1 We appreciate your thorough review of the manuscript and valuable comments which helped us to improve 2 the paper. We have revised the paper carefully and our point-by-point responses are detailed below. All the 3 revised parts have been highlighted in the manuscript.

Responds to Reviewer 1:

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6 7 O: "first, sections 3.3-3.5 contain the most valuable new information from the conducted 8 measurements. While these sections report the main findings, they lack clear scientific conclusions. 9 The few attempts toward this direction have often erroneous interpretations. For example, the authors 10 link low aerosol mass mean diameters to nucleation events (lines 3-5 on page 15200). It is true that less aged aerosols tend to have a smaller mass median diameter, but nucleated aerosols are very unlikely to 11 12 be responsible for the small mass median diameters. It is rather other way around: air masses with 13 lower mass median diameters tend to be younger/cleaner, making nucleation more probable. As 14 another example, I do not understand how aerosol hygroscopicity would contribute its growth (lines 15 12-13 on page 15200)." 16

A: Accepted. First of all, we re-organized the contents of section 3.3 according to the durations that were influenced by the conditions of planetary boundary layer, low free troposphere, new particle formation, in cloud and aged events. With each condition, the mass concentration, mass-resolved size distribution and their proportions on different size ranges were discussed in detail. We removed the discussion of the relationship between small aerosol mass mean diameters and nucleation events.

Secondly, in section 3.4, we polished the contents and improved interpretations about the back trajectory
 results to clearly show the conclusions.

Thirdly, in section 3.5, PMFs was employed to categorize OA into different types of organics. We re-run the
 dataset with both PMF and ME-2 to make sure the results are consistent. Results present BBOA should be
 categorized as one subtype of OA in fall, and some modification was also done in winter.

30 Q: "second, I am not fully satisfied with the selection of sites for the comparisons made in sections 3.1 31 and 3.2. Why this set of sites? I would have liked to see comparisons to aerosol chemistry at other 32 elevated sites all over the world, rather than picking up e.g. urban sites from here and there. One more 33 thing: while the first part of Table 1 contains short description of the type of site, the second part of 34 this table does not!"

A: We accepted these suggestions on site selections. At the early version of this MS, we only considered the
comparison of the MMDs of chemical species between our site and other sites with size distribution data.
A new Table1 has included more comparisons on chemistry of particles in different size ranges from more
elevated sites. Because the content of section 3.3 has been modified into five kinds of episodes for discussion,
table 2 and table 3 were replaced by figures accordingly (Figure 3, 4, and 5).

42 Q: "the introductions (section 1) have multiple problems. The first paragraph and beginning of the 43 second one in it are very difficult to understand. More specifically, it remains unclear how the 44 beginning of introduction motivates the research make in this manuscript. The introduction does not 45 state clear scientific goals for this work either. It is modest to stating that the purpose is to assess 46 regionally representative concentration levels and obtain seasonal variations."

- A: Accepted. The discussions in the Introduction have been improved, the scientific goals for this work were
 stated more clearly, which are to improve the understanding of regionally and seasonally representative
 chemical components in PM1, and related formation and transportation processes.
- Q: "Finally, no figures have been presented on the actual results (all the information is in Tables). This
 makes it very difficult for a reader to digest the results."

A: Accepted. Some unnecessary tables (Table 1, 2, 3 in old version) were deleted, and new tables (Table 1, 2
 in new version) were added. Figures (new Figures 2, 3, 4, and 5) were added to better show major results.

Some reorganization of figures has also been done (Figure 1 and Figure 2 in the old version were changed to Figure 6 and Figure 1, respectively).

Responds to Reviewer 2:

O: "1) Page 15193, line 9, what does it mean "cloud-nucleating ability"? "

A: We have changed "cloud-nucleating ability" to "aerosol-nucleating ability"

Q: "2) Page 15193, line 14, "S. Asia" what's this?"

A: it is South Asia, and it has been written in full name.

Q: "3) Page 15193, the second paragraph, it does not read well and needs to be re-organized."

A: It has been re-organized.

O: "4) Page 15196, section 2.2, the maximum size for SMPS is about 520 nm, the authors should explain how to calculate PM1.0 mass from SMPS measurements."

A: Accepted. The method of calculate PM1.0 mass from SMPS measurement has been explained in the new version on P3, L49-52.

Q: "5) Page 15196, section 2.3, the authors try to interpret how to differentiate the air mass from the PBL vs. from the FT, but it is hard to follow."

A: the definitions of PBL, FT, NPF, in cloud and aged have been re-interpreted in the revised paper on P4, L11-20.

Q: "6) Page15197, section 3.1, what are those criteria and objectives for the comparison? More interpretations are needed."

A: Also according to another reviewer's comment, a new table and some new comparison of chemistry from different elevated sites were summarized and discussed in the revised paper on section 3.1

Q: "7) Page 15199, the top paragraph, it is difficult to follow because of the language."

A: it has been modified.

O: "8) Page 15199, Section 3.3., the technical terms such as "fresh aerosols" and "aged aerosols" are usually used for those from primary sources. New particle formation events are of course subject to secondary origins of aerosols. The authors should consider revising the part."

A: it has been revised accordingly

O: "9) Page 15200, the second paragraph, no conclusion? For the third and fourth paragraph, the reviewer has difficult to link those piece-by-piece analyses with the conclusion."

A: Accepted. This part has been modified completely, the former version only focused on the comparison between aged and fresh air, the new version expended the contents to five different kinds of episodes, and the discussions were also changed accordingly.

Q: "10) Page 15201, Section 3.4, short distance is not defined?"

A: Accepted. It has been modified accordingly.

Q: "11) Section 3.5, what are new findings?"

A: we employed PMF to categorize organics in four seasons. At this stage, we re-analyze the AMS dataset with both PMF and ME-2 to double check the final results of resolved organics. BBOA was resolved in fall in the new version, and the values in table 3 have been updated accordingly. The final results show that the oxygenic organics aerosols occupied half of organics matters. Intensive influence from the field burning of agriculture residue in summer and coal combustion in winter can be concluded in this paper. These two sources should be concerned by policy makers to improve air quality in China.

Chemical Composition and Mass Size Distribution of PM₁ at an 2 **Elevated Site in Central East China** 3

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13 Abstract. Size-resolved aerosol chemical compositions were measured continuously for one and 14 half years from June 2010 to January 2012 with an aerosol mass spectrometer (AMS) to 15 characterize the mass and size distributions (MSDs) of major chemical components in submicron particles (approximately PM₁) at Mountain Tai, an elevated site in Central East China (CEC). The 16 annual mean mass concentrations of organic, sulfate, nitrate, ammonium and chloride were 11.2, 17 18 9.2, 7.2, 5.8 and 0.95 μ g m⁻³, respectively, which are much higher than those at most mountain sites in the USA and Europe, but lower than those at the nearby surface rural sites in China. A clear 19 20 seasonality was observed for all major components throughout the campaign with low 21 concentration in fall and high in summer, and was believed to be caused by seasonal variations in 22 planetary boundary layer (PBL) height, near surface pollutant concentrations and regional transport 23 processes. Air masses were classified into categories impacted by PBL, lower free troposphere 24 (LFT), new particle formation (NPF), in-cloud processes and polluted aerosols. Organics dominated 25 the PM₁ mass during the NPF episodes, while sulfate contributed most to PM₁ in cloud events. The 26 average MSDs of particles between 30–1000 nm during the entire campaign for organics, sulfate, 27 nitrate, and ammonium were approximately log-normal with mass median diameters (MMDs) of 28 539, 585, 542, and 545 nm, respectively. These values were slightly larger than those observed at 29 ground sites within the North China Plain (NCP), likely due to the relative aged and well-mixed 30 aerosol masses at Mt. Tai. There were no obvious differences in MMDs during the PBL, LFT, in cloud and polluted episodes, but smaller MMDs, especially for organics, were observed during the 31 32 NPF events. During the PBL, NPF and polluted episodes, organics accounted for major proportions at smaller modes, and reached to 70% at 100-200nm particles in the polluted events. In cloud 33 episodes, inorganics contributed 70% to the whole size range dominated by sulfate, which 34 contributed 40% to small particles (100-200nm), while organics occupied 20%, indicating that 35 36 sulfate is critical chemical component in cloud formation. Seven clusters of air masses were 37 classified based on 72-hour back trajectory analysis. The majority of the regionally-dispersed aerosols were found to be contributed from short distance mixed aerosols, mostly originated from 38 39 the south with organics and sulfate as major components. Air masses from long range transport 40 always brought clean and dry aerosols which resulted in low concentrations at the Mt. Tai. AMS-PMF (Positive Matrix Function) was employed to resolve the subtype of organics. Oxygenic 41 organics aerosols occupied 49%, 56%, 51% and 41% of OA in the four seasons respectively, 42 43 demonstrating that most OA were oxidized in summer due to strong photochemical reactions. 44 Biomass burning organics aerosols (BBOA) accounted for 34% of OA in summer mainly from field 45 burning of agriculture residues, and coal combustion organics aerosols (CCOA) accounted for 22% 46 of OA in winter from heating.

