Modeling kinetic partitioning of secondary organic aerosol and size distribution dynamics: representing effects of volatility, phase state, and particle-phase reaction

4

5 R. A. Zaveri¹, R. C. Easter¹, J. E. Shilling¹, and J. H. Seinfeld²

6 [1]{Atmospheric Sciences and Global Change Division, Pacific Northwest National7 Laboratory, Richland, WA, USA}

- 8 [2]{Division of Chemistry and Chemical Engineering and Division of Engineering and
- 9 Applied Science, California Institute of Technology, Pasadena, CA, USA}
- 10 Correspondence to: R. A. Zaveri (rahul.zaveri@pnnl.gov)
- 11

12 Abstract

13 This paper describes and evaluates a new framework for modeling kinetic gas-particle 14 partitioning of secondary organic aerosol (SOA) that takes into account diffusion and 15 chemical reaction within the particle phase. The framework uses a combination of: (a) an 16 analytical quasi-steady-state treatment for the diffusion-reaction process within the particle 17 phase for fast-reacting organic solutes, and (b) a two-film theory approach for slow- and non-18 reacting solutes. The framework is amenable for use in regional and global atmospheric 19 models, although it currently awaits specification of the various gas- and particle-phase 20 chemistries and the related physicochemical properties that are important for SOA formation. 21 Here, the new framework is implemented in the computationally efficient Model for 22 Simulating Aerosol Interactions and Chemistry (MOSAIC) to investigate the competitive growth dynamics of the Aitken and accumulation mode particles. Results show that the 23 timescale of SOA partitioning and the associated size distribution dynamics depend on the 24 25 complex interplay between organic solute volatility, particle-phase bulk diffusivity, and particle-phase reactivity (as exemplified by a pseudo-first-order reaction rate constant), each 26 of which can vary over several orders of magnitude. In general, the timescale of SOA 27 28 partitioning increases with increase in volatility and decrease in bulk diffusivity and rate 29 constant. At the same time, the shape of the aerosol size distribution displays appreciable

narrowing with decrease in volatility and bulk diffusivity and increase in rate constant. A proper representation of these physicochemical processes and parameters is needed in the next generation models to reliably predict not only the total SOA mass, but also its compositionand number-diameter distributions, all of which together determine the overall optical and cloud-nucleating properties.

6

7 1 Introduction

Submicron sized atmospheric aerosol particles are typically composed of ammonium, sulfate, 8 9 nitrate, black carbon, organics, sea salt, mineral dust, and water that are often internally mixed with each other in varying proportions. Depending on their dry state composition and overall 10 11 hygroscopicity, aerosol particles in the size range 0.03-0.1 µm (dry diameter) and larger may 12 act as cloud condensation nuclei (CCN) (Dusek et al., 2006; Gunthe et al., 2009, 2011) while 13 those larger than 0.1 µm (wet diameter) efficiently scatter solar radiation. Aerosol number and composition size distributions, therefore, together hold the key to determining its overall 14 15 climate-relevant properties.

16 Organic compounds constitute 20-90% of the submicron aerosol mass and are thought to play 17 a vital role in both the direct and indirect aerosol radiative forcing of climate (Kanakidou et 18 al., 2005). While primary organic aerosols (POA) from fossil fuel combustion and biomass burning are directly emitted into the submicron size range, the dominant source of organic 19 20 aerosols is secondary, which involves gas-to-particle conversion of many different volatile organic compounds (VOCs) of both anthropogenic and biogenic origin (Zhang et al., 2007). 21 22 Furthermore, biogenic VOCs are estimated to be the dominant source of secondary organic 23 aerosol (SOA), but their formation appears to be strongly influenced by anthropogenic 24 emissions (Weber et al., 2007; Hoyle et al., 2011; Shilling et al., 2013). Organic vapors are 25 also implicated in facilitating new particle formation initiated by sulfuric acid (Kulmala et al., 26 2004; Paasonen et al., 2010; Kuang et al., 2012) and are found to play a crucial role in the subsequent growth of the nanoparticles (Smith et al., 2008; Pierce et al., 2011, 2012; Riipinen 27 et al., 2011; Winkler et al., 2012). Thus, the majority of the optically- and CCN-active 28 particles are produced through the growth of smaller particles by condensation of SOA 29 species (Riipinen et al., 2012). It is therefore necessary that climate models be able to 30

accurately simulate not just the total organic mass loading, but also the evolution of aerosol
 number and composition size distributions resulting from SOA formation.

It is broadly understood that, in cloud-free air, SOA forms via three possible mechanisms: 1) 3 4 effectively irreversible condensation of very low volatility organic vapors produced by gasphase oxidation (Donahue et al., 2011; Pierce et al., 2011); 2) volume-controlled reversible 5 6 absorption of semi-volatile organic vapors into pre-existing particle organic phase according to Raoult's law (Pankow, 1994) or into preexisting particle aqueous phase according to 7 8 Henry's law (Carlton and Turpin, 2013); and 3) absorption of semi-volatile and volatile 9 organic vapors into pre-existing aerosol followed by particle-phase reactions to form effectively non-volatile products such as organic salts (Smith et al., 2010), oligomers, organic 10 acids, and other high molecular weight oxidation products (Gao et al., 2004; Kalberer et al., 11 2004; Heaton et al., 2007; Nozière et al., 2007; Ervens et al., 2010; Wang et al., 2010; Hall 12 and Johnston, 2011; Liu et al., 2012), hemiacetals (Kroll et al., 2008; Ziemann et al., 2012; 13 14 Shiraiwa et al., 2013), and organosulfates (Surratt et al., 2007; Zaveri et al., 2010). Recently, 15 Liu et al. (2014) presented an exact analytical solution to the diffusion-reaction problem in the 16 aqueous phase. While aqueous-phase chemistry in cloud droplets is also a potential source of SOA (Carlton et al., 2008; Ervens et al., 2008; Mouchel-Vallon et al., 2013), this route is not 17 considered in the present study. Several recent studies also indicate that the phase state of 18 19 SOA may be viscous semi-solids under dry and moderate relative humidity conditions (Virtanen et al., 2010; Vaden et al., 2011; Saukko et al., 2012), with very low particle-phase 20 bulk diffusivities (Abramson et al., 2013; Renbaum-Wolff et al., 2013). The timescales of 21 22 SOA partitioning (Shiraiwa and Seinfeld, 2012b) and the resulting aerosol size distributions from these three mechanisms can be quite different, and the particle phase state is expected to 23 24 modulate the growth dynamics as well.

25 Ripinen et al. (2011) analyzed the evolution of ambient aerosol size distributions with a 26 simplified model consisting of mechanisms #1 and #2 for liquid particles and concluded that both mechanisms were roughly equally needed to explain the observed aerosol growth. 27 28 Perraud et al. (2012) studied the gas-particle partitioning of organic nitrate vapors formed 29 from simultaneous oxidation of a-pinene by O₃ and NO₃ in a flow tube reactor. Their model 30 analysis suggested that, despite being semi-volatile, the organic nitrate species had effectively irreversibly condensed (mechanism #1) as their adsorbed layers were continuously "buried" 31 32 in presumably semi-solid particles by other incoming organic vapors. In a theoretical study,

1 Zhang et al. (2012) contrasted the aerosol size distributions produced by mechanisms #1 and 2 #2 for liquid particles and illustrated the roles of solute volatility and vapor source rate in 3 shaping the size distribution via mechanism #2. In another theoretical study, Shiraiwa and 4 Seinfeld (2012b) used the detailed multilayer kinetic flux model KM-GAP (Shiraiwa et al., 5 2012a; based on the PRA model framework of Pöschl-Rudich-Ammann, 2007) to investigate the effect of phase state on SOA partitioning. They showed that the timescale for gas-particle 6 7 equilibration via mechanism #2 increases from hours to days for organic aerosol associated 8 with semi-solid particles, low volatility, large particle size, and low mass loadings. More 9 recently, Shiraiwa et al. (2013a) studied SOA formation from photooxidation of dodecane in the presence of dry ammonium sulfate seed particles in an environmental chamber. Their 10 analysis of the observed aerosol size distribution evolution with the KM-GAP model revealed 11 the presence of particle-phase reactions (i.e., mechanism #3), which contributed more than 12 13 half of the SOA mass, with the rest formed via mechanism #2. Furthermore, the physical state of the SOA was assumed to be semi-solid with an average bulk diffusivity of 10⁻¹² cm² s⁻¹, 14 and the particle-phase reactions were predicted to occur mainly on the surface. 15

While valuable insights into the effect of phase state on SOA formation have emerged from 16 several recent studies, a comprehensive, quantitative analysis of the effects of organic solute 17 volatility, phase state, and particle-phase reaction on aerosol growth dynamics has not yet 18 19 been performed. Additionally, there is a lack of a kinetic SOA partitioning treatment for semisolids (with particle-phase chemical reactions) that is amenable for use in regional and global 20 atmospheric models. The present work addresses both these topics. The paper is structured as 21 22 follows. In Sect. 2, we examine the dynamics of diffusion and reaction in a spherical particle with an analytical solution to the problem. In Sect. 3, we extend the MOSAIC (Model for 23 24 Simulating Aerosol Interactions and Chemistry) aerosol model (Zaveri et al., 2008) to include 25 a new framework for kinetic gas-particle partitioning of SOA and evaluate it against a rigorous model based on the finite difference approach. The new framework uses a 26 combination of: (a) an analytical quasi-steady-state treatment for the diffusion-reaction 27 28 process within the particle phase for fast-reacting species, and (b) a two-film theory approach 29 for slow- and non-reacting organic solutes. The framework is amenable for eventual use in regional and global climate models, although it currently awaits specification of the actual 30 31 particle-phase reactions that are important for SOA formation. In Sect. 4, we apply the model 32 to evaluate the timescale of SOA partitioning and the associated evolution of the number and

composition size distributions for a range of solute volatilities, bulk diffusivities, and particle phase reaction rates. We close with a summary of our findings and their implications.

3 2 Dynamics of Diffusion and Reaction within a Particle

Consider an organic solute *i* that diffuses from the gas phase to a single spherical organic 4 5 aerosol particle and reacts irreversibly with a pseudo-first-order rate constant k_c (s⁻¹) as it diffuses inside the particle. This process is illustrated in Fig. 1 using three species (P_1 , P_2 , and 6 7 P_3) for simplicity. The organic solute P_1 diffuses and reacts to form a non-volatile species P_2 8 inside an organic particle (of radius R_p) that is initially composed of a non-volatile organic 9 species P₃. The solute's gas-phase concentrations far away from the particle (i.e., in the bulk gas-phase) and just above the particle surface are \overline{C}_g and \overline{C}_g^s (mol cm⁻³(air)), respectively. 10 The solute's particle-phase concentration just inside the particle surface and at any location in 11 the bulk of the particle are denoted as A^s and A (mol cm⁻³(particle)), respectively. The gas-12 and particle-phase diffusivities of the solute are D_g and D_b (cm² s⁻¹), respectively. 13

14 In this section we shall focus on the dynamics of diffusion and reaction *inside* the particle. In order to derive the timescales relevant to this problem, the particle, initially free of the organic 15 16 solute (i.e., at time t = 0), is assumed to be exposed to a constant concentration just inside the particle surface, A_i^s , at all times t > 0 (this assumption will be relaxed in Sect. 3 where we 17 will relate the temporally changing gas-phase concentration of the solute to its particle-phase 18 19 concentration). Assuming that the diffusive flux of the solute into the particle follows Fick's 20 law, the transient partial differential equation describing the particle-phase concentration $A_i(r,t)$ as a function of radius r and time t can be written as: 21

22
$$\frac{\partial A_i(r,t)}{\partial t} = D_{b,i} \frac{1}{r^2} \frac{\partial}{\partial r} \left(r^2 \frac{\partial A_i(r,t)}{\partial r} \right) - k_{c,i} A_i(r,t) .$$
(1)

The particle is assumed to be spherically symmetrical with respect to the concentration profiles of the organic solute in the particle at any given time, so the concentration gradient at the center of the particle (i.e., r=0) is always zero. These assumptions give rise to the following initial and boundary conditions:

27 I.C.:
$$A_i(r,0) = 0$$
, (2a)

28 B.C. 1:
$$A_i(R_p, t) = A_i^s$$
, (2b)

1 B.C. 2:
$$\frac{\partial A_i(0,t)}{\partial r} = 0$$
. (2c)

Equation (1) with conditions (2) can be analytically solved by first solving the pure diffusion problem in the absence of reaction (Carslaw and Jaeger, 1959; Crank, 1975) and then extending the solution to the case of first-order chemical reaction using the method of Danckwerts (1951) to yield the solution:

 $\frac{A_{i}(r,t)}{A_{i}^{s}} = \frac{R_{p}}{r} \frac{\sinh(q_{i}r/R_{p})}{\sinh(q_{i})} + \frac{2R_{p}}{\pi r} \sum_{n=1}^{\infty} \frac{(-1)^{n} n \sin(n\pi r/R_{p})}{(q_{i}/\pi)^{2} + n^{2}} \exp\left\{-\left(k_{c,i} + \frac{n^{2}\pi^{2}D_{b,i}}{R_{p}^{2}}\right)t\right\},$ (3)

7 where q_i is a dimensionless diffusion-reaction parameter defined as the ratio of the particle 8 radius R_p to the so-called reacto-diffusive length $\sqrt{D_{b,i}/k_{c,i}}$ (Pöschl et al., 2007):

9
$$q_i = R_p \sqrt{\frac{k_{c,i}}{D_{b,i}}}$$
 (4)

It should be noted that this solution assumes that R_p remains constant with time, so diffusion of additional material into the particle is relatively small (this assumption will also be relaxed in Sect. 3). It is also worth noting here that in glassy particles, the diffusion fronts of plasticzing agents (such as water) may move linearly inward, leading to a linear dependence on R_p instead of R_p^2 in Fickian diffusion (Zobrist et al. 2011).

Now, the timescale for Fickian diffusion of the dissolved solute *i* in the particle, τ_{da} , and the timescale for chemical reaction, τ_c (Seinfeld and Pandis, 2006) are defined as:

17
$$au_{da,i} = \frac{R_p^2}{\pi^2 D_{b,i}},$$
 (5)

18
$$au_{c,i} = \frac{1}{k_{c,i}}$$
 (6)

19 The model described by these equations has been applied to investigate mass transfer 20 limitation to the rate of SO_2 oxidation in cloud droplets (Schwartz and Freiberg, 1981; Shi and 21 Seinfeld, 1991), for which the droplets typically exceed 10 μ m diameter, with the aqueousphase diffusivity about 10^{-5} cm² s⁻¹. Here we apply this model to analyze the effects of particle-phase reactions in organic particles of sizes ranging from ~ 10^{-3} to 1 µm diameter, with D_b values ranging from $<10^{-18}$ to 10^{-5} cm⁻² s⁻¹ (Renbaum-Wolff et al., 2013). Since the actual particle-phase reactions of various organic species and the associated rate constants are still not well defined, we use a pseudo-first-order reaction as a proxy and vary its rate constant k_c over several orders of magnitude (10^{-5} to 10^{-1} s⁻¹) to examine its effect on the dynamics of particle growth.

8 The right-hand-side of Eq. (3) comprises two terms. The first term is the concentration profile 9 at steady state with the surface concentration, while the second term describes the temporal 10 evolution of the concentration profile. At steady state, the transient term disappears for $t \gg \tau_{da}$ and τ_c . Fig. 2 illustrates the relative effects of bulk diffusivity and reaction rate 11 12 constant on the temporal evolution of the diffusing solute concentration profiles within a particle of diameter $D_p = 0.1 \ \mu m$. The top row represents a liquid organic particle with a 13 rather high bulk diffusivity, $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$, with (a) no reaction ($k_c = 0$), and (b) a modest 14 reaction rate constant, $k_c = 5 \times 10^{-4} \text{ s}^{-1}$. In case (a), $\tau_{da} = 2.5 \text{ }\mu\text{s}$, and the solute attains a 15 uniform steady-state concentration profile across the particle radius in a little over 8 µs (i.e., 16 about $4\tau_{da}$). The temporal evolution of the concentration profiles in case (b) appears to be 17 identical to case (a) despite the presence of a chemical reaction, because τ_{da} is 2.5 µs but $\tau_c =$ 18 19 2000 s, i.e., diffusion occurs much more rapidly than reaction. In contrast, the bottom row represents a semi-solid organic particle, $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$, with (c) no reaction, and (d) $k_c =$ 20 5 x 10⁻⁴ s⁻¹. In case (c), τ_{da} = 2533 s (i.e., 42 min) and ~ 160 min is required for the solute to 21 attain a uniform steady state profile. In case (d), τ_{da} and τ_{c} are comparable, and as a result the 22 23 solute not only reaches the steady state sooner (in about 60 min) than in the no-reaction case, but also the steady state concentration profile is visibly non-uniform. This is a result of the 24 25 fact that there is sufficient time for appreciable amounts of the solute to be consumed by the 26 reaction as it diffuses towards the center of the particle.

Fig. 3 illustrates the steady state concentration profiles for a range of k_c values (from 10⁻⁵ to 0.1 s⁻¹) in a particle of diameter $D_p = 0.1 \,\mu\text{m}$ with four different D_b values: (a) 10⁻⁶ cm² s⁻¹, (b) 10⁻¹² cm² s⁻¹, (c) 10⁻¹³ cm² s⁻¹, and (d) 10⁻¹⁵ cm² s⁻¹. All together, these cases represent

twenty different combinations of τ_{da} and τ_c . In case (a), $\tau_{da} \ll \tau_c$ for all the k_c values 1 considered here, and as a result the steady state concentration profiles are essentially uniform 2 3 across the entire particle, with the consumption of the solute by chemical reaction occurring 4 uniformly across the entire volume of the particle. In case (b), even though the particle is considered to be a semi-solid with $D_b = 10^{-12} \text{ cm}^{-2} \text{ s}^{-1}$, τ_{da} and τ_c become comparable only 5 when $k_c = 0.1 \text{ s}^{-1}$ (and higher). However, slower reactions produce non-uniform steady state 6 concentration profiles in cases (c) and (d) for D_b values 10^{-13} cm² s⁻¹ and lower. In these 7 cases, most of the solute is consumed near the surface of the particle, with a concentration that 8 9 becomes progressively depleted towards the center of the particle as k_c increases. Thus, the 10 particle growth is volume reaction-controlled when the concentration profile is uniform and 11 tends to be surface reaction-controlled at the other extreme.

