



Modeling kinetic partitioning of SOA and size distribution dynamics

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Modeling kinetic partitioning of secondary organic aerosol and size distribution dynamics: representing effects of volatility, phase state, and particle-phase reaction

R. A. Zaveri¹, R. C. Easter¹, J. E. Shilling¹, and J. H. Seinfeld²

¹Atmospheric Sciences and Global Change Division, Pacific Northwest National Laboratory, Richland, WA, USA

²Division of Chemistry and Chemical Engineering and Division of Engineering and Applied Science, California Institute of Technology, Pasadena, CA, USA

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Correspondence to: R. A. Zaveri (rahul.zaveri@pnnl.gov)

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Abstract

This paper describes and evaluates a new formulation for modeling kinetic gas-particle partitioning of secondary organic aerosol (SOA) that takes into account diffusion and chemical reaction within the particle phase. The new formulation uses a combination of:

- (a) an analytical quasi-steady-state treatment for the diffusion-reaction process within the particle phase for fast-reacting organic solutes, and
- (b) a two-film theory approach for slow- and non-reacting solutes.

The formulation is amenable for use in regional and global atmospheric models, although it currently awaits specification of the actual species and particle-phase reactions that are important for SOA formation. Here, the formulation is applied within the framework of the computationally efficient Model for Simulating Aerosol Interactions and Chemistry (MOSAIC) to investigate the competitive growth dynamics of the Aitken and accumulation mode particles. Results show that the timescale of SOA partitioning and the associated size distribution dynamics depend on the complex interplay between organic solute volatility, particle-phase bulk diffusivity, and particle-phase reactivity (as exemplified by a pseudo-first-order reaction rate constant), each of which can vary over several orders of magnitude. In general, the timescale of SOA partitioning increases with increase in volatility and decrease in bulk diffusivity and rate constant. At the same time, the shape of the aerosol size distribution displays appreciable narrowing with decrease in volatility and bulk diffusivity and increase in rate constant. A proper representation of these physicochemical processes and parameters are needed in the next generation models to reliably predict not only the total SOA mass, but also its composition and number size distribution, all of which together determine its overall optical and cloud-nucleating properties.

1 Introduction

Submicron sized atmospheric aerosol particles are typically composed of ammonium, sulfate, nitrate, black carbon, organics, sea salt, mineral dust, and water that are of

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ten internally mixed with each other in varying proportions. Depending on their dry state composition and overall hygroscopicity, aerosol particles in the size range 0.03–0.1 μm (dry diameter) and larger may act as cloud condensation nuclei (CCN) (Dusek et al., 2006; Gunthe et al., 2009, 2011) while those larger than 0.1 μm (wet diameter) efficiently scatter solar radiation. Aerosol number and composition size distributions, therefore, together hold the key to determining its overall climate-relevant properties.

Organic compounds constitute 20–90 % of the submicron aerosol mass and are thought to play a vital role in both the direct and indirect aerosol radiative forcing of climate (Kanakidou et al., 2005). While primary organic aerosols (POA) from fossil fuel combustion and biomass burning are directly emitted into the submicron size range, the dominant source of organic aerosols is secondary, which involves gas-to-particle conversion of many different volatile organic compounds (VOCs) of both anthropogenic and biogenic origin (Zhang et al., 2007). Furthermore, biogenic VOCs are estimated to be the dominant source of secondary organic aerosol (SOA), but their formation appears to be strongly influenced by anthropogenic emissions (Weber et al., 2007; Hoyle et al., 2011; Shilling et al., 2013). Organic vapors are also implicated in facilitating new particle formation initiated by sulfuric acid (Kulmala et al., 2004; Paasonen et al., 2010; Kuang et al., 2012) and are found to play a crucial role in the subsequent growth of the nanoparticles (Smith et al., 2008; Pierce et al., 2011, 2012; Riipinen et al., 2011; Winkler et al., 2012). Thus, the majority of the optically- and CCN-active particles are produced through the growth of smaller particles by condensation of SOA species (Riipinen et al., 2012). It is therefore necessary that climate models be able to accurately simulate not just the total organic mass loading, but also the evolution of aerosol number and composition size distributions resulting from SOA formation.

It is broadly understood that, in cloud-free air, SOA forms via three possible mechanisms: (1) effectively irreversible condensation of very low volatility organic vapors produced by gas-phase oxidation (Donahue et al., 2011; Pierce et al., 2011); (2) volume-controlled reversible absorption of semi-volatile organic vapors into pre-existing aerosol according to Raoult's law (Pankow, 1994); and (3) absorption of semi-volatile

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and volatile organic vapors into pre-existing aerosol followed by particle-phase reactions to form effectively non-volatile products such as organic salts (Smith et al., 2010), oligomers and other high molecular weight oxidation products (Gao et al., 2004; Kalberer et al., 2004; Heaton et al., 2007; Nozière et al., 2007; Ervens et al., 2010; Wang et al., 2010; Hall and Johnston, 2011), hemiacetals (Kroll et al., 2008; Ziemann et al., 2012; Shiraiwa et al., 2013), and organosulfates (Surratt et al., 2007; Zaveri et al., 2010). While aqueous-phase chemistry in cloud droplets is also a potential source of SOA (Carlton et al., 2008; Ervens et al., 2008; Mouchel-Vallon et al., 2013), this route is not considered in the present study. Several recent studies also indicate that the phase state of SOA may be viscous semi-solids under dry and moderate relative humidity conditions (Virtanen et al., 2010; Vaden et al., 2011; Saukko et al., 2012), with very low particle-phase bulk diffusivities (Abramson et al., 2013; Renbaum-Wolff et al., 2013). The timescales of SOA partitioning (Shiraiwa and Seinfeld, 2012b) and the resulting aerosol size distributions from these three mechanisms can be quite different, and the particle phase state is expected to modulate the growth dynamics as well.

Riipinen et al. (2011) analyzed the evolution of ambient aerosol size distributions with a simplified model consisting of mechanisms #1 and #2 for liquid particles and concluded that both mechanisms were roughly equally needed to explain the observed aerosol growth. Perraud et al. (2012) studied the gas-particle partitioning of organic nitrate vapors formed from simultaneous oxidation of α -pinene by O_3 and NO_3 in a flow tube reactor. Their model analysis suggested that, despite being semi-volatile, the organic nitrate species had effectively irreversibly condensed (mechanism #1) as their adsorbed layers were continuously “buried” in presumably semi-solid particles by other incoming organic vapors. In a theoretical study, Zhang et al. (2012) contrasted the aerosol size distributions produced by mechanisms #1 and #2 for liquid particles and illustrated the roles of solute volatility and vapor source rate in shaping the size distribution via mechanism #2. In another theoretical study, Shiraiwa and Seinfeld (2012b) used the detailed multilayer kinetic flux model KM-GAP (Shiraiwa et al., 2012a) to investigate the effect of phase state on SOA partitioning. They showed that the timescale

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for gas-particle equilibration via mechanism #2 increases from hours to days for organic aerosol associated with semi-solid particles, low volatility, large particle size, and low mass loadings. More recently, Shiraiwa et al. (2013a) studied SOA formation from photooxidation of dodecane in the presence of dry ammonium sulfate seed particles in an environmental chamber. Their analysis of the observed aerosol size distribution evolution with the KM-GAP model revealed the presence of particle-phase reactions (i.e., mechanism #3), which contributed more than half of the SOA mass, with the rest formed via mechanism #2. Furthermore, the physical state of the SOA was assumed to be semi-solid with an average bulk diffusivity of $10^{-12} \text{ cm}^2 \text{ s}^{-1}$, and the particle-phase reactions were predicted to occur mainly on the surface.

While valuable insights into the effect of phase state on SOA formation have emerged from several recent studies, a comprehensive, quantitative analysis of the effects of organic solute volatility, phase state, and particle-phase reaction on aerosol growth dynamics has not yet been performed. Additionally, there is a lack of a kinetic SOA partitioning treatment for semi-solids (with particle-phase chemical reactions) that is amenable for use in regional and global atmospheric models. The present work addresses both these topics. The paper is structured as follows. In Sect. 2, we examine the dynamics of diffusion and reaction in a spherical particle with an analytical solution to the problem. In Sect. 3, we extend the MOSAIC (Model for Simulating Aerosol Interactions and Chemistry) aerosol model (Zaveri et al., 2008) to include a new formulation for kinetic gas-particle partitioning of SOA and evaluate it against a rigorous model based on the finite difference approach. The new formulation uses a combination of: (a) an analytical quasi-steady-state treatment for the diffusion-reaction process within the particle phase for fast-reacting species, and (b) a two-film theory approach for slow- and non-reacting organic solutes. The formulation is amenable for eventual use in regional and global climate models, although it currently awaits specification of the actual particle-phase reactions that are important for SOA formation. In Sect. 4, we apply the model to evaluate the timescale of SOA partitioning and the associated evolution of the number and composition size distributions for a range of solute volatilities,

bulk diffusivities, and particle-phase reaction rates. We close with a summary of our findings and their implications.

2 Dynamics of diffusion and reaction within a particle

Consider an organic solute i that diffuses from the gas phase to a single spherical organic aerosol particle and reacts irreversibly with a pseudo-first-order rate constant $k_c(\text{s}^{-1})$ as it diffuses inside the particle. This process is illustrated in Fig. 1 using three species (P_1 , P_2 , and P_3) for simplicity. The organic solute P_1 diffuses and reacts to form a non-volatile species P_2 inside an organic particle (of radius R_p) that is initially composed of a non-volatile organic species P_3 . The solute's gas-phase concentrations far away from the particle (i.e., in the bulk gas-phase) and just above the particle surface are \bar{C}_g and \bar{C}_g^s (mol cm^{-3} (air)), respectively. The solute's particle-phase concentration just inside the particle surface and at any location in the bulk of the particle are denoted as A^s and A (mol cm^{-3} (particle)), respectively. The gas- and particle-phase diffusivities of the solute are D_g and D_b ($\text{cm}^2 \text{s}^{-1}$), respectively.

In this section we shall focus on the dynamics of diffusion and reaction *inside* the particle. In order to derive the timescales relevant to this problem, the particle, initially free of the organic solute (i.e., at time $t = 0$), is assumed to be exposed to a constant concentration just inside the particle surface, A_i^s , at all times $t > 0$ (this assumption will be relaxed in Sect. 3 where we will relate the temporally changing gas-phase concentration of the solute to its particle-phase concentration). The transient partial differential equation describing the particle-phase concentration $A_i(r, t)$ as a function of radius r and time t can be written as (Schwartz and Freiberg, 1981):

$$\frac{\partial A_i(r, t)}{\partial t} = D_{b,i} \frac{1}{r^2} \frac{\partial}{\partial r} \left(r^2 \frac{\partial A_i(r, t)}{\partial r} \right) - k_{c,i} A_i(r, t). \quad (1)$$

The particle is assumed to be spherically symmetrical with respect to the concentration profiles of the organic solute in the particle at any given time, so the concentration

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gradient at the center of the particle (i.e., $r = 0$) is always zero. These assumptions give rise to the following initial and boundary conditions:

$$\text{I.C.: } A_i(r, 0) = 0, \quad (2a)$$

$$\text{B.C. 1: } A_i(R_p, t) = A_i^s, \quad (2b)$$

$$5 \quad \text{B.C. 2: } \frac{\partial A_i(0, t)}{\partial r} = 0. \quad (2c)$$

Equation (1) with conditions (Eq. 2) can be analytically solved with the method of separation of variables and Fourier series to obtain the following result:

$$\frac{A_i(r, t)}{A_i^s} = \frac{R_p}{r} \frac{\sinh(q_i r / R_p)}{\sinh(q_i)} + \quad (3)$$

$$10 \quad \frac{2R_p}{\pi r} \sum_{n=1}^{\infty} \frac{(-1)^n n \sin(n\pi r / R_p)}{(q_i / \pi)^2 + n^2} \exp \left\{ - \left(k_{c,i} + \frac{n^2 \pi^2 D_{b,i}}{R_p^2} \right) t \right\}$$

where q_i is a dimensionless diffusion-reaction parameter defined as:

$$q_i = R_p \sqrt{\frac{k_{c,i}}{D_{b,i}}}. \quad (4)$$

15 It should be noted that this solution assumes that R_p remains constant with time, so diffusion of additional material into the particle is relatively small (this assumption will also be relaxed in Sect. 3). Furthermore, the timescale for diffusion of the dissolved solute i in the particle, τ_{da} , and the timescale for chemical reaction, τ_c (Seinfeld and Pandis, 2006) are defined as:

$$\tau_{da, i} = \frac{R_p^2}{\pi^2 D_{b,i}}, \quad (5)$$

$$20 \quad \tau_{c,i} = \frac{1}{k_{c,i}}. \quad (6)$$

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The model described by these equations has been applied to investigate mass transfer limitation to the rate of SO₂ oxidation in cloud droplets (Schwartz and Freiberg, 1981; Shi and Seinfeld, 1991), for which the droplets typically exceed 10 μm diameter, with the aqueous-phase diffusivity about 10⁻⁵ cm² s⁻¹. Here we apply this model to analyze the effects of particle-phase reactions in organic particles of sizes ranging from ~ 10⁻³ to 1 μm diameter, with D_b values ranging from < 10⁻¹⁸ to 10⁻⁵ cm² s⁻¹ (Renbaum-Wolff et al., 2013). Since the actual particle-phase reactions of various organic species and the associated rate constants are still not well defined, we use a the pseudo-first-order reaction as a proxy and vary its rate constant k_c over several orders of magnitude (10⁻⁵ to 10⁻¹ s⁻¹) to examine its effect on the dynamics of particle growth.

The right-hand-side of Eq. (3) comprises two terms. The first term is the concentration profile at steady state with the surface concentration, while the second term describes the temporal evolution of the concentration profile. At steady state, the transient term disappears for $t \gg \tau_{da}$ and τ_c . Figure 2 illustrates the relative effects of bulk diffusivity and reaction rate constant on the temporal evolution of the diffusing solute concentration profiles within a particle of diameter $D_p = 0.1$ μm. The top row represents a liquid organic particle with a rather high bulk diffusivity, $D_b = 10^{-6}$ cm² s⁻¹, with (a) no reaction ($k_c = 0$), and (b) a modest reaction rate constant, $k_c = 5 \times 10^{-4}$ s⁻¹. In case (a), $\tau_{da} = 2.5$ μs, and the solute attains a uniform steady-state concentration profile across the particle radius in a little over 8 μs (i.e., about 4 τ_{da}). The temporal evolution of the concentration profiles in case (b) appears to be identical to case (a) despite the presence of a chemical reaction, because τ_{da} is 2.5 μs but $\tau_c = 2000$ s, i.e., diffusion occurs much more rapidly than reaction. In contrast, the bottom row represents a semi-solid organic particle, $D_b = 10^{-15}$ cm² s⁻¹, with (c) no reaction, and (d) $k_c = 5 \times 10^{-4}$ s⁻¹. In case (c), $\tau_{da} = 2533$ s (i.e., 42 min) and ~ 160 min is required for the solute to attain a uniform steady state profile. In case (d), τ_{da} and τ_c are comparable, and as a result the solute not only reaches the steady state sooner (in about 60 min) than in the no-reaction case, but also the steady state concentration profile is visibly non-uniform.

This is a result of the fact that there is sufficient time for appreciable amounts of the solute to be consumed by the reaction as it diffuses towards the center of the particle.

Figure 3 illustrates the steady state concentration profiles for a range of k_c values (from 10^{-5} to 0.1 s^{-1}) in a particle of diameter $D_p = 0.1 \mu\text{m}$ with four different D_b values: (a) $10^{-6} \text{ cm}^2 \text{ s}^{-1}$, (b) $10^{-12} \text{ cm}^2 \text{ s}^{-1}$, (c) $10^{-13} \text{ cm}^2 \text{ s}^{-1}$, and (d) $10^{-15} \text{ cm}^2 \text{ s}^{-1}$. All together, these cases represent twenty different combinations of τ_{da} and τ_c . In case (a), $\tau_{da} \ll \tau_c$ for all the k_c values considered here, and as a result the steady state concentration profiles are essentially uniform across the entire particle, with the consumption of the solute by chemical reaction occurring uniformly across the entire volume of the particle. In case (b), even though the particle is considered to be a semi-solid with $D_b = 10^{-12} \text{ cm}^2 \text{ s}^{-1}$, τ_{da} and τ_c become comparable only when $k_c = 0.1 \text{ s}^{-1}$ (and higher). However, slower reactions produce non-uniform steady state concentration profiles in cases (c) and (d) for D_b values $10^{-13} \text{ cm}^2 \text{ s}^{-1}$ and lower. In these cases, most of the solute is consumed near the surface of the particle, with a concentration that becomes progressively depleted towards the center of the particle as k_c increases. Thus, solute uptake is “volume-controlled” when the concentration profile is uniform and tends to be “surface-area-controlled” at the other extreme.

