Interactive comment on "Impact of the North Atlantic Oscillation on the variations of aerosol ground levels through local processes over Europe" by S. Jerez et al.

Anonymous Referee #2 Received and published: 31 July 2013

Referee Comments on the manuscript "Impact of the North Atlantic Oscillation on the variations of aerosol ground levels through local processes over Europe" by S. Jerez et al. submitted to ACPD.

Author's response

General comments:

This work presents an attempt to characterize the variations of aerosol ground levels over Europe caused by local processes and NAO patterns. In terms of scientific quality and significance, the subject of this paper is of scientific interest but the presentation of the work needs a number of revisions to make it suitable for publication with ACP. My main concern is focused on the following issues: the evaluation of the model results that the authors state it will be part of another paper, the use of NAO index from CPC (NOAA) without describing how it relates to the ECMWF data used for the MM5 that drive the air quality simulations and the lack of specific and detailed information in several parts of the text. For example, how can the authors discuss about the NAO patterns without including the Atlantic Ocean in their modeling domain?

We thank the reviewer for the time devoted to revise the manuscript and acknowledge the concerns raised. In particular, we have made an effort to argument better the model validation, to discuss the robustness of our results despite the data source and method employed for computing the NAO index at the time-scale analyzed here, and to amend the lack of required details. Please, find below the answers to every single comment.

The title of the manuscript reflects part of the contents of the paper, but since the results are based on non-realistic model simulations, I would suggest a slight change in the title: "Impact of the North Atlantic Oscillation on the variations of aerosol ground levels through multiannual model simulations over Europe using fixed anthropogenic emissions"

We agree with the reviewer that the original title was not entirely faithful to the manuscript contents and, after considering the reviewer's suggestion, decided to change to the following title: 'Impact of the North Atlantic Oscillation on European aerosol ground levels through local processes: a seasonal model-based assessment using fixed anthropogenic emissions'.

The specific comments that follow will help to clarify important issues and suggest a number of changes to be made in the text. My suggestion for publishing this work with Atmospheric Chemistry and Physics is to reconsider after the major revision comments have been addressed.

All the comments below have been carefully considered in the revision of the manuscript. We sincerely appreciate the feedback provided.

Specific comments:

Abstract:

I cannot see how the objectives described in the last sentence are met in the work presented here. How can we improve the predictability of climate-air quality interactions based on the findings of this work? This is not supported by the conclusions. I suggest that the abstract follows the conclusions more carefully and avoid using statements that cannot be substantiated.

We have removed the last sentence of the abstract. Its content is now only discussed in the conclusion section.

Introduction:

1. Page 13891, line15: The influence of radiation and temperature on gas-phase chemistry has been studied by numerous researchers worldwide. The authors should cite here more publications besides Katragkou et al. (2010).

Now we cited also the works by Katragkou et al. (2010); Forkel et al. (2012); Jiménez-Guerrero et al. (2012); Meier et al. (2012); Colette et al. (2013) and Hedegaard et al. (2013). Please find the detailed references in the revised version of the manuscript.

2. Page13892, line 29: How did the authors disregard the contribution from the large scale transport mechanisms? This is not adequately explained in the text. And why would they want to disregard a physical mechanism that strengthens the realistic representation of the atmospheric conditions?

We understand the reviewer's incredulity on this option and have now made an effort to further clarify these two points in the introduction:

1. We disregard the impact of the large-scale transport mechanisms by using constant climatological boundary conditions for the aerosol concentrations.

2. The rationale for that is to allow focusing/isolating on the role of the local processes, unlike previous works that have been focused on the role of the large-scale transport processes. Each contribution can go in one different direction, which makes the problem difficult to be faced from an integrated approach without having trying before to disentangle the role of each mechanism. For example, according to other authors, negative NAO phases would promote the large-scale advection of pollutants from North America into Europe, preventing at the same time the Affrican dust intrusions in south-western Europe, and also, according to our results, favoring cleaner atmospheres in southern Europe in winter by enhancing precipitation and wind there. Hence, in view of the complexity of the problem raising by the mix of scales and mechanisms, we find it interesting to isolate and assess each contribution separately. Naturally, we acknowledge that such option does not provide the most realistic framework (as we noted in the text) but helps improving our understanding of the underlying processes.

Section 2:

1. Page 13893: How can the authors discuss the NAO patterns without including the Atlantic Ocean in the modelling domain (only a small part)? Is the MM5 domain different from the one we see in Figs3-5? This should have been explained in the text.

Opposite to the CHIMERE air quality simulation, the MM5 climate simulation uses varying boundary conditions. They are obtain from ERA40 and updated every 6 hours. Hence, although the points/areas where the NAO index is defined are mostly out of the model domain, all the information about the circulation in the Atlantic ocean is given to the model through the domain boundaries every 6 hours. This way, the influence of the NAO on the climatic conditions in and, thereby, within the borders of the MM5 domain is actually captured (even though the MM5 domain does not include most of the Atlantic ocean). In fact, this can be observed in the patterns of Figures 3 and 4. We have now clarified this issue at the end of the 4th paragraph of section 2.1.

2. Page 13894, lines 5-9: The authors have used meteorological fields with coarse spatial resolution of 90km and these meteo fields have been interpolated to 0.2deg for the air quality model simulations. Even though the present work covers the European continent and most part of the Mediterranean Sea, the authors refer to publications on the Iberian Peninsula for the discussion on the skills of the modelling systems. This is quite misleading, as the basic question that arises from this part is how a 90km horizontal resolution can give reasonable results on rainfall patterns for the entire European

continent. In addition, there is no mention in the text if the model setup was exactly the same as in the 2 cited publications. I believe that all the above present a very weak point of the presented work. The authors have not thoroughly evaluated any part of their simulations (meteorological or air quality fields) and they state that this is part of an on-going paper. Yet, the results and conclusions of this paper depend entirely on the performance of the modelling systems, both for the NAO pattern and the atmospheric pollutants concentration.

We understand this complaint as it is certainly true that the two references provided to support the good performance of the climate modeling system focus only on the Iberian Peninsula. Partially because of that, we revise the patterns of the NAO-impact on the climate conditions obtained from the MM5 simulation in section 4.1. Since they are fairly in good agreement with the extensive literature on this regard (e.g. Osborn et al., 1999; Trigo and Palutikof, 2000; Trigo et al., 2002, 2004, 2008), where some of the authors have contributed significantly in the last 15 years by assessing NAO-climate links relative to the European continent, we are confident that our climate simulation captures well the signature of the NAO on the European climate, and as so it should be properly provided to the air-guality modeling system. Moreover, the ability of MM5 (MM5-CHIMERE) with identical or similar setups for reproducing the main meteorological (and air quality) features over different parts of Europe has been demonstrated in many works (e.g. Kotlarski et al., 2005; Monteiro et al., 2007; Flaounas et al., 2009; Renfrew et al., 2009; Péré et al., 2010; Pfeitffer and Zängl, 2010). These references have been now introduced in the revised manuscript. Besides, we do provide a partial validation of the air quality modeling system in Figure 1. This is not a general validation, but it does include the most relevant aspects for the purpose of this study, as argued in the text. A complete validation of the guality modeling system requires an additional long analysis, and corresponds to a separate independent paper that is still being elaborated.

Overall, the above mentioned arguments together with the coherence of the found signals, raises our confidence on the suitability of our modeling system for our experimental purpose.

As required, we now specify in the text that the MM5 model setup used here is the same as in Gomez et al., 2011. It is also used and validated (among others) in Jerez et al., 2013a. Please, see changes in the 3rd paragraph of section 2.1.

3. Page 13894, lines 17-22: What is the exact meaning of this sentence? That the long range transport is disregarded because of the climatological boundary conditions? If this is the case, I have to express my disagreement with this statement. Long-range transport occurs when atmospheric pollutants travel thousands of km away from their sources, i.e from N-Europe to N-Africa, and the domain shown in the figures can include part of such transport mechanisms. Especially, the last sentence has to be considered erroneous "the experimental design allows to better isolate and understand the role of the local processes, including the pollutants transport between different areas within our domain". I would suggest to clarify what is the meaning of "local processes" when simulating atmospheric pollutants in such coarse domain, since pollutant transport and transformation is included by default.

We use constant boundary conditions for all pollutants concentrations (and these can be different in each month). Hence, at the domain boundaries, the model sees every year the same concentrations regardless the NAO phase. Therefore, we are not taking into account the influence of changes in these concentrations due to changes in the NAO phase, i.e. the influence of the large-scale (planetary) transport mechanisms as they are governed by the NAO, as argued above (comment #2 of the introduction). We do model (by default, exactly) transports within the domain boundaries, which certainly involves a wide area where transports along large distances may occur, as pointed out by the reviewer. We have now carefully explained these features in the text, including a more precise definition (by giving an example) of what we call 'local processes'. Please see changes in the 4th paragraph of section 2.1.

4. Page 13895, lines 1-5: How does the model produces wind-blown dust and resuspended dust? This part is missing from the text. Especially since the African deserts are outside the modeling domain, I am not sure of the origin of the wind-blown dust that is shown in the results. A proper discussion should be included in the text.

For a detailed description of how the model produces wind-blown dust and resuspended dust, the reader is referred to Vautard et al. (2005); a paper devoted to the description of these processes. The developers of CHIMERE assume that the available concentration of dust (both wind-blown and resuspended) only depends on the wetness of the surface and the module of the 10-m wind. In this empirical view, the resuspension flux is governed by:

$$F = P f(w) u^{*1.43}$$

where f (w) is a function of the soil water content and P is a constant tuned in order to approximately close the PM10 mass. u* stands for the module of the 10-m wind.

On the other hand, regarding the fact that the modeling domain does not include the African deserts, note that the influence from outer regions is provided through the domain boundaries to the modeling system in the form of boundary conditions. Recall also the arguments to the comment #1 of section 2.

5. Page 13895, lines 6-7: The model-observation comparison has been performed using monthly, weekly, daily or hourly data? This information is not included in the text and it is important to understand how the correlation is above 0.7 most of the times. The authors should also consider giving the bias (not the standard deviation as in Fig.1). If the results are based on monthly PM10 and PM2.5 data, then there is no information on whether the model can capture the variability of the observations, since the lifetime of most aerosol species is a few days up to maximum one week. In that case, it should be made clear in the abstract and introduction that this is a seasonal analysis of the NAO impact. The 20 years simulation gives an overwhelming amount of data to handle, but I believe a more proper comparison should be included in the paper to convince the readers of the validity of the approach. Including the evaluation of the modelling approach in another publication when the model results support the findings in this work, is not appropriate.

