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# New approach to monitor transboundary particulate pollution over northeast Asia

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#### Abstract

A new approach to more accurately monitor and evaluate transboundary particulate matter (PM) pollution is introduced based on aerosol optical products from Korea's geostationary ocean color imager (GOCI). The area studied is northeast Asia includ-<sup>5</sup> ing eastern parts of China, the Korean peninsula and Japan, where GOCI has been monitoring since June 2010. The hourly multi-spectral aerosol optical data that were retrieved from GOCI sensor onboard geostationary satellite COMS (Communication, Ocean, and Meteorology Satellite) through Yonsei aerosol retrieval algorithm were first presented and used in this study. The GOCI-retrieved aerosol optical data are inte-<sup>10</sup> grated with estimated aerosol distributions from US EPA Models-3/CMAQ v4.5.1 model simulations via data assimilation technique, thereby making the aerosol data spatially continuous and available even for cloud contamination cells. The assimilated aerosol optical data are utilized to provide quantitative estimates of transboundary PM pollution from China to the Korean peninsula and Japan. For the period of 1 April to 31 May 2011

- this analysis yields estimates that AOD as a proxy for surface-level PM<sub>2.5</sub> or PM<sub>10</sub> during long-range transport events increased by 117–265 % compared to background average AOD at the four AERONET sites in Korea, and average AOD increases of 121 % were found when averaged over the entire Korean peninsula. The paper demonstrates that the use of multi-spectral AOD retrievals from geostationary satellites can improve astimates of transpoundary PM pollution. Such data will become more widely avail.
- estimates of transboundary PM pollution. Such data will become more widely available later this decade when new sensors such as GEMS (Geostationary Environment Monitoring Spectrometer) and GOCI-2 are scheduled to be launched.

#### 1 Introduction

Transboundary air pollution is an important issue in Asia (Arndt et al., 1998; Holloway et al., 2002; Yamaji et al., 2012). In particular, the long-range transport (LRT) of air pollutants from China (source) to the Korean peninsula and Japan (receptors) due



to persistent westerlies over northeast Asia has been a central issue in this region (Arndt et al., 1998; Song and Carmichael, 2001; Holloway et al., 2002; Yamaji et al., 2012). Transboundary air pollution is a serious issue not only in Asia, but in Europe and North America as well, causing a number of different international problems such as
acid precipitation, rising levels of hemispheric ground-level ozone, and regional visibility impairment (Levy II and Moxim, 1987; Arndt et al., 1998; Malcolm et al., 2000; Holloway et al., 2002, 2003; Brankov et al., 2003; Tulet et al., 2003; Park et al., 2004; Schaub et al., 2005; Wang et al., 2009; Yamaji et al., 2012). To manage transboundary air pollution between countries, the UNECE (United Nations Economic Commission for
Europe) under the Convention on Long-Range Transboundary Air Pollution (CLRTAP)

was launched to lessen air pollution through the long-range transport (LRT) in 1979 (UNECE, 2012). The EMEP (European Monitoring and Evaluation Programme) is a scientific effort established to support the CLRTAP.

Quantitative evaluation of transboundary air pollution remains a difficult task and
 <sup>15</sup> currently cannot be done based on observations alone due to a lack of explicit source signatures and lack of an appropriate observation system (Levy II and Moxim, 1987; Arndt et al., 1998; Malcolm et al., 2000; Holloway et al., 2002, 2003; Brankov et al., 2003; Tulet et al., 2003; Park et al., 2004; Schaub et al., 2005; Wang et al., 2009; UN-ECE, 2012; Yamaji et al., 2012). Most current evaluations of transboundary transport
 of pollutants are based on chemistry-transport model (CTM) simulations (Levy II and

- Moxim, 1987; Arndt et al., 1998; Malcolm et al., 2000; Holloway et al., 2002; Brankov et al., 2003; Tulet et al., 2003; Park et al., 2004; Schaub et al., 2005; Wang et al., 2009; Yamaji et al., 2012). However, model-based evaluations have significant limitations because of the large uncertainties in the emission rates of the air pollutants and
- <sup>25</sup> uncertainties related to the CTM and meteorological models (MM). For example, both atmospheric secondary particle formation and primary particle (e.g., mineral dust) generation/transport are difficult to correctly simulate (Song and Carmichael, 2001; Song et al., 2005; Volkamer et al., 2006; Robinson et al., 2007). These uncertainties are



a main reason for disagreement in source and receptor estimates made by different countries, which has led to distrust in the outcomes of model-based assessments.

Current efforts to reduce the uncertainty in estimates of long-range transport of pollutants by integrating models and observations through data assimilation are limited by

- <sup>5</sup> our current observational systems. Surface-based observations are too sparse to constrain the system, while current satellites are unable to retrieve aerosol concentrations. Instead, satellite retrievals of AOD are used to help constrain modeled AOD and the modeled relationship between AOD and surface-level PM (particulate matter) can then be used to estimate surface-level PM concentrations (Wang and Christopher, 2003; Liu
- et al., 2005; van Donkelaar et al., 2006; Schaap et al., 2009; Tsai et al., 2011). This approach is currently limited by the fact that AOD from polar-orbiting satellites produce retrievals only once per day (Kaufman et al., 1997; Diner et al., 2001) and current geo-stationary satellites lack "multi-spectral" aerosol data (Knapp et al., 2002; Wang et al., 2003). For example, the USA GOES series (GOES 8, 13), the Japanese GMS and MTSAT, and European SEVIRI have been meteorological sensors, with only one to two visible channels. The history of GEO satellite sensors for the retrieval of the aerosol
  - optical properties is summarized in Table 1.

This situation is changing as new geostationary satellite based sensors are becoming available and will be expanded by upcoming planned missions. The first hourly

- <sup>20</sup> multi-spectral aerosol data observed from a geostationary satellite sensor, GOCI (Geostationary Ocean Color Imager) is used in this paper to demonstrate how this new information can be used to improve our estimates of LRT of pollutants. Specifically, the level-1 radiance data from GOCI, onboard COMS (Communication, Ocean, and Meteorology Satellite), is converted into level-2 AOD products via the "Yonsei aerosol retrieval
- <sup>25</sup> algorithm" (the details of the algorithm will be discussed in the Methods section) (Lee et al., 2010). COMS/GOCI was launched in June 2010, and the COMS/GOCI data center began to release radiance data from April 2011, after an initial operation test (IOT) period. This satellite sensor was designed for monitoring ocean colors around the Korean peninsula, but it can also provide AOD data over a strategically important



region, i.e., northeast Asia. The domain covered by the GOCI sensor shown in Fig. 1 is  $2500 \text{ km} \times 2500 \text{ km}$  and includes eastern parts of China, the Korean peninsula, and Japan. The GOCI-retrieved AOD products were integrated with CMAQ-simulated AOD data to produce the more accurate aerosol data set used to investigate the impacts of transboundary PM pollution between the source and receptor countries in northeast Asia.