Since the timescale for diffusion varies as R_p^2 , the diffusion limitation to reaction also depends strongly on particle size. As shown in Fig. 4, the relative effects of particle size, bulk diffusivity, and reaction rate on the shape of the steady state concentration profiles are concisely captured in terms of the dimensionless parameter q, which is a function of R_p , k_c , and D_b (Eq. (4)). At low values of q (< 0.5), the steady state concentration profile is nearly uniform, but becomes increasingly non-uniform for q values of the order of unity and greater. While the temporal evolution of the radial concentration profile is highly informative, the

19 timescale to reach steady state, as well as the shape of the steady state profile, can be 20 conveniently quantified in terms of the average particle-phase concentration $\overline{A}(t)$. We 21 integrate the concentration profile given by Eq. (3) over the volume of the particle to obtain:

22
$$\frac{\overline{A}_{i}(t)}{A_{i}^{s}} = \frac{\int_{0}^{R_{p}} 4\pi r^{2} \frac{A_{i}(r,t)}{A_{i}^{s}} dr}{\frac{4}{3}\pi R_{p}^{3}} = Q_{i} - U_{i}(t), \qquad (7)$$

23 where

$$24 \qquad Q_i = 3\left(\frac{q_i \coth q_i - 1}{q_i^2}\right),\tag{8}$$

1
$$U_i(t) = \frac{6}{\pi^2} \sum_{n=1}^{\infty} \frac{\exp\left\{-\left(k_{c,i} + \frac{n^2 \pi^2 D_{b,i}}{R_p^2}\right)t\right\}}{(q_i / \pi)^2 + n^2}.$$
 (9)

Here, Q_i is the ratio of the average particle-phase concentration to the surface concentration 2 at steady-state, while $U_i(t)$ is the transient term, the value of which is always equal to Q_i at 3 t=0 and decreases exponentially to zero as $t \rightarrow \infty$. As noted earlier, the surface 4 concentration A_i^s is assumed to be constant in the analytical solution of Eq. (1). However, 5 since A_i^s can gradually change over time due to changes in the gas-phase concentration and 6 particle composition, it is more appropriate to refer to the steady state as quasi-steady state. 7 8 The timescale to reach a quasi-steady state (τ_{oss}) within the particle can then be defined as the e-folding time for the exponential decay of the unsteady state term U_i relative to the 9 quasi-steady state term Q_i . Thus, setting $U_i(\tau_{OSS}) = Q_i / e$, we get: 10

11
$$\sum_{n=1}^{\infty} \frac{\exp\left\{-\left(k_{c,i} + \frac{n^2 \pi^2 D_{b,i}}{R_p^2}\right) \tau_{QSS}\right\}}{\left(q_i / \pi\right)^2 + n^2} = \frac{1}{e} \times \frac{\pi^2}{2} \left(\frac{q_i \coth q_i - 1}{q_i^2}\right).$$
(10)

For a given set of values for D_p , D_b , and k_c , Eq. (10) can be numerically solved for τ_{QSS} with the bisection method.

We first examine the dependence of τ_{OSS} and Q on D_b and k_c for a particle of $D_p = 0.1 \ \mu m$ 14 (Fig. 5). The values of D_b are varied over fourteen orders of magnitude from 10^{-19} cm² s⁻¹ 15 (almost solid) to 10⁻⁵ cm² s⁻¹ (liquid water) to cover the full range of semi-solid and liquid 16 organic particles, and k_c values are varied over six orders of magnitude from of 10^{-6} s⁻¹ (very 17 slow reaction) to 1 s⁻¹ (practically instantaneous reaction). As seen in Fig. 5a, the contours of 18 τ_{oss} range from 1 µs for liquid particles to 1 day for highly viscous semi-solid particles. For 19 20 the semi-solid particles, there are two regions in the semi-solid zone as depicted by the gray dotted line. In the region above the dotted line, τ_{OSS} is sensitive only to the value of k_c and 21 decreases rapidly with increase in k_c . For instance, at $D_b = 10^{-19} \text{ cm}^2 \text{ s}^{-1}$, $\tau_{SS} = -1$ day for k_c 22 = 5 x 10⁻⁶ s⁻¹ but decreases to <1 min for $k_c = 10^{-2}$ s⁻¹. In the region below the dotted line, 23 τ_{OSS} is sensitive only to the value of D_b for both semi-solid and liquid particles. For example, 24

at $D_b = \sim 10^{-14} \text{ cm}^2 \text{ s}^{-1}$, τ_{OSS} remains constant at $\sim 1 \text{ min for } k_c$ values from 10^{-6} up to about 1 10^{-2} s⁻¹ (i.e., up to the dotted line) and only then becomes sensitive to reaction at higher values 2 of k_c . τ_{QSS} is sensitive to both k_c and D_b only in the relatively narrow envelope along the 3 4 dotted line itself. As seen in Fig. 5b, the values of Q are <0.001 for highly viscous semi-solid particles and high k_c values, while they approach unity as D_b increases and k_c decreases. 5 6 Note that the dotted line in Fig. 5a roughly corresponds to the contour for Q = 0.6 in Fig. 5b. Next, we examine the dependence of τ_{OSS} and Q on particle size. Fig. 6 shows τ_{OSS} vs. D_p 7 for D_b values ranging from 10^{-18} to 10^{-10} cm² s⁻¹ for (a) $k_c = 0$ s⁻¹, (b) $k_c = 10^{-3}$ s⁻¹, (c) k_c 8 0.01 s⁻¹, and (d) $k_c = 0.1$ s⁻¹. As seen in Fig. 6a, for any given D_b , τ_{QSS} increases by five 9 orders of magnitude as D_p increases from 0.003 to 1 µm. At the upper end, particles with D_b 10 $< 10^{-18}$ cm² s⁻¹ have τ_{QSS} about 10 min at $D_p = 0.003$ µm and increase to more than 10^4 min 11 at $D_p = 0.1 \ \mu\text{m}$. In contrast, particles with $D_b > 10^{-12} \ \text{cm}^2 \ \text{s}^{-1}$ have τ_{OSS} below 1 min 12 (indicated by the dotted gray line) for sizes up to 0.7 µm. From a practical standpoint, since 13 14 most ambient SOA particles are smaller than ~0.7 µm, concentration profiles of non-reacting solutes inside particles with $D_b > 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ may be assumed to be at steady-state. 15 However, significant diffusion limitation can exist for non-reacting solutes in particles with 16 $D_b < 10^{-12} {
m cm}^2 {
m s}^{-1}$ depending on their size. In stark contrast, for reacting solutes, au_{OSS} 17 asymptotically approaches a common maximum value for all values of D_b as the particle size 18 increases as shown in Fig. 6b,c,d. This maximum value of τ_{OSS} is about 7, 0.7, and 0.07 min 19 for $k_c = 10^{-3}$, 10^{-2} , and 0.1 s^{-1} , respectively. The typical timescale for changes in the bulk gas-20 phase concentrations due to transport and chemical reaction is of the order 10 min or more. 21 22 Thus, from a practical standpoint, the particle-phase concentration profiles of solutes reacting with $k_c > 10^{-2} \text{ s}^{-1}$ (for which $\tau_{QSS} \le 0.7 \text{ min}$) may be assumed to be at quasi-steady state in 23 particles of any size and any D_{h} value. 24

Fig. 7 illustrates variation of Q with D_p for the four cases shown in Fig. 6. At quasi-steady state, the particle-phase concentration profile for non-reacting solutes is always uniform (i.e., Q = 1) even though τ_{QSS} can differ significantly depending on the particle size and D_b value (Fig. 7a). For reacting solutes with k_c up to 0.1 s⁻¹, Q remains nearly equal to unity in 1 particles with $D_b > 10^{-10} \text{ cm}^2 \text{ s}^{-1}$ and D_p up to 1 µm. For $D_b < 10^{-10} \text{ cm}^2 \text{ s}^{-1}$, Q decreases as 2 D_p increases for a given D_b , while it increases as D_b increases for a given D_p .

In general, the above analysis indicates that: a) for a given D_p , a more reactive solute will reach quasi-steady state sooner and exhibit a more non-uniform concentration profile than a less reactive one, especially in particles with lower D_b than higher, and b) for a given set of values for k_c and D_b , a solute in smaller particles will reach quasi-steady state sooner and exhibit a more uniform quasi-steady state concentration profile than in larger particles.

8

9 **3** Kinetic Gas-Particle Partitioning Model

We shall now describe the development of a new framework for modeling kinetic partitioning 10 of SOA based on the insights gained from timescale analysis of the diffusion-reaction process 11 12 within the particle phase. The framework takes into account solute volatility, gas-phase 13 diffusion, interfacial mass accommodation, particle-phase diffusion, and particle-phase 14 reaction. However, instead of numerically resolving the concentration gradient inside the 15 particle (Shiraiwa et al., 2012a), which is computationally expensive and therefore impractical 16 for inclusion in 3-D Eulerian models, we use the analytical expressions of the quasi-steady state and transient behavior of the solute diffusing and reacting within the particle. 17

18 **3.1 Model Framework**

19 3.1.1 Single Particle Equations

We begin by relating the average particle-phase concentration of the solute \overline{A}_i (mol cm⁻³(particle)) to its average bulk gas-phase concentration $\overline{C}_{g,i}$ (mol cm⁻³(air)) over a single particle. Similar to the timescale for diffusion in the particle-phase (Eq. (5)), the timescale for the gas-phase concentration gradient outside the particle to reach a quasi-steady state (τ_{dg}) is given by (Seinfeld and Pandis, 2006):

25
$$au_{dg,i} = \frac{R_p^2}{\pi^2 D_{g,i}},$$
 (11)

1 where $D_{g,i}$ (cm² s⁻¹) is the gas-phase diffusivity. For a typical $D_{g,i}$ of 0.05 cm² s⁻¹, the value 2 of τ_{dg} is of the order 10⁻⁸ s or less for submicron size aerosols, which is much smaller than 3 the typical timescale for changes in the bulk gas-phase concentration in the ambient 4 atmosphere. We can therefore safely assume that the gas-phase concentration profile of the 5 solute around the particle is at quasi-steady state at any instant.

6 An ordinary differential equation describing the rate of change of \overline{A}_i due to mass transfer 7 between gas and a single particle with particle-phase reaction can then be written as:

8
$$\frac{dA_i}{dt} = \frac{3}{R_p} k_{g,i} \left(\overline{C}_{g,i} - C^s_{g,i} \right) - k_{c,i} \overline{A}_i,$$
 (12)

9 where $C_{g,i}^{s}$ (mol cm⁻³(air)) is the gas-phase concentration of the solute just outside the surface 10 of the particle, and $k_{g,i}$ (cm s⁻¹) is the gas-side mass transfer coefficient given as:

11
$$k_{g,i} = \frac{D_{g,i}}{R_p} f(Kn_i, \alpha_i).$$
 (13)

Here $f(Kn_i, \alpha_i)$ is the transition regime correction factor (Fuchs and Sutugin, 1971) to the Maxwellian flux as a function of the Knudsen number $Kn_i = \lambda_i / R_p$ (where λ_i is the mean free path) and the so-called mass accommodation coefficient, α_i , which is defined as the fraction ($0 \le \alpha_i \le 1$) of the incoming molecules that is incorporated into the particle surface:

16
$$f(Kn_i, \alpha_i) = \frac{0.75\alpha_i(1 + Kn_i)}{Kn_i(1 + Kn_i) + 0.283\alpha_i Kn_i + 0.75\alpha_i}.$$
 (14)

While the above correction factor was derived from a numerical solution of the Boltzmann diffusion equation for neutron transfer to a black sphere (i.e., representative of light molecules in a heavy background gas), its applicability for higher molecular weight trace gases in air has been experimentally confirmed (Seinfeld and Pandis, 2006, and references therein).

The timescale to achieve interfacial phase equilibrium between $C_{g,i}^{s}$ and the particle-phase concentration of *i* just inside the surface, A_{i}^{s} (mol cm⁻³(particle)) is at least (Seinfeld and Pandis, 2006):

where \overline{v}_i is the average speed of solute molecules in the gas phase. From kinetic theory of gases $\overline{v}_i = (8\Re T / \pi M_i)^{1/2}$ where \Re is the universal gas constant (8.314 X 10⁷ erg K⁻¹ mol⁻¹), T (K) is temperature, and M_i is the molecular weight of the solute. For representative values of $D_{b,i} \leq 10^{-5}$ cm² s⁻¹, $M_i = 100$ g mol⁻¹, T = 298 K, and α_i ranging from 0.1 to 0.001, the value of $\tau_{p,i}$ is of the order 10⁻⁶ s or less, which means it can be safely assumed that the interfacial phase equilibrium is achieved virtually instantaneously. We thus relate $C_{g,i}^s$ and A_i^s according to Raoult's law as:

9
$$C_{g,i}^{s} = \frac{A_{i}^{s}}{\sum_{j} A_{j}^{s}} C_{g,i}^{*}$$
 (16)

10 where $C_{g,i}^*$ is the effective saturation vapor concentration (mol cm⁻³(air)), and $\sum_j A_j^s$ is the 11 total particle-phase concentration of all the organic species at the surface. However, since the 12 surface concentrations of all the species are not always known, we use the total average 13 particle-phase concentration $\sum_j \overline{A}_j$ as an approximation for $\sum_j A_j^s$. Thus Eq. (12) is 14 rewritten in terms of A_i^s as:

15
$$\frac{d\overline{A}_{i}}{dt} = \frac{3}{R_{p}} k_{g,i} \left(\overline{C}_{g,i} - \frac{A_{i}^{s}}{\sum_{j} \overline{A}_{j}} C_{g,i}^{*} \right) - k_{c,i} \overline{A}_{i}.$$
(17)

16 A_i^s can be assumed to be equal to \overline{A}_i in liquid particles for a non-reactive or slowly reacting 17 solute that quickly attains a uniform concentration profile (as was previously shown in Fig. 18 2a,b). But, as discussed in the previous section, this equality may not hold for reactive and 19 non-reactive solutes in semi-solid particles. In such cases, Eq. (7) can be used to express A_i^s 20 in terms of \overline{A}_i as long as A_i^s does not change with time, because the analytical solution to Eq. 21 (1) assumes a constant A_i^s according to the boundary condition (2b). In practice, however, Eq. 1 (7) can be used if the timescale for changes in A_i^s is much greater than the timescale for the 2 solute to relax to its quasi-steady state profile inside the particle. With this caveat, we get:

$$3 \qquad \frac{d\overline{A}_{i}}{dt} = \frac{3}{R_{p}} k_{g,i} \left\{ \overline{C}_{g,i} - \frac{\overline{A}_{i}}{\sum_{j} \overline{A}_{j}} \frac{C_{g,i}^{*}}{(Q_{i} - U_{i}(t))} \right\} - k_{c,i} \overline{A}_{i}.$$

$$(18)$$

Note that Eq. (18) describes kinetic mass transfer of species *i* between bulk gas-phase and a 4 5 single particle, with chemical reaction within the particle phase, and includes mass-transfer 6 limitations due to gas-phase diffusion, interfacial mass accommodation, and particle-phase 7 diffusion. Previously, the mass accommodation coefficient (α) has been often used as a 8 tunable parameter to fit the observed kinetic limitation to mass-transfer during SOA 9 partitioning (Bowman et al., 1997; Saathoff et al., 2009; Parikh et al., 2011). However, α 10 does not correctly capture the mass transfer limitations due to diffusion and chemical 11 reaction occurring within the bulk of the particle. In the present framework, the interfacial 12 and bulk particle phase limitations to mass transfer are represented separately, with the 13 appropriate dependence for the latter on particle size.

In Eq. (18), the term $U_i(t)$ is to be evaluated at the "time since start." Equation (18) can 14 15 therefore only be used in a Lagrangian box model framework for a "closed system" where we 16 can specify an initial concentration of the solute vapor (at time t = 0), which then partitions to the particle-phase as a function of time. The solute vapor in the closed system is not subjected 17 to emissions, dilution, and loss due to gas-phase oxidation. In the case of no particle-phase 18 19 reaction, the solute vapor will eventually reach equilibrium with the particles. In the presence 20 of particle-phase reaction, the solute vapor concentration will eventually decay to zero. This is in stark contrast with the "general system" such as the ambient atmosphere and 3-D 21 22 atmospheric chemistry transport models where the solute vapor at a given location may 23 continuously change due to emissions, dilution, and gas-phase chemistry in addition to gas-24 particle partitioning. As a result, it is not possible to evaluate $U_i(t)$ in the general system, 25 because we cannot keep track of the "time since start" in the same sense as used in the transient analytical solution to Eq. (1). Therefore, based on the value of k_c and the associated 26 timescale for the particle-phase concentration profile to reach quasi steady state (τ_{QSS}), the 27 28 following two approximations to Eq. (18) are made for it to be applicable to the general 29 system.

1 Approximation 1: For Fast Reactions $(k_{c,i} \ge 0.01 \text{ s}^{-1})$

As discussed in the previous section (Fig. 6c), τ_{QSS} for a solute reacting with $k_{c,i} \ge 0.01 \text{ s}^{-1}$ is less than 1 min in particles with any D_b and of any size. Compared to the typical time step values of 5 min or greater in 3-D Eulerian models, the particle-phase concentration profile for solutes with $\tau_{QSS} \le 1$ min may be assumed to be at quasi-steady-state, and the term $U_i(t)$ can be safely neglected in Eq. (18) to yield:

7
$$\frac{d\bar{A}_{i}}{dt} = \frac{3}{R_{p}} k_{g,i} \left\{ \bar{C}_{g,i} - \frac{\bar{A}_{i}}{\sum_{j} \bar{A}_{j}} \frac{C_{g,i}^{*}}{Q_{i}} \right\} - k_{c,i} \bar{A}_{i} \quad \text{for } k_{c,i} \ge 0.01 \text{ s}^{-1}.$$
(19)

A similar equation was derived by Shi and Seinfeld (1991) for reactive mass transport of SO₂ (with Henry's law for absorption) in cloud droplets assuming quasi-steady state within the droplet phase. Now, as $k_c \rightarrow 0$, $Q \rightarrow 1$, and mass transfer is governed entirely by gas-phase diffusion and interfacial mass accommodation in Eq. (19). As a result, Eq. (19) tends to lose its ability to capture the resistance to mass transfer due to slow diffusion in the particle phase as $k_c \rightarrow 0$. Therefore, an alternate treatment for mass transfer is needed for slow reactions.