We fully agree with the reviewer's suggestion to better clarify the information on the timescale at which the validation was performed, as it was missed in the text (although it was specified in the caption of Figure 1). As the assessment focuses on interannual variations (note that our simulation is 30-year long, not 20-year long!), differentiating between winter (DJFM averages) and summer (JJAS averages), the validation presented here was done at the same time-scale, i.e. using separately DJFM-averaged series and JJAS-averaged series of the PM10 and PM2.5 concentration series. Although this validation considered temporal correlation, standard deviation ratio and bias, we chose the two former statistics for Figure 1 as they are the most relevant for our assessment: we need to ensure above all the reliability of the temporal evolution and the amplitude of our simulated series. Biases are always important, but they are likely canceled in our NAO-impact analysis based on composites, as pointed out in the text. Nonetheless, please note that the order of the biases obtained in the validation is also briefly commented in the text.

Hence, following the suggestions of the reviewer and taking into account the above arguments, we have specified (1) in the abstract, that this is a seasonal analysis aimed at capturing the NAO-impact on the interannual variations of the aerosol concentrations; and (2) in the text, that the validation presented here of the modeling system was thus performed at that time-scale.

Regarding the apprehension raised concerning the appropriateness of the validation provided, please recall the arguments provided above.

6. Section 2.2 is quite small in length and it does not need to be a separate sub-section. This can be part of section 2.1 (which will be named section 2). In this paragraph the authors say that they isolate the influence of climate variability on air quality by keeping the anthropogenic emissions fixed for 2005 during the 30-year run. Please explain the reason for choosing the year 2005 in the text. Since the climate variability is affected to a large extent by the feedback mechanisms between air quality and atmospheric conditions, it seems that this setup does not correspond to the real atmospheric conditions. It is rather a sensitivity model experiment that tests the model response (fixed anthropogenic influence) to NAO patterns. This should be clearly described in the text.

First, regarding the advisability of merging sections 2.1 and 2.2, we should acknowledge our previous own doubts. However, we do prefer to differentiate between the modeling system itself (whose reliability can be validated as it considers year-to-year varying emissions; section 2.1), and the way in which we used it for the purpose of this work (whose reliability can not be validated as it fixes emissions; section 2.2), although it can be certainly summarized in just one paragraph.

We agree with the reviewer that there is not a special reason for choosing the year 2005 emissions over Europe. Emissions are kept constant through the simulations in order to clarify all the analysis related to NAO patterns, as stated by the reviewer, and then the election of the year for emissions does not condition the simulations. The final reason for selecting 2005 emissions is that previous simulations by the group having 2005 emissions as reference (e.g. Jiménez-Guerrero et al., 2001; Jiménez-Guerrero et al., 2012; Jiménez-Guerrero et al., 2013) showed a good performance. Those results indicate a reliable behavior of the model for reproducing air quality climatologies over Europe, and therefore this year was also selected in this work.

Section 3:

Page13896, lines 23-25: The authors are using the NAO index provided by CPC (NOAA) and the use ERA40 or ECMWF analysis fields for the MM5 simulations. How well the 2 systems relate when it comes to the calculation of the NAO index? If the two datasets give very different indices then the results from this work cannot be justified as they compare NAO patterns that do not relate to the air quality simulations. This is a very important part of this work and should be handled with caution in section 3.

We acknowledge the convenience of this warning, which was also raised by the other reviewer. Indeed, the NAO index is not only affected by the use of different data sources, but also of different methods for its computing (i.e. station or PC based, location of the stations and fields from which the PCs are computed). However, at the seasonal time-scale analyzed here, these differences are fairly negligible. In particular, the dependence of the NAO on the reanalysis used for its computing is minimal, as these datasets are very similar at the monthly/seasonal timescales. This is clearly visible in the work by Greatbatch and Rong (2006). They show the summer NAO index computed from both the NCEP and the ERA40 reanalysis through the same methodology (PCA applied to the SLP field) (Figure R1). It can be observed that both series greatly agree in the overlapping period.

Since we did not find a similar plot for the winter NAO time series in the literature, we provide in Figure R2 the time series of different winter (DJFM-averaged) NAO indexes. Despite the apparent differences among them (due to the different methodologies and data sources employed for computing the NAO index in each case; we believe that the former cause plays the most important role), the correlations between all the series pairs is above 0.85 (Table R1). Besides, the years that would have been selected in each case as winter NAO⁺ or NAO⁻ years following our methodology mostly coincide (Table R2), which is even more important from a practical point of view, i.e. in order to support the validity of our approach/results.

Therefore, our results, at the time-scale analyzed here, are not really affected by the specific NAO index used. Hence, we have maintained the use of the NAO index provided by the CPC of the NOAA as it is the most commonly used nowadays.

Nonetheless, as the former conclusion may not work at other time-scales, we have included a comment in the manuscript reflecting the caveat raised by both reviewers in this regard.



Figure R1. Adapted from Greatbatch and Ping-ping Rong (2006). Principal component (PC) time series for the leading EOF computed from SLP averaged over July and August in each year (i.e. summer NAO index) using the NCEP-NCAR (dashed line) and the ERA-40 (solid line) reanalysis.



Figure R2. Winter (DJFM-averaged) NAO index time series obtained using different methods and datasets. In black: NAO index from the CPC of the NOAA (this is the NAO index series used in the manuscript). In green: PC-based Hurrell NAO index(*). In blue: station-based Hurrell NAO index(**). In red: as the station-based Hurrell NAO index but computed using ERA40 data. Note that the different methodologies seem also to imply differences in the pre and post-processing of the data. In particular, the black series looks to be normalized and varying between -1 and 1 while the others are not.

(*) http://climatedataguide.ucar.edu/sites/default/files/cas_data_files/asphilli/nao_pc_djfm_1.txt (**) http://climatedataguide.ucar.edu/sites/default/files/cas_data_files/asphilli/nao_station_djfm_1.txt

	NCEP-PC	HURRELL-PC	HURRELL-ST	ERA40-ST
NCEP-PC				
HURRELL-PC	0.86			
HURRELL-ST	0.88	0.93		
ERA40-ST	0.86	0.90	0.98	

Table R1. Temporal correlationbetween the different DJFM-averaged NAO index timeseries depicted in Figure R2.

Year	70	71	72	73	74	75	76	77	78	79	80	81	82	83	84	85	86	87	88	89	90	91	92	93	94	95	96	97	98	99
NCEP-PC																														
HURRELL-PC																														
HURRELL-ST																														
ERA40-ST																														
NAO > 70th pctl		NAC) < 30t	h pctl																										

Table R2. Years with winter (DJFM-averaged) NAO index above the 70th percentile (in pink) and below the 30th percentile (in blue) in the series depicted in Figure R2).

Section 4:

1. Page 13898, line 1: The phrase "essentially to evaluate the ability of our climate simulation" is not supported anywhere in this section. How is this evaluation performed? There is no comparison with observed or measured values and the above statement is not appropriate. There is no evidence that the differences shown in Fig.3 are representative of the actual atmospheric conditions. This is also where one of the main questions arise again: the NAO phases are calculated with the NOAA index but the atmospheric simulation is driven by ECMWF data. Are these comparable?

The NAO-impact on the European climatic conditions is well known and established elsewhere, as it is also proved the capacity of climate models to reproduce it. Hence, in section 4.1 we just want to certify that our simulation does reproduce the well-known NAOimpact patterns on the European climate, so that we can trust on it as provider of these climatic differences between NAO phases to CHIMERE. And, in fact, the patterns of Figures 3 and 4 agree with those available in the literature, as pointed out in the text. Hence, this proves "the ability of our climate simulation (that is driving the CHIMERE run) to simulate the well-known NAO-impact on the European atmospheric conditions". The alluded sentence is actually preceded and followed by these clarifications. Hopefully they help to avoid misunderstandings.

Please, recall also the arguments provided on comment #2 of section 2 and on the comment on section 3 related to the negligible influence of the choice of the NAO index in the framework of this study.

2. Page 13900, lines5-7: How is the enhanced DUST concentration in the Iberian peninsula (Fig.4c) related to the precipitation in the same area (Fig. 3e)? Please be more descriptive on the analysis of the results in cases like this one.

In a nutshell we explain the general 'rule' in the first paragraph of section 4.2 (before sections 4.2.1 and 4.2.2). In particular, it is said there:

'... patterns can be well explained by the diluting effect of stronger winds and scavenging processes by an enhanced precipitation and by the inhibitory effect of enhanced cloudiness as it implies reduced shortwave solar radiation (a main precursor for photochemistry and biogenic emitting activity).'

We try to avoid repeating this arguing for the sake of space. Besides, these general arguments are later recalled in the conclusions section.

3. Page 13901, line16: The SOA levels are shown in Fig.5f not 5d.

This has been updated with the new figures numbering.

Conclusions:

The results from this work, as discussed in this section, state that the aerosol concentrations are influenced by the changes in precipitation, temperature and wind fields. This is a result already known from the physics and chemistry of the underlying processes, without the need to perform a 30-year model simulation. I suggest that the authors focus their conclusions on the new findings of their work that are associated with the NAO phases and impacts on the European continent and the differences therein.

The 3rd paragraph of the conclusions section (the one alluded by the reviewer's comment) is devoted to explain the causes for the NAO-related variations in the levels of aerosols (which are summarized in the 2nd paragraph). These are primarily related to the NAO-impact on the precipitation, temperature and wind fields. As the influence of these fields on the aerosol concentrations is certainly well-known, we just recall it here in the framework of the links between the NAO-signals obtained for the climate conditions with the NAO-signals obtained for the aerosol concentrations. We agree, however, that this paragraph needed to be improved as it often lacked to mention the role of the NAO explicitly. This could certainly provoke misunderstandings, since the objective of this paragraph is not to elucidate the impact of climate on aerosol concentrations. Hence, we have emphasized it better in the new version of the manuscript, but still conserving the explanations provided for the link between climate (as it is controlled by the NAO) and aerosol concentrations.

On the other hand, the 30-year long simulations allowed us to assess variations at the interannual time-scale (i.e. changes in the seasonal means between years with opposite NAO phases) with statistical robustness, as pointed out at the end of section 2.2, instead of focusing on a small number of events. We thus think that the length of the simulation is really valuable.