47 Key words: Chemical Composition, Mass-size distribution, PM₁, Central East China

1 Introduction

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4 Atmospheric aerosol particles, especially the fine particles that have relatively long atmospheric 5 residence times, not only damage human health (Ramgolam et al., 2009), but also: (1) affect the 6 earth's radiative balance by scattering and absorbing solar radiation, (2) influence cloud albedo and precipitation (Charlson et al., 1992;Solomon et al., 2007), and (3) indirectly influence the earth's 7 8 radiative balance by serving as nuclei for cloud droplets (Twomey, 1974; Albrecht, 1989). Despite 9 of the growing recognition of their importance for human health and earth systems, the uncertainties 10 of magnitudes of these effects are still large (Solomon et al., 2007). In fact, the influence of regionally dispersed fine aerosols on weather-climate becomes more important than that 11 immediately influenced by local specific source and by coarser particles, whose chemical 12 composition and size distribution are two critical parameters. The size distributions of the principal 13 14 fine particle constituents also have important impacts on visibility (Watson, 2002), aerosol radiative 15 properties and aerosol-nucleating ability (Boucher et al., 2013). The mass-size distributions (MSDs) 16 of chemical components of fresh and aged aerosol are also critical for improving the simulation and 17 validation of the aerosol size distributions and their optical properties (Zhou et al., 2012). Differing from conditions at ground level, the atmosphere in mountainous areas tends to have lower 18 19 temperature, higher relative humidity (RH) and more intense solar radiation (Decesari et al., 20 2005;Li et al., 2011;Seinfeld et al., 2004). Tropospheric aerosols over mountains are most often 21 derived from long-range transport, and therefore samples from high elevations are often 22 representative of regional- to large-scale atmospheric conditions (Li et al., 2011; Wang et al., 2011).

23 As to the global scale, tropospheric aerosols are highly variable in time and space, in which the 24 aerosol chemical components over China are found to be much higher than those in majority areas 25 of the world, except for urban area in South Asia (Zhou et al., 2012). The sampling site in this study, Mountain Tai (Mt. Tai), is located in the Central East China with a peak elevation of ~1500 m 26 27 (a.s.l). As the tropospheric aerosols over mountains are most often representative of regional- to 28 large-scale mixed origins (Li et al., 2011; Wang et al., 2011), it is optimal for investigating the regionally dispersed pollution. The second feature of Mt. Tai is the high frequency with which 29 30 clouds envelop at the mountain peaks and provides a region where the interaction of aerosols with 31 clouds can be studied over extended periods of time. Although there are some recent studies that 32 investigated particulate concentrations at Mt. Tai, they have been limited on the measurement of organic molecular compositions (Fu et al., 2008), gas-phase total peroxides (Ren et al., 2009), role 33 34 of biogenic volatile organic compounds (Fu et al., 2010), the impact of open crop residual burning 35 on O₃, CO, black carbon (BC or called EC) and organic carbon (OC) (Yamaji et al., 2010), EC-OC and inorganic ions in PM₁₀ during spring time (Wang et al., 2011), individual particle analysis (Li 36 37 et al., 2011). These works were only limited on one or two seasons with daily sampling resolutions. 38 No systematic measurements with high time resolution of chemical components MSDs in fine 39 aerosol particles, covered four seasons were reported yet. It is essential to accurately and 40 objectively assess the characterization of various chemical components and their size distribution of 41 fine aerosols in high atmospheric level in CEC.

This paper presents a data set covering one and half years of measurement of the mass 42 43 concentrations and size distributions of selected chemical components in PM₁. The purpose was to 44 assess the regionally representative concentration levels of different aerosol chemical components, 45 and to obtain the seasonal variations. In terms of the high altitude of the site prone to be influenced by PBL, LFT, cloud events, through the identification of relative fresh (associated with new particle 46 47 formation event) and polluted aerosol, and aerosols in planet boundary layer (PBL), lower free 48 troposphere (LFT) and in cloud, the MSDs of organic, sulfate, nitrate and ammonium under 49 different conditions were estimated for better characterizing the aerosol chemical compositions of 50 well-mixed aerosol and also for model verification. As the site is often influenced by regional-scale transported pollutants, the chemical and size properties of PM₁ from different air masses are also 51 52 discussed. Finally, the secondary organic aerosol (SOA) fraction and primary organics aerosol (POA) were subtyped by using a positive matrix factorization method (PMF), the mass loading level and contribution of different types of organics were presented and discussed.

2 Experimental

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2.1 Campaign description

8 As part of the aerosol-cloud interaction campaign, supported by National Key Project of Basic 9 Research, the aerosol chemical composition analyses were conducted at the summit of Mt. Tai (36.251°N, 117.101°E), located in Shandong Province of China with the highest elevation (1534 m 10 a.s.l) in CEC and settled along the pathway of Asian continental outflow. Tai'an is the nearest small 11 12 city, 15 km away in the south with ~500 000 population while Jinan, the capital city of Shandong Province (population: 2.1 million) is 60 km away in the north. Because the elevation of Mt. Tai 13 station is close to the top of the planetary boundary layer, and the sampled aerosols are 14 representative of the region rather than the immediate locality, it is a suitable site for investigating 15 the regionally dispersed aerosol pollution over the heavily polluted CEC. 16

17 From June 2010 to January 2012, an aerosol mass spectrometer (Q-AMS, Aerodyne Research 18 Inc. Boston, MA, USA) was used to measure the mass concentrations (30 nm-1 µm) of organics, 19 sulfate, nitrate, ammonium and chloride in MS-mode with 5 minutes time resolution (Javne et al., 20 2000;Zhang et al., 2011). The sampling periods for the AMS study were from 24 June-15 August, 21 22 September-11 October in 2010, 26 March-20 April, 5 June-30 June, 22 October-1 December in 22 2011 and 22 December 2011-13 January 2012. As a result of interruptions due to calibration, 23 instrument failures and local burning events, 123 days of valid data were obtained. From July to 24 December in 2010, a scanning mobility particle sizer (SMPS) (TSI 3936, TSI Inc.) was also used to 25 monitor the number size distribution of aerosols over a size range of 10-680 nm at a time resolution 26 of 5 minutes (2.5 minutes per scan, two scans) (Zhang et al., 2011). From December 25, 2010 until 27 the end of the study, a tandem differential mobility particle sizer (TDMPS, IFT, Leibniz Institute 28 for Tropospheric Research, Germany) and an aerodynamic particle sizer (APS, model 3321, TSI 29 Inc., St Paul, USA) were used instead of the SMPS for particle number size distribution (PNSD) 30 measurements. Together these two instruments covered particles from 3 nm to 2.5 µm in diameter. 31 To correct the concentrations for bounce at the vaporizer and the partial transmission of particles by 32 the lens, a fixed particle-collection efficiency factor of 0.5 was used, which was obtained in studies 33 conducted in Beijing (Sun et al., 2010; Zhang et al., 2011; Zhang et al., 2012d), and a comparison of 34 results presented below indicates that a reasonable correction was conducted.

35 Detection limits (DLs) for the AMS mass concentration were evaluated based on the mass 36 spectra of particle-free ambient air, that is, air filtered with a HEPA filter. The DLs were defined as 37 three times the standard deviation of each species signal in the particle-free air. During this study, 38 the DLs for sulfate, nitrate, ammonium, organics and chloride for the 5 min averaged data were 39 $0.03, 0.017, 0.063, 0.101, and 0.01 \,\mu g \,m^{-3}$, respectively (Zhang et al., 2011).

40 The room temperature was controlled at 25 ± 3 °C, 40-60% for relative humidity (RH) and 41 atmospheric air was sampled through a PM₁₀ impactor, which was followed by a PM_{2.5} cyclone (the 42 flow rate was 16.7 L min⁻¹) and dried to a RH <30% with the use of an automatic aerosol dryer unit 43 (Tuch et al., 2009).

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45 2.2 Quality Assurance of the Data

To validate the assumed collection efficiency and demonstrate the data quality of our measurements, a plot of the NR-PM₁ mass concentration from the AMS against the reconstructed mass obtained with the SMPS/TDMPS is shown in Figure 1. Here, the SMPS/TDMPS dry mass concentrations were obtained by converting the measured SMPS/TDMPS number distributions to volume distributions which were then integrated into total volume, multiplied with the aerosol density and then SMPS/TDMPS mass was calculated. The aerosol density used in SMPS/TDMPS mass calculations was assumed to be that of the average composition of $(NH_4)_2SO_4$, NH_4NO_3 and organics, whose densities are 1.77, 1.72, and 1.3 g cm⁻³, respectively. As noted above, fixed collection efficiency (CE) of 0.5 was assumed for the entire campaign, which generally yielded a good correlation between the AMS and reconstructed <u>SMPS/TDMPS</u> mass data. The coefficient of determination (r²) of AMS mass versus <u>SMPS/TDMPS</u> mass was 0.744 with a slope of 0.984, demonstrating a good level of agreement in the masses obtained by the two methods.

