverses with the TP’s cooling source(Ding, 1994). Hence, the TP’s role in the
61 world’s largest monsoon system is explained.

62

63 Furthermore, classic studies (Flohn, 1957;Yeh et al., 1957;Luo and Yanai, 1984;Wu
64 and Zhang, 1998;Yanai et al., 1992;Hahn and Manabe, 1975;Webster et al., 1998;Xu
65 et al., 2010;Ye and Gao, 1979) also indicate that the rising warm and moist air from
66 the tropical oceans tends to be deflected predominantly to the right (carried along
67 the mid-latitude westerlies), once encountered with the sharply elevated plateau.
68 The deflected warm and moist air forms the well-known “south-westerly monsoonal
69 flows”, transporting water vapor down to the southeastern China, plausibly

70 explaining the abundant water resources in these areas (Xu et al., 2010;Zhao and
71 Chen, 2001;Xu et al., 2012) (see the small rectangle in the low reach of Yangtze
72 River basin in upper panel of Fig. 1). The lower southwesterly driving warm
73 and wet air transport from tropical oceans to these areas of southeastern China in
74 summer season could also induced by the conjunction of the TP and Eurasia
75 continental thermal forcing (Duan and Wu, 2005).

76

77 However, many environment resource surveys (Lu et al., 2005;Yao et al., 2012;Qiu,
78 2008) confirm that the TP itself contains a large amount of water resources, in the
79 forms of snowpacks, glaciers, lakes, rivers, and aquifers (the large rectangle over the
80 TP in upper panel of Fig. 1). The TP region contains one of the richest water
81 resources and constitutes one of the densest hydrological systems in the world. Xu
82 et al. (2008) identified the role of TP as the world water tower, and elucidated how
83 a hydrological cycle is completed over the plateau and its surrounding areas and
84 how atmosphere is able to supplement and reinforce the water that has been
85 continuously transported away from the TP. These studies certainly indicate that
86 despite the fact that a large amount of water vapor is deflected to the southeast
87 China, there must be appreciable amount of moist flows that are able to climb over
88 the TP, supplying and depositing a necessary amount of water on to the TP, to make
89 up the depleting surface flows.

90

91 In this study focusing on the climate mean in boreal summer, we investigate a
92 mechanism how a portion of moist air reaches over the TP to maintain the

93 atmospheric “water tower”, as shown with high vapor contents over the TP in lower
94 panel of Fig. 1. The mechanism depicts an understanding of dynamic and
95 thermodynamic processes forcing the moist air up to the plateau. In particular, a
96 coupling of two “dynamic pumps” with the CISK (Conditional Instability of the
97 Second Kind) mechanism similar to the typhoon’s thermal forcing, contiguous
98 horizontally but staggered vertically, are revealed. The two “water connected
99 pumps” will mutually support each other in such a way that they ladder and relay
100 the moist air over the elevated plateau.

101

102 **2. Data and method**

103 In this study, we used the reanalysis meteorology data of years 2000-2009 from the
104 Research Data Archive at the US NCEP (National Center for Atmospheric Research),
105 Computational and Information Systems Laboratory
106 (<http://dx.doi.org/10.5065/D6M043C6>) for all atmospheric variable analyses and the
107 cloud cover fraction data derived from the Chinese meteorological satellite FY-2F for
108 convective cloud analyses. Following the studies of Yanai (1961), Yanai and Johnson
109 (1993), Yanai and Tomita (1998), the apparent heat source (Q_1) and apparent moisture
110 sink (Q_2) are calculated. Atmospheric heat sources and moisture sinks are respectively
111 gauged with the Q_1 and Q_2 . As Q_1 includes Q_2 and radiative heating, here we concentrate
112 only on the collective effect of apparent heating (Q_1) over the TP. The heat source column (in
113 units of $W m^{-2}$) over the TP is obtained with both horizontal and vertical integration of Q_1
114 over the TP-area of 78°E-103°E and 28°N-38°N covering the most region with the
115 altitude of higher than 3000m (see the large TP-rectangle in upper panel of Fig. 1) to

116 form a one-dimensional variable representing the TP-thermal forcing. The correlation
117 coefficients between the TP-heat source column and the meteorological variables
118 (divergence, U-, V- and W-components of wind and vapor transport flux) are calculated
119 to build their horizontal and vertical distributions of correlations. Zonal, meridional and
120 vertical components of the correlation vector are respectively derived through the
121 correlation coefficients of the TP-heat source column to U-, V- and W-components of
122 vector of wind and vapor transport flux, indicating the variations in wind and vapor
123 transport flux induced by the TP-thermal forcing.

124

125 **3. Results**

126 **3.1. Elevated heat and wet islands over the TP**

127 The upper panel in Figure 2, respectively, depicts the vertical distribution in zonal
128 differences of air temperature and specific humidity averaged along 93-94°E around and
129 over the TP, and these differences are calculated respectively by subtracting air
130 temperature and specific humidity in summer (June, July and August) averaged over
131 2000-2009 from their zonal means in the northern hemisphere. A “warm-wet island”
132 elevated in the middle troposphere over the TP is identified from the positive differences
133 of air temperature and humidity over the TP (upper panel of Fig. 2). On average, the
134 urban temperature is 1-3°C warmer than surrounding rural environments (Voogt and Oke,
135 2003;Zhao et al., 2014), while air temperatures over the TP is 4~6°C and even up to 6°C
136 higher than its surrounding atmosphere at the same altitude in summer (upper panel of
137 Fig. 2). This heat island over the massive TP exceeds that of any urban agglomerations in
138 the world in both intensity and area.

139

140 A high total solar irradiance of 1688Wm^{-2} , 23% higher than the solar constant was
141 observed over the TP (Lu et al., 1995), as the plateau absorbs a large proportion of solar
142 radiation. Because the TP is the region with strong solar radiation exceeding the solar
143 constant in the world, air temperatures over the TP could be $4\sim 6^{\circ}\text{C}$ and even up to 10°C
144 higher than its surrounding atmosphere at the same altitude in summer (Yeh and Chen,
145 1992). The high solar radiation on the TP could result in a strong sensible heat exchange
146 in the surface layer. Air temperature is a measure of the sensible heat content of the air.
147 A good positive correlation between surface air temperature and vertical velocity at
148 500hPa over the TP (lower panel of Fig. 2) reflects an important role of the surface
149 sensible heating and its vertical transfer in building the heat and wet islands over the TP.
150 The surface heating from the plateau could trigger the air ascent driving the vertical water
151 vapor transport up to the free troposphere. Even if the surface heat fluxes from the
152 plateau have a negligible impact on the South Asian summer monsoon circulation
153 strength (Boos and Kuang, 2010), they could greatly impact the convective precipitation
154 over the TP. As shown in the upper panel of Figure 2 for the vertical structures of the
155 elevated heat and wet islands, a heat source column reaching the upper troposphere over
156 the TP could be visualized from the distribution of positive temperature differences with
157 two high cores, respectively, within near-surface layers and between 200 and 400 hPa
158 (upper panel of Fig. 2). Due to a monotonic decrease in surface sensible heating with the
159 increasing elevation, the "hollow heat island" with a warm core at 200-400 hPa could be
160 dominated by the latent heating released from the convective cloud and precipitation

