### 1 Ozone vegetation damage effects on gross primary productivity in the United States 2 3 Xu Yue and Nadine Unger 4 School of Forestry and Environmental Studies, Yale University, 195 Prospect Street, 5 New Haven, CT 06511 6 Correspondence to: Xu Yue (xu.yue@yale.edu) 7 8 **Abstract** 9 10 We apply an off-line process-based vegetation model (the Yale Interactive Terrestrial Biosphere model) to assess the impacts of ozone (O<sub>3</sub>) vegetation damage on gross 11 12 primary productivity (GPP) in the United States during the past decade (1998-2007). The 13 model's GPP simulation is evaluated at 40 sites of the North American Carbon Program 14 (NACP) synthesis. The ecosystem-scale model version reproduces interannual variability 15 and seasonality of GPP at most sites, especially in croplands. Inclusion of the O<sub>3</sub> damage 16 impact decreases biases of simulated GPP at most of the NACP sites. The simulation with the O<sub>3</sub> damage effect reproduces 64% of the observed variance in summer GPP and 17 42% on the annual average. Based on a regional gridded simulation over the U.S., 18 summertime average O<sub>3</sub>-free GPP is 6.1 g C m<sup>-2</sup> day<sup>-1</sup> (9.5 g C m<sup>-2</sup> day<sup>-1</sup> in the East of 19 95°W and 3.9 g C m<sup>-2</sup> day<sup>-1</sup> in the West). O<sub>3</sub> damage decreases GPP by 4-8% on average 20 in the eastern U.S. and leads to significant decreases of 11-17% in east coast hotspots. 21 22 Sensitivity simulations show that a 25% decrease in surface O<sub>3</sub> concentration halves the 23 average GPP damage to only 2-4% suggesting the substantial co-benefits to ecosystem 24 health that may be achieved via O<sub>3</sub> air pollution control. 25 26 Keywords: Ozone, gross primary productivity, photosynthesis, stomatal conductance 27

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#### 1 Introduction

The effects of tropospheric ozone (O<sub>3</sub>) damage on U.S. forests have been studied for half a century (Karnosky et al., 2007), but the impacts of O<sub>3</sub> on the North American carbon balance are still relatively poorly understood (Felzer et al., 2004; Huntingford et al., 2011). O<sub>3</sub> is a secondary pollutant produced in the troposphere during the photochemical oxidation of carbon monoxide, methane, and volatile organic compounds (VOCs) by the major tropospheric oxidant, the hydroxyl radical, in the presence of sunlight and nitrogen oxides. Fossil-fuel, biofuel and biomass burning since the industrial and agricultural revolutions have greatly increased the emissions of O<sub>3</sub> precursors and led to an approximate doubling of O<sub>3</sub> levels over the U.S. since the preindustrial. Deposition through stomatal uptake is an important sink for O<sub>3</sub> but damages photosynthesis, reduces plant growth and biomass accumulation, limits crop yields, and affects stomatal control over plant transpiration of water vapor between the leaf surface and atmosphere (Ainsworth et al., 2012; Hollaway et al., 2012).

Understanding the O<sub>3</sub> pollution influence on the North American forest sink is crucial to any effort to mitigate climethange by stabilizing atmospheric carbon dioxide (CO<sub>2</sub>) concentrations. Currently, North America is acting as a net source of CO<sub>2</sub> to the atmosphere (King et al., 2012). Sequestration of atmospheric CO<sub>2</sub> by forest ecosystems is a major control on atmospheric CO<sub>2</sub> abundance and its growth rate (Pan et al., 2011). Terrestrial ecosystems of North America absorb the equivalent of about 35% of North America's fossil fuel based CO<sub>2</sub> emissions, representing a source-to-sink ratio of nearly 3:1. Forestrowth in the U.S. is responsible for 30–70% of this North American CO<sub>2</sub> sink, which varies significantly from year to year (Pacala et al., 2001; Goodale et al., 2002; Pan et al., 2011; King et al., 2012). Worse still, there is evidence that the summer continental U.S. is more set than other world regions to climate forcing (Levy et al., 2008).

Experimental studies that examine O<sub>3</sub> impacts on plant productivity are typically performed for individual vegetation types, on the scale of sites, and within a limited time

period (e.g. Wittig et al., 2007; Feng et al., 2008; Lombardozzi et al., 2013). For example, based on measurements reported from over 100 studies, Wittig et al. (2007) estimated that chronic O<sub>3</sub> exposure depressed photosynthesis by 11% and stomatal conductance by 13% for several tree species at the ambient O<sub>3</sub> level of ~45 ppbv relative to that in O<sub>3</sub>-free air. The  $O_3$  damage effect is strongest for crops. With datasets from  $\sim 50$  peer-reviewed studies, Feng et al. (2008) estimated that elevated O<sub>3</sub> levels significantly decrease wheat photosynthetic rates by 20% and stomatal conductance by 22%. Emerging research has found that the O<sub>3</sub> vegetation damage effects may result in a loss of plant stomatal control, and a consequent decoupling of the stomatal response from photosynthesis inhibition (Lombardozzi et al., 2012a, 2012b, 2013).

Previous work has found that in the U.S. region during 1989-1993, O<sub>3</sub> pollution reduced net primary productivity (NPP) by 3-7% overall, and up to 13% in hotspots including the southeast and in the Midwest agricultural lands (Felzer et al., 2004, 2005). The indirect CO<sub>2</sub> radiative forcing due to the vegetation damage effects of anthropogenic O<sub>3</sub> increases since the industrial revolution may be as large as +0.4 Wm<sup>-2</sup> (Sitch et al., 2007), which is 25% of the magnitude of the direct CO<sub>2</sub> radiative forcing over the same period, and of similar magnitude to the direct O<sub>3</sub> radiative forcing. Through this perturbation of the carbon cycle, O<sub>3</sub> pollution affects the climate system on considerably longer timescales than its own atmospheric lifetime (Unger and Pan, 2012). Over the past decade since this previous assessment surface O<sub>3</sub> levels in most of the U.S. have decreased (Lefohn et al., 2010) due to domestic emission reductions following the implementation of air quality control legislation (Bloomer et al., 2010). However, increasing O<sub>3</sub> concentration is observed over western U.S. (Jaffe and Ray, 2007). Such a trend may in part be related to the inter-continental flow from Asia (Cooper et al., 2010) and the global increase in methane (Rigby et al., 2008).

The major goal of this study is to assess O<sub>3</sub> damage effects on gross primary productivity (GPP) in the U.S. for the recent decade 1998-2007 using a data-constrained vegetation model. In this work, we describe the implementation of a semi-mechanistic O<sub>3</sub> damage function (Sitch et al., 2007) into the Yale Interactive Terrestrial Biosphere model (YIBs)

that includes enzyme-kinetic biophysics (Unger et al., 2013). In the first stage of the study, we utilize eddy-derived GPP flux measurements at 40 sites across the U.S. and Canada that have been collated for the North American Carbon Program (NACP) sitelevel interim synthesis (Huntzinger et al., 2012; Schaefer et al., 2012; Barr et al., 2013; Ricciuto et al., 2013) to evaluate an off-line version of the vegetation model's site level GPP simulation and to assess the impact of surface O<sub>3</sub> damage at those sites. In the second stage of the study, the impacts of O<sub>3</sub> damage on GPP throughout the entire U.S. region are quantified using a regionally distremental configuration of the vegetation model.

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## 2 Methodology and data

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# 2.1 Vegetation biophysics

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Here, we apply an off-line version of the YIBs model that previously was implemented into the NASA Goddard Institute for Space Studies global chemistry-climate model (Unger et al., 2013). The off-line model can be run at the site-level or in distributed mode for a designated region. The vegetation biophysics module computes the photosynthetic uptake of CO<sub>2</sub> coupled with the transpiration of water vapor at the 1-hour physical integration time step of the off-line model. The vegetation biophysics calculates C3 and C4 photosynthesis using the well-established Michealis-Menten enzyme-kinetics leaf model of photosynthesis (Farquhar et al., 1980; von Caemmerer and Farquhar, 1981) and the stomatal conductance model of Ball and Berry (Collatz et al., 1991). The coupled photosynthesis/stomatal conductance leaf model has been widely used to project terrestrial biosphere responses to global change. The model is briefly summarized here for transparency and completeness. The leaf model assumes that the rate of net CO<sub>2</sub> assimilation  $(A_{net})$  in the leaves of C3 and C4 plants is limited by one of three processes: (i) the capacity of the ribulose 1,5-bisphosphate (RuBP) carboxylase-oxygenase enzyme (Rubisco) to consume RuBP  $(J_c)$ ; (ii) the capacity of the Calvin cycle and the thylakoid reactions to regenerate RuBP supported by electron transport  $(J_e)$ ; (iii) the capacity of starch and sucrose synthesis to consume triose phosphates and regenerate inorganic phosphate for photo-phosphorylation in C3 and phosphoenolpyruvate (PEP) limitation in

124 C4  $(J_s)$ .  $J_c$ ,  $J_e$ , and  $J_s$  are described as functions of the maximum carboxylation capacity

125  $(V_{cmax})$  at the optimal temperature, 25°C, and the internal leaf CO<sub>2</sub> concentration  $(C_i)$ . The

gross rate of carbon assimilation from photosynthesis (A) is given by:

$$A = \min(J_c, J_e, J_s) \tag{1}$$

128 Net carbon assimilation is given by:

$$A_{net} = A - R_d \tag{2}$$

where  $R_d$  is the rate of dark respiration:

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$$R_d = 0.015 \cdot V_{\text{cmax}} \tag{3}$$

Leaf stomata control the uptake of CO<sub>2</sub> versus the loss of H<sub>2</sub>O. At equilibrium, the

stomatal conductance of water vapor through the leaf cuticle (g<sub>s</sub> in mol [H<sub>2</sub>O] m<sup>-2</sup> s<sup>-1</sup>)

depends on the net rate of carbon assimilation:

$$g_s = m \frac{A_{net} \cdot RH}{c_s} + b = \frac{1}{r_s} \tag{4}$$

where m and b are the slope and intercept derived from empirical fitting to the Ball and

Berry stomatal conductance equations, RH is relative humidity,  $c_s$  is the  $CO_2$ 

138 concentration at the leaf surface, and  $r_s$  is the stomatal resistance to water vapor.