#### 2 Methods

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The procedure to produce GOCI-constrained AOD product combines the GOCI level-2 product with AOD predictions from the US EPA Model-3/CMAQ (Community Multiscale Air Quality) v4.5.1 model simulations. The CMAQ model simulations were per-10 formed over east Asia, with emission data prepared appropriately for the model domain and meteorological fields produced from the PSU/NCAR MM5 (the fifth generation Meteorological Model) simulations. The final AOD product is spatially continuous and available even for locations where GOCI retrievals are not available (e.g. cloud-contaminated points, sun-glint areas, desert, and snow-covered land). The use 15 of GOCI-retrievals overcomes the "temporal limitation" typical in analyses with LEO aerosol data and the assimilation overcomes the "spatial limitation" that has been typical in analyses based only on satellite data. The assimilated AOD data set using the GOCI data is utilized to monitor and evaluate transboundary particulate pollutions from China to the Korean peninsula in this study. 20

#### 2.1 Meteorology and chemistry-transport model simulations

The PSU/NCAR MM5 and US EPA Model-3/CMAQ model v4.5.1 were respectively used as the Meteorological Model (MM) and Chemistry-Transport Model (CTM) in this study (Byun and Schere, 2006; Park et al., 2011). The accuracy of the MM5 simulations was enhanced by four-dimensional data assimilation (FDDA) using the ob-



served meteorological data. Initial conditions (ICs) and boundary conditions (BCs) for the MM5 simulations were updated with the National Centers for Environmental Prediction (NCEP) reanalysis data. The latest remotely sensed surface wind data retrieved from the Advanced SCATterometer (ASCAT), which was launched on 19 October 2006

onboard the Meteorological Operational satellite-A (MetOp-A) was used for the Four Dimensional Data Assimilation (FDDA). Although it has limitations in retrieving high wind speeds, it is believed to be the most suitable measurement data for improving the meteorological modeling results (Bentanmy et al., 2008).

The schemes chosen for CMAQ model simulations were the same as those used

- in previous research (Park et al., 2011). The SAPRC-99 chemical mechanism and AERO4 aerosol module were selected for simulating gas-phase chemistry and atmospheric aerosol thermodynamic/dynamic processes, respectively. More details of the MM5 and CMAQ model simulations were described in a previous publication (Park et al., 2011).
- The CMAQ model simulations were carried out using the best available emission inventories over East Asia that have been sophisticated with, for example, the applications of monthly variations of anthropogenic emissions, the uses of an operational dust forecast model at the KMA (Korea Meteorological Administration) and the NCAR (National Center for Atmospheric Research) FINN (Fire Inventory from NCAR) biomass
- <sup>20</sup> burning emissions (Wiedinmyer et al., 2011). The efforts are summarized in Table 2. For the consideration of anthropogenic pollutants emissions, such as NO<sub>x</sub>, SO<sub>2</sub>, CO, black carbon, NMVOCs, etc, Intercontinental Chemical Transport Experiment-Phase B (INTEX-B), Regional Emission inventory in Asia (REAS), and Clean Air Policy Support System (CAPSS) were applied as the best suitable emission data over China, Japan,
- and the Korean peninsula, respectively. These emission inventories originally adjusted to the year 2006 were updated for the year 2011 modeling period, using the projection factors based on the Greenhouse gas and Air pollution INteractions and Synergies (GAINS) modeling emissions scenarios (http://gains.iiasa.ac.at/index.php/gains-asia). In the meanwhile, in order to consider isoprene and mono-terpene emissions over



the modeling domain, the isoprene emissions from Model of Emissions of Gases and Aerosols from Nature (MEGAN) and MOdel for Hydrocarbon emissions by the CANopy (MOHYCAN) were utilized for the CMAQ modeling input data (Müller et al., 2008). Mono-terpene emissions were also estimated from the isoprene emissions, using the

- ratios of mono-terpene to isoprene from Global Emissions Inventory Activity (GEIA). In particular, this modeling period (a spring season in northeast Asia) could not be explained without considering the effects of biomass burning and mineral dust emissions. FINNv1 is a biomass burning (BB) emission inventory annually provided from the year 2002 by NCAR (http://bai.acd.ucar.edu/Data/fire/). This FINNv1 has produced
- the BB emission data with a high horizontal resolution of ~ 1 km × 1 km over the entire globe, being applied for variable regions from local to global scales (Wiedinmyer et al., 2011). In this study, the FINNv1 was used for the BB emissions. The CMAQ v4.5.1 model does not include any dust modeling modules related to occurrence and transport of mineral dust events. However, without taking the influence of mineral dust events into
- account, accurate estimation of the aerosols effects observed by long-range transports in the Korean peninsula and Japan would not be possible. Therefore, the outbreak and transport of Asian mineral dust were considered by Asian Dust Aerosol Model (ADAM) simulations. The ADAM is currently operational model for forecasting Asian mineral dust events at the KMA and the details of the ADAM simulations were reported previ ously by e.g., Park and Lee (2004), Kim et al. (2011), and Park et al. (2011).

The domain of the CMAQ model simulation is also shown in Fig. 1. The CMAQ domain (blue color) covered the region of approximately 95–145° E, 20–50° N, which includes COMS/GOCI domain (red color) composed of northeast China, the Korean peninsula, and Japan. In this study, the CMAQ domain was chosen to be larger than

<sup>25</sup> the GOCI domain, since the air quality inside the GOCI domain is strongly influenced by the influxes from the outer areas with persistent westerlies in East Asia. The horizontal grid spacing of the CMAQ modeling was  $30 \text{ km} \times 30 \text{ km}$  and the vertical domain was distributed in 14 layers following  $\sigma$ -coordinate from 1000 hPa to 70 hPa. The MM5-CMAQ simulation period started at 00:00 UTC, 1 April to 23:00 UTC, 31 May 2011



(with a week-long CMAQ spin-up period), since the first batch of the level-1 radiance data from the COMS/GOCI data center was released for this period.

