Comparison of IASI water vapor retrieval with H_2O-Raman lidar in the framework of the Mediterranean HyMeX and ChArMEx programs

Patrick Chazette, Fabien Marnas, Julien Totems and Xiaoxia Shang

Laboratoire des Sciences du Climat et de l'Environnement (LSCE), UMR8212, Laboratoire mixte CEA-CNRS-UVSQ, CEA Saclay, 91191 Gif-sur-Yvette, France.

Abstract.

The Infrared Atmospheric Sounding Interferometer (IASI) is a spaceborne passive sensor of new generation mainly dedicated to meteorological applications. Operational Level-2 products are available via the European Organisation for the Exploitation of Meteorological Satellites (EUMETSAT) since several years. In particular, vertical profiles of water vapor measurements are retrieved from infrared radiances at the global scale. Nevertheless, the robustness of such products has to be checked because only few validations have been reported. For this purpose, the field experiments that were held during the HyMeX and ChArMEx international programs are a very good opportunity. A H_2O-Raman lidar was deployed on the Balearic Island of Menorca and operated continuously during ~6 and ~3 weeks during fall 2012 (Hydrological cycle in the Mediterranean eXperiment -HyMeX-) and summer 2013 (Chemistry-Aerosol Mediterranean Experiment -ChArMEx-), respectively. It measured simultaneously the water vapor mixing ratio and aerosol optical properties. This article does not aim to describe the IASI operational H_2O inversion algorithm, but to compare the vertical profiles derived from IASI onboard MetOp-A and the ground-based lidar measurements to assess the reliability of the IASI operational product for the water vapor retrieval in both the lower and middle troposphere. The links between water vapor contents and both the aerosol vertical profiles and the air mass origins are also studied. About 30 simultaneous observations, performed during nighttime in cloud free conditions, have been considered. For altitudes ranging from 2 to 7 km, root mean square errors (correlation) of ~0.5 g/kg (~0.77) and ~1.1 g/kg (~0.72) are derived between the operational IASI product and the available lidar profiles during HyMeX and ChArMEx, respectively. The values of both root mean square error and correlation are meaningful and show that the operational Level-2 product of the IASI-derived vertical water vapor mixing ratio can be considered for meteorological and climatic applications, at least in the framework of field campaigns.

1 Introduction

Satellite observations are powerful tools for meteorological forecasts. Their assimilation in models lead to an improvement on weather forecasts (e.g. *Collard and McNally*, 2009; *Bormann et al*., 2010). Among the main components of the atmospheric state, water vapor is an essential element, which plays a key role in frontogenesis, convection (e.g. *Held and Soden*, 2000), cloud formation and aerosol hydration (e.g. *Larson and Taylor*, 1983; *Rood et al*., 1987; *Randriamiarisoa et al*., 2006). In this way, it influences significantly the Earth climate and the atmospheric chemistry (e.g. IPCC, 2014). It is also an energy reservoir that exchanges with both the atmosphere and the surface through condensation and evaporation processes via the latent heat flux. Hence, for reliable weather forecasts, the vertical profile of the water vapor has to be precisely assessed.

During several decades, passive radiometers, such as those implemented onboard of the Televison InfraRed Operational Satellite (TIROS) from the National Oceanographic and Atmospheric Administration (NOAA), have allowed to retrieve temperature and moisture profiles with a vertical resolution of about 3 to 5 km in the troposphere, as defined by the instrumental weighting functions (e.g. *Susskind et al*., 1984; *Chedin et al*., 1985). A new generation of instruments has been launched on polar platforms satellites, such as

Interferometric Monitor for Greenhouse gases (IMG, e.g. Ogawa et al., 1994; Clerbaux et al., 1998), Tropospheric Emission Spectrometer (TES, e.g. *Shephard et al*., 2008; *Worden et al*., 2012), the Advanced Infrared Sounder (AIRS, *Chahine et al*., 1990; *Aumann and Miller,* 1995), and the Infrared Atmospheric Sounding Interferometer (IASI, e.g. *Clerbaux et al*., 2009; *Hilton et al*., 2012). Thanks to a larger number of spectral channels and an enhanced spectral resolution, these instruments lead to improved vertical resolutions down to 1 km and higher precision of both the atmospheric temperature and water vapor content retrieval.

We will focus our study on the reliability of the water vapor mixing ratio (WVMR) vertical profiles retrieved from the IASI spectrometer, which has been launched onboard the polar orbiting meteorological satellites MetOp (Meteorological Operational), which forms the space segment of the overall EUMETSAT Polar System (EPS).

Main mission of IASI is the operational meteorology (e.g. *Zhou et al*., 2009), although air-composition and climate applications are also well covered by the instrument as also discussed before launch (e.g. *Chazette et al*., 1998; *Clerbaux et al*., 1998) and now demonstrated (e.g. *Crevoisier et al*., 2013a; *Griffin et al*., 2013; *Grieco et al*., 2013). Hereafter we will only discuss the comparison between IASI-derived WVMR and the simultaneous measurements performed by a H_2O-Raman lidar deployed on the Balearic Island of Menorca in the framework of the Hydrological cycle in the Mediterranean eXperiment (HyMeX, http://www.hymex.org/, *Chazette et al*., 2013) and Chemistry-Aerosol Mediterranean Experiment (ChArMEx, http://www.mistrals-home.org).

For our concern, the IASI-derived WVMR operational Level-2 products have been available via the European Organisation for the Exploitation of Meteorological Satellites (EUMETSAT) for several years. In particular, vertical profiles of WVMR measurements are retrieved from infrared radiances at the global scale (e.g. *Carissimo et al*., 2005; *Schlüssel et* *al*., 2005; *Schneider and Hase*, 2011). The robustness of such products has to be checked, and the field experiments that were held during the HyMeX and ChArMEx international programs are a very good opportunity for that purpose. Few validation exercises have been conducted on the WVMR operational product. The main reason seems that for meteorological forecasting, the radiances are directly assimilated in the models (e.g. *Hilton et al*., 2009; *Hilton et al*., 2012; *Heilliette*, 2013; *Matricardi and McNally*, 2013; *Xu et al*., 2013). Nevertheless, the WVMR Level-2 product could have a great interest in order to help field experiment analyses. Moreover, few validations are available in the scientific literature. *Pougatchev et al*. (2009) used rawindsounding measurements to assess the error covariance matrix needed for the inversion algorithm. *Masiello et al*., (2013) argue that lidar measurements are excellent candidates for the validation of spaceborne sensors. They used different measurement techniques during the Convective and Orographically-induced Precipitation Study (COPS) campaign, and the comparisons were performed with a limited number of lidar profiles (6) during the same season. Such validations are very delicate because performing atmospheric measurements of WVMR with the required accuracy for satellite retrieval validation is a challenging issue due to the high spatio-temporal variability of atmospheric water vapor. The spatiotemporal coincidence between the ground-based and the spaceborne measurements has to be guaranteed to avoid important sampling errors.

In the following section, the Raman lidar system used for IASI WVMR comparisons and its technical specifications will be presented as well as the experimental sites used to conduct the validation during the Mediterranean project. The IASI derived WVMR product specifications will also be introduced. The third section will present the experimental comparisons. The statistical tools used to evaluate the WVMR products will be introduced and the experimental results obtained will be presented. Then, the influence of both the air mass origin and their aerosol content in the results will be discussed. Finally, the main results will be summarized in the conclusion.

2 Observations

The comparison between the WVMR ground-based lidar measurements and the IASI operational products took place in the framework of both HyMeX and ChArMEx Special Observation Periods during September-October 2012 and June-July 2013, respectively, on the Balearic island of Menorca. During HyMeX/IODA-MED (Innovative Observing and Data Assimilation systems for the MEDiterranean Weather), the Water vapor and Aerosol lidar (WALI) was located close to La Ciutadella (Western part of the island, 39°60'00" N and 3°50'20"E), while during ChArMEx it was deployed close to Mahon (Eastern part of the island, 39°53’12” N and 4°15’31” E). Hence, the WVMR vertical profiles derived from the IASI spaceborne spectrometer (Ether CNES/CNRS-INSU Ether web site http://www.pole-ether.fr) have been compared to the ones measured by WALI during nighttime for field experiment durations of 6 and 3 weeks for HyMeX and ChArMEx, respectively. The use of the Raman technique limits the range of daytime measurements (< 1 km), which are consequently not relevant for a validation purpose in the lower and middle troposphere.

2.1 The WALI Raman lidar

The WALI instrument uses an emitting wavelength of 354.7 nm and is designed to fulfill eye-safe conditions (Table 1). The instrument, its calibration and the associated errors are documented in *Chazette et al*. (2013) and will not be detailed here. During all the experiment the acquisition was performed for mean profiles of 1000 laser shots leading to a temporal sampling close to 1 minute. The UV pulse energy is ~60 mJ and the pulse repetition frequency is 20 Hz. It is equipped with four detection channels: an aerosol board including co-polarized and cross polarized channels with respect to the laser emission, a channel dedicated to the detection of the water vapor Raman signal at 407.5 nm and a fourth channel dedicated to the recording of the atmospheric nitrogen Raman signal at 386.6 nm.

The design of the WALI system leads to very good capabilities in terms of low altitude overlap and WVMR retrieval during nighttime. The absolute deviation from rawindsoundings is less than 0.5 g/kg (*Chazette et al*., 2013). The error on the WVMR reaches 11% in the marine boundary layer and decreases to 7% below 5 km range for a temporal averaging of 20 minutes and a vertical resolution of 15 m. Precision can deteriorate very quickly thereafter due to the decreasing Signal to Noise Ratio (SNR) with altitude. It is also worse during daytime, but measurements can be performed with the same uncertainty for altitude ranges below 1 km using a temporal averaging over ~1 hour. For the inter-comparisons presented in this paper, the chosen averaging time is 30 minutes, centered on the time value of the IASI profile to be compared, and the altitude range is from 0.3 to 7 km above the mean sea level (amsl). The original vertical and temporal resolutions are 15 m and 1 minute, respectively. The lidar profiles were smoothed for the comparison so that the vertical resolution used for this study is ~41 m.