161 processes over the TP in association with the vertical structure of air vapor in the wet
162 island over the TP (upper panel of Fig. 2).
163
164 The elevated land surface with a strong radiative heating could make the massive TP
165 “mesa” more favorable for initiating a large number of convective cells. These
166 convective cells over the plateau often give rise to precipitation over the TP and its
167 surroundings in the boreal summer(Xu et al., 2012;Sugimoto and Ueno, 2010). In fact,
168 the annual occurrences of convective clouds (cumulonimbus) over the TP are observed
169 with 2.5 times of the regional mean over the other areas of China (Xu et al., 2002), and
170 the TP region is regarded as a high frequency center of cumulonimbus or mesoscale
171 convective systems (MCSs) in China(Sugimoto and Ueno, 2012), which is also
172 confirmed by the mean distribution of convective clouds over the TP (see Section 3.3) in
173 the plateau low vortex region (upper panel of Fig. 4).

174

175 **3.2. Processes of water vapor transport upward the TP**

176 Based on the differences of temperature and humidity at a given pressure level of the
177 atmosphere over the TP and over adjacent non-elevated areas in boreal summer, the
178 vertical structures of heat source column and wet island on the TP are characterized in
179 Figure 2 (upper panel) with the particularly surprising "hollow heat island" between 200
180 and 400 hPa in the shape of "warm core" and "mushroom cloud" (high zonal air
181 temperature deviation) over the TP. The vertical structure of the elevated wet island over
182 the TP can also confirm that the large TP topography prevents dry and cool extratropical
183 air from “ventilating” the moist and warm tropics and subtropics (upper panel of Fig. 2).

184 It is particularly interesting that the TP "hollow heat island" structure is similar to the
185 warm core of Typhoon-CISK process (Charney and Eliassen, 1964; Smith, 1997) in the
186 company of the elevated wet island (upper panel of Fig. 2) and the meridional circulation
187 with strong convections (left upper panel of Fig. 3). The "CISK-like process" relaying
188 warm-moist air up to the TP in two ladders is identified between two couples of
189 tropospheric lower convergences (LC) and upper divergences (UD) corresponding to 1)
190 the LC in the South Asian monsoon regions and the UD over the southern TP-slopes as
191 well as 2) the LC on the TP main platform and the UD in the middle and upper
192 troposphere over herein (left upper panel of Fig. 3).

193

194 The strength of "heat source column Q_1 " could be represented by the atmosphere column
195 integration of apparent heat source Q_1 over the TP-region. The middle panel of Figure 3
196 presents the correlation vectors of the TP heat source column strength Q_1 over the TP to
197 the W- and V-wind components at the vertical sections around the TP averaged in July
198 of 2000-2009. In this study, In this study, zonal, meridional and a vertical components of
199 the correlation vector are derived through the correlation coefficients of the Q_1 to U-, V-
200 and W- wind (or transport flux) components, respectively, where the arrow length
201 denotes the correlation combination with a longer arrow implying a better correlation,
202 and the arrow direction means the direction of anomalous wind (or transport flux)
203 induced by the TP-thermal effect. Therefore, the middle panel of Figure 3 indicates that
204 the air ascent motions induced by the TP heating are profound over the TP during the
205 summer monsoon period. The large topography of TP with the "hollow heat island" can
206 force a water vapor pump with the strong upward air flows. A meridional circulation

207 produced by the thermal effect of "hollow heat island" and the mechanical impact of the
208 TP-topography can not only result in the Asian summer monsoon circulations but also
209 enhance the water vapor transport from the oceans crossing the Asian monsoon areas up
210 to the TP (middle panel of Fig. 3). The strong divergences of the South Asian High in the
211 upper troposphere are collocated with the near-surface convergences associated with the
212 plateau low vortex, which is a favorable pattern for vertical circulation enforcing a strong
213 water vapor uplift over the TP (left upper and middle panels of Fig. 3; upper panel of Fig.
214 4). The TP surface sensible heat and the latent heat release from the convective cloud and
215 precipitation may maintain the vertical circulation driving the vapor transport up into the
216 atmospheric "water tower" over the TP (lower panel of Fig. 2; Figs. 3-5). A water vapor
217 pump with cloud convective activities is motivated in the near-surface air convergences
218 over the TP, driven by the plateau heating (upper panel of Fig. 4; Fig. 6). The
219 atmospheric "water tower" is set up by the air pump forced with the TP heating [Xu et al.,
220 2008].

221

222 A coupling of two "dynamic pumps" with the CISK-like mechanism, contiguous
223 horizontally but staggered vertically, are revealed with the cooperative interaction of the
224 "heat source column" and the elevated wet islands over the roof of the world (see two
225 dotted rectangles in middle panel of Fig.3). This interaction could be achieved with a
226 positive feedback, when the forcing effect of the "heat source column" drives the water
227 vapor flows climbing up the TP in the vertical motion, in turn, and the phase changes of
228 water vapor to clouds and precipitation in the moist convection release latent heating
229 intensifying the "heat source column" and especially the "warm core" in the upper

230 troposphere associated with the South Asian High(Sugimoto and Ueno, 2012). The “heat
231 source column” could enhance convergences at lower levels and divergences at upper
232 levels in the troposphere for pushing the moist air up the TP (middle panel of Fig. 3; Fig.
233 4). There could be a mutual feedback between the UD on the southern plateau slopes and
234 the LC on the TP-platform through the dynamical interaction of the horizontally
235 contiguous UD and LC (right upper and middle panels of Fig. 3). The UD over the
236 southern TP-slopes and the LC on the TP-platform could be contributed by the water
237 vapor flow acceleration at the inflection point between the steep southern slopes and the
238 southern edge of TP-platform with the mechanical TP-impact on the air pump on the
239 platform (upper and middle panels of Fig. 3).

240

241 The two ladders of “CISK-like process” over the South Asian summer monsoon region
242 and the TP knit a close tie of vapor transport from tropical oceans to the atmospheric
243 “water tower” over the TP (Fig. 3). The South Asian summer monsoon precipitation is
244 produced in the first ladder of air vapor transport toward the TP atmosphere, which could
245 be attributed to the TP-topographical block at the steep southern slopes with less thermal
246 impact(Boos and Kuang, 2010). The second ladder resulting in convective cloud
247 precipitation over the large TP platform with less terrain obstacles for water vapor flows
248 is dominantly controlled by thermal forcing of the "hollow heat island" in a large
249 scale(Wu et al., 2012). The pump of the "hollow heat island" over the TP could not only
250 attract air vapor transport from tropical oceans to the TP but also intensify the dynamic
251 lift of air vapor on the southern slope of the TP for Asian summer monsoon (middle and
252 lower panels of Fig. 3). The dynamic structures of two couples of tropospheric LC and

253 UD with their interaction build up a meridional circulation in a two-ladder pump of moist
254 air along the plateau (left upper and middle panels of Fig. 3), which could also be
255 explained with the vertical distribution of apparent heat source Q_1 and apparent moisture
256 sink Q_2 around the TP (Fig. 5). In Figure 5 two couples of high Q_1 and Q_2 areas are
257 found between two couples of tropospheric LC and UD respectively on two ladders in the
258 process of water vapor transport up to the TP atmosphere (Fig. 3).

