

Abstract

 Evaluation of a regional air quality forecasting system for the Pacific 21 Northwest was carried out using a suite of surface and satellite observations. Wildfire events for the 2007 and 2008 fire seasons were simulated using the Air Information Report for Public Access and Community Tracking v.3 (AIRPACT-3) framework utilizing the Community Multi-scale Air Quality (CMAQ) model. Fire emissions were simulated using the BlueSky framework with fire locations determined by the Satellite Mapping Automated Reanalysis Tool for Fire Incident Reconciliation (SMARTFIRE). Plume rise was simulated using two different methods: the Fire Emission Production Simulator (FEPS) and the Sparse Matrix Operator Kernel Emissions (SMOKE) model. Predicted plume top heights were compared to the Cloud-Aerosol LIDAR with Orthogonal Polarization (CALIOP) instrument aboard the Cloud Aerosol LIDAR and Infrared Pathfinder Satellite Observation (CALIPSO) satellite. Carbon monoxide predictions were compared to the Atmospheric InfraRed Sounder (AIRS) instrument aboard the Aqua satellite. Horizontal distributions of column aerosol optical depth (AOD) were compared to retrievals by the Moderate Resolution Imaging Spectroradiometer (MODIS) instrument aboard the Aqua satellite. Model tropospheric nitrogen dioxide distributions were compared to retrievals from the Ozone Monitoring Instrument (OMI) aboard the Aura satellite. Surface ozone and PM2.5 predictions were compared to surface observations. The AIRPACT-3 model captured the location and transport direction of fire events well, but sometimes missed the timing of

 fire events and overall underestimated the PM2.5 impact of wildfire events at surface monitor locations. During the 2007 (2008) fire period the fractional biases (FB) of AIRPACT-3 for various pollutant observations included: average 24-hr PM2.5 FB=-33% (-27%); maximum daily average 8-hr ozone FB= -8% (+1%); AOD FB= -61% (-53%); total column CO FB= -10% (- 47 5%); and tropospheric column $NO₂$ FB= -39% (-28%). The bias in total column CO is within the range of expected error. Fractional biases of AIRPACT-3 plume tops were found to be -46% when compared in terms of above mean sea level (AMSL), but only -28% when compared in terms of above ground level (AGL), partly due to the under-estimation of AIRPACT-3 ground height in complex terrain that results from the 12-km grid-cell smoothing. We conclude that aerosol predictions were too low for locations 54 greater than \sim 100-300 km downwind from wildfire sources and that model predictions are likely under-predicting secondary organic aerosol (SOA) production due to a combination of very low VOC emission factors used in the United States Forest Service Consume model, an incomplete speciation of VOC to SOA precursors in SMOKE, and under-prediction by the SOA parameterization within CMAQ.

1 Introduction

1.1 MOTIVATION

The Pacific Northwest is home to a rural landscape that periodically

 experience large wildfires, especially during dry summers. Wildfire smoke and other particulate matter (PM) emitted into the atmosphere can cause severe health problems. Informing the public about upcoming poor air quality expected from fires requires a comprehensive knowledge of fire locations, land type being burned, terrain, wind direction, available moisture, timing, and other conditions. Reports generated by fire fighters are quickly provided to air quality managers by the United States Forest Service, but it is difficult to get an accurate assessment of wildfire conditions in remote locations with rough terrain, few access roads, and sparse air quality monitor distribution. Meteorological forecasts and chemistry transport models can be used to predict the air quality impacts of wildfire emissions, but the task is challenging (Simon et al., 2012). Satellite retrievals of air quality indicators provide a valuable asset that, when combined with surface measurements, can help to assess the validity of air quality models simulating large wildfire events. The analysis presented here utilizes multiple satellite products to evaluate simulations from the Air Information Report for Public Access and Community Tracking v.3 (AIRPACT-3) regional air quality model, which utilizes the BlueSky fire emissions framework and the Community Multi-scale Air Quality (CMAQ) model. As such, this work demonstrates how a suite of satellite products can be combined with in-situ observations to inform improvement of air quality forecast performance.

 The objective of this work is to report the level of performance and types of error that were found for modeled fire locations, plume heights, and

 pollutant concentrations simulated in AIRPACT-3 based on a combination of satellite products and surface pollutant observations. It is essential that 89 future AIRPACT versions accurately predict the impact of fires, given the very large fire seasons in recent history (e.g. 2012) and the expected increase of fire activity as the regional climate changes. We chose to use finalized activity reports to derive wildfire emissions, rather than forecast-mode data, so that we could focus on the emissions from known fire events and test the model's performance in a "best-case" scenario. We modeled wildfire events that occurred during the summers of 2007 and 2008 because of their interest to AIRPACT users, the extensive fire activity that occurred, and because satellite coverage throughout NASA's Afternoon Train (A-Train) of satellites was relatively complete. We focused on A-Train satellite data to keep 99 overpass times consistent (\sim 1:45 PDT) and because fire activity is best detected in the afternoon, when wildfires are most active. Simulations of the historically large fires that ignited in Idaho, Nevada, and Montana throughout July of 2007 provided great insight into AIRPACT-3 wildfire performance. In addition, the Northern California fires that ignited June 21, 2008 provided further valuable model information due to the technical challenge posed by the large fires that occurred on both sides of the southern boundary of the modeling domain.

1.2 FIRE ACTIVITY OF 2007 AND 2008

 The western US experienced abnormally dry winter and spring seasons in 2007, which led to a summer drought and extensive wildfire events in

 Idaho, Nevada, and Montana. Extreme temperatures and sparse precipitation during early summer 2007, coupled with lightning activity and several strong wind events, led to several expanding, long-lived fires. Precipitation events that started on August 17 slowed the expansion of wildfires and allowed fire fighters to contain many of the burning areas, though some fires continued to burn into September. The National Interagency Coordination Center (NICC) at the National Interagency Fire Center (NIFC; [http://www.predictiveservices.nifc.gov/\)](http://www.predictiveservices.nifc.gov/) reported that over 800,000 acres burned in Nevada during July 2007. By August 31 the Great Basin and Northern Rockies had wildfires that burned over 4 million acres, nearly twice the typical year-to-date area burned, with eight large fires or complexes having burned more than 100,000 acres each.

 The summer of 2008 was also dry but experienced significantly less fire activity across the US, except for California and parts of the southern U.S. Northern California, part of which is in the AIRPACT-3 domain, reported over 850,000 acres burned, which was nearly 9 times the 10-year average for that region. On June 20 - 21, 2008, widespread lightning started nearly one thousand fires in northern California and those in remote and difficult terrain burned for many days. Lightning storms in mid-August 2008 also caused numerous large fires in Idaho and Montana. The number of acres burned by state reported by the NICC NIFC is shown in Table 1 for 2007 and 131 2008. Analysis of O_3 and particulate matter enhancements at the Mt Bachelor Observatory by Wigder et al. (2013) identified 14 individual fire

plumes in 2008 and 6 in 2007.