#### 2.2 Remote-sensing data

## 2.2.1 COMS/GOCI data

- Sensors onboard low earth orbit (LEO) satellites, such as MODIS (Moderate Reso-5 lution Imaging Spectroradiometer) and MISR (Multi-angle Imaging Spectroradiometer), have ordinarily been utilized to obtain information on aerosol optical properties (AOPs), but their monitoring has taken place at most once per day. To more accurately monitor the transboundary LRT events of air pollutants, hourly (or highly temporallyresolved) monitoring data from geostationary earth orbit (GEO) sensors are needed 10 instead of the temporally-limited LEO measured data. As mentioned previously and also presented in Table 1, MSG, GOES, and MTSAT-1R sensors have been partly monitoring the AOPs over Europe, US, and Japan, respectively (Knapp et al., 2005; Popp et al., 2007; Kim et al., 2008), but these meteorological sensors have equipped with only one visible channel, which possibly makes the quality of the AOPs from 15 these meteorological sensors less accurate. In contrast, GOCI sensor onboard geostationary COMS has six-visible bands at 412, 443, 490, 555, 660, 680 nm, which can make advances by more accurately calculating the AOPs, compared to other
- meteorological GEO satellite sensors. In addition, atmospheric compensations, such
   as cloud-screening make the numerical accuracy of the AOPs even more improved,
   because the COMS/GOCI has 2 more bands at 745 nm and 865 nm (i.e., near IR).
   The GOCI sensor covers northeast China, the Korean peninsula, and Japan, and the
   products can be generated with 500 m × 500 m spatial resolution. GOCI level-1 data
   during daytime can be obtained 8 times from 09:00 a.m. to 04:00 p.m. LST (refer to
   http://kosc.kordi.re.kr/oceansatellite/coms-goci/introduction.kosc).

The hourly GOCI level-1 data are released somewhat slowly after the IOT and data calibration procedures. The GOCI level-1 radiance are then converted into level-2 AOD



data using a multi-spectral algorithm being developed by Jhoon Kim and coworkers at Yonsei University, Korea (Lee et al., 2010). The level-2 AOD data retrieved with a pixel size of  $6 \text{ km} \times 6 \text{ km}$  was used in this study. The retrieval algorithm was originally based on the NASA MODIS (i.e., LEO) algorithm, but it has evolved to date to include

- the aerosol models optimized for East Asia and to consider the geometry from GEO satellite instruments such as GOCI, GOCI-2, and GEMS (Geostationary Environment Monitoring Spectrometer) (both the GOCI-2 and GEMS sensors are scheduled to be launched in 2018). The Yonsei aerosol retrieval algorithm includes the capability to classify aerosol types based on absorption tests and size information (Kim et al., 2007;
- Lee et al., 2010). The evaluation of the current AOD data from the algorithm shows that the slope of the linear regression line is close to 1.03 over the ocean, but 1.15 over land when compared to values from AERONET. Although the detailed explanation and validation procedures of this Yonsei aerosol retrieval algorithm were given in previous publications (Lee et al., 2010, 2012), a brief introduction on the algorithm is made
- here. The retrievals of the AOPs were performed with top of atmosphere spectral reflectances. To obtain the AOPs over the regions including clouds, clouds are screened by spatial variability and threshold tests at each pixel. Even so, there are limitations in completely removing clouds and in perfectly retrieving surface reflectance, due to the lack of IR bands on GOCI. Therefore, the Yonsei aerosol retrieval algorithm eliminates
   clouds, utilizing Meteorological Imager (MI) that has 4 IR channels. The MI sensor is
- 20 Clouds, utilizing Meteorological imager (MI) that has 4 IR channels. The Mi sensor is also onboard COMS with GOCI. In addition, the Yonsei aerosol retrieval algorithm is composed of 2 different retrieval logics that are applying for "continent" including near coastal turbid water areas and "ocean", respectively.

Figure 2 presents one example: a comparison of the Terra/MODIS-retrieved and Aqua/MODIS-retrieved AODs with GOCI-retrieved AODs approximately at 10:00 LST and 13:00 LST on 10 April and 12 April, respectively. As shown in Fig. 2, there are agreements in the spatial AOD distributions from the 2 MODIS and GOCI sensors. As discussed previously, at most 2 snapshots of information per day from the LEO sensors (i.e. MODIS) can be obtained, whereas 8 snapshots per day from the GEO sensor (i.e.



GOCI) are available. This is an obvious "temporal" advantage that can be achieved from the GEO sensor. Furthermore, as shown in Fig. 2, the AOD data from the GOCI sensor shows a better spatial coverage, compared to those from the MODIS (see white bands along the north-south swath in the MODIS data). This lack of the MODIS data
<sup>5</sup> was caused by sun-glint over the ocean areas. But, such lack of the data is not found in the GOCI data. This is another "spatial" advantage that can be achieved from the GEO sensor.

#### 2.2.2 AERONET

Many studies have used the ground-based NASA AERONET data to verify modeling
 results and satellite monitoring data (e.g. Chin et al., 2002, 2004; Yumimoto and Takemura, 2011). The AERONET data measured by sun-photometers were utilized in this study to investigate the accuracy of CMAQ model simulations and GOCI AOP data. In this manuscript, the data obtained from 5 AERONET observation sites (i.e. Baengnyeong, Yonsei Univ., Gwangju GIST, Gosan, and Osaka) in the Korean penin sula and Japan was selected to analyze long-range transport effects over northeast Asia. The AERONET AOD v2.0 data used for comparison was obtained from the official NASA web site at http://aeronet.gsfc.nasa.gov. The only exception was the AERONET AOD data for the Yonsei Univ. (Seoul) site, which were not calibrated. Before use they were upgraded by eliminating some sporadic cloud contamination of the initial

high AOD data (AOD > 1.0) using the meteorological data from the Automatic Weather Stations (AWS) co-located with cimel sun-photometers. The uncertainty and possible anomalies in the cimel sun-photometer AOD data at the Yonsei Univ. site were also checked by the Laboratory for Terrestrial Physics at NASA Goddard Space Flight Center. The AOD uncertainties of 0.01 to 0.04 were estimated, mainly due to calibration.