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2.3 Data Separation of PBL, LFT, NPF, in cloud, polluted and Different Air-masses

10 Characterization of air masses sampled at the site from either PBL or LFT is important to the 11 subsequent analysis of the chemical data. In this study, the criteria developed by (Gallagher et al., 12 2011) were used to differentiate the data influenced by the PBL versus the LFT. Data during night 13 time (00:00-06:00) were considered to have been influenced by the LFT. The day time between 14 8:00 and 20:00 with enhanced water vapour and high aerosol concentration at the mountain top was considered to be a signal of PBL influence on the mountain. The NPF events were identified based 15 16 on the evolution of the size distributions and particle number concentrations following Kulmala's 17 definitions (Kulmala et al., 2004). The in-cloud events were picked up based on the meteorological data when the relative humidity is in range of 95%-100% with no rain event. The polluted events 18 19 were the periods with high mass concentration and high number concentration for accumulation 20 mode particles.

21 To characterize the long range transport pathways, 72-h back trajectories were computed every 6 22 h (00:00,6:00,12:00 and 18:00 UTC) using the HYSPLIT-4.8 (Hybrid Single Partical Lagrangian 23 Integrated Trajectories) model of the National Oceanic and Atmospheric Administration, USA (NOAA) (http://www.arl.noaa.gov/ready/hysplit4. html). The resolution of the terrain data in the 24 25 HYSPLIT model is 1°×1°, thus the real height of the mountain site has been smoothed. Therefore, 26 1500 m above the model ground level was chosen as trajectory start height, which is about 840 hPa 27 in the model and is a little lower than the pressure of the measurement site (845 hPa). TrajStat 28 (Wang et al., 2009), a program using trajectory statistical analysis data to identify potential source 29 pathways and locations from long-term air pollution measurements, was used to construct clusters 30 from the air-mass trajectories for long range transport pathways study. The Global Data 31 Assimilation System (GDAS) meteorological data archives of the Air Resource Laboratory, NOAA, 32 were used as the input.

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34 3 Results and Discussions35

36 3.1 Mass concentration of chemical components

38 The annual mean mass concentrations of organics, sulfate, nitrate, ammonium and chloride (OSNAC) in PM₁ were 11.2, 9.2, 7.2, 5.8 and 0.95 μ g m⁻³, respectively, totalling ~34.3 μ g m⁻³. This 39 PM₁ concentration is roughly half of the ground-based values at urban Beijing (~ 500 km northwest 40 of the Mt. Tai) (~76 µg m⁻³) (Zhang et al., 2012d), lower than that in sub-urban Gucheng station 41 (~52 µg m⁻³), ~38% of the sub-urban Tianjin level (~ 80 km southeast from Beijing), and lower 42 43 than those at several urban/rural sites in the Pearl River Delta in China (Xiao et al., 2011). It is higher than those in several European cities (10-30 µg m⁻³) (Lanz et al., 2007) and some field 44 campaigns conducted at various ground sites in urban areas, downwind of urban areas, and 45 46 rural/remote locations in the mid-latitudes of the Northern Hemisphere (Zhang et al., 2007), 47 comparable to that of Mexico City (Volkamer et al., 2006). The mass concentrations of chemical 48 components in PM₁ at Mt. Tai are about factor of 2-3 lower (for organics and sulfate), and slightly 49 lower (for nitrate, ammonium and chloride) than those at the near surface rural areas of China 50 (Zhou et al., 2012). The mass loading of Mt. Tai is much higher than those at other elevated sites 51 such as Whistler Mountain (Sun et al., 2009), Mt. Jungfraujoch (Cozic et al., 2008), Himalayan 52 station in Nepal (Decesari et al., 2010). The chemical components of PM_1 in spring for this study

- are about 45% of the previous research results during March-April in 2009 from filter results in
- PM₁₀ (Wang et al., 2011), which was carried out at the same site. But they are larger than those high altitude sites in Europe, Japan, India and USA in TSP, PM_{10} or $PM_{2.5}$ (Table 1).
- 4 Seasonally, the average concentrations of PM₁ in spring, summer, fall and winter were 30 μ g m⁻³, 55 μ g m⁻³, 18 μ g m⁻³ and 37 μ g m⁻³ respectively, which is similar with previous research for PM_{2.5} 5 (Zhou et al., 2009) at Mt. Tai (Table 1), but shows somewhat difference from typical seasonal 6 patterns of winter minimum and summer maximum at Mt. Jungfraujioch (Cozic et al., 2008). For 7 8 Mt. Tai, in summertime, plenty of VOCs, gas phase pollutants, active photochemistry, and stagnant 9 meteorological conditions on regional scale resulted in the high concentration of chemical species 10 within the boundary layer. With enhanced thermally driven convection, the vertical transport of ground pollutants influences the site largely. The minimum concentrations in fall would be related 11 12 to the weak vertical mixing, reduced emission from ground and the active horizontal regional 13 transport from clean places. In winter, although the site was more easily influenced by LFT for 14 weak vertical transport, the pollutants from coal combustion for heating would increase the mass 15 concentration of near surface aerosols and causing relative high concentration at the Mt. Tai 16 through vertical convection.
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Table1. The concentration of main aerosol chemical species at Mt. Tai compared to other Mountain stations. Data are provides in $\mu g \, m^{-3}$.

Location	Height (m)	Period	Size range	Organics	<mark>OC</mark>	EC	<mark>SOC</mark>	Sulfate	Nitrate	Ammonium	Reference
Mt. Tai, China	1534	Annual (2011)	PM ₁	11.2				9.2	7.2	5.8	This study
		Spring (2011)		<mark>8.6</mark>				7.3	8.8	5.6	
		Summer (2010,		16.4				20.1	8.3	11.0	
		2011) Fall (2010,		5.7				5.7	3.8	2.9	
		2011)									
		Winter (2011)	DM	11.6	12.0	2.2	7.0	8.7	9.6	6.8	
		March-April 2009	PM ₁₀		13.0	3.3	7.9	16.0	20.0	12.0	(Wang et 2011)
		From southly			12.0	2.7		21.0	23.0	<mark>16.0</mark>	
		From Northly			13.0	2.8		14.0	18.0	10.0	
		Sprig 2007	PM _{2.5}					12.8	<mark>5.8</mark>	<mark>5.6</mark>	(Zhou et 2009)
		Summer 2007						22.9	4.0	8.0	<mark>2009)</mark>
Mt. Hua, China	<mark>2060</mark>	March-April	PM ₁₀		<mark>5.9</mark>	1.4	2.1	13.0	5.0	2.5	(Wang et a
Mt. Waliguan,	<mark>3816</mark>	2009	TSP					0.2	0.2	0.3	2011)
China	3816	Oct-1994	1 SP					0.2	0.2	0.3	(Yang et a 1996)
Zhuzhang, China	<mark>3583</mark>	Jul.2004-	PM_{10}		<mark>3.1</mark>			1.6	0.5	0.2	(Zhou et a
Mt. Yulong,	3100	Mar.2005 Jan-Feb 2010	TSP					1.8	0.6	0.4	2012) (Zhang et a
China									0.0	<u></u>	2012b)
Lhasa, China	<mark>3363</mark>	Annual 2006	PM ₁₀		21.0	3.7					(Zhang et 2008)
Mt. Whistler,	<mark>2182</mark>	Spring 2006	PM ₁	1.1				<mark>0.6</mark>		0.2	(Sun et
Canada Mt. Junafraviash	2500	July-August	TCD	17				0.1	0.1	0.04	2009)
Mt. Jungfraujoch, Switzerland	<mark>3580</mark>	2005	TSP	1.7				<mark>0.1</mark>	<mark>0.1</mark>	0.04	(Henning al., 2003)
o witzeriana		July-August	PM_1	1.3				0.6	0.1	0.3	(Cozic et
NCO-P, Nepal	<mark>5079</mark>	2005 Apr 2006-May	PM ₁₀		1.5	0.2		0.7	0.3	0.2	2008) (Decesari
100 I, 10pu	<u></u>	2008			1.0	0.2		0.7	<u></u>	·	al., 2010)
			PM ₁								
Mt. Fuji, Japan	<mark>3776</mark>	Jun2001-Aug 2002	TSP					1.2	0.1	0.3	(Suzuki et 2008)
Mt. Abu, India	<mark>1680</mark>	Annual 2008	TSP		3.7	0.5	1.2	2.9	0.6	<mark>0.4</mark>	(Rastogi a
											Sarin, 2005;Ram
											al., 2008)
		Mar 2007-Feb	PM _{2.5}					<mark>3.9</mark>	0.3	1.2	(Kumar a
		2008 Mar 2007-Feb	PM _{2.5-}					0.5	1.1		Sarin, 2010
		<mark>2008</mark>	10								
Mt. Darjeeling, ndia	<mark>2200</mark>	Jan-Dec 2005	PM ₁₀					2.5	1.0	0.1	(Chatterjee al., 2010)
india		Jan-Dec 2005	PM _{2.5}					3.8	3.3	<mark>0.9</mark>	al., 2010)
Manora Peak,	<mark>1950</mark>	Winter 2004	TSP		8.7	1.1	2.3	2.6	0.5		(Rengarajar
ndia Yosemite NP,	1603	Jun-Sep 2002	PILS					1.0	0.3	0.4	et al., 2007) (Lee et
USA	1005	301-30p 2002	metho					1.0	0.5	0.4	(Lee et) 2008)
San Gorgonio,	1705	Apr-2003	d					0.6	3.2	<mark>0.9</mark>	
JSA		Jul-2003						1.3	1.3	1.0	

