

This manuscript represents the 6<sup>th</sup> installment of the IUPAC Task Group on Atmospheric Chemical Kinetic Data Evaluation on heterogeneous reactions involving aqueous particles. The uptake coefficients and adsorption parameters for various gas phase species on four aqueous surfaces have been considered including liquid water, deliquesced halide salts, other aqueous electrolytes, and aqueous sulphuric acid. An introduction into the various aspects of uptake kinetics is given to assist in understanding the evaluated parameters for description of the underlying kinetics. For each case the available experimental data is presented and a recommendation of parameters is given. This document represents a very good summary of the literature and an extensive evaluation considering that the authors placed the various experimental data into a unifying framework which made it necessary to reanalyze the laboratory data. This approach will allow parameterization of the heterogeneous kinetics in atmospheric chemistry models.

I appreciate the effort the authors took to describe the heterogeneous kinetics within a unifying framework. The introduction into the kinetics represents a nice summary of the corresponding papers by Pöschl, Rudich and Ammann [*Ammann and Pöschl, 2007; Pöschl et al., 2007*]. Although, I am not familiar with all experimental data sets I found the representation, discussion, and final evaluation, assisted by the applied theoretical framework, fair and comprehensive. Overall, I am very supportive of the publication of this manuscript in ACP. I have a couple of minor general and specific scientific comments the authors want to consider before publication of this document.

Regarding technical corrections: This is a long document that includes sections written by different authors, and thus exhibits in many places inconsistent formatting but also typos etc. I am giving some general comments on technical corrections below. However, instead of writing a very long list of individual page and line numbers to indicate repeated inconsistencies (in e.g. applied units etc.) or typos, for the over 350 ACPD pages, I have given these corrections as notes on top of the manuscript attached as pdf file to this review (see below). Also, specific scientific comments are indicated within the attached manuscript. I hope that this approach may speed up the correction process when using e.g. two screens/windows in parallel to find readily the places for making the corrections. In some places the corrections seem meticulous and only serve the authors to not miss out on instances when deciding to follow my suggestions.

#### General comments:

I advise to use SI units throughout this document as defined in the beginning of the manuscript. Pressures should be in Pa or hPa and diffusion coefficients in  $\text{atm cm}^2 \text{s}^{-1}$ , etc. I tried to mark all instances in which units have to be changed.

Many of the discussed uptake processes depend on Henry's law solubility which in turn depends, in cases, on the concentration of the solute or protons ( $\text{H}^+$ ). It has been shown that in aqueous electrolytic solutions the bulk concentration may not reflect the ion concentration at the gas-liquid interface (see e.g. [*Jungwirth and Tobias, 2006*] and others). This may not only affect the predicted kinetics but the overall interpretation of the underlying processes when going from gas to interface to bulk. I feel that this point may be worthwhile mentioning.

Regarding the uptake processes that involve aqueous sulphuric acid. In line with my previous point, gas uptake is affected by the Henry's law constant which in turn can depend on the proton concentration/activity,  $[\text{H}^+]$ . For a correct description of  $[\text{H}^+]$  at colder temperatures, the dissociation of  $\text{H}_2\text{SO}_4$  and  $\text{HSO}_4^-$  for supercooled temperatures must be known. Commonly the AIM model [*Carslaw et al., 1995; Clegg et al., 1998*] or other parameterizations are used to

obtain  $[H^+]$  at desired temperatures (e.g. for upper tropospheric / lower stratospheric, UT/LS, conditions). However, experimentally obtained  $HSO_4^-$  dissociation degrees, using supercooled aqueous  $H_2SO_4$  droplets [Knopf *et al.*, 2003; Myhre *et al.*, 2003], are significantly different than when predicted by e.g. AIM. The Pitzer ion interaction model by [Knopf *et al.*, 2003] reproduces the experimental findings and as a consequence yields significantly different  $H^+$  concentrations (in fact larger concentrations) compared to AIM predictions at colder temperatures. (For more discussion, see also comment and reply on our study [Clegg and Brimblecombe, 2005; Knopf *et al.*, 2005].) This in turn leads to a lower solubility of trace gases such as HCl or  $NH_3$ . The differences between AIM predictions and experimental data (dissociation degree, activity coefficients, water activity) are largest for UT/LS conditions. The change in solubility may affect, in cases, interpretation of the underlying kinetics/parameters and may be worthwhile to evaluate.

This compendium does not systematically evaluate the effect of organic coatings (mono- or multilayers of organic molecules) at the gas-liquid interface on the uptake of trace gases. For some cases (inconsistent), the authors mention very briefly the effect of a coating. I suggest discarding this information from the manuscript or if included, to evaluate all studies that investigated uptake of trace gases by water and aqueous solution surfaces coated by organic material. My feeling is that this would be beyond the scope of this manuscript. Furthermore, a coating or presence of an organic monolayer at the gas-liquid interface and its effect on the underlying kinetic parameters of the unifying framework may need more attention than it is given currently here (possibly inclusion of another layer with additional accommodation coefficients etc.). As is, the previous literature is not represented in a balanced way and thus I recommend omitting the few cases in which the authors mention the effect of organic coatings and focus solely on purely aqueous substrates.

Specific comments (see also remarks within manuscript pdf file):

After Eq. 6: Maybe I am missing the obvious, but I do not entirely understand lines 9-12. I think it would be helpful to explain from where comes the definition of  $k_{des}$ ? With B given, is it trivial to see that the characteristic time is  $1/k_{des}$ ? Which equilibrium is meant? I do not see readily how  $\gamma$  drops to zero with the information given. Please briefly elaborate on this section; this would be helpful for the reader.

After Eq. 16: "Equation (16) demonstrates that  $\gamma$  depends on the gas phase concentration of X, if  $K_{LangC}(X)[X]_g$  is similar to or larger than 1 (i.e., at high coverage)." I am not sure if a reader not as familiar with the resistance model can understand this statement. Maybe it would help to give the relationship between  $\gamma$  and  $\Gamma_s$ .

The summary sheets have inconsistent/missing formatting in terms of species phase state indication. E.g. " $H_2O(l)$ " and not " $H_2O(aq)$ " etc. Also, some reactions (and corresponding titles in the appendices) do not show the correct stoichiometric coefficients. Details are pointed out in manuscript file.

Table 1 could include the reacto-diffusive length as a parameter.

For the case of  $N_2O_5$  uptake by aqueous  $H_2SO_4$  our uptake study could be included [Knopf *et al.*, 2007] (if the authors feel it benefits the manuscript).  $N_2O_5$  uptake by aqueous  $H_2SO_4$  solution 60 and 80 wt% in concentration was determined using a rectangular channel flow

reactor (RCFR) coupled to CIMS correcting for changes in the  $\text{N}_2\text{O}_5$  profile due to uptake (i.e. accounting for diffusion effects) in flow and lateral direction.

### Technical Corrections:

There are many places in which technical corrections are needed for reasons of keeping consistency or correction of typos etc. Those are marked in the manuscript pdf-file.

A few general statements regarding technical corrections:

Keep document either in British or American English.

Keep format of "Experimental Data Table" and "Preferred Values Table" similar. E.g. "T/K" etc.

Keep concentration format consistent, i.e. "molecule  $\text{cm}^{-3}$ " and not " $\text{cm}^{-3}$ ", "molecules  $\text{cm}^{-3}$ ", or "atm".

I suggest using the mathematical symbol "times", i.e. "x" only when expressing the exponent as when giving concentrations but not in mathematical equations (this would indicate a cross product). For the latter application, please use a center dot.

When giving parameter ranges, make sure exponent and units are at the correct place. The formatting is very inconsistent throughout the manuscript. E.g.  $(1-5)\times 10^{-5}$  molecule  $\text{cm}^{-3}$ .

### **References**

Ammann, M., and U. Pöschl (2007), Kinetic model framework for aerosol and cloud surface chemistry and gas-particle interactions - Part 2: Exemplary practical applications and numerical simulations, *Atmos Chem Phys*, 7(23), 6025-6045.

Carlaw, K. S., S. L. Clegg, and P. Brimblecombe (1995), A Thermodynamic Model of the System  $\text{HCl-HNO}_3\text{-H}_2\text{SO}_4\text{-H}_2\text{O}$ , Including Solubilities of  $\text{HBr}$ , from  $< 200$  to  $328$  K, *J Phys Chem*, 99(29), 11557-11574.

Clegg, S. L., and P. Brimblecombe (2005), Comment on the "Thermodynamic dissociation constant of the bisulfate ion from Raman and ion interaction modeling studies of aqueous sulfuric acid at low temperatures", *J Phys Chem A*, 109(11), 2703-2706.

Clegg, S. L., P. Brimblecombe, and A. S. Wexler (1998), Thermodynamic model of the system  $\text{H}^+\text{-NH}_4^+\text{-SO}_4^{2-}\text{-NO}_3^-\text{-H}_2\text{O}$  at tropospheric temperatures, *J Phys Chem A*, 102(12), 2137-2154.

Jungwirth, P., and D. J. Tobias (2006), Specific ion effects at the air/water interface, *Chem Rev*, 106(4), 1259-1281.

Knopf, D. A., B. P. Luo, U. K. Krieger, and T. Koop (2003), Thermodynamic dissociation constant of the bisulfate ion from Raman and ion interaction modeling studies of aqueous sulfuric acid at low temperatures, *J Phys Chem A*, 107(21), 4322-4332.

Knopf, D. A., B. P. Luo, U. K. Krieger, and T. Koop (2005), Reply to "Comment on the 'Thermodynamic dissociation constant of the bisulfate ion from Raman and ion interaction modeling studies of aqueous sulfuric acid at low temperatures'", *J Phys Chem A*, 109(11), 2707-2709.

Knopf, D. A., L. M. Cosman, P. Mousavi, S. Mokamati, and A. K. Bertram (2007), A novel flow reactor for studying reactions on liquid surfaces coated by organic monolayers: Methods, validation, and initial results, *J Phys Chem A*, 111(43), 11021-11032.

Myhre, C. E. L., D. H. Christensen, F. M. Nicolaisen, and C. J. Nielsen (2003), Spectroscopic study of aqueous H<sub>2</sub>SO<sub>4</sub> at different temperatures and compositions: Variations in dissociation and optical properties, *J Phys Chem A*, 107(12), 1979-1991.

Pöschl, U., Y. Rudich, and M. Ammann (2007), Kinetic model framework for aerosol and cloud surface chemistry and gas-particle interactions - Part 1: General equations, parameters, and terminology, *Atmos Chem Phys*, 7(23), 5989-6023.

This discussion paper is/has been under review for the journal Atmospheric Chemistry and Physics (ACP). Please refer to the corresponding final paper in ACP if available.

# Evaluated kinetic and photochemical data for atmospheric chemistry: Volume VI – heterogeneous reactions with liquid substrates

M. Ammann<sup>1</sup>, R. A. Cox<sup>2</sup>, J. N. Crowley<sup>3</sup>, M. E. Jenkin<sup>4</sup>, A. Mellouki<sup>5</sup>,  
M. J. Rossi<sup>6</sup>, J. Troe<sup>7</sup>, and T. J. Wallington<sup>8</sup>

<sup>1</sup>Laboratory of Radiochemistry and Environmental Chemistry, OFLB 103, Paul Scherrer Institut, 5232 Villigen, Switzerland

<sup>2</sup>Centre for Atmospheric Science, Dept. of Chemistry, University of Cambridge, Lensfield Road, Cambridge CB2 1EP, UK

<sup>3</sup>Max-Planck-Institut für Chemie, Division of Atmospheric Chemistry, Hahn-Meitner-Weg 1, 55128 Mainz, Germany

<sup>4</sup>Atmospheric Chemistry Services, Okehampton, Devon EX20 1FB, UK

<sup>5</sup>ICARE (Institut de Combustion, Aérothermique, Réactivité et Environnement)/OSUC, CNRS (Centre National de la Recherche Scientifique) – UPR3021, 1C, Avenue de la recherche scientifique, 45071 Orléans cedex 02, France

<sup>6</sup>Laboratory of Atmospheric Chemistry, OFLA 008, Paul Scherrer Institut, 5232 Villigen, Switzerland

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

<sup>7</sup>Institute of Physical Chemistry, University of Göttingen, Tammannstr. 6, 37077 Göttingen, Germany

<sup>8</sup>Ford Motor Company, Research and Advanced Engineering, Mail Drop RIC-2122, Dearborn, Michigan 48121-2053, USA

Received: 3 September 2012 – Accepted: 5 September 2012 – Published: 17 December 2012

Correspondence to: M. Ammann (markus.ammann@psi.ch)

Published by Copernicus Publications on behalf of the European Geosciences Union.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Abstract

This article, the sixth in the ACP journal series, presents data evaluated by the IUPAC Task Group on Atmospheric Chemical Kinetic Data Evaluation. It covers the heterogeneous processes involving liquid particles present in the atmosphere, for which uptake coefficients and adsorption parameters have been presented on the IUPAC website since 2009. The article consists of an introduction and guide to the evaluation, giving a unifying framework for parameterisation of atmospheric heterogeneous processes. We provide summary sheets containing the recommended uptake parameters for the evaluated processes. The experimental data on which the recommendations are based are provided in data sheets in separate appendices for the four surfaces considered: liquid water, deliquesced halide salts, other aqueous electrolytes and sulfuric acid.

## 1 Introduction

Since 2005 the IUPAC Task Group on Atmospheric Chemical Kinetic Data Evaluation has extended its evaluation work to produce recommendations for essential parameters, which can be used to calculate the rates of heterogeneous reactions of trace gases in the atmosphere. The results of this work have been added to the IUPAC website over the last 4 years, together with the updated kinetic data for gas phase reactions. Following our policy of publication of our updated evaluations in Atmospheric Chemistry and Physics, we first presented evaluations of heterogeneous processes involving solid particles in Volume V (Crowley et al., 2010). Volume VI of the series extends our work to cover heterogeneous processes involving liquid particles present in the atmosphere. This is done with a view to widening the dissemination and enhancing the accessibility of this evaluated material to the scientific community. The evaluation also fosters usage of a unified terminology throughout the scientific community (Kolb et al., 2010). Last but not least, it will enable appropriate incorporation of kinetic parameters of heterogeneous processes into atmospheric models.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The article consists of summary sheets containing the recommended values for the parameters describing adsorption, solvation and reaction, for which sufficient experimental information exists to allow a recommendation. Uncertainties to the kinetic parameters, as well as parameters describing equilibria and liquid phase diffusion are listed in the individual datasheets only. The summary tables are followed by a guide to the datasheets, which outlines the physico-chemical basis underlying the quantitative parameters describing atmospheric heterogeneous processes, and which provides definitions of terms employed. Finally, we present four appendices containing the data sheets for uptake of a range of  $O_x$ ,  $HO_x$ ,  $NO_x$ ,  $SO_x$ , organic, and halogen-containing trace gases, on liquid surfaces of water (Appendix A1), deliquesced halide salts (Appendix A2), other aqueous electrolytes (Appendix A3) and sulfuric acid (Appendix A4). As adopted in Volume V, the data sheets follow a similar format to those used for gas phase reactions, providing details of published experimental information upon which the recommendations are based, a table of preferred values and their reliability, and comments on the state of knowledge leading to the preferred values. The data sheets include relevant references.

---

**Evaluated kinetic and photochemical data for atmospheric chemistry**M. Ammann et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)



## 2 Summary sheets

### Preferred values for uptake on water surfaces – Appendix A1

ID	Species/reaction	$\gamma_{ss}$	$\alpha_s$	$\alpha_b$	$k_b/M^{-1} s^{-1}$	T/K
VI.A1.1	O <sub>3</sub>			$>10^{-3}$		298
VI.A1.2	H <sub>2</sub> O <sub>2</sub>			$1.3 \times 10^{-6} \exp(3230/T)$		260–292
VI.A1.3	NO <sub>2</sub>			$2 \times 10^{-2}$	$8.1 \times 10^9 \exp(-1630/T) M^{-1} s^{-1}$	276–293
VI.A1.4	NO <sub>3</sub>			$1.3 \times 10^{-2}$	$k^1 = 23 s^{-1}$	273
VI.A1.5	N <sub>2</sub> O <sub>5</sub>	$2.7 \times 10^{-5} \exp(1800/T)$				265–300
VI.A1.6	NH <sub>3</sub>			0.1		260–300
VI.A1.7	$\text{HNO}_2(\text{g}) + \text{H}_2\text{O}(\text{l}) \rightarrow \text{NO}(\text{aq})$ $+ \text{NO}_2(\text{aq}) + \text{H}_2\text{O}(\text{l})$			0.05	$k_{-} = 3.1 \times 10^{14} \exp(-9090/T) M^{-1} s^{-1}$ $k_{-} = 1.7 \times 10^8 M^{-1} s^{-1}$	273–300 273–300
VI.A1.8	HNO <sub>3</sub>			$7.5 \times 10^{-5} \exp(2100/T)$		268–300
VI.A1.9	CH <sub>3</sub> SO <sub>3</sub> H			1		260–300
VI.A1.10	SO <sub>2</sub> + H <sub>2</sub> O (aq) ⇌ H <sub>2</sub> SO <sub>3</sub> (aq)		1	$0.18 \times (1 - \theta)$		260–300
VI.A1.11	HCl			$4.4 \times 10^{-6} \exp(2898/T)$		272–294
VI.A1.12	HBr			$1.3 \times 10^{-8} \exp(4290/T)$		263–281
VI.A1.13	HI			$6.35 \times 10^{-9} \exp(4519/T)$		262–278
VI.A1.14	ClNO + H <sub>2</sub> O(l) → H <sup>+</sup> (aq) + Cl <sup>-</sup> (aq) + HNO <sub>2</sub> (aq)	0.01				273–293
VI.A1.15	ClNO <sub>2</sub>	$3 \times 10^{-6}$		$9 \times 10^{-3}$		273–298
VI.A1.16	ClONO <sub>2</sub> + H <sub>2</sub> O(l) → HOCl + HONO <sub>2</sub>			0.11		273–290
VI.A1.17	OH			$>0.1$		275–310
VI.A1.18	H <sub>2</sub> O			$3.6 \times 10^{-5} \exp(2370/T)$		258–298
VI.A1.19	SO <sub>2</sub> + H <sub>2</sub> O <sub>2</sub> (aq) ⇌ H <sub>2</sub> SO <sub>4</sub> (aq)			0.11	$2.1 \times 10^6 / (0.1 + [H^+]) M^{-2} s^{-1}$	268–298
VI.A1.20	HO <sub>2</sub>			$>0.5$		270–300

## Preferred values of uptake coefficients on aqueous halide salt surfaces – Appendix A2

ID	Species/reaction	$\gamma_{ss}$	$\alpha_s$	$k_s$	$\alpha_b$	$k_b/M^{-1} s^{-1}$	T/K
VI.A2.0	$O_3 + I^- (aq) \rightarrow \text{products}$				>0.1	$(2.4 \pm 1.3) \times 10^9$	293
VI.A2.0	$O_3 + Br^- (aq) \rightarrow \text{products}$			<i>see datasheet</i>		$6.3 \times 10^8 \exp(-4450/T)$	273–298
VI.A2.1	OH	$\gamma_{gs} = 0.04 \times ([Cl^-]/M)$			>0.1		298
VI.A2.2	HO <sub>2</sub>	0.1			>0.5		290–300
VI.A2.4	$NO_2 + NaCl \rightarrow NOCl + NaNO_3$					200	298
VI.A2.5	$NO_3 + Cl^- (aq) \rightarrow NO_3^- (aq) + Cl (aq)$				$1.3 \times 10^{-2}$	$2.76 \times 10^6$	273
VI.A2.5	$NO_3 + Br^- (aq) \rightarrow NO_3^- (aq) + Br (aq)$				$1.3 \times 10^{-2}$	$1.02 \times 10^8$	273
VI.A2.5	$NO_3 + I^- (aq) \rightarrow NO_3^- (aq) + I (aq)$				$1.3 \times 10^{-2}$	$4.6 \times 10^9$	273
VI.A2.6	N <sub>2</sub> O <sub>5</sub>	0.02					260–300
VI.A2.7	HNO <sub>3</sub>				>0.5		298
VI.A2.8	HOCl	$< 2 \times 10^{-4}$					296
VI.A2.9	$ClNO_2 + Cl^- \rightarrow$				0.01	$> 10^7$	274
VI.A2.9	$ClNO_2 + Br^- \rightarrow$				0.01		274
VI.A2.10	$ClONO_2 + Cl^- \rightarrow Cl_2 + HNO_3$				0.11		273–290
VI.A2.10	$ClONO_2 + Br^- \rightarrow ClBr + HNO_3$				0.11		273–290
VI.A2.11	$HOBr + Cl^-/Br^- (aq) + H^+ \rightarrow \text{products}$				0.6	$k_{ter} > 5.6 \times 10^9 M^{-2} s^{-1}$	296
VI.A2.12	$HOI + Cl^-/Br^- \rightarrow IBr/ICl + HO^-$				>0.1		298
VI.A2.13	$BrCl + Br^- (aq) \rightarrow \text{Products}$				0.33		270–280
VI.A2.14	$ICl + Br^- \rightarrow IBr + Cl^-$				0.016	$> 10^3$	298
VI.A2.15	$IBr + Cl^-/Br^- \rightarrow \text{products}$				0.016		298
VI.A2.16	$Cl_2 + Br^- (aq) \rightarrow BrCl (aq) + Cl^- (aq)$		1	<i>see datasheet</i>	1	$4.58 \times 10^{14} \exp(-2866/T)$	260–295

## Preferred values for uptake on other aqueous electrolyte surfaces – Appendix A3

ID	Species/reaction	$\gamma_{ss}$	$\alpha_s$	$\alpha_b$	$k_b/M^{-1} s^{-1}$	T/K
VI.A3.5	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (ammonium sulphate)	$\gamma_r = 0.244 - 7.9 \times 10^{-4} T(K)$		0.03	$1.0 \times 10^5$	298
VI.A3.6	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (ammonium bisulphate)			0.04	$1.5 \times 10^5$	298
VI.A3.7	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (Na <sup>+</sup> /NH <sub>4</sub> <sup>+</sup> nitrate)			0.04	$1.5 \times 10^5 \left\{ 1 - \frac{k_2[NO_3^-]}{k_2[NO_3^-] + k_3[H_2O]} \right\}$	298
5 VI.A3.8	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (humic acid)	$3.1 \times 10^{-5} \exp(0.046 \times RH(\%))$				298
VI.A3.8	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (malonic acid)			0.035	$1.0 \times 10^5$	298
VI.A3.8	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (succinic acid)			0.035	$3.0 \times 10^4$	298
VI.A3.8	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O (glutaric acid)			0.035	$5.0 \times 10^4$	298
VI.A3.9	HO <sub>2</sub> + NH <sub>4</sub> HSO <sub>4</sub> (aq) → products			>0.5	<i>see datasheet</i>	293
VI.A3.10	HO <sub>2</sub> + (NH <sub>4</sub> ) <sub>2</sub> SO <sub>4</sub> (aq) → products			>0.5	<i>see datasheet</i>	290–300

# Preferred values for uptake on sulfuric acid surfaces – Appendix A4

ID	Species	$\gamma_{ss}$	$\alpha_s$	$k_s$	$\alpha_b$	$k_b/M^{-1} s^{-1}$	T/K
VI.A4.0	O <sub>3</sub>	<10 <sup>-6</sup>					200–220
VI.A4.1	OH				1		220–298
VI.A4.2	HO <sub>2</sub>				0.8		220–300
VI.A4.3	H <sub>2</sub> O <sub>2</sub>	7.8 × 10 <sup>-4</sup>					270–300
VI.A4.4	NO	<5 × 10 <sup>-6</sup>					200–298
VI.A4.5	NO <sub>2</sub>	<10 <sup>-6</sup>					200–298
VI.A4.6	NO <sub>3</sub>	<1 × 10 <sup>-3</sup>					200–270
VI.A4.7	HNO <sub>2</sub>	$\gamma_{gs} < 10^{-5}$			>0.05	320	220–300
VI.A4.8	HNO <sub>3</sub>		>0.8		>0.8		188–223
VI.A4.9	HO <sub>2</sub> NO <sub>2</sub>				>0.2		205–225
VI.A4.10	NH <sub>3</sub> (>50 wt % H <sub>2</sub> SO <sub>4</sub> )	1.0					265–300
VI.A4.11	N <sub>2</sub> O <sub>5</sub> + H <sub>2</sub> O	$\gamma_f = [(7353/T) - 24.83]^{-1}$			2.3 × 10 <sup>-5</sup> exp(2240/T)		240–300
VI.A4.12	H <sub>2</sub> CO(g)	$\Gamma_{b,poly} = 0.001$			0.04	$k_{H_2O} = 7800 \exp(-1910/T) (1+870[H^+]) M^{-1} s^{-1}$	200–300
VI.A4.13	CH <sub>2</sub> C(O)OONO <sub>2</sub> (PAN)				>0.008		
VI.A4.14	HCl		>0.9		$k_{sol} = 7 \times 10^9 / \eta$ (cP) $k_{des} = 8.7 \times 10^7 \exp(-5000/T)$		190–300
VI.A4.15	ClO ([H <sub>2</sub> O] = 6.7 × 10 <sup>-4</sup> mbar)	$\ln \gamma_{ss} = -9.361 + (3.22 \pm 1.34) \times (1000/T - 3.840)$					240–295
VI.A4.16	HOBr						
VI.A4.17	HBr						
VI.A4.18	HOI						
VI.A4.19	(CH <sub>3</sub> ) <sub>2</sub> CO						
VI.A4.20	CHBr <sub>3</sub>						
VI.A4.21	H <sub>2</sub> SO <sub>4</sub>				>0.7		200–300
VI.A4.22	HOCl + HCl → H <sub>2</sub> O + Cl <sub>2</sub>				$\alpha_b$ (HOCl) = 1 $\alpha_b$ (HCl) = 1	80 [H <sup>+</sup> ] T/η 1 × 10 <sup>7</sup>	190–273
VI.A4.23	HOCl + HBr → BrCl + H <sub>2</sub> O						228
VI.A4.24	ClONO <sub>2</sub> + HCl → HONO <sub>2</sub> + Cl <sub>2</sub>	$\Gamma_{ss} = 66.12 \exp(-1374/T) \times H_{ClONO_2} \times M_{HCl}$	1.0			see datasheet for $\Gamma_b$	190–280
VI.A4.25	ClONO <sub>2</sub> + H <sub>2</sub> O → HONO <sub>2</sub> + HOCl				1.0	$k_{H_2O} = 1.95 \times 10^{10} \exp(2800/T) s^{-1}$ $k_{H_2O} = 1.22 \times 10^{12} \exp(6200/T) M^{-1} s^{-1}$ $\exp(154 - 1.63W) \exp(-(38500 - 478W)/T)$	190–280
VI.A4.26	HOBr + HCl → BrCl + H <sub>2</sub> O				1.0		200–230
VI.A4.27	HOBr + HBr → Br <sub>2</sub> + H <sub>2</sub> O				0.3		228
VI.A4.28	BrONO <sub>2</sub>	$\Gamma_b = 0.11 + \exp(29.2 - 0.40 \text{ wt \% } H_2SO_4)$			0.8		210–300
VI.A4.29	HONO + HCl → ClNO + H <sub>2</sub> O		1	see datasheet		25[H <sup>+</sup> ] T/η	200–300
VI.A4.30	HONO + HBr → BrNO + H <sub>2</sub> O				1.0		200–300
VI.A4.31	HOCl						
VI.A4.32	CH <sub>3</sub> SO <sub>3</sub> H				1.0		296
VI.A4.33	N <sub>2</sub> O <sub>5</sub> + HCl → HONO <sub>2</sub> + ClNO <sub>2</sub>				0.11	see datasheet	200–220

## 3 Guide to the data sheets

The heterogeneous processes considered in this evaluation involve chemical and/or physical interactions between trace gases and liquid atmospheric condensed phase material. Broadly, this material falls into 2 categories: (i) large water droplets (diameter, d > 1 μm), and (ii) fine liquid submicron (d < 1 μm) aerosol particles. The interaction can

be reversible (physisorption at the surface or dissolution), reactive, catalytic or a combination of all or some of these operating in parallel or sequentially, and can depend strongly on ambient conditions such as temperature or relative humidity.

This publication covers only interactions between gases and liquid surfaces including aqueous droplets, salt solutions (e.g., halide and sulphate) and sulphuric acid. The parameterisation of a heterogeneous process depends on the nature of the surface but here we present the general framework for reaction kinetics for both solid and liquid substrates.

A proper description of the interaction of a trace gas with a surface would include transport to and accommodation at the surface, followed by a number of competitive or parallel processes such as desorption back to the gas phase, reaction with the substrate surface or with other trace gases on the surface, and diffusion into and reaction in the particle bulk. The rates and efficiencies of these processes are controlled by surface and bulk-phase rate coefficients, local reactant concentrations, diffusion coefficients in the condensed phase, and solubilities. Each of these controlling factors may change with temperature and composition. This convolution of equilibrium and kinetic parameters for phase change and reaction renders the description of atmospheric heterogeneous processes highly complex.

Only rarely are all the individual steps controlling reaction rates of heterogeneous processes known. A quasi steady state resistance model has frequently been used to describe the uptake of gas species to surfaces, and to relate the experimentally observable net probability of uptake ( $\gamma$ , see below) to the physical parameters that control it. In this model, a linear combination of flux resistances (decoupled and normalized fluxes) by analogy to resistances in electric circuits, is used to describe the uptake process (Ammann et al., 2003; Hanson, 1997; Jayne et al., 1990; Pöschl et al., 2007), and we will refer to these formulations below where appropriate. Models involving solution of coupled differential rate equations of mass transport and chemical reactions using continuum flow formulations (Winkler et al., 2004) and gas kinetic formulations (Behr et al., 2004; Flückiger and Rossi, 2003; Pöschl et al., 2007; Pfrang et al., 2010; Shiraiwa

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



et al., 2009, 2010) have recently been described. The latter circumvent steady state approximations that are not easy to assess and are preferable in complex situations.

### 3.1 Description of heterogeneous kinetics

#### 3.1.1 Definition of the uptake coefficient

5 The most widely used approach to describe the kinetics of heterogeneous processes is to use the uptake coefficient,  $\gamma$ , which is the net probability that a molecule X undergoing a gas-kinetic collision with a surface is actually taken up at the surface. This approach links the processes at the interface and beyond with an apparent first order loss of X from the gas phase:

$$10 \quad \frac{d[X]_g}{dt} = -k_r[X]_g = -\gamma \frac{\bar{c}}{4} [SS]_g [X]_g \quad (1)$$

$[X]_g$  denotes the concentration of X in the gas phase (molecule  $\text{cm}^{-3}$ ),  $[SS]_g$  is the specific surface area of the condensed phase ( $\text{cm}^{-1}$ ) (i.e., surface area of condensed phase per unit volume of gas phase),  $k_r$  is the first order rate coefficient for loss, and  $\bar{c}$  is the mean thermal velocity of X ( $\text{cm s}^{-1}$ ).  $\gamma$  often depends on time, as uptake may be limited by adsorption equilibrium on the surface, by a limited number of reactant molecules on the surface, by solubility, or by a limited number of reactant molecules in the bulk of the particles.  $\gamma$  may also depend on the gas phase concentration of X. Therefore,  $\gamma$  is not a constant, and  $\gamma$  values reported from laboratory measurements often cannot be transferred directly to other conditions. This must be done through a proper parameterization of the processes at the interface and beyond, as outlined below. Only a systematic variation of conditions in laboratory studies enables decon-  
20 volution of measured uptake coefficients into these underlying processes.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



### 3.1.2 Gas phase diffusion

If  $\gamma$  is large,  $[X]_g$  may become depleted near the particle surface. In this case, a correction factor is commonly applied, such that Eq. (1) can still be used to relate the net flux into the particle phase with the overall loss from the gas phase.  $\gamma$  in Eq. (1) is then replaced by  $\gamma_{\text{eff}}$ , which is given by  $\gamma_{\text{eff}} = C_{\text{diff}}\gamma$ , where  $C_{\text{diff}}$  is a correction factor. Under appropriate steady state assumptions, this relation can also be expressed in the form of a resistor formulation (Finlayson-Pitts and Pitts, 2000; Hanson et al., 1994; Schwartz, 1986):

$$\frac{1}{\gamma_{\text{eff}}} = \frac{1}{\Gamma_{\text{diff}}} + \frac{1}{\gamma} \quad (2)$$

Approximate formulas for  $C_{\text{diff}}$  or  $\Gamma_{\text{diff}}$  have been derived in the literature (Brown, 1978; Fuchs and Sutugin, 1970; Pöschl et al., 2007; Seinfeld and Pandis, 1998) for various geometries, such as suspended aerosol particles, but also for cylindrical flow tubes. They are a function of the geometry (e.g., particle diameter,  $d_p$ ) and the diffusion coefficient of X,  $D_g$ , and  $\gamma$ . Therefore, temperature and bath gas dependent diffusion coefficients for X are required. Diffusion may also limit uptake rates to particles in the atmosphere, requiring a similar treatment for uptake kinetics.

### 3.1.3 Surface accommodation and kinetics of adsorption

A molecule colliding with the surface of solid or liquid condensed matter can undergo elastic or inelastic scattering processes that involve collisions with one or a few surface atoms or molecules on time scales of up to  $10^{-12}$  s. Elastic and inelastic scattering result in reflection back to the gas phase. The molecule can also undergo adsorption, in which case it accommodates into a weakly bound state, which may involve hydrogen bonds, charge transfer, or Van der Waals interactions. Adsorbed molecules may leave the surface through thermally activated desorption, which results in lifetimes on the surface of typically between nanoseconds and seconds at atmospheric temperatures.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Thermally desorbing or scattered molecules can be distinguished through molecular beam experiments (Morris et al., 2000; Nathanson et al., 1996). The adsorption state as defined above is usually referred to as physisorption in the surface science literature (Masel, 1996). Physisorption should be differentiated from chemisorption, which involves breaking of chemical bonds or significant distortion of electronic structure of adsorbate and substrate. As discussed below, chemisorption is considered to be a surface reaction. We recommend use of the term “surface accommodation coefficient”,  $\alpha_s$ , for the probability of adsorption on a clean surface. Note that the symbol S and terms such as sticking probability or adsorption coefficient are used in the literature (Ammann et al., 2003; Carslaw and Peter, 1997; Davidovits et al., 1995; Garrett et al., 2006; Hanson, 1997; Jayne et al., 1990; Pöschl et al., 2007; Tabazadeh and Turco, 1993; Vieceli et al., 2005), however these terms often lack an unequivocal definition. The term thermal accommodation coefficient,  $\alpha_t$ , has been used to describe the probability that a molecule accommodates to the thermal energy of the substrate upon adsorption in a single collision model, when a molecule collides with the surface (Li et al., 2001; Vieceli et al., 2005; Winkler et al., 2004; Worsnop et al., 2002). A more detailed discussion of the terminology issues has been given by Kolb et al. (2010).

In spite of the fact that Langmuir type adsorption is strictly only applicable to solid surfaces with fixed, periodically arranged surface sites, we also adopt this concept for liquid surfaces to represent the competition for limited surface area. The concept of Langmuir type adsorption assumes that molecules can only adsorb on free surface sites, so that the rate of adsorption ( $\text{cm}^{-2} \text{s}^{-1}$ ) is given by

$$J_{\text{ads}} = \frac{\alpha_s \bar{c}}{4} (1 - \theta) [X]_{\text{g}} \quad (3)$$

Note that this equation has been corrected from the corresponding equation in Crowley et al. (2010) that contained a typo.

$\theta$  denotes the fractional surface coverage, which is related to the surface concentration,  $[X]_{\text{S}}$ , via  $[X]_{\text{S}} = \theta \times N_{\text{max}}$ , where  $N_{\text{max}}$  is the maximum number of molecules

that can be adsorbed per  $\text{cm}^2$ . The reverse process, thermally activated desorption, is usually parameterised by a first order rate expression:

$$J_{\text{des}} = k_{\text{des}}[X]_{\text{s}} \quad (4)$$

$k_{\text{des}}$  denotes the desorption rate constant ( $\text{s}^{-1}$ ). In the absence of surface reaction or transfer to the bulk, the uptake coefficient as a function of time is given by

$$\gamma(t) = \alpha_{\text{s}} e^{-Bt} \quad (5)$$

with

$$B = \alpha_{\text{s}} \frac{\bar{c}[X]_{\text{g}}}{4N_{\text{max}}} + k_{\text{des}} \quad (6)$$

Equation (5) indicates that for low coverages, where  $k_{\text{des}} = \alpha_{\text{s}}(\bar{c}[X]_{\text{g}})/4N_{\text{max}}$  (applies for many atmospherically relevant adsorption processes), the characteristic time to reach equilibrium is given by  $1/k_{\text{des}}$ . After equilibrium has been established (see below),  $\gamma$  drops to zero. Therefore, at all concentrations, very high time resolution is necessary to observe an uptake coefficient equal to  $\alpha_{\text{s}}$ , and laboratory experiments have to be carefully evaluated to judge whether initial uptake coefficients (normally reported as  $\gamma_0$ ) may correspond to  $\alpha_{\text{s}}$  or are lower limits to  $\alpha_{\text{s}}$ . In the data sheets we tabulate reported values of  $\gamma_0$  and provide a preferred value for  $\alpha_{\text{s}}$  if appropriate and as discussed in the associated comments. Equation (5) may also be used to parameterize the temperature dependence of the uptake coefficient observed at a given time or averaged over a given time interval.