14 Approximation 2: For Slow Reactions ($k_{c,i} < 0.01 \text{ s}^{-1}$)

For $k_{c,i} < 0.01 \text{ s}^{-1}$ (or $\tau_{QSS} > 1 \text{ min}$), we use the classical two-film theory of mass transfer 15 between the gas and particle phases. The two-film theory was originally introduced by Lewis 16 17 and Whitman (1924) and has been widely used to model mass transfer in two phase systems, with and without chemical reactions (Astarita, 1967; Doraiswamy and Sharma, 1984; Bird et 18 19 al., 2007). Fig. 8 shows the schematic of the two-film model which assumes that the 20 concentration gradients in the gas and particle phases are confined in the respective 21 hypothetical "films" adjacent to the interface. The gas- and particle-side film thicknesses are denoted by δ_{g} and δ_{p} (cm), respectively, and the respective mass transfer coefficients (cm s⁻ 22 ¹) are defined as $k_g = D_g / \delta_g$ and $k_p = D_b / \delta_p$. The overall gas-side mass transfer coefficient 23 K_g (cm s⁻¹) is then given by (see Appendix A for the derivation): 24

25
$$\frac{1}{K_{g,i}} = \frac{1}{k_{g,i}} + \frac{1}{k_{p,i}} \left(\frac{C_{g,i}^*}{\sum_j \overline{A}_j} \right).$$
 (20)

1 The ordinary differential equation describing the rate of change of \overline{A}_i due to gas-particle mass 2 transfer and particle-phase reaction can then be written in terms of the overall driving force 3 as:

$$4 \qquad \frac{d\bar{A}_{i}}{dt} = \frac{3}{R_{p}} K_{g,i} \left\{ \bar{C}_{g,i} - \frac{\bar{A}_{i}}{\sum_{j} \bar{A}_{j}} C_{g,i}^{*} \right\} - k_{c,i} \bar{A}_{i} \quad \text{for } k_{c,i} < 0.01 \text{ s}^{-1}.$$
(21)

5 A similar equation was derived by Zaveri (1997) for reactive mass transport of SO₂ (with Henry's law for absorption) in cloud droplets assuming quasi-steady state within the droplet 6 7 phase. The advantage of the two-film model formulation is that the diffusion limitations from 8 both the gas and particle sides are represented in the overall mass transfer coefficient, and can 9 therefore be used to model mass transfer of slow-reacting solutes. The gas-side mass transfer coefficient (k_g) is already known from Eq. (13) where $\delta_g = R_p$. However, the particle-side 10 film thickness, δ_p , and therefore k_p , are not readily known. In a general system, the bulk gas-11 and particle-phase concentrations of a reactive semi-volatile solute tend to reach a quasi-12 13 steady state when the net source rate of the solute in the gas phase is relatively steady. Since both Eqs. (19) and (21) describe the same process, they should predict identical gas- and 14 particle-phase concentrations at quasi-steady-state. Thus, setting $d\overline{A}_i / dt = 0$ in both Eqs. (19) 15 and (21) and equating the expressions for $(\overline{A}_i / \overline{C}_{g,i})$ resulting from each of them yields the 16 general expressions for δ_p and k_p in terms of D_b , k_c , and R_p (see Appendix B for the 17 18 derivation):

19
$$\delta_{p,i} = R_p \left(\frac{1 - Q_i}{q_i \coth q_i - 1} \right), \tag{22}$$

20
$$k_{p,i} = \frac{D_{b,i}}{R_p} \left(\frac{q_i \coth q_i - 1}{1 - Q_i} \right).$$
 (23)

21 For the limiting case of a non-reactive solute, $k_c \rightarrow 0$, $q \rightarrow 0$, $Q \rightarrow 1$ and Eq. (23) reduces to:

22
$$k_{p,i} = 5 \frac{D_{b,i}}{R_p}$$
. (24)

1 3.1.2 Polydisperse Aerosol Equations

We now extend the <u>closed system</u> box model Eq. (18) for a single particle to a polydisperse aerosol in a sectional framework. For a given size section *m*, with number concentration N_m (cm⁻³(air)) and particle radius $R_{p,m}$ (cm), we define $\overline{C}_{a,i,m}$ (mol cm⁻³(air)) as the total average concentration of solute *i* in size section *m*:

6
$$\bar{C}_{a,i,m} = \frac{4}{3} \pi R_{p,m}^3 N_m \bar{A}_{i,m}$$
 (25)

7 Multiplying Eq. (18) by $(4\pi R_{p,m}^3 N_m / 3)$ gives:

$$8 \qquad \frac{d\bar{C}_{a,i,m}}{dt} = 4\pi R_{p,m}^2 N_m k_{g,i,m} \left\{ \bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{(Q_i - U_i(t))} \right\} - k_{c,i} \bar{C}_{a,i,m},$$
(26)

9 where $S_{i,m}$ is the saturation ratio:

10
$$S_{i,m} = \frac{C_{g,i}^*}{\sum_j \overline{C}_{a,j,m}}.$$
 (27)

11 The corresponding equation governing the gas-phase concentration of solute *i* is:

12
$$\frac{d\bar{C}_{g,i}}{dt} = -\sum_{m} \left[4\pi R_{p,m}^2 N_m k_{g,i,m} \left\{ \bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{(Q_i - U_i(t))} \right\} \right].$$
 (28)

Similarly, the particle-phase and gas-phase equations for polydisperse aerosols in the general
 system are as follows.

15 Approximation 1: For $k_{c,i} \ge 0.01 \text{ s}^{-1}$

16
$$\frac{d\overline{C}_{a,i,m}}{dt} = 4\pi R_{p,m}^2 N_m k_{g,i,m} \left(\overline{C}_{g,i} - \overline{C}_{a,i,m} \frac{S_{i,m}}{Q_i}\right) - k_{c,i} \overline{C}_{a,i,m}, \qquad (29)$$

17
$$\frac{d\bar{C}_{g,i}}{dt} = -\sum_{m} \left[4\pi R_{p,m}^2 N_m k_{g,i,m} \left(\bar{C}_{g,i} - \bar{C}_{a,i,m} \frac{S_{i,m}}{Q_i} \right) \right].$$
(30)

18 **Approximation 2: For** $k_{c,i} < 0.01 \text{ s}^{-1}$

19
$$\frac{dC_{a,i,m}}{dt} = 4\pi R_{p,m}^2 N_m K_{g,i,m} \left(\overline{C}_{g,i} - \overline{C}_{a,i,m} S_{i,m} \right) - k_{c,i} \overline{C}_{a,i,m}, \qquad (31)$$

$$1 \qquad \frac{d\bar{C}_{g,i}}{dt} = -\sum_{m} \left[4\pi R_{p,m}^2 N_m K_{g,i,m} \left(\bar{C}_{g,i} - \bar{C}_{a,i,m} S_{i,m} \right) \right].$$
(32)

The proposed framework, described by Eqs. (29) through (32), is relatively simple and amenable for use in regional and global aerosol models, although it presently awaits specification of the actual particle-phase chemical reactions that are important for SOA formation.

6 We have implemented both the closed system and general system frameworks in the 7 computationally efficient, multicomponent aerosol box-model MOSAIC and adapted the 8 existing semi-implicit Euler method solver to numerically integrate the set of coupled 9 ordinary differential equations for any number of solutes *i* over any number of size bins *m* 10 (Zaveri et al., 2008). Sectional growth in MOSAIC is calculated using the two-moment 11 approach of Simmel and Wurzler (2006). The closed system framework is to be used in the 12 box-model version only while the general system framework can be used in both box- and 3-13 D Eulerian models. The complete solution to these equations may be labeled as "semi-14 numerical," because the particle-phase diffusion-reaction process is represented analytically 15 while the set of ordinary differential equations themselves are integrated numerically.

16 MOSAIC already performs kinetic partitioning of inorganic gases (H₂SO₄, HNO₃, HCl, and 17 NH₃) to size-distributed particles and predicts liquid water associated with inorganic species 18 as a function of relative humidity. While the focus of the present work is on kinetic 19 partitioning of organic gases to particulate organic phase, the new framework can be readily 20 adapted to kinetically partition water soluble organic gases into the particulate aqueous phase 21 if that is the only liquid phase in the particle. However, additional research is needed to 22 extend the present framework to mixed inorganic-organic particles that experience liquid-23 liquid phase separation (i.e., coexistence of separate aqueous and organic phases; You et al., 24 2012).

25 3.2 Model Validation

We shall now validate the new framework in MOSAIC against a "fully numerical" finitedifference solution to Eq. (1) with a flux type boundary condition that includes mass transfer of the solute between the gas phase and the particle surface. The volume of the spherical particle is resolved with multiple layers, and diffusion and reaction of the solute species through these layers are integrated numerically. We used 300 uniformly-spaced layers in the present exercise. The finite difference model is conceptually similar to the KM-GAP model (Shiraiwa et al., 2012a), but does not include reversible adsorption at the surface and heat transfer processes. The finite-difference solution is used as a benchmark here because it rigorously solves Eq. (1) and does not assume the surface concentration to remain constant with time.

6 For validation purposes, we consider a monodisperse semi-solid aerosol composed of nonvolatile organic species P_3 (molecular weight 100 g mol⁻¹ and density 1 g cm⁻³), with initial 7 particle diameter $D_p = 0.2 \,\mu\text{m}$, particle number concentration $N = 5000 \,\text{cm}^{-3}$, and bulk 8 diffusivity $D_b = 10^{-15}$ cm² s⁻¹. For simplicity, the molecular weight and density of the 9 condensing solute (P_1) and its reaction product species (P_2) are also assumed to be 100 g mol⁻¹ 10 and 1 g cm⁻³, respectively. The three species (P_1 , P_2 , and P_3) are assumed to form an ideal 11 solution that participates in the absorption of P₁ according to Raoult's law. Model validation 12 13 is demonstrated below for both closed and general systems.

14 3.2.1 Closed System

15 In three separate closed system cases, the initial monodisperse aerosol was exposed to the solute (P₁) gas concentration of 2 μ g m⁻³ with volatility $C_g^* = 10$, 100, and 1000 μ g m⁻³. Fig. 9 16 compares the solution given by MOSAIC (Eqs. (26) and (28)) with the finite-difference 17 model solution for gas-phase concentration decay due to kinetic gas-particle partitioning for 18 particle-phase reaction rate constants k_c ranging from 0 to 0.1 s⁻¹. When $k_c = 0$, the gas-phase 19 concentration reaches an equilibrium value that depends on the solute volatility, while in other 20 cases it decays to zero at different rates as governed by the particle-phase reaction rate 21 constant and diffusion limitation. MOSAIC is able to reproduce the finite difference results 22 23 quite well, although small deviations can be seen during the initial portions of the gas decay for $k_c \le 10^{-4}$ s⁻¹ and $C_g^* = 10$ and 100 µg m⁻³. The following metrics were used to quantify the 24 accuracy of MOSAIC relative to the finite difference (FD) model: 25

26 Mean Normalized Bias, MNB =
$$\overline{\left(C_{g,1}^{\text{MOSAIC}} - C_{g,1}^{\text{FD}}\right)/C_{g,1}^{\text{FD}}}$$
, (33)

27 Mean Normalized Gross Error, MNGE =
$$\left| \overline{C_{g,1}^{\text{MOSAIC}} - C_{g,1}^{\text{FD}}} \right| / \overline{C_{g,1}^{\text{FD}}}$$
, (34)

28 Maximum Normalized Gross Error, maxNGE = max $\left(\left| C_{g,1}^{\text{MOSAIC}} - C_{g,1}^{\text{FD}} \right| / C_{g,1}^{\text{FD}} \right).$ (35)

1 These metrics were calculated using the model outputs at 5 min intervals for the 10 h long simulations. However, negligibly small gas-phase concentrations ($< 0.05 \ \mu g \ m^{-3}$) towards the 2 latter part of the simulations (where applicable) were excluded in the calculations of the 3 4 metrics. The results are displayed in Table 1. The MNB and MNGE are comparable in magnitude and range from ~0.1% to ~10%, with values greater than ~5% seen only for C_g^* = 5 10 µg m⁻³. The large maxNGE values (>20%) seen for $C_g^* = 10 \mu \text{g m}^{-3}$ occur as the gas-phase 6 concentrations approach zero. Overall, the agreement between the two models is quite good 7 8 for the closed system.

9 3.2.2 General System

10 In three separate general system cases, the initial monodisperse aerosol was exposed to solute P₁ with $C_g^* = 10$, 100, and 1000 µg m⁻³ at a constant gas-phase source rate of $\gamma = 0.1$ µg m⁻³ 11 h^{-1} in each case. The initial gas-phase concentration of P₁ was zero in each case. Fig. 10 12 compares the evolution of the gas-phase concentration of P₁ predicted by MOSAIC (Eqs. 13 (29)-(32)) and the finite-difference model. The particle-phase reaction rate constant k_c ranged 14 from 0 to 0.1 s⁻¹. When $k_c = 0$, the gas-phase concentration of P₁ increases almost linearly 15 16 with time upon reaching quasi-equilibrium with the particle phase. For $k_c > 0$, the gas-phase 17 concentration of P₁ remains constant after the initial build up as the source rate is balanced by 18 the loss rate due to particle-phase diffusion and reaction. This quasi-steady state gas-phase concentration level depends on the combination of C_g^* , D_b , and k_c . For $C_g^* = 10 \ \mu g \ m^{-3}$, the 19 time required to establish quasi-steady state between gas and particle phases ranges from less 20 than 1 h at $k_c = 0.1 \text{ s}^{-1}$ to more than 20 h at $k_c = 10^{-4} \text{ s}^{-1}$. The time to reach quasi-equilibrium 21 (for non-reactive solutes) and quasi-steady state (for reactive solutes) increases as the value of 22 C_g^* increases. Approximations 1 and 2 in MOSAIC are able to capture both the initial "spin-23 up" phase, when the gas-phase concentration builds up, as well as the later phase where the 24 concentration remains in quasi-equilibrium or quasi-steady state. Furthermore, for $k_c = 10^{-3} \text{ s}^{-1}$ 25 ¹, Approximation 1 (black dotted line in Fig. 10) gives nearly identical results as 26 Approximation 2 for all three C_g^* values, indicating that the transition from one to the other 27 does not cause a sudden change in the behavior of the solution. Approximation 1 predicts 28 29 faster gas uptake than the finite difference model for slow reactions while Approximation 2

predicts slower gas uptake than the finite difference model for fast reactions (not shown), especially for low volatility solutes ($C_g^* = \sim 10 \ \mu g \ m^{-3}$). A combination of Approximations 1 and 2 is thus needed to cover the full range of possible k_c values.

The normalized gross errors in MOSAIC are relatively large during the spin-up phase where the gas-phase concentrations are very small. In a 3-D Eulerian model application, the spin-up phase occurs at the beginning of the simulation and is usually discarded. Here, we discard the first two hours of spin-up in each simulation to avoid small gas-phase concentrations when calculating the bias and error metrics, shown in Table 2. Both MNB and MNGE are generally less than ~3%. The maxNGE values ranged between 0.3 and 8.5%. The overall performance of MOSAIC for the general system is excellent.

11 **3.3 Future Considerations**

While the general system framework is amenable for eventual use in regional and global climate models, it currently awaits specification of the various gas and particle-phase chemistries important for SOA formation. The following issues must be must be taken into consideration when specifying the various physical and chemical details in the model and evaluating it using laboratory and field observations.

17 First, the present framework uses a pseudo-first order (PFO) reaction for a condensing solute as a proxy for second order chemical reactions that may occur within a particle. The 18 assumption of PFO reaction for the condensing solute is valid when the pre-existing bulk 19 20 reactant species is uniformly distributed with the depth of the particle, e.g., when the reaction 21 timescale for the reactant species is much longer than that for diffusion. The issue arises when the reaction time scale is much shorter than that for diffusion such that the bulk reactant 22 species is not homogeneously distributed depth-wise (Berkemeier et al., 2013). In such cases, 23 it may be possible to parameterize the PFO reaction rate constant for the condensing solute in 24 25 terms of its second order rate constant multiplied by the volume average concentration of the pre-existing reactant solutes in the particle phase. The detailed finite difference model using 26 27 second order reactions can be used to provide guidance for improving and validating the parameterized reactions in the semi-numerical framework. 28

Second, while the present framework allows particles of different sizes and composition to have different bulk diffusivities, it cannot explicitly treat the potential variation of diffusivity within a given particle of complex morphology. Examples include black carbon or solid ammonium sulfate particles coated with organics as well as particles with non-ideal internal mixtures of hydrophobic and hydrophilic organics. The diffusion-reaction process inside such complex and potentially non-spherical particles will again have to be parameterized based on the average bulk properties, with possible guidance from more detailed finite difference models where applicable.

7 Third, as mentioned earlier, the new framework can be readily adapted to kinetically partition 8 water soluble organic gases into the particulate aqueous phase if that is the only liquid phase 9 in the particle. However, additional work is needed to extend the present framework to mixed 10 inorganic-organic particles in which water and organics may form separate liquid phases (You 11 et al., 2012).