259

260 The convective clouds and precipitation of the plateau low vortex or cyclone are
261 triggered by the plateau heating. The CISK-like process is found to play an important role
262 in the local low vortex development for the TP-precipitation (Qiao and Zhang, 1994).

263 The good correlations of the strength of “heat source column” Q_1 to the total water vapor
264 and to the net transport flux divergence over the TP (two lower panels of Fig. 4) further
265 interpret a large scale effect of “CISK-like mechanism” with a positive feedback among
266 the heat source column, the vertical convection and the water vapor supply for the
267 atmospheric “water tower” over the TP. The two ladder “CISK-like mechanism” is a key
268 process attracting water vapor toward the TP for building the TP’s “water tower” in
269 Asian water cycle. To further discover the process initiating the upward transport of
270 water vapor flows over the TP, the lag correlations of the TP’s heat source column Q_1 at
271 10 prior days to the divergences and the meridional circulation are analyzed in the lower
272 panel of Figure 3, which reflect that the plateau heating could initiate and trigger the
273 vertical circulations for the “hollow heat island” process with a leading effect of the heat
274 source column on water vapor transport toward the TP.

275

276 3.3 Cloud distribution over the TP

277 The TP region is identified as a frequent occurrence center of MCSs in China (Sugimoto
278 and Ueno, 2012). In association with Asian summer monsoons, the summertime
279 convective clouds bring the precipitation over the TP and its surroundings (Xu et al.,
280 2012; Sugimoto and Ueno, 2010). To further clarify the atmospheric “water tower” over
281 the TP in Asian water cycle, Figure 6 presents the spatial distribution of total cloud cover
282 over the TP and its surrounding area averaged in July 2008.

283

284 During the Asian summer monsoon period, the dense cloud covers existed over the
285 regions from the Bay of Bengal, South Asian monsoon region to the southern TP (Fig. 6).
286 As characterized with the correction vectors of the column heat source over the TP to the
287 moisture transport over and around the TP (middle panel of Fig. 3), two convergence
288 zones of moisture transport fluxes ($\nabla \cdot qV < 0$) are found on two ladders over the
289 plateau’s southern slopes and main platform during the moisture transport from the
290 oceans up to the TP, resulting in these regions of dense cloud covers shown in Figure 6,
291 It is noteworthy that the high cloud amounts are zonally concentrated between the steep
292 southern plateau slopes and the shear line of the plateau low vortex over the TP (upper
293 panel of Fig. 4; Fig. 6) with the monthly mean cloud cover fractions up to 90%, which
294 could resulted from the “CISK-like mechanism” for building the TP’s atmospheric
295 “water tower” (Fig. 3). Over the large TP platform with relatively plain terrain, the
296 monthly mean cloud covers of around 45% are mostly observed on the central-eastern
297 region with the less cloud covers over the northwestern TP, depending on the moisture
298 transport across the TP. The plateau low vortex over the TP and the southward air flows

299 with less moisture on the north of the shear line could lead to the less cloud covers in the
300 northwestern platform of TP (upper panel of Fig. 4).

301 .

302 The observed cloud distribution over the TP confirms that the “CISK-like mechanism”
303 is an important mechanism sustaining the atmospheric “water tower” over the TP.

304 Connecting with the cloud and precipitation in the atmospheric “water tower”, the
305 plausible hydrological cycles could be realized between tropical oceans and the TP.

306

307 **4. Conclusions and discussions**

308 The present analyses clearly indicate that the TP presents itself as a “warm-wet island”.

309 The surface heating over the plateau leads to a low-pressure center causing flow
310 convergence at low levels of the plateau and triggers vertical motion subsequently. This
311 convective system will result in plateau clouds and precipitation, which would explain
312 abundant water storage in the atmosphere over the TP and its surrounding regions.

313

314 The classic Asian summer monsoon theory elucidated an “air pump” mechanism in
315 relation to the TP. The warm-moist air from the low-latitude oceans is drawn toward the
316 plateau by this air pump. Our analysis on relationship of the “heat source column” over
317 the TP and warm-moist air transport in the present study further reveals a CISK-like
318 mechanism on water vapor suction up the plateau. An appreciable portion of warm-moist
319 air converges at the foot of the south rim of the plateau. The convergence of the warm-
320 moist air ascends along the plateau’s slope and diverges at about the altitude of the
321 plateau top. This divergence flow enforces the convergence at the heated low-pressure

322 center over the TP and feeds in the convective system with warm-moist air, which results
323 in the clouds and precipitations for the atmospheric water tower over the TP.
324
325 These dynamic and thermodynamic processes depict a coupling of two CISK
326 type systems, both with convergence at low levels and divergence at upper levels, but the
327 systems are horizontally contiguous as well as vertically staggered. The two systems
328 display a mutually supportive mechanism with the mechanical and thermal TP-impact
329 between the southern slopes and the platform of the TP in the interaction region marked
330 in Figure 7. It is this coupling that ladders the moist air up to the plateau building the
331 atmospheric “water tower” over the TP.

332

333 In this study, the mean climate of air vapor transport to the TP is investigated based on
334 the summertime averages over the past years, and two ladder “CISK-like mechanism” is
335 identified as a key process sustaining the atmospheric “water tower” over the TP. The role
336 of intraseasonal variability, synoptic-scale system activities and diurnal variation in the
337 atmospheric heat source and moisture over the TP (Sugimoto et al., 2008; Fujinami and
338 Yasunari, 2004) will be considered in future study on the warm-moist air transport up to
339 the plateau. It should be emphasized that considering the quality of reanalysis data over
340 and around the TP, a comparison between NCEP/NCAR and some other reanalysis
341 datasets such as JRA-25, ERA-Interim, or MERRA is necessary in further work.

342 Furthermore, the two CISK type system revealed from this observational analysis need to
343 be further studied with numerical models to understand the mechanism to work.

344

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452 Figure Captions

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454 Figure 1. Geographical distribution of water sources in glaciers (snowpacks), rivers
455 and lakes over China with white, green and light blue colors respectively. Two major
456 lake groups are marked by two red rectangles in the TP and Eastern China (upper
457 panel). Column vapor content (10^{-2}g cm^{-2}) over 500hPa in summer averaged over
458 2000-2009 (lower panel).

