

# 1 **Modeling analysis of the seasonal characteristics of haze formation in** 2 **Beijing**

3 X. Han<sup>1</sup>, M. Zhang<sup>1,\*</sup>, J. Gao<sup>2</sup>, S. Wang<sup>2</sup>, and F. Chai<sup>2</sup>

4 <sup>1</sup> State Key Laboratory of Atmospheric Boundary Layer Physics and Atmospheric Chemistry, Institute of  
5 Atmospheric Physics, Chinese Academy of Sciences, Beijing, China

6 <sup>2</sup> Chinese Research Academy of Environmental Sciences, Beijing, China

7  
8 \* *Correspondence to:* M. Zhang (mgzhang@mail.iap.ac.cn)

9  
10 **Abstract.** The air quality modeling system RAMS-CMAQ, coupled with an aerosol optical property  
11 scheme, was applied to simulate the meteorological field, major aerosol components (sulfate, nitrate,  
12 ammonium, black carbon, organic carbon, dust, and sea salt), and surface visibility over the North China  
13 Plain (NCP) in 2011. The modeled results in February and July 2011 were selected and analyzed to obtain  
14 an in-depth understanding of the haze formation mechanism in Beijing for different seasons. The  
15 simulation results showed that the visibility was below 10 km for most regions of the NCP, and dropped  
16 to less than 5 km over the megacities of Beijing and Tianjin, the whole of Hebei province, and the  
17 northwest part of Shandong province during pollution episodes in February and July. The heavy mass  
18 concentration of PM<sub>2.5</sub> ranged from 120  $\mu\text{g m}^{-3}$  to 300  $\mu\text{g m}^{-3}$ , and was concentrated in the areas with low  
19 visibility. The haze formation mechanism in Beijing in winter was different from that in summer. The  
20 mass concentration of PM<sub>2.5</sub> was higher, and the components more complicated in winter. While the mass  
21 concentration of PM<sub>2.5</sub> in summer was lower than that in winter, the mass concentrations of hygroscopic  
22 inorganic salts were comparable with those in winter, and the relative humidity was, as expected, higher.  
23 Therefore, the water uptake of hygroscopic aerosols played a key role in summer. Moreover, the analysis  
24 showed that the influence of the PM<sub>2.5</sub> mass burden on visibility was very weak when its value was larger  
25 than 100  $\mu\text{g m}^{-3}$ . Only when the mass burden of PM<sub>2.5</sub> decreased to a certain threshold interval did the  
26 visibility increase rapidly. This indicates that when emission reduction measures are taken to control haze  
27 occurrence, the mass burden of PM<sub>2.5</sub> must be cut to below this threshold interval. The relationship  
28 between the threshold of haze occurrence and the relative humidity in Beijing was fitted by an

29 exponential function, and the resulting fitting curves could provide a new theoretical basis to understand  
30 and control haze formation in Beijing.

## 32 **1. Introduction**

33 The emission of air pollutants have increased significantly because of the economic growth, rapid  
34 population expansion, and urbanization in the North China Plain (NCP). Beijing has a population of over  
35 20 million and is the political, economic, and cultural center of China, of which it is also the capital. This  
36 megacity is located at the northern tip of the NCP and surrounded by high mountains in its northern and  
37 western boundaries. Beijing has suffered from air quality deterioration in the past decade because of  
38 strong local emissions (Sun et al., 2006) and long-range transport from the surrounding urban areas  
39 (Zhang et al., 2012) to the east and south of the NCP: such areas include Tianjin, Shijiazhuang, and a  
40 number of cities where economic development is most active in Hebei Province. The air pollution in  
41 Beijing is easily aggravated by its special geographic position, when stable weather appears or the south  
42 wind dominates. Although the SO<sub>2</sub> emission in Beijing in the last five years has been decreased by  
43 various measures prescribed by the current legislation on emission controls (Lu et al., 2010; Zhang et al.,  
44 2006), the mass burden of particulate pollutants remains at a high level (Hao et al., 2013; Zhang et al.,  
45 2013), causing serious environmental issues and associated health effects.

46 Atmospheric haze is caused by visibility deterioration to less than 10 km when the relative humidity  
47 does not exceed 90% (Wu et al., 2007) through light extinction by aerosol particles. As a result of the high  
48 level of aerosol loadings, wide spread haze cloud caused by serious air pollution has occurred more  
49 frequently over this region in the past decade (Ma et al., 2010; Tao et al., 2012; Wang et al., 2013; Zhao,  
50 et al., 2011). A number of studies have investigated the long-term variation features of haze days in  
51 Beijing and the NCP. Quan et al. (2011) collected monitoring data and summarized the haze day  
52 occurrence trend over NCP for the past 56 years. They also analyzed the effect of high aerosol loadings  
53 on haze formation by conducting a field measurement campaign and found the important role for  
54 hygroscopic growth of aerosols during the haze period. Yu et al. (2010) analyzed aerosol optical  
55 properties during haze days in the past seven years, and compared the features of single-scattering albedo  
56 and asymmetry factors during haze days with those during dust days in Beijing. This study also found that  
57 fine-mode particles were dominant in aerosol size distributions during haze days.

58 Numerous methods have been used to investigate the chemical and physical properties of aerosols

59 during haze occurrences in Beijing in different seasons. Li et al. (2013) identified the aerosol size  
60 distribution and chemical composition from ground-based remote sensing measurements during haze  
61 days in winter. Li et al. (2010) detected the aerosol components by using transmission electron  
62 microscopy with energy-dispersive X-ray spectrometry during a haze episode in summer, and determined  
63 the influence of carbonaceous aerosols. Liu et al. (2013) and Zhao et al. (2013) conducted intensive field  
64 experiments to identify the aerosol components of fine particles and discussed the constituent features of  
65  $PM_{2.5}$  during the haze periods in autumn and winter, respectively. Wang et al. (2006) compared the  
66 characteristics of aerosol components during dust, haze, and clean days. These previous works have  
67 provided abundant information on the physical and chemical properties of aerosols during haze days.  
68 However, the complex mechanism of haze formation over Beijing and its surrounding regions requires  
69 further study. Various influencing factors, including the meteorological field, key aerosol components,  
70 and microphysical properties, should be comprehensively considered in investigating the relationship  
71 between aerosols and surface visibility. Moreover, the seasonal similarities and differences of the haze  
72 formation mechanism in Beijing remain unclear because most of these studies were generally focused on  
73 the pollution periods in the same season.

74 In the present study, an air quality modeling system called the Regional Atmospheric Modeling  
75 System–Community Multi-Scale Air Quality (RAMS–CMAQ) coupled with an aerosol optical property  
76 scheme was applied to simulate the meteorological field, the mass burden of the major aerosol  
77 components (sulfate, nitrate, ammonium, black carbon (BC), organic carbon (OC), dust, and sea salt), and  
78 the surface visibility over NCP in 2011. The simulation results in February and July 2011 were selected  
79 for analysis. This study aims to discuss the contributions of various influencing factors to visibility  
80 deterioration and to compare the differences of the winter and summer haze formation mechanisms.

## 81 **2. Methodology**

82 The air quality modeling system RAMS–CMAQ was applied to concurrently simulate the  
83 atmospheric and land processes that affecting the transport, transformation, and deposition of aerosols and  
84 their precursors. The major component of this modeling system was CMAQ (version 4.7), developed by  
85 the US Environmental Protection Agency for assessing the effect of multiple pollutants, including  
86 tropospheric ozone and other oxidants, aerosols, and acid deposition (Byun and Schere, 2006; Eder and  
87 Yu, 2006; Eder et al., 2009; Mathur et al., 2008). The gas-phase chemistry mechanism was updated to the  
88 expanded version CB05 (Sarwar et al., 2008). The thermodynamic equilibrium between inorganic aerosol

89 species and gas-phase concentrations was treated by ISORROPIA (Nenes et al., 1999). Regional  
90 Particulate Model (Binkowski and Shankar, 1995) was used to describe the processes of aerosol dynamics  
91 in CMAQ; such processes include new particle production, coagulation, and condensation (Bhave et al.,  
92 2004; Yu et al., 2013). The formation of secondary organic aerosol (SOA) was mainly treated by the  
93 CB05 mechanism, which was extended to allow for production of SOA from anthropogenic and biogenic  
94 precursors. In the CB05, the SOA formation was modeled by forming semi-volatile products in volatile  
95 organic compounds (VOCs) reactions. The semi-volatile products were partitioned between the gas and  
96 aerosol phase according to the ambient conditions, such as temperature, relative humidity, vapor pressure,  
97 existing aerosol particles. The aerosol particles in the modeling system were divided into three modes,  
98 namely, Aitken, accumulation, and coarse modes (dust and sea salt). All modes were assumed to follow  
99 the log normal distribution. The aerosol components, the geometric standard deviation, and the geometric  
100 mean radius of each mode are listed in Table 1. The numerical prediction model RAMS was coupled with  
101 CMAQ in the offline method to provide CMAQ with a meteorological field. A general description of  
102 RAMS and its capabilities has provided by Cotton et al. (2003). RAMS can describe the boundary layer  
103 and the underlying surface effect, which is important for capturing air pollutants and haze occurrence.  
104 The background meteorological fields and sea surface temperature were obtained from the European  
105 Centre for Medium-Range Weather Forecasts reanalysis datasets ( $1^\circ \times 1^\circ$  spatial resolution) and were  
106 based on weekly mean values and observed monthly snow cover information, respectively.

107 The anthropogenic emissions of precursors and primary aerosols ( $\text{NO}_x$ ,  $\text{SO}_2$ , VOCs, BC, OC,  $\text{PM}_{2.5}$ ,  
108 and  $\text{PM}_{10}$ ) were obtained from the monthly-based emission inventory in China for 2010. This emission  
109 inventory has a spatial resolution of  $0.25^\circ \times 0.25^\circ$  and included four categories, namely, power, industry,  
110 residential, and transport (Lu et al., 2011). The nitrogen oxides and ammonia from soil were adopted from  
111 the Global Emissions Inventory Activity  $1^\circ \times 1^\circ$  monthly global inventory (Benkovitz et al., 1996). The  
112 monthly mean inventory of the Global Fire Emissions Database Version 2 (Randerson et al., 2007) was  
113 used to provide the biomass burning emissions from forest wildfires, savanna burning, and slash-and-burn  
114 agriculture. The online mechanisms introduced by Han et al. (2004) and Gong (2003) for capturing dust  
115 and sea salt emissions, respectively, were included in the modeling system.

116 A scheme of aerosol optical properties was added to the modeling system to estimate the aerosol  
117 extinction coefficient. This scheme contains a parameterization (Ghan and Zaveri, 2007) to efficiently  
118 simplify Mie theory calculation while maintaining sufficient accuracy. Briefly speaking, the lognormal

119 distribution in each mode can be expressed as:

$$120 \quad \frac{dN}{d \ln D} = \frac{N}{(2\pi)^{1/2} \ln \sigma_g} \exp\left(-\frac{(\ln D - \ln D_p)^2}{2 \ln^2 \sigma_g}\right) \quad (1)$$

121 where  $N$  is the number concentration of aerosol particles,  $\sigma_g$  is the geometric standard deviation,  $D$  is the  
122 particle diameter, and  $D_p$  is the geometric mean diameter. If the refractive index and  $\sigma_g$  are given and the  
123  $N$  is set as a normalized value, the aerosol optical properties can be calculated by Mie theory under  
124 several size distributions with different  $D_p$ . The values of the specific optical properties under these size  
125 distributions can be fitted by the Chebyshev polynomials with just five fitting coefficients. Subsequently,  
126 the fitting coefficient table can be constructed with all possible values of refractive index and  $\sigma_g$ . The  
127 scheme also applies Kohler theory (Pruppacher and Klett, 1997) and the Maxwell–Garnett mixing rule  
128 (Chuang et al., 2002) to describe the effects of water uptake and internal mixture, respectively. A detailed  
129 description of this scheme can be found in Han et al. (2011). The visibility can be obtained by using the  
130 following equation:

$$131 \quad VIS = 3.912/\beta \quad (2)$$

132 where  $VIS$  is the horizontal visibility, and  $\beta$  is the aerosol extinction coefficient (Seinfeld and Pandis,  
133 1998). This modeling system was used to simulate the mass concentration and optical properties of key  
134 aerosols in previous studies on aerosol effects on the climate and environment in China (Han et al., 2013;  
135 Han et al., 2011; Zhang et al., 2005, 2006, 2007).

136 For the simulation over NCP, a coarse domain covering most of East Asia with a horizontal grid  
137 distance of 64 km and a total area of 6654 km × 5440 km with a two-way nested inner domain was  
138 established (Han et al., 2011). The inner domain (Fig. 1) had 94 × 90 grid cells and a 16 km resolution on  
139 a rotated polar stereographic map projection centered at (116 °E, 40 °N). This domain included all major  
140 regions in the NCP, namely, the megacities of Beijing and Tianjin and Hebei, Shandong, and Shanxi  
141 provinces. A total of 15 vertical levels, nearly half of which were concentrated in the lowest 2 km, were  
142 used to improve the simulation of the atmospheric boundary layer. The positions of the measurement  
143 stations applied to model evaluation are marked on Fig. 1, as are the district areas of four major cities:  
144 Beijing, Tianjin, Shijiazhuang, and Jinan.

145

### 146 **3. Model evaluation**

147 In this section, the model simulations are compared with the observations. Meteorological drivers are

148 an important factor in aerosol and visibility simulation. The wind vector, temperature, and relative  
149 humidity are inherently related to aerosol transport, scavenging, and water uptake. Thus, the monitoring  
150 data from the surface stations of the Chinese National Meteorological Center (CNMC;  
151 (<http://cdc.cma.gov.cn/home.do>) were collected to evaluate the performance of the meteorological field  
152 simulation. The CNMC has 726 measurement stations that evenly distributed throughout mainland China,  
153 and has been providing long-term surface observations of several meteorological variables since 1  
154 January 1951 (Feng et al., 2004).

