

Aqueous phase oligomerization of methyl vinyl ketone through photooxidation

Part 1: Aging processes of oligomers

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Abstract

It has recently been established that unsaturated water soluble organic compounds (UWSOC) might efficiently form oligomers in polluted fogs and wet aerosol particles, even for weakly soluble ones like methyl vinyl ketone (MVK). The atmospheric relevance of these processes is explored by means of multiphase process model studies (in a companion paper). In the present study, we investigate the aging of these aqueous phase MVK-oligomers formed *via* $\cdot\text{OH}$ -oxidation, as well as their ability to form secondary organic aerosol (SOA) upon water evaporation. The comparison between aqueous phase composition and aerosol composition after nebulization of the corresponding solutions shows similar trends for oligomer formation and aging. The measurements reveal that oligomer aging leads to the formation of organic diacids. Quantification of the SOA mass formed after nebulization is performed, and the obtained SOA mass yields seem to depend on the spectral irradiance of the light used to initiate the photochemistry. Investigating a large range of initial MVK concentrations (0.2 – 20 mM), the results show that its $\cdot\text{OH}$ -oxidation undergoes competition between functionalization and oligomerization that is dependent on the precursor concentration. At high initial MVK concentrations ($\geq 2\text{mM}$), oligomerization prevails over functionalization, while at lower initial concentrations, oligomerization is not the major process, and functionalization dominates, resulting in small carbonyls, dicarbonyls and monoacids. The atmospheric implications of these processes are discussed.

1 **1. Introduction**

2 Organic aerosol plays an important role in many atmospheric processes and has an important
3 impact on climate and human health. Globally, about 20 % of the organic aerosol mass is
4 emitted directly (Kanakidou et al., 2005; Spracklen et al., 2011), which conversely indicates
5 the relevance of aerosol formed secondarily by transformation of organic gas phase species,
6 *i.e.* secondary organic aerosol (SOA). The most commonly studied mechanism of SOA
7 formation is the oxidation of volatile organic compounds (VOCs), which can lead to the
8 formation of less volatile species that subsequently partition into the condensed phase
9 (Donahue et al., 2011; Kanakidou et al., 2005; Kroll and Seinfeld, 2008; Hallquist et al.,
10 2009). Nevertheless, the oxidation of VOCs also results in more water soluble products that
11 readily partition into the aqueous phase (Blando and Turpin, 2000; Ervens et al., 2011;
12 Epstein et al., 2013). Due to further reactivity in the liquid phase, higher molecular weight and
13 less volatile compounds can be formed, which can remain at least in part in the condensed
14 phase upon water evaporation, thus leading to additional secondary organic aerosol formation
15 through aqueous phase reactions (aqSOA) (El Haddad et al., 2009; Carlton et al., 2009;
16 Ervens et al. 2011; Ortiz-Montalvo et al., 2012). In particular, Lee et al. (2012) observed a
17 significant enhancement of organic mass during the initial stage of oxidation of cloud water
18 organics, that they explained by functionalizing dissolved volatile organics *via* hydroxyl
19 radical ($\cdot\text{OH}$) oxidation. Aqueous phase processes can be very different from those in the gas
20 phase, thus leading to aqSOA with likely very different physical and chemical properties
21 (Ervens et al. 2011; Ortiz-Montalvo et al., 2012). These differences can explain that the
22 oxidation state of SOA formed during dry smog chamber experiments is significantly lower
23 than that of ambient SOA (Kroll and Seinfeld, 2008; Aiken et al., 2008; De Carlo et al., 2008;
24 Ng et al., 2010; Lee et al., 2012).

25 Volkamer et al. (2007) suggested that chemical processes in the aqueous phase of hygroscopic
26 particles (wet aerosol) can efficiently contribute to aqSOA mass. Besides, wet aerosol
27 provides higher precursor concentrations than in cloud and fog water droplets and reside in
28 the atmosphere over hours or days (Ervens et al., 2011), suggesting a significant role for
29 aqSOA formation in wet aerosol, in particular, in regions with high relative humidity (Carlton
30 and Turpin, 2013) and hygroscopic aerosol. Isoprene has the largest global atmospheric
31 emissions (estimated at $\sim 600 \text{ Tg yr}^{-1}$ Guenther et al., 2006) of all non-methane VOCs. Its key
32 oxidation products, *i.e.* methacrolein (MACR) and hydroperoxides (Kroll et al., 2006) are
33 known to contribute directly to the formation of SOA in the atmosphere. Methyl vinyl ketone

1 (MVK) is the other main gas-phase oxidation products of isoprene (yielding from 32 to 44 %,
2 Lee et al., 2005, Kroll et al., 2006). Unlike MACR, MVK does not lead to the formation of
3 SOA during its gas phase photooxidation (Kroll et al. 2006; Surratt et al 2006), likely because
4 of the lack of an aldehydic hydrogen which precludes the formation of acidic products such as
5 2,3-dihydroxymethacrylic acid (*i.e.* 2-methylglyceric acid: 2-MG) for further particle-phase
6 esterification reactions (Surratt et al., 2006). However, these results were obtained in smog
7 chamber experiments performed under dry conditions where aqueous phase processes were
8 excluded.

9 The photooxidation of carbonyl compounds has been studied in the aqueous phase, and their
10 ability to form oligomers and potentially aqSOA was shown (Altieri et al., 2006 and 2008;
11 Carlton et al., 2006 and 2007; Perri et al., 2009; El Haddad et al., 2009; Tan et al., 2009, 2010
12 and 2012; Zhang et al., 2010; Zhao et al., 2012; Liu et al., 2012; Ortiz-Montalvo et al., 2012;
13 Lim et al., 2013; Renard et al., 2013; Kameel et al., 2013; Kameel et al., 2014; Daumit et al.,
14 2014). In particular, Renard et al. (2013) showed that $\cdot\text{OH}$ oxidation of MVK in the aqueous
15 phase proceeds via a radical mechanism leading to oligomers which molecular masses
16 increase (up to m/z 1200) with the precursor initial concentration from 2 to 20 mM. At lower
17 precursor initial concentrations (e.g. 0.2 mM of MVK), Renard et al. (2013) observed lower
18 weight molecular compounds (up to m/z 300), and Zhang et al., (2010) observed the
19 formation of small oxidized compounds (such as acetic, formic, pyruvic acids, methylglyoxal,
20 formaldehyde), thus suggesting a competition between functionalization and oligomerization
21 at these low initial concentrations.

22 Liu et al. (2012) showed the ability of the generated oligomers to form SOA after water
23 evaporation. It is thus likely that the atmospheric impact of MVK reactivity, and especially its
24 ability to form SOA, is very different under dry and humid conditions.

25 The aim of the present study is to investigate the aging of the oligomers formed through
26 aqueous phase photooxidation of MVK. We determine the SOA chemical composition during
27 the formation and aging of the aqueous phase oligomers and we revisit the corresponding
28 SOA mass yields. A large range of initial precursor concentrations (from 0.2 to 20 mM) is
29 investigated in order to study the competition between functionalization and oligomerization.

30

1 **2. Experimental**

2 A photoreactor was used to simulate the aqueous phase photooxidation of MVK. $\cdot\text{OH}$
3 radicals were generated from H_2O_2 photolysis (Table 1). The liquid phase was analyzed using
4 a variety of analyzers for qualitative and quantitative characterization of the solution (detailed
5 in section 2.2).

6 For aerosol generation, aliquots of the solution were sampled from the photoreactor at specific
7 reaction times, then nebulized and dried prior to aerosol characterization using a scanning
8 mobility particle sizer (SMPS) and a high resolution time-of-flight aerosol mass spectrometer
9 (AMS) (Figure 1). Each experiment (aqueous phase photooxidation and aerosol generation)
10 was repeated at least once.

11 **2.1 Photoreactor**

12 The photoreactor set-up used was based on the one described by Renard et al., (2013). It was
13 a 450 cm³ Pyrex thermostated photoreactor, equipped with a 1000 Watt Xenon arc lamp
14 (LOT-Oriel, LSH 601) and a glass filter (ASTM 490 AM 0). The resulting spectral irradiance
15 into the reactor is compared to that of the sun at sea level for a 48.3° zenith angle Figure S1.

16 All experiments were performed at 25°C and started with irradiation of UHQ water (18.2 MΩ
17 cm, Millipore), then H_2O_2 (Acros, 30 %, non-stabilized) was introduced, and after 10 min of
18 H_2O_2 photolysis, MVK (Sigma Aldrich, 99 %) was introduced at time 0.

19 Tan et al. (2010) and Renard et al. (2013) have shown the important impact of initial
20 concentrations on oligomer formation. The experiments were thus carried out with various
21 MVK initial concentrations (Table 1), *i.e.*, 0.2, 0.5, 2, 5 and 20 mM (corresponding to 9.6 to
22 960 mgC L⁻¹). Considering MVK as a proxy for UWSOC, this concentration range is
23 comprised in the range of the estimated total UWSOC concentrations from fog droplets to wet
24 aerosol (Renard et al., 2013).

25 The 50 cm³ gas phase head space of the photoreactor was opened to ambient air for a few
26 seconds during each sampling. We verified in control experiments that this procedure induced
27 insignificant losses of MVK from the solution.

28 The initial H_2O_2 concentrations were chosen in order to obtain a ratio ($\frac{[\text{H}_2\text{O}_2]_0}{[\text{MVK}]_0} = 20$), in order
29 to favor $\cdot\text{OH}$ reaction with MVK rather than with H_2O_2 by more than 90 %. Under these
30 conditions, $\cdot\text{OH}$ concentrations were estimated in the range $(2 - 6) \times 10^{-14}$ M (*organic aerosol*
31 (*LV-OOA*) elementary information SI2), which falls in the range of the estimated values for

1 *OH concentrations in cloud and fog droplets (Herrmann et al., 2010; Ervens and Volkamer,
2 2010 and Arakaki et al., 2013).

3 **2.2 Aqueous phase characterization**

4 Aliquots of the solution sampled from the photoreactor were analyzed for qualitative structure
5 elucidation of the oligomers using ultra-performance liquid chromatography mass
6 spectrometry (UPLC-ESI-MS); and for quantitative studies of the concentrations of i) MVK
7 and H₂O₂ by liquid chromatography coupled to UV detection (UHPLC-UV), iii) carboxylic
8 acids by ion chromatography-mass spectrometry (IC-ESI-MS), and iv) oligomers using
9 preparative liquid chromatography associated to total organic carbon (TOC) analyses.

10 *2.2.1 UPLC-ESI-MS analyses*

11 Aliquots of the solution sampled from the photoreactor were analyzed for organic species
12 using an ultra-high performance liquid chromatographic system coupled to a time of flight
13 mass spectrometer equipped with an electrospray source and an ion mobility cell (Synapt-G2
14 HDMS, Waters). The mass spectrometer was used in its resolution mode (V-mode), up to
15 18,000 FWHM (Full width at half maximum) at m/z 400. The mass accuracy was < 5 ppm,
16 and allowed for the determination of elemental composition of organic species (Renard et al.,
17 2013 and Renard et al., 2014), using the I-FIT software. The I-FIT isotope predictive filtering
18 is a strategy to reduce the number of proposed elemental compositions using algorithms to
19 estimate the number of carbon, oxygen (or sulfur) atoms in an unknown molecule based on
20 the mass of the molecular ion and the relative intensity of the 1st and 2nd isotopes (Hobby,
21 2005).

22 All parameters used are detailed in Renard et al., (2013). Briefly, the chromatographic
23 separations were carried out on an UPLC column (Waters, HSS T3 C18, 2.1 * 100 mm – 1.8
24 µm) at 40°C. The mobile phases consisted in (A) 0.1 % formic acid (Biosolve, 99 %) in water
25 and (B) acetonitrile (Biosolve, ULC/MS). The gradient elution was performed at a flow rate
26 of 600 µL min⁻¹ using 5 to 95 % of B within 7 min and held at 95 % of B for 1.5 min. The
27 sample injection volume was 10 µL.

28 During each chromatographic run, leucine enkephalin (Waters, 2 ng µL⁻¹, C₂₈H₃₇N₅O₇) was
29 used for lock-mass correction to obtain accurate masses for each organic component eluting
30 from the column. Optimum ESI conditions were found using a 0.5 kV capillary voltage, 40 V
31 sample cone voltage, 450°C desolvation temperature, 120°C source temperature, 20 L h⁻¹
32 cone gas flow rate and 800 L h⁻¹ desolvation gas flow rate.

1 All products were detected as their protonated molecules ($[M + H]^+$) or sodium adducts ($[M +$
2 $Na]^+$) in the positive mode, and their deprotonated molecules ($[M - H]^-$) in the negative
3 mode. Data were collected from m/z 50 to 1800 in both ionization modes.

4 2.2.2 UHPLC-UV analyses

5 An ultra-high performance liquid chromatographic (UHPLC) system (ThermoScientific,
6 Accela 600 auto sampler and Accela 600 pump) coupled to a diode array detector
7 (ThermoScientific, Accela 600 PDA detector) was used to monitor the concentrations of
8 MVK and H_2O_2 sampled from the photoreactor. The chromatographic separation was
9 performed using a column (ThermoScientific, Hypersil GOLD, 100 x 2.1 mm - 1.9 μ m) at
10 40°C and a flow rate of 300 μ L min^{-1} . The mobile phase was water/acetonitrile (98:2) (v/v)
11 and the injection volume was set to 2 μ L. The spectra were recorded from 200 to 360 nm.

12 Under these conditions, H_2O_2 has a retention time of 0.5 minutes and is chromatographically
13 separated from MVK which has a retention time of 1.8 minutes. The UV spectrum of aqueous
14 H_2O_2 exponentially increases with decreasing wavelength, it becomes intense below 300 nm.
15 Aqueous solutions of MVK show an intense absorption band (K-band; $\pi \rightarrow \pi^*$ transition) that
16 peaks at 211 nm and a weak absorption band (R-band; $n \rightarrow \pi^*$ transition) that peaks at 308
17 nm. The chromatograms were monitored at 270, 229 nm and 211 nm and the peak areas were
18 found to be directly proportional to both the H_2O_2 and the MVK concentrations in the range
19 of the studied concentrations: at 211 nm for low MVK concentrations ($[MVK] \leq 2$ mM), at
20 229 nm for $[H_2O_2] < 40$ mM and $2 < [MVK] \leq 20$ mM), and at 270 nm for higher
21 concentrations of H_2O_2 (up to 400 mM).

22 2.2.3 IC-ESI-MS analyses

23 Quantification of organic acids in the solutions was performed with an ion chromatography
24 system (Dionex ICS3000) driven by Chromeleon[®] software (6.80 version), composed of a
25 gradient pump (Dionex SP-5), an autosampler (Dionex AS40), a conductivity detector
26 (Dionex, CD25) and coupled to a quadrupole mass spectrometer (Thermo Scientific Surveyor
27 MSQ) operated in the negative electrospray ionization (ESI) mode, with nitrogen gas (gas
28 flow: 6 L h^{-1} , 40 psi, temperature 500°C; Capillary voltage 3,5 kV; Sample cone voltage 75
29 V). A electrolytic suppressor (Dionex, 4 mm ASRS 300) operated in external water mode (7
30 mL min^{-1}) was placed before the conductivity cell. An additional peristaltic pump was used
31 during measurements to wash the entrance cone of the mass spectrometer with water at a flow
32 rate of 0.4 mL min^{-1} . The chromatographic separations were carried out on a column (Dionex,

1 IonPac AS11-HC, 4 x 250 mm) coupled to a guard column (Dionex, AG11-HC, 4 x 50 mm).
2 A 25 μL sample was injected automatically using a 25 μL loop injection valve. The analysis
3 was performed at 35°C, with a flow rate set at 0.8 mL min^{-1} . Eluent A (Ultra High Quality
4 water) and eluent B (100 mM NaOH) were flushed with purified helium gas for 30 min and
5 kept under nitrogen atmosphere during the procedure. Separation was carried out using the
6 following gradient (min, B %): 0, 1 %; 12, 5 %; 30, 19 %; 40, 40 %, 50, 1 %. The analytes
7 were monitored using the selected ion-monitoring (SIM) mode, and signal areas (counts min^{-1})
8 of each peak were used for quantification.