Figures 3 to 5:

The quality of these figures that present the main findings of this work is not acceptable. It is very difficult to see the details in each plot and see how the text is supported by the figures. The authors should leave 4-6 panels (maximum) in each figure and make sure that the details are easily discernible.

Following this requirement, the former Figure 3 is now split into the current Figures 3 and 4, and the former Figures 4 and 5 are now split into the current Figures 5, 6, 7 and 8.

Technical corrections:

1. Please replace "non-antropogenic: with "non-anthropogenic" everywhere in the text.

Replaced.

2. Abstract, line 13: replace the word "rebounds" with "influences" or "affects".

Replaced.

3. Abstract, line 16: please rephrase the part that reads "of this later" as the meaning is not clear. What is this later?

The sentence has been removed.

4. Introduction, p13891, line 7: please replace the word "paramount" with a more modest one.

Changed.

5. Section 2.1, page 13893: The sentence "This resolution enhances from previous works. . .." must be rephrased as the verb "enhance" is not appropriate.

Amended.

6. Page13895, line26: Please rephrase the "it arises mandatory" with "it becomes mandatory".

Done.

REFERENCES (not included in the manuscript):

Vautard, R., Honoré, C., Beekmann, M., and Rouil, L., 2005: Simulation of ozone during the august 2003 heat wave and emission control scenarios. *Atmos. Environ.*, 39(16):2957-2967).

1	Impact of the North Atlantic Oscillation on the variations of
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20 Date: May September 2013

21 Abstract

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This contribution assesses non-anthropogenic variations in ground-level aerosol 23 concentrations over Europe associated to changes in the phase of the North-24 25 Atlantic Oscillation (NAO). The North Atlantic Oscillation (NAO) controls a large amount of the European climate variability with asymmetric impacts in both time and 26 27 space. Here we investigate how the local atmospheric processes (without 28 considering large inter-continental transport mechanisms), as they are governed by 29 the NAO, affect the levels of various aerosol species using simulated data under 30 constant emissions, which are fixed to the 2005 levels in order to avoid anthropogenic-induced signals. In particular, we analyze interannual variations at 31 32 the seasonal time-scale and focus on the ground-level. The results show Based on simulated data and focusing on how the local atmospheric processes (without 33 34 considering large-scale mechanisms) governed by the NAO affect the levels of 35 various aerosol species, this study highlights that positive NAO phases favor 36 increased aerosols concentrations levels in southern (northern) regions in winter (summer), while negative NAO phases enhance them in northern (southern) regions in 37 38 winter (summer). The underlying processes are Variations are up to and over 100% for most aerosols, being clearly related to the NAO -impact on precipitation and wind, as 39 40 they act to clean the atmosphere through removal and dispersion processes, and to the NAO -impact on the radiation balance (i.e. cloudiness) as it rebounds on affects the 41 42 biogenic emitting activity and on the oxidative capacity of the atmosphere. Differences for 43 all the species studied (natural inert, secondary inorganic and organic aerosols) are up to 5 μ g m⁻³, reaching 10 and 20 μ g m⁻³ for PM10 and PM2.5 respectively, which 44 represents variations about 20-40% in their mean levels between opposite NAO 45 phases. Beyond deepening on the understanding of fundamental interactions 46 47 between climate and air guality, these results provide a basis for improving the 48 potential predictability of this later since much work is being done in order to gain 49 accuracy in the NAO predictions.

50 1 Introduction

51

52 Air pollution is a major environmental and health problem affecting industrialized and developing countries around the world. Main detrimental consequences of the exposure of 53 humans to photochemical and particulate matter pollution include respiratory difficulties, 54 especially for sensitive people (Pope et al., 2006). The results of the APHEIS study 55 56 (Ballester et al., 2008) indicate that atmospheric pollution causes the premature death of over 310.000 dwellers in the European Union each year. The most serious problems in 57 58 Europe are related with particulate matter with a diameter of less than 10 micrometers (PM10), nitrogen dioxide and ozone. In this sense, worldwide epidemiological studies 59 show a consistent increase in cardiac and respiratory morbidity and mortality from 60 exposure to air pollution (e.g. Pope et al., 2009). Besides, ecosystems are also affected, 61 62 with losses of agricultural crops and damages in aguatic and terrestrial ecosystems having been reported (Van Dingenen et al., 2009). In order to anticipate when and why episodes 63 64 of air pollution arise and how they can be abated, reliable estimations of air pollution levels and a better understanding of the chemico-physical processes behind them are crucial of 65 66 paramount importance.

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68 Air pollution levels depend on both emissions (either natural or anthropogenic) and the atmospheric conditions steering and transforming them through processes related to 69 70 chemistry, transport and removal (Jacob and Winner, 2009). Precipitation provokes wet 71 deposition, hence favoring the removal of airborne pollutants, and wind favors their 72 transport and dispersion, although enhanced winds also promote the formation of marine 73 aerosols over the water mass areas. On the other hand, radiation levels and temperature 74 play a major role on gas-phase chemistry through the photolysis of primary and secondary pollutants (Katragkou et al., 2010; Forkel et al., 2012; Jiménez-Guerrero et al., 2012; 75 Meier et al., 2012; Colette et al., 2013; Hedegaard et al., 2013, just to mention some 76 77 recent works on the topic), determining also the vegetation activity and thereby 78 modifying natural emissions.

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Despite the great climate heterogeneity and variability that characterizes Europe, just a few large-scale teleconnection modes control a large amount of it (Trigo et al., 2008). In particular, many studies establish the fundamental role of the North Atlantic Oscillation (NAO) at this regard, affecting especially western Europe with an asymmetric impact

between northern and southern areas (Osborn et al., 1999; Wanner et al., 2001; Trigo et 84 85 al., 2002, 2008). The NAO pattern consists of a meridional gradient in Sea Level Pressure 86 (SLP) over the North Atlantic with centers roughly around the Azores Islands (high pressures) and near Iceland (low pressures). During its positive phases (NAO⁺), such a 87 88 dipole in SLP is enhanced with the consequent intensification of westerly winds in northern Europe associated with an intensified jet stream. This configuration promotes also the 89 90 steering of most storms over northern regions, advecting humid air from the ocean, hence 91 favoring cloudiness and precipitation. Contrary and consistently, NAO⁺ reduces cloudiness 92 and precipitation in southern Europe. Conversely, the intense zonal circulation promoted by the negative NAO phases (NAO⁻) enhances the westerly flow penetrating in southern 93 94 Europe through the Iberian Peninsula (Jerez et al., 2013b) and extending until eastern 95 Europe and allows cyclones to follow more southern paths within the European continent 96 (Trigo, 2006). Most studies have focused their analysis on NAO-climate links during the 97 winter season, where the SLP gradient is stronger and the climatic impact of this large-98 scale mode of variability is extensive extends to Eurasia (Wanner et al., 2001; Trigo et al., 2002) but also North America, Greenland, the Arctic and North Atlantic oceans (Hurrell 99 100 and van Loon, 1997). Readers looking for comprehensive reviews of the dynamics of the 101 NAO and associated impacts should consult the books by Hurrell et al. (2003); Vicente-102 Serrano and Trigo (2011). However, recent works have shown that the summer NAO still 103 plays an important role on shaping the climate of northern and central Europe (albeit less 104 relevant than in winter) including parts of the Mediterranean basin (Folland et al., 2009; 105 Bladé et al., 2012).

106

107 These NAO-related impacts on the atmospheric fields are bound to inflict an important impact on air pollution levels (Dayan et al., 2008; Sanchez-Lorenzo et al., 2008; Chiacchio 108 109 and Wild, 2010), particularly on aerosols concentrations. This influence should arise 110 through both local processes and large-scale air pollutants transport. For instance, 111 Eckhardt et al. (2003) showed that NAO⁺ enhances northward transport of pollutants 112 from Europe to Arctic regions, mainly in winter and spring, compared to NAO⁻ 113 phases. At the same time, NAO⁺ promotes African dust intrusions into southwestern European regions (which occurs mostly in the summertime) as the westerly 114 115 winds associated to NAO⁻ events prevent subtropical air masses to reach the 116 European mid latitudes; although, on the other hand, this NAO⁻ related large-scale 117 zonal winds also favor the transport of pollutant from North America into Europe

(Moulin et al., 1997; Dayan et al., 2011; Christoudias et al., 2012; Cusack et al., 2012; 118 Pey et al., 2013). But the signature of the NAO in the climatic conditions (and 119 120 thereby on air pollution levels) does not restrict to changes in the large-scale circulation patterns, as these have also repercussions on other atmospheric 121 122 variables such as precipitation and temperature as commented above, with a potential subsequent impact on air quality through local processes such as aerosol 123 wet deposition. However, few studies have been devoted to explore air pollution 124 levels from a climatic perspective, even less disentangling between local and large-125 scale mechanisms, therefore the contribution of the local NAO-controlled processes 126 on the climatology of air pollution levels is still hardly established., either from 127 Europe to Artic regions (Eckhardt et al., 2003) or from North America and Africa-128 towards Europe (Moulin et al., 1997; Davan et al., 2011; Christoudias et al., 2012). 129 130 However, since few studies are devoted to explore air pollution levels from a 131 climatic perspective, the impact of the NAO in such a framework is still hardly 132 established.

133

Hence, the objective of the present study is to **deepen on elucidate** the signature of the 134 NAO in terms of mean concentration of aerosols in a region covering the entire 135 136 Mediterranean basin from north Africa to north Europe with the focus on elucidating the influence of the small scale processes, as well as the associated underlying 137 138 mechanisms. To achieve this goal, we use a numerical simulation of the atmospheric chemical composition that spans three decades of the recent past without considering 139 variations in the anthropogenic emissions, thus allowing to isolate the natural variations in 140 141 the aerosol levels. Besides, tThe simulation was designed to disregard the contribution from the long-range large scale transport by using constant climatological boundary 142 143 conditions for the aerosol concentrations. This design does not provide a realistic picture, but allows to improve our understanding on the role of the local underlying 144 145 mechanisms, as it remains unmasked by the large-scale inter-continental advective phenomena. Hence, we focus on the role played by local processes as they are governed 146 147 by the NAO, restricting the evaluation to the ground-level.

148

The structure of this work is as follows. Section 2 describes the modeling system and the
experimental set-up. Section 3 provides the methodology. Section 4 presents the results.
Finally, Section 5 summarizes and discusses the main conclusions.