 The analysis presented here includes results for two separate time periods: July 3 – August 22, 2007 and June 22 – August 27, 2008, which were chosen to include the largest annual fire events in the AIRPACT-3 domain. Details about each reported fire complex that burned during the analysis period are given in Table S1 of the Supplementary Materials. Fire events during the analysis periods that included at least one reported fire over 5,000 acres of burn area are shown in Fig. 1 (Fig. S1 includes labels for fire complex names).

2 Methods

2.1 AIRPACT-3 AIR QUALITY MODELING SYSTEM

 The AIRPACT-3 modeling system (Chen et al., 2008; Herron-Thorpe et al., 2010, 2012) simulates air quality in the Pacific Northwest with the CMAQ v4.6 chemical transport model (Byun and Schere, 2006). Area and non-road mobile emissions are from the 2002 EPA NEI, projected to 2005 using the EPA's Economic Growth Analysis System (EGAS) software; on-road mobile emissions are based on the EPA MOBILE v6.2; anthropogenic emissions for Canada are from the 2000 Greater Vancouver Regional District (GVRD) inventory; and biogenic emissions are obtained from the Biogenic Emissions Inventory System version 3 (BEIS-3). The AIRPACT-3 base emissions are spatially and temporally allocated using the Sparse Matrix Operator Kernel Emissions (SMOKE) v2.4 model while all fire emissions are processed with

 the SMOKE v2.7 model. The AIRPACT-3 domain includes a 95 x 95 grid of 12 km x 12 km cells using 21 layers from the surface to the lower stratosphere. The version of CMAQ used includes the SAPRC-99 chemical kinetic mechanism, the ISOROPIA inorganic aerosol equilibrium module, and the Secondary Organic Aerosols Model (SORGAM). Meteorology inputs for AIRPACT-3 were derived from forecasts by Mass and colleagues [\(http://www.atmos.washington.edu/mm5rt/;](http://www.atmos.washington.edu/mm5rt/) Mass et al., 2003) and preprocessed for CMAQ using the Meteorology Chemistry Interface Processor (MCIP). The Mesoscale Model v5 (MM5; Mass et al., 2003) was used for the year 2007 simulations while the Weather Research and Forecasting (WRF; Skamarock et al., 2005) model was used for the year 2008 simulations. Model of OZone And Related Tracers, version 4 (MOZART-4; Emmons et al., 2010) simulations produced at the National Center for Atmospheric Research (NCAR) were used as chemical boundary conditions around the AIRPACT-3 domain (Emmons et al., 2010; Herron-Thorpe et al., 2012). The MOZART-4 simulations included the assimilation of satellite CO column v4 retrievals from the Measurement Of Pollution In The Troposphere (MOPITT) instrument, a gas-correlation radiometer on-board the NASA Terra satellite (Deeter et al., 2010). The MOZART-4 emissions are the same as those used in Wespes et al. (2012), which include anthropogenic emissions based on the inventory developed by D. Streets for the NASA ARCTAS experiment (http://bio.cgrer.uiowa.edu/arctas/emission.html) and biomass burning emissions from FINN (Fire Inventory from NCAR, Wiedinmyer et al., 2011).

 Fire location, area, and emissions were calculated using BlueSky v3.1 data [\(http://www.airfire.org/bluesky\)](http://www.airfire.org/bluesky), which utilizes United States Forest Service fire reports and hotspot detects reported by the Hazard Mapping System (HMS) together in the Satellite Mapping Automated Reanalysis Tool for Fire Incident Reconciliation (SMARTFIRE; Larkin et al., 2009 and Raffuse et al., 2009). SMARTFIRE reports wildfire locations (Larkin et al., 2009; Strand et al., 2012), but is ultimately limited by the accuracy and completeness of the satellite detects and USFS reports filed. Air quality forecasts use the fire locations reported over the past 48-hours and assume them to persist throughout the simulation. However, the fire reports used in this model reanalysis are from the final SMARTFIRE archive, as distinct from the information reported in near real-time, which allows us to scrutinize the model performance independent of the near real-time fire reporting system.

 For this analysis, the BlueSky framework (Larkin et al., 2009; Raffuse et al., 2009) was operated in default mode, which includes the use of the Consume v3 (Ottmar et al., 2009), Fuel Characteristic Classification System v1 (FCCS; Riccardi et al., 2007), and Fire Emission Production Simulator v1 (FEPS; Anderson et al., 2004) software programs provided by the USFS. FCCS v1 provides vegetation type and corresponding fuels (Fig. 1) at 1-km resolution based on Bailey ecoregions and satellite-derived cover type, which provides input to Consume. Consume was developed empirically using a variety of vegetation types and fire conditions, providing fuel consumption and emissions by combustion phase (smoldering or flaming) data to FEPS.

 FEPS calculates the heat released and the individual pollutant emissions, based on combustion efficiency of the burn. The default behavior of BlueSky classifies fuels as "dry", unless otherwise reported by SMARTFIRE. This can result in large over-predictions during events that don't consume most available fuels, but generally it is reasonable to assume that fire activity occur in areas with dry fuels. A summary of the fire-related model pathways used for AIRPACT-3 is shown in Fig. 2.

 Two plume rise methods were used in this analysis, resulting in two sets of AIRPACT-3 model results. The first method uses the SMOKE-ready files created by BlueSky, which include hourly information, to explicitly set 211 the plume rise to what FEPS predicts. The second set of model simulations were performed using methods that bypassed the FEPS plume rise algorithm and instead converted standard BlueSky output to create daily input files for SMOKE. It is important to note that the two plume rise methods used are based upon the same heat flux and smoldering/flaming emissions ratios but results differ in two ways: 1) whereas FEPS plume rise method allocates all 217 smoldering emissions to the surface layer, the SMOKE plume rise method allows for smoldering emissions to be allocated throughout multiple layers near the surface; and 2) whereas FEPS plume rise method does not utilize meteorology or surface elevation when predicting flaming plume heights, the SMOKE plume rise method computes flaming plume heights as a function of buoyancy using the heat content predicted by BlueSky, modeled meteorology, and modeled terrain heights (Pouliot et al., 2005).