155 The comparative results of the daily average temperature, relative humidity, wind speed, and  
156 maximum wind direction at eight stations in February and July are shown in Figs. 2 to 5. The modeled  
157 temperature, relative humidity, and wind speed were in good agreement with the observations at nearly all  
158 stations. A persistent underestimation of wind speed by the models was found at the Wutaishan and  
159 Taishan sites. The modeled wind speed presented in Fig. 4 was obtained by converting the output values  
160 of the first layer (90 m to 200 m) to near-surface wind (~10 m) according to Monin–Obukhov similarity  
161 theory (Ding et al., 2001). These two sites are located on a mountainside at elevations of 2208 and 1533  
162 m, respectively. Thus, the underestimation may be attributed to the different elevations between the  
163 simulation and the observation. As shown in Fig. 5, the modeled wind directions did not coincide well  
164 with the observed data. A direct comparison is difficult to achieve because of the difference in time  
165 resolutions between the site measurements (10 min, average) and the model output (1 h). Nevertheless,  
166 the variation trends of the modeled and observed wind directions are similar at most sites, as shown in Fig.  
167 5. The monthly modeled precipitation over the NCP is compared to the observations of surface  
168 monitoring data from 87 CNMC sites in Fig. 6. The modeling system generally performed well capturing  
169 the distribution patterns and seasonal variation features of precipitation in the megacities of Beijing and  
170 Tianjin, and Inner Mongolia, Hebei, and Shandong provinces. However, the modeled results  
171 underestimated the precipitation in North Beijing and the northern part of Hebei province in July, which  
172 was perhaps the source of the error of wet deposition estimation. The relative humidity was also  
173 underestimated in the second half of February at the Miyun and Tanggu sites, and for almost the whole of  
174 July at Miyun as shown in Fig. 3. Underestimation most often happened when the relative humidity  
175 exceeded 70%, implying that the model did not accurately deal with high relative humidity over this  
176 region. The comparison of the modeled and observed precipitation and relative humidity in Beijing are  
177 discussed in detail below.

178 The modeled hourly NO<sub>2</sub>, O<sub>3</sub>, PM<sub>2.5</sub>, and visibility in February and July were also compared with the  
179 observed data provided by the Chinese Research Academy of Environmental Sciences (CRAES). This  
180 dataset comprised the real-time mass burden of air pollutants in Beijing (Gao et al., 2012). The results are  
181 shown in Figs. 7 and 9. The statistical parameters, including means, standard deviations, and correlation  
182 coefficients between the observations and simulations are listed in Table 2. These metrics were used to  
183 evaluate model performance, following the work of Yu et al. (2006). The model efficiently captured the  
184 daily variation of the pollutant gases and the high mass burden of PM<sub>2.5</sub> in these two months as shown in  
185 Fig. 7. Table 2 shows that most of the correlation coefficients were higher than 0.6, and the means and  
186 standard deviations of the simulations were also similar to those of the observations. However, the  
187 correlation coefficients of PM<sub>2.5</sub> and NO<sub>2</sub> were lower than 0.5 in July. From Fig.7 and the means and  
188 standard deviations in Table 2, we see that the model generally overestimated PM<sub>2.5</sub> in the middle of July,  
189 and the fluctuation range of modeled NO<sub>2</sub> was larger than that of the observation results. The comparison  
190 of modeled and observed daily precipitation in Beijing is given in Fig. 8 (the observation data was  
191 collected from CNMC. The modeled precipitation in July was clearly lower than that of observation in the  
192 middle of July. This may result in weaker wet deposition which can cause an overestimation of the  
193 aerosol burden in Beijing. For the simulated NO<sub>2</sub>, the larger diurnal variation was perhaps caused by the  
194 uncertainties from the related gas-phase chemical scheme in CB05. The modeled visibility also agreed  
195 well with the observations, particularly for visibility lower than 10 km, suggesting that the model could  
196 provide reasonable simulation during haze occurrence. The means and standard deviations of the modeled  
197 visibility were quite similar to those of observations. Meanwhile, continuous haze was found in the  
198 modeled and observed results in the middle of July, as shown in Fig. 9. This phenomenon indicates that  
199 although the model overestimated the mass burden of PM<sub>2.5</sub>, the visibility simulation during this period  
200 remained reliable. We also collected the hourly observation data of relative humidity from CRAES to  
201 evaluate the model performance in Beijing, and the comparison is shown in Fig. 10. Even though the  
202 model underestimated the relative humidity at the Miyun and Tangshan sites, the simulation results  
203 closely followed the observations in Beijing. The model only just overestimated the relative humidity  
204 when its value was lower than about 30%. This evaluation indicates that the model more reliably  
205 simulates relative humidity in Beijing, than in Miyun and Tangshan.

206 The modeled daily average mass concentrations of the major aerosol components were compared  
207 with the observed data from the CRAES measurements, as shown in Fig. 11. The observed data lacked

208 information for the first half of February and a number of days in July because of instrument failure.  
209 Although the magnitudes of the mass concentrations between the simulation and the observation did not  
210 exactly coincide, the modeled results broadly reproduced the peaks of the observed data from February 20  
211 to February 23 and from July 20 to July 23; the modeled results also followed the seasonal variation  
212 features. For instance, the modeled and observed carbonaceous aerosols were both high in February and  
213 low in July. The model demonstrated an obvious systematic underestimation of organic carbon in these  
214 two months, as shown in Fig. 11. Numerous studies have reported that such a phenomenon is a common  
215 issue in regional chemistry and transport models (Heald et al., 2005; Koch et al., 2007). The simulation  
216 error was primarily due to the uncertainties in the estimation of VOCs and primary organic aerosol  
217 emissions and the formation mechanism of secondary organic aerosol (Kroll et al., 2006; Henze and  
218 Seinfeld, 2006; Yu et al., 2007). However, this discrepancy did not significantly affect the accuracy of the  
219 visibility simulation. Therefore, these evaluations suggest that the modeling system can reasonably  
220 simulate the meteorological field, the mass burdens of major aerosol components, and the surface  
221 visibility in February and July 2011.

222

## 223 **4. Results and discussions**

### 224 4.1 Distribution features of aerosol concentration and visibility

225 As shown in Figs. 7 to 11, two typical heavy air pollution episodes occurred over the NCP from  
226 February 20 to February 23 and from July 20 to July 23. These two periods were selected to investigate  
227 the distribution features of pollutants and visibility over the NCP during occurrences of heavy pollution in  
228 different seasons. Fig. 12 presents the horizontal distributions of the daily average mass concentration of  
229  $PM_{2.5}$  and surface wind field over the NCP from February 20 to February 23 and from July 20 to July 23.  
230 The heavy mass burden of  $PM_{2.5}$  (over  $120 \mu g m^{-3}$ ) was mainly concentrated in the megacities of Beijing  
231 and Tianjin, the whole area of Hebei province and northwest part of Shandong province. The mass  
232 concentration of  $PM_{2.5}$  in February, which exceeded  $200 \mu g m^{-3}$  in Beijing, Tianjin, Shijiazhuang, and  
233 Jinan, was higher than that in July. The high mass burden of  $PM_{2.5}$  appeared in the same regions as those  
234 in February broadly ranging from  $75 \mu g m^{-3}$  to  $200 \mu g m^{-3}$  in July, and rarely exceeded  $200 \mu g m^{-3}$  over  
235 the entire NCP. The high mass burden of  $PM_{2.5}$  in Beijing generally appeared when the NCP was  
236 dominated by the south wind field, which transported air masses from the polluted regions in the south.  
237 The heavy  $PM_{2.5}$  mass burden may have been transported to Northeast China and the Bohai Sea by the



238 strong south wind from February 22 to February 23 and on July 23, respectively, thereby increasing the  
239 mass concentration of  $\text{PM}_{2.5}$  by  $45 \mu\text{g m}^{-3}$  to  $125 \mu\text{g m}^{-3}$  in these two regions, as shown in Fig. 12.

240 Fig. 12 also presents the horizontal distributions of daily average visibility and surface relative  
241 humidity over the NCP. The data shows that haze cloud spread throughout the NCP during each pollution  
242 episode. The visibility in most parts of Hebei and Shandong was generally less than 8 km, decreasing to 3  
243 to 5 km in the four urban areas in February and July. The distribution patterns of visibility broadly  
244 followed those of the  $\text{PM}_{2.5}$  mass burden, and the deterioration in visibility mainly appeared in the regions  
245 where the heavy  $\text{PM}_{2.5}$  mass burden was concentrated. The visibility generally decreased to 3 to 5 km  
246 when the mass concentration of  $\text{PM}_{2.5}$  exceeded  $200 \mu\text{g m}^{-3}$  in February. However, similar values of  
247 visibility also appeared in July when the mass concentration of  $\text{PM}_{2.5}$  was in the range 120 to  $200 \mu\text{g m}^{-3}$ .  
248 Such a phenomenon was apparent on July 23 when the visibility over the entire Bohai Sea ranged from 3  
249 5 km, and the mass concentration of  $\text{PM}_{2.5}$  was maintained between 120 and  $200 \mu\text{g m}^{-3}$ . These  
250 differences may be due to the strong extinction of soluble particles caused by the high relative humidity  
251 (exceeding 70%) in July, as shown in Fig. 12. This feature is discussed in detail below.

#### 252 4.2 Meteorological factors, major aerosol components and their contributions to extinction in Beijing

253 Figs. 13(a) to 13(h) present the time series of regional average surface wind speed, and relative  
254 humidity, visibility, as well as the mass concentrations of  $\text{PM}_{2.5}$ , sulfate, nitrate, ammonium, BC, and OC  
255 in Beijing in February and July 2011. The averages of these variables during the haze days in February  
256 and July are shown in Table 3. The mass burden of  $\text{PM}_{2.5}$  was the most important factor influencing  
257 changes of visibility, as it is generally inversely correlated with the variation of visibility. The mass  
258 concentrations of the three kinds of inorganic salt, namely, nitrate, sulfate, and ammonium, suggest that  
259 they were the three major aerosol components of  $\text{PM}_{2.5}$  in Beijing, as shown in Figs. 13(e) to 13(f). The  
260 mass burden of organic carbon was comparable with that of nitrate, so the total mass burden of organic  
261 matter should be larger than that of nitrate. If not, then nitrate should be the main particulate pollutant  
262 during winter because the mass burden of nitrate was higher than those of the other components in  
263 February. Although the diurnal variation of nitrate concentration was significant in July, the daily  
264 maximum nitrate concentration was still larger than that of sulfate during the nighttime. These findings  
265 suggest that the emission from the transportation sector was the major source of secondary particles in  
266 Beijing. The mass burden of carbonaceous aerosols was high in February and low in July. In addition to  
267 the diffusion conditions, the strong emissions of coal and biomass burning were the main reasons for the

268 high values of carbonaceous aerosols during the winter (Zhang et al., 2014).

269 As shown in Table 3, July had a greater number of haze days than February, and the average  
270 visibility during haze days in July was lower than that in February. These features indicate that air  
271 pollution was more serious in July than in February. However, the average mass concentration of PM<sub>2.5</sub> in  
272 July during haze days was obviously lower than that in February. In addition, Table 3 shows that the  
273 relatively low value of PM<sub>2.5</sub> mass concentration in July was primarily due to the small quantities of  
274 carbonaceous aerosol burden. Meanwhile, the total mass burden of nitrate, sulfate, and ammonium was  
275 higher in July than in February. Therefore, deterioration in visibility was caused by the simultaneous  
276 occurrence of the high mass burden of soluble particles and high relative humidity in July. The difference  
277 in the haze formation mechanism during winter and summer was associated with the different PM<sub>2.5</sub>  
278 particle composition and the ambient relative humidity.

279 Figs. 13(g) to 13(h) present the time series of the regional average contribution ratios of sulfate,  
280 nitrate, ammonium, BC, OC, and other components (dust, sea salt, and unspecified anthropogenic mass)  
281 to the total surface extinction in Beijing in February and July. The monthly mean of these contribution  
282 ratios are shown in Table 4. The contribution ratios were calculated by subtracting the extinction  
283 coefficient with and without each aerosol component when estimated the aerosol optical properties by  
284 using the scheme introduced in Section 2. The inorganic salts nitrate, sulfate, and ammonium,  
285 significantly contributed to the surface extinction in Beijing, which was ~70% in February and over 80%  
286 in July. Carbonaceous aerosol had a 20% and 5% contribution in February and July, respectively, whereas  
287 other aerosol components had around a 10% contribution. These ratios generally followed the magnitude  
288 of the mass concentrations. Except for the diurnal variation of nitrate in July, the contribution ratios of  
289 each aerosol component did not significantly change when the mass concentration of PM<sub>2.5</sub> exceeded ~50  
290  $\mu\text{g m}^{-3}$ . In contrast, when the mass concentration of PM<sub>2.5</sub> decreased to less than 50  $\mu\text{g m}^{-3}$ , the  
291 contribution ratios of carbonaceous aerosol and other components obviously increased. A higher mass  
292 concentration of PM<sub>2.5</sub> corresponded to a higher contribution ratio of the three inorganic salts. This  
293 feature confirmed that nitrate, sulfate, and ammonium were the major aerosol components influencing  
294 haze formation in Beijing.

#### 295 4.3 Haze occurrence threshold in Beijing

296 It can be seen from Fig. 13 that the mass concentration of PM<sub>2.5</sub> was closely inversely correlated  
297 with visibility. However, when the mass concentration of PM<sub>2.5</sub> was located in different mass value

298 intervals, the influence on the visibility was not consistent. Fig. 14(a) shows the time series of the  
299 regional mean mass concentration of  $PM_{2.5}$  and visibility in Beijing from 23 to 25 July. The air quality for  
300 these three days improved over time and the visibility continuously increased. The mass concentration of  
301  $PM_{2.5}$  decreased from  $260 \mu g m^{-3}$  on 23 July to  $20 \mu g m^{-3}$  on 25 July. For the convenience, the time taken  
302 for the decrease was divided into Period A, B and C, as shown in the figure. Period A, the mass  
303 concentration of  $PM_{2.5}$  changed from  $260 \mu g m^{-3}$  to  $120 \mu g m^{-3}$ , decreasing by about  $140 \mu g m^{-3}$  while the  
304 visibility increased by less than 5 km; in Period B, the mass concentration of  $PM_{2.5}$  changed from  $120 \mu g$   
305  $m^{-3}$  to  $50 \mu g m^{-3}$ , decreasing by about  $70 \mu g m^{-3}$  and the visibility increased obviously, from 5km to about  
306 20 km; finally, in Period C, the mass concentration of  $PM_{2.5}$  changed from  $35 \mu g m^{-3}$  to  $20 \mu g m^{-3}$ ,  
307 decreasing by only  $15 \mu g m^{-3}$  and the visibility increased dramatically by 60 km. The above analysis  
308 indicated that even though emission reduction measures were taken to dramatically decrease the  $PM_{2.5}$   
309 mass burden, the improvement of visibility would still be weak if the mass concentration of  $PM_{2.5}$   
310 remained at a high level. Only when the mass concentration of  $PM_{2.5}$  decreased to the certain value range  
311 did the visibility improve. The visibility was calculated by using Formula (2) in the modeling system. The  
312 aerosol extinction coefficient should be the key factor influencing the visibility. From Fig. 14(b), the  
313 variation of the extinction coefficient and mass concentration of  $PM_{2.5}$  were quite similar. We therefore  
314 deduced that when the value of extinction coefficient became small, the value of visibility could change  
315 dramatically with only a micro-variation of the extinction coefficient. This should be the main factor  
316 causing the drastic change of visibility during Period C. Therefore, strictly speaking it was necessary to  
317 distinguish the atmospheric haze and the atmospheric pollution. That is, an improvement of air quality  
318 with decreasing mass concentrations of pollutants did not mean that the haze disappeared. If the  
319 occurrence of haze was controlled by decreasing mass concentration of  $PM_{2.5}$  in the atmosphere, a  
320 reasonable solution was to set a haze occurrence threshold interval, corresponding to the values of mass  
321 concentration of  $PM_{2.5}$  when the visibility reached 10 km under different ambient conditions. Only by  
322 strictly keeping the mass concentration of  $PM_{2.5}$  below this threshold did the visibility improve.  
323 Otherwise, even with the emission reduction measures taken when a heavy pollution event appears, the  
324 improvement of visibility would be very small if the mass concentration of  $PM_{2.5}$  failed to fall below the  
325 values of haze occurrence threshold. Furthermore, the specific value of the threshold is closely related to  
326 the pollutant characteristics, meteorological conditions and other factors.