12

13 4 Results and Discussion

14 We now apply the updated MOSAIC model to a series of polydisperse aerosol scenarios to investigate the influence of particle-phase reactions, phase state, and solute volatility on SOA 15 partitioning timescale and the evolution of aerosol size distribution. While the exact 16 mechanism(s) responsible for the growth of newly formed particles (1 to 10 nm range) is still 17 18 unknown, it is suspected to occur via effectively irreversible condensation of very low 19 volatility organic species which can overcome the strong Kelvin effect (Pierce et al., 2011). 20 In the present study, we focus on the competitive growth dynamics of the Aitken and 21 accumulation mode particles, as might result after the newly formed particles have grown up 22 to Aitken mode sizes. The Kelvin effect and coagulation are neglected for simplicity. Fig. 11 shows the initial aerosol number and volume size distributions used for this exercise. Again, 23 this pre-existing aerosol is assumed to be composed of non-volatile organic species (P_3) of 24 molecular weight 100 g mol⁻¹ and density 1 g cm⁻³. The entire size distribution, consisting of 25 an Aitken mode and an accumulation mode, is discretized over 1000 logarithmically-spaced 26 27 size bins (lower boundary of the smallest bin = $0.008 \ \mu m$ and the upper boundary of the largest bin = 1 μ m). The total number concentration of particles in the Aitken mode is 6223 28 cm⁻³ while that in the accumulation mode is 1139 cm⁻³; the total aerosol mass concentration is 29 2 µg m⁻³. Fig. 11 also shows the condensational sink $k_{CS,i,m} = 4\pi R_{p,m}^2 N_m k_{g,i,m}$ for each size bin 30 m as a function of D_p . For this particular size distribution, the sum of k_{CS} over all the size 31

bins in the Aitken mode is equal to that in the accumulation mode, so that there is no initial bias in the condensation rate of the solute species towards either mode merely due to differences in the initial condensational sink rates for the two modes. Both closed and general systems scenarios are examined.

5 4.1 Closed System

6 A set of closed system simulations was performed in which the initial organic aerosol was separately exposed to the solute gas (P₁) with three different C_{σ}^* values: 10, 100, and 1000 µg 7 m^{-3} (molecular weight = 100 g mol⁻¹), with an initial gas-phase concentration of 6 µg m⁻³ in 8 9 each case. For each solute volatility case, the effect of aerosol phase state was examined using four different D_b values: 10⁻⁶, 10⁻¹², 10⁻¹³, and 10⁻¹⁵ cm² s⁻¹. In all cases, k_c was set at 0.01 s⁻¹ 10 so that τ_{ss} was always less than ~0.7 min across the entire size distribution. In each case, the 11 12 simulation was run until the gas-phase solute was completely absorbed and reacted to form a non-volatile product in the particle phase. Again, the molecular weight and density of the 13 product species (P₂) were assumed to be 100 g mol⁻¹ and 1 g cm⁻³, respectively, and all three 14 species $(P_1, P_2, and P_3)$ were assumed to form an ideal solution that participated in the 15 16 absorption of P1 according to Raoult's law. An additional set of reference simulations were performed for two extreme scenarios: 1) instantaneous particle-phase reaction (i.e., $k_c \rightarrow \infty$), 17 18 which is equivalent to solving the non-volatile solute condensation case (i.e., mechanism #1), 19 and 2) no particle-phase reaction ($k_c = 0$), which is referred to as Raoult's law partitioning (i.e., mechanism #2). In the latter case, the initial gas-phase concentrations for the different 20 C_g^* subcases were increased such that 6 µg m⁻³ of solute was partitioned into the particle 21 22 phase at steady state (i.e., at equilibrium) in each case.

23 4.1.1 Reference Cases

We shall first discuss the results of the closed system reference cases. Fig. 12 shows the gasphase decay and the corresponding temporal evolution of aerosol size distribution and mass fraction of newly formed SOA for the instantaneous particle-phase reaction case. Here, gasparticle partitioning is independent of the particle phase state and is governed entirely by gasphase diffusion limitation. Vapor concentration is completely depleted in about 1 h, and aerosol size distribution evolution displays the well-known narrowing characteristics as the small particles grow faster (more precisely, have greater $d \ln Dp / dt$) than the large ones (Zhang et al., 2012). Consequently, the mass fraction of the newly formed SOA in smaller particles is much higher than in the larger ones. Note that in the SOA mass fraction panel, the left-most point on each line with mass fraction ≈ 1 corresponds to the smallest initial particles $(D_n = 0.008 \ \mu m \ at t = 0)$.

6 In contrast, aerosol evolution due to Raoult's law partitioning depends on both solute 7 volatility and particle phase state. Fig. 13 shows the gas-phase concentration decay and the 8 corresponding aerosol size distribution and SOA mass fraction evolution for the less volatile solute with $C_g^* = 10 \ \mu g \ m^{-3}$. The effect of phase state is illustrated with two bulk diffusivities: 9 $D_b = 10^{-6}$ and 10^{-15} cm² s⁻¹. In the case with liquid particles ($D_b = 10^{-6}$ cm² s⁻¹), there is 10 negligible resistance to mass transfer within the particle (refer to Fig. 6a), and as a result the 11 12 vapor concentration rapidly decreases during the first 1 h and reaches a steady state in about 7.5 h. In the first ~ 20 min, the size distribution exhibits the narrowing of the Aitken mode 13 14 similar to that seen in gas-phase diffusion-limited growth, although not as intense. The SOA 15 mass fraction reaches up to 0.97 in small particles while it is only about 0.25 in the large particles. However, as the vapor concentration decreases further, the peak of the size 16 17 distribution begins to decrease and the width broadens due to evaporation from small particles 18 while the large particles continue to grow (Zhang et al., 2012). The SOA mass fraction in 19 small particles decreases to 0.75, while it gradually increases to 0.75 in the large particles. 20 The vapor concentration remains steady while this inter-particle mass transfer (via the gas 21 phase) occurs over a relatively longer period (~480 h) until the entire aerosol size distribution reaches equilibrium. 22

Similar behavior is seen in the case with semi-solid particles ($D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$), although 23 24 the timescale over which it occurs is relatively longer due to much higher particle-phase 25 diffusion limitation. While the vapor concentration declines rapidly in the beginning (e-26 folding timescale of 16.5 h), it takes about 175 h to reach the steady state and more than 400 h 27 for the aerosol size distribution to reach equilibrium. Also, because the particle-phase 28 diffusion limitation is much less in small particles than the large ones (refer to Fig. 6a), the 29 Aitken mode exhibits more intense narrowing and a higher peak (at about 1 h) than seen in 30 liquid particles. Then, again, as the vapor concentration decreases further, the width broadens 31 and the peak decreases due to evaporation of small particles while the large ones continue to grow more slowly. The final aerosol size distribution and SOA mass fraction across the size
 spectrum are identical (within numerical errors) to those obtained in the liquid particle case.

Fig. 14 shows the results for the more volatile solute with $C_g^* = 1000 \ \mu g \ m^{-3}$. In the case with 3 liquid particles ($D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$), the vapor concentration reaches the steady state in just 20 4 min (vs. 7.5 h for $C_g^* = 10 \ \mu g \ m^{-3}$) while it takes nearly 400 h (vs. 175 h for $C_g^* = 10 \ \mu g \ m^{-3}$) 5 in the case with semi-solid particles ($D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$). Again, the final aerosol size 6 distribution and SOA mass fraction solutions at equilibrium are identical to those obtained for 7 the $C_g^* = 10 \ \mu g \ m^{-3}$ cases, but their temporal evolutions are quite different. In the case with 8 9 liquid particles, the width of the aerosol size distribution does not narrow and the peak height 10 remains the same as the particles grow. This is because the small particles quickly attain a 11 quasi-equilibrium state with the more volatile solute. Consequently, the SOA mass fraction in 12 the small particles quickly reaches the equilibrium value of 0.75 (instead of overshooting as seen for $C_g^* = 10 \ \mu g \ m^{-3}$) while the larger particles catch up just a bit more slowly. The entire 13 size distribution reaches equilibrium within 1 h. 14

In the case with semi-solid particles, the Aitken mode size distribution narrows (similar to that seen in Fig. 13a) in the first few minutes, but broadens back within 30 min. Again, the SOA mass fraction in small particles quickly reaches the equilibrium value of 0.75, while it still takes ~480 h for the large particles in the spectrum to reach equilibrium due to the significant diffusion limitation in the particle phase.

20 4.1.2 Reactive Partitioning Cases

We now present results for the closed-system reactive partitioning cases with $k_c = 0.01 \text{ s}^{-1}$. 21 Fig. 15 shows vapor concentration decay for each of the three solute volatility cases ($C_g^* = 10$, 22 100, and 1000 μ g m⁻³) for D_b values ranging from 10⁻⁶ to 10⁻¹⁵ cm² s⁻¹. It also shows a plot of 23 the e-folding timescale (τ_g) for the decay as a function of D_b for the different volatilities. 24 Each plot includes the reference case of instantaneous reaction for comparison. Unlike in 25 Raoult's law partitioning, the vapor concentration always decays to zero in reactive 26 partitioning and the decay rate slows down with increase in C_{o}^{*} . The vapor decay rate also 27 slows down with decrease in D_b and it is especially sensitive to D_b in semi-solid particles. 28

Fig. 16 illustrates the effects of the different C_g^* and D_b values on the final aerosol size 1 2 distribution. The final results for the reference cases of instantaneous reaction and Raoult's law partitioning are also shown for easy comparison. In the case of $C_g^* = 10 \ \mu g \ m^{-3}$, the 3 Aitken mode exhibits significant narrowing for all values of D_{h} . The narrowing becomes 4 more pronounced for $D_b < 10^{-13} \text{ cm}^2 \text{ s}^{-1}$ with the shape of the entire size distribution for $D_b =$ 5 10⁻¹⁵ cm² s⁻¹ being nearly identical to that for the instantaneous reaction reference case. 6 Further decrease in D_b will produce even more narrowing. Since there is negligible particle-7 phase diffusion limitation for $D_b > 10^{-10} \text{ cm}^2 \text{ s}^{-1}$ ($Q \approx 1$; Fig. 7c), the size distribution of liquid 8 9 aerosol narrows because its initial evolution (in the case of low volatility solutes) resembles that of gas-phase diffusion-limited growth, and the particle-phase reaction rate is fast enough 10 to transform the absorbed solute to a non-volatile product before it can evaporate. For D_b < 11 10^{-13} cm² s⁻¹, the steep gradient in Q across the size distribution results in significantly lower 12 13 surface concentrations over small semi-solid particles compared to the large ones. The small 14 semi-solid particles therefore grow even faster than the large ones compared to the 15 corresponding liquid aerosol case, producing relatively more intense narrowing of the size distribution. 16

As the solute C_g^* increases to 100 and 1000 µg m⁻³, liquid particles tend to attain quasi-17 equilibrium with the gas phase relatively faster than the solute reacts within the particle. As a 18 result, the final size distributions for $D_b \leq 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ progressively resemble that of the 19 Raoult's law partitioning case. However, significant narrowing is still seen for $D_b = 10^{-15} \text{ cm}^2$ 20 s^{-1} due to the steep gradient in Q across the size distribution, which causes the small semi-21 22 solid particles to grow much faster than the large semi-solid ones when compared to the 23 corresponding liquid aerosol case where $Q \approx 1$ across the entire size distribution. In general, 24 the final size distribution shape tends to be closer to that for instantaneous reaction case for lower C_g^* and D_b values and higher k_c values, while it tends to be closer to that for Raoult's 25 law partitioning for higher C_g^* and D_b and lower k_c . 26

Fig. 17 illustrates the influence of C_g^* and D_b values on the final SOA mass fraction size distribution. Curves for the two reference cases are also included for comparison. In the case of $C_g^* = 10 \ \mu g \ m^{-3}$, the curves for all D_b values are similar to that of the instantaneous 1 reference case due to appreciable narrowing of the size distribution. But as C_g^* increases, the 2 SOA mass fraction curves progressively become more uniform for $D_b = 10^{-6}$ cm² s⁻¹ while 3 they remain non-uniform for $D_b < 10^{-12}$ cm² s⁻¹ for particles with $D_p > 0.2$ µm. In all C_g^* 4 cases, the SOA mass fraction curves for $D_b = 10^{-15}$ cm² s⁻¹ closely resemble the instantaneous 5 reaction case.

6 4.2 General System

A set of general system simulations was performed in which the initial organic aerosol was separately exposed to solutes with $C_g^* = 10, 100$, and $1000 \ \mu g \ m^{-3}$ at a moderate but constant gas-phase source rate of $\gamma = 0.6 \ \mu g \ m^{-3} \ h^{-1}$ in each case. The effect of aerosol phase state was examined using two different D_b values: 10^{-6} and $10^{-15} \ cm^2 \ s^{-1}$. For each combination of C_g^* and D_b values, the effect of particle-phase reaction was examined for $k_c = 0.01, 0.1, 1$, and $\infty \ s^{-1}$. Each simulation was 12 h long.

Fig. 18 shows the time evolutions of total SOA mass concentration for liquid particles (D_{h} = 13 10^{-6} cm² s⁻¹) with different solute C_g^* values and the corresponding final aerosol size 14 distributions at t = 12 h. In the case with $C_g^* = 10 \ \mu g \ m^{-3}$, the SOA formation rate is 15 essentially the same for $k_c \ge 0.01 \text{ s}^{-1}$, with a total of about 7 µg m⁻³ SOA formed at the end of 16 12 h. Appreciable narrowing of the Aitken mode size distribution occurs for $k_c = 0.01 \text{ s}^{-1}$, 17 which is qualitatively similar to the closed system results for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ shown 18 19 previously in Fig. 16a. Higher k_c values produce even more intense narrowing of the Aitken 20 mode and the shapes are practically indistinguishable from that for instantaneous reaction. As C_{g}^{*} increases, the solute vapor tends towards quasi-equilibrium with the particle phase for low 21 22 k_c values. As a result, the SOA formation rate slows down and the Aitken mode shapes for $k_c = 0.01 \text{ s}^{-1}$ qualitatively tend to resemble that of Raoult's law partitioning in the closed 23 24 system shown previously in Fig. 16b,c. But as k_c increases, the mass transfer becomes progressively more gas-phase diffusion limited, which results in faster growth of the smaller 25 26 particles and, therefore, increasing narrowing of the Aitken mode.

Fig. 19 shows the results for semi-solid particles ($D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$). It is seen that the presence of significant particle-phase diffusion limitation slows down the SOA formation rates, especially with increasing C_g^* and decreasing k_c . The marked size-dependence of the diffusion limitation also gives rise to more intense narrowing of the size distribution than seen in the corresponding liquid particle cases.

6 In the absence of particle-phase reaction (i.e., $k_c = 0$, not shown in the figures) only ~1.2 µg m⁻³ SOA is formed in both the liquid and semi-solid aerosol cases after 12 h when $C_g^* = 10$ 7 μ g m⁻³ while negligibly small amounts of SOA are formed for higher C_g^* values. Overall, the 8 9 growth characteristics seen in the general system cases considered here are qualitatively 10 similar to the closed system results, although significant differences between them can occur 11 if the vapor source rate is appreciably different than the one used in the present study. For 12 instance, if the vapor source rate is very small, then the growth characteristics will tend 13 towards Raoult's law partitioning. In contrast, if the vapor source rate is very high, then the 14 growth will tend to become gas-phase diffusion limited.

15

16 **5** Summary and Implications

17 We have extended the computationally efficient MOSAIC aerosol model (Zaveri et al., 2008) to include a new framework for kinetic SOA partitioning that takes into account solute 18 19 volatility, gas-phase diffusion, interfacial mass accommodation, particle-phase diffusion, and 20 particle-phase reaction. The framework uses a combination of: (a) an analytical quasi-steady-21 state treatment for the diffusion-reaction process within the particle phase for fast-reacting organic solutes such that the timescales (τ_{OSS}) for their particle-phase concentrations to reach 22 23 quasi-steady state are shorter than 1 min, and (b) a two-film theory approach for slow- and 24 non-reacting organic solutes. The updated MOSAIC model was successfully validated against 25 a benchmark finite-difference solution of the diffusion-reaction problem. MOSAIC already 26 predicts liquid water associated with inorganic spcies, and the new framework can be readily 27 adapted to kinetically partition water soluble organic gases into the particulate aqueous phase 28 if that is the only liquid phase in the particle. Additional work is needed to treat mass transfer of gas-phase species to mixed inorganic-organic particles that experience liquid-liquid phase 29 30 separation (You et al., 2012). The proposed framework is amenable for use in regional and global atmospheric models, although it currently awaits specification of the various gas- and
 particle-phase chemistries and the related physicochemical properties that are important for
 SOA formation.

4 In the present study, we have applied the model to evaluate the effects of solute volatility (C_{\circ}^{*}) , particle-phase bulk diffusivity (D_{b}) , and particle-phase chemical reaction, as 5 exemplified by the pseudo-first-order rate constant (k_c) , on kinetic SOA partitioning. We 6 7 focus on the competitive growth dynamics of the Aitken and accumulation mode particles due 8 to condensation while the Kelvin effect and coagulation are neglected for simplicity. Our 9 analysis shows that the timescale of SOA partitioning and the associated evolution of aerosol number and composition size distributions depend on the complex interplay between C_g^* , D_b , 10 and k_c , each of which can vary over several orders of magnitude. The key findings and their 11 12 implications are summarized below.

13 a) In the case of instantaneous particle-phase reaction $(k_c \rightarrow \infty)$, SOA partitioning is

mathematically equivalent to irreversible condensation of non-volatile organic vapors (C_g^* = 0; mechanism #1). Mass transfer is gas-phase diffusion limited, which produces the well-known narrowing of the aerosol size distribution as small particles grow faster than the large ones (Zhang et al., 2012).

b) In the case of non-reactive reversible absorption of semi-volatile and volatile organic 18 vapors by Raoult's law ($k_c = 0$; mechanism #2), the final partitioning across the size 19 20 distribution is volume-controlled (Zhang et al., 2012) and the partitioning timescale increases with decrease in C_g^* and D_b (Shiraiwa and Seinfeld, 2012b). In the absence of 21 22 the Kelvin effect and coagulation, the mole fraction of SOA across the final size 23 distribution at equilibrium is identical. As a result, the size distribution simply shifts along the diameter axis while its shape (mode widths and peak heights) remains unchanged. 24 25 However, in a closed system, this mechanism may produce temporary narrowing of the size distribution as small particles reach quasi-equilibrium faster than the large ones 26 (Zhang et al., 2012). The narrowing is especially pronounced if the pre-existing particles 27 are highly viscous semi-solids ($D_b < 10^{-12} \text{ cm}^2 \text{ s}^{-1}$) and the initial gas-phase concentration 28 29 is appreciably higher than the solute vapor volatility. Also, while the vapor concentration 30 may reach a steady-state relatively quickly, the timescale for the "narrowed" aerosol size

1 distribution to relax back to its final (equilibrium) shape can be of the order few minutes 2 to days, depending on the values of D_b and C_g^* .