Appropriate photosynthesis parameters for the local vegetation type are taken from

140 (Friend and Kiang, 2005) and the Community Land Model (Oleson et al., 2010) with

updates from Bonan et al. (2011) (Table 1). In both the site-level and distributed models,

we apply these model PFT-specific photosynthesis parameters and do not tune or

calibrate to the local vegetation properties. The model calculates evapotranspiration as a

144 function of the stomatal conductance. However, we do not consider the feedback of the

145 changes in evapotranspiration to the boundary-layer meteorology because we use

prescribed meteorological variables from reanalyses in the simulations.

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148 The canopy radiative transfer scheme assumes a closed canopy and layers the canopy for

light stratification using an adaptive number of layers (typically 2-16) (Friend and Kiang,

150 2005). Each canopy layer distinguishes sunlit and shaded regions for which the direct and

diffuse photosynthetically active radiation (PAR) is computed (Spitters et al., 1986). The

coupled photosynthesis and stomatal conductance equations are solved analytically using

a cubic function of  $A_{net}$ .  $C_i$  is calculated explicitly at the leaf level. Scaling of the leaf to

canopy level is through stratification of canopy light levels and leaf area profiles. The photosynthetic uptake of CO<sub>2</sub> is accumulated into a carbon reserve pool, from which other processes may allocate uses.

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## 2.1.1 O<sub>3</sub> damage effect on photosynthesis

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O<sub>3</sub> oxidizes cellular membranes and photosynthetic tissues when it enters leaves through stomata, leading to reductions in photosynthesis and GPP. O<sub>3</sub> damage inhibits stomatal conductance, which is closely related to the photosynthetic rate, resulting in a reduction in transpiration. A semi-mechanistic parameterization is employed to estimate the O<sub>3</sub> damage effects to both photosynthesis and stomatal conductance (Sitch et al., 2007). The exposure to O<sub>3</sub> leads to reductions in photosynthesis:

$$166 A' = F \cdot A_{net} (5)$$

where F is the reduction fraction calculated as

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$$F = 1 - a \cdot U_{>0.3T}$$
 (6)

- where a is the  $O_3$  sensitivity coefficient derived from observations. Two cases are examined: high and low  $O_3$  sensitivity following Sitch et al. (2007).  $U_{>O3T}$  is the
- instantaneous leaf uptake of  $O_3$  flux above a plant function type (PFT)-specific threshold
- 172 of O3T (Table 1),

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$$U_{>O3T} = \max[(F_{O3} - O3T), 0] \tag{7}$$

here  $F_{O3}$  is the O<sub>3</sub> flux entering the leaf through the stomata,

$$F_{O3} = \frac{\left[O_3\right]}{r_b + \kappa \cdot r_s} \tag{8}$$

- where  $[O_3]$  is the  $O_3$  concentration at the top of the canopy,  $r_b$  is the boundary layer
- 177 resistance. The stomatal resistance to  $O_3$  is calculated based on stomatal resistance to
- water  $r_s$  with a ratio constant  $\kappa=1.67$ . From Equation (4), the decrease in  $A_{net}$  reduces the
- stomatal conductance *g*<sub>s</sub> proportionally,

The  $r_s$  and  $g_s$  are the O<sub>3</sub>-damaged stomatal resistance and conductance, respectively. When the plant is exposed to [O<sub>3</sub>] (Equation 8), the excess O<sub>3</sub> flux entering leaves (Equation 7) causes F<1 (Equation 6), decreasing  $A_{net}$  (Equation 5) while increasing the stomatal resistance (Equation 9). The latter will act to reduce the O<sub>3</sub> uptake flux (Equation 8) to protect the plant. Thus, the scheme considers associated changes in both photosynthetic rate and stomatal conductance. When photosynthesis is inhibited by O<sub>3</sub>, the stomatal conductance decreases accordingly to resist more air passing through the stomata, resulting in a decline of the oxidant fluxes inside leaves, as described through equations (5)-(9). Consequently, this coupled scheme represents the equilibrium state between the CO<sub>2</sub> demand for vegetation growth and the protection against O<sub>3</sub> damage by plant. The parameters for the scheme, including the O<sub>3</sub> damage threshold and sensitivity coefficients, were originally derived based on the calibration of the MOSES vegetation model. Since the MOSES model employs the (almost) identical Farquhar-Ball-Berry photosynthesis/stomatal conductance scheme as in the YIBs model, it is appropriate to adopt the same parameters as those derived in Sitch et al. (2007) (Table 1). Evaluation of the YIBs simulated O<sub>3</sub>-induced GPP response with available field and laboratory measurements across a range of PFTs in Section 3.4 indicates that our assumption is reasonable.

# 2.1.2 Vegetation structure

The YIBs vegetation model simulates eight PFTs, using either C3 or C4 photosynthesis (Table 1). We apply two different sets of land cover and leaf area index (LAI) in the simulations. The first set is the PFT-specified vegetation cover fraction and LAI retrieved by the Moderate Resolution Imaging Spectroradiometer (MODIS, Knyazikhin et al., 1998). The value on a specific day is linearly interpolated from the monthly means of the nearest two months based on the distance of this day to the middle dates of those two months. The second set uses LAI from the Global Modeling and Assimilation Office (GMAO) Modern Era-Retrospective Analysis (MERRA) dataset. The MERRA LAI is assimilated based on radiance data retrieved by over 20 satellites (Rienecker et al., 2011) and is available on daily scale from 1980 onwards. Since the MERRA LAI dataset does

not provide PFT-specific information, the actual site-level PFT is assumed for the site level simulations. For the regional distributed simulations, the land cover is prescribed to the gridded International Satellite Land-Surface Climatology Project (ISLSCP, Hall et al., 2006).

# 2.1.3 Meteorological forcing

For the site-level simulations, we use hourly *in situ* measurements of surface meteorological variables, including surface air temperature, specific humidity, wind speed, surface pressure, and CO<sub>2</sub> concentrations. There are some missing values in the measurements due to occasional instrument failure. We gap-fill the site-based observations with that from the MERRA-land data (Reichle et al., 2011), which is interpolated to each site based on the site location.

For the distributed simulations, the off-line YIBs model uses hourly MERRA-land data climatic variables including: surface air temperature, specific humidity, wind speed, surface pressure, precipitation, direct PAR, and diffuse PAR, and soil temperature and soil moisture at 6 soil depths. The original data resolution of  $0.5^{\circ} \times 0.667^{\circ}$  by latitude and longitude is degraded to  $1^{\circ} \times 1.333^{\circ}$  due to current disk space limitation.

# 2.1.4 Surface [O<sub>3</sub>]

Hourly and daily maximum 8-hour average surface [O<sub>3</sub>] representative of the present day climate (~ 2005) are taken from previous simulations using NASA Model-E2 (Shindell et al., 2013). The global model has 2°×2.5° latitude by longitude horizontal resolution with 40-vertical layers extending to 0.1 hPa. The gas-phase chemistry and aerosol modules are fully integrated, so that these components interact with each other and with the physics of the climate model (Bell et al., 2005; Shindell et al., 2006; Unger, 2011; Shindell et al., 2013). The model surface O<sub>3</sub> is validated using measurements from 73 Clean Air Status and Trends Network (CASTNET) sites operated by the Unitied States Environmental Protection Agency (EPA) (http://epa.gov/castnet/javaweb/ozone.html) and ~1200

monitor sites managed by the EPA AIRDATA (<a href="http://www.epa.gov/airdata/">http://www.epa.gov/airdata/</a>). These sites are operated on the county level scale. The CASTNET provides hourly [O<sub>3</sub>] at rural sites from 1996-2005. The AIRDATA network provides daily maximum 8-hour average (MDA8) [O<sub>3</sub>], covering both urban and rural regions. We use AIRDATA data for the year 2005.

#### 2.2 Simulations

#### 2.2.1 Site-level runs

We configure a site-level version of the YIBs model for the 40 eddy covariance flux tower sites described in detail in the NACP synthesis (Fig. S1 and Appendix Table A, Schaefer et al., 2012). Measurements are available for a wide range of time periods across the different sites ranging from the minimum of 1 year at Fermi Lab (US-IB1) and the maximum of 15 years at Harvard Forest (US-HA1). These sites cover a range of different vegetation types including: evergreen needleleaf forest, deciduous broadleaf forest, grasslands, croplands, closed shrublands, mixed forests, permanent wetlands, and woody savannas. Table S1 summarizes how the NACP vegetation types are mapped onto the 8 model PFTs. For the site-level simulations, we assume C4 photosynthetic pathway for all cropland sites, which are mainly corn (Schaefer et al., 2012).