9 2.2.4 TOC analyses

10 TOC measurements were associated to preparative liquid chromatography to separate the
11 oligomers from the small and/or volatile reactants and reaction products in the liquid samples,
12 in order to measure the oligomer mass yields in experiment A (see section 3.2.3). A total
13 organic carbon / total nitrogen (TOC/TN) analyzer (Analytik Jena, N/C2100S) with the non-
14 purgeable organic carbon (NPOC) method was used to quantify the produced oligomers in our
15 liquid samples.

16 The NPOC method consists in pre-purging samples with oxygen and pre-acidifying (at pH=2
17 with HCl) to remove the inorganic carbon and purgeable organic carbon. TOC is measured by
18 injecting the sample into a heated combustion tube (800°C) with an oxidation catalyst. The
19 CO₂ produced is measured by a non-dispersive infrared (NDIR) gas analyzer. TN is measured
20 in parallel using chemiluminescence detection (CLD).

21 2.3 Particle generation and characterization

22 For aerosol generation, 35 mL of the solution was sampled at specific reaction times (Table
23 1), and nebulized using an atomizer (TSI, 3079) with a flow rate of 3.5 L min^{-1} (Figure 1).
24 The generated droplet flow was led through a silica gel diffusion dryer and diluted with
25 filtered ambient air (at 5 L min^{-1} , using a HEPA capsule filter). A small fraction of the sample
26 ($\approx 0.4 \text{ L min}^{-1}$) was passed through a Nafion dryer (Permapure, MD-110), before entering a
27 small 100 mL glass mixing chamber and the on-line analytical devices. The obtained relative
28 humidity was constant during all experiments at ca. 15 % measured at the entrance of the
29 AMS (Figure 1). The nebulization time for each sample was 30 min and, to ensure constant
30 and reproducible aerosol generation, only the last 15 min of nebulization were employed for
31 data analysis. To avoid memory effects, before each nebulization experiment, the system was
32 flushed by nebulizing UHQ water for 30 min.

1 The number size distribution was measured using a scanning mobility particle sizer (SMPS),
2 (Grimm, SMPS+C) consisting of a differential mobility analyzer (L-DMA) with a
3 condensation particle counter (Grimm, CPC, 5.403). The analyzed particle size ranged from
4 11 to 1083 nm (scanned within 6 min and 43 s).

5 A high resolution time-of-flight aerosol mass spectrometer was used to measure the bulk
6 chemical composition of the non-refractory submicron particulate matter (De Carlo et al.
7 2006; Canagaratna et al., 2007). The instrument was used under standard conditions
8 (vaporizer at 600°C and electron ionization at 70 eV), in the high sensitivity V-mode (up to
9 2000 at m/z 200). Each measurement point was averaged for 2 min and 40 s (MS- and PToF-
10 cycle, 40 s each, 2 cycles per run). Mass spectra of filtered air (using a HEPA capsule filter)
11 were taken prior each series of nebulizing experiments in order to adjust the m/z 44 entry of
12 the fragmentation table due to gas phase CO₂.

13 The standard fragmentation table with the corrected air fragment column for our carrier gas
14 and the default values of relative ionization efficiency were used in the AMS data analysis
15 (Squirrel 1.51H and the software PIKA 1.10H).

16 **3. Results**

17 We investigated MVK-[•]OH oxidation in the aqueous phase and subsequent SOA formation
18 upon water evaporation, by nebulizing the liquid solution. During this process, oligomer
19 formation and aging was monitored, and a qualitative and a quantitative study of the SOA
20 formed after nebulization was performed. For these investigations, a large range of initial
21 MVK concentrations (0.2 – 20 mM) was explored.

23 **3.1 Evidence for oligomer formation and aging**

24 During MVK-[•]OH oxidation, the aqueous phase composition was monitored and compared to
25 the composition of the corresponding nebulized solutions.

26 *3.1.1 Aqueous phase analyses*

27 For each experiment, the solution was directly monitored using UPLC-ESI-MS and UHPLC-
28 UV for reaction times up to 150 min (Table 1). This time was higher than the complete
29 consumption of MVK in order to study the formation of oligomers and their aging processes,
30 as illustrated in Figure 2.

1 During experiment B (*i.e.* $[MVK]_0 = 5\text{mM}$), after 5 minutes of reaction, no significant
2 formation of high molecular weight compounds (HMWC) was observed (Figure 2b), whereas
3 after 10 minutes of reaction, mass spectra show that oligomer systems were formed on the
4 whole range of the investigated m/z (50-1200), with a regular pattern of 70.042 amu, which
5 corresponds to the exact mass of MVK. At 50 minutes of reaction, the maximum of
6 oligomerization was reached (Figure 2c). At this time, we observed several series of MVK-
7 oligomers, corresponding to several initiator radicals identified by Renard et al. (2013) under
8 the same conditions. As an example, the molecular structure of the most intense series is
9 given in Figure 2c and is highlighted in red in the mass spectrum. At that time, 90 % of MVK
10 was consumed. Finally, the intensities of all the oligomer series decreased simultaneously for
11 all masses with no change in the oligomer pattern up to 90 minutes. From this reaction time,
12 the mass spectra show a collapse of the regular pattern in both negative (Figure 2d, 150 min)
13 and positive modes, certainly corresponding to a drastic aging process in which oligomers
14 formed smaller molecules. This hypothesis is confirmed by a more global approach, using the
15 SMPS and the AMS analysis of the SOA formed after nebulization of the solutions.

16 3.1.2 Aerosol composition of SOA generated after nebulization of the solutions

17 Under similar conditions, we verified in a previous study (Liu et al., 2012) that nebulization
18 of the reacted solutions and subsequent aerosol particle drying processes induced negligible
19 chemical transformations of the oligomers compared to the aqueous phase composition. It was
20 thus meaningful to compare the compositions of aqueous phase and SOA after nebulization.

21 The AMS mass spectra (Figure 3) show two dominant fragments, at m/z 43 and m/z 44
22 (corresponding to $\text{C}_2\text{H}_3\text{O}^+$ and CO_2^+ fragments). The time profiles of the AMS total organic
23 mass and both fragments clearly show (Figure 3) the same three-step kinetic behavior as the
24 one described above in section 3.1.1 for the corresponding solutions (Figure 2). Until 10 min
25 of reaction, the intensity of the AMS total mass remains low (Figure 3a) and the mass
26 spectrum at 5 min (Figure 3b) is not significantly different from the one obtained by
27 nebulizing an aqueous solution containing the reactants before reaction, with m/z fragments
28 lower than 100. Then, the total mass increases to reach a maximum at 50 min (Figure 3c), an
29 order of magnitude higher than at 5 min. The mass spectrum is dominated by the m/z 43
30 fragment (Figure 3c). This observation is likely due to fragmentation (by electronic impact of
31 the AMS, Ng et al., 2010) of oligomers containing repetitive carbonyl functions such as those
32 identified in the aqueous phase (see the example of the molecular structure in Figure 2c, and
33 in Renard et al., 2013). Finally, the intensity of both the total organic mass and that of m/z 43

1 fragment decrease, the one of m/z 44 increases, and they both dominate the AMS mass
2 spectrum with the same intensity at the end of our investigation (150 min, Figure 3d).
3 Furthermore, comparing the AMS mass spectra between 50 and 150 min at higher masses
4 (m/z 100 – 200) (Figure 4), it is clear that at 50 min of reaction, the mass spectrum contains
5 more fragments in this range, than at 150 min. It is thus likely that the oligomers are being
6 significantly photooxidized through a fragmentation mechanism that forms smaller acidic
7 compounds (as observed by Aljawhary et al. (2013) for different precursors), and it confirms
8 the oligomer aging process suggested in Figure 2. After 50 min, oligomer fragmentation
9 prevails over oligomer formation.

10 For the quantitative study (section 3.2), we used the data provided by the SMPS analysis.
11 Note that the overall collection efficiencies (CE) of the AMS in our experiments varied from
12 0.07 to 0.21 (related to the SMPS signal). These low CE values (compared to chamber studies
13 or ambient aerosols) can be due to particle bounce at the vaporizer surface before
14 volatilization and to the shape and size-dependent transmission of the aerodynamic lens. As a
15 result, the studied compounds did not volatilize sufficiently fast at standard AMS vaporizer
16 temperatures to be fully detected (Liu et al., 2007; Docherty et al., 2013; Miyakawa et al.,
17 2013). In addition to these effects, it is possible that our low CE values were also due to the
18 particle size range (50 nm - 150 nm mass distribution), as the lowest part of this size range
19 corresponds to the region where the AMS transmission curve varies greatly (Liu et al., 2007).
20 This effect is confirmed by the fact that our lowest values for CE (0.07) were obtained for the
21 lowest MVK initial concentrations (0.2 – 2 mM) where the smallest particles were formed (50
22 nm mass distribution).

23 **3.2 Quantitative study of SOA**

24 The oligomer formation, mass yields and aging were observed by the quantitative analyses
25 performed with the SMPS measurements of the nebulized solutions.

26 *3.2.1 SOA mass*

27 For experiment B, Figure 5a shows a continuously increasing number size distribution with
28 reaction time from 5 to 150 min, with an increasing mode size during the two first kinetic
29 steps (up to 50 min), and a decreasing mode size during the third one, which corresponds to
30 oligomer aging. In order to determine the particle mass concentrations, we used the method
31 described by Kuwata et al. (2012) (eq. 1) to determine the density (ρ_{org}) of the SOA generated
32 in our system at each reaction time t ,

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$$\rho_{\text{org, } t} \text{ (g cm}^{-3}\text{)} = \frac{12+1 \times (\text{H/C})_t + 16 \times (\text{O/C})_t}{7+5 \times (\text{H/C})_t + 4.15 \times (\text{O/C})_t} \quad (\text{eq. 1})$$

where $(\text{O/C})_t$ and $(\text{H/C})_t$ are elemental ratios at reaction time t , as determined by the AMS analysis of the SOA formed in our system. These ratios extend to the same ranges as those used by Kuwata et al. (2012), and the resulting particle densities are reported in Table 2 and Table 3. In particular, Table 2 shows a substantial change in the H/C (decrease) and O/C (increase) after 50 min of reaction (t_{max} , for which the maximum SOA mass is reached) for experiment B, denoting the oligomer aging and inducing an increase of the aerosol density.

Using these particle densities, the total mass concentrations were determined, and the time evolution of the resulting distribution particle mass concentrations is shown in Figure 5b for experiment B. The blank signal was determined prior to each individual experiment by nebulizing pure water samples and was subtracted in the results for the mass calculation. At the initial reaction time (0 min), the particle size distribution was determined by nebulizing an aqueous mixture of the reactants (using experiment B concentrations), it showed a mass concentration ($11.0 \pm 1.4 \mu\text{g m}^{-3}$) not statistically different from the one obtained by nebulizing pure water (assuming a density of 1.1 g cm^{-3}). This confirms that the reactants are too volatile to form substantial amounts of organic aerosol by nebulization of the solution prior to reaction.

Confirming the UPLC-ESI-MS aqueous phase analyses and the AMS results, a three-step kinetic behavior is also observed on the SMPS total mass concentrations (Figure 5b and Figure 6). A slow increase is observed during the first step (0 – 10 min). Then oligomerization takes place corresponding to a fast increase of the SMPS mass, until 50 min. Finally, after this maximum of oligomerization, a significant decrease of the SMPS mass is observed. This decrease may be related to the decrease in the particle size (Figure 5a), which can be due to the decrease of the oligomer size, by fragmentation of the oligomers. It is thus likely that the oligomer aging forms more volatile compounds that the SMPS does not measure. The high correlation between the total aerosol mass concentration and the consumed MVK observed in Figure 6 from 0 to 50 min, allows for the determination of the SOA mass yield, as discussed in section 3.2.3.

1 3.2.2 Influence of initial MVK concentrations

2 The influence of the initial aqueous phase concentration of MVK on the SOA formation was
3 investigated over a wide range, *i.e.* from 0.2 to 20 mM (Table 1). Not surprisingly, Figure 7
4 shows that the total aerosol mass concentration increases with increasing initial MVK
5 concentration. This observation is in very good agreement with the influence of MVK initial
6 concentration on the oligomerization process observed in the aqueous phase by Renard et al.
7 (2013). For experiments D and E, corresponding to the lowest initial MVK concentrations, the
8 SMPS and AMS signals were low, and they could be influenced by water impurities, whereas
9 no such influence was observed for experiments A, B and C. This is why the signal obtained
10 from the blank experiments was subtracted only for experiments D and E in Figure 7.
11 Moreover, Figure 7 clearly shows a different kinetic behavior of the SOA mass concentration
12 from the lowest initial concentration experiments (D and E), compared to the three highest
13 ones (experiments A, B and C). For experiments A, B and C, the SOA mass concentration
14 increases rapidly, reaches a maximum, and then decreases, while for experiments D and E, the
15 signal slowly increases and does not reach a maximum. This particular evolution may be due
16 to different chemical mechanisms occurring at different initial concentrations. We
17 hypothesized the predominance of oligomerization at 2 mM initial concentration and above,
18 this is further discussed in section 4.

19 The continuous increase of the particle number (shown in Figure 5a for experiment B) with
20 reaction time was observed for all initial concentrations (experiments A to E), whereas the
21 decrease of the size mode (in the number size distributions, after t_{max}) was observed for the
22 three highest initial concentrations only (experiments A, B and C) and not for experiments D
23 and E, *i.e.* only during oligomer aging.

24 Furthermore, studies on relevant mixtures of UWSOC are in progress and could contribute to
25 this chemistry.

26 3.2.3 SOA mass yields

27 The SOA mass yields, Y_t , were calculated at each reaction time step t from eq. 2. $Y_t = \frac{[SOA]_t}{\Delta[MVK]_t}$
28 (eq. 2)

29 Where $\Delta[MVK]_t$ is the consumed [MVK] in mg L⁻¹ at reaction time t ; and $[SOA]_t$ is the
30 formed SOA mass at reaction time t , in mg per L of evaporated water. This term takes into
31 account the SOA mass (M_{SMPS}) measured by the SMPS at time t (in $\mu\text{g m}^{-3}$), the atomizer flow

1 ($F_{atomizer}$ in $L m^{-3}$), the dilution (f_{dil}), and the transmission efficiency in our nebulizing system
2 (T_{eff} in %) (see Table S1).

$$3 \quad [SOA]_t = \frac{M_{SMPs} \times T_{eff}}{F_{atomizer} \times f_{dil} \times 1000} \quad (\text{eq. 3})$$

4 The yields obtained at t_{max} for experiments A, B and C are shown in Table 3. Although the
5 total SOA mass (at t_{max}) increases linearly with the initial concentration for these three
6 experiments, the yields are statistically identical as well as their H/C and O/C ratios. Due to
7 the very large uncertainties affected to our yield determinations (see below), it is not possible
8 to use these data (Table 2 and 3) to provide any interpretation on the possible effect of initial
9 concentrations on the yields. In contrast, the O/C and H/C ratios clearly show statistically
10 stable values when the total particle mass increases from 100 to 900 $\mu g/m^3$ (Table 3). It is thus
11 likely that the total mass loading does not influence the relative oxygenation of the SOA
12 produced (at t_{max}) under our experimental conditions.