2.2 AQUA-MODIS AOD

 The Aqua satellite was launched in May 2002 carrying the Moderate Resolution Imaging Spectroradiometer (MODIS) as part of NASA's Afternoon- Train (A-Train) of Earth Observing Satellites (EOS). The Aqua-MODIS retrievals provide aerosol information at nearly the same time as the other A- Train instruments, allowing coincident multi-species analyses, as presented in this analysis. Aqua MODIS reliably retrieves Aerosol Optical Depth (AOD; τ) for much of the globe on a daily basis with a nadir footprint of 10 km. Algorithms described by Remer et al. (2005) are used to interpolate the 470 nm and 660 nm retrievals to provide a 550-nm AOD product (MYD04_L2 v5.1; Land_and_Ocean) where only the highest quality data (Quality Flag=3) is used. Typical AOD values at a clean site are below 0.3, while values over 236 1.0 are indicative of multiple scattering caused by high aerosol loading (i.e. heavy haze, biomass burning, or dust events). The maximum AOD values 238 historically retrieved by MODIS are \sim 5.0, but these are rare events. MODIS AOD error is not reported for each pixel but studies have validated the an error of 15%, which is influenced by unique aerosol composition, varied land 241 cover color, cloud fringes, and snow cover at high elevations (Levy et al., 2007 and Drury et al., 2008). MODIS AOD retrievals are useful in areas with no clouds but they have been shown to be biased low compared to AERONET and MISR (Kahn et. al, 2010 and Eck et. al, 2013).

 All MODIS AOD retrievals used in this analysis were projected to the AIRPACT-3 grid by using the pixel with the closest proximity to the center of

 each AIRPACT-3 grid-cell. This method gives a more detailed map than would otherwise be calculated using weighted spatial interpolation, and is suitable here since the MODIS spatial resolution is finer than AIRPACT-3. AIRPACT-3-simulated aerosol distributions were generated for all modeled aerosol species: nitrates, sulfates, ammonium, elemental carbon (EC), organic particulates, and coarse mode aerosols. AOD was calculated from AIRPACT-3 simulated aerosol species concentrations and size distributions using algorithms developed by Binkowski and Roselle (2003). This method uses the simulated aerosol total volume concentration for the Aitken and accumulation mode aerosols and their associated Mie extinction efficiencies to calculate AOD per modeled layer, which is then integrated vertically through the troposphere to yield the reported model AOD. An accurate approximation method from Evans and Fournier (1990) was used to calculate the Mie extinction efficiency factors. AIRPACT-3 grid-cells that did not have corresponding high-quality MODIS retrievals were omitted from the analysis.

2.3 AIRS CO

 In addition to MODIS, the Aqua satellite includes the Atmospheric Infra-Red Sounder (AIRS), which provides information about weather and trace gases. The AIRS instruments are an infrared spectrometer and a visible light/near-infrared photometer. The AIRS total column carbon monoxide level-2 v5 product used in this analysis (AIRX2RET) provides data reported on the Advanced Microwave Sounding Unit (AMSU) ground footprint, which varies from 36 km x 36 km to 50 km x 50 km. AIRS level-2

 v5 data includes 7 trapezoidal layers of CO mixing ratio in the troposphere and an averaging kernel matrix for the full 9-layer profile available in the support product files. In this study the AIRPACT-3 profiles were convolved with the AIRS averaging kernels as discussed in Olsen et al. (2007) and Maddy and Barnet (2008), and the total column CO values were then interpolated to the original AIRPACT-3 projection using a Delaunay triangulation scheme. The AIRS averaging kernel slightly reduces the AIRPACT-3 total column CO, with some loss of information in the lower troposphere and enhanced middle troposphere sensitivity (Herron-Thorpe et al., 2012). AIRS typically has only 1 degree of freedom in the troposphere, with its greatest sensitivity to the mid-troposphere. Thus AIRS retrievals likely underestimate total column CO for fire plumes contained within a shallow boundary layer. However, the convolution of the model with the AIRS averaging kernels should address potential comparison problems. The typical reported error in the AIRS CO product varies by layer, with moderate 285 error $(\sim 45\%)$ throughout the middle and upper troposphere and even larger 286 error $(\sim]60\%)$ in the lower troposphere. However, large CO values (e.g. 287 greater than 2.3E+18 molec./cm²), as the case with large fire plumes, are typically associated with very low errors (10-20%) throughout the layers.

2.4 OMI TROPOSPHERIC NO²

 The Aura satellite successfully joined the A-Train in July 2004, carrying multiple instruments that retrieve information about atmospheric chemistry. Although tropospheric ozone retrieved by the Ozone Monitoring Instrument

293 (OMI) is typically not precise enough for this wildfire analysis, the 294 tropospheric $NO₂$ columns provided by the Tropospheric Emission Monitoring 295 Internet Service (TEMIS; [http://www.temis.nl/airpollution/no2.html\)](http://www.temis.nl/airpollution/no2.html) are of 296 significant value. The Derivation of OMI tropospheric $NO₂$ (DOMINO) 297 algorithms calculate air mass factors (AMF), a priori profiles, stratospheric 298 $NO₂$, and ghost columns from the daily global Tracer Model v4 (TM4), which 299 is driven with meteorological fields from the European Centre of Medium-300 Range Forecasts (ECMWF) (Boersma et al., 2011). The product provides 301 tropospheric NO₂ column retrievals with a 13 km x 24 km footprint at nadir 302 with increasing footprint size as the observation moves off-nadir. A pixel's 303 "ghost column" (below cloud) is estimated from the a priori profile for the 304 pixel and OMI's retrieval of $NO₂$ above the cloud cover pressure level, with 305 vertical sensitivity defined by the averaging kernel. The sum of the OMI 306 ghost column and tropospheric column can be compared to a model column 307 for an estimate of model performance. However, when the model $NO₂$ profile 308 is convolved with the averaging kernel, the ghost column is no longer 309 required. Typical reported errors in the DOMINO product are lowest $(\sim 25\%)$ 310 where there is a large signal (e.g. over $2E+15$ molec/cm²) but errors are 311 typically much higher (\sim 50%) when the signal is considerably less.

 Since OMI's NO₂ averaging kernel shows decreasing sensitivity as the vertical profile approaches the surface, the result of applying the averaging 314 kernel to AIRPACT-3 $NO₂$ allows for essentially a "free troposphere" comparison with OMI. In this study we used OMI pixels with low cloud

 fraction (<35%) and convolved all AIRPACT-3 profiles with the OMI averaging kernel. AIRPACT-3 cells that fall within the spatial boundaries of each OMI pixel were averaged and interpolated, effectively reducing the resolution of the model results to equal that of the co-located OMI pixel, and then both were interpolated to the original AIRPACT-3 projection using a Delaunay triangulation scheme. This method works well for most areas but can lead to inconsistencies over areas with complex terrain (Herron-Thorpe 323 et al., 2010). Comparisons of CMAQ $NO₂$ to satellite retrievals also have inherent uncertainty associated with the rapid conversion of NOx to PAN and nitrate (Alvarado et al., 2010 and Akagi et al., 2012).