3.2 Relatively aged chemical components with summer maximum found at the elevated site

4 Figure 2 summarized the average mass size distributions (MSDs) of organics, sulfate, nitrate and 5 ammonium on annual and seasonal basis. The MSDs of these four chemical species in particles 6 with diameters between 30 and 1000 nm were found to be approximately lognormal. The annually 7 mass median diameters (MMDs) for bulk organics, sulfate, nitrate, and ammonium in entire observational period were quite similar with 538, 585, 540 and 541 nm, respectively. The majorities 8 9 of particle mass are with diameter larger than 100 nm. The standard deviations (σ_{g}) of the fitted 10 MSDs were generally smaller than 2, showing that accumulation mode particles significantly impact the MSDs for each chemical component. Seasonally, maximum MMDs for almost all 11 chemical components were found in summer and fall, minimum in spring. During summer-half year. 12 13 high temperature, high relative humidity and strong radiation enhanced oxidation reactions, under the stagnant meteorological condition, the aerosol will stay longer at the atmosphere, and the aged 14 15 pollutants are prone to be transported to the site via the vertical convection process, which results in 16 more aged aerosols. In the winter time, organics and nitrate display smaller MMDs than sulfate, 17 indicating the organics and nitrate are fresher than sulfate. The local emission from coal combustion and traffic is the main sources contributing to organics and nitrate, and the aged sulfate is from the 18 19 regional transport. In addition, the smallest MMDs were found in springtime, which was due to the transport of relatively fresh aerosols from the north. Consequently, the relative larger MMDs were 20 21 also found at this high elevated site than those at several ground sites in the NCP, including the suburban Gucheng station and sub-urban Tianjin (Zhang, 2011), and urban Beijing (Zhang et al., 22 23 2012d), showing the relative aged and regionally-dispersed fine aerosols received compared with 24 those from ground-based measurement.

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3.3 PBL, LFT, NPF, in-cloud and polluted episodes

3.3.1 Mass concentration of chemical components for different episodes

In order to further explore the controlling factors on aerosols at this site, episodes influenced by planet boundary layer (PBL), lower free troposphere (LFT), new particle formation (NPF), enveloped by cloud (in-cloud) and polluted aerosols were classified following methods mentioned in section 2.3 throughout the entire campaign. Overall, the site was influenced by PBL and LFT at same frequency in spring and winter, while the site was influenced more by PBL than LFT in summer and fall. Cloud and polluted events occurred more frequently in summer and fall, and NPF events were observed mostly in spring.

Figure 3 shows the average mass concentration of PM_1 during these episodes. It is shown that 37 38 the highest with 66 µg m⁻³ was from the polluted episode, and lowest was during NPF events from 39 large contribution of ultrafine particles. The mass loading during PBL influence was slight higher than that during LFT (42 µg m⁻³) and in cloud (40µg m⁻³). This LFT mass level is about 2-10 times 40 41 higher than that monitored during night time at Mount. JFJ (Cozic et al., 2008), NCO-P (Decesari et 42 al., 2010) and Murodo (Kido et al., 2001). The relative high concentration at the free troposphere of 43 Mt. Tai indicated that although the boundary layer decented lower than the site, almost 80% 44 pollutants transported by the vertical convections stayed at low free troposphere or residual layer, 45 and the pollutants from ground have impacted the air quality of the mountain site seriously.

The proportions of organics, sulfate, nitrate, ammonium and chloride in different episodes for PM₁ are presented in Figure 4 by pie charts, which show that organics and sulfate contribute distinct between in-cloud and NPF events. During the new particle formation events, the percentage of organics and sulfate was 37% and 22% respectively. Conversely, sulfate dominated the aerosols with 38% and organics contributed less with 24% in cloud episodes. The higher measured sulfate mass concentrations in the cloud residuals were also observed at the top of East Peak mountain in Cape San Juan (Allan et al., 2008). Most sulfate aerosols in the atmosphere are secondary sulfate

1 formed by the oxidation of gaseous precursors (SO₂), followed by particle formation through nucleation and condensation processes. There are several pathways for sulfate formation such as 2 3 liquid-phase reactions inside cloud droplets or oxidation of SO₂ with OH via gaseous phase 4 reactions (Calvo et al., 2012). As the saturation vapour of sulphuric acid is very low, it is not found 5 in the gaseous phase in the troposphere but rather condenses rapidly to form droplets of a sulfuric 6 acid solution. Under normal atmospheric conditions, these particles are partially or totally 7 neutralised by ammonia (NH₃) and in the process, depending on relative humidity, they may 8 become solids (Wang et al., 2008). In this case, ammonium sulfate is more stable in the atmosphere. 9 Once there is sufficient water vapour supplying, it is possible that some of the mass could have 10 been added in the cloud through aqueous processes, the sulfate-dominant particles were easier to be activated as CCN. Under the polluted background and limitation of water vapour, still some 11 interstitial aerosol remained in the cloud and was measured by the AMS. The higher contribution of 12 13 organics in NPF indicates its significant role for the particle formation and growth process. The importance of organics in the atmospheric formation was emphasized in several researches 14 15 (Kulmala et al., 2013). Similar phenomenon was also observed and discussed in more details at 16 ground site campaign in Beijing (Zhang et al., 2011). In terms of nitrate, ammonium and chloride, 17 no dramatic variations on their proportions of PM_1 were observed in these episodes, meaning their minor roles at the different kind of episodes. 18 19

3.3.2 MSDs of chemical components during different episodes

The MSDs of PM₁ (Σ OSNA), chemical species and their proportions at different size ranges during PBL, LFT, NPF, in cloud and polluted episodes are plotted in Figure 4 and Figure 5. Totally, the MSDs of PM₁ for these episodes displayed accumulation mode with 600-700nm MMD, except smaller MMD (550nm) with wider size distribution for NPF. The standard deviation of the fitted MSD for NPF (σ =2.1) was larger than other episodes (σ =1.7-1.8) for the impact of small particles. Same MSDs for other events were found, indicating they may originate from similar sources or evolution process again.

29 Since NPF events being the significant sources of the aerosols, the links between MSDs of NPF 30 and polluted can be employed to investigate the evolution of particles. Statistically results from all 31 the NPF events and polluted events showed the comparatively small MMDs (~373–459 nm) for 32 organics during NPF events at Mt. Tai and the larger MMDs (~473-792 nm) during polluted episodes. The nitrate and ammonium also displayed relative small MMDs during NPF than during 33 34 polluted episodes. No obvious differences of sulfate MMDs between NPF and polluted were found, 35 meaning sulfate was more aged than organics, nitrate and ammonium in NPF. This phenomenon suggests sulfate may originate from regional polluted area, while organics, nitrate and ammonium 36 37 from local sources. The organics increasing factors in MMDs between NPF and polluted air were 38 larger than that of an urban ground site in Beijing (\sim 38-61%), a suburban site in Wuging (\sim 27-32%), 39 and a rural site in Gucheng (8-42%). The increasing factors for sulfate, nitrate and ammonium in 40 Beijing were all lower than that at Mt. Tai. The increasing factors of MMD from NPF to polluted 41 aerosol at Mt. Tai are smaller than those in less polluted areas. At a regional site in Pittsburgh, USA, 42 the increasing factors for organics, sulfate, nitrate and ammonium from a nucleation event to polluted air were 165%, 200%, 29%, 160%, respectively. These data suggest the background level 43 44 of aerosol would influence the evolution of MMDs, and the MMDs at a certain degree, i.e. the 45 pollution level and particle aging.

