

We thank the two anonymous referees and Dr. Berresheim for their comments. We respond (in italics) to each point separately below. When appropriate, the responses also list all the relevant changes made in the revised manuscript.

Anonymous Referee #1:

This is a well structured paper which builds on previous radical measurements made in Polar regions, particularly at the South Pole, and highlights that the elevated oxidative capacity reported at the SP extends to other regions of Antarctica. The authors highlight the dominance of HONO as an OH source, but demonstrate inconsistencies in the HONO measurements and the observed radical concentrations. Using this approach, the authors conclude that the LOPAP technique, used to make the HONO observations, may suffer from an artefact under these conditions; similar conclusions are reached, using an alternative approach, in a companion paper. Inconsistencies in the NO:NO₂ ratio and peroxy radical concentration observed are also highlighted. This paper is well suited for publication in Atmospheric, Chemistry and Physics and I only have some minor comments and clarifications listed below which should be addressed prior to publication:

1) Pg 15005, ln 25: Repetition of ln 11.

The note about local time in line 25 has been removed.

2) Pg 15006, ln 28: this manuscript would benefit from brief description of the role of the ‘radical quencher (NO₂)’ here.

Added:

“NO₂ used as a scavenger removes not only the OH radicals, but also peroxy radicals converting them into HO₂NO₂ and RO₂NO₂ nitrates”.

3) Pg 15007, ln 9: the manuscript would benefit from expanding briefly on the two OH measurements modes here – their purpose, did the two modes agree?

Added:

“Ratio of the signals with the short and the long conversion times may be used as an indicator of an artificial OH formation in the reactor [Kukui et al., 2008].”

Pg 15008, ln 12 – 15: Did the humidity change from point of humidification to end of the calibrator? Were any changes in [H₂O] accounted for?

The [H₂O] was controlled with the humidity sensor at the entrance into the photolysis reactor. We added on line 10:

“...the humidity measurements in the photolysis reactor”

Pg 15008, ln 20: The modelled HO₂:RO₂ ratio is dependent on the [CO] and [CH₄] assumed. As neither CO nor CH₄ were measured, what is the level of uncertainty in the radical ratio and overall [RO₂] determined from estimating these values?

As discussed in Section 3.2 the uncertainty of about 10% (1σ) for the HO₂/RO₂ ratio was estimated using uncertainties of the measurements at Dome C. For the uncertainty of 10% was adopted for [CO] (Section 3.2). For [CH₄] the uncertainty of 5% was used.

Pg 15009, ln 7: Why was the uncertainty greater during the night time?

During the night time the relative uncertainty was higher due to the lower measured signals.

Pg 15009, ln 9: Please provide the LOD for RO₂ also.

Added:

“The lower limits of detection for OH and RO₂ radicals at signal-to-noise ratio of 3 and a 2 minute integration time were 5×10^5 molecule cm^{-3} and 2×10^6 molecule cm^{-3} , respectively.”

Pg 15010, ln 26: this interference in the HONO measurements of 10 – 20 pptv needs to be followed by the [HO₂NO₂] assumed by Legrand et al. How does this estimated [HO₂NO₂] compare with the concentration calculated by the 0D model?

The corresponding paragraph is modified:

“Legrand et al. (this issue) report tests done both in the field and in the lab that tend to suggest an overestimation of HONO measurements in the range of 10 to 20 pptv due to the presence of HO₂NO₂ in the range of 50-100 pptv in the cold atmosphere at Dome C.. This range of HO₂NO₂ mixing ratios is in agreement with the median [HO₂NO₂] of 80 pptv estimated from RO₂ and NO₂ levels measured at Dome C (see Section 3.2). Also, as discussed by Legrand et al. (this issue), similar levels of HO₂NO₂ were previously observed in Antarctica.”

Pg 15011, ln 14: I believe ‘Sect 2.2’ should be ‘Sect 3.2’

Corrected

Pg 15012: Following on from the description of the 1D model used to estimate the vertical distribution of HONO, I think it would be pertinent to state at this point the respective measurement heights of the radicals and HONO and, if different, the change in [HONO] estimated by the 1D model between the two heights.