327 A sensitivity test was conducted to evaluate the mass concentration threshold of  $PM_{2.5}$  above which

328 haze occurred in Beijing. First, the mass ratio of each aerosol component to the total mass burden of all  
329 aerosol particles was calculated from the results of the model simulation at every grid point. Then, the  
330 sensitivity test was conducted by using several possible values of the total aerosol burden and following  
331 the same ratio of each aerosol component at the same grid points to identify the mass concentration  
332 threshold of PM<sub>2.5</sub> when the visibility decreased to 10 km under different relative humidity. The values of  
333 relative humidity were: 70%, 75%, 80%, 85%, 88%, 89%, and 90%. Lower values of relative humidity  
334 were disregarded because the water uptake of soluble particles was insignificant when the relative  
335 humidity was less than 70%. Values of relative humidity higher than 90% indicate that light fog occurred  
336 as expressed by Wu et al. (2007). Figs. 15(a) and 15(b) present the time series of the regional average  
337 threshold of haze occurrence under different values of relative humidity from the sensitivity test in  
338 February and July in Beijing. The threshold changed significantly with the variation in relative humidity,  
339 and its declining trend increased with increasing relative humidity. The range of the mass concentration  
340 threshold reached 30  $\mu\text{g m}^{-3}$  when the relative humidity changed from 70% to 90%. Conversely, the  
341 threshold generally maintained a small change ( $<5 \mu\text{g m}^{-3}$ ) when the relative humidity was fixed. This  
342 indicates that if the aerosol components did not have a dramatic variation in Beijing, a relatively fixed  
343 haze occurrence threshold could be determined.

344 However, the mass concentration threshold on July 29 increased by approximately 10  $\mu\text{g m}^{-3}$  under  
345 the same relative humidity, as shown in Fig. 15(b). Further analysis showed that this phenomenon might  
346 be related to the variation of aerosol in accumulation mode. Figs. 15(c) to 15(h) present the time series of  
347 the regional average mass ratios and contribution ratios of the three particle modes to the total aerosol  
348 burden and the total extinction, respectively. As seen in Figs. 15(e) and 15(h), the extinction contribution  
349 of accumulation mode particles was above 97% due to the high mass concentration ratio and extinction  
350 efficiency. However, the extinction contribution of accumulation mode particles on 29 July decreased by  
351 about 5%. Therefore, we deduced there should be a high correlation between the haze occurrence  
352 threshold and extinction contribution of accumulation mode particles. Furthermore, it can be seen from  
353 Fig. 15(j) that the visibility on 29 July rose rapidly from less than 5 km to more than 20 km. Thus, it can  
354 be deduced that a weather process beneficial to pollutant scavenging eliminated the mass concentration of  
355 accumulation mode particles efficiently, i.e. the new particles in the atmosphere were eliminated  
356 immediately before coagulation or condensation in this period. The mass concentration ratio and  
357 extinction contribution ratio of Aitken mode particles increased by about 10% and 5%, respectively. This

358 is the major reason for the decreasing extinction contribution ratio of accumulation mode particles.  
359 Similar weather processes also occurred on 2–3, 5–10, and 25 July. During these periods the mass  
360 concentration of accumulation mode particles decreased significantly. However, different from the  
361 condition on 29 July, the mass concentration ratio and extinction contribution of Aitken mode particles  
362 did not change, but the mass concentration ratio of coarse mode particles increased dramatically.  
363 Although the mass concentration of the coarse mode particles accounted for 10%–20% of the total aerosol,  
364 its extinction contribution was below 1% in most periods. Therefore, except for some special cases (e. g.,  
365 the dust event), the influence of coarse mode particles on the extinction was weaker than other modes.  
366 This was also the main reason for the lack of significant variation of the extinction contribution ratio of  
367 accumulation mode particles in these three processes.

368 Generally speaking, besides relative humidity, the haze occurrence threshold is also sensitive to the  
369 extinction contribution ratio of accumulation mode particles. From the above analysis, the mass  
370 concentration ratio of accumulation mode particles generally remained at a high level and the fluctuation  
371 range was small during the heavy pollution episode in Beijing. The variation range of haze occurrence  
372 threshold was less than  $5 \mu\text{g m}^{-3}$  when the relative humidity was fixed. The increase of the haze  
373 occurrence threshold due to the variation of the extinction contribution of accumulation mode particles  
374 only appeared when the mass burden of Aitken mode particles increased in clean periods. The reason for  
375 this phenomenon is that the extinction efficiency of Aitken mode particles is far smaller than that of  
376 accumulation mode particles. Therefore, more Aitken mode particles are needed to form the haze.  
377 However, smaller particles generally existed during the clean period as shown by the simulation results,  
378 which means the haze did not appear. Thus, the influence of the extinction contribution of accumulation  
379 mode particles on haze occurrence threshold can be neglected in Beijing. The relative humidity should be  
380 the only impact factor which needs to be considered.

381 The monthly means of the threshold of haze occurrence are shown in Table 5. It can be seen from  
382 Table 5 that when the relative humidity changed from 70% to 90%, the threshold interval increased from  
383  $52 \mu\text{g m}^{-3}$  to  $83 \mu\text{g m}^{-3}$ . Within a certain relative humidity range, the average monthly thresholds were  
384 similar in February and July. Here, the relationship between the haze occurrence threshold and relative  
385 humidity was fitted using:

$$386 \quad RH = a + b \times \exp(c \times M) \quad (3)$$

387 where  $RH$  represents relative humidity;  $M$  represents the  $\text{PM}_{2.5}$  mass concentration threshold;  $a$ ,  $b$  and  $c$

388 represent fitting parameters, and their values are listed in Table 6. The fitting curve is shown in Fig. 16.  
389 From Table 6 it can be seen that the values of  $R^2$  were all higher than 0.9, indicating that their relationship  
390 can be described well by Formula 3. Therefore, the fitting curves given in Fig. 16 can be used to capture  
391 the haze occurrence threshold in Beijing. Below 90% relative humidity, when  $PM_{2.5}$  mass concentration  
392 increases beyond the corresponding values on the curve, the haze should appear. Furthermore, the  
393 analysis in this study also indicated that the haze occurrence can be efficiently controlled by strictly  
394 restricting the  $PM_{2.5}$  mass concentration near or below the fitting curve. Otherwise, even a very large  
395 decrease of the  $PM_{2.5}$  mass burden would not reduce the possibility of haze occurrence.

396

## 397 **5. Conclusions**

398 In this study, the air quality modeling system RAMS-CMAQ, coupled with an aerosol optical  
399 property scheme, was used to simulate the meteorological field, the mass concentration of aerosols, and  
400 the surface visibility over the NCP in 2011. The modeling system provided reliable simulation results.  
401 The distribution patterns and time series of related meteorological factors and aerosol characteristic in  
402 February and July 2011 were analyzed to elucidate the seasonal variation features of the haze formation  
403 mechanism in Beijing and its surrounding regions. In addition, a sensitivity test was conducted to  
404 investigate the  $PM_{2.5}$  mass concentration threshold of haze occurrence in Beijing under distinct conditions.  
405 The results are summarized as follows:

406 (1) The simulation results showed that the high mass burden of  $PM_{2.5}$  over the NCP was mainly  
407 concentrated in Beijing and Tianjin megacities, the whole area of Hebei province and the northwest part  
408 of Shandong province. The daily average mass concentration of  $PM_{2.5}$  over these regions was generally  
409 over  $120 \mu\text{g m}^{-3}$  during the pollution episodes in February and July. The worst air quality over the NCP  
410 was found in Beijing because of the heavy daily average mass burden of  $PM_{2.5}$ , which exceeded  $300 \mu\text{g}$   
411  $\text{m}^{-3}$  in February. The south wind that carried pollutants from the southern regions was an important  
412 source of the heavy aerosol loading in Beijing. In addition to the horizontal diffusion, the vertical  
413 convection also played an important role in the pollutant scavenging in Beijing.

414 (2) The distribution patterns of visibility generally followed those of the  $PM_{2.5}$  mass burden. The  
415 daily average visibility below 10 km covered most regions of the NCP during the pollution episodes in  
416 February and July and was below 5 km over the urban areas. The daily average relative humidity rarely  
417 exceeded 90%, suggesting that the haze cloud could spread throughout the NCP when the pollution

418 episode appeared in both winter and summer.

419 (3) The simulation results showed that nitrate, sulfate, and ammonium were the three major aerosol  
420 components and the main causes of the visibility deterioration in Beijing. The mass burdens of these three  
421 inorganic salts were obviously higher than those of other aerosols, and their total contribution ratios to  
422 surface extinction reached 70% in February and 85% in July. Nitrate was also the first and second major  
423 contributor to surface extinction in February and July, respectively, implying that the emission from the  
424 transportation sector was the major source of secondary particles in Beijing. Carbonaceous aerosols  
425 accounted for 15% extinction in February and below 5% extinction in July. This indicates that the  
426 pollution status and emission sources were more complicated during winter in Beijing.

427 (4) The haze formation mechanism in Beijing in winter was obviously different from that in summer.  
428 Firstly, the mass concentration of  $PM_{2.5}$  in winter was relatively higher and the components were  
429 complicated. The ratios of inorganic salts and carbonaceous aerosols were generally balanced. Therefore,  
430 the high mass concentration of  $PM_{2.5}$  and diverse aerosol components were the major reasons of the  
431 serious haze occurrence in winter. While the mass concentration of  $PM_{2.5}$  in summer was lower than that  
432 in winter, the ratio of hygroscopic inorganic salts, including sulfate, nitrate and ammonium, increased and  
433 their mass concentrations were even higher than those in winter. With higher relative humidity, serious  
434 haze may still form on an equal level as winter even though the mass concentration of  $PM_{2.5}$  is lower. The  
435 water uptake of hygroscopic components played a key role. This indicated that it is important to apply  
436 emission reduction measures based on the specific pollution and meteorological characteristics in  
437 different seasons. In this way, the possibility of haze occurrence can be effectively decreased.

438 (5) From analysis, it was found that even though the mass concentration of  $PM_{2.5}$  was closely  
439 inversely correlated with visibility, the influencing effect was diverse when the mass concentration of  
440  $PM_{2.5}$  was located in different intervals. When the mass concentration of  $PM_{2.5}$  was larger than  $100 \mu g$   
441  $m^{-3}$ , the influence of its variation on visibility was very weak. Only when the mass concentration of  $PM_{2.5}$   
442 was below a certain interval could its decrease make the visibility increase rapidly. Therefore, it was  
443 reasonable to set a haze occurrence threshold interval, and this was chosen to be the value of the mass  
444 concentration of  $PM_{2.5}$  when the visibility exceeded 10 km under different ambient conditions. If the mass  
445 concentration of  $PM_{2.5}$  failed to fall below the values of the haze occurrence threshold, the improvement  
446 of visibility would still be very weak when the emission reduction measures are taken.

447 (6) Through the sensitivity experiment, this study estimated the haze occurrence threshold interval in

448 Beijing, and discussed related impact factors. Generally speaking, if the components of PM<sub>2.5</sub> did not have  
449 a dramatic variation in Beijing, the haze occurrence threshold was only sensitive to the extinction  
450 contribution ratio of the accumulation mode particles and relative humidity. Considering that the variation  
451 of the extinction contribution ratio of the accumulation mode particles generally occurs in clean periods,  
452 the relative humidity should be the only impact factor. Finally, the relationship between the threshold of  
453 haze occurrence and relative humidity in winter and summer in Beijing was fitted by the exponential  
454 function, and these fitting curves could form a new theoretical basis for the further understanding and  
455 control of the haze formation in Beijing. As the analysis in this study, the fitting function could be applied  
456 to diagnose the haze events under ordinary conditions in Beijing. However, this function is not suitable  
457 for the haze over other regions, or in the event of uncommon pollution episodes (e.g. the dust storm) in  
458 Beijing, because the ratio of major aerosol components and particle size distribution would be different.  
459 Further study therefore remains necessary for other regions and various pollution features.

460

## 461 **Acknowledgments**

462 This study was supported by the National Basic Research Program of China (2014CB953802), the  
463 “Strategic Priority Research Program (B)” of the Chinese Academy of Sciences (XDB05030105,  
464 XDB05030102, XDB05030103), and the National Natural Science Foundation of China (41105106 and  
465 41205123).