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460 Figure 2. Vertical sections of the temperature ($^{\circ}\text{C}$; filled contours) and specific
461 humidity (g/kg; contour lines) differences relative to the zonal means along 93-94 $^{\circ}\text{E}$
462 in summer averaged over 2000-2009. The plateau section is marked with soil color
463 (upper panel). A scatter plot of surface air temperature and vertical velocity at
464 500hPa in the TP region in July of 2000-2009 (lower panel).

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467 Figure 3. Vertical sections of wind vectors and divergences (filled contours) for
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470 Vertical sections of the correlations of the daily TP heat source column Q_1 to the
471 divergences (filled contours) and the correction vectors of daily Q_1 to V- and W-
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473 circulations and the uplifting vapor transport denoted by blue dash lines and black
474 arrows, respectively (middle panel). Vertical sections of the lag-correlations of TP
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482 Figure 4. Correlation vectors of the TP heat source column strength Q_1 to the
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486 The TP region with the altitude of higher than 3000m is shaded in yellow contour
487 (upper panel). Correlations of the heat source strength Q_1 , total water vapor q and

488 net vapor transport flux divergence (div) in the TP air column in summer of 2000-
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491 Figure 5. The vertical distributions of apparent heat source Q_1 (filled contours) and
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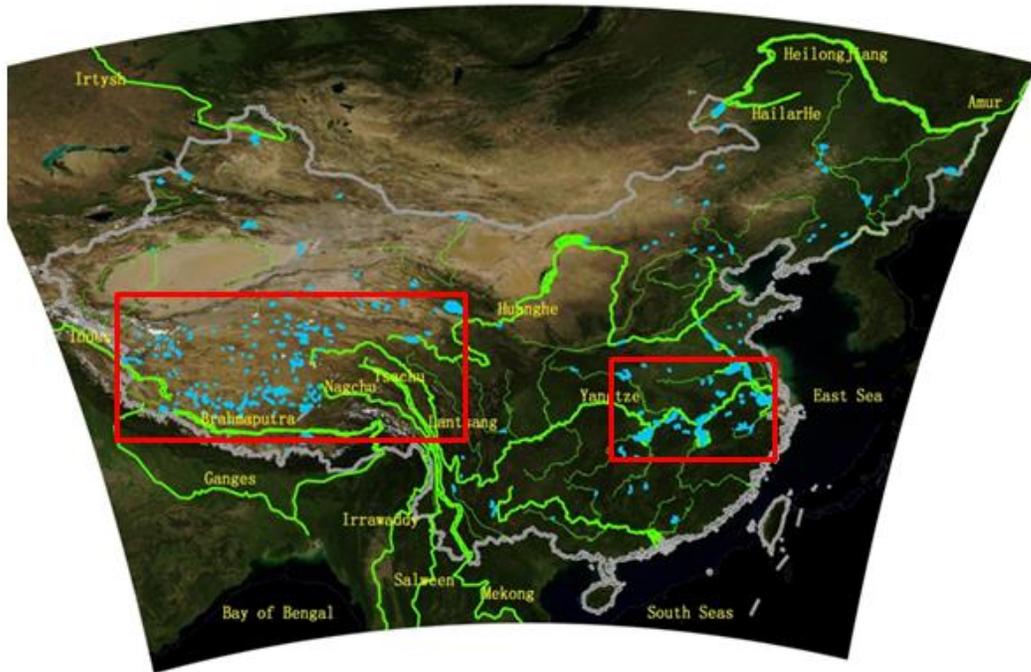
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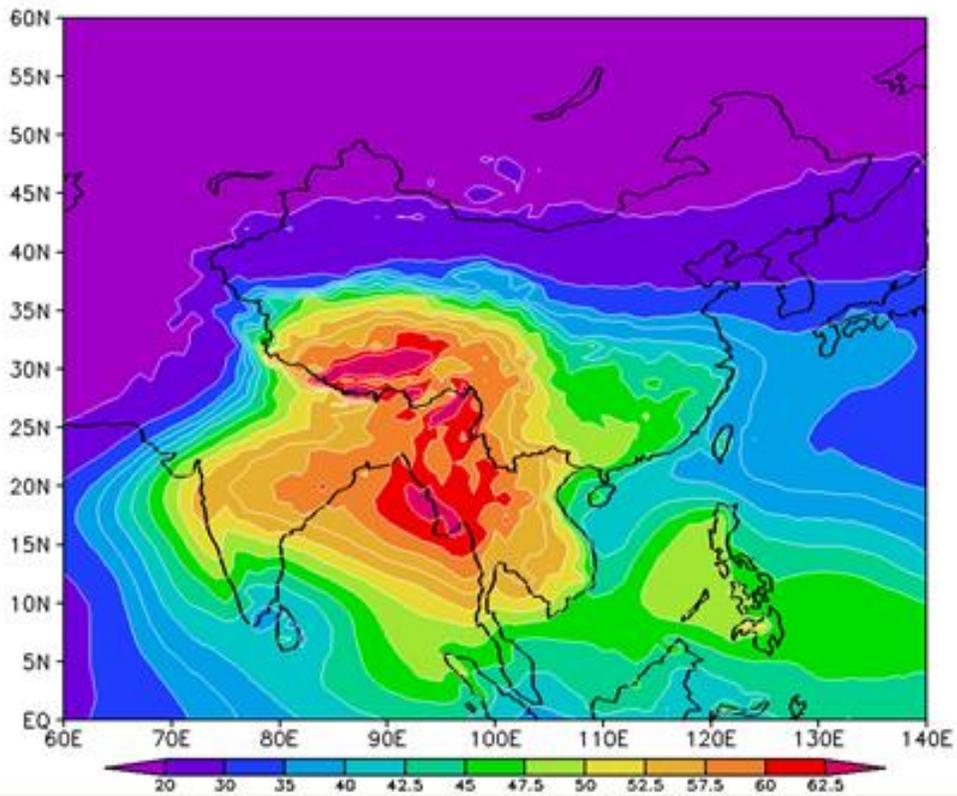
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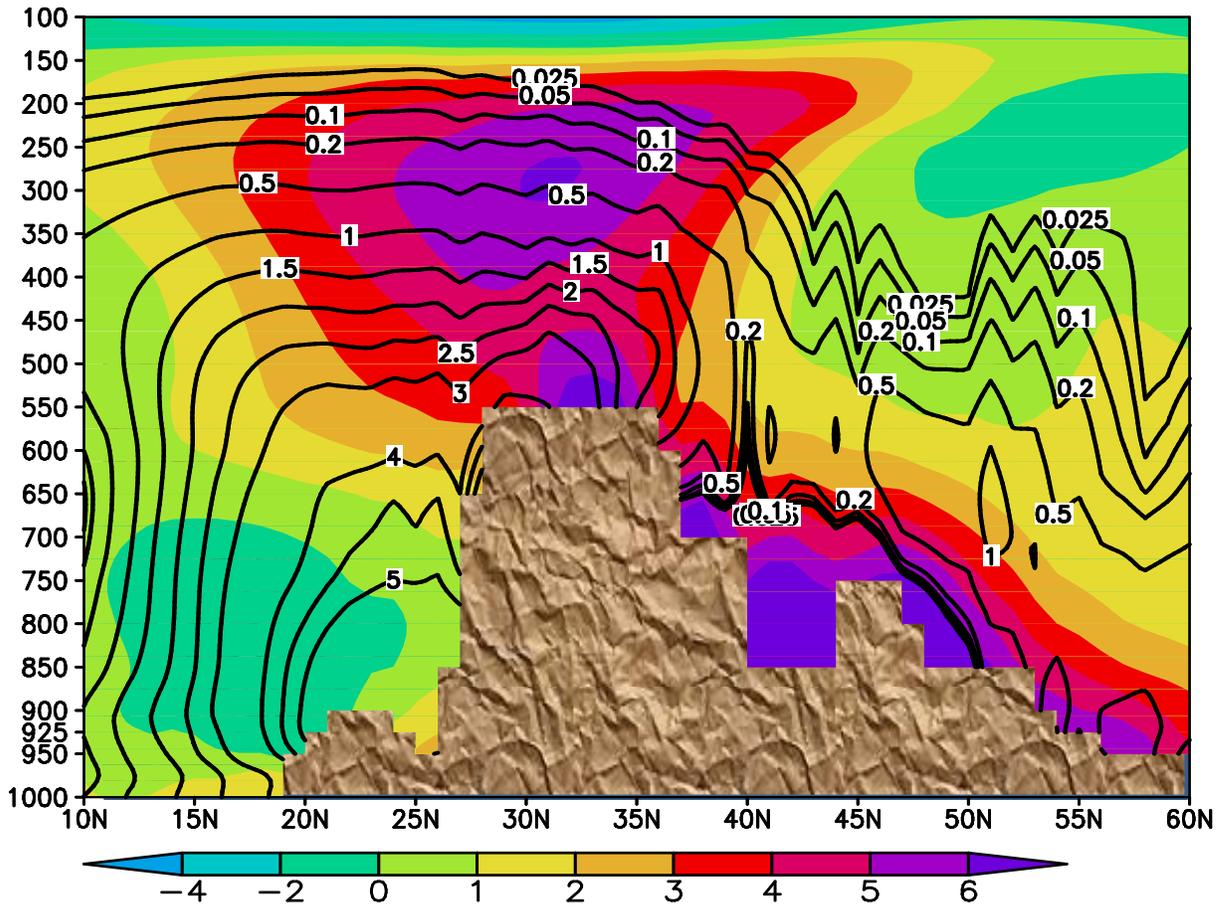