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- 153
- 154 **2 Data**
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- 156 **2.1 Air-quality modeling system**
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158 The modeling system consists of a climatic version of the Fifth-Generation Pennsylvania 159 State University - National Center for Atmospheric Research Mesoscale Model (MM5) 160 (Grell et al., 1994) driven by ERA40 reanalysis (Uppala et al., 2005), when available, or ECMWF analysis data when not (in both cases without nudging), coupled off-line to the 161 CHIMERE chemistry transport model (Bessagnet et al., 2004; Rouil et al., 2009). 162 MELCHIOR2 gas-phase mechanism is implemented within CHIMERE (Derognat et al., 163 164 2003). The chemistry transport model includes aerosol and heterogeneous chemistry, distinguishing among different chemical aerosol components, namely nitrate, sulfate, 165 ammonium, elemental and organic carbon with three subcomponents (primary, secondary 166 anthropogenic and secondary biogenic) and marine aerosols. Unspecified primary 167 anthropogenic aerosols and aerosol water are additionally kept as separate components. 168 The model considers the thermodynamic equilibrium using the ISORROPIA model (Nenes 169 et al., 1998). Last, the aerosol microphysical description for CHIMERE is based on a 170 171 sectional aerosol module including 6 bins from 10 nm to 40 µm using a geometrical 172 progression.

173

The CHIMERE domain considered in the modeling system covers all the Mediterranean basin extending to northern Europe having a spatial resolution of 0.2 degrees in the horizontal, which is around 25 km at the European latitudes considered, and eight vertical levels unevenly spaced up to 550 hPa. This resolution **is higher than usual in climate runs enhances from previous works (e.g.** Carvalho et al., 2010; Katragkou et al., 2010; Jiménez-Guerrero et al., 2011; Juda-Rezler et al., 2012; Manders et al., 2012).

180

The fields from MM5 (provided with a spatial resolution of 90 km) are bilinearly interpolated to the CHIMERE working grid. Detailed descriptions of the climate modeling system used and its skill to realistically reproduce the main regional climatic features of the climate in the target domain, such as the temperature and precipitation annual cycles or the

interaction between the large-scale circulation and the orography, which largely modulates 185 the rainfall patterns having an important influence on air guality, can be found in Gomez-186 187 Navarro et al. (2011) and Jerez et al. (2013a). The model setup used here is the same as in the former reference; in the second, it is evaluated in the context of a multi-188 189 physics ensemble of present-day climate simulations. Although both works focus on the Iberian Peninsula, there are extensive literature demonstrating the ability of 190 MM5 for reproducing diverse meteorological features along Europe (e.g. Kotlarski et 191 al., 2005; Renfrewet al., 2009; Pfeiffer and Zängl, 2010). Moreover, Section 4.1 192 193 provides a revision of the accuracy of the patterns of the NAO impact on climate as they are obtained from the climate simulation driven CHIMERE. 194

195

Boundary conditions of gas-phase pollutants concentration for CHIMERE the 196 197 chemistry transport model are based on the global chemistry model LMDz-INCA2 198 (Szopa et al., 2009) developed by the Laboratoire des Sciences du Climat et 199 l'Environnement (LSCE). A detailed description of the Interactive Chemistry and Aerosol (INCA) model is presented in Hauglustaine et al. (2004) and Folberth et al. (2006). For 200 201 aerosols, boundary conditions are taken from the GOCART model (Chin et al., 2002). These boundary conditions consist of constant mMonthly mean data that are 202 203 interpolated in the horizontal and vertical dimensions to force the major chemical concentrations at the boundaries of the CHIMERE domain. Although the influence of 204 205 using The use of constant climatological boundary conditions prevents on ground level concentrations is largely overwhelmed by local processes (Jiménez-Guerrero et al., 206 2012), it should be acknowledged that, in particular, this hampers to capture the 207 208 interannual variability of the NAO -impact on the aerosol concentration levels at the domain boundaries, thus avoiding the evaluation of large-scale inter-continental transport 209 210 mechanisms related to the NAO phase. On the contrary, and despite the pollutants transport between different areas within our extensive domain is modeled by 211 212 default, this experimental design allows to largely isolate the role of more local processes (for instance, how the NAO influences the concentration of aerosols 213 214 through its impact on local precipitation patterns). On the contrary, this experimental design allows to better isolate and understand the role of the local 215 216 processes, including the pollutants transport between different areas within ourdomain. At this regard, it should be noted that the design of the driving MM5 climate 217 218 simulation does not follow this approach. The climate boundary conditions at the

219 MM5 domain boundaries are updated every 6 hours according to the ECMWF data 220 specified above. Hence, although the MM5 domain does not cover, in particular, 221 most of the Atlantic ocean, the influence of the NAO on the climate conditions 222 simulated within its borders should be actually captured.

223

Year-to-year varying anthropogenic emissions are derived from the EMEP database on a 224 monthly basis (Vestreng et al., 2009). Natural emissions depend on climate conditions, 225 and consequently they are modeled according to the MM5 meteorological outputs. 226 227 However, the levels of air pollutants are estimated without considering possible changes 228 on vegetation, land use or any feedback from the chemical compounds to the 229 meteorological fields. Biogenic emissions were generated dynamically using MEGAN (Model of Emissions of Gases and Aerosols from Nature) (Guenther et al., 2006) with the 230 231 parametrization form of the canopy environment model. The model estimates hourly 232 isoprene, monoterpene, and other BVOC emissions based on plant functional type and as 233 a function of temperature and ground level shortwave radiation.

234

Beyond the several works supporting the ability of similar MM5-CHIMERE systems 235 to reproduce the main air quality features over Europe (e.g. Monteiro et al., 2007; 236 237 Flaounas et al., 2009; Péré et al., 2010), our This modeling system has been validated 238 by comparing a simulation covering the period 1990-2010 with EMEP observations 239 (Tørseth et al., 2012). A Although a thorough evaluation of the modeling approach is not included in the main objectives of this work, and that task constitutes the focus of an 240 on-going paper. In any case, we acknowledge that it is opportune to here we show 241 242 shortly that the simulated series of aerosols concentrations correlates acceptably and captures a large fraction of the interannual variability of the observational series at the 243 244 working time-scale (i.e. seasonal). These two aspects are the most relevant for the 245 assessment performed below, while systematic biases would not represent a major 246 concern as they should be largely canceled when computing the composites of the NAO 247 -impact (see next Section). For the sake of brevity, we provide here the results for the simulated concentration of aerosols under several two size-thresholds, namely PM10 and 248 PM2.5 (Figure 1). For both cases, correlations between simulated and observational 249 250 seasonal-averaged series are overwhelmingly above 0.8, in most cases above 0.9, both 251 in winter and summer. As well Also, it is possible to verify that biases in the standard 252 deviation of the simulated series are largely negligible in comparison to their mean values

253 (being orders of magnitude smaller). It is also worth mentioning that normalized biases are 254 always found to stay below 30% the worst (not shown), which is in the top range expected 255 (Pay et al., 2010). This These results guarantees the phase accordance (timing) 256 between the simulated and observational series, their similar amplitude and, also, the 257 quantitative accuracy of the simulated climatologies, hence making us confident on the 258 accuracy suitability of the modeling system for the purpose of this study.

259

260 2.2 Experimental set-up

261

262 In order to isolate the influence of climate variability on air quality, it arises becomes mandatory to avoid the signals derived from human policies. During the last decades, 263 264 these policies have committed European countries to strongly reduce the emissions of contaminants to the atmosphere, which has caused strong non-natural trends in the 265 266 observational series of almost all aerosols species (Vestreng et al., 2009; Tørseth et al., 267 2012). Hence, the impact of the NAO has been evaluated here based on a 30-year long MM5-CHIMERE simulation for which emissions were fixed at their 2005 mean levels. The 268 simulated period (1970-1999) is not particularly significant in itself, but its length is-269 270 relevant as it supports the robustness of our results from a climatic perspective.

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- 272

273 3 Methodology

274

This assessment focuses on several aerosol families and species, namely natural inert aerosols (sea salt aerosols, SALT, and wind-blown and resuspended dust, DUST), secondary inorganic aerosols (sulfate, $SO_4^{2^\circ}$, nitrate, NO_3^{-} , and ammonium, NH_4^{+}), organic matter (OM), with particular attention to secondary organic aerosols (SOA), and elemental carbon (EC). The total concentrations of PM10 and PM2.5 are also studied.

280

The analysis is performed at the seasonal timescale for winter (December-to-March averages) and summer (June-to-September averages). This decision is based on previous tests performed at the monthly timescale showing the strongest responses in those months, with very similar structures within in each season considered but very different in each one of them. The remaining months from the two transitional seasons depicted intermediate and softer signals and thus are not included in being disregarded from

the rest of the analysis.

288

Seasonal means (DJFM and JJAS averages) of the monthly NAO index provided by 289 the Climate Prediction Center (CPC) from the National Ocean and Atmospheric 290 291 Administration (NOAA) are used to define classes of positive and negative NAO 292 phases (NAO⁺ and NAO⁻ respectively) throughout the study period. Following a a 293 common approach adopted in previous studies (e.g. Suski and Ridgway, 2007; 294 Gouveia and Trigo, 2008), NAO⁺ (NAO⁻) phases are defined as those with the NAO 295 index above (below) the 70th (30th) percentile of the NAO index time series considered (Figure 2). The use of these non-subjective thresholds in the definition 296 297 of NAO⁺ and NAO⁻ phases ensures the evenness between the number of events of each class (within the 30-year simulated period, 1970-1999, we retain 9 events of 298 299 each class in each season) and prevents the unbalancing influence of potential 300 long-term trends in the NAO index time series. On the other hand, it should be 301 noticed that different data sources and/or methods for computing the NAO index may provided different values of it. In particular, we are aware that the CPC/NOAA 302 303 NAO index is not based on the ECMWF data used to drive MM5 in our modeling system. However, we confirmed that, at the time-scale assessed here, this influence 304 is negligible (Greatbatch and Rong, 2006). Hence, the choice of the CPC/NOAA NAO 305 index for the present study is based on its easy access and its up-to-date 306 307 computation approach, which in fact makes it widely used nowadays in studies focused on the recent past (e.g. Bladé et al., 2012; Jerez et al., 2013b; Pey et al., 308 309 2013).