Since the timescale for diffusion varies as R_p^2 , the diffusion limitation to reaction also depends strongly on particle size. As shown in Fig. 4, the relative effects of particle size, bulk diffusivity, and reaction rate on the shape of the steady state concentration profiles are concisely captured in terms of the dimensionless parameter q , which is a function of R_p , k_c , and D_b (Eq. 4). At low values of q (< 0.5), the steady state concentration profile is nearly uniform, but becomes increasingly non-uniform for q values of the order of unity and greater.

While the temporal evolution of the radial concentration profile is highly informative, the timescale to reach steady state, as well as the shape of the steady state profile, can be conveniently quantified in terms of the average particle-phase concentration $\bar{A}(t)$. We integrate the concentration profile given by Eq. (3) over the volume of the particle

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to obtain:

$$\frac{\bar{A}_i(t)}{A_i^s} = \frac{\int_0^{R_p} 4\pi r^2 \frac{A_i(r,t)}{A_i^s} dr}{\frac{4}{3}\pi R_p^3} = Q_i - U_i(t), \quad (7)$$

where

$$Q_i = 3 \left(\frac{q_i \coth q_i - 1}{q_i^2} \right), \quad (8)$$

$$U_i(t) = \frac{6}{\pi^2} \sum_{n=1}^{\infty} \frac{\exp \left\{ - \left(k_{c,i} + \frac{n^2 \pi^2 D_{b,i}}{R_p^2} \right) t \right\}}{(q_i/\pi)^2 + n^2}. \quad (9)$$

Here, Q_i is the ratio of the average particle-phase concentration to the surface concentration at steady-state, while $U_i(t)$ is the transient term, the value of which is always equal to Q_i at $t = 0$ and decreases exponentially to zero as $t \rightarrow \infty$. As noted earlier, the surface concentration A_i^s is assumed to be constant in the analytical solution of Eq. (1). However, since A_i^s can gradually change over time due to changes in the gas-phase concentration and particle composition, it is more appropriate to refer to the steady state as quasi-steady state. The timescale to reach a quasi-steady state (τ_{QSS}) within the particle can then be defined as the e-folding time for the exponential decay of the unsteady state term U_i relative to the quasi-steady state term Q_i . Thus, setting $U_i(\tau_{QSS}) = Q_i/e$, we get:

$$\sum_{n=1}^{\infty} \frac{\exp \left\{ - \left(k_{c,i} + \frac{n^2 \pi^2 D_{b,i}}{R_p^2} \right) \tau_{QSS} \right\}}{(q_i/\pi)^2 + n^2} = \frac{1}{e} \times \frac{\pi^2}{2} \left(\frac{q_i \coth q_i - 1}{q_i^2} \right). \quad (10)$$

For a given set of values for D_p , D_b , and k_c , Eq. (10) can be numerically solved for τ_{QSS} with the bisection method.

We first examine the dependence of τ_{QSS} and Q on D_b and k_c for a particle of $D_p = 0.1 \mu\text{m}$ (Fig. 5). The values of D_b are varied over fourteen orders of magnitude from $10^{-19} \text{cm}^2 \text{s}^{-1}$ (almost solid) to $10^{-5} \text{cm}^2 \text{s}^{-1}$ (liquid water) to cover the full range of semi-solid and liquid organic particles, and k_c values are varied over six orders of magnitude from 10^{-6}s^{-1} (very slow reaction) to 1s^{-1} (practically instantaneous reaction). As seen in Fig. 5a, the contours of τ_{QSS} range from 1 μs for liquid particles to 1 day for highly viscous semi-solid particles. For the semi-solid particles, there are two regions in the semi-solid zone as depicted by the gray dotted line. In the region above the dotted line, τ_{QSS} is sensitive only to the value of k_c and decreases rapidly with increase in k_c . For instance, at $D_b = 10^{-19} \text{cm}^2 \text{s}^{-1}$, $\tau_{SS} \sim 1$ day for $k_c = 5 \times 10^{-6} \text{s}^{-1}$ but decreases to < 1 min for $k_c = 10^{-2} \text{s}^{-1}$. In the region below the dotted line, τ_{QSS} is sensitive only to the value of D_b for both semi-solid and liquid particles. For example, at $D_b \sim 10^{-14} \text{cm}^2 \text{s}^{-1}$, τ_{QSS} remains constant at ~ 1 min for k_c values from 10^{-6} up to about 10^{-2}s^{-1} (i.e., up to the dotted line) and only then becomes sensitive to reaction at higher values of k_c . τ_{QSS} is sensitive to both k_c and D_b only in the relatively narrow envelope along the dotted line itself. As seen in Fig. 5b, the values of Q are < 0.001 for highly viscous semi-solid particles and high k_c values, while they approach unity as D_b increases and k_c decreases. Note that the dotted line in Fig. 5a roughly corresponds to the contour for $Q = 0.6$ in Fig. 5b.

Next, we examine the dependence of τ_{QSS} and Q on particle size. Figure 6 shows τ_{QSS} vs. D_p for D_b values ranging from 10^{-18} to $10^{-10} \text{cm}^2 \text{s}^{-1}$ for (a) $k_c = 0 \text{s}^{-1}$, (b) $k_c = 10^{-3} \text{s}^{-1}$, (c) $k_c = 0.01 \text{s}^{-1}$, and (d) $k_c = 0.1 \text{s}^{-1}$. As seen in Fig. 6a, for any given D_b , τ_{QSS} increases by five orders of magnitude as D_p increases from 0.003 to $1 \mu\text{m}$. At the upper end, particles with $D_b < 10^{-18} \text{cm}^2 \text{s}^{-1}$ have τ_{QSS} about 10 min at $D_p = 0.003 \mu\text{m}$ and increase to more than 10^4 min at $D_p = 0.1 \mu\text{m}$. In contrast, particles

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with $D_b > 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ have τ_{QSS} below 1 min (indicated by the dotted gray line) for sizes up to $0.7 \mu\text{m}$. From a practical standpoint, since most ambient SOA particles are smaller than $\sim 0.7 \mu\text{m}$, concentration profiles of non-reacting solutes inside particles with $D_b > 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ may be assumed to be at steady-state. However, significant diffusion limitation can exist for non-reacting solutes in particles with $D_b < 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ depending on their size. In stark contrast, for reacting solutes, τ_{QSS} asymptotically approaches a common maximum value for all values of D_b as the particle size increases (Fig. 6b–d). This maximum value of τ_{QSS} is about 7, 0.7, and 0.07 min for $k_c = 10^{-3}$, 10^{-2} , and 0.1 s^{-1} , respectively. Thus, from a practical standpoint, concentration profiles of solutes reacting with $k_c > 10^{-2} \text{ s}^{-1}$ may be assumed to be at quasi-steady state in particles with any D_b and of any size.

Figure 7 illustrates variation of Q with D_p for the four cases shown in Fig. 6. At quasi-steady state, the particle-phase concentration profile for non-reacting solutes is always uniform (i.e., $Q = 1$) even though τ_{QSS} can differ significantly depending on the particle size and D_b value (Fig. 7a). For reacting solutes with k_c up to 0.1 s^{-1} , Q remains nearly equal to unity in particles with $D_b > 10^{-10} \text{ cm}^2 \text{ s}^{-1}$ and D_p up to $1 \mu\text{m}$. For $D_b < 10^{-10} \text{ cm}^2 \text{ s}^{-1}$, Q decreases as D_p increases for a given D_b , while it increases as D_b increases for a given D_p .

In general, the above analysis indicates that: (a) for a given D_p , a more reactive solute will reach quasi-steady state sooner and exhibit a more non-uniform concentration profile than a less reactive one, especially in particles with lower D_b than higher, and (b) for a given set of values for k_c and D_b , a solute in smaller particles will reach quasi-steady state sooner and exhibit a more uniform quasi-steady state concentration profile than in larger particles.

3 Kinetic gas-particle partitioning model

We shall now describe the development of a new formulation for modeling kinetic partitioning of SOA based on the insights gained from timescale analysis of the diffusion-reaction process within the particle phase. The formulation takes into account solute volatility, gas-phase diffusion, interfacial mass accommodation, particle-phase diffusion, and particle-phase reaction. However, instead of numerically resolving the concentration gradient inside the particle (Shiraiwa et al., 2012a), which is computationally expensive and therefore impractical for inclusion in 3-D Eulerian models, we use the analytical expressions of the quasi-steady state and transient behavior of the solute diffusing and reacting within the particle.

3.1 Model formulation

3.1.1 Single particle equations

We begin by relating the average particle-phase concentration of the solute \bar{A}_i (mol cm⁻³ (particle)) to its average bulk gas-phase concentration $\bar{C}_{g,i}$ (mol cm⁻³(air)) over a single particle. Similar to the timescale for diffusion in the particle-phase (Eq. 5), the timescale for the gas-phase concentration gradient outside the particle to reach a quasi-steady state (τ_{dg}) is given by (Seinfeld and Pandis, 2006):

$$\tau_{dg,i} = \frac{R_p^2}{\pi^2 D_{g,i}}, \quad (11)$$

where $D_{g,i}$ (cm² s⁻¹) is the gas-phase diffusivity. For a typical $D_{g,i}$ of 0.05 cm² s⁻¹, the value of τ_{dg} is of the order 10⁻⁸ s or less for submicron size aerosols, which is much smaller than the typical timescale for changes in the bulk gas-phase concentration in the ambient atmosphere. We can therefore safely assume that the gas-phase concentration profile of the solute around the particle is at quasi-steady state at any instant.

An ordinary differential equation describing the rate of change of \bar{A}_i due to mass transfer between gas and a single particle with particle-phase reaction can then be written as:

$$\frac{d\bar{A}_i}{dt} = \frac{3}{R_p} k_{g,i} (\bar{C}_{g,i} - C_{g,i}^s) - k_{c,i} \bar{A}_i, \quad (12)$$

5 where $C_{g,i}^s$ (mol cm⁻³(air)) is the gas-phase concentration of the solute just outside the surface of the particle, and $k_{g,i}$ (cm s⁻¹) is the gas-side mass transfer coefficient given as:

$$k_{g,i} = \frac{D_{g,i}}{R_p} f(Kn_i, \alpha_i). \quad (13)$$

10 Here $f(Kn_i, \alpha_i)$ is the transition regime correction factor (Fuchs and Sutugin, 1971) to the Maxwellian flux as a function of the Knudsen number $Kn_i = \lambda_i/R_p$ (where λ_i is the mean free path) and mass accommodation coefficient, α_i , that accounts for interfacial mass transport limitation:

$$f(Kn_i, \alpha_i) = \frac{0.75\alpha_i(1 + Kn_i)}{Kn_i(1 + Kn_i) + 0.283\alpha_i Kn_i + 0.75\alpha_i}. \quad (14)$$

15 Assuming that $C_{g,i}^s$ and the particle-phase concentration of i just inside the surface, A_i^s (mol cm⁻³ (particle)), are related by Raoult's law equilibrium, we can write:

$$C_{g,i}^s = \frac{A_i^s}{\sum_j A_j^s} C_{g,i}^* \quad (15)$$

where $C_{g,i}^*$ is the effective saturation vapor concentration (mol cm⁻³ (air)), and $\sum_j A_j^s$ is the total particle-phase concentration of all the organic species at the surface. However,

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since the surface concentrations of all the species are not always known, we use the total average particle-phase concentration $\sum_j \bar{A}_j$ as an approximation for $\sum_j A_j^s$. Thus Eq. (12) is rewritten in terms of A_i^s as:

$$\frac{d\bar{A}_i}{dt} = \frac{3}{R_p} k_{g,i} \left(\bar{C}_{g,i} - \frac{A_i^s}{\sum_j \bar{A}_j} C_{g,i}^* \right) - k_{c,i} \bar{A}_i. \quad (16)$$

A_i^s can be assumed to be equal to \bar{A}_i in liquid particles for a non-reactive or slowly reacting solute that quickly attains a uniform concentration profile (as was previously shown in Fig. 2a and b). But, as discussed in the previous section, this equality may not hold for reactive and non-reactive solutes in semi-solid particles. In such cases, Eq. (7) can be used to express A_i^s in terms of \bar{A}_i as long as A_i^s does not change with time, because the analytical solution to Eq. (1) assumes a constant A_i^s according to the boundary condition (Eq. 2b). In practice, however, Eq. (7) can be used if the timescale for changes in A_i^s is much greater than the timescale for the solute to relax to its quasi-steady state profile inside the particle. With this caveat, we get:

$$\frac{d\bar{A}_i}{dt} = \frac{3}{R_p} k_{g,i} \left\{ \bar{C}_{g,i} - \frac{\bar{A}_i}{\sum_j \bar{A}_j} \frac{C_{g,i}^*}{(Q_i - U_i(t))} \right\} - k_{c,i} \bar{A}_i. \quad (17)$$

Note that Eq. (17) describes kinetic mass transfer of species i between bulk gas-phase and a single particle, with chemical reaction within the particle phase, and includes mass-transfer limitations due to gas-phase diffusion, interfacial accommodation, and particle-phase diffusion. Here, the term $U_i(t)$ is to be evaluated at the “time since start.” Equation (17) can therefore only be used in a Lagrangian box model framework for a “closed system” where we can specify an initial concentration of the solute vapor (at time $t = 0$), which then partitions to the particle-phase as a function of time. The solute vapor in the closed system is not subjected to emissions, dilution, and loss due

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to gas-phase oxidation. In the case of no particle-phase reaction, the solute vapor will eventually reach equilibrium with the particles. In the presence of particle-phase reaction, the solute vapor concentration will eventually decay to zero. This is in stark contrast with the “general system” such as the ambient atmosphere and 3-D atmospheric chemistry transport models where the solute vapor at a given location may continuously change due to emissions, dilution, and gas-phase chemistry in addition to gas-particle partitioning. As a result, it is not possible to evaluate $U_i(t)$ in the general system, because we cannot keep track of the “time since start” in the same sense as used in the transient analytical solution to Eq. (1). Therefore, based on the value of k_c and the associated timescale for the particle-phase concentration profile to reach quasi steady state (τ_{QSS}), the following two approximations to Eq. (17) are made for it to be applicable to the general system.

Approximation 1: for fast reactions ($k_{c,i} \geq 0.01 \text{ s}^{-1}$)

As discussed in the previous section (Fig. 6c), τ_{QSS} for a solute reacting with $k_{c,i} \geq 0.01 \text{ s}^{-1}$ is less than 1 min in particles with any D_b and of any size. Compared to the typical time step values of 5 min or greater in 3-D Eulerian models, the particle-phase concentration profile for solutes with $\tau_{QSS} \leq 1 \text{ min}$ may be assumed to be at quasi-steady-state, and the term $U_i(t)$ can be safely neglected in Eq. (17) to yield:

$$\frac{d\bar{A}_i}{dt} = \frac{3}{R_p} k_{g,i} \left\{ \bar{C}_{g,i} - \frac{\bar{A}_i}{\sum_j \bar{A}_j} \frac{C_{g,i}^*}{Q_i} \right\} - k_{c,i} \bar{A}_i \text{ for } k_{c,i} \geq 0.01 \text{ s}^{-1}. \quad (18)$$

A similar equation was derived by Shi and Seinfeld (1991) for reactive mass transport of SO_2 (with Henry’s law for absorption) in cloud droplets assuming quasi-steady state within the droplet phase. Now, as $k_c \rightarrow 0$, $Q \rightarrow 1$, and mass transfer is governed entirely by gas-phase diffusion in Eq. (18). As a result, Eq. (18) tends to lose its ability to capture the resistance to mass transfer due to slow diffusion in the particle phase as $k_c \rightarrow 0$. Therefore, an alternate treatment for mass transfer is needed for slow reactions.