2.5 CALIOP AEROSOL DETECTION

 The Cloud Aerosol LIDAR and Infrared Pathfinder Satellite Observation (CALIPSO) satellite successfully joined the A-Train in April 2006, carrying the Cloud-Aerosol LIDAR with Orthogonal Polarization (CALIOP) instrument as its main payload. CALIOP transmits a linearly polarized laser pulse and then detects the light that is reflected back. Determining the aerosol type from this space-based LIDAR depends on the attenuated backscatter, altitude, location, surface type, and the volume depolarization (ratio of the perpendicular backscatter to the parallel backscatter of the laser light retrieved). Detailed information about the CALIOP data is in the CALIPSO Users Guide [\(http://www](http://www-calipso.larc.nasa.gov/resources/calipso_users_guide/)[calipso.larc.nasa.gov/resources/calipso_users_guide/\)](http://www-calipso.larc.nasa.gov/resources/calipso_users_guide/). The laser beam 338 diameter of CALIOP is \sim 90 meters at the Earth's surface, combined with a

 horizontal resolution along scan that varies from 333 m (surface) to 1 km (8.5 km to 20 km altitude). The v3.01 CALIOP level-2 Vertical Feature Mask (Liu et al., 2005; Mielonen et al., 2009; and Winker et al., 2009) product available from the NASA Langley Research Center Atmospheric Science Data Center was used to evaluate AIRPACT-3 plume top height performance. We evaluated plume top heights above mean sea level (AMSL) and above ground level (AGL), so that discrepancies in terrain height could be evaluated. For this analysis, we consider AGL plume heights to be relative to the ground level reported by the respective dataset.

2.6 DAILY REMOTE SENSING ACTIVITY

 In addition to the methods described above, we also assessed overall fire conditions using MODIS true-color imagery of smoke plumes with markers for hot-spot locations, available from the Land Atmosphere Near Real-time Capability for EOS (LANCE; USA subset 1; [http://lance](http://lance-modis.eosdis.nasa.gov/imagery/subsets/index.php?project=fas)[modis.eosdis.nasa.gov/imagery/subsets/index.php?project=fas\)](http://lance-modis.eosdis.nasa.gov/imagery/subsets/index.php?project=fas). A daily remote sensing log of the LANCE-MODIS imagery and corresponding remote sensing comparisons, derived from the AIRPACT-3 FEPS plume-rise scenario, was also compiled (Tables S2 – S5). Each fire region that was significantly over the signal-to-noise threshold was counted and tallied in the daily 358 remote-sensing log for AOD and tropospheric $NO₂$ comparisons. AIRS resolution did not allow us to identify "distinguishable events" and were not tallied. The horizontal footprint and sensitivity of each remote sensing instrument varies, thus distinguishable events counted in the log ranged

from strong isolated fires to large areas with numerous mixed plumes.

2.7 MODEL PERFORMANCE STATISTICS AND GROUND-SITE SELECTION

 Definitions of the model performance statistics used are shown in Table 2. Guidance on the treatment of negative values in satellite products suggests that long-term studies (e.g. with time-averaging) should retain the negative values so that no artificial bias is introduced for clean conditions (see [http://modis-atmos.gsfc.nasa.gov/MOD04_L2/format.html\)](http://modis-atmos.gsfc.nasa.gov/MOD04_L2/format.html). However, we were interested in short-term pollution events and chose to discard negative OMI and MODIS values. This approach helped us avoid spurious fractional statistics because it allowed little signal from the variance in "unpolluted" satellite retrievals and focused our statistics on "polluted" events. To assess the model performance for wildfire impacts, the ground- site analysis presented here uses combinations of 140 U.S. surface monitor locations where AIRPACT-3 predicted more than double the normal surface PM2.5 levels sometime during the analysis as an indicator of wildfire impacts. Surface monitor datasets that were excluded from the analysis had one or more of the following problems: no quality-controlled hourly dataset was available, the site was primarily indicative of urban emissions, the site was in Canada (AIRPACT-3 has no wildfire emissions in the Canadian part of the domain), or the site exhibited no distinguishable increase in surface PM2.5 during fire events. The 2007 analysis period had 67 qualified PM2.5 sites and 10 qualified ozone sites; while the 2008 analysis period had 82 qualified PM2.5 sites and 18 qualified ozone sites. The primary analysis of AOD,

385 tropospheric column $NO₂$, and total column CO includes all 140 site locations. For the purpose of generating model performance statistics, we assessed model performance at these discrete site locations rather than across the entire domain. This was done so that surface monitor observations and satellite retrievals could be compared more consistently, and so that the randomness of the location of usable retrievals did not skew our results spatially or with urban signatures. A more selective rural-sites-only subset includes 43 locations with no possible influence of transported urban pollution in the remote sensing records. This rural-sites-only subset is used for the "matched-threshold" analysis to help determine model performance for fire- polluted cases by only including instances where AIRPACT-3 and the 396 monitor/retrieval in question both surpassed a threshold value: 10 μ q/m³ for 397 the average 24-hr surface PM2.5, 0.3 for AOD, $1.0E+15$ molecules/cm² for 398 tropospheric column $NO₂$, and 1.9E+18 molecules/cm² for total column CO.

All surface monitor comparisons in this analysis (Fig. S2) were made using hourly data from the EPA Air Quality System

[\(http://www.epa.gov/ttn/airs/airsaqs/detaildata/downloadaqsdata.htm\)](http://www.epa.gov/ttn/airs/airsaqs/detaildata/downloadaqsdata.htm),

except for data from Mt. Bachelor Observatory (MBO) in the Oregon Cascade

mountains, which is not an AQS reporting site. The Mt. Bachelor Observatory

has been used to collect air quality data since 2004, including near-

405 continuous observations of CO, $O₃$, aerosol scattering and meteorological

- parameters, and various other chemical species during intensive campaigns.
- MBO is located at coordinates 43.98° N, 121.69° W at an elevation of 2.7

 km. The site has been used to investigate long-range transport of Asian pollution and biomass burning, regional wildfires, and other events including stratospheric intrusions (Weiss-Penzias et al., 2006; Ambrose et al., 2011; 411 Wigder et al., 2013). AIRPACT-3 $PM_{2.5}$ and carbon monoxide concentrations were extracted from the layer corresponding to a height of 2.7 km AMSL in the model for comparisons to Mt. Bachelor Observatory to account for the discrepancy in model surface height.

3 Results

3.1 AIRPACT-3 compared to AIRS, MODIS, and OMI

 Remote sensing of atmospheric gases and aerosols is limited by cloud conditions and the source signal strength at the relevant infrared/visible/UV 419 wavelengths. Maps of AOD, tropospheric $NO₂$ column, and total carbon monoxide column for analysis days in 2007 (2008) with favorable remote- sensing conditions are shown in Figs. 3, 4 (5, 6) for the SMOKE plume-rise scenario (see Figs. S3-S8 for the FEPS plume-rise scenario).