### 3.1.4 Adsorption equilibrium

The kinetics of adsorption may be relevant for extracting  $\alpha_{\text{s}}$  from laboratory measurements, but only rarely represents a rate limiting step of loss of gas phase species under

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





atmospheric conditions. Here, adsorption equilibrium is of wider importance as it may be used to estimate gas surface partitioning which defines the concentration of surface species available for surface reaction or for transfer to the bulk underneath (e.g., in liquid water clouds or aerosols). Thus adsorption of surface-active gases on aqueous solutions (Donaldson and Anderson, 1999) or of gases that form a relatively stable surface complex as intermediate to solvation (Jayne et al., 1990) or to a surface reaction (Donaldson and Valsaraj, 2010) has been described using this concept. Adsorption equilibrium is established when  $J_{\text{ads}} = J_{\text{des}}$  and can be described by:

$$\theta = \frac{N}{N_{\text{max}}} = \frac{K_{\text{LangC}}[X]_{\text{g}}}{1 + K_{\text{LangC}}[X]_{\text{g}}} \quad (7)$$

$$K_{\text{LangC}}[X]_{\text{g}} = \frac{\alpha_{\text{s}} \bar{c}}{4k_{\text{des}} N_{\text{max}}} \quad (8)$$

Equation (7) is an isotherm commonly referred to as the Langmuir isotherm, which allows derivation of the partition coefficient,  $K_{\text{LangC}}$ , from measurements of surface coverage (molecule  $\text{cm}^{-2}$ ) as a function of trace gas concentration or pressure. Equation (8) relates the partition coefficient to the kinetic parameters that determine the equilibrium on a molecular level. Therefore, measurement of  $K_{\text{LangC}}$  also provides constraints on  $\alpha_{\text{s}}$ .  $k_{\text{des}}$  is likely the main driver of the temperature dependence of the partition coefficient.

The Langmuir isotherm appears to be a reasonable approximation for adsorption characteristics on many model surfaces used in laboratory studies of adsorption of atmospheric trace gases onto both solid and liquid surfaces, in the latter case mainly for less soluble gases. We have adopted this formalism for representation of evaluated data, unless there is a gross departure from the simple picture. Note also that an expression similar to that of Eq. (7) can be used to express the equilibrium between the bulk of an aqueous solution and the surface for surface active solutes (Donaldson and

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Anderson, 1999). From a thermodynamic point of view, the equilibrium between the gas and the surface can also be expressed using a dimensionless partition coefficient,  $K_p^0$ :

$$[X]_s/[X]_g \cdot A/V = \exp(-\Delta G_{\text{ads}}^0/RT) \equiv K_p^0 \quad (9)$$

5 where  $A/V$  is the area-to-volume ratio ( $\text{cm}^{-1}$ ) of an ideal gas adsorbed at the surface ( $\sim 1.7 \times 10^7 \text{ cm}^{-1}$ ; Kemball, 1946; Kemball and Rideal, 1946; Donaldson et al., 2012) and  $\Delta G_{\text{ads}}^0$  is the free energy of adsorption.  $A/V$  defines the standard state for the adsorbed phase, which corresponds to a molar area of  $A = 3.74 \times 10^7 \text{ m}^2 \text{ mol}^{-1}$ .  $\Delta G_{\text{ads}}^0$  is related as usual via the Gibbs equation to the enthalpy and entropy of adsorption:

$$10 \quad -RT \ln K_p^0 = \Delta H_{\text{ads}}^0 - T \Delta S_{\text{ads}}^0 \quad (10)$$

The relation between  $K_p^0$  and  $K_{\text{LangC}}$  (the “C” of LangC refers to the fact that units of concentration in molecules  $\text{cm}^{-3}$  are used) is equivalent to

$$K_{\text{LangC}} = K_p^0 \cdot V/A \cdot 1/N_{\text{max}} \text{ in units of } \text{cm}^3 \text{ molecule}^{-1} \quad (11)$$

The use of gas pressures results in:

$$15 \quad K_{\text{LangP}} = K_p^0 \frac{V}{A} \frac{1}{k_b T N_{\text{max}}} \quad (12)$$

in units of  $\text{Pa}^{-1}$  or  $\text{atm}^{-1}$  or  $\text{Torr}^{-1}$  or  $\text{mbar}^{-1}$ . A modified analysis using only the linear regime of the adsorption isotherm is sometimes possible. The partition coefficient ( $K_{\text{linC}}$ ) once again has different units:

$$K_{\text{linC}} = K_p^0 \cdot V/A \quad (13)$$

20  $K_{\text{linC}}$  has units of molecule  $\text{cm}^{-2}/\text{molecule cm}^{-3}$  (or  $\text{cm}$ ) if concentrations (in molecule  $\text{cm}^{-3}$ ) are used, and units of molecule  $\text{cm}^{-2} \text{ Pa}^{-1}$  or molecule  $\text{cm}^{-2} \text{ Torr}^{-1}$

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

if pressure is used, then denoted as  $K_{\text{linP}}$ . Fractional surface coverages can be calculated from each different form of the partition coefficient via:

$$\Theta_X = K_0^p \frac{V}{A} \frac{p_x}{k_b T N_{\text{max}}} = K_0^p \frac{V}{A} \frac{[X]_g}{N_{\text{max}}} = K_{\text{linC}} \frac{[X]_g}{N_{\text{max}}} = K_{\text{linP}} \frac{p_x}{N_{\text{max}}} = K_{\text{LangC}} [X]_g = K_{\text{LangP}} p_x \quad (14)$$

For the purpose of comparing partition coefficients and deriving preferred expressions for calculating equilibrium surface coverages, a single form of the partition coefficient is required. In principle  $K_p^0$  would be the best choice; the disadvantage is that it is difficult to extract from the various studies if experimental surface to volume ratios ( $S_{\text{exp}}/V_{\text{exp}}$ ) are not known, or if  $N_{\text{max}}$  has to be chosen arbitrarily and not from the experiment itself. Therefore, for practical purposes (e.g., using the constants to calculate surface coverages) reporting consistently in the form of  $K_{\text{linC}}$  has the advantage that no inherent assumption about  $N_{\text{max}}$  has to be made to derive partitioning from the experiments at low trace gas pressures. The tabulated values of partition coefficients are therefore presented as  $K_{\text{linC}}$ . The accompanying notes in each data sheet provide the original expressions and the values of  $N_{\text{max}}$  and  $V/A$  used to calculate  $K_{\text{linC}}$ .

If data are available from the low coverage (linear) part of the isotherm these are used preferably, as the influence of lateral interactions is reduced, and  $N_{\text{max}}$ , which is often difficult to obtain experimentally due to adsorbate-adsorbate interactions (Jedlovsky et al., 2006), is not required. The Van't Hoff equation (i.e., the differential form of the Gibbs equation):

$$\frac{d(\ln K_p^0)}{d(1/T)} = \frac{-\Delta H^0}{R} \quad (15)$$

describes the temperature dependence of  $K$ , but the entropy and enthalpy of adsorption derived from this analysis can be coverage dependent and this needs to be considered when comparing results from different experiments.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



### 3.1.5 Surface reactions

Parameterisation of uptake resulting from reaction of trace gases on a surface (reactive uptake) requires knowledge of the mechanism of the reaction. Generally there are two types of reaction. First, those in which the molecule arriving at the surface reacts with a constituent from the bulk phase to form either involatile products (e.g., stable hydrates) or volatile products which partition back to the gas phase. Second, those in which the molecule arriving at the surface reacts with a second species, which is present on the surface in the adsorbed state to form either involatile or volatile products. In both types the kinetics depend on the surface concentration of the second reactant, which must be included in any parameterisation of the reactive uptake coefficient. The two mechanisms, which are commonly used to describe the kinetics of surface reactions, are often referred to as Langmuir-Hinshelwood (LH) and Eley-Rideal mechanisms (ER). It should be noted that some reports in the literature use the term Langmuir-Rideal for the latter and that the reverse abbreviation RE has also been used (see also IUPAC Gold Book). They differ in that the LH mechanism involves adsorption of the arriving molecule at available surface sites prior to a bimolecular reaction. The ER mechanism involves direct collision induced reaction of the arriving gaseous molecule with a reactant molecule present on the surface. The LH mechanism thus may involve competition of both reactant molecules for available surface sites. Moreover on the surface of bulk liquids, the second reactant molecule may originate in the gas phase or the bulk liquid phase. For parameterisation of the uptake coefficients we adopt the approach presented by Ammann et al. (2003), which builds on earlier studies (Elliott et al., 1991; Mozurkewich, 1993; Tabazadeh and Turco, 1993; Carslaw and Peter, 1997).

#### Bimolecular reaction between two surface species

The parameterisation for the reactive uptake coefficient,  $\gamma$ , for gas phase species X reacting with surface species, Y, after adsorption on the surface (LH mechanism) is

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



given by:

$$\Gamma_s = \frac{4k_s[Y]_s K_{\text{LangC}}(X) N_{\text{max}}}{\bar{c} (1 + K_{\text{LangC}}(X) [X]_g)} \quad (16)$$

Here  $[Y]_s$  is the surface concentration (molecules  $\text{cm}^{-2}$ ) of species Y and  $k_s$  is the surface reaction rate coefficient (units of  $\text{cm}^2 \text{molecule}^{-1} \text{s}^{-1}$ ).  $1/\Gamma_s$  can be considered the resistance for the surface reaction. Note that  $\Gamma_s$  represents a normalized rate and is therefore not restricted to values smaller than one. If Y is a volatile molecule also present in the gas phase, its equilibrium surface concentration,  $[Y]_s$ , can be calculated using an appropriate adsorption isotherm. Values of  $k_s K_{\text{LangC}}(X)$  can be determined experimentally from measurements of  $\gamma$  as a function of surface coverage  $[Y]_s$ . Equation (16) demonstrates that  $\gamma$  depends on the gas phase concentration of X, if  $K_{\text{LangC}}(X)[X]_g$  is similar to or larger than 1 (i.e., at high coverage). This is especially important when interpreting data from laboratory experiments performed using gas-phase reactant concentrations which lead to significant surface coverage. In the data sheets, we provide preferred values for  $\alpha_s$  and  $k_s$ .

## Direct gas surface reaction

Parameterisation for the reactive uptake coefficient,  $\gamma$ , for gas phase species X directly reacting with surface species, Y, upon collision (ER mechanism) is given by:

$$\gamma_{\text{gs}} = \gamma_{\text{gs}}^{\theta=1}(X) \times \frac{[Y]_s}{N_{\text{max}}^Y} \quad (17)$$

Here  $\gamma_{\text{gs}}$  is the elementary reaction probability that a gas phase molecule X colliding with surface component Y reacts with it.  $N_{\text{max}}^Y(Y)$  denotes the maximum coverage of Y for a volatile species Y in equilibrium with the gas phase.  $\gamma$  can be calculated for a

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



given gas phase concentration of Y if values of  $K_{\text{LangC}}(\text{Y})$  and  $\gamma_{\text{gs}}$  are available. In the absence of other rate limiting reactions, values of  $\gamma_{\text{gs}}$  can be determined experimentally from measurements of  $\gamma$  as a function of surface coverage near saturation, or by extrapolation using an appropriate adsorption isotherm. In the ER mechanism,  $\gamma$  does not depend on the gas phase concentration of X, whereas in the LH type mechanism (at high  $[\text{X}]_{\text{g}}$ ), it does. Also the temperature dependence, driven by the temperature dependence of  $K_{\text{LangC}}(\text{X})$  in the LH case, is substantially different (Pöschl et al., 2007; Ammann and Pöschl, 2007).

### 3.1.6 Exchange with the bulk and bulk accommodation

Previously, the process of transfer of a gas molecule from the gas phase into the bulk of a liquid has been considered a quasi-elementary process, and the mass accommodation coefficient has been defined as the probability that a molecule colliding with the surface is actually taken up into the bulk, mostly in relation with liquids. However, in many cases it is necessary to decouple this process into adsorption on the surface and surface to bulk transfer, i.e., dissolution (Davidovits et al., 1995; Hanson, 1997).

With the aim to clearly differentiate uptake into the bulk from surface accommodation, the term bulk accommodation is recommended, and the corresponding coefficient as bulk accommodation coefficient,  $\alpha_{\text{b}}$  (Kolb et al., 2010). In the absence of surface reactions,

$$\alpha_{\text{b}} = \alpha_{\text{s}}(k_{\text{sb}}/(k_{\text{sb}} + k_{\text{des}})) \quad (18)$$

In Eq. (18),  $k_{\text{sb}}$  denotes the surface to bulk transfer rate coefficient in units of  $\text{s}^{-1}$ . If the adsorption equilibrium is established much faster than transfer to the bulk, the two processes can be expressed in the form of separated resistances:

$$\frac{1}{\alpha_{\text{b}}} = \frac{1}{\alpha_{\text{s}}} + \frac{1}{\Gamma_{\text{sb}}} \quad \text{with} \quad \Gamma_{\text{sb}} = \frac{\alpha_{\text{s}}k_{\text{sb}}}{k_{\text{des}}} \quad (19)$$

Equations (18) and (19) also allow parameterization of the temperature dependence of  $\alpha_b$ , which has been used as a proof of the nature of bulk accommodation as a coupled process (Davidovits et al., 1995).

### 3.1.7 Solubility equilibrium

If the rates of transfer from the surface to the bulk and back are equal, solubility equilibrium is established:

$$K_{cc} = K_{cp}RT = [X]_{b,eq}/[X]_g \approx H \quad (20)$$

$K_{cc}$  denotes the solubility equilibrium constant in dimensionless units,  $K_{cp}$  in pressure/molarity based units (Sander, 1999). For infinitely dilute solutions,  $K_{cc}$  or  $K_{cp}$  becomes equivalent to the Henry's Law constant:  $K_{cp} \approx H$  in units of  $\text{M atm}^{-1}$ . This equilibrium can be related to the kinetic parameters via (Hanson, 1997; Pöschl et al., 2007):

$$K_{cc} = \frac{k_{sb}}{k_{bs}} \frac{k_{ads}}{k_{des}} = \frac{k_{sb}}{k_{bs}} \frac{\alpha_s(1-\theta)\bar{c}}{4k_{des}} \quad (21)$$

$k_{bs}$  denotes the rate coefficient associated with bulk to surface transfer, i.e., desolvation. As evident from this expression, solubility becomes linked to the surface coverage, adding the constraint that Henry's Law is only valid for sufficiently dilute conditions also at the surface. Examples of deviations have been shown for  $\text{SO}_2$  (Jayne et al., 1990) at high pressures or for surface active organics (Djikaev and Tabazadeh, 2003). A more detailed discussion of such cases has been given by Ammann and Pöschl (2007).

For the majority of cases in this evaluation, the link of solubility to adsorption is either not relevant or has not been considered in the experiments. The most direct measurement of  $H$  is expressed in Eq. (20), i.e., through measurement of both liquid phase and gas phase concentrations at equilibrium. Values from such measurements are used when available to evaluate kinetic parameters. Note that we quote values of  $H$  from the literature but do not evaluate  $H$  measurements themselves.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



### 3.1.8 Solubility limited uptake into the bulk

In the absence of surface or bulk reactions, uptake into the bulk of liquid particles proceeds until the solubility equilibrium is reached. Under quasi-steady state conditions, the uptake coefficient can be described by Eq. (22):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{sol}}} \quad \text{with} \quad \Gamma_{\text{sol}} = \frac{4HRT}{\bar{c}\sqrt{\pi}} \sqrt{\frac{D_l}{t}} \quad (22)$$

$H$  denotes the Henrys Law coefficient ( $\text{M atm}^{-1}$ ),  $R$  the gas constant ( $1 \text{ atm mol}^{-1} \text{ K}^{-1}$ ) and  $D_l$  the liquid phase diffusion coefficient ( $\text{cm}^2 \text{ s}^{-1}$ ) (Schwartz, 1986). Solubility limited uptake can be used to measure the product  $H(D_l)^{0.5}$ , through the time dependence of the observed  $\gamma$ . If independent measurements of the solubility are available, a value for the diffusion coefficient can be obtained. Otherwise, the diffusion coefficient can be estimated from viscosity data to get an estimate for  $H$  (see also Sect. 2.2).

### 3.1.9 Reactive uptake into the bulk

For atmospheric trace gas uptake to the bulk of liquid (usually aqueous) particles, in the absence of a surface reaction,

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_b} \quad \text{with} \quad \Gamma_b = \frac{4HRT}{\bar{c}} \sqrt{D_l k'_b} \quad (23)$$

$k'_b$  is the pseudo first-order bulk reaction rate coefficient. For small droplets the concept of the reacto-diffusive length is important. This is the characteristic distance beyond the surface of the particle in which the trace gas is lost by reaction and is given by:

$$l = \sqrt{\frac{D_l}{k'_b}} \quad (24)$$



For spherical particles, Eq. (23) can be modified to account for this with:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{c}{4HRT \sqrt{D_1 \cdot k'_b} [\coth(r/l) - (l/r)]} \quad (25)$$

This is the basis of the “Framework” paper (Hanson et al., 1994) for modelling uptake rates to stratospheric aerosol. Note that for a molecule that can dissociate in the aqueous phase, an effective solubility,  $H^*$ , is used whereby:

$$H^* = H \times \left( \frac{1 + K_a [H^+]_e}{K_w} \right) \quad (26)$$

$K_w$  is the autoprotolysis constant of  $H_2O$ ,  $K_a$  is the acid dissociation constant of the trace gas in water and  $[H^+]_e$  is the equilibrium hydrogen ion concentration at the droplet surface.

In the limiting case for small liquid particles or mixed phase particles when the volume of the liquid phase is such that  $l < r$ , the reactants fill the entire liquid phase and uptake rate becomes volume limited. In this case the diffusion restriction in the bulk term in Eq. (23) disappears. If  $\alpha_b \gg \Gamma_b$  this leads to a linear dependence of the uptake coefficient on the volume/surface ratio of the liquid phase. For spherical geometry Eq. (25) simplifies to:

$$\gamma = \frac{4HRT k'_b \left( \frac{V_l}{S_a} \right)}{\bar{c}} \quad (27)$$

where  $V_l$  and  $S_a$  are the volume and surface area of the liquid phase, respectively.

### 3.1.10 Coupled processes on the surface and in the bulk

When processes occur on both the surface and in the bulk, adsorption and transfer to the bulk have to be separated. Following previous derivations (Ammann et al., 2003;

Davidovits et al., 1995; Hanson, 1997; Jayne et al., 1990; Pöschl et al., 2007; Shi et al., 1999), coupling a Langmuir-Hinshelwood type surface reaction with a reaction in the bulk leads to:

$$\frac{1}{\gamma} = \frac{1}{\alpha_s} + \frac{1}{\Gamma_s + \left(\frac{1}{\Gamma_{sb}} + \frac{1}{\Gamma_b}\right)^{-1}} \quad (28)$$

5 In the case of an [Eley-Rideal](#) reaction which occurs without prior adsorption or surface to bulk transfer, the resistance acts in parallel to adsorption, leading to an expression proposed by Pöschl et al. (2007):

$$\frac{1}{\gamma} = \frac{1}{\gamma_{gs} + \left(\frac{1}{\alpha_s} + \frac{1}{\Gamma_{sb}} + \frac{1}{\Gamma_b}\right)^{-1}} \quad (29)$$

10 Note that while Eq. (29) is consistent with the resistor diagram of Hu et al. (1995), it is not consistent with their expression for  $1/\gamma$ . In follow-up papers, e.g., Shi et al. (1999), Eq. (29) was used.

## 3.2 Liquid surfaces considered

### 3.2.1 [Sulfuric acid](#)

15 Stratospheric aerosols consist mainly of sub-micron size liquid particles containing aqueous concentrated  $\text{H}_2\text{SO}_4$  (50 to 80 wt%) at low temperature (185–260 K). The water content depends in a predictable way on the relative humidity, which is defined by the temperature and the absolute water concentration. Substrates with well-characterised surface area and acid content are relatively straightforward to investigate experimentally, either as bulk substrates or as an aerosol of sub-micron particles of defined size distribution. The latter form minimises the limiting effects of gas-phase diffusion on the uptake rates, which can be a problem when bulk surfaces are used and  $\gamma$  values are large.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



In spite of the chemical simplicity of this substrate, parameterization of uptake and reaction of trace gases on and within sulfuric acid is fairly complex due to the strong dependence of physical and acidic properties of sulfuric acid on composition, which itself is a strong function of humidity and temperature. The thermodynamic basis for describing sulfuric acid composition for atmospheric conditions has been given in detail by Carslaw et al. (1995).  $\text{H}_2\text{SO}_4$  content is mostly given in units of weight percent (wt %), then referred to as *wt*, or mole fraction *X*, which can be interconverted using

$$X = wt / (wt + (100 - wt)98/18) \quad (30)$$

Shi et al. (2001) present a parameterization for the water activity:

$$a_w = \exp[(-69.775X - 18253.7X^2 + 31072.2X^3 - 25668.8X^4)(1/T - 26.9033/T^2)] \quad (31)$$

Since many of the relevant reactions in sulfuric acid solutions are acid catalyzed Shi et al. also suggest a parameterization of acid activity in units of  $\text{mol L}^{-1}$ , which also extends the Carslaw et al. model to dilute solutions:

$$a_{\text{H}^+} = \exp[60.51 - 0.095 wt + 0.0077 wt^2 - 1.61 \times 10^{-5} wt^3 - (1.76 + 2.52 \times 10^{-4} wt^2)T^{0.5} + (-805.89 + 253.05 wt^{0.076})/T^{0.5}] \quad (32)$$

An inherent physical property influencing the rate of stratospheric heterogeneous processes is the viscosity,  $\eta$ , in units of cP. Williams and Long (1995) provide direct measurements of viscosity for a range of relevant sulfuric acid compositions and temperatures. Again, combining with data from dilute solutions, Shi et al. extended the fit of Williams and Long to an empirical equation originally derived by Eicher and Zwolinski (1971) to more dilute solutions:

$$\eta = AT^{-1.43} \exp(448 \text{ K} / (T - T_0)) \quad (33)$$

$$A(\text{Matm}^{-1} \text{K}^{1.43}) = 169.5 + 5.18 wt - 0.0825 wt^2 + 3.27 \times 10^{-3} wt^3 \quad (34)$$

$$T_0(K) = 144.11 + 0.166 wt - 0.015 wt^2 + 2.18 \times 10^{-4} wt^3 \quad (35)$$

The viscosity calculated from these parameterisations can be used to obtain an estimate for the diffusion coefficient using the Stokes-Einstein equation:

$$D_1 = cT/\eta \quad (36)$$

The constant  $c$  has been either determined from fits to experimentally measured diffusion coefficients or has been estimated. Values are given in the individual datasheets.

### 3.2.2 Aqueous aerosols

Aerosol particles in the troposphere frequently consist of water-soluble electrolyte salts. Over marine regions the aerosol is dominated by sea salt aerosol, with NaCl as the main constituent. In continental regions ammonium sulphate, ammonium bisulphate, and ammonium nitrate are major inorganic constituents of the aerosol. The phase of the particles depends on relative humidity. At low humidity, below efflorescence RH, the particles can contain solid crystals. At higher humidity ( $\geq 40\%$ ) the particles exist as aqueous droplets and this is the predominant form in the troposphere. The dependence of the particle composition (e.g. water content) on relative humidity and concentrations of ionic constituents (and selected soluble organics) can be calculated using the Extended Aerosol Thermodynamics Model (AIM) (Wexler and Clegg, 2002) (<http://www.aim.env.uea.ac.uk/aim/aim.php>). Aqueous tropospheric particles may also contain a large range of organic solutes. Smaller carboxylic acids lead to comparable solutions, and they have also been included in the AIM. As far as they have been used in trace gas uptake studies, they are dealt with in the category "other aqueous electrolytes" in the present evaluation. For larger organic solutes, we note however that the physical state is not always well defined as they may form highly viscous semi-solid phases or even glasses at low humidity or low temperature (Koop et al., 2011; Zobrist et al., 2008). These substrates are not covered by the present evaluation.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

---

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Liquid substrate surfaces with well characterised surface area and electrolyte content are relatively straightforward to prepare, either as bulk substrates in wetted wall or droplet train configuration, or as an aerosol of sub-micron particles of defined size distribution. In the atmosphere mixed phase particles may occur. The rate parameters given in this evaluation can be used for uptake into particles which have predominantly liquid (aqueous) phase, at least near the surface. In this case the spherical geometry applies but the size threshold for volume limited uptake (see Eq. 27) may be affected. This may be the preferred parameterisation for small multicomponent aerosol particles in the troposphere. For particles of predominantly solid phase we refer to recommendations in ACP V (Crowley et al., 2010). Under most ambient conditions the aqueous phase diffusion coefficient for most trace gases for the substrates used in the present evaluation is close to  $10^{-5} \text{ cm}^2 \text{ s}^{-1}$ . Individual values required for parameterisation of  $\gamma$  are listed in the datasheets. Exceptions are sulfuric acid solutions at low temperature dealt with in the previous subsection or highly viscous solutions (Zobrist et al., 2008).

### 3.3 Surface areas

$\gamma$  is usually obtained from observed loss rates of gas-phase species and calculated collision rates with the surface. For the latter the surface area available for uptake/reaction ( $A$ ) is required. For liquid surfaces this generally poses no problems as the surface is “smooth” at the molecular level and the geometric surface area ( $A_{\text{geom}}$ ) is generally applicable and readily defined. The use of aerosol particles rather than bulk surfaces in principle provides a better mimic of atmospheric conditions for experimental surface area and may provide confirmation of particle size dependent uptake coefficients when reaction and diffusion compete.

### 3.4 Methods used in the study of heterogeneous processes

A number of experimental techniques have been developed for the study of heterogeneous processes. Most methods rely on the determination of loss rates or time de-

pendent concentration changes of gas-phase species in contact with a surface. These include low pressure, coated surface laminar flow tube reactors and Knudsen cells for bulk surfaces and films, and droplet train reactors, aerosol flow tubes and static aerosol chambers for dispersed surfaces. Surface adsorbed reactants and products have frequently been observed using surface-sensitive techniques, such as reflectance infrared spectroscopy (DRIFTS, RAIRS), and these have in a few cases been applied to kinetics studies. A list of the related abbreviations used in the data sheets is given in Table 2 below.

### 3.5 Organisation of the datasheets

The basic structure of the heterogeneous datasheets is similar to the well established datasheets on homogeneous gas-phase reactions. For each heterogeneous interaction on a particular substrate we give a list of experimentally measured values of the uptake coefficients ( $\gamma$ ), accommodation coefficients ( $\alpha$ ), partition coefficients ( $H$ ,  $K_{\text{linC}}$ ), and diffusion coefficients ( $D_l$ ), where appropriate for physical uptake. For reactive uptake reported values for rate coefficients for reaction in the liquid phase, including their temperature and composition dependence where known. The parameters are distinguished according to the definitions given above (e.g.,  $\alpha_s$ ,  $\alpha_b$ ,  $\gamma_0$ ,  $\gamma_{\text{SS}}$ ,  $H^*$ ,  $k_r$ , etc.). Also given are references and pertinent information about the technique, reactant concentration, substrate and other conditions in linked comments.

Recommended values for the uptake coefficient,  $\gamma$  and/or for parameters needed for its calculation (e.g.  $\alpha_s$ ,  $\alpha_b$ ,  $k'_r$ ,  $H$ ,  $K_{\text{linC}}$ ,  $N_{\text{max}}$ ,  $D_l$ ,  $\gamma_{\text{GS}}$  etc.) using an expression specified in the text, are provided. Recommended temperature dependencies are given in Arrhenius form when appropriate, together with our estimate of the uncertainties in the parameters. Comments on the state of the data set and justification of the recommendations are then presented, followed by a list of references cited. A list of the symbols and description of the parameters used in the datasheets is given in Table 1 below. The datasheets available at the website (<http://www.iupac-kinetic.ch.cam.ac.uk>) also con-

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

tain figures where appropriate to illustrate comparison between our recommendation and available experimental data.

The ordering of reacting species within each surface category follows that adopted for gas phase reactions, i.e.  $O_x$ ,  $HO_x$ ,  $NO_x$ ,  $SO_x$ , organics, and halogenated species.

5 The website also offers search options to access individual datasheets.

### 3.6 Citation

The citation for recommendations in the data sheets and summary tables is: ACP journal reference (e.g., Ammann, M., Cox, R. A., Crowley, J. N., Jenkin, M. E., Mellouki, A., Rossi, M. J., Troe, J., and Wallington, T. J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume VI – heterogeneous reactions with liquid substrates, Atmos. Chem. Phys. Discuss., 12, 32109–32472, 2012) + web-  
10 site address: IUPAC Task Group on Atmospheric Chemical Kinetic Data Evaluation, <http://www.iupac-kinetic.ch.cam.ac.uk>.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table 1.** Parameters used to describe heterogeneous reactions.

Parameter	Description	Units	Notes
$\gamma$ ( $\gamma_{ss}$ , $\gamma_0$ )	Net uptake coefficient	–	a
$\Gamma_b$	Limiting uptake coefficient for bulk reaction (liquid)		
$\Gamma_{sol}$	Limiting uptake coefficient for dissolution (liquid)		
$\Gamma_{diff}$	Limiting uptake coefficient for gas-phase diffusion		
$\Gamma_{sb}$	Limiting uptake coefficient for surface to bulk transfer		
$\alpha_s$	Surface accommodation coefficient	–	b
$\alpha_t$	Thermal accommodation coefficient		
$\gamma_{gs}$	Elementary gas-surface reactive uptake coefficient (or $\gamma_{ER}$ )	–	
$k_s$	Surface reaction rate coefficient	$\text{cm}^2 \text{ molecule}^{-1} \text{ s}^{-1}$	
$\alpha_b$	Bulk accommodation coefficient		c
$\varepsilon$	Evaporation coefficient		
$K_{inc}$	Gas-Surface partition coefficient(liquid surfaces)	$\text{cm}^{-2}/\text{cm}^{-3}$	d
$K_v$	Gas-Volume partition coefficient(liquid surfaces)	$\text{cm}^{-3}/\text{cm}^{-3}$	e
$k_r'$	bulk liquid phase reaction rate coefficient	$\text{cm}^3 \text{ s}^{-1}$	
$H$	Solubility (Henry)	$\text{mol}/\text{L}\cdot\text{atm}$	
$D_l$	Liquid phase diffusion coefficient	$\text{cm}^2 \text{ s}^{-1}$	
$D_g$	Gas phase diffusion coefficient	$\text{cm}^2 \text{ s}^{-1}$	
$N$	Surface coverage	$\text{molecule cm}^{-2}$	
$N_{max}$	Surface coverage at saturation	$\text{molecule cm}^{-2}$	

## Notes

- 5 a: As  $\gamma$  can be time dependent, subdivisions are necessary, whereby  $\gamma_0$  is the experimentally observed, initial (frequently maximum) uptake coefficient; and  $\gamma_{ss}$  is the experimentally observed, steady state uptake coefficient.
- b: The probability (per collision) that a gas phase molecule impinging on the solid surface resides on the surface for a finite time ( $\gg 10^{-12}$  s).
- 10 c: The probability (per collision) that a gas phase molecule impinging on the liquid surface enters the liquid.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



d: The partition coefficient that describes the gas-surface partitioning at equilibrium. As described above, the Langmuir isotherm is most commonly used in various units. We report the partitioning coefficient in the limit of low coverage (linear dependence of coverage on gas concentration) where the units are as given above.

- 5 e: The partition coefficient that describes the distribution of trace gas between the gas phase and condensed phase volumes. For a liquid particle this is the solubility in the volume of the particle.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Table 2.** Techniques used to study heterogeneous reactions.

Methods	
CWFT	Coated Wall Flow Tube
WWFT	Wetted Wall Flow Tube
RWFT	Rotating Wetted Wall Flow Tube
CRFT	Coated Rod Flow Tube
AFT	Aerosol Flow Tube
Knudsen	Knudsen Reactor
DT	Droplet Train
LJ	Liquid Jet
BC	Bubble Column
PBFT	Packed Bed Flow Tube
Detection	
AMS	Aerosol Mass Spectrometry
ATR	Attenuated Total Reflectance spectroscopy
APS	Aerosol Particle Sizer
DRIFTS	Diffuse Reflectance Infra-red Fourier Transform Spectroscopy
DMA	Differential Mobility Analyser
RF	Resonance Fluorescence
CL	Chemi-Luminescence
GC	Gas Chromatography
MS	Mass Spectrometry
MBMS	Molecular Beam Sampling MS
CIMS	Chemical Ionisation Mass Spectrometry
UV-Vis	Ultra-Violet-Visible Spectroscopy
SR	Static Reactor
SMPS	Scanning Mobility Particle Sizer
TDL	Tunable Diode Laser absorption spectroscopy
TIR	Transmission Infra Red spectroscopy
TEM	Transmission Electron Microscopy
RC	Counting of decays of radioactive isotopes

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

*Acknowledgements.* The Chairman and members of the Subcommittee wish to express their appreciation to I.U.P.A.C. for the financial help which facilitated the preparation of this evaluation. We also acknowledge financial support from the following organisations: EU Framework Program 6 and 7, ACCENT network of excellence; the European Science Foundation (INTROP program), the UK Natural Environmental Research Council; the Swiss National Science Foundation (Projects 200020\_125204 and 200021\_130175), and the Deutsche Forschungsgemeinschaft (SFB 357).

## References

- Ammann, M., Pöschl, U., and Rudich, Y.: Effects of reversible adsorption and Langmuir-Hinshelwood surface reactions on gas uptake by atmospheric particles, *Phys. Chem. Chem. Phys.*, 5, 351–356, 2003.
- Ammann, M. and Pöschl, U.: Kinetic model framework for aerosol and cloud surface chemistry and gas-particle interactions – Part 2: Exemplary practical applications and numerical simulations, *Atmos. Chem. Phys.*, 7, 6025–6045, doi:10.5194/acp-7-6025-2007, 2007.
- Behr, P., Terziyski, A., and Zellner, R.: Reversible gas adsorption in coated wall flow tube reactors, Model simulations for Langmuir kinetics, *Z. Phys. Chem. Int. J. Res. Phys. Chem. Chem. Phys.*, 218, 1307–1327, 2004.
- Brown, R. L.: Tubular flow reactor with first-order kinetics, *J. Res. Natl. Bureau St.*, 83, 1–8, 1978.
- Carslaw, K. S., Clegg, S. L., and Brimblecombe, P.: A thermodynamic model of the system HCl-HNO<sub>3</sub>-H<sub>2</sub>SO<sub>4</sub>-H<sub>2</sub>O, including solubilities of HBr, from <200 to 328 K, *J. Phys. Chem.*, 99, 11557–11574, 1995.
- Carslaw, K. S. and Peter, T.: Uncertainties in reactive uptake coefficients for solid stratospheric particles. 1. Surface chemistry, *Geophys. Res. Lett.*, 24, 1743–1746, 1997.
- Compilation of Henry's Law Constants for Inorganic and Organic Species of Potential Importance in Environmental Chemistry (Version 3): <http://www.henrys-law.org>, 1999.
- Crowley, J. N., Ammann, M., Cox, R. A., Hynes, R. G., Jenkin, M. E., Mellouki, A., Rossi, M. J., Troe, J., and Wallington, T. J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume V – heterogeneous reactions on solid substrates, *Atmos. Chem. Phys.*, 10, 9059–9223, doi:10.5194/acp-10-9059-2010, 2010.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Davidovits, P., Hu, J. H., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: Entry of gas molecules into liquids, *Faraday Discuss.*, 100, 65–81, 1995.
- Djikaev, Y. S. and Tabazadeh, A.: Effect of adsorption on the uptake of organic trace gas by cloud droplets, *J. Geophys. Res.-Atmos.*, 108, art. no. 4689, 2003.
- 5 Donaldson, D. J. and Anderson, D.: Adsorption of atmospheric gases at the air-water interface. 2. C-1-C-4 alcohols, acids, and acetone, *J. Phys. Chem. A*, 103, 871–876, 1999.
- Donaldson, D. J. and Valsaraj, K. T.: Adsorption and Reaction of Trace Gas-Phase Organic Compounds on Atmospheric Water Film Surfaces: A Critical Review, *Env. Sci. Technol.*, 44, 865–873, 2010.
- 10 Donaldson, D. J., Ammann, M., Bartels-Rausch, T., and Pöschl, U.: Standard states and thermochemical kinetics in heterogeneous atmospheric chemistry, *J. Phys. Chem. A*, 116, 6323–6316, 2012.
- Eicher, L. D. and Zwolinski, B. J.: High-Precision Viscosity of Supercooled Water and Analysis of Extended Range Temperature Coefficient, *J. Phys. Chem.*, 75, 2016–2024, 1971.
- 15 Finlayson-Pitts, B. J. and Pitts, J. N., Jr.: *Chemistry of the upper and lower atmosphere*, Academic Press, San Diego, CA, 2000.
- Flückiger, B. and Rossi, M. J.: Common precursor mechanism for the heterogeneous reaction of D<sub>2</sub>O, HCl, HBr, and HOBr with water ice in the range 170–230 K: Mass accommodation coefficients on ice, *J. Phys. Chem. A*, 107, 4103–4115, 2003.
- 20 Fuchs, N. A. and Sutugin, A. G.: *Highly dispersed aerosols*, Ann Arbor Science Publishers, Inc., Ann Arbor, London, 1970.
- Garrett, B. C., Schenter, G. K., and Morita, A.: Molecular Simulations of the Transport of Molecules across the Liquid/Vapor Interface of Water, *Chem. Rev.*, 106, 1355–1374, 2006.
- Hanson, D. R., Ravishankara, A. R., and Solomon, S.: Heterogeneous Reactions In Sulfuric Acid Aerosols – A Framework For Model-Calculations, *J. Geophys. Res.-Atmos.*, 99, 3615–3629, 1994.
- 25 Hanson, D. R.: Surface specific reactions on liquids, *J. Phys. Chem. B*, 101, 4998–5001, 1997.
- Hu, J. H., Shi, Q., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: Reactive uptake of Cl<sub>2</sub>(g) and Br<sub>2</sub>(g) by aqueous surfaces as a function of Br- and I-ion concentration: the effect of chemical reaction at the interface, *J. Phys. Chem.*, 99, 8768–8776, 1995.
- 30 Huff, A. K. and Abbatt, J. P. D.: Gas-Phase Br<sub>2</sub> production in heterogeneous reactions of Cl<sub>2</sub>, HOCl and BrCl with halide-ice surfaces, *J. Phys. Chem. A*, 104, 7284–7293, 2000.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Jayne, J. T., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: Uptake of SO<sub>2</sub>(g) by Aqueous Surfaces as a Function of pH – the Effect of Chemical-Reaction at the Interface, *J. Phys. Chem.*, 94, 6041–6048, 1990.
- 5 Kolb, C. E., Cox, R. A., Abbatt, J. P. D., Ammann, M., Davis, E. J., Donaldson, D. J., Garrett, B. C., George, C., Griffiths, P. T., Hanson, D. R., Kulmala, M., McFiggans, G., Pöschl, U., Riipinen, I., Rossi, M. J., Rudich, Y., Wagner, P. E., Winkler, P. M., Worsnop, D. R., and O'Dowd, C. D.: An overview of current issues in the uptake of atmospheric trace gases by aerosols and clouds, *Atmos. Chem. Phys.*, 10, 10561–10605, doi:10.5194/acp-10-10561-2010, 2010.
- 10 Koop, T., Bookhold, J., Shiraiwa, M., and Pöschl, U.: Glass transition and phase state of organic compounds: dependency on molecular properties and implications for secondary organic aerosols in the atmosphere, *Phys. Chem. Chem. Phys.*, 13, 19238–19255, 2011.
- Li, Y. Q., Davidovits, P., Shi, Q., Jayne, J. T., Kolb, C. E., and Worsnop, D. R.: Mass and thermal accommodation coefficients of H<sub>2</sub>O(g) on liquid water as a function of temperature, *J. Phys. Chem. A*, 105, 10627–10634, 2001.
- 15 Masel, R. I.: Principles of adsorption and reaction on solid surfaces, 1st ed., John Wiley & Sons, New York, 804 pp., 1996.
- Morris, J. R., Behr, P., Antman, M. D., Ringeisen, B. R., Splan, J., and Nathanson, G. M.: Molecular beam scattering from supercooled sulfuric acid: Collisions of HCl, HBr, and HNO<sub>3</sub> with 70 wt% D<sub>2</sub>SO<sub>4</sub>, *J. Phys. Chem. A*, 104, 6738–6751, 2000.
- 20 Nathanson, G. M., Davidovits, P., Worsnop, D. R., and Kolb, C. E.: Dynamics and kinetics at the gas-liquid interface, *J. Phys. Chem.*, 100, 13007–13020, 1996.
- Pfrang, C., Shiraiwa, M., and Pöschl, U.: Coupling aerosol surface and bulk chemistry with a kinetic double layer model (K2-SUB): oxidation of oleic acid by ozone, *Atmos. Chem. Phys.*, 10, 4537–4557, doi:10.5194/acp-10-4537-2010, 2010.
- 25 Pöschl, U., Rudich, Y., and Ammann, M.: Kinetic model framework for aerosol and cloud surface chemistry and gas-particle interactions – Part 1: General equations, parameters, and terminology, *Atmos. Chem. Phys.*, 7, 5989–6023, doi:10.5194/acp-7-5989-2007, 2007.
- Santschi, C. and Rossi, M. J.: The heterogeneous interaction of HOCl with solid KBr substrates: The catalytic role of adsorbed halogens, *Phys. Chem. Chem. Phys.*, 7, 2599–2609, 2005.
- 30 Schwartz, S. E.: Mass transport considerations pertinent to aqueous phase reactions of gases in liquid water clouds, in: NATO ASI Ser., edited by: Jaeschke, W., Springer-Verlag, New York, 415–471, 1986.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Seinfeld, J. H. and Pandis, S. N.: Atmospheric chemistry and physics: From air pollution to climate change, John Wiley & Sons, Inc., New York, Chichester, Weinheim, Brisbane, Singapore, Toronto, 1998.