The measurement heights for radicals (3m) and HONO (1m) are given in Sections 2.1 and 2.2, respectively. The 1D calculated HONO concentrations are presented in Figure 10.

Pg 15016, ln 16: ‘60%’ should be ‘78%’ according to fig 6.

Corrected:

“The variability of $P_{\text{HO}_2+\text{NO}}$ and P_{HONO} then explain ~80% of the variability of OH.”

Pg 15016, ln 22: remove comma after ‘etc.’

Corrected.

Pg 15019, ln 9: change to ‘reduced by factors of 2, 4..’

Corrected.

Pg 15019, ln 18: change ‘0.25’ to ‘4’

Corrected.

Section 3.5: Along with the comparison of HONO calculated with the 1D model and estimated from the radical budget, this section would benefit from a discussion on the expected HO₂NO₂ diurnal profile – what profile does the 0D model estimate for example? Does this profile suggest that the correction needing to be applied will vary diurnally?

The calculated 0D model diurnal profile of HO₂NO₂ (Figure A) does not exhibit any clear diurnal variation. We therefore think that, in absence of HO₂NO₂ measurements, further discussions based on these calculations would be too speculative.

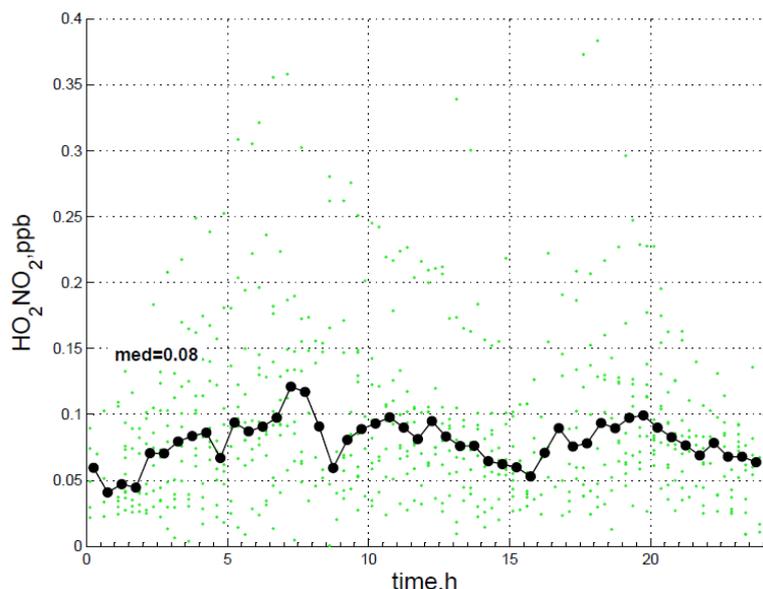


Figure A. *Calculated diurnal profile of HO₂NO₂*

Pg 15022, ln 2: I can't find a reference to the proposed gas-phase source of HONO from HO₂.H₂O + NO₂ in Legrand et al. As highlighted by H Berresheim, the impact of this reaction as a sink for HO₂ needs to be discussed in the manuscript given that the recycling of HO₂ to OH is the second most important OH source.

This hypothetical HONO source would contribute for 10-20% of the HONO production from the reaction OH+NO and would result in less than 1% of the measured HONO. The discussion of HONO formation via reaction HO₂(H₂O)+NO₂ is added in the revised version of Legrand et al. (Section 5) For information we copy the corresponding text at the end of this document in the Appendix A.

Estimated with the rate constants presented in Legrand et al. (see Appenix A) the contribution of this reaction to the HO₂ losses is less than 1%.

We add the note about this reaction in Section 3.2:

“The contribution of the reaction of HO₂(H₂O) with NO₂ (Li et al., 2014) to the RO₂ losses is estimated with the rate constants given in Legrand et al. (this issue) to be less than 1%.”

Pg 15022, ln 25: change to ‘the peak calculated..’

Corrected

Pg 15022, ln 26: How exactly was P(O₃) calculated? Were losses of NO₂ that did not result in ozone production (e.g. OH + NO₂) considered? I think an equation that highlights the reactions considered needs to be added here.