For each site, a group of six sensitivity simulations are performed (Table 2). We conduct the first four runs using different combinations of meteorological and vegetation forcings, to assess the sensitivity of the results to local versus renalaysis meterological forcing and LAI (Table 2). Two, METmerra\_LAImodis and METmerra\_LAImerra, use hourly meteorology from MERRA-land reanalyses alone. The other two, METsite\_LAImodis and METsite\_LAImerra, use site-based meteorology with gap-filled MERRA reanalysis. Simulations use two datasets of LAI: (1) METmerra\_LAImerra and METsite\_LAImerra use LAI from the MERRA-land reanalyses, which provide non PFT-specific LAI that we assign to the local PFT type at each site (Table A), while (2) METmerra\_LAImodis and METsite LAImodis use PFT-specific LAI retrieved by the MODIS. We perform two

additional site-level simulations, which use the same forcings as that for METsite Imerra but with the impact of O<sub>3</sub> uptake on photosynthesis. These two experiments, METsite\_LAImerra\_HO3 and METsite\_LAImerra\_LO3, use either high or low O<sub>3</sub> sensitivity as defined by the coefficient *a* in Table 1.

To quantify the performance of the vegetation model, we estimate the  $\chi^2$  for each site following the method described in Schaefer et al. (2012),

$$\chi^2 = \frac{1}{n} \sum_{i=1}^n \left(\frac{r_i}{\varepsilon_i}\right)^2 \tag{10}$$

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$$r_i = (GPP_{si} - GPP_{oi}) \tag{11}$$

is the difference between the pair of simulated and observed GPPs.  $\varepsilon_i$  are the observation certainties. n is the length of observations (e.g. the number of days for the daily variables). The lower the  $\chi^2$ , the smaller the model biases. If  $\chi^2 < 1$ , the simulation bias is on average smaller than the measurement uncertainty, indicating a good performance of the model. Here, we define a reasonable performance of  $\chi^2 < 4$ , when the residual is less than twice the measurement uncertainty. We also calculate the root mean square error (RMSE) as follows:

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$$RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (GPP_{si} - GPP_{oi})^2}$$
 (12)

We validate the simulated O<sub>3</sub> damage effect with measurements from literature. Field and laboratory experiments may have different [O<sub>3</sub>] compared to the ambient level we used complicating the validation. As a result, we perform 14 additional sensitivity simulations for each of NACP sites. All tests use meteorological and vegetation forcings the same as METsite\_LAImerra (Table 2), except for the different [O<sub>3</sub>] and O<sub>3</sub> sensitivity. These experiments are divided into two groups, 7 in each, using either low or high O<sub>3</sub> sensitivity. In each group, simulations are performed with constant [O<sub>3</sub>] at 20, 40, 60, 80, 100, 120, 140 ppby, respectively. We do not include diurnal and seasonal variations of [O<sub>3</sub>] in these sensitivity simulations as that in METsite LAImerra for two reasons. First,

field measurements for the  $O_3$  vegetation damage are usually performed with fixed  $[O_3]$  during the growth season (e.g. Ishii et al., 2004; Zhang et al., 2012). Second, the diurnal cycles and seasonality of  $[O_3]$  are very different for different sites (Bloomer et al., 2010), making it difficult to apply a uniform temporal cycle for all the NACP sites. The reductions in GPP at these simulations are compared with results from field measurements at the corresponding  $[O_3]$  level.

#### 2.2.2 Distributed run over U.S. region

A gridded version of the YIBs model at  $1^{\circ}\times1.333^{\circ}$  latitude by longitude horizontal resolution for the U.S. region is driven with MERRA meteorological forcings for the period 1998-2007. In the distributed model, vegetation cover types are from the ISLSCP and LAI is form the MERRA-land reanalysis. We assign the MERRA LAI to the corresponding PFT types defined by ISLSCP (Fig. S2). The 18 ISLSCP land types are converted to 8 PFTs used in the model (Table S1). Some of the ISLSCP land types, such as the deciduous needleleaf forest, are not represented in the YIBs model. However, the coverage of these types is very small in the U.S. (Fig. S2) and will not influence the regional simulation after the conversion to the model types. For the regional simulation, we assume that the total crop area in each crop grid cell is split 50% C3 and 50% C4 to account for the dominance of both soybean (C3) and corn (C4) crops in the central and northern U.S. agricultural regime. We perform 2 simulation cases with high and low O<sub>3</sub> damage sensitivity. Finally, to understand how the O<sub>3</sub> vegetation damage effect may respond to possible future changes in [O<sub>3</sub>], we perform four additional sensitivity experiments with  $\pm 25\%$  changes in [O<sub>3</sub>] for each O<sub>3</sub> sensitivity

# 3. Results

#### 3.1 Evaluation of O<sub>3</sub>-free GPP at NACP sites

- We compare the monthly mean LAI from MERRA and MODIS at each NACP site (Fig.
- 1). For each site, the MERRA LAI is averaged for the period when GPP measurements

are available. The two datasets show similar annual cycles at several sites but are inconsistent for 7 out of 20 evergreen sites (CA-Ca1, CA-Ca2, CA-Ca3, CA-NS1, US-Me2, US-Me3, and US-Me5) and 2 out of 5 shrubland sites (US-SO2 and US-Ton). In addition, for grasslands and croplands, the datasets exhibit different seasonality, with MERRA LAI near zero during the cold season in contrast to the MODIS LAI that is positive all year round. It must be emphasized that the MERRA and MODIS LAI represent the average state in the retrieval product grid cells and as such may not represent the local LAI for the actual PFT species at each site. The local site LAI values are not available for comparison. For example, for most of the evergreen NACP sites, the MERRA and MODIS LAI indicate large seasonality.

The long-term monthly mean O<sub>3</sub>-free GPP from the simulation METsite LAImerra is compared with observations at individual NACP sites (Fig. 2). Both the observation and simulation are averaged over the measurement period at each site. The simulations capture the magnitude and seasonality of GPP for most sites especially for deciduous broating and shrubland. The largest model overestimate (factor of 3-8) occurs at CA-SJ1 and CA-SJ2. For the grassland sites US-ARM (in Great Plains) and US-Var (in California), the model maximum GPP occurs in summer (July), 2-3 months later than in the measurements (April). This incorrect model seasonality is a result of the MERRA LAI (compare Fig. 1) that does not begin to increase rapidly until after May and is not consistent with the local LAI at the site. In reality, California grasslands exhibit rapid growth in spring then mature and die after April or May (Chiariello, 1989). The grasslands in the Great Plains may have up to six different phenological groups, including some species active in spring (e.g. in US-ARM) while some others peak in summer (e.g. in US-Shd) (Henebry, 2003). The correlation coefficient between simulated annual mean GPP and observations at all 40 sites is 0.65. The corre is higher (0.81) for summer (June-August, Fig. S3). The annual GPP averaged over all sites is 3.8 g C m<sup>-2</sup> day<sup>-1</sup>, 27% higher than the observational average (3.0 g C m<sup>-2</sup> day<sup>-1</sup>).

Among the 40 NACP sites, 22 have reasonable performance with  $\chi^2$  < 4 (Fig. 3a). The simulation is the best ( $\chi^2$  =1.2) at the evergreen needleleaf site US-Dk3 (Fig. 2). For the

top 10 sites with the lowest biases, 8 are based in evergreen needleleaf forests. These sites usually have multiple years of measurements and provide good samples for testing the consistency between simulations and observations. Simulations at 4 broadleaf forest, 4 cropland, and 3 shrubland sites have  $\chi^2 < 4$ , and the latter usually has low GPP whose peak is not higher than 10 g C m<sup>-2</sup> day<sup>-1</sup> (Fig. 2). Compared with the 24 land surface models in Schaefer et al. (2012), the YIBs model shows significant improvement at the crop PFT sites ( $\chi^2 < 4.1$  vs.  $\chi^2 > 6$ ). YIBs simulates GPP with  $\chi^2 < 4$  at 22 sites in total compared to 16 sites for the ensemble simulations in Schaefer et al. (2012). YIBs GPP simulation is weaker ( $\chi^2 > 4$ ) at 18 sites including 9 needleleaf forest, 2 broadleaf, and 2 shrubland PFTs. The common feature of the biases at these sites is the overestimation of peak GPP during summer (e.g. CA-SJ1, CA-SJ2, CA-Mer, Fig. 2). It is possible that the model does not represent the full realism of the biophysical processes accurately. However, we assert that the most likely cause of the model overestimate is the uniform application of model PFT-specific photosynthesis parameters that are not tuned to local site level vegetation parameters and, for instance, do not take into account plant specia and age. Similar to the multi-model results in Schaefer et al. (2012), YIBs performance is weakest at the 5 grassland sites. In this case, the bias is mainly due to the delayed LAI seasonality in the MERRA satellite dataset (Figs. 1 and 2). In general, application of the remotely sensed LAI is a source of error because the gridded satellite data may not represent the local site changes in plant growth and phenology, especially for vegetation types with low biomass. The limitation of the satellite LAI spatial resolution implies that the model is unable to resolve GPP variability for sites in close proximity. For example, sites CA-SJ1, CA-SJ2, and CA-SJ3 are located close to each other. Simulations at these sites have similar magnitude in simulated GPP while observations show distinct variability between the sites.