As the MSDs of chemical species between PBL and LFT were very similar, the LFT's MSDs in Figure 5 was omitted. By investigating the MSDs and percentage of chemical species in different size mode, organics and sulfate were found to present different roles, while nitrate and ammonium displayed relative stable patterns. During PBL, NPF and polluted episodes, organics accounted for major proportions at small mode, and organics was about 70% at 100-200 nm particles in polluted events. Under episodes influenced by PBL and NPF, half of chemical species was organics for the particles around 100 nm, 20% from sulfate, another 20% from nitrate and 10% from ammonium. The larger the particle was, the more the sulfate contributed to the particles. In cloud episodes, sulfate contributed most fractions (60%) to the small particles (100-200nm), and dominated the particles through the whole size range of PM₁. Organics occupied less with 20% at small particles. The percentage of organics increased to maximum at 200-300nm, and fractions of nitrate and ammonium also increased accordingly. Totally, more than 70% of inorganics contributed to the PM₁ in cloud episodes. Due to the hygroscopicity of ammonium sulfate and ammonium nitrate, they were likely grown up by water uptake and partially participated in the cloud formation.

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3.4 Transport pathways and associated chemical component changes

11 Seven main back-trajectory clusters were identified by using TrajStat and the HYSPLIT-4.8 model (Figure 6). Based on the distances of air masses transported, cluster 1, 2, and 3 were defined as 12 13 short pathways, cluster 4 and 6 were mediums, and cluster 5 and 7 were considered as long 14 transport pathway. Cluster-1 (~17% of the total) represents the shortest transport pathway, which 15 was from the south. These trajectories passed over Xuzhou in Anhui Province, an area that suffers 16 from serious pollution, mainly due to biomass burning (Wang et al., 2002; Woo et al., 17 2003; Suthawaree et al., 2010). Tai' an, a prefecture-level city, located on the southern flank of Mt. 18 Tai, is a likely source for pollution when the airflow followed this pathway. Trajectories in Cluster-19 2 (~16%) started over the Yellow Sea and passed over the eastern side of the Shandong Peninsula. 20 Air masses following this relatively short path would bring both marine aerosol and ground-level 21 air pollutants to Mt. Tai. Cluster-3 (another $\sim 25\%$) was from the north/northeast of Mt. Tai. 22 starting at Chengde in Hebei Province and then passing through Bohai sea and Shandong Peninsula. 23 The three short distance clusters account for $\sim 60\%$ of all the air masses, showing the majority of 24 regionally-dispersed aerosol received at Mt. Tai was from short distance mixed aerosol, mostly 25 from its south.

26 Cluster-4 (~7%) shows the airflow from north of China to the NCP, including Tianjin and Hebei 27 Province. These trajectories are representative of regional-scale transport path. Trajectories in 28 Cluster-5 (20%) originate over remote areas in north/north-west China with a transport path above 29 3000 m. These air masses swept through Hebei province to the site at high wind speed, which has 30 assumed delivering most of aerosol in FT. Cluster-6 (~11%) represents regional transport from the 31 west of Mt. Tai. Cluster-7 (4%) represents group of trajectories from the further northwest with 32 longer distances. The air masses in these four clusters account for another ~40% of all transport, 33 and bring less polluted aerosol from north. These can also be supported by the sums of the mass concentrations of the OSNAC of 51, 59 and 48 μ g m⁻³ for the first three clusters aerosols, and 36, 34 16, 34 and 14 μ g m⁻³ for the rest clusters respectively. The mass loadings for Cluster 5 and 7 were 35 the lowest of seven clusters. Even though they are still much higher than those at Whistler Mt. (1.2 36 μ g m⁻³), they are close to a similar level to what has been observed at Korea and Japan (11–13 μ g 37 m⁻³) (Zhang et al., 2007). The MMDs, standard deviations and mass concentrations of chemical 38 39 components corresponding to the seven clusters are summarized in Table 2. Only mass 40 concentration but no size information for chloride was presented because no mass fragmentation of 41 chloride was selected in time-of-flight (TOF) mode during the campaign. The relative smaller 42 MMDs were also found for chemical components associated with long-distance transported 43 aerosols (cluster 4-7), suggesting the relative fresh aerosols. While the larger MMDs for short 44 clusters indicated that the aerosols were more aged.

Organics and sulfate were the two largest fractions, accounting for $\sim 30\%$ each for the $\Sigma OSNAC$ 45 46 in the airflow associated with the first three short-distant paths (cluster 1, 2 and 3) from the south (Table 2), demonstrating again the more pollutants from Pan-Yangtze river delta area. The nitrate 47 was about 20% with mean concentrations of 9.0, 11.2, 10.9 µg m⁻³ for cluster 1, 2 and 3 48 respectively. The ammonium and chloride were the two smallest fractions with $\sim 17\%$ and 2%. 49 50 respectively. The much higher organics were found for aerosol travelled long-distance from north 51 relative to the short-distance moving aerosols from south. About 40% organic were found in 52 aerosols associated with cluster 4, 5, 6 and 7.

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Air mass	Organics			Sulfate			Nitrate			Ammonium			Chloride
	MMD	$\sigma_{\rm g}$	Mass	MMD	$\sigma_{\rm g}$	Mass	MMD	$\sigma_{\rm g}$	Mass	MMD	σ_{g}	mass	Mass
Cluster-1	545	1.7	16.8	570	1.7	15.6	582	1.7	9.0	542	1.7	8.7	0.9
Cluster-2	533	1.6	18.2	584	1.7	17.6	581	1.6	11.2	553	1.7	10.2	1.0
Cluster-3	511	1.7	16.4	528	1.8	11.4	514	1.7	10.9	497	1.8	8.1	1.2
Cluster-4	464	1.7	12.8	517	1.7	9.1	470	1.7	7.4	478	1.7	5.8	0.8
Cluster-5	458	1.6	6.8	515	2.0	3.7	462	2.1	2.9	440	2.0	2.5	0.5
Cluster-6	417	1.8	12.7	464	2.0	6.8	435	1.8	8.0	437	1.9	5.4	1.0
Cluster-7	426	2.0	5.2	459	1.9	3.4	357	2.3	2.5	395	2.2	2.1	0.4

Table2. Mass median diameters (MMD, nm, standard deviations (σ_g) and mass concentration (Mass, $\mu g m^{-3}$) for chemical components in samples grouped by air-mass trajectory cluster

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3.5 Composition of OA

9 To further investigate the secondary OA, AMS-PMF modelling of OA spectra (Aiken et al., 2009; Ulbrich et al., 2009) was used to identify presumptive sources for the organic aerosol during 10 11 each of four seasons at Mt Tai. Different types of organics were resolved based on specific massspectral profiles, and the mass concentrations are presented in Table 3. The mass spectra for 12 13 Hydrocarbon-like OA (HOA) is characterized by hydrocarbon ions of the general form C_nH_{2n+1} and C_nH_{2n-1} , including $C_3H_7^+$ (m/z 43), $C_4H_7^+$ (m/z 55), and $C_4H_9^+$ (m/z 57). A signal from m/z 60, a 14 15 tracer ion for biomass burning organic aerosol (BBOA) (Alfarra et al., 2006; Aiken et al., 2009), 16 which can be attributable to POA (Jimenez et al., 2009), was resolved in the mass spectra from spring, summer and fall. In winter, another factor with high m/z 43 and m/z 60 was identified as the 17 18 coal-combustion organic aerosol (CCOA) (Sun et al., 2013). The mass concentration of subtype 19 organics and their percentage in organic matter are present in Table 3. HOA, BBOA and CCOA can be considered as primary organic matter (POA) directly emitted into the atmosphere. The POA 20 21 dominated OA during wintertime with 59% of organics, and accounting for 51%, 44% and 48% of 22 OA in spring, summer, fall and winter, respectively. Same result was reported at a ground site in urban site Beijing in winter (Sun et al., 2013). BBOA was derived from AMS database by PMF 23 24 model in spring, summer and fall. June every year in China is the high season for harvest of wheat, 25 and also the period when straw burning takes place seriously in some part of the country (Qu et al., 26 2012). Previous study mentioned that field burning of wheat straws in the North China Plain during May–June 2006 in urban areas such as Beijing (Li et al., 2007). Study based on organic molecular 27 28 compositions of Mt. Tai also identified levoglucosan in summer (Fu et al., 2008). It is deemed that 29 agriculture residues burning, crop residues burning are the main sources of BBOA influenced the site. On the other hand, as a tourist spot, burning incense is a feature of Mt. Tai. Although some 30 31 special incense burning events including weekends, holidays, and traditional festivals have been deleted in the datasets, incense burning probably is another BBOA source influence Mt. Tai on 32 normal days, which needs further investigations. Assuming the emission of incense burning was 33 34 constant, more BBOA was measured in summer than spring and fall, which suggests intensive emission from field burning of agriculture residues in summer. In winter, 37% of organics was 35 identified as HOA and 22% as CCOA. Coal combustion in China has been found to emit a large 36 37 quantity of carbonaceous aerosols, contributing 70% of total emitted PM2.5 (Zhang et al., 2012a). It

is widely accepted that HOA is mainly associated with combustion-related emissions, e.g., diesel exhaust (Sun et al., 2013). The relatively high concentration of HOA in winter should result from the low ambient temperature, low atmospheric oxidants, and most of fresh traffic exhaust was transported to the site with less oxidizations.

