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# Long term characterization of organic and elemental carbon in the PM<sub>2.5</sub> fraction: the case of Athens, Greece

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2000 air sampler (R&P Co), operating at 15.0 and 16.7 L min<sup>-1</sup>, respectively. Sampling duration was 24 h starting at 14:00 LT.

Samples were collected on quartz fiber filters (4.7 cm, Whatman QMA). Filters were prebaked at 550 °C for 4 h to remove any absorbed organic material and were stored and transported to and from the field into petrislides. After sampling they were stored until analysis. Filter blanks and blank field samples were also prepared and analyzed.

For the estimation and correction of the amount of adsorbed organic vapor on the sampling filter, a tandem filter collection method was applied (e.g. Turpin et al., 2000; Kirchstetter et al., 2001; Yttri et al., 2007). Two quartz fiber filters from the same lot were used in series, the front filter collects essentially 100 % of the particulate matter (and some gas-phase material) and the backup filter is exposed only to gas-phase compounds. The measured carbon content of the backup filter can be subtracted from that of the front filter to give an estimate of the particulate phase OC (Turpin et al., 2000; Novakov et al., 2005; Sillanpää et al., 2005). Tandem filters were collected for a selected number of samples, using the aforementioned Partisol samplers. In our analyses, the backup filters indicated a positive OC artifact which was always smaller than 5 %, thus no correction was applied.

## 2.3 Mass and chemical analyses

### 2.3.1 Aerosol mass

PM mass was determined by weighing the filters before and after sampling according to the US EPA RFPS-1298–126 method, using a Mettler Toledo MX5 microbalance (1 µg sensitivity). The filters (samples and blanks) remained for 48 h under controlled conditions (RH 40 ± 5 %, T 20 ± 3 °C), prior to each weighing.

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### 2.3.2 OC – EC analysis

All quartz filters (samples and blanks) were analyzed by a thermal optical transmission technique, using a Sunset Laboratory Inc. (Oregon) carbon analyzer. A punch of 1 cm<sup>2</sup> was removed from the filter and loaded into the analyzer as described by Bougiatioti et al. (2013) and Novakov et al. (2005). OC and EC were determined using the EUSAAR-2 protocol (Cavalli et al., 2010). Briefly, the first phase subsumes 4 stages of heating at 200, 300, 450 and 650 °C at He atmosphere, while, during the second 4-step temperature program, the sample is heated at 500, 550, 700 and 850 °C at He/O<sub>2</sub> atmosphere. The detection limit of the analysis was 0.26 and 0.05 µg C cm<sup>-2</sup> for OC and EC, respectively. The reported results were blank-corrected.

### 2.3.3 WSOC analysis

For the determination of the Water Soluble Organic Carbon (WSOC) concentrations, punches of 2 cm<sup>2</sup> were removed from the filter and placed in an ultrasonic bath for 45 min using 15 mL of nanopure water. The sample extract was filtered using syringe filters (PALL IC Acrodisc (PES), 0.45 mm, 13 mm). The solutions obtained after the filtration were analyzed for total organic carbon (TOC) using a Shimadzu TOC-V CSH Total Organic Carbon Analyzer. Total water soluble organic carbon is calculated by extracting inorganic water soluble carbon from total water soluble carbon (TOC = TC – IC) and is labeled as WSOC. The analytical detection limit for WSOC was 80 ppb. All data were corrected for the obtained blank levels.

### 2.3.4 Ion chromatography

Punches of 2 cm<sup>2</sup> from the quartz filters were extracted in ultrasonic bath with 10 mL of nanopure water, for 45 min, and were then filtered using syringe filters (PALL IC Acrodisc (PES), 0.45 mm, 13 mm) to remove any insoluble species. The acquired filtered solutions were analyzed by ion chromatography (IC) for the determination of

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the main ionic species concentrations (anions:  $\text{Cl}^-$ ,  $\text{Br}^-$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ ,  $\text{PO}_4^{3-}$ ,  $\text{C}_2\text{O}_4^{2-}$  and cations:  $\text{NH}_4^+$ ,  $\text{K}^+$ ,  $\text{Na}^+$ ,  $\text{Mg}^{2+}$ ,  $\text{Ca}^{2+}$ ). Anions were determined using a Dionex-500 ion chromatograph equipped with an Ion Pac AS4A-SC column and an AG4A-SC pre-column, with an ASRS-300 suppressor. Anions' separation was conducted with isocratic elution of  $\text{NaHCO}_3$  (3.4 mM)/ $\text{Na}_2\text{CO}_3$  (3.6 mM) as an eluent and a flow of 1.5 mL min<sup>-1</sup>. For the determination of cations an Ion Pac CS12A column and a CG12A guard column was used, with a CSRS-300 suppressor, under isocratic elution of 20 mM MSA (Methanesulphonic acid) at a flow rate of 1.0 mL min<sup>-1</sup>. The detection limit of the analysis was 20, 12, 40, 12 and 40 ppb for  $\text{NH}_4^+$ ,  $\text{K}^+$ ,  $\text{Na}^+$ ,  $\text{Mg}^{2+}$  and  $\text{Ca}^{2+}$ , respectively, while the corresponding detection limit for all anions ( $\text{Cl}^-$ ,  $\text{Br}^-$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ ,  $\text{PO}_4^{3-}$ ,  $\text{C}_2\text{O}_4^{2-}$ ) was 20 ppb. The reported concentrations were corrected for blanks.