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We compare  $R^2$ , RMSE, and  $\chi^2$  for the different sensitivity experiments in order to ascertain which combination of meteorological and LAI forcings best reproduces the measured GPP over North America (Table 3 and Fig. 4). CA-Let, CA-NS1, US-Var, CA-SJ1, and CA-SJ2 are excluded from the analysis because of the excessive bias at those sites (Fig. 3a). The average  $R^2$  increases while RMSE decreases when MERRA

reanalyses are substituted with site-based meteorology, or the MERRA LAI is used instead of MODIS LAI (Table 3). The choice of LAI forcing has the most significant impact on YIBs simulation performance consistent with previous work that demonstrated the dominance of phenology over meteorology in controlling local terrestrial carbon exchange (Desai et al., 2008; Puma et al., 2013). Using MODIS LAI, YIBs has a total  $\chi^2$  of 9.2 that shows an average reduction of 4.7 (52%) with MERRA LAI (Table 3 and Fig. 4). Applying the site meteorology relative to MERRA meteorological forcings offers smaller improvements. For example, the total  $\chi^2$  value decreases by 5% in METsite\_LAImodis compared with that in METmerra\_LAImodis and 15% in METsite LAImerra relative to that in METmerra (Table 3).

# 3.2 Evaluation of modeled surface [O<sub>3</sub>]

We validate summertime surface O<sub>3</sub> simulated by the NASA Model-E2 chemistry-climate model with observations from the CASTNET and AIRDATA (Fig. 5). High O<sub>3</sub> level appears in the eastern U.S. due to anthropogenic emissions and in the mountainous western U.S. due to high elevation. The model generally captures this spatial pattern with a correlation coefficient of 0.39 against observations over the selected 73 CASTNET sites (Figs. 5a-b). The simulation overestimates the  $O_3$  level by ~4 ppbv (12%) in the East and ~1 ppbv (3%) in the West. The CASTNET sites are located in rural sites, which usually have lower [O<sub>3</sub>] than that in urban areas, except for some megacities where the excessive NO<sub>x</sub> emissions result in lower O<sub>3</sub> level (Gregg et al., 2003). Therefore, we also compare the simulated MDA8 [O<sub>3</sub>] with monitored at ~1200 AIRDATA sites, which covers both urban and rural regions (Fig. 5c). In the eastern U.S., the model captures high [O<sub>3</sub>] centers around Michigan, Indiana, and Ohio states and that along the northeast coast. In the West, the simulation reproduces high  $[O_3]$  over mountain regions and in California. On average, the simulated MDA8  $[O_3]$  is lower by  $\sim 0.5$  ppbv (1%) in the East and  $\sim 3.5$ ppby (7%) in the West. The correlation coefficient between simulations and observations is as high as 0.51 (Fig. 5d).

# 3.3 O<sub>3</sub> damage effects at NACP sites

426 427 We must apply the simulated O<sub>3</sub> to quantify the O<sub>3</sub> vegetation damage at the NACP sites 428 because the sites do not monitor local  $[O_3]$ . The summer average  $[O_3]$  is 30-50 ppbv at 24 429 U.S. sites (Fig. 6a). The O<sub>3</sub> damage effect is relatively stronger at sites with both high O<sub>3</sub>-430 free GPP and ambient [O<sub>3</sub>] (Fig. 6d). The most significant damages are predicted at US-431 MMS (DBF) and US-Dk3 (ENF) sites, where the GPP reductions are 5-14% depending 432 on the low or high O<sub>3</sub> sensitivity (Fig. 6d). At these two sites, the high stomatal conductance (4.0 and 3.4 mm s<sup>-1</sup>, Fig. 6b) and ambient [O<sub>3</sub>] (both 43 ppby, Fig. 6a) result 433 in the largest O<sub>3</sub> stomatal flux (both ~0.3 mmol m<sup>-2</sup> d<sup>-1</sup>, Fig. 6c) among the 24 sites. The 434 435 lowest O<sub>3</sub> damage (1-2% GPP reduction) appears in the 3 shrub sites, US-Ton, US-SO<sub>2</sub>, 436 and US-Los, although mean [O<sub>3</sub>] there is as high as 43 ppbv. The main reason for the limited O<sub>3</sub> damage is the low stomatal conductance (average 1.4 mm s<sup>-1</sup>, Fig. 6b) related 437 to the small O<sub>3</sub>-free GPP (average 4.6 g C m<sup>-2</sup> d<sup>-1</sup>, Fig. 6d). Similarly, the O<sub>3</sub> damage for 438 439 C3 grass is as low as 1-2%, although the GPP of this plant is highly sensitive to O<sub>3</sub> (Table 440 1). For needleleaf and broadleaf forest sites, the average site-level O<sub>3</sub> damage effects are 441 estimated to be 2-5% and 3-9% respectively with differences between these ecosystem types predominantly driven by differences in sensitivity to O<sub>3</sub>. The four C4 crop sites, 442 US-Ne1, US-IB1, US-Ne2, and US-Ne3, exhibit the highest O<sub>3</sub>-free GPP but show only 443 444 moderate O<sub>3</sub> damage effects (GPP reductions of 4-6%, Fig. 6d). This result is driven by 445 low ambient [O<sub>3</sub>] at the C4 crop sites (average 32 ppbv, Fig. 6a) in combination with the 446 reduced C4 stomatal conductance (higher water use efficiency) relative to C3 plants (average 3.2 mm s<sup>-1</sup>, Fig. 6b). Indeed, the C4 photosynthetic pathway has been observed 447

Inclusion of  $O_3$  damage effect improves the site-level simulations (Figs. 3b-c). For 36 out of the 40 sites, the  $\chi^2$  of simulated GPP decreases when considering vegetation responses to  $O_3$ , and the improvement is better when higher  $O_3$  sensitivity is applied. At these sites, for example, CA-TP4, US-Dk3, US-MMS, and US-PFa, the reduced GPP at peak seasons is closer to measurements (Fig. S4), leading to smaller biases for simulations. On average, the  $\chi^2$  decreases by 3-8% at these sites, depending on the  $O_3$  sensitivity in the simulation (Fig. 4). Finally, the simulated annual GPP averaged over all NACP sites

to offer protection against O<sub>3</sub> damage (Heagle et al., 1989; Rudorff et al., 1996).

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changes from 3.8 g C m<sup>-2</sup> day<sup>-1</sup> to 3.6 g C m<sup>-2</sup> day<sup>-1</sup> with the high O<sub>3</sub> sensitivity simulation case, closer to the observations of 3.0 g C m<sup>-2</sup> day<sup>-1</sup>. The bias-correction from O<sub>3</sub> damage is much smaller relative to the effect of phenology (Fig. 4). Moreover, the O<sub>3</sub>-induced damage does not improve the GPP correlation between observations and simulations, which remains similar at ~0.8 (for 40 sites) with and without O<sub>3</sub> effects (Fig. S3).

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### 3.4 Evaluation of simulated O<sub>3</sub> vegetation damage against field and laboratory data

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We compare the simulated O<sub>3</sub> damage effect with field and laboratory measurements from the published literature (Fig. 7). In total, 14 additional sensitivity experiments are performed with different levels of [O<sub>3</sub>] at each NACP site (see section 2.2.1). GPP reductions increase accordingly with the increasing  $[O_3]$  (Fig. 7). For a given  $[O_3]$ , the  $O_3$ damage effect is strongest for C4 crops (despite the lower  $g_s$ : $A_{net}$  ratio) but weakest for shrubland. YIBs simulates reasonable O<sub>3</sub> damage to GPP for all model PFTs compared to the meta-analyses of Wittig et al. (2007) and Lombardozzi et al. (2013). Field studies in shrubland are limited. Zhang et al. (2012) investigated the responses of four shrub species to  $[O_3]=70$  ppbv and found large reductions in net photosynthesis of 50-60%. The average  $O_3$ -free  $A_{net}$  of those shrub species was 8-16 g [C] m<sup>-2</sup> s<sup>-1</sup>, much higher than even the gross photosynthesis (A) of 6 g [C] m<sup>-2</sup> s<sup>-1</sup> at the shrub NACP sites, likely because the latter are located in dry and/or cold areas (Fig. S1). The YIBs simulated O<sub>3</sub> vegetation damage effects for C4 plants are in good agreement with field measurements from Taylor et al. (2002) and Grantz et al. (2012). In the case of C3 grass and C3 crop, the model simulates consistent GPP reduction percentages with observations from Feng et al. (2008) for wheat, Foot et al. (1996) for colluna vulgaris, and Mulchi et al. (1992) for soybean. However, these O<sub>3</sub> damage results are all >50% less than for available measurements in rice crops (Ishii et al., 2004; Ainsworth et al., 2008), suggesting that rice may have much higher O<sub>3</sub> sensitivity than other C3 plants. In the U.S. rice plantation area is much smaller than that of soybean and corn. Therefore, we adopt the O<sub>3</sub> sensitivity parameters for C3/C4 plants shown in Table 1 for the distributed regional simulations.

## 3.5 O<sub>3</sub> vegetation damage effect on GPP in U.S. region

High values of simulated summertime GPP (including O<sub>3</sub> damage effect) appear east of 95°W in the U.S. (Fig. 8a), because the land surface there is covered by crops and forests. A high center of GPP (> 10 g C m<sup>-2</sup> day<sup>-1</sup>) appears over cropland in the north central U.S. In the western U.S., the coverage of grass and shrub and the low water availability (low precipitation and soil moisture) over semi-arid regions lead to low carbon assimilation rate. The regional gridded simulated GPP reproduces the JJA growing season average NACP site-level fluxes with a correlation coefficient of 0.62 for 32 sites below 50°N (Fig. 8a). The correlation is lower than the 0.84 estimated for the site-level simulation METsite LAImerra at the same sites and the same season. Since the meteorological forcings and LAI are similar, the difference in land cover, ISLSCP versus site definitions (Fig. S1 and Fig. S2), accounts for the discrepancy between regional and site-level

simulations.

