

Exploring the Severe Winter Haze in Beijing: Impact of Synoptic Weather, Regional Transport and Heterogeneous Reactions

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26 **Abstract**

27 Extreme haze episodes repeatedly shrouded Beijing during the winter of 2012–2013, causing
28 major environmental and health problems. To better understand these extreme events, we
29 performed a model-assisted analysis of the hourly observation data of PM_{2.5} and its major
30 chemical compositions. The synthetic analysis shows that, (1) the severe winter haze was driven
31 by stable synoptic meteorological conditions over northeastern China, rather than by an abrupt
32 increase in anthropogenic emissions. (2) Secondary species, including organics, sulfate, nitrate,
33 and ammonium, were the major constituents of PM_{2.5} during this period. (3) Due to the dimming
34 effect of high loading of aerosol particles, gaseous oxidant concentrations decreased
35 significantly, suggesting a reduced production of secondary aerosols through gas phase
36 reactions. Surprisingly, the observational data reveals an enhanced production rate of secondary
37 aerosols, suggesting an important contribution from other formation pathways, most likely
38 heterogeneous reactions. These reactions appeared to be more efficient in producing secondary
39 inorganics aerosols than organic aerosols resulting in a strongly elevated fraction of inorganics
40 during heavily polluted periods. (4) Moreover, we found that high aerosol concentration was a
41 regional phenomenon. The accumulation process of aerosol particles occurred successively
42 from southeast cities to Beijing. The apparent sharp increase in PM_{2.5} concentration of up to
43 several hundred $\mu\text{g m}^{-3}$ per hour recorded in Beijing represented rapid recovery from an
44 interruption to the continuous pollution accumulation over the region, rather than purely local
45 chemical production. This suggests that regional transport of pollutants played an important
46 role during these severe pollution events.

47

48 1 Introduction

49 Severe haze episodes in the winter of 2012–2013 engulfed Beijing, as well as other cities in
50 southeastern China, causing one of the worst atmospheric pollution events in history. With
51 hourly fine particle (PM_{2.5}) concentrations up to ~900 µg/m³, outdoor exposure caused adverse
52 health effects (Nel, 2005; Pöschl, 2005; Peplow, 2014), including severe respiratory system
53 related symptoms and deceases (Cao et al., 2014; Ouyang, 2013). Meanwhile the visibility was
54 reduced down to 100 m, which disrupted traffic with canceled flights and closed highways. The
55 government had to adopt emergency response measures to deal with these pollution episodes
56 (<http://english.sina.com/china/p/2013/0113/548263.html>). In addition to massive amounts of
57 primary particulate matter, high emissions in China provided plenty of gas pollutants to serve
58 as precursors for secondary aerosols (Zhang et al., 2009). Densely distributed mega-cities (i.e.,
59 city clusters) have worsened this situation, contributing to regional air pollution. Once the
60 regional pollution is formed, the advection becomes less effective in scavenging local pollutants
61 (no clean air from upwind). Thus, the regional pollution is more persistent compared with air
62 pollution within a specific city. Moreover, cities within this region could not eliminate their
63 pollution solely by reducing local emissions (Chan and Yao, 2008; Cheng et al. 2008a).

64 These extreme haze episodes attracted great scientific interest. The visibility impairment has
65 been attributed to scattering and absorption of solar radiation by aerosol particles (mostly PM_{2.5})
66 and their hygroscopic growth under high relative humidity (Cheng et al., 2006; Cheng et al.,
67 2008b,c). Regional transport of pollutants was found to contribute considerably to
68 concentrations of PM_{2.5} (Z. Wang et al., 2014; L. T. Wang et al., 2014), dust (Yang et al., 2013;
69 Y. Wang et al., 2014), and SO₂ (Yang et al., 2013) in Beijing. Atmospheric dynamic processes
70 during hazy conditions were different from clean conditions, with a significant two-way
71 feedback between PM_{2.5} and boundary layer evolution (Z. Wang et al., 2014). Secondary
72 inorganic aerosol species were suggested to be the major contributor to severe haze, based on
73 off-line PM_{2.5} analysis (Quan et al., 2014), and on-line non-refractory PM₁ analysis by an
74 Aerosol Chemical Speciation Monitor (Sun et al., 2014). In addition, some studies described
75 unusual atmospheric phenomena taking place under heavily polluted conditions, such as
76 extremely low ozone concentration (less than 5 ppb) in the absence of diurnal variation (Zhao
77 et al., 2013) and the synergistic oxidation of SO₂ and NO₂ (He et al., 2014). These findings
78 suggest need for a better understanding on the haze formation mechanisms.

79 In this study, we address the following questions for the winter haze episodes aforementioned:
80 (1) the relative importance of enhanced emission versus meteorology; (2) the cause of the sharp
81 PM_{2.5} increase during the haze episodes in Beijing, whether it was mainly driven by an
82 extremely rapid local chemical production or by regional transport; and (3) the dominant
83 chemical mechanisms of haze formation.

84

85 **2 Experimental Methods**

86 On-line ambient observation was conducted from 1–31 Jan. 2013 on the campus of Tsinghua
87 University. The observation site is situated on the rooftop of the Environmental Science
88 Building (40°00' 17" N, 116°19' 34" E), approximately 10 m above ground. Tsinghua
89 University is located in the northwest part of urban Beijing, close to the North 4th Ring Road,
90 without any major pollution sources nearby. All observation data are hourly unified data.

91 Mass concentrations of fine (PM_{2.5}) and coarse (PM_{2.5-10}) particles were simultaneously
92 measured based on the β-ray absorption method by a PM-712 Monitor (Kimoto Electric Co.,
93 Ltd., Japan), which was equipped with a US-EPA PM₁₀ inlet and a PM_{2.5} virtual impactor
94 (Kimoto Electric Co., Ltd., 2012; Kaneyasu et al., 2014). Dehumidification was achieved with
95 the hygroscopic growth correction formula:

$$96 \text{ Dehumidified PM}_{2.5} \text{ mass conc.} = \text{Measured PM}_{2.5} \text{ mass conc.} \times \frac{1}{1 + 0.010 \times e^{\frac{6.000 \text{ RH}}{100}}} \quad (1)$$

97 where the 0.010 and 6.000 are localized coefficients, and RH is relative humidity in %. All
98 PM_{2.5} hereinafter refer to the dehumidified PM_{2.5} data.

99 A Sunset Model 4 semi-continuous carbon analyzer (Beaverton, OR, USA) was used to
100 measure hourly organic carbon (OC) and elemental carbon (EC) concentrations in PM_{2.5}. A
101 NIOSH (National Institute for Occupational Safety and Health) temperature protocol was used
102 and the calculation discrepancy under high ambient concentrations was corrected accordingly
103 (G.J. Zheng et al., 2014). Organic matter (OM) was estimated as 1.6*OC, based on previous
104 results (Zhang et al., 2014; Xing et al., 2013). The use of fixed OM/OC ratio requires caveats
105 because the ratio might change due to the variable oxidation degree of OM under different
106 conditions.

107 Hourly sulfate and nitrate concentrations in PM_{2.5} were measured using an ACSA-08 Monitor
108 (Kimoto Electric Co., Ltd., Japan). The ACSA-08 Monitor measured nitrates using a ultra-
109 violet spectrophotometric method, and quantified sulfates with the BaSO₄-based turbidimetric
110 method after addition of BaCl₂ dissolved in polyvinyl pyrrolidone solution (Kimoto et al., 2013).
111 Ammonium was predicted under the assumption that it existed as NH₄NO₃ and (NH₄)₂SO₄ (He
112 et al., 2012), which might be an overestimation based on the non-refractory PM1 results (Sun
113 et al., 2014). Thus the predicted ammonium given here should be regarded as an upper limit.

114 An automatic meteorological observation instrument (Milos520, VAISALA Inc., Finland) was
115 used to obtain meteorological parameters, including atmospheric pressure, temperature, RH,
116 wind speed, and wind direction. Specific humidity was calculated from these measured
117 parameters (<http://www.srh.noaa.gov/epz/?n=wxcalc>).

118 SO₂ and NO₂ concentrations in Beijing, and PM_{2.5} concentrations in other cities were acquired
119 from the Atmospheric Environment Monitoring Network (Tang et al., 2012). Daily averaged
120 solar radiation reaching ground data were downloaded from the China Meteorological Data
121 Sharing Service System (<http://cdc.cma.gov.cn>). Planetary boundary layer (PBL) height was
122 simulated with the Weather Research & Forecasting (WRF) Model (B. Zheng et al., 2014).

123

124 **3 General Characteristics of Beijing Winter Haze**

125 Primary atmospheric pollutant in Beijing during the winter of 2012–2013 was PM_{2.5}, which
126 constituted about 70% of PM₁₀. This ratio increased when PM_{2.5} pollution became worse (Fig.
127 1b). Monthly average PM_{2.5} concentration reached 121.0 µg/m³ in Jan. 2013, and hourly PM_{2.5}
128 concentrations peaked at 855.10 µg/m³, which was the highest ever reported in Beijing (Zhao
129 et al., 2009; Zhao et al., 2011; Zhao et al., 2013; Zhang et al., 2014). The severe PM_{2.5} pollution
130 lasted nearly the whole month, characterized by frequent and long-lasting pollution episodes.
131 Here, we define an episode as a set of continuous days with daily PM_{2.5} averages exceeding 75
132 µg/m³. In total, four episodes were identified in Jan. 2013 (Fig. 1a): 4–8 Jan. (Episode I), 10–
133 16 Jan. (Episode II), 18–23 Jan. (Episode III), and 25–31 Jan. (Episode IV). Maximum episode-
134 averaged PM_{2.5} concentrations reached 245.4 µg/m³ in Episode II (see Table 1 for comparative
135 information on Episodes I to III; Episode IV was not included because of missing data). In
136 addition to the high average concentrations, these episodes were frequent (intervals between

137 episodes were all ~1 day) and long-lasting (5–7 days) compared with typical durations (5 days)
138 and frequencies (1–3 days) of previous Beijing winter haze episodes (Jia et al., 2008).

139 Another unique feature of the PM_{2.5} mass concentrations during this winter haze period was
140 their dramatic hourly fluctuation. The maximum daily variation was 778.6 µg/m³ on 12 Jan.
141 Hourly PM_{2.5} changes of over 100 µg/m³ (increases or decreases) were observed over 40 times
142 during this haze period. Hourly increases or decreases could reach up to 351.8 µg/m³ and -217.7
143 µg/m³, respectively. Causes of these sharp transitions are discussed in Section 5.