#### 2.3 AOD calculations from CMAQ model simulations

Concentrations of nitrate, sulfate, ammonium, organic aerosols (OAs), black carbon (BC), sea-salt aerosols, and mineral dust are provided as results of the CMAQ model simulations. This aerosol composition can be converted into the AOPs including AOD by considering the mass extinction efficiencies (MEEs) depending on the scattering and absorbing characteristics for each aerosol composition and the hygroscopicity. First, the absorption, scattering, and extinction coefficients ( $\sigma_{abs}$ ,  $\sigma_{scat}$ , and  $\sigma_{ext}$ ) were calculated by the following formulas:

$$\sigma_{\text{ext}}(\lambda) = \sigma_{\text{abs}}(\lambda) + \sigma_{\text{scat}}(\lambda)$$

10

5

$$\sigma_{\rm abs}\left(\lambda:{\rm Mm}^{-1}\right) = \sum_{i}^{n} \sum_{j}^{m} \left\{ (1-\omega_{ij}(\lambda)) \cdot \beta_{ij}(\lambda) \cdot f_{ij}({\rm RH}) \cdot [C]_{ij} \right\}$$

$$\sigma_{\text{scat}}\left(\lambda:\text{Mm}^{-1}\right) = \sum_{i}^{n} \sum_{j}^{m} \{\omega_{ij}(\lambda) \cdot \beta_{ij}(\lambda) \cdot f_{ij}(\text{RH}) \cdot [C]_{ij}\}$$
(3)

where, *i* and *j* represent the particulate species *i* at the size-bin or particle-mode *j*;  $f_{ij}(RH)$  the hygroscopic enhancement factor as a function of relative humidity (RH); and  $\omega_{ij}(\lambda)$  and  $\beta_{ij}(\lambda)$  the single scattering albedo and MEE of particulate species *i* at the size-bin or particle-mode *j*, respectively. The terms related to the MEEs ( $\beta_{ij}$ ) and the hygroscopicity ( $f_{ij}(RH)$ ) have been differently defined by several studies, e.g., d'Almeida et al. (1991), Malm et al. (1994), Chin et al. (2002), Malm and Hand (2007), and Park et al. (2011). According to the previous investigations, the way to compute the extinction coefficients ( $\sigma_{ext}$ ) has been mainly classified into 2 categories. The first method is to use empirical equations (e.g. Malm et al., 1994; Malm and Hand, 2007;



(1)

(2)

Song et al., 2008; Park et al., 2011). The other method is to use the equations obtained based on the Mie theory (e.g. d'Almeida et al., 1991; Chin et al., 2002).

In this study, based on the particulate composition and meteorological parameters obtained from the MM5-CMAQ model simulations, the AOD was calculated at a wave-<sup>5</sup> length of 550 nm (AOD<sub>550</sub>), using the conversion algorithms of Chin et al. (2002), Park et al. (2011), and others (d'Almeida et al., 1991; Malm et al., 1994; Hess et al., 1998; Malm and Hand, 2007):

$$AOD_{550} = \int_{0}^{z} \sigma_{\text{ext}} (550 \,\text{nm}) \,\text{d}z$$

where *z* represents the vertical height. For example, the CMAQ-simulated AOD products calculated from Park et al. (2011) usually showed underestimations, when compared to the AERONET AOD data. However, when Chin et al. (2002)'s conversion algorithm was applied, the AOD values were usually larger than those calculated by Park et al. (2011)'s method. These results will be discussed further in a later section.

#### 2.4 Data assimilation

- In several previous studies, optimal interpolation (OI) technique with a Kalman gain matrix has been employed to improve the predictions of AOD and aerosol mass distributions with the combination of the modeling results and monitoring data (Collins et al., 2003; Yu et al., 2003; Adhikary et al., 2008; Chung et al., 2010; Park et al., 2011). For examples, Adhikary et al. (2008) showed some enhancements in the estimations of the AOD values over Asia by combining the STEM modeled and MODIS retrieved data
- the AOD values over Asia by combining the STEM-modeled and MODIS-retrieved data through the OI technique. The accuracy of the AOD values was also greatly improved after the application of the OI method, using CMAQ-simulated and MODIS-retrieved AOD data (AOD<sub>m</sub> and AOD<sub>o</sub> in Eq. 5) (Park et al., 2011).

$$AOD_{m}^{'} = AOD_{m} + K[AOD_{o} - H(AOD_{m})]$$

(4)

(5)

## $K = BH^T (HBH^T + O)^{-1}$

where, H denotes a linear operator for interpolation from the modeling grid to the satellite-retrieving grid.

- In this study, the OI technique was again utilized for the assimilation of the AOD data from the CMAQ model simulations and the GOCI sensor. The Kalman gain matrix (*K* in Eq. 6) was obtained by the covariance of modeled and observed fields (*B* and *O* in Eq. 6) which is a function of free parameters. The yearly-fixed free parameters were mainly used in several previous studies (Collins et al., 2001; Yu et al., 2003; Adhikary et al., 2008; Chung et al., 2010), whereas Park et al. (2011) proposed to apply to seasonal
- variations of free parameters. Detail explanations for the OI technique including the selections of the free parameters were discussed in Park et al. (2011).

#### 2.5 Long-range transport of particulate matters

In this study, transboundary transports of the AOD plumes from the China continent to the Korean peninsula and Japan have been focused. However, defining the LRT events over northeast Asia is not easy, mainly due to complexities in air mass movements. We considered large-scale AOD plumes that caused the peak AOD values larger than 0.4 at the receptor sites (peak AOD filter). Such LRT AOD plumes were further traced back via NOAA HYSPLIT (HYbrid Single-Particle Lagrangian Integrated Trajectory) model analysis (Draxler and Hess, 1998) from the points of times when the peak AOD values

- occurred. The HYSPLIT model simulations were performed with meteorological data, which were archived from GDAS (Global Data Assimilation System) of NCEP (National Centers for Environmental Prediction). When the backward trajectories from the occurring points of the peak AOD values had passed through the China continent within
- 48–72 h, the AOD plumes were finally determined as LRT plumes (backward trajectory filter). A LRT case shown in Fig. 3 was decided through both the filters. Figure 4 shows an example of non-LRT case.



(6)

In order to investigate the LRT of particulate matters over northeast Asia, the AOD values instead of PM concentrations were used. In other words, the AOD has been regarded as a proxy quantity to PM concentration in this study. In this context, relations between PM<sub>10</sub> (or PM<sub>2.5</sub>) and AOD can be important. The relationships are affected <sup>5</sup> by various parameters, such as height of PBL (planetary boundary layer), RH, vertical distribution of aerosols, seasonality, diurnal variation, geographical information, particle compositions, and size distributions (Wang and Christopher, 2003; Engel-Cox et al., 2004; Liu et al., 2005; van Donkelaar et al., 2006; Choi et al., 2009; Schaap et al., 2009; Tsai et al., 2011). Several studies have estimated two dimensional PM<sub>10</sub> from the relationships between ground-based PM<sub>10</sub> and satellite-retrieved AOD (Wang and Christopher, 2003; Engel-Cox et al., 2004; Liu et al., 2009; Tsai et al., 2004; Liu et al., 2009; Tsai et al., 2004; Liu et al., 2009; Tsai et al., 2011).

#### 3 Results and discussions

<sup>15</sup> The hourly assimilated AOD data was used to identify and analyze the PM LRT events over northeast Asia. Based on this, the variations of the AOD values during LRT events in comparison with those during non-LRT events were analyzed on a more quantitative basis.