2.2 The MetOp /IASI satellite data

MetOp (Meteorological Operational) consists of a series of three polar heliosynchronous orbiting satellites, to be flown successively for more than 14 years, from 2006. This series forms the space segment of the overall EUMETSAT Polar System (EPS). EPS is the European contribution to the Initial Joint Polar System agreement (IJPS), an agreement between EUMETSAT and NOAA. MetOp flies in a Low Earth orbit at an altitude of 817 km corresponding to local 'morning', while the US is responsible for 'afternoon' coverage (*Klaes et al*., 2007). MetOp-A (launched on 19 October 2006) and MetOp-B (launched on 17 September 2012) provide detailed observations of the global atmosphere, oceans and continents. MetOp-C is due to be launched in 2017. The series provides data for both operational meteorology and climate studies. A combination of passive remote sensing instruments offers the capability to observe the Earth by day and night, as well as under cloudy conditions. The most innovative and one of the key instruments on MetOp is the Michelson interferometer IASI. Three IASI instruments were developed for MetOp by CNES (Centre National d'Etudes Spatiales) in cooperation with EUMETSAT. They are built to provide temperature and moisture measurements with unprecedented accuracy and resolution, and additionally to provide information for the monitoring of atmospheric trace gases.

The bandwidth of IASI is divided into 8461 spectral channels between 645 and 2760 cm^{-1} with a mean spectral resolution of 0.5 cm^{-1} after apodization. IASI scans across-track in 30 successive elementary fields of view (EFOV), each composed of 4 instantaneous fields of view (IFOV) of 0.8225° leading to a footprint of 12 km diameter at sub-satellite point. The footprint dimension increases from 20 to 39 km along-track directions to the swath edge, respectively (*Cayla*, 1993). The swath width on the ground is approximately 2200 km, which provides global Earth coverage twice per day.

Operational products from EPS/MetOp are generated in the EPS Core Ground Segment. The IASI Level-2 processing development targeted the generation of temperature and humidity profile information, the associated surface information and the retrieval of some trace gas species: CO, O_3, CH_4, N_2O and CO_2. The vertical temperature and water-vapor profiles are currently distributed on a 90-level grid extending between 0.005 and 1050 hPa (*August et al*., 2012). Note that the operational product uses a statistical approach to retrieve the geophysical parameters. Other approaches use a physical scheme and give access to a better vertical resolution (e.g. *Amato et al*., 2009; *Masiello et al*., 2013). Nevertheless, the goal of this paper is to provide quantitative elements of validation for the operational product using the statistical approach.

Both the temperature and moisture of the troposphere and lower stratosphere are derived under cloud-free conditions with a vertical resolution of 1-2 km in the lower troposphere; a horizontal resolution of 25 km, and an accuracy of 1 K and 10%, respectively. The number of independent pieces of information which are determined in the moisture profiles is in the order of 10. The sensitivity to the lower troposphere is lower and leads to larger error beneath 3 km, although ~80% of moisture is contained in this layer. For the WVMR retrieval, the IASI weighting functions are generally maximum above 700 hPa.

For the comparison presented hereafter, we considered the 12 closest IASI pixels from the lidar ground-based station. The mean values and the associated standard deviations are then calculated if the number of relevant IASI-derived WVMR profiles are at least equal to 6.

3 Comparison between the IASI and WALI water vapor products

Here we assess the representativeness of IASI in terms of atmospheric moisture content considering both vertical profiles and integrated values to evaluate the potentiality of these products to be used for meteorological studies purposes. The relevant IASI coincidences are established before a comparison with the Raman lidar WALI separately for the two time periods of field experiments.

3.1 Coincidences

Figure 1 gives the temporal evolution of the WVMR vertical profiles above Menorca during the two time periods on which field experiments were conducted. The water vapor contents are highly variable and highlight contrasted atmospheric situations, which are of interest for comparison to IASI-derived WVMR. On the same figure are given the satellite overpass times for which comparisons are relevant. We have identified 30 coincidences in cloud-free conditions with available IASI profiles, during September-October 2012 and June-July 2013. Note that the presence of high aerosol content is also classified as a cloudy condition. For each time period, the coincidences are identified by their number in a chronological way hereafter called coincidence number (CN). All the coincidences are reported in Table 2 and Table 3 for the time periods of HyMeX and ChArMEx, respectively (15 CN each). The distance between the central pixel of IASI and the lidar ground-based station (*D*), and the number of relevant IASI pixels (*N*) are also indicated.

3.2 WVMR vertical profiles

The WVMR vertical profiles for the whole retained atmospheric situations are shown in Figure 2 and Figure 3 for HyMeX and ChArMEx, respectively. The coincident WVMR simulated from the European Centre for Medium-Range Weather Forecasts (ECMWF) operational analysis are also plotted on the figures. The 9 closest model grids from the ground-based lidar station are considered to compute both the mean and the standard deviation vertical profiles. The meteorological fields have been provided by ECMWF and have been obtained from the ESPRI/IPSL data server for a horizontal resolution of 0.5°.

The statistical indicators used to evaluate the relevance of the IASI-derived WVMR (r_{iasi}) with respect to lidar observations (r_{lidar}) are the Root Mean Square Error (RMSE) and the (Pearson) correlation (COR). They are often used to evaluate model performances as in *Boylan and Russell* (2006) and can be written as

$$RMSE = \sqrt{\frac{1}{N_t}\sum_{i=1}^{N_t}(r_{iasi} - r_{lidar})^2} \quad (1)$$

$$COR = \frac{\sum_{i=1}^{N_t}(r_{lidar}-\overline{r_{lidar}})(r_{iasi}-\overline{r_{iasi}})}{\sqrt{\sum_{i=1}^{N_t}(r_{lidar}-\overline{r_{lidar}})^2 \sum_{i=1}^{N_t}(r_{iasi}-\overline{r_{iasi}})^2}} \quad (2)$$

where N_t is the total number of coincidences and the overbar terms are averages. The vertical values for both r_{IASI} and r_{lidar} are used at the IASI-L2 pressure level gird.

Figure 4 gives the vertical profiles of both RMSE and COR for the two time periods. The statistical indicators have been computed between WALI and IASI, and WALI and ECMWF data.

During the first time period (fall time), the lidar and modelled profiles are in better agreement with a mean RMSE and COR of 0.42 g/kg and 77% (between 0.5 and 7 km), respectively, whereas between the lidar and IASI vertical profiles these values reach ~0.6 g/kg and 70%, respectively. The shapes of RMSE and COR against altitude are however very similar. It is not surprising because ECMWF analyses are made by assimilating the IASI radiances (e.g. *Hilton et al*., 2012) in addition to the rawinsounding performed in Palma de Mallorca (100 km Southwest of Menorca). In the planetary boundary layer (PBL) more discrepancy could be encountered due to local effects.

An opposite behavior happens in terms of RMSE for the second time period (summer time) where the IASI-derived WVMR (*RMSE* = 1.64 g/kg) is better than that of the model (*RMSE* = 2.04 g/kg) when compared to the Raman lidar. It is mainly true below 2 km. Nevertheless, the correlation is better between the lidar and the ECMWF analyses (0.82) than between the lidar and IASI (0.59).

In the free troposphere, where the IASI weighted functions mostly have their maxima, the agreement is better according to the RMSE and COR indicators. This agreement is higher for the HyMeX time period and might be due to a lesser influence of the aerosol layers. For this period *RMSE* is lower than 0.5 g/kg and *COR* is ~77%, to be compared to ~1.1 g/kg and ~72% during the ChArMEx time periods. Below 2 km, the agreement is degraded as expected: RMSE is between ~2 and 3 g/kg and the COR value tends to 0. Table 4 summarizes the results for different atmospheric layers between 0.5 and 7 km. Such results are consistent with those of *Schneider and Hase* (2011) who used rawinsoundings as validation tools for the IASI WVMR Level-2 operational products. With the exception of the PBL, they found a correlation coefficient of ~0.80.

3.3 Water vapor integrated content

When considering the Water vapor integrated content (WVIC) between 0.5 and 7 km, the agreement between lidar- and IASI-derived moisture is within a standard deviation between 0.18 and 0.25 g/cm^2. Figure 5 illustrates this agreement: the IASI-derived WVIC exhibits a bias lower than 0.15 g/cm^2 compared to the one retrieved from WALI. In fact, the WVIC retrieved from IASI is in the range value (between 0.5 and 2 g/cm^2) for the HyMeX time period (fall 2012), but it is mostly underestimated by ~10% during the ChArMEx time period (summer 2013). The slopes of the regressions are 0.89 and 0.81 for the HyMeX and ChArMEx time periods, respectively. Note that during the HyMeX time period (fall 2012), the agreement between the lidar and IASI profiles is better, even in the general shape.

When compared to ECMWF analyses, standard deviations with respect to WALI are close to 0.17 and 0.45 g/cm^2 for the two previous time periods, respectively. With respect to previous IASI cross-comparisons, results are not degraded during the HyMeX fall period but significantly worse over the ChArMEx summer period where the slope of the linear fit is close to 0.70. Such discrepancy (underestimation) may be due to an incorrect consideration of the instrumental error in the variance/covariance matrix needed for the assimilation process (e.g. *Wang et al*., 2013). The error on the contribution to the IASI radiances may be linked to local heating associated to the aerosol presence not being taken into account in the model, as for all spaceborne infrared sensors (e.g. *Pierangélo et al*., 2004). This point is not within the topic of this paper and has to be further investigated.

4 Influence of the air mass origins - aerosol as air mass tracer

Among all 30 coincidences, the origins of air masses are very different and can be characterized using simultaneously several aerosol optical properties and air mass back trajectories. The Raman lidar WALI offers the capability to retrieve fundamental aerosol optical properties (*Chazette et al.*, 2013): the vertical profiles of the volume depolarization ratio (VDR) to identify the presence of dust-like aerosols, the aerosol extinction coefficient (AE) to locate in altitude the scattering layers, the equivalent backscatter to extinction ratio (BER) which is proportional to the single scattering albedo, and the aerosol optical thickness (AOT) characterizing the aerosol column burden. The inversion process used both the N_2-Raman and elastic channels at 355 nm and is described in various papers as *Royer et al*. (2011) or *Chazette et al*. (2012) where the related uncertainties are assessed. Hence, using the aerosol optical properties described above, coupled with air mass back trajectory analysis, the air masses influencing the IASI-derived WVMR can be identified.