310

In order to assess the NAO -impact, we have followed a common approach adopted 311 in previous studies (e.g. Suski and Ridgway, 2007; Gouveia and Trigo, 2008) and 312 defined classes of positive (NAO⁺ \geq 70th percentile) and negative (NAO⁻ \leq 30th 313 percentile) NAO phases using the seasonal means of the monthly NAO index-314 provided by the Climate Prediction Center (CPC) from the National Ocean and 315 Atmospheric Administration (NOAA). This procedure ensures the balance between 316 317 the number of events of each class that will be considered. Within the 30-year-318 simulated period (1970-1999), we obtain 9 winters and summers of each NAO class 319 (Figure 2). The NAO -impact on the assessed magnitudes is then evaluated through 320 composites showing the differences in the mean fields between positive and negative NAO

321 phases. Several significance controls are applied to these differences ensuring both their 322 statistical robustness and physical consistence, so that they are just considered when 323 fitting the following criteria: (1) being statistically significant above the 90% confidence 324 level, and (2) being supported by statistically significant temporal correlations (above the 325 90% confidence level) between the NAO and the corresponding chemical component series. Statistical significance is evaluated by performing two-tailed t-tests for the null 326 hypothesis of equal means or zero correlation respectively (Snedecor and Cochran, 1989). 327 Moreover, the signal-to-noise ratio, defined as the ratio between the absolute value of the 328 NAO⁺ minus NAO⁻ differences in a specific magnitude and the standard deviation of the 329 330 whole seasonal series of such a magnitude, is considered as a relative measure of the importance of the NAO -impact. In particular we have blurred the areas where the signal-331 332 to-noise ratio is below the unit. In the areas not blurred, the NAO -impact exceeds one standard deviation of the series and hence we are likely moving from one tail to the other 333 334 of the probability density function describing the variability in the levels of the assessed 335 magnitude when the NAO phase changes.

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- 337

338 **4 Results**

339

4.1 Revisiting the NAO -impact on atmospheric conditions

341

Several works have been devoted to explore the role and signature of the NAO in the 342 343 European climate (e.g. Hurrell, 1995; Hurrell and van Loon., 1997; Hurrell et al., 2003; Trigo et al., 2002, 2008). Indeed, the reported NAO -impacts on the climatic variables 344 motivated the present study. Other works have shown a good capacity of the general 345 circulation models to reproduce the large-scale patterns of this climatic impact over Europe 346 particularly for winter (Osborn et al., 1999; Osborn, 2011; Hurrell et al., 2003) but also for 347 348 summer (Bladé et al., 2012). Hence, this Section does not intend to provide novel insights, 349 but essentially to evaluate the ability of our climate simulation (that is driving the CHIMERE 350 run) to simulate the extensively reported NAO -impact on the European atmospheric conditions. This fulfills two relevant requirements, namely to (1) guarantee that it is 351 352 effectively able to reproduce the known NAO -impacts on the European climate, and (2) 353 provide an appropriate context for the interpretation of the following results of the NAO 354 -impact on aerosols concentrations, since the analysis is specifically focused on those

355 variables driving air pollution levels.

356

357 Figure 3 further confirms the expected NAO -impact on the wind field, showing asymmetric 358 responses in winter and summer. In winter, NAO⁺ (NAO⁻) promotes a windy westerly flow 359 in northern (southern) Europe (Figures 3a,c), with the largest significant differences in the wind speed (above 10%) appearing northward in the western sector of the domain (Figure 360 3c). Conversely, in summer, NAO⁻ phases have are associated with a stronger westerly 361 flow in northern Europe than NAO⁺ phases (Figure 3b,d), with differences in wind module 362 363 exceeding 5% and extending more eastward than in winter. Windier conditions associated 364 to the NAO⁻ phase prevail also in the south-west of the domain in summer, mainly in the 365 surroundings of the Iberian Peninsula (differences about 5%) (Figure 3d).

366

367 Consequences in precipitation, cloudiness and temperature derived from the former 368 conditions are also provided in Figure 34, as these variables play a key role on the levels 369 of air pollutants in general, and aerosols in particular. Obtained patterns are in good agreement with previous works for winter (e.g. Osborn, 2011) or summer (Bladé et al., 370 2012) relating them to the westerly winds advection of humid air from the Atlantic, which 371 favors the formation of clouds and enhances precipitation. NAO⁺ minus NAO⁻ differences 372 in precipitation (Figure 3e,f 4a,b) are larger in winter than in summer, when they are 373 374 overall negative and affect mainly southern Europe ranging west-to-east from 40% (50 375 mm/month) to 20% (20 mm/month). In summer, these differences reach 20% (up to 30 mm/month) over large sectors of Europe, being negative in northern Europe while positive. 376 and less important, in the southern affected areas. Consistently, the composites for the 377 378 column integrated cloud water (a variable representative of cloudiness) (Figure 3g,h 4c,d) show negative values in southern Europe in winter (resembling the west-to-east gradient 379 380 of the precipitation signal, with differences ranging from 30 to 10%) and northern areas in 381 summer (differences around 10%), and positive values in north-western areas in winter 382 (i.e. NAO⁺ enhancing cloudiness there, around 20%). Regarding the NAO -impact on 383 mean 2-meter temperature, the most important signal consists of positive NAO⁺ minus 384 NAO⁻ differences (up to 3 K) largely spread over northern Europe in winter (Figure **3**; **4**e). Smaller negative (positive) differences (up just to 1 K) appear also over some 385 southernmost (northern) areas in winter (summer) (Figure 3i, j 4e, f). 386

387

388 These results support the ability of the climate simulation to reproduce the expected

389 responses to the NAO phase, providing a meaningful representation of the atmospheric 390 conditions governing during positive and negative NAO events that will be used to 391 understand the NAO -impact on aerosol concentration levels assessed in the next Section. 392

393 4.2 NAO -impact on mean ground-level aerosols concentrations

394

NAO⁺ minus NAO⁻ composites of mean ground-level concentration of the various aerosol 395 species are provided in Figures 4 and 5 to 8. In general, the positive phase of NAO 396 397 enhances aerosol concentrations in southern Europe in winter and in northern Europe in summer. These positive NAO⁺ minus NAO⁻ differences match generally well the sectors 398 399 characterized by negative signals in wind speed, precipitation and cloudiness (Figures 3 400 and 4). Thus such patterns can be well explained by the diluting effect of stronger winds 401 and scavenging processes by an enhanced precipitation and by the inhibitory effect of 402 enhanced cloudiness as it implies reduced shortwave solar radiation (a main precursor for 403 photochemistry and biogenic emitting activity).

404

405 4.2.1 Winter signals

406

407 As a general rule, in winter, the regions holding the largest signals are predominantly 408 located in the Mediterranean basin, particularly over the Iberian Peninsula, northern Italy 409 and the Balkans (Figures 4 and 5 to 8, left column). There, results show increases in the mean concentration of all aerosols species during positive NAO phases exceeding the 410 50% of the mean levels during negative NAO phases, even doubling these latter in some 411 412 occasions. In absolute values (although these should be taken with care due to our experimental design not accounting for real emission levels neither for long-range 413 414 transports), the largest variations between NAO phases reach 5 µg m⁻³ and affect the inorganic compounds (DUST, SO_4^{2-} , NH_4^{-} and NO_3^{-}), while the change in the concentration 415 416 of carbonaceous material is just about 1 µg m⁻³. It is worth stressing the slight but still 417 significant negative patches of NAO⁺ minus NAO⁻ differences in the British Islands and in north-eastern Europe that appear in the patterns of OM and EC, indicating a distinct 418 response of this compounds to the same NAO phase depending on the latitude. 419

420

421 The NAO -impact on NO₃⁻ concentration levels is negligible over most of Europe. However, 422 the removal effect of the NAO⁻ phase associated with enhanced precipitation in northern Italy during NAO⁻ phases (not counteracting by enhanced temperatures) seems to
play a key cleaning role at this regard in this NO₃⁻-strong emitting area.

425

426 Contrary to the rest of aerosol species, enhanced winds promote the formation of SALT 427 (sea salt aerosol emissions depends on the cube of the wind speed), thus leading to higher concentrations. This is mainly observed over the water mass areas, with NAO⁺ 428 leading to 40% higher SALT concentrations in the surroundings of the British Islands. 429 which matches with the NAO⁺ enhanced winds observed in this area (Figure 3c), but 430 diminishing them in the southern Mediterranean (Figure 45a). Singular positive signals 431 appear also in the Gulf of Genoa (north-westward of Italy) and the strait of Gibraltar 432 (southward of Spain). Regarding these three latter cases, it can be roughly appreciated in 433 434 Figure 3a that NAO⁺ promotes windier conditions over the mentioned areas (longer arrows) than NAO⁻, although these differences in wind speed do not remain after applying 435 436 the significativity controls as they are not reflected in Figure 3c.

437

The NAO -impact on winter air quality is appreciable in the various size-classes of particulate matter. Figure **5g,h 8a,c** shows that NAO⁺ enhances 20-40% the concentrations of both PM10 and PM2.5 in the southern European regions (where the PM climatologies present the higher values and the number of exceedances of the limit values for the protection to human health are more frequent), and also, but to a smaller extent, in some northernmost areas. Therefore, the winter NAO phase has a clear impact on the air quality-related human healthy risk in this season.

445

446 **4.2.2 Summer signals**

447

In summer, the highest signals within our domain appear in the British Islands, northern
France, Belgium, Netherlands, northern Germany and northern Poland (Figures 4 and 5 to
8, right column). It is also interesting to see the recurrent positive signals appearing over
Italy also in this season, although it should be acknowledged the patchy nature of them.

452

As in winter, differences in the aerosol concentrations between NAO phases are up to 100%, demonstrating the profound impact of the NAO in this area/season. Likewise, the obtained NAO⁺ related increments are, in absolute values, about (even over) 5 μ g m⁻³ for each inorganic species but SALT (which shows smaller signals, Figure **45**b). The NAO⁺-

related increases in the concentration of the organic and carbonaceous aerosols are also a bit smaller (when they are expressed in μ g m⁻³).

459

460 It is worth stressing the signal in mean SOA levels (Figure 57d), which can be doubly 461 related to the reducing effect of NAO⁺ in precipitation (Figure 3f 4b), as it prevents wet 462 deposition, and to the increased temperatures during NAO⁺ enhancing the biogenic 463 emissions (such as isoprene and monoterpenes), which leads to higher levels of biogenic 464 SOA.

465

The NO_3^- signal appears again mainly restricted to the stronger emitting-areas, being located in this season between northern France, Belgium and Germany (Figure 4j 6d). There, the reduced precipitation during NAO^+ events doubles the concentrations of nitrate in comparison to the levels during NAO^- phases.