144 The variation of chemical composition with PM_{2.5} pollution level, and among episodes, was
145 also explored. We classified PM_{2.5} pollution into 4 categories according to the Air Quality Index
146 (http://kjs.mep.gov.cn/hjbhzbz/bzwb/dqhjbh/jcgfffbz/201203/t20120302_224166.htm?COLLCC=2906016564&) (Fig. 1b): clean (PM_{2.5} ≤ 35 µg/m³), slightly polluted (35 < PM_{2.5} ≤
148 115 µg/m³), polluted (115 < PM_{2.5} ≤ 350 µg/m³), and heavily polluted (PM_{2.5} > 350 µg/m³), where
149 PM_{2.5} refers to the hourly concentration. Under this classification, the slightly polluted, polluted,
150 and heavily polluted levels generally correspond to small, moderate, and large PM_{2.5} peaks in
151 Fig. 1b. **Mean percentile compositions of major components in PM_{2.5} under different pollution
152 levels were shown in Fig. 2a.** With increasing pollution level, the EC fraction decreased slightly,
153 OC fraction decreased significantly, while sulfate and nitrate contributions increased sharply
154 (Fig. 2a). It suggests that secondary inorganic aerosol species become more important during
155 polluted periods concerning their contribution to the PM_{2.5}. A similar trend was observed for
156 NR-PM1 (Sun et al., 2014) and off-line samples (Cheng et al., submitted to Atmos. Environ.).
157 On average, OC, EC, nitrate, and sulfate comprised 21%, 3%, 19% and 22% of PM_{2.5} (Fig. 2b).
158 Good correlations with PM_{2.5} were observed for OC, EC and nitrate (R² > 0.8 for these three
159 species) for all data in Jan. 2013, while for sulfate the correlation became weaker, reflecting
160 larger episodic variations (Fig. 2b). In Episode III, NO₂ exceeded SO₂ by 50% (Table 1),
161 generally in accordance with previous studies (Meng et al., 2009). In contrast, concentration of
162 SO₂ exceeded NO₂ in Episodes I and II. Compared with Episode II, Episode I was much drier,
163 which is unfavorable to the sulfate formation. The relatively high SO₂ but low NO₂
164 concentrations in Episodes I and II may indicate the significance of stationary sources (coal
165 combustion, etc.) in local emissions or regional SO₂-rich air masses transported to Beijing.

166

167 4 Emission Enhancement vs. Synoptic Conditions

168 Haze episodes were much more severe and frequent in winter 2013 than in 2012. One possible
169 explanation is that there was an abrupt emission enhancement during 2013. However, we didn't
170 find such change in the emission inventory (<http://www.meicmodel.org/>). Annual average
171 emissions of primary PM_{2.5}, SO₂ and NO_x show slight differences between 2013 and 2012
172 (1.2%, -1.3% and 0.8%, respectively) for the Beijing-Tianjin-Hebei region. The changes of
173 monthly averaged emissions in Jan. were higher than the annual average changes in rates, i.e.,
174 2.1%, 1.5% and 2.5% for primary PM_{2.5}, SO₂, and NO_x, respectively; but they are still not
175 significant compared to the changes in pollutant concentrations. Thus, we suspect that these
176 haze episodes arose from the unfavorable synoptic conditions in Jan. 2013.

177 The relative importance of enhanced emission versus unfavorable meteorology in PM_{2.5}
178 concentration of Jan. 2013 was estimated by model simulations with three scenarios (Fig. 3).
179 Base scenario *a* was designed to simulate the actual situation, i.e., with both input emission
180 inventory and meteorology for Jan. 2013. In scenarios *b* and *c*, Jan. 2012 meteorology and Jan.
181 2012 emission inventory data were used, respectively. Since the original WRF-CMAQ
182 (Weather Research and Forecasting - Community Multiscale Air Quality) modeling system
183 cannot reproduce the observed concentrations under heavily polluted conditions (B. Zheng et
184 al. 2014), a revised WRF-CMAQ system with enhanced heterogeneous reactions (Wang et al.,
185 2012) was adopted to improve the model performance. The revised model could effectively
186 capture the measured concentrations of total PM_{2.5} (with normalized mean biases (NMB) being
187 0.4 %) and its different chemical compositions for both clean and heavily polluted haze days
188 (B. Zheng et al., 2014). Details of the model configuration, modifications, and validation are
189 described in B. Zheng et al. (2014).

190 As expected, the influence of emission difference was negligible (Fig. 3a and 3c). For the whole
191 simulation domain of the North China Plain (NCP), both simulation with Jan. 2012 meteorology
192 (Scenario *c*) and Jan. 2013 meteorology (Scenario *a*) resulted in similar PM_{2.5} concentration
193 ranges (~50 to ~500 µg/m³) and spatial distributions. Difference of PM_{2.5} concentration at any
194 site was within ± 10 µg/m³ (Fig. 3e). Simulation results of Scenario *a* and *c* were not only
195 similar in average concentration levels, but also in temporal variations. For example, in Beijing,
196 simulated hourly PM_{2.5} concentration results under this two scenarios presented not only similar

197 concentration (being $279.1 \pm 170.2 \mu\text{g}/\text{m}^3$ and $278.8 \pm 168.9\mu\text{g}/\text{m}^3$, respectively) but also
198 excellent correlation with R^2 reaching 0.97.

199 In contrast, stable synoptic conditions in Jan. 2013, which favored accumulation of emitted
200 pollutants, were essential to the formation of the severe regional haze. Under the same emission
201 level, changing the meteorological conditions from 2012 to 2013 resulted in a monthly average
202 $\text{PM}_{2.5}$ increase of 10–40 $\mu\text{g}/\text{m}^3$ in the Beijing area, and up to 120 $\mu\text{g}/\text{m}^3$ over the whole NCP
203 (Fig. 3a, b, d). This suggests that the severe haze episodes in Jan. 2013 were most likely due to
204 unfavorable meteorology, rather than an abrupt increase in emissions (Fig. 3d, e).

205 Figure 4 compares peak $\text{PM}_{2.5}$ concentrations in the NCP region during Episodes II to IV and
206 their corresponding surface weather maps, together with surface weather map from a clean hour
207 (Fig. 4g). During severe haze episodes, the regional pollution covered most of Hebei Province
208 and northern Henan Province. In general, Shandong Province was less polluted, except during
209 Episode IV. Beijing borders this polluted region, with mountains to the northwest. Surface
210 weather maps from polluted periods were generally characterized by a weak high-pressure
211 center (1034–1037 hPa) northeast of Beijing, which could result in low surface wind speed and
212 prevent the influx of northwest clean air (Xu et al., 2011; Zhao et al., 2013). During the peak
213 hours of Episode II, Beijing was located near a low-pressure trough, where air masses from
214 south, west and northeast converged. During Episode III, Beijing was located in a saddle
215 between two pairs of high- and low-pressure centers, which also led to enhanced stability. In
216 contrast, weather patterns for the clean hours were characterized by strong high-pressure centers
217 (up to 1046 hPa) northwest of Beijing, i.e., the Siberian Anticyclone. With sharp pressure
218 gradient, synoptic conditions produce effective convection and strong northerly winds, bringing
219 dry and clean air masses into Beijing.

220 Local meteorology, controlled by synoptic conditions, could have “deterministic impacts” on
221 air pollution levels (Xu et al., 2011). Compared with the clean periods, the polluted periods
222 were associated with significantly lower wind speed and PBL, and higher temperature and RH
223 (Fig. 5). Besides changes in the average level, diurnal pattern of temperature in polluted periods
224 could also differ from clean periods, with diminished overnight (0:00 to 6:00 a.m.) temperature
225 drop.

226

227 5 Local Chemical Production vs. Regional Transport

228 As shown in Fig. 6, Episode II consists of several sharp-increase events, in which PM_{2.5}
229 concentrations increased by over 400 µg/m³ within 1–3 hours (maximum mass growth rate up
230 to 351.8 µg/m³/h). Earlier studies have attributed this dramatic rate of increase to fast local
231 chemical production (Y. Wang et al., 2014). However, we found that the apparent rapid changes
232 are more likely to be caused by the regional transport of clean/polluted air masses. In winter,
233 the Siberian Anticyclone could bring clean air masses into NCP (Jia et al., 2008; Liu et al., 2013)
234 while southerly winds refill the areas with polluted air masses. The transition between clean
235 and polluted air masses may result in an apparent sharp build-up of particle concentrations. In
236 other words, these events reflected interruption and rapid recovery of pollution from adjacent
237 areas, rather than merely local chemical production.

238 The impact of transport is supported by the temporal variations in the regional distribution of
239 PM_{2.5} concentrations, the surface weather maps, and the specific humidity (Fig. 6 and 7). The
240 first evidence is that these sharp PM_{2.5} build-up events were unique to Beijing among all the 8
241 cities around/in the NCP (Fig. 6). Chengde and Zhangjiakou are situated to the north of NCP
242 with mountains in between (Fig. 6a). Among the NCP cities, Beijing is located at the northern
243 tip, with mountains to the north and west shielding the city (Fig. 6 (a2)). When conditions favor
244 transport of clean air from north or northwest (i.e. with the advent of a cold air current), Beijing
245 is the first one among NCP cities to **be scavenged**, which resulted in a sharp drop of PM_{2.5}
246 concentrations. In this case, PM_{2.5} levels in Beijing became similar to the upwind cities, i.e.,
247 Chengde and Zhangjiakou (yellow solid circles; Fig. 6(b1)). However, these cold air currents
248 were too weak to go further, leaving the rest NCP cities unaffected. Not surprisingly, the
249 influence of these weak cold air currents soon receded and the polluted air parcels were
250 transported back to Beijing, which lead to a sharp increase in the PM_{2.5} level similar as the rest
251 NCP cities (e.g., Shijiazhuang, Baoding, Tianjin, Langfang, and Tangshan) (yellow solid circles;
252 Fig. 6 (b2 and b3)).

253 In accordance with the above description, surface weather maps showed that the sharp PM_{2.5}
254 increase/decrease events in Beijing during Jan. 2013 were always accompanied with quick
255 transition between low/high pressure systems. As shown in Fig. 7b, the two sharp drops in PM_{2.5}
256 concentration on 11 and 12 Jan. corresponded to a weak high-pressure system developed in the
257 mountains northwest of Beijing, which brought clean air mass into the city. When the high-

258 pressure systems diminished, a low-pressure system developed southwest of Beijing, and the
259 air mass in Beijing was again affected by the regional background pollution, resulting in a sharp
260 increase in PM_{2.5} concentration.

261 The observed variation of the specific humidity, an indicator for the origin of air masses (Jia et
262 al., 2008), also supports our explanation (Fig. 7a). Air masses from the south were usually
263 warmer and wetter than the northern air masses, thus possessing a higher specific humidity.
264 During the rapid changes of PM_{2.5}, the trend of specific humidity nicely followed the variations
265 of PM_{2.5} (Fig. 7a, pink and yellow rectangles marked periods), which reflected the quick
266 transition of air parcel origins. It has been suggested that the decrease of PBL height will
267 compress air pollutants into a shallow layer, resulting in elevated pollution levels (Liu et al.,
268 2013). However, our results indicated that the compression was not really happening. Rather,
269 the decrease of PBL height hindered the vertical mixing of pollutants, resulting in a faster
270 accumulation and higher concentrations. As shown in Fig. 7a, the time lag between variations
271 in PBL and its effects on PM_{2.5} concentration is a clear evidence demonstrating that the PBL
272 was not “compressing” air pollutants into a shallower layer. Otherwise, concurrent increase in
273 PM_{2.5} will be found during the decrease of PBL height.