 On July 22, 2007, AIRPACT-3 under-predicted AOD related to fires in Montana, southern Idaho, and Nevada (Figs. 3, S3). AIRPACT-3 also under-425 predicted tropospheric column $NO₂$ in Nevada and Montana on July 22, 2007 426 but the largest modeled fires were not observed via remote sensing, in central Idaho near the Montana border, likely due to mismatch in timings of fire emissions and satellite detections. August 12, (Figs. 4, S4) and August 18 (Figs. S5a, S5b) show typical AIRPACT-3 comparisons during the largest

 fire periods in 2007. AIRPACT-3 under-estimated the fire-generated pollutants from N. California on June 29, 2008 (Figs. 5, S6) and missed pollutants transported from outside of the domain. AIRPACT-3 did better predicting fires in N. California on July 11, 2008 (Figs. S7a, S7b) but continued to miss fire-generated pollutants from outside of the domain. This is especially evident in Nevada when fire-generated AOD originating from south of the AIRPACT-3 domain is observed but not predicted, suggesting that boundary conditions derived from the MOZART-4 simulations under- predict the influence of fires from outside the domain. AIRPACT-3 did well predicting an interesting transport case on July 20, 2008 but over-predicted 440 the near-source pollutants in N. California/S. Oregon while under-predicting the transported aerosol from within the domain and over-predicting the transported CO from within the domain (Figs. 6, S8). In general, fire locations and air quality impacts were predicted well near fire sources, but AOD predictions were often too low in regions beyond 100 km downwind of large fires. Furthermore, AIRPACT-3 did not predict the observed fire impacts in Nevada that were transported from south of the domain.

 The Daily AOD Log for 2007 (2008) discussed in Table S2 (S4) notes that there were 44 (64) days in the period analyzed that confidently showed MODIS AOD due to fires: of the 176 (108) total discernible events, 8% (6%) were observed but not predicted, 37% (32%) were under-predicted, 30% (31%) were predicted well, 20% (18%) were over-predicted, and 5% (13%) were predicted but not observed. We found that the magnitude of predicted

 AOD that extended to large distances from sources inside the domain was under-predicted for 13% (31%) of discernible events. Additionally, we found that the magnitude of predicted AOD from sources outside the domain was under-predicted during 8 (27) of the 44 (64) days. There were also 2 (3) days where MODIS AOD clearly showed aerosol loading retained from the 458 previous day that were not predicted. The Daily $NO₂$ Log for 2007 (2008) in Table S3 (S5) also notes that there were 31 (44) days in the period analyzed 460 that confidently showed tropospheric $NO₂$ due to fires: of the 122 (76) total discernible events, 0% (4%) were observed but not predicted, 23% (13%) were under-predicted, 21% (30%) were predicted well, 48% (37%) were over-predicted, and 8% (16%) were predicted but not observed. There was 464 also one day (July 1, 2008) where OMI clearly showed tropospheric $NO₂$ loading retained from the previous day that was not predicted.

 Overall, AIRPACT was biased low for all analyzed pollutants for both the 2007 and 2008 timelines. In comparison, for non-fire periods across the whole domain, AIRPACT tends to over-estimate long-term average PM2.5 469 levels by \sim 3% (Chen et al., 2008). The 2007 (2008) fractional biases of the SMOKE plume rise scenario for all 140 sites were -61% (-53%) for AOD, - 471 39% (-28%) for tropospheric column $NO₂$, and -10% (-5%) for total column 472 CO. The FEPS plume rise scenario changed results by a few percent with fractional biases of -66% (-58%), -38% (-26%), and -13% (-7%), respectively (Table 3). In comparison, the fractional biases for the matched-threshold analysis of the SMOKE plume rise scenario for all 43 rural sites

(where both the model and satellite retrieval were greater than 0.3 AOD,

1.9E+18 VCD CO, or 1.0E+15 VCD NO2) were -101% (-105%), -98% (-

93%), and -10% (-9%), respectively. The fractional biases for the matched-

479 threshold analysis of the FEPS plume rise scenario were -117% (-125%), -

97% (-90%), and -18% (-12%), respectively (Table 4). The biases in total

column CO are within the reported retrieval error, and thus are not

482 significant. The low tropospheric $NO₂$ biases were greater in magnitude than

483 the reported retrieval errors, and mostly driven by the lack of $NO₂$ coming in

from south of the domain. The low AOD biases were much greater in

magnitude than the expected retrieval error, indicating persistent problems

with AIRPACT-3 aerosol predictions.

3.2 CALIPSO PLUME TOP HEIGHT COMPARISON

 CALIOP retrievals were compared to AIRPACT aerosols across the model domain when CALIPSO passed over the Idaho and California wildfire smoke plumes during the analysis periods of 2007 and 2008, respectively. There were many instances where both AIRPACT-3 and CALIOP showed the presence of fire-related aerosol pollution at similar heights. In 2007 (2008), CALIOP retrievals showed aerosol pollution over 328 (383) unique AIRPACT grid cells across Nevada, Idaho, and Canada (California, Oregon, Washington, and Canada), while 218 (281) and 219 (275) of those grid cells had AIRPACT-3 aerosol pollution in the SMOKE and FEPS plume rise scenarios.

498 There was moderate linear correlation $(r^2=0.41$ for FEPS plume rise;

 r^2 =0.50 for SMOKE plume rise) between AIRPACT-3 and CALIPSO plume top heights AMSL, when both showed the presence of an aerosol subtype (Fig. 7). On average, in 2007 (2008) the AIRPACT-3 FEPS plume-rise scenario under-predicted plume top heights AMSL by 3.1±2.3 km (2.5±1.5 km), while the SMOKE plume-rise scenario under-predicted plume top heights AMSL by 3.1 ± 2.0 km (2.2 ±1.6 km). There were many instances in which AGL comparisons were reasonable but dissimilar terrain heights resulted in large under-predictions in plume top heights AMSL. The horizontal resolution of AIRPACT smoothes the surface elevation in complex terrain so that it is consistently lower relative to CALIOP retrievals, and is a large source of uncertainty when evaluating AIRPACT plume tops. We found smaller linear 510 correlation (r^2 =0.18 for FEPS plume rise; r^2 =0.24 for SMOKE plume rise) between AIRPACT-3 and CALIPSO plume tops heights AGL (Table 5 and Fig. 7). On average, though, in 2007 (2008) the AIRPACT-3 FEPS plume-rise 513 scenario under-predicted plume top heights AGL by 1.4 ± 2.3 km (1.0 ± 1.2) km) while the SMOKE plume-rise scenario under-predicted plume top heights AGL by 1.5±1.9 km (0.9±1.3 km). This is consistent with a national study using a similar modeling structure, where CMAQ plume heights were under-517 predicted by ~20%, relative to CALIOP retrievals (Raffuse et al., 2012).