On average, the simulated summer GPP (including the high  $O_3$  damage effect) is 9.5 g C m<sup>-2</sup> day<sup>-1</sup> in the eastern U.S. and 3.9 g C m<sup>-2</sup> day<sup>-1</sup> in the western U.S., giving a mean value of 6.1 g C m<sup>-2</sup> day<sup>-1</sup> for the U.S. region. The total carbon uptake is estimated to be  $4.43 \pm 0.18$  Pg C during the summer growing season, accounting for 57-60% of the annual average value of  $7.59 \pm 0.25$  Pg C over the 1998-2007 period. Our estimate of annual carbon uptake is consistent with previous published estimates. For example, Xiao et al. (2010) upscaled site-level GPP flux to continental scale with a regression tree approach based on both NACP fluxes and remote-sensing variables. They estimated that the total GPP in U.S. ranges from 6.91 to 7.33 Pg C per year during 2000-2006. Using the same observations but with a process-based biogeochemical model, Chen et al. (2011) estimated a range of 7.02-7.78 Pg C per year for 2000-2005, which is even closer to our estimate.

We calculate both O<sub>3</sub> stomatal flux (Fig. 8b) and the resultant damage on GPP (Fig. 9) in the U.S. region for the 1998-2007 period. High O<sub>3</sub> stomatal flux is predicted in the eastern U.S. due to co-location of the high GPP (medium to high stomatal conductance)

and the substantial ambient [O<sub>3</sub>]. On average, the summertime O<sub>3</sub> plant uptake is 117 µmol m<sup>-2</sup> day<sup>-1</sup>, with 207 µmol m<sup>-2</sup> day<sup>-1</sup> in the eastern U.S. and 59 µmol m<sup>-2</sup> day<sup>-1</sup> in the western U.S. Following the O<sub>3</sub> stomatal flux, the largest mean GPP reductions are predicted for the eastern U.S. growing season, in the range 4-8% depending on the O<sub>3</sub> sensitivity applied in the simulations (Fig. 9). Locally, reduction fraction reaches as high as 11-17% in areas with high [O<sub>3</sub>] pollution, such as Michigan, Indiana, Ohio, and states along the northeast coast. Despite high surface [O<sub>3</sub>] over mountainous elevated areas in the West (Fig. 5), impacts on GPP are limited due to the low stomatal conductance and low photosynthetic rate there. The Pacific northwestern forests are an exception, with a moderate GPP reduction of 1-7%. On average, the total summer GPP is reduced by 2-5% due to O<sub>3</sub> damage effects in the U.S. Similar reduction fractions are predicted for the annual GPP.

U.S. surface  $[O_3]$  exhibits a decreasing trend over the past 2 decades, especially in the eastern U.S., due to precursor emission controls (Lefohn et al., 2010). However, the community continues to debate how surface  $[O_3]$  will respond to future emissions and climate change. On the one hand, surface  $[O_3]$  may decline by the mid  $21^{st}$  century due to large reductions in regional anthropogenic precursor emissions (Wu et al., 2008). On the other hand, climate change effects alone may increase local surface  $[O_3]$  due to the warmer, drier, and more stable environment (Leibensperger et al., 2008; Wu et al., 2008). Due to the uncertainty in future surface  $[O_3]$  projections, our strategy here is to perform four additional sensitivity experiments with  $\pm 25\%$  changes in  $[O_3]$  for each  $O_3$  sensitivity case. Increases of 25% in  $[O_3]$  may reduce GPP in the eastern U.S. by 6-11%, with a maximum local reduction of 25% for the high  $O_3$  sensitivity case (Fig. 10d). The damage magnitude with low  $O_3$  sensitivity (Fig. 10b) mimics the present-day estimate with high  $O_3$  sensitivity (Fig. 9b). In contrast, the  $O_3$  damage to the eastern U.S. GPP is as low as 2-4% in response to 25% decreases in  $[O_3]$  (Figs. 10a, c), suggesting a substantial cobenefit to ecosystem-health of  $O_3$  precursor emissions control.

## 4. Discussion and Conclusions

We have performed an updated assessment of O<sub>3</sub> vegetation damage effects on GPP in the U.S. for the 1998-2007 period using the YIBs vegetation model. The semimechanistic parameterization of O<sub>3</sub> inhibition on photosynthesis proposed by Sitch et al. (2007) has been implemented into this process-based vegetation model. The simulated O<sub>3</sub> damage effects are consistent with laboratory and field measurements reported in previously published studies. We evaluated the simulated O<sub>3</sub>-free and O<sub>3</sub>-damaged GPP with in situ measurements from 40 NACP sites. The O<sub>3</sub>-free and O<sub>3</sub>-damaged GPP simulations capture the seasonality and interannual variability of GPP at most sites. The model GPP biases are lowest at forest and cropland sites but highest at grassland sites. Model GPP is highly sensitive to choice of LAI forcing. Simulations that apply MERRA LAI generally perform better (show lower biases) than those with MODIS LAI. In response to the simulated ambient [O<sub>3</sub>] of 30-50 ppbv, simulated GPP decreases by 1-14% at the NACP sites, depending on the O<sub>3</sub> sensitivity and PFT types. Maximum reductions of 5-14% occur in two forest sites, where both O<sub>3</sub>-free GPP and ambient [O<sub>3</sub>] are relatively high. Inclusion of the O<sub>3</sub> damage offers only a small improvement to the simulated annual average GPP at NACP sites (from 3.8 g C m<sup>-2</sup> day<sup>-1</sup> to 3.6 g C m<sup>-2</sup> day<sup>-1</sup> 1) such that the model still overestimates the observational average of 3.0 g C m<sup>-2</sup> day<sup>-1</sup>. The model GPP overestimate is most likely related to the use of generic PFT-specific photosynthesis parameters and the satellite prescribed LAI that may not represent the local site LAI. In this work, we assumed a coupled response between photosynthesis and stomatal conductance. Emerging research has found that the O<sub>3</sub> vegetation damage effects can actually result in a loss of plant stomatal control, and a consequent decoupling of the stomatal response from photosynthesis inhibition (Lombardozzi et al., 2012a, b). Treatment of this decoupled response in the YIBs model would lead to a higher level of O<sub>3</sub> flux entering leaves, thus causing stronger damage. Interestingly, this mechanism would therefore provide a way to improve the simulated GPP overestimates. That said, other studies have suggested that the O<sub>3</sub> damage effect is limited by carbon-nitrogen interactions (Ollinger et al., 2002; Kvalevag and Myhre, 2013).

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Regional distributed simulations for the U.S. yield a summertime GPP (with high sensitivity O<sub>3</sub> damage) of 6.1 g C m<sup>-2</sup> day<sup>-1</sup> (9.5 g C m<sup>-2</sup> day<sup>-1</sup> in the eastern U.S. and 3.9

g C m<sup>-2</sup> day<sup>-1</sup> in the western U.S.). The total carbon uptake was estimated to be 4.43  $\pm$ 0.18 Pg C for the summer, accounting for 57-60% of the annual value of  $7.59 \pm 0.25$  Pg C over the 1998-2007 period. Carbon assimilation rate is suppressed by 4-8% on average in the summertime eastern U.S. with maximum local damage of 11-17% in states close to the Great Lakes and along the eastern coast. When [O<sub>3</sub>] is decreased by 25%, O<sub>3</sub> damage to GPP is only 2-4% in the eastern U.S., indicating substantial improvements to vegetation health and carbon assimilation rate. Previously, Felzer et al. (2004) found annual average O<sub>3</sub>-induced NPP reductions of 3-7% over the U.S. for 1989-1993 and simulated the largest reductions in states close to the Great Lakes and along the East Coast, where the high O<sub>3</sub> sensitivity of crops makes the dominant contribution. Our study examined O<sub>3</sub> damage effects a decade later than Felzer et al. (2004) but gives consistent results. Qualitatively, this consistency between decades may be explained by the offsetting influences of (i) surface O<sub>3</sub> reductions due to air quality control legislation and (ii) GPP increases due to CO<sub>2</sub>-fertilization and rising temperatures. Felzer et al. (2004) estimated a maximum local NPP reduction of 34%, which is double the maximum of 17% in our analyses. Furthermore, Felzer et al. (2004) found widespread reductions of > 6% in the Midwest where there is almost no O<sub>3</sub> impact in this study (Fig. 9). Differences between the studies are mostly likely driven by the use of different vegetation cover and LAI datasets, and the use of a semi-mechanistic flux-based uptake in this study versus the concentration-based uptake method elsewhere.