470

SALT depicts slight variations associated to the NAO phase in summer (around 0.5 µg m⁻³,
Figure 45b). However, contrary to the winter analysis, significant impact areas are now
located over land areas and not over the sea, namely in eastern Iberia and Italy and in
northeastern Europe, where the results show positive NAO⁺ minus NAO⁻ differences.

475

476 Last, PM2.5 and PM10 show differences up to 10 and 20 μg m⁻³ respectively (higher
477 concentrations during NAO⁺), mainly concentrated in the northernmost areas of the
478 domain and representing variations up to 20-40% between NAO phases (Figure 5h,j 8b,d)
479 as observed for winter.

480

481

482 **5 Conclusions**

483

This study establishes the strong impact that the NAO-related local atmospheric processes have on mean ground-level aerosol concentrations over Europe. For that we use a 30-year long air quality simulation with a spatial resolution of 25 km over the target region in which the masking influence of human policies aimed at reducing emissions has been intentionally omitted. Moreover, this simulation allows isolating the influence of the local processes, i.e. those taking place within the boundaries of the domain, as the boundary

490 conditions for the aerosols concentration levels did not vary from year to year in our 491 experimental design. In this sense, it must be underlined that our results should not be 492 considered deterministic, i.e. while the whole real picture does include varying 493 emissions and the effect of long-range pollutant transport, our study, aimed at 494 getting a better understanding of the natural and local processes, is narrowly 495 focused on them.

496

497 The results show impacts with asymmetries in both time (i.e. between seasons) and space 498 (i.e. between northern and southern areas). In winter, higher ground-level concentrations 499 of all aerosol species (except for sea salt) are observed around the Mediterranean basin 500 during the positive NAO phases, while these signals are northward shifted in the summer 501 season. These differences involve variations up to and over 100% in the mean levels of 502 each species; about 20-40% for PM10 and PM2.5. Eventually, softer signals of opposite 503 sign (i.e. NAO⁻ enhancing the ground-level concentration of aerosols) are observed in 504 northern (southern) areas in winter (summer).

505

The causes for these NAO-related variations in the levels of aerosols have to be sought in 506 a multiplicity of climatic factors varying between NAO positive/negative phases, namely: -507 508 The main climate-related mechanisms include (1) increased/reduced temperature 509 (particularly in northern Europe), (2) different distribution of the precipitation patterns 510 across Europe, (3) increased changes in the photolysis of primary and secondary pollutants due to changes in lower cloudiness, and (4) the cleaning effect of enhanced 511 winds (Katragkou et al., 2010). According to the results of this work and based on the 512 513 established relationship between meteorological fields and air-quality (e.g. Wu et al., 2008; Katragkou et al., 2010; Jiménez-Guerrero et al., 2012; Manders et al., 2012), the 514 515 NAO impact on climate supports the NAO impact on aerosol concentrations as follows. Changes in precipitation drives the NAO impact on modification in the 516 517 concentration of most aerosol components (in both northern and southern Europe), since 518 a with the decrease in the precipitation modeled during NAO⁺ phases leadings to a 519 regional increase in the levels of secondary inorganic aerosols and mineral matter (e.g. Jiménez-Guerrero et al., 2012; Manders et al., 2012). As well, the weaker winds 520 521 associated to the NAO⁺ events favor the increase of particulate matter in polluted 522 regions such as large cities or entire industrial regions (e.g. Po valey in northern Italy and 523 the Rhine-Ruhr area in northern Germany and Holland). Also, the enhanced oxidative

524 capacity of the atmosphere with high temperatures causes SO₂ gas-phase emissions to turn into the particulate phase, thus increasing sulphate concentrations as observed 525 526 during NAO⁺ phases in southern (northern) regions in winter (summer). Last, the levels of secondary organic aerosols (SOA) are conditioned by the dependence of 527 528 biogenic emissions on the climatological patterns of variability. In this sense, SOA over Europe is mainly driven by the warming-induced increase in biogenic emitting activity. 529 Although vegetation is kept invariable in the simulation analyzed here, MEGAN 530 estimations of these emissions strongly depends on shortwave radiation and temperature 531 (Guenther et al., 2006), which are substantially conditioned by the cloudiness and thereby 532 533 by the NAO. Accordingly, the lower cloudiness associated to NAO⁺ phases mostly in southern Europe arises as a main driver for the secondary conversion of aerosols. 534

535

536 These results deepen on the knowledge between the climatic conditions and air quality levels, highlighting that the great dependence of the European climate on the NAO phase 537 538 has associated strong natural variations in the aerosols concentration levels. Additionally, 539 we are confident that these results can provide the basis for inferring future air quality scenarios from either future projections or short-to-medium range forecasts of the NAO. 540 541 Although the potential predictability of the NAO or other large-scale climatic indexes is still moderate (Gámiz-Fortis et al., 2002; Saunders and Quian, 2002) and future projections of 542 543 the NAO differ much from one experiment to the other (Hurrell et al., 2003), there are 544 realistic expectations that this would be largely improved in the near future (e.g. Brands et 545 al., 2012).

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547

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553 **References**

554

Ballester, F., Medina, S., Boldo, E., Goodman, P., Neuberger, M., Iñiguez, C., and Künzli,
N.: Reducing ambient levels of fine particulates could substantially improve health: a
mortality impact assessment for 26 European cities. Journal of Epidemiology &
Community Health, 62: 98-105, 2008.

559

Bessagnet, B., Hodzic, A., Vautard, R., Beekmann, M., Cheinet, S., Honore, C., Liousse,
C., and Rouil, L.: Aerosol modelling with CHIMERE - Preliminary evaluation at the
continental scale, Atmospheric Environment, 38:2803-2817, 2004.

563

Bladé, I., Liebmann, B., Fortuny, D., and van Oldenborgh, G.J.: Observed and simulated
impacts of the summer NAO in Europe: implications for projected drying in the
Mediterranean region, Climate Dynamics, 39(3-4):709-727, 2012.

567

568 Brands, S., Manzanas, R., Gutiérrez, J.M., and Cohen, J.: Seasonal Predictability of 569 Wintertime Precipitation in Europe Using the Snow Advance Index, Journal of Climate, 570 25(12):4023-4028, 2012.

571

572 Carvalho, A., Monteiro, A., Solman, S., Miranda, A. I., and Borrego, C.: Climate-driven
573 changes in air quality over Europe by the end of the 21st century, with special reference
574 to Portugal, Environmental Science and Policy, 13:445-458, 2010.

575

576 Chiacchio, M., and Wild, M.: Influence of NAO and clouds on long-term seasonal 577 variations of surface solar radiation in Europe, Journal of Geophysical Research, 115, 578 D00D22, doi:10.1029/2009JD012182, 2010.

579

580 Chin, M., Ginoux, P., Kinne, S., Torres, O., Holben, B.N., Duncan, B.N., Martin, R.V., Lo-

581 gan, J.A., Higurashi, A., Nakajima, T., 2002: Tropospheric Aerosol Optical Thickness

582 from the GOCART Model and Comparisons with Satellite and Sun Photometer Measure-583 ments, Journal of Atmospheric Sciences, 59:461–483, 2002.

584

585 Christoudias, T., Pozzer, A., and Lelieveld, J.: Influence of the North Atlantic Oscillation on 586 air pollution transport, Atmospheric Chemistry and Physics, 12:869-877, 2012.

- Colette, A., Bessagnet, B., Vautard, R., Szopa, S., Rao, S., Schucht, S., Klimont, Z.,
 Menut, L., Clain, G., Meleux, F., Curci, G., and Rouïl, L.: European atmosphere in
 2050, a regional air quality and climate perspective under CMIP5 scenarios,
 Atmospheric Chemistry and Physics, 13:7451-7471, 2013.
- 592

587

593 Cusack, M., Alastuey, A., Pérez, N., Pey, J., and Querol, X.: Trends of particulate 594 matter (PM 2.5) and chemical composition at a regional background site in the 595 Western Mediterranean over the last nine years (2002–2010), Atmospheric 596 Chemistry and Physics, 12(18):8341-8357, 2012.

597

598 Dayan, U., Ziv, B., Shoob, T., and Enzel, Y.: Suspended dust over southeastern 599 Mediterranean and its relation to atmospheric circulations, International Journal of 600 Climatology, 28(7):915-924, 2008.

601

Dayan, U.: Impacts of the NAO on atmospheric pollution in the Mediterranean Basin, in
Advances in Global Change Research: Hydrological, Socioeconomic and Ecological
Impacts of the North Atlantic Oscillation in the Mediterranean Region, S.M. VicenteSerrano and R.M. Trigo Ed., Springer, 171-181, 2011.

606

Derognat, C., Beekmann, M., Baeumle, M., Martin, D., and Schmidt, H.: Effect of biogenic
volatile organic compound emissions on tropospheric chemistry during the atmospheric
pollution over the Paris Area (ESQUIF) campaign in the Ile-de-France region, Journal of
Geophysical Research, 108, 8560, doi:10.1029/2001JD001421, 2003.

611

Eckhardt, S., Stohl, A., Beirle, S., Spichtinger, N., James, P., Forster, C., Junker, C.,
Wagner, T., Platt, U., and Jennings, S.G.: The North Atlantic Oscillation controls air
pollution transport to the Arctic, Atmospheric Chemistry and Physics, 3(5):1769-1778,
2003.

616

Flaounas, E., Coll, I., Armengaud, A., and Schmechtig, C.: The representation of dust
 transport and missing urban sources as major issues for the simulation of PM
 episodes in a Mediterranean area, Atmospheric Chemistry and Physics, 9: 8091-

620 **8101, 2009.**

621

Fiore, A.M., Levy, H., and Jaffe, D.A.: North American isoprene influence on
 intercontinental ozone pollution, Atmospheric Chemistry and Physics, 11(4):1697 1710, 2011.

625

Folberth, G., Hauglustaine, D.A., Lathiere, J., and Brocheton, J.: Interactive chemistry in
the Laboratoire de Meteorologie Dynamique general circulation model: model
description and impact analysis of biogenic hydrocarbons on tropospheric chemistry,
Atmospheric Chemistry and Physics, 6:2273-2319, 2006.

630

Forkel, R., Werhahn, J., Hansen, A. B., McKeen, S., Peckham, S., Grell, G., and
Suppan, P.: Effect of aerosol-radiation feedback on regional air quality - a case
study with WRF/Chem, Atmospheric Environment, 53:202-211, 2012.