### 3 Results and discussion

#### 3.1 Mass concentrations of OC – EC and observed seasonal trends

The daily concentration levels of OC range from 0.1 to 8.5  $\mu\text{g m}^{-3}$  (average:  $2.1 \pm 1.3 \mu\text{g m}^{-3}$ ), contributing on average about  $11 \pm 3\%$  to the total PM<sub>2.5</sub> mass. Accordingly, the mass concentration of EC ranges from 0.01 to 3.33  $\mu\text{g m}^{-3}$  (average:  $0.54 \pm 0.39 \mu\text{g m}^{-3}$ ), contributing  $3 \pm 1\%$  to the total PM<sub>2.5</sub>. OC and EC values from previous studies in Athens and other European sites are presented in Table 1. It should be noted that most of these studies refer to shorter time periods (4–11 months), and for different site types. Compared to OC and EC concentrations in Athens back in 2003, this study reveals levels 3–4 times lower, while later studies revealed similar levels (or slightly higher depending on the site type) in the range of 1.5–4  $\mu\text{g m}^{-3}$  for OC and 0.4–1.8  $\mu\text{g m}^{-3}$  for EC. OC concentrations from several studies in southern Europe cities (mainly western Mediterranean) during the period 2002–2006 are in the range 3–9  $\mu\text{g m}^{-3}$  while EC values are in the range 0.7–1.7  $\mu\text{g m}^{-3}$ . Concentrations in an urban

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site in Istanbul during the period 2008–2009 appear much higher compared to the values of this study. Regional background concentration levels at the Finokalia site (Crete Island, 2004–2006) are on average 1.8 and 0.27  $\mu\text{g m}^{-3}$ , for OC and EC, respectively.

Total carbon is calculated as the sum of OC and EC ( $\text{TC} = \text{OC} + \text{EC}$ ), with OC accounting for  $79 \pm 4\%$  of the total carbon, constituting the predominant carbon contributor, in accordance with findings in other European sites (e.g. Lonati et al., 2007; Sanchez de la Campa et al., 2009). However, the contribution of TC to PM in this study ( $14 \pm 3\%$ ) is lower than what is reported in the literature for other urban/urban background locations (e.g. about 19% in Sanchez de la Campa et al. (2009); 27–28% in Lonati et al. (2007); 29% in Perrone et al. (2011); 25–46% in Viana et al. (2006)).

Conversion of OC to OM (Organic Matter) is recognized as one of the most critical factors of uncertainty in mass closure calculations (Yttri et al., 2007). Turpin and Lim (2001) suggest the use of different conversion factors per site type:  $1.6 \pm 0.2$  for urban areas, 1.9–2.3 for aged aerosols, and 2.2–2.6 for biomass burning. Taking into account that our site is an urban background station also affected by long range transport (e.g. Sciare et al., 2005; Kanakidou et al., 2011; Pateraki et al., 2012; Remoundaki et al., 2013), we used the value of 2.1 ( $\text{OM} = 2.1 \cdot \text{OC}$ ). On average, OM accounts for  $24 \pm 6\%$  of  $\text{PM}_{2.5}$  which is consistent with the value (22%) provided by Remoundaki et al. (2013) for Athens, while lower to those reported in other European sites (e.g. Sillanpää et al., 2005; Sanchez de la Campa et al., 2009).

In Fig. 1 the daily concentrations of OC and EC are presented for the whole sampling period (May 2008–April 2013). It is observed that OC and EC present significant covariance for most of the period, with maxima mainly during the colder months. The seasonal variability in OC and EC is clearly demonstrated in Fig. 2, which reveals higher concentrations from November to March (the coldest period of the year), due to emissions from residential heating and low altitude temperature inversions that trap local sources of particulate OC near the surface (e.g. Yttri et al., 2007). The minimum OC concentrations appears in late spring and early autumn, when there is no domestic heating and photochemistry is slowed down compared to summer. In the case of

EC, minimum concentrations are found during the warmest months of the year, when there is less fuel combustion in Athens (period of summer vacations) and thus lower production of primary EC.

The summer vs. winter contrast is further shown in Fig. 3, on an annual basis. Winter refers to December–February period of the December year, while summer to June–August. The EC summer levels after 2008 appear slightly reduced compared to 2008, coinciding the economic recession period in Greece, which led to reduction of traffic emissions (Vrekoussis et al., 2013). In winter, there is an obvious increase in the last two years (2011–2012), attributed to the massive turn of Athens residents in wood burning for main domestic heating. The same behavior in winter is reproduced also by OC. The summer to winter ratio of EC mass concentration presents a significant declining trend from 1.21 in 2008 to 0.59 in 2012, illustrating the combined effect of summer and winter changes in EC emissions, as related to the economic crisis (see below).

### 3.2 OC/EC concentration ratios

The daily OC/EC ratios range from 0.5 to 29.3, with an average of  $4.7 \pm 1.3$  for the 5 year period. These ratios are generally comparable with results from other Mediterranean sites (see Table 1), however, on a seasonal basis the OC/EC ratios ( $[\text{OC}/\text{EC}]_{\text{winter}} = 4.6$ ,  $[\text{OC}/\text{EC}]_{\text{summer}} = 6.3$ ) are higher than those reported for other urban background sites around the world ( $[\text{OC}/\text{EC}]_{\text{winter}} = 2.4\text{--}3.5$ ,  $[\text{OC}/\text{EC}]_{\text{summer}} = 1.3\text{--}3.9$ ) as reviewed by Lonati et al. (2008). Higher ratios are expected as a result of secondary production from gas to particle conversion of VOCs during long-range transport of polluted air masses (Pio et al., 2007).

In more detail, the seasonal cycle of OC/EC ratio (Fig. 4) presents maximum values during summer (6.3), when photo oxidation reactions lead to the production of secondary aerosol (increase of OC concentrations) (e.g. Viana et al., 2007) and reduction of heating (decrease of EC concentration). On January (the coldest month of the year), the high OC/EC ratio could be attributed to the elevated OC concentrations, mainly

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due to fuel combustion for heating, that coincide with a comparatively smaller increase in the production of EC.

Overall, the high values of OC/EC ratio ( $> 2$ ) year round, indicate the clear impact of additional processes like the production of secondary organic carbon. To estimate the primary and secondary origins of carbonaceous aerosol, the correlation between the OC and EC mass concentrations is investigated (e.g. Turpin and Huntzicker, 1995). In the present study a statistically significant correlation is found between OC and EC for the whole period (not shown,  $R^2 = 0.49$ , slope = 2.33,  $n = 1365$ ,  $p < 0.001$ ), revealing that OC and EC emission rates are proportional to each other (e.g. Na et al., 2004). A comparable correlation coefficient was reported by Koulouri et al. (2008) for another Eastern Mediterranean, though remote coastal, location (Finokalia, Crete). The above suggest that OC and EC fractions are emitted by common primary sources at a regional scale.