Shi, Q., Davidovits, P., Jayne, J. T., Worsnop, D. R., and Kolb, C. E.: Uptake of gas-phase ammonia. 1. Uptake by aqueous surfaces as a function of pH, *J. Phys. Chem. A*, 103, 8812–8823, 1999.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: Kinetic model for reaction of ClONO<sub>2</sub> with H<sub>2</sub>O and HCl and HOCl with HCl in sulfuric acid solutions, *J. Geophys. Res.-Atmos.*, 106, 24259–24274, 2001.

Shiraiwa, M., Garland, R. M., and Pöschl, U.: Kinetic double-layer model of aerosol surface chemistry and gas-particle interactions (K2-SURF): Degradation of polycyclic aromatic hydrocarbons exposed to O<sub>3</sub>, NO<sub>2</sub>, H<sub>2</sub>O, OH and NO<sub>3</sub>, *Atmos. Chem. Phys.*, 9, 9571–9586, doi:10.5194/acp-9-9571-2009, 2009.

Shiraiwa, M., Pfrang, C., and Pöschl, U.: Kinetic multi-layer model of aerosol surface and bulk chemistry (KM-SUB): the influence of interfacial transport and bulk diffusion on the oxidation of oleic acid by ozone, *Atmos. Chem. Phys.*, 10, 3673–3691, doi:10.5194/acp-10-3673-2010, 2010.

Tabazadeh, A. and Turco, R. P.: A Model For Heterogeneous Chemical Processes On the Surfaces of Ice and Nitric-Acid Trihydrate Particles, *J. Geophys. Res.-Atmos.*, 98, 12727–12740, 1993.

Vieceli, J., Roeselova, M., Potter, N., Dang, L. X., Garrett, B. C., and Tobias, D. J.: Molecular Dynamics Simulations of Atmospheric Oxidants at the Air-Water Interface: Solvation and Accommodation of OH and O<sub>3</sub>, *J. Phys. Chem. B*, 109, 15876–15892, 2005.

Wexler, A. S. and Clegg, S. L.: Atmospheric aerosol models for systems including the ions H<sup>+</sup>, NH<sub>4</sub><sup>+</sup>, Na<sup>+</sup>, SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, Cl<sup>-</sup>, Br<sup>-</sup>, and H<sub>2</sub>O, *J. Geophys. Res.-Atmos.*, 107, art. no. 4207, 2002.

Williams, L. R. and Long, F. S.: Viscosity of Supercooled Sulfuric-Acid-Solutions, *J. Phys. Chem.*, 99, 3748–3751, 1995.

Winkler, P. M., Vrtala, A., Wagner, P. E., Kulmala, M., Lehtinen, K. E. J., and Vesala, T.: Mass and thermal accommodation during gas-liquid condensation of water, *Phys. Rev. Lett.*, 93, art. no. 075701, 2004.

Worsnop, D. R., Morris, J. W., Shi, Q., Davidovits, P., and Kolb, C. E.: A chemical kinetic model for reactive transformations of aerosol particles, *Geophys. Res. Lett.*, 29, art. no. 1996, 2002.

Zobrist, B., Marcolli, C., Pedernera, D. A., and Koop, T.: Do atmospheric aerosols form glasses?, Atmos. Chem. Phys., 8, 5221–5244, doi:10.5194/acp-8-5221-2008, 2008.

Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

ACPD

12, 32109–32472, 2012

---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

---

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

32143



## Datasheets

### Appendix A1

#### Uptake on liquid water surfaces

##### VI.A1.1

##### 5 O<sub>3</sub> + H<sub>2</sub>O(l) → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
$\alpha_b > 2 \times 10^{-3}$ ( $\pm 20\%$ )	276	Utter et al. (1992)	WWFT-CLD (a)
$\alpha_b > 2 \times 10^{-3}$ (2.89 M > $[I^-]$ > 0.36 M)	282	Magi et al. (1997)	DT-MS (b)
$> 2.0 \times 10^{-2}$	298	Schütze and Herrmann (2002)	(c)

#### Comments

10 (a) Wetted-wall flow reactor with flowing film of liquid water (0.2 mm thick). The ozone concentration was  $10^{11}$  molecule  $\text{cm}^{-3}$ . No uptake of O<sub>3</sub> was observed into pure deionized water, but  $\gamma$  increased as a function of the concentration of the trapping agent (e.g., Na<sub>2</sub>SO<sub>3</sub>) to values in the  $10^{-2}$  range. A conservative lower limit for the accommodation coefficient of  $2 \times 10^{-3}$  was obtained.

15 (b) Uptake on a train of aqueous drops (80 to 150  $\mu\text{m}$ ) in a laminar flow tube coupled to an ion trap MS with [O<sub>3</sub>] monitored at m/e 46 (NO<sub>2</sub><sup>+</sup>) after titration by NO to NO<sub>2</sub>. No measurable uptake of O<sub>3</sub> on pure water. The lower limit for  $\alpha_b$  was found from the intercept of a plot of  $1/\gamma$  vs.  $a(I^-)^{-1/2}$  when NaI was used as scavenger.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





(c) Uptake onto static single drop (2–3 mm in diameter) containing NaI as scavenger. The product  $I_3^-$  was monitored by time-resolved UV/Vis absorption spectroscopy. Absorbance-time profiles at two different wavelengths (288, 353 nm) were used to derive uptake coefficients. The tabulated value of  $\alpha_b$  results from the intercept of  $1/\gamma$  vs.  $a(t^-)^{-\frac{1}{2}}$  plots.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	$>10^{-3}$	298
$D_1$ (cm <sup>2</sup> s <sup>-1</sup> )	$1.85 \times 10^{-5}$	
$H$ (M atm <sup>-1</sup> )	$1.15 \times 10^{-2} \exp(2560/T - 8.6)$	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298
$\Delta \log(k^{II})$	$\pm 0.5$	

### Comments on preferred values

In the absence of fast loss processes, the uptake of O<sub>3</sub> on pure water is reversible (Staehelin and Hoigné, 1982). The low solubility of O<sub>3</sub> means that uptake rapidly saturates and net O<sub>3</sub> uptake is observable under conditions of the work of Utter et al. (1992). Therefore, the studies which derived an estimate for the bulk accommodation coefficient  $\alpha_b$  used a scavenger (SO<sub>3</sub><sup>2-</sup> or I<sup>-</sup>) to obtain  $\alpha_b$  from the intercept of  $1/\gamma$  vs.  $a(t^-)^{-\frac{1}{2}}$  plots. We prefer the lower limit obtained by Magi et al. (1997) from a relatively large data set. The data by Schütze and Herrmann (2002) are more strongly affected by diffusion, and the model to take that into account required several simplifications. In addition, Schütze and Herrmann found evidence for limitation of product formation via HOI +

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$2\text{I}^- \rightarrow \text{I}_3^- + \text{OH}^-$ , while Magi et al. directly determined the net loss of  $\text{O}_3$  from the gas phase. Note that for more concentrated iodide solutions, we prefer a value of  $\alpha_b > 0.1$ . Since halogenide ions may play a role in the bulk accommodation of ozone, we refrain from transferring the high bulk accommodation coefficient to dilute solutions.

5 The preferred solubility is the expression compiled by Chameides (1984). The room temperature value of the diffusion coefficient is from Matrozov et al. (1976).

## References

Chameides, W. L.: J. Geophys. Res., 89, 4739–4755, 1984.

10 Magi, L., Schweitzer, F., Pallares, C., Cherif, S., Mirabel, Ph., and George, C.: J. Phys. Chem. A, 101, 4943–4949, 1997.

Matrozov, V. I., Kashtanov, S. A., Stepanov, A. M., and Tregubov, B. A.: J. Appl. Chem.-USSR, 49, 1111–1114, 1976.

Staelin, J. and Hoigné, J.: Environ. Sci. Technol., 16, 676–681, 1982.

Schütze, M. and Herrmann, H.: Phys. Chem. Chem. Phys., 4, 60–67, 2002.

15 Utter, R. G., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4973–4979, 1992.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A1.2

### H<sub>2</sub>O<sub>2</sub> + H<sub>2</sub>O(l) → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
0.32 ± 0.02	260	Worsnop et al. (1989)	DT-TDL (a)
0.18 ± 0.02	273		
0.08 ± 0.02	292		

#### 5 Comments

(a) Droplets (pH = 7) with interaction times from 0.7 to 14 ms. Initial H<sub>2</sub>O<sub>2</sub> densities were in the range 1.5 to 15 × 10<sup>12</sup> molecule cm<sup>-3</sup>.

#### Preferred values

Parameter	Value	T/K
$\alpha_b$	1.3 × 10 <sup>-6</sup> exp (3230/T)	260–292
<i>Reliability</i>		
$\Delta(E/R)$	±500 K	260-292

#### 10 Comments on preferred values

We adopt the single dataset of Worsnop et al. (1989) as the basis of our recommendation. The strong negative temperature dependence of  $\alpha_b$  was suggested to derive

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

from an activation energy necessary to overcome a barrier between a precursor state and the solvated state of  $\text{H}_2\text{O}_2$ . The title reaction therefore represents a complex reaction involving a preequilibrium whose temperature dependence affects the overall T-dependence of the uptake.

## 5 References

Worsnop, D. R., Zahniser, M. S., Kolb, C. E., Gardner, J. A., Watson, L. R., Van Doren, J. M., Jayne, J. T., and Davidovits, P.: J. Phys. Chem., 93, 1159–1172, 1989.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.3



## Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$10^{-7}$ (pure water and 0.05 M NaOH)	273	Lee and Tang (1988)	(a)
$(1.5 \pm 0.6) \times 10^{-3}$ ( $9 \times 10^{14} \text{ cm}^{-3} \text{ NO}_2$ )	298	Ponche et al. (1993)	DT-IC (b)
$(8.7 \pm 0.6) \times 10^{-5}$	298	Msibi et al. (1993)	CWFT (c)
$(2.4 \pm 1.4) \times 10^{-4}$ ( $9 \times 10^{14} \text{ cm}^{-3} \text{ NO}_2$ )	298	Mertes and Wahner (1995)	LJ (d)
$(1.2 \pm 0.4) \times 10^{-3}$ ( $5 \times 10^{15} \text{ cm}^{-3} \text{ NO}_2$ )	298		
$(9.5 \pm 1.5) \times 10^{-7}$ ( $3\text{--}6 \times 10^{14} \text{ cm}^{-3} \text{ NO}_2$ )	298	Kleffmann et al. (1998)	Bubbler-IR/IC (e)
$< 5 \times 10^{-4}$	273	Cheung et al. (2000)	DT-MS (f)
$\leq 10^{-7}$ (pH 2–12)	298	Gutzwiller et al. (2002)	
		Ammann et al. (2005)	WWFT-CLD (g)
$\approx 10^{-6}$ ( $6 \times 10^{14} \text{ cm}^{-3} \text{ NO}_2$ , 0.01 M NaOH)	288	Komiyama and Inoue (1980)	Bubbler-UV (h)
<i>Accommodation coefficient: <math>\alpha_b</math></i>			
$(6.3 \pm 0.7) \times 10^{-4}$ ( $> 10^{-2} \text{ M}$ hydroquinone)	273	Lee and Tang (1988)	(a)
$(4.2 \pm 0.9) \times 10^{-4}$ ( $5 \times 10^{-2} \text{ M}$ ascorbate)	298	Msibi et al. (1993)	CWFT (c)
$(2.4 \pm 1.8) \times 10^{-2}$ (deliquesced dihydroxyphenolates)	296	Sosedova et al. (2009)	AFT-RC (i)
<i><math>HK_2^{1/2}</math></i>			
$70 \text{ M}^{-1/2} \text{ atm}^{-1} \text{ s}^{-1/2}$	295	Lee and Schwartz (1981)	(j)
$134 \text{ M}^{-1/2} \text{ atm}^{-1} \text{ s}^{-1/2}$	284	Cape et al. (1993)	(k)
$71 \pm 10 \text{ M}^{-1/2} \text{ atm}^{-1} \text{ s}^{-1/2}$	293	Cheung et al. (2000)	(f)
$92 \text{ M}^{-1/2} \text{ atm}^{-1} \text{ s}^{-1/2}$	295	Park and Lee (1988)	(l)
<i>Solubility</i>			
$2.0 \times 10^{-2} \text{ M atm}^{-1}$	298	Andrew and Hanson (1960)	(m)
$2.2 \times 10^{-2} \text{ M atm}^{-1}$	288	Komiyama and Inoue (1980)	(h)
$(7.0 \pm 0.5) \times 10^{-3} \text{ M atm}^{-1}$	295	Lee and Schwartz (1981)	(i)
$(5.5 \pm 0.6) \times 10^{-2} \text{ M atm}^{-1}$	284	Cape et al. (1993)	(k)
$(1.4 \pm 0.2) \times 10^{-2} \text{ M atm}^{-1}$	293	Cheung et al. (2000)	(f)
$(2.3_{-0.9}^{+0.3}) \times 10^{-2} \text{ M atm}^{-1}$	276		

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Parameter	Temp./K	Reference	Technique/Comments
<i>Liquid phase rate constant</i>			
$(1.0 \pm 0.1) \times 10^8 \text{ M}^{-1} \text{ s}^{-1}$	295	Lee and Schwartz (1981)	(j)
$(8.4 \pm 1.5) \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$	295	Park and Lee (1988)	(m)
$(6.0 \pm 2.0) \times 10^6 \text{ M}^{-1} \text{ s}^{-1}$	284	Cape et al. (1993)	(k)
$(3.0 \pm 0.9) \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$	293	Cheung et al. (2000)	(f)
$(2.2 \pm 0.6) \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$	276		

## Comments

- (a) Uptake experiment involving a stopped flow of  $\text{NO}_2$  (initial concentration  $10^{15} \text{ cm}^{-3}$ ) in humidified He carrier gas exposed to a circulating fluid sample of either pure water, water with 0.05 M NaOH, or water containing  $(1 - 7) \times 10^{-2} \text{ M}$  hydroquinone sodium salt. The uptake coefficients were corrected for gas phase diffusion.
- (b) Gas uptake into a monodisperse (70  $\mu\text{m}$  to 110  $\mu\text{m}$  diameter) droplet train. The pH of water was adjusted to 4, 5.6 and 11. Uptake was determined by detecting nitrate by ion chromatography in the collected aqueous phase after addition of 0.16 M  $\text{H}_2\text{O}_2$  to convert all nitrite to nitrate. The uptake coefficient corrected for gas phase diffusion was independent of pH. The value for the diffusion coefficient used was  $0.192 \text{ cm}^2 \text{ s}^{-1}$  in the gas phase and  $1 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$  in the liquid phase. The solubility was assumed to be  $1.2 \times 10^{-2} \text{ M atm}^{-1}$ . The observed uptake coefficient  $\gamma_{\text{obs}}$  was larger than expected from solubility limited uptake, independent of pH. Impurities might have been present as a liquid phase sink for  $\text{NO}_2$  to drive reaction limited uptake.
- (c) A flow of 6l per min of air at 97% relative humidity containing about 0.5 ppm  $\text{NO}_2$  at atmospheric pressure passed an annular glass reactor, coated with either deionised water or a mixture of 0.05 M ascorbic acid and 0.05 M NaOH (pH 9.3).  $\text{NO}_2$  was measured using a luminol based  $\text{NO}_2$  analyzer.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- (d) Liquid jet (50  $\mu\text{m}$  diameter) in a reactor at 298 K. Experiments were done with pure water, with 0.01 M NaOH, 33.4 g/l triethanolamine and 1 g/l NaAsO<sub>2</sub> to vary the liquid phase sink for NO<sub>2</sub>. Nitrite and nitrate were measured spectrophotometrically. NO<sub>2</sub> uptake was independent of pH or added sink, but was much larger than solubility limited uptake. The authors therefore proposed a surface reaction proceeding with the uptake coefficients listed in the table.
- (e) NO<sub>2</sub> was circulated through a thermostated bubbler containing 125 to 500 ml of water. NO<sub>2</sub> and HNO<sub>2</sub> were measured using tunable diode laser absorption spectrometry. Nitrate and nitrite in the liquid were measured using ion chromatography, showing that they were initially formed at equimolar amounts. The first order rate constant for NO<sub>2</sub> decay did not depend on the NO<sub>2</sub> concentration, indicating overall first-order behavior.
- (f) Experiments were performed with both a droplet train and a bubble train apparatus. The droplet train was operated with 10<sup>13</sup>–10<sup>16</sup> NO<sub>2</sub> molecule cm<sup>-3</sup>. Uptake of NO<sub>2</sub> was not detectable in the droplet train experiment, leading to an upper limit of  $\gamma < 5 \times 10^{-4}$  at 273 K. In the bubble train flow reactor, the NO<sub>2</sub>(g) concentration was monitored by QMS. Loss of NO<sub>2</sub> from the gas phase was fitted with a model of NO<sub>2</sub> uptake considering both solubility limitation and reaction in the bulk. This led to both, independent values for H and the second order liquid phase rate constant.
- (g) The wetted wall flow tube was operated at room temperature and with 80 ppb of NO<sub>2</sub> in N<sub>2</sub>. The uptake coefficient listed in the table was obtained as blank values for pure water and water with pH adjusted to between 2 and 12 as part of a study to explore the reactivity of NO<sub>2</sub> with hydroxysubstituted aromatics.
- (h) NO<sub>2</sub> (10 to 2000 ppm) was passed either over a flat liquid surface (both gas and liquid stirred) or through a bubbler. NO<sub>2</sub> concentration in the gas phase and nitrate and nitrite were measured using UV absorption. The mass transfer characteristics

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

were determined with  $\text{CO}_2$ . Nitrate and nitrite concentrations were equal over the whole  $\text{NO}_2$  pressure range. The  $\text{NO}_2$  absorption rate reported over the full range was proportional to the square of the  $\text{NO}_2$  pressure at high and proportional to  $\text{NO}_2$  at low  $\text{NO}_2$  pressures as expected for mass transport limitation. The absorption rate reported for about  $6 \times 10^{14} \text{ cm}^{-3}$  is converted to an uptake coefficient and listed in the table.

- (i) Aerosol particles were produced from nebulizing aqueous solutions of dihydroxybenzoic acid and hydroquinone sodium salts.  $^{13}\text{N}$  labeled  $\text{NO}_2$  was used as reactant, and uptake to particles was monitored by counting radioactive decays of  $^{13}\text{N}$  associated with particles after the flow reactor with residence times between 1 and 60 s at ambient pressure and 40 % relative humidity. The observed uptake coefficients in the range of a few  $10^{-3}$  could be well explained by bulk reaction limited uptake, which allowed derivation of the estimate for  $\alpha_b$  listed in the table.
- (j)  $\text{NO}_2$  at partial pressures of 100 ppb to 800 ppm was passed through 10 to 70 ml of water in a fritted bubbler in a thermostated vessel.  $\text{NO}_2$  was measured with a chemiluminescence detector. The ionic products in water were determined with a conductivity detector online. Nitrate and nitrite were measured offline with a colorimetric method. The mass transport characteristics were calibrated using  $\text{CO}_2$ , leading to mass transfer times between 1.7 and 5.3 s.
- (k)  $\text{NO}_2$  at partial pressures of 10 to 100 ppb was passed through 1.0 l of water or aqueous solution in a fritted bubbler in a thermostated vessel. The loss of  $\text{NO}_2$  was measured with a chemiluminescence detector. The mixing time in the reactor was 60 s based on experiments with  $\text{CO}_2$ . The authors caution that the distribution of  $\text{NO}_2$  in the reactor may have not been uniform, leading to an overestimate of the rate constants.



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

- (l) Investigation of  $\text{HNO}_2$  decomposition kinetics in water in a bubbler type apparatus. The kinetics of the hydrolysis reaction was indirectly derived from the evolution of NO and  $\text{NO}_2$  measured with a chemiluminescence detector.
- 5 (m)  $\text{NO}_2$  (200 ppm to 10 %) was passed through water in a bubbler type reactor. The solubility listed in the table was derived from observations of  $\text{HNO}_2$  formation as a function of  $\text{NO}_2$  pressure.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	$2 \times 10^{-2}$	273–298
$k/\text{M}^{-1} \text{s}^{-1}$	$8.1 \times 10^9 \exp(-1630/T)$	276–293
$H/\text{M atm}^{-1}$	$4.4 \times 10^{-6} \exp(-2350/T)$	273–298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.7$	273–298
$\Delta \log(k)$	$\pm 0.08$	273–298
$\Delta \log(H)$	$\pm 0.06$	273–298
$\Delta \log(E/R)$	$\pm 50$	273–298

### Comments on preferred values

10 Uptake of  $\text{NO}_2$  to pure water is driven by the low solubility and slow second order hydrolysis in bulk water. The study by Cheung et al. (2000) has covered the  $\text{NO}_2$  concentration range down to atmospherically relevant levels and most carefully elaborated the associated mass transfer issues, allowing them to obtain an independent estimate for the solubility at 293 K. They also estimate a value at 276 K using arguments about

the expected solubility dependence presented by Schwartz and White (1981). The recommended expression for the temperature dependence is based on these two values.

Cheung et al. also show that their value for  $Hk^{1/2}$  is consistent with most earlier studies covering the lower concentration range, especially the one by Lee and Schwartz (1981). The disagreement with Cape et al. (1993) is likely due to inhomogeneous distribution of  $\text{NO}_2$  in their reactor, as cautioned by the authors. Due to strong coupling between the liquid phase rate constant and the solubility, we calculate an average value of  $Hk^{1/2}$  extracted from the Cheung et al. (2000), Lee and Schwartz (1981) and Park and Lee (1988) and use the solubility as recommended to arrive at the expression for the rate constant.

Bambauer et al. (1994) report an experiment in a cloud chamber, in which the cloud droplet seeds were NaCl, which seemed to be inconsistent with the second-order reaction of  $\text{NO}_2$  in water. This and other evidence for apparent first order uptake with surprisingly high uptake coefficient can be discussed as due to reactions with impurities (Ponche et al., 1993; Msibi et al., 1993; Mertes and Wahner, 1995) rather than being representative of bulk accommodation limitation. Yabushita et al. (2009) suggest the formation of a surface complex with chloride ions to explain enhanced  $\text{NO}_2$  hydrolysis in mM chloride solutions (see also data sheet VI.A2.4). Due to the low solubility of  $\text{NO}_2$ , a strong aqueous phase scavenger is required to get into a bulk accommodation limited kinetic regime. While the CWFT study by Msibi et al. (1993) was likely affected by evaporation, Lee and Tang (1988) show evidence for the transition from reaction limited to bulk accommodation limited kinetics. Sosedova et al. (2009) used an aerosol flow tube approach, which is practically free of gas phase diffusion effects, to obtain a bulk accommodation coefficient of 0.02 for deliquesced sodium salts of hydroquinone and dihydroxybenzoate, which is used as a basis for the recommendation for  $\alpha_b$ .

The recommended values can be used to calculate uptake coefficients for the reaction of  $\text{NO}_2$  with dilute aqueous solutions in absence of significant aqueous phase scavengers. In this case,  $1/\alpha_b$  is not rate limiting, and the uptake coefficients can be

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

calculated using

$$\gamma = \frac{4HRT}{\bar{c}} \sqrt{D_1 \cdot k'}, \quad k' = H_{\text{NO}_2} p_{\text{NO}_2} k$$

This leads to uptake coefficients of e.g.  $6 \times 10^{-9}$  at 10 ppb  $\text{NO}_2$  due to hydrolysis. Under many conditions, the reacto-diffusive length may be large compared to the diameter of the droplet, so that the geometry correction needs to be applied to this expression. On shorter time scales, it needs to be carefully evaluated whether or not  $1/\Gamma_{\text{sol}}$  is rate limiting.

## References

- Ammann, M., Rossler, E., Strekowski, R., and George, C.: Phys. Chem. Chem. Phys., 7, 2513–2518, 2005.
- Andrew, S. P. S. and Hanson, D.: Chem. Eng. Sci., 14, 105–114, 1961.
- Bambauer, A., Brantner, B., Paige, M., and Novakov, T.: Atmos. Environ., 28, 3225–3232, 1994.
- Cape, J. N., Storeton-West, R. L., Devine, S. F., Beatty, R. N., and Murdoch, A.: Atmos. Environ., 27A, 2613–2621, 1993.
- Cheung, J. L., Li, Y. Q., Boniface, J., Shi, Q., Davidovits, P., Worsnop, D. R., Jayne, J. T., and Kolb, C. E.: J. Phys. Chem. A, 104, 2655–2662, 2000.
- Gutzwiller, L., George, C., Rössler, E., and Ammann, M.: J. Phys. Chem. A, 106, 12045–12050, 2002.
- Kleffmann, J., Becker, K. H., and Wiesen, P.: Atmos. Environ., 32, 2721–2729, 1998.
- Kleffmann, J., Becker, K. H., and Wiesen, P.: J. Chem. Soc. Faraday Trans., 94, 3289–3292, 1998.
- Komiyama, H. and Inoue, H.: Chem. Eng. Sci., 35, 154–161, 1980.
- Lee, Y.-N. and Schwartz, S. E.: J. Phys. Chem., 85, 840–848, 1981.
- Lee, J. H. and Tang, I. N.: Atmos. Environ., 22, 1147–1151, 1988.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Mertes, S. and Wahner, A.: J. Phys. Chem., 99, 14000–14006, 1995.  
Park, J.-Y. and Lee, Y.-N.: J. Phys. Chem., 92, 6294–6302, 1988.  
Ponche, J. L., George, C., and Mirabel, P.: J. Atmos. Chem., 16, 1–21, 1993.  
Kleffmann, J., Kurtenbach, R., Becker, K. H., and Wiesen, P.: Faraday Discuss., 100,  
5 121–127, 1995.  
Kleffmann, J., Becker, K. H., and Wiesen, P.: Atmos. Environ., 32, 2721–2729, 1998.  
Shen, C. H. and Rochelle, G. T.: Environ. Sci. Technol., 32, 1994–2003, 1998.  
Shen, C. H. and Rochelle, G. T.: J. Air Waste Man. Assoc., 49, 332–338, 1999.  
Sosedova, Y., Rouvière, A. L., Gäggeler, H. W., and Ammann, M.: J. Phys. Chem. A,  
10 113, 10979–10987, 2009.  
Yabushita, A., Enami, S., Sakamoto, Y., Kawasaki, M., Hoffmann, M. R., and Colussi,  
A. J.: J. Phys. Chem. A, 113, 4844–4848, 2009.

---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

---

M. Ammann et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## VI.A1.4

### NO<sub>3</sub> + H<sub>2</sub>O (aq) → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
$\alpha_b$			
$>4 \times 10^{-2}$	273 ± 1	Rudich et al. (1996)	WWFT-AS (a)
$>2 \times 10^{-3}$	293	Thomas et al. (1998)	(b)
$4.2_{-1.7}^{+2.2} \times 10^{-3}$	293 ± 1	Schütze et al. (2005)	(c)
$H$ (M atm <sup>-1</sup> )			
0.6 ± 0.3	273 ± 1	Rudich et al. (1996)	WWFT-AS (a)
1.8	293	Thomas et al. (1998)	(b)
0.2 ± 0.1	293 ± 1	Schütze et al. (2005)	(c)

#### 5 Comments

(a) Flow tube operated at 12–23 mbar. NO<sub>3</sub> ( $2\text{--}10 \times 10^{11}$  molecule cm<sup>-3</sup>) was formed by the thermal dissociation of N<sub>2</sub>O<sub>5</sub> and detected by diode laser absorption at 662 nm over a 12.6 m pathlength. A lower limit for  $\alpha_b$  was estimated from reactive uptake measurements (see datasheet VI.A2.05). The dependence of uptake coefficient,  $\gamma$ , on Cl<sup>-</sup> concentrations was combined with a literature value (Exner et al., 1992) for the rate coefficient for Cl<sup>-</sup> with NO<sub>3</sub> ( $2.76 \times 10^6$  M<sup>-1</sup> s<sup>-1</sup> at 273 K) to determine a value of  $HD_1^{0.5}$  of  $(1.9 \pm 0.4) \times 10^{-3}$  M atm<sup>-1</sup> cm s<sup>-0.5</sup>. Assuming a value of  $D_1$  of  $(1.0 \pm 0.5) \times 10^{-5}$  cm<sup>2</sup> s<sup>-1</sup> enabled a solubility of NO<sub>3</sub> in H<sub>2</sub>O to be derived.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

(b) Uptake of  $\text{NO}_3$  (generated by mixing  $\text{NO}$  with  $\text{O}_3$  at  $\approx 400^\circ\text{C}$ ) to 0.1 M chloride solution was monitored using three serial coiled glass denuders (2 mm id). Nitrate in solution was determined following reduction to  $\text{NO}_2^-$  and the Saltzman reaction. Large diffusion limitations result in only a lower limit to  $\gamma$ .

(c) Single droplet ( $\approx 7\text{ mm}^3$ ) suspended from a pipette in a flow tube (10 mbar He) with UV-Vis absorption spectroscopy for concentration measurement in both gas and aqueous phases ( $\text{NO}_3^-$  at 235 nm,  $\text{NO}_3$  using the 662 nm feature).  $\text{NO}_3$  was generated by reacting  $\text{NO}_2$  with  $\text{O}_3$  at 393 K to keep the  $\text{N}_2\text{O}_5$  level low ( $\approx 5\%$ ).  $\text{HNO}_3$  was formed at  $\approx$  the same concentration as  $\text{NO}_3$ . Uptake of  $\text{NO}_3$  to the droplet was monitored by nitrate anion absorption. A value of  $HD_1^{0.5}k^{0.5} = (1.9 \pm 0.2)\text{ M atm}^{-1}\text{ cm s}^{-0.5}$  was derived.

### Preferred values

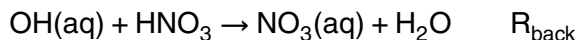
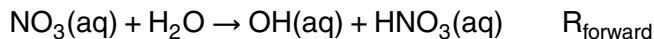
Parameter	Value	T/K
$\alpha_b$	$1.3 \times 10^{-2}$	273
$k_{\text{H}_2\text{O}} (\text{M}^{-1}\text{ s}^{-1})$	23	273
$H (\text{M atm}^{-1})$	0.6	273
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.5$	273
$\Delta \log(k_{\text{H}_2\text{O}})$	$\pm 0.5$	273
$\Delta \log(H)$	$\pm 0.5$	273

### Comments on preferred values

The accommodation coefficient,  $\alpha_b$ , was derived from uptake of  $\text{NO}_3$  to salt solutions as described in datasheet VI.A2.05.

The solubility,  $H$ , of  $\text{NO}_3$  has been determined on several occasions, the more recent results suggest that it is low, with numbers of  $(0.6 \pm 0.3)$  and  $(1.8 \pm 1.5) \text{ M}^{-1} \text{ atm}^{-1}$  derived by Rudich et al. (1996) and Thomas et al. (2005) compared to e.g.  $12 \text{ M}^{-1} \text{ atm}^{-1}$  reported by Chameides (1986). We prefer the results of Rudich et al., in order to maintain an internally consistent set of parameters for modelling  $\text{NO}_3$  uptake to both pure water and halide solutions (see VI.A2.05).

Rudich et al. (1996) observed reactive uptake of  $\text{NO}_3$  in the absence of halide ions (or other detectable impurities) and attributed this to the hydrolysis of  $\text{NO}_3$ :



with a rate constant of  $k_{\text{H}_2\text{O}} = 23_{-13}^{+30} \text{ M}^{-1} \text{ s}^{-1}$  for the forward reaction. Similarly, Schütze et al. (2005) observed reactive uptake and formation of nitrate ions in the interaction of  $\text{NO}_3$  with a pure  $\text{H}_2\text{O}$  droplet with an uptake coefficient of  $\sim 10^{-4}$  implying a hydrolysis rate coefficient of  $670_{-580}^{+2300} \text{ M}^{-1} \text{ s}^{-1}$  if nitrate (i.e. aqueous phase  $\text{HNO}_3$ ) was formed only as shown above. Even the lower limits of these rate constants are incompatible (by orders of magnitude) with the observations of Thomas et al. (1998), who derive an upper limit of  $0.5 \text{ M}^{-1} \text{ s}^{-1}$  for  $k_{\text{H}_2\text{O}}$ .

Whilst Rudich et al. made efforts to eliminate reaction with impurities in their water film and ruled out gas-phase reactions as being responsible for  $\text{NO}_3$  loss, Thomas et al. hypothesise that impurities, a surface (rather than bulk) reaction of  $\text{NO}_3$  or non-laminar flow in the liquid film of Rudich et al. could contribute to observation of an apparent, large hydrolysis rate constant. On the other hand, calculations of the equilibrium constant for reaction of  $\text{NO}_3$  with water, Rudich et al. combined with the measured rate constant for the back reaction (Katsamura et al., 1991) result in  $k_{\text{H}_2\text{O}} = 6 \text{ M}^{-1} \text{ s}^{-1}$  at 298 K, roughly consistent with the observations.

Whilst recognising that there is very large uncertainty associated with values of each of  $\alpha_b$ ,  $H$  and  $k_{\text{H}_2\text{O}}$ , we adopt the results of Rudich et al. (1996) so that the following

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



expression, when used to calculate an uptake coefficient for  $\text{NO}_3$  on pure water, returns a value of  $\gamma \sim 2 \times 10^{-4}$ , which is consistent with their experimental observations.

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1 k'_{\text{H}_2\text{O}})^{0.5}} \right\}^{-1}$$

with  $D_1 \sim 1 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$ .

5 The aqueous phase products of  $\text{NO}_3$  hydrolysis are suggested to be OH and  $\text{HNO}_3$ . Clearly, further work on the uptake and reaction of  $\text{NO}_3$  on water is required to reduce uncertainties associated with this process.