#### 3.1 Long-range transport of aerosols from China to Korea

A representative snapshot of a LRT event of particulate matter from China to the Korean peninsula, Japan, and the northwestern Pacific Ocean is presented in Fig. 3. It can be seen that a huge AOD plume (AOD > 0.6) was generated in the greater Beijing area and Shandong province on 10 April and was then transported over long-distances to the Korean peninsula following a low pressure trough. GOCI retrievals with one-hour intervals from 09:00 to 16:00 LST (8 snapshots per day) on 10 April show high aerosol



loadings (AOD > 1.2) over the Shandong peninsula and Shanghai regions that extend over Korea. The plumes extend farther out over the Yellow Sea and slowly shift to lower latitudes as time progresses. In Fig. 3, only two-hour resolution results are presented. It is important to note that GOCI AOD values are not retrieved in regions where clouds

- exist. As shown in Fig. 3, cloud-contaminated pixels appear with white color, which indicates that no data is available at these pixels. The GOCI retrievals when assimilated into the CMAQ model produce complete spatial distributions where AOD is estimated in all pixels (see the third column of Fig. 3). Therefore, the assimilated AOD fields are much improved over the predicted (a priori) CMAQ AOD distributions (first column of
- Fig. 3). The a priori CMAQ AOD values have a tendency to be underestimated, compared to the GOCI-retrieved AOD values, as shown in Fig. 3. This underestimation is believed to be related to uncertainties in the emission rates of gaseous and particulate pollutants and the CMAQ model simulations, as well as the uncertainties in the assumptions made during the AOD estimation from the CMAQ modeling results.
- In this analysis, the AOD calculation via Chin et al. (2002)'s conversion method was used. However, several other estimations were also tested (e.g. Malm et al., 1994; Chin et al., 2004; Park et al., 2011). One example is shown in Fig. 5 and Table 3, which were produced by Park et al. (2011)'s conversion method. It was found that the AOD values were underestimated, because the MEE values of Park et al. (2011)'s conversion method. It was found that the AOD values were underestimated, because the MEE values of Park et al. (2011)'s conversion method.
- sion method were reported to be in general smaller than those of Chin et al. (2002)'s method. After assimilation both the magnitudes and spatial distributions are significantly improved. The AOD LRT event shown in Fig. 3 is only an example. Seven to fourteen small- and large-scale AOD LRT events occurred over the two month period (from 1 April to 31 May 2011).
- <sup>25</sup> On the other hand, the non-LRT events were also identified. During the non-LRT periods, air masses were generally stationary. Although the influx of air pollutants emitted from the highly polluted China regions to the Korean peninsula and Japan was significantly reduced, the stagnant weather conditions could seriously deteriorate air quality. Figure 4 shows a case of the non-LRT events. Noticeable aerosol plumes di-



rectly transported from China to the Korean peninsula and Japan could not be identified. Some AOD plumes were generated in the southern parts of China, and were then long-distance transported over the northwestern Pacific Ocean. This case was classified as non-LRT case, since the AOD plumes were not directly transported to the 5 Korean peninsula and/or Japan.

In addition, a high AOD plume (in the second column in Fig. 4) was transported through southern Russia (between 09:00 and 11:00 LST). This plume appears to be thin clouds (i.e. cloud optical depth, COD). This can be confirmed partly via the Terra/MODIS AOD data at 10:30 on 12 April. The Terra/MODIS AOD data did not show any AOD plume over the same region (refer to the third and the fourth rows in Fig. 2). Since the GOCI AOD algorithm has limitation in completely eliminating clouds, thin clouds are sometimes considered as aerosol plumes. Another possibility for this plume is that aerosols could be generated through Siberian forest fires. In order to check

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this possibility, the global MOZART (Model for OZone And Related chemical Tracers)
 model simulation data were investigated (http://www.acd.ucar.edu/wrf-chem). But, no
 Siberian fire significantly affected aerosol plumes on/before 12 April 2011.

The AOD LRT events were investigated at the five AERONET sites (Holben et al., 1998) in the Korean peninsula and in Japan shown in Fig. 1. The AERONET AOD values observed at ground sites are regarded as "ground truth" because they are not interferent but the surface reflectance (Holber et al., 1000). Dub suit at al., 2000). Chin at

- interfered by the surface reflectance (Holben et al., 1998; Dubovik et al., 2000; Chin et al., 2002, 2004; Park et al., 2011), and due to the high accuracy (~0.01–0.02) of the fully calibrated data (Eck et al., 1999). Figure 6 shows the comparison of the GOCI-retrieved/assimilated AOD products (red horizontal bars and blue solid line in Fig. 6, respectively) with the AERONET AOD data (black bars in Fig. 6). The GOCI-retrieved
- <sup>25</sup> AOD values agree well with the AERONET values (e.g., R = 0.84 see Table 5). However both AERONET and GOCI data show large time periods with no retrievals, pointing out the need for models to provide additional information. The assimilated AOD product is well-correlated with the AERONET and GOCI-retrieved AOD values and is much more accurate than the a priori CMAQ AOD product when compared with the



AERONET AOD values. This situation is summarized in Table 5, where the errors and biases after the assimilation become much smaller than those before the assimilation, indicating that the assimilation clearly improves the quality of the AOD product. The assimilated product provides a continuous estimate of the time series of AOD at every location, which the observations alone cannot (the red horizontal bars in Fig. 6 only cover limited time periods).

#### 3.2 Estimation of the impacts of long-range transport

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The grey-shaded areas in Fig. 6 represent the periods during which the receptors were affected by the PM LRT events. The PM LRT events were determined based on the
 meteorological fields generated from MM5 simulations together with comprehensive backward trajectory analysis. For example, the Baengnyeong site was affected by the small- and large-scale LRT events fourteen times over the two month period. Most of the high (and wide) AOD peaks are due to the large-scale LRT events from China, and most of the small/sharp spikes are usually caused by local pollutions under stagnant
 meteorological conditions. The grey-shaded periods in Fig. 6 usually start and end, when the large-scale LRT AOD plumes arrive in and depart from the AERONET sites.