In estimation of the ozone production rate the reactions of NO₂ with OH and RO₂ were neglected compared to the photolysis of NO₂. We add appropriate note:

“As seen in Figure 11, the peak calculated ozone production rate (P(O₃)) is about 0.3 ppbv h⁻¹ during daytime (using the measurements of RO₂ at 3 m, NO at 4 m above the snowpack and assuming P(O₃) equal to NO₂ production rate in the reaction of RO₂ with NO).”

Pg 15022, ln 28: It would be more appropriate to compare the P(O₃) calculated with the diurnal O₃ profile observed during the OPALE campaign rather than an earlier campaign at Dome C.

We agree but the detailed analysis of the ozone data gained at Dome C during the 2011-2012 campaign will be presented in a future publication (in preparation). At present, we can only compare our estimate with the value derived from ozone data gained over the 2007/08 summer season.

Table 2: Add a line after the reactions listed to separate from the summed radical budget section.

Corrected

Table 2: How was the 1σ uncertainty estimated?

Corrected

“1σ uncertainty estimated with accounting for measurement uncertainties”

Figure 7: A comment about the negative intercepts in the correlations between observations and model with zero HONO is needed.

The negative intercepts is the result of the difference in the modelled and measured diurnal profiles mentioned in Section 3.4 This difference is better visible on the Figure 8 from the diurnal profiles of the M/O ratios.

Dr. Berresheim

Overall, this is a very well written paper with strong enhancements in current knowledge of HO_x/RO_x/NO_x chemistry in the Antarctic troposphere and thorough comparison with previous measurements by Mauldin et al. at South Pole. The crucial role of HONO (and/or HNO₄) has been nicely confirmed.

1. However, the only major issue I have is that the paper is strongly entangled with the paper by Legrand et al. (same issue), and that in my opinion it needs to be somewhat "disentangled" to stand on its own. This pertains specifically to sub-chapter 3.5, the "Comparison with 1D model" where it seems that the reader has to flip-flop back and forth between the results of both papers in order to follow (and accept) the arguments made in the discussion (e.g., how much interference exactly due to HNO₄? how crucial is snow temperature? how justified is a turbulent vs. molecular flux approach for HONO vertical transport?). I suggest that every time Kukui et al. use a result from Legrand et al. it should be clearly outlined, and indeed explained, in their discussion even if it means repeating some of the context from Legrand et al.

The following additions were made to provide more information from the work of Legrand et al.

Section 2.2 Last paragraph:

“Legrand et al. (this issue) report tests done both in the field and in the lab that tend to suggest an overestimation of HONO measurements in the range of 10 to 20 pptv due to the presence of HO₂NO₂ in the range of 50-100 pptv in the cold atmosphere at Dome C.. This range of HO₂NO₂ mixing ratios is in agreement with the median [HO₂NO₂] of 80 pptv estimated from RO₂ and NO₂ levels measured at Dome C (see Section 3.2). Also, as discussed by Legrand et al. (this issue), similar levels of HO₂NO₂ were previously observed in Antarctica.”

Section 3.5 1st paragraph:

“According to Legrand et al. (this issue), about 100 pptv of HO₂NO₂ may result in interference equivalent to about 15 pptv of HONO.”

Section 3.5 Paragraph 2.

“Under temperature conditions encountered at Dome C the HONO/NO_x ratio ranged from 0.57 during the day (at -25°C) and 0.3 at night (at -35°C).”

2. Also, I cannot quite agree with the authors' conclusions (p. 15022, line 20) that the model agrees well with night time HONO, as this seems to be not convincing for the 0-6 hours period shown in Fig. 10

The above statement about night-time agreement concerned the HONO estimated from the OH+RO₂ budget (the blue line) which is in good agreement with 1D calculations for the period 0-6 hours. Note, however, that in accordance with suggestions of Referee 3 the Figure 10 has been modified to include the HONO derived from the budget with PSS estimated NO₂. The corresponding comments in the text have been modified as well.

3. A third point I would like to make is the authors' brief acknowledgement of the recent work by Li et al. (2014) of which they mention only in passing that the HO₂(H₂O) reaction with NO₂ would be negligible at Antarctic temperatures as a source for HONO. However, they neglect to recognize that this reaction could - on the other hand - be a significant sink for HO₂, and this should have implications for the model results.