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The current work has used an off-line approach. Yet, the O<sub>3</sub>-vegetation-meteorology system is strongly coupled. For instance, plant productivity itself controls the emission of isoprene, a major O<sub>3</sub> precursor. The O<sub>3</sub>-induced modification to stomatal conductance may inhibit evapotranspiration, leading to changes in canopy temperature, precipitation, soil moisture, and other surface hydrology and meteorology (Bernacchi et al., 2007; vanLoocke et al., 2012). In future work, we will study O<sub>3</sub> vegetation damage effects using YIBs embedded within a fully coupled global chemistry-climate model framework in order to account for these feedbacks including altered canopy energy fluxes and partitioning between latent and sensible heat that drive regional climate and hydrology. In addition, future work will exploit recent extensive meta-data analyses (Lombardozzi et

612 al., 2013; Wittig et al., 2007) to refine the ozone damage parameterization in YIBs 613 including the decoupled modification of photosynthesis and stomatal conductance. 614 Acknowledgements 615 The authors are grateful to B. Felzer and an anonymous reviewer for assistance in 616 617 evaluating this paper. Data for this analysis was provided by the North American Carbon 618 Program Site Synthesis. Funding for this research was provided by Yale University. This project was supported in part by the facilities and staff of the Yale University Faculty of 619 Arts and Sciences High Performance Computing Center. 620

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Table 1. Parameters for vegetation model and  $\mathrm{O}_3$  damage scheme

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PFT <sup>a</sup>	TDA	GRAC3	GRAC4	SHR	DBF	ENF	TRF	CI	RO
Carboxylation	C3	C3	C4	С3	C3	C3	C3	C3 <sup>b</sup>	C4 <sup>b</sup>
$V_{max 25}$ (µmol m <sup>-2</sup> s <sup>-1</sup> )	33	43	24	25	30	43	75	40	40
m	9	11	5	9	9	9	9	11	5
<i>b</i> (mmol m <sup>-2</sup> s <sup>-1</sup> )	2	8	2	2	2	2	2	8	2
O3T (nmol m <sup>-2</sup> s <sup>-1</sup> )	1.6	5	5	1.6	1.6	1.6	1.6	5	5
a  (high)  (mmol <sup>-1</sup> m <sup>-2</sup> )	0.1	1.4	0.735	0.1	0.15	0.075	0.15	1.4	0.735
<i>a</i> (low) (mmol <sup>-1</sup> m <sup>-2</sup> )	0.03	0.25	0.13	0.03	0.04	0.02	0.04	0.25	0.13

<sup>&</sup>lt;sup>a</sup> Plant function types (PFTs) are tundra (TDA), C3 grassland (GRAC3), C4 savanna/grassland (GRAC4), shrubland (SHR), deciduous broadleaf forest (DBF), evergreen needleleaf forest (ENF), tropical rainforest (TRF), and cropland (CRO).

<sup>&</sup>lt;sup>b</sup> For site-level simulations, we consider CRO as C4 plant. For regional simulation, we consider half CRO as C3 plants (soybean) and the rest C4 plant (corn).

**Table 2.** Description of the site-level simulations.

		Meteorology		Vegetation (LAI)		
ID	Simulations <sup>a</sup>	Site	MERRA	MODIS	MERRA	Incl. O <sub>3</sub> <sup>c</sup>
1	METmerra_LAImodis		Yes	Yes		
2	METsite_LAImodis <sup>b</sup>	Yes	Yes	Yes		
3	METmerra_LAImerra		Yes		Yes	
4	METsite_LAImerra <sup>b</sup>	Yes	Yes		Yes	
5	METsite_LAImerra_LO3 b	Yes	Yes		Yes	Low d
6	METsite_LAImerra_HO3 b	Yes	Yes		Yes	High <sup>d</sup>

<sup>&</sup>lt;sup>a</sup> The name of each simulation is composed of at least two words. The prefix indicates the source of meteorological forcings. The suffix or the second word indicates the sources of vegetation forcings.

<sup>&</sup>lt;sup>b</sup> For simulations with prefix 'SITE' use site-based meteorological forcings, which are gap-filled with MERRA-land reanalyses.

<sup>&</sup>lt;sup>c</sup> Ambient [O<sub>3</sub>] is applied at each site.

Low and high indicate the sensitivity of GPP to  $[O_3]$  defined by the coefficient a in Table 1.

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**Table 3.** Statistics <sup>a</sup> for site-level simulations.

				$R^2$		RMSE			$\chi^2$		
I	ID	Simulations	min	max	mean	min	max	mean	min	max	total
1	1	METmerra_LAImodis	0.14	0.85	0.63	1.2	6.4	3.0	1.4	37.6	9.4
2	2	METsite_LAImodis	0.16	0.85	0.67	1.1	5.8	2.9	1.1	39.6	8.9
3	3	METmerra_LAImerra	0	0.88	0.66	1.3	4.2	2.4	1.2	16.6	4.8
4	4	METsite_LAImerra	0	0.87	0.68	1.0	4.2	2.3	1.2	13.4	4.1
5	5	METsite_LAImerra_LO3	0	0.88	0.69	1.0	4.1	2.3	1.1	13.0	3.9
6	6	METsite_LAImerra_HO3	0	0.88	0.69	1.0	4.1	2.2	1.0	12.3	3.7
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919 a Statistics include minimum and maximum values of  $R^2$ , RMSE, and  $\chi^2$  for 35 NCAP 920 sites with  $\chi^2 < 16$  (Fig. 3a). We also calculate the mean values of  $R^2$  and RMSE for these 921 sites. We calculate the total  $\chi^2$  (shown as red bars in Fig. 4) using all the available 922 observations over all sites.

925 Appendix
926 Table A. Descriptions of NACP sites in Canada <sup>a</sup>

Site	PFT <sup>b</sup>	Description	Longitude	Latitude	Period
CA-Ca1	ENF	Campbell River	125.3°W	49.9°N	1998-2006
CA-Ca2	ENF	Campbell River	125.3°W	49.9°N	2001-2006
CA-Ca3	ENF	Campbell River	124.9°W	49.5°N	2002-2006
CA-Gro	MF	Groundhog River	82.2°W	48.2°N	2004-2006
CA-Let	GRA	Lethbridge Grassland	112.9°W	49.7°N	2001-2007
CA-Mer	WET	Eastern Peatland	75.5°W	45.4°N	1999-2006
CA-NS1	ENF	UCI Chronosequence	124.9°W	49.5°N	2001-2005
CA-Oas	DBF	BERMS	106.2°W	53.6°N	1997-2006
CA-Obs	ENF	BERMS	105.1°W	54.0°N	2000-2006
CA-Ojp	ENF	BERMS	104.7°W	53.9°N	2000-2006
CA-Qfo	ENF	Quebec	74.3°W	49.7°N	2004-2006
CA-SJ1	ENF	BERMS	104.7°W	53.9°N	2002-2005
CA-SJ2	ENF	BERMS	104.6°W	53.9°N	2004-2006
CA-SJ3	ENF	BERMS	104.6°W	53.9°N	2005-2006
CA-TP4	ENF	Turkey Point	80.4°W	42.7°N	2003-2007
CA-WP1	WET	Western Peatland	112.5°W	55.0°N	2004-2007
<b>US-ARM</b>	GRA c	Southern Great Plains	97.5°W	36.6°N	2003-2007
US-Dk3	ENF	Duke Forest	79.1°W	36.0°N	1998-2005
US-Ha1	DBF	Harvard Forest	72.2°W	42.5°N	1992-2006
US-Ho1	ENF	Howland Forest	68.7°W	45.2°N	1996-2004
US-IB1	CRO	Fermi Lab	88.2°W	41.9°N	2006
US-IB2	GRA	Fermi	88.2°W	41.8°N	2005-2006
US-Los	WET	Lost Creek	90.0°W	46.1°N	2001-2006
US-MMS	DBF	Morgan Monroe State Forest	86.4°W	39.3°N	1999-2006
US-MOz	DBF	Missouri Ozark	92.2°W	38.7°N	2005-2007
US-Me2	ENF	Metolius	121.6°W	44.5°N	2002-2007
US-Me3	ENF	Metolius	121.6°W	44.3°N	2004-2005
US-Me5	ENF	Metolius	121.6°W	44.4°N	2000-2002
US-NR1	ENF	Niwot Ridge	105.5°W	40.0°N	1999-2007
US-Ne1	CRO	Mead	96.5°W	41.2°N	2002-2005
US-Ne2	CRO	Mead	96.5°W	41.2°N	2003-2005
US-Ne3	CRO	Mead	96.4°W	41.2°N	2002-2005
US-Pfa	MF	Park Falls	90.3°W	45.9°N	1997-2004
US-SO2	CSH	Sky Oaks	116.6°W	33.4°N	1999-2006
US-Shd	GRA	Shidler	96.7°W	36.9°N	1998-1999
US-Syv	MF	Sylvania Wilderness Area	89.3°W	46.2°N	2002-2006
US-Ton	WSA	Tonzi Ranch	121.0°W	38.4°N	2002-2007
US-UMB	DBF	UMBS	84.7°W	45.6°N	1999-2006
US-Var	GRA	Varia Ranch	121.0°W	38.4°N	2001-2007
US-WCr	DBF	Willow Creek	90.1°W	45.8°N	1999-2006

<sup>&</sup>lt;sup>a</sup> Site information is adopted from Schaefer et al. (2012), except that the operational time span listed here is only for the period when measurements of GPP are available.

<sup>&</sup>lt;sup>b</sup> PFT names are: evergreen needleleaf forest (ENF), deciduous broadleaf forest (DBF), grasslands (GRA), croplands (CRO), closed shrublands (CSH), mixed forests (MF), permanent wetlands (WET), and woody savannas (WSA).

<sup>c</sup> The land type at US-ARM is cropland in Schaefer et al. (2012). However, the site is covered by cattle pasture and wheat fields (<a href="https://www.arm.gov/sites/sgp">https://www.arm.gov/sites/sgp</a>), which are more like C3 grassland.

# Figure Captions

- 940 Fig. 1. Comparison of monthly mean leaf area index (LAI, m<sup>2</sup> m<sup>-2</sup>) from the Modern Era-
- 941 Retrospective Analysis (MERRA) reanalysis (red solid lines) and the Moderate
- Resolution Imaging Spectroradiometer (MODIS) (blue dashed lines) at each NACP site.
- The name and location (longitude, latitude) are shown in the title. The land types include
- evergreen needleleaf forest (ENF), deciduous broadleaf forest (DBF), shrublands (SHR),
- grasslands (GRA), and croplands (CRO).