3.3 SURFACE CONCENTRATION RESULTS

 From July 3 to Aug. 22, 2007 (June 22 to Aug. 27, 2008) the daily 24- 520 hr average $PM_{2.5}$ was averaged across 67 (82) sites and the maximum daily 8-hr average ozone was averaged across 10 (18) sites for modeled and

 measured concentrations. The "all sites" comparison (Fig. 8) shows that maximum daily 8-hr surface ozone was generally under-predicted by 2 – 8 ppb in 2007, which might be expected with simulations of ozone in the presence of aerosols (Alvarado and Prinn, 2009). The maximum daily 8-hr ozone was nearly matched in 2008. In general, AIRPACT-3 predicted changes in ozone that were similar to what was observed across the region. The timeline also shows that AIRPACT-3 generally under-predicted daily 529 surface PM2.5 averages by 2 - 5 μ g/m³ and followed the measured curve 530 closely except for gross over-prediction of surface $PM_{2.5}$ concentrations from August 14 – 16, 2007 and July 12 – 13, 2008.

3.4 PM2.5 NAAQS COMPARISONS

 AIRPACT-3 daily 24-hr PM2.5 was assessed from a policy standpoint 534 for both the daily (35 μ g/m³) and annual (12 μ g/m³) National Ambient Air Quality Standards (NAAQS) threshold values. For each site, we calculated the number of days when both the model results and the observations showed PM_{2.5} concentrations greater than the NAAQS. We tallied the number of these days during the analysis period, for 67 sites in 2007 and 82 sites in 2008. For the FEPS plume-rise scenario we found: 97.7% of the data pairs were in agreement, with values less than the daily threshold; 0.2% of the data pairs were in agreement, with values higher than the daily threshold; 0.3% of the data pairs included observations higher than the daily threshold, with no such model prediction; and 1.8% of the data pairs included model predictions higher than the daily threshold, with no such observation. The

 SMOKE plume-rise scenario reduced the number of model predictions that were higher than the daily threshold, with no such observation, by 27% (or 1.3% of the total data pairs).

 In terms of the annual threshold, the FEPS plume-rise scenario showed that: 90.7% of the data pairs were in agreement, with values less than the annual threshold; 1.8% of the data pairs were in agreement, with values higher than the annual threshold; 4% of the data pairs included observations higher than the annual threshold, with no such model prediction; and 3.5% of the data pairs included model predictions higher than the annual threshold, with no such observations. The SMOKE plume-rise scenario increased the number of data pairs that were in agreement, with values higher than the annual threshold, by 17% (2.1% of the total data pairs). 557 Further details of the $PM_{2.5}$ NAAQS comparison are in Table 6 and Figs. S9-S10.

3.5 MT. BACHELOR OBSERVATORY COMPARISON

 Hourly observed and predicted AIRPACT-3 values for PM, carbon monoxide, and ozone at Mt. Bachelor Observatory during the 2008 California wildfires (Fig. 9) show how AIRPACT-3 generally does with medium-range transport of wildfire emissions. There is evidence of model under-prediction, especially in the FEPS plume-rise scenario, but the SMOKE plume-rise scenario resulted in over-prediction of CO for most fire events. There was generally good agreement of the timing of pollution events but occasionally the timing was off by a day, as occurred on August 8-9 (Fig. 9). Note that PM

568 for AIRPACT-3 in the Mt. Bachelor analysis is reported as $PM_{2.5}$ but the observations are of sub-micron aerosols converted from scattering observations using the method described in Wigder et al., (2013), which can have large uncertainty when there is significant variance in the aerosol size distribution (Akagi et al., 2012).

 On July 20, 2008, there was a large transport event that carried pollutants northwest from the fires in California until reaching the coast of Oregon where the plume was diverted inland to the northeast, sweeping across Oregon (Figs. 6, S8, S11). MBO measurements of sub-micron PM 577 were between 80 and 120 μ g/m³ from midnight to noon, and between 20 and \pm 45 µg/m³ for the proceeding 24 hours. AIRPACT-3 predictions of carbon 579 monoxide and $PM_{2.5}$ were well timed with monitor observations, but the AIRPACT-3 FEPS plume-rise scenario consistently under-predicted CO and PM concentrations during the event while the SMOKE plume rise scenario did better on average but still under-predicted PM. The event did not have emissions from outside the domain that significantly contributed to the plume, but some model aerosols were clearly lost to the domain boundary. However, the aerosol transported out of the boundary was not enough to explain well-predicted carbon monoxide combined with 30%-50% under- predictions in PM. There was a smaller event with similar comparisons between observations and predictions on July 25, 2008 as well. Throughout the 2008 MBO analysis dates, AIRPACT-3 generally under-predicted aerosols when CO was predicted well and over-predicted CO when aerosols were

 predicted well. This is consistent with other observations that show 592 AIRPACT-3 PM2.5/CO ratios to be low at locations greater than \sim 100 km from the fire location. Observations on July 20, July 25, and August 9 594 resulted in PM1/CO ratios of \sim 0.3 ug/m³/ppbv, higher than the ratios observed for fires in closer proximity to MBO, which has been previously interpreted to indicate SOA formation during plume transport (Wigder et al., 2013).

 The remote sensing comparison of the unique event on July 20, 2008 confirmed a consistent negative bias in predicted transported aerosols, even where CO in the SMOKE plume-rise scenario agreed well with AIRS. MODIS observed AOD values as high as 1.2 directly northwest of MBO, with lower values near 0.4 directly over the site. AIRPACT-3 only predicted AOD of 0.1 to 0.4 through the region of the large plume over those same regions around MBO (Fig. 6). AIRS also retrieved good quality carbon monoxide columns west of MBO, in the more concentrated part of the plume, showing a model 606 under-bias of \sim 10%. Tropospheric NO₂ columns over the transported portion of the plume were below the signal to noise threshold of OMI.

4 Discussion

 AIRPACT-3 correctly predicted which regions were impacted by fires in Idaho, Montana, Nevada, California, and Oregon during the summers of 2007 and 2008. This is reflected in the comparisons to AIRS carbon monoxide, 612 OMI tropospheric $NO₂$, and MODIS AOD, which all exhibited good

 spatiotemporal correlation to AIRPACT-3. General model performance was quite similar between the two years, which suggests that the differences from using MM5 in 2007 and WRF in 2008 did not have a significant effect on 616 the chemical transport modeling during the fire events.

 The SMOKE plume-rise scenario exhibited the best comparisons, with 618 average fractional biases at \sim 2 p.m. for AOD, tropospheric column NO₂ and total column CO found to be -61%, -39%, and -10% during the 2007 fire period, respectively; while during the 2008 fire period the average fractional biases were -53%, -28%, and -5% respectively. Surface concentrations of 622 PM $_{2.5}$ were also reasonable, especially in the SMOKE plume rise scenarios, which lifted some of the surface emissions aloft and constrained large plume top heights. The fractional bias of daily average 24-hr PM2.5 was found to be approximately -30% during both fire periods. Fractional biases of AIRPACT-3 plume tops were found to be -46% above mean sea level (AMSL), but only -28% above ground level (AGL), partly due to the under-estimation of AIRPACT-3 elevation in complex terrain. Underestimation of plume heights, which affects transport, may be partly responsible for under- prediction in transported aerosols. However, the under-prediction of SOA in model simulations is likely the largest source of model error, especially when we consider that other species, such as CO, were not under-predicted by such large magnitudes.