274

275 **6 Formation of Secondary Aerosols**

276 Compared with clean conditions, the hazy days are characteristic of weaker radiation and higher
277 RH. The RH depends on the synoptic conditions while the radiation reduction is due to the
278 direct radiative effects of aerosol particles (Crutzen and Birks, 1982; Ramanathan and
279 Carmichael, 2008; Ramanathan et al., 2001; Cheng et al., 2008b; Wendisch et al., 2008).
280 Secondary aerosols (inorganic and organic) are major components in fine particles in China
281 (Yang et al., 2011). In this section, we will evaluate the impact of changes in radiation and RH
282 on the formation of secondary aerosols.

283 **To evaluate the role of chemical productions, we analyzed the EC-scaled concentrations for**
284 **individual compounds. The purpose of using EC-scaled concentration is to eliminate the**
285 **influence of different dilution/mixing conditions on the variation of observed pollutant**
286 **concentrations. The observed variations of pollutant concentrations are not only controlled by**
287 **the chemical reactions but also subject to the influence of boundary layer developments. For**

288 the same emission rate and chemical production rate, different mixing conditions will result in
289 different level of air pollutants. It is thus highly uncertain to conclude a stronger/weaker
290 chemical production based on purely concentrations data without considering the boundary
291 layer effect. Since EC is an aerosol species coming from only primary emission and quite
292 inertial to chemical reactions, its variations well reflect the influence of atmospheric physical
293 processes (dilution/mixing effect). The ratio of other species to EC will to a large extent
294 eliminate the variations due to mixing/dilution and better represent the contribution from
295 chemical reactions.

296 **6.1 Weakened Importance of Photochemistry**

297 The radiative reduction imposed by aerosol particles is particularly strong during haze episodes
298 because of extremely high particle concentrations. Take Beijing for example, during haze
299 episodes, the amount of solar radiation reaching the ground was significantly lower (e.g., down
300 to 2.77 MJ/m²/day, 13 Jan.) than clean days (averaging 9.36 ± 0.60 MJ/m²/day for all the six
301 clean days), rendering high photochemical activity impossible. The reduction of radiation
302 intensities will change the atmospheric photochemistry and oxidant concentrations (hydroxyl
303 radical (OH) and ozone (O₃)), which will consequently change the production and aging of
304 secondary organic aerosols (SOA) (Hallquist et al., 2009; Jimenez et al., 2009).

305 As the haze pollution spread over most of the NCP, a weakening of photochemistry was
306 expected on the regional scale, which is confirmed by both observations and model simulations.
307 Extremely low ozone concentration (less than 10 ppb) in the absence of diurnal variation was
308 observed during heavy pollution episodes for all of the three major cities in Jing-Jin-Ji Area
309 (i.e., Beijing, Tianjin and Shijiazhuang) in Jan. 2013 (Y. Wang et al., 2014). Similar
310 phenomenon was observed before in another heavy pollution episode in winter Beijing (Zhao
311 et al., 2013). In accordance with the observed low ozone concentration, model simulations also
312 showed a regional-scale reduction in the concentrations of ozone and OH (Fig. 8). Average
313 daytime concentrations of oxidants were significantly lower during polluted periods than clean
314 periods. For most areas in the NCP, O₃ and OH dropped from 12~44 ppbV and 0.004 ~ 0.020
315 pptV to less than 12 ppbV and 0.004 pptV, respectively, as the air quality changed from clean
316 to heavily polluted conditions. This regional drop in oxidant concentrations demonstrates the
317 impact of air pollution on the photochemistry.

318 Ozone and OH radicals are known as crucial oxidants in the formation of secondary organics
319 aerosols (SOA) (Jimenez et al. 2009). Weakened photochemistry is therefore expected to reduce
320 the SOA production and concentrations. To have a semi-quantitative estimation on the
321 contribution of photochemistry, secondary organic carbon (SOC) was estimated (Fig. 9a) using
322 the EC-tracer method (Lim and Turpin, 2002). Briefly, SOC was estimated using these formulae:

$$323 \quad \text{Primary OC} = \text{EC} * (\text{OC/EC})_{\text{pri}} + \text{N} \quad (2)$$

$$324 \quad \text{SOC} = \text{OC} - \text{Primary OC} \quad (3)$$

325 The basic assumptions and underlying principles of this method are discussed in Lim and
326 Turpin (2002) and Lin et al. (2009). Only daytime (7:00~18:00) carbonaceous aerosol data were
327 used here to exclude possible interference from day-night source variations (such as the heavy-
328 duty diesel truck traffic which is allowed only during nighttime in Beijing). In our study, data
329 pairs with the lowest 10% percentile of ambient OC/EC ratios were used to estimate the primary
330 OC/EC ratio (Fig. 9a). York regression (York et al., 2004) was used to estimate the intercept N
331 and the slope, i.e., values of (OC/EC) pri, according to Saylor et al. (2006). Our analysis shows
332 that SOC constituted ~28% of total OC, consistent with earlier studies in the winter of 2009-
333 2012 (~30%, Cheng et al., 2011; Sun et al., 2013b).

334 High concentration of aerosol particles can reduce solar radiation and atmospheric
335 photochemistry. Since SOC is a product of photochemical reactions, we would expect a reduced
336 SOC production rate under heavily polluted conditions. This is confirmed by the measured SOC
337 concentrations shown in Fig. 9. Here again the EC-scaled SOC was used to account for the
338 different boundary layer effect (dilution/mixing) on the aerosol concentrations. Both SOC/EC
339 and the accumulated SOC/EC (afternoon – morning values) decrease when it changed from
340 clean to heavily polluted periods. The accumulated SOC/EC is used to better represent the
341 production during the daytime.

342 Reduction in photochemistry-related PM_{2.5} production is further supported by model simulation
343 results. In our model configurations, the photolysis rate is calculated online using simulated
344 aerosols and ozone concentrations (B. Zheng et al., 2014). As a result, with the enhanced PM_{2.5}
345 concentration, the photolysis rate will be reduced, and so will the concentrations of
346 photochemical oxidants (Fig. 8) and secondary aerosol particles. During the haze events, this

347 effect can be counteracted by the enhanced heterogeneous reactions and it is difficult to unravel
348 them from the measurement data.

349 In order to demonstrate the influence of reduced photochemistry, we adopted the original WRF-
350 CMAQ model setup and excluded the enhanced heterogeneous reactions. In this case, only gas
351 phase oxidations are counted for the formation of sulfate and organics (aqueous-phase reactions
352 in the original WRF-CMAQ only happen in clouds and don't apply for the aerosol phase) (B.
353 Zheng et al., 2014) and their simulated concentrations will directly reflect the influence of
354 reduced photochemistry. As shown in Table S1, the simulated $PM_{2.5}/EC$ ratios decreased from
355 16.05 to 11.72 when the pollution level changed from the clean to the heavily polluted case,
356 reflecting the reduced gas-phase photochemical production. Note that $PM_{2.5}$ concentration is
357 normalized by EC to counteract the influence of reduced boundary layer. Otherwise, the
358 reduced boundary layer itself could lead to a tremendous increase in the pollutant concentration
359 under heavily polluted conditions, and thus cover the real effect of reduced photochemistry.

360 The simulated individual components of $PM_{2.5}$ also reflected the influence of photochemistry.
361 As shown in Table S1, although primary organic matter (POM) to EC ratios kept nearly constant
362 during all pollution levels, the normalized secondary species all showed a decreasing trend,
363 reflecting the reduced photochemical production. SOA/EC , SO_4^{2-}/EC , and NO_3^-/EC ratios
364 decreased by 53.3%, 51.9% and 28.6%, respectively from clean to heavily polluted periods. For
365 the formation of NO_3^- , two heterogeneous reactions have been included in the original WRF-
366 CMAQ model and therefore the NO_3^-/EC shows relatively less reduction than SOA/EC and
367 SO_4^{2-}/EC .

368

369 **6.2 Enhanced Heterogeneous Chemistry**

370 Unlike OM, relative contributions of sulfate and nitrate to $PM_{2.5}$ were increasing during the
371 haze events (Fig. 2). Again, we used their ratios to EC to account for the boundary layer effect.
372 An increasing trend of SO_4^{2-}/EC and NO_3^-/EC ratios was found (Column 1 in Fig. 10) from
373 clean periods (3.03 and 3.33, respectively) to heavily polluted periods (6.35 and 5.89,
374 respectively), suggesting enhanced chemical productions. The SOR and NOR (molar ratio of
375 sulfate or nitrate to sum of sulfate and SO_2 or nitrate and NO_2) have been used as indicators of

376 secondary transformation (Sun et al., 2006). The fact that SOR and NOR increased much more
377 rapidly than SO₂ and NO₂ as pollutions became more severe (Column 4 in Fig. 10), is another
378 evidence of elevated secondary formations of sulfate and nitrate during severe haze events.

379 Both gas-phase and heterogeneous reactions could contribute to the formation of sulfate and
380 nitrate from SO₂ and NO₂, and thus elevating the SOR and NOR. Sulfate is formed through
381 oxidation of SO₂ by gas-phase reactions with OH (Stockwell and Calvert, 1983; Blitz et al.,
382 2003) and stabilized Criegee intermediate (which is formed by O₃ and alkenes) (Mauldin et al.,
383 2012), and by heterogeneous reactions with dissolved H₂O₂ or with O₂ under the catalysis of
384 transition metal (Seinfeld and Pandis, 2006). Nitrate formation is dominated by the gas-phase
385 reaction of NO₂ with OH during daylight, and the heterogeneous reactions of nitrate radical
386 (NO₃) during nighttime (Seinfeld and Pandis, 2006). **Since gas phase production of secondary
387 aerosols is expected to decrease under heavily polluted periods (Section 6.1), the increase of
388 SO₄²⁻/EC and NO₃⁻/EC ratios is a clear evidence for the dominant contribution from other
389 pathways, most probably from the heterogeneous reactions.**

390 **If we assume heterogeneous chemistry as answer to the high SO₄²⁻ and NO₃⁻ concentrations,
391 there is a problem because heterogeneous chemistry still requires oxidation by oxidizing agents,
392 e.g. OH, O₃, etc., which were indeed significantly reduced (Section 6.1). Our explanation for
393 this puzzle is that despite of reduced oxidant concentrations, the aerosol volume/surface
394 increases so much (due to elevated aerosol concentration and the accompanied high RH, Fig.
395 1) that it is enough to compensate its influence, and moreover, leads to a net increase in the
396 formation of secondary aerosols.**