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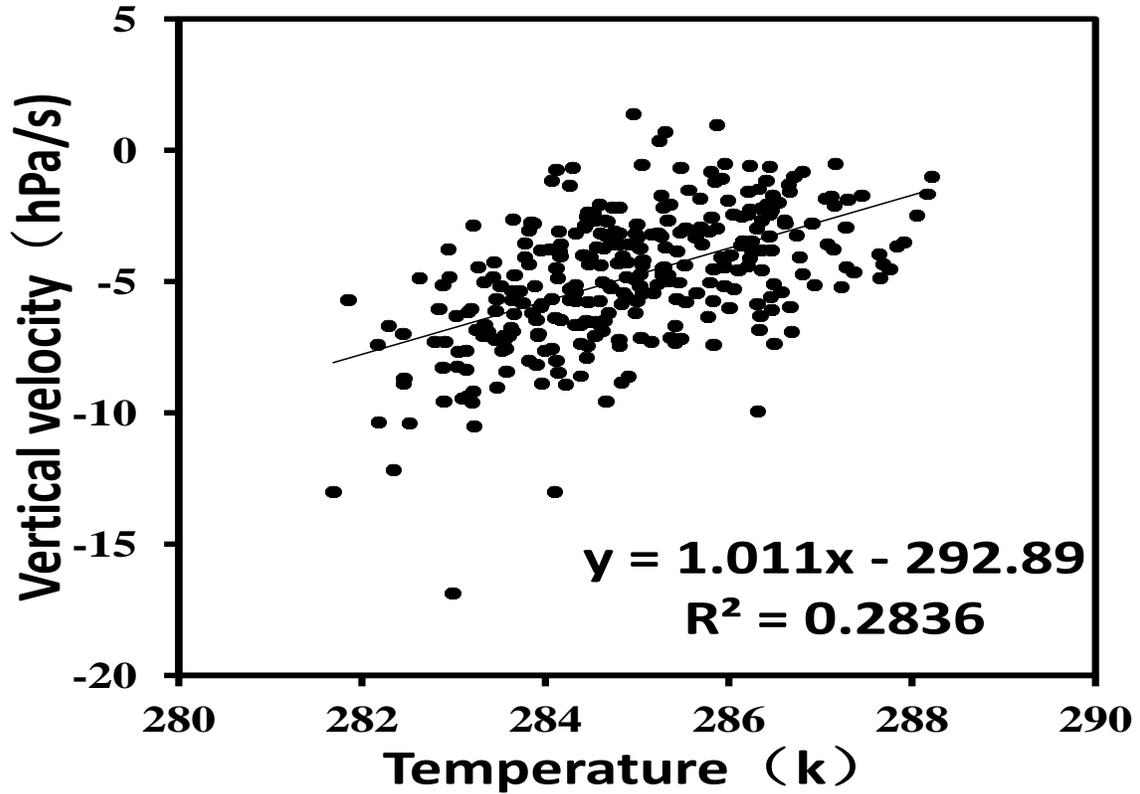


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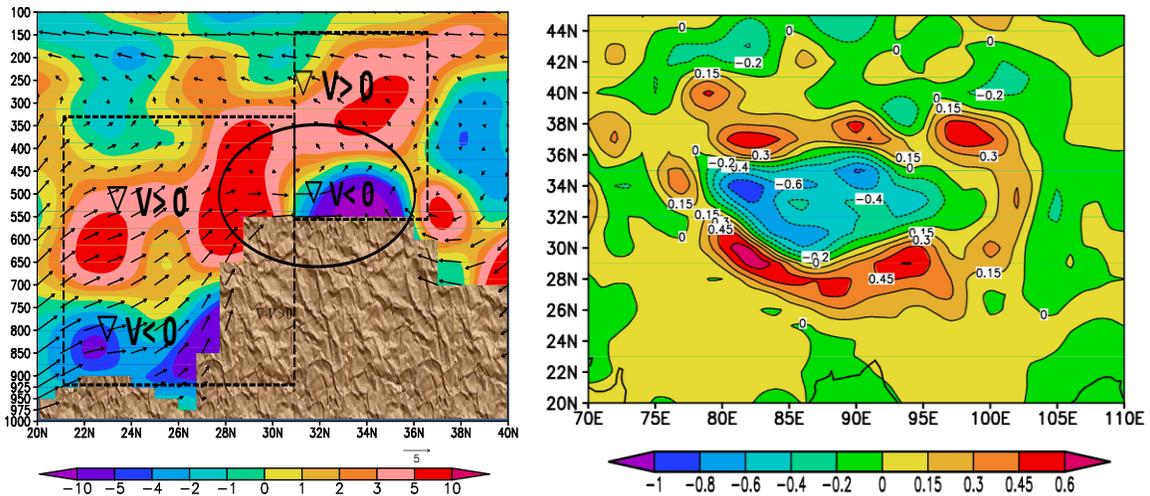