13 Although the particle mass loadings (M_{SMPs}) were accurately measured, our yield
14 determinations were affected by large uncertainties due to the estimation of the transmission
15 efficiency in our nebulizing system (see supplementary information 1). In order to confirm
16 these yields' values, another method was tested for experiment A at 90 min of reaction (*i.e.*
17 close to t_{max}). Preparative chromatography was performed using UPLC, where small
18 molecules were separated from the oligomers using a divert valve (at retention times lower
19 than 2 min.). The solution containing oligomers was accumulated, concentrated and analyzed
20 using a TOC analyzer. From the carbon mass, we deduced the total mass using the H/C and
21 O/C ratios given by the AMS. The yield was then directly calculated from the total mass of
22 sample (in $mg L^{-1}$) divided by the mass of consumed MVK at the same reaction time. A yield
23 of 59 ± 5 % (in mass) was obtained with this method at 90 min of reaction, thus statistically
24 similar from the one obtained by the nebulizing method (70 ± 50 %) at t_{max} .

25 These yields are significantly higher than those obtained by Liu et al. (2012) who obtained
26 yields up to 9.9 % under similar experimental conditions as ours. It is important to note that
27 these values were obtained assuming all the particle densities were $1 g cm^{-3}$ in Liu et al.
28 (2012), and also the transmission efficiency of the nebulizing system was calibrated with
29 NaCl solutions. However, it is likely that succinic acid or ammonium nitrate are more
30 adequate for the calibration, and we show (in the supplementary information 1) that the
31 transmission efficiency of NaCl solutions are significantly different from the two other
32 solutions. The nebulizing system was slightly different, with a teflon bag in Liu et al. (2012)
33 that could enable i) larger amounts of wall losses for organic particles as compared to the

1 system presented here; but ii) longer particle residence times, leaving more time for gas-
2 particle equilibrium than in our system. However, the control experiment using preparative
3 chromatography confirms the high yield value obtained here, independent on the nebulizing
4 system and its calibration. The different yields obtained here as compared to the study by Liu
5 et al. (2012) may be due to the different irradiation Xe lamp used: 300 W (with a pyrex filter)
6 in Liu et al., 2012, and 1000 W (with a ASTM 490 AM 0 filter) in the present study. The
7 influence of the lamp spectra on SOA mass yields of other systems (gas phase photooxidation
8 of biogenic and anthropogenic precursors) have been previously observed in atmospheric
9 simulation chambers (Bregonzio-Rozier et al., 2014). We verified, using a spectroradiometer
10 (SR-501, LOT-Oriel), that the spectral irradiance of the 300W and the 1000W Xe lamps at λ
11 ≥ 400 nm represent respectively half and twice the solar irradiance intensity at sea level, for a
12 48.3° zenith angle (Figure S1). Due to the high variability of the irradiance in the atmosphere
13 at $\lambda \geq 400$ nm (as shown by the Tropospheric Ultraviolet and Visible Radiation Model
14 http://cprm.acd.ucar.edu/Models/TUV/Interactive_TUV/), both lamps can be seen as
15 representative of the natural irradiance in this wavelength range. However at 300 nm, the
16 spectral irradiance of the 1000W Xe lamp is 7 and 9 times higher than that of the direct solar
17 irradiance (for a 48.3° zenith angle) and the 300W Xe lamp respectively (Figure S1). This
18 part of the spectrum is essential for photochemistry, and may induce different photochemical
19 processes: we verified that we observed the same series of oligomers as in Liu et al. (2012),
20 but with different relative intensities (Figure S2). The different spectral irradiance of the lights
21 used at 300 nm may be the reason for the different yields obtained, but it needs to be
22 confirmed by a thorough study of the influence of the spectral irradiance (in the UV) on the
23 oligomer mass yields.

24 It is interesting to note that the yields and densities obtained in the present study are in the
25 same range as those of a similar study with a different precursor, *i.e.* glycolaldehyde and a
26 different irradiation system even more intense in the UV (*i.e.* a 254-nm mercury lamp) (Ortiz-
27 Montalvo et al., 2012). They reported aqSOA yields for oxidation products of glycolaldehyde
28 (1 mM) which decrease gradually with reaction time from about 120% to 50%; while the
29 calculated densities increase from 1.3 to 1.6 g cm⁻³.

30 **4. Discussion**

31 This section discusses the results obtained on the aging of the reaction products of MVK-[•]OH
32 experiments and the related oxidation processes in two ways. First, the competition between

1 functionalization and oligomerization and the subsequent aging according to the initial
2 concentration of MVK is discussed; and second, a Van Krevelen diagram (i.e, H/C vs O/C
3 ratios) of the obtained SOA is presented and compared to previous studies.

4 **4.1 MVK functionalization versus oligomerization: influence of initial concentration**

5 In order to compare our work with the aging of organic aerosol in the atmosphere compiled
6 by Ng et al. (2010), we used the AMS data, and in particular, the two dominant fragments,
7 m/z 44 (CO_2^+) and m/z 43 (mostly $\text{C}_2\text{H}_3\text{O}^+$), observed in our study. More precisely, we
8 focused on the ratio of m/z 44 and m/z 43 to total organic aerosol, f_{44} and f_{43} , respectively. In
9 Ng et al. (2010), low volatility oxygenated organic aerosol (LV-OOA) has higher f_{44} than
10 semi-volatile oxygenated organic aerosol (SV-OOA) which in turn has higher f_{43} values.
11 Despite the very different conditions, it is interesting to compare our aging data with those
12 compiled from field studies and other aqueous phase experiments. Figure 8 compares our f_{44}
13 versus f_{43} values to those provided in the compilation by Ng et al. (2010) as well as the
14 nebulization data by Lee et al., (2011a). It is clear from this figure that the SOA composition
15 and its evolution highly depend on the initial MVK concentration. For the three highest initial
16 concentrations (experiments A, B and C), oligomerization takes place with the formation of
17 oligomers containing repetitive carbonyl functions such as those identified in the aqueous
18 phase (Figure 2c), inducing an important increase of f_{43} and a simultaneous decrease of f_{44} ,
19 roughly up to t_{max} . After t_{max} , most of the initial MVK is consumed, slowing down the
20 oligomerization process, and an important decrease of f_{43} and a simultaneous increase of f_{44} is
21 observed, likely due to oligomer aging, as detailed (for experiment B) in section 3.1. At lower
22 initial MVK concentrations (experiments D and E), oligomerization seems much less
23 important and oxidation is the dominant process, as evidenced by the continuous increase of
24 f_{44} . It can thus be suggested that, at these lower initial concentrations, functionalization
25 dominates over oligomerization, and the aerosol is mainly composed of low-volatility organic
26 acid and not of MVK-oligomers.

27 This observation is strengthened by the comparison of our results with those of previous
28 studies. Zhang et al., (2010) performed aqueous phase $\cdot\text{OH}$ oxidation of MVK (0.2 mM initial
29 concentration), and observed the formation of functionalization products, *i.e.* formaldehyde,
30 glyoxal, methylglyoxal, pyruvic, oxalic, formic, acetic, and malonic acids. Furthermore, for
31 experiments D and E ($[\text{MVK}]_0 \leq 0.5$ mM), our f_{44} - f_{43} plots are similar to those obtained using
32 a similar set-up, starting from pinonic acid, glyoxal and glyoxylic acid at similar and higher
33 initial concentrations (Lee et al., 2011a; Lee et al., 2011b). In particular, starting at 3 mM of

1 glyoxal, Lee et al., 2011b obtained similar f_{44} - f_{43} plots as our experiments performed at much
2 lower initial concentrations ($[MVK]_0 \leq 0.5$ mM), thus showing that the concentration is not
3 the only important parameter in oligomerization processes, but the chemical nature of the
4 precursor is also fundamental.

5 MVK oligomerization occurs *via* saturation of the vinyl group (Renard et al., 2013). The
6 resulting radical monomer is stabilized by the resonance effect with the adjacent carbonyl
7 group and this stabilization decreases the enthalpy of polymerization (Odian 2004) and hence
8 facilitates the oligomerization in the aqueous phase compared to other molecules.

9

10 **4.2 Oligomer aging processes**

11 The Van Krevelen diagram (Figure 9) shows a significant increase of O/C and a significant
12 decrease of H/C with reaction time after t_{max} . When oligomerization is the dominating
13 process, almost no changes are observed in the van Krevelen diagram: the H/C and O/C
14 values are confined in a restricted circle until t_{max} . The atomic ratios for H/C and O/C of MVK
15 and the oligomers (with a degree of polymerization (DP) of 5) identified by Renard et al.
16 (2013) are also reported in this diagram. MVK, oligomers and the nebulized solutions until
17 t_{max} are confined in a circle that highlights the similarity of their structures.

18 After that time, the values of O/C (H/C respectively) increase (decrease respectively) out of
19 the circle, thus denoting an oligomer aging process. Changes in functionality of organic
20 aerosol are traced in this diagram along a line, which slope is -0.6. A very similar slope value
21 (-0.5) was interpreted by Ng et al. (2011) as a COOH group addition to the site of a C-C bond
22 cleavage, thus suggesting that the oligomer aging process proceeds *via* fragmentation. This is
23 also suggested by the time evolution of the particle number size distributions (Figure 5a: see
24 section 3.2.1). While a continuous increase of the particle number with reaction time was
25 observed for all initial concentrations (experiments A to E), a significant decrease of the size
26 mode was observed after t_{max} for the highest concentrations only (experiments A, B and C),
27 *i.e.* during oligomer aging. This size mode decrease was also correlated with a decreasing
28 total mass (Figure 5b). These observations, added to the fact that f_{44} increases during oligomer
29 aging, indicate that the oligomer aging proceeds *via* fragmentation processes that generate
30 smaller (or more volatile) and more acidic compounds.

31 The formation of carboxylic acids in the aqueous phase was monitored during the course of
32 the reaction. In good agreement with Zhang et al. (2010), small volatile monocarboxylic acids
33 such as acetic, formic and pyruvic acids were formed as primary reaction products from MVK

1 reactivity. We further observed the formation of diacids as secondary or tertiary reaction
2 products, such as oxalic, malonic, succinic (Figure 10) malic and tartaric acids (not
3 quantified). Finally, the formation of these diacids started at t_{\max} , and was correlated to the
4 increase of the AMS m/z 44 (CO_2^+) signal observed from the nebulized solutions. It is thus
5 likely that the oligomer aging proceed *via* fragmentation (by $\cdot\text{OH}$ oxidation and/or
6 photolysis), leading to the formation of smaller partially oxidized products (*i.e.* hydroxyacids
7 or ketoacids such as those identified by Jaoui et al., 2006), which, in turn are oxidized into
8 stable diacids.

9 **5. Atmospheric implications**

10 Considering the results obtained here on oligomer formation and aging from MVK at varying
11 initial concentrations together with those obtained by previous studies on the identification of
12 the low-molecular-weight compounds products of the reaction (Zhang et al., 2010), a general
13 scheme of the potential atmospheric fate of MVK in the aqueous phase is shown on Figure 11.
14 MVK $\cdot\text{OH}$ -oxidation undergoes kinetic competition between functionalization and
15 oligomerization, depending on the precursor initial concentration. At 2 mM of MVK and
16 above this concentration, oligomerization dominates over functionalization. At these
17 concentrations, $\cdot\text{OH}$ -oxidation of MVK forms oligomers that are SV-OOA, with low O/C
18 (lower than 0.50) and high f_{43} . Oligomers are then fragmented, *via* unidentified intermediates
19 that have the properties of LV-OOA (with increasing O/C and decreasing H/C, Figure 9)
20 which then form organic diacids. For lower initial MVK concentrations (< 2 mM),
21 oligomerization is not the major process, and functionalization dominates, ending into small
22 carbonyls, dicarbonyls and acids that were identified by Zhang et al., (2010) (Figure 11).

23 Among the atmospherically relevant alkenoic alcohols, acids, ketones and aldehydes (*i.e.*
24 UWSOC), although MVK is one of the most abundant species, it is one of the most volatile
25 ($P_{\text{sat}} = 10^{-5} - 0.1$ atm, with $P_{\text{sat}(\text{MVK})} = 0.1$ atm at 25°C ; Asher and Pankow 2006), and one of
26 the least soluble compounds ($K_{\text{H}} = 1 - 10^3$ M atm $^{-1}$, with $K_{\text{H}(\text{MVK})} = 41$ M atm $^{-1}$ at 25°C ; Iraci
27 et al., 1999). The atmospheric impacts of the processes shown here should thus be very
28 limited for MVK alone. However, the oligomerization mechanism undergone by MVK occurs
29 *via* saturation of the vinyl group (Renard et al., 2013), and the resulting radical monomer is
30 stabilized by the resonance effect with the adjacent carbonyl group thus decreasing the
31 enthalpy of polymerization facilitating the oligomerization in the aqueous phase compared to
32 other molecules. More generally, conjugation of the C=C with substituents such as the

1 benzene ring (styrene and α -methylstyrene), and alkene double bond (butadiene and isoprene),
2 the carbonyl linkage (acrylic acid, methyl acrylate, methyl methacrylate), and the nitrile group
3 (acrylonitrile) similarly leads to stabilization of the monomer and decreases enthalpies of
4 polymerization (O'dian 2004). It is thus likely that a large number of atmospherically relevant
5 molecules can follow the same process either in the bulk or at the wet aerosol interface
6 (Kameel et al., 2013; Kameel et al., 2014). In this context, our results suggest that this class of
7 compounds can impact the aerosol composition, and contribute to aqSOA formation upon
8 water evaporation. The corresponding aqSOA mass yields seem to depend on the spectral
9 irradiance of the light used to initiate the photochemistry, but further studies are needed to
10 confirm this point. Finally, the aging of the oligomers formed could be an explanation (at least
11 in part) for the presence of the diacids (such as oxalic, malonic and succinic acids) observed in
12 the ambient aerosol (Legrand et al., 2007; Kawamura et al., 2010). In Part 2 of this study, the
13 atmospheric relevance of these processes is explored by means of multiphase box model
14 studies.