397 **A simplified case study could show how aerosol volume/surface increases could compensate
398 the effect of oxidant reduction, and even lead to a net increase in the formation of secondary
399 aerosols. Take sulfate for example, the production rate of sulfate (S(VI)) through heterogeneous
400 reactions can be estimated by:**

$$401 \quad dC_{S(VI)}/dt \approx k[S(IV)(aq)]*[oxidants(aq)]*V_{aerosol} \quad (4)$$

402 **in which $C_{S(VI)}$ is the sulfate concentration, k is the effective rate coefficient, $[S(IV)(aq)]$ is the
403 $S(IV)$ concentration in the aqueous phase of aerosols, $[oxidants(aq)]$ is the concentration of
404 oxidants in the aqueous phase of aerosols, and $V_{aerosol}$ is the volume concentration of
405 humidified aerosol at ambient RH.**

406 Equation (4) shows that the oxidants and V_{aerosol} are both essential for the heterogeneous
 407 reactions. From the clean to the heavily polluted case, O_3 is reduced by 80%, dropping from >
 408 $50\mu\text{g}/\text{m}^3$ to $< 10\mu\text{g}/\text{m}^3$ (Y. Wang et al., 2014). Based on our model simulation results, H_2O_2
 409 concentration also dropped significantly from ~ 78 ppbV to ~ 11 ppbV. Thus we assume an
 410 upper limit of 90% reduction in [oxidants (aq)]. V_{aerosol} depends on the dry aerosol
 411 concentrations V_{dry} and its hygroscopic growth factor (GF) of particle size, which is a function
 412 of RH. Assuming a constant aerosol dry density, then V_{dry} is proportional to the mass
 413 concentration. From clean to heavily polluted case, average $\text{PM}_{2.5}$ mass concentration
 414 increased by 25 time, changing from $18\mu\text{g}/\text{m}^3$ to $450\mu\text{g}/\text{m}^3$ while average RH increased from
 415 dry ($\sim 20\%$) to $\sim 70\%$. Thus we have:

$$416 \frac{[\text{oxidants (aq)}]_{\text{HP}}}{[\text{oxidants (aq)}]_{\text{Clean}}} * \frac{(V_{\text{aerosol}})_{\text{HP}}}{(V_{\text{aerosol}})_{\text{Clean}}} = \frac{[\text{oxidants (aq)}]_{\text{HP}}}{[\text{oxidants (aq)}]_{\text{Clean}}} * \frac{(V_{\text{dry}})_{\text{HP}}}{(V_{\text{dry}})_{\text{Clean}}} * (GF_{\text{HP/Clean}})^3 \approx 0.1 * 25 * (1.1)^3 = 3.33 \quad (5)$$

417 where HP and Clean indicated heavily polluted and clean periods, respectively. A GF of 1.1
 418 was taken from previous measurements in Beijing (Meier et al., 2009).

419 Equation (5) shows that the increase of aerosol volume concentrations could sufficiently
 420 compensate the effect of oxidant reduction, resulting in a net increase of sulfate production.

421 Similarly, for NO_3^- , the influence of oxidant reduction could also be compensated by the
 422 increase of aerosol volume concentrations. There might be other oxidants associated with
 423 heterogeneous reactions, such as O_2 (especially under the catalysis of mineral metals) and
 424 other oxidants existed in aerosol phase such as Organic Peroxides (Seinfeld et al., 2006).

425 In accordance with above discussions, both observation and model simulation supported the
 426 importance of heterogeneous reactions. Observed SOR and NOR showed an obvious
 427 dependence on RH (Fig. 11). Both SOR and NOR were constant under dry conditions (RH <
 428 50%) (Fig. 11 a, b) while started increasing when RH >50%, resulting in average values
 429 around 0.34 and 0.28 at RH 70%–80%, respectively. This suggests important contributions
 430 from heterogeneous reactions with abundant aerosol water under wet conditions (Sun et al.,
 431 2013a). The observed SOR value was high compared with previously reported values of 0.24
 432 (Wang et al., 2006) and 0.29 (Zhao et al., 2013) during hazy days in Beijing. The NOR value
 433 for this study was higher than for spring hazy days in 2001–2004 (0.22; Wang et al., 2006),
 434 but significantly lower than for the hazy episode in Jan. 2010 (0.51; Zhao et al., 2013). Our

435 model simulation results (B. Zheng et al., 2014) also supported the importance of
436 heterogeneous chemistry in sulfate and nitrate productions (Figure R1). With the addition of
437 the heterogeneous reactions, the revised CMAQ showed much better performance in the
438 polluted periods (B. Zheng et al., 2014), which demonstrated the importance of heterogeneous
439 reaction in the production of secondary aerosols.

440 Concerning the SOA formation, the contribution of heterogeneous reactions might be possible,
441 but it should be much less significant than for sulfate and nitrate. For $RH > 50\%$, SO_4^{2-}/EC and
442 NO_3^-/EC ratios rose significantly (Fig. 11d) while SOC/EC ratios remained constant (Fig. 9b).
443 By using HOA (hydrocarbon-like organic aerosol) instead of EC, Sun et al. (2013a) found
444 similar phenomena. Apparently, SOC doesn't have a heterogeneous formation pathway as
445 effective as those of sulfate and nitrate.

446

447 **7 Conclusion**

448 The severe haze pollution during Jan. 2013 was not a Beijing-localized phenomenon. Rather, it
449 was the result of local pollutants superposed on background regional pollution, which affected
450 the whole NCP. Although pollutant emissions were high, there was no abrupt enhancement in
451 2013. The occurrence of the severe winter haze resulted from stable synoptic meteorological
452 conditions over a large area of northeastern China. Surface weather maps from hazy periods
453 were characterized by a weak high-pressure center northeast of Beijing, while the termination
454 of a haze episode was always accompanied by the Siberian Anticyclone (Xu et al., 2011; Jia et
455 al., 2008; Liu et al., 2013). Atmospheric chemistry and physics during severe haze pollutions
456 are illustrated in a conceptual model (Fig. 12). With the onset of stable synoptic conditions, RH
457 rises, primary pollutants begin to accumulate and regional pollution begins to form. If the stable
458 conditions last long enough, $PM_{2.5}$ build-up occurs, and as a consequence, solar radiation is
459 reduced at the ground level. This inhibits surface temperature fluctuation, making easier the
460 formation of inversed layer and rendering the atmosphere into a more stable condition.
461 Meanwhile, photochemical activity is weakened under low solar radiation, and secondary
462 aerosol formation via this pathway becomes less important. However, under high RH,
463 heterogeneous reactions may play a more important role, especially those associated with the
464 aerosol aqueous phase. This results in the rapid build-up of secondary aerosols, especially

465 sulfates and nitrates, enhancing $PM_{2.5}$ pollution. The accumulation of aerosol particles
466 terminates with the incursion of a strong cold front, usually the Siberian Anticyclone.

467 Our analysis also reveals that the regional transport can be a key process controlling the
468 variations of local air pollutant concentrations. Take the sharp increases of aerosols
469 concentrations on 11-13 Jan. for example, Beijing pollution was temporarily flushed away by
470 strong winds associated with the arrival of a weak cold air current, as its influence weakened,
471 the polluted regional air mass readily reoccupied the Beijing area, resulting in an apparent rapid
472 build-up of $PM_{2.5}$. This was supported by data on the $PM_{2.5}$ levels around Beijing, specific
473 humidity and PBL height, as well as surface weather maps. Our results reveal that the apparent
474 formation rate (the rate of change in $PM_{2.5}$ or other air pollutants) is not only due to chemical
475 reactions but also controlled by the regional transport along with other processes. It requires
476 caveats to derive a real chemical production rate based on a single-site measurement. Our
477 results also show a clear impact of regional transport on the local air pollution, suggesting the
478 importance of regional-scale emission control measures in the local air quality management of
479 Beijing.

480

481

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Table 1. General information on severe haze episodes in January 2013

		Episode I			Episode II			Episode III			January		
		Ave.	min	max	Ave.	min	max	Ave.	min	Max	Ave.	min	max
Meteoro- logy Data	Temperature (°C)	-4.63	-11.10	1.40	-2.79	-8.30	2.80	-1.26	-6.60	5.00	-2.94	-12.50	5.00
	RH(%)	31.16	13.50	58.50	56.59	27.20	77.60	55.05	27.10	79.70	47.97	13.50	88.30
	WS(m/s)	2.10	0.90	4.40	2.08	1.00	3.40	1.96	0.90	3.60	2.18	0.90	4.50
PM_{2.5} and PM₁₀	PM _{2.5} (µg/m ³)	112.50	11.00	311.50	245.37	21.10	855.10	167.66	35.40	387.30	161.77	4.40	855.10
	PM ₁₀ (µg/m ³)	152.17	28.80	411.00	327.17	31.60	1157.50	214.03	41.50	479.80	223.53	13.90	1157.50
	PM _{2.5} /PM ₁₀ (%)	0.69			0.75			0.79			0.70		
Gas Data	NO ₂	76.39			109.44			95.86			86.09		
	SO ₂	79.73			123.35			63.86			77.54		