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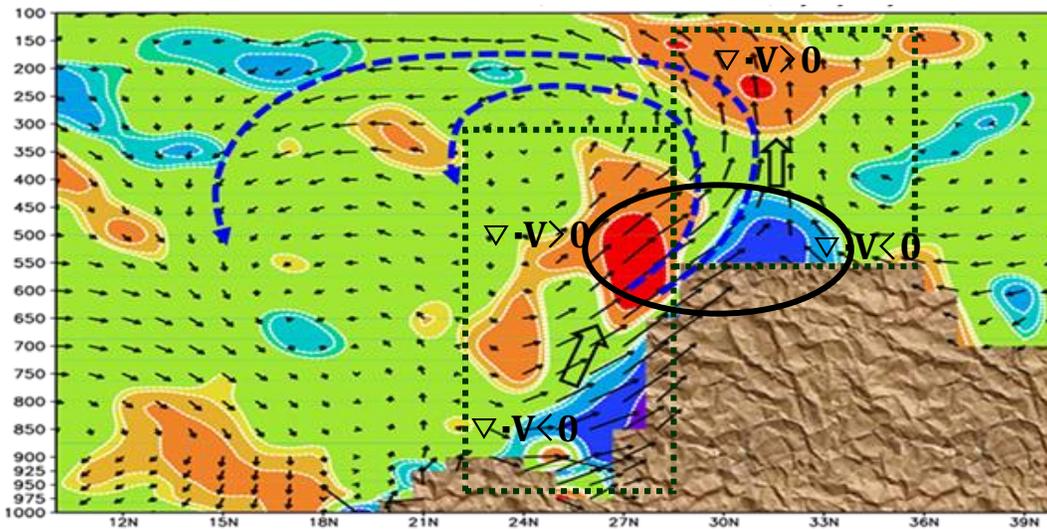
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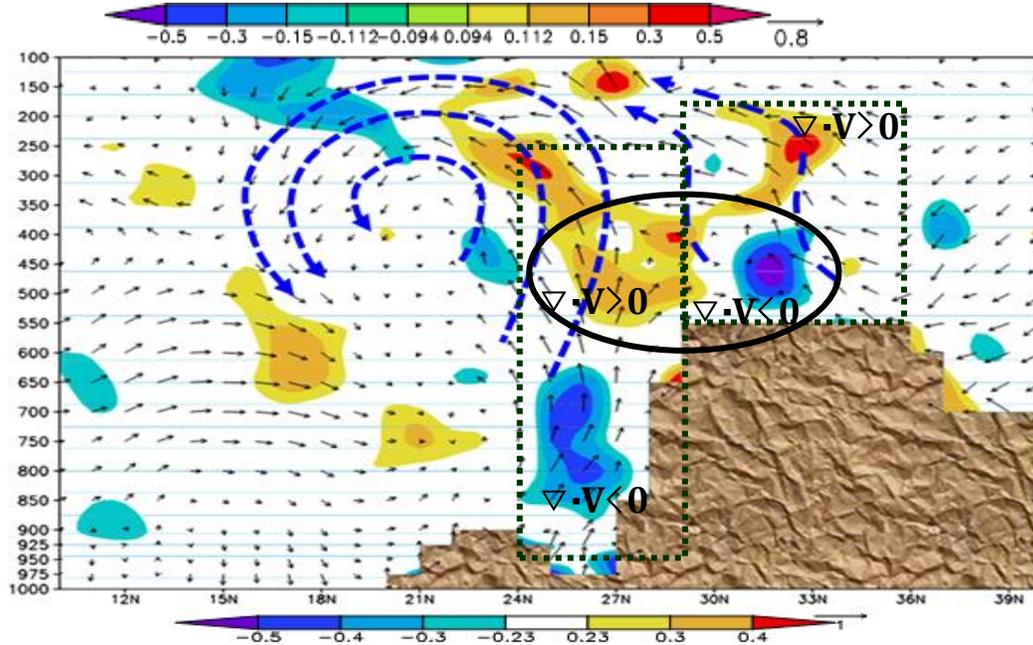
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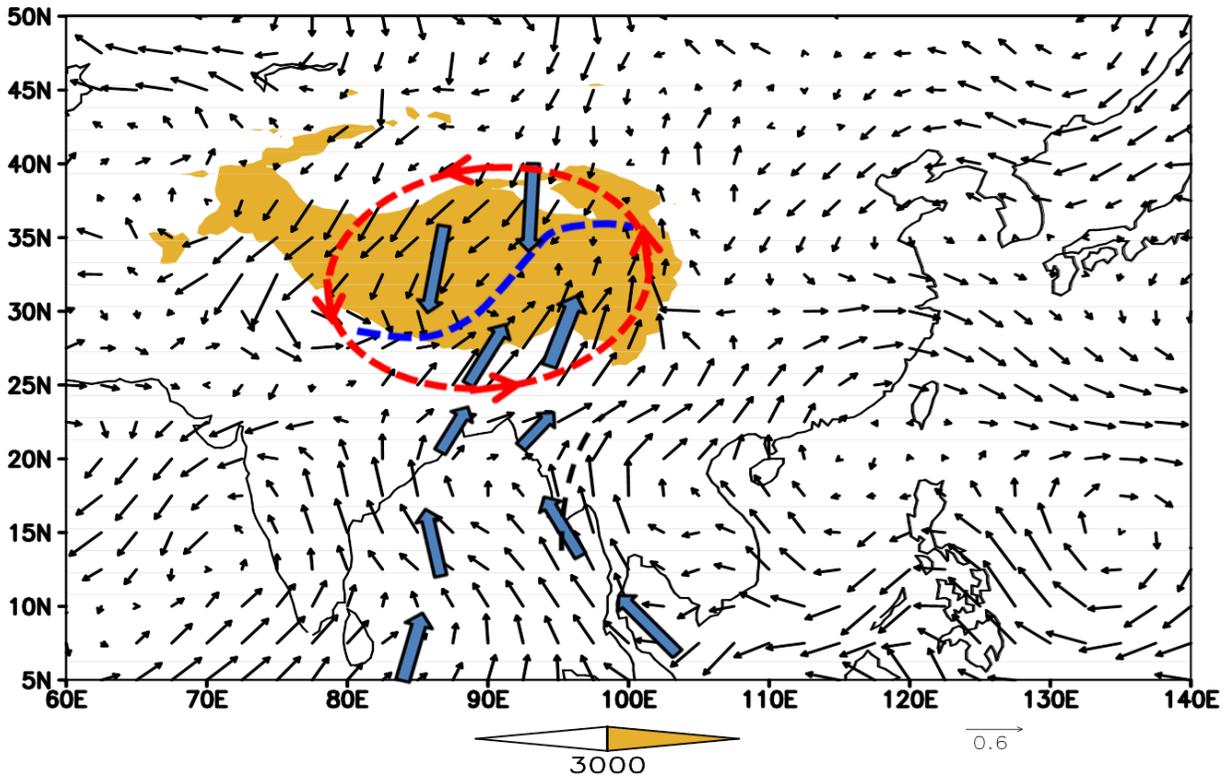