Approximation 2: for slow reactions ($k_{c,i} < 0.01 \text{ s}^{-1}$)

For $k_{c,i} < 0.01 \text{ s}^{-1}$ (or $\tau_{\text{QSS}} > 1 \text{ min}$), we use the classical two-film theory of mass transfer between the gas and particle phases. The two-film theory was originally introduced by Lewis and Whitman (1924) and has been widely used to model mass transfer in two phase systems, with and without chemical reactions (Astarita, 1967; Doraiswamy and Sharma, 1984; Bird et al., 2007). Figure 8 shows the schematic of the two-film model which assumes that the concentration gradients in the gas and particle phases are confined in the respective hypothetical “films” adjacent to the interface. The gas- and particle-side film thicknesses are denoted by δ_g and δ_b (cm), respectively, and the respective mass transfer coefficients (cm s^{-1}) are defined as $k_g = D_g/\delta_g$ and $k_b = D_b/\delta_b$. The overall gas-side mass transfer coefficient K_g (cm s^{-1}) is then given by (see Appendix A for the derivation):

$$\frac{1}{K_{g,i}} = \frac{1}{k_{g,i}} + \frac{1}{k_{b,i}} \left(\frac{C_{g,i}^*}{\sum_j \bar{A}_j} \right). \quad (19)$$

The ordinary differential equation describing the rate of change of \bar{A}_i due to gas-particle mass transfer and particle-phase reaction can then be written in terms of the overall driving force as:

$$\frac{d\bar{A}_i}{dt} = \frac{3}{R_p} K_{g,i} \left\{ \bar{C}_{g,i} - \frac{\bar{A}_i}{\sum_j \bar{A}_j} C_{g,i}^* \right\} - k_{c,i} \bar{A}_i \text{ for } k_{c,i} < 0.01 \text{ s}^{-1}. \quad (20)$$

A similar equation was derived by Zaveri (1997) for reactive mass transport of SO_2 (with Henry's law for absorption) in cloud droplets assuming quasi-steady state within the droplet phase. The advantage of the two-film model formulation is that the diffusion limitations from both the gas and particle sides are represented in the overall mass transfer coefficient, and can therefore be used to model mass transfer of slow-reacting

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solutes. The gas-side mass transfer coefficient (k_g) is already known from Eq. (13) where $\delta_g = R_p$. However, the particle-side film thickness, δ_b , and therefore k_b , are not readily known. In a general system, the bulk gas- and particle-phase concentrations of a reactive semi-volatile solute tend to reach a quasi-steady state when the net source rate of the solute in the gas phase is relatively steady. Since both Eqs. (18) and (20) describe the same process, they should predict identical gas- and particle-phase concentrations at quasi-steady-state. Thus, setting $d\bar{A}_i/dt = 0$ in both Eqs. (18) and (20) and equating the expressions for $(\bar{A}_i/\bar{C}_{g,i})$ resulting from each of them yields the general expressions for δ_b and k_b in terms of D_b , k_c , and R_p (see Appendix B for the derivation):

$$\delta_{b,i} = R_p \left(\frac{1 - Q_i}{q_i \coth q_i - 1} \right), \quad (21)$$

$$k_{b,i} = \frac{D_{b,i}}{R_p} \left(\frac{q_i \coth q_i - 1}{1 - Q_i} \right). \quad (22)$$

For the limiting case of a non-reactive solute, $k_c \rightarrow 0$, $q \rightarrow 0$, $Q \rightarrow 1$ and Eq. (22) reduces to:

$$k_{b,i} = 5 \frac{D_{b,i}}{R_p}. \quad (23)$$

3.1.2 Polydisperse aerosol equations

We now extend the *closed system* box model Eq. (17) for a single particle to a polydisperse aerosol in a sectional framework. For a given size section m , with number concentration N_m ($\text{cm}^{-3}(\text{air})$) and particle radius $R_{p,m}$ (cm), we define $\bar{C}_{a,i,m}$ ($\text{mol cm}^{-3}(\text{air})$) as the total average concentration of solute i in size section m :

$$\bar{C}_{a,i,m} = \frac{4}{3} \pi R_{p,m}^3 N_m \bar{A}_{i,m}. \quad (24)$$

Multiplying Eq. (17) by $(4\pi R_{p,m}^3 N_m / 3)$ gives:

$$\frac{d\bar{C}_{a,i,m}}{dt} = 4\pi R_{p,m}^2 N_m k_{g,i,m} \left\{ \bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{(Q_i - U_i(t))} \right\} - k_{c,i} \bar{C}_{a,i,m}, \quad (25)$$

where $S_{i,m}$ is the saturation ratio:

$$S_{i,m} = \frac{C_{g,i}^*}{\sum_j \bar{C}_{a,j,m}}. \quad (26)$$

5 The corresponding equation governing the gas-phase concentration of solute i is:

$$\frac{d\bar{C}_{g,i}}{dt} = - \sum_m \left[4\pi R_{p,m}^2 N_m k_{g,i,m} \left\{ \bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{(Q_i - U_i(t))} \right\} \right]. \quad (27)$$

Similarly, the particle-phase and gas-phase equations for polydisperse aerosols in the general system are as follows.

Approximation 1: for $k_{c,i} \geq 0.01 \text{ s}^{-1}$

$$10 \quad \frac{d\bar{C}_{a,i,m}}{dt} = 4\pi R_{p,m}^2 N_m k_{g,i,m} \left(\bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{Q_i} \right) - k_{c,i} \bar{C}_{a,i,m}, \quad (28)$$

$$\frac{d\bar{C}_{g,i}}{dt} = - \sum_m \left[4\pi R_{p,m}^2 N_m k_{g,i,m} \left(\bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{Q_i} \right) \right]. \quad (29)$$

Approximation 2: for $k_{c,i} < 0.01 \text{ s}^{-1}$

$$\frac{d\bar{C}_{a,i,m}}{dt} = 4\pi R_{p,m}^2 N_m K_{g,i,m} (\bar{C}_{g,i} - \bar{C}_{a,i,m} S_{i,m}) - k_{c,i} \bar{C}_{a,i,m}, \quad (30)$$

$$\frac{d\bar{C}_{g,i}}{dt} = - \sum_m \left[4\pi R_{p,m}^2 N_m K_{g,i,m} (\bar{C}_{g,i} - \bar{C}_{a,i,m} S_{i,m}) \right]. \quad (31)$$

5 The proposed formulation, described by Eq. (28) through Eq. (31), is relatively simple and amenable for use in regional and global aerosol models, although it presently awaits specification of the actual particle-phase chemical reactions that are important for SOA formation.

10 We have implemented both the closed system and general system formulations in the computationally efficient, multicomponent aerosol box-model MOSAIC and adapted the existing semi-implicit Euler method solver to numerically integrate the set of coupled ordinary differential equations for any number of solutes i over any number of size bins m (Zaveri et al., 2008). Sectional growth in MOSAIC is calculated using the two-moment approach of Simmel and Wurzler (2006). The closed system formulation is to
15 be used in the box-model version only while the general system formulation can be used in both box- and 3-D Eulerian models. The complete solution to these equations may be labeled as “semi-numerical,” because the particle-phase diffusion-reaction process is represented analytically while the set of ordinary differential equations themselves are integrated numerically.

20 3.2 Model validation

We shall now validate the new formulation in MOSAIC against a “fully numerical” finite-difference solution to Eq. (1) with a flux type boundary condition that includes mass transfer of the solute between the gas phase and the particle surface. The volume of the spherical particle is resolved with multiple layers, and diffusion and reaction of the

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solute species through these layers are integrated numerically. We used 300 uniformly-spaced layers in the present exercise. The finite difference model is conceptually similar to the KM-GAP model (Shiraiwa et al., 2012a), but does not include reversible adsorption at the surface and heat transfer processes. The finite-difference solution is used as a benchmark here because it rigorously solves Eq. (1) and does not assume the surface concentration to remain constant with time.

For validation purposes, we consider a monodisperse semi-solid aerosol composed of non-volatile organic species P_3 (molecular weight 100 g mol^{-1} and density 1 g cm^{-3}), with initial particle diameter $D_p = 0.2 \mu\text{m}$, particle number concentration $N = 5000 \text{ cm}^{-3}$, and bulk diffusivity $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$. For simplicity, the molecular weight and density of the condensing solute (P_1) and its reaction product species (P_2) are also assumed to be 100 g mol^{-1} and 1 g cm^{-3} , respectively. The three species (P_1 , P_2 , and P_3) are assumed to form an ideal solution that participates in the absorption of P_1 according to Raoult's law. Model validation is demonstrated below for both closed and general systems.

3.2.1 Closed system

In three separate closed system cases, the initial monodisperse aerosol was exposed to the solute (P_1) gas concentration of $2 \mu\text{g m}^{-3}$ with volatility $C_g^* = 10, 100$, and $1000 \mu\text{g m}^{-3}$. Figure 9 compares the solution given by MOSAIC (Eqs. 24 and 25) with the finite-difference model solution for gas-phase concentration decay due to kinetic gas-particle partitioning for particle-phase reaction rate constants k_c ranging from 0 to 0.1 s^{-1} . When $k_c = 0$, the gas-phase concentration reaches an equilibrium value that depends on the solute volatility, while in other cases it decays to zero at different rates as governed by the particle-phase reaction rate constant and diffusion limitation. MOSAIC is able to reproduce the finite difference results quite well, although small deviations can be seen during the initial portions of the gas decay for $k_c \leq 10^{-4} \text{ s}^{-1}$ and $C_g^* = 10$ and $100 \mu\text{g m}^{-3}$. The following metrics were used to quantify the accuracy of

MOSAIC relative to the finite difference (FD) model:

$$\text{Mean Normalized Bias, MNB} = \left(C_{g,1}^{\text{MOSAIC}} - C_{g,1}^{\text{FD}} \right) / C_{g,1}^{\text{FD}}, \quad (32)$$

$$\text{Mean Normalized Gross Error, MNGE} = \left| C_{g,1}^{\text{MOSAIC}} - C_{g,1}^{\text{FD}} \right| / C_{g,1}^{\text{FD}}, \quad (33)$$

$$\text{Maximum Normalized Gross Error, maxNGE} = \max \left(\left| C_{g,1}^{\text{MOSAIC}} - C_{g,1}^{\text{FD}} \right| / C_{g,1}^{\text{FD}} \right). \quad (34)$$

These metrics were calculated using the model outputs at 5 min intervals for the 10 h long simulations. However, negligibly small gas-phase concentrations ($< 0.05 \mu\text{g m}^{-3}$) towards the latter part of the simulations (where applicable) were excluded in the calculations of the metrics. The results are displayed in Table 1. The MNB and MNGE are comparable in magnitude and range from $\sim 0.1\%$ to $\sim 10\%$, with values greater than $\sim 5\%$ seen only for $C_g^* = 10 \mu\text{g m}^{-3}$. The large maxNGE values ($> 20\%$) seen for $C_g^* = 10 \mu\text{g m}^{-3}$ occur as the gas-phase concentrations approach zero. Overall, the agreement between the two models is quite good for the closed system.

3.2.2 General system

In three separate general system cases, the initial monodisperse aerosol was exposed to solute P_1 with $C_g^* = 10, 100, \text{ and } 1000 \mu\text{g m}^{-3}$ at a constant gas-phase source rate of $\gamma = 0.1 \mu\text{g m}^{-3} \text{ h}^{-1}$ in each case. The initial gas-phase concentration of P_1 was zero in each case. Figure 10 compares the evolution of the gas-phase concentration of P_1 predicted by MOSAIC (Eqs. 28–31) and the finite-difference model. The particle-phase reaction rate constant k_c ranged from 0 to 0.1 s^{-1} . When $k_c = 0$, the gas-phase concentration of P_1 increases almost linearly with time upon reaching quasi-equilibrium with the particle phase. For $k_c > 0$, the gas-phase concentration of P_1 remains constant after the initial build up as the source rate is balanced by the loss rate due to particle-phase diffusion and reaction. This quasi-steady state gas-phase concentration

level depends on the combination of C_g^* , D_b , and k_c . For $C_g^* = 10 \mu\text{g m}^{-3}$, the time required to establish quasi-steady state between gas and particle phases ranges from less than 1 h at $k_c = 0.1 \text{ s}^{-1}$ to more than 20 h at $k_c = 10^{-4} \text{ s}^{-1}$. The time to reach quasi-equilibrium (for non-reactive solutes) and quasi-steady state (for reactive solutes) increases as the value of C_g^* increases. Approximations 1 and 2 in MOSAIC are able to capture both the initial “spin-up” phase, when the gas-phase concentration builds up, as well as the later phase where the concentration remains in quasi-equilibrium or quasi-steady state. Approximation 1 predicts faster gas uptake than the finite difference model for slow reactions while Approximation 2 predicts slower gas uptake than the finite difference model for fast reactions (not shown), especially for low volatility solutes ($C_g^* \approx 10 \mu\text{g m}^{-3}$). A combination of Approximations 1 and 2 is thus needed to cover the full range of possible k_c values.

The normalized gross errors in MOSAIC are relatively large during the spin-up phase where the gas-phase concentrations are very small. In a 3-D Eulerian model application, the spin-up phase occurs at the beginning of the simulation and is usually discarded. Here, we discard the first two hours of spin-up in each simulation to avoid small gas-phase concentrations when calculating the bias and error metrics, shown in Table 2. Both MNB and MNGE are generally less than $\sim 3\%$. The maxNGE values ranged between 0.3 and 8.5 %. The overall performance of MOSAIC for the general system is excellent.

The general system framework in MOSAIC is thus amenable for eventual use in regional and global climate models, although it currently awaits specification of the actual particle-phase reactions that are important for SOA formation. There are, however, some limitations of this framework that must be taken into consideration when specifying the actual physical and chemical details and evaluating them using laboratory and field observations. While the framework allows particles of different sizes and composition to have different bulk diffusivities, it cannot explicitly treat the potential variation of diffusivity within a given particle of complex morphology. Examples include black carbon or solid ammonium sulfate particles coated with organics as well as particles

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with non-ideal internal mixtures of hydrophobic and hydrophilic organics. The diffusion-reaction process inside such complex and potentially non-spherical particles will have to be parameterized based on the average bulk properties, with possible guidance from more detailed finite difference models where applicable.

4 Results and discussion

We now apply the updated MOSAIC model to a series of polydisperse aerosol scenarios to investigate the influence of particle-phase reactions, phase state, and solute volatility on SOA partitioning timescale and the evolution of aerosol size distribution. While the exact mechanism(s) responsible for the growth of newly formed particles (1 to 10 nm range) is still unknown, it is suspected to occur via effectively irreversible condensation of very low volatility organic species which can overcome the strong Kelvin effect (Pierce et al., 2011). In the present study, we focus on the competitive growth dynamics of the Aitken and accumulation mode particles, as might result after the newly formed particles have grown up to Aitken mode sizes. The Kelvin effect and coagulation are neglected for simplicity. Figure 11 shows the initial aerosol number and volume size distributions used for this exercise (note that in this and subsequent figures, $\text{dlog } D_p$ means $\text{dlog}_{10} D_p$). Again, this pre-existing aerosol is assumed to be composed of non-volatile organic species (P_3) of molecular weight 100 g mol^{-1} and density 1 g cm^{-3} . The entire size distribution, consisting of an Aitken mode and an accumulation mode, is discretized over 1000 logarithmically-spaced size bins (lower boundary of the smallest bin = $0.008 \mu\text{m}$ and the upper boundary of the largest bin = $1 \mu\text{m}$). The total number concentration of particles in the Aitken mode is 6223 cm^{-3} while that in the accumulation mode is 1139 cm^{-3} ; the total aerosol mass concentration is $2 \mu\text{g m}^{-3}$. Figure 11 also shows the condensational sink $k_{\text{CS},i,m} = 4\pi R_{p,m}^2 N_m k_{g,i,m}$ for each size bin m as a function of D_p . For this particular size distribution, the sum of k_{CS} over all the size bins in the Aitken mode is equal to that in the accumulation mode, so that there is no initial bias in the condensation rate of the solute species towards either mode merely due to

differences in the initial condensational sink rates for the two modes. Both closed and general systems scenarios are examined.

4.1 Closed system

A set of closed system simulations was performed in which the initial organic aerosol was separately exposed to the solute gas (P_1) with three different C_g^* values: 10, 100, and $1000 \mu\text{g m}^{-3}$ (molecular weight = 100 g mol^{-1}), with an initial gas-phase concentration of $6 \mu\text{g m}^{-3}$ in each case. For each solute volatility case, the effect of aerosol phase state was examined using four different D_b values: 10^{-6} , 10^{-12} , 10^{-13} , and $10^{-15} \text{ cm}^2 \text{ s}^{-1}$. In all cases, k_c was set at 0.01 s^{-1} so that τ_{SS} was always less than ~ 0.7 min across the entire size distribution. In each case, the simulation was run until the gas-phase solute was completely absorbed and reacted to form a non-volatile product in the particle phase. Again, the molecular weight and density of the product species (P_2) were assumed to be 100 g mol^{-1} and 1 g cm^{-3} , respectively, and all three species (P_1 , P_2 , and P_3) were assumed to form an ideal solution that participated in the absorption of P_1 according to Raoult's law. An additional set of reference simulations were performed for two extreme scenarios: (1) instantaneous particle-phase reaction (i.e., $k_c \rightarrow \infty$), which is equivalent to solving the non-volatile solute condensation case (i.e., mechanism #1), and (2) no particle-phase reaction ($k_c = 0$), which is referred to as Raoult's law partitioning (i.e., mechanism #2). In the latter case, the initial gas-phase concentrations for the different C_g^* subcases were increased such that $6 \mu\text{g m}^{-3}$ of solute was partitioned into the particle phase at steady state (i.e., at equilibrium) in each case.

4.1.1 Reference cases

We shall first discuss the results of the closed system reference cases. Figure 12 shows the gas-phase decay and the corresponding temporal evolution of aerosol size distribution and mass fraction of newly formed SOA for the instantaneous particle-phase

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reaction case. Here, gas-particle partitioning is independent of the particle phase state and is governed entirely by gas-phase diffusion limitation. Vapor concentration is completely depleted in about 1 h, and aerosol size distribution evolution displays the well-known narrowing characteristics as the small particles grow faster (more precisely, have greater $d\ln D_p/dt$) than the large ones (Zhang et al., 2012). Consequently, the mass fraction of the newly formed SOA in smaller particles is much higher than in the larger ones. Note that in the SOA mass fraction panel, the left-most point on each line with mass fraction ≈ 1 corresponds to the smallest initial particles ($D_p = 0.008 \mu\text{m}$ at $t = 0$).