15

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25

1 **References**

- 2 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S.,
3 Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn, A.,
4 Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R., Prevot,
5 A. S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U. and Jimenez, J. L.: O/C
6 and OM/OC Ratios of Primary, Secondary, and Ambient Organic Aerosols with High-
7 Resolution Time-of-Flight Aerosol Mass Spectrometry, *Environmental Science &*
8 *Technology*, 42(12), 4478–4485, doi:10.1021/es703009q, 2008.
- 9 Aljawhary, D., Lee, A. K. Y. and Abbatt, J. P. D.: High-resolution chemical ionization mass
10 spectrometry (ToF-CIMS): application to study SOA composition and processing,
11 *Atmospheric Measurement Techniques*, 6(11), 3211–3224, doi:10.5194/amt-6-3211-2013,
12 2013.
- 13 Altieri, K. E., Carlton, A. G., Lim, H.-J., Turpin, B. J. and Seitzinger, S. P.: Evidence for
14 Oligomer Formation in Clouds: Reactions of Isoprene Oxidation Products, *Environmental*
15 *Science & Technology*, 40(16), 4956–4960, doi:10.1021/es052170n, 2006.
- 16 Altieri, K. E., Seitzinger, S. P., Carlton, A. G., Turpin, B. J., Klein, G. C. and Marshall, A. G.:
17 Oligomers formed through in-cloud methylglyoxal reactions: Chemical composition,
18 properties, and mechanisms investigated by ultra-high resolution FT-ICR mass
19 spectrometry, *Atmospheric Environment*, 42(7), 1476–1490,
20 doi:10.1016/j.atmosenv.2007.11.015, 2008.
- 21 Arakaki, T., Anastasio, C., Kuroki, Y., Nakajima, H., Okada, K., Kotani, Y., Handa, D.,
22 Azechi, S., Kimura, T., Tshako, A. and Miyagi, Y.: A General Scavenging Rate Constant
23 for Reaction of Hydroxyl Radical with Organic Carbon in Atmospheric Waters,
24 *Environmental Science & Technology*, 130718140737000, doi:10.1021/es401927b, 2013.
- 25 Asher, W. E. and Pankow, J. F.: Vapor pressure prediction for alkenoic and aromatic organic
26 compounds by a UNIFAC-based group contribution method, *Atmospheric Environment*,
27 40(19), 3588–3600, doi:10.1016/j.atmosenv.2005.12.004, 2006.
- 28 Blando, J. D. and Turpin, B. J.: Secondary organic aerosol formation in cloud and fog
29 droplets: a literature evaluation of plausibility, 2000, 34(10), 1623–1632,
30 doi:10.1016/S1352-2310(99)00392-1, 2000.
- 31 Brégonzio-Rozier, L., Siekmann, F., Giorio, C., Pangui, E., Morales, S. B., Temime-Roussel,
32 B., Gratien, A., Michoud, V., Ravier, S., Tapparo, A., Monod, A. and Doussin, J.-F.:
33 Gaseous products and Secondary Organic Aerosol formation during long term oxidation of
34 isoprene and methacrolein, *Atmospheric Chemistry and Physics Discussions*, 14(16),
35 22507–22545, doi:10.5194/acpd-14-22507-2014, 2014.
- 36 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang, Q.,
37 Onasch, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R.,
38 Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P. and
39 Worsnop, D. R.: Chemical and microphysical characterization of ambient aerosols with the
40 aerodyne aerosol mass spectrometer, *Mass Spectrometry Reviews*, 26(2), 185–222,
41 doi:10.1002/mas.20115, 2007.
- 42 Carlton, A. G. and Turpin, B. J.: Particle partitioning potential of organic compounds is
43 highest in the Eastern US and driven by anthropogenic water, *Atmospheric Chemistry and*
44 *Physics*, 13(20), 10203–10214, doi:10.5194/acp-13-10203-2013, 2013.

- 1 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S., Reff, A., Lim, H.-J. and Ervens, B.:
2 Atmospheric oxalic acid and SOA production from glyoxal: Results of aqueous
3 photooxidation experiments, *Atmospheric Environment*, 41(35), 7588–7602,
4 doi:10.1016/j.atmosenv.2007.05.035, 2007.
- 5 Carlton, A. G., Turpin, B. J., Lim, H.-J., Altieri, K. E. and Seitzinger, S.: Link between
6 isoprene and secondary organic aerosol (SOA): Pyruvic acid oxidation yields low volatility
7 organic acids in clouds, *Geophysical Research Letters*, 33(6), doi:10.1029/2005GL025374,
8 2006.
- 9 Carlton, A. G., Wiedinmyer, C. and Kroll, J. H.: A review of Secondary Organic Aerosol
10 (SOA) formation from isoprene, *Atmospheric Chemistry and Physics*, 9(14), 4987–5005,
11 2009.
- 12 Daumit, K. E., Carrasquillo, A. J., Hunter, J. F. and Kroll, J. H.: Laboratory studies of the
13 aqueous-phase oxidation of polyols: submicron particles vs. bulk aqueous solution,
14 *Atmospheric Chemistry and Physics*, 14(19), 10773–10784, doi:10.5194/acp-14-10773-
15 2014, 2014.
- 16 DeCarlo, P. F., Dunlea, E. J., Kimmel, J. R., Aiken, A. C., Sueper, D., Crouse, J., Wennberg,
17 P. O., Emmons, L., Shinzuka, Y., Clarke, A. and others: Fast airborne aerosol size and
18 chemistry measurements with the high resolution aerosol mass spectrometer during the
19 MILAGRO Campaign, *Atmos. Chem. Phys. Discuss*, 7(18), 269–18, 2007.
- 20 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C.,
21 Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R. and Jimenez, J. L.:
22 Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer,
23 *Analytical Chemistry*, 78(24), 8281–8289, doi:10.1021/ac061249n, 2006.
- 24 Docherty, K. S., Jaoui, M., Corse, E., Jimenez, J. L., Offenberg, J. H., Lewandowski, M. and
25 Kleindienst, T. E.: Collection Efficiency of the Aerosol Mass Spectrometer for Chamber-
26 Generated Secondary Organic Aerosols, *Aerosol Science and Technology*, 47(3), 294–309,
27 doi:10.1080/02786826.2012.752572, 2013.
- 28 Donahue, N. M., Epstein, S. A., Pandis, S. N. and Robinson, A. L.: A two-dimensional
29 volatility basis set: 1. organic-aerosol mixing thermodynamics, *Atmospheric Chemistry
30 and Physics*, 11(7), 3303–3318, doi:10.5194/acp-11-3303-2011, 2011.
- 31 Epstein, S. A., Tapavicza, E., Furche, F. and Nizkorodov, S. A.: Direct photolysis of carbonyl
32 compounds dissolved in cloud and fog~droplets, *Atmospheric Chemistry and Physics*,
33 13(18), 9461–9477, doi:10.5194/acp-13-9461-2013, 2013.
- 34 Ervens, B., Turpin, B. J. and Weber, R. J.: Secondary organic aerosol formation in cloud
35 droplets and aqueous particles (aqSOA): a review of laboratory, field and model studies,
36 *Atmospheric Chemistry and Physics*, 11(21), 11069–11102, doi:10.5194/acp-11-11069-
37 2011, 2011.
- 38 Ervens, B. and Volkamer, R.: Glyoxal processing by aerosol multiphase chemistry: towards a
39 kinetic modeling framework of secondary organic aerosol formation in aqueous particles,
40 *Atmospheric Chemistry and Physics*, 10(17), 8219–8244, doi:10.5194/acp-10-8219-2010,
41 2010.
- 42 Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P. I., Geron, C. and others:
43 Estimates of global terrestrial isoprene emissions using MEGAN (Model of Emissions of
44 Gases and Aerosols from Nature), *Atmospheric Chemistry and Physics*, 6(11), 3181–3210,
45 2006.

- 1 Haddad, I. E., Liu, Y., Nieto-Gligorovski, L., Michaud, V., Temime-Roussel, B., Quivet, E.,
2 Marchand, N., Sellegri, K. and Monod, A.: In-cloud processes of methacrolein under
3 simulated conditions—Part 2: Formation of secondary organic aerosol, *Atmospheric*
4 *Chemistry and Physics*, 9(14), 5107–5117, 2009.
- 5 Hallquist, M., Wenger, J., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen,
6 J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. V., Herrmann, H.,
7 Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A.,
8 Maenhaut, W., McFiggans, G. B., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J.
9 H., Surratt, J. D., Szmigielski, R. and wild: The formation, properties and impact of
10 secondary organic aerosol: current and emerging issues, *Atmospheric Chemistry and*
11 *Physics*, 9(14), 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- 12 Herrmann, H., Hoffmann, D., Schaefer, T., Brüner, P. and Tilgner, A.: Tropospheric
13 Aqueous-Phase Free-Radical Chemistry: Radical Sources, Spectra, Reaction Kinetics and
14 Prediction Tools, *ChemPhysChem*, 11(18), 3796–3822, doi:10.1002/cphc.201000533,
15 2010.
- 16 Hobby, K.: A novel method of isotope prediction applied to elemental composition analysis,
17 2005.
- 18 Iraci, L. T., Baker, B. M., Tyndall, G. S. and Orlando, J. J.: Measurements of the Henry's law
19 coefficients of 2-methyl-3-buten-2-ol, methacrolein, and methylvinyl ketone, *Journal of*
20 *atmospheric chemistry*, 33(3), 321–330, 1999.
- 21 Jaoui, M., Corse, E., Kleindienst, T. E., Offenberg, J. H., Lewandowski, M. and Edney, E. O.:
22 Analysis of Secondary Organic Aerosol Compounds from the Photooxidation of d -
23 Limonene in the Presence of NO X and their Detection in Ambient PM 2.5, *Environmental*
24 *Science & Technology*, 40(12), 3819–3828, doi:10.1021/es052566z, 2006.
- 25 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H.,
26 DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I.
27 M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V.
28 A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J.,
29 Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J.,
30 E., Dunlea, J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K.,
31 Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo,
32 D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J.
33 Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C.,
34 Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E.,
35 Baltensperger, U. and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere,
36 *Science*, 326(5959), 1525–1529, doi:10.1126/science.1180353, 2009.
- 37 Kameel, F. R., Hoffmann, M. R. and Colussi, A. J.: OH Radical-Initiated Chemistry of
38 Isoprene in Aqueous Media. Atmospheric Implications, *The Journal of Physical Chemistry*
39 *A*, 117(24), 5117–5123, doi:10.1021/jp4026267, 2013.
- 40 Kameel, F. R., Riboni, F., Hoffmann, M. R., Enami, S. and Colussi, A. J.: Fenton Oxidation
41 of Gaseous Isoprene on Aqueous Surfaces, *The Journal of Physical Chemistry C*,
42 140725063829009, doi:10.1021/jp505010e, 2014.
- 43 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C.,
44 Dingenen, R. V., Ervens, B., Nenes, A., Nielsen, C. J. and others: Organic aerosol and
45 global climate modelling: a review, *Atmospheric Chemistry and Physics*, 5(4), 1053–1123,
46 2005.

- 1 Kawamura, K., Kasukabe, H. and Barrie, L. A.: Secondary formation of water-soluble organic
2 acids and α -dicarbonyls and their contributions to total carbon and water-soluble organic
3 carbon: Photochemical aging of organic aerosols in the Arctic spring, *Journal of*
4 *Geophysical Research*, 115(D21), doi:10.1029/2010JD014299, 2010.
- 5 Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C. and Seinfeld, J. H.: Secondary Organic
6 Aerosol Formation from Isoprene Photooxidation, *Environmental Science & Technology*,
7 40(6), 1869–1877, doi:10.1021/es0524301, 2006.
- 8 Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and
9 evolution of low-volatility organics in the atmosphere, *Atmospheric Environment*, 42(16),
10 3593–3624, doi:10.1016/j.atmosenv.2008.01.003, 2008.
- 11 Kuwata, M., Zorn, S. R. and Martin, S. T.: Using Elemental Ratios to Predict the Density of
12 Organic Material Composed of Carbon, Hydrogen, and Oxygen, *Environmental Science &*
13 *Technology*, 46(2), 787–794, doi:10.1021/es202525q, 2012.
- 14 Lee, A. K. Y., Hayden, K. L., Herckes, P., Leaitch, W. R., Liggio, J., Macdonald, A. M. and
15 Abbatt, J. P. D.: Characterization of aerosol and cloud water at a mountain site during
16 WACS 2010: secondary organic aerosol formation through oxidative cloud processing,
17 *Atmospheric Chemistry and Physics*, 12(15), 7103–7116, doi:10.5194/acp-12-7103-2012,
18 2012.
- 19 Lee, A. K. Y., Herckes, P., Leaitch, W. R., Macdonald, A. M. and Abbatt, J. P. D.: Aqueous
20 OH oxidation of ambient organic aerosol and cloud water organics: Formation of highly
21 oxidized products: AQUEOUS OXIDATION OF AMBIENT ORGANICS, *Geophysical*
22 *Research Letters*, 38(11), n/a–n/a, doi:10.1029/2011GL047439, 2011a.
- 23 Lee, A. K. Y., Zhao, R., Gao, S. S. and Abbatt, J. P. D.: Aqueous-Phase OH Oxidation of
24 Glyoxal: Application of a Novel Analytical Approach Employing Aerosol Mass
25 Spectrometry and Complementary Off-Line Techniques, *The Journal of Physical*
26 *Chemistry A*, 115(38), 10517–10526, doi:10.1021/jp204099g, 2011b.
- 27 Lee, W., Baasandorj, M., Stevens, P. S. and Hites, R. A.: Monitoring OH-initiated oxidation
28 kinetics of isoprene and its products using online mass spectrometry, *Environmental*
29 *science & technology*, 39(4), 1030–1036, 2005.
- 30 Legrand, M., Preunkert, S., Oliveira, T., Pio, C. A., Hammer, S., Gelencsér, A., Kasper-Giebl,
31 A. and Laj, P.: Origin of C2–C5 dicarboxylic acids in the European atmosphere inferred
32 from year-round aerosol study conducted at a west-east transect, *J. Geophys. Res.*,
33 112(D23), D23S07, doi:10.1029/2006JD008019, 2007.
- 34 Lim, Y. B., Tan, Y. and Turpin, B. J.: Chemical insights, explicit chemistry, and yields of
35 secondary organic aerosol from OH radical oxidation of methylglyoxal and glyoxal in the
36 aqueous phase, *Atmospheric Chemistry and Physics*, 13(17), 8651–8667, doi:10.5194/acp-
37 13-8651-2013, 2013.
- 38 Liu, P. S. K., Deng, R., Smith, K. A., Williams, L. R., Jayne, J. T., Canagaratna, M. R.,
39 Moore, K., Onasch, T. B., Worsnop, D. R. and Deshler, T.: Transmission Efficiency of an
40 Aerodynamic Focusing Lens System: Comparison of Model Calculations and Laboratory
41 Measurements for the Aerodyne Aerosol Mass Spectrometer, *Aerosol Science and*
42 *Technology*, 41(8), 721–733, doi:10.1080/02786820701422278, 2007.
- 43 Liu, Y., Siekmann, F., Renard, P., El Zein, A., Salque, G., El Haddad, I., Temime-Roussel,
44 B., Voisin, D., Thissen, R. and Monod, A.: Oligomer and SOA formation through aqueous