Estimated with the rate constants presented in Legrand et al. (see Appendix A) the contribution of this reaction to the HO₂ losses is less than 1%.

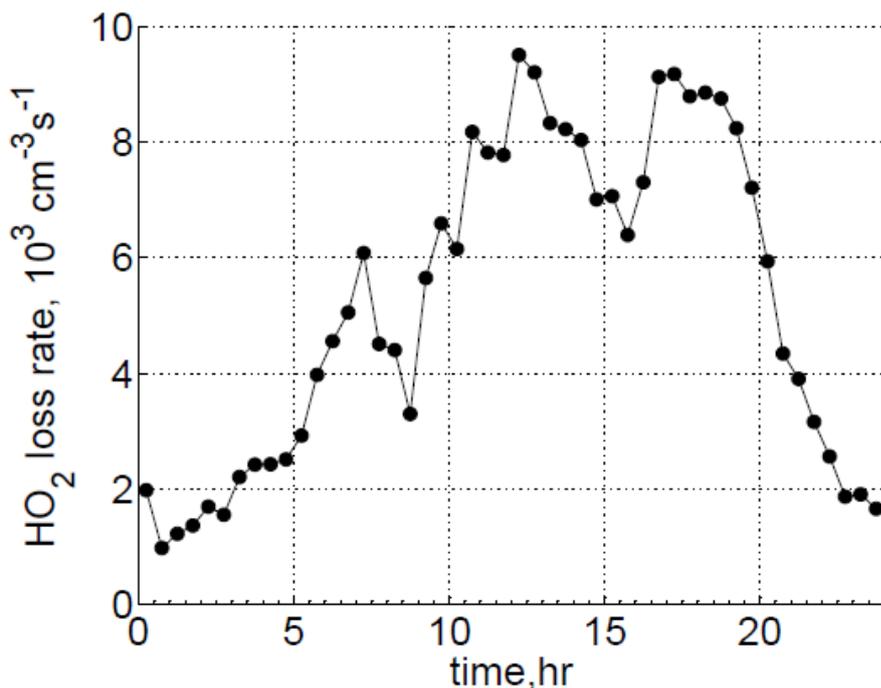


Figure B. Loss of HO₂ via HO₂(H₂O)+NO₂. (compare with Figure 5 of the manuscript)

We add the note about this reaction in Section 3.2:

“The contribution of the reaction of HO₂(H₂O) with NO₂ (Li et al., 2014) to the RO₂ losses is estimated with the rate constants given in Legrand et al. (this issue) to be less than 1%.”

Other than these three comments I have only some minor comments:

pp. 15007-8: With respect to instrument calibration please discuss

a) whether low ambient H₂O concentrations presented any difficulties, and

At temperatures and humidity encountered at Dome C the time of conversion of SO₃ to H₂SO₄ in the conversion reactor was typically 0.5 ms during the day and at maximum of 1ms at night. This time is short enough compared with 4 ms and 20 ms used for two different OH measurements modes. According to the reactor model, the sensitivity varied less than 5% for the range of temperatures and humidity at Dome C. Absence of any artefact related to low H₂O was also confirmed by an absence of variability of the ratio of the OH signals for the two different OH measurement modes.

b) Where and how far away the NO, NO₂ exhaust flows were disposed and whether there could have been interferences with the measurements (also with respect to CO, CH₄ for RO₂) ?

The following text has been added at the end of Section 2.1:

“To avoid possible contamination of ambient air by the SO₂, NO and NO₂ reactants added to the chemical conversion reactor, a trap was set up at the pumps exhaust by using two 100 L cylinders containing zeolites. The cylinders were refilled several times during measurements. Flexible exhaust tube of 30 m length was always placed downwind from the container. When the exhaust tube was intentionally placed upwind and close to the radicals sampling point no effect on radical measurements was detected. Also, no influence of the exhaust on the measurements of NO_x and HONO could be noticed.”

p. 15011, line 13: Where and how have these been "estimated in section 2.2" ? Please explain.