### 3.3 Water soluble organic carbon

Water soluble organic carbon (WSOC) is considered a major component of carbonaceous aerosol as it serves as a proxy for secondary organic aerosols due to its highly oxidized and soluble nature (Park et al., 2013).

The WSOC concentration in  $PM_{2.5}$  ranged from 0.03 to  $10.6 \mu\text{g m}^{-3}$ , with an average of  $1.5 \pm 0.9 \mu\text{g m}^{-3}$ , and no significant differences between summer ( $1.7 \pm 0.8 \mu\text{g m}^{-3}$ ) and winter ( $1.6 \pm 1.2 \mu\text{g m}^{-3}$ ). These levels are similar to those reported for the remote-coastal site of Finokalia, Crete, by Bougiatioti et al. (2011), but lower than those reported for Istanbul by Theodosi et al. (2010) ( $PM_{10}$  in urban site, see Table 1). The concentration of WSOC presents a statistically significant correlation with OC (Fig. 6) ( $R^2 = 0.87$ ,  $n = 1000$ ,  $p < 0.001$ ), indicative of their common sources. Moreover, WSOC follows the seasonal trend of OC (Fig. 7) throughout the year, showing maximum values during winter and mid-summer, when WSOC is either emitted from primary combustion sources, such as traffic, and biomass burning or is produced through atmospheric processing of volatile organic species (Park et al., 2013).

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WSOC usually consists of organic compounds that derive from the atmospheric oxidation of VOCs and gas particle conversion (secondary OC), while the oxygenation of particulate OC may also occur in the particulate phase or via multiphase chemistry by reaction of strong oxidants such as ozone or hydroxy radicals. As a result, the WSOC/OC ratio decreases in urban locations and near fossil fuel sources (~ 20 %) and increases at more remote sites due to aerosol aging (Pio et al., 2007).

In this study, the WSOC accounts, on average, for about  $63 \pm 7\%$  of the OC, which is in agreement with the value (67 %) reported by Bougiatioti et al. (2011) for Finokalia, Crete, but higher than the value (42 %) reported by Lonati et al. (2007) for Milan. The above indicate the dominant role of aerosol aging processes at the regional scale. Finally, the WSOC/OC ratio shows maximum values during summer (Fig. 7), when there is intense photo oxidation that leads to the production of secondary organic species.

The water insoluble organic carbon (WIOC) has been calculated through subtraction of WSOC from the total OC concentration. On an annual basis WIOC contributes about  $39 \pm 7\%$  to OC. Although WIOC shows no significant correlation with EC on a daily basis, their seasonal cycles show covariance (Fig. 7), indicating that part of the production of WIOC can be attributed to primary sources such as fossil fuel combustion.

The concentration of non-sea salt sulfate ( $\text{nss-SO}_4^{2-}$ ) has been also estimated following the procedure described in Sciare et al. (2005) and Pio et al. (2007). The average value for the whole period is  $3.0 \pm 2.0 \mu\text{g m}^{-3}$ , which is in agreement with the findings from other studies in Athens (Theodosi et al., 2011; Remoundaki et al., 2013). WSOC is weakly correlated to  $\text{nss-SO}_4^{2-}$  ( $R^2 = 0.15$ ,  $n = 1017$ ,  $p < 0.001$ ), suggesting that secondary organic products, formed through formation pathways similar to those of sulfate, are not the dominant constituents of WSOC. Additionally, the low correlation coefficient of WSOC with  $\text{nss-K}^+$  ( $R^2 = 0.05$ ,  $n = 1023$ ,  $p < 0.001$ ) propounds limited production of WSOC from biomass burning. Finally, the higher correlation coefficient of WSOC with EC ( $R^2 = 0.35$ ,  $n = 993$ ,  $p < 0.001$ ) indicates that there is some contribution to WSOC from primary anthropogenic sources rather than biogenic. The correlation between WSOC and  $\text{nss-SO}_4^{2-}$  slightly improves during summer (summer:  $R^2 = 0.26$ ;

winter:  $R^2 = 0.03$ ); while, on the contrary, the dependence of WSOC to EC is improving in winter (summer:  $R^2 = 0.29$ ; winter:  $R^2 = 0.44$ ).

Based on the above, the primary fossil fuel and wood combustion sources appear to contribute more to the winter WSOC concentrations, whereas during summer, sources like oxidation processes of secondary products are more efficient in generating WSOC.

### 3.4 Sources of carbonaceous aerosol

The independent contribution of biomass burning and fossil fuel combustion emissions to organic matter concentrations could be quantitatively estimated through the potassium–OC and potassium–EC correlations (e.g. Satsangi et al., 2012). In general, biomass burning contributes significantly to the atmospheric concentrations of potassium ( $K^+$ ) and the latter is applied as an inorganic tracer for aerosol derived from this source (e.g. Currie et al., 1994; Pio et al., 2008; Schmidl et al., 2008), while fossil fuel combustion leads to minor production of potassium (e.g. Yang et al., 2005a, b).

To utilize  $K^+$  as a representative tracer of pure biomass burning emissions, we have applied a correction for the contribution from sea salt. Sea salt potassium ( $ss-K^+$ ) was calculated in accordance with the composition of sea water (Sciare et al., 2005; Pio et al., 2007; Bougiatioti et al., 2013) and was then subtracted from the total  $K^+$  concentration, to calculate the remaining non-sea salt potassium ( $nss-K^+$ ). The  $nss-K^+$  concentrations were ranged up to  $1.8 \mu\text{g m}^{-3}$ , with a 5 year average of  $0.17 \pm 0.19 \mu\text{g m}^{-3}$ , comparable to the values reported by Koulouri et al. (2008).