In the case of reactive partitioning (finite k_c ; mechanism #3), the size distribution 3 4 experiences permanent narrowing (Shiraiwa et al., 2013a), which can be especially pronounced for low values of C_g^* (~10 µg m⁻³ and less) and D_b (< 10⁻¹³ cm² s⁻¹) and high 5 values of k_c (~0.01 s⁻¹ and higher). As C_g^* and D_b increase and k_c decreases, the 6 narrowing reduces and the final size distribution tends to resemble that produced by 7 mechanism #2. But unlike in mechanism #2, the gas-phase concentration of the solute 8 eventually decays to zero and the partitioning timescale increases with increase in C_g^* and 9 decrease in D_b and k_c . The partitioning timescale and the shape of the size distribution 10 are especially sensitive to the phase state when D_b is about 10^{-13} cm² s⁻¹ or less. At D_b = 11 10^{-15} cm² s⁻¹ and $k_c = 0.01$ s⁻¹, the decay timescale ranges from 1 h for $C_g^* = 10 \ \mu g \ m^{-3}$ to 12 about 3 days for $C_g^* = 1000 \ \mu g \ m^{-3}$. Consequently, for intermediate volatility solutes (C_g^* 13 $> 1000 \ \mu g \ m^{-3}$) to partition in appreciable amounts to semi-solid SOA via particle-phase 14 reactions, their k_c values need to be > 0.1 s⁻¹. 15

d) From a practical standpoint, the particle-phase concentration profiles of a solute (with any 16 C_g^*) reacting with $k_c > 0.01 \text{ s}^{-1}$ may be assumed to be at steady-state in particles of any 17 size and any phase state. Furthermore, for $k_c \leq 0.1 \text{ s}^{-1}$ and $D_b \geq 10^{-10} \text{ cm}^2 \text{ s}^{-1}$, the 18 19 particle-phase reaction occurs uniformly through the entire volume of submicron particles. At higher k_c or lower D_b values, the particle-phase concentration profile becomes 20 21 increasingly non-uniform (i.e., depleted towards the center of the particle) as the particle size increases. As a result, particle-phase reactions in large semi-solid particles occur 22 primarily near the surface while in smaller particles the same reactions may still occur 23 through the entire volume. These differences in the diffusion-reaction dynamics across the 24 25 size distribution, and its dependence on the particle phase state, together control the SOA 26 partitioning timescale and the size distribution evolution.

e) Observations of the evolution of the size distribution can provide valuable clues about the
underlying mechanisms of SOA formation (Riipinen et al., 2011; Shiraiwa et al., 2013a).

- 1 However, all three mechanisms, under certain combinations of C_g^* , D_b , and k_c values,
- 2 can produce similar looking aerosol number size distributions. A concerted experimental
- 3 strategy is therefore necessary to properly constrain these and other key model parameters
- 4 and effectively evaluate the next generation of SOA models that treat phase state
- 5 thermodynamics, particle-phase diffusion and particle-phase reactions.
- 6 f) A proper representation of these physicochemical processes and parameters is needed to
 7 reliably predict not only the total SOA mass, but also its composition- and number-
- 8 diameter distributions, which together determine the overall optical and cloud-nucleating9 properties.

Future model development work entails implementation of comprehensive gas-phase VOC oxidation mechanisms and the key particle-phase reactions that form organic salts, oligomers, hemiacetals, organosulfates, and other high molecular weight oxidation products, which constitute a significant fraction of SOA. At the same time, a computationally efficient treatment for phase transition thermodynamics (including liquid-liquid phase separation) is needed to provide the combined feedbacks of ambient temperature, relative humidity, and particle composition on the bulk diffusivity and reactivity of the absorbed organic solutes.

17

18 Appendix A: Overall Gas-side Mass Transfer Coefficient K_g

19 Flux F (mol cm⁻² s⁻¹) of species *i* across the gas-particle interface can be written in multiple 20 ways depending on the choice of the mass transfer coefficient:

- 21 Gas-side mass transfer coefficient: $F_i = k_{g,i} (\overline{C}_{g,i} C_{g,i}^s),$ (A1)
- 22 Particle-side mass transfer coefficient: $F_i = k_{p,i} (A_i^s \overline{A}_i),$ (A2)
- 23 Overall gas-side mass transfer coefficient: $F_i = K_{g,i}(\overline{C}_{g,i} S'_i \overline{A}_i)$. (A3)
- In Eq. (A3) the term $(\overline{C}_{g,i} S'_i \overline{A}_i)$ is the overall driving force for mass transfer between the bulk gas-phase and the average bulk particle phase, where

26
$$S'_{i} = \frac{C^{*}_{g,i}}{\sum_{j} \overline{A_{j}}}$$
 (A4)

In the above equations, k_g (cm s⁻¹) is the gas-side mass transfer coefficient, k_p (cm s⁻¹) is the particle-side mass transfer coefficient, and K_g (cm s⁻¹) is the overall gas-side mass transfer coefficient.

4 We can rewrite Eq. (A3) as:

5
$$\frac{1}{K_{g,i}} = \frac{\overline{C}_{g,i} - S'_i \overline{A}_i}{F_i} = \frac{\left(\overline{C}_{g,i} - C^s_{g,i}\right) + \left(C^s_{g,i} - S'_i \overline{A}_i\right)}{F_i}.$$
 (A5)

6 Applying Raoult's law at the interface, we get:

$$7 C_{g,i}^s = S_i' A_i^s (A6)$$

8 Combining Eqs. (A5) and (A6), we get:

9
$$\frac{1}{K_{g,i}} = \frac{\left(\overline{C}_{g,i} - C_{g,i}^{s}\right)}{F_{i}} + \frac{S_{i}'\left(A_{i}^{s} - \overline{A}_{i}\right)}{F_{i}}.$$
 (A7)

10 Combining Eqs. (A1), (A2), and (A7), we can relate the overall gas-side mass transfer 11 coefficient to gas-side and particle-side mass transfer coefficients as:

12
$$\frac{1}{K_{g,i}} = \frac{1}{k_{g,i}} + \frac{S'_i}{k_{p,i}}.$$
 (A8)

13 Finally, replacing the flux term in Eq. (19) with Eq. (A3) gives

14
$$\frac{d\overline{A}_{i}}{dt} = \frac{3}{R_{p}} K_{g,i} \left\{ \overline{C}_{g,i} - \frac{\overline{A}_{i}}{\sum_{j} \overline{A}_{j}} C_{i}^{*} \right\} - k_{c,i} \overline{A}_{i}.$$
(A9)

15

16 Appendix B: Particle-side Mass Transfer Coefficient *k*_p

17 As noted in the main paper, the particle-side film thickness δ_p , and therefore k_p and K_g , are 18 not readily known. We estimate these parameters by assuming that under quasi-steady state 19 conditions, the analytical solution (Eq. (19)) and the two-film theory (Eq. (21)) give the same 20 results. Under quasi-steady state conditions, Eq. (19) becomes:

21
$$\frac{d\overline{A}_{i}}{dt} = \frac{3}{R_{p}} k_{g,i} \left\{ \overline{C}_{g,i} - \frac{S'_{i} \overline{A}_{i}}{Q_{i}} \right\} - k_{c,i} \overline{A}_{i} = 0.$$
(B1)

1 Rearranging Eq. (B1), we have:

$$2 \qquad \frac{\overline{A}_i}{\overline{C}_{g,i}} = \left(\frac{S'_i}{Q_i} + \frac{k_{c,i}R_p}{3k_{g,i}}\right)^{-1}.$$
 (B2)

3 Similarly, assuming quasi-steady-state for Eq. (21), we get:

$$4 \qquad \frac{dA_i}{dt} = \frac{3}{R_p} K_{g,i} \left\{ \overline{C}_{g,i} - S'_i \overline{A}_i \right\} - k_{c,i} \overline{A}_i = 0.$$
(B3)

5 Rearranging Eq. (B3), we have:

$$6 \qquad \frac{\overline{A}_i}{\overline{C}_{g,i}} = \left(S'_i + \frac{k_{c,i}R_p}{3K_{g,i}}\right)^{-1}.$$
(B4)

7 With our assumption that the two approaches produce the same quasi-steady state solutions,

8 the left-hand-sides of Eqs. (B2) and (B4) are equal, so equating their right-hand-sides gives:

9
$$\frac{S'_i}{Q_i} + \frac{k_{c,i}R_p}{3k_{g,i}} = S'_i + \frac{k_{c,i}R_p}{3K_{g,i}}.$$
 (B5)

10 Substituting the expression for $K_{g,i}$ from Eq. (A8) in Eq. (B5), and simplifying the resulting 11 equation for $k_{p,i}$ yields:

12
$$k_{p,i} = \frac{k_{c,i}R_p}{3} \left(\frac{Q_i}{1-Q_i}\right).$$
 (B6)

13 Substituting the expression for Q_i from Eq. (8) in Eq. (B6), we get:

14
$$k_{p,i} = \frac{k_{c,i}R_p}{q_i^2} \left(\frac{q_i \coth q_i - 1}{1 - Q_i}\right).$$
 (B7)

15 Using $q_i^2 = R_p^2 k_{c,i} / D_{b,i}$ in Eq. (B7) gives:

16
$$k_{p,i} = \frac{D_{b,i}}{R_p} \left(\frac{q_i \coth q_i - 1}{1 - Q_i} \right).$$
 (B8)

17 The particle-side film thickness is then expressed as:

18
$$\delta_{p,i} = R_p \left(\frac{1 - Q_i}{q_i \coth q_i - 1} \right).$$
(B9)

1 Note that as $k_{c,i} \to 0$, $q_i \to 0$, and $Q_i \to 1$. Taylor's series expansion of Eq. (B8) yields:

2
$$k_{p,i} = \frac{D_{b,i}}{R_p} \frac{\left(\frac{1}{3} - \frac{q^2}{45} + \frac{2q^4}{945} - ...\right)}{\left(\frac{1}{15} - \frac{2q^2}{315} + ...\right)}$$
 (B10)

3 Thus, in the limiting case of non-reacting solute ($k_{c,i} = 0$), Eq. (B10) reduces to:

4
$$k_{p,i} = 5 \frac{D_{b,i}}{R_p}$$
. (B11)

5

6 Acknowledgements

7 RAZ thanks Drs. M. Shiraiwa (Max Planck Institute for Chemistry), A. Zelenyuk, M. 8 Shrivastava (Pacific Northwest National Laboratory), J. Wang (Brookhaven National 9 Laboratory), M. West (University of Illinois at Urbana-Champaign), and A. S. Wexler 10 (University of California, Davis) for stimulating discussions. This research was supported by 11 the US Department of Energy (DOE) Atmospheric System Research (ASR) Program under 12 contract DE-AC06-76RLO 1830 at Pacific Northwest National Laboratory (PNNL). 13 Participation of JHS in this research was supported by DOE grant DE-SC0006626 and National Science Foundation grant AGS-1057183. PNNL is operated for the US DOE by 14 15 Battelle Memorial Institute.

1 References

- Abramson, E., Imre, D., Beránek, J., Wilson, J., and Zelenyuk, A.: Experimental
 determination of chemical diffusion within secondary organic aerosol particles, Phys. Chem.
- 4 Chem. Phys., 15, 2983-2991, doi:10.1039/c2cp44013j, 2013.
- 5 Astarita, G.: Mass Transfer with Chemical Reaction. Elsevier, New York, 1967.
- Bird, R. B., Stewart, W. E., and Lightfoot, E. N.: *Transport Phenomena* (revised 2nd ed.).
 John Wiley & Sons, Inc., Hoboken, NJ, 2007.
- 8 Berkemeier, T., Huisman, A. J., Ammann, M., Shiraiwa, M., Koop, T., and Pöschl, U.:
- 9 Kinetic regimes and limiting cases of gas uptake and heterogeneous reactions in atmospheric
- 10 aerosols and clouds: a general classification scheme, Atmos. Chem. Phys., 13, 6663-6686,
- 11 doi:10.5194/acp-13-6663-2013, 2013.
- Bowman, F. M., Odum, J. R., and Seinfeld, J. H.: Mathematical model for gas-particle partitioning of secondary organic aerosols, Atmos. Environ., 31(23), 3921-3931, 1997.
- Carlton, A. G., and Turpin, B. J.: Particle partitioning potential of organic compounds is
 highest in the Eastern US and driven by anthropogenic water, Atmos. Chem. Phys., 13,
 10203-10214, doi:10.5194/acp-13-10203-2013, 2013.
- Carlton, A. G., Turpin, B. I., Altieri, K. E., Seitzinger, S. P., Mathur, R., Roselle, S. J., and
 Weber, R. J.: CMAQ model performance enhanced when in-cloud secondary organic aerosol
- 19 is included: comparisons of organic carbon predictions with measurements, Environ. Sci.
- 20 Technol., 42, 8798-8802, 2008.
- Carslaw, H. S., and Jaeger, J. C.: Conduction of Heat in Solids, 2nd ed., Clarendon Press,
 Oxford, 1959.
- Crank, J.: The Mathematics of Diffusion, 2nd ed., Oxford University Press Inc., New York,
 1975.
- Danckwerts, P. V.: Absorption and simultaneous diffusion and chemical reaction into
 particles of various shapes and into falling drops, Trans. Faraday Soc., 47, 1014-1023, 1951.
- 27 Donahue, N. M., Trump, E. R., Pierce, J. R., and Riipinen, I.: Theoretical constraints on pure
- vapor-pressure driven condensation of organics to ultrafine particles, Geophys. Res. Lett., 38,
- 29 L16801, doi:10.1029/2011GL048115, 2011.

- 1 Doraiswamy, L. K., and Sharma, M. M.: Heterogeneous Reactions: Analysis, Examples, and
- 2 Reactor Design, Vol. 2.: Fluid-Fluid-Solid Reactions. Wiley-Interscience, New York, 1984.
- 3 Dusek, U., Frank, G. P., Hildebrandt, L., Curtius, J., Schneider, J., Walter S., Chand, D.,
- 4 Drewnick, F., Hings, S., Jung, D., Borrmann, S., and Andreae, M. O.: Size matters more than
- 5 chemistry for cloud-nucleating ability of aerosol particles, Science, 312, 1375-1378,
- 6 doi:10.1126/science.1125261, 2006.
- 7 Ervens, B., Carlton, A. G., Turpin, B. J., Altieri, K. E., Kreidenweis, S. M., and Feingold, G.:
- 8 Secondary organic aerosol yields from cloud-processing of isoprene oxidation products,
- 9 Geophys. Res. Lett., 35, L02816, doi:10.1029/2007GL031828, 2008.
- 10 Ervens, B. and Volkamer, R.: Glyoxal processing by aerosol multiphase chemistry: towards a
- 11 kinetic modeling framework of secondary organic aerosol formation in aqueous particles,
- 12 Atmos. Chem. Phys., 10, 8219-8244, doi:10.5194/acp-10-8219-2010, 2010.
- Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J., Yoo, K.
 Y., Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle phase acidity
 and oligomer formation in secondary organic aerosol, Environ. Sci. Technol., 38(24), 6582-
- 16 6589, doi:10.1021/es049125k, 2004.
- Gunthe, S. S., King, S. M., Rose, D., Chen, Q., Roldin, P., Farmer, D. K., Jimenez, J. L.,
 Artaxo, P., Andreae, M. O., Martin, S. T., and Pöschl, U.: Cloud condensation nuclei in
 pristine tropical rainforest air of Amazonia: size-resolved measurements and modeling of
 atmospheric aerosol composition and CCN activity, Atmos. Chem. Phys., 9, 7551-7575,
 doi:10.5194/acp-9-7551-2009, 2009.
- Gunthe, S. S., Rose, D., Su, H., Garland, R. M., Achtert, P., Nowak, A., Wiedensohler, A.,
 Kuwata, M., Takegawa, N., Kondo, Y., Hu, M., Shao, M., Zhu, T., Andreae, M. O., and
 Pöschl, U.: Cloud condensation nuclei (CCN) from fresh and aged air pollution in the
 megacity region of Beijing, Atmos. Chem. Phys., 11, 11023-11039, doi:10.5194/acp-1111023-2011, 2011.
- Hall, W. A. and Johnston, M. V.: Oligomer content of α-pinene secondary organic aerosol,
 Aerosol Sci. Technol, 45(1), 37–45, 2011.
- 29 Heaton, K. J., Dreyfus, M. A., Wang, S., and Johnston, M. V.: Oligomers in the early stage of
- 30 biogenic secondary organic aerosol formation and growth, Environ. Sci. Technol., 41, 6129-
- 6139, 2007.