466

467

468

469

470

471

472

473

474

475

476



## References

- Benkovitz, C., Schultz, M., Pacyna, J., Tarrason, L., Dignon, J., Voldner, E., Spiro, P., Logan, J., and Graedel, T.: Global gridded inventories of anthropogenic emissions of sulfur and nitrogen, *J. Geophys. Res.*, 101, 29239-29254, 1996.
- Bhave, P. V., Roselle, S. J., Binkowski, F. S., Nolte, C. G., Yu, S. C., Gipson, G. L., and Schere, K. L.: CMAQ Aerosol Module Development: Recent Enhancements and Future Plans, Paper presented at 3rd Annual CMAS Models-3 Users' Conference, Commun. Model., and Anal. Syst. Cent., Chapel Hill, N.C., 18-20 October, 2004.
- Binkowski, F. S. and Shankar, U.: The regional particulate model, 1, Model description and preliminary results, *J. Geophys. Res.*, 100, 26191-26209, 1995.
- Byun, D. and Schere, K.: Review of the governing equations, computation algorithms, and other components of the Models-3 Community Multiscale Air Quality (CMAQ) modeling system. *Appl. Mech. Rev.*, 59, 51-77, 2006.
- Chuang, C., Penner, J., Prospero, J., Grant, K., Rau, G., and Kawamoto, K.: Cloud susceptibility and the first aerosol indirect forcing: sensitivity to black carbon and aerosol concentrations, *J. Geophys. Res.*, 107, doi: 10.1029/2000JD000215, 2002.
- Cotton, W., Pielke, R., Walko, G., Liston, G., Tremback, C., Jiang, H., McAnelly, R., Harrington, J., Nicholls, M., Carrio, G., and McFadden, J.: RAMS 2001: current status and future directions, *Meteorol. Atmos. Phys.*, 82, 5-29, 2003.
- Ding, F., Pal Arya, S., and Lin, Y.: Large-eddy simulations of the atmospheric boundary layer using a new subgrid-scale model-II. Weakly and moderately stable cases, *Environ. Fluid Mech.*, 1, 49-69, 2001.
- Eder, B., Kang, D., Mathur, R., Pleim, J., Yu, S. C., Otte, T., and Pouliot, G.: A performance evaluation of the national air quality forecast capability for the summer of 2007, *Atmos. Environ.*, 43, 2312-2320, 2009.
- Eder, B. and Yu, S. C.: An evaluation of model performance of EPA models-3/CMAQ, *Atmos. Environ.*, 40, 4811-4824, 2006.
- Feng, S., Hu, Q., and Qian, W.: Quality control of daily meteorological data in China, 1951-2000: a new dataset. *Int. J. Climatol.*, 24, 853-870, 2012.
- Gao, J., Zhang, Y., Wang, S., Chi, F., and Chen, Y.: Study on the Characteristics and Formation of a Multi-Day Haze in October 2011 in Beijing, *Res. Environ. Sci. (in Chinese)*, 25, 1201-1207, 2012.
- Ghan, S., and Zaveri, R.: Parameterization of optical properties for hydrated internally mixed aerosol, *J. Geophys. Res.*, 112, D10201, doi:10.1029/2006JD007927, 2007.
- Gong, S. L.: A parameterization of sea-salt aerosol source function for sub- and super-micron particles, *Global Biogeochem. Cy.*, 17, doi: 10.1029/2003GB002079, 2003.
- Han, X., Zhang, M., Han, Z., Xin, J., and Liu, X.: Simulation of aerosol direct radiative forcing with RAMS-CMAQ in East Asia, *Atmos. Environ.*, 45, 6576-6592, 2011.
- Han, X., Zhang, M., Tao, J., Wang, L., Gao, J., Wang, S., Chai, F.: Modeling aerosol impacts on atmospheric visibility in Beijing with RAMS-CMAQ, *Atmos. Environ.*, 72, 177-191, 2013.
- Han, Z., Ueda, H., Matsuda, K., Zhang, R., Arao, K., Kanai, Y., and Hasome, H.: Model study on particle size segregation and deposition during Asian dust events in March 2002, *J. Geophys. Res.*, 109, doi: 10.1029/2004JD004920, 2004.
- Hao, J. and Wang, L.: Improving Urban Air Quality in China: Beijing Case Study, *Air Waste Manage. Assoc.*, 62, 1298-1305, 2012.

520 Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Liao, H., and Weber,  
521 R. J.: A large organic aerosol source in the free troposphere missing from current models, *Geophys.*  
522 *Res. Lett.*, 32, L18809, doi:10.1029/2005GL023831.

523 Henze, D. K. and Seinfeld, J.H.: Global secondary organic aerosol from isoprene oxidation, *Geophys. Res.*  
524 *Lett*, 33:L09812, doi:10.1029/2006GL025976, 2006.

525 Koch, D., Bond, T. C., Streets, D., Unger, N., van der Werf, G. R.: Global impact of aerosols from  
526 particular source regions and sectors, *J. Geophys. Res.*, 112, D02205, doi:10.1029/2005JD007024,  
527 2007.

528 Kroll, J. H., Ng, N. L., Murphy, S. M., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol  
529 formation from isoprene photooxidation, *Environ. Sci. Technol.*, 40, 1869-1877, 2006.

530 Li, W., Shao, L., Buseck, P.: Haze types in Beijing and the influence of agricultural biomass burning,  
531 *Atmos. Chem. Phys.*, 10, 8119-8130, 2010.

532 Li, Z., Gu, X., Wang, L., Li, D., Li, K., Dubovik, O., Schuster, G., Goloub, P., Zhang, Y., Li, L., Xie, Y.,  
533 Ma, Y., and Xu, H.: Aerosol physical and chemical properties retrieved from ground-based remote  
534 sensing measurements during heavy haze days in Beijing winter, *Atmos. Chem. Phys.*, 13,  
535 5091-5122, 2013.

536 Liu, X., Li, J., Qu, Y., Han, T., Hou, L., Gu, J., Chen, C., Yang, Y., Liu, X., Yang, T., Zhang, Y., Tian, H.,  
537 and Hu, M.: Formation and evolution mechanism of regional haze: a case study in the megacity  
538 Beijing, China, *Atmos. Chem. Phys. Discuss.*, 12, 16259-16292, 2012.

539 Lu, Z., Streets, D. G., Zhang, Q., Wang, S., Carmichael, G. R., Cheng, Y. F., Wei, C., Chin, M., Diehl, T.,  
540 and Tan, Q.: Sulfur dioxide emissions in China and sulfur trends in East Asia since 2000, *Atmos.*  
541 *Chem. Phys.*, 10, 6311–6331, 2010.

542 Lu, Z, Zhang, Q, Streets, D.: Sulfur dioxide and primary carbonaceous aerosol emissions in China and  
543 India, 1996–2010. *Atmospheric Chemistry and Physics*, 11: 9839-9864, 2011.

544 Ma, J., Xu, X., Zhao, C., Yan, P.: A review of atmospheric chemistry research in China: Photochemical  
545 smog, haze pollution, and gas-aerosol interactions. *Adv. Atmos. Sci.*, 29, 1006-1026, 2010.

546 Mathur, R., Yu, S. C., Kang, D., and Schere, K. L.: Assessment of the winter-time performance of  
547 developmental particulate matter forecasts with the Eta-CMAQ modeling system, *J. Geophys. Res.*,  
548 113, D02303, doi:10.1029/2007JD008580, 2008.

549 Nenes, A., Pandis, S., and Pilinis, C.: Continued development and testing of a new thermodynamic  
550 aerosol module for urban and regional air quality models, *Atmos. Environ.*, 33, 1553-1560, 1999.

551 Pruppacher, H. and Klett, J.: *Microphysics of Clouds and Precipitation*. Springer, New York. p. 954, 1997.

552 Quan, J., Zhang, Q., He, H., Liu, J., Huang, M., and Jin, H.: Analysis of the formation of fog and haze in  
553 North China Plain (NCP), *Atmos. Chem. Phys.*, 11, 8214-8250, 2011.

554 Randerson, J. T., van der Werf, G. R., Giglio, L., Collatz, G. J., and Kasibhatla, P. S.: Global Fire  
555 Emissions Database, Version 2 (GFEDv2.1), Data set available on-line (<http://daac.ornl.gov/>) from  
556 Oak Ridge National Laboratory Distributed Active Archive Center, Oak Ridge, Tennessee, USA,  
557 doi:10.3334/ORNLDAAAC/849, 2007.

558 Sarwar, G., Luecken, D., Yarwood, G., Whitten, G., and Carter, W.: Impact of an updated carbon bond  
559 mechanism on predictions from the CMAQ modeling system: preliminary assessment, *J. Appl.*  
560 *Meteorol. Clim.*, 47, 3-14, 2008.

561 Seinfeld, J. and Pandis, S.: *Atmospheric Chemistry and Physics*. Wiley, New York, USA, 1998.

562 Sun, Y., Zhuang, G., Tang, A., Wang, Y., and An, Z.: Chemical characteristics of PM<sub>2.5</sub> and PM<sub>10</sub> in  
563 Haze-Fog episodes in Beijing, *Environ. Sci. Technol.*, 40, 3148-3155, 2006.

564 Tao, M., Chen, L., Su, L., and Tao, J.: Satellite observation of regional haze pollution over the North  
565 China Plain, *J. Geophys. Res.*, 117, D12203, DOI: 10.1029/2012JD017915, 2012.

566 Wang, X., Sun, M., Yang, T., and Wang, Z.: Interdecadal change in frequency of dust-haze episodes in  
567 North China Plain, *Clim. Environ. Res.* (in Chinese), 18, 165-170, 2013.

568 Wang, Y., Zhuang, G., Sun, Y., An, Z.: The variation of characteristics and formation mechanisms of  
569 aerosols in dust, haze, and clear days in Beijing, *Atmos. Environ.*, 40, 6579-6591, 2006.

570 Wu, D., Bi, X., Deng, X., Li, F., Tan, H.: Effect of atmospheric haze on the deterioration of visibility over  
571 the Pearl River Delta, *Acta Geogr. Sin.* (in Chinese), 21, 215-223, 2007.

572 Yu, S. C., Bhave, P. V., Dennis, R. L., and Mathur, R.: Seasonal and regional variations of primary and  
573 secondary organic aerosols over the continental United States: semi-empirical estimates and model  
574 evaluation, *Environ. Sci. Technol.*, 41, 4690-4697, 2007.

575 Yu, S. C., Eder, B., Dennis, R., Chu, S. H., and Schwartz, S.: New unbiased symmetric metrics for  
576 evaluation of air quality models, *Atmos. Sci. Lett.*, 7, 26-34, 2006.

577 Yu, S. C., Mathur, R., Pleim, J., Wong, D., Gilliam, R., Alapaty, K., Zhao, C., and Liu, X.: Aerosol  
578 indirect effect on the grid-scale clouds in the two-way coupled WRF-CMAQ: model description,  
579 development, evaluation and regional analysis, *Atmos. Chem. Phys. Discuss.*, 13, 25649-25739,  
580 2013.

581 Yu X, Zhu B, Yin Y, Yang, J., Li, Y., and Bu, X.: A comparative analysis of aerosol properties in dust and  
582 haze-fog days in a Chinese urban region. *Atmos. Res.*, 99, 241-247, 2011.

583 Zhang, A., Qi, Q., Jiang, L., Zhou, F., and Wang, J.: Population exposure to PM<sub>2.5</sub> in the urban area of  
584 Beijing, *PloS one*, 8, e63486, 2013.

585 Zhang, J., Miao, H., Ouyang, Z., and Wang, X.: Ambient air quality trends and driving factor analysis  
586 since 1980's in Beijing, *Acta Sci. Circumstantiae* (in Chinese), 26, 1886-1892, 2006.

587 Zhang, J., Sun, Y., Liu, Z., Ji, D., Hu, B., Liu, Q., and Wang, Y.: Characterization of submicron aerosols  
588 during a month of serious pollution in Beijing, 2013, *Atmos. Chem. Phys.*, 14, 2887-2903, 2014.

589 Zhang, J., Zhu, T., Zhang, Q., Li, C., Shu, H., Ying, Y., Dai, Z., Liu, X., Liang, A., and Shen, H.: The  
590 impact of circulation patterns on regional transport pathways and air quality over Beijing and its  
591 surroundings, *Atmos. Chem. Phys.*, 11, 5031-5053, 2012.

592 Zhang, M., Gao, L., Ge, C., and Y.: Simulation of nitrate aerosol concentrations over East Asia with the  
593 model system RAMS-CMAQ, *Tellus B*, 59, 372-380, 2007.

594 Zhang, M., Xu, Y., Zhang, R., and Han, Z.: Emission and concentration distribution of black carbon  
595 aerosol in East Asia during springtime, *Chin. J. Geophys.*, 48, 55-61, 2005.

596 Zhang, M., Uno, I., Zhang, R., Han, Z., Wang, Z., and Pu, Y.: Evaluation of the Models-3 Community  
597 Multi-scale Air Quality (CMAQ) modeling system with observations obtained during the TRACE-P  
598 experiment: comparison of ozone and its related species, *Atmos. Environ.*, 40, 4874-4882, 2006.

599 Zhao, P., Zhang, X., Xu, X., and Zhao, X.: Long-term visibility trends and characteristics in the region of  
600 Beijing, Tianjin, and Hebei, China, *Atmos. Res.*, 101, 711-718, 2011.

601 Zhao, X., Zhao, P., Xu, J., Meng, W., Pu, W., Dong, F., He, D., and Shi, Q.: Analysis of a winter regional  
602 haze event and its formation mechanism in the North China Plain, *Atmos. Chem. Phys.*, 13,  
603 5685-5696, 2013.

604

605

606

607

**Table 1** Aerosol size distribution parameters in RAMS-CMAQ.

Mode	Aerosol components	$\sigma^a$	$r^b$ , $\mu\text{m}$
Aitken	ASO <sub>4</sub> <sup>c</sup> , ANO <sub>3</sub> <sup>d</sup> , ANH <sub>4</sub> <sup>e</sup> , BC <sup>f</sup> , OC <sup>g</sup>	1.7	0.015
Accumulation	ASO <sub>4</sub> , ANO <sub>3</sub> , ANH <sub>4</sub> , BC, OC, Dust, Sea salt	2.0	0.150
Coarse dust	Dust	3.0	0.300
Coarse sea salt	Sea salt	3.5	0.300

608 <sup>a</sup>  $\sigma$  is the geometric standard deviation.609 <sup>b</sup>  $r$  is the mode radius.610 <sup>c</sup> ASO<sub>4</sub> represents sulfate aerosol.611 <sup>d</sup> ANO<sub>3</sub> represents nitrate aerosol.612 <sup>e</sup> ANH<sub>4</sub> represents ammonium aerosol.613 <sup>f</sup> BC represents black carbon.614 <sup>g</sup> OC represents organic carbon.