 Fire emissions generated from south of the domain were not well represented in AIRPACT-3 chemical boundary conditions derived from

 MOZART-4; a few events in 2008 appeared to be significantly affected by those under-predictions in boundary condition concentrations. This is consistent with the analysis of Pfister et al. (2011) that showed FINN emission factors were too low in the 2008 California fire simulations due to a misclassification of fuel type. MOZART-4 showed general agreement with the 641 background values of CO and O_3 (Fig. 9), but missed the high values expected from fires due to the coarse model resolution and the underestimation of fire emissions and plume height. Thus AIRPACT-3 model performance would benefit from revised methods to better represent fire influence on AIRPACT-3 boundary conditions.

 Comparisons of AIRPACT-3 plumes with CALIOP show that the dynamics of plume dispersion in the model are greatly affected by errors in surface terrain and vertical plume distribution and their interaction with the wind profiles. There is also evidence that the underestimation of terrain height in AIRPACT-3 and the overestimation of plume-top heights AGL could be compensating errors in some of the FEPS plume rise scenarios.

652 AIRPACT-3 tropospheric $NO₂$ was generally under-predicted, but there were occasionally what appeared to be large overestimates of tropospheric NO₂ over active fire regions (Figs. S5a, S5b, S7a, S7b). It is important to 655 note that these large tropospheric $NO₂$ predictions shown are a direct result of our application of the OMI averaging kernel, which weights the upper troposphere with a factor greater than one. In most cases, the plumes are low enough to the ground that the averaging kernel causes a net reduction in

659 AIRPACT-3 tropospheric $NO₂$ columns. However, in cases where FEPS 660 considerably over-predicted plume top height, the modeled tropospheric $NO₂$ column convolved with the averaging kernel caused a spike much higher than that of the original AIRPACT-3 results. The effect still occurs in the SMOKE plume rise scenario, though there are fewer extreme instances. 664 Furthermore, the OMI tropospheric $NO₂$ algorithms have large errors over 665 wildfires due to a combination of the a priori profiles used that assume $NO₂$ is concentrated near the surface, the high aerosol loadings emitted, and issues with comparisons over complex terrain (Boersma et al., 2011).

 Collectively, the results of this analysis show that AIRPACT-3 can over- predict surface fire emissions and occasionally under-predict fire emissions aloft which, coupled with discrepancies in modeled surface elevation, significantly affects the ability of AIRPACT-3 to accurately predict downwind surface concentrations of transported pollutants in complex terrain. Our analysis shows that AIRPACT-3 CO performs quite well when compared to surface concentrations (Fig 9) and AIRS total column retrievals (Figs. 5, 6, S5a, S5b). This is in contrast to the frequent underestimates of transported aerosols that were evident in AIRPACT-3 predictions of surface PM2.5 (Fig. 9) and AOD (Figs. 4, 5, 6, S5a). Satellite comparisons clearly show that when modeled CO across the domain is largely in close agreement with observations, aerosol performance systematically degrades with distance from the fire source. Akagi et al. (2011) and Yokelson et al. (2013) suggest that the emission factors for VOCs used in CONSUME-3 (Hardy, 1996 and

 Ward et al., 1989) should be much higher. This underestimation in VOC emissions further exacerbates known under-predictions of SOA in CMAQ, which can be a significant fraction of the total PM2.5 for plumes transported large distances (Wigder et al., 2013; Strand et al., 2012; Hu et al., 2008; Heilman et al., 2013) and is highly variable (Jolleys et al., 2012; Yokelson et al., 2009; Vakkari et al., 2014).

5 Conclusions & Future Work

 In general, AIRPACT-3 over-predicts pollutant concentrations due to near-source surface emissions from fires and under-predicts concentrations associated with long-range transport both from within the domain and outside the domain. Most fire locations are captured by the BlueSky SMARTFIRE tool, but there are occasionally fires predicted that are poorly timed or are missed. Our analysis suggests that total fire emissions in the domain are, overall, modestly under-predicted. Although we have shown that AIRPACT-3 chemical boundary conditions largely underestimate fire- emissions from outside the domain, this problem does not explain most under-predictions that occur at ground sites. The under-predictions are instead likely due to a combination of some or all of the following: 1) underestimates of area burned in the SMARTFIRE feed; 2) underestimates of fuel mass, especially in shrub-lands and other vegetation types that have sparse woody fuels but are classified with zero dead woody fuels in the FCCS; 3) underestimates of VOC emissions in the Consume model; 4) under-predictions of SOA production in CMAQ, thus causing under-predictions of PM

 in plumes that travel large distances; and 5) terrain height in the AIRPACT-3 model is too smooth in mountainous areas, causing problems with the elevation of emissions and dynamics of transport. Under-predictions in fire size also scale directly with under-predictions in plume top heights, since heat content of a fire is directly proportional to the total fuel in Consume, which adds uncertainty to predictions of transport.

 The high-resolution MODIS AOD retrievals provided considerable insight into AIRPACT aerosol performance. We also feel that alternative retrieval algorithms better suited for fire plume conditions might address some of the errors associated with AIRS and OMI trace gas comparisons. Furthermore, we recognize that coupling fire dynamics with meteorological 716 simulations, such as in the WRF-Fire framework (Coen et al., 2013; Kochanski et al., 2013; Mandel et al., 2011) may be the best method for forecasts once WRF-Fire simulations can be generated fast enough. We have recently updated the system to AIRPACT-4, which includes 4 km x 4 km horizontal grid cells and the SMOKE plume-rise method, in addition to updated BlueSky software which includes higher resolution fuel loading in FCCS and an updated SMARTFIRE (v2). Canadian fires within the model domain will be included, starting in 2015, but AIRPACT-4 would still benefit by having chemical boundary conditions that accurately represent smoke originating from outside the AIRPACT domain. Planned updates to the AIRPACT vertical layer spacing in the middle troposphere should also help model performance during fire emissions transport events.

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1069 **8 Tables & Figures**

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Table 1: Total annual fires and acres burned by state

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1073 NIFC Sources:
1074 http://www.nif

1074 http://www.nifc.gov/fireInfo/fireInfo_stats_YTD2007.html

1075 http://www.nifc.gov/fireInfo/fireInfo_stats_YTD2008.html

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1078 Table 2: Definitions of Model Comparison Statistics (Chen et al., 2008)

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Table 3: Summary of FEPS plume-rise scenario comparisons (SMOKE plumerise scenario shown in parentheses when different) from July 3 to August 23, 2007 (top) and June 22 to August 27, 2008 (bottom).