42 At Mt. Tai, one can still find other large fraction of oxygenated organic aerosol (OOA) that is 43 considered to be secondary generally (Herndon et al., 2008;Volkamer et al., 2006), in which two

subtypes of OOA can be distinguished by large fraction of CO_2^+ (m/z 44) for low-volatility oxidized 1 organic aerosol (LV-OOA) and by $C_{3}H_{7}^{+}$ (m/z 43), CO_{2}^{+} (m/z 44), $C_{3}H_{3}O^{+}$ (m/z 55) and $C_{3}H_{5}O^{+}$ 2 (m/z 57) for semi-volatile oxidized organic aerosol (SV-OOA). LV-OOA is strongly correlated with 3 non-volatile secondary species such as sulfate and has a high O:C, mainly attributable to regional, 4 5 heavily aged OA; SV-OOA has a higher correlation with semi-volatile species such as ammonium 6 nitrate and ammonium chloride and has a lower O:C, consistent with less-photo chemically fresh 7 OA (Jimenez et al., 2009). SV-OOA and LV-OOA together contributed 49%, 55%, 51% and 41% 8 from spring to winter. More OOA was observed in summer which attributed to the strong 9 photochemical reaction, sufficient oxidization process along with regional disperse in the CEC. 10 This identification about relative contributions of POA and SOA is similar to a 2-yr filter 11 measurement from 16 sites in various regions of China, which reported that the secondary organic 12 carbon contributed ~55% and 60% for urban and rural aerosol, respectively(Zhang et al., 2012c).

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Table3. Seasonal averaged mass concentrations (unit: $\mu g m^{-3}$) and standard deviation with percentage (in parentheses) of specific types of organic aerosols from Positive matrix factorization analysis

Sub-type organics ^a	Spring	Summer	<mark>Fall</mark>	Winter
Primary OA				
HOA	2.3±1.2 (27%)	<mark>1.6±1.8</mark> (10%)	<mark>0.60±0.38</mark> (11%)	<mark>5.6±1.6</mark> (37%)
BBOA	<mark>2.1±1.1</mark> (24%)	5.1±10.7 (34%)	1.1±0.65 (19%)	
CCOA			<mark>1.1±0.61</mark> (19%)	3.3±2.2 (22%)
Secondary OA				
SV-OOA		3.1±4.2 (21%)		
LV-OOA	4.2±2.3 (49%)	5.3±6.4 (35%)	2.9±1.59 (51%)	6.1±3.56 (41%)

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18 19 20 ^aAbbreviations: HOA = hydrocarbon-like organic aerosol, BBOA = biomass-burning organic aerosol, CCOA= coal combustion organic aerosol, semi-volatile oxidized organic aerosol = SV-OOA, and low-volatility oxidized organic aerosol = LV-OOA.

22 4 Summary 23

24 The mass concentrations and size distributions of chemical components in PM₁ were characterized 25 in situ with an AMS at the summit of Mt. Tai (~1500 m a.s.l) from June 2010 to January 2012.

26 The mass concentration of organics, sulfate, nitrate, ammonium and chloride (OSNAC) in PM₁ 27 at Mt. Tai is higher than those at most mountain sites in the USA and Europe, but comparable with 28 previous research results at Mt. Tai. Seasonally, high concentrations of all chemical components in ~1500 m high level were found in summer with minimum in fall, slightly different from typical 29 30 seasonal patterns in various aerosol components at ground-based measurement in China ("spring 31 minimum"). This seasonal pattern was controlled by the development of PBL, horizontal 32 transportation and local emissions.

33 The MSDs for organics, sulfate, nitrate and ammonium were approximately log-normal with 34 more than 99% mass fraction attributable to the particle with diameter larger than 100 nm. Different 35 from the seasonality of mass concentrations, maximum MMDs for almost all chemical components were found in summer and secondary high values in fall with minimum in spring, indicating the 36 37 summer-half of the year was more conducive for the formation of aged fine aerosols. The larger 38 MMDs are found at Mt. Tai than those in ground sites, showing the relative aged and well-mixed 39 aerosol observed. The smaller MMDs were found for organics of the NPF compared to the aged 40 aerosols with MMDs of ~473–792 nm. Compared with increasing factor of organics between NPF

41 and polluted events at Mt. Tai, small factors have been reported at an urban ground site in Beijing

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(~38-61%), a suburban site in Wuqing (~27-32%), and a rural site in Gucheng (8-42%). Larger
increasing factors for sulfate, nitrate and ammonium were found in less polluted site in Europe and
USA. The MMDs of chemical species could indicate the polluted level and particle aging to a
certain degree.
Five kinds of episodes influenced by PBL, LFT, NPF, in-cloud and polluted aerosols at Mt.

6 Tai were classified. The highest mass concentration was observed during polluted aerosols with 7 large MMD and lowest in NPF with relatively small MMD. The LFT was 80% of PBL, indicating 8 most of pollutants in PBL could be transported to the low free troposphere. For chemical species, 9 organics dominated the PBL, NPF and polluted events and sulfate dominated in cloud episodes. 10 There were no obvious variations on proportions of nitrate and ammonium in five kinds of episodes. 11 In cloud, inorganics contributed 70% to the PM₁ at the whole size range, and sulfate dominated with 40% to the small mode particles, while organics is 20% of small mode particles. The larger 12 13 quantities of inorganics in particles make it easier to add into the cloud through water uptake.

14 The air-mass back trajectories were grouped into seven clusters. Three of the clusters regional 15 exemplified transport from the north, east, and south of Mt. Tai; two medium length clusters were 16 from the northwest and west of site, and two long-range clusters showed airflow from the northwest 17 of the site. Shorter transport pathways were corresponded with higher aerosol mass concentrations. The analysis of transport showed that the air quality at Mt. Tai was impacted by the pollution from 18 19 the Shandong Peninsula, the NCP and areas immediately north of Suzhou. The air masses 20 represented by clusters 5 and 7 brought clean, dry air from the northwest, which swept pollution 21 away from the site.

22 At Mt. Tai, one can still find the large fraction (normally half) of total OA can be attributable 23 to oxygenated organic aerosol (OOA) that is considered to be secondary generally. Extra high 24 proportions of SOA found in summer, presenting the intensive solar radiation and oxidized process. 25 Regionally, heavily aged OA are observed with >40% OOA. Hydrocarbon-like OA (HOA), 26 biomass burning organic aerosol (BBOA) and coal-combustion organic aerosol (CCOA) were 27 considered as primary organic matter (POA) directly emitted into the atmosphere, and accounted 28 together for 51%, 44%, 48% and 59% of the total measured organics in spring, summer, fall and 29 winter, respectively. Considerable amount of BBOA in summer and CCOA in winter were found 30 with 34% and 22% of OA respectively.