In contrast, aerosol evolution due to Raoult's law partitioning depends on both solute volatility and particle phase state. Figure 13 shows the gas-phase concentration decay and the corresponding aerosol size distribution and SOA mass fraction evolution for the less volatile solute with $C_g^* = 10 \mu\text{g m}^{-3}$. The effect of phase state is illustrated with two bulk diffusivities: $D_b = 10^{-6}$ and $10^{-15} \text{cm}^2 \text{s}^{-1}$. In the case with liquid particles ($D_b = 10^{-6} \text{cm}^2 \text{s}^{-1}$), there is negligible resistance to mass transfer within the particle (refer to Fig. 6a), and as a result the vapor concentration rapidly decreases during the first 1 h and reaches a steady state in about 7.5 h. In the first ~ 20 min, the size distribution exhibits the narrowing of the Aitken mode similar to that seen in gas-phase diffusion-limited growth, although not as intense. The SOA mass fraction reaches up to 0.97 in small particles while it is only about 0.25 in the large particles. However, as the vapor concentration decreases further, the peak of the size distribution begins to decrease and the width broadens due to evaporation from small particles while the large particles continue to grow (Zhang et al., 2012). The SOA mass fraction in small particles decreases to 0.75, while it gradually increases to 0.75 in the large particles. The vapor concentration remains steady while this inter-particle mass transfer (via the gas phase) occurs over a relatively longer period (~ 480 h) until the entire aerosol size distribution reaches equilibrium.

Similar behavior is seen in the case with semi-solid particles ($D_b = 10^{-15} \text{cm}^2 \text{s}^{-1}$), although the timescale over which it occurs is relatively longer due to much higher

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particle-phase diffusion limitation. While the vapor concentration declines rapidly in the beginning (e-folding timescale of 16.5 h), it takes about 175 h to reach the steady state and more than 400 h for the aerosol size distribution to reach equilibrium. Also, because the particle-phase diffusion limitation is much less in small particles than the large ones (refer to Fig. 6a), the Aitken mode exhibits more intense narrowing and a higher peak (at about 1 h) than seen in liquid particles. Then, again, as the vapor concentration decreases further, the width broadens and the peak decreases due to evaporation of small particles while the large ones continue to grow more slowly. The final aerosol size distribution and SOA mass fraction across the size spectrum are identical (within numerical errors) to those obtained in the liquid particle case.

Figure 14 shows the results for the more volatile solute with $C_g^* = 1000 \mu\text{g m}^{-3}$. In the case with liquid particles ($D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$), the vapor concentration reaches the steady state in just 20 min (vs. 7.5 h for $C_g^* = 10 \mu\text{g m}^{-3}$) while it takes nearly 400 h (vs. 175 h for $C_g^* = 10 \mu\text{g m}^{-3}$) in the case with semi-solid particles ($D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$). Again, the final aerosol size distribution and SOA mass fraction solutions at equilibrium are identical to those obtained for the $C_g^* = 10 \mu\text{g m}^{-3}$ cases, but their temporal evolutions are quite different. In the case with liquid particles, the width of the aerosol size distribution does not narrow and the peak height remains the same as the particles grow. This is because the small particles quickly attain a quasi-equilibrium state with the more volatile solute. Consequently, the SOA mass fraction in the small particles quickly reaches the equilibrium value of 0.75 (instead of overshooting as seen for $C_g^* = 10 \mu\text{g m}^{-3}$) while the larger particles catch up just a bit more slowly. The entire size distribution reaches equilibrium within 1 h.

In the case with semi-solid particles, the Aitken mode size distribution narrows (similar to that seen in Fig. 13a) in the first few minutes, but broadens back within 30 min. Again, the SOA mass fraction in small particles quickly reaches the equilibrium value of 0.75, while it still takes ~ 480 h for the large particles in the spectrum to reach equilibrium due to the significant diffusion limitation in the particle phase.

4.1.2 Reactive partitioning cases

We now present results for the closed-system reactive partitioning cases with $k_c = 0.01 \text{ s}^{-1}$. Figure 15 shows vapor concentration decay for each of the three solute volatility cases ($C_g^* = 10, 100, \text{ and } 1000 \mu\text{g m}^{-3}$) for D_b values ranging from 10^{-6} to $10^{-15} \text{ cm}^2 \text{ s}^{-1}$. It also shows a plot of the e-folding timescale (τ_g) for the decay as a function of D_b for the different volatilities. Each plot includes the reference case of instantaneous reaction for comparison. Unlike in Raoult's law partitioning, the vapor concentration always decays to zero in reactive partitioning and the decay rate slows down with increase in C_g^* . The vapor decay rate also slows down with decrease in D_b and it is especially sensitive to D_b in semi-solid particles.

Figure 16 illustrates the effects of the different C_g^* and D_b values on the final aerosol size distribution. The final results for the reference cases of instantaneous reaction and Raoult's law partitioning are also shown for easy comparison. In the case of $C_g^* = 10 \mu\text{g m}^{-3}$, the Aitken mode exhibits significant narrowing for all values of D_b . The narrowing becomes more pronounced for $D_b < 10^{-13} \text{ cm}^2 \text{ s}^{-1}$ with the shape of the entire size distribution for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$ being nearly identical to that for the instantaneous reaction reference case. Further decrease in D_b will produce even more narrowing. Since there is negligible particle-phase diffusion limitation for $D_b > 10^{-10} \text{ cm}^2 \text{ s}^{-1}$ ($Q \approx 1$; Fig. 7c), the size distribution of liquid aerosol narrows because its initial evolution (in the case of low volatility solutes) resembles that of gas-phase diffusion-limited growth, and the particle-phase reaction rate is fast enough to transform the absorbed solute to a non-volatile product before it can evaporate. For $D_b < 10^{-13} \text{ cm}^2 \text{ s}^{-1}$, the steep gradient in Q across the size distribution results in significantly lower surface concentrations over small semi-solid particles compared to the large ones. The small semi-solid particles therefore grow even faster than the large ones compared to the liquid aerosol case, producing relatively more intense narrowing of the size distribution.

As the solute C_g^* increases to 100 and $1000 \mu\text{g m}^{-3}$, liquid particles tend to attain quasi-equilibrium with the gas phase relatively faster than the solute reacts within the

particle. As a result, the final size distributions for $D_b \leq 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ progressively resemble that of the Raoult's law partitioning case. However, significant narrowing is still seen for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$ due to the steep gradient in Q across the size distribution, which causes the small semi-solid particles to grow much faster than the large ones.

In general, the final size distribution shape tends to be closer to that for instantaneous reaction case for lower C_g^* and D_b values and higher k_c values, while it tends to be closer to that for Raoult's law partitioning for higher C_g^* and D_b and lower k_c .

Figure 17 illustrates the influence of C_g^* and D_b values on the final SOA mass fraction size distribution. Curves for the two reference cases are also included for comparison.

In the case of $C_g^* = 10 \mu\text{g m}^{-3}$, the curves for all D_b values are similar to that of the instantaneous reference case due to appreciable narrowing of the size distribution. But as C_g^* increases, the SOA mass fraction curves progressively become more uniform for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ while they remain non-uniform for $D_b < 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ for particles with $D_p > 0.2 \mu\text{m}$. In all C_g^* cases, the SOA mass fraction curves for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$ closely resemble the instantaneous reaction case.

4.2 General system

A set of general system simulations was performed in which the initial organic aerosol was separately exposed to solutes with $C_g^* = 10, 100, \text{ and } 1000 \mu\text{g m}^{-3}$ at a moderate but constant gas-phase source rate of $\gamma = 0.6 \mu\text{g m}^{-3} \text{ h}^{-1}$ in each case. The effect of aerosol phase state was examined using two different D_b values: 10^{-6} and $10^{-15} \text{ cm}^2 \text{ s}^{-1}$. For each combination of C_g^* and D_b values, the effect of particle-phase reaction was examined for $k_c = 0.01, 0.1, 1, \text{ and } \infty \text{ s}^{-1}$. Each simulation was 12 h long.

Figure 18 shows the time evolutions of total SOA mass concentration for liquid particles ($D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$) with different solute C_g^* values and the corresponding final aerosol size distributions at $t = 12 \text{ h}$. In the case with $C_g^* = 10 \mu\text{g m}^{-3}$, the SOA formation rate is essentially the same for $k_c \geq 0.01 \text{ s}^{-1}$, with a total of about $7 \mu\text{g m}^{-3}$ SOA

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formed at the end of 12 h. Appreciable narrowing of the Aitken mode size distribution occurs for $k_c = 0.01 \text{ s}^{-1}$, which is qualitatively similar to the closed system results for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ shown previously in Fig. 16a. Higher k_c values produce even more intense narrowing of the Aitken mode and the shapes are practically indistinguishable from that for instantaneous reaction. As C_g^* increases, the solute vapor tends towards quasi-equilibrium with the particle phase for low k_c values. As a result, the SOA formation rate slows down and the Aitken mode shapes for $k_c = 0.01 \text{ s}^{-1}$ qualitatively tend to resemble that of Raoult's law partitioning in the closed system shown previously in Fig. 16b and c. But as k_c increases, the mass transfer becomes progressively more gas-phase diffusion limited, which results in faster growth of the smaller particles and, therefore, increasing narrowing of the Aitken mode.

Figure 19 shows the results for semi-solid particles ($D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$). It is seen that the presence of significant particle-phase diffusion limitation slows down the SOA formation rates, especially with increasing C_g^* and decreasing k_c . The marked size-dependence of the diffusion limitation also gives rise to more intense narrowing of the size distribution than seen in the corresponding liquid particle cases.

In the absence of particle-phase reaction (i.e., $k_c = 0$, not shown in the figures) only $\sim 1.2 \mu\text{g m}^{-3}$ SOA is formed in both the liquid and semi-solid aerosol cases after 12 h when $C_g^* = 10 \mu\text{g m}^{-3}$ while negligibly small amounts of SOA are formed for higher C_g^* values. Overall, the growth characteristics seen in the general system cases considered here are qualitatively similar to the closed system results, although significant differences between them can occur if the vapor source rate is appreciably different than the one used in the present study. For instance, if the vapor source rate is very small, then the growth characteristics will tend towards Raoult's law partitioning. In contrast, if the vapor source rate is very high, then the growth will tend to become gas-phase diffusion limited.

5 Summary and implications

We have extended the computationally efficient MOSAIC aerosol model (Zaveri et al., 2008) to include kinetic SOA partitioning that takes into account solute volatility, gas-phase diffusion, interfacial mass accommodation, particle-phase diffusion, and particle-phase reaction. The new model formulation uses a combination of: (a) an analytical quasi-steady-state treatment for the diffusion-reaction process within the particle phase for fast-reacting organic solutes such that the timescales (τ_{QSS}) for their particle-phase concentrations to reach quasi-steady state are shorter than 1 min, and (b) a two-film theory approach for slow- and non-reacting organic solutes. The updated MOSAIC model was successfully validated against a benchmark finite-difference solution of the diffusion-reaction problem. The new formulation is amenable for use in regional and global atmospheric models, although it currently awaits specification of the actual species and particle-phase reactions that are important for SOA formation.

In the present study, we have applied the model to evaluate the effects of solute volatility (C_g^*), particle-phase bulk diffusivity (D_b), and particle-phase chemical reaction, as exemplified by the pseudo-first-order rate constant (k_c), on kinetic SOA partitioning. We focus on the competitive growth dynamics of the Aitken and accumulation mode particles due to condensation while the Kelvin effect and coagulation are neglected for simplicity. Our analysis shows that the timescale of SOA partitioning and the associated evolution of aerosol number and composition size distributions depend on the complex interplay between C_g^* , D_b , and k_c , each of which can vary over several orders of magnitude. The key findings and their implications are summarized below.

- a. In the case of instantaneous particle-phase reaction ($k_c \rightarrow \infty$), SOA partitioning is mathematically equivalent to irreversible condensation of non-volatile organic vapors ($C_g^* = 0$; mechanism #1). Mass transfer is gas-phase diffusion limited, which produces the well-known narrowing of the aerosol size distribution as small particles grow faster than the large ones (Zhang et al., 2012).

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- b. In the case of non-reactive reversible absorption of semi-volatile and volatile organic vapors by Raoult's law ($k_c = 0$; mechanism #2), the final partitioning across the size distribution is volume-controlled (Zhang et al., 2012) and the partitioning timescale increases with decrease in C_g^* and D_b (Shiraiwa and Seinfeld, 2012b). In the absence of the Kelvin effect and coagulation, the mole fraction of SOA across the final size distribution at equilibrium is identical. As a result, the size distribution simply shifts along the diameter axis while its shape (mode widths and peak heights) remains unchanged. However, in a closed system, this mechanism may produce temporary narrowing of the size distribution as small particles reach quasi-equilibrium faster than the large ones (Zhang et al., 2012). The narrowing is especially pronounced if the pre-existing particles are highly viscous semi-solids ($D_b < 10^{-12} \text{ cm}^2 \text{ s}^{-1}$) and the initial gas-phase concentration is appreciably higher than the solute vapor volatility. Also, while the vapor concentration may reach a steady-state relatively quickly, the timescale for the "narrowed" aerosol size distribution to relax back to its final (equilibrium) shape can be of the order few minutes to days, depending on the values of D_b and C_g^* .
- c. In the case of reactive partitioning (finite k_c ; mechanism #3), the size distribution experiences permanent narrowing (Shiraiwa et al., 2013a), which can be especially pronounced for low values of C_g^* ($\sim 10 \mu\text{g m}^{-3}$ and less) and D_b ($< 10^{-13} \text{ cm}^2 \text{ s}^{-1}$) and high values of k_c ($\sim 0.01 \text{ s}^{-1}$ and higher). As C_g^* and D_b increase and k_c decreases, the narrowing reduces and the final size distribution tends to resemble that produced by mechanism #2. But unlike in mechanism #2, the gas-phase concentration of the solute eventually decays to zero and the partitioning timescale increases with increase in C_g^* and decrease in D_b and k_c . The partitioning timescale and the shape of the size distribution are especially sensitive to the phase state when D_b is about $10^{-13} \text{ cm}^2 \text{ s}^{-1}$ or less. At $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$ and $k_c = 0.01 \text{ s}^{-1}$, the decay timescale ranges from 1 h for $C_g^* = 10 \mu\text{g m}^{-3}$ to about 3 days for $C_g^* = 1000 \mu\text{g m}^{-3}$. Consequently, for interme-

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diate volatility solutes ($C_g^* > 1000 \mu\text{g m}^{-3}$) to partition in appreciable amounts to semi-solid SOA via particle-phase reactions, their k_c values need to be $> 0.1 \text{ s}^{-1}$.

- d. From a practical standpoint, the particle-phase concentration profiles of a solute (with any C_g^*) reacting with $k_c > 0.01 \text{ s}^{-1}$ may be assumed to be at steady-state in particles of any size and any phase state. Furthermore, for $k_c \leq 0.1 \text{ s}^{-1}$ and $D_b \geq 10^{-10} \text{ cm}^2 \text{ s}^{-1}$, the particle-phase reaction occurs uniformly through the entire volume of submicron particles. At higher k_c or lower D_b values, the particle-phase concentration profile becomes increasingly non-uniform (i.e., depleted towards the center of the particle) as the particle size increases. As a result, particle-phase reactions in large semi-solid particles occur primarily near the surface while in smaller particles the same reactions may still occur through the entire volume. These differences in the diffusion-reaction dynamics across the size distribution, and its dependence on the particle phase state, together control the SOA partitioning timescale and the size distribution evolution.
- e. Observations of the evolution of the size distribution can provide valuable clues about the underlying mechanisms of SOA formation (Riipinen et al., 2011; Shiraiwa et al., 2013a). However, since all three mechanisms, under certain combinations of C_g^* , D_b , and k_c values, can produce similar looking aerosol number size distributions, observations of the size distributions of the products of particle-phase reactions are also needed to effectively parameterize and evaluate the next generation of SOA models that treat phase state thermodynamics, particle-phase diffusion and particle-phase reactions.
- f. A proper representation of these physicochemical processes and parameters are needed to reliably predict not only the total SOA mass, but also its composition and number size distribution, which together determine its overall optical and cloud-nucleating properties.

Future model development work entails implementation of comprehensive gas-phase VOC oxidation mechanisms and the key particle-phase reactions that form organic salts, oligomers, hemiacetals, organosulfates, and other high molecular weight oxidation products, which constitute a significant fraction of SOA. At the same time, a computationally efficient treatment for phase transition thermodynamics is needed to provide the combined feedbacks of ambient temperature, relative humidity, and particle composition on the bulk diffusivity and reactivity of the absorbed organic solutes.