The background average AOD values calculated excluding the AOD values during the LRT events are presented by dashed lines at the five AERONET sites in Fig. 6. The average impact of the transboundary PM pollution over the entire two month period (i.e. over both "the non-LRT + LRT periods") was calculated, using the following formula over the two month period:

$$\overline{AOD}_{avg} = \frac{\sum_{i=1}^{N} \{\overline{AOD}_{LRT} \times x_{LRT} + \overline{AOD}_{non-LRT} \times (1 - x_{LRT})\}}{N}$$

$$\overline{\text{AOD}}_{\text{inc}} = \Delta \overline{\text{AOD}} = \overline{\text{AOD}}_{\text{LRT}} - \overline{\text{AOD}}_{\text{non-LRT}}$$
15884



(7)

(8)

where  $AOD_{avg}$ ,  $AOD_{non-LRT}$  and  $AOD_{LRT}$  represent the average AOD values during the entire period, the average AOD values calculated during the non-LRT and LRT periods, respectively.  $x_{LRT}$  is the fraction of the LRT events over the two month period. *N* is the total number of the grid cells being considered (for the calculation at an AERONET site,

 $^{5}$  N = 1). After that, the average AOD increase ( $\overline{AOD}_{inc}$ ) during the two month period was calculated by Eq. (8). The results for the analysis for the AERONET sites are summarized in Table 4. From this analysis, the average AOD increase of 117–265 % is found at the five AERONET sites.

The average AOD increases due to the transboundary PM pollution over the Korean peninsula were estimated in a similar manner. The Korean peninsula (receptor) was divided into three sub-regions (Northern Korea (A), Central Korea (B), and Southern Korea (C)) as shown in Fig. 1 to determine sub-regional differences in the degree of impact from the LRT events. The average AOD increases (AOD<sub>inc</sub>) calculated using Eqs. (7) and (8) were 128%, 121%, and 112% over the sub-regions (A), (B), and
 15 (C), respectively, showing that northern Korea experienced the largest impact of the transboundary PM pollution from China during this period. The overall average estimated AOD increase over the entire Korean peninsula during the two month period was 121%. The daily variations of the average AOD values over the three sub-regions

(A), (B), (C), and the entire Korean peninsula are shown in Fig. 7.

#### 20 4 Summary and conclusions

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In this study, a new approach was introduced to more accurately monitor transboundary LRT events of particulate pollutants over northeast Asia, combining the CMAQsimulated AOD with the GOCI-retrieved AOD via an OI technique (data assimilation). The PM LRT phenomena from China (source) to the Korean peninsula and Japan (receptor) usually take place by persistent westerlies in northeast Asia. The hourly GOCI-retrieved AOD with high spatial resolution of 500 m × 500 m was provided in or-



der to improve the accuracy of the CMAQ-simulated AOD through data assimilation, which produced a two-dimensional set of AOD data over the entire modeling domain. Using the assimilated AOD data, the LRT and the non-LRT events were distinguished, and the impact of the LRT events were estimated. During the LRT events, the aver-

- <sup>5</sup> age AOD increases of 117–265 % were found at the 5 AERONET sites in Korea and Japan, compared to the average background AOD value. In addition, the average AOD increase of 121 % over the entire Korean peninsula was also estimated. This strongly indicates that the noticeable amounts of particulate matters emitted/formed from China were transported into the Korean peninsula through the LRT processes.
- <sup>10</sup> The analysis framework presented here that utilizes new GOCI-retrievals of AOD provides a powerful means to better document and estimate the contributions of LRT to PM pollution and this has important policy implications. AOD retrievals from geo-stationary platforms greatly increase the amount of observations that go into such estimates and greatly reduce the uncertainties that model adds to the analysis. With
- <sup>15</sup> higher confidence and accuracy in the estimates it is easier for policy makers to discuss transboundary PM pollution over a region like northeast Asia. Looking forward, new observation systems are being designed that will provide more comprehensive aerosol information. For example, the Korean GEMS and the second version of GOCI (GOCI-2) sensors are scheduled to be launched in 2018. GEMS will become the world's first
- <sup>20</sup> geostationary environmental monitoring sensor that will be able to provide approximately 10 snapshots per day of information on air pollutants over entire China, Korea, Japan, southeast Asia, and parts of India, Russia and Central Asia, with a grid resolution of 5 km × 5 km. GEMS will be equipped with visible and UV channels. Thus, aerosol optical properties (like AOD) as well as gas data (e.g. tropospheric NO<sub>2</sub>, SO<sub>2</sub>,
- O<sub>3</sub>, CH<sub>2</sub>O, and glyoxal) will also be retrieved. GOCI-2 will again monitor the ocean colors over a similar domain as GOCI-1, but with 13 visible/NIR channels and a smaller grid resolution of 0.25 km × 0.25 km (Ahn, 2009). The AOD distributions obtained by applying the methodology discussed in this paper with the GEMS and GOCI-2 data will be greatly improved. Ultimately, the GEO satellite sensor-derived hourly production of



more accurate AOPs together with other pollutant information over east/northeast Asia is expected to greatly improve our ability to predict and to investigate air quality and climate forcing in east/northeast Asia in the future.

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# **Table 1.** Brief descriptions of aerosol optical property retrievals from some geostationary satellites.

| Satellite &<br>Sensor | GMS5 <sup>1</sup><br>VISSR <sup>7</sup>           | MTSAT-1R <sup>2</sup><br>JAMI <sup>8</sup> | MSG <sup>3</sup><br>SEVIRI <sup>9</sup> | GOES <sup>4</sup><br>GOES-8                                   | COMS⁵<br>GOCI <sup>10</sup>         | GOES-R<br>ABI <sup>11</sup>           | MP-<br>GEOSAT <sup>6</sup><br>GEMS <sup>12</sup> &<br>ABI &GOCI-<br>2 |
|-----------------------|---|--|---|---|-------------------------------------|---------------------------------------|---|
| Sensor type           | MET <sup>13</sup><br>sensor                       | MET sensor                                 | MET sensor                              | MET sensor  | Ocean color<br>imager               | MET sensor                            | Environ. &<br>Ocean color<br>imager                                   |
| Temporal resolution   | 1 h   | 30 min                                     | 15 min                                  | 15 min  | 1 h                                 | TBD <sup>17</sup>                     | 1 h (GEMS)  |
| Spatial resolution    | <sup>a</sup> 1.25 km/ <sup>b</sup> 5 kn           | n 1.25 km/5 km                             | 1 km/3 km                               | 1 km/4 km   | 0.5 km/0.5 km                       | 0.5 km/2 km                           | 0.025 km/1 kn   |
| Number of<br>bands    | 1 VIS <sup>14</sup> /3<br>IR <sup>15</sup>        | 1 VIS/4 IR                                 | 2 VIS/2<br>NIR/8 IR                     | 1 VIS/4 IR  | 6 VIS/2<br>NIR <sup>16</sup>        | 2 VIS/4<br>NIR/10 IR                  | 13 VIS/NIR  |
| Covering region       | East Asia &<br>West Pacific<br>Ocean              | East Asia &<br>West Pacific<br>Ocean       | Central<br>Europe                       | Eastern US  | Northeast<br>Asia                   | US                                    | East, South-<br>east & Cen-<br>tral Asia                              |
| Launch year           | March 1995  | February<br>2005                           | August 2002                             | April 1994  | June 2010                           | Schedule to<br>be launched<br>in 2015 | Schedule to<br>be launched<br>in 2018                                 |
| References            | Wang et al.<br>(2003);<br>Masuda et<br>al. (2002) | Kim et al.<br>(2008)                       | Popp et al.<br>(2007)                   | Christopher<br>et al. (2002);<br>Knapp et al.<br>(2002, 2005) | Lee et al.<br>(2010);<br>this study | Laszlo et al.<br>(2008)               |   |