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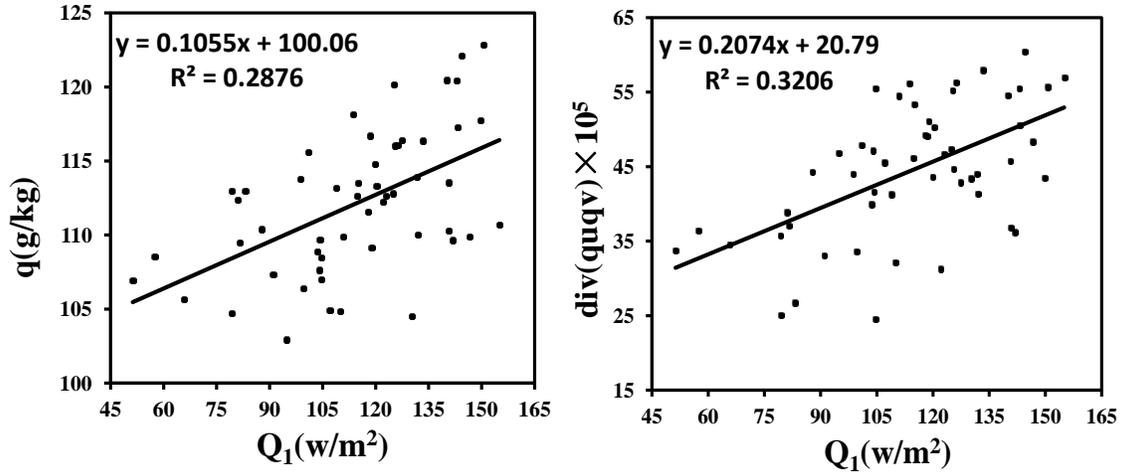
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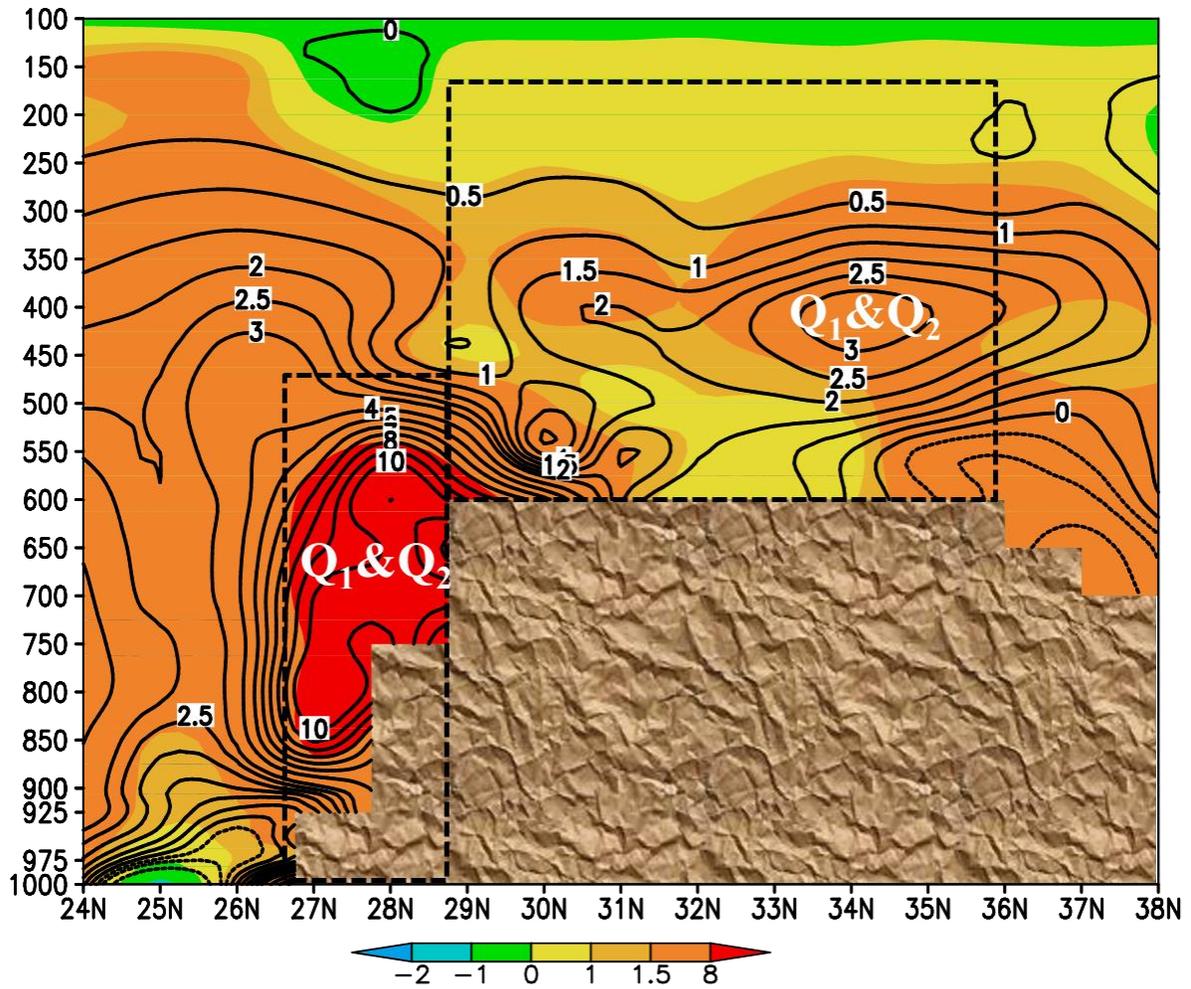
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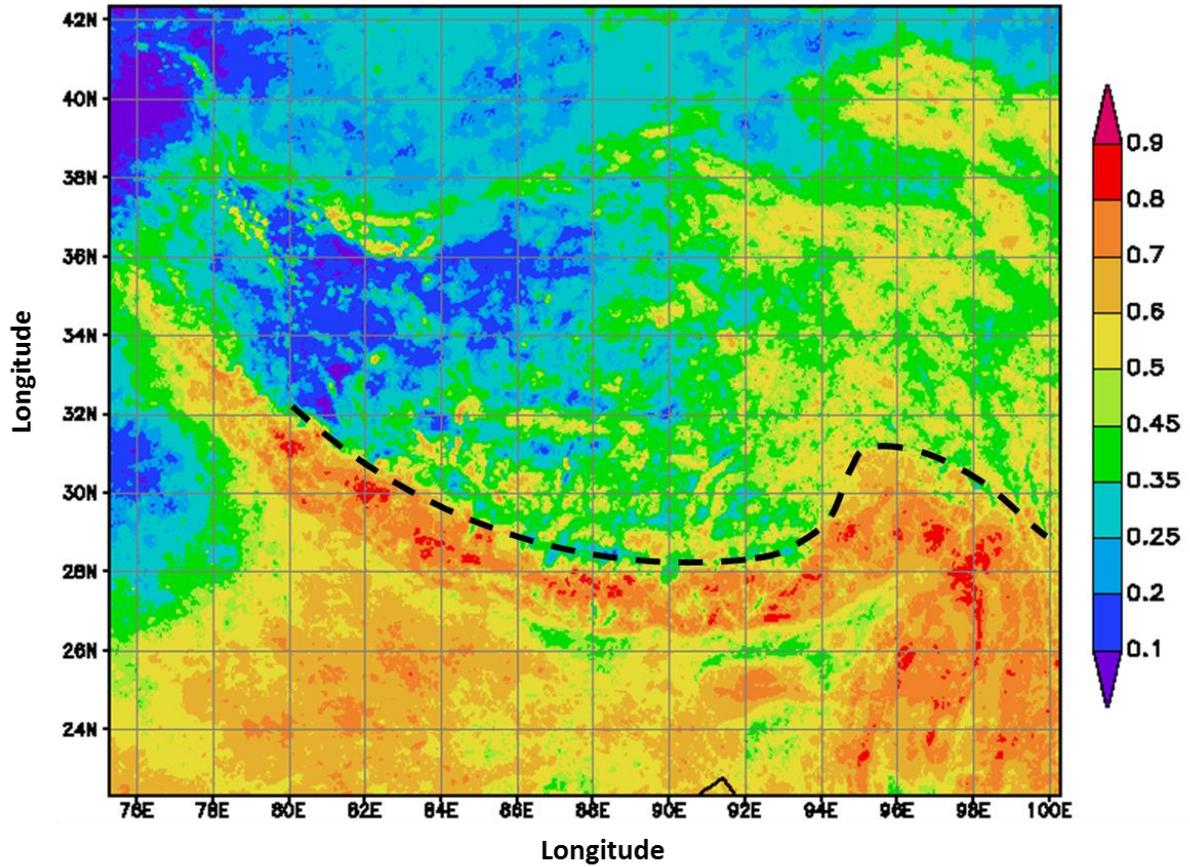