## References

Chameides, W. L.: J. Geophys. Res. 91, 14571–14572, 1986.

10 Exner, M., Herrmann, H., and Zellner, R.: Ber. Bunsen. Ges. Phys. Chem. Chem. Phys 96, 470–477, 1992.

Rudich, Y., Talukdar, R. K., Ravishankara, A. R., and Fox, R. W.: J. Geophys. Res., 101, 21023–21031, 1996.

Schütze, M. and Herrmann, H.: J. Atmos. Chem. 52, 1–18, 2005.

15 Thomas, K., Volz-Thomas, A., Mihelcic, D., Smit, H. G. J., and Kley, D.: J. Atmos. Chem., 29, 17–43, 1998.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.5

### $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$ (water droplets) $\rightarrow$ products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{\text{ss}}</math>, <math>\gamma_0</math></i>			
$\gamma = 0.040 \pm 0.005$	282	Van Doren et al. (1990)	DT-TDLAS (a)
$\gamma = 0.061 \pm 0.004$	271		
$>0.005$	293	Kirchner et al. (1990)	LJ-IC (b)
$\gamma = 0.030 \pm 0.002$	262	George et al. (1994)	DT-FTIR/HPLC (c)
$\gamma = 0.013 \pm 0.008$	277		
$\gamma = 0.018 \pm 0.003$	262–278	Schweitzer et al. (1998)	DT-FTIR/MS (d)
$\gamma = 0.011 \pm 0.012 - 0.006$	293	Schütze and Herrmann (2002)	DT-UV/Vis (e)

#### 5 Comments

(a) Fast train of monodisperse 200  $\mu\text{m}$   $\text{H}_2\text{O}$  droplets traversing a flow tube with TD-LAS detection. The temperature of the droplet was controlled by varying partial pressure of  $\text{H}_2\text{O}$  in the carrier gas and hence evaporation rate. The temperature reported is the mean temperature in the vicinity of the droplet. Total pressures = 5.02 Torr at 271 K and 9.26 Torr at 282 K.  $[\text{N}_2\text{O}_5] \sim 2 \times 10^{14}$  molecule  $\text{cm}^{-3}$ .  $\gamma$  determined from measured fractional uptake with and without droplets present, with (small,  $<10\%$ ) correction for gas phase diffusion and distortion of molecular velocity distribution.

(b) Interaction of  $\text{N}_2\text{O}_5$  (116 ppm in 1 bar air, measured by UVA) with a 90 mm dia. liquid jet of pure water with contact time 0.03–1.0 ms. Uptake coefficients were deduced by comparing  $\text{NO}_3^-$  concentrations in the absorbing liquid, which are measured as a function of contact time, with numerical solutions of the convective

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



diffusion equation.  $\gamma$  is regarded as a lower limit due to its sensitivity to the value of  $D_g(\text{N}_2\text{O}_5)$  and uncertainties in the velocity distribution of the liquid jet.

- (c) Fast train of water droplets (80–150  $\mu\text{m}$  diameter) inside a flow tube at total pressures of 27–80 mbar synthetic air with 4–16 ms contact time. The temperature of the droplet was controlled by varying partial pressure of  $\text{H}_2\text{O}$  in the carrier gas and hence evaporation rate.  $\text{N}_2\text{O}_5$  (20 to 2000 ppm) monitored by FTIR and the extent of uptake was followed by measuring the  $\text{NO}_3^-$  concentration in the collected droplets using HPLC.
- (d) Experimental configuration (c). FTIR absorption and ion-trap mass spectrometry used for detection of changes in gas phase  $\text{N}_2\text{O}_5$  on exposure to droplets. No products were detected. The uptake coefficients were independent of temperature in the stated range.
- (e) Uptake onto static single drop monitored by time-resolved UV/Vis absorption spectroscopy in the range 240 to 800 nm. Uptake coefficient determined from absorbance-time profiles of the product  $\text{NO}_3^-$  at two different wavelengths (302, 345 nm) after diffusion correction.  $[\text{N}_2\text{O}_5] = (2.3 \text{ to } 4.6) \times 10^{13} \text{ molecule cm}^{-3}$  at a total pressure (He) of 100 mbar.

### Preferred values

Parameter	Value	T/K
$\gamma$	$2.7 \times 10^{-5} \exp(1800/T)$	265–300
<i>Reliability</i>		
$\Delta \log(\gamma)$	$\pm 0.3$	298
$\Delta (E/R)$	$\pm 1000 \text{ K}$	265–300

## Comments on preferred values

Except for the liquid jet experiment of Kirchner et al. (1990), the reported studies all used similar experimental configurations involving uptake onto large water droplets, but with different methods used for measurement of uptake rate, i.e. loss of  $\text{N}_2\text{O}_5$  from gas phase (Van Doren et al., 1990; Schweitzer et al., 1998) and accumulation of hydrolysis product  $\text{NO}_3^-$  in the droplets (George et al., 1992; Schütze and Herrmann, 2002). Corrections for effects of gas phase diffusion on uptake rates were necessary in all cases, and are a particular source of uncertainty in the liquid drop experiment. The values of  $\gamma$  from the different studies are broadly consistent with each other, and show a negative temperature dependence, although the  $\gamma$  values of Van Doren et al. (1990) are higher than in the other studies. There are potential sources of systematic error in these studies but there is no obvious explanation of the differences in the results. The recommended expression for  $\gamma$  is a least squares fit to all the cited data from the droplet train experiments, plotted in Arrhenius form, with large error limits reflecting the possible systematic errors.

It has been suggested the measured uptake coefficient represents the mass accommodation coefficient for  $\text{N}_2\text{O}_5$  on an aqueous surface, which by definition gives the maximum rate of reactive uptake. Maximum uptake coefficients for  $\text{N}_2\text{O}_5$  on aqueous electrolyte aerosols (e.g. NaCl, malonic acid,  $\text{H}_2\text{SO}_4$ ) at room temperature and high RH often exceed those measured on pure water by a factor of  $\sim 2$  to 3. The mass accommodation coefficient for  $\text{N}_2\text{O}_5$  at 298 K is thus likely to be at least 0.03.

## References

- George, C., Ponche, J. L., Mirabel, Ph., Behnke, W., Scheer, V., and Zetzsch, C.: J. Phys. Chem., 98, 8780–8784, 1994.
- Kirchner, W., Welter, F., Bongartz, A., Kames, J., Schweighofer, S., and Schurath, U.: J. Atm. Chem., 19, 427–449, 1990.
- Schütze, M. and Herrmann, H.: Phys. Chem. Chem. Phys., 4, 60–67, 2002.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Schweitzer, F., Mirabel, Ph., and George, C.: J. Phys. Chem. A, 102, 3942–3952, 1998.  
Van Doren, J. M., Watson, L. R., Davidovits, P., Worsnop, D. R., Zahniser, M. S. and  
Kolb, C. E.: J. Phys. Chem., 94, 3265–3269, 1990.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A1.6

### NH<sub>3</sub> + H<sub>2</sub>O (water droplets) → Products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{SS}</math>, <math>\gamma_0</math></i>			
0.040 <sup>+0.03</sup> <sub>-0.005</sub>	299	Bongartz et al. (1995)	LJ-IC (a)
0.097 ± 0.009	290	Ponche et al. (1993)	DT-FTIR/HPLC (b)
0.1 (pH = 0.5–3.0)	291	Shi et al. (1999)	DT-FTIR/MS (c)
0.06 (pH = 8)	291		
0.006 (pH = 10–13)	291		
0.35 (pH = 0.5–3.0)	260		

#### Comments

(a) Uptake of NH<sub>3</sub> in a liquid jet of dilute HCl solution with continuous measurement of NH<sub>4</sub><sup>+</sup> in the collected jet waters. Contact time of 0.1 ms to 0.8 ms virtually free of saturation effects in the liquid phase. The  $\gamma$  values for air (+He) as carrier gas agreed very well with those measured in He ( $\gamma = 3.5 \times 10^{-2}$  in air vs.  $4.0 \times 10^{-2}$  in He).

(b) Gas interaction with monodisperse (70  $\mu$ m to 110  $\mu$ m diameter) droplet train propagating axially in a low pressure (46 mbar to 93 mbar) flow tube. Uptake determined by ion chromatographic analysis of collected aqueous phase was independent of pH in the range 4.2 to 11.5 but dependent on the contact time and pressure. The value listed in the table has been corrected for gas phase diffusion. The measured average uptake coefficient  $\gamma_{\text{obs}}$  was  $(2.4 \pm 1.0) \times 10^{-2}$  at 93 mbar N<sub>2</sub> and 290 K, obtained by extrapolation to zero interaction time.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

(c) Uptake experiment of gas phase  $\text{NH}_3$  on train of droplets whose size was in the range 150–300  $\mu\text{m}$  entrained in a flowing mixture of helium and water vapour (between 2.9 and 23.3 mbar). The reaction time was between 2 and 15 ms, the ammonia concentration was in the range  $10^{13}$  to  $2 \times 10^{14} \text{ cm}^{-3}$  and was monitored using a VUV lamp emitting at  $\lambda = 121.6 \text{ nm}$ . The theoretical framework included gas phase diffusion of ammonia, mass accommodation, dissolution, bulk and surface reaction whose relative contributions changed with pH. The effective Henry's law solubility was measured in the range 264 to 350 K:  $\log(H) = -3.221 + 1396/T$ . The uptake coefficient was dependent on pH, falling rapidly from  $\sim 0.08$  at  $\text{pH} = 7$  to 0.006 at  $\text{pH} = 10$ . At low pH there is a negative temperature dependence of  $\gamma$ , which is attributed to accommodation controlled uptake. At high pH,  $\text{NH}_3$  is less soluble and uptake controlled by a surface complex is suggested. Co-deposition with  $\text{SO}_2$  enhanced the ammonia uptake to values corresponding to acidic solutions.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.1	260–300
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	260–300

### Comments on preferred values

The uptake coefficients on pure water, measured using the droplet train technique are somewhat larger than onto dilute HCl in a liquid jet. Both methods require correction for gas phase diffusion and for saturation effects. On balance, the DT experiment gives more reliable data and we therefore base the recommendation on the results of Ponche et al. (1993) and Shi et al. (1999), who also determined the temperature dependence

at pH = 1. The value is higher than usually observed for the mass accommodation coefficient,  $\alpha_b$ , for uptake on pure water, and the strong dependence of  $\gamma$  on pH (Shi et al., 1999), indicates a direct chemical interaction of  $\text{NH}_3$  at the surface, e.g. by protonation.

## References

- 5 Bongartz, A., Schweighofer, S., Roose, C., and Schurath, U.: J. Atmos. Chem., 20, 35–58, 1995.
- Ponche, J. L., George, C., and Mirabel, Ph.: J. Atmos. Chem., 16, 1–21, 1993.
- Shi, Q., Davidovits, P., Jayne, J. T., Worsnop, D. R., and Kolb, C. E.: J. Phys. Chem. A, 103, 8812–8823, 1999.

---

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

---

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Bulk accommodation coefficients: <math>\alpha_b</math></i>			
$>5 \times 10^{-3}$	299	Kirchner et al. (1990)	LJ-IC (a)
$4.3 \times 10^{-3}$ (wet $\text{Na}_2\text{CO}_3$ )	298	Msibi et al. (1993)	CWFT (b)
$(5_{-1}^{+2}) \times 10^{-2}$	297	Bongartz et al. (1994)	LJ-IC (c)
$(5_{-1}^{+4}) \times 10^{-2}$	245		DT-IC (c)
$4 \times 10^{-3} < \alpha_b < 4 \times 10^{-2}$	278	Mertes and Wahner (1995)	LJ (d)
<i>Solubility: <math>H</math></i>			
$223 \pm 11$	273	Park and Lee (1988)	Bubbler-CLD (e)
$121 \pm 6$			
$60 \pm 3$			
$38 \pm 2$			
$H^* = 194 \pm 25$ (8.7 wt % $\text{H}_2\text{SO}_4$ )	269	Becker et al. (1996)	Static-TDLAS/IC (f)
$H^* = 117 \pm 20$ (8.7 wt % $\text{H}_2\text{SO}_4$ )	278		
$H^* = 73 \pm 10$ (8.7 wt % $\text{H}_2\text{SO}_4$ )	285		
$H^* = 50 \pm 8$ (8.7 wt % $\text{H}_2\text{SO}_4$ )	291		
$H^* = 34 \pm 5$ (8.7 wt % $\text{H}_2\text{SO}_4$ )	298		
<i>Rate constants, <math>k_1</math> [<math>\text{M}^{-1} \text{s}^{-1}</math>]</i>			
3.46	283	Park and Lee (1988)	Bubbler-CLD (e)
13.4	295		
28.6	303		
<i>Rate constants, <math>k_{-1}</math> [<math>\text{M}^{-1} \text{s}^{-1}</math>]</i>			
$1.98 \times 10^8$	283	Park and Lee (1988)	Bubbler-CLD (e)
$1.58 \times 10^8$	295		
$1.67 \times 10^8$	303		

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments

- (a) Interaction of  $\text{HNO}_2$  (60 ppm) with a liquid jet of pure water. The kinetics were studied by ion chromatographic analysis of the liquid jet nitrite content. The gas phase diffusion coefficient value used was  $0.13 \text{ cm}^2 \text{ s}^{-1}$ , the value used for the solubility of  $\text{HNO}_2$  in water was  $60 \text{ M atm}^{-1}$ . A kinetic model taking into account diffusion in the gas phase as well as the radially resolved flow velocity within the jet was used to simulate the rate of appearance of nitrite in the aqueous phase. Agreement with the data was not very good likely due to uncertainties in the velocity distribution of the liquid jet. Taking into account dissociation of  $\text{HNO}_2$  in the aqueous phase did not improve the fits. Thus, the bulk accommodation coefficient is regarded as a lower limit.
- (b) A flow of 6 l per min of air at 97 % relative humidity containing about 0.5 ppm  $\text{NO}_2$  at atmospheric pressure passed an annular glass reactor, coated with a mixture of  $\text{Na}_2\text{CO}_3$  and glycerol. Uptake of nitrous acid detected by wet extraction of reactor segments.
- (c) Two experimental flow techniques, an improved version of the liquid jet and the droplet train technique, were used. Both techniques involve the measurement of the concentration of  $\text{NO}_2^-$  in the collected liquid droplets of  $10 \mu\text{m}$  to  $100 \mu\text{m}$  diameter at residence times of 3.7 ms to 16 ms at total pressures of 40 mbar to 80 mbar. Temperature refers to droplet surface temperature of  $(245 \pm 5) \text{ K}$ . A 20 % uncertainty in  $D_g$  leads to the stated error intervals.
- (d) Liquid jet at 298 K. Experiments were done with pure water, with 0.01 M NaOH, 33.4 g/l triethanolamine and 1 g/l  $\text{NaAsO}_2$  to vary the liquid phase sink.  $\text{HNO}_2$  was measured by DOAS in a White cell.  $\text{HNO}_2$  was produced indirectly by reaction of  $\text{NO}_2$  with the glass walls upstream of the jet leading to concentrations of  $10^{13}$  to  $10^{14} \text{ cm}^{-3}$  in presence of a factor of 50 more  $\text{NO}_2$ . Nitrite and nitrate were measured spectrophotometrically.  $\text{HNO}_2$  uptake was independent of pH or added

32169

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



sink. Within the short residence time of the jet, the liquid phase remained far from equilibrium. A kinetic model was used to obtain the likely range for  $\alpha_b$  based on the observed loss of  $\text{HNO}_2$  from the gas phase (lower limit) and the appearance of nitrite in the aqueous phase (upper limit), respectively.

(e) A thermostated pyrex reactor with a fritted gas inlet. The solubility of  $\text{HNO}_2$  was measured by using low concentration  $\text{HNO}_2$  solutions ( $10^{-6}$  M) and observing the purge behavior with a chemiluminescence detector.  $\text{HNO}_2$  decomposition experiments were performed at higher solution concentrations  $>10^{-5}$  M to allow observation of  $\text{NO}$  and  $\text{NO}_2$  purging from the bubbler apparatus. The rate constants listed in the table for the forward (decomposition) ( $k_1$ ) and backward reaction ( $k_{-1}$ ) were obtained by fitting a kinetic model to the purge data. The mass transfer properties of the bubbler were calibrated using  $\text{CO}_2$ .

(f) Solubility of  $\text{HNO}_2$  was measured in a 11-L Pyrex glass reactor directly by monitoring both gas phase composition by tunable diode laser spectrometry and liquid phase by ion chromatography. Formation of  $\text{NO}_2$  due to  $2\text{HNO}_2 \rightleftharpoons \text{NO} + \text{NO}_2 + \text{H}_2\text{O}$  was observed. Since the  $\text{NO}_2$  concentration equilibrated with time, this reaction was included into calculating the effective solubility.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.05	273–300
$H^*$	$4.2 \times 10^{-6} \exp(4873/T)$ $(1 + 5.9 \times 10^{-4} \exp(-1760(1/T - 1/298)))/[H^+]$	273–300
$k_1$ ( $M^{-1} s^{-1}$ )	$3.1 \times 10^{14} \exp(-9090/T)$	273–300
$k_{-1}$ ( $M^{-1} s^{-1}$ )	$1.7 \times 10^8$	273–300
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.7$	273–300
$\Delta \log(k)$	$\pm 0.3$	273–300
$\Delta E/R$ [K]	$\pm 50$	273–300

## Comments on preferred values

Most of the attempts to measure the bulk accommodation coefficient of  $HNO_2$  into water or dilute aqueous solutions were done with a liquid jet experiment at relatively high pressure. Uncertainties arise from not understanding precisely the radial flow velocity distribution within the jet, which critically determines whether or not uptake runs into solubility equilibrium. The other significant issue is that especially at atmospheric pressure, transfer of the gas into the liquid is strongly affected by gas phase diffusion. We therefore adopt the value from the Bongartz et al. (1994) study performed at lower pressure than the others.

The solubility of  $HNO_2$  in water has been directly measured by Park and Lee (1988). Their data are also in line with the extrapolated solubility from those by Becker et al. (1996), who measured the solubility as a function of sulphuric acid composition (0.3 wt % and above). The preferred expression is consistent with that presented in data sheet VI.A4.7 for 0 wt %  $H_2SO_4$  and agrees well with the temperature dependence measured by Park and Lee. The acid dissociation constant used in this expression is

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

that determined by Park and Lee, which is identical to that calculated from the standard free energies of formation of  $\text{HNO}_2(\text{g})$  and  $\text{HNO}_2(\text{aq})$  (Schwartz and White, 1981). We note that the acid dissociation constant determined by Riordan et al. (2005) was somewhat lower ( $1.6 \times 10^{-3}$  ( $\text{pK} = 2.8$ ) at 298 K).

The rate constants for the decomposition reaction of  $\text{HNO}_2$  ( $k_1$ ) and the reverse process ( $k_{-1}$ ) has been adopted from Park and Lee (1988); for  $k_{-1}$  the average of the available rate constants at three temperatures have been taken as preferred value. They are somewhat above those of earlier studies (Schwartz and White, 1981, and references therein), but better represent the low concentrations expected under atmospheric conditions. The preferred value for  $k_1$  is much lower than that in  $\text{H}_2\text{SO}_4$  at about 50 wt % ( $320 \text{ M}^{-1} \text{ s}^{-1}$ , see data sheet VI.A4.7). Given that the back reaction is significant, an expression for  $\gamma$  based on the resistance model is not appropriate.

## References

- Becker, K. H., Kleffmann, J., Kurtenbach, R., and Wiesen, P.: J. Phys. Chem., 100, 14984–14990, 1996.
- Bongartz, A., Kames, J., Schurath, U., George, C., Mirabel, P., and Ponche, J. L.: J. Atmos. Chem., 18, 149–169, 1994.
- Kirchner, W., Welter, F., Bongartz, A., Kames, J., Schweighofer, S., and Schurath, U.: J. Atmos. Chem., 10, 427–449, 1990.
- Mertes, S. and Wahner, A.: J. Phys. Chem., 99, 14000–14006, 1995.
- Msibi, I. M., Shi, J. P., and Harrison, R. M.: J. Atmos. Chem., 17, 339–351, 1993.
- Park, J.-Y. and Lee, Y.-N.: J. Phys. Chem., 92, 6294–6302, 1988.
- Ponche, J. L., George, C., and Mirabel, P.: J. Atmos. Chem., 16, 1–21, 1993.
- Riordan, E., Minogue, N., Healy, D., O'Driscoll, P., and Sodeau, J. R.: J. Phys. Chem. A, 109, 779–786, 2005.
- Schwartz, S. E. and White, W. H.: in: Advances in Environmental Science and Engineering, edited by: Pfafflin, J. R. and Ziegler, E. N., Gordon and Breach Science Publishers, NY, 1–45, 1981.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

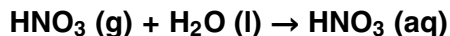
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.8



### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$0.19 \pm 0.02$	268	Van Doren et al. (1990)	DT-IR (a)
$0.071 \pm 0.02$	293		
$>0.01$	298	Kirchner et al. (1990)	LJ-IC (b)
$0.11 \pm 0.01$	298	Ponche et al. (1993)	DT-IC (c)
0.03	298	Schütze and Herrmann (2002)	(d)

### 5 Comments

(a) Experiments performed at room temperature flow with the water vapor pressure adjusted to the vapour pressure of water at the experimental temperature.  $\text{HNO}_3$  was detected downstream of the flow tube using IR absorption. The uptake coefficients were corrected for gas phase diffusion. The data show significant negative temperature dependence within the range indicated in the table. Given that under the experimental conditions the near surface region of the droplet does not saturate, it is argued that the uptake coefficients derived are actually bulk accommodation coefficients equivalent to the rate limiting step being solvation.

10  
15  
(b) Interaction of  $\text{HNO}_3$  (60 ppm) in 1 bar of synthetic air with a liquid jet of pure water. The kinetics were studied by ion chromatographic analysis of the liquid jet nitrate content. The uptake coefficient was attributed to a bulk accommodation coefficient and is regarded as a lower limit due to uncertainties in the velocity distribution of the liquid jet. The gas phase diffusion coefficient value used was

$0.132 \text{ cm}^2 \text{ s}^{-1}$ , the value used for the effective solubility of  $\text{HNO}_3$  in water was  $2.45 \times 10^6 \text{ M}^2 \text{ atm}^{-1}$ .

(c) Gas uptake into a monodisperse (70  $\mu\text{m}$  to 110  $\mu\text{m}$  diameter) droplet train. The pH of water was adjusted to 9.5 and 11.0. Uptake determined by chemical analysis (ion chromatography) of collected aqueous phase. The uptake coefficient corrected for gas phase diffusion was independent of pH, of the contact time and of the  $\text{HNO}_3$  concentration in the range 10 to 100 ppm. The value for the diffusion coefficient used was  $0.160 \text{ cm}^2 \text{ s}^{-1}$  in the gas phase and  $2.6 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$  in the liquid phase. The solubility was assumed to be  $2.1 \times 10^5 \text{ M atm}^{-1}$ . The measured average uptake coefficient  $\gamma_{\text{obs}}$  was  $(5.0 \pm 0.9) \times 10^{-2}$  at 67 mbar  $\text{N}_2$  and 298 K.

(d) A drop of water (2–3 mm diameter) was exposed to  $\text{HNO}_3$  ( $3\text{--}5 \times 10^{13} \text{ cm}^{-3}$ ) in a flow tube. The drop was probed by UV spectroscopy to monitor the appearance of nitrate in the aqueous phase. A model was used to determine the effect of gas phase diffusion on the measured uptake coefficient. The value for the gas phase diffusion coefficient was  $0.124 \text{ cm}^2 \text{ s}^{-1}$  in  $\text{H}_2\text{O}$  at 298 K and  $0.534 \text{ atm cm}^2 \text{ s}^{-1}$  in He at 298. The value used for the effective solubility of  $\text{HNO}_3$  in water was  $2.1 \times 10^5 \text{ M atm}^{-1}$ . The measurements indicate that the uptake coefficient given in the table is a lower limit.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	$7.5 \times 10^{-5} \exp(2100/T)$	268–300
Reliability		
$\Delta \log(\alpha_b)$	$\pm 0.7$	
$\Delta E/R$ [K]	$\pm 1000$	

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred values

The experiments on  $\text{HNO}_3$  uptake to pure water all resulted in large uptake coefficients, and all studies concluded that bulk accommodation rather than solubility or reaction in the liquid phase (dissociation) was rate limiting. They all carry a significant uncertainty due to gas phase diffusion, depending on geometry and pressure. The high pressure liquid jet experiment by Kirchner et al. (1990) was most sensitively affected by the magnitude of the diffusion coefficient and in addition contained an uncertainty with respect to the liquid flow regime. Schütze and Herrmann (2002) also used several simplifications and suggest that their uptake coefficient represents a lower limit to the bulk accommodation coefficient. Also, the diffusion correction commonly applied for high uptake coefficients to the droplet train experiments as reported by Van Doren et al. (1990) has recently been challenged (Morita et al., 2003; Hanson et al., 2004; Garrett et al., 2006; Davidovits et al., 2006). Van Doren's experiments at high temperature were measured at higher total pressure than those at lower temperature. The two droplet train studies, in spite of determining uptake via loss of gas phase  $\text{HNO}_3$  in one case (Van Doren et al., 1990) and the appearance of nitrate in the aqueous phase in the other (Ponche et al., 1993), but both using the same method to correct for diffusion, agree fairly well with each other. The recommended expression for  $\alpha_b$  is a least square fit to the data by Ponche et al. and Van Doren et al. in Arrhenius form, with large error limits to account for potential systematic errors.

## References

- Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Chem. Rev., 106, 1323–1354, 2006.
- Garrett, B. C., Schenter, G. K., and Morita, A.: Chem. Rev., 106, 1355–1374, 2006.
- Hanson, D. R., Sugiyama, M., and Morita, A.: J. Phys. Chem. A, 108, 3739–3744, 2004.
- Kirchner, W., Welter, F., Bongartz, A., Kames, J., Schweighofer, S., and Schurath, U.: J. Atmos. Chem., 10, 427–449, 1990.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Morita, A., Sugiyama, M., and Koda, S.: J. Phys. Chem. A, 107, 1749–1759, 2003.  
Ponche, J. L., George, C., and Mirabel, P.: J. Atmos. Chem., 16, 1–21, 1993.  
Schütze, M. and Herrmann, H.: Phys. Chem. Chem. Phys., 4, 60–67, 2002.  
Van Doren, J. M., Watson, L. R., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and  
5 Kolb, C. E.: J. Phys. Chem., 94, 3265–3269, 1990.

---

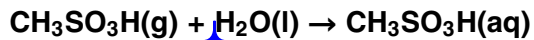
**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**M. Ammann et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)



## VI.A1.9



### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficient: <math>\alpha_b</math></i>			
$0.15 \pm 0.01$	264	De Bruyn et al. (1994)	DT (a)
$0.11 \pm 0.02$	279		
$0.16 \pm 0.04$	261	Schweitzer et al. (1998)	DT-MS (b)
$0.11 \pm 0.01$	283		
$0.89 \pm 0.04$ (7–15 wt % $\text{H}_2\text{SO}_4$ )	296	Hanson (2005)	AFT-CIMS (c)

### Comments

(a) Monodispersed droplets of 50 to 200  $\mu\text{m}$  in diameter Uptake coefficients were corrected for gas phase diffusion using  $D_g = 0.066 \text{ atm cm}^2 \text{ s}^{-1}$  in  $\text{H}_2\text{O}$  and  $0.299 \text{ atm cm}^2 \text{ s}^{-1}$  in He. The measured uptake coefficients were independent of pH adjusted with NaOH or HCl and independent of NaCl content up to 3.5 M. Since the observed uptake coefficient were time-independent, uptake was considered limited by bulk accommodation.

(b) Monodispersed droplets of 80 to 150  $\mu\text{m}$  in diameter  $\text{CH}_3\text{SO}_3\text{H}$  was admitted to the reactor at  $10^{13} \text{ cm}^{-3}$ ; its concentration at the reactor exit was measured by an ion-trap mass spectrometer. Uptake coefficients were corrected for gas phase diffusion using  $D_g = 0.13 \text{ cm}^2 \text{ s}^{-1}$  in  $\text{H}_2\text{O}$  and  $0.37 \text{ cm}^2 \text{ s}^{-1}$  in He and assuming the  $T$  dependence following  $T^{1.75}$ . The corrected uptake coefficients were inde-

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



pendent of time and NaCl content up to 2 M. It was considered limited by bulk accommodation in all cases.

(c) Uptake to sulphuric acid aerosol was studied in a laminar flow reactor coupled to CIMS detection using  $\text{HNO}_3$  as source of primary ions. Sulphuric acid particles were generated by homogeneous nucleation from supersaturated vapour leading to a lognormal particle size distribution within 50–120 nm, with a few  $10^4$  particles per  $\text{cm}^3$ , characterised by a differential mobility analyzer. Concentrations of  $\text{CH}_3\text{SO}_3\text{H}$  were  $3 \times 10^{10} \text{ cm}^{-3}$  in the flow tube. The measured uptake coefficients were corrected for gas phase diffusion using the Fuchs-Sutugin correction factor. The diffusion coefficient was directly measured based on the observed wall loss rates in absence of aerosol particles. It did not significantly depend on humidity. Its average value was  $0.0786 \text{ atm cm}^2 \text{ s}^{-1}$ .

### Preferred values

Parameter	Value	<i>T/K</i>
$\alpha_b$	1	260–300
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	260–290

### Comments on preferred values

The two droplet train studies of  $\text{CH}_3\text{SO}_3\text{H}$  uptake to water and dilute aqueous solutions agree very well and report a bulk accommodation coefficient around 0.1. However, the aerosol flow tube study by Hanson (2005) found a value for  $\alpha_b$  not different from 1. Since this experiment was much less affected by gas phase diffusion, we adopt this result for our recommendation for  $\alpha_b$ . Solubility of  $\text{CH}_3\text{SO}_3\text{H}$  is high enough to allow bulk accommodation limited uptake into the droplets over the experimental gas-particle interaction times of all studies. While  $\alpha_b$  was apparently independent of pH

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



and NaCl concentration, Schweitzer et al. (1998) and De Bruyn et al. (1994) found a temperature dependence, which they interpreted in terms of nucleation of a critical cluster as rate limiting step for the solvation rate leading to an average cluster of  $\text{CH}_3\text{SO}_3\text{H} \times 0.5 \text{H}_2\text{O}$ . This would, however, indicate a quite low ability of this very soluble species ( $H = 8.7 \times 10^{11} \text{ M/atm}$ , Brimblecombe and Clegg, 1988) and strong acid ( $\rho K_a = -2$ , Serjeant and Dempsey, 1979) to hydrogen bond to water.

## References

Brimblecombe, P. and Clegg, S. L.: J. Atmos. Chem., 7, 1–18, 1988.

De Bruyn, W. J., Shorter, J. A., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and

Kolb, C. E.: J. Geophys. Res., 99, 16927–16932, 1994.

Hanson, D. R.: J. Phys. Chem. A, 109, 6919–6927, 2005.

Schweitzer, F., Magi, L., Mirabel, P., and George, C.: J. Phys. Chem. A, 102, 593–600, 1998.

Serjeant, E. P. and Dempsey, B.: Ionization Constants of Organic Acids in Solution, IUPAC Chemical Data Series No. 23, Pergamon Press, Oxford, UK, 1979.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

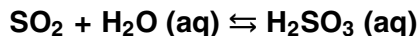
Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
$\gamma, \gamma_0$ 0.054 ± 0.006 (init. droplet pH = 11.5) $\gamma_0 = 0.12$	295	Gardner et al. (1987, 1989)	DT-TDL (a)
$\gamma_0 = 0.11 \pm 0.2$ (init. droplet pH = 11.5)	260–292	Worsnop et al. (1989)	DT-TDL (b)
$8.0 \times 10^{-3}$ (droplet pH = 0 to 2) 0.11 (droplet pH > 6)	283	Jayne et al. (1990)	DT-TDL (c)
0.06 (+0.14, -0.03)	300	Weiler et al. (1990)	LJ-IC (d)
$\gamma_0 = 0.13 \pm 0.01$	298	Ponche et al. (1993)	DT-IC (e)
0.028 ± 0.010 (pH = 13.2)	293.5	Shimono and Koda (1996)	(f)

### Comments

(a) Uptake of  $\text{SO}_2$  into a fast moving train of water droplets propagated axially in a flow tube. The droplet diameter varied from 80 to 180  $\mu\text{m}$  and the residence time of the individual droplet ranged from 2.2 to 12.8 ms. The total pressure was 28 mbar ( $\text{He} + \sim 22 \text{ mbar } \text{H}_2\text{O}$ ).  $[\text{SO}_2]$  measured by diode-laser adsorption in a long path (728 cm) White cell; range:  $(1.5\text{--}15.0) \times 10^{12}$  molecule  $\text{cm}^{-3}$ . Surface saturation effects were shown to be absent at  $[\text{SO}_2]$  and initial pH used. Measurements of aqueous  $[\text{S(IV)}]$  in exposed droplets confirmed uptake coefficient values determined from  $\text{SO}_2$  loss. Correction for gas phase diffusion to a static drop gives the cited value of  $\gamma_0$ , which can be considered equal to the bulk accommodation coefficient.

(b) Experiment designed to determine the temperature-dependent bulk accommodation coefficients of  $\text{SO}_2$  and  $\text{H}_2\text{O}_2$  on aqueous surfaces, using oxidation of  $\text{HSO}_3^{2-}$

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



to overcome solubility limitation of S(IV) uptake. A fast moving train of 200  $\mu\text{m}$  diameter water droplets propagated in a flow tube either in axial or transverse direction with  $[\text{SO}_2]_0 = 1 \times 10^{13}$  molecule  $\text{cm}^{-3}$ . The interaction time ranged from 2 to 12 ms (axial) or from 0.5 to 2 ms (transverse), with a temperature range 260–292 K at pressures from 13.3 to 66.5 mbar (He + H<sub>2</sub>O) for the axial configuration, and 3.1 to 13.3 mbar for the transverse geometry. Uptake coefficients were corrected for gas diffusion. At low pH,  $\gamma$  decreased with droplet exposure time due to reduced solubility of S(IV), and increased with pH. Temperature dependence of the initial uptake coefficient  $\gamma_0$  was given by the expression:  $\gamma_0 / (\gamma_0 - 1) = A \exp(-\Delta E/RT)$  with  $A = 4 \times 10^{-2}$  and  $\Delta E = -2 \pm 5$  kJ mol<sup>-1</sup>. The value of  $\gamma_0$  equates to the bulk accommodation coefficient  $\alpha_b$ .

(c) details as (b). Uptake measured for a range of initial pH in the bulk droplets from 0 to 12 and at  $[\text{SO}_2]_0$  of  $10^{13}$  and  $10^{15}$  molecule  $\text{cm}^{-3}$ . Corrections for gas phase diffusion and additional acidification of the droplet due to SO<sub>2</sub> adsorption were applied to the uptake kinetics. The  $\gamma_0$  values displayed in the table are corrected values obtained at minimal droplet-gas interaction time (2 ms) and minimum SO<sub>2</sub> density ( $10^{13}$  molecule  $\text{cm}^{-3}$ ). The observed uptake rates over the pH range were significantly greater than predicted on the basis of the known rate of SO<sub>2</sub> reaction in bulk liquid water at pH > 5 and known Henry solubility at low pH. A mechanism involving formation of a surface complex HSO<sub>3</sub><sup>-</sup>-H<sup>+</sup> was proposed. Jayne et al. (1990) derive kinetic and thermodynamic parameters governing these surface interactions.

(d) Liquid jet of water (100  $\mu\text{m}$  diameter); H<sub>2</sub>O<sub>2</sub> was added in order to rapidly oxidize SO<sub>2</sub> to sulfate. The uptake kinetics of typically 100 ppm SO<sub>2</sub> in synthetic air was measured by analysis of the sulfate concentration in the collected jet waters using ion chromatography. Retrieval of  $\gamma$  using gas diffusion model to define gas-surface collision rate. The uptake was found to be dependent on [SO<sub>2</sub>] due to a change in surface pH. Addition of NH<sub>3</sub> to the SO<sub>2</sub>/air mixture minimized the pH dependence.

- (e) Uptake into droplet train of  $[\text{SO}_2] = 2.8$  to 28 ppmv at 50 mbar total pressure. Dilute aqueous  $\text{H}_2\text{O}_2$  is added to the condensed phase to rapidly oxidize dissolved  $\text{SO}_2$  to  $\text{H}_2\text{SO}_4$ . Uptake monitored by chemical analysis of the collected aqueous sulphate by ion chromatography. The gas/liquid contact time varied between 3.7 and 45 ms. The observed uptake coefficient  $\gamma$  is strongly dependent on the initial pH of the droplet in the range 4 to 11 and, at  $\text{pH} < 11$ , on the interaction time. Cited value of  $\gamma_0$  which is independent of pH is corrected for gas phase diffusion and represents the mass accommodation coefficient. (The uncorrected  $\gamma_0$  at  $t = 0$  was  $0.060 \pm 0.008$ ).
- (f) Uptake measurement of  $\text{SO}_2$  on counterflowing liquid  $\text{H}_2\text{O}$  using LIF detection of  $\text{SO}_2$  excited at 224.34 nm. The range of  $[\text{SO}_2]$  was between  $3.3 \times 10^{11}$  and  $1.7 \times 10^{14}$  molecule  $\text{cm}^{-3}$  in the temperature range 280–305 K at an average contact time of 37 ms and a total pressure range of 37–131 mbar.  $\gamma$  was determined from the concentration dependence of  $\text{SO}_2$  in the impinging flow field and was found to strongly depend on the pH of the condensed phase.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$$K_1 = \frac{[\text{H}^+][\text{HSO}_3^-]}{[\text{SO}_2 \cdot \text{H}_2\text{O}]}$$

$$\text{and } K_2 = \frac{[\text{H}^+][\text{SO}_3^{2-}]}{[\text{HSO}_3^-]}$$

5 which gives in turn:

$$[\text{S(IV)}]_{\text{total}} = p(\text{SO}_2)H \times \left\{ 1 + \frac{K_1}{[\text{H}^+]} + \frac{K_1K_2}{[\text{H}^+]^2} \right\}$$

$$= p(\text{SO}_2) \cdot H^*$$

This leads to the mole fractions and total dissolved S(IV) being a strong function of pH. Thus the effective Henry's solubility constant,  $H^*$ , decreases with pH. Over the pH range typical of atmospheric droplets (pH 2 to 6), most dissolved  $\text{SO}_2$  is in the form of  $\text{HSO}_3^-$ .