615

616

617

618

619

620

621

622

623

624

625

626

627

628

629

630

631

632

633

634

635

636

637

638

639

640

641

642

643

644

645

646 **Table 2** Statistical summary of the comparisons of PM<sub>2.5</sub>, O<sub>3</sub>, NO<sub>2</sub>, and visibility between simulation and observations in  
 647 Beijing

		$N^a$	$C_{\text{mod}}^b$	$C_{\text{obs}}^c$	$\sigma_{\text{mod}}^d$	$\sigma_{\text{obs}}^e$	$R^f$
PM <sub>2.5</sub>	Feb	665	133.05	127.5	102.41	129.7	0.76
	Jul	663	112.78	89.92	74.09	72.98	0.43
O <sub>3</sub>	Feb	621	15.74	16.10	14.04	14.04	0.78
	Jul	630	56.59	48.75	38.44	36.21	0.74
NO <sub>2</sub>	Feb	672	33.94	44.18	17.59	24.98	0.75
	Jul	626	25.29	24.08	17.01	11.62	0.42
Visibility	Feb	672	10.97	10.78	6.53	7.27	0.76
	Jul	744	8.06	9.64	5.74	6.48	0.65

648 <sup>a</sup>Number of samples.  
 649 <sup>b</sup>Total mean of observations.  
 650 <sup>c</sup>Total mean of simulations.  
 651 <sup>d</sup>Standard deviation of observations.  
 652 <sup>e</sup>Standard deviation of simulations.  
 653 <sup>f</sup>Correlation coefficient between observation and simulation.

654  
655  
656  
657  
658  
659  
660  
661  
662  
663  
664  
665  
666  
667  
668  
669  
670  
671  
672  
673  
674  
675  
676  
677  
678  
679  
680

681 **Table 3** The number of haze days in February and July in Beijing. Also shown are the regional and temporal average  
 682 surface wind speed ( $\text{m s}^{-1}$ ), visibility (km), relative humidity (%), and mass concentrations ( $\mu\text{g m}^{-3}$ ) of sulfate, nitrate,  
 683 ammonium, BC, OC, and  $\text{PM}_{2.5}$  during the haze days in February and July, respectively, in Beijing.

variables	February	July
Number of haze days	7	13
Wind speed	3.13	3.41
Visibility	6.22	5.73
Relative humidity	55.80	74.32
Sulfate	37.99	52.32
Nitrate	54.78	48.37
Ammonium	30.15	33.60
BC	13.29	4.31
OC	19.51	5.16
$\text{PM}_{2.5}$	174.26	148.32

684  
 685  
 686  
 687  
 688  
 689  
 690  
 691  
 692  
 693  
 694  
 695  
 696  
 697  
 698  
 699  
 700  
 701  
 702  
 703  
 704  
 705  
 706  
 707  
 708  
 709  
 710  
 711  
 712  
 713

714  
715  
  
716  
717  
718  
719  
720  
721  
722  
723  
724  
725  
726  
727  
728  
729  
730  
731  
732  
733  
734  
735  
736  
737  
738  
739  
740  
741  
742  
743  
744  
745  
746  
747  
748  
749  
750  
751  
752  
753  
754

**Table 4** Regional and monthly average extinction contribution ratios (%) of sulfate, nitrate, ammonium, BC, OC, and other aerosols (dust, sea salt, and unspecified anthropogenic mass) in February and July, respectively, in Beijing.

	Sulfate	Nitrate	Ammonium	BC	OC	Others
Feb	22.73	29.69	17.13	5.05	13.22	12.18
Jul	39.31	24.77	21.88	0.33	3.96	9.74

755  
756  
  
757  
758  
759  
760  
761  
762  
763  
764  
765  
766  
767  
768  
769  
770  
771  
772  
773  
774  
775  
776  
777  
778  
779  
780  
781  
782  
783  
784  
785  
786  
787  
788  
789  
790  
791  
792  
793  
794  
795

**Table 5** Regional and monthly average mass concentration threshold of PM<sub>2.5</sub> (μg m<sup>-3</sup>) under different relative humidities from the sensitivity test in Beijing.

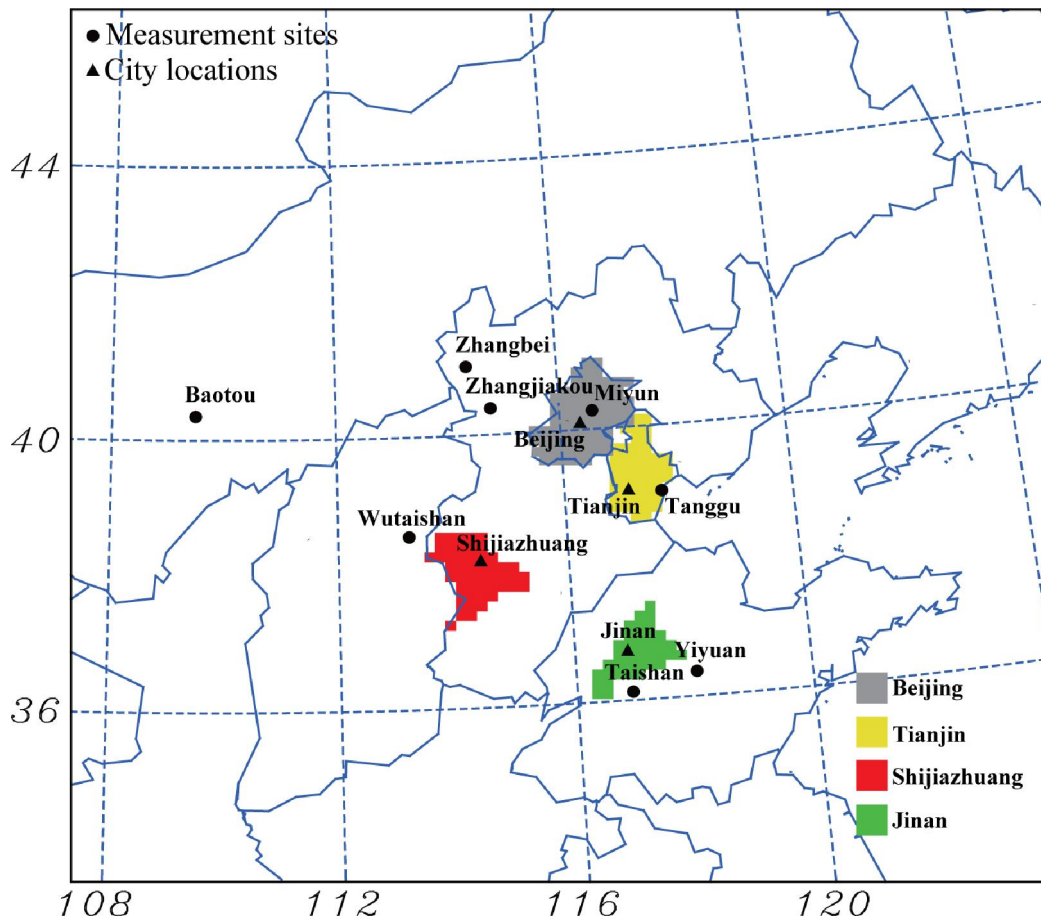
	70%	75%	80%	85%	88%	89%	90%
Feb	82.08	78.01	72.34	64.54	58.27	55.87	53.22
Jul	83.08	78.64	72.46	64.01	57.35	54.82	52.10



796  
797  
  
798  
799  
800  
801  
802  
803  
804  
805  
806  
807  
808  
809  
810  
811  
812  
813  
814  
815  
816  
817  
818  
819  
820  
821  
822  
823  
824  
825  
826

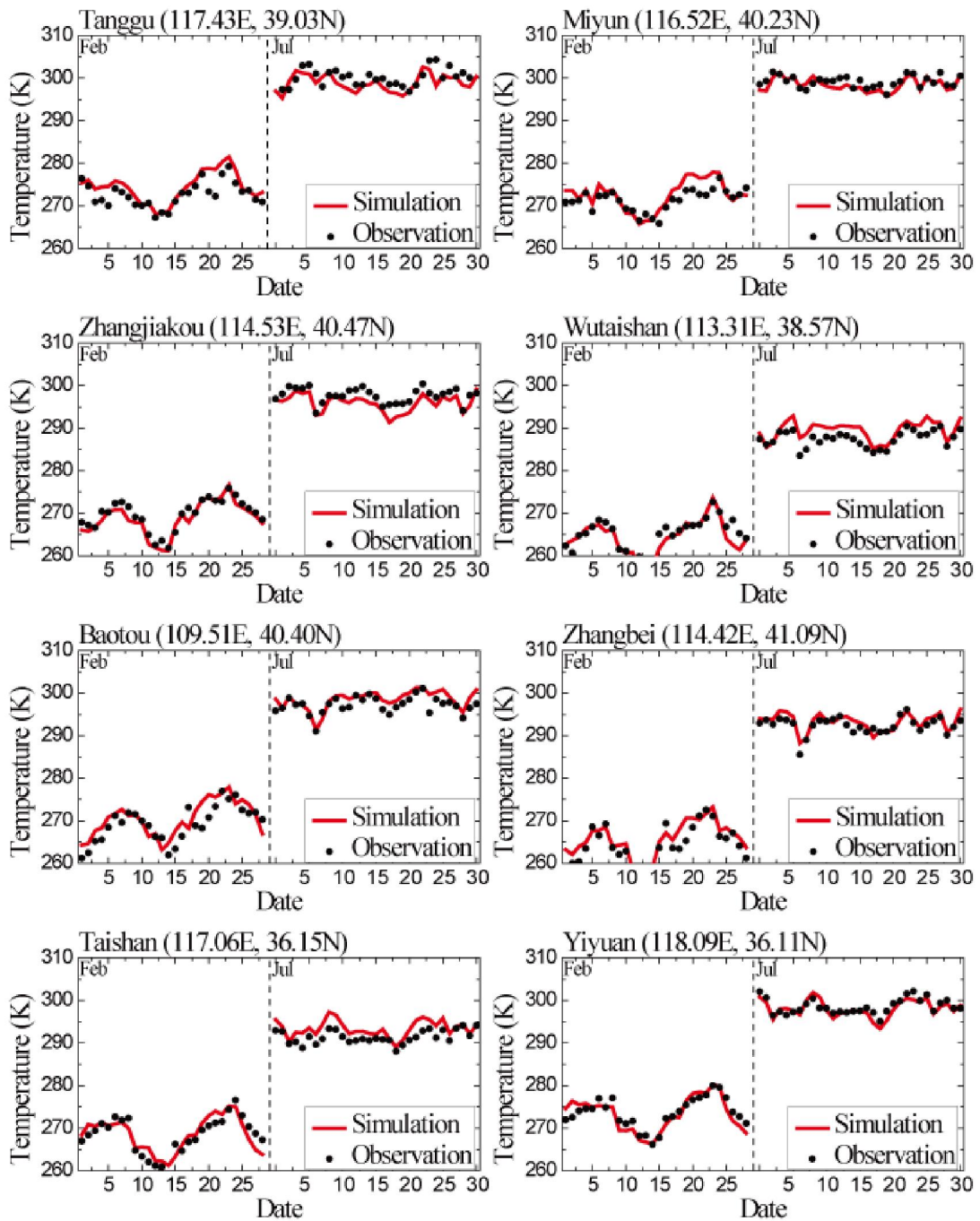
**Table 6** Parameters of the exponential fit for the regional and monthly average PM<sub>2.5</sub> mass concentration thresholds of haze occurrence in Beijing.

	<i>a</i>	<i>b</i>	<i>c</i>	R-square
Feb	96.3276	-0.4859	0.0486	0.9997
Jul	96.8810	0.7367	0.0483	0.9997



827  
828  
829  
830  
831  
832  
833  
834  
835  
836  
837  
838  
839  
840  
841  
842

**Fig. 1.** Geographic location of API monitoring cities and CNMC measurement stations in the model domain. The gray, yellow, red, and green areas represent the districts of Beijing, Tianjin, Shijiazhuang, and Jinan, respectively.



843  
 844 **Fig. 2.** Observed (black circles) and modeled (solid red lines) daily average temperatures (K) at 8 stations in February  
 845 and July 2011.  
 846

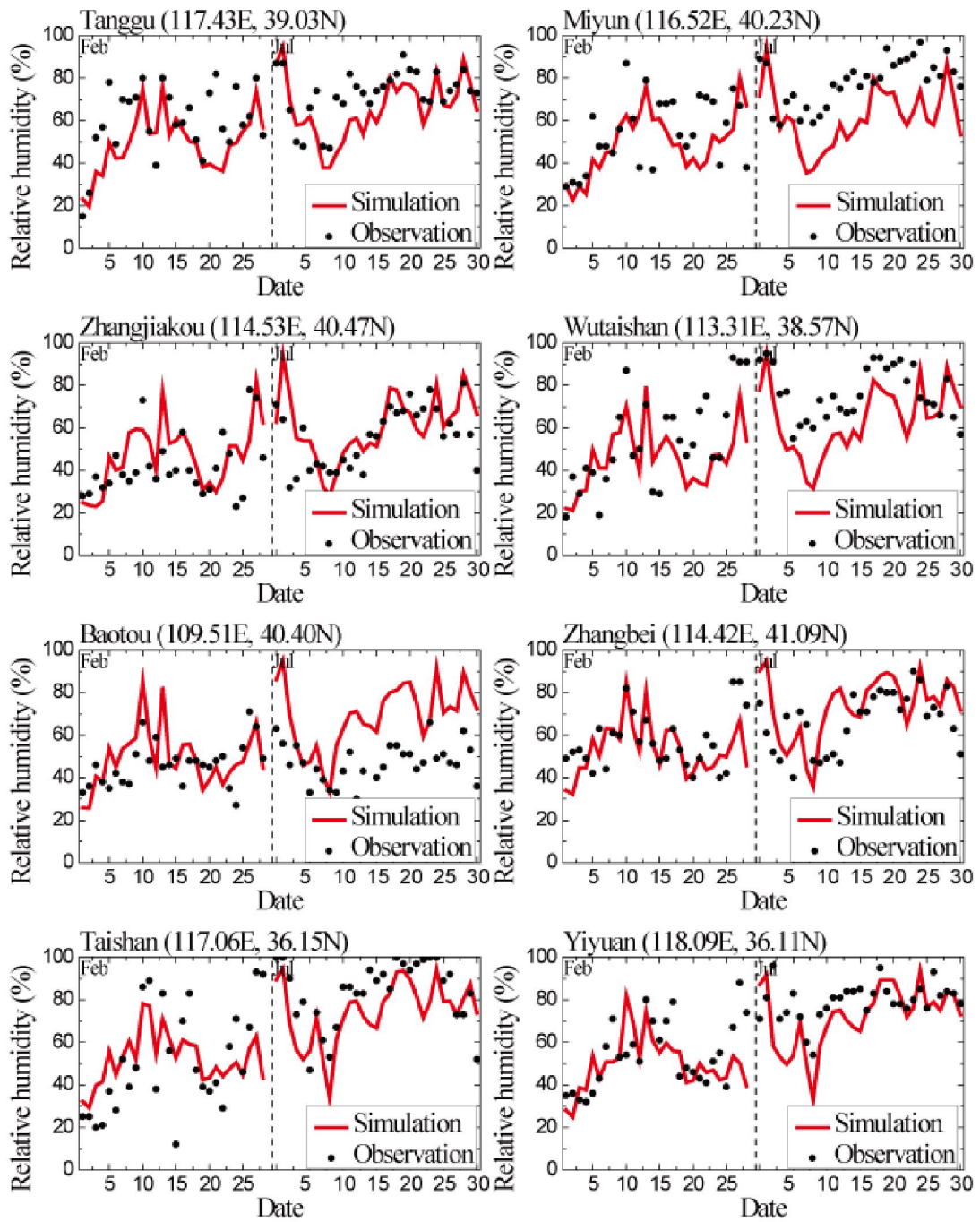


Fig. 3. Same as Fig. 2 but for relative humidity (%).