4.1 Aerosol optical properties

As in Figure 1, the dates of the relevant IASI coincidences are highlighted in white dotted lines in Figure 6. This figure represents the temporal evolution of the vertical profile of VDR for HyMeX and ChArMEx time periods, respectively. In general, the relevant coincidences do not occur during the major dust events where the VDR is maximal (in brown on the figure), likely because the dust plume is classified as cloud: it is sufficiently thick to significantly influence the brightness temperature used to invert the IASI infrared spectrum. Nevertheless, other sources of aerosol may affect the IASI measurements. Thereby, the BER is also an important parameter to identify the aerosol types (e.g. *Cattrall et al*., 2005) as it is linked to their chemical composition. It is given in Figure 7 as a column average and presents a strong variability, ranging from ~0.01 sr^{-1} for pollution aerosol (e.g. *Raut and Chazette*, 2009) to ~0.04 sr^{-1} for marine aerosol (e.g. *Flamant et al*., 2000). The intermediate values are for aerosol mixing, dust aerosols (e.g. *Mattis et al*., 2002; *Chazette et al*., 2007) or long-range transport pollution aerosols (e.g. *Chazette et al*., 2012). The AE and VDR vertical profiles are also given Figure 8 and Figure 9 for the coincidences of the two time periods. They often show strong heterogeneities with respect to altitude which are directly related to the vertical profiles of WVMR given Figure 2 and Figure 3, respectively. All the vertical structures encountered have to be investigated to compare the WVMR-derived from IASI and WALI. The aerosol atmospheric content in terms of AOT is also very different from one observation to another because it ranges from 0.04 (very clean air) to ~0.4 (polluted air and/or dust event). Hence, the coincidences are very diverse for an inter-comparison exercise, and allow evaluating the IASI-derived WVMR retrieval for very distinct atmospheric situations and aerosol contents.

4.2 Air mass back trajectories

Air mass back trajectories have been computed to determine the corresponding aerosol transport routes using the NOAA Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT) model (*Draxler and Rolph*, 2003) with 3-hourly archived meteorological data provided from the US National Centers for Environmental Prediction (NCEP) Global Data Assimilation System (GDAS) at the horizontal resolution of 0.5°. The altitudes of the trajectory starting points (1, 2 and 4 km) were selected primarily from the lidar observations of aerosol layer heights highlighted in Figure 8 and Figure 9. The air mass back trajectories are shown Figure 10, Figure 11 and Figure 12 for the 3 starting points and for each time period. One path was drawn on 72 hours for each coincidence between IASI and WALI measurements. The air mass origins are very variable during the time periods for all starting point altitudes. There are two major contributions to the air masses passing over Menorca, the first one from the Sahara and the second one from the Atlantic Ocean.

4.3 Discussion

The summary of our conclusion about the origins of air mass revealed by the shape of the WVMR vertical profile is given in Table 2 and Table 3 for the HyMeX and ChArMEx time periods, respectively. Depolarizing layers (DL) and residual pollution layers (RPL) are specifically identified.

The atmospheric situations observed during the coincidences present significantly high moisture content (WVMR > 5 g/kg in the free troposphere and WVIC close to 2 g/cm^2) for 4 (5) cases during the HyMeX (ChArMEx) time period, which correspond to CN = 1, 2, 6 and 10 (2, 5, 6, 7 and 8). Such situations are generally well represented by the Level-2 product of the IASI operational ground segment, excepted for CN = 10 during the HyMeX time period where the IASI product overestimates the WVMR by ~25%. In this case, the air mass came from Morocco and brought moisture with dust aerosols between ground and ~4 km above the mean sea level (amsl). Nevertheless, it must be noted that Saharan air masses are often associated with higher moisture content and the agreement between IASI- and WALI-derived WVMR is generally better for these air masses because of the smoother transitions in vertical structures due to higher moisture content in these layers.

The major discrepancies are observed for the drier air masses (WVIC less or close to 1 g/cm^2) which present a strong vertical gradient of WVMR, generally between the PBL and the free troposphere. Such a gradient is not reproduced from IASI measurements due to its insufficient vertical resolution. Note that the dry air masses observed during the field campaigns originated from the Atlantic and had a small AOT (< 0.2).

For the other coincidences, the agreement between IASI and WALI is good. The median value of the atmospheric aerosol content (AOT ~ 0.2) is similar during the two time periods and cannot explain the observed differences between them. During the ChArMEx time period several coincidences are associated with very clean air (*AOT* < 0.1) situations. Furthermore, the differences observed in the WALI/ECMWF comparison cannot be explained by the presence of an aerosol layer. These discrepancies seem linked to a seasonal role, which could be due to an incorrect consideration of the sea surface temperature in the model.

The air mass origin plays a major role through the shape of the original vertical structure which can be kept during several days along the transport. As discussed by *Kim et al*. (2004), the larger amount of water vapor in an aerosol layer contributes to a higher radiative heating, increasing the potential temperature and static stability of the layer. This may help to maintain the structure of the layer for a longer period of time. Note that the vertical structures observed during our two field campaigns are not uncommon in the atmosphere (e.g. *Chazette et al*., 2001; *Kim et al*., 2009). All this suggests a need for an increased vertical resolution of infrared spaceborne sounders, and then for the improvement of their spectral resolution (e.g. *Crevoisier et al*., 2013b).

5 Conclusion

Following the international field campaigns HyMeX (Hydrological cycle in the Mediterranean eXperiment) and ChArMEx (Chemistry-Aerosol Mediterranean Experiment) in fall 2012 and summer 2013, respectively, 30 relevant coincidences between the ground-based lidar WALI (Water vapor and Aerosol lidar) and the spaceborne instrument Infrared Atmospheric Sounding Interferometer (IASI) have been selected to conduct a comparison exercise of the water vapor mixing ratio (WVMR) vertical profile retrieval. The general result is in good agreement between the two instruments. Two statistical indicators generally used to evaluate model performances have been considered: the Root Mean Square Error (RMSE) and the (Pearson) correlation (COR). In the middle troposphere (2-7 km amsl) the COR value is ~77 and 72%, and the RMSE is lower than 0.5 and 1.1 g/kg for the fall and summer periods, respectively. Discrepancies are higher in the planetary boundary layer (PBL) because the weighted functions of IASI do not correctly sample this layer close to the ground (*RMSE* ~ 1.6 g/kg and *COR* < 0.4). Considering the water vapor integrated content within the altitude range of 0.3 and 7 km amsl, the standard deviation between IASI and WALI are 0.18 and 0.25 g/cm^2 for the fall and summer periods, respectively. The disagreement is higher during summer time and we may suspect the presence of aerosol layers and/or contrasted vertical atmospheric structures to be responsible for this bias.

During coincidences, we note that the integrated atmospheric aerosol content has been found with aerosol optical thickness between 0.04 and 0.4 associated with various particle types (pollution, marine or dust aerosols), as identified from both the lidar-derived backscatter to extinction ratio and air mass back trajectories. The aerosol optical thickness does not significantly affect the results of the intercomparison. The divergence on the WVMR vertical profile is mainly due to the existence of sharp transitions which mainly occurs between the PBL and the free troposphere. The agreement is generally better for Saharan air masses because of the smoother transitions in vertical structures due to higher moisture content in these layers (~5 g/kg). Our results calls for an improvement of the spectral resolution of the Fourier transform spectrometer IASI. Such consideration is being studied for the next generation, the so-called IASI-NG. Moreover, the synergetic use of microwave measurements is capable of improving the water vapor retrievals, especially in the PBL. An upcoming version (6) of the operational IASI Level-2 processor with synergistic use of the Advanced Microwave Sounding Unit (AMSU) and the Microwave Humidity Sounder (MHS) data is scheduled. It was reported at the last International TOVS Study Conference hold to Jeju Island (South Korea, 26 March - 1 April 2014) to contain substantial improvements of the profiles when compared with the European Centre for Medium-Range Weather Forecasts (ECMWF) analysis, in particular in the lower levels and for the entire water vapor profiles.

The approach presented in this study can be applied to the next generation of IASI operational water vapor products.

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Table 1: Main technical characteristics of the WALI instrument.

Laser	Nd:Yag
Energy	60 mJ at 355 nm
Frequency	20 Hz
Reception channels	Elastic total 354.67 nm Elastic $\perp$ 354.67 nm Raman-N_2 386.63 nm Raman-H_2O 407.5 nm
Reception diameters	15 cm
Field of view	~2.3 mrad
Full overlap	~300 m
Detector	Photomultiplier tubes
Filter bandwidths	0.2 - 0.3 nm
Vertical sampling	0.75 m (analog) 15 m (photon counting)
Vertical resolution used for this study	~ 40 m
Acquisition system	PXI technology at 200 MHz

Table 2: List of the coincidence numbers (CN) and description of the associated aerosol layer synoptic origin during the HyMeX experiment (2012 fall period). *D* represents the distance between the ground-based lidar and the center of the 12 selected IASI pixels. *N* is the number of available averaged IASI profiles. The wettest (driest) coincidences are in bold (italic). The presences of dry layer (DL) and residual particle layer (RPL) are indicated.