Following the usual formalism for the resistance model the net uptake coefficient of  $\text{SO}_2$  to an aqueous surface is given by:

$$Y_{\text{net}} = \left( \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{sol}} + \Gamma_b} \right)^{-1}$$

15 In the absence of chemical reaction in solution uptake is solubility limited, and the time-dependent resistance term for uptake into the diffusion-limited shell in the droplet is given by:

$$\Gamma_{\text{sol}} = \frac{4RTH^*}{c} \left( \frac{D_l}{t} \right)^{1/2} \quad (1)$$

20 In laboratory experiments the effect of aqueous chemistry on  $H^*$  was found to lead to a strong dependence of the uptake coefficient on pH. In addition, acidity due to the  $\text{H}^+$

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





from hydrolysis of SO<sub>2</sub> taken up into the surface layer, leading to lower effective surface pH, needs to be accounted for in the uptake kinetics. Assuming that gas-phase SO<sub>2</sub> is in equilibrium with SO<sub>2</sub>(aq) at the liquid surface and neglecting the second dissociation involving K<sub>2</sub>, the actual [H<sup>+</sup>] at the surface is given by:

$$[H^+] = H' \pm \left( H'^2 + K_w + K_1[SO_2] \right)^{1/2}$$

where

$$H' = \frac{1}{2} \left( [H^+]_0 - \frac{K_w}{[H^+]_0} \right)$$

Literature values for the Henry's constant, *H*, for SO<sub>2</sub>, and for the dissociation constants *K<sub>w</sub>* (= 10<sup>14</sup> M<sup>2</sup>) for water, and *K<sub>1</sub>* and *K<sub>2</sub>*, for bisulphite and sulphite ions respectively, are taken from Goldberg and Parker (1985):

<i>H</i> /M bar <sup>-1</sup>	3.42 × 10 <sup>-5</sup> exp(3133/ <i>T</i> )	273–300
<i>K<sub>1</sub></i> /M	1.4 × 10 <sup>-2</sup>	298
<i>K<sub>2</sub></i> /M	6.5 × 10 <sup>-8</sup>	298
<i>D<sub>1</sub></i> (SO <sub>2</sub> ,aq)/cm <sup>2</sup> s <sup>-1</sup>	8 × 10 <sup>-6</sup>	283

These are used to compute the effective solubility *H*<sup>\*</sup> for a specified pH, taking into account the effect of dissolved SO<sub>2</sub> (as described in Worsnop et al., 1989, and Ponche et al., 1993).

In alkaline droplets (e.g. pH = 11) surface saturation of SO<sub>2</sub> is negligible and *γ* is time independent. At lower pH *γ* decreases with time due to saturation of the surface layer. These effects demand careful attention to the design of experiments and the model used to interpret the results. Notwithstanding this, the results cited in the table from different experimental systems, are in reasonably good agreement, especially when the uncertainties and assumptions made are taken into account. Thus after correction for gas phase diffusion and exposure time, the values reported for the initial uptake coefficient at pH > 6, lie in the range 0.06 to 0.13.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Jayne et al. (1990) investigated the time and pH dependence of  $\gamma$  in detail. They found that in the high pH range the uptake is governed by the rate of reaction of  $\text{SO}_2$  with  $\text{H}_2\text{O}$  to form  $\text{HSO}_3^-$ . In this case  $\gamma_t$  at fixed pH/exposure time is given by the following expression:

$$\Gamma_b = \frac{4H^*RT(D_1k^1)^{1/2}}{c} \quad (2)$$

where  $k^1$  is the pseudo first order rate constant for reaction of  $\text{SO}_2$  with  $\text{H}_2\text{O}$ .

However, if the value of  $k^1$  in bulk aqueous solution measured by Eigen and co-workers ( $= 3.4 \times 10^6 \text{ s}^{-1}$ ;  $(D_1/k^1)^{0.5} = 15 \text{ nm}$ ) is employed in Eq. (2), the predicted  $\gamma$  is substantially lower than the observed values. The discrepancy also extends to low pH when uptake is limited by Henry's law solubility. Jayne et al. (1990) suggest that this discrepancy in uptake rate could be due to a rapid surface reaction to form an  $\text{HSO}_3^- - \text{H}^+$  surface complex, which is in equilibrium with gas-phase  $\text{SO}_2$ . This complex allows a more rapid entry of S(IV) into the bulk as  $\text{HSO}_3^-$  compared to transfer and reaction of  $\text{SO}_2$  molecules.

We interpret the surface complex as suggested by Jayne et al. as the adsorbed species as defined within the framework outlined in the *guide to the heterogeneous data sheets* for this Evaluation (IUPAC, 2012); see also Ammann and Pöschl (2007) for a complete time dependent modelling of this uptake process. The presence of a surface adsorbed partially hydrated  $\text{SO}_2$  species is also in line with second harmonic generation (SHG) experiments reported by Donaldson et al. (1995). The equilibrium constant denoted by  $A_{\text{eq}}^*$  by Jayne et al. is equal to  $K_{\text{LinC}}$  used here; similarly, the kinetic parameters  $k_{i-g} = k_{\text{des}}$  and  $k_{i-l} = k_{\text{sb}}$ . As discussed in detail by Jayne et al., it is difficult to constrain these parameters from the available data of the time dependence of the uptake coefficients.  $K_{\text{LinC}}$  was obtained from the slope of the time dependent uptake coefficient. Because the surface accommodation coefficient is not directly accessible

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

and

$$K_{\text{LinC}} = \frac{\alpha_s \bar{c}}{4k_{\text{des}}}$$

only the ratio of  $\alpha_s$  and  $k_{\text{des}}$  is constrained by the experiment at low  $\text{SO}_2$  pressures. The bulk accommodation coefficient is linked to these parameters via

$$\alpha_b = \alpha_s (1 - \theta) \frac{k_{\text{sb}}}{k_{\text{sb}} + k_{\text{des}}}$$

Where the surface coverage is given by  $\theta = \frac{K_{\text{LangC}}[\text{X}]_g}{1 + K_{\text{LangC}}[\text{X}]_g}$ , with  $K_{\text{LangC}} = K_{\text{LinC}}/N_{\text{max}}$ , which leads to the pressure dependence of  $\alpha_b$ .

In absence of an additional loss process of the surface complex at the surface itself, and assuming that transfer of  $\text{SO}_2$  into the bulk is driven by solubility only for acidic conditions, the uptake coefficient (Eq. 25 in guide to heterogeneous reaction data sheets) simplifies to:

$$\frac{1}{\gamma} \approx \frac{1}{\alpha_s} + \frac{1}{\Gamma_{\text{sb}}} + \frac{1}{\Gamma_{\text{sol}}} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{sol}}} = \frac{1 + K_{\text{LangC}}[\text{X}]_g}{\alpha_s} + \frac{1}{\Gamma_{\text{sol}}} = \frac{1}{\alpha_s} + \frac{\alpha_s \bar{c}}{4k_{\text{des}} N_{\text{max}}} [\text{X}]_g + \frac{1}{\Gamma_{\text{sol}}}$$

This is equivalent to the expression used by Jayne et al. to fit the time and pressure dependent data at low pH. We have adopted a value of  $\alpha_s = 1$ , with a large uncertainty.

Reducing  $\alpha_s$  by a factor of 10 leads to adjustments of  $k_{\text{sb}}$  and  $k_{\text{des}}$  of similar magnitude. These parameters describe quite well the experimental uptake coefficients observed over a range of pH and contact times in the droplet train experiments. Specifically, they lead to a value of  $\alpha_b$  of 0.11 at low  $\text{SO}_2$  pressure, consistent also with the other experiments.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

- Donaldson, D. J., Guest, J. A., and Goh, M. C.: The Journal of Physical Chemistry, 99, 9313–9315, 1995.
- Eigen, M., Kustin, K., and Maass, G.: Z. Phys. Chem. N. F., 30, 130–136, 1961.
- 5 Gardner, J. A., Watson, L. R., Adewuyi, Y. G., Davidovits, P., Zahniser, M. S., Worsnop, D., and Kolb, C. E.: J. Geophys. Res., 92, 10887–10895, 1987.
- Gardner, J. A., Watson, L. R., Adewuyi, Y. G., Van Doren, J. M., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: Biogenic Sulfur in the Environment, Am. Chem. Soc. Symp. Ser., vol. 393, edited by: Saltzman, E. S. and Cooper, W. J., 504, ACS
- 10 Washington D.C., 1989.
- Jayne, J. T., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 94, 6041–6048, 1990.
- Jayne, J. T., Gardner, J. A., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Geophys. Res., 95, 20559–20563, 1990b.
- 15 Li, L., Chen, M. Z., Zhang, Y. H., Zhu, T., Li, J. L., and Ding, J.: Atmos. Chem. Phys. 6, 2453–2464, 2006.
- Ponche, J. L., George, C., and Mirabel, Ph.: J. Atm. Chem., 16, 1–21, 1993.
- Shimono, A. and Koda, S.: J. Phys. Chem., 100, 10269–10276, doi:10.1021/jp9537670, 1996.
- 20 Welter, F., Schweighoefer, S., and Schurath, U.: The Proceedings of Eurotrac Symposium 90', edited by: Borrell, P. et al., 335, SPB Academic Publishing, 1991.
- Worsnop, D. R., Zahniser, M. S., Kolb, C. E., Gardner, J. A., Watson, L. R., Van Doren, J. M., Jayne, J. T., and Davidovits, P.: J. Phys. Chem., 93, 1159–1172, 1989.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.11

### HCl + H<sub>2</sub>O(l) → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
$(1.77^{+2.4}_{-1.7}) \times 10^{-1}$	274	Van Doren et al. (1990)	DT-LAS (a)
$(6.4^{+1.2}_{-0.8}) \times 10^{-2}$	294		
0.01–0.02	296.7	Kirchner et al. (1990)	LJ-HPLC (b)
0.24 ± 0.02	263	Schweitzer et al. (2000)	DT-MS (c)
0.13 ± 0.01	281		
0.24 ± 0.02	273	Li et al. (2002)	DT-MS (d)
0.10 ± 0.01	293		

#### Comments

(a) Droplets (200 μm diameter) crossing a laminar flow tube in transverse geometry. Trace gas analysis is by IR diode laser absorption at total pressures of 6.7–26.6 mbar. [HCl] was typically =  $7.3 \times 10^{12}$  molecule cm<sup>-3</sup>. Values listed are corrected for gas phase diffusion and distortion of velocity vector due to rapid uptake.

(b) The interaction of HCl (about 223 ppm) in synthetic air with a liquid jet of pure water. The uptake kinetics was studied by analyzing the liquid water content for Cl<sup>-</sup> using ion chromatography. The jet diameter was 90 μm, the jet length between 0.2 and 6 mm leading to contact times of 0.03 to 1.0 ms. The uptake coefficient was interpreted as a bulk accommodation coefficient. The results are regarded as lower limits due to uncertainties in the velocity distribution of the liquid jet. The

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



bulk accommodation coefficient showed a weak dependence on the mean speed of the jet ( $10\text{--}750\text{ cm s}^{-1}$ ) and was independent of the relative humidity (0 to 48 %) of the gas phase.

- (c) Droplets of  $80\text{--}150\text{ }\mu\text{m}$  diameter with a gas-liquid interaction time of  $0\text{--}20\text{ ms}$ . The HCl concentrations were in the range  $10^{12}\text{--}10^{14}\text{ molecule cm}^{-3}$ , with most experiments performed at  $10^{13}\text{ molecule cm}^{-3}$ . Rate of uptake was time-independent as well as independent of pH in the range  $7\text{--}14$ . Diffusion-corrected uptake coefficient is interpreted as a bulk accommodation coefficient  $\alpha_b$  having a significant negative temperature dependence.
- (d) Droplet train experiment with ethylene glycol/water mixtures. Droplet size was in the range  $70\text{ to }300\text{ }\mu\text{m}$ ,  $2.7\text{--}25.3\text{ }\mu\text{mbar}$  of (mostly) He in flow tube with trace gas species concentrations in the range  $2 \times 10^{13}\text{ to }3 \times 10^{14}\text{ molecule cm}^{-3}$ . The value of  $\alpha_b$  was independent of the interaction time in the range  $2.5\text{ to }17\text{ ms}$  and had a significant negative temperature dependence.  $\alpha_b$  varies continuously with increasing mole fraction of  $\text{H}_2\text{O}$  from a high (pure glycol) to a low (pure  $\text{H}_2\text{O}$ ) value which has been successfully modelled using surface tension data for glycol/ $\text{H}_2\text{O}$  solutions of variable composition.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	$7.4 \times 10^{-2}$	298
$\alpha_b$	$4.4 \times 10^{-6} \exp(2898/T)$	272–294
<i>Reliability</i>		
$\Delta \log(\alpha)$	$\pm 0.2$	298
$\Delta(E/R)$	$\pm 500\text{ K}$	272–294

## Comments on preferred values

Van Doren et al. (1990) interpret their diffusion corrected uptake coefficients as mass accommodation coefficients. The effective Henry's law constant for aqueous dissolution of HCl is large enough to make the characteristic phase relaxation time significantly longer than the time scale for the droplet train uptake experiment, typically one to a few ms such that there is no surface saturation of the gas-water interface. A surprising result is that both the absolute values of  $\alpha_b$  as well as its strong negative temperature dependence were found for HCl, HNO<sub>3</sub>, N<sub>2</sub>O<sub>5</sub>, H<sub>2</sub>O<sub>2</sub> and others that were investigated using the same experimental technique. This strongly points towards physical solvation as the rate-determining process whose absolute rate is not accessible as the bulk accommodation coefficient only describes the overall (complex) process and is therefore not an elementary reaction parameter.

The preferred values are taken from Van Doren et al. (1990), Schweitzer et al. (2000) and Li et al. (2002), despite the rather large uncertainty in the Van Doren et al. (1990) study. This fact is due in part to the use of the droplet train apparatus in transverse geometry which affords a restricted range of gas-liquid contact times and therefore a low dynamic range of the measurements. All three studies were performed using the same experimental technique, and the individual  $\alpha_b$  values are within the experimental uncertainties of the cited studies. The activation energy of  $\alpha_b$  is  $E = -24.1$  kJ/Mol.

The liquid jet technique used by Kirchner et al. (1990) that combines the measurement of the HCl uptake kinetics with the numerical simulation of the convective diffusion towards the thin H<sub>2</sub>O jet resulted in a best fit value  $0.02 \geq \alpha_b \geq 0.01$  of the bulk accommodation coefficient which is in agreement with the recommendation when regarded as a lower limiting value. The sensitivity of the experimental technique used on  $\alpha_b$  is low in the range  $1 \geq \alpha_b \geq 0.05$  and seems most suited for the measurement of values  $\alpha_b \leq 0.01$ . Moreover, the careful study of the pressure dependence of  $\gamma$  in the Li et al. (2002) study enabled an easy separation of  $\alpha_b$  from the measured uptake coefficients  $\gamma_{\text{obs}}$ .

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

Kirchner, W., Welter, F., Bongartz, A., Kames, J., Schweighofer, S., and Schurath, U.: J. Atmos. Chem., 10, 427–449, 1990.

5 Li, Y. Q., Zhang, H. Z., Davidovits, P., Jayne, J. T., Kolb, C. E., and Worsnop, D. R.: J. Phys. Chem. A, 106, 1220–1227, 2002.

Schweitzer, F., Mirabel, Ph., and George, C.: J. Phys. Chem. A, 104, 72–76, 2000.

Van Doren, J. M., Watson, L. R., Davidovits, P., Worsnop, D. R., Zahniser, M., and Kolb, C. E.: J. Phys. Chem., 94, 3265–3269, 1990.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





HBr + H<sub>2</sub>O(liq) → products

## Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
0.16	263	Schweitzer et al. (2000)	DT-MS (a)
$6.8 \times 10^{-2}$	281		
$0.26 \pm 0.02$	273	Li et al. (2002)	DT-MS (b)
$0.14 \pm 0.02$	283		

## 5 Comments

(a) 80–150  $\mu\text{m}$  diameter droplets with a gas-liquid interaction time of 0–20 ms. The HBr concentrations were in the range  $10^{12}$ – $10^{14}$  molecule  $\text{cm}^{-3}$ , with most experiments performed at  $10^{13}$  molecule  $\text{cm}^{-3}$ . Rate of uptake was time-independent as well as independent on pH in the range 7–14. Diffusion-corrected uptake coefficient is interpreted as a bulk accommodation coefficient  $\alpha_b$  having a significant negative temperature dependence.

(b) Droplet train experiment with ethylene glycol/water mixtures. Droplet size was in the range 70 to 300  $\mu\text{m}$ , 2.6–25 mbar of (mostly) He in flow tube with trace gas species concentrations in the range  $2 \times 10^{13}$  to  $3 \times 10^{14}$  molecule  $\text{cm}^{-3}$ . The value of  $\alpha_b$  was found to be independent of the interaction time in the range 2.5 to 17 ms, and temperature.  $\alpha_b$  varied continuously with increasing mole fraction of H<sub>2</sub>O from a high (pure glycol) to a low (pure H<sub>2</sub>O) value which has been successfully modelled using surface tension data for glycol/H<sub>2</sub>O solutions of variable composition.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_b$	$2.3 \times 10^{-2}$	298
$\alpha_b$	$1.3 \times 10^{-8} \exp(4290/T)$	263–281
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298
$\Delta(E/R)$	$\pm 600$ K	263–281

## Comments on preferred values

Although both Schweitzer et al. (2000) and Li et al. (2002) present bulk accommodation coefficients with a strong negative temperature dependence, there is disagreement of a factor of between 3 and 5 in the absolute values of  $\alpha_b$ . We have taken the data of Schweitzer et al. (2000) as the basis of our recommended values owing in part to the larger number of data taken within a larger temperature range.

## References

- Li, Y. Q., Zhang, H. Z., Davidovits, P., Jayne, J. T., Kolb, C. E., and Worsnop, D. R.: J. Phys. Chem. A, 106, 1220–1227, 2002.
- Schweitzer, F., Mirabel, Ph., and George, C.: J. Phys. Chem. A, 104, 72–76, 2000.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.13

### HI + H<sub>2</sub>O(liq) → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
$8 \times 10^{-2}$	278	Schweitzer et al. (2000)	DT-MS (a)
0.19	262		

#### 5 Comments

(a) Uptake experiment on fast droplet train of 80–150  $\mu\text{m}$  diameter with a gas-liquid interaction time of 0–20 ms. The HI concentrations were in the range  $10^{12}$ – $10^{14}$  molecule  $\text{cm}^{-3}$ , with most experiments performed at  $10^{13}$  molecule  $\text{cm}^{-3}$ . The rate of uptake was time-independent thus non-saturating as well as independent of pH in the range 7–14. The diffusion-corrected uptake coefficient is interpreted as a bulk accommodation coefficient  $\alpha_b$  with a significant negative temperature dependence.

10

## Preferred values

Parameter	Value	T/K
$\alpha_b$	$2.4 \times 10^{-2}$	298
$\alpha_b$	$6.35 \times 10^{-9} \exp(4519/T)$	262–278
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298
$\Delta(E/R)$	$\pm 300$ K	262–278

## Comments on preferred values

The uptake data were interpreted as bulk accommodation coefficients. Saturation effects of HI on the surface of the H<sub>2</sub>O drop were absent because the uptake kinetics on neutral water and 1 M NaOH aqueous solution were identical. The resulting activation energy is  $E = -37.6$  kJ/mol, the Arrhenius plot had a correlation coefficient of 0.9697.

## References

Schweitzer, F., Mirabel, Ph., and George, C.: J. Phys. Chem. A, 104, 72–76, 2000.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.14



### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
$\gamma$				
$5.8 \times 10^{-3}$	H <sub>2</sub> O film	273–293	Scheer et al. (1997)	WWFT/DT-FTIR/IC (a)
$1.2 \times 10^{-2}$	H <sub>2</sub> O droplets			
$1.2 \times 10^{-2}$	1 M NaOH(aq)			

### 5 Comments

(a) The uptake kinetics has been measured using the two complementary techniques of the wetted-wall flow-tube at atmospheric and reduced (89–90 mbar) pressure and the droplet train technique, with combined FTIR and HPLC detection. A slight negative temperature dependence of  $\gamma$  in the range indicated was observed. The rate of uptake of ClNO on aqueous NaCl (up to 1 M) and HCl (up to 1 M) was the same as on pure water. The hydrolysis products were HNO<sub>2</sub> and HCl, detected as NO<sub>2</sub><sup>-</sup> and Cl<sup>-</sup> in solution.

### Preferred values

Parameter	Value	T/K
$\gamma$	0.01	273–293
<i>Reliability</i>		
$\Delta \log(\alpha)$	±0.5	298

32197

ACPD

12, 32109–32472, 2012

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred values

The recommendation accepts the value of Scheer et al. (1997). The results show that ClNO is reactive towards liquid water surfaces. The products are HCl and HNO<sub>2</sub>; HNO<sub>2</sub> was released to the gas phase, but the efficiency of transfer to the gas-phase depended  
5 on the droplet acidity.

## References

Scheer, V., Frenzel, A., Behnke, W., Zetzsch, C., Magi, L., George, C., and Mirabel, Ph.: J. Phys. Chem. A, 101, 9359, 1997.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.15

### ClNO<sub>2</sub> + H<sub>2</sub>O(l) → products

#### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
$\gamma, \gamma_{ss}, \gamma_0$				
$\gamma < 10^{-5}$	H <sub>2</sub> O droplets	280	George et al. (1995)	DT-FTIR/IC (a)
$\gamma_{ss} = 4.8 \times 10^{-6}$	H <sub>2</sub> O film	291	Behnke et al. (1997)	WWFT-FTIR (b)
$\gamma_0 = 3.4 \times 10^{-6}$	H <sub>2</sub> O film	278	Frenzel et al. (1998)	WWFT-FTIR (c)
$\gamma_0 = 9.0 \times 10^{-6}$		282		
$\gamma_0 = 4.5 \times 10^{-6}$		287		
$\gamma_0 = (4.8 \pm 0.1) \times 10^{-6}$		291		
$\gamma_{ss} = (2.6 \pm 1.8) \times 10^{-6}$	H <sub>2</sub> O film	274	Fickert et al. (1998)	WWFT-MS (d)
<i>accommodation coefficient:</i>				
$\alpha_l = (9 \pm 4) \times 10^{-3}$				

#### 5 Comments

(a) Uptake study using the droplet train technique at typically 13 to 40 mbar total pressure. The trace gas density is monitored at the entry of the flow tube by FTIR. The amounts of Cl<sup>-</sup> and NO<sub>3</sub><sup>-</sup> in the droplets measured by ion chromatography was below the detection limit, therefore only an upper limit of  $\gamma$  is given. The continuous flow of ClNO<sub>2</sub> diluted in synthetic air at concentrations of up to 2500 ppm was generated by passing N<sub>2</sub>O<sub>5</sub> through a 30 cm long wetted wall flow tube interacting with a 3 M NaCl aqueous solution.

(b) Variable length (10–80 cm) wetted-wall flow tube at 1 bar of synthetic air. The liquid film thickness was between 0.8 and 1.2 mm at a surface speed of 2–10 cm s<sup>-1</sup>.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The gas flow rate was in the range 200–400 mL min<sup>-1</sup> at 291 K resulting in an average linear flow velocity of 15–30 cm s<sup>-1</sup>. The initial [ClNO<sub>2</sub>] was measured upstream, the unreacted ClNO<sub>2</sub> was measured downstream of the WWFT using FTIR. Between pure H<sub>2</sub>O and 1 M NaCl  $\gamma$  decreases by more than a factor of ten.

(c) Experimental details as (b). The uptake coefficients were obtained by fitting the initial measured time/position-dependent concentration profiles using a complex reaction mechanism including gas and liquid phase diffusion. The formation of ClNO<sub>2</sub> was also observed when Cl<sub>2</sub> interacted with the H<sub>2</sub>O film.

(d) Uptake study in a wetted-wall tubular flow reactor on a falling film of pure water and aqueous alkali halide salt solutions. The ClNO<sub>2</sub> reactant and products was monitored using a differentially-pumped MS with electron impact ionization. The value of  $\gamma_{ss}$  remains unchanged upon addition of 1 M NaCl whereas it increased upon addition of 0.1 M OH<sup>-</sup>. Analysis of uptake rates measured under gas-phase diffusion controlled conditions (0.5 M KOH film) gave the cited value for the accommodation coefficient on aqueous surfaces,  $\alpha_b = (9 \pm 4) \times 10^{-3}$ .

### Preferred values

Parameter	Value	T/K
$\alpha_b$	$9 \times 10^{-3}$	273–298
$\gamma_{ss}$	$3 \times 10^{-6}$	273–298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.5$	298



## Comments on preferred values

The results show consistently that ClNO<sub>2</sub> is non reactive towards liquid water surfaces. The recommended accommodation coefficient is based on the measurement of Fickert et al. (1998). The reactive uptake is controlled by hydrolysis which is very slow in pure water. The products have not been definitely identified.

## References

Behnke, W., George, C., Scheer, V., and Zetzsch, C.: J. Geophys. Res., 102, 3795–3804, 1997.

Fickert, S., Helleis, F., Adams, J., Moortgat, G. K., and Crowley, J. N.: J. Phys. Chem. A, 102, 10689–10696, 1998.

Frenzel, A., Scheer, V., Sikorski, R., George, C., Behnke, W., and Zetzsch, C.: J. Phys. Chem. A, 102, 1329–1337, 1998.

George, C., Behnke, W., Scheer, V., Zetzsch, C., Magi, L., Ponche, J. L., and Mirabel, Ph.: Geophys. Res. Lett., 22, 1505–1508, 1995.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

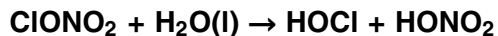
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.16



## Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
$\gamma$				
$0.027 \pm 0.0025$	pure water	275–285	Deiber et al. (2004)	DFT-MS (a)

5 **Comments**

(a) Uptake rates measured onto 200  $\mu\text{m}$  pure water droplets following loss of reactant in conventional droplet train apparatus. Droplet temperature controlled by evaporative cooling with adjustment of  $p(\text{H}_2\text{O})$ . Uptake coefficient determined with a simple correction for diffusion effects. No gas phase products were detected.

10 **Preferred values**

Parameter	Value	T/K
$\alpha_b$	0.11	273–290
$H(k^1)^{1/2}$ ( $\text{M atm}^{-1} \text{s}^{-1/2}$ )	$1.2 \times 10^5$	273–290
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.15$	273–290

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred values

The cited work is the only study of reactive uptake of ClONO<sub>2</sub> on liquid water substrates, all of the other reported studies used solid substrates (ice) or sulphuric acid solutions. These studies showed that uptake led to HOCl and HNO<sub>3</sub> formation by hydrolysis. No gas phase products were observed from uptake on water droplets due to the high solubility of the products. The measured uptake coefficient was independent of temperature over the small range investigated, but increased significantly when Br<sup>-</sup> was present in solution (see data sheet for ClONO<sub>2</sub> + Cl<sup>-</sup>/Br<sup>-</sup>). This was interpreted in terms of decreasing chemical lifetime of ClONO<sub>2</sub>(aq) due to reaction with Br<sup>-</sup>. This allowed evaluation of the accommodation coefficient by extrapolation of the uptake coefficients corrected for gas phase diffusion effects, to high [Br<sup>-</sup>] using the resistance model:

$$\frac{1}{\gamma} - \frac{1}{\gamma_{\text{diff}}} = \left\{ \frac{1}{\alpha_b} + \frac{c}{4HRT(D_1 k^1)^{0.5}} \right\}^{-1} \quad \text{where } k^1 = k^{II}[\text{Br}^-]$$

This gave  $\alpha_b = (0.108 \pm 0.011)$  at 274.5 K. The recommended accommodation coefficient is based on this analysis. The value of  $H\sqrt{k^1}$  can also be derived using the resistance model, assuming that the reduction in uptake coefficient into water compared to the accommodation limited rate, is due to slow rate of hydrolysis in solution. A value of  $D_1 = 5.0 \times 10^{-6} \text{ cm}^2 \text{ s}^{-1}$  was used for this analysis. There are no reported values of  $H$  or  $D_1$ .

## References

Deiber, G., George, C., Le Calve, S., Schweitzer, F., and Mirabel, Ph.: Atmos. Chem. Phys., 4, 1291–1299, 2004,  
<http://www.atmos-chem-phys.net/4/1291/2004/>.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



OH (g) + H<sub>2</sub>O (l) → products

## Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$\gamma_{\text{ss}} > 3.5 \times 10^{-3}$	275	Hanson et al. (1992)	WWFT-LIF (a)
$(4.2 \pm 2.8) \times 10^{-3}$ pH 5.6	293	Takami et al. (1998)	(b)
$(8.2 \pm 2.6) \times 10^{-3}$ pH 1			
$(1.2 \pm 0.3) \times 10^{-2}$ pH 11			
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
1	281–312	Takami et al. (1998)	(b)

## 5 Comments

(a) OH [(0.5 – 3) × 10<sup>11</sup> molecule cm<sup>-3</sup>] was generated through the reaction H + NO<sub>2</sub> and was monitored using LIF. Deionized water was flowed down the vertical flow tube at 274.5 K and 12 mbar total pressure. The value listed in the table was corrected for gas phase diffusion. The diffusion coefficient for OH in H<sub>2</sub>O has been taken to be the value of the self-diffusion coefficient for H<sub>2</sub>O.

(b) A carrier gas flow impinging on a flowing water surface at about 100 mbar total pressure. The uptake coefficient was derived from the OH gradient above the liquid measured using laser-induced fluorescence and after correction for gas phase diffusion. The uptake coefficient was found to decrease during the first 300 ms. The observed pH dependence was in agreement with bulk reaction limited uptake due to reaction with HSO<sub>4</sub><sup>-</sup>, OH<sup>-</sup>, O<sup>-</sup> and self reaction, simulated with a kinetic model of diffusion and reaction in the bulk liquid. For the highest uptake

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



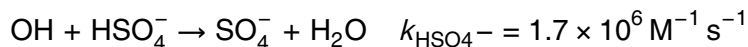
coefficients, best agreement between model simulation and data was obtained for  $\alpha_b = 1$ . A slight negative temperature dependence of the uptake coefficient was observed.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.1	275–310
$H^*$	$\exp(1010/T)(1 + 1.2 \times 10^{-12} \times 10^{\text{pH}})$	275–310
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	undetermined	
$\Delta \log(H^*)$	$\pm 0.7$	275–310

### Comments on preferred values

The available studies of the interaction of OH with liquid water agree that the bulk accommodation coefficient is large, likely close to 1. In both experiments, gas phase diffusion affected the observed kinetics. Hanson et al. (1992) likely observed uptake driven by the self reaction of OH in the bulk, if the Henry's Law coefficient is on the order of  $100 \text{ M atm}^{-1}$ . Takami et al. (1998) come to a similar conclusion and explain their pH dependent data by applying a kinetic model for the bulk liquid using the following reactions (Buxton et al., 1988):



The expression for the effective Henry's Law constant given above is adopted

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

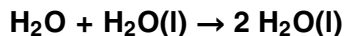
from Takami et al. (1998), with the temperature dependence as suggested by Hanson et al. (1992). The temperature dependence of the rate constant of the OH self reaction in water above tropospheric temperatures has been investigated in detail by Elliot et al. (1990) and Janik et al. (2007).

At low pH (due to the fast reaction with  $\text{HSO}_4^-$ ) and high pH (self reaction and increase of effective Henry's Law constant), the uptake coefficients were high enough to constrain the value of  $\alpha_b$  close to 1. Given the degrees of freedom to adjust the simulations, a lower limit to  $\alpha_b$  of 0.1 is recommended. The negative temperature dependence of the observed uptake coefficient was likely due to the interplay of temperature dependencies of the effective Henry's Law constant and the diffusion coefficient, rather than  $\alpha_b$ . High values for  $\alpha_b$  were predicted by molecular dynamics simulations (Roeselova et al., 2003).

## References

- Buxton, G. V., Greenstock, C. L., Helman, W. P., and Ross, A. B.: J. Phys. Chem. Ref. Data, 17, 513–886, 1988.
- Elliot, A. J., McCracken, D. R., Buxton, G. V., and Wood, N. D.: J. Chem. Soc. Faraday Trans., 86, 1539–1547, 1990.
- Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.
- Janik, I., Bartels, D. M., and Jonah, C. D.: J. Phys. Chem. A, 111, 1835–1843, 2007.
- Takami, A., Kato, S., Shimono, A., and Koda, S.: Chem. Phys., 231, 215–227, 1998.
- Roeselova, M., Jungwirth, P., Tobias, D. J., and Gerber, R. B.: J. Phys. Chem. B, 107, 12690–12699, 2003.

## VI.A1.18



## Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math>, <math>\alpha_t</math></i>			
$\alpha_b > 0.5$	263–298	Eames et al. (1997)	(a)
$\alpha_b > 0.1$ (dynamic conditions)	255–372	Marek and Straub (2001)	(b)
$\alpha_b < 0.1$ (stagnant conditions)			
$\alpha_b = 0.17 \pm 0.03$ ( $\text{H}_2^{17}\text{O}$ on liquid)	280	Li et al. (2001)	DT (c)
$\alpha_b = 0.32 \pm 0.04$	258		
$\alpha_t = 1.0 \pm 0.1$ ( $\text{D}_2\text{O}$ H/D exchange)	265–280		
$\alpha_b = 1.0$ (Ag-seeded $\mu\text{m}$ -size droplets)	250–290	Winkler et al. (2004)	(d)
$\alpha_b = 0.3$	295	Cappa et al. (2005)	LJ-MS (e)
$\alpha_b = 0.62 \pm 0.09$	295	Smith et al. (2006)	(f)
$\alpha_b = 0.18 \pm 0.08$	273	Zientara et al. (2008)	(g)
$\alpha_b = 0.13 \pm 0.02$	293		
$\alpha_t = 1.0 \pm 0.1$	273–293		
$\alpha_b = 0.57 \pm 0.06$ ( $\text{D}_2\text{O}$ )	295	Drisdell et al. (2008)	(h)

## 5 Comments

- (a) Literature review in which most of the early experimental results on the evaporation coefficient ( $\varepsilon$ ) of  $\text{H}_2\text{O}$  from liquid water surfaces were evaluated.
- (b) Review of literature for liquid  $\text{H}_2\text{O}$  up until 2001.
- (c) Droplet diameters 70–130  $\mu\text{m}$  or 150–300  $\mu\text{m}$ . Experiments conducted at various carrier gas pressures and for different carrier gases.

10

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- (d) Cloud expansion chamber study. Experimental water droplet growth curves monitored by Mie scattering of laser light.
- (e) Measurement of the evaporation rate from a 2.5  $\mu\text{m}$  radius  $\text{H}_2\text{O}$  microjet. The conditions were chosen so as to avoid recondensation of water vapour.
- 5 (f) The change in volume-averaged temperature of an evaporating train of microdroplets of 6–8  $\mu\text{m}$  radius was measured using Raman emission excited at 514.5 nm to obtain evaporation rates. The numerical fit used a model explicitly taking into account the  $\nabla T$ -gradient of the droplet.
- 10 (g) The temporal droplet radius evolution of a levitated water droplet of around 8  $\mu\text{m}$  diameter was observed in an electrostatic balance using angle-resolved Mie scattering at atmospheric pressure of  $\text{N}_2$  or air on evaporation.
- (h) Same as (f), but for  $\text{D}_2\text{O}$ .

### Preferred values

Parameter	Value	$T/\text{K}$
$\alpha_b$	$3.6 \times 10^{-5} \exp(2370/T)$	258–298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	258–298

### 15 Comments on preferred values

The uptake of gas phase water to liquid water has seen a long history of experiment, theory and debate due to its significance in cloud droplet and aerosol particle growth. The most recent evaluation of experimental and theoretical results within the

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





atmospheric science community on this subject has been compiled by Davidovits et al. (2006), Garrett et al. (2006) and updated by Kolb et al. (2010) and Davidovits et al. (2011). This debate has been initiated by the conflicting results between H<sub>2</sub>O uptake experiments to a train of droplets (Li et al., 2001) and one cloud droplet growth experiment (Winkler et al., 2004) and involved consideration of mass transfer aspects in the gas phase but also questions around energy dissipation in the condensed phase. We are not reiterating a detailed discussion here.