852  
853  
854  
855  
856  
857  
858  
859

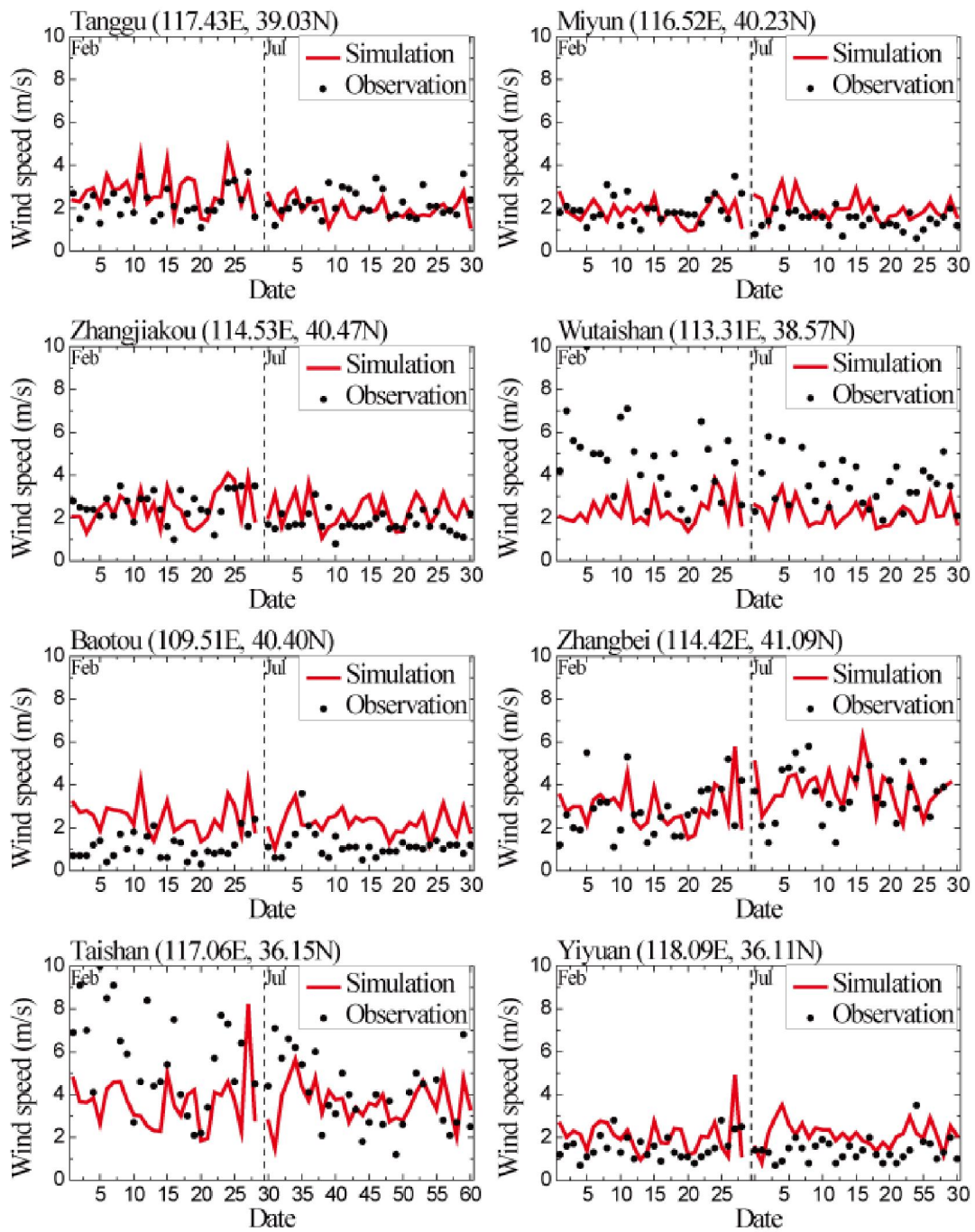


Fig. 4. Same as Fig. 2 but for wind speed ( $\text{m s}^{-1}$ ).

860  
861  
862  
863  
864  
865  
866  
867

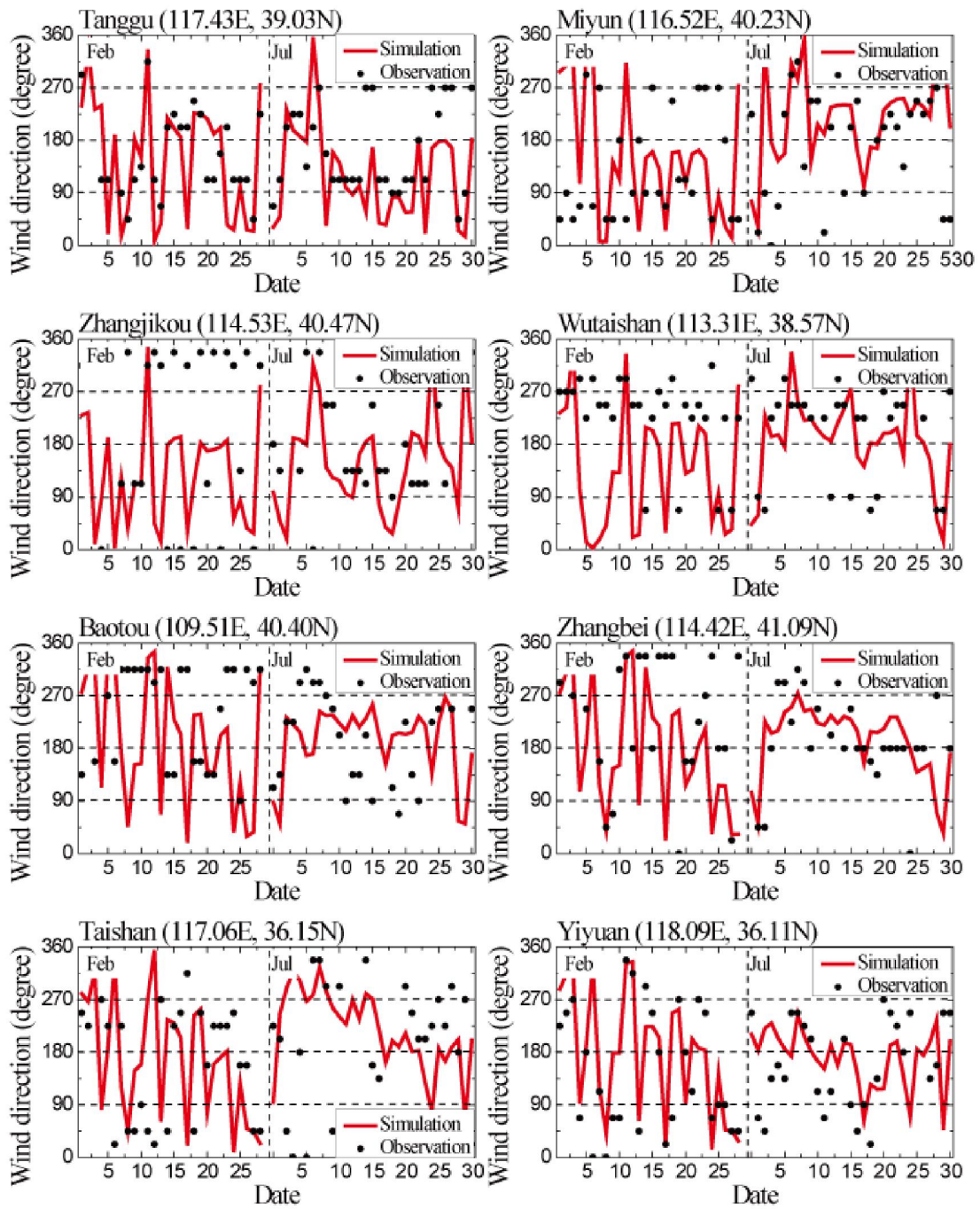
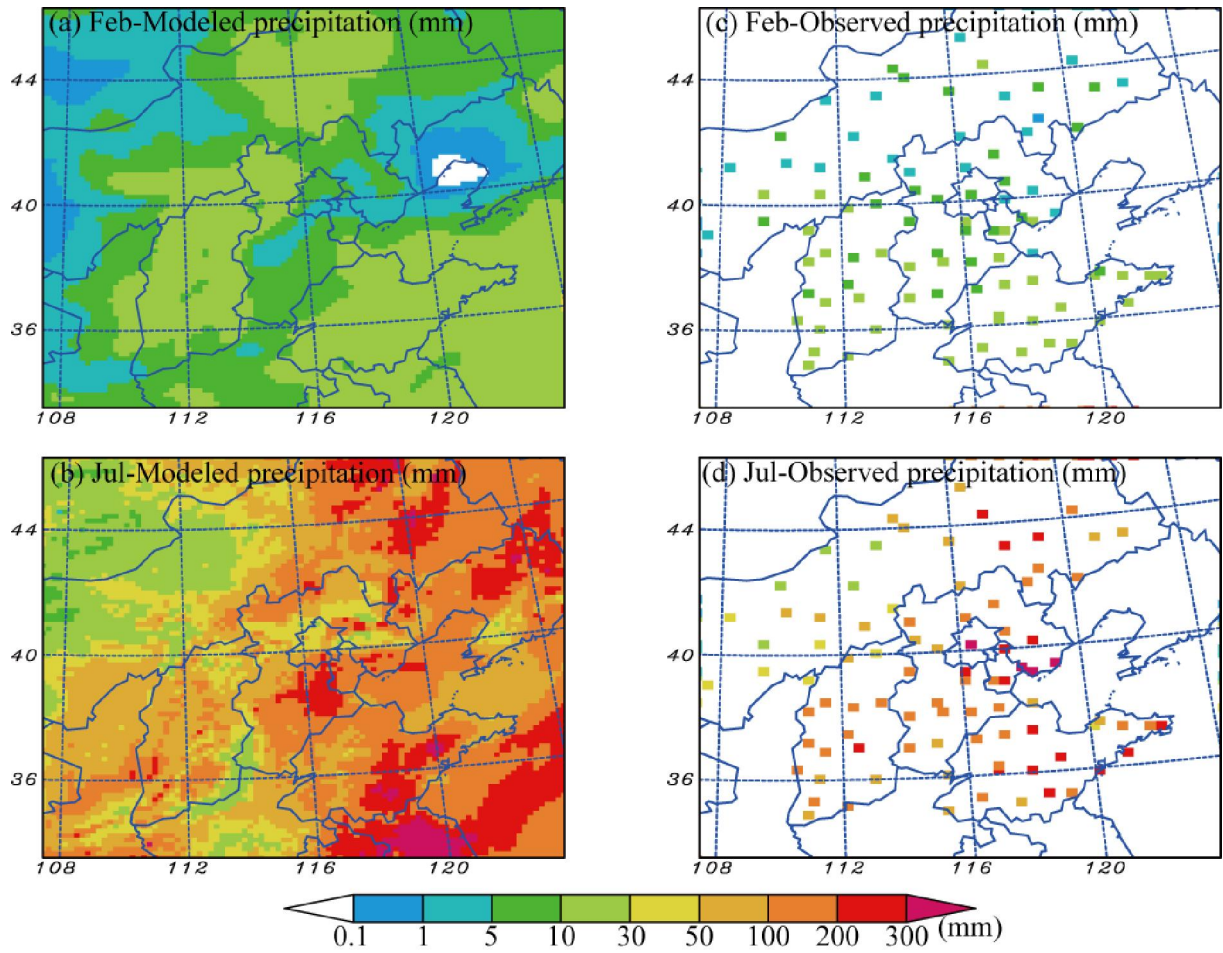
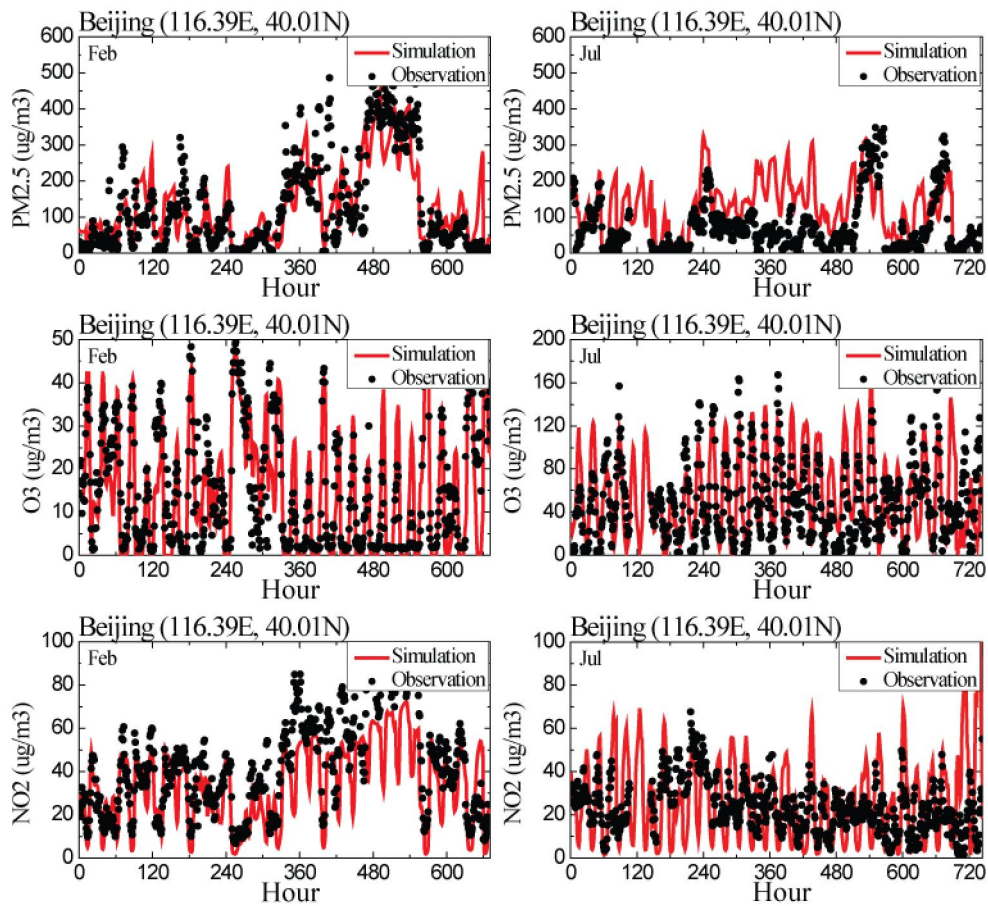


Fig. 5. Same as Fig. 2 but for wind direction (°).