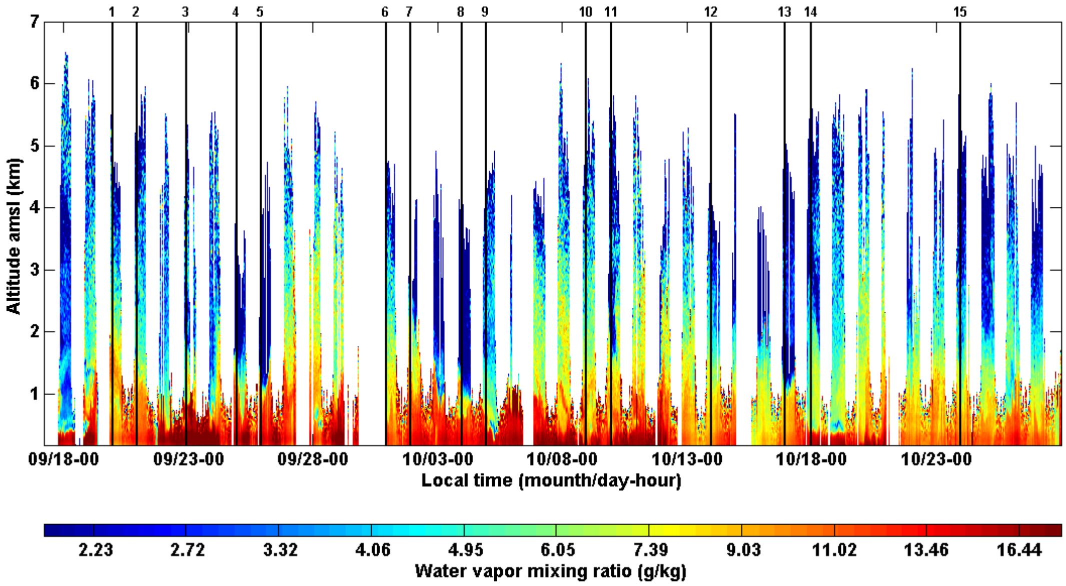
(CN) Month/Day - LT	*D* (km)	*N*	Observation
(1) 09/19 - 23:03	0.6	8	Atlantic-Spain origin, RPL below 2 km DL between 3 and 4 km, air mass off Eastern African coast
(2) 09/20 - 22:42	2.4	12	Atlantic - Spain origin DL between 2 and 4 km, air mass along Eastern African coast Subsidence between ~0 and 3 km above Southern Spain
(3) 09/22 - 21:59	19.3	12	Saharan origin between 1 and 5 km (Algeria-Morocco) DL between 1 and 4 km, air mass along Eastern African coast Strong ascent from ~0 to 3 km above Morocco
(4) 09/24 - 22:59	2.8	6	Atlantic - Spain origin with a RPL below 1.5 km (from Valencia) Strong subsidence from ~4 to 0 km
(5) 09/25 - 22:39	3.9	12	Atlantic-Southern Spain origin Below 1 km, RPL from Gibraltar with strong subsidence from ~3 km
(6) 09/30 - 22:33	55.2	6	Between 2 and 5 km, Saharan, France and Spain origin Below 1 km, RPL from Gibraltar (petrochemistry) Strong subsidence from 8 to 4 km above Moroccan sea coast
(7) 10/01 - 22:15	8.3	11	Northern Spain - Southwestern France origin RPL below 3 km from Valencia-Barcelona
(8) 10/03 - 23:12	5.2	12	Spain origin RPL below 2 km from Valencia coast
(9) 10/04 - 22:51	2.6	12	Saharan (Morocco) and Southern - Spain origin DL between 2 and 4 km
(10) 10/08 - 23:09	2.1	9	Tropical Atlantic - Spain origin DL below 4 km, mainly between 2 and 4 km May be Saharan air masses off African west coast
(11) 10/09 - 22:48	6.5	12	Tropical Atlantic - Spain origin Below 2 km, RPL from Valencia-Barcelona coast
(12) 10/13 - 23:06	8.5	11	Atlantic - Northern Spain origin RPL from Barcelona coast
(13) 10/16 - 22:03	45.8	6	Atlantic-Spain origin Strong subsidence from 4 to 1 km over Spain RPL from Valencia-Barcelona coast
(14) 10/17 - 23:24	5.2	11	Saharan origin with a strong DL below 3 km Strong subsidence from 4 to 0 km over Sahara
(15) 10/23 - 22:57	3.3	9	Spain origin - Long passage over the Mediterranean sea No significant aerosol layer

Table 3: Same as Table 2 for the ChArMEx experiment (2013 summer period).

(CN) Month/Day - LT	*D* (km)	*N*	Observation
(1) 06/10 - 21:59	5.8	12	Spain - Southern France origin Strong subsidence from ~6 to 2 km
(2) **06/11 - 23:18**	7.2	12	Atlantic - Spain origin RPL from Barcelona coast
(3) 06/12 - 22:57	3.5	12	Atlantic - Spain origin Strong subsidence from ~8 to 5 km
(4) 06/13 - 22:36	5.0	12	Atlantic-Southern France origin RPL from Perpignan coast Dry air mass at ~3 km (drying over the Pyrenees)
(5) **06/14 - 22:15**	2.2	12	France origin below 1 km Tropical Atlantic - Southern Spain origin above 1 km RPL from Valencia coast
(6) **06/15 - 21:54**	4.4	12	Mediterranean origin below 1 km Atlantic (off Moroccan coast) - Southern Spain origin above 1 km DL between ~2.5 and 4 km
(7) **06/16 - 23:15**	11.2	12	Mediterranean origin below 1 km Morocco-Algeria origin above 1 km DL between 1 and 5 km
(8) **06/17 - 22:54**	4.0	10	Atlantic - Morocco origin with a small DL ~0-7 km. Likely dust uptake between 0 and 4 km
(9) 06/20 - 23:33	47.1	12	Spain origin RPL from Valencia coast
(10) 06/22 - 22:51	3.0	12	Atlantic - Spain - Southwestern France origin Strong subsidence between ~9 and 4 km RPL from Barcelona coast
(11) *06/24 - 22:09*	8.2	12	Atlantic - Spain - Southwestern France origin Strong subsidence between ~7 and 0 km Dry air mass at ~2.5 km RPL from Barcelona coast
(12) 06/25 - 23:27	29.0	12	Atlantic - Spain - Southwestern France origin Strong subsidence between ~8 and 4 km RPL from Perpignan coast
(13) *06/27 - 22:48*	14.3	6	Northern Atlantic - France origin RPL from Perpignan coast
(14) 06/30 - 23:24	19.4	12	Northwestern Atlantic - France origin
(15) 07/02 - 22:42	4.2	10	France - Spain - Morocco origin with a DL between 3 and 5 km Strong wind-shear - RPL from Montpellier and Barcelona coasts

Table 4: Scores on the WVMR retrieval for the inter-comparisons between WALI and IASI (WALI-IASI), and WALI and ECMWF (WALI-ECMWF). The results are given for different atmospheric layers in terms of COR and RMSE for the two time periods.

Altitude range (km)	COR		RMSE (g/kg)	
September-October 2012				
	WALI-IASI	WALI-ECMWF	WALI_IASI	WALI-ECMWF
0.5-2.0	0.37	0.73	1.42	1.15
2.0-5.0	0.77	0.81	0.66	0.55
5.0-7.0	0.78	0.73	0.25	0.26
0.5-7.0	0.70	0.77	0.78	0.65
June-July 2013				
	WALI-IASI	WALI-ECMWF	WALI_IASI	WALI-ECMWF
0.5-2.0	0.15	0.74	1.80	2.42
2.0-5.0	0.70	0.91	1.34	1.16
5.0-7.0	0.75	0.77	0.67	0.66
0.5-7.0	0.59	0.82	1.28	1.43



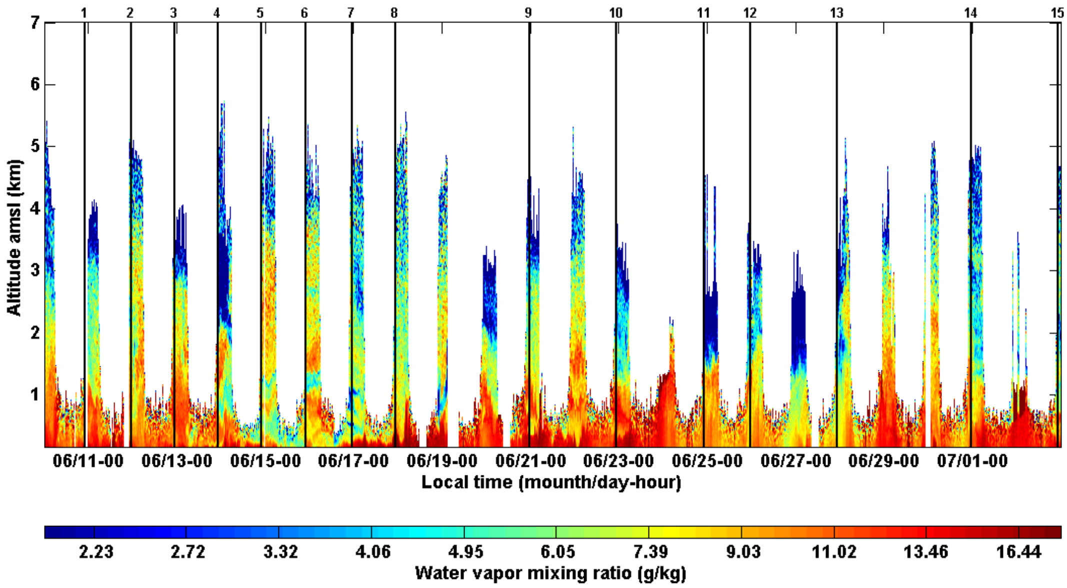


Figure 1: Time localization of the IASI profiles used for the inter-comparisons (vertical black lines identified with the CN at the top) with respect to the temporal WALI lidar WVMR (in g/kg) retrieval evolution as a function of altitude (in km) during the HyMeX (up) and ChArMEx (down) periods. Lidar profiles are given with a high temporal resolution of 5 minutes averaging. The color bar ranges from low water vapor mixing-ratio (blue) to high ones (red).