Corrected section 2.2 => Section 3.2

Anonymous Referee #3:

First-time measurements of OH and RO₂ concentrations were reported for Dome C. High radical concentration measurements provide important evidence for a near surface oxidation layer over Antarctica due to snow emissions. Concurrent measurements of O₃, HONO, NO, NO₂, HCHO and H₂O₂ are used in a photochemical steady state model to analyze the radical budgets and the sources and sinks of OH and RO₂. While the reported observations are valuable and the procedures of measurements and modeling results are described and analyzed well, I cannot say that I agree with the way that the conclusions of this paper is presented. The caveats for the conclusions are not apparent enough that a casual reader can easily misunderstand the results.

It will be better to clearly acknowledge the three problems in the analysis: (1) HONO measurement appears to have a (large) high bias; (2) NO₂ measurement appears to have a (large) high bias; (3) The radical budget balance is off by a factor of 3 (comparing daily median net sources of OH and RO₂ with net radical losses in Table 2).

Although the biases of (1) and (2) are described, most of the modeling analyses assume either (1) or (2) but not both. There lies a problem that needs to be rectified before publication.

The conclusions on how much HONO is “real” (line 16-22, P. 15001) is based on modelling analysis using observed NO₂. If the high bias of NO₂ measurement is removed, as in one model simulation (x0, NO₂pss), the amount of HONO needed to explain observed OH and RO₂ will be (much?) smaller than 25%. In fact, looking the modelling results, this is the simulation that has OH and RO₂ concentrations in better agreement with the observations than most of the other simulations.

The discussion on the high bias of observed NO₂ in P. 15017 and 15018 is very informative. I think the discussion should be presented before the discussion of radical budget analysis since NO and NO₂ govern the cycling and loss of radicals. Either there is a large measurement bias or there is some significant unknown chemistry. If there is significant chemistry missing that can alter NO₂/NO ratio by a factor of 3, the model simulated radical budgets without this chemistry cannot be correct. One has to question any conclusions drawn by the analysis. On the other hand, if it is measurement bias, which the authors seem to suggest being the case and I agree, the radical budget analysis should focus on the case not using the biased NO₂ measurements (use NO₂pss instead).

For problem (3), I think it's important to show if the radical budget is balanced for the x0-NO₂pss case. Reducing HONO to 25% works (for the budget balance) because the radical loss estimate using observed NO is high. I think that a self-consistent balanced radical budget is a prerequisite before the modeling results can be used to draw useful conclusions.

The analysis based on Figure 10 is an excellent idea. However, the observed NO₂ was used in the analysis. It would be more instructive to see how much HONO is required when calculated NO₂pss is used.

While I support the publication of this paper, I think that the caveats for the analysis results need be clearly stated. I further suggest that substantial changes be made to restructure the paper to put more emphasis on the model results that account for both biases of (1) and (2).

We agree that possibility of NO₂ overestimation should be better introduced and emphasised in the manuscript. It may be noted, however, that in contrast to HONO measurements for which the interference from HNO₄ was identified, the reason for the observed large NO₂/NO ratio is not yet clear.

We have introduced the following changes:

- *The budget calculations with PSS NO₂ are presented in Table 2;*
- *The Figure 10 shows two HONO profiles calculated with the measured and the PSS estimated NO₂;*
- ***The changes in the article text were made:***

1) Abstract:

“To explain the observations of radicals in this case an additional source of OH equivalent to about (25-35)% of measured photolysis of HONO is required. Even with a factor of 5 reduction in the concentrations of HONO, the photolysis of HONO represents the major primary radical source at Dome C. To account for a possibility of an overestimation of NO₂ observed at Dome C the calculations were also performed with NO₂ concentrations estimated by assuming steady state NO₂/NO ratios. In this case the net radical production from the photolysis of HONO should be reduced by a factor of 5 or completely neglected based on the photochemical budget of OH or OD modelling, respectively.”

2) Section 2.2. In the description of the NO_x measurements we emphasize the uncertainty related to the too large NO₂/NO ratio.

“The NO₂ and NO measurements are discussed in Frey et al. (this issue). The ratios of NO₂ to NO observed at Dome C during 2011-2012 campaign were up to 3 times larger than in 2009-2010 (Frey et al., 2013) and significantly larger, up to 7 times, than the ratios estimated assuming photochemical steady-state (PSS) conditions for NO_x. It is suggested that some part of this inconsistency may be due to unknown interference leading to an overestimation of the NO₂ concentrations (Frey et al., this issue).”