868  
 869  
 870  
 871  
 872  
 873  
 874  
 875

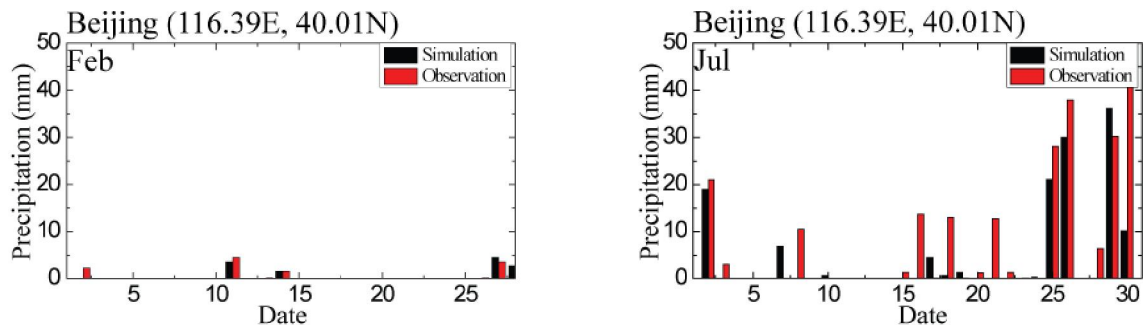


**Fig. 6.** The horizontal distributions of the monthly precipitation (mm) from simulation and surface observation data in February and July 2011 over the NCP.



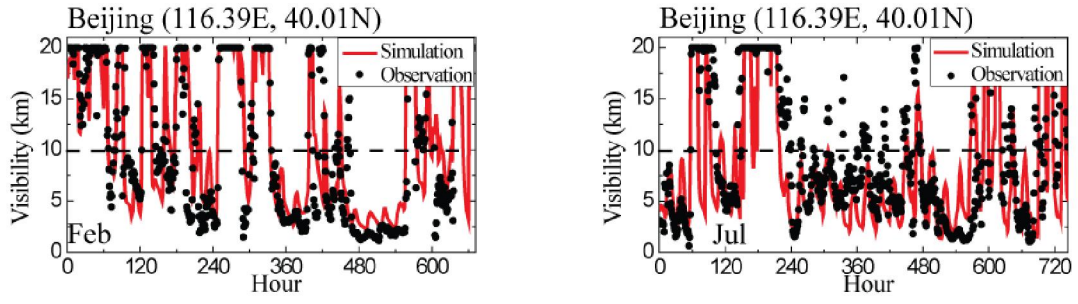
891  
 892 **Fig. 7.** Observed (black circles) and modeled (red solid lines) hourly mass concentrations ( $\mu\text{g m}^{-3}$ ) of  $\text{PM}_{2.5}$ ,  $\text{O}_3$ , and  $\text{NO}_2$   
 893 in February and July 2011 at Beijing.  
 894





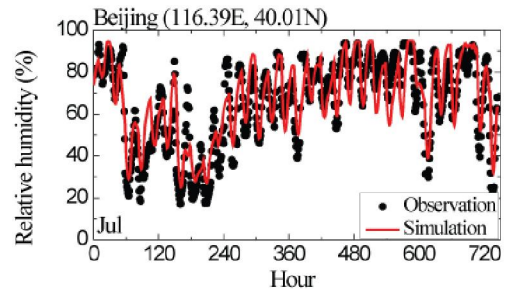
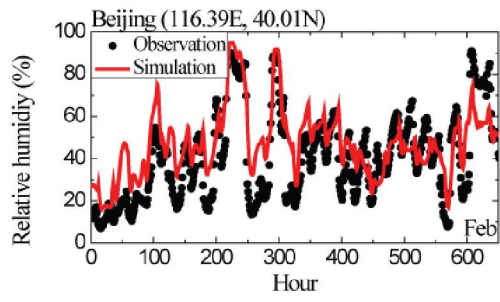
**Fig. 8.** Observed and modeled daily precipitation (mm) in February and July 2011 in Beijing

907  
908  
909  
910  
911  
912  
913  
914  
915  
916  
917  
918  
919  
920  
921  
922  
923  
924  
925  
926  
927  
928  
929  
930  
931



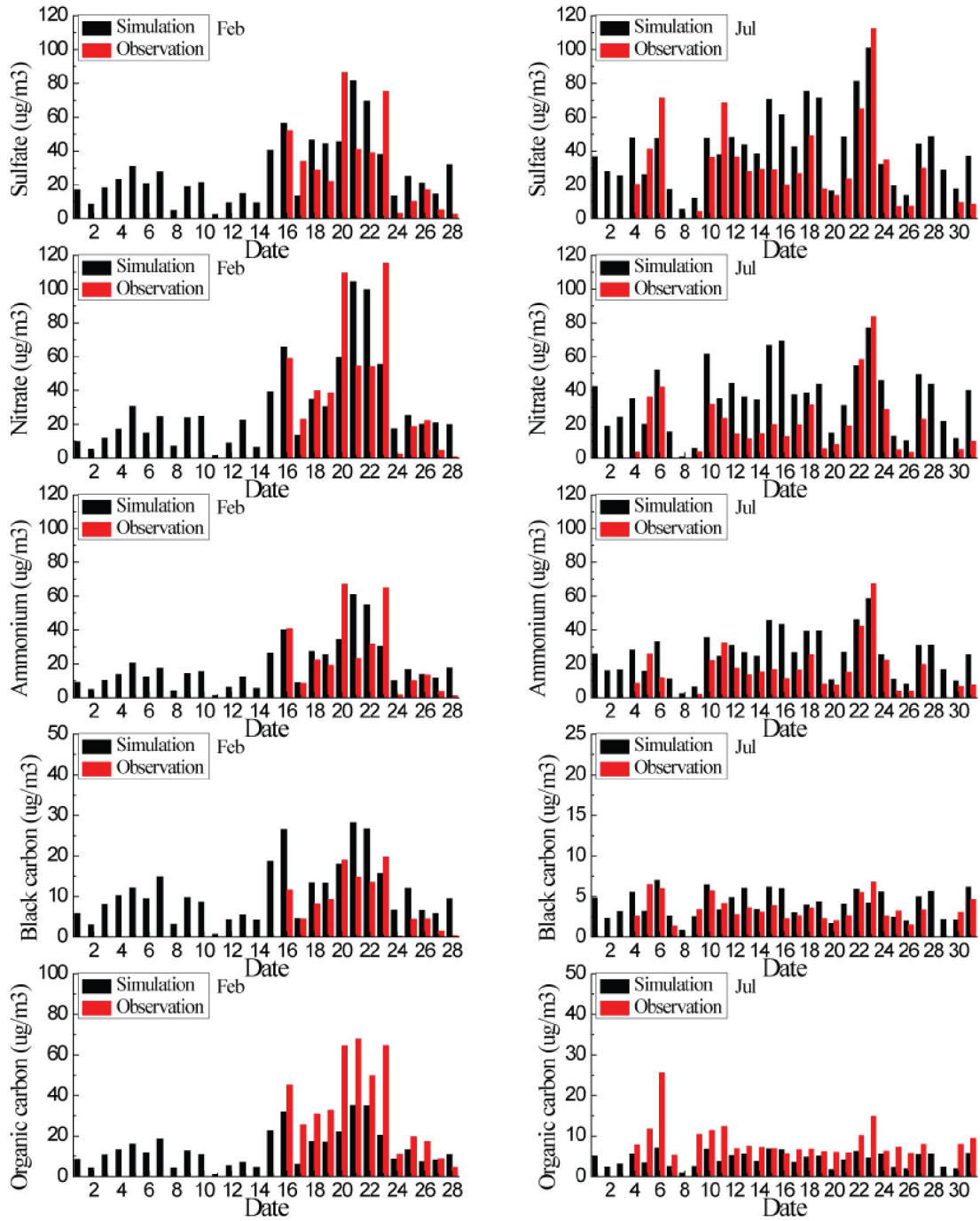
**Fig. 9.** Observed (black circles) and modeled (solid red lines) hourly visibility (km) in February and July 2011 in Beijing.

932  
 933  
 934  
 935  
 936  
 937  
 938  
 939  
 940  
 941  
 942  
 943  
 944  
 945  
 946  
 947  
 948  
 949  
 950  
 951  
 952  
 953  
 954  
 955  
 956

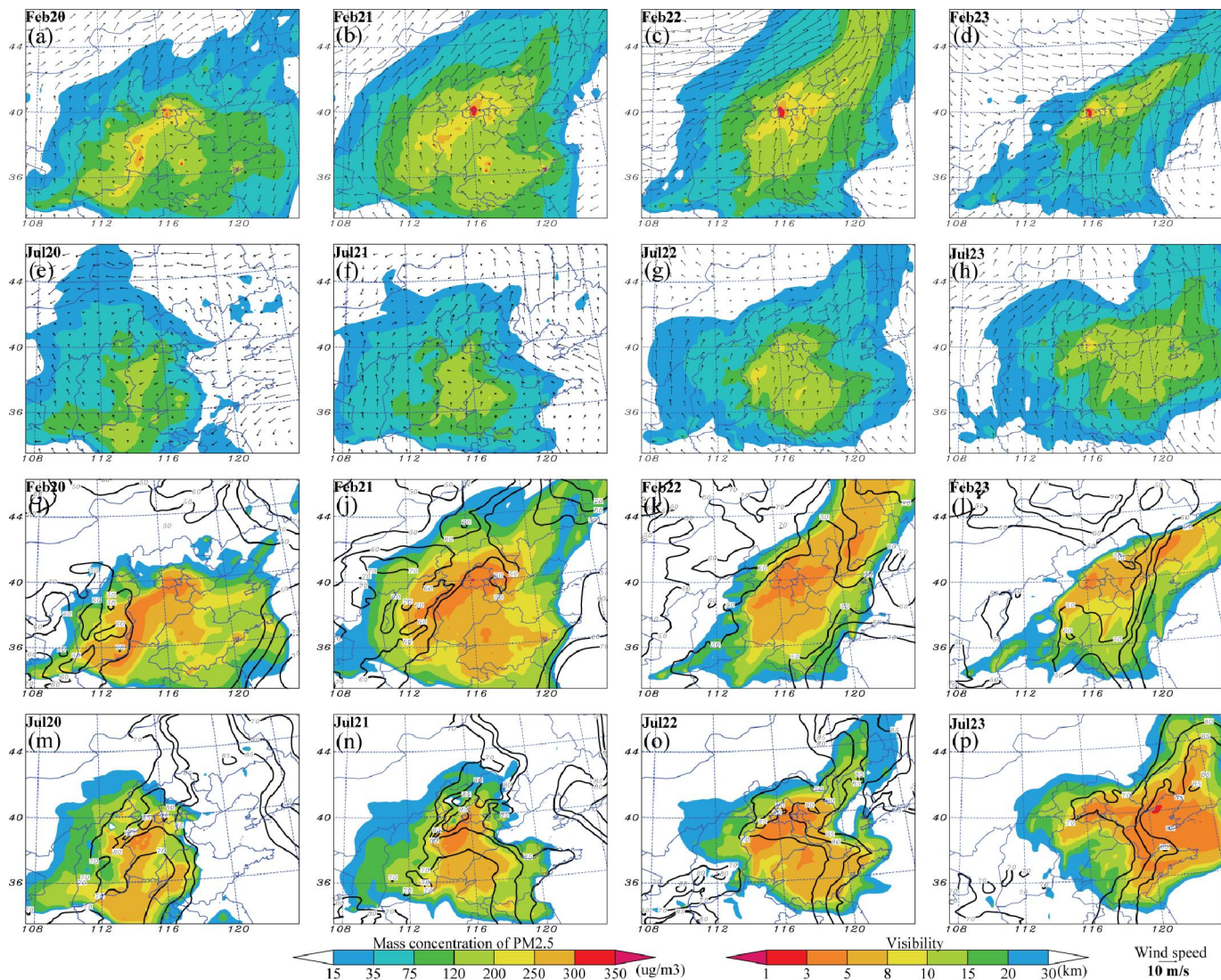


**Fig. 10.** Observed (black circles) and modeled (solid red lines) hourly relative humidity (%) in February and July 2011 in Beijing.