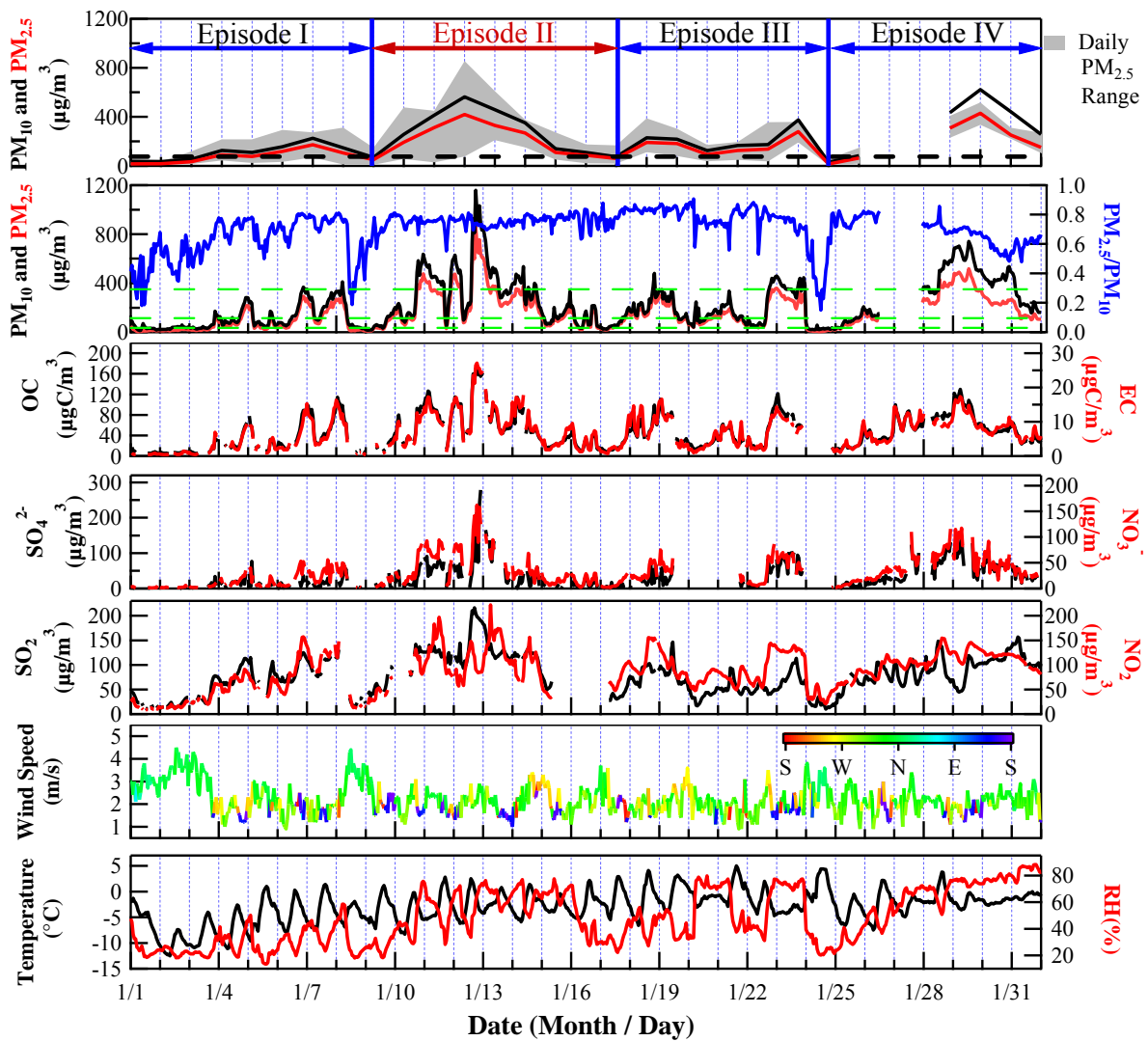


Figure 1. Time series of PM₁₀, PM_{2.5}, and its major components (OC, EC, SO₄²⁻ and NO₃⁻), and meteorological data (wind speed, wind direction, temperature and relative humidity) for January 2013.

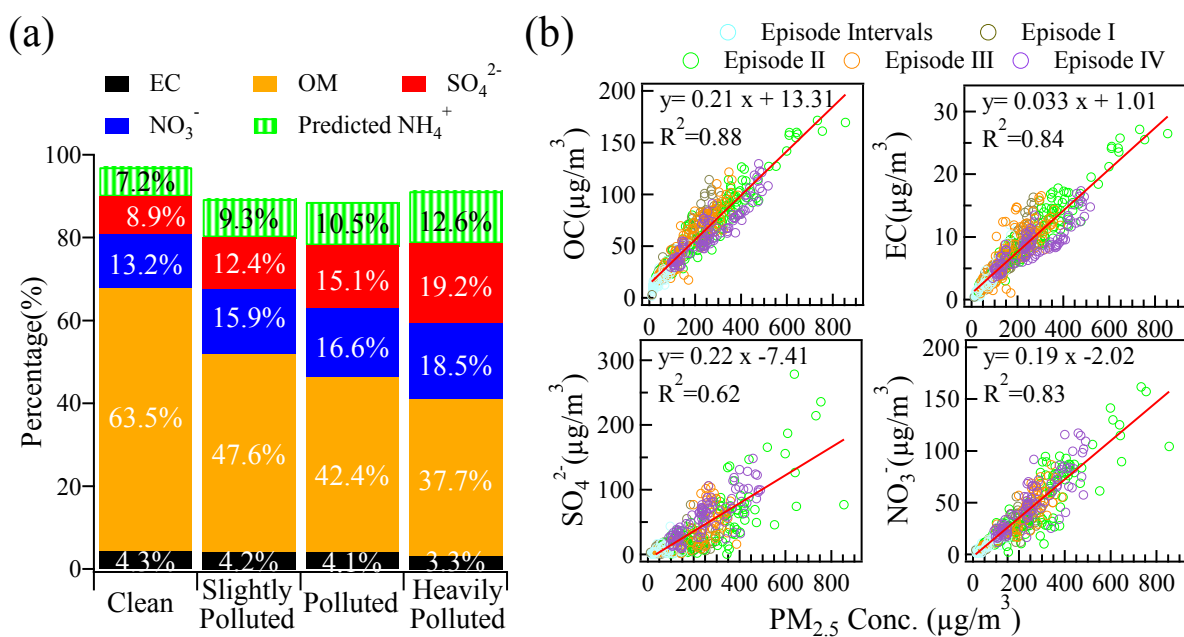


Figure 2. Major composition of PM_{2.5} with respect to pollution level. (a) Mean percentile composition and (b) hourly concentration of individual species plotted against PM_{2.5} mass concentration. Values showed in (a) were derived as average of ratios.

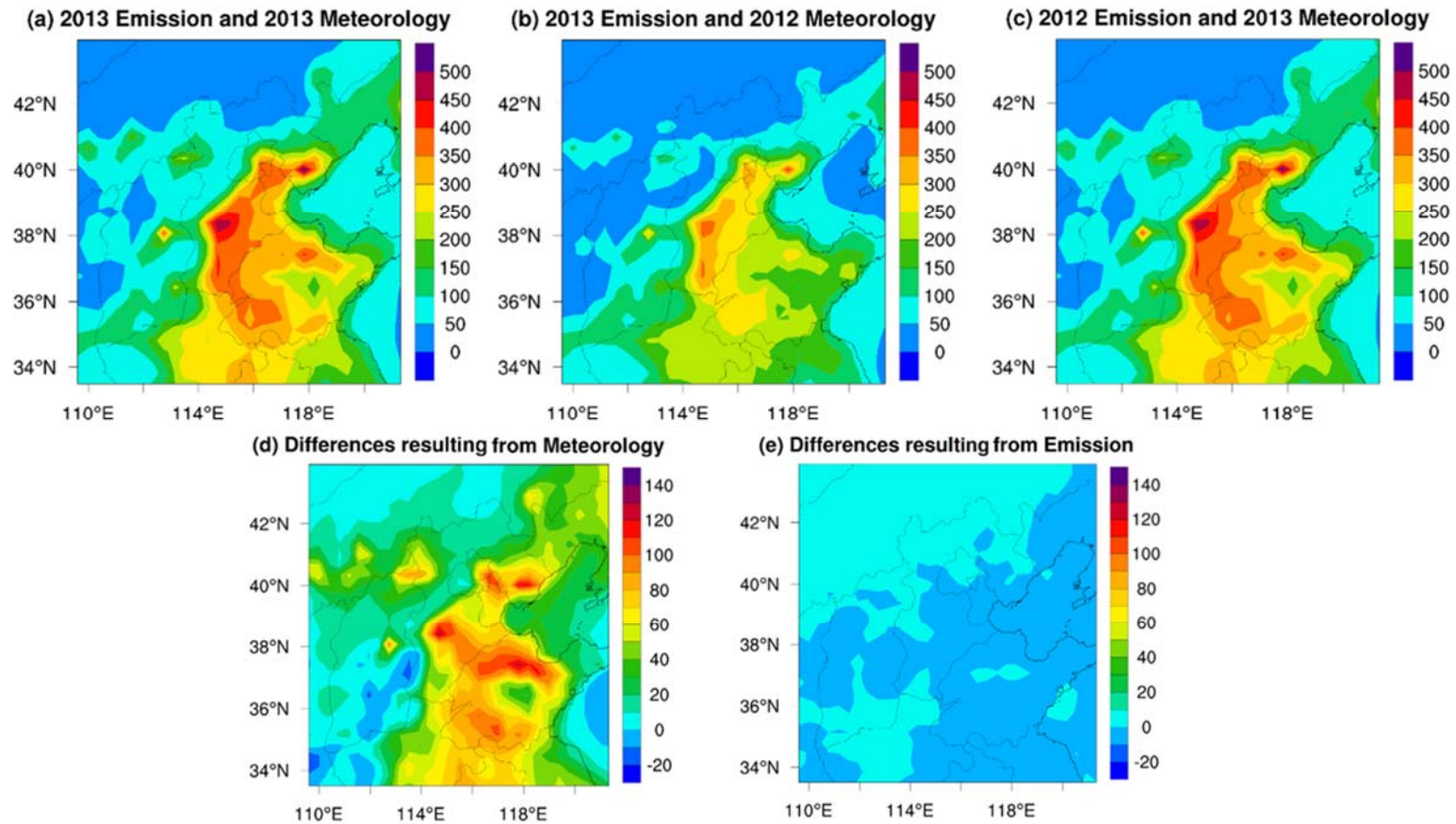


Figure 3. Revised WRF-CMAQ simulated monthly-averaged $PM_{2.5}$ concentration ($\mu g/m^3$) under different scenarios. (a) Base scenario. Actual Jan. 2013 emission and Jan. 2013 meteorology data were used. (b) Jan. 2012 meteorology data were used, and (c) Jan. 2012 emissions were used. The different $PM_{2.5}$ concentrations ($\mu g/m^3$) caused by meteorology (d; equivalent to a–b) and emission (e, equivalent to a–c) are also shown.