3) Section 3.2. In the discussion of the radical losses we discuss the results with PSS NO₂ presented in Table 2.

“As the losses of OH and RO₂ via the reactions with NO₂ may be overestimated due to unknown interference in the NO₂ measurements (Frey et al., this issue) we also present in the Table 2 (values in parenthesis) the radical losses for [NO₂] estimated assuming PSS conditions for NO_x. In this case the net daytime radical losses are dominated by the radical cross reactions RO₂+RO₂ (25) and OH+RO₂ (20).”

4) Section 3.3. In the discussion of the radical budget:

“As shown in Table 2 the assumption of steady-state NO₂ concentrations lead to a significant overestimation of the net radical production for RO₂ and (RO₂+OH) even when neglecting net OH production by the photolysis of HONO. For OH budget, neglecting the net OH production by the HONO photolysis would lead to an underestimation of the OH production.”

5) Section 3.4 At the end of the section in the description of the OD modelling:

“By assuming the PSS derived NO₂ concentrations, the balance for the OH radical budget is achieved by reducing PHONO by a factor of 5, while for RO₂ and the sum of RO₂ and OH the radical production is overestimated even if the net source from HONO photolysis is neglected.”

6) Section 3.5 1D modelling. We have added to the Figure 10 the HONO derived from the OH budget with PSS NO₂ and added the following comment to the text:

“The HONO mixing ratio-time profiles calculated with the 1D model are compared in Figure 10 with the HONO profiles resulting from analysis of the radical budgets. The levels of HONO derived from the OH budget with measured NO₂ are about 10 pptv higher than the

HONO values obtained using PSS NO₂ concentrations. In both cases the HONO mixing ratios derived from the OH budget are in reasonable agreement with [HONO] predicted by the 1D model (within 5 pptv)."

7) Section 4. We have modified the conclusion to better emphasise the possibility of NO₂ bias. We have added the following text:

"The conclusions based on the radical budget analysis and 0D modelling using the measured concentrations of NO and NO₂ may be significantly biased because the chemical mechanism derived from the available field observations at Dome C is inconsistent with observed large ratios of [NO₂] to [NO]. Assuming that measured NO₂ mixing ratios were overestimated due to unknown interference and using instead [NO₂] estimated assuming steady-state results in lower radical losses and, hence, stronger overestimation of the radical production. In this case, based on the analysis of the radical budgets the observed concentrations of OH radicals are consistent with the levels of HONO corresponding to about (15-20)% of the measured values, while for the sum of the radicals the radical production is overestimated even neglecting the net OH source from the photolysis of HONO. Based on 0D modelling steady-state derived NO₂, the measured OH concentrations are in agreement with steady-state HONO mixing ratios (about 1-2 pptv), while the concentrations of RO₂ radicals are overestimated by about 50% even neglecting the net radical production by the photolysis of HONO.

Hence, in both cases corresponding to the measured or the PSS derived concentrations of NO₂ the calculations, 0D modelling or budget analysis, overestimate the OH and RO₂ concentrations. If this inconsistency is due to an overestimation of the concentrations of HONO, the degree of the overestimation depends on the concentrations of NO₂ used in the calculations. Using the measured NO₂ results in an overestimation of HONO by a factor of 3-4. If the concentrations of NO₂ are estimated assuming steady-state conditions the net radical production from the HONO photolysis should be reduced by a factor of 5 or completely neglected based on the budget of OH or 0D modelling, respectively. "

Other comments:

(1) Line 22-23, P15002, photolysis of H₂O₂ from snow emissions is a primary source. Photolysis of H₂O₂ formed from two HO₂ is not a primary radical source.

"primary" is replaced by "net"

(2) Line 1-9, P15009, some estimates on how the measurement uncertainties affect the budget uncertainties would be useful.

The uncertainties presented in Table 2 or in Figure 9 were estimated from the measurement uncertainties (including OH and RO₂). The comment b) to the Table 2 was modified.