Appendix A

Overall gas-side mass transfer coefficient K_g

Flux F ($\text{mol cm}^{-2} \text{s}^{-1}$) of species i across the gas-particle interface can be written in multiple ways depending on the choice of the mass transfer coefficient:

$$\text{Gas-side mass transfer coefficient: } F_i = k_{g,i} (\bar{C}_{g,i} - C_{g,i}^s), \quad (\text{A1})$$

$$\text{Particle-side mass transfer coefficient: } F_i = k_{b,i} (A_i^s - \bar{A}_i), \quad (\text{A2})$$

$$\text{Overall gas-side mass transfer coefficient: } F_i = K_{g,i} (\bar{C}_{g,i} - S'_i \bar{A}_i). \quad (\text{A3})$$

In Eq. (A3) the term $(\bar{C}_{g,i} - S'_i \bar{A}_i)$ is the overall driving force for mass transfer between the bulk gas-phase and the average bulk particle phase, where

$$S'_i = \frac{C_i^*}{\sum_j \bar{A}_j}. \quad (\text{A4})$$

In the above equations, k_g (cm s^{-1}) is the gas-side mass transfer coefficient, k_b (cm s^{-1}) is the particle-side mass transfer coefficient, and K_g (cm s^{-1}) is the overall gas-side mass transfer coefficient.

Substituting the expression for Q_i from Eq. (8) in Eq. (B6), we get:

$$k_{b,i} = \frac{k_{c,i} R_p}{q_i^2} \left(\frac{q_i \coth q_i - 1}{1 - Q_i} \right). \quad (\text{B7})$$

Using $q_i^2 = R_p^2 k_{c,i} / D_{b,i}$ in Eq. (B7) gives:

$$k_{b,i} = \frac{D_{b,i}}{R_p} \left(\frac{q_i \coth q_i - 1}{1 - Q_i} \right). \quad (\text{B8})$$

5 The particle-side film thickness is then expressed as:

$$\delta_{b,i} = R_p \left(\frac{1 - Q_i}{q_i \coth q_i - 1} \right). \quad (\text{B9})$$

Note that as $k_{c,i} \rightarrow 0$, $q_i \rightarrow 0$, and $Q_i \rightarrow 1$. Taylor's series expansion of Eq. (B8) yields:

$$k_{b,i} = \frac{D_{b,i}}{R_p} \frac{\left(\frac{1}{3} - \frac{q^2}{45} + \frac{2q^4}{945} - \dots \right)}{\left(\frac{1}{15} - \frac{2q^2}{315} + \dots \right)}. \quad (\text{B10})$$

Thus, in the limiting case of non-reacting solute ($k_{c,i} = 0$), Eq. (B10) reduces to:

$$10 \quad k_{b,i} = 5 \frac{D_{b,i}}{R_p}. \quad (\text{B11})$$

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References

- Abramson, E., Imre, D., Beránek, J., Wilson, J., and Zelenyuk, A.: Experimental determination of chemical diffusion within secondary organic aerosol particles, *Phys. Chem. Chem. Phys.*, 15, 2983–2991, doi:10.1039/c2cp44013j, 2013.
- 5 Astarita, G.: *Mass Transfer with Chemical Reaction*, Elsevier, New York, 1967.
- Bird, R. B., Stewart, W. E., and Lightfoot, E. N.: *Transport Phenomena* (revised 2nd end.), John Wiley & Sons, Inc., Hoboken, NJ, 2007.
- Carlton, A. G., Turpin, B. I., Altieri, K. E., Seitzinger, S. P., Mathur, R., Roselle, S. J., and Weber, R. J.: CMAQ model performance enhanced when in-cloud secondary organic aerosol
- 10 is included: comparisons of organic carbon predictions with measurements, *Environ. Sci. Technol.*, 42, 8798–8802, 2008.
- Donahue, N. M., Trump, E. R., Pierce, J. R., and Riipinen, I.: Theoretical constraints on pure vapor-pressure driven condensation of organics to ultrafine particles, *Geophys. Res. Lett.*, 38, L16801, doi:10.1029/2011GL048115, 2011.
- 15 Doraiswamy, L. K. and Sharma, M. M.: *Heterogeneous Reactions: Analysis, Examples, and Reactor Design*, Vol. 2.: Fluid-Fluid-Solid Reactions, Wiley-Interscience, New York, 1984.
- Dusek, U., Frank, G. P., Hildebrandt, L., Curtius, J., Schneider, J., Walter S., Chand, D., Drewnick, F., Hings, S., Jung, D., Borrmann, S., and Andreae, M. O.: Size matters more than chemistry for cloud-nucleating ability of aerosol particles, *Science*, 312, 1375–1378,
- 20 doi:10.1126/science.1125261, 2006.
- Ervens, B. and Volkamer, R.: Glyoxal processing by aerosol multiphase chemistry: towards a kinetic modeling framework of secondary organic aerosol formation in aqueous particles, *Atmos. Chem. Phys.*, 10, 8219–8244, doi:10.5194/acp-10-8219-2010, 2010.
- Ervens, B., Carlton, A. G., Turpin, B. J., Altieri, K. E., Kreidenweis, S. M., and Feingold, G.: Secondary organic aerosol yields from cloud-processing of isoprene oxidation products, *Geophys. Res. Lett.*, 35, L02816, doi:10.1029/2007GL031828, 2008.
- 25 Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J., Yoo, K. Y., Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle phase acidity and oligomer formation in secondary organic aerosol, *Environ. Sci. Technol.*, 38, 6582–6589, doi:10.1021/es049125k, 2004.
- 30 Gunthe, S. S., King, S. M., Rose, D., Chen, Q., Roldin, P., Farmer, D. K., Jimenez, J. L., Artaxo, P., Andreae, M. O., Martin, S. T., and Pöschl, U.: Cloud condensation nuclei in

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pristine tropical rainforest air of Amazonia: size-resolved measurements and modeling of atmospheric aerosol composition and CCN activity, *Atmos. Chem. Phys.*, 9, 7551–7575, doi:10.5194/acp-9-7551-2009, 2009.

Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A., Kuwata, M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and Pöschl, U.: Cloud condensation nuclei (CCN) from fresh and aged air pollution in the megacity region of Beijing, *Atmos. Chem. Phys.*, 11, 11023–11039, doi:10.5194/acp-11-11023-2011, 2011.

Hall, W. A. and Johnston, M. V.: Oligomer content of α -pinene secondary organic aerosol, *Aerosol Sci. Tech.*, 45, 37–45, 2011.

Heaton, K. J., Dreyfus, M. A., Wang, S., and Johnston, M. V.: Oligomers in the early stage of biogenic secondary organic aerosol formation and growth, *Environ. Sci. Technol.*, 41, 6129–6139, 2007.

Hoyle, C. R., Boy, M., Donahue, N. M., Fry, J. L., Glasius, M., Guenther, A., Hallar, A. G., Huff Hartz, K., Petters, M. D., Petäjä, T., Rosenoern, T., and Sullivan, A. P.: A review of the anthropogenic influence on biogenic secondary organic aerosol, *Atmos. Chem. Phys.*, 11, 321–343, doi:10.5194/acp-11-321-2011, 2011.

Kalberer, M., Paulsen, D., Sax, M., Steinbacher, M., Dommen, J., Prevot, A. S. H., Fisseha, R., Weingartner, E., Frankevich, V., Zenobi, R., and Baltensperger, U.: Identification of polymers as major components of atmospheric organic aerosols, *Science*, 303, 1659–1662, 2004.

Kuang, C., Chen, M., Zhao, J., Smith, J., McMurry, P. H., and Wang, J.: Size and time-resolved growth rate measurements of 1 to 5 nm freshly formed atmospheric nuclei, *Atmos. Chem. Phys.*, 12, 3573–3589, doi:10.5194/acp-12-3573-2012, 2012.

Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a review of observations, *J. Aerosol Sci.*, 35, 143–176, 2004.

Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of low-volatility organics in the atmosphere, *Atmos. Environ.*, 42, 3593–3624, 2008.

Lewis, W. K. and Whitman, W. G.: Principles of gas absorption, *Ind. Eng. Chem.*, 16, 1215–1220, 1924.

Mouchel-Vallon, C., Bräuer, P., Camredon, M., Valorso, R., Madronich, S., Herrmann, H., and Aumont, B.: Explicit modeling of volatile organic compounds partitioning in the atmo-

spheric aqueous phase, *Atmos. Chem. Phys.*, 13, 1023–1037, doi:10.5194/acp-13-1023-2013, 2013.

Nozière, B., Dziedzic, P., and Córdoba, A.: Formation of secondary light-absorbing “fulvic-like” oligomers: a common process in aqueous and ionic atmospheric particles?, *Geophys. Res. Lett.*, 34, L21812, doi:10.1029/2007GL031300, 2007.

Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flen-
tje, H., Birmili, W., Wiedensohler, A., Hörrak, U., Metzger, A., Hamed, A., Laaksonen, A.,
Facchini, M. C., Kerminen, V.-M., and Kulmala, M.: On the roles of sulphuric acid and low-
volatility organic vapours in the initial steps of atmospheric new particle formation, *Atmos.*
Chem. Phys., 10, 11223–11242, doi:10.5194/acp-10-11223-2010, 2010.

Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in the
atmosphere, *Atmos. Environ.*, 28, 185–188, 1994.

Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Yu, Y., Alexander, M. L., Zelenyuk, A.,
Imre, D., Chang, W. L., Dabdub, D., Pankow, J. F., and Finlayson-Pitts, B. J.: Nonequilibrium
atmospheric secondary organic aerosol formation and growth, *P. Natl. Acad. Sci. USA*, 109,
2836–2841, doi:10.1073/pnas.1119909109, 2012.

Pierce, J. R., Riipinen, I., Kulmala, M., Ehn, M., Petäjä, T., Junninen, H., Worsnop, D. R., and
Donahue, N. M.: Quantification of the volatility of secondary organic compounds in ultrafine
particles during nucleation events, *Atmos. Chem. Phys.*, 11, 9019–9036, doi:10.5194/acp-
11-9019-2011, 2011.

Pierce, J. R., Leaitch, W. R., Liggio, J., Westervelt, D. M., Wainwright, C. D., Abbatt, J. P. D.,
Ahlm, L., Al-Basheer, W., Cziczo, D. J., Hayden, K. L., Lee, A. K. Y., Li, S.-M., Russell, L. M.,
Sjostedt, S. J., Strawbridge, K. B., Travis, M., Vlasenko, A., Wentzell, J. J. B., Wiebe, H. A.,
Wong, J. P. S., and Macdonald, A. M.: Nucleation and condensational growth to CCN sizes
during a sustained pristine biogenic SOA event in a forested mountain valley, *Atmos. Chem.*
Phys., 12, 3147–3163, doi:10.5194/acp-12-3147-2012, 2012.

Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, M., Sellier, M., Murray, B. J.,
Shilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of α -pinene secondary organic
material and implications for particle growth and reactivity, *P. Natl. Acad. Sci.*, 110, 8014–
8019, doi:10.1073/pnas.1219548110, 2013.

Riipinen, I., Pierce, J. R., Yli-Juuti, T., Nieminen, T., Häkkinen, S., Ehn, M., Junninen, H., Lehti-
palo, K., Petäjä, T., Slowik, J., Chang, R., Shantz, N. C., Abbatt, J., Leaitch, W. R., Ker-
minen, V.-M., Worsnop, D. R., Pandis, S. N., Donahue, N. M., and Kulmala, M.: Organic

ACPD

13, 28631–28694, 2013

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- condensation: a vital link connecting aerosol formation to cloud condensation nuclei (CCN) concentrations, *Atmos. Chem. Phys.*, 11, 3865–3878, doi:10.5194/acp-11-3865-2011, 2011.
- Riipinen, I., Yli-Juuti, T., Pierce, J. R., Petäjä, T., Worsnop, D. R., Kulmala, M., and Donahue, N. M.: The contribution of organics to atmospheric nanoparticle growth, *Nat. Geosci.*, 5, 453–458, 2012.
- Seinfeld, J. H. and Pandis, S. N.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change*, 2nd edn., John Wiley & Sons, Inc., Hoboken, NJ, 2006.
- Schwartz, S. E. and Freiberg, J. E.: Mass-transport limitation to the rate of reaction of gases in liquid droplets: application to oxidation of SO₂ in aqueous solutions, *Atmos. Environ.*, 15, 1129–1144, 1981.
- Shi, B. and Seinfeld, J. H.: On mass-transport limitation to the rate of reaction of gases in liquid droplets, *Atmos. Environ.*, 22, 2491–2499, 1991.
- Shilling, J. E., Zaveri, R. A., Fast, J. D., Kleinman, L., Alexander, M. L., Canagaratna, M. R., Fortner, E., Hubbe, J. M., Jayne, J. T., Sedlacek, A., Setyan, A., Springston, S., Worsnop, D. R., and Zhang, Q.: Enhanced SOA formation from mixed anthropogenic and biogenic emissions during the CARES campaign, *Atmos. Chem. Phys.*, 13, 2091–2113, doi:10.5194/acp-13-2091-2013, 2013.
- Shiraiwa, M. and Seinfeld, J. H.: Equilibration timescale of atmospheric secondary organic aerosol partitioning, *Geophys. Res. Lett.*, 39, L24801, doi:10.1029/2012GL054008, 2012.
- Shiraiwa, M., Pfrang, C., Koop, T., and Pöschl, U.: Kinetic multi-layer model of gas-particle interactions in aerosols and clouds (KM-GAP): linking condensation, evaporation and chemical reactions of organics, oxidants and water, *Atmos. Chem. Phys.*, 12, 2777–2794, doi:10.5194/acp-12-2777-2012, 2012.
- Shiraiwa, M., Yee, L. D., Schilling, K. A., Loza, C. L., Craven, J. S., Zuend, A., Ziemann, P. J., and Seinfeld, J. H.: Size distribution dynamics reveal particle-phase chemistry in organic aerosol formation, *P. Natl. Acad. Sci. USA*, 110, 11746–11750, doi:10.1073/pnas.1307501110, 2013a.
- Shiraiwa, M., Zuend, A., Bertram, A. K., and Seinfeld, J. H.: Gas-particle partitioning of atmospheric aerosols: interplay of physical state, non-ideal mixing and morphology, *Phys. Chem. Chem. Phys.*, 15, 11441–11453, 2013b.
- Simmel, M. and Wurzler, S.: Condensation and activation in sectional cloud microphysical models, *Atmos. Res.*, 80, 218–236, 2006.

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- Smith, J. N., Dunn, M. J., VanReken, T. M., Iida, K., Stolzenburg, M. R., McMurry, P. H., and Huey, L. G.: Chemical composition of atmospheric nanoparticles formed from nucleation in Tecamac, Mexico: evidence for an important role for organic species in nanoparticle growth, *Geophys. Res. Lett.*, 35, L04808, doi:10.1029/2007GL032523, 2008.
- 5 Surratt, J. D., Kroll, J. H., Kleindienst, T. E., Edney, E. O., Claeys, M., Sorooshian, A., Ng, N. L., Offenberg, J. H., Lewandowski, M., Jaoui, M., Flagan, R. C., and Seinfeld, J. H.: Evidence for organosulfates in secondary organic aerosol, *Environ. Sci. Technol.*, 41, 517–527, 2007.
- Tolocka, M., Jang, M., Ginter, J. M., Cox, F. J., Kamens, R. M., and Johnston, M. V.: Formation of oligomers in secondary organic aerosol, *Environ. Sci. Technol.*, 38, 1428–1434, 2004.
- 10 Vaden, T. D., Imre, D., Beránek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation kinetics and phase of laboratory and ambient secondary organic aerosol, *P. Natl. Acad. Sci. USA*, 108, 2190–2195, 2011.
- Virtanen, A., Joutsensaari, J., Koop, T., Kannosto, J., Yli-Pirila, P., Leskinen, J., Makela, J. M., Holopainen, J. K., Poschl, U., Kulmala, M., Worsnop, D. R., and Laaksonen, A.: An amorphous solid state of biogenic secondary organic aerosol particles, *Nature*, 467, 824–827, 2010.
- 15 Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V., and Zhang, R.: Atmospheric nanoparticles formed from heterogeneous reactions of organics, *Nat. Geosci.*, 3, 238–242, 2010.
- 20 Weber, R. J., Sullivan, A. P., Peltier, R. E., Russell, A., Yan, B., Zheng, M., de Gouw, J., Warneke, C., Brock, C., Holloway, J. S., Atlas, E. L., and Edgerton, E.: A study of secondary organic aerosol formation in the anthropogenic-influenced southeastern United States, *J. Geophys. Res.-Atmos.*, 112, D13302, doi:10.1029/2007jd008408, 2007.
- Winkler, P. M., Ortega, J., Karl, T., Cappellin, L., Friedli, H. R., Barsanti, K., McMurry, P. H., and Smith, J. N.: Identification of the biogenic compounds responsible for size-dependent nanoparticle growth, *Geophys. Res. Lett.*, 39, L20815, doi:10.1029/2012GL053253, 2012.
- 25 Zaveri, R. A.: Development and evaluation of a comprehensive tropospheric chemistry model for regional and global applications, Ph.D. thesis, Virginia Polytech. Inst., and State Univ., Blacksburg, Virginia, 1997.
- 30 Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for Simulating Aerosol Interactions and Chemistry (MOSAIC), *J. Geophys. Res.*, 113, D13204, doi:10.1029/2007JD008782, 2008.