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- 5 References
- 6

- Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M., Docherty,
 K. S., Sueper, D., Kimmel, J. R., and Worsnop, D. R.: Mexico City aerosol analysis during
 MILAGRO using high resolution aerosol mass spectrometry at the urban supersite (T0)-Part 1:
- MILAGRO using high resolution aerosol mass spectrometry at the urban supersite (10)-Part 1:
 Fine particle composition and organic source apportionment, Atmos. Chem. Phys., 9, 6633-6653,
 2009.
- Albrecht, B. A.: Aerosols, cloud microphysics, and fractional cloudiness, Science, 245, 1227-1230,
 1989.
- Alfarra, M. R., Paulsen, D., Gysel, M., Garforth, A. A., Dommen, J., Prévôt, A. S. H., Worsnop, D.
 R., Baltensperger, U., and Coe, H.: A mass spectrometric study of secondary organic aerosols
 formed from the photooxidation of anthropogenic and biogenic precursors in a reaction chamber,
 Atmos. Chem. Phys., 6, 5279-5293, 2006.
- Allan, J. D., Baumgardner, D., Raga, G. B., Mayol-Bracero, O. L., Morales-García, F., GarcíaGarcía, F., Montero-Martinez, G., Borrmann, S., Schneider, J., and Mertes, S.: Clouds and
 aerosols in Puerto Rico-a new evaluation, Atmos. Chem. Phys., 8, 1293-1309, 2008.
- Boucher, O., Randall, D., Artaxo, P., Bretherton, C., Feingold, G., Forster, P., Kerminen, V.-M.,
 Kondo, Y., Liao, H., Lohmann, U., Rasch, P., Satheesh, S. K., Sherwood, S., Stevens, B., and
 Zhang, X. Y.: Clouds and Aerosols, in: Climate Change 2013: The Physical Science Basis.
 Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel
 on Climate Change, edited by: Stocker, T. F., D. Qin, G.-K. Plattner, M. Tignor, S. K. Allen, J.
 Boschung, A. Nauels, Y. Xia, V. Bex and P. M. Midgley, Cambridge University Press, New
 York, 2013.
- Calvo, A. I., Alves, C., Castro, A., Pont, V., Vicente, A. M., and Fraile, R.: Research on aerosol
 sources and chemical composition: Past, current and emerging issues, Atmos. Res., 120-121, 1 28, 2012.
- Charlson, R. J., Schwartz, S. E., Hales, J. M., Cess, R. D., Coakley, j. J. A., Hansen, J. E., and
 Hofmann, D. J.: Climate forcing by anthropogenic aerosols, Science, 255, 423-430, 1992.
- Chatterjee, A., Adak, A., Singh, A. K., Srivastava, M. K., Ghosh, S. K., Tiwari, S., Devara, P. C.,
 and Raha, S.: Aerosol chemistry over a high altitude station at northeastern Himalayas, India,
 PloS one, 5, e11122, 2010.
- Cozic, J., Verheggen, B., Weingartner, E., Crosier, J., Bower, K., Flynn, M., Coe, H., Henning, S.,
 Steinbacher, M., and Henne, S.: Chemical composition of free tropospheric aerosol for PM1 and
 coarse mode at the high alpine site Jungfraujoch, Atmos. Chem. Phys., 8, 407-423, 2008.
- Decesari, S., Facchini, M. C., Fuzzi, S., McFiggans, G. B., Coe, H., and Bower, K. N.: The water soluble organic component of size-segregated aerosol, cloud water and wet depositions from Jeju
 Island during ACE-Asia, Atmos. Environ., 39, 211-222, 2005.
- Decesari, S., Facchini, M., Carbone, C., Giulianelli, L., Rinaldi, M., Finessi, E., Fuzzi, S., Marinoni,
 A., Cristofanelli, P., and Duchi, R.: Chemical composition of PM 10 and PM 1 at the highaltitude Himalayan station Nepal Climate Observatory-Pyramid (NCO-P)(5079 m asl), Atmos.
 Chem. Phys., 10, 4583-4596, 2010.
- Fu, P., Kawamura, K., Okuzawa, K., Aggarwal, S. G., Wang, G., Kanaya, Y., and Wang, Z.:
 Organic molecular compositions and temporal variations of summertime mountain aerosols over
 Mt. Tai, North China Plain, J. Geophys. Res., 113, D19107, doi:10.1029/2008JD009900, 2008.
- Fu, P., Kawamura, K., Kanaya, Y., and Wang, Z.: Contributions of biogenic volatile organic
 compounds to the formation of secondary organic aerosols over Mt. Tai, Central East China,
 Atmos. Environ., 44, 4817-4826, 2010.

- Gallagher, J. P., McKendry, I. G., Macdonald, A. M., and Leaitch, W. R.: Seasonal and diurnal variations in aerosol concentration on Whistler Mountain: Boundary layer influence and synoptic-scale controls, J. Appl. Meteorol., 50, 2210-2222, 2011.
- Henning, S., Weingartner, E., Schwikowski, M., Gäggeler, H., Gehrig, R., Hinz, K. P., Trimborn,
 A., Spengler, B., and Baltensperger, U.: Seasonal variation of water soluble ions of the aerosol
 at the high alpine site Jungfraujoch (3580 m asl), J. Geophys. Res., 108, 4030,
 doi:4010.1029/2002JD002439, 2003.
- 8 Herndon, S. C., Onasch, T. B., Wood, E. C., Kroll, J. H., Canagaratna, M. R., Jayne, J. T., ..., and
 9 Worsnop, D. R.: Correlation of secondary organic aerosol with odd oxygen in Mexico City,
 10 Geophys. Res. Lett., 35, L15804, 2008.
- Jayne, J. T., Leard, D. C., Zhang, X., Davidovits, P., Smith, K. A., Kolb, C. E., and Worsnop, D. R.:
 Development of an aerosol mass spectrometer for size and composition analysis of submicron
 particles, Aerosol.Sci.TechnoL., 33, 49-70, 2000.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., and Worsnop, D. R.: Evolution of Organic
 Aerosols in the Atmosphere, Science, 326, 1525, DOI: 10.1126/science.1180353, 2009.
- Kido, M., Osada, K., Matsunaga, K., and Iwasaka, Y.: Diurnal variation of ionic aerosol species and
 water soluble gas concentrations at a high elevation site in the Japanese Alps, J. Geophys.
 Res., 106, 17335-17345, 2001.
- Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W.,
 and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a review of
 observations, J. Aerosol. Sci., 35, 143-176, 2004.
- Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T.,
 Sipilä, M., Schobesberger, S., and Rantala, P.: Direct observations of atmospheric aerosol
 nucleation, Science, 339, 943-946, 2013.
- Kumar, A., and Sarin, M.: Atmospheric water-soluble constituents in fine and coarse mode aerosols
 from high-altitude site in western India: Long-range transport and seasonal variability, Atmos.
 Environ., 44, 1245-1254, 2010.
- Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and Prévôt, A. S. H.:
 Source apportionment of submicron organic aerosols at an urban site by factor analytical
 modelling of aerosol mass spectra, Atmos. Chem. Phys., 7, 1503-1522, 2007.
- Lee, T., Yu, X. Y., Kreidenweis, S. M., Malm, W. C., and Collett, J. L.: Semi-continuous
 measurement of PM2.5 ionic composition at several rural locations in the United States, Atmos.
 Environ., 42, 6655-6669, 2008.
- Li, W. J., Zhang, D. Z., Shao, L. Y., Zhou, S. Z., and Wang, W. X.: Individual particle analysis of
 aerosols collected under haze and non-haze conditions at a high-elevation mountain site in the
 North China plain, Atmos. Chem. Phys., 11, 11733-11744, 2011.
- Li, X., Wang, S., Duan, L., Hao, J., Li, C., Chen, Y., and Yang, L.: Particulate and trace gas
 emissions from open burning of wheat straw and corn stover in China, Environ. Sci. TechnoL.,
 41, 6052-6058, 2007.
- Qu, C., Li, B., Wu, H., and Giesy, J. P.: Controlling air pollution from straw burning in China calls
 for efficient recycling, Environ. Sci. TechnoL., 46, 7934-7936, 2012.
- Ram, K., Sarin, M., and Hegde, P.: Atmospheric abundances of primary and secondary
 carbonaceous species at two high-altitude sites in India: Sources and temporal variability, Atmos.
 Environ., 42, 6785-6796, 2008.
- Ramgolam, K., Favez, O., Cachier, H., Gaudichet, A., Marano, F., Martinon, L., and Baeza-Squiban,
 A.: Size-partitioning of an urban aerosol to identify particle determinants involved in the
 proinflammatory response induced in airway epithelial cells, Part. Fibre. Toxicol., 6, 1-12, 2009.
- Rastogi, N., and Sarin, M.: Long-term characterization of ionic species in aerosols from urban and
 high-altitude sites in western India: Role of mineral dust and anthropogenic sources, Atmos.
- 50 Environ., 39, 5541-5554, 2005.