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- 947 **Fig. 2.** Comparison of long-term monthly mean GPP (g [C] m<sup>-2</sup> day<sup>-1</sup>) from observations
- 948 (blue points with error bars indicating one standard deviation) and the O<sub>3</sub>-free simulations
- 949 (black lines with shadings indicating one standard deviation) averaged over measurement
- 950 period at each site. The simulation, METsite\_LAImerra, is driven with meteorological
- 951 forcings from MERRA and site measurements. The LAI is from MERRA. The name and
- location (longitude, latitude) are shown in the title. The land types include evergreen
- 953 needleleaf forest (ENF), deciduous broadleaf forest (DBF), shrublands (SHR), grasslands
- 954 (GRA), and croplands (CRO).

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- 956 Fig. 3. The calculated (a)  $\chi^2$  of GPP and changes in  $\chi^2$  after the inclusion of  $O_3$  damage
- 957 impact with (b) low and (c) high O<sub>3</sub> sensitivity at each site. The sites are sorted according
- 958 to the values of  $\chi^2$  in (a). The land cover definitions are: GRA, Grasslands; CRO,
- 959 Croplands; ENF, Evergreen Needleleaf Forest; DBF, Deciduous Broadleaf Forest; SHR,
- Shrubland. See section 2.2.1 for the definition of  $\chi^2$ .

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- 962 Fig. 4. The calculated average  $\chi^2$  of GPP over NACP sites for 6 different simulations as
- listed in Table 2. The blue bars are results for all 40 NACP sites. The red bars are results
- 964 excluding sites CA-Let, CA-NS1, US-Var, CA-SJ1, and CA-SJ2, where the simulated
- 965 site-level  $\chi^2$  is larger than 16 as shown in Fig. 3a.

- 967 Fig. 5. Validation of simulated summertime surface (a, b) diurnal mean and (c, d) daily
- 968 maximum 8-hour average O<sub>3</sub> with in situ measurements from (a, b) the EPA Clean Air
- 969 Status and Trends Network (CASTNET) and (c, d) the AIRDATA. For (b) and (d), the

blue points indicate sites East of 95°W and the red ones for West of 95°W. The correlation coefficients between simulations and observations are shown in (b) and (d).

Please refer to Fig. S5 for separate results from GCM and observations.

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**Fig. 6.** Simulated summertime (a) surface [O<sub>3</sub>], (b) stomatal conductance, (c) ozone stomatal flux, and (d) damages to GPP at different O<sub>3</sub> sensitivity for 24 U.S. sites. The sites are sorted according to the simulated O<sub>3</sub>-free GPP in (d). For each site, the result is averaged over the period when the site provides GPP measurements. The land cover definitions are: GRA, Grasslands; CRO, Croplands; ENF, Evergreen Needleleaf Forest; DBF, Deciduous Broadleaf Forest; SHR, Shrubland.

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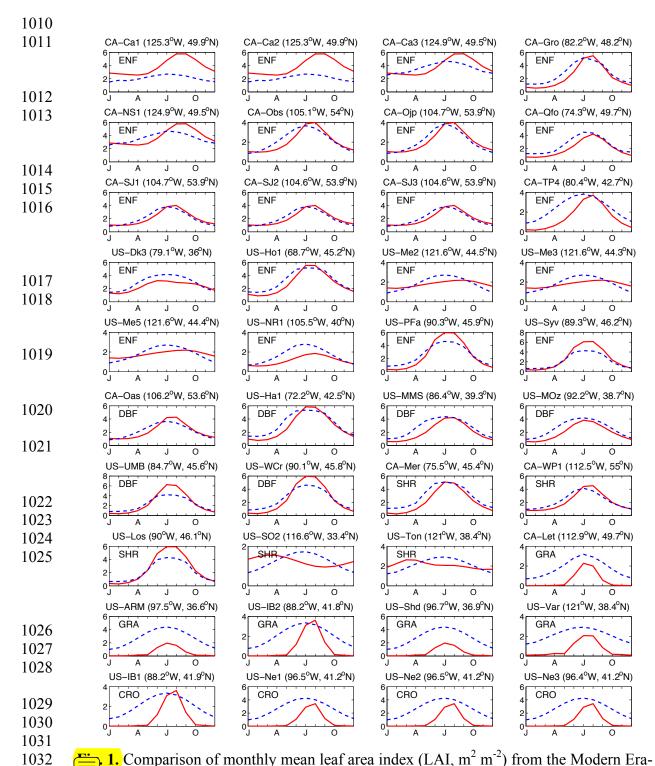
Fig. 7. Changes in GPP for all and individual PFTs in the presence of different levels of [O<sub>3</sub>] as simulated by the vegetation model. Simulations are performed at 40 NACP sites with a fixed  $[O_3]$  for either low or high  $O_3$  sensitivity. The short blue lines show the damages ranging from low to high O<sub>3</sub> sensitivity, with the blue points indicating the average reductions. The simulation results are averaged for all the sites or for the sites with the same PFT. The number of sites used for average is shown in the title bracket of each subplot. The solid squares with lines show the results (mean plus uncertainty) based on measurements reported by multiple literatures. These measurements are collected by Lombardozzi et al. (2013) for all PFTs, Wittig et al. (2007) for evergreen needleleaf forest (ENF) and deciduous broadleaf forest (DBF). For C4 grass or crop (CRO C4), we have collected data from Taylor et al. (2002) for spartina alterniflora and Grantz et al. (2012) for sugarcane hybrids. For C3 grass or crop (GRA C3), we have collected data from Feng et al. (2008) for wheat, Foot et al. (1996) for colluna vulgaris, Mulchi et al. (1992) for soybean, and Ishii et al. (2004) and Ainsworth (2008) for rice. The values for rice are denoted in green and others in red. The author initials are indicated for the corresponding studies.

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**Fig. 8.** Simulated summertime (a) O<sub>3</sub>-exposed GPP and (b) O<sub>3</sub> stomatal flux over the U.S. The simulated GPP is overlaid with *in situ* measurements from NACP. The simulations

are performed with land cover from ISLSCP and meteorological forcings from MERRA reanalyses. Please refer to Fig. S6 for separate results from simulations and observations. Fig. 9. Simulated reduction fraction in summer GPP in the U.S. due to (a) low and (b) high O<sub>3</sub> sensitivity for 1998-2007. Fig. 10. Simulated changes in summer GPP due to (a, c) 25% reduction or (b, d) 25% increase in [O<sub>3</sub>] for (a, b) low or (c, d) high O<sub>3</sub> sensitivity. 



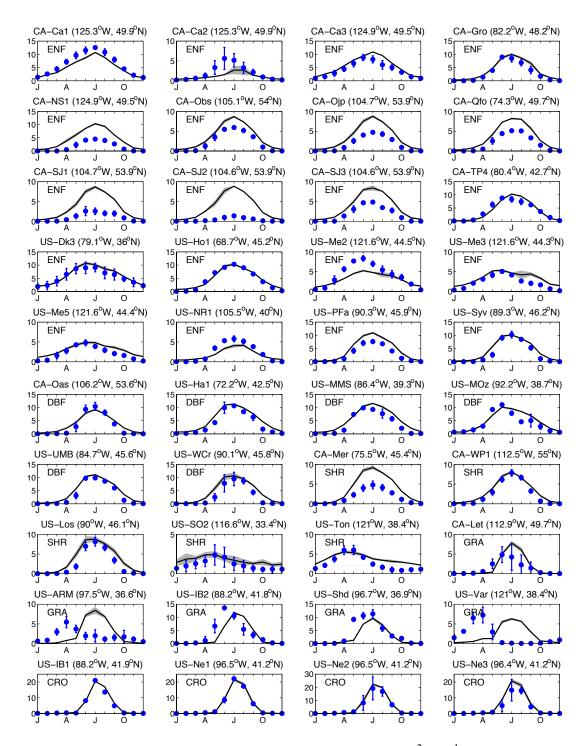
1. Comparison of monthly mean leaf area index (LAI, m<sup>2</sup> m<sup>-2</sup>) from the Modern Era-Ketrospective Analysis (MERRA) reanalysis (red solid lines) and the Moderate Resolution Imaging Spectroradiometer (MODIS) (blue dashed lines) at each NACP site. The name and location (longitude, latitude) are shown in the title. The land types include evergreen needleleaf forest (ENF), deciduous broadleaf forest (DBF), shrublands (SHR), grasslands (GRA), and croplands (CRO).