In this study, no significant correlations are calculated between  $nss-K^+$  with OC and EC (neither do they improve on a seasonal basis), indicating that during the whole period, biomass burning does not consist the main source of organic carbon and that the dominant primary source of EC is mainly fossil fuel combustion (such as traffic). The  $nss-K^+/OC$  ratio ranges between 0.03 and 2.4, with an average of  $0.11 \pm 0.04$ , while  $nss-K^+/EC$  ratios vary from 0.01 to as high as 7.4, with an average of  $0.45 \pm 0.14$ . Relatively high  $K^+/EC$  ratios for biomass burning (range: 0.21–0.46) and low ratios for

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the regional context. The increased values in Athens compared to Finokalia, can be attributed to an additional EC fraction due to traffic and a constant fractional input of OC added to the fossil fuel road transport emitted aerosol, in agreement with Theodosi et al. (2011). According to Pio et al. (2007) and Theodosi et al. (2011) the simple use of the same primary OC/EC factor to estimate secondary carbon aerosol in rural and remote European areas in the warm season, is adequate. Thus, from April to October in Athens, we estimated the contribution of primary and secondary organic aerosols in the organic carbon concentrations using Eq. (2) and applying an OC/EC<sub>pri</sub> ratio equal to 1, since traffic is the main active source during this time period (e.g. Lonati et al., 2008).

During this period, SOC ranges from 0.06 to 5.2  $\mu\text{g m}^{-3}$ , with an average of  $1.5 \pm 0.9 \mu\text{g m}^{-3}$ , constituting about  $75 \pm 6\%$  of organic carbon in Athens. It is shown that SOC is a major contributor to OC during the warm season in the area, as also indicated by Bougiatioti et al. (2013) and Lonati et al. (2007), underlying the secondary nature of fine particles (Pateraki et al., 2012). In absolute values, the average in Athens is lower than the one ( $2.7 \mu\text{g m}^{-3}$ ) reported by Sanchez de la Campa et al. (2009) in an urban background station in Spain. The monthly means of SOC (Fig. 9) in the warm period presents a maximum in summer (July/August), when there is intense photochemical activity that leads to the production of secondary organic carbon. WSOC, which is considered to be predominantly a secondary atmospheric product, also follows the seasonal variability of SOC (Fig. 9), a conclusion strengthened by the significant correlation between daily SOC and WSOC values ( $R^2 = 0.53$ , slope = 0.69,  $n = 591$ ,  $p < 0.001$ ). Contrary, the correlation between SOC and nss-SO<sub>4</sub><sup>2-</sup> was much weaker ( $R^2 = 0.22$ , slope = 1.00,  $n = 785$ ,  $p < 0.001$ ), indicating less common sources between SOC and nss-SO<sub>4</sub><sup>2-</sup>. Finally, it is worth noting the decreasing trend of SOC during the five year period of this study, accompanied by a decrease in WSOC and nss-SO<sub>4</sub><sup>2-</sup>. This downward tendency of the aforementioned concentrations can be attributed to the decrease in fuel consumption and the concurrent reduction of private car use, as a result of the economic crisis in Greece.



quantify the rapidly changing air pollution characteristics in the studied area, which is a mandate for future air quality policies.

The main conclusions arising from this study are summarized below:

- OC and EC concentrations in Athens back in 2003, were 3–4 times higher compared to this study, while later studies revealed similar (or slightly higher values depending on the site type). Average OC and EC concentration is  $2.1 \pm 1.3 \mu\text{g m}^{-3}$  and  $0.54 \pm 0.39 \mu\text{g m}^{-3}$ , respectively. OC and EC present covariance for most of the sampling period while maximum values are observed during the coldest months of the year, when there is increased fuel combustion. The aforementioned outputs reveal that OC and EC emission rates are proportional to each other, suggesting hence that OC and EC fractions are emitted by common primary sources at a regional scale.
- On a seasonal basis, summer EC is formed by predominant regional sources ( $53 \pm 12\%$  of EC), such as fossil fuel combustion (traffic), whereas during winter, the prevailing result of local sources such as fuel combustion is highlighted ( $71 \pm 8\%$  of EC). EC during summertime is decreasing through the years, since there is a simultaneous reduction in traffic due to the economic crisis in Greece. On the contrary, EC concentration during winter presents a significant increase since 2011, which can be attributed to the selection of wood as the major fuel source for domestic heating. OC concentration does not present a clear seasonal pattern.
- Total carbon for the whole sampling period (sum of OC and EC) constitutes a significant part of  $\text{PM}_{2.5}$  aerosol fraction ( $\text{TC}/\text{PM} = \sim 14\%$ ), while OC accounts for  $79 \pm 4\%$  of total carbon, constituting the predominant carbon contributor. Additionally, the estimated POM accounts for  $24 \pm 6\%$  of the collected fine aerosol.
- The total of five-year samples provides an average OC/EC ratio of  $4.7 \pm 1.3$ , while on a seasonal basis the OC/EC ratios are higher than those reported for other

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**Table 1.** PM<sub>2.5</sub> mass, OC, EC, WSOC concentrations and OC/EC ratios reported in bibliography compared to this study.