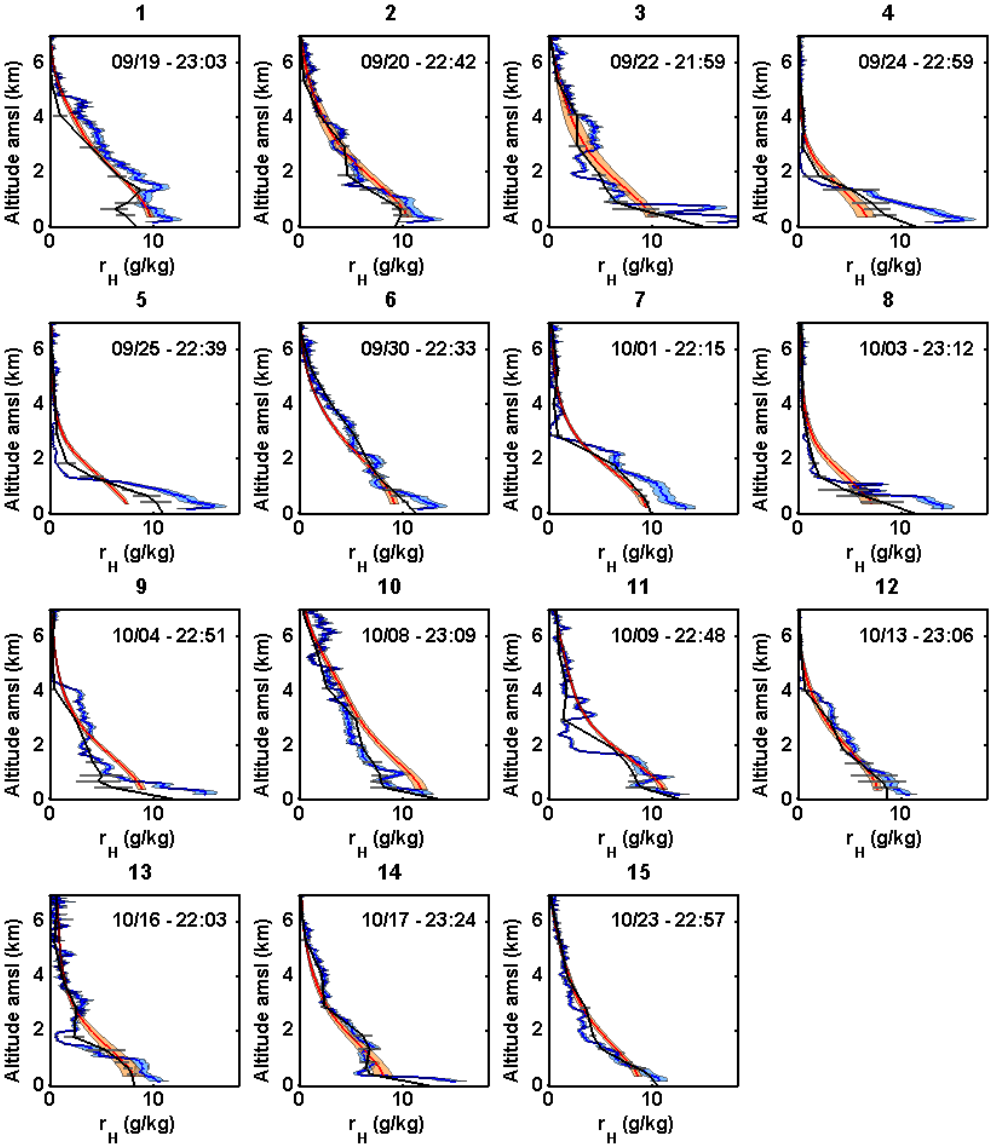


Figure 2: Comparisons of WVMR vertical profile retrieval as a function of altitude between: IASI (red), WALI lidar (blue) and ECMWF analysis (black), over Menorca during the HyMeX experiment (September and October 2012). The date and time of the IASI and WALI measurements are also given for each panel of individual profile in the form month/day HH:MM. The CN is given at the top of each figure.

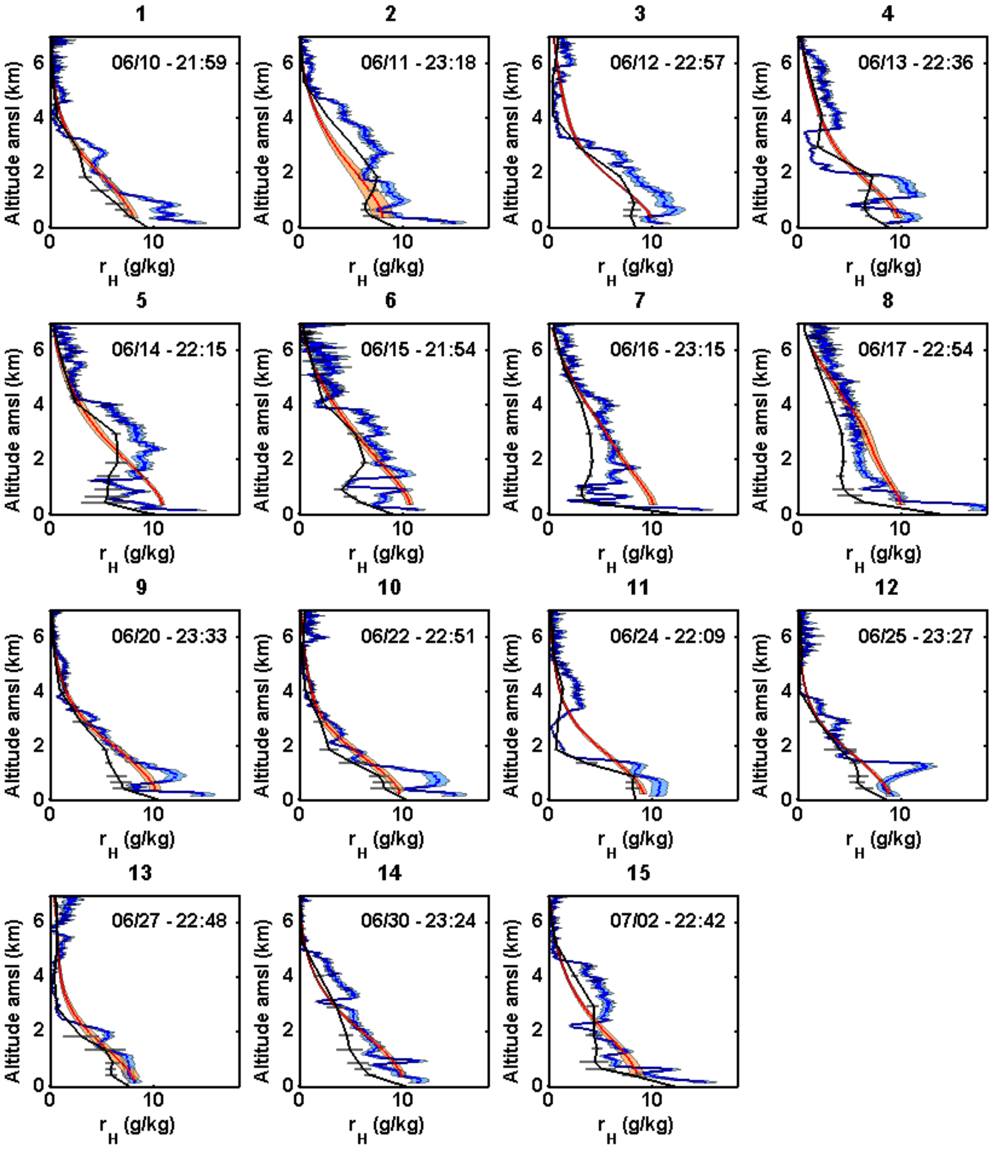


Figure 3: Same as Figure 2 for the ChArMEx experiment (June and July 2013).

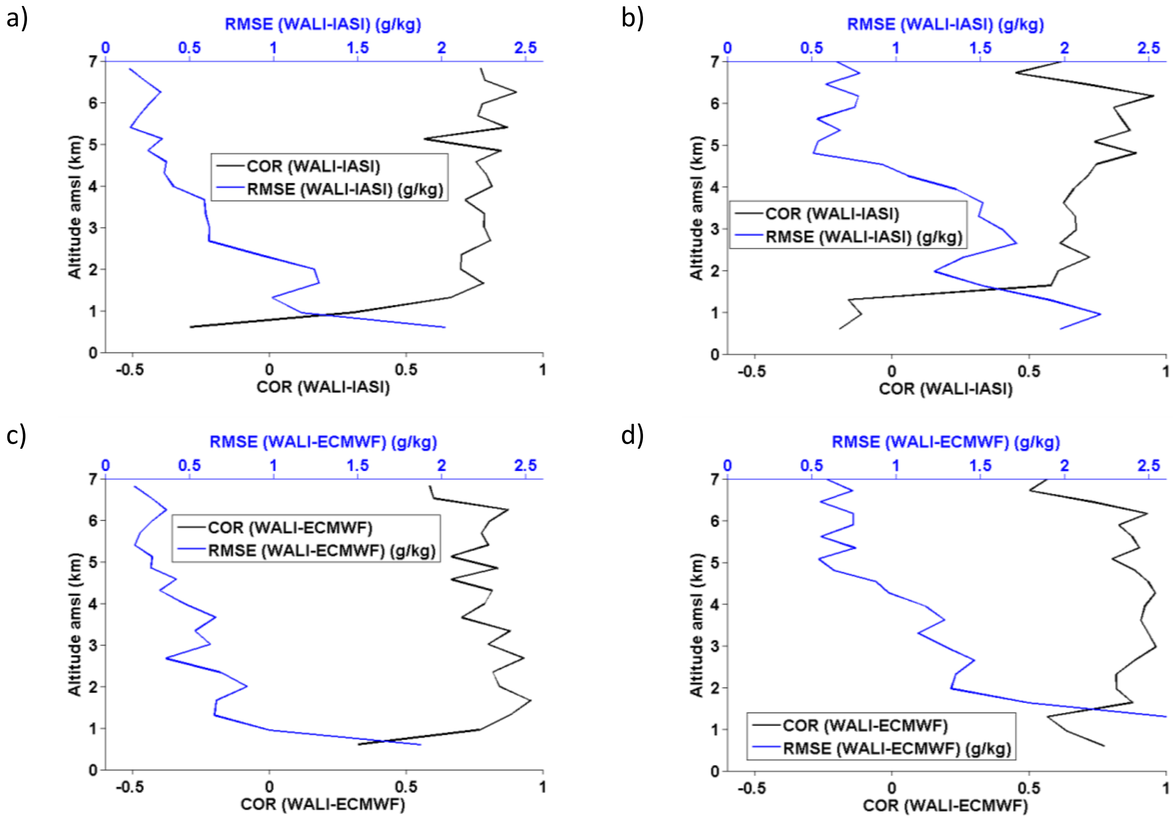


Figure 4: Evolution as a function of altitude of the RMSE in g/kg (blue) and COR (black) between IASI and WALI WVMR retrievals for HyMeX (a) and ChArMEx (b) periods, and between WALI WVMR retrieval and ECMWF analysis for HyMeX (c) and ChArMEx (d) periods.