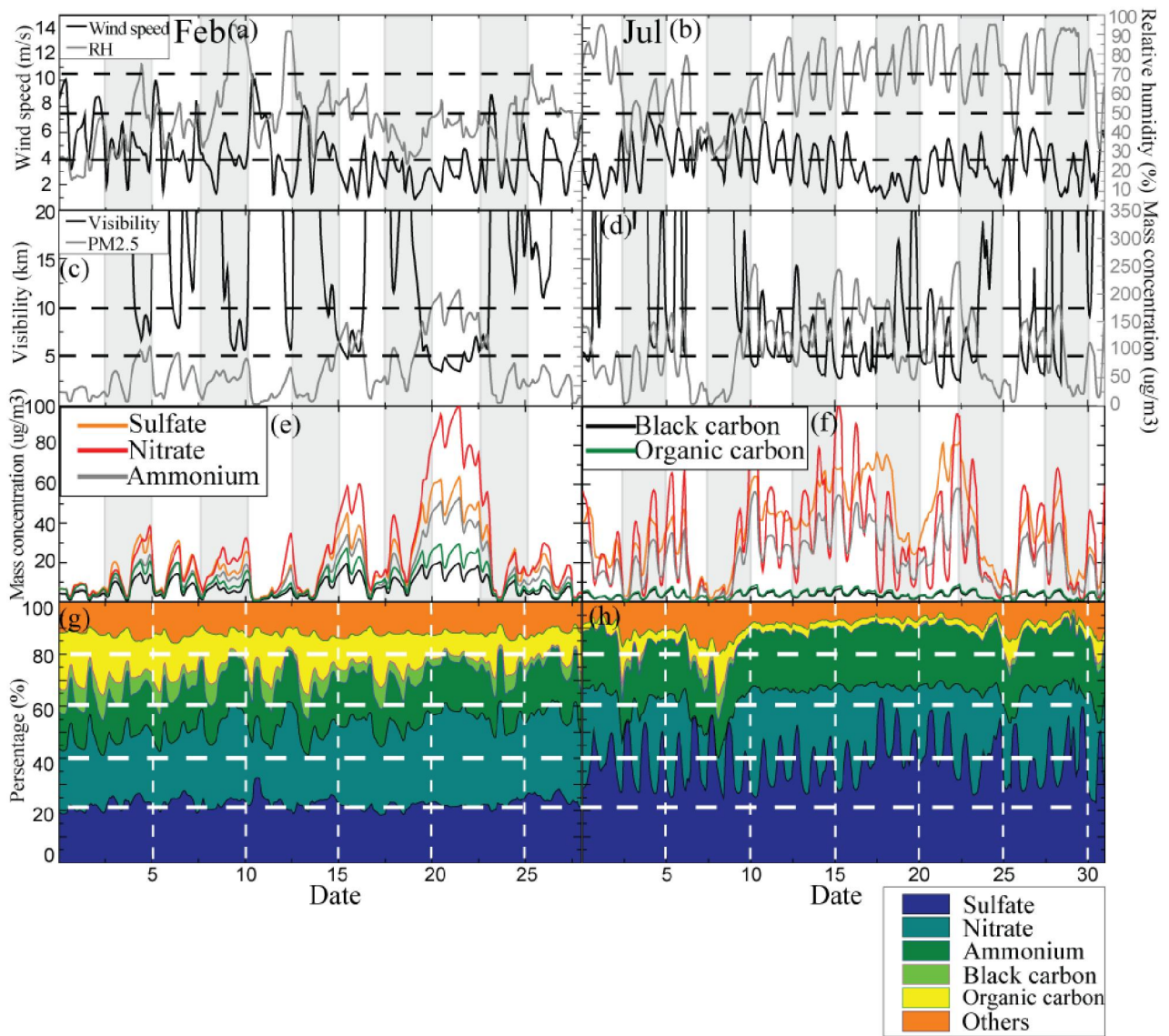
957  
958  
959  
960  
961  
962  
963  
964  
965  
966  
967  
968  
969  
970  
971  
972  
973  
974  
975  
976  
977  
978  
979  
980  
981



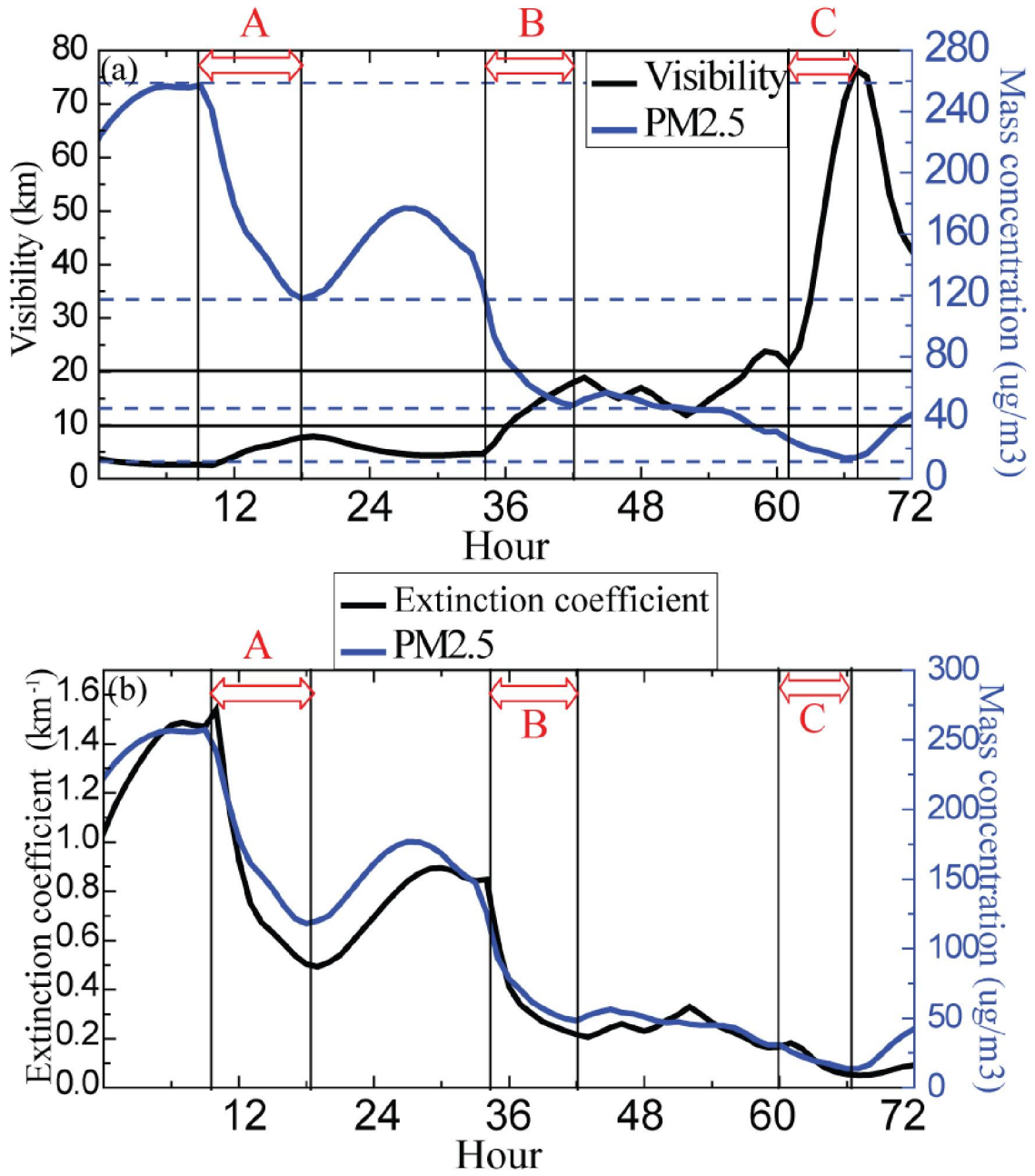
982  
 983 **Fig. 11.** Observed (red bars) and modeled (black bars) daily average mass concentrations ( $\mu\text{g m}^{-3}$ ) of sulfate, nitrate,  
 984 ammonium, black carbon, and organic carbon in February and July 2011 in Beijing.  
 985  
 986  
 987  
 988  
 989



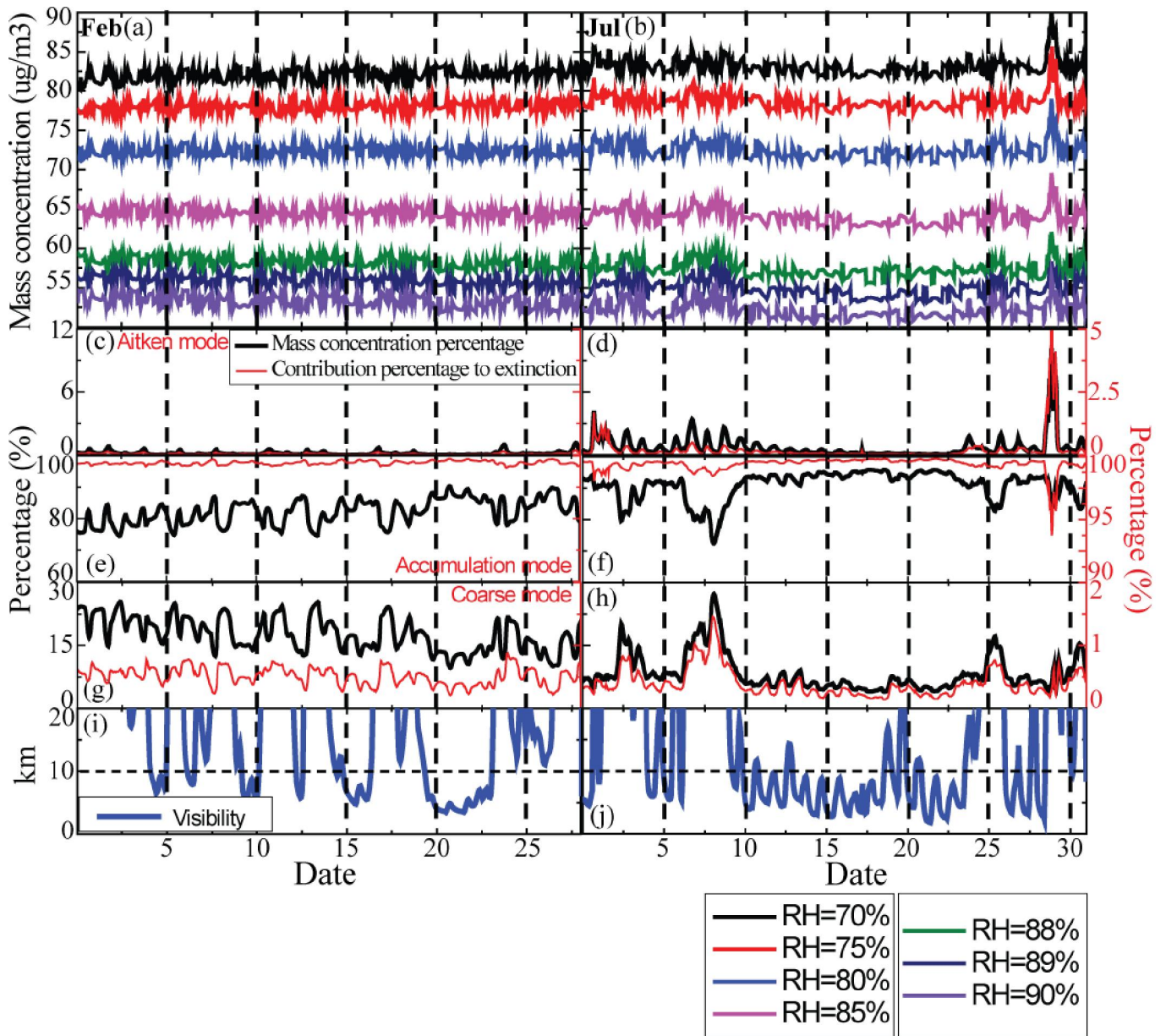
990  
 991 **Fig. 12.** The horizontal distributions of the daily average mass concentration of PM<sub>2.5</sub> ( $\mu\text{g m}^{-3}$ ; a-h) and visibility (km; i-p)  
 992 from February 20 to 23 and July 20 to 23 over the NCP. Also shown are the wind field (arrows) and the relative humidity  
 993 (%; black contour lines).



1002  
 1003 **Fig. 13.** Time series of the regional average surface wind speed ( $\text{m s}^{-1}$ ), relative humidity (%), visibility (km), and  $\text{PM}_{2.5}$   
 1004 mass concentrations ( $\mu\text{g m}^{-3}$ ) in February and July in Beijing (a-d). Also shown are the mass concentrations ( $\mu\text{g m}^{-3}$ ; e-f)  
 1005 and extinction contribution ratios (g-h) of sulfate, nitrate, ammonium, BC, OC, and other aerosols (dust, sea salt, and  
 1006 unspecified anthropogenic mass).

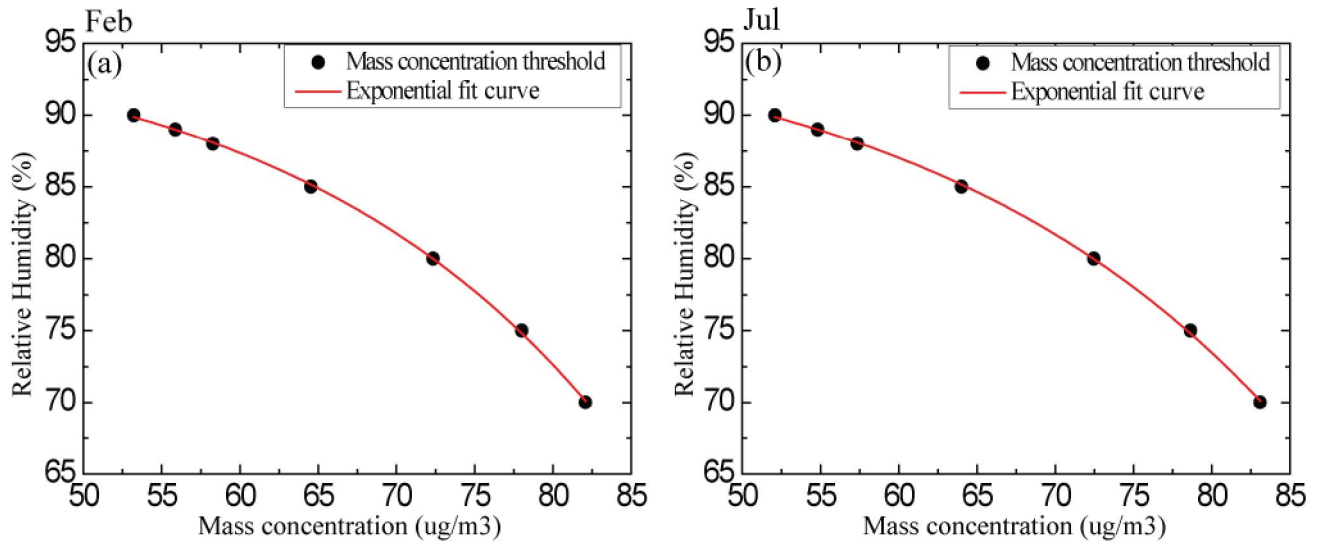


1013  
 1014 **Fig. 14.** The time series of the regional average visibility (km) vs. mass concentration of PM<sub>2.5</sub> (a), and extinction  
 1015 coefficient vs. mass concentration of PM<sub>2.5</sub> (b) from July 23 to 25 in Beijing.  
 1016  
 1017  
 1018  
 1019  
 1020  
 1021



1022  
 1023 **Fig. 15.** Time series of the regional average mass concentration threshold of  $PM_{2.5}$  under different relative humidity (%)  
 1024 from the sensitivity tests in February and July in Beijing. Also shown are the mass ratios and extinction contribution  
 1025 ratios (%) of the Aitken, accumulation, and coarse modes, and visibility (km).  
 1026  
 1027  
 1028  
 1029  
 1030  
 1031  
 1032





**Fig. 16.** The relationship between regional and monthly average  $\text{PM}_{2.5}$  mass concentration threshold of haze occurrence and relative humidity in Beijing. Also shown are the exponential fitted curves.

1033  
 1034  
 1035  
 1036  
 1037  
 1038  
 1039