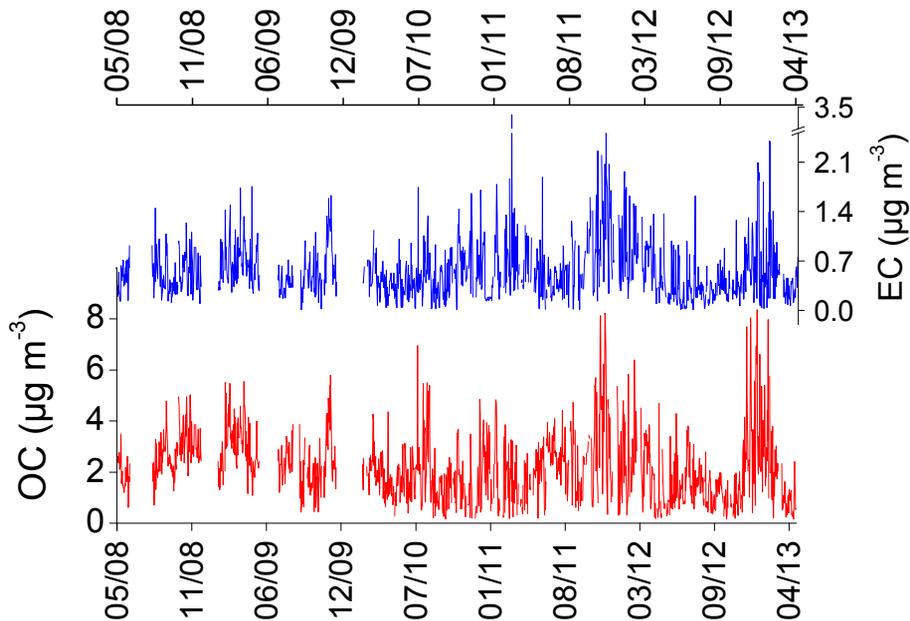
Location site*	Sampling period***	PM <sub>2.5</sub> mass ( $\mu\text{g m}^{-3}$ )	OC ( $\mu\text{g m}^{-3}$ )	EC ( $\mu\text{g m}^{-3}$ )	WSOC ( $\mu\text{g m}^{-3}$ )	OC/EC
Athens, Urb-BG, Current study	May 2008–Apr 2013	20 ± 11	2.±1.3	0.54 ± 0.39	–	4.7 ± 1.3
El Arenosillo (Spain), Urb-BG <sup>1</sup>	2 Jun 2005–30 Jun 2006	21	2.9	1.1	–	2.9
Barcelona, Urb-BG <sup>2</sup>	27 Jul–01 Sep 2004	16.4	3.7	0.8	1.6	4.6
Barcelona, Urb-BG <sup>2</sup>	16 Nov–16 Dec 2004	25.8	6.7	1.2	2.1	5.6
Duisburg, Urb-BG <sup>3</sup>	Oct–Nov 2002	14.7	–	1.32	–	–
Prague, Urb-BG <sup>3</sup>	Nov 2002–Jan 2003	29.6	–	1.69	–	–
Amsterdam, Urb-BG <sup>3</sup>	Jan–Mar 2003	25.4	–	1.37	–	–
Helsinki, Urb-BG <sup>3</sup>	Mar–May 2003	8.3	–	0.7	–	–
Barcelona, Urb-BG <sup>3</sup>	Mar–May 2003	20	–	1.52	–	–
Athens, Urb-BG <sup>3</sup>	Jun–Jul 2003	25.3	–	1.67	–	–
Salento's pen. (Lecce), Urb-BG <sup>4</sup>	Mar–Dec 2007	19 ± 9	5.6 ± 2.8	1.5 ± 12	–	–
Milan, Urb-BG <sup>5,6</sup>	Aug 2002–Dec 2003, spor	24.5	9.2 ± 7.2	1.4 ± 0.4	–	–
Madrid, Urb-BG <sup>7</sup>	Jun 2009	–	2.71	0.94	–	–
Madrid, Urb-BG <sup>7</sup>	Jan–Feb 2010	–	1.2	1.53	–	–
Athens, Urban <sup>8</sup>	Jan–Aug 2003	–	6.8	2.2	–	–
Istanbul, Urban <sup>9</sup>	Jul 2008–Jun 2009	–	6.65	2.92	–	–
Athens, Urban <sup>10</sup>	Feb–Dec 2010, spor	20	2.43	0.99	–	2.92 ± 0.6
Athens, Suburb <sup>11</sup>	29 Feb–2 May and 4 Jun–5 Aug 2008	23.8	4.1	0.48	–	–
Athens, Traf-Ind <sup>11</sup>	29 Feb–2 May and 4 Jun–5 Aug 2008	24.7	1.46	1.8	–	–
Athens, Coa-BG <sup>11</sup>	29 Feb–2 May and 4 Jun–5 Aug 2008	19.2	2.39	0.44	–	–
Thessaloniki, Urb-Ind <sup>12</sup>	Jun–Sep 2007	–	6.4	2.91	–	2.2
Thessaloniki, Urb-Ind <sup>12</sup>	Dec 2006–Mar 2007	–	8.73	2.93	–	2.98
Thessaloniki, Urb-Traf <sup>12</sup>	Jun–Sep 2007	–	7.66	2.64	–	2.9
Thessaloniki, Urb-Traf <sup>12</sup>	Dec 2006–Mar 2007	–	8.07	1.82	–	4.43
Finokalia (Crete), Remote <sup>13</sup>	Jul 2004–Feb 2007, spor	–	–	–	–	5.1
Finokalia (Crete), Remote <sup>14</sup>	Jul 2004–Jul 2006, spor	–	1.8 ± 1.4	0.27 ± 0.18	–	–
Finokalia (Crete), Remote <sup>15</sup>	10 Jul–23 Jul 2007	–	2.58 ± 1.12	0.36 ± 0.25	1.74 ± 0.86	–
Finokalia (Crete), Remote <sup>15</sup>	24 Jul–6 Aug 2007	–	3.04 ± 1.55	0.32 ± 0.17	1.93 ± 0.84	–
Akrotiri (Chania, Crete), Suburban <sup>16</sup>	Oct 2009	–	2.3 ± 1.2	0.8 ± 0.2	–	2.8 ± 1.0
El Campus (Spain), Rural <sup>1</sup>	2 Jun 2005–30 Jun 2006	21	3	0.6	–	4.7
French Riviera, Coastal <sup>17</sup>	5–29 May 2007	–	2.07–5.37	0.13–0.33	0.02–0.08	–
Montseny, Reg-BG <sup>18</sup>	2002–2007	–	2	0.2	–	–
C. de Belver, Mallorca, Sub-BG <sup>19</sup>	8 Jan 2004–29 Feb 2005, spor	20	2.9	0.5	–	–

\* Urb-BG: Urban Background, Urb-Ind: Urban Industrial, Urb-Traf: Urban Traffic, Traf-Ind: Traffic Industrial, Sub-BG: Suburban Background, Coa-BG: Coastal Background.

\*\* PM<sub>10</sub> sampling.

\*\*\* Spor: sporadically.