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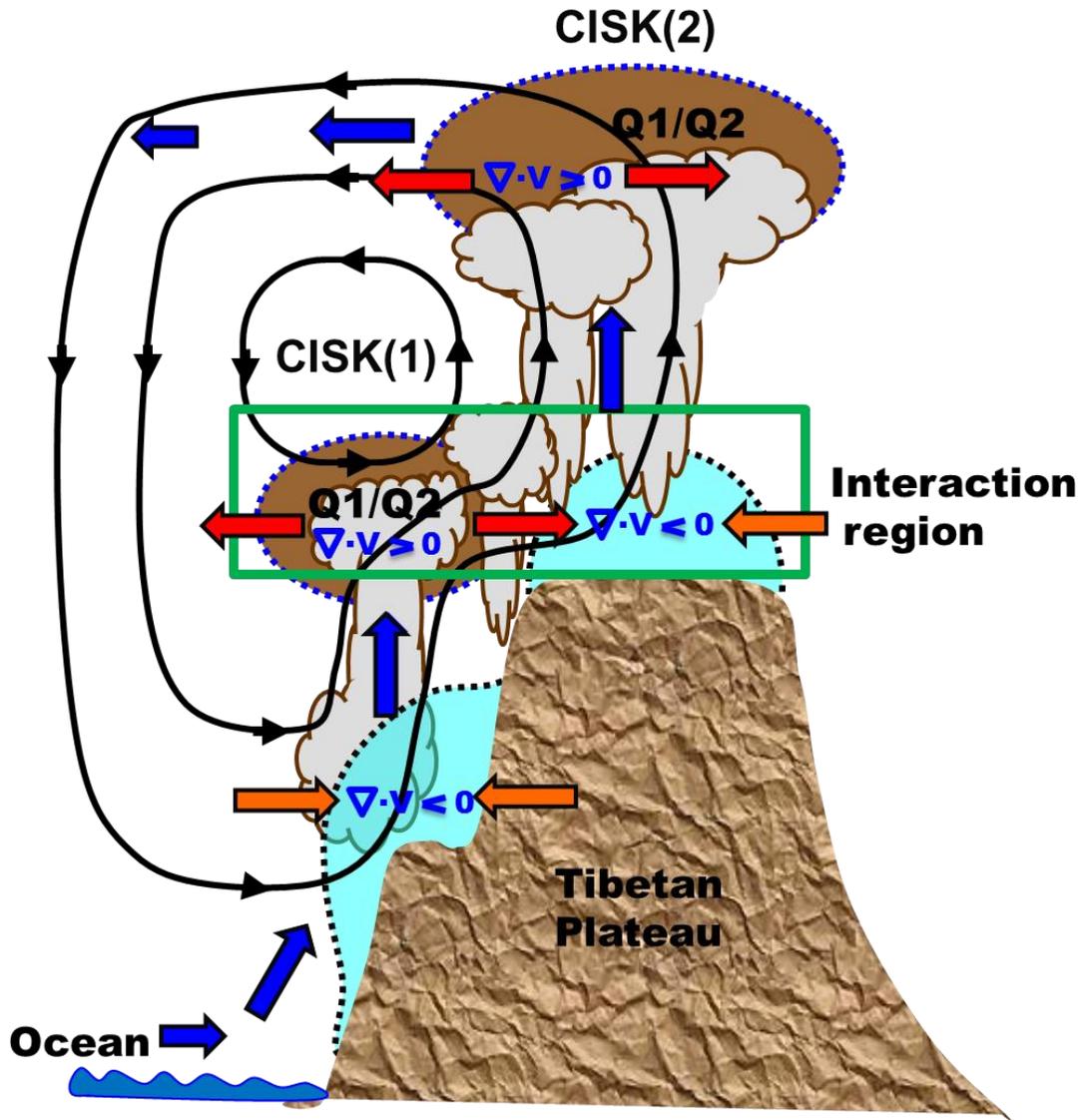
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