634

Folland, C.K., Knight, J., Linderholm, H.W., Fereday, D., Ineson, S., Hurrell, J.W.: The
summer North Atlantic Oscillation: past, present, and future, Journal of Climate,
22:1082-1103, 2009.

638

Gámiz-Fortis, S. R., Pozo-Vázquez, D., Esteban-Parra, M.J., and Castro-Díez, Y.: Spectral
characteristics and predictability of the NAO assessed through Singular Spectral
Analysis, Journal of Geophysical Research, 107, 4685, doi:10.1029/2001JD001436,
2002.

643

Gomez-Navarro, J.J., Montavez, J.P., Jerez, S., Jimenez-Guerrero, P., Lorente-Plazas, R.,
González-Rouco, J.F., and Zorita, E.: A regional climate simulation over the Iberian
Peninsula for the last millennium, Climate of the Past, 7(2):451-472, 2011.

647

Gouveia, C., and Trigo, R.M.: Infuence of climate variability on wheat production in
Portugal, in *geoENV VI - Geostatistics for Environmental Applications*, A. Soares, M.J.
Pereira and R. Dimitrakopoulos Ed., Springer, 335-345, 2008.

651

652 Greatbatch, R.J., and Rong, P.P.: Discrepancies between different Northern 653 Hemisphere summer atmospheric data products, Journal of Climate, 19:1261-

654 1273, 2006. 655 656 Grell, G.A., Dudhia, J., and Stauffer, D.R.: A description of the fifth-generation Penn State/NCAR Mesoscale Model (MM5), NCAR Tech. Note 398+STR, Natl. Cent. For 657 658 Atmos. Res, Boulder, Colo., 1994. 659 Guenther, A., Karl, T., Harley, P., Wiedinmyer, C., Palmer, P.I., and Geron, C.: Estimates of 660 global terrestrial isoprene emissions using MEGAN (Model of Emissions of Gases and 661 662 Aerosols from Nature), Atmospheric Chemistry and Physics, 6:3181-3210, 2006. 663 Hauglustaine, D.A., Hourdin, F., Jourdain, L., Filiberti, M.A., Walters, S., Lamargue, J.-F., 664 665 and Holland, E.A.: Interactive chemistry in the Laboratoire de Meteorologie Dynamigue 666 general circulation model: Description and background tropospheric chemistry 667 evaluation, Journal of Geophysic Research, 109, D04314, doi:10.1029/2003JD003957, 668 2004. 669 Hedegaard, G.B., Christensen, J.H., and Brandt, J.: The relative importance of 670 impacts from climate change vs. emissions change on air pollution levels in the 671 672 21st century, Atmospheric Chemistry and Physics, 13:3569-3585, 2013. 673 674 Hurrell, J.: Decadal trends in North Atlantic Oscillation and relationship to regional temperature and precipitation, Science, 269: 676-679, 1995. 675 676 677 Hurrell, J., and Van Loon, H.: Decadal variations in climate associated with the North Atlantic Oscillation, Climatic Change, 36: 301-326, 1997. 678 679 680 Hurrell, J., Kushnir, Y., Ottersen, G., and Visbeck, M.: The North Atlantic Oscillation -681 Climatic Significance and Environmental Impact, Geophysical Monograph Series, No. 682 134, American Geophysical Union, 2003. 683 Jacob, D.J., and Winner, D.A.: Effect of climate change on air guality, Atmospheric 684 685 Environment, 43: 51-63, 2009. 686 687 Jerez, S., Montavez, J.P., Jiménez-Guerrero, P., Gomez-Navarro, J.J., Lorente-Plazas, R.,

and Zorita, E.: A multi-physics ensemble of present-day climate regional simulations
over the Iberian Peninsula, Climate Dynamics, 40, 3023-3046, 2013a.

690

Jerez, S., Trigo, R.M., Vicente-Serrano, S.M., Pozo-Vázquez, D., Lorente-Plazas, R.,
Lorenzo-Lacruz, J., Santos-Alamillos, F., Montávez, J.P.: The impact of the North
Atlantic Oscillation on renewable energy resources in southwestern Europe, Journal of
Applied Meteorology and Climatology, doi:10.1175/JAMC-D-12-0257.1, 2013b.

- 695
- Jiménez-Guerrero, P., Gómez-Navarro, J.J., Jerez, S., Lorente-Plazas, R., García-Valero,
 J.A., and Montávez, J.P.: Isolating the effects of climate change in the variation of
 secondary inorganic aerosols (SIA) in Europe for the 21st century (1991-2100),
 Atmospheric Environment, 45(4):1059-1063, 2011.
- 700

Jiménez-Guerrero, P., Montávez, J.P., Gómez-Navarro, J.J., Jerez, S., and LorentePlazas, R.: Impacts of climate change on ground level gas-phase pollutants and
aerosols in the Iberian Peninsula for the late XXI century, Atmospheric Environment,
55:483-495, 2012.

705

Juda-Rezler, K., Reizer, M., Huszar, P., Krüger, B., Zanis, P., Syrakov, D., Katragkou, E.,
Trapp, W., Melas, D., Chervenkov, H., Tegoulias, I., and Halenka, T.: Modelling the
effects of climate change on air quality over Central and Eastern Europe: concept,
evaluation and projections, Climate Research, 53:179-203, 2012.

710

Katragkou, E., Zanis, P., Tegoulias, I., Melas, D., Kruger, B.C., Huszar, P., Halenka, T., and
Rauscher, S.: Decadal regional air quality simulations over Europe in present climate:
Near surface ozone sensitivity to external meteorological forcing, Atmospheric Chemistry
and Physics, 10(11):805-821, 2010.

715

Kotlarski, S., Block, A., Böhm, U., Jacob, D., Keuler, K., Knoche, R., Rechid, D., and
Walter, A.: Regional climate model simulations as input for hydrological
applications: evaluation of uncertainties. Advances in Geosciences, 5:119-125,
2005.

720

721 Leibensperger, E.M., Mickley, L.J., Jacob, D.J., Barret, S.R.H.: Intercontinental

influence of NOx and CO emissions on particulate matter air quality, Atmospheric
 Environment, 45(19): 3318-3324, 2011.

724

Manders, A.M.M., van Meijgaard, E., Mues, A.C., Kranenburg, R., van Ulft, L.H., and
Schaap, M.: The impact of differences in large-scale circulation output from climate
models on the regional modeling of ozone and PM, Atmospheric Chemistry and Physics,
12:9441-9458, 2012.

729

Meier, J., Tegen, I., Heinold, B., and Wolke, R.: Direct and semi-direct radiative
effects of absorbing aerosols in Europe: Results from a regional model,
Geophysical Research Letters, 39:L09802, 2012.

733

Monteiro, A., Miranda, A. I., Borrego, C., Vautard, R., Ferreira, J., and Perez, A. T.:
 Long-term assessment of particulate matter using chimere model, Atmospheric
 Environment, 41,:7726-7738, 2007.

737

Moulin, C., Lambert, C.E., Dulac, F., and Dayan, U.: Control of atmospheric export of dust
from North Africa by the North Atlantic Oscillation, Nature, 387, 691-694, 1997.

740

Nenes, A., Pandis, S.N., and Pilinis, C.: ISORROPIA: A new thermodynamic equilibrium
model for multiphase multicomponent inorganic aerosols, Aquatic Geochemistry, 4:123152, 1998.

744

Osborn, T.J., Briffa, K.R., Tett, S.F.B., Jones, P.D., and Trigo, R.M.: Evaluation of the North
Atlantic Oscillation as simulated by a coupled climate model, Climate Dynamics,
15(9):685-702, 1999.

748

Osborn, T.J.: Variability and Changes in the North Atlantic Oscillation Index, Advances inGlobal Change Research, 46:9-22, 2011.

751

Pay, M.T., Piot, M., Jorba, O., Gassó, S., Gonçalves, M., Basart, S., Dabdub, D., JiménezGuerrero, P., and Baldasano, J.M.: A full year evaluation of the CALIOPE-EU air quality
modeling system over Europe for 2004, Atmospheric Environment, 44:3322-3342, 2010.

Péré, J.C., Mallet, M., Pont, V., and Bessagnet, B.: Evaluation of an aerosol optical
 scheme in the chemistry-transport model CHIMERE, Atmospheric Environment,
 44:3688-3699, 2010.

759

Pey, J., Querol, X., Alastuey, A., Forastiere, F., and Stafoggia, M.:African dust
 outbreaks over the Mediterranean Basin during 2001–2011: PM 10 concentrations,
 phenomenology and trends, and its relation with synoptic and mesoscale
 meteorology, Atmospheric Chemistry and Physics, 13(3):1395-1410, 2013.

764

Pfeiffer, A., and Zängl, G.: Validation of climate-mode MM5-simulations for the
 European Alpine Region, Theoretical and Applied Climatology, 101(1-2):93-108,
 2010.

768

Pope, C.A., and Dockery, D.W.: Health effects of fine particulate air pollution: lines that
 connect. Journal of the Air & Waste Management Association, 56: 709-742, 2006.

771

Pope, C.A., Ezzati, M., and Dockery, D.W.: Fine-particulate air pollution and life
expectancy in the United States. New England Journal of Medicine, 360: 376-386, 2009.

Pozzer, A., Zimmermann, P., Doering, U.M., van Aardenne, J., Tost, H., Dentener, F.,
 Janssens-Maenhout, G., and Lelieveld, J.: Effects of business-as-usual
 anthropogenic emissions on air quality, Atmospheric Chemistry and Physics,
 12:6915-6937, 2012.

779

Renfrew, I.A., Petersen, G.N., Sproson, D.A.J., Moore, G.W.K., Adiwidjaja, H., Zhang,
 S., and North, R.: A comparison of aircraft-based surface-layer observations over
 Denmark Strait and the Irminger Sea with meteorological analyses and QuikSCAT
 winds, Quarterly Journal of the Royal Meteorological Society, 135(645):2046-2066,
 2009.

785

Rouil, L., Honore, C., Vautard, R., Beekmann, M., Bessagnet, B., Malherbe, L., Meleux, F.,
Dufour, A., Elichegaray, C., Flaud J-M, Menut, L., Martin, D., Peuch, A., Peuch, V.-H.,
and Poisson, N.: PREV'AIR, an operational forecasting and mapping system for air
quality in Europe, Bulletin of the American Meteorology Society, 90:73-83, 2009.