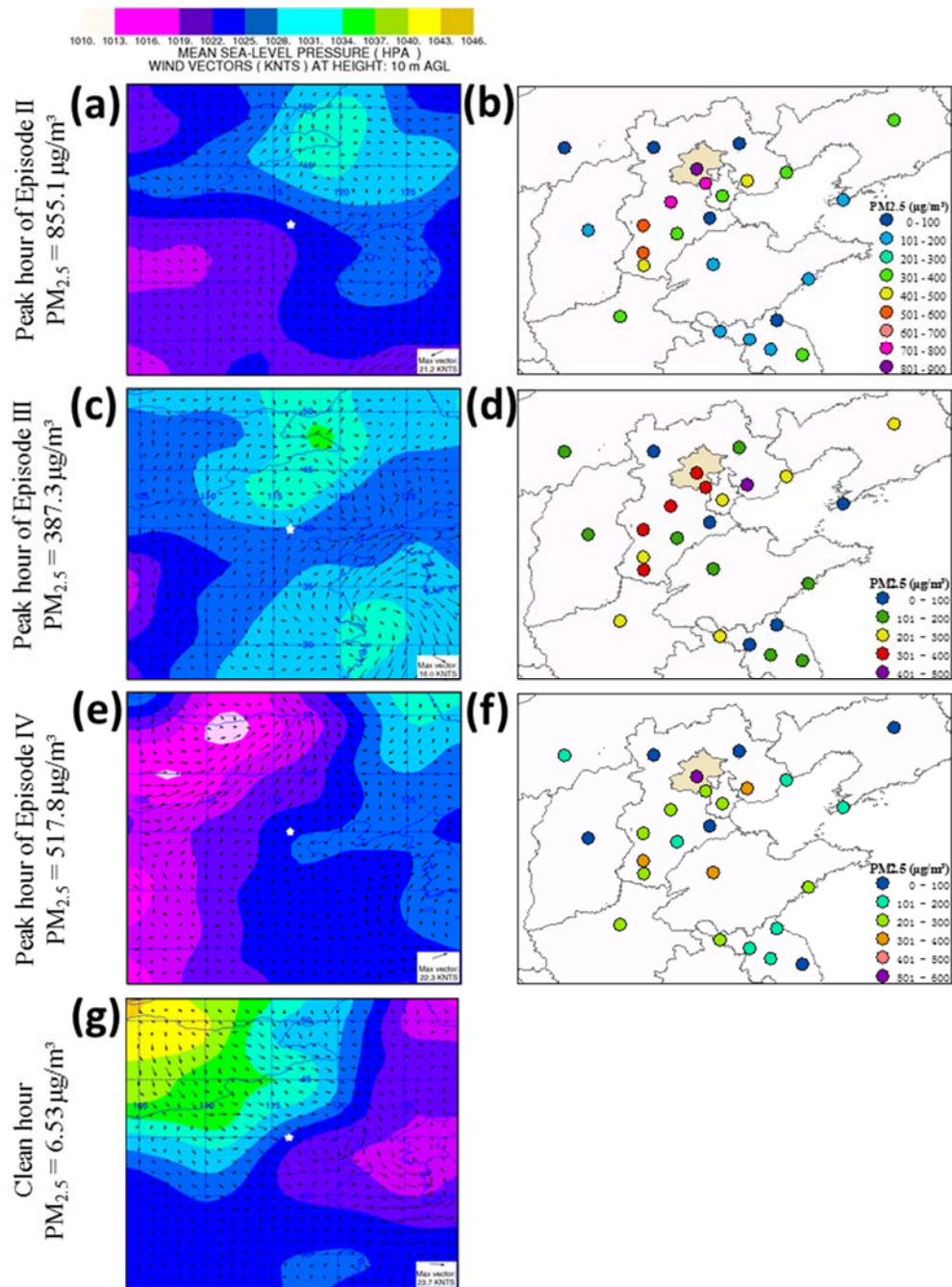


Figure 4. Surface weather maps (a, c, e, g) and PM_{2.5} concentrations (b, d, f) of the North China Plain on 12 Jan. LT 18:00 (a, b), 18 Jan. LT 20:00 (c, d), 29 Jan. LT 13:00 (e, f), and 1 Jan. LT 8:00 (g). The location of Beijing is indicated as a white star on the weather maps, and as the shaded area on the PM_{2.5} concentration maps. PM_{2.5} concentrations in Beijing at the four selected time points are also shown on the left for reference.

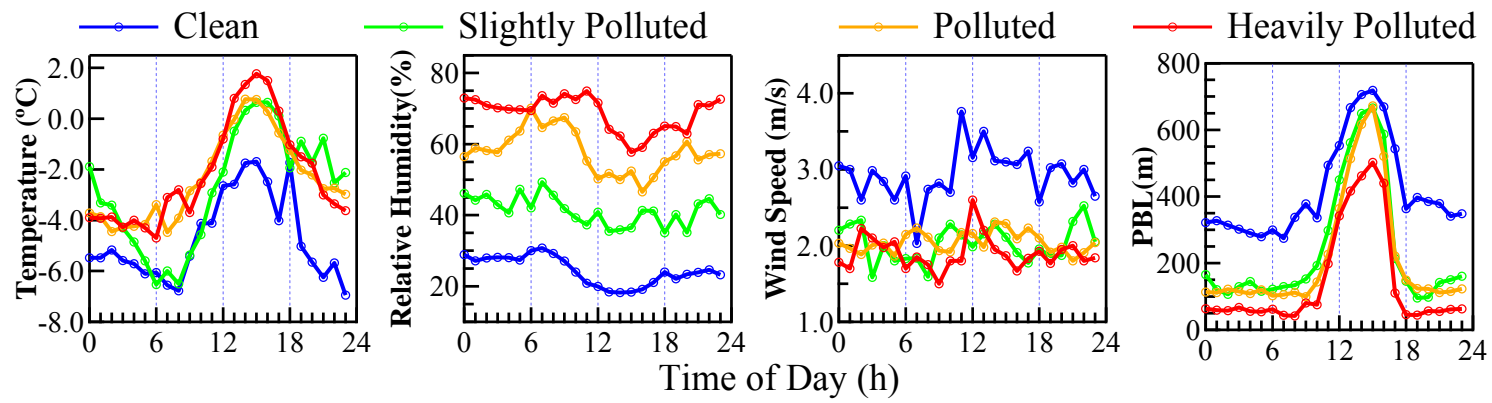


Figure 5. Mean diurnal variation in meteorological parameters for different pollution levels.

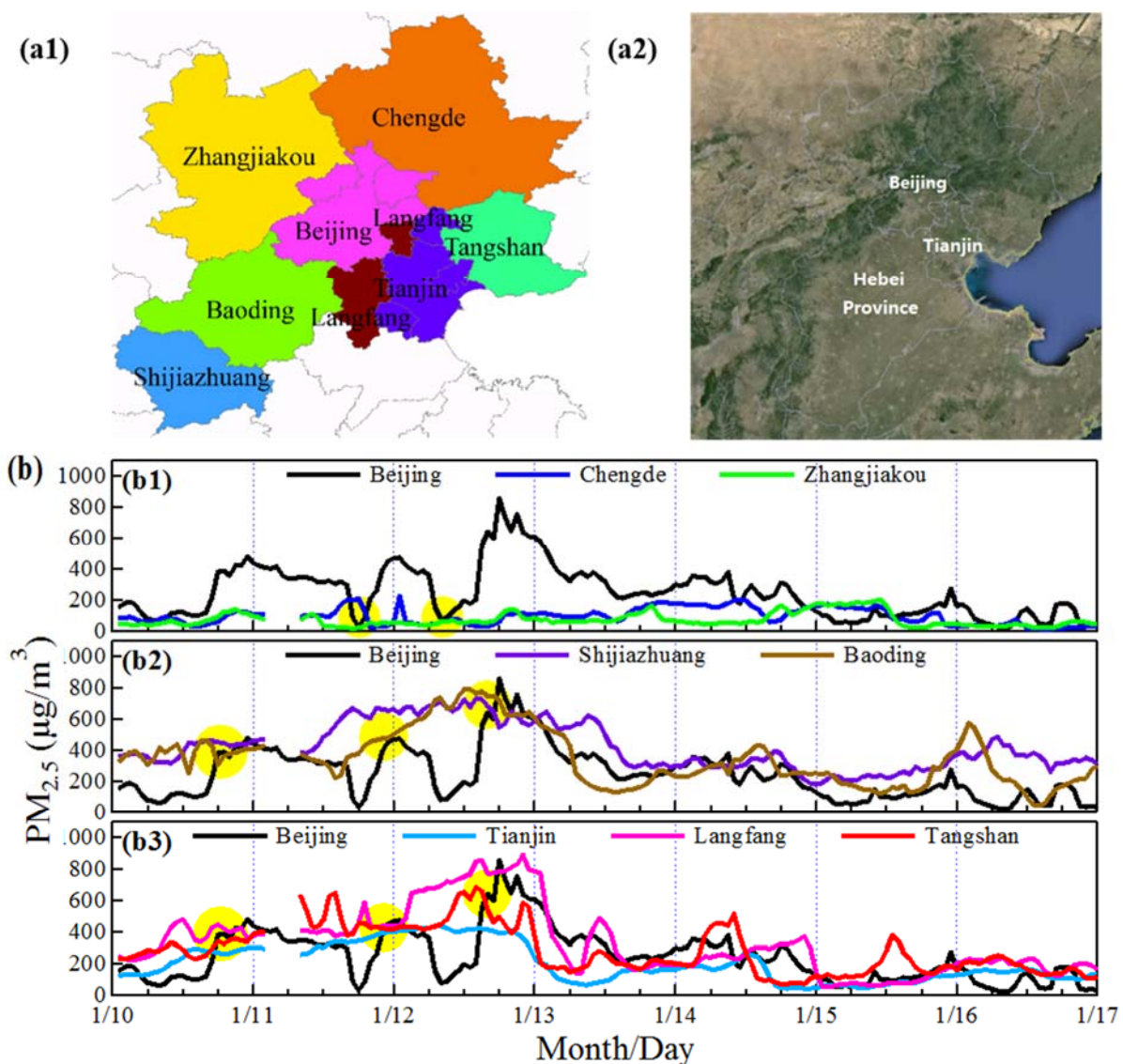


Figure 6. (a1) The location of all cities shown below, and (a2) topographic map around Beijing. (b) PM_{2.5} concentrations of Beijing and its (b1) northern cities, (b2) southwest cities, and (b3) southeast cities for the period 10–17 Jan. 2013. Yellow solid circles indicated the time periods when the sharp drops (b1) and sharp increases (b2 and b3) of PM_{2.5} concentration occurred.