<sup>1</sup> Sanchez de la Campa et al. (2009) <sup>2</sup> Viana et al. (2007) <sup>3</sup> Sillanpää et al. (2006) <sup>4</sup> Perrone et al. (2011) <sup>5</sup> Lonati et al. (2008) <sup>6</sup> Lonati et al. (2007) <sup>7</sup> Mirante et al. (2014) <sup>8</sup> Grivas et al. (2012) <sup>9</sup> Theodosi et al. (2010) <sup>10</sup> Remoundaki et al. (2013) <sup>11</sup> Pateraki et al. (2012) <sup>12</sup> Terzi et al. (2010) <sup>13</sup> Bougiatioti et al. (2013)<sup>14</sup> Koulouri et al. (2008) <sup>15</sup> Bougiatioti et al. (2011) <sup>16</sup> Kopanakis et al. (2012) <sup>17</sup> Piazzola et al. (2012) <sup>18</sup> Pey et al. (2009a) <sup>19</sup> Pey et al. (2009b).



**Figure 1.** Daily OC and EC concentrations at Penteli station, Athens, for the period May 2008–April 2013.

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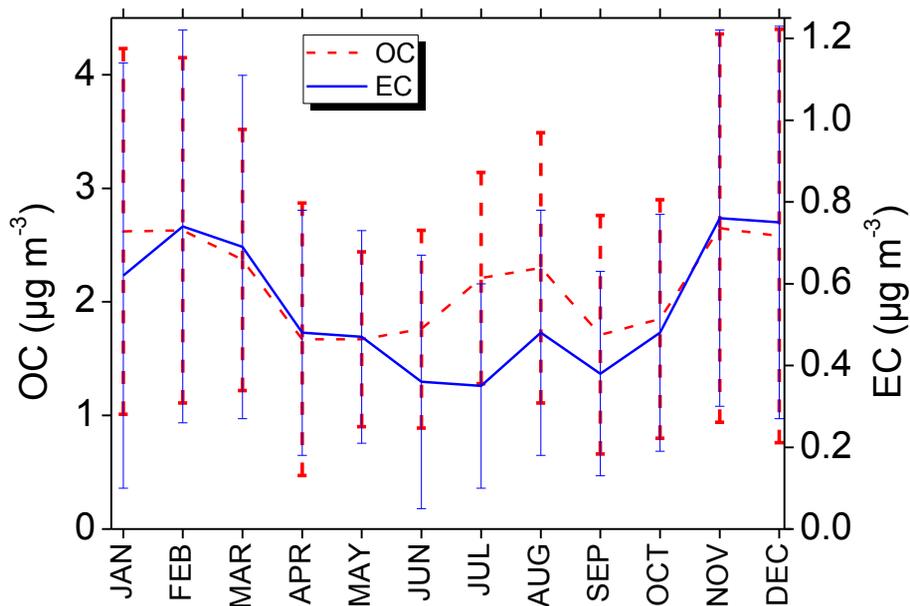
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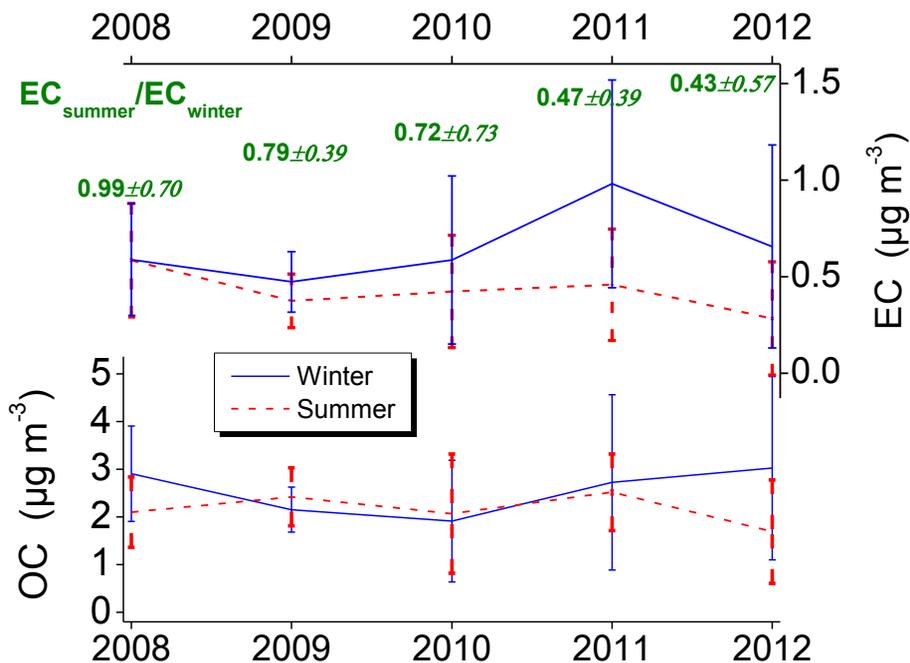


**Figure 2.** Average seasonal cycle of OC and EC concentrations calculated from daily values, for the period May 2008–April 2013.

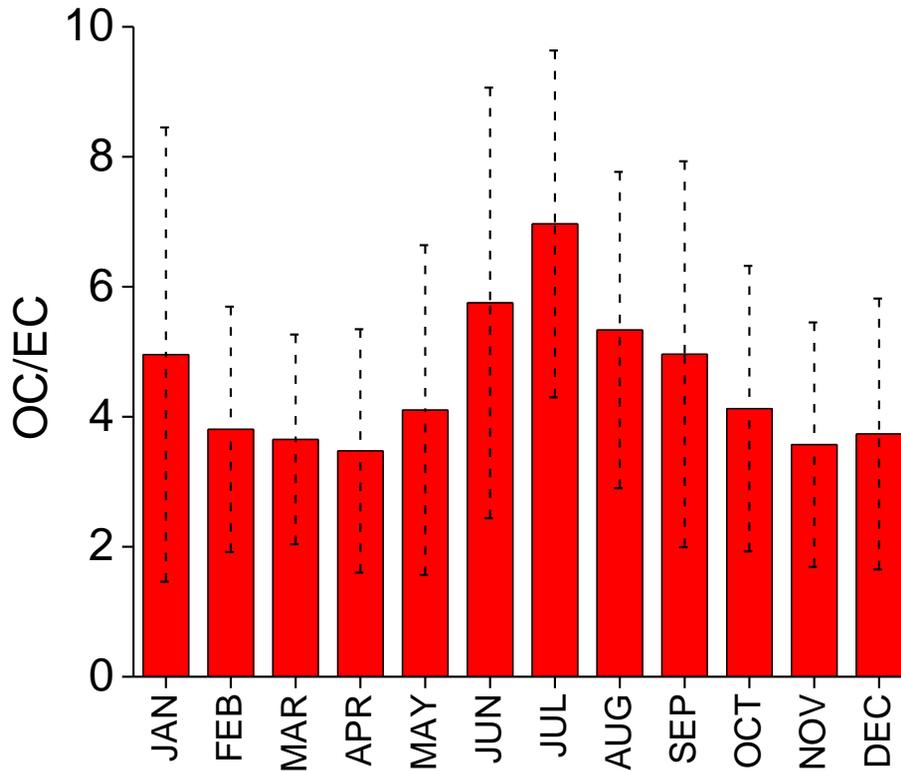
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**Figure 3.** Average winter and summer OC and EC concentrations calculated from daily values and;  $\text{EC}_{\text{summer}}$  to  $\text{EC}_{\text{winter}}$  median ratios, for the period May 2008–April 2013.