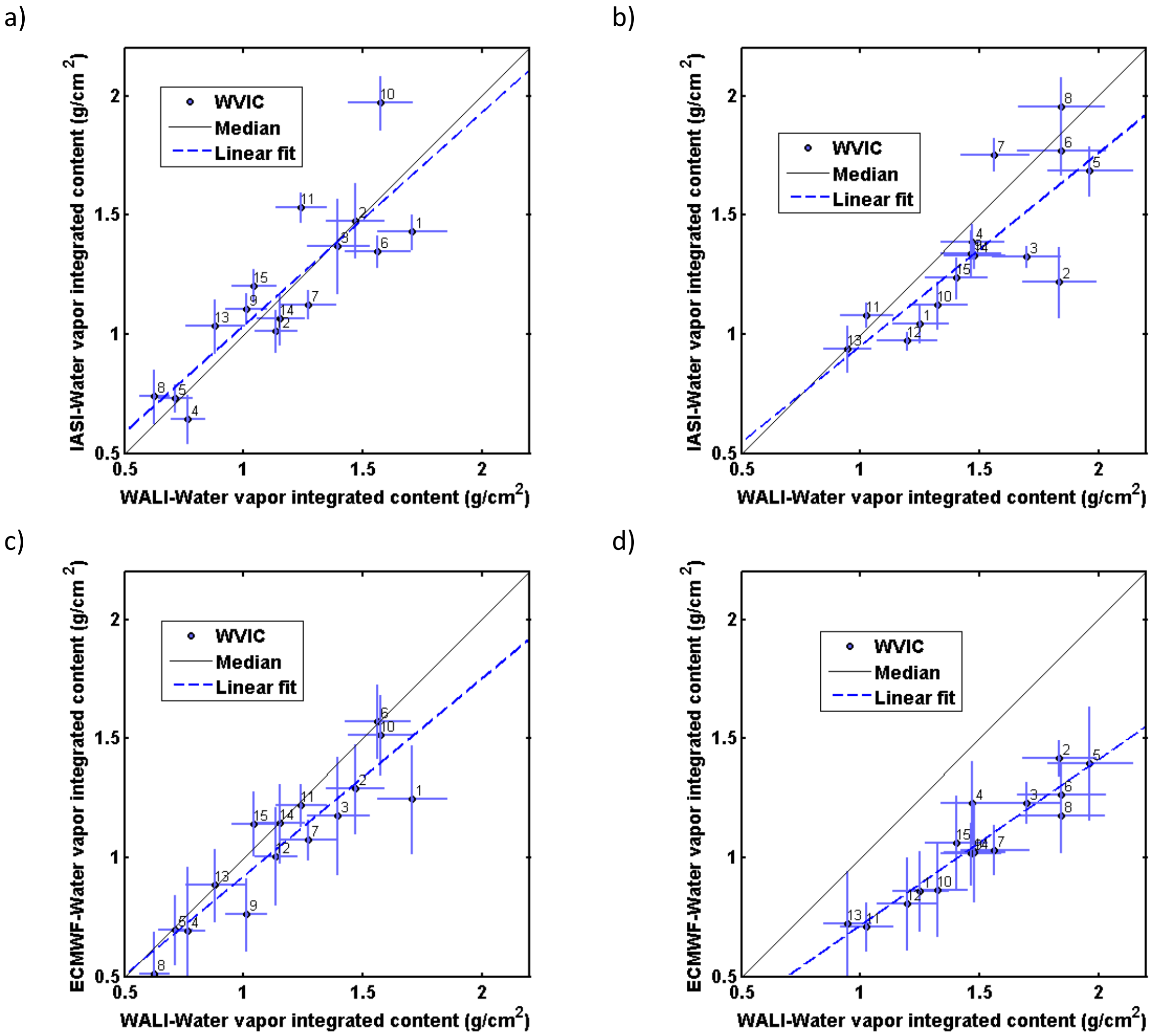
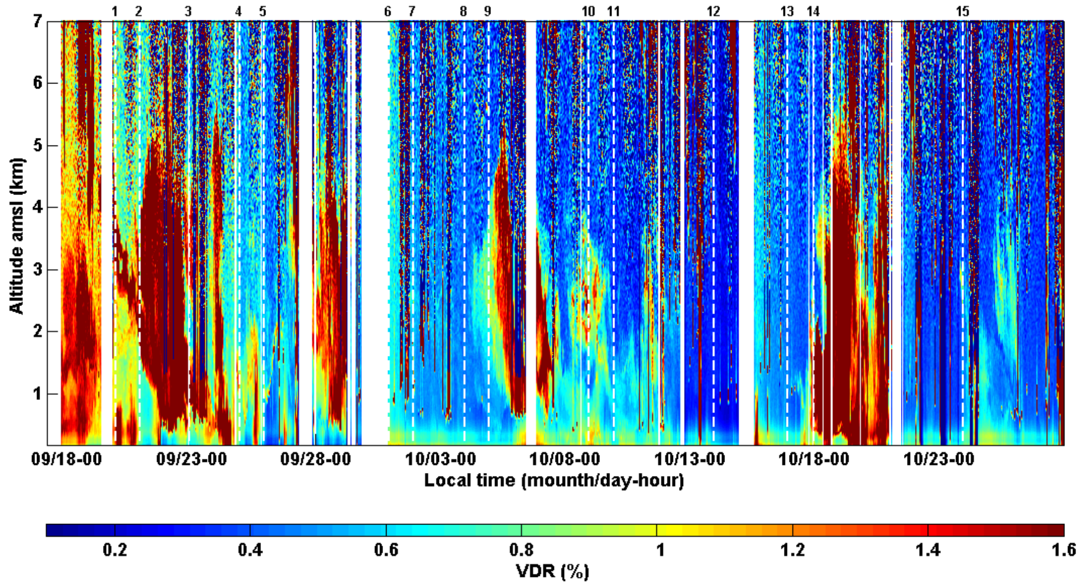


Figure 5: Water Vapor Integrated Content (WVIC in g/cm^2) as measured by WALI lidar against IASI WVIC retrieval for HyMeX (a) and ChArMEx (b) periods and against ECMWF WVIR for HyMeX (c) and ChArMEx periods (d). Line styles are given in each individual figure legend.



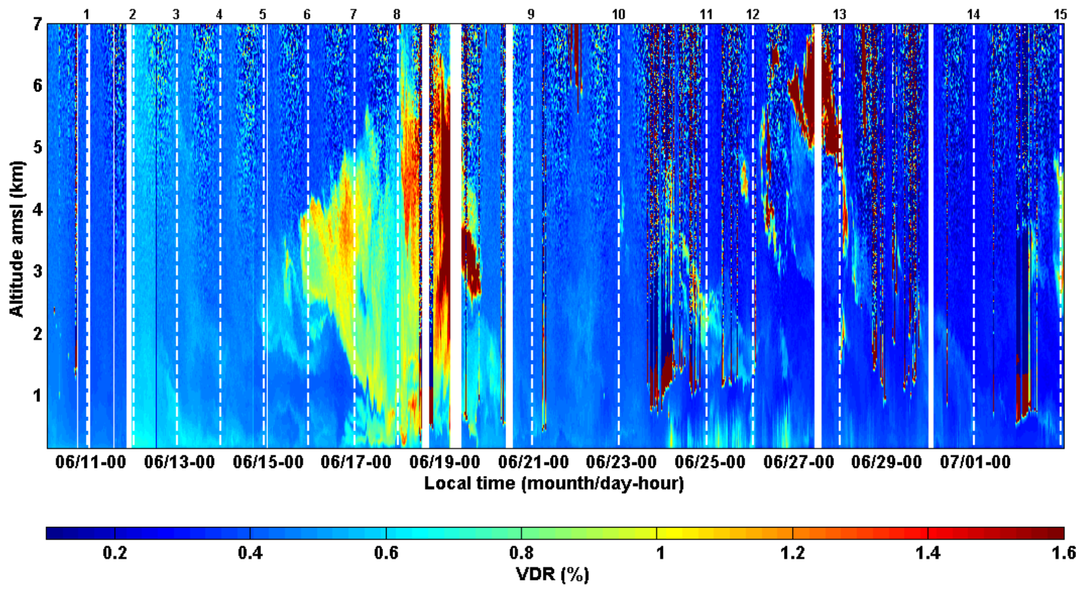


Figure 6: Volume Depolarization Ratio (VDR in %) evolution as a function of altitude (in km) during HyMeX (up) and ChArMEx (down) experiments comparisons. The vertical white dotted lines identify the coincidences with the CN at the top. The color bar ranges from low VDR (blue) to high ones (red).

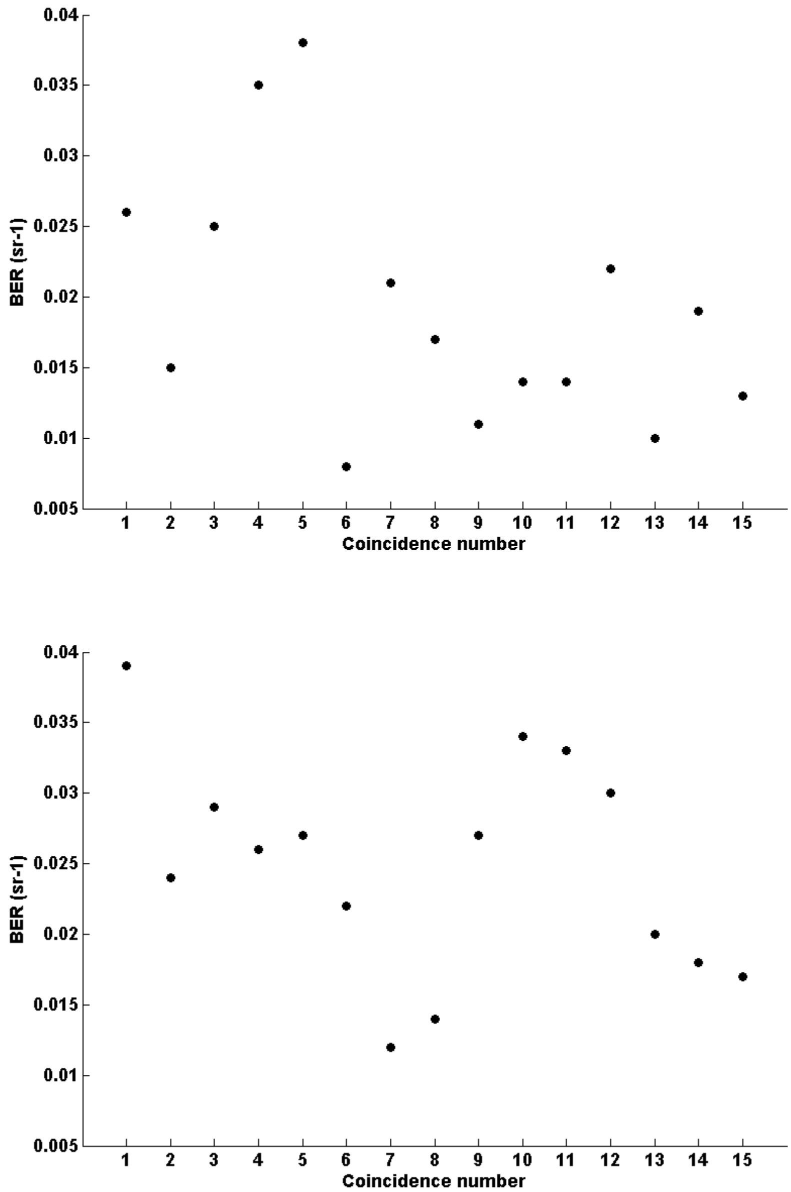


Figure 7: Column Backscatter to Extinction Ratio (BER in sr^{-1}) derived from WALI temporal evolution for the different inter-comparisons exercises during HyMeX (up) and ChArMEx (down) experiments. Abscissa represents the CN.