"b) 1σ uncertainty estimated with accounting for measurement uncertainties"

(3) Figure 1, please show the time series of NO, NO₂, NO₂pss. These species are critical for the discussion of radical budgets.

The NO_x profiles are presented in the accompanying article of Frey et al. (this issue). Here we present only median values collected in Table 1. We added to the Table 1 the PSS estimated NO₂ median and range mixing ratios.

(4) Line 5-7, P15013, please show a figure of OH dependence on J(O₁D). If J(O₁D)+H₂O is not a significant primary radical source (Table 2), why is there such a square-root dependence?

In this paper we present the dependence of OH on $J(\text{NO}_2)$ which is found to be quasi-linear, but the correlation is weak. We show then that the correlation of OH is linear and strong with the production rates of OH from the photolysis of HONO and from the reaction HO_2+NO . These dependences are in agreement with our conclusion about significance of these two sources at Dome C.

Concerning the dependence on $J(\text{O}^1\text{D})$ we state in Section 3.1: “The relationship of [OH] with $J(\text{O}^1\text{D})$ was close to a power-law dependence with an exponent of ~ 0.5 in agreement with a typical close to quadratic dependence of $J(\text{O}^1\text{D})$ on $J(\text{NO}_2)$ observed at Dome C.”

The dependence of OH on $J(\text{O}^1\text{D})$ is shown below. As we do not see how this dependence with very large scattering can be helpful for the discussion of sources and sinks of the radicals we suggest that it should not be included in the article.

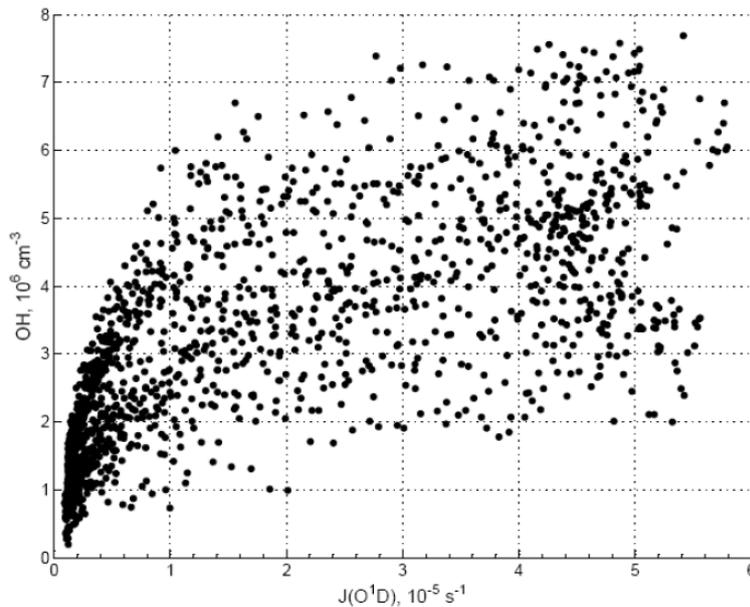


Figure C. Dependence of OH on $J(\text{O}^1\text{D})$.

(5) Figure 4, HONO produced from $\text{OH}+\text{NO}+\text{M} \rightarrow \text{HONO} + \text{M}$ should be kept in the simulations of all x0 cases. Removing gas-phase produced HONO introduces an artificial radical sink, which can be large for high NO conditions and may explain some of the model bias, which seems to worsen with increasing NO.

In fact, the case x0 corresponds to the model with calculated PSS HONO not constrained with measured HONO. The corresponding correction is made in the figure captures for Figures 4,7,8 and in the text.

(6) Line 28, P. 15015, an average of 80 pptv HNO_4 is much higher than 40-60 pptv observed at South Pole. It would be easy to see if HNO_4 and the corresponding NO observations are compared between Dome C and South Pole.

“similar” is replaced by “somewhat higher than”.

Note also that up to 150 pptv of HO_2NO_2 was observed between the ground and 50 m elevation over the Antarctic plateau (Slusher et al., J. Geophys. Res. Atmos., 115, D07304, doi:10.1029/2009JD012605, 2010)

(7) Line 16-17, P. 15016, the sum of two variabilities is $> 100\%$. Please comment on what it means.