The preferred values refer to pure water and to experimental conditions that only marginally depart from equilibrium and where the growth rates due to condensation of water vapor are accordingly low. H/D isotope exchange experiments on a train of liquid water droplets essentially yield  $\alpha_s = 1.0$ , when we do not differentiate between surface accommodation and thermal accommodation. The differing kinetic behaviour of oxygen isotope exchange vs. H/D exchange reported by Li et al. (2001) provides convincing evidence for precursor mediated bulk accommodation as proposed by Davidovits (1991) that is consistent with the negative temperature dependence of  $\alpha_b$  observed in the droplet train experiments. Davidovits et al. (2004) in their comparison between Li et al. (2001) and Winkler et al. (2004) point out the agreement on the measured value of the thermal accommodation coefficient ( $0.85 \leq \alpha_t \leq 1.0$ ) but state the disagreement of the results on  $\alpha_b$ . A potential reason may lie in the vastly differing growth rates applied in the two experiments. We allow for sufficient uncertainty in the absolute value of  $\alpha_b$  to take into account possible systematic errors in the analysis of the experiments discussed in the above mentioned reviews. Similar negative temperature dependence of  $\alpha_b$  and values between 0.1 and 1 have also been found for other highly soluble trace gases to liquid substrates as well as for the uptake of H<sub>2</sub>O on ice (see datasheet V.A1.6, Crowley et al., 2010).

Since condensation and evaporation rates are equal at equilibrium, information about  $\alpha_b$  can also be retrieved from evaporation rate measurements. These were therefore also considered a source of relevant data. Eames et al. (1997) propose that the most probable value of the evaporation coefficient  $\varepsilon$  (numerically equal to  $\alpha_b$ ) is unity across

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

all experimental conditions. In contrast, Marek and Straub (2001) conclude in their review that the condensation and evaporation coefficients are different from each other, and a decline of both coefficients with increasing temperature and pressure is derived, at variance with theoretical predictions and the principle of microscopic reversibility.

5 The data of Zientara et al. (2008) obtained on a stationary droplet are consistent with a negative  $T$  dependence of  $\alpha_p$  and  $\alpha_t = 1$ . Other evaporation rate measurements were inconclusive about the temperature dependence (Smith et al., 2006; Drisdell et al., 2008). Although the emphasis of the work of Cappa et al. (2005) was placed on obtaining (H/D) isotope fractionation ratios, the cooling of the microjet along the axis was used to derive  $E_a = 55 \text{ kJ/Mol}$  for the evaporation rate of  $\text{H}_2\text{O}$  which is approximately 10 kJ/Mol larger than  $\Delta H_{\text{vap}}$ . Supporting TST (transition state) calculations by Cappa et al. (2007) were also consistent with a negative  $T$  dependence of the evaporation coefficient but with a different slope than in the experiments. They noted a temperature dependence of the pre-exponential factor, though. Therefore, most evaporation rate measurements are also in reasonable agreement with our recommended uptake parameters.

## References

Cappa, C. D., Drisdell, W. S., Smith, J. D., Saykally, R.J., and Cohen, R. C.: J. Phys. Chem. B, 109, 24391–24400, 2005.

20 Cappa, C. D., Smith, J. D., Drisdell, W. S., Saykally, R. J., and Cohen, R. C.: J. Phys. Chem. C, 111, 7011–7020, 2007.

Davidovits, P., Jayne, J. T., Duan, S. X., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 95, 6337–6340, 1991.

Davidovits, P., Worsnop, D. R., Jayne, J. T., Kolb, C. E., Winkler, P., Vrtala, A., Wagner, P. E., Kulmala, M., Lehtinen, K. E. J., Vesala, T., and Mozurkewich, M.: Geophys. Res. Lett., 31, L22111, doi:10.1029/2004GL020835, 2004.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

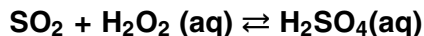
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Chem. Rev., 106, 1323–1354, 2006.
- Davidovits, P., Kolb, C. E., Williams, L. R., Jayne, J. T., and Worsnop, D. R.: Chem. Rev., 111, doi:10.1021/cr100360b, 2011.
- 5 Drisdell, W. S., Cappa, C. D., Smith, J. D., Saykally, R. J., and Cohen, R. C.: Atmos. Chem. Phys., 8, 6699–6706, 2008,  
<http://www.atmos-chem-phys.net/8/6699/2008/>.
- Eames, I. W., Marr, N. J., and Sabir, H.: Int. J. Heat Mass Transfer, 40, 2963–2973, 1997.
- 10 Garrett, B. C., Schenter, G. K., and Morita, A.: Chem. Rev., 106, 1355–1374, 2006.
- Li, Y. Q., Davidovits, P., Shi, Q., Jayne, J. T., Kolb, C. E., and Worsnop, D. R.: J. Phys. Chem. A, 105, 10627–10634, 2001.
- Kolb, C. E., Cox, R. A., Abbatt, J. P. D., Ammann, M., Davis, E. J., Donaldson, D. J., Garrett, B. C., George, C., Griffiths, P. T., Hanson, D. R., Kulmala, M., McFiggans, G.,
- 15 Pöschl, U., Riipinen, I., Rossi, M. J., Rudich, Y., Wagner, P. E., Winkler, P. M., Worsnop, D. R., and O'Dowd, C. D.: Atmos. Chem. Phys., 10, 10561–10605, 2010,  
<http://www.atmos-chem-phys.net/10/10561/2010/>.
- Marek, R. and Straub, J.: Int. J. Heat Mass Transfer, 44, 39–53, 2001.
- Smith, J. D., Cappa, C. D., Drisdell, W. S., Cohen, R. C., and Saykally, R. J.: J. Amer. Chem. Soc., 128, 12892–12898, 2006.
- 20 Winkler, P. M., Vrtala, A., Wagner, P. E., Kulmala, M., Lehtinen, K. E. J., and Vesala, T.: Phys. Rev. Lett., 93, 075701, doi:10.1103/PhysRevLett.93.075701, 2004.
- Zientara, M., Jakubczyk, D., Kolwas, K., and Kolwas, M.: J. Phys. Chem. A, 112, 5152–5158, 2008.



### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_0(\text{SO}_2)</math>, <math>\gamma(\text{H}_2\text{O}_2)</math></i>			
0.054 ± 0.005 (droplet pH = 11.5)	295	Gardner et al. (1987, 1989)	DT-TDL (a)
$\gamma_0 = (0.11 \pm 0.2)$ (droplet pH = 11.5)	273	Worsnop et al. (1989)	DT-TDL (b)
$\gamma(\text{H}_2\text{O}_2) = (0.18 \pm 0.2)$	273		
$(8.0 \pm 2.0) \times 10^{-3}$ pH = 0.25, 1 M $\text{H}_2\text{O}_2$ droplets	283	Jayne et al. (1990b)	DT-TDL (c)
0.020 ± 0.006 pH = 2, 1 M $\text{H}_2\text{O}_2$ droplets			
0.035 ± 0.004 pH = 4, 1 M $\text{H}_2\text{O}_2$ droplets			
0.06 + 0.14 – 0.03	300	Welter et al. (1990)	LJ-IC (d)
0.11 ± 0.01	298	Ponche et al. (1993)	DT-IC (e)

### Comments

- (a) The droplet size varied from 80 to 180  $\mu\text{m}$  and the residence time of the individual droplet ranged from 2.2 to 12.8 ms.  $[\text{SO}_2]$  ranged between 1.5 and  $15 \times 10^{12}$  molecule  $\text{cm}^{-3}$  and the total pressure was 27.9 mbar (He+ $\text{H}_2\text{O}$ ).
- (b) Experiment to determine the temperature-dependent mass accommodation coefficients of  $\text{SO}_2$  and  $\text{H}_2\text{O}_2$  on aqueous surfaces. A fast moving train of 200  $\mu\text{m}$  diam. water droplets propagated in a flow tube either in axial or transverse direction with  $[\text{SO}_2]_0 = 10^{13}$  molecule  $\text{cm}^{-3}$ . The interaction time ranged from 2 to 12 ms (axial) or from 0.5 to 2 ms (transverse), with a temperature range 260–292 K at pressures from 13.3 to 66.5 mbar (He+ $\text{H}_2\text{O}$ ) for the axial configuration, and 3.1 to 13.3 mbar for the transverse geometry. Uptake coefficients were corrected for gas diffusion with gas diffusion coefficients measured for  $\text{SO}_2$  and  $\text{SO}_2$  in  $\text{H}_2\text{O}$ , Ar and He in the same study. For  $\text{SO}_2$  at low pH,  $\gamma$  decreased with droplet

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

exposure time due to saturation, and increased with pH. No saturation effects were observed for H<sub>2</sub>O<sub>2</sub> uptake on the timescale of the experiments. Variation of droplet temperature by changing the water vapor content of the experimental atmosphere was used to determine the effect of surface temperature on the uptake rate. Temperature dependence of the initial uptake coefficient for SO<sub>2</sub> and H<sub>2</sub>O<sub>2</sub> was given by the expression:  $\gamma_0/(\gamma_0 - 1) = A \exp(-\Delta E/RT)$  with  $A = 4 \times 10^{-2}$  and  $\Delta E = -2 \pm 5 \text{ kJ mol}^{-1}$  for  $\gamma_0(\text{SO}_2)$ , and  $A = 3 \times 10^{-6}$  and  $\Delta E = -26 \pm 7 \text{ kJ mol}^{-1}$  for  $\gamma_0(\text{H}_2\text{O}_2)$  (see also data sheet VI.A1.1). The value of  $\gamma_0$  in both cases equates to the accommodation coefficient  $\alpha_b$ .

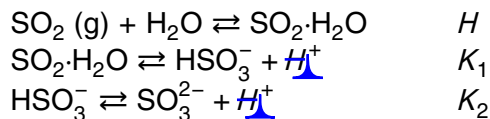
- (c) details as (b). Uptake of SO<sub>2</sub> measured for a range of initial pH in the bulk droplets from 0 to 8 and with H<sub>2</sub>O<sub>2</sub> content up to 1 M. The observed uptake coefficients over the pH and [H<sub>2</sub>O<sub>2</sub>] range were consistent with the reaction rate of H<sub>2</sub>O<sub>2</sub> with S(IV) measured in bulk solution.
- (d) Liquid jet of water (100 μm diameter); H<sub>2</sub>O<sub>2</sub> was added in order to rapidly oxidize SO<sub>2</sub> to sulfate. The uptake kinetics of typically 100 ppm SO<sub>2</sub> in synthetic air was measured by analysis of the sulfate concentration in the collected jet waters using ion chromatography.  $\gamma$  was retrieved by applying a gas diffusion model to correct the loss rate of SO<sub>2</sub>. The uptake was found to be dependent on [SO<sub>2</sub>] due to a change in surface pH. Addition of NH<sub>3</sub> to the SO<sub>2</sub>/air mixture minimized the pH dependence.
- (e) Uptake monitored by chemical analysis of the collected aqueous phase by ion chromatography. The gas/liquid contact time varied between 3.7 and 45 ms. Dilute aqueous H<sub>2</sub>O<sub>2</sub> is added to the condensed phase in order to rapidly oxidize dissolved SO<sub>2</sub> to H<sub>2</sub>SO<sub>4</sub>. The observed uptake coefficient is strongly dependent on the pH of the droplet in the range 4 to 11 and on the interaction time.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.11	268–298
$k/M^{-2} s^{-1}$	$2.1 \times 10^6 / (0.1 + [H^+])$	283
$k^l/s^{-1}$	see Eq. (ii)	
$D_1 (SO_2.aq)/cm^2 s^{-1}$	$8 \times 10^{-6}$	283
<i>Reliability</i>		
$\Delta(\log \alpha_b)$	$\pm 0.1$	268–298
$\Delta(\log k^l)$	$\pm 0.3$	268–298

## Comments on preferred values

When  $SO_2$  enters aqueous solution droplets it reacts with water molecules and an equilibrium is set up between the forms of S(IV): hydrated  $SO_2$ ,  $HSO_3^-$  and  $SO_3^{2-}$  via the following equilibria:



where 
$$H = \frac{[SO_2 \cdot H_2O]}{p(SO_2)}$$

$$K_1 = \frac{[H^+][HSO_3^-]}{[SO_2 \cdot H_2O]}$$

$$\text{and } K_2 = \frac{[\text{H}^+][\text{SO}_3^{2-}]}{[\text{HSO}_3^-]}$$

which gives in turn:

$$[\text{S(IV)}]_{\text{total}} = p(\text{SO}_2)H \times \left\{ 1 + \frac{K_1}{[\text{H}^+]} + \frac{K_1K_2}{[\text{H}^+]^2} \right\}$$

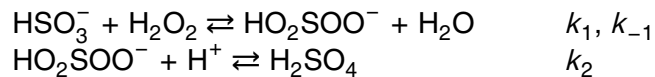
$$= p(\text{SO}_2) \cdot H^*$$

This leads to the equilibrium mole fractions and total dissolved S(IV) being a strong function of pH. Thus the effective Henry's solubility constant,  $H^*$ , decreases with pH. Literature values for the Henry's constant,  $H$ , for  $\text{SO}_2$ , and for the dissociation constants  $K_w (=10^{14} \text{ M}^2)$  for water, and  $K_1$  and  $K_2$ , for bisulphite and sulphite ions respectively, are taken from Goldberg and Parker (1985):

$H/\text{M bar}^{-1}$	$3.42 \times 10^{-5} \exp(3133/T)$	273–300
$K_1/\text{M}$	$1.4 \times 10^{-2}$	298
$K_2/\text{M}$	$6.5 \times 10^{-8}$	298

These are used to compute the effective solubility  $H^*$  for a specified pH, taking into account the effect of dissolved  $\text{SO}_2$  (as described in Worsnop et al., 1989, and Ponche et al., 1993).

Over the pH range typical of atmospheric droplets (pH 2 to 6), most dissolved S(IV) is in the form of  $\text{HSO}_3^-$ . In the presence of  $\text{H}_2\text{O}_2$  the  $\text{HSO}_3^-$  is oxidised to sulphuric acid by the following mechanism:



The rate law for this reaction in bulk solution has been given by Martin and

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Damschen (1981) as:

$$-\frac{d[S(IV)]}{dt} = \frac{d[S(VI)]}{dt} = k[H^+][H_2O_2][HSO_3^-]$$

where

$$k(M^{-2}s^{-1}) = \frac{2.1 \times 10^6}{0.1 + [H^+]} \quad \text{at 283 K} \quad (i)$$

5 Taking account of the equilibria following entry of  $SO_2$  into solution the pseudo first order rate constant for aqueous reaction of  $HSO_3^-$  with  $H_2O_2$ ,  $k^1$ , for a given pH is given by:

$$k^1(s^{-1}) = \frac{H}{H^*} kK_1[H_2O_2] \quad (ii)$$

and the reactive uptake coefficient for  $SO_2$  into aqueous droplets containing  $H_2O_2$  is given by the expression:

$$\Gamma_b = \frac{4H^*RT}{c} (D_1k^1)^{1/2} \quad (iii)$$

The overall uptake of  $SO_2$  is a convolution of this reactive uptake, which is time independent, with the time-dependent  $SO_2$  uptake due to solubility without added  $H_2O_2$ , which is liquid diffusion controlled and also contains a term to represent the surface adsorption of  $SO_2$  molecules (see data sheet for  $SO_2 + H_2O(aq)$ : VI.A1.10).

15 Jayne et al. demonstrated that at  $pH < 3$  the measured uptake coefficient both in magnitude and pH dependence, at contact times of 2–10 ms, was in reasonable accord with the bulk  $S(IV) + H_2O_2$  reaction rate, given by the expressions above. Above  $pH 3$  the uptake of  $SO_2$  increases but in a complex manner due to interaction of rapidly varying diffusion controlled solubility and  $S(IV) + H_2O_2$  reaction kinetics. No evidence was found for a surface reaction and furthermore the solubility of  $H_2O_2$  is high, and so equilibrium between gas phase and bulk phase is established rapidly The maximum

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





overall reactive uptake of SO<sub>2</sub> is limited by the value of α<sub>b</sub> of 0.11. The rate coefficient for oxidation, k<sup>1</sup>, reduces with increasing pH due to the rate determining reaction of the intermediate HO<sub>2</sub>SOO<sup>-</sup> ion with H<sup>+</sup> (as represented in Eq. i) but the overall reaction rate increases with pH since there is more S(IV) available for reaction (Eq. ii). Because the reactive uptake is liquid diffusion controlled the reaction rate enters as the square root in (Eq. iii).

The uptake coefficients can be calculated from:

$$\gamma_{\text{rxn}} = \left( \frac{1}{\alpha_b} + \frac{1}{\Gamma_b} \right)^{-1}$$

The recommended value for α<sub>b</sub> is based on the measurements of Worsnop et al. (1989) and Ponche et al. (1993). Note that this value is only recommended for low SO<sub>2</sub> pressures. For higher SO<sub>2</sub> pressures, the parameterization explained on the data sheet for SO<sub>2</sub> + H<sub>2</sub>O(aq) = VI.A1.10 should be used. The rate constant for bulk phase oxidation reaction of S(IV) by H<sub>2</sub>O<sub>2</sub> (Eq. i) is taken from Martin and Damschen (1981). This parameterisation gives a good fit to the laboratory measurement of Jayne et al. (1990) at [H<sub>2</sub>O<sub>2</sub>] = 1 M and Ponche et al. (1993) at [H<sub>2</sub>O<sub>2</sub>] = 0.16 M.

## References

Gardner, J. A., Watson, L. R., Adewuyi, Y. G., Davidovits, P., Zahniser, M. S., Worsnop, D., and Kolb, C. E.: J. Geophys. Res., 92, 10887–10895, 1987.

Gardner, J. A., Watson, L. R., Adewuyi, Y. G., Van Doren, J. M., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: Biogenic Sulfur in the Environment, Am. Chem. Soc. Symp. Ser., vol. 393, edited by: Saltzman, E. S. and Cooper, W. J., 504, ACS Washington D.C., 1989.

Goldberg, R. N. and Parker, V. B.: J. Res. Nat. Bur. Stand., 90, 341–358, 1985.

Jayne, J. T., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 94, 6041–6048, 1990a.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Jayne, J. T., Gardner, J. A., Davidovits, P., Worsnop, D. R., Zahniser, M. S. and Kolb, C. E.: J. Geophys. Res., 95, 20559–20563, 1990b.

Martin, L. R. and Damschen, D. E.: Atmos. Environ., 15, 1615–1621, 1981.

Penkett, S. A., Jones, B. M. R., Brice, K. A., and Eggleton, A. E. J.: Atmos. Environ., 13, 123–137, 1979.

Ponche, J. L., George, C., and Mirabel, Ph.: J. Atm. Chem., 16, 1–21, 1993.

Welter, F., Schweighoefer, S., and Schurath, U.: The Proceedings of Eurotrac Symposium 90', edited by: Borrell, P. et al., 335, SPB Academic Publishing, 1991.

Worsnop, D. R., Zahniser, M. S., Kolb, C. E., Gardner, J. A., Watson, L. R., Van Doren, J. M., Jayne, J. T., and Davidovits, P.: J. Phys. Chem., 93, 1159–1172, 1989.

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A1.20

### HO<sub>2</sub> + H<sub>2</sub>O (aq) → products

#### Experimental data

	Temp./K	Reference	Technique/Comments
Accommodation coefficients: $\alpha$ >0.01	275	Hanson et al. (1992)	WWFT-LIF (a)

#### 5 Comments

(a) Uptake of HO<sub>2</sub> ( $5 - 30 \times 10^{10}$  molecule cm<sup>-3</sup>) to a film (0.2 mm thick) of de-ionised water, or water containing  $10^{-3}$  M CuSO<sub>4</sub>. HO<sub>2</sub> was formed in the reaction of F with H<sub>2</sub>O<sub>2</sub> and detected as OH after reaction with NO. HO<sub>2</sub> uptake was limited by diffusion through the 1 Torr of He bath gas. Levels of HO<sub>2</sub> were sufficiently low to neglect loss due to gas-phase self reaction. Addition of CuSO<sub>4</sub> had no effect on HO<sub>2</sub> loss rates.

#### Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.5	270–300
<i>Reliability</i>		
$\Delta(\log \alpha_b)$	undetermined	

32219

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

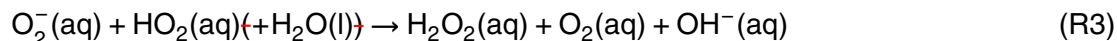
Interactive Discussion



## Comments on preferred values

The sole experimental study of the uptake of HO<sub>2</sub> to pure water returned a lower limit to the accommodation coefficient of  $\alpha_b > 0.01$ . This is consistent with molecular dynamics calculations of the HO<sub>2</sub>-water interaction (Morita et al., 2004) which suggest that the accommodation coefficient could be unity and also with more sensitive experiments on other aqueous surfaces such as (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> and NH<sub>4</sub>HSO<sub>4</sub> (see datasheets VI.A3.09 and VI.A3.10). We thus prefer a value of  $\alpha_b$  of  $>0.5$ .

The uptake of HO<sub>2</sub> in aqueous solution with pH  $> 5$ , is presently believed to be driven by self-reaction and acid-base dissociation of HO<sub>2</sub> (pKa  $\sim 4.7$ ) with formation of H<sub>2</sub>O<sub>2</sub> (Reactions R2, R3). In the presence of transition metal ions (TMI) the reaction of HO<sub>2</sub> and especially O<sub>2</sub><sup>-</sup> (Reaction R4) can be important:



If a first-order loss process for HO<sub>2</sub> or O<sub>2</sub><sup>-</sup> in the aqueous phase dominates (e.g. reaction with TMI such as Cu(II)), and assuming equal rates of reaction throughout the particle, the uptake coefficient can be calculated from the expression below:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{\bar{c}}{4H^{\text{eff}}RT\sqrt{k_{\text{TMI}}[\text{TMI}]D_1}}$$

$H^{\text{eff}} = H^{\text{HO}_2} (1 + K_{\text{eq}}/[\text{H}^+])$ ,  $K_{\text{eq}} = 2.1 \times 10^{-5} \text{ M}$  at 298 K (Jacob, 2000),  $H^{\text{HO}_2} = 9.5 \times 10^{-6} \exp(5910/T)$  (Hanson et al., 1992) and  $D_1 = [1 \times 10^{-5}(T/298)]/(1.09 \times 10^8 \exp(-0.0687T) + 0.873) \text{ cm}^2 \text{ s}^{-1}$  (Schwartz, 1984; Thornton et al., 2008) where the

32220

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



denominator in the  $D_1$  term was derived from a fit to the water viscosity data of Hallett (1963).

According to the reaction scheme above, in the absence of TMI, the rates of loss of aqueous-phase  $\text{HO}_2$  are quadratically dependent on  $[\text{HO}_2]_{\text{aq}}$  and  $[\text{O}_2^-]_{\text{aq}}$  and are thus strongly dependent on the gas-phase concentration of  $\text{HO}_2$ . At low  $\text{HO}_2$  concentrations (e.g. as found in the atmosphere) the liquid phase reactions become rate limiting and  $\gamma$  is expected to be much smaller as observed in dilute solutions by Mozurkewich et al. (1987) and the simple formalism above breaks down. Thornton and Abbatt (2005) suggest that the rate of loss of  $\text{HO}_2$  from the gas-phase (in molecule  $\text{cm}^{-3} \text{s}^{-1}$ ) is best described by a system in thermodynamic (Henry's law) equilibrium so that (Thornton et al., 2008):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{3\bar{c}N_A}{8000 (H^{\text{eff}}RT)^2 k_{\text{aq}}[\text{HO}_2]r}$$

$k_{\text{aq}}$  can be calculated from the rate coefficients for Reaction (R2) ( $k_2$ ) and Reaction (R3) ( $k_3$ ) (Bielski et al., 1985) and the pH:

$$k_{\text{aq}} = \frac{k_2 + \left(\frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}}\right) k_3}{\left(1 + \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}}\right)^2}$$

This formalism predicts that the loss of  $\text{HO}_2$  to particles is favoured by high  $\text{HO}_2$  mixing ratios, low temperatures (higher solubility) and low pH. At low concentrations of  $\text{HO}_2$  (whereby the self reaction and reaction with  $\text{O}_2^-$  are inefficient), values of  $\gamma$  of  $<0.005$  are calculated, which are however much less than the uptake coefficients of Taketani et al. (2008) who investigated the uptake of low concentrations of  $\text{HO}_2$  to  $(\text{NH}_4)_2\text{SO}_4$  particles. In addition, as discussed by Hanson et al. (1992) and Thornton and Abbatt (2005), there is considerable uncertainty (factor of 2.5) associated with the solubility of  $\text{HO}_2$  ( $H^{\text{HO}_2}$ ) and its temperature dependence. The above schemes also do

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



not account for the RH dependence of uptake observed for  $(\text{NH}_4)_2\text{SO}_4$  particles (see datasheet VI.A3.10).

Until these apparent discrepancies have been resolved by further experiments, we make no recommendation for parameterising  $\text{HO}_2$  uptake to aqueous aerosol. We refer to recent publications for a more detailed description of the effect of different parameterisation schemes (Thornton et al., 2008; Macintyre and Evans, 2011).

## References

Bielski, B. H. J., Cabelli, D. E., Arudi, R. L., and Ross, A. B.: J. Phys. Chem. Ref. Data, 14, 1041–1100, 1985.

Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.

Jacob, D. J.: Atmos. Env., 34, 2131–2159, 2000.

Macintyre, H. L. and Evans, M. J.: Atmos. Chem. Phys., 10, 7409–7414, 2010, <http://www.atmos-chem-phys.net/10/7409/2010/>.

Morita, A., Kanaya, Y., and Francisco, J. S.: J. Geophys. Res.-Atmos., 109, 2004.

Mozurkewich, M., McMurry, P. H., Gupta, A., and Calvert, J. G.: J. Geophys. Res.-Atmos., 92, 4163–4170, 1987.

Schwartz, S. E.: J. Geophys. Res.-Atmos., 89, 1589–1598, 1984.

Taketani, F., Kanaya, Y., and Akimoto, H.: J. Phys. Chem. A, 112, 2370–2377, doi:10.1021/jp0769936, 2008.

Thornton, J. and Abbatt, J. P. D.: J. Geophys. Res.-Atmos., 110, D08309, doi:10.1029/2004JD005402, 2005.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Appendix A2

### Uptake on deliquesced halide salts

#### VI.A2.0

#### O<sub>3</sub> + Cl<sup>-</sup>/Br<sup>-</sup>/I<sup>-</sup> (aq) → products

#### 5 Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake, Accommodation coefficients: <math>\gamma</math>, <math>\alpha</math></i>			
$\gamma_{ss} > 2 \times 10^{-3}$ ( $\pm 20\%$ ) (SnCl <sub>2</sub> solution)	276	Utter et al. (1992)	WWFT (a)
$\alpha_b = 0.10$	277	Hu et al. (1995)	DT-MS (b)
$1/\gamma_1 = (85_{-10}^{+20})\alpha(I^-)^{-\frac{1}{2}} + (10 \pm 5)$ with [I <sup>-</sup> ] between 0.5 and 3.0 M			
$\gamma_{ss} = 3.7 \times 10^{-3}$ ( $a(I^-) = 0.36$ M)	282	Magi et al. (1997)	DT-MS (c)
$\gamma_{ss} = 1.16 \times 10^{-2}$ ( $a(I^-) = 2.89$ M)	282		
$\alpha_b > 0.1$			
$\gamma_{ss} < 1.0 \times 10^{-4}$ (unbuffered and buffered (pH = 7.2) NaCl aerosol)	300	Abbatt and Waschewsky (1998)	AFT-CIMS (d)
$\alpha_b > 2.0 \times 10^{-2}$ (NaI solution)	298	Schütze and Herrmann (2002)	(e)
$\alpha_b = 0.6_{-0.5}^{+0.4}$ (deliquesced KI aerosol, rh = 72 %)	293	Rouvière et al. (2010)	AFT (f)
$\gamma_0 = (1.1 \pm 0.2) \times 10^{-2}$ (excess I <sup>-</sup> )			
$\gamma_0 = 4.4 \times 10^{-4}$			
(mixed KI/NaCl aerosol with [I <sup>-</sup> ] = 0.9 M, [Cl <sup>-</sup> ] = 6.1 M)			
$1/\gamma_s = (1.4 \pm 0.2) \times 10^{-7} [O_3] + 0.06 \times 10^7$	273	Oldridge and Abbatt (2011)	CWFT (g)
$\gamma_1 = (7.7 \pm 1.8) \times 10^{-8}$	273		

#### Comments

(a) Uptake of O<sub>3</sub> (10<sup>11</sup> molecule cm<sup>-3</sup>) to a liquid film (20 μm thick). The total pressure was between 13.6 to 14.6 mbar with half of it being H<sub>2</sub>O vapor. No uptake of O<sub>3</sub> was observed into pure deionized water, but  $\gamma$  increased as a function of the con-

ACPD

12, 32109–32472, 2012

Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



centration of the trapping agent (e.g.,  $\text{SnCl}_2$ ) to values in the  $10^{-2}$  range. The measured loss rate was limited by the diffusion of  $\text{O}_3$  to the wetted wall when  $\gamma$  was larger than 0.02 in the presence of solutions of scavengers  $\text{Na}_2\text{SO}_3$ ,  $\text{Na}_2\text{S}_2\text{O}_3$  and  $\text{SnCl}_2$ .

- (b) Droplet diameters in the range 120–250  $\mu\text{m}$ . The concentration of  $\text{O}_3$  varied from  $5 \times 10^{12}$  to  $1 \times 10^{14}$  molecule  $\text{cm}^{-3}$  at a gas residence time between 2 to 15 ms. Detection of the reactant gas is by residual gas electron-impact quadrupole mass spectrometry. The partial pressure of  $\text{H}_2\text{O}$  in the reacting zone is in the range 2.9 to 23.3 mbar corresponding to temperatures  $-10$  to  $20^\circ\text{C}$  using Kr as an internal standard.  $\text{H}_2\text{O}$  vapor was mixed with variable amounts of He. The iodide ion concentration was in the range 0.5 to 3.0 M.
- (c) Uptake on a train of aqueous NaI droplets (80 to 150  $\mu\text{m}$ ) in a laminar flow tube operated in the range 275–293 K. The  $\text{O}_3$  concentration in the flow tube was measured at  $m/e$  46 ( $\text{NO}_2^+$ ) after titration by NO and conversion to  $\text{NO}_2$ . No measurable uptake of  $\text{O}_3$  on pure water was observed. The uptake into aqueous NaI was limited by reaction of  $\text{O}_3$  with  $\text{I}^-$  as shown by the dependence of  $\gamma_{\text{ss}}$  on the ionic strength of the solution. From the intercept and linearity of the plot of  $1/\gamma$  vs.  $a(\text{I}^-)^{-1/2}$  Magi et al. (1997) derive a lower limit of  $\alpha_b$  of 0.1.
- (d) The NaCl aerosol had a bimodal distribution (large average diameter: 2–4  $\mu\text{m}$ , small average diameter:  $<1 \mu\text{m}$ ) with typical surface areas and particle number densities in the range of  $1 - 6 \times 10^{-3} \text{ cm}^2/\text{cm}^3$  and  $1 - 4 \times 10^4$  particles/ $\text{cm}^3$ , respectively. The value of  $\gamma$  is an upper limit and is given for unbuffered and NaCl aerosol buffered at  $\text{pH} = 7.2$  using  $\text{NaH}_2\text{PO}_4/\text{Na}_2\text{HPO}_4$ . A slow wall-catalyzed reaction occurred when working with HCl-acidified NaCl aerosol at  $\text{pH} = 0.3$  resulting in  $\text{Cl}_2$ . This reaction presumably occurs on NaCl aerosol particles adhering to the walls of the flow tube.



(e) Uptake onto a static single drop (2–3 mm in diameter) containing NaI ( $2 \times 10^{-2}$  to 4 M) was monitored by time-resolved UV/Vis absorption spectroscopy in the range 240 to 800 nm. The time-dependent values of the optical density of  $I_3^-$  were analyzed to derive uptake coefficients  $\gamma$  in the range  $2.0 \times 10^{-5}$  to  $1.5 \times 10^{-2}$ .

(f) Aerosol flow tube experiment at atmospheric pressure using deliquesced KI and KI/NaCl aerosol. Aerosol mode ranged from 70 to 573 nm and  $O_3$  concentration detected by CLD from 70 to 300 ppb. Iodide excess for pseudo first-order conditions could not be maintained in all experiments such that numerical data analysis had to be applied in order to retrieve the bulk accommodation coefficient  $\alpha_b$  and the second-order rate constant  $k^{II}$  for the liquid phase reaction  $O_3 + I^-$ . The remainder of the parameters for data retrieval is effectively identical to the work of Magi et al. (1997).

(g) CWFT study at 120 mbar of air coupled to  $I^-$ -CIMS detection of  $O_3$ ,  $H_2O$  and  $Br_2$ .  $O_3$  was monitored at 254 nm and was in excess relative to bromide. The detection limit for most negative ions was in the range  $10^9$ – $10^{10}$  molecule  $cm^{-3}$ . The liquid sample was positioned in a container lying on the bottom of the horizontal flow tube and consisted of an aqueous solution of 0.55 M NaCl and 8.5 mM KBr at pH = 1.97 exposed to  $O_3$  at 273 K. Uptake coefficients were derived from the rate of appearance of  $Br_2$  as product assuming no other reactive losses other than the reaction  $2Br^- + O_3 + 2H^+ \rightarrow Br_2 + O_2 + H_2O$ .

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Preferred values

Parameter	Value for Iodide ion as a reactant	T/K
$\alpha_b$	>0.1	298
$k^{\text{II}}/\text{M}^{-1} \text{s}^{-1}$	$(2.4 \pm 1.3) \times 10^9$ (average $[\text{I}^-]$ )	293
	$(3.2 \pm 1.5) \times 10^8$	275
$D_{\text{I}}/\text{cm}^2 \text{s}^{-1}$	$1.85 \times 10^{-5}$	293
$H/M \text{atm}^{-1}$	$1.15 \times 10^{-2} \exp(2560(1/T - 1/298))$	273–293
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	undetermined	298
$\Delta \log(k^{\text{II}})$	$\pm 0.2$	298

Parameter	Value for Bromide ion as a reactant	T/K
$K_{\text{LangC}}/\text{cm}^3 \text{molecule}^{-1}$	$10^{-13}$	233–273
$k_b^{\text{II}}/\text{M}^{-1} \text{s}^{-1}$	$6.3 \times 10^8 \exp(-4450/T)$	273–298
$D_{\text{I}}$	$8.9 \times 10^{-6} \text{cm}^2 \text{s}^{-1}$	273
$H/M \text{atm}^{-1}$	$1.15 \times 10^{-2} \exp(2560(1/T - 1/298))$	273
<i>Reliability</i>		
$\Delta \log(k^{\text{II}})$	$\pm 0.3$	273–298
$\Delta \log K_{\text{LangC}}$	$\pm 0.5$	233–273

## Comments on preferred values

5  $\text{O}_3 + \text{I}^-$

The datasets of Hu et al. (1995), Magi et al. (1997), Rouvière et al. (2010) concur in that the reaction proceeds without a significant surface component and can be de-

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

scribed by the following resistor expression:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT \sqrt{k^{II} a(I^-) D_1}}$$

These studies indicate that  $\alpha_b > 0.1$ .

There is excellent agreement in the retrieved parameters  $\alpha_b$  and  $k^{II}$  of Rouvière et al. (2010) and Magi et al. (1997) despite the different types of experiments. The liquid phase rate constant  $k^{II}$  obtained in both studies agreed with literature values (Garland et al. (1980) derive  $2 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$  at 298 K) and the values reported by Magi et al. (1997) are recommended. The value of  $k^{II} = 4 \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$  at 277 K derived by Hu et al. (1995) seems to be larger by at least a factor of two when compared to literature values for  $\text{O}_3 + \text{I}^-$ .

The effects of salting out (decrease of  $H$  with increasing scavenger concentration) on  $k^{II}$  is within the experimental uncertainty of the data and has been ignored in the data evaluation. The values for  $k^{II}$  displayed in the Table therefore correspond to an average iodide concentration.

~~The uptake coefficient of  $\text{O}_3$  on deliquescent aerosol containing iodide decreases with increasing coverage of saturated fatty acids depending on the physical and chemical properties of the surfactant (Rouvière and Ammann, 2010).~~

$\text{O}_3 + \text{Br}^-$

The experiments of Oldridge and Abbatt (2011), the only ones reported in the literature on the interaction of  $\text{O}_3$  with liquid aqueous bromide solutions at 273 K, are interpreted in terms of both a surface as well as a bulk reactivity component of excess  $\text{O}_3$  reacting with a constant supply of surface bromide in terms of two limiting cases. The surface reaction is invoked, because at low  $\text{O}_3$  concentration, the uptake coefficient was about an order of magnitude higher than predicted by bulk reaction alone, and because  $\gamma$  falls off with increasing  $\text{O}_3$  concentration similar to other surface reactions analyzed in terms

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

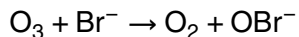
Printer-friendly Version

Interactive Discussion



of Langmuir-Hinshelwood kinetics. A concurrent bulk and surface reaction agrees, at least qualitatively, with the laboratory study of Clifford and Donaldson (2007) on the interaction of O<sub>3</sub> with bromide at the air-aqueous interface in which the authors have studied the occurrence of a pH change at the interface using a surface-adsorbed pH indicator.

The reaction of O<sub>3</sub> with aqueous bromide proceeds via:



We recommend using the bulk reaction rate constant measured by Haag and Hoigné (1983) over the temperature range of interest here. At low pH (pH = 2 in the experiments by Oldridge and Abbatt), the likely fate of OBr<sup>-</sup> is the reaction of HOBr with bromide or chloride in sea salt to form Br<sub>2</sub> or BrCl, respectively.