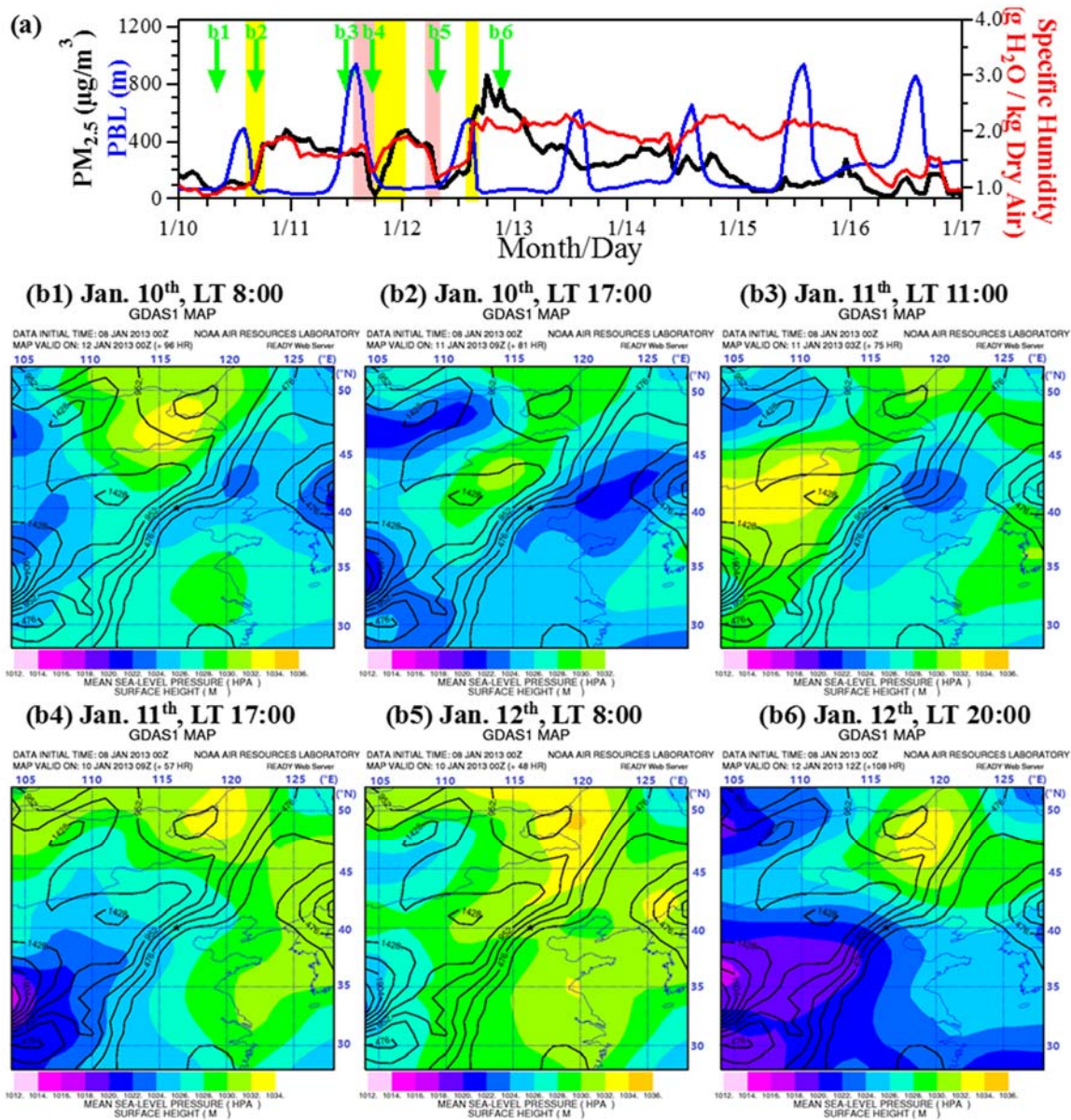


Figure 7. Evidence for regional transport of pollutants as a major factor contributing to sharp concentration increases in Beijing. (a) PM_{2.5} concentration, PBL height, and specific humidity in Beijing for 10–17 Jan. 2013. Pink and yellow rectangles indicated the sharp drop and sharp increase periods of PM_{2.5}, respectively. Note how nicely specific humidity and PM_{2.5} followed each other during these periods. (b) Weather patterns before and after the sharp increases events. Corresponding time point of b1 to b6 was indicated by arrows in (a). The topography map (elevation) is also shown for reference. Location of Beijing was indicated by the black star in center of each graph.

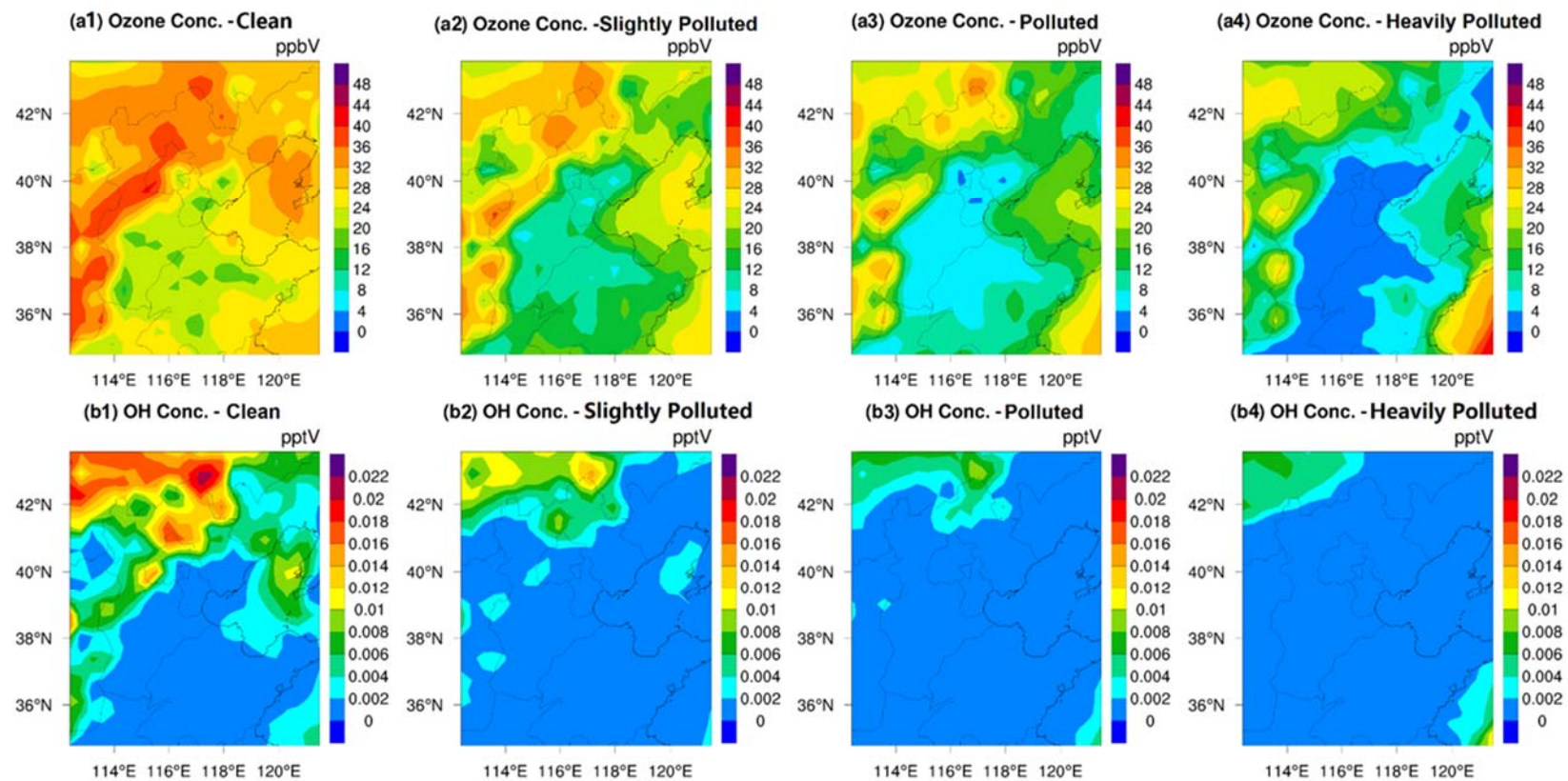


Figure 8. Revised WRF-CMAQ simulated regional distribution of daytime (7:00 ~ 18:00) concentration of (a) O₃ (ppbV) and (b) OH (pptV) at different pollution level.

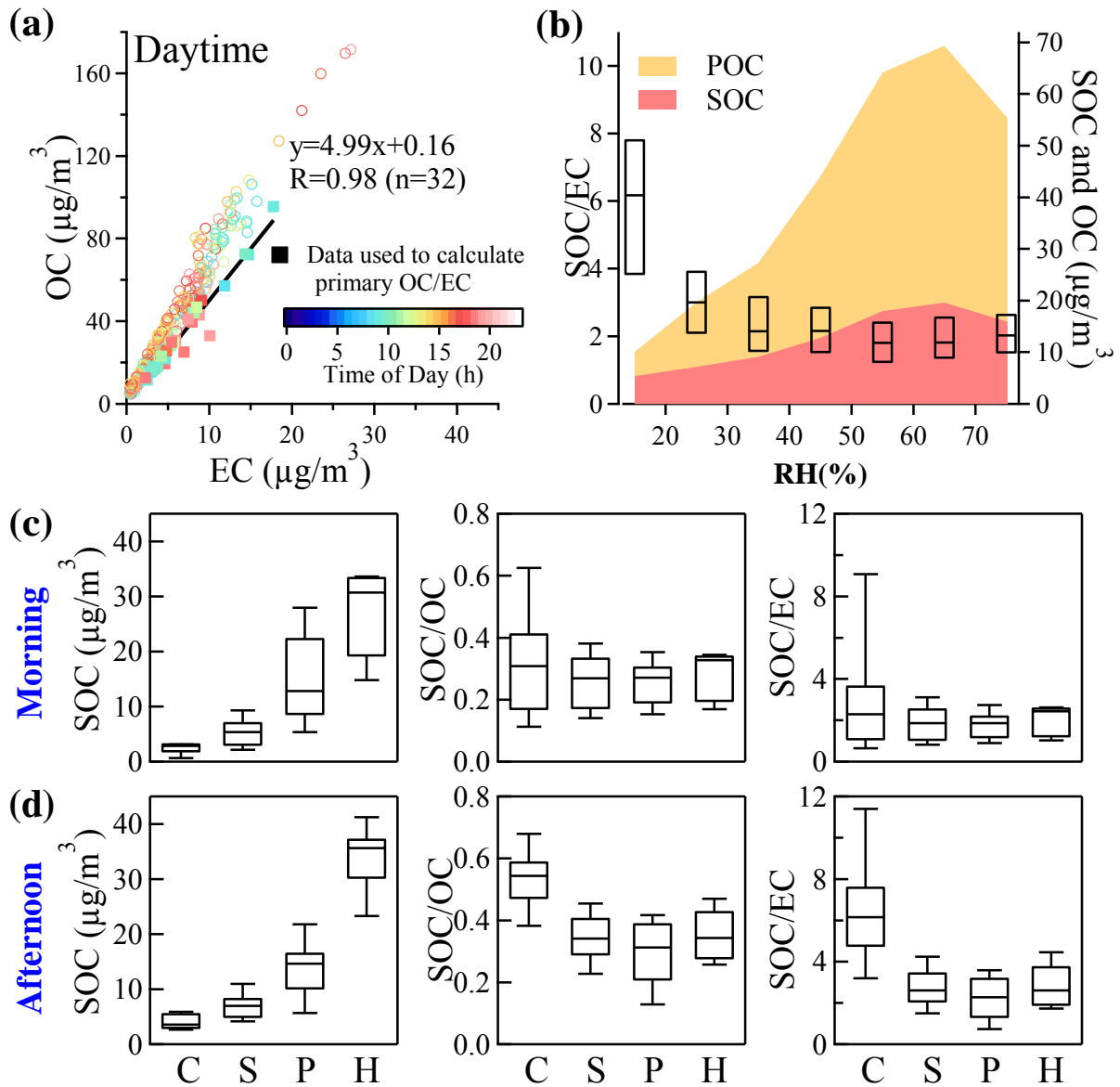


Figure 9. Evaluation of SOC formation. (a) Estimation of SOC with EC-tracer method. Squares indicate data used to calculate primary OC/EC, while open circles indicate other OC/EC data. (b) Change of SOC, OC and SOC/EC with RH. Data points shown in a and b referred to hourly concentrations in daytime (7:00~18:00). (c)-(d) Variation of SOC, SOC/OC and SOC/EC (c) in the morning (7:00~12:00) and (d) in the afternoon (13:00~18:00) with pollution level. “C”, “S”, “P”, “H” refer to “clean”, “slightly polluted”, “polluted” and “heavily polluted”, respectively. In the box-whisker plots, the boxes (b, c, d) and whiskers (c, d) indicated the 95th, 75th, 50th (median), 25th and 5th percentiles, respectively.