Corrected (please, see also the comment to this by referee 1):

“The variability of P_{HO_2+NO} or P_{HONO} then explain $\sim 80\%$ of the variability of OH.”

Appendix A:

About the reaction HO₂(H₂O)+NO₂ from Legrand et al. (this issue)

Another gas-phase source of HONO was recently proposed by Li et al. (2014) via reaction of HO₂(H₂O) complex with NO₂:



Reaction of HO₂(H₂O) complex with NO₂ was first suggested by Sander and Peterson (1984) to explain the observation of a linear dependence of the effective rate constant of the reaction of HO₂ with NO₂ on the concentration of water vapour in the temperature range 275-298 K. Assuming reaction mechanism (2-4) Sander and Peterson (1984) derived temperature dependence for the effective third-order rate constant of the reaction HO₂+NO₂+H₂O, $k_{4}^{\text{III}}(\text{T})$, with $k_{4}^{\text{III}}(\text{T})$ representing the product $k_4 \times K_3$, where k_4 is the bimolecular rate constant for reaction HO₂(H₂O) with NO₂ and K_3 is equilibrium constant for reaction (3). The possible contribution of reaction (4) to form HONO at Concordia was evaluated by assuming a unity yield of HONO for the reaction (4). The rate constant $k_4(\text{T})$ in the temperature range 275-298 K was estimated from the $k_{4}^{\text{III}}(\text{T})$ data of Sander and Peterson (1984) using recent recommendations for $K_3(\text{T})$ and $k_2(\text{T})$ from Sander et al. (2011): $k_4(\text{T}) = k_{4}^{\text{III}}(\text{T}) / K_3(\text{T}) \times k_2(\text{T}) / k_2(\text{T})^{\text{Sander}}$, where $k_2(\text{T})^{\text{Sander}}$ are data from Sander and Peterson (1984). The values of $k_4(\text{T})$ at low temperatures encountered at Concordia were obtained by extrapolating the $k_4(\text{T})/k_2(\text{T})$ data from Sander and Peterson (1984) and assuming a logarithmic dependence of $k_4(\text{T})/k_2(\text{T})$ on $1/\text{T}$, similar to reaction of HO₂(H₂O) with HO₂ (Sander et al., 2011). The resulting dependence ($k_4(\text{T})/k_2(\text{T}) = 10^{-1505.3/\text{T}(\text{K})+5.4}$) predicts significantly lower water enhancement effect at low temperature ($k_4/k_2=0.12$ at 240K compared to 2.2 at 298K). Using these k_4 values and observations of OH, NO, HO₂, NO₂ and H₂O, the low temperatures encountered at Concordia make negligible the formation of HONO from the reaction (4). This hypothetical HONO source would contribute for 10-20% of the HONO production from the reaction OH+NO and would result in less than 1% of the measured HONO.

Li, X., Rohrer, F., Hofzumahaus, A., Brauers, T., Häseler, R., Bohn, B., Broch, S., Fuchs, H., Gomm, S., Holland, F., Jäger, J., Kaiser, J., Keutsch, F. N., Lohse, I., Lu, K., Tillmann, R., Wegener, R., Wolfe, G. M., Mentel, T. F., Kiendler-Scharr, A., and Wahner, A.: Missing Gas-

Phase Source of HONO Inferred from Zeppelin Measurements in the Troposphere, *Science*, 344, 292-296, 2014.

Sander, S.P., and Peterson, A.E.: Kinetics of the reaction $\text{HO}_2 + \text{NO}_2 + \text{M} \rightarrow \text{HO}_2\text{NO}_2 + \text{M}$, *J. Phys. Chem.*, 88, 166-1571, 1984.

Sander, S.P., J. Abbatt, J.R. Barker, J.B. Burkholder, R.R. Friedl, D.M. Golden, R.E. Huie, C. E. Kolb, M.J. Kurylo, G.K. Moortgat, V.L. Orkin and P.H. Wine : "Chemical Kinetics and Photochemical Data for Use in Atmospheric Studies, Evaluation No. 17," JPL Publication 10 June 2011, Jet Propulsion Laboratory, Pasadena, <http://jpldataeval.jpl.nasa.gov>, 2011.