- 1 phase photooxidation of methacrolein and methyl vinyl ketone, *Atmospheric Environment*,
2 49, 123–129, doi:10.1016/j.atmosenv.2011.12.012, 2012.
- 3 Miyakawa, T., Matsuzawa, R., Katayama, M. and Takegawa, N.: Reconsidering Adhesion
4 and Bounce of Submicron Particles Upon High-Velocity Impact, *Aerosol Science and*
5 *Technology*, 47(5), 472–481, doi:10.1080/02786826.2013.763895, 2013.
- 6 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H. and Worsnop, D.
7 R.: Changes in organic aerosol composition with aging inferred from aerosol mass spectra,
8 *Atmospheric Chemistry and Physics*, 11(13), 6465–6474, doi:10.5194/acp-11-6465-2011,
9 2011.
- 10 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H.,
11 Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt,
12 L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H., Dinar, E., Rudich, Y.
13 and Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric
14 datasets from Aerosol Mass Spectrometry, *Atmospheric Chemistry and Physics*, 10(10),
15 4625–4641, doi:10.5194/acp-10-4625-2010, 2010.
- 16 Odian, G. G.: Principles of polymerization, Wiley, Hoboken, N.J. [online] Available from:
17 <http://public.eblib.com/EBLPublic/PublicView.do?ptiID=469767> (Accessed 10 October
18 2014), 2004.
- 19 Ortiz-Montalvo, D. L., Lim, Y. B., Perri, M. J., Seitzinger, S. P. and Turpin, B. J.: Volatility
20 and Yield of Glycolaldehyde SOA Formed through Aqueous Photochemistry and Droplet
21 Evaporation, *Aerosol Science and Technology*, 46(9), 1002–1014,
22 doi:10.1080/02786826.2012.686676, 2012.
- 23 Perri, M. J., Seitzinger, S. and Turpin, B. J.: Secondary organic aerosol production from
24 aqueous photooxidation of glycolaldehyde: Laboratory experiments, *Atmospheric*
25 *Environment*, 43(8), 1487–1497, doi:10.1016/j.atmosenv.2008.11.037, 2009.
- 26 Renard, P., Reed Harris, A. E., Rapf, R. J., Ravier, S., Demelas, C., Coulomb, B., Quivet, E.,
27 Vaida, V. and Monod, A.: Aqueous Phase Oligomerization of Methyl Vinyl Ketone by
28 Atmospheric Radical Reactions, *The Journal of Physical Chemistry C* [online] Available
29 from: <http://pubs.acs.org/doi/abs/10.1021/jp5065598> (Accessed 23 October 2014), 2014.
- 30 Renard, P., Siekmann, F., Gandolfo, A., Socorro, J., Salque, G., Ravier, S., Quivet, E.,
31 Clément, J.-L., Traikia, M., Delort, A.-M., Voisin, D., Vuitton, V., Thissen, R. and Monod,
32 A.: Radical mechanisms of methyl vinyl ketone oligomerization through aqueous phase
33 OH-oxidation: on the paradoxical role of dissolved molecular oxygen, *Atmospheric*
34 *Chemistry and Physics*, 13(13), 6473–6491, doi:10.5194/acp-13-6473-2013, 2013.
- 35 Spracklen, D. V., Jimenez, J. L., Carslaw, K. S., Worsnop, D. R., Evans, M. J., Mann, G. W.,
36 Zhang, Q., Canagaratna, M. R., Allan, J., Coe, H., McFiggans, G., Rap, A. and Forster, P.:
37 Aerosol mass spectrometer constraint on the global secondary organic aerosol budget,
38 *Atmospheric Chemistry and Physics*, 11(23), 12109–12136, doi:10.5194/acp-11-12109-
39 2011, 2011.
- 40 Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A.,
41 Szmigielski, R., Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C. and Seinfeld, J.
42 H.: Chemical Composition of Secondary Organic Aerosol Formed from the Photooxidation
43 of Isoprene, *The Journal of Physical Chemistry A*, 110(31), 9665–9690,
44 doi:10.1021/jp061734m, 2006.

- 1 Tan, Y., Carlton, A. G., Seitzinger, S. P. and Turpin, B. J.: SOA from methylglyoxal in clouds
2 and wet aerosols: Measurement and prediction of key products, *Atmospheric Environment*,
3 44(39), 5218–5226, doi:10.1016/j.atmosenv.2010.08.045, 2010.
- 4 Tan, Y., Lim, Y. B., Altieri, K. E., Seitzinger, S. P. and Turpin, B. J.: Mechanisms leading to
5 oligomers and SOA through aqueous photooxidation: insights from OH radical oxidation
6 of acetic acid and methylglyoxal, *Atmospheric Chemistry and Physics*, 12(2), 801–813,
7 doi:10.5194/acp-12-801-2012, 2012.
- 8 Tan, Y., Perri, M. J., Seitzinger, S. P. and Turpin, B. J.: Effects of Precursor Concentration
9 and Acidic Sulfate in Aqueous Glyoxal–OH Radical Oxidation and Implications for
10 Secondary Organic Aerosol, *Environmental Science & Technology*, 43(21), 8105–8112,
11 doi:10.1021/es901742f, 2009.
- 12 Volkamer, R., San Martini, F., Molina, L. T., Salcedo, D., Jimenez, J. L. and Molina, M. J.: A
13 missing sink for gas-phase glyoxal in Mexico City: Formation of secondary organic
14 aerosol, *Geophysical Research Letters*, 34(19), doi:10.1029/2007GL030752, 2007.
- 15 Zhang, X., Chen, Z. M. and Zhao, Y.: Laboratory simulation for the aqueous OH-oxidation of
16 methyl vinyl ketone and methacrolein: significance to the in-cloud SOA production,
17 *Atmospheric Chemistry and Physics*, 10(19), 9551–9561, doi:10.5194/acp-10-9551-2010,
18 2010.
- 19 Zhao, R., Lee, A. K. Y. and Abbatt, J. P. D.: Investigation of Aqueous-Phase Photooxidation
20 of Glyoxal and Methylglyoxal by Aerosol Chemical Ionization Mass Spectrometry:
21 Observation of Hydroxyhydroperoxide Formation, *The Journal of Physical Chemistry A*,
22 116(24), 6253–6263, doi:10.1021/jp211528d, 2012.

Exp. Name	[MVK] ₀ (mM)	[H ₂ O ₂] ₀ (mM)	MVK 90% Consumption time (min)	Aqueous phase characterization	Aerosol Characterization after nebulization
A	20	400	120		
B	5	100	50		
C	2	40	30	UPLC-ESI-MS UHPLC-UV IC-ESI-MS ^a	SMPS AMS
D	0.5	10	25		
E	0.2	4	20		

Table 1: Experimental conditions of [•]OH-oxidation of MVK in the aqueous phase. All experiments are performed at 25°C. Time 0 corresponds to MVK injection in the photoreactor (Samples taken at 0, 5, 10, 15, 30, 50, 75, 105, 150 min).

Reaction Time (min)	D_p (nm) ^a	N ($\times 10^3 \text{ cm}^{-3}$) ^a	H/C ^b	O/C ^b	ρ_{org} (g cm^{-3}) ^{b,c}	M_{SMPS} ($\mu\text{g m}^{-3}$) ^{a,d}	Yield (%) ^{a,d,e}
5	28.3 ± 0.7	90 ± 6	1.55 ± 0.02	0.28 ± 0.05	1.0 ± 0.2	10 ± 10	70 ± 80
15	38 ± 1	530 ± 30	1.55 ± 0.01	0.26 ± 0.01	1.1 ± 0.1	100 ± 20	90 ± 60
50	44 ± 1	860 ± 40	1.55 ± 0.01	0.29 ± 0.01	1.1 ± 0.1	240 ± 50	80 ± 50
105	37 ± 1	1200 ± 50	1.47 ± 0.01	0.42 ± 0.01	1.2 ± 0.2	220 ± 40	70 ± 40
150	35 ± 1	1310 ± 50	1.38 ± 0.01	0.57 ± 0.01	1.4 ± 0.2	200 ± 30	60 ± 40

Table 2: Characteristics of the aerosol formed from nebulized MVK-solutions at different reaction times for experiment B.

Values are the average of ^a three consecutive SMPS measurements, and ^b five consecutive O/C and H/C AMS measurements, and uncertainties represent the corresponding standard deviation of these averages.

^c Particle densities are calculated using the method by Kuwata et al. (2012), the associated uncertainties include the accuracy of $\pm 12\%$ stated by these authors for eq.1.

^d Mass concentration values include the corresponding aerosol densities.

^e Yield values and associated uncertainties include dilution + transmission efficiency determined for the nebulizing system (see supplementary information 1).

Exp. Name	[MVK] ₀ (mM)	t _{max} (min) ^a	ρ _{org} (g cm ⁻³) ^b	M _{SMPS} (μg m ⁻³) ^b	Yield (%) ^{b,c}	H/C ^b	O/C ^b
A	20	105	1.1 ± 0.1	900 ± 200	70 ± 50	1.54 ± 0.01	0.30 ± 0.01
B	5	50	1.1 ± 0.1	240 ± 50	80 ± 50	1.55 ± 0.01	0.29 ± 0.01
C	2	30	1.1 ± 0.1	100 ± 20	80 ± 60	1.55 ± 0.01	0.28 ± 0.01

Table 3: Overview of aerosol properties and mass yields for different initial MVK concentrations.

^a Values are given at t_{max}, corresponding to the maximum aerosol mass concentration.

^b All values and associated uncertainties are calculated as indicated in Table 2.

^c Including dilution + transmission efficiency (TE) in the nebulizing system; TE: 17.2 (±7.9) for [MVK]₀ = 2 - 20mM determined from NH₄NO₃ (see supplementary information 1).

1 **Figure 1:** Schematic overview of the aerosol generation setup
2

3 **Figure 2:** (a): Time profiles of MVK concentrations during the reaction (experiment B); and
4 the corresponding evolution of ESI-MS spectra for m/z 50 - 1200 at (b): 5 min, (c): 50 min
5 and (d): 150 min. Highlighted in red in (c) is the most intense peaks of the main series of
6 oligomers together with its related molecular structure.
7

8 **Figure 3:** Time profiles of the AMS total organic mass (red), ion fragments $C_2H_3O^+$ at m/z 43
9 (blue) and CO_2^+ m/z 44 (green) for nebulized solutions from experiment B (**a**), and the
10 corresponding evolution of AMS mass spectra (for m/z 0 - 100) for nebulized solutions
11 sampled after 5 min (**b**), 50 min (**c**) and 150 min (**d**). Values are averages of 5 consecutive
12 AMS runs, error bars represent their standard deviations.
13

14 **Figure 4:** Comparison of AMS mass spectra (in the range 100 - 200 amu) of the nebulized
15 solutions sampled from experiment B after 50 and 150 min of reaction.
16

17 **Figure 5:** Evolution of particle number (a) and mass (b) size distributions for nebulized
18 solutions sampled at different reaction times for experiment B.
19

20 **Figure 6:** Time profiles of the total aerosol mass (black diamonds) from the nebulized
21 solutions, and consumed MVK in the aqueous phase (red circles) for experiment B. ^a Values
22 represent averages of three consecutive SMPS measurements for each reaction time
23 considering the corresponding density (Table 1). Error bars represent the sum of the standard
24 deviation of these averages and the uncertainties of the density calculation. ^bMVK
25 concentrations were determined by means of UHPLC-UV, with an uncertainty of ± 2 %.
26

27 **Figure 7:** Influence of the initial MVK concentration on the evolution of the total SOA mass
28 obtained from the nebulized solutions. Values represent averages of three consecutive SMPS
29 measurements for each reaction time considering the corresponding density (Table 3). Error
30 bars represent the sum of the standard deviation of these averages and the uncertainties of the
31 density calculation. For the lowest initial concentrations (experiments D and E), blank signals
32 were subtracted, and a density of 1.1 g cm^{-3} was assumed.

1 **Figure 8:** Fractions f_{44} versus f_{43} for the nebulized solutions from experiments A, B, C, D and
2 E, as measured by the AMS are compared to ambient air LV-OOA and SV-OOA from the
3 compilation by Ng et al. (2010) (dashed black triangle), and are also compared to the
4 nebulization data by Lee et al. (2011a) (dotted grey triangle). For our experiments, the signal
5 from blank experiments was subtracted. The data are shown for experiments A, B, and C from
6 10 to 150 min of reaction, and for experiments D and E from 15 to 150 min of reaction.

7
8 **Figure 9:** Van Krevelen diagram for the nebulized solutions of experiment C, as measured by
9 the AMS. Only experiment C is shown for clarity as the data for experiments A and B are
10 stacked together with the data shown. ^a Red diamonds represent the elemental ratios of
11 oligomers with a Degree of Polymerization = 5 for the ten most abundant oligomer series
12 identified by Renard et al. (2013). The black dotted circle highlights compounds with similar
13 structures. ^b Ng et al., 2011, for details see text section 4.2.

14
15 **Figure 10 :** Time profiles of the quantified diacids in the solutions as measured by IC-ESI-
16 MS for experiment B ($[MVK]_0 = 5 \text{ mM}$).

17
18 **Figure 11:** general scheme of the potential atmospheric fate of MVK in the aqueous phase. X-
19 Axis denotes volatility (\log_{10} of C^* at 298K), y-axis denotes oxidation state, approximated by
20 the O/C ratios (Jimenez et al., 2009).

21

Table 1: Experimental conditions of OH-oxidation of MVK in the aqueous phase. All experiments were performed at 25°C. Time 0 corresponds to MVK injection in the photoreactor (Samples taken at 0, 5, 10, 15, 30, 50, 75, 105, 150 min).

Exp. Name	[MVK] ₀ (mM)	[H ₂ O ₂] ₀ (mM)	MVK 90% Consumption time (min)	Aqueous phase characterization	Aerosol Characterization after nebulization
A	20	400	90	UPLC-ESI-MS UPLC-UV IC-ESI-MS	SMPS AMS
B	5	100	50		
C	2	40	30		
D	0.5	10	25	UPLC-ESI-MS UPLC-UV	SMPS AMS
E	0.2	4	20		

Table 2: Characteristics of the aerosol formed from nebulized MVK-solutions at different reaction times for experiment B.

Values are the average of ^a three consecutive SMPS measurements, and ^b five consecutive O/C and H/C AMS measurements, and uncertainties represent the corresponding standard deviation of these averages.

^c Particle densities are calculated using the method by Kuwata et al. (2012), the associated uncertainties include the accuracy of $\pm 12\%$ stated by these authors for eq.1.

^d Mass concentration values include the corresponding aerosol densities.

^e Yield values and associated uncertainties include dilution + transmission efficiency determined for the nebulizing system (see Table S1).

Reaction Time (min)	DP (nm) ^a	N ($\times 10^3 \text{ cm}^{-3}$) ^a	H/C ^b	O/C ^b	ρ^{org} (g cm^{-3}) ^{b,c}	M _{SMPS} ($\mu\text{g m}^{-3}$) ^{a,d}	Yield (%) ^{a,d,e}
5	28.3 \pm 0.7	90 \pm 6	1.55 \pm 0.02	0.28 \pm 0.05	1.0 \pm 0.2	10 \pm 10	70 \pm 80
15	38 \pm 1	530 \pm 30	1.55 \pm 0.01	0.26 \pm 0.01	1.1 \pm 0.1	100 \pm 20	90 \pm 60
50	44 \pm 1	860 \pm 40	1.55 \pm 0.01	0.29 \pm 0.01	1.1 \pm 0.1	240 \pm 50	80 \pm 50
105	37 \pm 1	1200 \pm 50	1.47 \pm 0.01	0.42 \pm 0.01	1.2 \pm 0.2	220 \pm 40	70 \pm 40
31 150	35 \pm 1	1310 \pm 50	1.38 \pm 0.01	0.57 \pm 0.01	1.4 \pm 0.2	200 \pm 30	60 \pm 40

Table 3: Overview of aerosol properties and mass yields for different initial MVK concentrations.

^a Values are given at t_{max} , corresponding to the maximum aerosol mass concentration.

^b All values and associated uncertainties are calculated as indicated in Table 2.

^c Including dilution + transmission efficiency (TE) in the nebulizing system; TE: $17.2 (\pm 7.9)$ for $[\text{MVK}]_0=2\text{-}20\text{mM}$ determined from NH_4NO_3 .