<sup>1</sup> GMS: Geostationary Meteorological Satellite <sup>2</sup> MTSAT-1R: Multi-functional Transport Satellite <sup>3</sup> MSG: Meteosat Second Generation <sup>4</sup> GOES: Geostationary Operational Environmental Satellite <sup>5</sup> COMS: Communication, Ocean, and Meteorological Satellite <sup>6</sup> MP-GEOSAT: Multi Purpose GEOstationary SATellite <sup>7</sup> VISSR: Visible and Infra-Red Spin Scan Radiometer <sup>8</sup> JAMI: Japanese Advanced Meteorological Imager <sup>9</sup> SEVIRI: Spinning Enhanced Visible and Infra-Red Imager <sup>10</sup> GOCI: Geostationary Ocean Color Imager <sup>11</sup> ABI: Advanced Baseline Imager <sup>12</sup> GEMS: Geostationary Environment Monitoring Spectrometer <sup>13</sup> MET: Meteorological <sup>14</sup> VIS: VISible <sup>15</sup> IR: Infra-Red <sup>16</sup> NIR: Near Infra-Red <sup>17</sup> TBD: To Be Determined. <sup>a</sup> Resolution for the visible channels <sup>b</sup> Resolution for the infra-red channels.



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| Region                   | Emission inventory  | Species  |
|--------------------------|---|--|
| China and North<br>Korea | INTEX-B <sup>1</sup>  | SO <sub>2</sub> , NO <sub>x</sub> , NMVOCs, CO, CO <sub>2</sub> , BC, OC   |
|                          | EDGAR <sup>2</sup>  | NH <sub>3</sub>  |
| South Korea              | CAPSS <sup>3</sup><br>INTEX-B   | $SO_2$ , $NO_x$ , NMVOCs, CO, $NH_3$<br>$CO_2$ , BC, OC, Chemical split factors of<br>NMVOCs   |
| Japan                    | REAS <sup>4</sup>   | $SO_2$ , $NO_x$ , NMVOCs, CO, $CO_2$ , $NH_3$ , BC, OC   |
| Entire domain            | GAINS <sup>5</sup> emission<br>scenarios<br>FINN <sup>6</sup> fire emissions<br>MEGAN <sup>7</sup> +MOHYCAN <sup>8</sup><br>GEIA <sup>9</sup> | Inter-annual variations of CO, $NO_x$ ,<br>VOC, SO <sub>2</sub> , NH <sub>3</sub> , CO <sub>2</sub> , CH <sub>4</sub> , PM <sub>2.5</sub> <sup>a</sup><br>CO, NO, NO <sub>2</sub> , SO <sub>2</sub> , NH <sub>3</sub> , NMVOCs,<br>BC, OC, PM <sub>2.5</sub> , PM <sub>10</sub><br>Isoprene<br>Ratios of mono-terpene to isoprene <sup>b</sup> |

 Table 2. Emission inventories used in the CMAQ model simulations.

<sup>1</sup> INTEX-B: Intercontinental Chemical Transport Experiment-Phase B <sup>2</sup> EDGAR: Emission Database for Global Atmospheric Research <sup>3</sup> CAPSS: Clean Air Policy Support System <sup>4</sup> REAS: Regional Emission inventory in Asia <sup>5</sup> GAINS: Greenhouse gas and Air pollution INteractions and Synergies <sup>6</sup> FINN: Fire Inventory from NCAR <sup>7</sup> MEGAN: Model of Emissions of Gases and Aerosols from Nature <sup>8</sup> MOHYCAN: MOdel for Hydrocarbon emissions by the CANopy <sup>9</sup> GEIA: Global Emissions Inventory Activity. <sup>a</sup> Emission increase (or decrease) rates from 2006 to 2011 for the air pollutants were obtained from the GAINS emission scenarios. <sup>b</sup> Mono-terpene emissions are not included in MEGAN+MOHYCAN emissions. Therefore, the ratios of mono-terpene to isoprene in the biogenic emissions were obtained from the GEIA emission inventory.



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**Table 3.** Comparison analysis of the assimilated AOD for LRT and non-LRT events at five AERONET sites during the entire modeling period of 1 April to 31 May 2011.

|                     | 1              | lon-LR     | T event | S    |                | LRT  | events |      | Data fraction of                        | AOD increase                      | Ratio of AOD increase             |  |
|---------------------|----------------|------------|---------|------|----------------|------|--------|------|---|-----------------------------------|-----------------------------------|--|
|                     | N <sup>1</sup> | $\sigma^2$ | Avg.    | Max. | N <sup>1</sup> | σ²   | Avg.   | Max. | LRT events $(x_{LRT} \text{ in Eq. 7})$ | during LRT<br>events <sup>3</sup> | during LRT<br>events <sup>4</sup> |  |
| Baengnyeong         | 849            | 0.11       | 0.16    | 0.95 | 606            | 0.36 | 0.59   | 2.15 | 0.42                                    | 0.43                              | 2.59                              |  |
| Yonsei_Univ (Seoul) | 821            | 0.15       | 0.22    | 0.77 | 634            | 0.29 | 0.49   | 2.13 | 0.44                                    | 0.28                              | 1.28                              |  |
| Gwangju_GIST        | 1024           | 0.13       | 0.20    | 1.10 | 431            | 0.31 | 0.47   | 1.89 | 0.30                                    | 0.27                              | 1.36                              |  |
| Gosan               | 967            | 0.12       | 0.17    | 1.09 | 488            | 0.24 | 0.41   | 1.53 | 0.34                                    | 0.24                              | 1.37                              |  |
| Osaka               | 1247           | 0.11       | 0.16    | 1.16 | 208            | 0.21 | 0.39   | 1.23 | 0.14                                    | 0.23                              | 1.41                              |  |
| Total               |                | 0.12       | 0.18    | 1.16 |                | 0.31 | 0.49   | 2.15 | 0.33                                    | 0.31                              | 1.68                              |  |

The calculations were conducted with the assimilated AOD products. The AOD data assimilations were carried out with the GOCI-retrieved AOD products (retrieved via Yonsei algorithm, Lee et al., 2010) and the CMAQ-simulated AOD products (calculated via Park et al.'s conversion algorithm, Park et al., 2011). <sup>1</sup> The number of data for non-LRT or LRT events for the entire modeling period; <sup>2</sup> Standard deviation; <sup>3</sup> AOD increased by LRT events above the average background AOD calculated during non-LRT periods; <sup>4</sup> The ratios were estimated by  $((\overline{AOD}_{LRT} - \overline{AOD}_{non-LRT})/\overline{AOD}_{non-LRT})$ . Here,  $\overline{AOD}_{LRT}$  and  $\overline{AOD}_{non-LRT}$  indicate the average AOD calculated during the LRT and non-LRT events, respectively.