The full expression for the uptake coefficient of O<sub>3</sub> to aqueous bromide solutions based on the resistor model including both surface and bulk reaction is:

$$\frac{1}{\gamma} = \frac{1}{\alpha_s} + \frac{1}{\Gamma_s + \left( \frac{1}{\Gamma_{sb}} + \frac{1}{\Gamma_b} \right)^{-1}}$$

Since the bulk accommodation coefficient is likely to be >0.1 and the observed steady state uptake coefficients are 10<sup>-6</sup> at maximum, surface accommodation and surface to bulk transfer are not limiting ( $\alpha_s \gg \gamma$ ;  $\Gamma_{sb} \gg \Gamma_b$ ) and the expression simplifies to:

$$\gamma \approx \Gamma_s + \Gamma_b$$

For the surface reaction, the expression for a Langmuir-Hinshelwood reaction is:

$$\Gamma_s = \frac{4k_s^{\text{II}}[\text{Br}^-]_s K_{\text{LangC}} N_{\text{max}}}{\bar{c} \left( 1 + K_{\text{LangC}} [\text{O}_3]_g \right)}$$

32228

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The surface concentration of bromide ions can be estimated by normalizing the bulk mole fraction to the surface density of H<sub>2</sub>O molecules (10<sup>15</sup> molecules cm<sup>-2</sup>), leading to [Br<sup>-</sup>]<sub>s</sub> ≈ 2.9 × 10<sup>12</sup> molecules cm<sup>-2</sup> for the 8.5 mM solution of Oldridge and Abbatt. Since the surface to volume ratio was not varied in these experiments, only the product  $k_s^{\text{II}} \times N_{\text{max}}$  is constrained by the data for which a value of 0.03 s<sup>-1</sup> is obtained for the best fit. Assuming that  $N_{\text{max}}$  is about 10<sup>15</sup> molecules cm<sup>-2</sup>,  $k_s^{\text{II}}$  would be a few 10<sup>-17</sup> cm<sup>2</sup> molecule<sup>-1</sup>. In turn, the value for  $K_{\text{LangC}}$  is well constrained by the inverse O<sub>3</sub> concentration dependence, and the value found by Oldridge and Abbatt is adopted.

The contribution owing to uptake in the bulk reaction can be calculated by:

$$\Gamma_b = \frac{4HRT}{\bar{c}} \sqrt{D_1 k_b^{\text{II}}}$$

In the absence of independent measurements of the solubility of O<sub>3</sub> in halide solutions, we recommend using the expression as compiled by Chameides (1984). The diffusion coefficient given in the table for 273 K has been extrapolated from its room temperature value via the Stokes-Einstein relation. The parameterization reproduces the experimental data by Oldridge and Abbatt fairly well, even though it overestimates  $\gamma$  at high concentration, because the surface reaction doesn't fall off as quickly. Note that the reacto-diffusive length for deliquesced sea-salt is at least a few tens of  $\mu\text{m}$ , making the correction necessary for small spherical particles. However, for the atmospherically relevant concentration range (<10<sup>12</sup> molecules cm<sup>-3</sup>), the shape independent surface reaction term would dominate.

Br<sub>2</sub> has been observed as gas-phase product resulting from the interaction of O<sub>3</sub> with frozen and aqueous bromide surfaces (Oum et al., 1998; Hunt et al., 2004).

## References

- Abbatt, J. P. D. and Waschewsky, G. C. G.: J. Phys. Chem. A, 102, 3719, 1998.  
Chameides, W. L.: J. Geophys. Res., 89, 4739–4755, 1984.

32229

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Clifford, D. and Donaldson, D. J.: J. Phys. Chem. A, 111, 9809, 2007.
- Garland, J. A., Elzermann, A. W., and Penkett, S. A.: J. Geophys. Res., 85, 7488, 1980.
- Haag, W. R. and Hoigné, J.: Environ. Sci. Technol., 17, 261–267, doi:10.1021/es00111a004, 1983.
- 5 Hu, J., Shi, Q., Davidovits, P., Worsnop, D. R., Zahniser, M., and Kolb, C. E.: J. Phys. Chem., 95, 8768, 1995.
- Hunt, S. W., Roeselova, M., Wang, W., Wingen, L. M., Knipping, E. M., Tobias, D. J., Dabdub, D., and Finlayson-Pitts, B. J.: J. Phys. Chem. A, 108, 11559, 2004.
- Magi, L., Schweitzer, F., Pallares, C., Cherif, S., Mirabel, Ph., and George, C.: J. Phys. Chem. A, 101, 4943, 1997.
- 10 Oldridge, N. W. and Abbatt, J. P. D.: J. Phys. Chem. A, 115, 2590, 2011.
- Oum, K. W., Lakin, M. J., and Finlayson-Pitts, B. J.: Geophys. Res. Lett., 25, 3923–3926, 1998.
- Rouvière, A., Sosedova, Y., and Ammann, M.: J. Phys. Chem. A, 114, 7085, 2010.
- 15 Rouvière, A. and Ammann, M.: Atmos. Chem. Phys., 10, 11489, 2010.
- Schütze, M. and Herrmann, H.: Phys. Chem. Chem. Phys., 4, 60, 2002.
- Utter, R. G., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4973, 1992.

## VI.A2.1

### OH + aqueous sea salt aerosol → products

#### Experimental data

Parameter	RH/%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\approx 0.2$	82	297	Knipping et al. (2000)	(a)
$> 0.1$	70–80	298	Laskin et al. (2006)	(b)

#### Comments

(a) NaCl aerosol was produced using a nebulizer and dispersed in a chamber. Formation of  $\text{Cl}_2$  was observed with atmospheric pressure chemical ionization MS in presence of  $0.63$  to  $3.4 \times 10^{14} \text{ cm}^{-3}$  of  $\text{O}_3$  and  $254 \text{ nm}$  light in the chamber.  $\text{Cl}_2$  formation was simulated with a kinetic model that was extended in the follow up study by Knipping and Dabdub (2002). The uptake coefficient given in the table is the result of fitting the simulation to experimental data based on a parameterization of a surface reaction between OH and  $\text{Cl}^-$ . Parallel uptake of OH into the bulk of the particles (with  $\alpha_b = 0.1$ ) and aqueous phase chemistry therein was considered as well.

(b) NaCl aerosol was produced using a nebulizer followed by drying, resulting in  $0.9 \mu\text{m}$  mean diameter particles that were deposited on 300 mesh gold grids. The grids were exposed to OH from  $\text{O}_3$  photolysis in an ambient temperature, atmospheric pressure flow cell. The humidity was first raised to 85 % RH to safely induce deliquescence and then kept between 70 and 80 % for the experiments. OH concentrations were around a few  $10^9 \text{ cm}^{-3}$ . Offline particle analysis using

SEM/EDX was used to determine Cl loss as a function of time using Na as reference. Wet particle sizes were inferred from dry particle sizes. The uptake coefficient is considered a lower limit as gas phase diffusion limitations existed.

### Preferred values

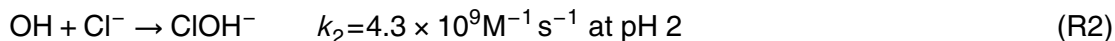
Parameter	Value	T/K
$\alpha_b$	>0.1	298
$\gamma_{gs}$	$0.04 \times ([Cl^-]/M)$	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	undetermined	
$\Delta \log(\gamma_{ER})$	$\pm 1$	298

### Comments on preferred values

The only direct experimental investigation of the reaction of OH with deliquesced NaCl provides a lower limit to the uptake coefficient (Laskin et al., 2006). The chamber experiment by Knipping et al. (2000) provided evidence for formation of  $Cl_2$  from the reaction of OH with deliquesced NaCl due to an overall reaction (R1).



In the bulk aqueous phase, the initial reaction of OH is believed to be (Finlayson-Pitts, 2003)



Known bulk aqueous phase kinetics was unable to explain  $Cl_2$  formation in the study by Knipping et al. (2000) and later by Knipping and Dabdub (2002). Therefore, an Eley-Rideal type surface reaction was proposed to bring experiment and simulation into agreement, which is recommended in parameterized form as a function of the

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





chloride concentration. Note that in this case the surface process acts parallel to bulk accommodation and bulk reaction:

$$\gamma = \gamma_{\text{gs}} + \left( \frac{1}{\alpha_{\text{b}}} + \frac{1}{\Gamma_{\text{b}}} \right)^{-1} \quad \text{with} \quad \Gamma_{\text{b}} = \frac{4HRT}{\bar{c}} \sqrt{D_1 k'_{\text{b}}}$$

where  $k_{\text{b}}^{\text{I}}$  is the first order rate coefficient in the bulk aqueous phase, e.g., due to Reaction (R2). However, in many cases multiple reactions may occur in parallel in the bulk. For the Henry's Law constant it is suggested to use the value recommended for pure water (see data sheet VI.A1.17, this evaluation).

Evidence for the formation of  $\text{OH}^-$  was provided by Laskin et al. (2003) and by Shaka et al. (2007), who demonstrated formation of  $\text{OH}^-$  in deliquesced  $\text{MgCl}_2$ . For  $\alpha_{\text{b}}$ , we recommend the same lower limit as for dilute aqueous solutions.

On dry sea salt, measured uptake coefficients are on the order of  $10^{-2}$  as reviewed by Rossi (2003) and Finlayson-Pitts (2003). Recently, the humidity dependence of the uptake coefficient of OH with sea salt was observed to be consistent with the humidity dependent uptake of OH on  $\text{MgCl}_2$  and  $\text{CaCl}_2$ , the low deliquescing components of sea salt (Park et al., 2008, 2009), consistent with a much higher OH loss at the surface of hydrated halides.

## References

- Finlayson-Pitts, B. J.: Chem. Rev., 103, 4801–4822, doi:10.1021/cr020653t, 2003.  
Knipping, E. M. and Dabdub, D.: J. Geophys. Res., 107, art. no. 4360, 2002.  
Knipping, E. M., Lakin, M. J., Foster, K. L., Jungwirth, P., Tobias, D. J., Gerber, R. B., Dabdub, D., and Finlayson-Pitts, B. J.: Science, 288, 301–306, 2000.  
Laskin, A., Gaspar, D. J., Wang, W., Hunt, S. W., Cowin, J. P., Colson, S. D., and Finlayson-Pitts, B. J.: Science, 301, 340–344, doi:10.1126/science.1085374, 2003.  
Laskin, A., Wang, H., Robertson, W. H., Cowin, J. P., Ezell, M. J., and Finlayson-Pitts, B. J.: J. Phys. Chem. A, 110, 10619–10627, 2006.

32233

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Park, J. H., Ivanov, A. V., and Molina, M. J.: J. Phys. Chem. A, 112, 6968–6977, 2008.
- Park, J.-H., Christov, C. I., Ivanov, A. V., and Molina, M. J.: Geophys. Res. Lett., 36, art. no. L02802, doi:10.1029/2008GL036160, 2009.
- Rossi, M. J.: Chem. Rev., 103, 4823–4882, 2003.
- 5 Shaka, H., Robertson, W. H., and Finlayson-Pitts, B. J.: Phys. Chem. Chem. Phys., 9, 1980–1990, 2007.

**HO<sub>2</sub> + Cl<sup>-</sup>/Br<sup>-</sup>/I<sup>-</sup> (aq) → products****Experimental data**

	Temp./K	Reference	Technique/Comments
Uptake coefficients: $\gamma$			
0.1 ± 0.03 (NaCl, RH 53–75 %)	296 ± 2	Taketani et al. (2008)	AFT (a)
0.1 ± 0.04 (SSS, RH 53–75 %)	296 ± 2	Taketani et al. (2009)	AFT (b)
0.1 ± 0.03 (NS, RH 53–75 %)			
0.07 ± 0.03 (KCl, RH 75 %)			
Accommodation coefficients: $\alpha_b$			
0.65 ± 0.17	296 ± 2	Taketani et al. (2008)	AFT (a)
0.55 ± 0.19	296 ± 2	Taketani et al. (2009)	AFT (b)

5 **Comments**

(a) Uptake of HO<sub>2</sub> ( $\sim 10^8$  molecule cm<sup>-3</sup>) to aqueous NaCl particles (mean surface area weighted diameter of 80–110 nm) at RH between 45 and 75 %. HO<sub>2</sub> was generated by the photolysis of H<sub>2</sub>O in air and detected as OH (by LIF) following conversion in reaction with NO. In experiments to determine  $\alpha_b$  the particles contained CuSO<sub>4</sub> ( $\sim 0.5$  M) to scavenge HO<sub>2</sub>. The authors originally reported:  $\gamma(\text{NaCl}) = 0.11 \pm 0.03$ ,  $0.09 \pm 0.02$  and  $0.10 \pm 0.02$  at RH = 53, 63 and 75 %, respectively.

(b) Same experimental set up as (a). SSS = synthetic sea-salt, NS = natural seawater. The authors originally reported:  $\gamma(\text{SSS}) = 0.07 \pm 0.03$ ,  $0.12 \pm 0.04$  and  $0.13 \pm 0.04$  at RH = 35, 50 and 75 %, respectively.  $\gamma(\text{NS}) = 0.10 \pm 0.03$ ,  $0.11 \pm 0.02$  and  $0.10 \pm 0.03$  at RH = 35, 50 and 75 %, respectively. In experiments

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



to determine  $\alpha_b$  aqueous KCl particles contained  $\text{CuSO}_4$  ( $\sim 0.5$  M) to scavenge  $\text{HO}_2$ .

## Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.5	290–300
$\gamma$	0.1	290–300
<i>Reliability</i>		
$\Delta \log \alpha$	undetermined	
$\Delta \log \gamma$	$\pm 1$	290–300

## 5 Comments on preferred values

Taketani et al. (2008) observed exponential  $\text{HO}_2$  decay to aqueous NaCl particles with  $\gamma$  equal to  $0.65 \pm 0.17$  when  $\text{CuSO}_4$  was present, defining the lower limit to  $\alpha_b$ . In the absence of  $\text{CuSO}_4$   $\gamma$  was independent of RH and close to 0.1. Taketani et al. (2009) used synthetic sea salt and natural sea salt to derive similar results. Our preferred values for uptake of  $\text{HO}_2$  to aqueous sea-salt particles are based on this data-set. As the mechanism for  $\text{HO}_2$  loss in the particle remains unknown (see below) we add substantial uncertainty to the preferred value of  $\gamma$ .

The uptake of  $\text{HO}_2$  in aqueous solution with  $\text{pH} > 5$ , is presently believed to be driven by self-reaction and acid-base dissociation of  $\text{HO}_2$  ( $\text{pK}_a \sim 4.7$ ) with formation of  $\text{H}_2\text{O}_2$  (Reactions R2, R3). In the presense of transition metal ions (TMI) the reaction of  $\text{HO}_2$  and especially  $\text{O}_2^-$  (Reaction R4) can be important:



32236

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

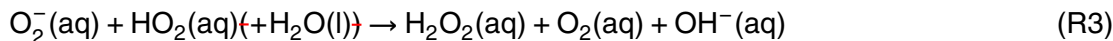
Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





If a first-order loss process for  $\text{HO}_2$  or  $\text{O}_2^-$  in the aqueous phase dominates (e.g. reaction with TMI such as  $\text{Cu}(\text{II})$ ), and assuming equal rates of reaction throughout the particle, the uptake coefficient can be calculated from the expression below:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{\bar{c}}{4H^{\text{eff}}RT \sqrt{k_{\text{TMI}}[\text{TMI}]D_1}}$$

$H^{\text{eff}} = H^{\text{HO}_2} (1 + K_{\text{eq}}/[\text{H}^+])$ ,  $K_{\text{eq}} = 2.1 \times 10^{-5} \text{ M}$  at 298 K (Jacob, 2000),  $H^{\text{HO}_2} = 9.5 \times 10^{-6} \exp(5910/T)$  (Hanson et al., 1992) and  $D_1 = (1 \times 10^{-5}(T/298))/(1.09 \times 10^8 \exp(-0.068T) + 0.873) \text{ cm}^2 \text{ s}^{-1}$  (Schwartz, 1984; Thornton et al., 2008) where the denominator in the  $D_1$  term was derived from a fit to the water viscosity data of Hallett (1963).

According to the reaction scheme above, the rates of loss of aqueous-phase  $\text{HO}_2$  are quadratically dependent on  $[\text{HO}_2]_{\text{aq}}$  and  $[\text{O}_2^-]_{\text{aq}}$  and are thus strongly dependent on the gas-phase concentration of  $\text{HO}_2$  in the absence of TMI. At low  $\text{HO}_2$  concentrations (e.g. as found in the atmosphere) the liquid phase reactions become rate limiting and  $\gamma$  is expected to be much smaller as observed in dilute solutions by Mozurkewich et al. (1987) and the simple formalism above breaks down. Thornton and Abbatt (2005) suggest that the rate of loss of  $\text{HO}_2$  from the gas-phase (in molecule  $\text{cm}^{-3} \text{ s}^{-1}$ ) is best described by a system in thermodynamic (Henry's law) equilibrium so that (Thornton et al., 2008):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{3\bar{c}N_A}{8000 (H^{\text{eff}}RT)^2 k_{\text{aq}}[\text{HO}_2]r}$$

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


$k_{\text{aq}}$  can be calculated from the rate coefficients for Reactions (R2) ( $k_2$ ) and (R3) ( $k_3$ ) (Bielski et al., 1985) and the pH:

$$k_{\text{aq}} = \frac{k_2 + \left( \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}} \right) k_3}{\left( 1 + \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}} \right)^2}$$

This formalism predicts that the loss of  $\text{HO}_2$  to particles is favoured by high  $\text{HO}_2$  mixing ratios, low temperatures (higher solubility) and low pH. At low concentrations of  $\text{HO}_2$  (where the self reaction and reaction with  $\text{O}_2^-$  are slow), values of  $\gamma$  of  $<0.005$  are calculated, which are however much less than the uptake coefficients of Taketani et al. (2008, 2009) who investigated the uptake of low concentrations of  $\text{HO}_2$  to aqueous salt particles. In addition, as discussed by Hanson et al. (1992) and Thornton and Abbatt (2005), there is considerable uncertainty (factor of 2.5) associated with the solubility of  $\text{HO}_2$  ( $H^{\text{HO}_2}$ ) and its temperature dependence. Until these apparent discrepancies and uncertainties have been resolved by further experiments, we make no recommendation for parameterising  $\text{HO}_2$  uptake to halide containing aqueous aerosol. The hypothesis of a fast disproportionation reaction  $2\text{HO}_2 \rightarrow \text{H}_2\text{O}_2 + \text{O}_2$  in aqueous solution advanced by Thornton and Abbatt (2005) is inconsistent with the measured value of  $\gamma$  for  $\text{HO}_2$  loss on both  $\text{NaCl}$  and  $(\text{NH}_4)_2\text{SO}_4$  aerosol when extrapolated to the  $\text{HO}_2$  concentrations lower by a factor of 500 used in the work of Taketani et al. (2008). We suggest that a value close to 0.1 is appropriate for marine environments where sea salt aerosols will be deliquescent. We refer to recent publications for a more detailed description of the effect of different parameterisation schemes (Thornton et al., 2008; Macintyre and Evans, 2011).

Experiments conducted using dry salt surfaces at room temperature reveal a much lower uptake coefficient than those reported for aqueous particles (Taketani et al., 2008, 2009) with values ranging from  $1.8\text{--}2.3 \times 10^{-3}$  (Loukhovitskaya et al., 2009, on  $\text{NaCl}$ ,  $\text{NaBr}$  and SSS bulk surfaces),  $7.5 \times 10^{-3}$  (Antsupov, 1988),  $12\text{--}13 \times 10^{-3}$  (Remorov, 2002; Gershenson, 1995, on  $\text{NaCl}$ ) and  $18 \times 10^{-3}$  (Gershenson, 1995, on  $\text{KCl}$ ).

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Although disagreeing considerably on absolute values of the uptake coefficient (perhaps reflecting different modes of sample presentation and/or possible sample contamination) they all return strong, negative dependencies of  $\gamma$  on the temperature and show that at most temperatures  $\gamma$  is independent of the  $\text{HO}_2$  concentration, the exception being at the highest temperatures studied by Remorov (2002), where they observed that  $\gamma$  decreased at lower  $[\text{HO}_2]$ . Remorov et al. (2002) (bulk NaCl) and Loukhovitskaya et al. (2009) (bulk  $\text{MgCl}_2 \cdot 6\text{H}_2\text{O}$ ) both observed a decrease in  $\text{HO}_2$  uptake when the RH was increased. This is in stark contrast to the observation of Taketani et al. (2008, 2009) who observed that  $\gamma$  increased when the RH was increased. In combination with the observation of  $\text{H}_2\text{O}_2$  as sole product (at 0.5 yield, Loukhovitskaya et al., 2009), these observations indicate that, on solid particles, the uptake is driven by reversible surface accommodation of  $\text{HO}_2$  and subsequent self reaction independent of the presence of halide ions.

**References**

- Antsupov, E.: Sov. J. Chem. Phys., 7, 1082, 1988.
- Bielski, B. H. J., Cabelli, D. E., Arudi, R. L., and Ross, A. B.: J. Phys. Chem. Ref. Data, 14, 1041–1100, 1985.
- Gershenson, Y. M., Grigorieva, V. M., Ivanov, A. V., and Remorov, R. G.: Faraday Discuss., 100, 83–100, 1995.
- Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.
- Jacob, D. J.: Atmos. Env., 34, 2131–2159, 2000.
- Loukhovitskaya, E., Bedjanian, Y., Morozov, I., and Le Bras, G.: Phys. Chem. Chem. Phys., 11, 7896–7905, 2009.
- Macintyre, H. L. and Evans, M. J., Atmos. Chem. Phys., 10, 7409–7414, 2010, <http://www.atmos-chem-phys.net/10/7409/2010/>.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Morita, A., Kanaya, Y., and Francisco, J. S.: J. Geophys. Res.-Atmos., 109, art. no. D09201, doi:10.1029/2003jd004240, 2004.
- Mozurkewich, M., McMurry, P. H., Gupta, A., and Calvert, J. G.: J. Geophys. Res.-Atmos., 92, 4163–4170, 1987.
- 5 Remorov, R. G., Gershenson, Y. M., Molina, L. T., and Molina, M. J.: J. Phys. Chem. A, 106, 4558–4565, 2002.
- Schwartz, S. E.: J. Geophys. Res.-Atmos., 89, 1589–1598, 1984.
- Taketani, F., Kanaya, Y., and Akimoto, H.: Atmospheric Environment, 43, 1660–1665, doi:10.1016/j.atmosenv.2008.12.010, 2009.
- 10 Taketani, F., Kanaya, Y., and Akimoto, H.: J. Phys. Chem. A, 112, 2370–2377, 2008.
- Thornton, J. and Abbatt, J. P. D.: J. Geophys. Res.-Atmos., 110, D08309, doi:10.1029/2004JD005402, 2005.



NO<sub>2</sub> (g) + aqueous sea salt aerosol → products

## Experimental data

Parameter	RH/%	T/K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma < 10^{-4}$	50–80	292	Behnke et al. (1996)	(a)
$\gamma < 10^{-4}$	75	298	Abbatt and Waschewsky (1998)	AFT-CIMS (b)
$(3.7 \pm 0.5) \times 10^{-4}$	90	279	Harrison and Collins (1998)	AFT-CLD (c)
$(2.8 \pm 0.5) \times 10^{-4}$	85	279		
$\approx 10^{-4}$ (1 mM NaCl)		298	Yabushita et al. (2009)	(d)
<i>Solubility, H (Matm<sup>-1</sup>)</i>				
$(5.5 \pm 0.6) \times 10^{-2}$		284	Cape et al. (1993)	(e)
<i>Liquid phase rate constants</i>				
$k_{R1}^H = (1.8 \pm 0.7) \times 10^6 \text{ M}^{-1} \text{ s}^{-1}$ (30 g l <sup>-1</sup> NaCl)		284	Cape et al. (1993)	(e)
$k_{R2}^H = 1.4 \times 10^{-6} \text{ M s}^{-1}$	80	293	Karlsson and Ljungström (1995)	AFT-IC (f)
$k_{R2}^H = 1.7 \times 10^{-6} \text{ M s}^{-1}$		283		
$k_{R2}^I = 4.4 \times 10^{-6} \text{ M s}^{-1}$		273		

## 5 Comments

(a) Deliquesced NaCl aerosol (70  $\mu\text{g m}^{-3}$ ) was exposed to 58 ppb of NO<sub>2</sub> in a chamber, leading to about 10 % of NO<sub>2</sub> loss after about three hours in the dark, which was used to estimate the uptake coefficient listed in the table, based on the size distribution measured by SMPS. Switching on simulated sunlight did not result in Cl atoms expected from photolysis NOCl, but rather in OH radicals presumably from photolysis of HONO formed via hydrolysis of NOCl.

(b) Uptake experiment on liquid NaCl in atmospheric pressure aerosol flow tube using CIMS detection. The NaCl aerosol produced with a nebulizer had a bimodal distribution (large average diameter: 2–4  $\mu\text{m}$ , small: <1  $\mu\text{m}$ ) with typical

32241

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



surface areas and particle number densities in the range of  $1\text{--}6 \times 10^{-3} \text{ cm}^2/\text{cm}^3$  and  $1 \text{ to } 4 \times 10^4 \text{ particles}/\text{cm}^3$ , respectively. The solution nebulized was either unbuffered NaCl solution, or solution adjusted to  $\text{pH} = 0.3$  with HCl, or to  $\text{pH} = 7.2$  with  $\text{NaH}_2\text{PO}_4/\text{Na}_2\text{HPO}_4$ . No  $\text{NO}_2$  loss could be detected between  $3 \times 10^{12}$  and  $2 \times 10^{14}$  molecules/ $\text{cm}^3$ . The value of  $\gamma$  is an upper limit based on the detection limit.

- (c) Uptake experiment on dry and deliquesced NaCl aerosol in atmospheric pressure flow tube at 2 ppm of  $\text{NO}_2$ .  $\text{NO}_2$  loss was monitored using a chemiluminescence detector. A nylon filter and a filter impregnated with sodium carbonate was used to prevent  $\text{HNO}_3$  and HONO, respectively, from being detected by the CLD.  $\text{NO}_2$  decay rates were determined as a function of the aerosol surface to volume ratio determined by SMPS, from which the uptake coefficient was derived. The values for the conditions of deliquesced NaCl aerosol are listed in the table. Surprisingly, the uptake coefficient for dry NaCl particles at 50% relative humidity was even higher,  $10^{-3}$ .
- (d) Microdroplets were generated by nebulization of aqueous halogenide solutions into ppmv  $\text{NO}_2(\text{g})/\text{N}_2$  ( $1\text{--}5 \times 10^{14} \text{ cm}^{-3}$ ) mixtures at atmospheric pressure, leading to interaction times of about 1 ms. Droplets were then expanded into a mass spectrometer. The rate of appearance of nitrate, from which the uptake coefficient was estimated, had a maximum in the  $10^{-3} \text{ M}$  range for chloride, bromide and iodide, which is listed in the table. The uptake coefficient for these solutions was 4 orders of magnitude larger than that on neat water.
- (e)  $\text{NO}_2$  at partial pressures of 10 to 100 ppb was passed through 1.0l of aqueous solution in a fritted bubbler in a thermostatted vessel. The loss of  $\text{NO}_2$  was measured with a chemiluminescence detector. Apart from pure water, the aqueous solutions used were  $30 \text{ g l}^{-1}$  NaCl, two synthetic sea salt mixtures at comparable concentration, and authentic coastal sea-water. The solubility derived from the

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

kinetic experiments was independent of the type of solution. However, the rate constants for the solutions were higher, up to a factor of 10 for sea-water. The authors caution that the distribution of  $\text{NO}_2$  in the reactor may have not been uniform, leading to an overestimate of the rate constants.

- (f) The flow reactor consisted of six 10 l halocarbon wax coated glass tubes connected in series. Sea salt aerosol was produced by nebulizing synthetic sea water and bringing to charge equilibrium with a  $^{85}\text{Kr}$  source. The particle size distribution was measured with an electrostatic classifier coupled to an aerosol electrometer. The geometric mean number diameter was between 100 and 120 nm. Nitrogen dioxide concentrations ranged from 50 to 500 ppbv and were monitored with a chemiluminescence detector. Particles were collected on a filter after the flow reactor and after a  $\text{KMnO}_4$  coated denuder removing gas phase  $\text{NO}_2$ . Experiments at humidities below the deliquescence of the particles indicated a lower limit for the uptake coefficient of  $3 \times 10^{-4}$  with negative humidity dependence. The first order loss rate constants for chloride at 80 % relative humidity listed in the table were derived from experiments at 190 ppb  $\text{NO}_2$ . The authors ascribed this loss entirely to the reaction  $2 \text{NO}_2 + \text{Cl}^- \rightarrow \text{NO}_3^- + \text{NOCl}$ .

### Preferred values

Parameter	Value	T/K
$k_{\text{R2}}^{\text{II}}/\text{M}^{-1} \text{s}^{-1}$	200	298
<i>Reliability</i>		
$\Delta \log(k)$	$\pm 0.7$	

## Comments on preferred values

While a significant number of studies have investigated the reaction of NO<sub>2</sub> with solid NaCl and other sea salt proxies (Rossi, 2003) and references therein) and also the significant impact of humidity on these reactions, only a few studies have addressed the uptake to deliquesced sea salt particles representative of marine boundary layer conditions. The only aerosol flow tube study that covers relevant low NO<sub>2</sub> concentrations is that by Abbatt and Waschewsky (1998), which reports an upper limit for  $\gamma$  consistent with the results from the chamber experiment by Behnke et al. (1996). The high uptake coefficients obtained by Harrison and Collins (1998) are likely the result of an experimental artifact, e.g., due to the fact that particles were deposited on a filter before analysis of the gas phase, leading to additional residence time during which they were effectively exposed to NO<sub>2</sub>. These high uptake coefficients are not used to recommend a value for the bulk accommodation coefficient,  $\alpha_b$ .

Apart from the hydrolysis reaction (R1), a possible fate of NO<sub>2</sub> at low concentration is Reaction (R2):



Karlsson and Ljungström (1995) determined removal rates of chloride in an aerosol flow tube experiment of  $1.7 \times 10^{-6} \text{ M s}^{-1}$  at 190 ppb NO<sub>2</sub> at 293 K, which increased with decreasing temperature. Taking the Henry's Law constant for pure water of  $1.4 \times 10^{-2} \text{ M atm}^{-1}$  (Cheung et al., 2000) and a chloride concentration of 4.4 M, this would lead to a liquid phase rate constant for Reaction (R2) of  $200 \text{ M}^{-1} \text{ s}^{-1}$ , which is the basis for the recommendation. Karlsson and Ljungström observed NOCl in a separate bubbler experiment with higher NO<sub>2</sub> concentration. Behnke et al. (1996) hypothesised that NOCl hydrolysed to form HONO in their chamber experiment.

The NO<sub>2</sub> loss rates measured by Cape et al. (1993) in a 0.5 M NaCl solution at 283 K between 10 and 100 ppb of NO<sub>2</sub> were not different from those in pure water and

32244

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



ascribed entirely to the hydrolysis reaction. Cape et al. observed  $\text{NO}_2$  removal rates enhanced by up to a factor of 10 in authentic sea water, which they ascribed either to catalysis of the hydrolysis reaction or to an unknown reaction with another species. The magnitude of the hydrolysis rate constant reported by Cape et al. was likely affected by non-uniform distribution of reactants in the reactor. We therefore do not provide a recommendation for Reaction (R1) here, but suggest using the rate constant recommended for pure water on data sheet VI.A1.3,  $3.0 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$ , also for deliquesced sea salt aerosol.

Bambauer et al. (1994) determined the rate of nitrate formation in cloud chamber experiments, in which the cloud droplets contained mM amounts of NaCl. The reaction order was about 1, in disagreement with Cape et al. (1993). An uptake coefficient was not extracted, nor was the possibility explored that Reaction (R2) could have interfered with the hydrolysis of  $\text{NO}_2$ . For droplets of similar composition, Yabushita et al. (2009) observed a very large uptake coefficient and suggest that  $\text{NO}_2$  hydrolysis is enhanced in presence of millimolar concentrations of halogenide ions through formation of a charge transfer complex  $\text{NO}_2\text{-X}^-$  at the solution – air interface, which could account for the discrepancy between the Bambauer et al. and Cape et al. studies that were otherwise discussed in terms of different mixing regimes. Given that the  $\text{NO}_2$  concentrations in the experiment by Yabushita et al. were three orders of magnitude above atmospheric, the extent of such a surface process for deliquesced sea salt particles remains uncertain, and no recommendation is given.

The value reported for the solubility by Cape et al. (1993) is affected by the same issues as the rate constants mentioned above, as they are interconnected through  $Hk_2^{1/2}$ , which is the parameter directly retrieved from their experiments. As Cape et al. did not see a difference in solubility between pure water and chloride solution, we recommend using the solubility for pure water reported by Cheung et al. (2000).

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

- Abbatt, J. P. D. and Waschewsky, G. C. G.: J. Phys. Chem. A, 102, 3719–3725, 1998.
- Bambauer, A., Brantner, B., Paige, M., and Novakov, T.: Atmos. Environ., 28, 3225–3232, 1994.
- 5 Behnke, W., Elend, M., George, C., Kruger, H.-U., Scheer, V., and Zetzsch, C.: in: Heterogeneous and Liquid-Phase Processes, edited by: Warneck, P., Springer-Verlag, Heidelberg, 153–162, 1996.
- Cape, J. N., Storetonwest, R. L., Devine, S. F., Beatty, R. N., and Murdoch, A.: Atmos. Environ., 27a, 2613–2621, 1993.
- 10 Cheung, J. L., Li, Y. Q., Boniface, J., Shi, Q., Davidovits, P., Worsnop, D. R., Jayne, J. T., and Kolb, C. E.: J. Phys. Chem. A, 104, 2655–2662, 2000.
- Harrison, R. M. and Collins, G. M.: J. Atmos. Chem., 30, 397–406, 1998.
- Karlsson, R. and Ljungström, E.: J. Aerosol Sci., 26, 39–50, 1995.
- Rossi, M. J.: Chem. Rev., 103, 4823–4882, 2003.
- 15 Yabushita, A., Enami, S., Sakamoto, Y., Kawasaki, M., Hoffmann, M. R., and Colussi, A. J.: J. Phys. Chem. A, 113, 4844–4848, 2009.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

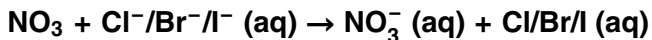
Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Experimental data

Parameter	Aqueous solution	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients, <math>\gamma</math></i>				
$0.9 - 3.5 \times 10^{-3}$	$\text{I}^- (5 \times 10^{-6} - 0.4 \text{ M})$	$273 \pm 1$	Rudich et al. (1996a)	WWFT-AS (a)
$0.8 - 6.0 \times 10^{-3}$	$\text{Cl}^- (0.01 - 0.5 \text{ M})$	$273 \pm 1$	Rudich et al. (1996b)	WWFT-AS (b)
$1.6 - 5.5 \times 10^{-3}$	$\text{Br}^- (0.1 - 0.01 \text{ M})$			
$1.1 - 2.0 \times 10^{-3}$	$\text{Cl}^- (0.125 - 1 \text{ M})$	$293 \pm 1$	Schütze et al. (2005)	(c)
$> 2 \times 10^{-3}$	$\text{Cl}^- (0.1 \text{ M})$	293	Thomas et al. (1998)	(d)
<i>Accommodation coefficient, <math>\alpha_b</math></i>				
$> 4 \times 10^{-2}$	$\text{I}^- (5 \times 10^{-6} - 0.4 \text{ M})$	$273 \pm 1$	Rudich et al. (1996a)	(a)
$> 2 \times 10^{-3}$	$\text{Cl}^- (0.1 \text{ M})$	293	Thomas et al. (1998)	(d)
$4.2_{-1.7}^{+2.2} \times 10^{-3}$	$\text{Cl}^- (0.125 - 1 \text{ M})$	$293 \pm 1$	Schütze et al. (2005)	(c)

## 5 Comments

(a) Flow tube operated at 12–23 mbar.  $\text{NO}_3$  ( $2 - 10 \times 10^{11}$  molecule  $\text{cm}^{-3}$ ) was formed by the thermal dissociation of  $\text{N}_2\text{O}_5$  and detected by diode laser absorption at 662 nm over a 12.6 m pathlength. The salt solution had a pH of 5–6.5. The dependence of the experimental  $\text{NO}_3$  uptake coefficient on the  $\text{I}^-$  concentration was used to derive a value of  $H^2 D_1 k$  of  $(1.75 \pm 0.5) \times 10^4 \text{ Matm}^{-2} \text{ cm}^2 \text{ s}^{-2}$ . Using a previously determined value for  $HD_1^{0.5}$  (Rudich et al., 1996b, this datasheet) this leads to an aqueous-phase rate constant for  $\text{NO}_3 + \text{I}^-$  of  $(4.6 \pm 0.5) \times 10^9 \text{ M}^{-1} \text{ s}^{-1}$  at 273 K. A lower limit for  $\alpha_b$  was estimated from the maximum values of the uptake coefficients.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

- (b) Same set-up as in comment (a). The dependence of  $\gamma$  on salt concentrations was combined with a literature value (Exner et al., 1992) for the rate coefficient for  $\text{Cl}^-$  with  $\text{NO}_3$  ( $2.76 \times 10^6 \text{ M}^{-1} \text{ s}^{-1}$  at 273 K) to determine (for chloride) a value of  $HD_1^{0.5}$  of  $(1.9 \pm 0.4) \times 10^{-3} \text{ M atm}^{-1} \text{ cm s}^{-0.5}$ . Assuming a value of  $D_1$  enabled a solubility of  $\text{NO}_3$  in  $\text{H}_2\text{O}$  of  $(0.6 \pm 0.3) \text{ M}^{-1} \text{ atm}^{-1}$  to be derived.
- (c) Single droplet ( $\approx 7 \text{ mm}^3$ ) suspended from a pipette in a flow tube (10 mbar He) with UV-Vis absorption spectroscopy for concentration measurement in both gas and aqueous phases ( $\text{NO}_3^-$  at 235 nm,  $\text{NO}_3$  using the 662 nm feature).  $\text{NO}_3$  was generated by reacting  $\text{NO}_2$  with  $\text{O}_3$  at 393 K to keep the  $\text{N}_2\text{O}_5$  level low ( $\approx 5\%$ ).  $\text{HNO}_3$  was formed at  $\approx$  the same concentration as  $\text{NO}_3$ . Uptake of  $\text{NO}_3$  to the droplet was monitored by nitrate anion absorption. A value of  $HD_1^{0.5} k^{0.5} = (1.9 \pm 0.2) \text{ M atm}^{-1} \text{ cm s}^{-0.5}$  was derived.
- (d) Uptake of  $\text{NO}_3$  (generated by mixing  $\text{NO}$  with  $\text{O}_3$  at  $\approx 400^\circ\text{C}$ ) to 0.1 M chloride solution was monitored using three serial coiled glass denuders (2 mm id). Nitrate in solution was determined following reduction to  $\text{NO}_2^-$  and the Saltzman reaction. Large diffusion limitations result in only a lower limit to  $\gamma$ .