**Figure 4.** Average seasonal cycle of the OC/EC ratio calculated from daily values, for the period May 2008–April 2013.

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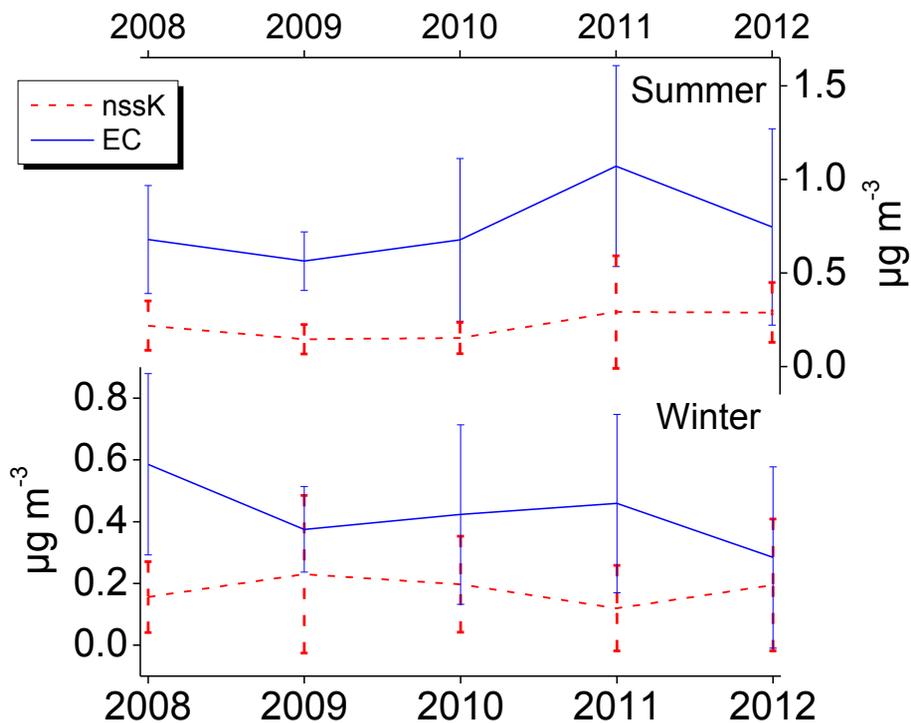
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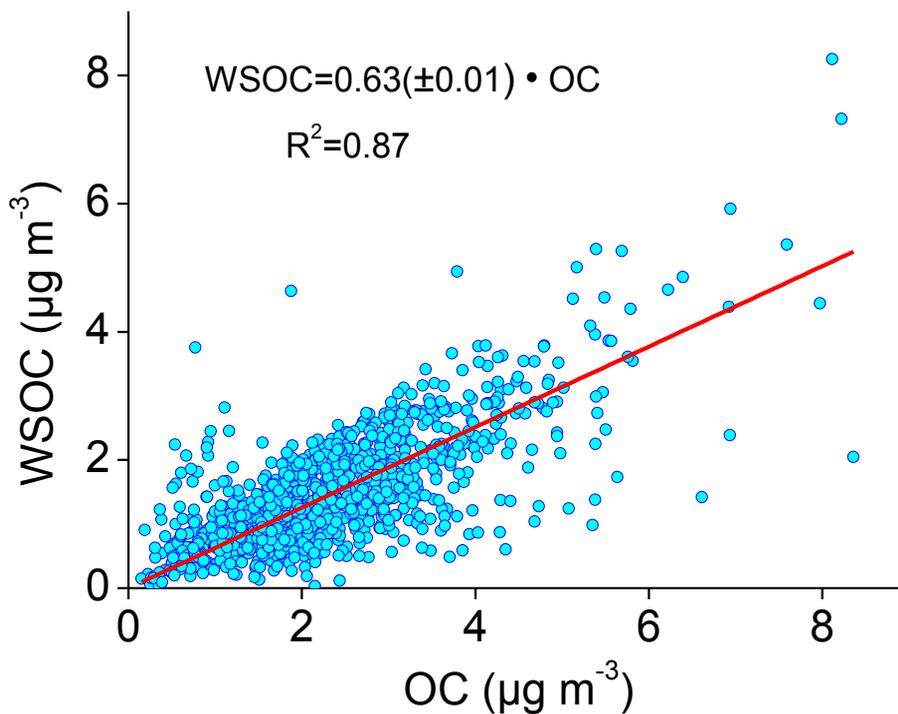
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**Figure 5.** Average summer and winter nss-K<sup>+</sup> and EC concentrations calculated from daily values, for the period May 2008–April 2013.