- Ren, Y., Ding, A., Wang, T., Shen, X., Guo, J., Zhang, J., Wang, Y., Xu, P., Wang, X., and Gao, J.:
 Measurement of gas-phase total peroxides at the summit of Mount Tai in China, Atmos.
 Environ., 43, 1702-1711, 2009.
- Rengarajan, R., Sarin, M., and Sudheer, A.: Carbonaceous and inorganic species in atmospheric
 aerosols during wintertime over urban and high altitude sites in North India, J. Geophys. Res.,
 112, D21307, doi:21310.21029/22006JD008150, 2007.
- Seinfeld, J. H., Carmichael, G. R., Arimoto, R., Conant, W. C., Brechtel, F. J., Bates, T. S., Cahill,
 T. A., Clarke, A. D., Doherty, S. J., and Flatau, P. J.: ACE-ASIA-Regional climatic and atmospheric chemical effects of Asian dust and pollution, B. Am. Meteorol. Soc., 85, 367-380, 2004.
- Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Averyt, K. B., Tignor, M., and Miller,
 H. L.: The physical science basis, Contribution of working group I to the fourth assessment
 report of the intergovernmental panel on climate change, 235-337, 2007.
- Sun, J. Y., Zhang, Q., Canagaratna, M. R., Zhang, Y. M., Ng, N. L., Sun, Y. L., Jayne, J. T., Zhang,
 X. C., Zhang, X. Y., and Worsnop, D. R.: Highly time-and size-resolved characterization of
 submicron aerosol particles in Beijing using an Aerodyne Aerosol Mass Spectrometer, Atmos.
 Environ., 44, 131-140, 2010.
- Sun, Y. L., Zhang, Q., Macdonald, A. M., Hayden, K., Li, S. M., Liggio, J., Liu, P. S. K., Anlauf, K.
 G., Leaitch, W. R., and Steffen, A.: Size-resolved aerosol chemistry on Whistler Mountain,
 Canada with a high-resolution aerosol mass spectrometer during INTEX-B, Atmos. Chem. Phys.,
 9, 3095-3111, 2009.
- Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.: Aerosol composition, sources and processes during wintertime in Beijing, China, Atmos. Chem. Phys., 13, 4577-4592, 2013.
- Suthawaree, J., Kato, S., Okuzawa, K., Kanaya, Y., Pochanart, P., Akimoto, H., Wang, Z., and Kajii,
 Y.: Measurements of volatile organic compounds in the middle of Central East China during
 Mount Tai Experiment 2006 (MTX2006): observation of regional background and impact of
 biomass burning, Atmos. Chem. Phys., 10, 1269-1285, 2010.
- Suzuki, I., Hayashi, K., Igarashi, Y., Takahashi, H., Sawa, Y., Ogura, N., Akagi, T., and Dokiya, Y.:
 Seasonal variation of water-soluble ion species in the atmospheric aerosols at the summit of Mt.
 Fuji, Atmos. Environ., 42, 8027-8035, 2008.
- Tuch, T. M., Haudek, A., Müller, T., Nowak, A., Wex, H., and Wiedensohler, A.: Design and
 performance of an automatic regenerating adsorption aerosol dryer for continuous operation at
 monitoring sites, Atmos. Meas. Tech., 2, 1143-1160, 2009.
- 35 Twomey, S.: Pollution and the planetary albedo, Atmos. Environ., 8, 1251-1256, 1974.
- Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of
 organic components from Positive Matrix Factorization of aerosol mass spectrometric data,
 Atmos. Chem. Phys., 9, 2891-2918, 2009.
- Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T.,
 Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from anthropogenic air
 pollution: Rapid and higher than expected, Geophys. Res. Lett., 33, L17811,
 doi:10.1029/2006GL026899, 2006.
- Wang, G., Li, J., Cheng, C., Hu, S., Xie, M., Gao, S., Zhou, B., Dai, W., Cao, J., and An, Z.:
 Observation of atmospheric aerosols at Mt. Hua and Mt. Tai in central and east China during
 spring 2009-Part 1: EC, OC and inorganic ions, Atmos. Chem. Phys, 11, 4221-4235, 2011.
- Wang, J., Hoffmann, A. A., Park, R. J., Jacob, D. J., and Martin, S. T.: Global distribution of solid
 and aqueous sulfate aerosols: Effect of the hysteresis of particle phase transitions, J. Geophys.
 Res., 113, D11206, doi:11210.11029/12007JD009367, 2008.
- Wang, T., Cheung, T. F., Li, Y. S., Yu, X. M., and Blake, D. R.: Emission characteristics of CO,
 NOx, SO2 and indications of biomass burning observed at a rural site in eastern China, J.
 Geophys. Res., 107, D12, 4157, 10.1029/2001JD000724, 2002.

 Wang, Y. Q., Zhang, X. Y., and Draxler, R. R.: TrajStat: GIS-based software that uses various trajectory statistical analysis methods to identify potential sources from long-term air pollution measurement data, Environ. Model. Soft., 24, 938-939, 2009.

- Woo, J., Streets, D. G., Carmichael, G. R., Tang, Y., Yoo, B., Lee, W., Thongboonchoo, N.,
 Pinnock, S., Kurata, G., and Uno, I.: Contribution of biomass and biofuel emissions to trace gas
 distributions in Asia during the TRACE-P experiment, J. Geophys. Res., 108, 2156-2202, 2003.
- Xiao, R., Takegawa, N., Zheng, M., Kondo, Y., Miyazaki, Y., Miyakawa, T., Hu, M., Shao, M.,
 Zeng, L., and Gong, Y.: Characterization and source apportionment of submicron aerosol with
 aerosol mass spectrometer during the PRIDE-PRD 2006 campaign, Atmos. Chem. Phys., 11,
 6911-6929, 2011.
- Yamaji, K., Li, J., Uno, I., Kanaya, Y., Irie, H., Takigawa, M., Komazaki, Y., Pochanart, P., Liu, Y.,
 and Tanimoto, H.: Impact of open crop residual burning on air quality over Central Eastern
 China during the Mount Tai Experiment 2006 (MTX2006), Atmos. Chem. Phys., 10, 7353-7368,
 2010.
- Yang, D., Yu, X., Fang, X., Wu, F., and Li, X.: A study of aerosol at regional background stations
 and baseline station, J. Appl. Meteorol., 7, 396–405, 1996
- Zhang, H., Wang, S., Hao, J., Wan, L., Jiang, J., Zhang, M., Mestl, H. E. S., Alnes, L. W. H.,
 Aunan, K., and Mellouki, A. W.: Chemical and size characterization of particles emitted from
 the burning of coal and wood in rural households in Guizhou, China, Atmos. Environ, 51, 94–99,
 2012a.
- Zhang, N., Cao, J., Ho, K., and He, Y.: Chemical characterization of aerosol collected at Mt.
 Yulong in wintertime on the southeastern Tibetan Plateau, Atmos. Res., 107, 76-85, 2012b.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R.,
 Takami, A., Middlebrook, A. M., and Sun, Y. L.: Ubiquity and dominance of oxygenated species
 in organic aerosols in anthropogenically-influenced Northern Hemisphere midlatitudes, Geophys.
 Res. Lett., 34, L13801, 2007.
- Zhang, X. Y., Wang, Y. Q., Zhang, X. C., Guo, W., Niu, T., Gong, S. L., Yin, Y., Zhao, P., Jin, J.
 L., and Yu, M.: Aerosol monitoring at multiple locations in China: contributions of EC and dust to aerosol light absorption, Tellus. B., 60, 647-656, 2008.
- Zhang, X. Y., Wang, Y. Q., Niu, T., Zhang, X. C., Gong, S. L., Zhang, Y. M., and Sun, J. Y.:
 Atmospheric aerosol compositions in China: spatial/temporal variability, chemical signature,
 regional haze distribution and comparisons with global aerosols, Atmos. Chem. Phys., 12, 779doi:10.5194/acp-12-779-2012, 2012c.
- Zhang, Y. M.: Characterization of sub-micron aerosol and its change processes in BIV (Beijing and
 its vicinity) region, PhD, Chinese Academy of Meteorological Sciences, Beijing, 2011.
- Zhang, Y. M., Zhang, X. Y., Sun, J. Y., Lin, W. L., Gong, S. L., Shen, X. J., and Yang, S.:
 Characterization of new particle and secondary aerosol formation during summertime in Beijing,
 China, Tellus. B., 63, 382-394, 2011.
- Zhang, Y. M., Sun, J. Y., Zhang, X. Y., Shen, X. J., Wang, T. T., and Qin, M. K.: Seasonal characterization of components and size distributions for submicron aerosols in Beijing, Sci.
 China. Earth. Sci., 1-11, 2012d.
- Zhou, C. H., Gong, S., Zhang, X. Y., Liu, H. L., Xue, M., Cao, G. L., An, X. Q., Che, H. Z., Zhang,
 Y. M., and Niu, T.: Towards the improvements of simulating the chemical and optical properties
 of Chinese aerosols using an online coupled model-CUACE/Aero, Tellus. B., 64, 1-20, 2012.
- Zhou, Y., Wang, T., Gao, X., Xue, L., Wang, X., Wang, Z., Gao, J., Zhang, Q., and Wang, W.:
 Continuous observations of water-soluble ions in PM2. 5 at Mount Tai (1534 m asl) in centraleastern China, J. Atmos. Chem., 64, 107-127, DOI 10.1007/s10874-010-9172-z, 2009.
- 49 50
- 51

⁴ Watson, J.: Visibility: Science and Regulation, J. Air Waste Manage., 52, 628-713, 2002.

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Figure Captions

- 4 Fig. 1. Correlation between AMS and (reconstructed) SMPS mass concentrations
- 5 Fig. 2. Mass size distributions of chemical species in annually and seasonally scale
- 6 **Fig. 3.** Mass concentrations of chemical components in different episodes
- 7 **Fig. 4.** The MSDs of PM_1 in different episodes
- 8 Fig. 5. The MSDs and proportions of chemical species in different episodes
- 9 Fig. 6. Averaged mass concentrations, mass-size distribution and percentage of chemical components associating
- 10 with different air-mass trajectory clusters
- 11





Fig.2. 1 2





2 <mark>Fig.3.</mark>



Fig.4.







<mark>Fig. 6</mark>



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