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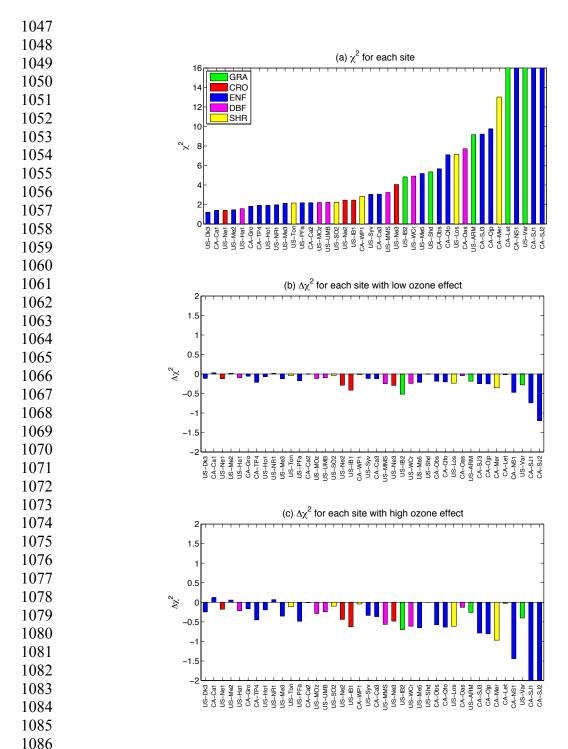
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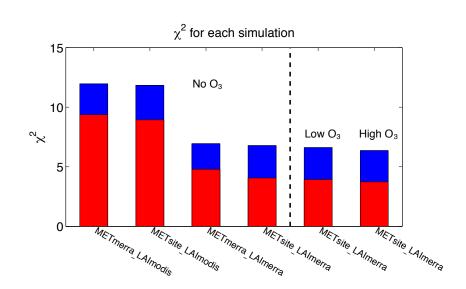
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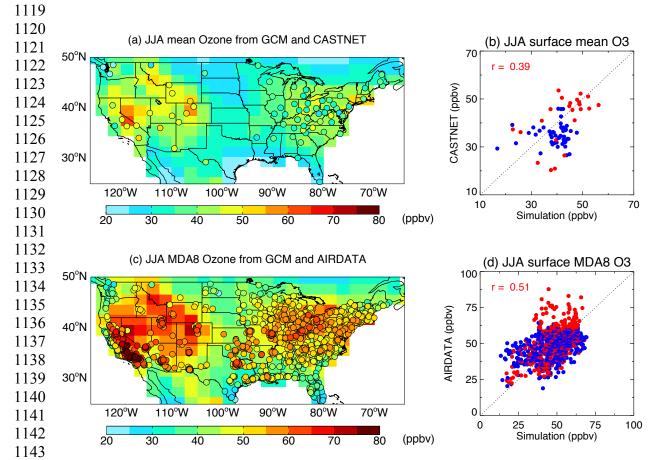
**Fig. 2.** Comparison of long-term monthly mean GPP (g [C] m<sup>-2</sup> day<sup>-1</sup>) from observations (blue points with error bars indicating one standard deviation) and the O<sub>3</sub>-free simulations (black lines with shadings indicating one standard deviation) averaged over measurement period at each site. The simulation, METsite\_LAImerra, is driven with meteorological forcings from MERRA and site measurements. The LAI is from MERRA. The name and location (longitude, latitude) are shown in the title. The land types include evergreen needleleaf forest (ENF), deciduous broadleaf forest (DBF), shrublands (SHR), grasslands (GRA), and croplands (CRO).



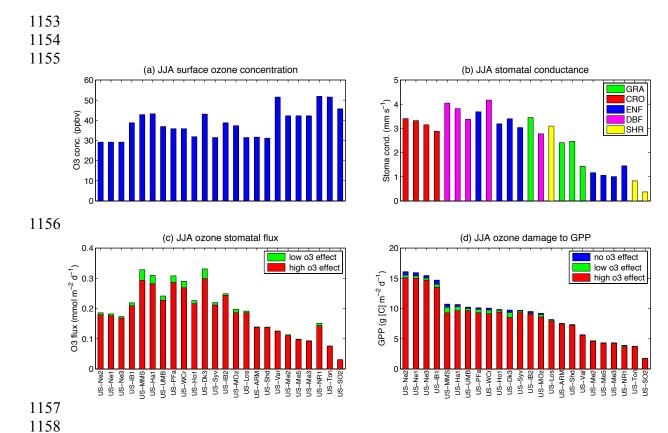
**Fig. 3.** The calculated (a)  $\chi^2$  of GPP and changes in  $\chi^2$  after the inclusion of O<sub>3</sub> damage impact with (b) low and (c) high O<sub>3</sub> sensitivity at each site. The sites are sorted according to the values of  $\chi^2$  in (a). The land cover definitions are: GRA, Grasslands; CRO, Croplands; ENF, Evergreen Needleleaf Forest; DBF, Deciduous Broadleaf Forest; SHR, Shrubland. See section 2.2.1 for the definition of  $\chi^2$ .



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**Fig. 5.** Validation of simulated summertime surface (a, b) diurnal mean and (c, d) daily maximum 8-hour average O<sub>3</sub> with *in situ* measurements from (a, b) the EPA Clean Air Status and Trends Network (CASTNET) and (c, d) the AIRDATA. For (b) and (d), the blue points indicate sites East of 95°W and the red ones for West of 95°W. The correlation coefficients between simulations and observations are shown in (b) and (d). Please refer to Fig. S5 for separate results from GCM and observations.



**Fig. 6.** Simulated summertime (a) surface [O<sub>3</sub>], (b) stomatal conductance, (c) ozone stomatal flux, and (d) damages to GPP at different O<sub>3</sub> sensitivity for 24 U.S. sites. The sites are sorted according to the simulated O<sub>3</sub>-free GPP in (d). For each site, the result is averaged over the period when the site provides GPP measurements. The land cover definitions are: GRA, Grasslands; CRO, Croplands; ENF, Evergreen Needleleaf Forest; DBF, Deciduous Broadleaf Forest; SHR, Shrubland.

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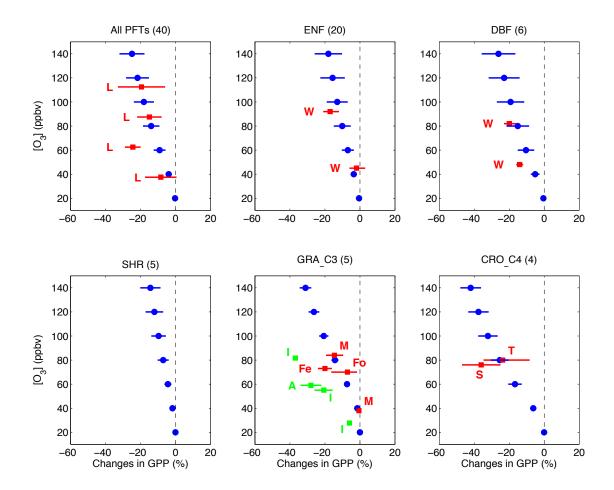
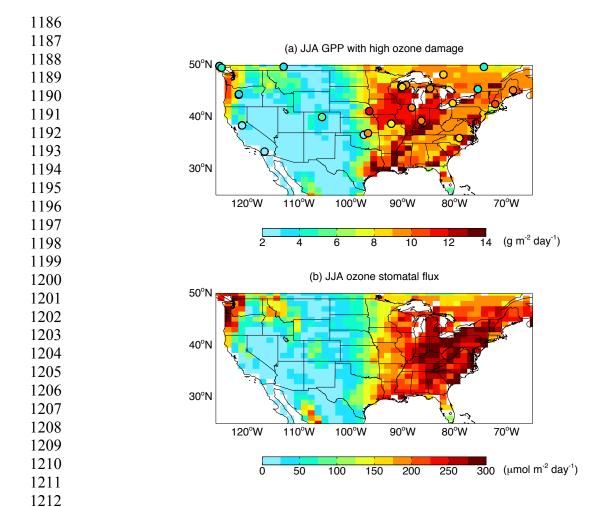
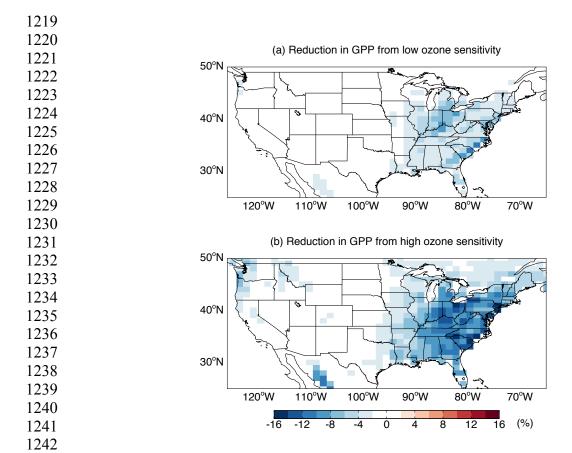


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**Fig. 9.** Simulated reduction fraction in summer GPP in the U.S. due to (a) low and (b) high  $O_3$  sensitivity for 1998-2007.

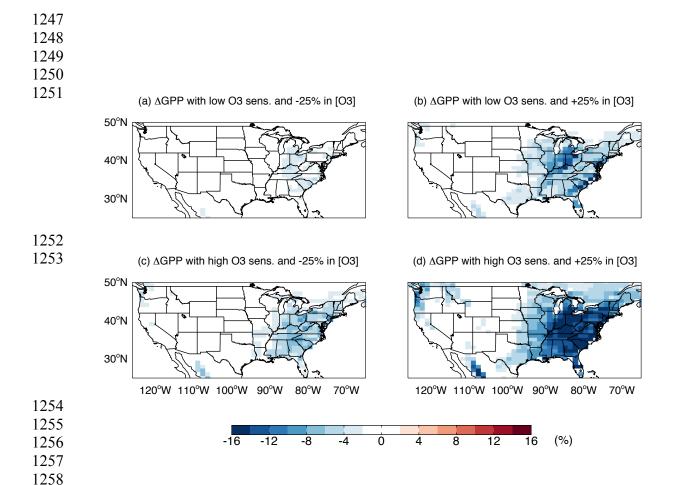


Fig. 10. Simulated changes in summer GPP due to (a, c) 25% reduction or (b, d) 25% increase in  $[O_3]$  for (a, b) low or (c, d) high  $O_3$  sensitivity.