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- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M. R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K., DeCarlo, P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimojo, A., Hatakeyama, S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen, J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.: Ubiquity and dominance of oxygenated species in organic aerosols in anthropogenically-influenced Northern Hemisphere mid-latitudes, *Geophys. Res. Lett.*, 34, L13801, doi:10.1029/2007GL029979, 2007.
- Zhang, X., Pandis, S. N., and Seinfeld, J. H.: Diffusion-limited vs. quasi-equilibrium aerosol growth, *Aerosol Sci. Tech.*, 46, 874–885, doi:10.1080/02786826.2012.679344, 2012.
- Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol formation, *Chem. Soc. Rev.*, 41, 6582–6605, 2012.

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Table 1. Bias and error statistics for MOSAIC predictions for the closed system simulations.

k_c (s^{-1})	$C_g^* = 10 \mu g m^{-3}$			$C_g^* = 100 \mu g m^{-3}$			$C_g^* = 1000 \mu g m^{-3}$		
	MNB (%)	MNGE (%)	maxNGE (%)	MNB (%)	MNGE (%)	maxNGE (%)	MNB (%)	MNGE (%)	maxNGE (%)
0	4.5	4.5	7.7	0.3	0.3	0.4	0.03	0.03	0.1
10^{-4}	8.5	11.3	19.4	−1.7	1.7	3.1	−0.3	0.3	0.4
10^{-3}	10.0	11.3	25.7	−1.3	1.3	3.2	−0.1	0.1	0.2
10^{-2}	−1.3	1.3	3.4	3.4	3.4	4.2	−0.3	0.3	1.0
10^{-1}	−4.3	4.3	10.7	−2.2	2.6	7.7	0.5	0.7	1.1

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Table 2. Bias and error statistics for MOSAIC predictions for the general system simulations.

k_c (s^{-1}) (s^{-1})	$C_g^* = 10 \mu g m^{-3}$			$C_g^* = 100 \mu g m^{-3}$			$C_g^* = 1000 \mu g m^{-3}$		
	MNB (%)	MNGE (%)	maxNGE (%)	MNB (%)	MNGE (%)	maxNGE (%)	MNB (%)	MNGE (%)	maxNGE (%)
0	0.8	0.8	8.5	0.3	0.3	2.0	0.03	0.03	0.3
10^{-4}	−1.0	2.2	6.4	0.6	0.6	1.8	0.1	0.1	0.2
10^{-3}	−3.0	3.1	5.8	−0.7	1.2	3.1	0.1	0.1	0.2
10^{-2}	−3.2	3.2	5.8	−2.3	2.3	4.8	−0.2	0.2	0.8
10^{-1}	−2.4	2.4	5.0	−0.2	1.4	2.8	1.0	1.0	1.6

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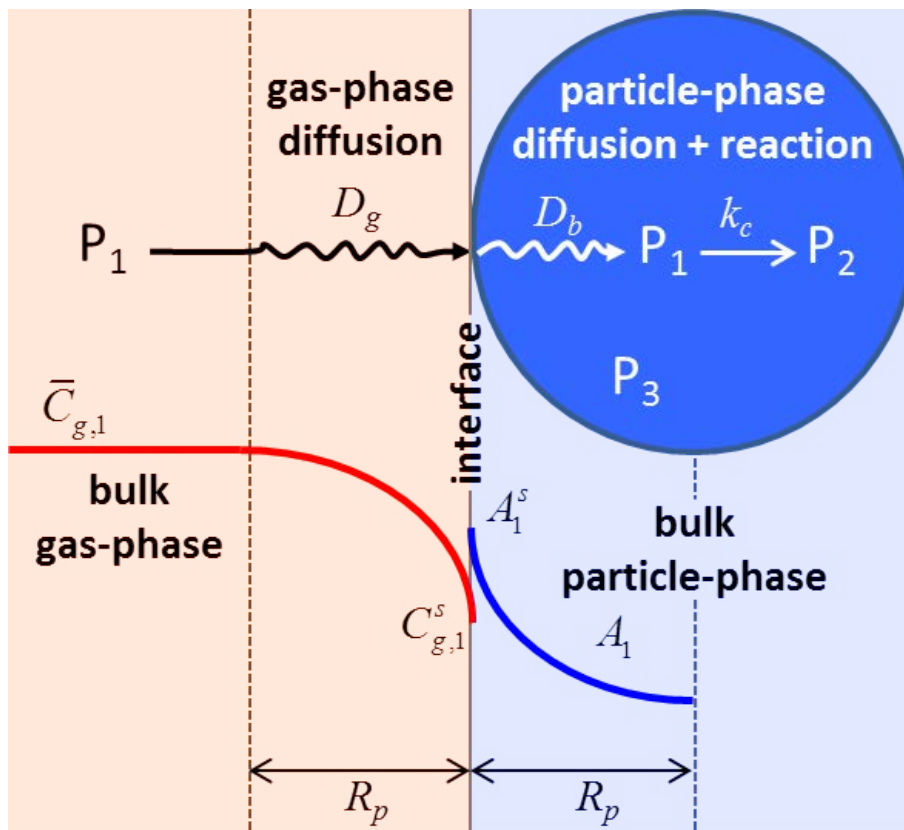


Fig. 1. Schematic of the gas-particle mass transfer process, with both diffusion and reaction occurring inside the particle phase.

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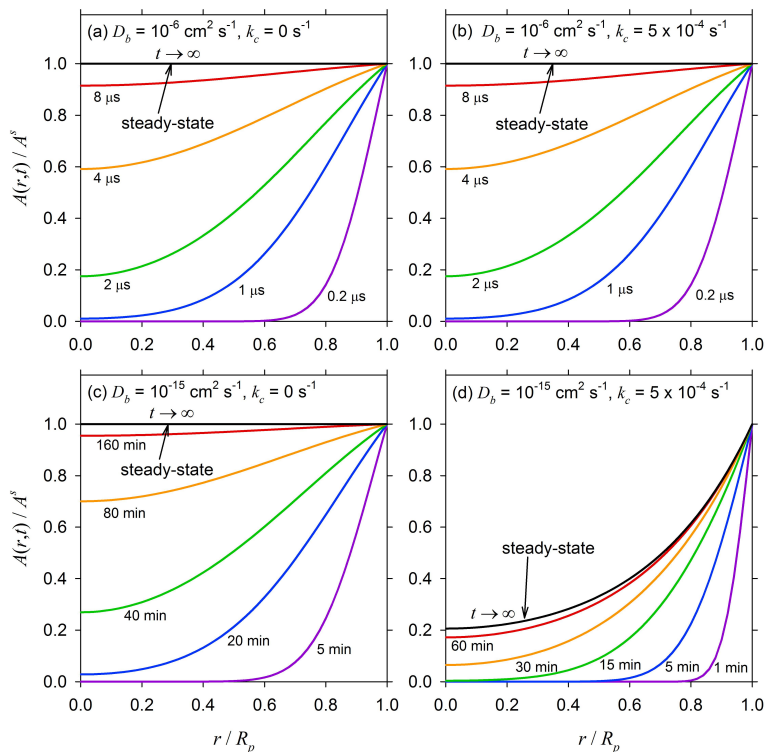


Fig. 2. Normalized transient concentration ($A(r,t)/A^s$) profiles as a function of normalized radius (r/R_p) for a particle of diameter $R_p = 0.05 \mu\text{m}$ for different values of bulk phase diffusivity and first-order reaction rate constants: **(a)** $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$, $k_c = 0 \text{ s}^{-1}$; **(b)** $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$, $k_c = 5 \times 10^{-4} \text{ s}^{-1}$; **(c)** $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$, $k_c = 0 \text{ s}^{-1}$; and **(d)** $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$, $k_c = 5 \times 10^{-4} \text{ s}^{-1}$.

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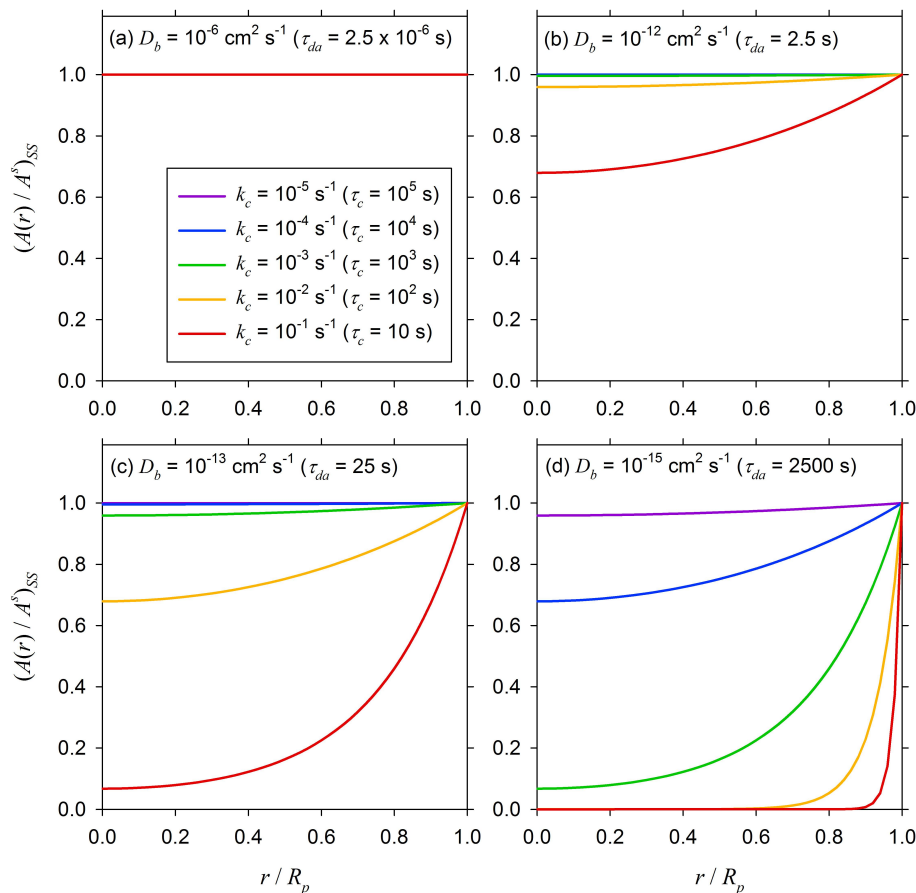


Fig. 3. Normalized steady-state concentration $(A(r)/A^s)_{SS}$ profiles as a function of normalized radius (r/R_p) for a particle of diameter $R_p = 0.05 \mu\text{m}$ and a range of k_c values for (a) $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ (b) $D_b = 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ (c) $D_b = 10^{-13} \text{ cm}^2 \text{ s}^{-1}$ (d) $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$.

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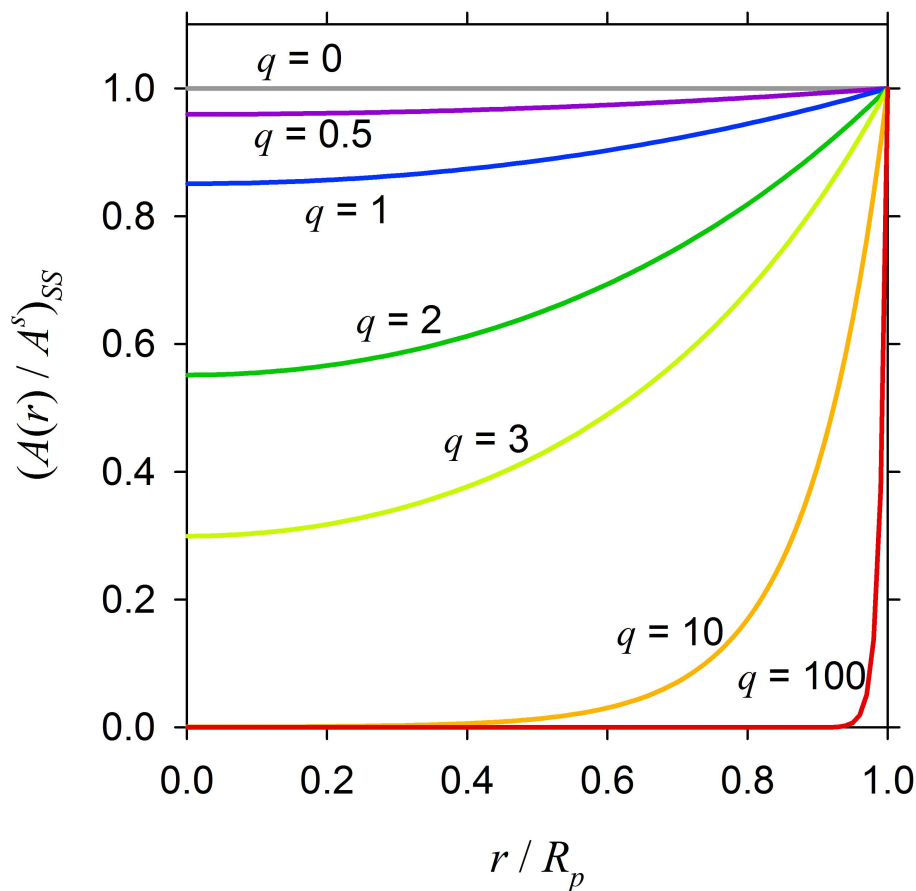


Fig. 4. Normalized steady state concentration $(A(r)/A^s)_{ss}$ profiles as a function of normalized radius (r/R_p) for different values of dimensionless diffuso-reactive parameter q .

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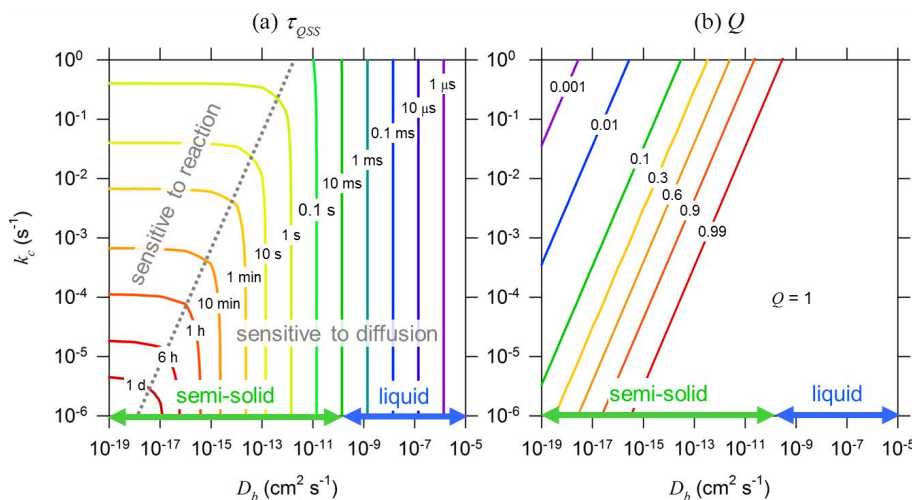


Fig. 5. (a) Contour plots of: (a) particle-phase quasi-steady state timescale (τ_{QSS}), and (b) quasi-steady state parameter $Q = (\bar{A}/A^s)_{QSS}$ as functions of first-order rate constant (k_c) and bulk diffusion coefficient (D_b) for a species diffusing and reacting within semi-solid and liquid particles of diameter $D_p = 0.1 \mu\text{m}$.