Exp. Name	$[\text{MVK}]_0$ (mM)	t_{max} (min) ^a	ρ_{org} (g cm^{-3}) ^b	M_{SMPS} ($\mu\text{g m}^{-3}$) ^b	Yield (%) ^{b,c}	H/C ^b	O/C ^b
A	20	105	1.1 ± 0.1	900 ± 200	70 ± 50	1.54 ± 0.01	0.30 ± 0.01
B	5	50	1.1 ± 0.1	240 ± 50	80 ± 50	1.55 ± 0.01	0.29 ± 0.01
C	2	30	1.1 ± 0.1	100 ± 20	80 ± 60	1.55 ± 0.01	0.28 ± 0.01

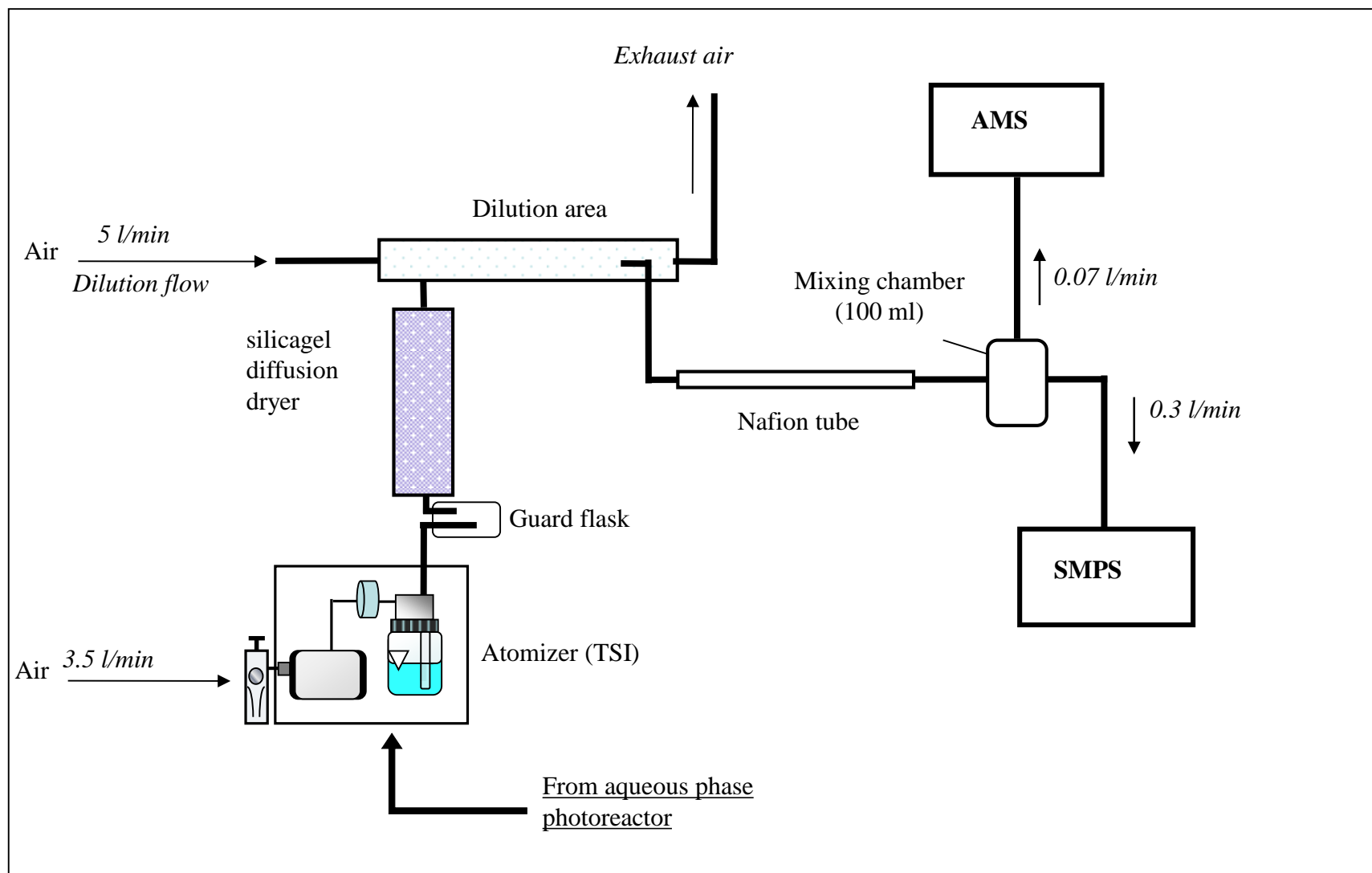


Figure 1: Schematic overview of the aerosol generation setup.

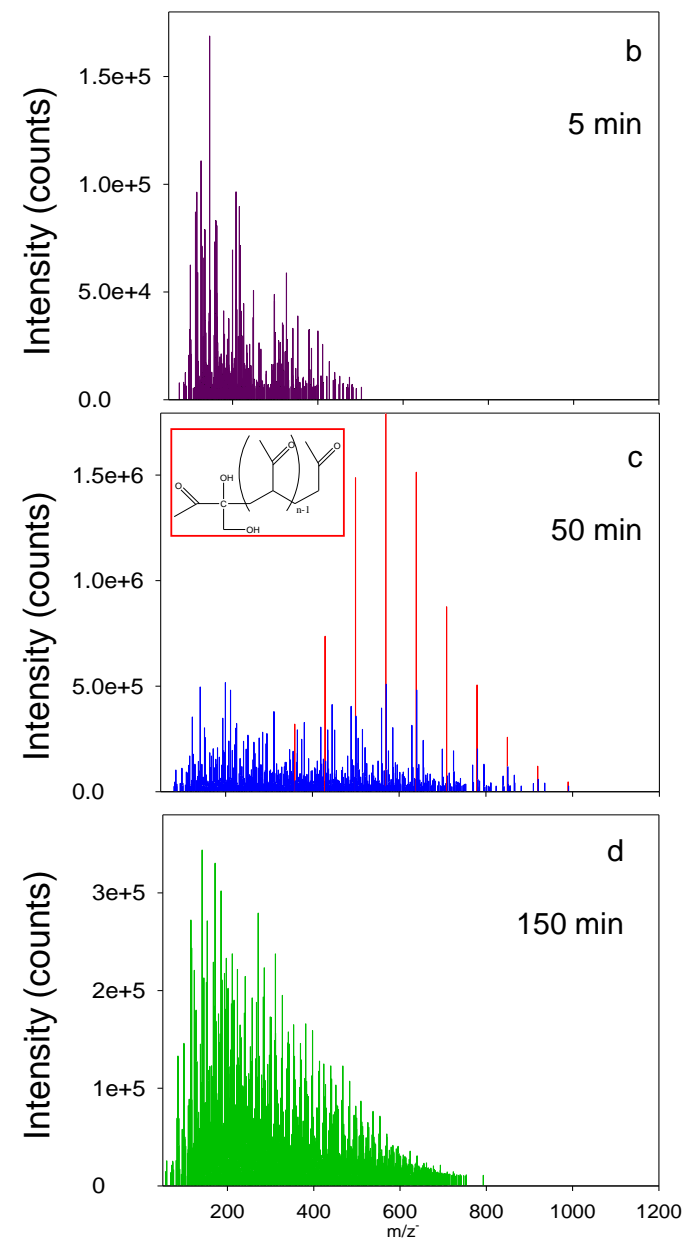
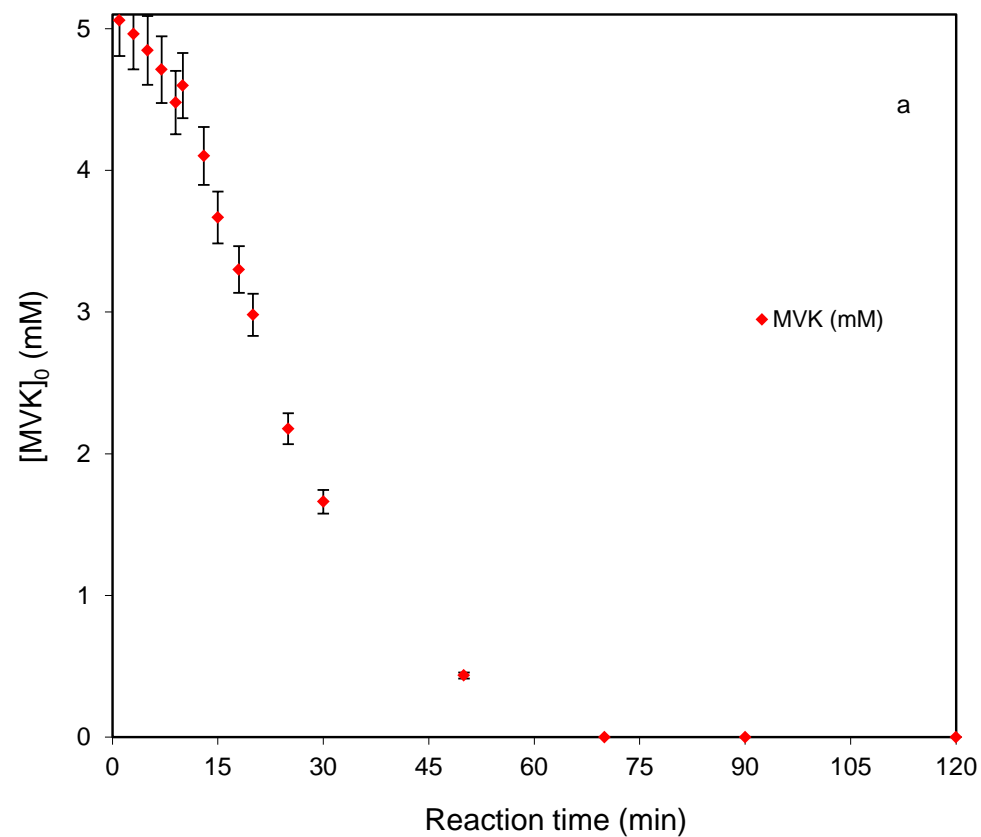


Figure 2: (a): Time profiles of MVK concentrations during the reaction (experiment B); and the corresponding evolution of ESI-MS⁻ spectra for m/z 50 - 1200 at (b): 5 min, (c): 50 min and (d): 150 min. Highlighted in red in (c), the most intense peaks of the main series of oligomers together with its related molecular structure.

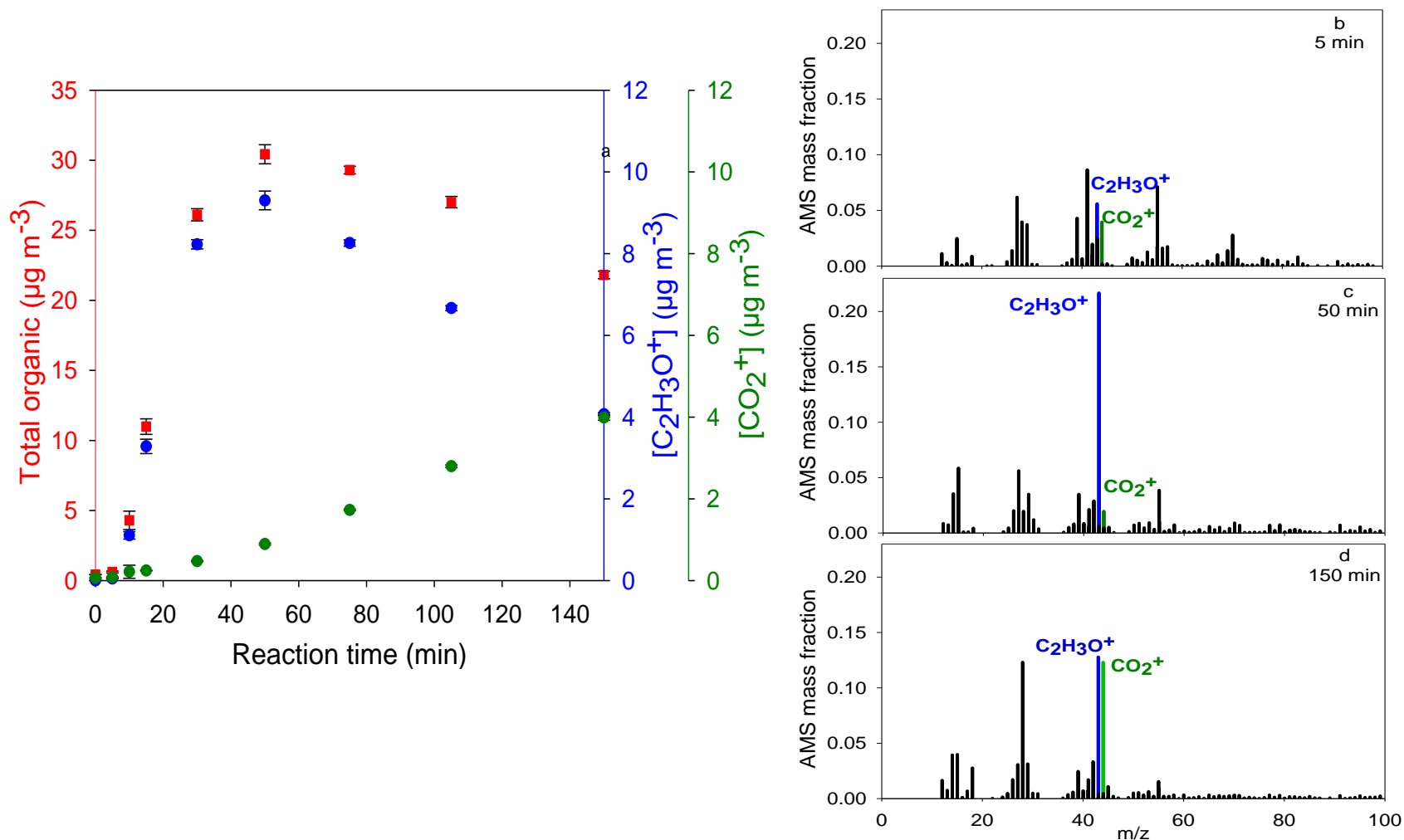


Figure 3: Time profiles of the HR-ToF-AMS total organic mass (red), ion fragments $C_2H_3O^+$ at m/z 43 (blue) and CO_2^+ m/z 44 (green) for nebulized solutions from experiment B (a), and the corresponding evolution of HR-ToF-AMS mass spectra (for m/z 0 - 100) for nebulized solutions sampled after 5 min (b), 50 min (c) and 150 min (d). Values are averages of 5 consecutive HR-ToF-AMS-runs, error bars represent their standard deviations. The AMS-mass fraction is the signal intensity contribution of each fragment to the total signal.