July 3 - Aug. 22, 2007

June 22 -Aug. 27, 2008

Table 4: Summary of FEPS plume-rise scenario matched threshold comparison (SMOKE plume-rise scenario shown in parentheses when different) from July 3 to August 23, 2007 (top) and June 22 to August 27, 2008 (bottom). "Matched Threshold" refers to both model and observation values being removed from the analysis if either is below the threshold in combination with satellite statistics using rural sites only.

Species	A24-hr PM2.5 $(\mu g/m^3)$	AOD	Tot. Col. CO $(E+18$ molec./cm ²)	Trop. Col. $NO2$ $(E+15$ molec./cm ²)
Source	EPA AQS	MODIS	AIRS	OMI
Threshold	10	0.3	1.9	1.0
Paired Points	555	150	356	599
Correlation (r)	0.4(0.5)	0.0	0.3(0.4)	0.2
Measured Mean	16.8	0.5	2.1	1.7
Mean Bias	$5.9(-0.1)$	-0.3	$-0.3(-0.2)$	-1.1
Mean Error	19.1 (12.1)	0.4	0.4	1.2
Normalized Mean Bias (%)	$24(-3)$	-66 (-47)	$-15(-8)$	-59 (-60)
Normalized Mean Error (%)	104 (70)	77 (84)	17 (19)	68
Fractional Bias (%)	$-38(-36)$	$-117(-101)$	$-18(-10)$	-97 (-98)
Fractional Error (%)	$80(-10)$	123 (115)	19	101 (102)

July 3 - Aug. 22, 2007

June 22 -Aug. 27, 2008

Table 5: Plume top height model comparisons with CALIOP for the FEPS plume-rise scenario (SMOKE plume-rise scenario shown in parentheses when different). Please note that some plumes contribute multiple paired points.

Table 6: PM2.5 National Ambient Air Quality Standards summary for both 2007 and 2008 fire periods analyzed per site per day for the FEPS plume-rise scenario (change due to SMOKE plume-rise scenario shown in parentheses).

Fig. 1: Fire events with individual burn areas greater than 5000 acres during the analysis periods of 2007 (orange) and 2008 (red). Total fuel loading derived from the FCCS v1 is also shown for the AIRPACT-3 domain.

FCCS Fuels and Modeled Fires

Fig. 2: Fire-related modeling pathways used in the AIRPACT-3 simulations.

Fire-Related Model Pathways

Fig. 3: AOD (left column), tropospheric NO2 columns (middle column), and total carbon monoxide columns (right column) for July 22, 2007 (\sim 2 p.m. LST) with NASA EOS retrieval (top row), AIRPACT-3 with SMOKE plume rise (middle row), and differences (bottom row). Grey color indicates no or lowquality data from the satellite retrieval and exclusion from analysis. Values greater than the color scale maximum are shown as pink in the AIRPACT-3 and NASA EOS maps. Values outside the range of the difference color scales are shown as saturated blue/red.

NO₂ Difference: AIRPACT - OMI (E+15 molec./cm²) CO Difference: AIRPACT - AIRS (E+18 molec./cm²)

AOD Difference: AIRPACT - MODIS

Fig. 4: AOD (left column), tropospheric $NO₂$ columns (middle column), and total carbon monoxide columns (right column) for August 12, 2007 (\sim 2 p.m. LST) with NASA EOS retrieval (top row), AIRPACT-3 with SMOKE plume rise (middle row), and differences (bottom row). Grey color indicates no or lowquality data from the satellite retrieval and exclusion from analysis. Values greater than the color scale maximum are shown as pink in the AIRPACT-3 and NASA EOS maps. Values outside the range of the difference color scales are shown as saturated blue/red.

Fig. 5: AOD (left column), tropospheric $NO₂$ columns (middle column), and total carbon monoxide columns (right column) for June 29, 2008 (\sim 2 p.m. LST) with NASA EOS retrieval (top row), AIRPACT-3 with SMOKE plume rise (middle row), and differences (bottom row). Grey color indicates no or lowquality data from the satellite retrieval and exclusion from analysis. Values greater than the color scale maximum are shown as pink in the AIRPACT-3 and NASA EOS maps. Values outside the range of the difference color scales are shown as saturated blue/red.

Fig. 6: AOD (left column), tropospheric $NO₂$ columns (middle column), and total carbon monoxide columns (right column) for July 20, 2008 (\sim 2 p.m. LST) with NASA EOS retrieval (top row), AIRPACT-3 with SMOKE plume rise (middle row), and differences (bottom row). Grey color indicates no or lowquality data from the satellite retrieval and exclusion from analysis. Values greater than the color scale maximum are shown as pink in the AIRPACT-3 and NASA EOS maps. Values outside the range of the difference color scales are shown as saturated blue/red. Mt Bachelor is shown as a black triangle near central Oregon.

 -1.0 -0.8 -0.6 -0.4 -0.2 0.0 0.2 0.4 0.6 0.8 1.0 **AOD Difference: AIRPACT - MODIS**

 -1.0 -0.8 -0.6 -0.4 -0.2 -0.0 -0.2 -0.4 -0.6 -0.8 -1.0 -1.0 -0.8 -0.6 -0.4 -0.2 -0.0 -0.2 -0.4 -0.6 -0.8 -1.0 NO $_2$ Difference: AIRPACT - OMI (E+15 molec./cm²) CO Difference: AIRPACT - AIRS (E+18 molec./cm²)

Fig. 7: AIRPACT-3 vs. CALIOP plume top heights for 2007 (red) and 2008 (blue) when CALIPSO passed over the Idaho and California wildfires, respectively (\sim 2 p.m. LST). Plume top heights above sea level (left) and above ground level (right) are shown for both the FEPS plume rise (open circle) and SMOKE plume rise (solid dot) scenarios. Note that plume top heights are only shown for locations where both CALIOP and AIRPACT-3 showed an aerosol plume.

Fig. 8: July 3 to August 22, 2007 (top) Daily 24-hr average $PM_{2.5}$ averaged across 67 sites (a) and Max Daily 8-hr average ozone averaged across 10 sites (b); June 22 to August 27, 2008 (bottom) Daily 24-hr average PM2.5 averaged across 82 sites (c) and Max Daily 8-hr average ozone averaged across 18 sites (d) from. Model simulations are shown in red with squares (FEPS plume rise) and orange dotted (SMOKE plume rise) while observations are shown in dotted blue with diamonds.

Fig. 9: Particulate Matter (top), carbon monoxide (middle), and ozone (bottom) at Mt. Bachelor Observatory for July 12 to August 21, 2008. AIRPACT-3 model simulations are shown in red (FEPS plume rise) and orange (SMOKE plume rise), MOZART-4 model simulations are shown in black, and observations are shown in dotted blue. Note that aerosols for AIRPACT-3 are reported as $PM_{2.5}$ and observed aerosols are sub-micron aerosols converted from scattering observations using the method described in Wigder et al., (2013).