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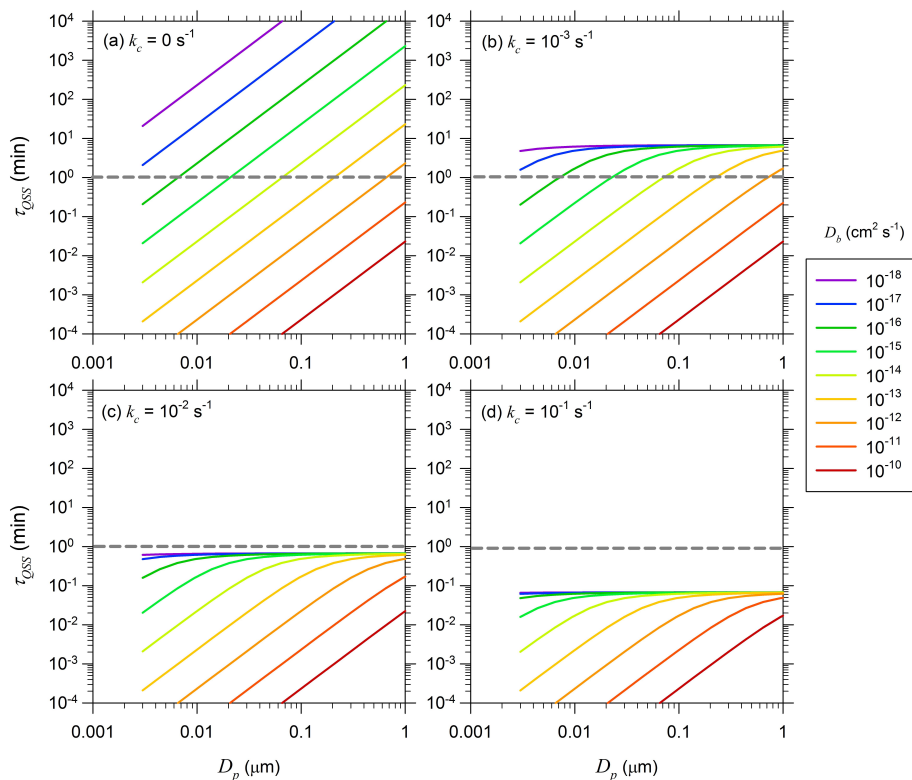


Fig. 6. Dependence of τ_{QSS} on particle diameter D_p for D_b values ranging from 10^{-18} to $10^{-10} \text{ cm}^2 \text{ s}^{-1}$: **(a)** $k_c = 0 \text{ s}^{-1}$; **(b)** $k_c = 10^{-3} \text{ s}^{-1}$; **(c)** $k_c = 10^{-2} \text{ s}^{-1}$; and **(d)** $k_c = 10^{-1} \text{ s}^{-1}$.

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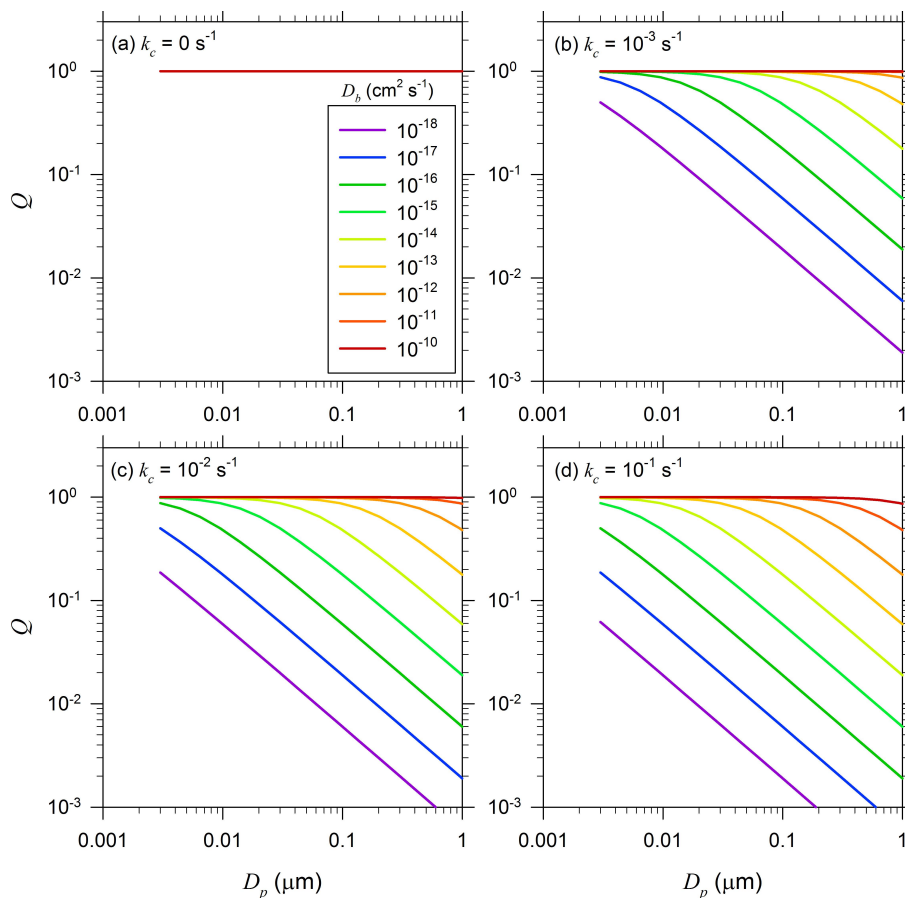


Fig. 7. Dependence of Q on particle diameter D_p for D_b values ranging from 10^{-10} to $10^{-18} \text{ cm}^2 \text{ s}^{-1}$: **(a)** $k_c = 0 \text{ s}^{-1}$; **(b)** $k_c = 10^{-3} \text{ s}^{-1}$; **(c)** $k_c = 10^{-2} \text{ s}^{-1}$; and **(d)** $k_c = 10^{-1} \text{ s}^{-1}$.

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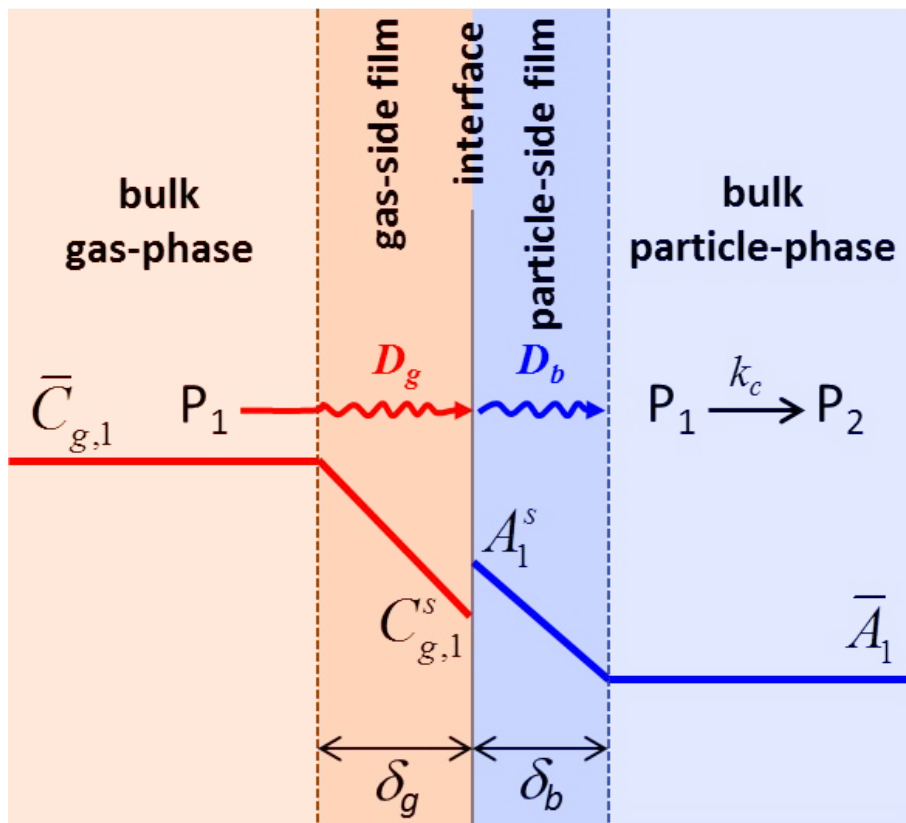


Fig. 8. Schematic of the two-film theory.

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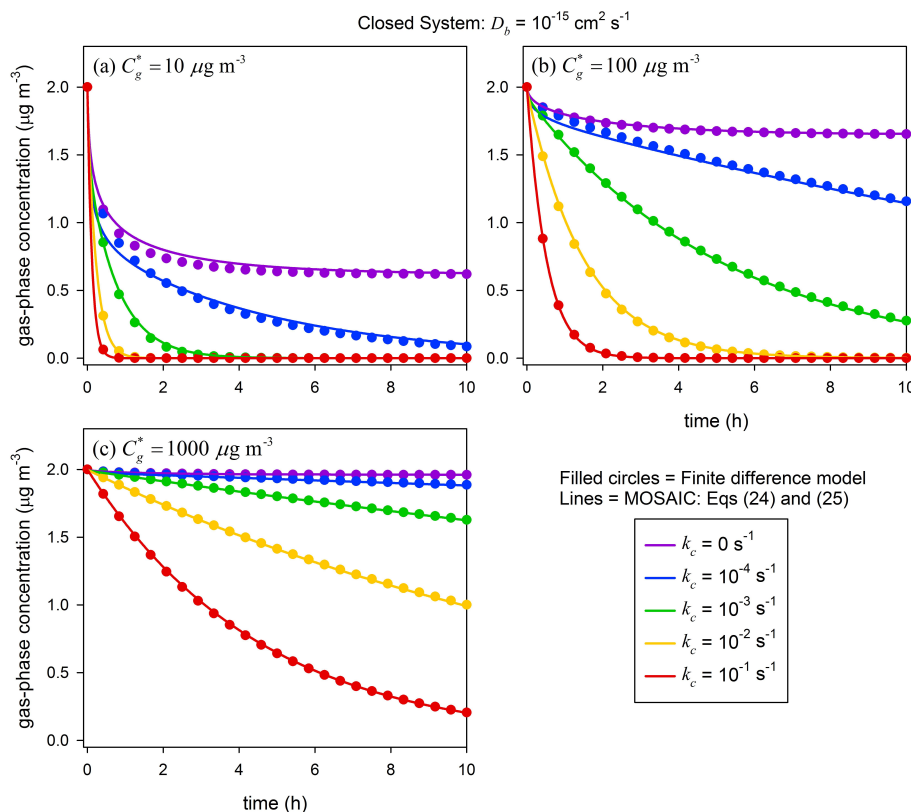


Fig. 9. Comparison of MOSAIC (lines) and finite difference model (filled circles) solutions for gas-phase concentration decay due to kinetic gas-particle partitioning to particles with initial $D_p = 0.2 \mu\text{m}$, $N = 5000 \text{ cm}^{-3}$, $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$ and k_c ranging from 0 to 0.1 s^{-1} for three solute volatilities: **(a)** $C_g^* = 10 \mu\text{g m}^{-3}$, **(b)** $C_g^* = 100 \mu\text{g m}^{-3}$, and **(c)** $C_g^* = 1000 \mu\text{g m}^{-3}$.

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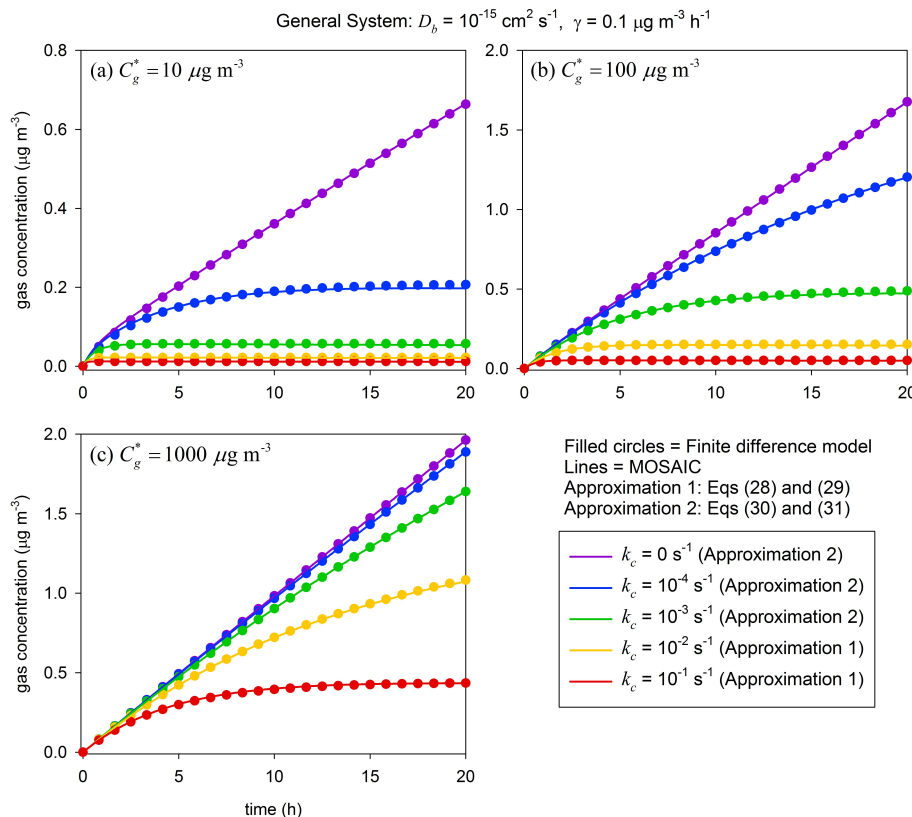


Fig. 10. Comparison of MOSAIC (lines) and finite difference model (filled circles) solutions for gas-phase concentration evolution in a general system due to kinetic gas-particle partitioning to particles with initial $D_p = 0.2 \mu\text{m}$, $N = 5000 \text{ cm}^{-3}$, $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$, $\gamma = 0.1 \mu\text{g m}^{-3} \text{ h}^{-1}$, and k_c ranging from 0 to 0.1 s^{-1} for three solute volatilities: **(a)** $C_g^* = 10 \mu\text{g m}^{-3}$, **(b)** $C_g^* = 100 \mu\text{g m}^{-3}$, and **(c)** $C_g^* = 1000 \mu\text{g m}^{-3}$.

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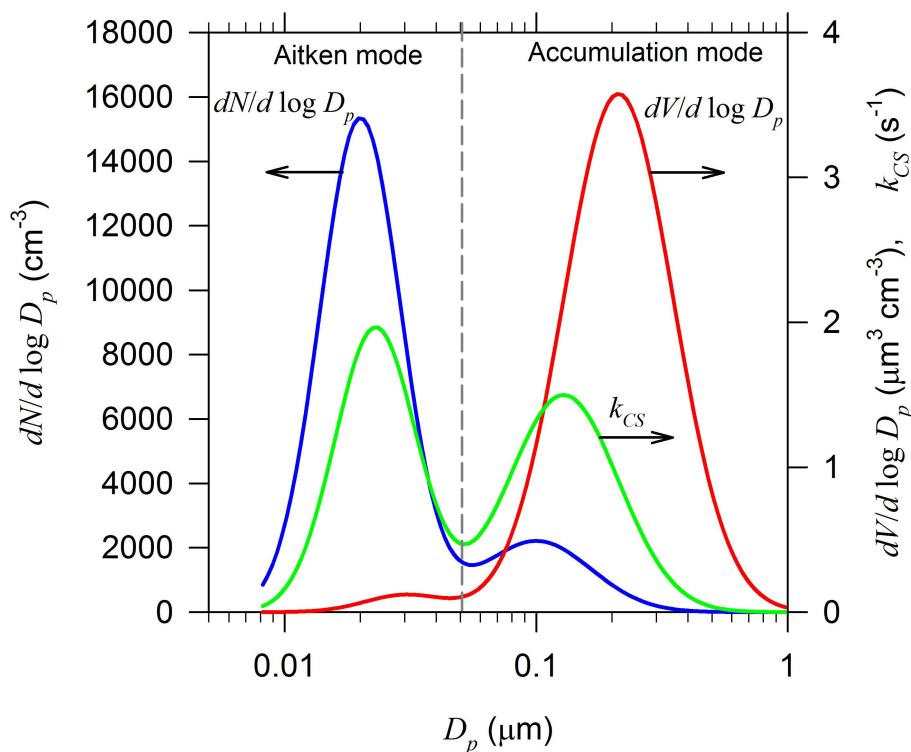


Fig. 11. Initial aerosol number and volume size distributions along with the condensational sink k_{CS} . The dashed line demarcates the Aitken mode from the accumulation mode and the initial condensation sink is such that the sum of k_{CS} over all the size bins in the Aitken mode is equal to that in the accumulation mode.

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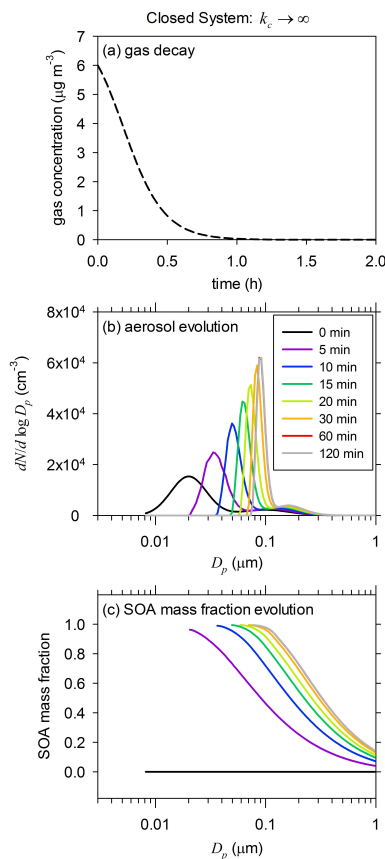


Fig. 12. Results for the instantaneous reaction reference case ($k_c \rightarrow \infty$; equivalent to non-volatile solute condensation): **(a)** gas-phase concentration decay; **(b)** temporal evolution of aerosol size distribution; and **(c)** temporal evolution of the mass fraction of newly formed SOA.