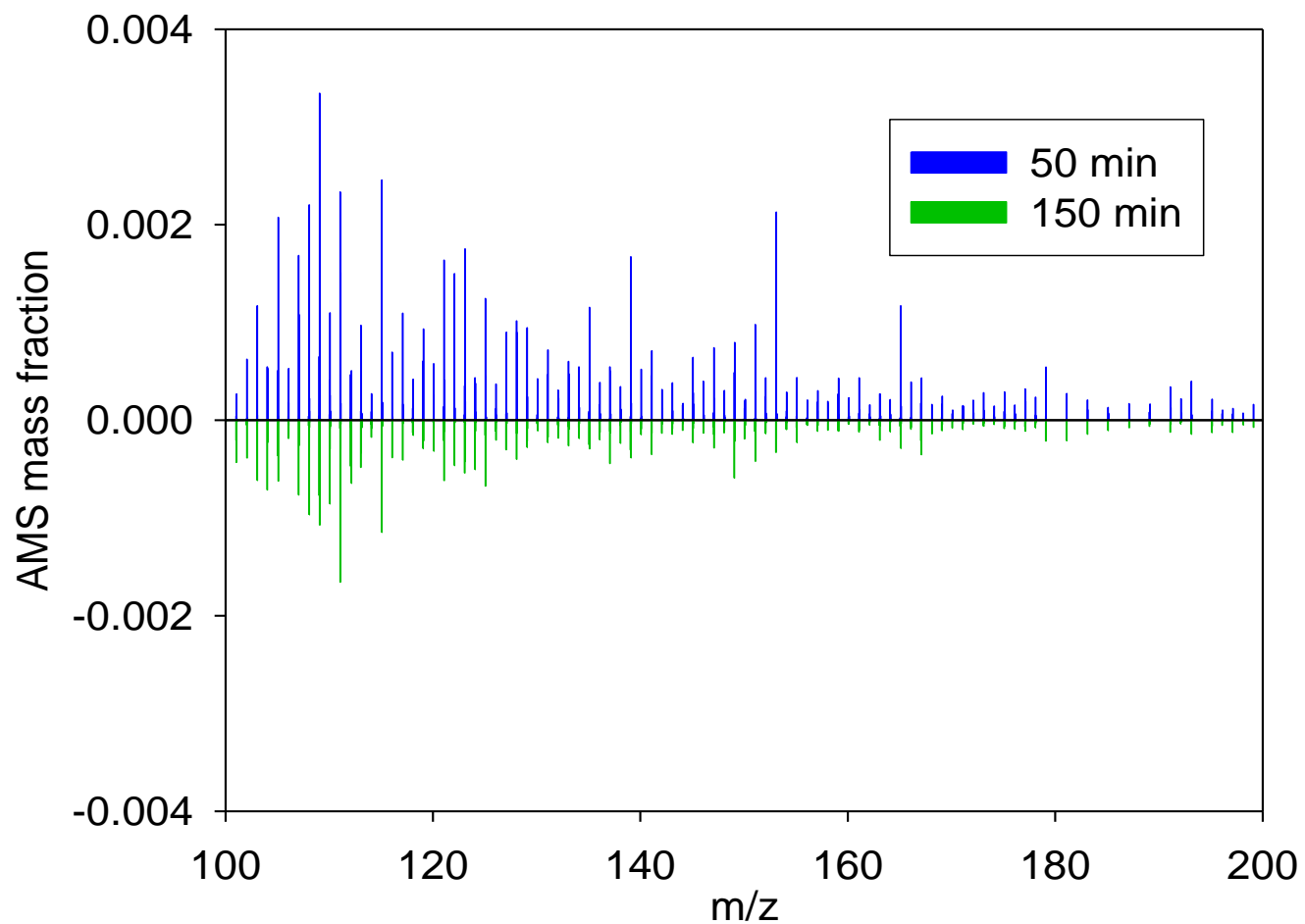


Figure 4: Comparison of HR-ToF-AMS mass spectra (in the range 100 - 200 amu) of the nebulized solutions sampled from experiment B after 50 and 150 min of reaction in the aqueous phase. The AMS-mass fraction is the signal intensity contribution of each fragment to the total signal.

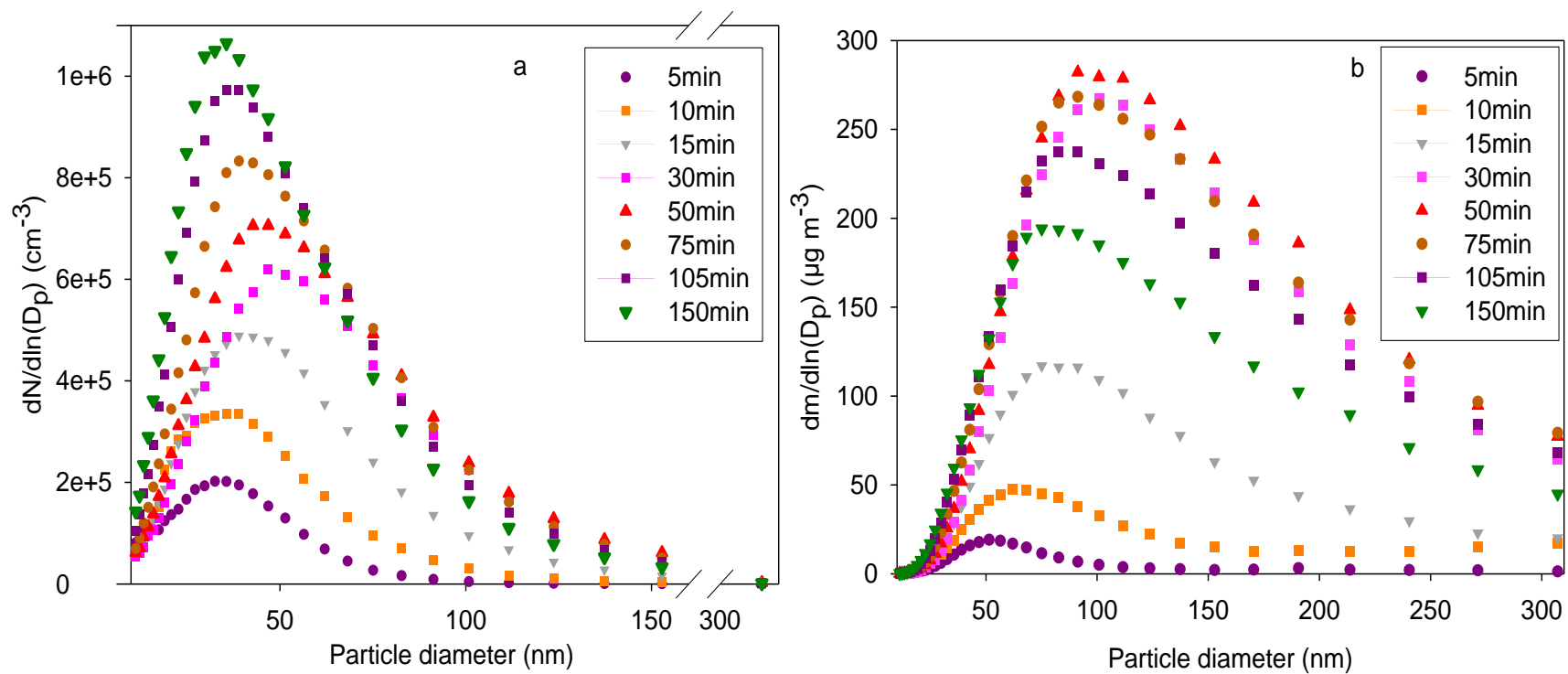


Figure 5: Evolution of particle number (a) and mass (b) size distributions for nebulized solutions sampled at different reaction times for experiment B, measured by means of SMPS. The standard deviations are smaller than the width of the symbols

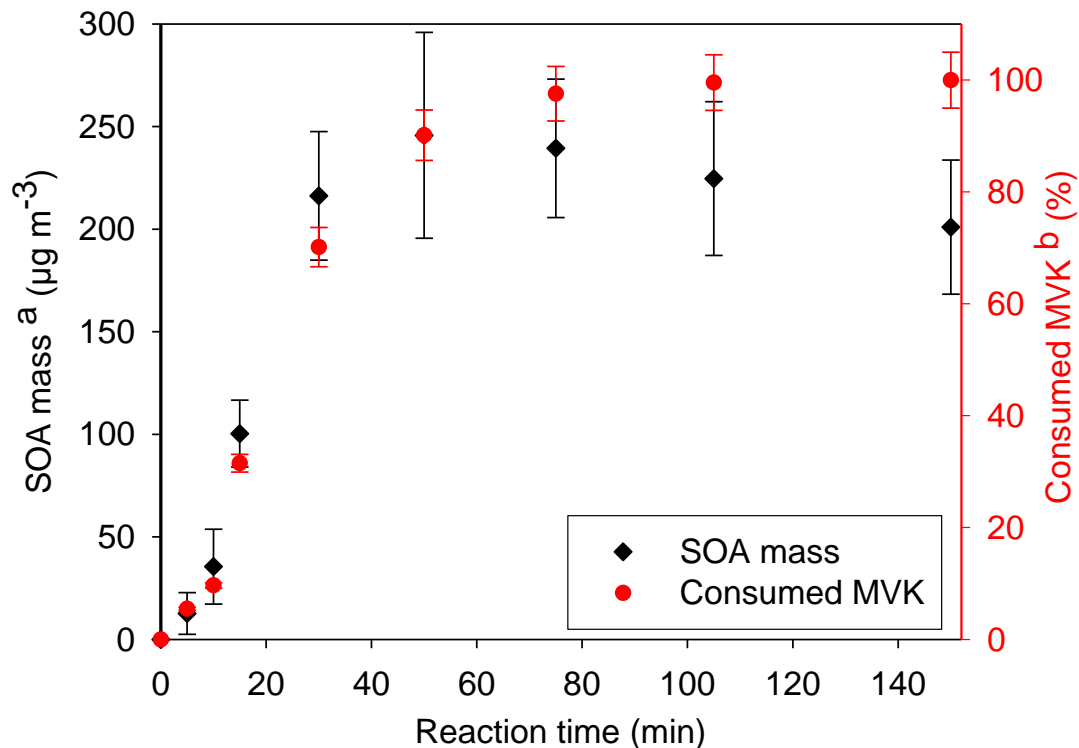


Figure 6: Time profiles of the total aerosol mass (black diamonds) from the nebulized solutions, and consumed MVK in the aqueous phase (red circles) for experiment B. ^a Values represent averages of three consecutive SMPS-measurements for each reaction time considering the corresponding density (Table 1). Error bars represent the sum of the standard deviation of these averages and the uncertainties of the density calculation. ^b % in concentration of MVK. These concentrations were determined by means of UHPLC-UV, with an uncertainty of ± 2 %;

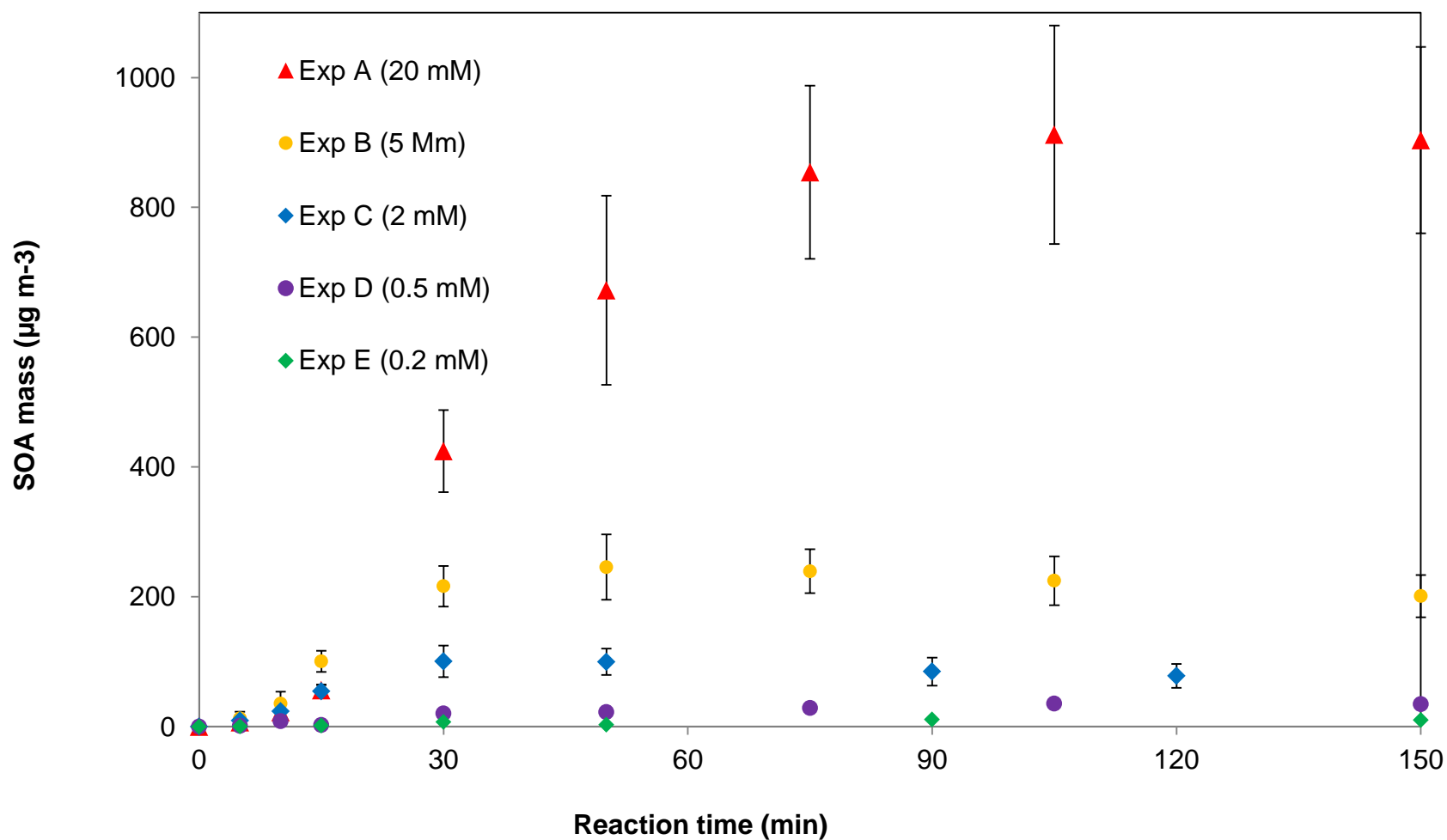


Figure 7: Influence of the initial MVK concentration on the evolution of the total SOA mass obtained from the nebulized solutions. Values represent averages of three consecutive SMPS-measurements for each reaction time considering the corresponding density (Table 3). Error bars represent the sum of the standard deviation of these averages and the uncertainties of the density calculation. For the lowest initial concentrations (experiments D and E), blank signals were subtracted, and a density of 1.1 g cm^{-3} was assumed.