- Hoyle, C. R., Boy, M., Donahue, N. M., Fry, J. L., Glasius, M., Guenther, A., Hallar, A. G.,
 Hartz, K. H., Petters, M. D., Petaja, T., Rosenoern, T., and Sullivan, A. P.: A review of the
 anthropogenic influence on biogenic secondary organic aerosol, Atmos. Chem. Phys., 11,
 321–343, doi:10.5194/acp-11-321-2011, 2011.
- 5 Kalberer, M., Paulsen, D., Sax, M., Steinbacher, M., Dommen, J., Prevot, A. S. H., Fisseha,
- 6 R., Weingartner, E., Frankevich, V., Zenobi, R., and Baltensperger, U.: Identification of
- polymers as major components of atmospheric organic aerosols, Science, 303, 1659-1662,
 2004.
- 9 Kuang, C., Chen, M., Zhao, J., Smith, J., McMurry, P. H., and Wang, J.: Size and time-
- 10 resolved growth rate measurements of 1 to 5 nm freshly formed atmospheric nuclei, Atmos.
- 11 Chem. Phys., 12, 3573–3589, doi:10.5194/acp-12-3573-2012, 2012.
- 12 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili,
- 13 W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a
- 14 review of observations, J. Aerosol Sci., 35, 143-176, 2004.
- Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and
 evolution of low-volatility organics in the atmosphere, Atmos. Environ., 42(16), 3593–3624,
 2008.
- Lewis, W. K. and Whitman, W. G.: Principles of gas absorption, Ind. Eng. Chem., 16(12),
 1215-1220, 1924.
- Liu, A. T., Zaveri, R. A., and Seinfeld, J. H.: Technical Note: Analytical solution for transient
 partitioning and reaction of a condensing vapor species in a droplet, Atmos. Environ., 89,
 651-654, 2014.
- Liu, Y., Monod, A., Tritscher, T., Praplan, A. P., DeCarlo, P. F., Temime-Roussel, B., Quivet,
 E., Marchand, N., Dommen, J., and Baltensperger, U.: Aqueous phase processing of
 secondary organic aerosol from isoprene photooxidation, Atmos. Chem. Phys., 12, 5879–
 5895, doi:10.5194/acp-12-5879-2012, 2012.
- Mouchel-Vallon, C., Brauer, P., Camredon, M., Valorso, R., Madronich, S., Herrmann, H.,
 and Aumont, B.: Explicit modeling of volatile organic compounds partitioning in the
 atmospheric aqueous phase, Atmos. Chem. Phys., 13, 1023-1037, doi:10.5194/acp-13-10232013, 2013.

- 1 Nozière, B., Dziedzic, P., and Córdova, A.: Formation of secondary light-absorbing "fulvic-
- 2 like" oligomers: A common process in aqueous and ionic atmospheric particles?, Geophys.
- 3 Res. Lett., 34, L21812, doi:10.1029/2007GL031300, 2007.
- 4 Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flentje,
- 5 H., Birmili, W., Wiedensohler, A., Hõrrak, U., Metzger, A., Hamed, A., Laaksonen, A.,
- 6 Facchini, M. C., Kerminen, V.-M., and Kulmala, M.: On the roles of sulphuric acid and low-
- 7 volatility organic vapours in the initial steps of atmospheric new particle formation, Atmos.
- 8 Chem. Phys., 10, 11223–11242, doi:10.5194/acp-10-11223-2010, 2010.
- 9 Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in the
- 10 atmosphere, Atmos. Environ., 28, 185-188, 1994.
- 11 Parikh, H. M., Carlton, A. G., Vizuete, W., and Kamens, R. M.: Modeling secondary organic
- 12 aerosol using a dynamic partitioning approach incorporating particle aqueous-phase
- 13 chemistry, Atmos. Environ., 45, 1126-1137, doi:10.1016/j.atmosenv.2010.11.027, 2011.
- Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Yu, Y., Alexander, M. L., Zelenyuk,
 A., Imre, D., Chang, W. L., Dabdub, D., Pankow, J. F., and Finlayson-Pitts, B. J.:
 Nonequilibrium atmospheric secondary organic aerosol formation and growth, Proc. Nat.
 Acad. Sci. USA, 109(8), 2836-2841, doi: 10.1073/pnas.1119909109, 2012.
- Pierce, J. R., Riipinen, I., Kulmala, M., Ehn, M., Petäjä, T., Junninen, H., Worsnop, D. R.,
 and Donahue, N. M.: Quantification of the volatility of secondary organic compounds in
 ultrafine particles during nucleation events, Atmos. Chem. Phys., 11, 9019–9036,
 doi:10.5194/acp-11-9019-2011, 2011.
- 22 Pierce, J. R., Leaitch, W. R., Liggio, J., Westervelt, D. M., Wainwright, C. D., Abbatt, J. P.
- 23 D., Ahlm, L., Al-Basheer, W., Cziczo, D. J., Hayden, K. L., Lee, A. K. Y., Li, S.-M., Russell,
- L. M., Sjostedt, S. J., Strawbridge, K. B., Travis, M., Vlasenko, A., Wentzell, J. J. B., Wiebe,
- 25 H. A., Wong, J. P. S., and Macdonald, A. M.: Nucleation and condensational growth to CCN
- 26 sizes during a sustained pristine biogenic SOA event in a forested mountain valley, Atmos.
- 27 Chem. Phys., 12, 3147-3163, doi:10.5194/acp-12-3147-2012, 2012.
- 28 Pöschl, U., Rudich, Y., and Ammann, M.: Kinetic model framework for aerosol and cloud
- 29 surface chemistry and gas-particle interactions Part 1: General equations, parameters, and
- 30 terminology, Atmos. Chem. Phys., 7, 5989–6023, doi:10.5194/acp-7-5989-2007, 2007.

- 1 Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, M., Sellier, M., Murray, B. J.,
- 2 Shilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of α-pinene secondary organic
- 3 material and implications for particle growth and reactivity, Proc. Nat. Acad. Sci. USA,
- 4 110(20), 8014-8019, doi:10.1073/pnas.1219548110, 2013.
- 5 Riipinen, I., Pierce, J. R., Yli-Juuti, T., Nieminen, T., Häkkinen, S., Ehn, M., Junninen, H.,
- 6 Lehtipalo, K., Petäjä, T., Slowik, J., Chang, R., Shantz, N. C., Abbatt, J., Leaitch, W. R.,
- 7 Kerminen, V.-M., Worsnop, D. R., Pandis, S. N., Donahue, N. M., and Kulmala, M.: Organic
- 8 condensation: a vital link connecting aerosol formation to cloud condensation nuclei (CCN)
- 9 concentrations, Atmos. Chem. Phys., 11, 3865–3878, doi:10.5194/acp-11-3865-2011, 2011.
- 10 Riipinen, I., Yli-Juuti, T., Pierce, J. R., Petäjä , T., Worsnop, D. R., Kulmala, M., and
- 11 Donahue, N. M.: The contribution of organics to atmospheric nanoparticle growth, Nature
- 12 Geoscience, 5, 453-458, 2012.
- 13 Saathoff, H., Naumann, K. -H., Mohler, O., Jonsson, A. M., Hallquist, M., Kiendler-Scharr,
- A., Mentel, Th. F., Tillmann, R., and Schurath, U.: Temperature dependence of yields of
 secondary organic aerosols from the ozonolysis of a-pinene and limonene, Atmos. Chem.
- 16 Phys., 9, 1551-1577, doi:10.5194/acp-9-1551-2009, 2009.
- Seinfeld, J. H. and Pandis, S. N.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change* (2nd ed.). John Wiley & Sons, Inc., Hoboken, NJ, 2006.
- Schwartz, S. E. and Freiberg, J. E.: Mass-transport limitation to the rate of reaction of gases in
 liquid droplets: application to oxidation of SO₂ in aqueous solutions. Atmos. Environ., 15,
 1129-1144, 1981.
- Shi, B. and Seinfeld, J.H: On mass-transport limitation to the rate of reaction of gases in
 liquid droplets. Atmos. Environ., 22, 2491-2499, 1991.
- 24 Shilling, J. E., Zaveri, R. A., Fast, J. D., Kleinman, L., Alexander, M. L., Canagaratna, M. R.,
- 25 Fortner, E., Hubbe J. M., Jayne, J. T., Sedlacek, A., Setyan, A., Springston, S., Worsnop, D.
- 26 R., and Zhang, Q.: Enhanced SOA formation from mixed anthropogenic and biogenic
- emissions during the CARES campaign, Atmos. Chem. Phys., 13, 2091-2113, 2013.
- 28 Shiraiwa, M., Pfrang, C., Koop, T., and Pöschl, U.: Kinetic multi-layer model of gas-particle
- 29 interactions in aerosols and clouds (KM-GAP): linking condensation, evaporation, and
- 30 chemical reactions of organics, oxidants and water. Atmos. Chem. Phys., 12, 2777-2794,
- 31 doi:10.5194/acp-12-2777-2012, 2012a.

- 1 Shiraiwa, M., and Seinfeld, J. H.: Equilibration timescale of atmospheric secondary organic
- 2 aerosol partitioning, Geophys. Res. Lett., 39, L24801, doi:10.1029/2012GL054008, 2012b.
- 3 Shiraiwa, M., Yee, L. D., Schilling, K. A., Loza, C. L., Craven, J. S., Zuend, A., Ziemann, P.
- 4 J., and Seinfeld, J. H.: Size distribution dynamics reveal particle-phase chemistry in organic
- 5 aerosol formation, Proc. Nat. Acad. Sci. USA, 110(29), 11746-11750,
- 6 doi: 10.1073/pnas.1307501110, 2013a.
- 7 Shiraiwa, M., Zuend, A., Bertram, A. K., and Seinfeld, J. H.: Gas-particle partitioning of
- 8 atmospheric aerosols: interplay of physical state, non-ideal mixing and morphology, Phys.
- 9 Chem. Chem. Phys., 15, 11441-11453, 2013b.
- 10 Simmel, M., and Wurzler, S.: Condensation and activation in sectional cloud microphysical
- 11 models, Atmos. Res., 80, 218-236, 2006.
- 12 Smith, J. N., Dunn, M. J., VanReken, T. M., Iida, K., Stolzenburg, M. R., McMurry, P. H.,
- 13 and Huey, L. G.: Chemical composition of atmospheric nanoparticles formed from nucleation
- 14 in Tecamac, Mexico: Evidence for an important role for organic species in nanoparticle
- 15 growth, Geophys. Res. Lett., 35, L04808, doi:10.1029/2007GL032523, 2008.
- 16 Surratt, J. D., Kroll, J. H., Kleindienst, T. E., Edney, E. O., Claeys, M., Sorooshian, A., Ng,
- 17 N. L., Offenberg, J. H., Lewandowski, M., Jaoui, M., Flagan, R. C., and Seinfeld, J. H.:
- 18 Evidence for organosulfates in secondary organic aerosol, Environ. Sci. Technol., 41, 517-
- 19 527, 2007.
- Tolocka, M., Jang, M., Ginter, J. M., Cox, F. J., Kamens, R. M., and Johnston, M. V.:
 Formation of oligomers in secondary organic aerosol, Environ. Sci. Technol., 38, 1428-1434,
 2004.
- Vaden, T. D., Imre, D., Beránek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation kinetics
 and phase of laboratory and ambient secondary organic aerosol, Proc. Nat. Acad. Sci. USA,
 108, 2190-2195, 2011.
- Virtanen, A., Joutsensaari, J., Koop, T., Kannosto, J., Yli-Pirila, P., Leskinen, J., Makela, J.
 M., Holopainen, J. K., Poschl, U., Kulmala, M., Worsnop, D. R., and Laaksonen, A.: An
 amorphous solid state of biogenic secondary organic aerosol particles, Nature, 467, 824-827,
 2010.

- Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V., and Zhang, R.: Atmospheric
 nanoparticles formed from heterogeneous reactions of organics, Nature Geoscience, 3, 238 242, 2010.
- Weber, R. J., Sullivan, A. P., Peltier, R. E., Russell, A., Yan, B., Zheng, M., de Gouw, J.,
 Warneke, C., Brock, C., Holloway, J. S., Atlas, E. L., and Edgerton, E.: A study of secondary
 organic aerosol formation in the anthropogenic-influenced southeastern United States, J.
- 7 Geophys. Res. Atmos., 112, D13302, doi:10.1029/2007jd008408, 2007.
- 8 Winkler, P. M., Ortega, J., Karl, T., Cappellin, L., Friedli, H. R., Barsanti, K., McMurry, P.
- 9 H., and Smith, J. N.: Identification of the biogenic compounds responsible for size-dependent
- 10 nanoparticle growth, Geophys. Res. Lett., 39, L20815, doi:10.1029/2012GL053253, 2012.
- 11 You, Y., Renbaum-Wolff, L., Carreras-Sospedra, M., Hanna, S. J., Hiranuma, N., Kamal, S.,
- 12 Smith, M. L., Zhang, X., Weber, R. J., Shilling, J. E., Dabdub, D., Martin, S. T., and Bertram,
- 13 A. K.: Images reveal that atmospheric particles can undergo liquid-liquid phase separations,
- 14 Proc. Nat. Acad. Sci. USA, 109(33), 13188-13193, doi: 10.1073/pnas. 1206414109, 2012.
- Zaveri, R. A.: Development and evaluation of a comprehensive tropospheric chemistry model
 for regional and global applications, Ph.D. thesis, Virginia Polytech. Inst. and State Univ.,
 Blacksburg, Virginia, 1997.
- Zaveri, R. A., Easter, R. C., Fast, J. D., and Peters, L. K.: Model for Simulating Aerosol
 Interactions and Chemistry (MOSAIC), J. Geophys. Res., 113, D13204,
 doi:10.1029/2007JD008782, 2008.
- Zaveri, R. A., Berkowitz, C. M., Brechtel, F. J., Gilles, M. K., Hubbe, J. M., Jayne, J. T.,
 Kleinman, L. I., Laskin, A., Madronich, S., Onasch, T. B., Pekour, M., Springston, S. R.,
 Thornton, J. A., Tivanski, A. V., and Worsnop, D. R.: Nighttime chemical evolution of
 aerosol and trace gases in a power plant plume: Implications for secondary organic nitrate and
 organosulfate aerosol formation, NO₃ radical chemistry, and N₂O₅ heterogeneous hydrolysis,
 J. Geophys. Res., 115, D12304, doi:10.1029/2009JD013250, 2010.
- 27 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Allan, J. D., Coe, H., Ulbrich, I., Alfarra, M.
- 28 R., Takami, A., Middlebrook, A. M., Sun, Y. L., Dzepina, K., Dunlea, E., Docherty, K.,
- 29 DeCarlo, P. F., Salcedo, D., Onasch, T., Jayne, J. T., Miyoshi, T., Shimono, A., Hatakeyama,
- 30 S., Takegawa, N., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S.,
- 31 Demerjian, K., Williams, P., Bower, K., Bahreini, R., Cottrell, L., Griffin, R. J., Rautiainen,

- J., Sun, J. Y., Zhang, Y. M., and Worsnop, D. R.: Ubiquity and dominance of oxygenated
 species in organic aerosols in anthropogenically-influenced Northern Hemisphere mid latitudes, Geophys. Res. Lett., 34, L13801, doi:10.1029/2007GL029979, 2007.
- Zhang, X., Pandis, S. N., and Seinfeld, J. H.: Diffusion-limited versus quasi-equilibrium
 aerosol growth, Aerosol Science and Technology, 46, 874-885,
 doi:10.1080/02786826.2012.679344, 2012.
- Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic
 aerosol formation, Chem. Soc. Rev., 41(19), 6582-6605, 2012.
- 9 Zobrist, B., Soonsin, V., Luo, B. P., Krieger, U. K., Marcolli, C., Peter, T., and Koop, T.:
- 10 Ultra-slow water diffusion in aqueous sucrose glasses, Phys. Chem. Chem. Phys., 13, 3514–
- 11 3526, doi:10.1039/C0CP01273D, 2011.
- 12

	$C_g^* = 10 \ \mu g \ m^{-3}$			$C_g^* = 100 \ \mu g \ m^{-3}$			$C_g^* = 1000 \ \mu g \ m^{-3}$		
k_c	MNB	MNGE	maxNGE	MNB	MNGE	maxNGE	MNB	MNGE	maxNGE
(s^{-1})	(%)	(%)	(%)	(%)	(%)	(%)	(%)	(%)	(%)
0	4.5	4.5	7.7	0.3	0.3	0.4	0.03	0.03	0.1
10-4	8.5	11.3	19.4	-1.7	1.7	3.1	-0.3	0.3	0.4
10 ⁻³	10.0	11.3	25.7	-1.3	1.3	3.2	-0.1	0.1	0.2
10 ⁻²	-1.3	1.3	3.4	3.4	3.4	4.2	-0.3	0.3	1.0
10-1	-4.3	4.3	10.7	-2.2	2.6	7.7	0.5	0.7	1.1

Table 1. Bias and Error Statistics for MOSAIC Predictions for the Closed System
 Simulations.

	$C_g^* = 10 \ \mu g \ m^{-3}$			$C_g^* = 100 \ \mu g \ m^{-3}$			$C_g^* = 1000 \ \mu g \ m^{-3}$		
k_c (s ⁻¹)	MNB (%)	MNGE (%)	maxNGE (%)	MNB (%)	MNGE (%)	maxNGE (%)	MNB (%)	MNGE (%)	maxNGE (%)
0	0.8	0.8	8.5	0.3	0.3	2.0	0.03	0.03	0.3
10-4	-1.0	2.2	6.4	0.6	0.6	1.8	0.1	0.1	0.2
10-3	-3.0	3.1	5.8	-0.7	1.2	3.1	0.1	0.1	0.2
10 ⁻²	-3.2	3.2	5.8	-2.3	2.3	4.8	-0.2	0.2	0.8
10-1	-2.4	2.4	5.0	-0.2	1.4	2.8	1.0	1.0	1.6

Table 2. Bias and Error Statistics for MOSAIC Predictions for the General System
 Simulations.

1 Figure Captions

2

Figure 1. Schematic of the gas-particle mass transfer process, with both diffusion and reaction
occurring inside the particle phase.

5

6 Figure 2. Normalized transient concentration $(A(r,t)/A^s)$ profiles as a function of 7 normalized radius (r/R_p) for a particle of diameter $R_p = 0.05 \ \mu\text{m}$ for different values of bulk 8 phase diffusivity and first-order reaction rate constants: (a) $D_b = 10^{-6} \ \text{cm}^2 \ \text{s}^{-1}$, $k_c = 0 \ \text{s}^{-1}$; (b) 9 $D_b = 10^{-6} \ \text{cm}^2 \ \text{s}^{-1}$, $k_c = 5 \ \text{x} \ 10^{-4} \ \text{s}^{-1}$; (c) $D_b = 10^{-15} \ \text{cm}^2 \ \text{s}^{-1}$, $k_c = 0 \ \text{s}^{-1}$; and (d) $D_b = 10^{-15} \ \text{cm}^2$ 10 $\ \text{s}^{-1}$, $k_c = 5 \ \text{x} \ 10^{-4} \ \text{s}^{-1}$.