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Table 4. Comparison analysis of the assimilated AOD for LRT and non-LRT events at five AERONET sites during the entire modeling period of 1 April to 31 May 2011.

|                     | ١              | Non-LR     | T event | S    |                | LRT        | events |      | Data fraction of<br>LRT events      | AOD increase<br>during LRT | Ratio of AOD increase<br>during LRT |
|---------------------|----------------|------------|---------|------|----------------|------------|--------|------|-------------------------------------|----------------------------|-------------------------------------|
|                     | N <sup>1</sup> | $\sigma^2$ | Avg.    | Max. | N <sup>1</sup> | $\sigma^2$ | Avg.   | Max. | ( <i>x</i> <sub>LRT</sub> in Eq. 7) | events <sup>3</sup>        | events <sup>4</sup>                 |
| Baengnyeong         | 849            | 0.12       | 0.22    | 0.98 | 606            | 0.48       | 0.79   | 2.79 | 0.42                                | 0.57                       | 2.65                                |
| Yonsei_Univ (Seoul) | 821            | 0.21       | 0.30    | 1.13 | 634            | 0.37       | 0.66   | 2.93 | 0.44                                | 0.35                       | 1.17                                |
| Gwangju_GIST        | 1024           | 0.16       | 0.27    | 1.22 | 431            | 0.40       | 0.63   | 2.50 | 0.30                                | 0.36                       | 1.30                                |
| Gosan               | 967            | 0.12       | 0.22    | 1.14 | 488            | 0.31       | 0.54   | 1.95 | 0.34                                | 0.32                       | 1.47                                |
| Osaka               | 1247           | 0.11       | 0.21    | 0.90 | 208            | 0.22       | 0.48   | 1.26 | 0.14                                | 0.27                       | 1.27                                |
| Total               |                | 0.15       | 0.24    | 1.22 |                | 0.40       | 0.65   | 2.93 | 0.33                                | 0.40                       | 1.67                                |

The calculations were conducted with the assimilated AOD products. The AOD data assimilations were carried out with the GOCI-retrieved AOD products (retrieved via Yonsei algorithm, Lee et al., 2010) and the CMAQ-simulated AOD products (calculated via Chin et al.'s conversion algorithm, Chin et al., 2002). <sup>1</sup> The number of data for non-LRT or LRT events for the entire modeling period; <sup>2</sup> Standard deviation; <sup>3</sup> AOD increased by LRT events above the average background AOD calculated during non-LRT periods; <sup>4</sup> The ratios were estimated by

((AOD<sub>LRT</sub> – AOD<sub>non-LRT</sub>)/AOD<sub>non-LRT</sub>). Here, AOD<sub>LRT</sub> and AOD<sub>non-LRT</sub> indicate the average AOD calculated during the LRT and non-LRT events, respectively.

**Table 5.** Statistical analysis among AERONET, GOCI-retrieved, CMAQ-simulated, and assimilated AODs for the modeling period of 1 April to 31 May 2011.

| AOD (via Chin et al.'s conversion |                       |       | vs. A             | ERONET            |        |                  |
|-----------------------------------|-----------------------|-------|-------------------|-------------------|--------|------------------|
| algorithm, Chin et al., 2002)     | <i>N</i> <sup>1</sup> | $R^2$ | RMSE <sup>3</sup> | MNGE <sup>4</sup> | $MB^5$ | NMB <sup>6</sup> |
| CMAQ                              | 946                   | 0.47  | 0.31              | 46.52             | -0.09  | -14.51           |
| GOCI                              | 499                   | 0.84  | 0.19              | 42.42             | -0.03  | -15.15           |
| Assimilation                      | 695                   | 0.63  | 0.28              | 35.75             | -0.04  | -6.50            |







**Fig. 1.** Domains of CMAQ model simulations (in blue color) and COMS/GOCI sensor (in red color). Also shown are five AERONET sites in the receptor areas (Korea and Japan). For the sake of analytical convenience, three sub-regions (A), (B), and (C) in the Korean peninsula are defined, each of which represents northern, central, and southern parts of the Korean peninsula, respectively.





Fig. 2. Snapshots of (a) Terra and Aqua/MODIS-retrieved AODs and (b) GOCI-retrieved AOD over northeast Asia at 10:00 LST (Terra/MODIS) and 13:00 LST (Aqua/MODIS) on 10 April and 12 April.





**Fig. 3.** Spatial and temporal distributions of **(a)** CMAQ-simulated AOD (calculated via Chin et al. (2002)'s conversion algorithm), **(b)** GOCI-retrieved AOD (retrieved via Yonsei aerosol retrieval algorithm, Lee et al., 2010), and **(c)** assimilated AOD with two-hour resolution over northeast Asia at 09:00, 11:00, 13:00, and 15:00 LST on 10 April 2011.





**Fig. 4.** As Fig. 3, for a non-LRT event on 12 April 2011. This non-LRT case was determined, because the main AOD plumes were transported into over northwestern Pacific Ocean, not over the Korean peninsula and/or Japan.











**Fig. 6.** Time-series analysis of AERONET AOD (black bars) and two AOD products (red horizontal bars indicate the GOCI-retrieved AOD products; blue lines represent the assimilated AOD products) over the modeling period of 1 April to 31 May 2011 at five AERONET sites: Baengnyeong, Yonsei\_Univ (Seoul), Gwangju\_GIST, Gosan\_SNU,, and Osaka. The averaged background AODs estimated from the assimilated AOD values during non-LRT events are presented by blue dashed lines. Also, the periods of LRT events are displayed with grey-shaded areas.





**Fig. 7.** Time-series plots of the assimilated average AODs (calculated via Chin et al. (2002)'s conversion algorithm) over Northern Korea **(A)**, Central Korea **(B)**, Southern Korea **(C)**, and the entire Korean peninsula, shown in Fig. 1, for the modeling period of 1 April to 31 May 2011.