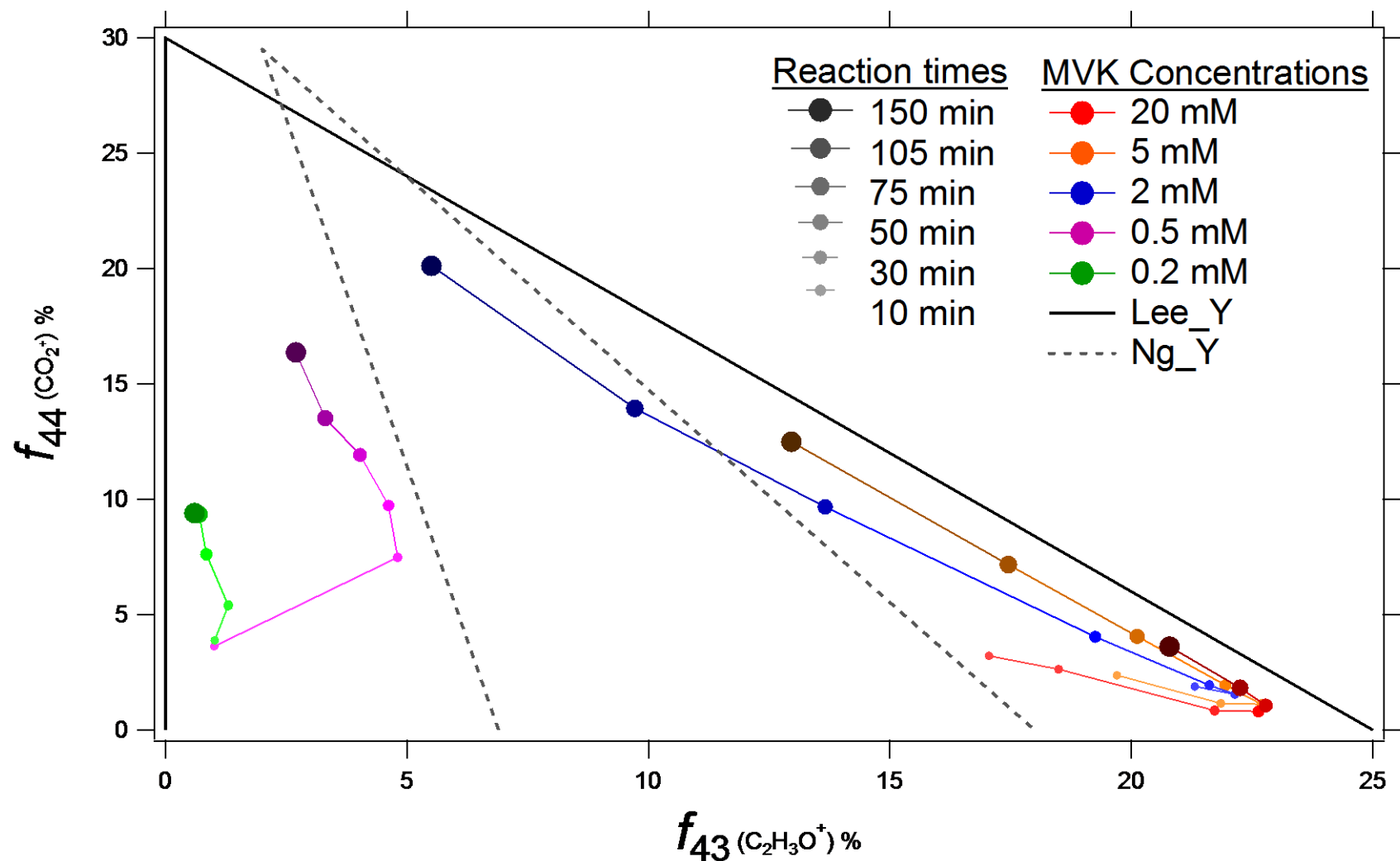


Figure 8: Fractions f_{44} versus f_{43} for the nebulized solutions from experiments A, B, C, D and E, as measured by the HR-ToF-AMS are compared to ambient air LV-OOA and SV-OOA from the compilation by Ng et al. (2010) (dashed black triangle), and are also compared to the nebulization data by Lee et al. (2011a) (dotted grey triangle). For our experiments, the signal from blank experiments was subtracted. The data are shown for experiments A, B, and C from 10 to 150 min of reaction, and for experiments D and E from 15 to 150 min of reaction. The gas/particle partitioning in our set up may alter the relative oxygenation of the OA produced, especially at low initial concentrations.

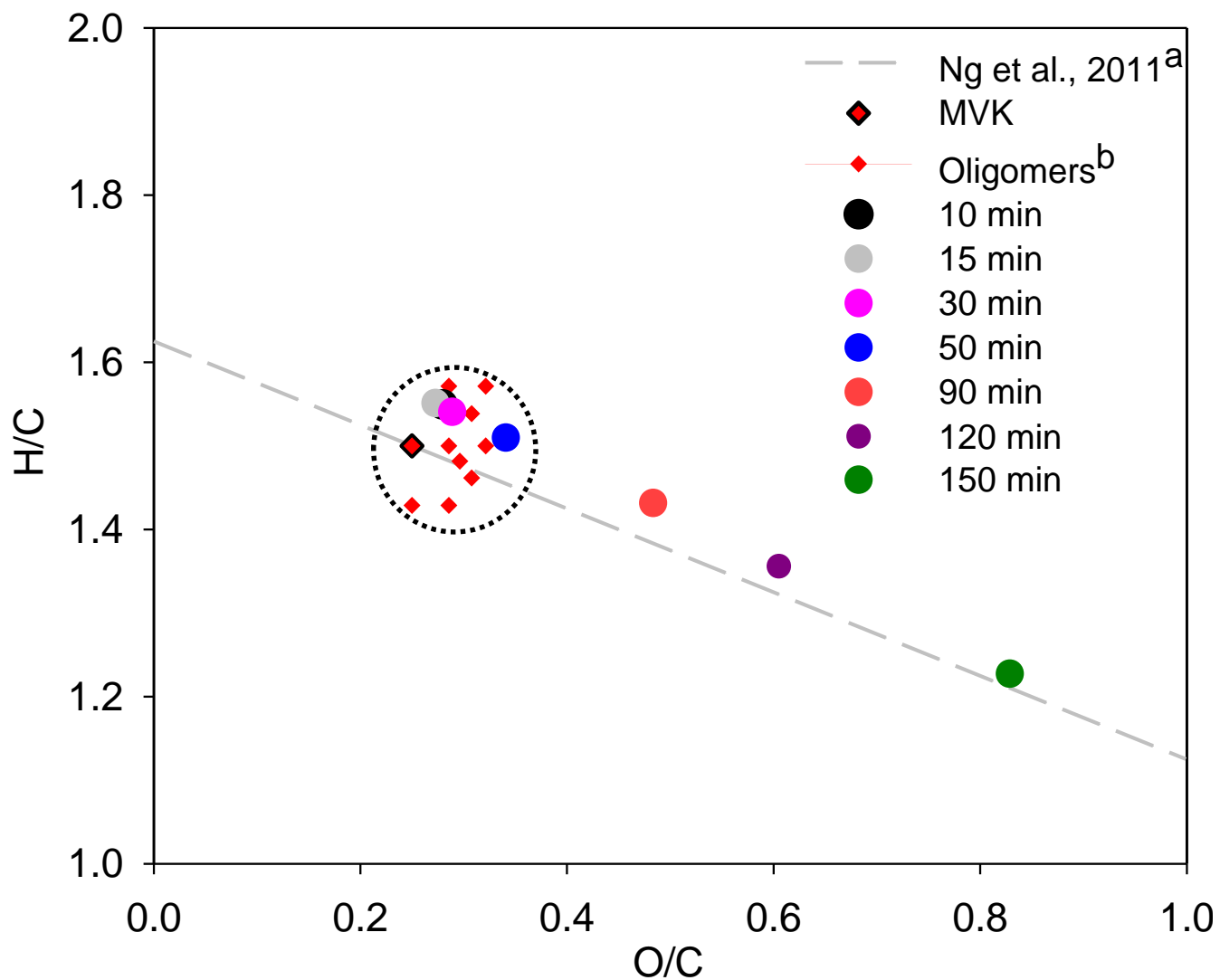
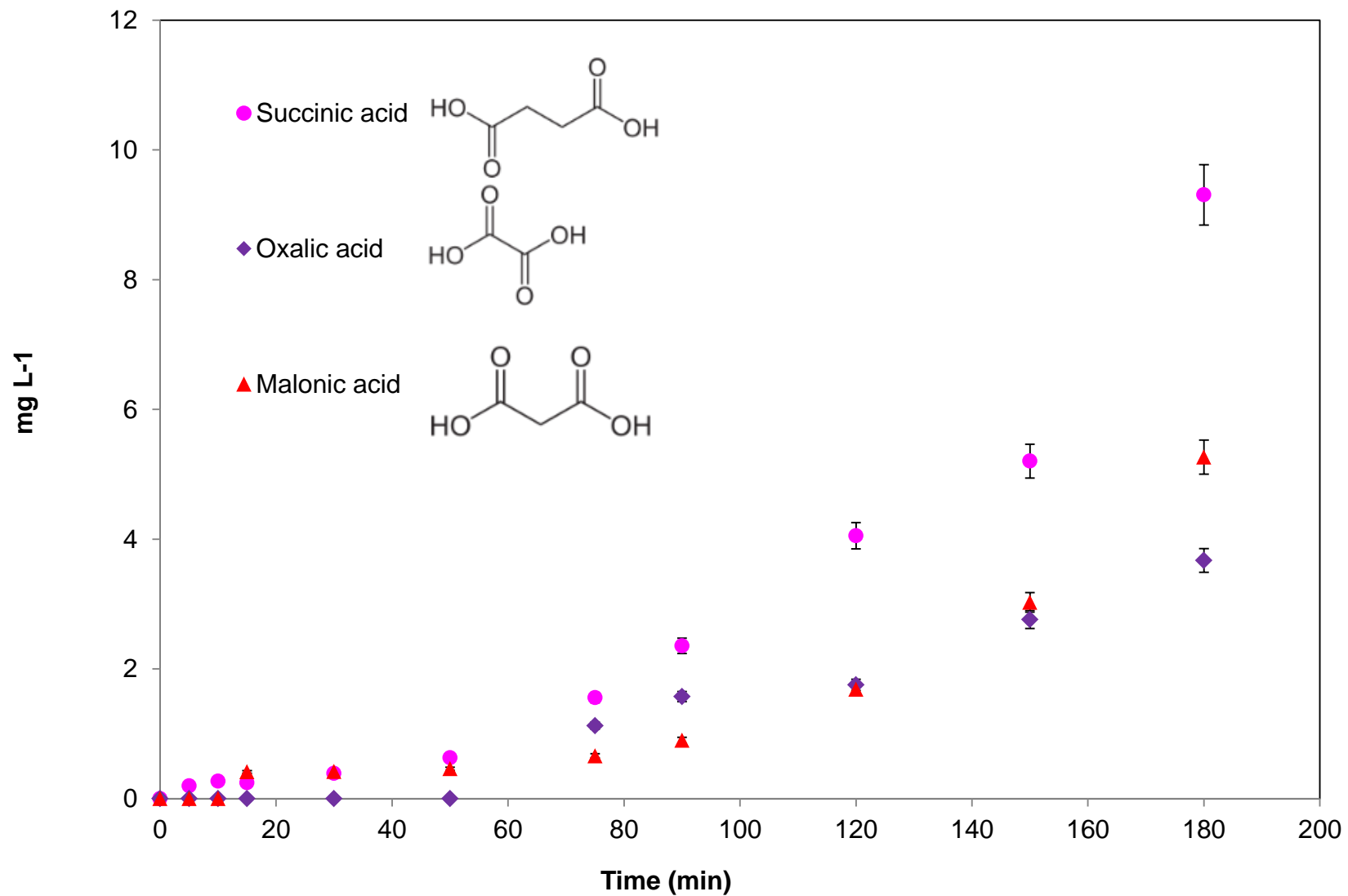
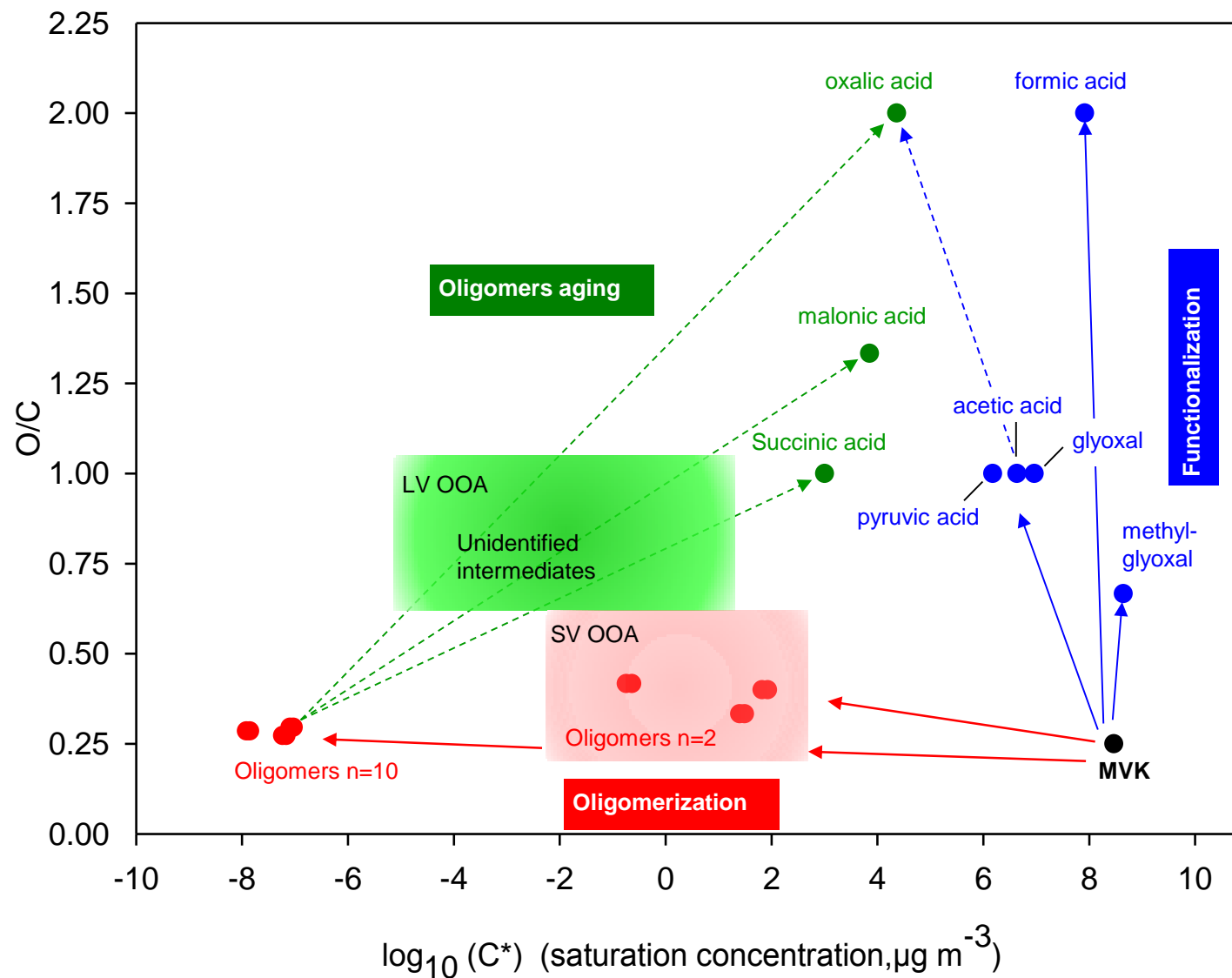


Figure 9: Van Krevelen diagram for the nebulized solutions of experiment C, as measured by the AMS. Only experiment C is shown for clarity as the data for experiments A and B are stacked together with the data shown. ^a Red diamonds represent the elemental ratios of oligomers with a Degree of Polymerization = 5 for the ten most abundant oligomer series identified by Renard et al. (2013). The black dotted circle highlights compounds with similar structures. ^b Ng et al., 2011, for details see text section 4.2.



42 **Figure 10** : Time profiles of the quantified diacids in the solutions as measured by IC-ESI-MS for experiment B ($[MVK]_0 = 5$ mM).



43 **Figure 11:** Potential atmospheric fate of MVK in the aqueous phase. X-Axis denotes volatility (\log_{10} of C^* at 298K), y-axis denotes oxidation state, approximated by O/C (Jimenez et al., 2009).