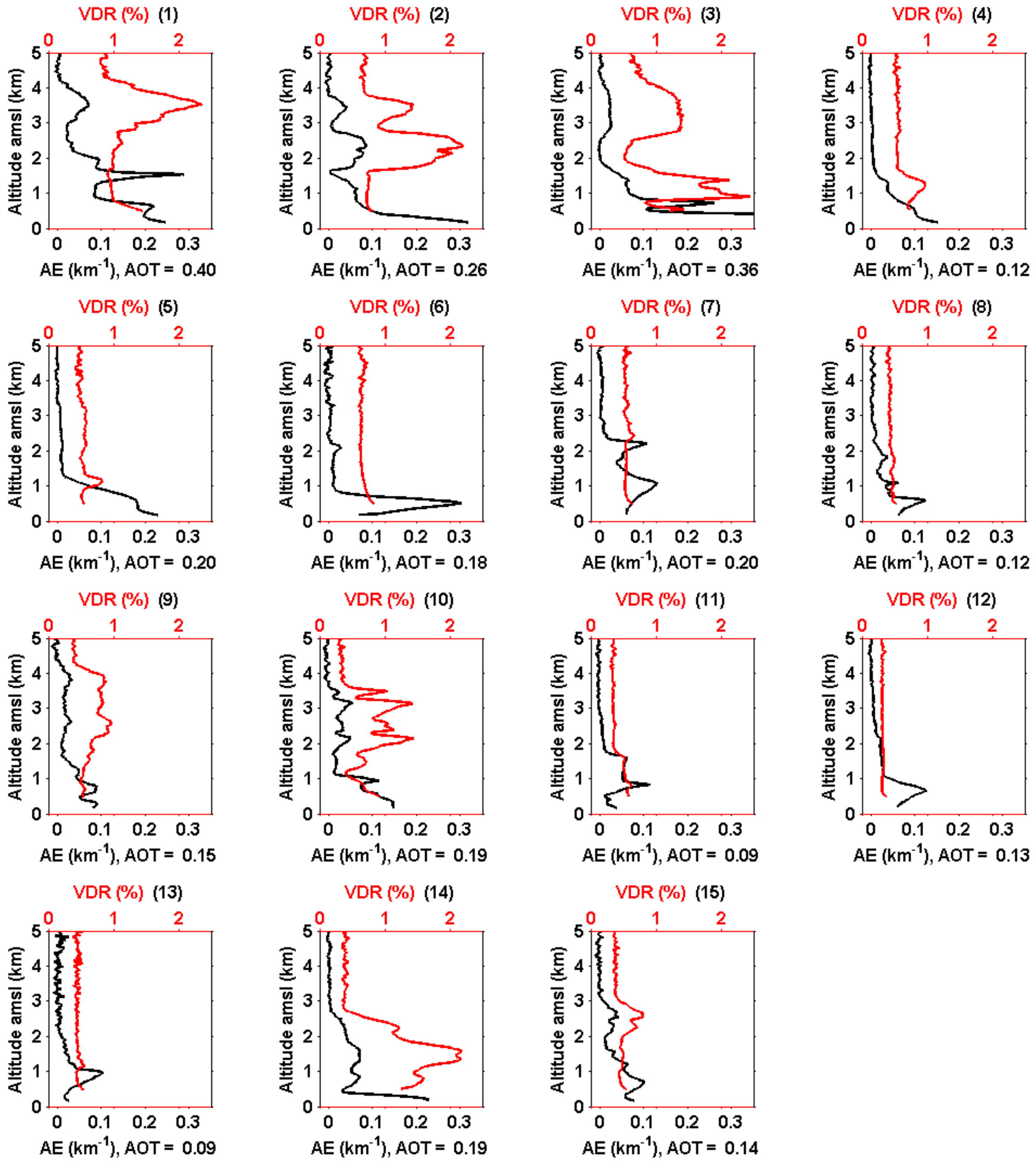


Figure 8: Aerosol extinction coefficient (AE in km^{-1}, black lines) and Volume Depolarization Ratio (VDR in %, red lines) as a function of altitude for the inter-comparisons cases during HyMeX experiment. For each individual case, the Atmospheric Optical Thickness (AOT) is also reported. The CN is given in black at the top of each figure.

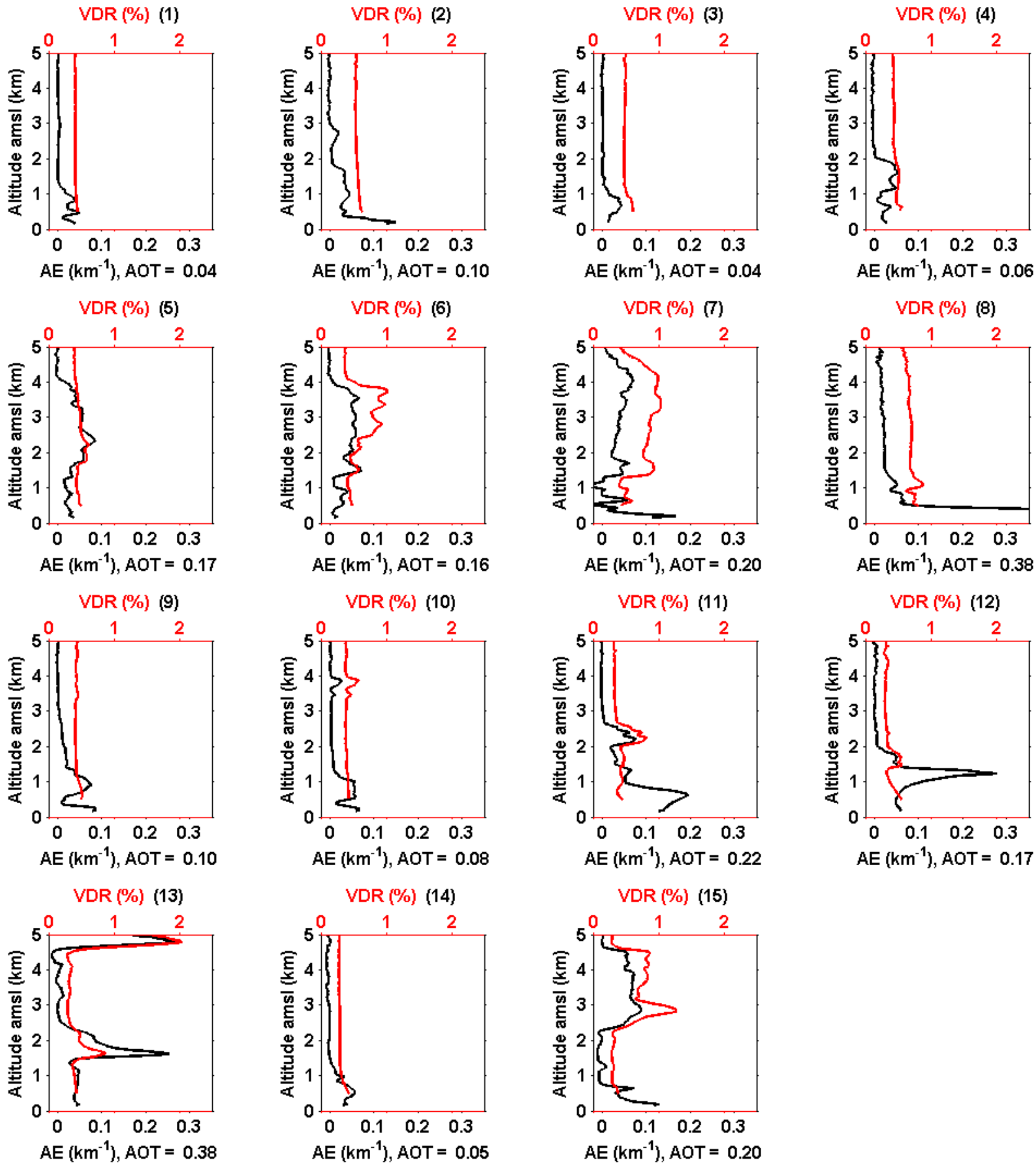


Figure 9: Same as Figure 8 for the ChArMEx experiment (June and July 2013).