- 790
- Sanchez-Lorenzo, A., Calbó, J., and Martin-Vide, J.: Spatial and temporal trends in
 sunshine duration over Western Europe (1938-2004), Journal of Climate, 21(22):60896098, 2008.
- 794

Saunders, M. A., and Qian, B.: Seasonal predictability of the winter NAO from north
Atlantic sea surface temperatures, Geophysical Research Letters, 29, 2049,
doi:10.1029/2002GL014952, 2002.

- 798
- Snedecor, G.W., and Cochran, W.G.: Statistical Methods, 8th Edition, ed. Iowa State
 University Press, 1989.

801

- Suski, C.D., and Ridgway, M.S.: Climate and body size influence nest survival in a fish with parental care, Journal of Animal Ecology, 76(4):730-739, 2007.
- 804
- Szopa, S., Foret, G., Menut, L., and Cozic, A.: Impact of large scale circulation on
 European summer surface ozone and consequences for modelling forecast,
 Atmospheric Environment, 43:1189-1195, 2009.
- 808
- Tørseth, K., Aas, W., Breivik, K., Fjæraa, A.M., Fiebig, M., Hjellbrekke, A.G., Lund Myhre,
 C., Solberg, S., and Yttri, K.E.: Introduction to the European Monitoring and Evaluation
 Programme (EMEP) and observed atmospheric composition change during 1972–2009,
 Atmospheric Chemistry and Physics, 12:5447-5481, 2012.
- 813

Trigo, I.F.: Climatology and interannual variability of storm-tracks in the Euro-Atlantic sector: a comparison between ERA-40 and NCEP/NCAR reanalyses, Climate Dynamics, 26(2):127-143, 2006.

817

Trigo, R.M., Osborn, T.J., and Corte-Real, J.M.: The North Atlantic Oscillation influence on
Europe: climate impacts and associated physical mechanisms, Climate Research,
20(1):9-17, 2002.

821

Trigo, R.M., Valente, M.A., Trigo, I.F., Miranda, P.M.A., Ramos, A.M., Paredes, D., and García-Herrera, R.: The impact of north Atlantic wind and cyclone trends on European precipitation and significant wave height in the Atlantic, Annals of the New York Academy
of Sciences, 1146:212-234, 2008.

826

827 Uppala, S.M., Källberg, P.W., Simmons, A.J., Andrae, U., Bechtold, V., Fiorino, M., Gibson, 828 J.K., Haseler, J., Hernandez, A., Kelly, G.A., Li, X., Onogi, K., Saarinen, S., Sokka, N., Allan, R.P., Andersson, E., Arpe, K., Balmaseda, M.A., Beljaars, A.C.M., Van De Berg, 829 L., Bidlot, J., Bormann, N., Caires, S., Chevallier, F., Dethof, A., Dragosavac, M., Fisher, 830 M., Fuentes, M., Hagemann, S., Hólm, E., Hoskins, B.J., Isaksen, L., Janssen, P.A.E.M., 831 832 Jenne, R., Mcnally, A.P., Mahfouf, J.-F., Morcrette, J.-J., Rayner, N.A., Saunders, R.W., Simon, P., Sterl, A., Trenberth, K.E., Untch, A., Vasiljevic, D., Viterbo, P., and Woollen, J.: 833 The ERA-40 re-analysis, Quarterly Journal of the Royal Meteorological Society, 834 835 131:2961-3012, 2005.

836

Van Dingenen, R., Dentener, F. J., Raes, F., Krol, M.C., Emberson, L., and Cofala, J.: The
global impact of ozone on agricultural crop yields under current and future air quality
legislation, Atmospheric Enviroment, 43: 604-618, 2009.

840

Vestreng, V., Ntziachristos, L., Semb, A., Reis, S., Isaksen, I. S. A., and Tarrasón, L.:
Evolution of NOx emissions in Europe with focus on road transport control measures,
Atmospheric Chemistry and Physics, 9:1503-1520, 2009.

844

Vicente-Serrano, S.M., and Trigo, R.M., Ed.: Advances in Global Change Research.
Hydrological, Socioeconomic and Ecological Impacts of the North Atlantic Oscillation in
the Mediterranean Region, Springer, 2011.

848

Wanner, H., Brönnimann, S., Casty, C., Gyalistras, D., Luterbacher, J., Schmutz, C.,
Stephenson D.B., and Xoplaki, E.: North Atlantic Oscillation - concepts and studies,
Surveys on Geophysics, 22: 321-382, 2001.

852

Wu, S., Mickley, L. J., Jacob, D. J., Rind, D., and Streets, D. G.: Effects of 2000–2050
changes in climate and emissions on global tropospheric ozone and the policyrelevant background surface ozone in the United States, Journal of Geophysical
Research, 113:D18312, 2008.

857 List of captions

858

Figure 1. Evaluation of the MM5-CHIMERE air quality modeling system described in Section 2.1. Gray shaded colors depict the simulated climatologies in μ g m⁻³ of PM10 (up) and PM2.5 (bottom) in winter (DJFM averages, left) and summer (JJAS averages, right). The comparison of the simulated series with EMEP observations is provided by symbols: their color shows the difference in the standard deviation of the series between the simulation and the observations (in μ g m⁻³), and their shape informs on the magnitude of the temporal correlation between the simulated and the observed series.

866

Figure 2. Winter (**up** top) and summer (bottom) NAO series for the period 1970-1999. The former are DJFM averages (therefore resulting in 29 values) and the latter are JJAS averages (resulting in 30 values). The red (blue) lines denote the 70th (30th) percentile value of each series, **having been highlighted highlighting** the years/values with a NAO index above (bellow) that percentiles, i.e. those selected as NAO⁺ (NAO⁻) events for the composites analysis.

873

Figure 3. NAO impact in winter (left) and summer (right) on atmospheric conditions. (a,b) depict mean 10m-wind direction during NAO⁺ (red) and NAO⁻ (blue) phases (being the arrows length proportional to the wind speed). (c,d) provide the NAO⁺ minus NAO⁻ composites for the mean 10m-wind module (in m/s). These differences are represented only if they are statistically significant at the 90% level, dots blur the values not representing a signal-to-noise ratio above 1, and contours depict differences expressed as percentage.

881

Figure 4. As in Figure 3c,d, here for (a,b) precipitation (in mm/month), (c,d)
integrated cloud water (in mm) and (e,f) 2m-temperature (in K).

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Figure 5. NAO impact in winter (left) and summer (right) on mean ground levels of natural inert aerosols: NAO⁺ minus NAO⁻ composites for (a,b) SALT and (c,d) DUST. Differences in µg m⁻³ are represented only if they are statistically significant at the 90% level, dots blur the values not representing a signal-to-noise ratio above 1, and contours depict differences expressed as percentage. Period considered: 1970-

890 **1999.**

891

Figure 6. As in Figure 5, here for secondary inorganic aerosols: (a,b) SO_4^{2-} , (c,d) NH_4^+ and (e,f) NO_3^- .

894

Figure 7. As in Figure 5, here for organic materials: (a,b) OM, (c,d) SOA and (e,f) EC.
896

897 Figure 8. As in Figure 5, here for particulate matter: (a,b) PM10 and (c,d) PM2.5.

898

899 Figure 3. NAO-impact in winter (left) and summer (right) on the atmospheric-900 conditions. (a,b) depict mean 10m-wind direction during NAO⁺ (red) and NAO⁻ (blue) 901 phases (being the arrows length proportional to the wind speed). The rest of panels 902 provide the NAO⁺ minus NAO⁻ composites for mean (c,d) 10m-wind module (in m/s), (e,f) precipitation (in mm/month), (g,h) integrated cloud water (in mm) and (e,f) 2m-903 904 temperature (in K). Differences are represented only if they are statistically-905 significant at the 90% level, dots blur the values not representing a signal-to-noise 906 ratio above 1, and contours depict differences expressed as percentage.

907

908 Figure 4. NAO-impact on mean ground-level aerosols concentrations: NAO⁺ minus 909 NAO⁻ composites for (a,b) SALT, (c,d) DUST, (e,f) SO₄²⁻, (g,h) NH₄⁺ and (i,j) NO₃⁻ in 910 winter (left) and summer (right). Differences in µg m⁻³ are represented only if they 911 are statistically significant at the 90% level, dots blur the values not representing a 912 signal-to-noise ratio above 1, and contours depict differences expressed as-913 percentage.

914

915 Figure 5. As Figure 4 for (a,b) OM, (c,d) SOA, (e,f) EC, (g,h) PM10 and (i,j) PM2.5.



Figure 1. Evaluation of the MM5-CHIMERE air quality modeling system described in Section 2.1. Gray shaded colors depict the simulated climatologies in μ g m⁻³ of PM10 (up) and PM2.5 (bottom) in winter (DJFM averages, left) and summer (JJAS averages, right). The comparison of the simulated series with EMEP observations is provided by symbols: their color shows the difference in the standard deviation of the series between the simulation and the observations (in μ g m⁻³), and their shape informs on the magnitude of the temporal correlation between the simulated and the observed series.



Figure 2. Winter (top) and summer (bottom) NAO series for the period 1970-1999. The former are DJFM averages (therefore resulting in 29 values) and the latter are JJAS averages (resulting in 30 values). The red (blue) lines denote the 70th (30th) percentile value of each series, highlighting the years/values with a NAO index above (bellow) that percentiles, i.e. those selected as NAO⁺ (NAO⁻) events for the composites analysis.



Figure 3. NAO impact in winter (left) and summer (right) on atmospheric conditions. (a,b) depict mean 10m-wind direction during NAO⁺ (red) and NAO⁻ (blue) phases (being the arrows length proportional to the wind speed). (c,d) provide the NAO⁺ minus NAO⁻ composites for the mean 10m-wind module (in m/s). These differences are represented only if they are statistically significant at the 90% level, dots blur the values not representing a signal-to-noise ratio above 1, and contours depict differences expressed as percentage.



Figure 4. As in Figure 3c,d, here for (a,b) precipitation (in mm/month), (c,d) integrated cloud water (in 934 mm) and (e,f) 2m-temperature (in K).



Figure 5. NAO impact in winter (left) and summer (right) on mean ground levels of natural inert aerosols: NAO⁺ minus NAO⁻ composites for (a,b) SALT and (c,d) DUST. Differences in µg m⁻³ are represented only if they are statistically significant at the 90% level, dots blur the values not representing a signal-to-noise ratio above 1, and contours depict differences expressed as percentage. Period considered: 1970-1999.



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