11

Figure 3. Normalized steady-state concentration $(A(r)/A^s)_{ss}$ profiles as a function of normalized radius (r/R_p) for a particle of diameter $R_p = 0.05 \ \mu\text{m}$ and a range of k_c values for (a) $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ (b) $D_b = 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ (c) $D_b = 10^{-13} \text{ cm}^2 \text{ s}^{-1}$ (d) $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$.

15

16 Figure 4. Normalized steady state concentration $(A(r)/A^s)_{ss}$ profiles as a function of 17 normalized radius (r/R_p) for different values of dimensionless diffuso-reactive parameter q.

18

Figure 5. (a) Contour plots of: (a) particle-phase quasi-steady state timescale (τ_{QSS}), and (b) quasi-steady state parameter $Q = (\overline{A} / A^s)_{QSS}$ as functions of first-order rate constant (k_c) and bulk diffusion coefficient (D_b) for a species diffusing and reacting within semi-solid and liquid particles of diameter $D_p = 0.1 \,\mu\text{m}$.

23

Figure 6. Dependence of τ_{QSS} on particle diameter D_p for D_b values ranging from 10^{-10} to 10^{-10} $k_c = 0 \text{ s}^{-1}$; (b) $k_c = 10^{-3} \text{ s}^{-1}$; (c) $k_c = 10^{-2} \text{ s}^{-1}$; and (d) $k_c = 10^{-1} \text{ s}^{-1}$. 1 Figure 7. Dependence of Q on particle diameter D_p for D_b values ranging from 10^{-10} to 10^{-18} 2 cm² s⁻¹: (a) $k_c = 0$ s⁻¹; (b) $k_c = 10^{-3}$ s⁻¹; (c) $k_c = 10^{-2}$ s⁻¹; and (d) $k_c = 10^{-1}$ s⁻¹.

3

4 Figure 8. Schematic of the two-film theory.

5

Figure 9. Comparison of MOSAIC (lines) and finite difference model (filled circles) solutions for gas-phase concentration decay in a closed system due to kinetic gas-particle partitioning to particles with initial $D_p = 0.2 \,\mu\text{m}$, $N = 5000 \,\text{cm}^{-3}$, $D_b = 10^{-15} \,\text{cm}^2 \,\text{s}^{-1}$ and k_c ranging from 0 to 0.1 s⁻¹ for three solute volatilities: (a) $C_g^* = 10 \,\mu\text{g m}^{-3}$, (b) $C_g^* = 100 \,\mu\text{g m}^{-3}$, and (c) $C_g^* = 100 \,\mu\text{g m}^{-3}$.

11

Figure 10. Comparison of MOSAIC (lines) and finite difference model (filled circles) solutions for gas-phase concentration evolution in a general system due to kinetic gas-particle partitioning to particles with initial $D_p = 0.2 \,\mu\text{m}$, $N = 5000 \,\text{cm}^{-3}$, $D_b = 10^{-15} \,\text{cm}^2 \,\text{s}^{-1}$, $\gamma = 0.1$ $\mu\text{g m}^{-3} \,\text{h}^{-1}$, and k_c ranging from 0 to 0.1 s⁻¹ for three solute volatilities: (a) $C_g^* = 10 \,\mu\text{g m}^{-3}$, (b) $C_g^* = 100 \,\mu\text{g m}^{-3}$, and (c) $C_g^* = 1000 \,\mu\text{g m}^{-3}$.

17

Figure 11. Initial aerosol number and volume size distributions along with the condensational sink k_{CS} . The dashed line demarcates the Aitken mode from the accumulation mode and the initial condensation sink is such that the sum of k_{CS} over all the size bins in the Aitken mode is equal to that in the accumulation mode.

22

Figure 12. Results for the instantaneous reaction reference case ($k_c \rightarrow \infty$; equivalent to nonvolatile solute condensation): (a) Gas-phase concentration decay; (b) temporal evolution of aerosol size distribution; and (c) temporal evolution of the mass fraction of newly formed SOA.

Figure 13. Results for kinetic SOA partitioning due to Raoult's law ($k_c = 0 \text{ s}^{-1}$) for $C_g^* = 10$ µg m⁻³: (a) Gas-phase concentration decay for $D_b = 10^{-6}$ and $10^{-15} \text{ cm}^2 \text{ s}^{-1}$, (b) aerosol evolution for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$; (c) SOA mass fraction evolution for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$; (d) aerosol evolution for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$; and (e) SOA mass fraction evolution for $D_b = 10^{-15}$ $\text{cm}^2 \text{ s}^{-1}$. In both cases, the final (i.e., equilibrium) concentration of the newly formed SOA is 6 $\mu \text{g m}^{-3}$.

7

8 Figure 14. Same as Fig. 13, except $C_g^* = 1000 \ \mu g \ m^{-3}$.

9

Figure 15. Gas-phase concentration decay due to kinetic SOA partitioning with particle-phase reaction ($k_c = 0.01 \text{ s}^{-1}$) for bulk diffusivities ranging from 10⁻⁶ to 10⁻¹⁵ cm² s⁻¹ and three gas volatilities: (a) $C_g^* = 10 \ \mu g \ m^{-3}$; (b) $C_g^* = 100 \ \mu g \ m^{-3}$; and (c) $C_g^* = 1000 \ \mu g \ m^{-3}$. Each plot also shows gas-phase concentration decay for the reference case of instantaneous reaction (black line, $k_c \rightarrow \infty$). In each case, the final concentration of the newly formed SOA is 6 μg m⁻³. Panel (d) shows the plot of gas-phase concentration decay timescale (τ_g) as a function of D_b for the different gas volatilities.

17

Figure 16. Initial (dashed line) and final (solid lines) aerosol number size distribution due to 18 Raoult's law gas-particle partitioning coupled with particle-phase reaction ($k_c = 0.01 \text{ s}^{-1}$) for 19 bulk diffusivities ranging from 10^{-6} to 10^{-15} cm² s⁻¹ and three gas volatilities: (a) $C_g^* = 10 \ \mu g$ 20 m⁻³; (b) $C_g^* = 100 \ \mu g \ m^{-3}$; and (c) $C_g^* = 1000 \ \mu g \ m^{-3}$. Panel (d) shows the final size 21 distributions for the two reference cases: instantaneous reaction (black line; $k_c \rightarrow \infty$) and 22 Raoult's law partitioning (gray line; $k_c = 0$) for any D_b and $C_g^* > 0$. As illustrated in Fig. 15, 23 the time required to reach the final state differs significantly for different cases, but the final 24 SOA formed in each case is $6 \mu g m^{-3}$. 25

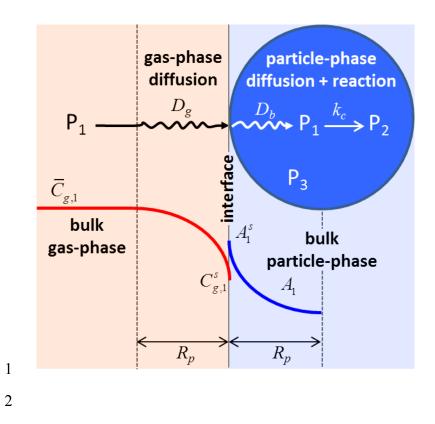
Figure 17. Final size distributions of the newly formed SOA mass fraction for different D_b values and: (a) $C_g^* = 10 \ \mu g \ m^{-3}$; (b) $C_g^* = 100 \ \mu g \ m^{-3}$; and (c) $C_g^* = 1000 \ \mu g \ m^{-3}$. Each panel also shows the reference plots for instantaneous reaction (black line; $k_c \rightarrow \infty$) and for Raoult's law partitioning (gray line; $k_c = 0 \ s^{-1}$) for any D_b and $C_g^* > 0$.

5

Figure 18. Temporal evolution of total SOA mass concentration (left column) and aerosol size distribution (right column) at t = 12 h for $D_b = 10^{-6}$ cm² s⁻¹, $\gamma = 0.6 \ \mu g \ m^{-3} \ h^{-1}$, $k_c = 0.01$ to ∞ s⁻¹, and three different solute volatilities: (a, b) $C_g^* = 10 \ \mu g \ m^{-3}$; (c, d) $C_g^* = 100 \ \mu g \ m^{-3}$; and (e, f) $C_g^* = 1000 \ \mu g \ m^{-3}$.

10

11 Figure 19. Same as Fig. 18, except $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$.



- 3 Figure 1. Schematic of the gas-particle mass transfer process, with both diffusion and reaction
- 4 occurring inside the particle phase.

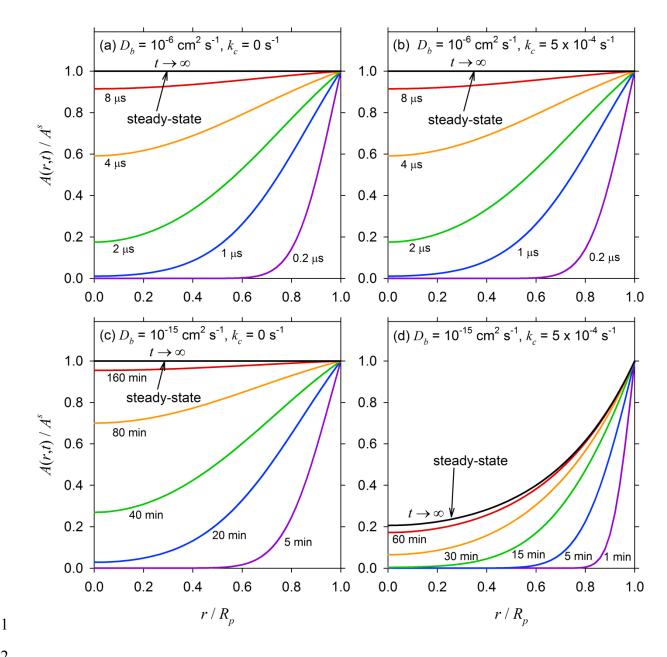


Figure 2. Normalized transient concentration $(A(r,t)/A^s)$ profiles as a function of 3 4 normalized radius (r/R_p) for a particle of diameter $R_p = 0.05 \mu m$ for different values of bulk phase diffusivity and first-order reaction rate constants: (a) $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$, $k_c = 0 \text{ s}^{-1}$; (b) 5 $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}, \ k_c = 5 \text{ x} \ 10^{-4} \text{ s}^{-1}; \ \text{(c)} \ D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}, \ k_c = 0 \text{ s}^{-1}; \ \text{and} \ \text{(d)} \ D_b = 10^{-15} \text{ cm}^2$ 6 s^{-1} , $k_c = 5 \times 10^{-4} s^{-1}$. 7

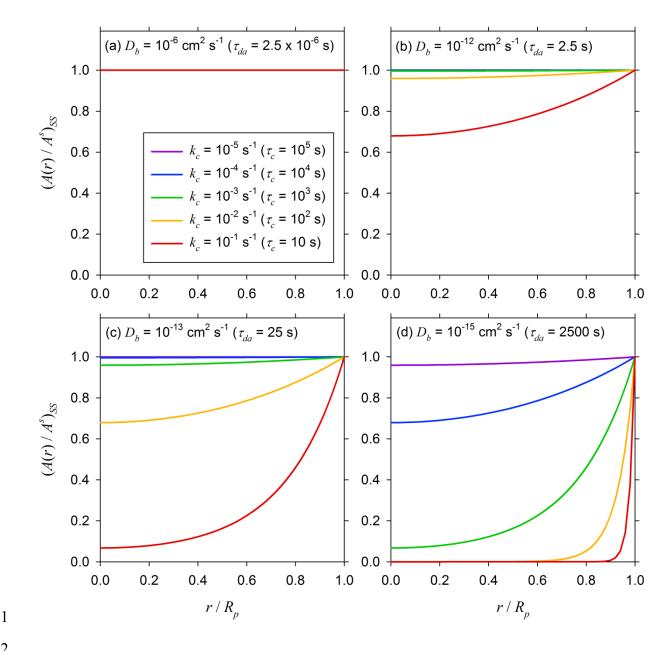


Figure 3. Normalized steady-state concentration $(A(r)/A^s)_{ss}$ profiles as a function of normalized radius (r/R_p) for a particle of diameter $R_p = 0.05 \ \mu m$ and a range of k_c values for (a) $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ (b) $D_b = 10^{-12} \text{ cm}^2 \text{ s}^{-1}$ (c) $D_b = 10^{-13} \text{ cm}^2 \text{ s}^{-1}$ (d) $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$.

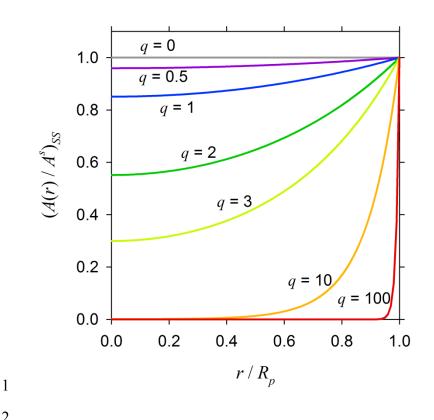
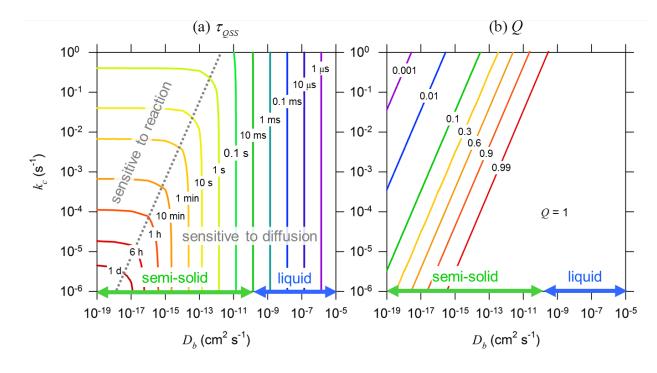


Figure 4. Normalized steady state concentration $(A(r)/A^s)_{ss}$ profiles as a function of normalized radius (r/R_p) for different values of dimensionless diffuso-reactive parameter q.



2

Figure 5. (a) Contour plots of: (a) particle-phase quasi-steady state timescale (τ_{QSS}), and (b) quasi-steady state parameter $Q = (\overline{A} / A^s)_{QSS}$ as functions of first-order rate constant (k_c) and bulk diffusion coefficient (D_b) for a species diffusing and reacting within semi-solid and liquid particles of diameter $D_p = 0.1 \,\mu\text{m}$.

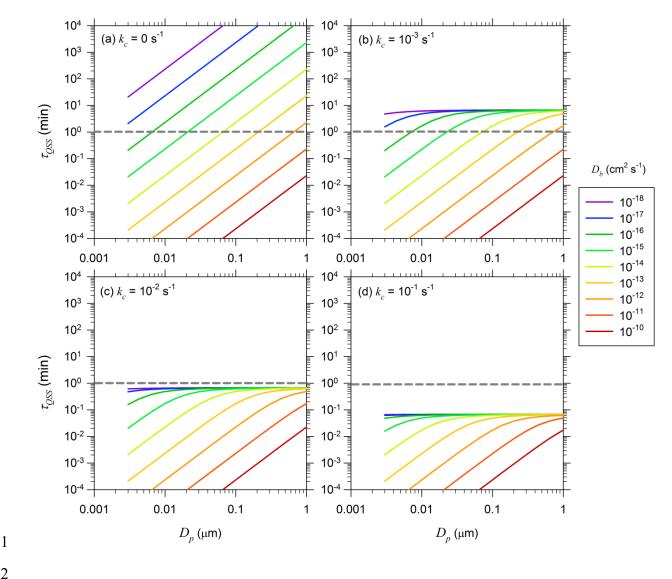


Figure 6. Dependence of τ_{QSS} on particle diameter D_p for D_b values ranging from 10^{-10} to 10^{-10} ¹⁸ cm² s⁻¹: (a) $k_c = 0$ s⁻¹; (b) $k_c = 10^{-3}$ s⁻¹; (c) $k_c = 10^{-2}$ s⁻¹; and (d) $k_c = 10^{-1}$ s⁻¹.

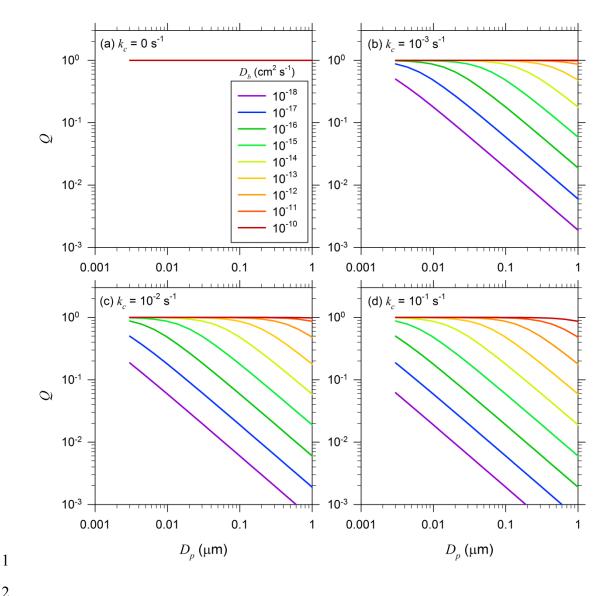
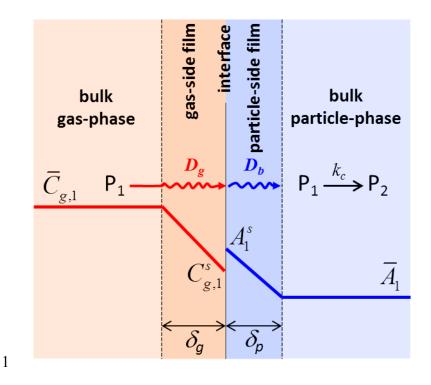
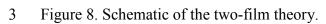


Figure 7. Dependence of Q on particle diameter D_p for D_b values ranging from 10^{-10} to 10^{-18} cm² s⁻¹: (a) $k_c = 0$ s⁻¹; (b) $k_c = 10^{-3}$ s⁻¹; (c) $k_c = 10^{-2}$ s⁻¹; and (d) $k_c = 10^{-1}$ s⁻¹.