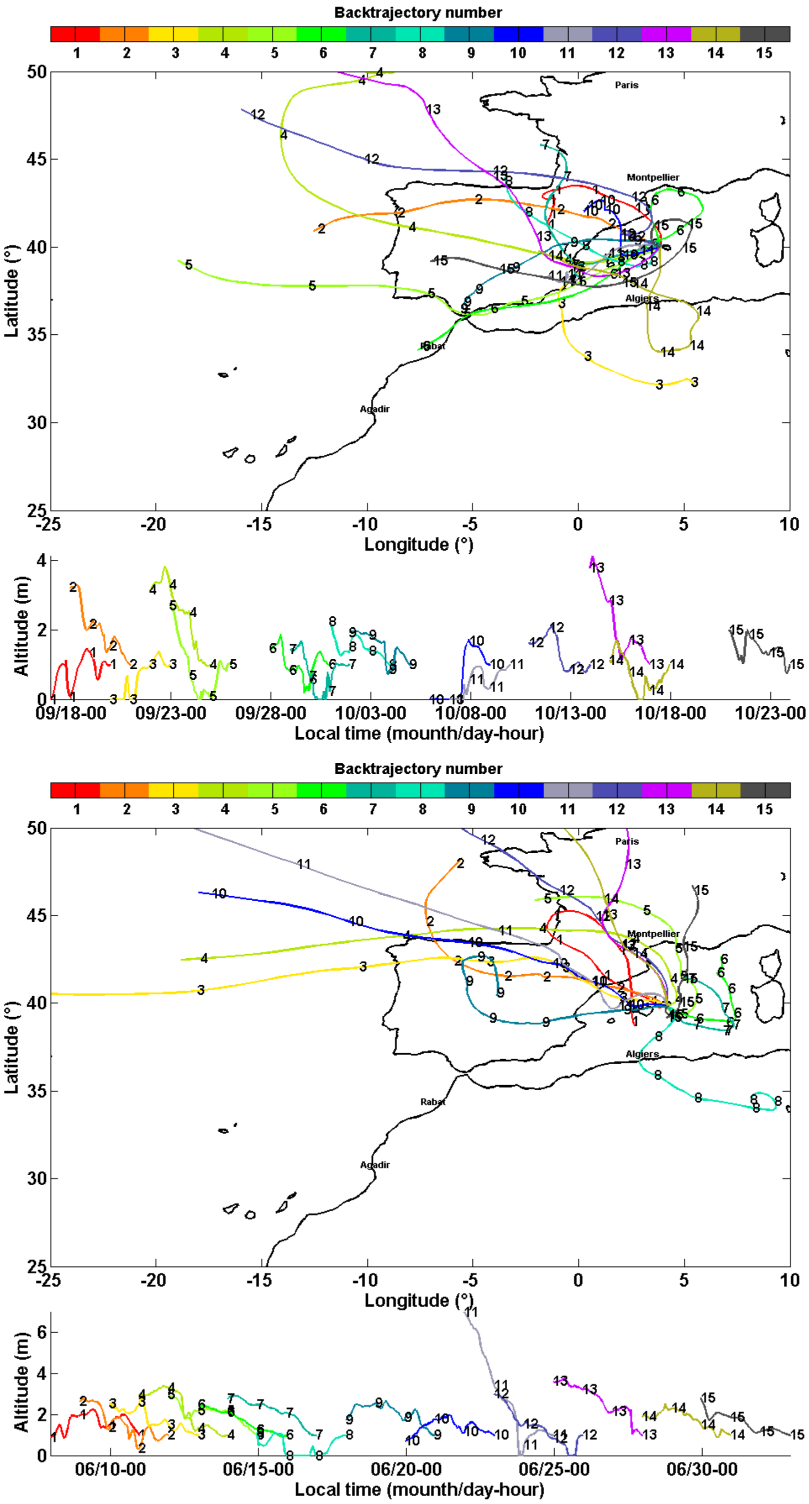
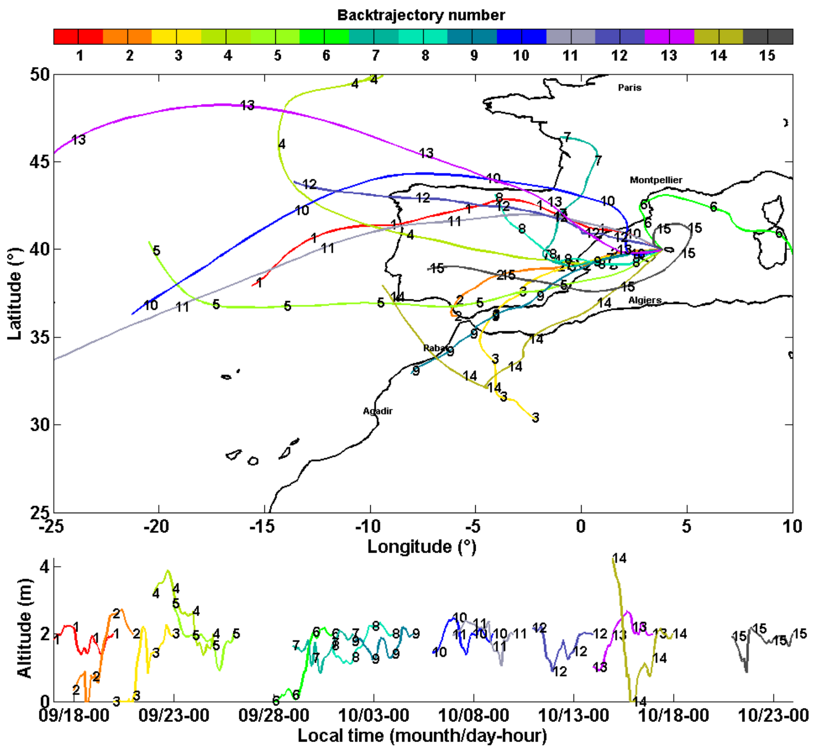


Figure 10: Back trajectories for each CN identified by the number shown on the curves. They have been computed using the Hysplit model (courtesy of NOAA Air Resources Laboratory; http://www.arl.noaa.gov). The end locations of the air masses are for the sites of Ciutadella

(up) and Mahon (down) for the HyMeX and ChArMEx time periods, respectively, at the altitudes of 1 km amsl.



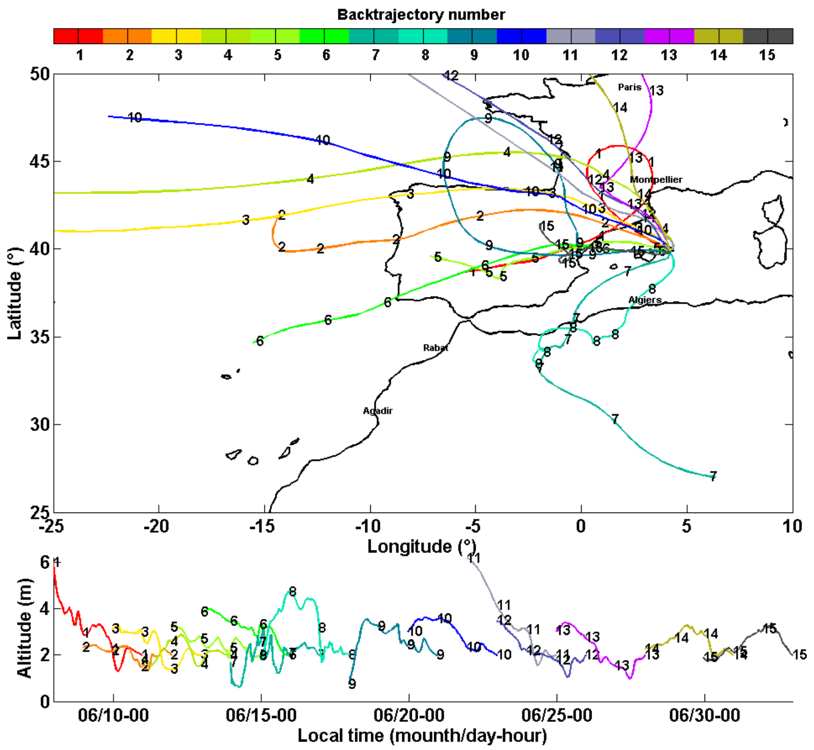
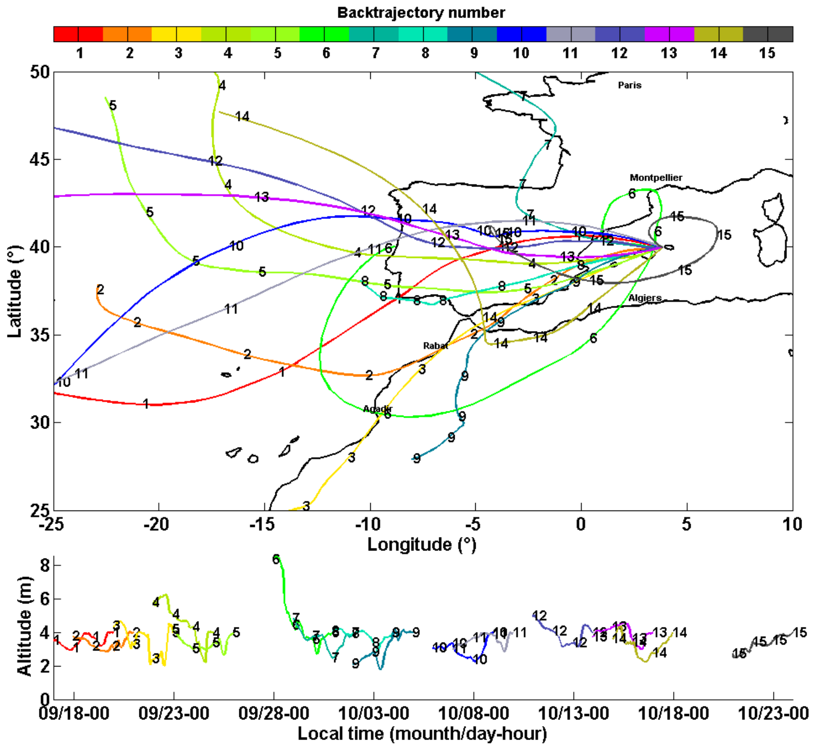


Figure 11: Same as Figure 10 for 2 km amsl.



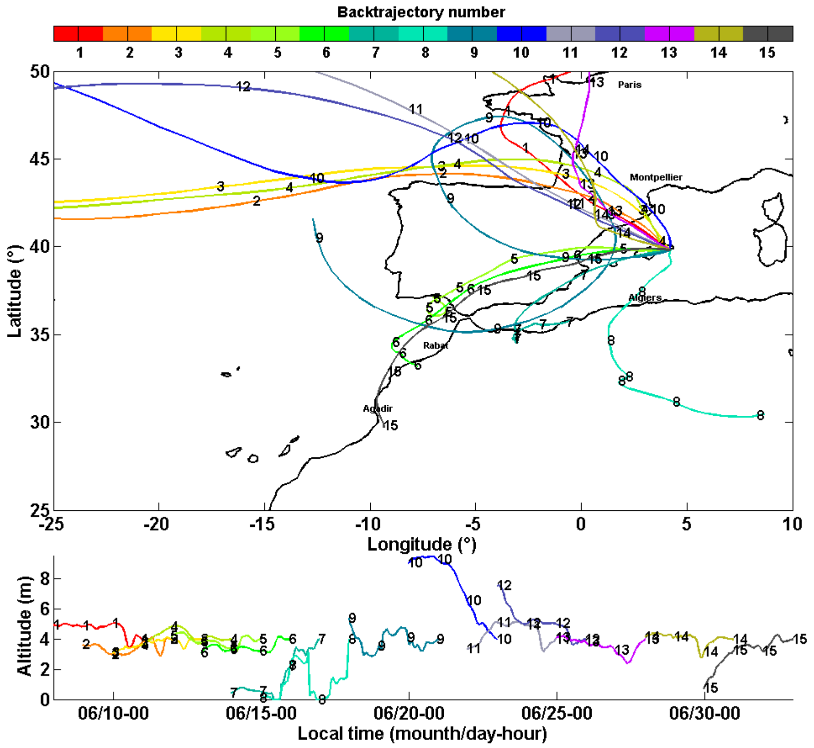


Figure 12: Same as Figure 10 for 4 km amsl.