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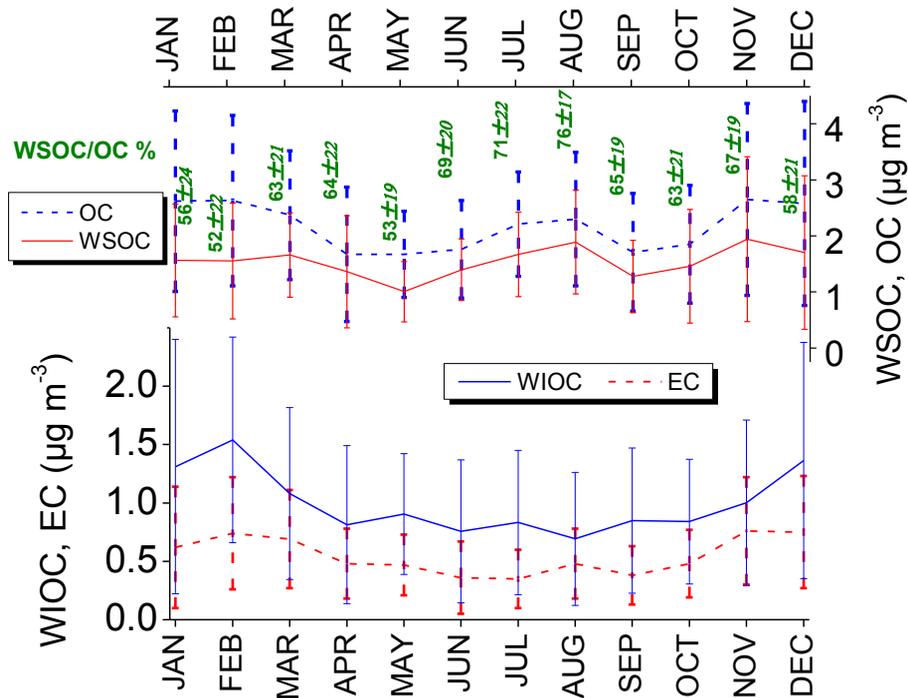


**Figure 6.** Scatter plot of the daily OC and WSOC concentrations, for the period May 2008–April 2013.

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**Figure 7.** Average seasonal cycle of OC, EC, WSOC, WIOC concentrations and the WSOC/OC ratio calculated from daily values, for the period May 2008–April 2013.

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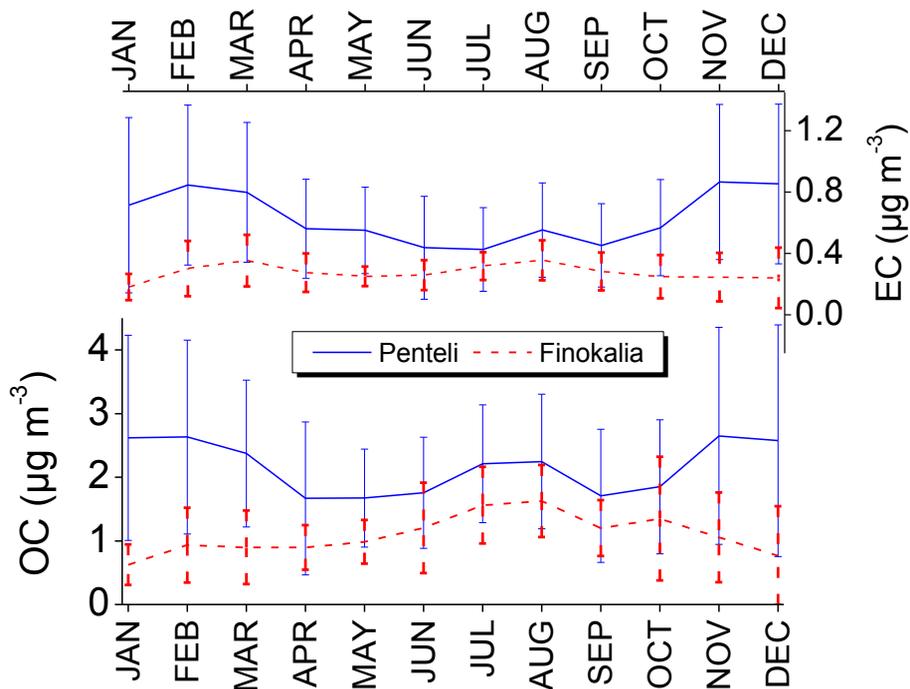
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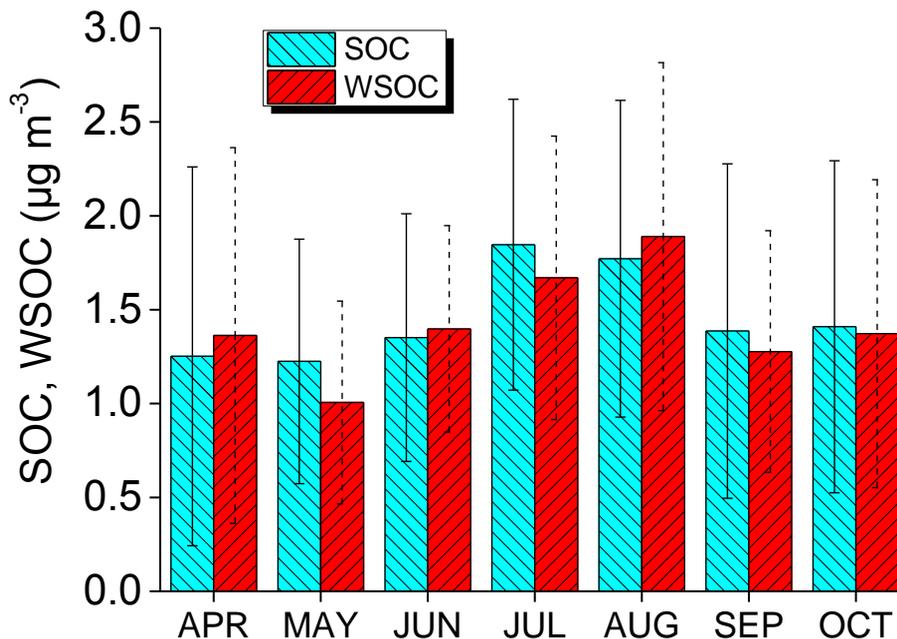
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**Figure 8.** Average seasonal cycles of OC and EC concentrations at Penteli, Athens and Finokalia, Crete Island, calculated from daily values, for the period May 2008–April 2013.

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**Figure 9.** Average monthly variability (from April to October) of WSOC and estimated SOC concentrations, calculated from daily values, for the period May 2008–April 2013.