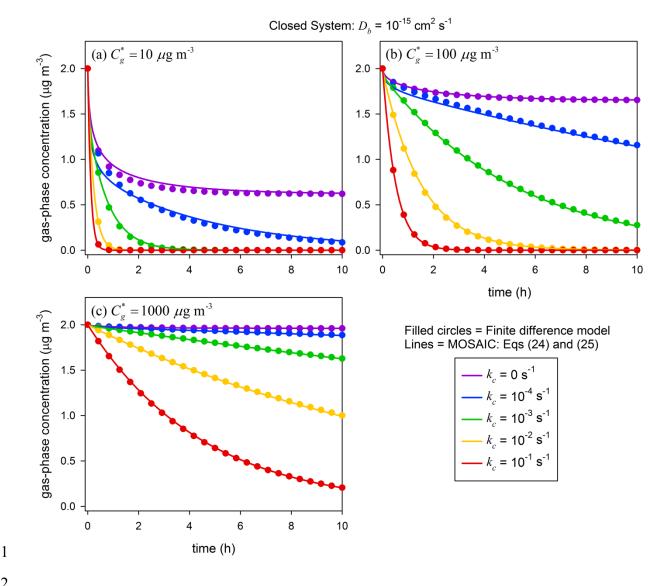
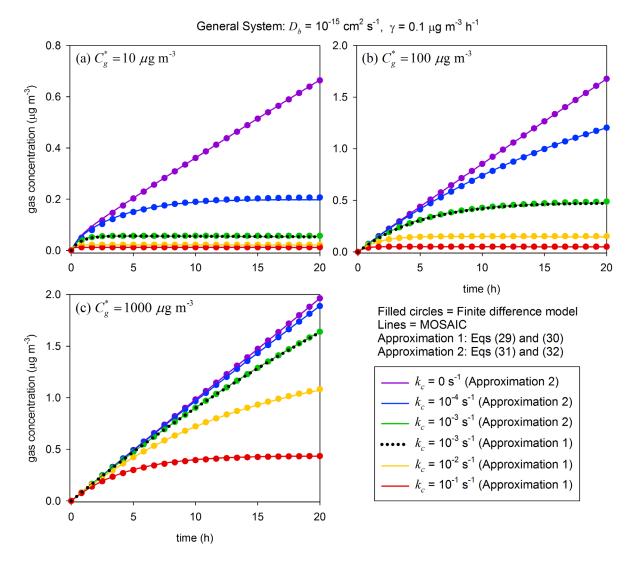




Figure 9. Comparison of MOSAIC (lines) and finite difference model (filled circles) solutions for gas-phase concentration decay in a closed system due to kinetic gas-particle partitioning to particles with initial $D_p = 0.2 \ \mu\text{m}$, $N = 5000 \ \text{cm}^{-3}$, $D_b = 10^{-15} \ \text{cm}^2 \ \text{s}^{-1}$ and k_c ranging from 0 to 0.1 s⁻¹ for three solute volatilities: (a) $C_g^* = 10 \ \mu g \ m^{-3}$, (b) $C_g^* = 100 \ \mu g \ m^{-3}$, and (c) $C_g^* =$ $\mu g\ m^{\text{-3}}.$



2

Figure 10. Comparison of MOSAIC (lines) and finite difference model (filled circles) solutions for gas-phase concentration evolution in a general system due to kinetic gas-particle partitioning to particles with initial $D_p = 0.2 \,\mu\text{m}$, $N = 5000 \,\text{cm}^{-3}$, $D_b = 10^{-15} \,\text{cm}^2 \,\text{s}^{-1}$, $\gamma = 0.1$ $\mu\text{g m}^{-3} \,\text{h}^{-1}$, and k_c ranging from 0 to 0.1 s⁻¹ for three solute volatilities: (a) $C_g^* = 10 \,\mu\text{g m}^{-3}$, (b) $C_g^* = 100 \,\mu\text{g m}^{-3}$, and (c) $C_g^* = 1000 \,\mu\text{g m}^{-3}$.

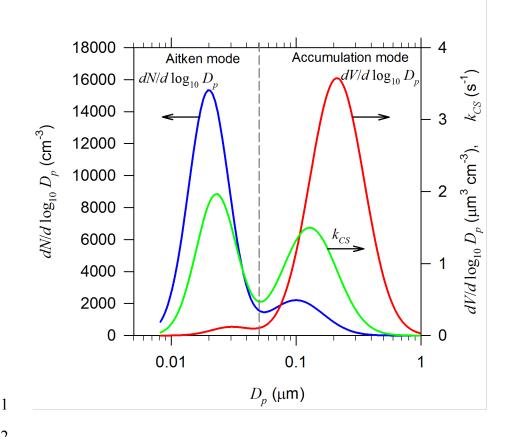
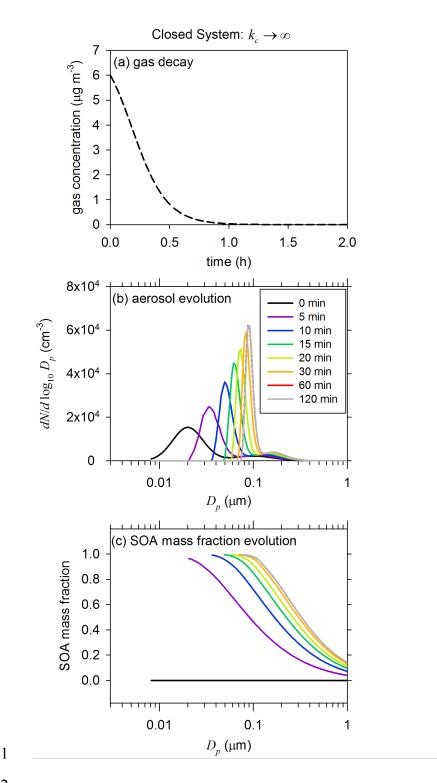


Figure 11. Initial aerosol number and volume size distributions along with the condensational sink k_{CS} . The dashed line demarcates the Aitken mode from the accumulation mode and the initial condensation sink is such that the sum of k_{CS} over all the size bins in the Aitken mode is equal to that in the accumulation mode.



2

Figure 12. Results for the instantaneous reaction reference case ($k_c \rightarrow \infty$; equivalent to nonvolatile solute condensation): (a) Gas-phase concentration decay; (b) temporal evolution of aerosol size distribution; and (c) temporal evolution of the mass fraction of newly formed SOA.

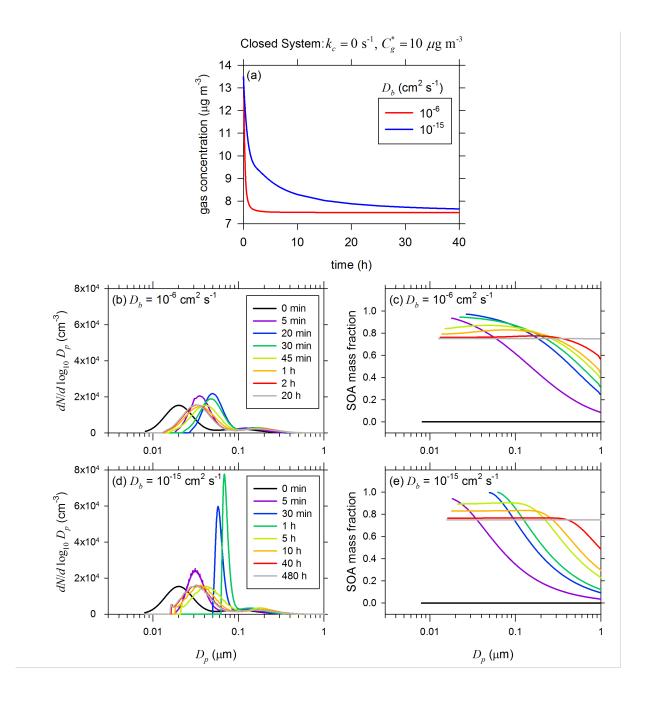
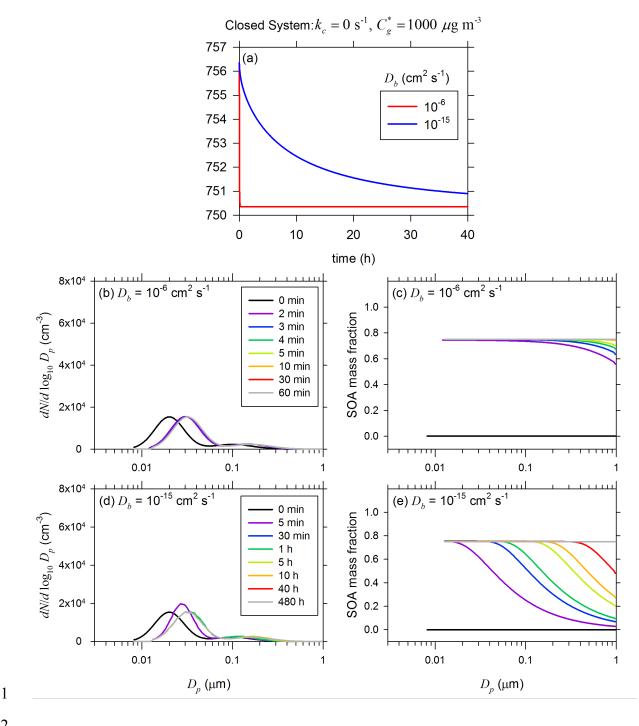




Figure 13. Results for kinetic SOA partitioning due to Raoult's law ($k_c = 0 \text{ s}^{-1}$) for $C_g^* = 10$ $\mu \text{g m}^{-3}$: (a) Gas-phase concentration decay for $D_b = 10^{-6}$ and $10^{-15} \text{ cm}^2 \text{ s}^{-1}$, (b) aerosol evolution for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$; (c) SOA mass fraction evolution for $D_b = 10^{-6} \text{ cm}^2 \text{ s}^{-1}$; (d) aerosol evolution for $D_b = 10^{-15} \text{ cm}^2 \text{ s}^{-1}$; and (e) SOA mass fraction evolution for $D_b = 10^{-15}$ $m^2 \text{ s}^{-1}$. In both cases, the final (i.e., equilibrium) concentration of the newly formed SOA is 6 $\mu \text{g m}^{-3}$.



3 Figure 14. Same as Fig. 13, except $C_g^* = 1000 \ \mu g \ m^{-3}$.

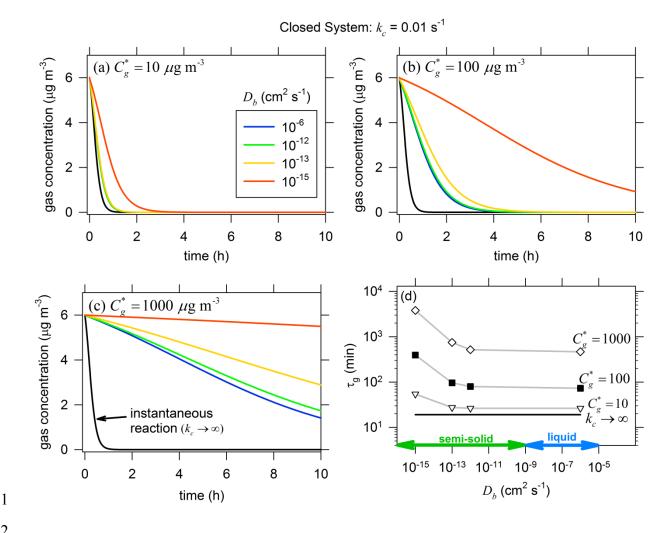
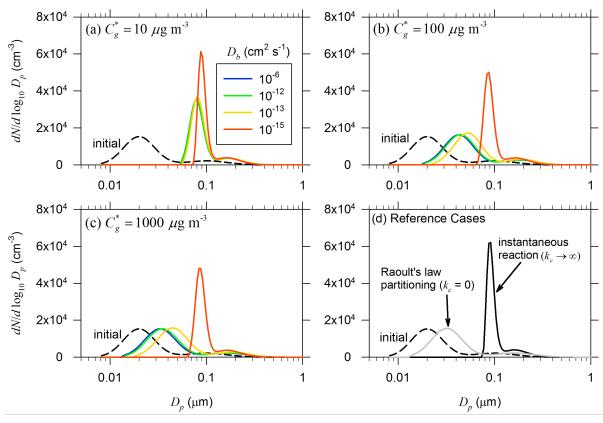




Figure 15. Gas-phase concentration decay due to kinetic SOA partitioning with particle-phase 3 reaction ($k_c = 0.01 \text{ s}^{-1}$) for bulk diffusivities ranging from 10^{-6} to $10^{-15} \text{ cm}^2 \text{ s}^{-1}$ and three gas 4 volatilities: (a) $C_g^* = 10 \ \mu g \ m^{-3}$; (b) $C_g^* = 100 \ \mu g \ m^{-3}$; and (c) $C_g^* = 1000 \ \mu g \ m^{-3}$. Each plot 5 6 also shows gas-phase concentration decay for the reference case of instantaneous reaction (black line, $k_c \rightarrow \infty$). In each case, the final concentration of the newly formed SOA is 6 µg 7 m⁻³. Panel (d) shows the plot of gas-phase concentration decay timescale (τ_g) as a function of 8 9 D_b for the different gas volatilities.



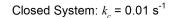


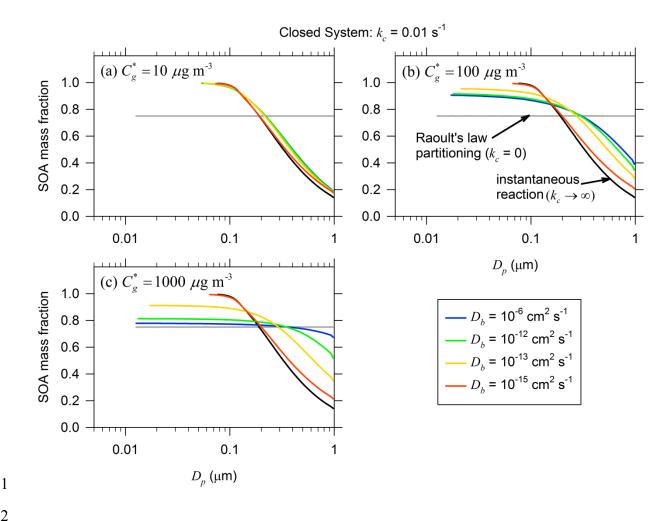
Figure 16. Initial (dashed line) and final (solid lines) aerosol number size distribution due to Raoult's law gas-particle partitioning coupled with particle-phase reaction ($k_c = 0.01 \text{ s}^{-1}$) for bulk diffusivities ranging from 10⁻⁶ to 10⁻¹⁵ cm² s⁻¹ and three gas volatilities: (a) $C_g^* = 10 \ \mu\text{g}$ m⁻³; (b) $C_g^* = 100 \ \mu\text{g}$ m⁻³; and (c) $C_g^* = 1000 \ \mu\text{g}$ m⁻³. Panel (d) shows the final size distributions for the two reference cases: instantaneous reaction (black line; $k_c \rightarrow \infty$) and Raoult's law partitioning (gray line; $k_c = 0$) for any D_b and $C_g^* > 0$. As illustrated in Fig. 15, the time required to reach the final state differs significantly for different cases, but the final

11

10

SOA formed in each case is $6 \ \mu g \ m^{-3}$.

12





3 Figure 17. Final size distributions of the newly formed SOA mass fraction for different D_b values and: (a) $C_g^* = 10 \ \mu \text{g m}^{-3}$; (b) $C_g^* = 100 \ \mu \text{g m}^{-3}$; and (c) $C_g^* = 1000 \ \mu \text{g m}^{-3}$. Each panel 4 also shows the reference plots for instantaneous reaction (black line; $k_c \rightarrow \infty$) and for 5 Raoult's law partitioning (gray line; $k_c = 0 \text{ s}^{-1}$) for any D_b and $C_g^* > 0$. 6

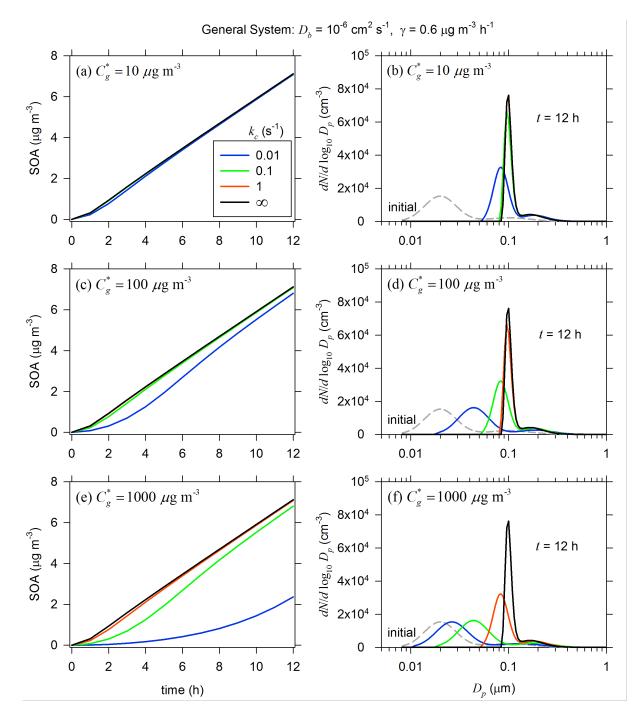




Figure 18. Temporal evolution of total SOA mass concentration (left column) and aerosol size distribution (right column) at t = 12 h for $D_b = 10^{-6}$ cm² s⁻¹, $\gamma = 0.6 \ \mu g \ m^{-3} \ h^{-1}$, $k_c = 0.01$ to ∞ s⁻¹, and three different solute volatilities: (a, b) $C_g^* = 10 \ \mu g \ m^{-3}$; (c, d) $C_g^* = 100 \ \mu g \ m^{-3}$; and (e, f) $C_g^* = 1000 \ \mu g \ m^{-3}$.