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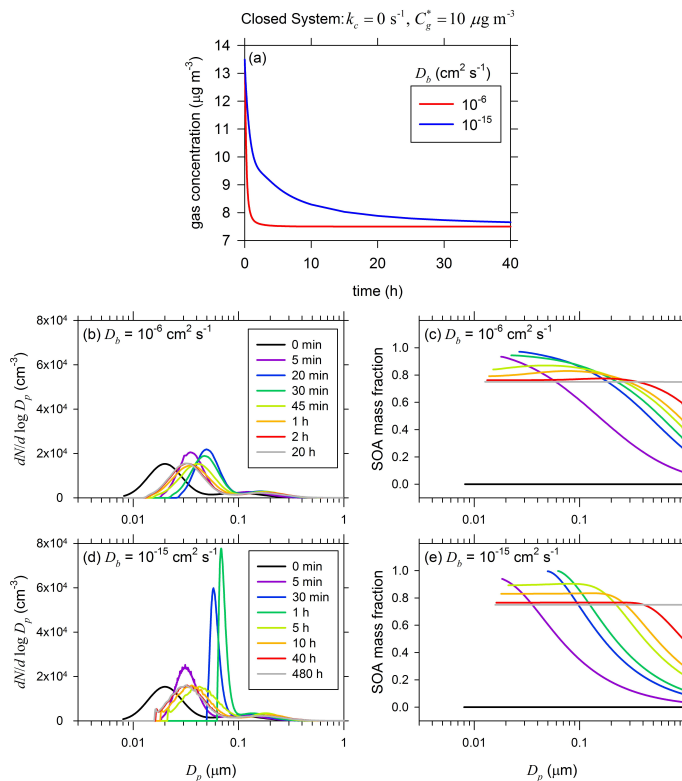


Fig. 13. Results for kinetic SOA partitioning due to Raoult's law ($k_c = 0 \text{ s}^{-1}$) for $C_g^* = 10 \text{ μg m}^{-3}$: **(a)** gas-phase concentration decay for $D_b = 10^{-6}$ and $10^{-15} \text{ cm}^2 \text{ s}^{-1}$, **(b)** aerosol evolution for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$; **(c)** SOA mass fraction evolution for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$; **(d)** aerosol evolution for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$; and **(e)** SOA mass fraction evolution for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$. In both cases, the final (i.e., equilibrium) concentration of the newly formed SOA is 6 μg m^{-3} .

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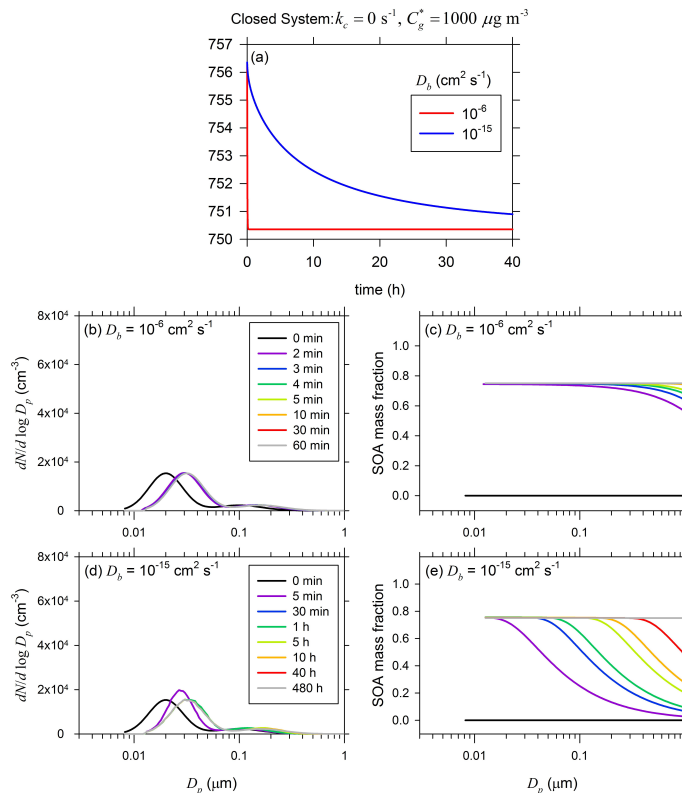


Fig. 14. Same as Fig. 13, except $C_g^* = 1000 \mu\text{g m}^{-3}$.

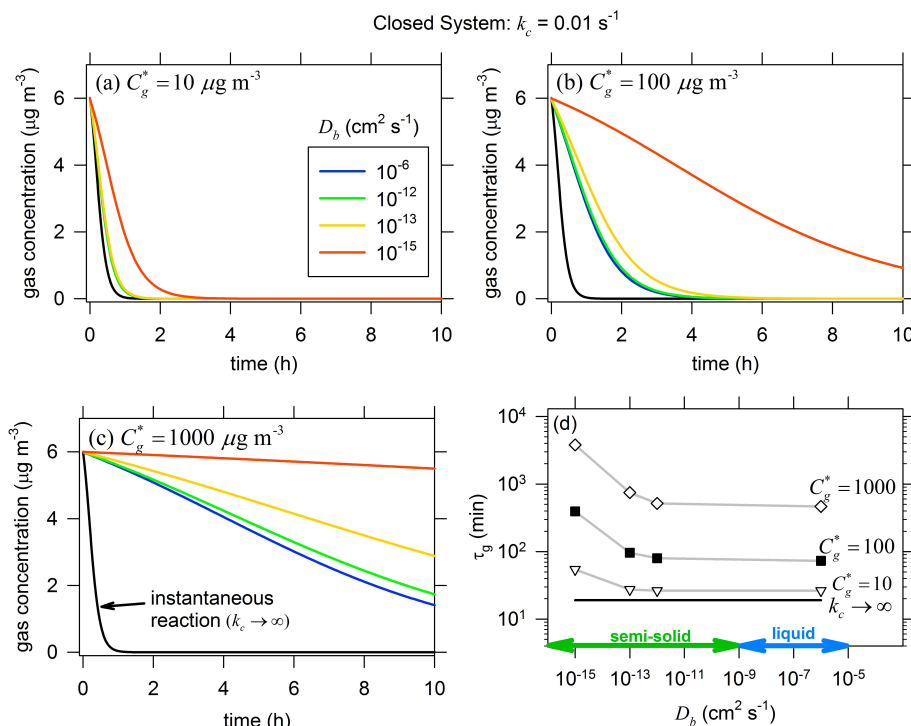


Fig. 15. Gas-phase concentration decay due to kinetic SOA partitioning with particle-phase reaction ($k_c = 0.01 \text{ s}^{-1}$) for bulk diffusivities ranging from 10^{-6} to $10^{-15} \text{ cm}^2 \text{ s}^{-1}$ and three gas volatilities: **(a)** $C_g^* = 10 \mu\text{g m}^{-3}$; **(b)** $C_g^* = 100 \mu\text{g m}^{-3}$; and **(c)** $C_g^* = 1000 \mu\text{g m}^{-3}$. Each plot also shows gas-phase concentration decay for the reference case of instantaneous reaction (black line, $k_c \rightarrow \infty$). In each case, the final concentration of the newly formed SOA is $6 \mu\text{g m}^{-3}$. Panel **(d)** shows the plot of gas-phase concentration decay timescale (τ_g) as a function of D_b for the different gas volatilities.

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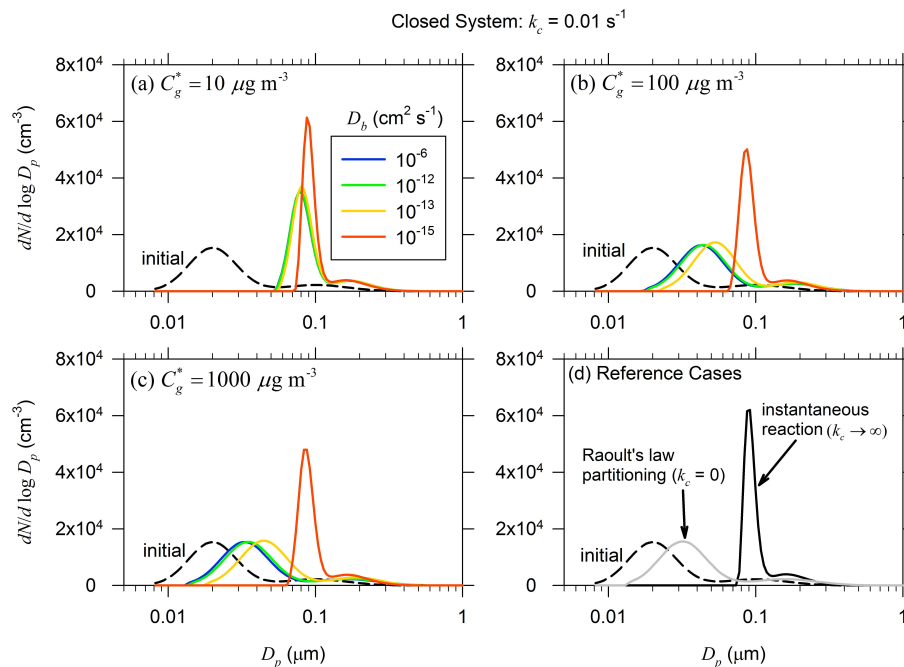


Fig. 16. Initial (dashed line) and final (solid lines) aerosol number size distribution due to Raoult's law gas-particle partitioning coupled with particle-phase reaction ($k_c = 0.01 \text{ s}^{-1}$) for bulk diffusivities ranging from 10^{-6} to $10^{-15} \text{ cm}^2 \text{ s}^{-1}$ and three gas volatilities: **(a)** $C_g^* = 10 \mu\text{g m}^{-3}$; **(b)** $C_g^* = 100 \mu\text{g m}^{-3}$; and **(c)** $C_g^* = 1000 \mu\text{g m}^{-3}$. Panel **(d)** shows the final size distributions for the two reference cases: instantaneous reaction (black line; $k_c \rightarrow \infty$) and Raoult's law partitioning (gray line; $k_c = 0$) for any D_b and $C_g^* > 0$. As illustrated in Fig. 12, the time required to reach the final state differs significantly for different cases, but the final SOA formed in each case is $6 \mu\text{g m}^{-3}$.

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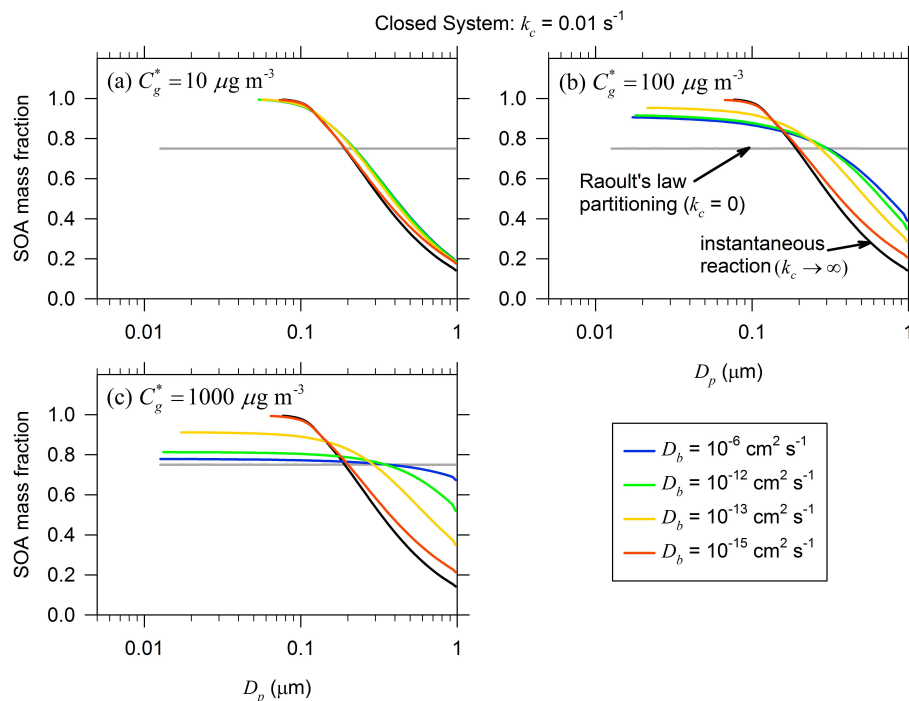


Fig. 17. Final size distributions of the newly formed SOA mass fraction for different D_b values and: **(a)** $C_g^* = 10 \mu\text{g m}^{-3}$; **(b)** $C_g^* = 100 \mu\text{g m}^{-3}$; and **(c)** $C_g^* = 1000 \mu\text{g m}^{-3}$. Each panel also shows the reference plots for instantaneous reaction (black line; $k_c \rightarrow \infty$) and for Raoult's law partitioning (gray line; $k_c = 0 \text{ s}^{-1}$) for any D_b and $C_g^* > 0$.

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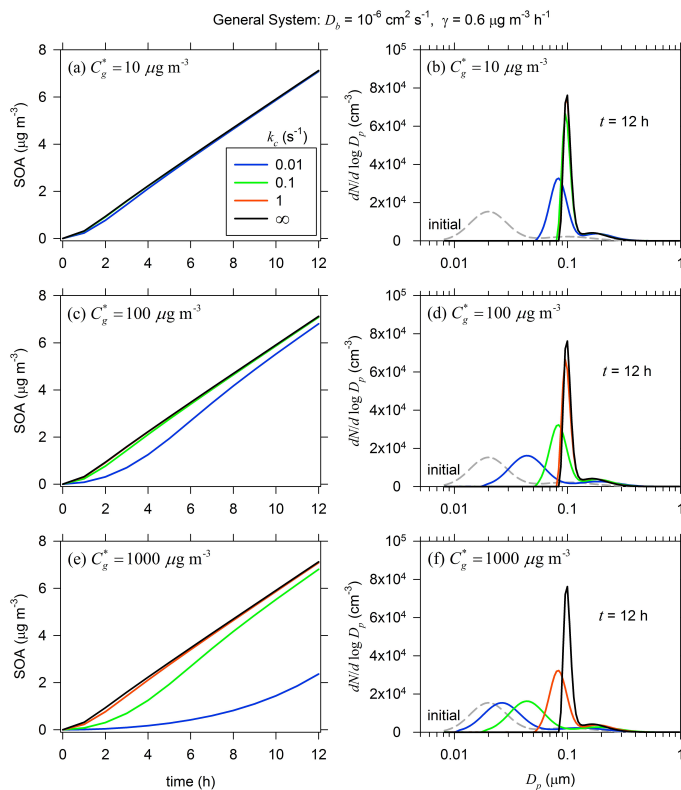


Fig. 18. Temporal evolution of total SOA mass concentration (left column) and aerosol size distribution (right column) at $t = 12 \text{ h}$ for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$, $\gamma = 0.6 \mu\text{g m}^{-3} \text{ h}^{-1}$, $k_c = 0.01$ to $\infty \text{ s}^{-1}$, and three different solute volatilities: **(a, b)** $C_g^* = 10 \mu\text{g m}^{-3}$; **(c, d)** $C_g^* = 100 \mu\text{g m}^{-3}$; and **(e, f)** $C_g^* = 1000 \mu\text{g m}^{-3}$.

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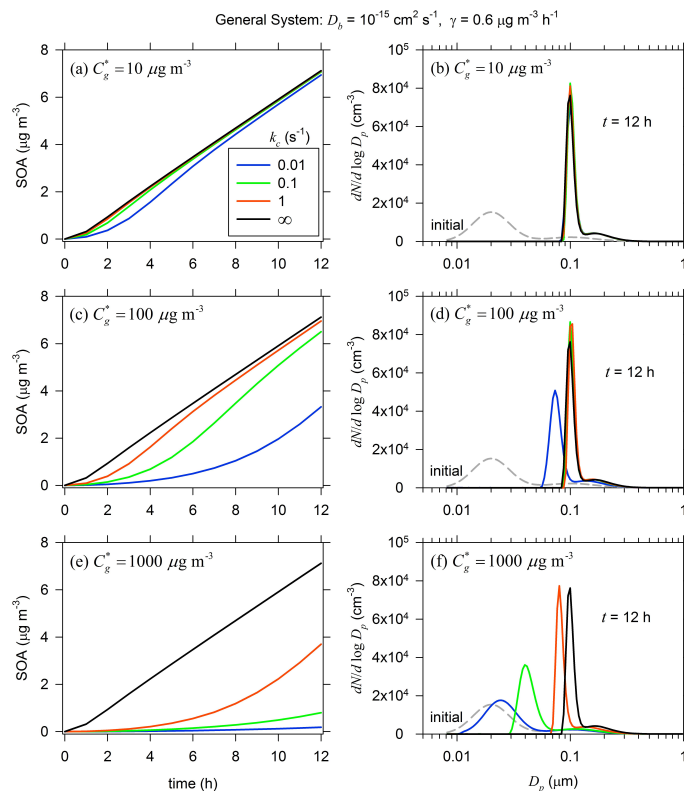


Fig. 19. Same as Fig. 18, except $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$.