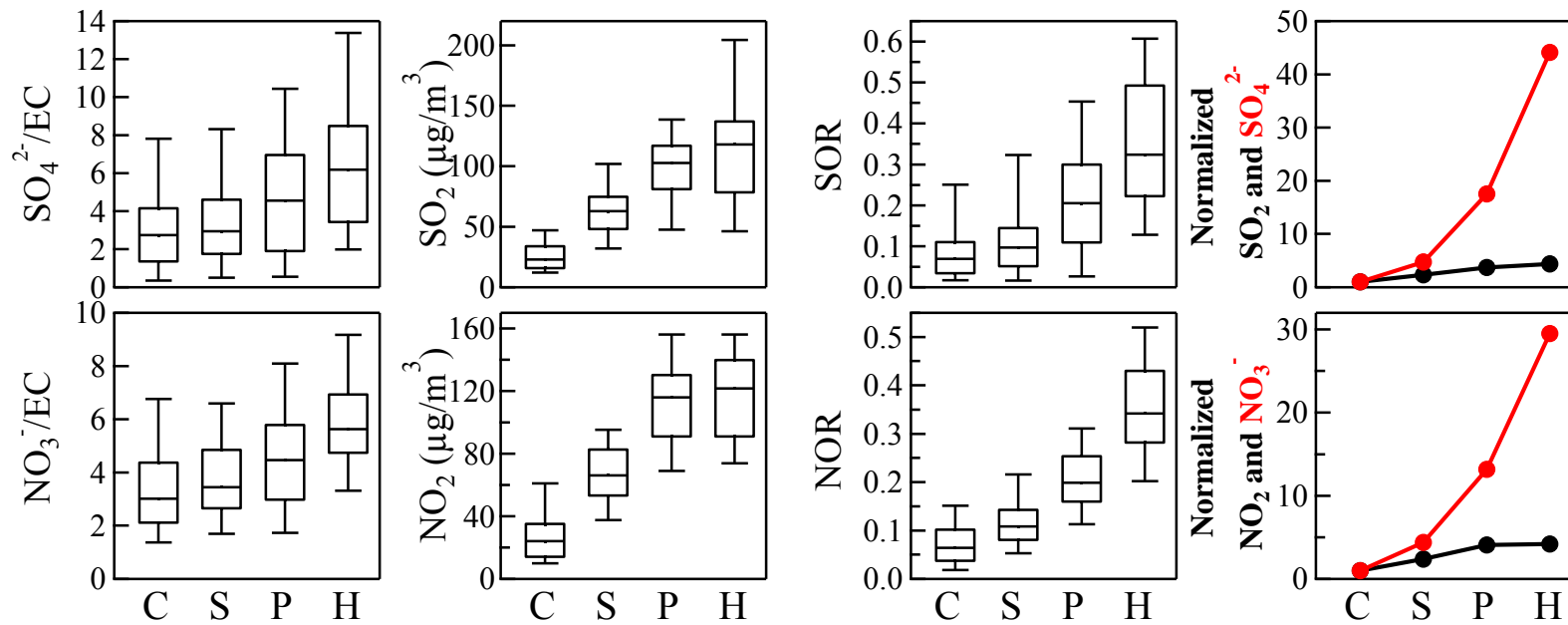


Figure 10. Variation of $\text{SO}_4^{2-}/\text{EC}$, NO_3^-/EC , SO_2 , NO_2 , SOR and NOR with pollution level. “C”, “S”, “P”, “H” refer to “clean”, “slightly polluted”, “polluted” and “heavily polluted”, respectively. Normalized X in Column 4 refers to the average concentration of X in any pollution level, scaled by its average concentration during clean periods. In the box-whisker plots, the boxes and whiskers indicated the 95th, 75th, 50th (median), 25th and 5th percentiles, respectively.

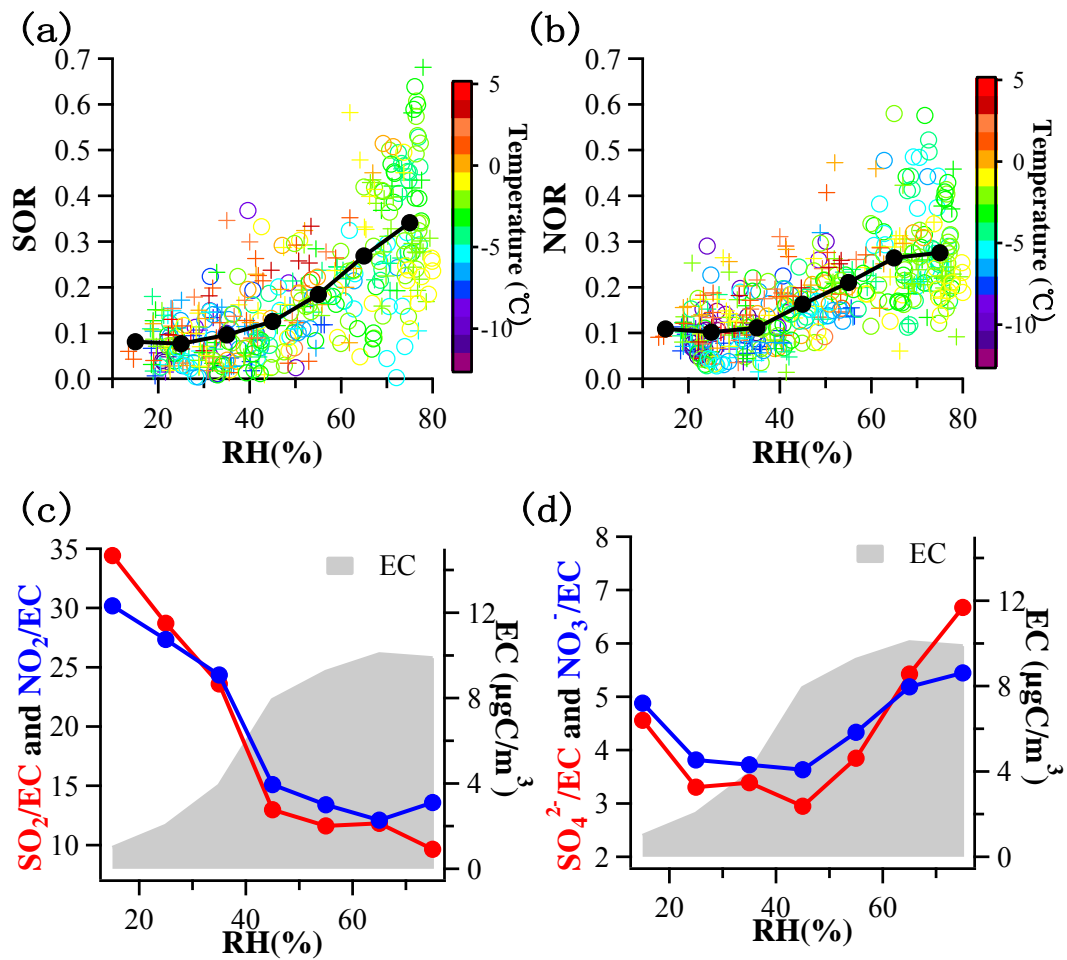


Figure 11. Importance of heterogeneous chemistry in sulfate and nitrate formation. (a-b) Hourly SOR and NOR plotted against RH, colored with temperature. (c-d) EC-scaled precursors (SO₂ and NO₂) and products (SO₄²⁻ and NO₃⁻) plotted against RH. EC concentrations at different RH levels were shown for reference.

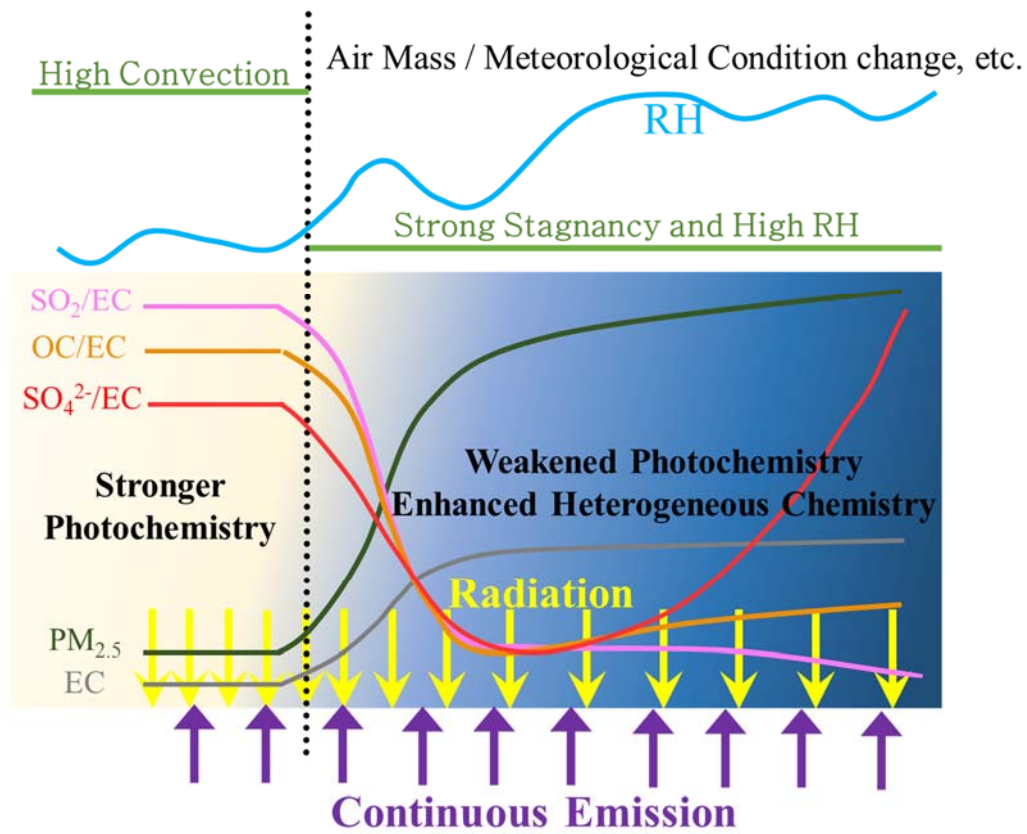


Figure 12. Conceptual model of atmospheric chemistry during the heavy pollutions. The dotted black line indicated the meteorology changed from convection-favoring condition to stagnant condition.