### Preferred values

Parameter	Value	T/K
$\alpha_b$	$1.3 \times 10^{-2}$	273
$k_{\text{Cl}} (\text{M}^{-1} \text{s}^{-1})$	$2.76 \times 10^6$	273
$k_{\text{Br}} (\text{M}^{-1} \text{s}^{-1})$	$1.02 \times 10^8$	273
$k_1 (\text{M}^{-1} \text{s}^{-1})$	$4.6 \times 10^9$	273
<i>Reliability</i>		
$\Delta \log(\gamma)$	$\pm 0.5$	273



## Comments on preferred values

The detailed studies (Rudich et al., 1996b; Schütze et al., 2005) on the uptake of  $\text{NO}_3$  to aqueous chloride solutions of various ionic strength use very different methods and return values of the common observable ( $HD_1^{0.5}k^{0.5}$ ) which differ by  $\approx 70\%$ , which can be considered to be in reasonable agreement. Some of the discrepancy can be attributed to different  $\text{NO}_3$  solubilities ( $H$ ) at 273 and 291 K. The derivation of the accommodation coefficient is not exact in either Rudich et al. (1996a) nor in Schütze et al. (2005) and the results differ by at least a factor of 6. The need to correct data due to aqueous nitrate formed from large impurity levels of  $\text{HNO}_3$  in the work of Schütze et al. (2005) leads us to prefer the results of Rudich et al. (1996). The accommodation coefficient chosen is slightly lower than the lower limit given by Rudich et al., as this better represents their dataset within the parameterisation listed below.

The rate coefficient for the aqueous phase reaction of  $\text{NO}_3$  with  $\text{Cl}^-$  was taken from Exner et al. (1992), though there is some disagreement concerning this number (Herrmann, 2003). We adopt this rate coefficient and the solubility derived by Rudich et al. (1996b) as it allows us to present an internally consistent set of parameters based on observations of the composite  $HD_1^{0.5}k^{0.5}$  term. The rate coefficient listed for the reaction of  $\text{NO}_3$  with  $\text{Br}^-$  was taken from the ratio  $k_{\text{Cl}}/k_{\text{Br}} = 37 \pm 4$  as reported by Rudich et al. (1996b). The reliability of the parameters listed above cannot be given individually, but an assessment of the overall uncertainty in  $\gamma$  (calculated as below) is given.

Parameterisation of the net uptake coefficient for interaction of  $\text{NO}_3$  with halide aerosol requires components for the hydrolysis term and the reaction with the halide ions (see expression below). There is large uncertainty association with the hydrolysis rate constant (VI.A1.04), though we note that for sea-salt aerosol (5M NaCl), neither bromide, iodide nor hydrolysis contribute significantly to  $\text{NO}_3$  loss when compare to chloride. The net uptake coefficient can be calculated from:

$$\gamma = \left\{ \frac{1}{\alpha} + \frac{c}{4HRT(D_1k')^{0.5}} \right\}^{-1}$$
 where  $k' = k_{\text{Cl}} [\text{Cl}^-] + k_{\text{Br}} [\text{Br}^-] + k_{\text{I}} [\text{I}^-] + k'_{\text{H}_2\text{O}}$  and  $[X^-]$  are the activities of the respective anions.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



We recommend use of  $D_1 = 1 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$ ,  $H = 0.6 \text{ M atm}^{-1}$  and  $k_{\text{H}_2\text{O}} = 23$  (see VI.A1.04).

There are some studies of the uptake of  $\text{NO}_3$  to solid salt surfaces (Seisel et al., 1997, 1999; Gratpanche et al., 1999; Gerschenzon et al., 1999). All suggest larger uptake coefficients to bromide surfaces compared to chloride, in line with the larger uptake coefficient listed above. There is also some indirect evidence of halogen atom release to the gas-phase (Seisel et al., 1997).

## References

Chameides, W. L.: J. Geophys. Res., 91, 14571–14572, 1986.

Exner, M., Herrmann, H., and Zellner, R.: Ber. Bunsen. Ges. Phys. Chem. Chem. Phys., 96, 470–477, 1992.

Gershenzon, M. Y., Il'in, S., Fedotov, N. G., and Gershenzon, Y. M.: J. Atmos. Chem., 34, 119–135, 1999.

Gratpanche, F. and Sawerysyn, J. P.: J. Chim. Phys. Phys.-Chim. Biol., 96, 213–231, 1999.

Herrmann, H.: Chem. Rev., 103, 4691–4716, 2003.

Rudich, Y., Talukdar, R. K., Imamura, T., Fox, R. W., and Ravishankara, A. R.: Chem. Phys. Lett., 261, 467–473, 1996a.

Rudich, Y., Talukdar, R. K., Ravishankara, A. R., and Fox, R. W.: J. Geophys. Res., 101, 21023–21031, 1996b.

Schütze, M. and Herrmann, H.: J. Atmos. Chem., 52, 1–18, 2005.

Seisel, S., Caloz, F., Fenter, F. F., van den Bergh, H., and Rossi, M. J.: Geophys. Res. Lett., 24, 2757–2760, 1997.

Seisel, S., Fluckiger, B., Caloz, F., and Rossi, M. J.: Phys. Chem. Chem. Phys., 1, 2257–2266, 1999.

Thomas, K., Volz-Thomas, A., Mihelcic, D., Smit, H. G. J., and Kley, D.: J. Atmos. Chem., 29, 17–43, 1998.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.6

### $\text{N}_2\text{O}_5 + \text{Cl}^-/\text{Br}^- (\text{aq}) \rightarrow \text{products}$

#### Experimental data

Parameter	Aqueous solution	Temp./K	Reference	Technique/ Comments
<i>Uptake coefficients, <math>\gamma</math></i>				
$(3.9 \pm 0.13) \times 10^{-2}$	NaCl (1 M)	263	George et al. (1994)	DT-HPLC (a)
$(3.7 \pm 1.2) \times 10^{-2}$		268		
$(1.4 \pm 0.3) \times 10^{-2}$		273		
$(1.4 \pm 0.08) \times 10^{-2}$		278		
$(3.2 \pm 0.2) \times 10^{-2}$	NaCl (1.7–5.1 M)	291	Behnke et al. (1997)	Aerosol chamber (b)
$(1.8 \pm 0.3) \times 10^{-2}$	NaCl (1 M) NaBr (0.1–1 M) NaI (0.1–1 M)	262–278	Schweitzer et al. (1998)	DT-MS (c)
$0.64 \times 10^{-2}$	NaCl (RH = 30 %)	295 ± 2	Stewart et al. (2004)	AFT-CIMS (d)
$0.9 \times 10^{-2}$	NaCl (RH = 50 %)			
$1.04 \times 10^{-2}$	NaCl (RH = 70 %)			
$0.78 \times 10^{-3}$	NaCl (RH = 80 %)			
$1.6 \times 10^{-2}$	natural sea-salt (RH = 30 %)			
$2.8 \times 10^{-2}$	natural sea-salt (RH = 50 %)			
$1.3 \times 10^{-2}$	natural sea-salt (RH = 70 %)			
$3.1 \times 10^{-2}$	natural sea-salt (RH = 80 %)			
$(2.2 \pm 0.4) \times 10^{-2}$	synthetic sea-salt (RH = 50 %)	295	Thornton and Abbatt (2005)	AFT-CIMS (e)
$(3.0 \pm 0.8) \times 10^{-2}$	synthetic sea-salt (RH = 65 %)			
$(2.4 \pm 0.5) \times 10^{-2}$	synthetic sea-salt (RH = 70 %)			

#### 5 Comments

(a)  $\text{N}_2\text{O}_5$  made in-situ by reacting NO with  $\text{O}_3$ . Droplet train flow tube operated at 27–80 mbar He with 80–150  $\mu\text{m}$  droplets. Trace gas concentration measured by FTIR at entrance to flow tube, nitrate content of droplets analysed by HPLC to derive uptake coefficients. This required knowledge of the relative efficiency of nitrate and nitryl-chloride products.

10

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(b) Teflon aerosol smog chamber at 1 atm. pressure. Initial concentration of  $\text{N}_2\text{O}_5$  determined by FTIR. Subsequent to reaction of  $\text{N}_2\text{O}_5$  with aerosol,  $\text{NO}_3$  (in equilibrium with gas-phase  $\text{N}_2\text{O}_5$ ) was photolysed and  $\text{NO}_x$  was analysed to indirectly derive total un-reacted  $\text{N}_2\text{O}_5$ .  $\text{N}_2\text{O}_5$  taken up to the aerosol was calculated from the difference in initial and final concentrations. The amount of  $\text{ClNO}_2$  was determined by its photolytic conversion to Cl atoms (determined by hydrocarbon consumption). Aerosol number and size distribution (average diameter  $\approx 150$  nm) were obtained using a DMA-CPC. Values for the uptake coefficient supersede those reported in short communications by Behnke et al. (1991, 1992, 1993).

(c)  $\text{N}_2\text{O}_5$  made in-situ by reacting NO with  $\text{O}_3$ . 80–150  $\mu\text{m}$  droplets. Gas analysed by ion-trap MS and FTIR (for  $\text{ClNO}_2$  formation).

(d) Uptake of  $\text{N}_2\text{O}_5$  (100–700 ppbv at atmospheric pressure) to particles of aqueous aerosol of pure NaCl or natural sea-salt with diameters of  $\sim 60$ –250 nm at RH between 30 and 80%.  $\text{N}_2\text{O}_5$  was detected as the change in NO signal (monitored with a CLD) following thermal dissociation to  $\text{NO}_3$  and titration with NO. The measured uptake coefficients were strongly influenced by droplet size, indicating volume limited uptake. Once corrected for diffusive-reactive effects, the uptake coefficients (listed in the table) were independent of RH.

(e) Uptake of  $\text{N}_2\text{O}_5$  (8–30 ppbv at atmospheric pressure) to particles of aqueous aerosol of synthetic sea-salt (containing  $\text{Na}^+$ ,  $\text{K}^+$ ,  $\text{Mg}^{2+}$ ,  $\text{Ca}^{2+}$ ,  $\text{Cl}^-$  and  $\text{Br}^-$ ) with surface area weighted particle radii of 90 to 150 nm.  $\text{N}_2\text{O}_5$  was detected as  $\text{NO}_3^-$  using  $\text{I}^-$  primary ions. No significant influence of particle size (varied e.g. for RH = 50% from 85 to 134 nm) on  $\gamma$  was observed. ~~The presence of (an estimated) monolayer of hexanoic acid reduced the uptake coefficient (at RH = 70%) to  $(8 \pm 4) \times 10^{-3}$  but had no effect at RH = 50%.~~

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Preferred values

Parameter	Value	T/K
$\gamma$	0.02	260–300
$k_3/k_2$	450	298
<i>Reliability</i>		
$\Delta \log(\gamma)$	$\pm 0.3$	260–300
$\Delta \log(k_3/k_2)$	$\pm 0.3$	298

## Comments on preferred values

The preferred values of  $\gamma$  are independent of RH and temperature and refer to the uptake of  $\text{N}_2\text{O}_5$  to pure water-halide solutions. The presence of nitrate and/or organic components can reduce  $\gamma$  (datasheets VI.A3.7 and VI.A3.8). Indeed, Thornton and Abbatt (2005) argue that the particle size dependence of  $\gamma$  observed by Stewart et al. (2004) was not entirely due to reacto-diffusive length considerations but also to use of high gas-phase  $\text{N}_2\text{O}_5$  resulting in high particle nitrate content, which suppresses the uptake of  $\text{N}_2\text{O}_5$ . Both Stewart et al. (2004) and Thornton and Abbatt (2005) showed that the uptake coefficient on sea-salt dried to below the crystallisation RH was much lower. For aqueous particles with  $\text{RH} > 50\%$ ,  $\gamma$  is independent of chloride or bromide concentration or relative humidity. Within experimental scatter and the range covered there is also no dependence of  $\gamma$  on the temperature, with datasets on synthetic salt surfaces at 295 K giving the same uptake coefficient as NaCl and NaCl/NaBr containing aqueous solutions at  $\sim 270$  K.

The observation that the uptake coefficient is insensitive to the aqueous composition (content of chloride, bromide or iodide) and that the yield of  $\text{ClNO}_2$  following uptake of  $\text{N}_2\text{O}_5$  to chloride solutions of concentration  $\geq 1$  M approaches unity (Behnke et al.,

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

1997; Schweitzer et al., 1998; Roberts et al., 2009) has led to the following mechanism being proposed, with dissociation of  $\text{N}_2\text{O}_5$  (Reaction R1) the rate limiting step.



From the reaction scheme above, the yield of  $\text{ClNO}_2$  is defined by competition between hydrolysis (with rate coefficient  $k_2$ ) and reaction with chloride anions (rate coefficient  $k_3$ ) so that:

$$\frac{[\text{ClNO}_2]}{\Delta [\text{N}_2\text{O}_5]} = \left( \frac{k_2 [\text{H}_2\text{O}]}{k_3 [\text{Cl}^-]} + 1 \right)^{-1}$$

10 Schweitzer et al. (1999) reported unity ( $1.00 \pm 0.14$ ) yield of  $\text{ClNO}_2$  per  $\text{N}_2\text{O}_5$  taken up to 1 M  $\text{Cl}^-$  solution. Within the error bounds this is consistent with the results of Behnke et al. (1999) who used the WWFT method to derive yields of  $\text{ClNO}_2$  from 0.4 (at their lowest, non-zero chloride concentration) to  $>0.9$  at 2 M chloride and above. From their data they calculated  $k_3/k_2$  (at 291 K) =  $836 \pm 32$ . On synthetic sea-salt (50 % RH)

15 Thornton and Abbatt (2005) derived a lower limit to the  $\text{ClNO}_2$  yield of 50 %. Roberts et al. (2009), measured chloride molarity dependent yields of  $\text{ClNO}_2$  over the range 0.02 to 0.5 M  $\text{Cl}^-$ . A number of chloride containing substrates were examined including  $(\text{NH}_4)\text{HSO}_4$ ,  $(\text{NH}_4)_2\text{SO}_4$ , water, oxalic acid, sea-salt. The yield of  $\text{ClNO}_2$  depended only on the chloride concentration, though there may have been evidence for a slightly

20 enhanced yield at low pH. They calculated  $k_3/k_2$  (at 297 K) =  $450 \pm 100$ . Bertram and Thornton (2009) examined the effect of chloride and nitrate concentrations on the  $\text{ClNO}_2$  yield from the uptake of  $\text{N}_2\text{O}_5$  to mixed nitrate/chloride particles and derived  $k_3/k_2$  (at 298 K) =  $483 \pm 175$ .

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

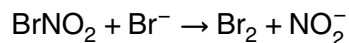
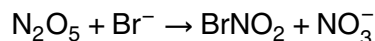
Printer-friendly Version

Interactive Discussion

Our preferred value for the value of  $k_3/k_2$  (at 298 K) is based on the most detailed study (Roberts et al., 2009). Roberts et al. (2009) suggest that  $k_2$  (but not the ion recombination reaction,  $k_3$ ) is likely to have a large barrier and thus be slower at lower temperatures, which would result in a larger  $k_3/k_2$  and thus larger ClNO<sub>2</sub> yield for a given Cl<sup>-</sup> concentration. The experimental datasets are however not precise enough to clarify if the larger  $k_3/k_2$  value of Behnke et al. (1997) is due to use of slightly lower temperatures in their experiments or if it is related to use of various ionic strength solutions and thus Cl<sup>-</sup> activities. We note also that, at low pH the yield of ClNO<sub>2</sub> can be depleted due to conversion to Cl<sub>2</sub> (Roberts et al., 2008). This is covered in datasheet VI.A2.9 dealing with ClNO<sub>2</sub> uptake.

Schweitzer et al. (1999) report BrNO<sub>2</sub> and Br<sub>2</sub> formation when using NaBr solutions and I<sub>2</sub> when using NaI solution.

There have been a number of studies on the uptake of N<sub>2</sub>O<sub>5</sub> to dry salt surfaces (Finlayson-Pitts, 1989; Leu et al., 1995; Fenter et al., 1996; Msibi et al., 1998; Koch et al., 1999; Hoffmann et al., 2003). These studies also report high yields of ClNO<sub>2</sub> when N<sub>2</sub>O<sub>5</sub> reacts with a chloride surface (Finlayson-Pitts et al., 1996; Fenter et al., 1996; Koch et al., 1999; Hoffmann et al., 2003). The reaction with bromide leads to Br<sub>2</sub> formation, presumably via formation of BrNO<sub>2</sub>, which can further react with surface bromide (Fenter et al., 1996):



## References

Behnke, W., Krüger, H.-U., Scheer, V., and Zetzsch, C.: J. Aerosol Sci., 22, 609–612, 1991.

Behnke, W., Krüger, H.-U., Scheer, V., and Zetzsch, C.: J. Aerosol. Sci., 23, S933–S936, 1992.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Behnke, W., Scheer, V., and Zetzsch, C.: J. Aerosol Sci., 24, S115–S116, 1993.

Behnke, W., George, C., Scheer, V., and Zetzsch, C.: J. Geophys. Res., 102, 3795–3804, 1997.

Bertram, T. H. and Thornton, J. A.: Toward a general parameterization of N<sub>2</sub>O<sub>5</sub> reactivity on aqueous particles: the competing effects of particle liquid water, nitrate and chloride, Atmos. Chem. Phys., 9, 8351–8363, doi:10.5194/acp-9-8351-2009, 2009.

Fenter, F. F., Caloz, F., and Rossi, M. J.: J. Phys. Chem., 100, 1008–1019, 1996.

Finlayson-Pitts, B. J., Ezell, M. J., and Pitts, J. N. J.: Nature, 337, 241–244, 1989.

George, C., Ponche, J. L., Mirabel, P., Behnke, W., Scheer, V., and Zetzsch, C.: J. Phys. Chem., 98, 8780–8784, 1994.

Hoffman, R. C., Gebel, M. E., Fox, B. S., and Finlayson-Pitts, B. J.: Phys. Chem. Chem. Phys., 5, 1780–1789, 2003.

Koch, T. G., Vandenbergh, H., and Rossi, M. J.: Phys. Chem. Chem. Phys., 1, 2687–2694, 1999.

Leu, M. T., Timonen, R. S., Keyser, L. F., and Yung, Y. L.: J. Phys. Chem., 99, 13203–13212, 1995.

Msibi, I. M., Li, Y., Shi, J. P., and Harrison, R. M.: J. Atmos. Chem., 18, 291–300, 1994.

Roberts, J. M., Osthoff, H. D., Brown, S. S., and Ravishankara, A. R.: Science, 321, 1059–1059, 2008.

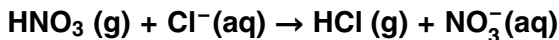
Roberts, J. M., Osthoff, H. D., Brown, S. S., Ravishankara, A. R., Coffman, D., Quinn, P., and Bates, T.: Geophys. Res. Lett., 36, L20808, doi:10.1029/2009GL040448, 2009.

Schweitzer, F., Mirabel, P., and George, C.: J. Phys. Chem. A, 102, 3942–3952, 1998.

Stewart, D. J., Griffiths, P. T., and Cox, R. A.: Atmos. Chem. Phys., 4, 1381–1388, 2004, <http://www.atmos-chem-phys.net/4/1381/2004/>.

Thornton, J. A. and Abbatt, J. P. D.: J. Phys. Chem. A, 109, 10004–10012, 2005.





### Experimental data

Parameter	RH/%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{ss}</math>, <math>\gamma_0</math></i>				
$\gamma > 0.2$ (deliquesced NaCl)	75	298	Abbatt and Waschewsky (1998)	AFT-CIMS (a)
$\gamma = 0.5 \pm 0.2$ (deliquesced sea salt)	55	300	Guimbaud et al. (2002)	AFT-RC (b)
$\gamma_0 = (4.9 \pm 2.7) \times 10^{-3}$ (deliquesced NaCl)	80	298	Tolocka et al. (2004)	AFT-AMS (c)
$\gamma_0 = 0.15 \pm 0.05$ (deliquesced NaCl)	60	298	Saul et al. (2006)	AFT-AMS (d)
$\gamma_0 = 0.15 \pm 0.05$ (deliquesced NaCl/MgCl <sub>2</sub> )	60	298		
$\gamma = 0.21 - 0.11$ (deliquesced NaCl)	55–80	298	Liu et al. (2007)	(e)
$\gamma = 0.25 - 0.12$ (NaCl/MgCl <sub>2</sub> )	55–80			
$\gamma = 0.27 - 0.12$ (sea salt)	55–80			
$\gamma = 0.5^{+0.5}_{-0.2}$ (deliquesced NaCl)	60	296	Stemmler et al. (2008)	AFT-RC/IC (f)

### 5 Comments

(a) NaCl aerosol was produced using a nebulizer. The particle sizes were measured using an optical particle counter and were in the range of 2 to 4  $\mu\text{m}$ . The  $\text{HNO}_3$  concentration (about  $10^{13} \text{ cm}^{-3}$ ) was measured downstream of the flowtube using chemical ionization mass spectrometry. The uptake coefficient was calculated from the measured apparent first order loss rate constant including correction for diffusion in the gas phase, leading to the lower limit given in the table.

(b) Sea salt aerosol was produced by nebulizing a sea salt solution. The particle size distribution (with a mode around 70 nm) was monitored by SMPS. Experiments were performed at 55 % RH.  $\text{HNO}_3$  (in the concentration range of  $10^{11}$  to  $10^{13} \text{ cm}^{-3}$ ) was labeled with the short-lived radioactive tracer  $^{13}\text{N}$ . Both, the loss of  $\text{HNO}_3(\text{g})$  and appearance of aerosol phase product was monitored. The up-

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

take coefficient (with only minor correction by diffusion) was independent of gas phase concentration up to  $10^{12} \text{ cm}^{-3}$  and decreased at higher concentration due to depletion of chloride in the particulate phase. The Aerosol Inorganic Model (AIM) was used to confirm that the low concentration data were obtained under conditions far from equilibrium of the  $\text{HNO}_3\text{-(g,aq)-HCl-(g,aq)}$  system.

- (c) NaCl aerosol produced using a nebulizer, from which a monodisperse aerosol (100–220 nm diameter) was selected using a differential mobility analyzer. The size distribution was measured using SMPS. The product aerosol was sampled into a real-time laser ablation (193 nm) single particle time-of-flight mass spectrometer operating in negative-ion mode, from which the mole fraction of chloride and nitrate was obtained. Removal of chloride from the particulate phase was used to derive an uptake coefficient for  $\text{HNO}_3$ . The reported uptake coefficient was independent of  $\text{HNO}_3$  concentration between  $10^{12}$  and  $10^{13} \text{ cm}^{-3}$ . The small uptake coefficients were explained by slow formation of HCl in the aqueous phase that appeared to limit uptake of  $\text{HNO}_3$ , the main argument being the observed size dependence that could not be due to gas phase diffusion or liquid phase diffusion.
- (d) Same experimental approach as in (c). Additionally, mixtures of NaCl with 0.114 mole fraction of  $\text{MgCl}_2$  were used. The differences in the setup included changes to the flow reactor design that led to systematic errors in the quantification of the actual  $\text{HNO}_3$  concentration in the reactor of the previous study. With this revised setup, the observed chloride mole fractions as a function of exposure time at  $2 \times 10^{12} \text{ cm}^{-3}$  of  $\text{HNO}_3$  was consistent with those in the previous study. However, the uptake coefficient derived using the correct  $\text{HNO}_3$  concentration led to uptake coefficients between 0.1 and 0.2, still depending on size. Therefore, it is still suggested that formation of HCl in the aqueous phase is limiting uptake of  $\text{HNO}_3$ .
- (e) NaCl aerosol produced using a nebulizer and then deposited onto TEM grids (formvar on nickel grids) using an impactor. Experiments were performed by im-

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



pinging a laminar jet of humidified  $\text{HNO}_3$  ( $10^{11} \mu\text{m}^{-3}$ ) in  $\text{N}_2$  onto the sample. The loss of  $\text{HNO}_3$  over the sample was measured by a chemiluminescence detector. Chemical processing of individual particles was monitored using offline scanning electron microscopy connected with energy dispersive X-ray analysis. The dry particle diameter was  $0.9 \mu\text{m}$ , the wet particle size was estimated from the dry diameter measured in the SEM and hygroscopic growth curves in the literature. The dynamics of the trace gas – particle on the substrate interaction was accompanied by CFD to support the relationship between the measured loss rate, the particle density and the uptake coefficient on individual particles. It showed that the net uptake coefficient was largely controlled by gas phase diffusion and was therefore dependent on the (wet) particle diameter. The decrease of the uptake coefficient with increasing RH above the efflorescence point is explained by the decreasing chloride concentration in the particle, assuming that the formation of HCl is the rate limiting factor. A correction is not applied to account for gas phase diffusion and to report a true uptake coefficient at the particle surface.

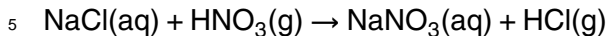
- (f) ~~Same experimental approach as in (b) with only slight modifications. The main aim was to investigate the effects of surfactants on the  $\text{HNO}_3$  uptake. In the absence of surfactants, the uptake coefficients were consistent with those reported by Guimbaud et al. (2002). Offline, ion chromatographic analysis of filter samples confirmed the complete displacement of reacted chloride. The revised error analysis included a discussion of the mixing time in the flow tube and re-equilibration in the denuder system.~~

## Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.5	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298

## Comments on preferred values

The available studies agree that the acid displacement



drives the uptake of  $\text{HNO}_3$  to sea salt solutions and their proxies, deliquesced  $\text{NaCl}$  or mixtures of  $\text{NaCl}$  with  $\text{MgCl}_2$ , and that the uptake is fast. This is in line with many studies on solid halogenide salts, where the uptake coefficient of  $\text{HNO}_3$  showed a strong dependence on the amount of surface adsorbed water, providing evidence that chloride ion (partially or fully hydrated) is required to drive acid displacement (see Rossi, 2005, and references therein). The Abbatt and Waschewsky (1998) and the Liu et al. (2007) data were affected to a varying degree by gas phase diffusion, so that those reported uptake coefficients should be considered as lower limits. Regarding the absolute values of the uptake coefficient, the Saul et al. (2006) data based on an improved calibration of the  $\text{HNO}_3$  concentration are considered to supersede the Tolocka et al. (2004) data. Therefore, the available studies agree fairly well. From the very high solubility of  $\text{HNO}_3$ , the very fast rates for its dissociation in the aqueous phase, and the very fast protonation of chloride, we conclude that solubility, reaction or diffusion cannot limit the rate of uptake of  $\text{HNO}_3$ , which was suspected to be responsible for the size and humidity dependence by Saul et al. , Liu et al. and also earlier by ten Brink (1998). We therefore recommend the bulk accommodation coefficient observed by Guimbaud et al. (2002)

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



and Stemmler et al. (2008) as a lower limit. With NaCl solution in excess,  $\text{HNO}_3$  – protons are nearly completely displaced into the gas phase as HCl at equilibrium due to the divergent behaviour of the activity coefficients of chloride and nitrate at high ionic strength (Brimblecombe and Clegg, 1988).

## 5 References

Brimblecombe, P. and Clegg, S. L.: J. Atmos. Chem., 7, 1–18, 1988.

Guimbaud, C., Arens, F., Gutzwiller, L., Gäggeler, H. W., and Ammann, M.: Atmos. Chem. Phys., 2, 249–257, 2002,  
<http://www.atmos-chem-phys.net/2/249/2002/>.

10 Liu, Y., Cain, J. P., Wang, H., and Laskin, A.: J. Phys. Chem. A, 111, 10026–10043, 2007.

Rossi, M. J.: Chem. Rev., 103, 4823–4882, 2003.

Saul, T. D., Tolocka, M. P., and Johnston, M. V.: J. Phys. Chem. A, 110, 7614–7620, 2006.

15 Stemmler, K., Vlasenko, A., Guimbaud, C., and Ammann, M.: Atmos. Chem. Phys., 8, 5127–5141, 2008,  
<http://www.atmos-chem-phys.net/8/5127/2008/>.

ten Brink, H. M.: J. Aerosol Sci., 29, 57–64, 1998.

20 Tolocka, M. P., Saul, T. D., and Johnston, M. V.: J. Phys. Chem. A, 108, 2659–2665, 2004.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A2.8

### HOCl + Cl<sup>-</sup>/Br<sup>-</sup> (aq) → products

#### Experimental data

Parameter	Aqueous solution	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients, <math>\gamma</math></i> (0.39–1.79) × 10 <sup>-3</sup>	Natural salt (pH ≈ -0.7 to -1, RH 75%–85%)	296	Pratte and Rossi (2006)	AFT-MS (a)
<2 × 10 <sup>-4</sup>	Cl <sup>-</sup> and RSS (pH ≈ -0.7 to -1, RH 75%–85%)			

#### 5 Comments

(a) Atmospheric pressure flow tube. Aerosols (≈200–300 nm diameter, 1.7–11 × 10<sup>-4</sup> cm<sup>2</sup> cm<sup>-3</sup> total surface area density) were made from acidified salt solutions (0.034 M Cl<sup>-</sup>) at pH of 1. The salts used were pure NaCl, re-crystallised sea-salt (RSS) and natural sea-salt (NSS). The composition of the aerosol (3–3.8 M Cl<sup>-</sup>, 2.5–6.4 M H<sub>2</sub>SO<sub>4</sub>) was varied by adjusting the relative humidity. No uptake of HOCl was observed for NaCl and RSS.

#### Preferred values

Parameter	Value	T/K
$\gamma$	<2 × 10 <sup>-4</sup>	296
<i>reliability</i>		
$\Delta \log(\gamma)$	undetermined	

32262

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred values

Pratte and Rossi (2006) observed a low uptake coefficient for HOCl to natural sea-salt solutions, with a negative dependence of  $\gamma$  on relative humidity (i.e. with a positive dependence on ion concentrations). Uptake to pure chloride solution with the same composition was not observed.

The observation of Pratte and Rossi (2006), of measurable uptake onto aqueous NSS, but not to aqueous RSS (presumably with similar chloride/bromide composition) is difficult to understand. No products were observed in that study, so it is not possible to determine which reaction was driving the uptake. Pratte and Rossi also speculate that the high  $\text{H}_2\text{SO}_4$  concentration may inhibit the uptake.

Inserting values for the termolecular rate constant for reaction of HOCl with  $\text{H}^+/\text{Cl}^-$  ( $k_{\text{ter}} = 1.5 \times 10^4 \text{ M}^{-2} \text{ s}^{-1}$ , Wang and Margerum, 1994) the solubility of HOCl ( $H_{\text{HOCl}} = 6.6 \times 10^2 \text{ M atm}^{-1}$ , Huthwelker et al., 1995) a value for the liquid phase diffusion coefficient of  $2 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$  and an accommodation coefficient ( $\alpha_b$ ) of between 1 and 0.01 into the expression below gives rise to values of  $\gamma$  between  $\approx 0.1$  and 0.01 at 290 K for 2 M  $\text{Cl}^-$  and a pH of 2.

$$\gamma = \left\{ \frac{1}{\alpha} + \frac{\bar{c}}{4HRT(D_l k_{\text{ter}}[\text{H}^+][\text{Cl}^-])^{0.5}} \right\}^{-1}$$

Uptake coefficients of this magnitude should have been observable. Note that the forward reaction forms  $\text{Cl}_2$  ( $\text{HOCl} + \text{H}^+ + \text{Cl}^- \rightarrow \text{Cl}_2 + \text{H}_2\text{O}$ ) which may hydrolyse back to HOCl. The equilibrium constant  $K = [\text{HOCl}][\text{H}^+][\text{Cl}^-]/[\text{Cl}_2]$  is circa  $1 \times 10^{-3} \text{ M}^2$  at 25 °C, so that, a pH of 2 and 2 M  $\text{Cl}^-$  result in an equilibrium ratio of  $\text{HOCl}/\text{Cl}_2 \approx 0.05$ .

Experiments on solid salt surfaces (Santschi and Rossi, 2005; Huff and Abbatt, 2000) revealed no reactivity for HOCl on chloride surfaces, but high reactivity for pure bromide and mixed bromide/chloride surfaces. The uptake of HOCl to a frozen, mixed chloride/bromide surface at 248 K has been shown to result in formation of mainly  $\text{Br}_2$  (and little BrCl) as gas-phase product. This presumably arises via conversion of BrCl

32263

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

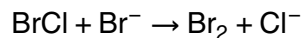
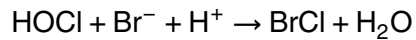
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(initially formed) with surface bromide.



- 5 Until further data is available, we prefer to recommend an upper limit to  $\gamma$  for reaction of HOCl with chloride/bromide containing aqueous particles.

## References

Huff, A. K. and Abbatt, J. P. D.: J. Phys. Chem., 104, 7284–7293, 2000.

10 Huthwelker, T., Clegg, S. L., Peter, T., Carslaw, K., and Luo, B. P.: J. Atmos. Chem., 21, 81–95, 1995.

Pratte, P. and Rossi, M. J.: Phys. Chem. Chem. Phys., 8, 3988–4001, 2006.

Santschi, C. and Rossi, M. J.: Phys. Chem. Chem. Phys., 7, 2599–2609, 2005.

Wang, T. X. and Margerum, D. W.: Inorg. Chem., 33, 1050–1055, 1994.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.9

### CINO<sub>2</sub> + Cl<sup>-</sup>/Br<sup>-</sup> → products

#### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{ss}</math>, <math>\gamma_0</math></i>				
$4.8 \times 10^{-6}$	4.6 M NaCl	300	Behnke et al. (1997)	WWFT-FTIR (a)
$(1.4 \pm 0.2) \times 10^{-5}$	0.5 and 1 mM NaBr, HBr	275, 291	Frenzel et al. (1998)	WWFT-FTIR (b)
$\gamma_0 = 9.0 \times 10^{-6}$		282		
$\gamma_{ss} = (6.7 \pm 1.3) \times 10^{-6}$	10 <sup>-4</sup> M NaBr	275–288	Schweitzer et al. (1998)	WWFT-FTIR/MS (c)
$\gamma_{ss} = (8.2 \pm 1.0) \times 10^{-4}$	10 <sup>-2</sup> M NaBr			
$\gamma_{ss} = (8.8 \pm 2.6) \times 10^{-4}$	1 M NaBr			
$\gamma_{ss} = (1 - 10) \times 10^{-4}$	10 <sup>-4</sup> to 1 M NaBr aqueous film	274	Fickert et al. (1998)	WWFT-MS (d)
$(6 \pm 2) \times 10^{-3}$	0.05 to 6 M NaCl aqueous film; pH < 2	298	Roberts et al. (2008)	WWFT-CRDS/CIMS (e)

#### 5 Comments

(a) Variable length (10–80 cm) wetted-wall flow tube at 1 bar of synthetic air. The liquid film thickness was between 0.8 and 1.2 mm at a surface speed of 2–10 cm s<sup>-1</sup>. The gas flow rate was in the range 200–400 mL min<sup>-1</sup> at 291 K resulting in an average linear flow velocity of 15–30 cm s<sup>-1</sup>. The initial [CINO<sub>2</sub>] was measured upstream, the unreacted CINO<sub>2</sub> was measured downstream of the WWFT using FTIR. Between pure H<sub>2</sub>O and 1 M NaCl  $\gamma$  decreases by more than a factor of ten.

(b) Uptake study in a variable length wetted-wall flow tube at P = 1 atm of synthetic air. Experimental details may be found in the data sheet for CINO<sub>2</sub> + H<sub>2</sub>O under (b). The uptake coefficients were obtained by fitting the measured initial time/position-dependent concentration profiles in the range 0 to 7 s using a complex reaction mechanism including gas phase and liquid diffusion. Observed reaction products

32265

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



included  $\text{BrNO}_2$ ,  $\text{Br}_2$ ,  $\text{NO}_2$  in the gas phase and  $\text{Cl}^-$ ,  $\text{NO}_2^-$  and  $\text{NO}_3^-$  in the condensed phase

- (c) Uptake experiment in a wetted-wall flow tube equipped with FTIR/long-path absorption and ion-trap mass spectrometry for gas phase detection of  $\text{ClNO}_2$  and products. The uptake coefficients were independent of temperature in the stated range and scaled linearly when plotted as  $1/\gamma$  vs.  $[\text{NaX}]^{-1}$  with  $\text{X} = \text{I}^-$ . The main gas phase product was  $\text{Br}_2$ .
- (d) Uptake study in a wetted-wall tubular flow reactor on a falling film of pure water and aqueous alkali halide salt solutions at 14 to 18 Torr of He as a carrier gas. The reactant was monitored using a differentially-pumped MS. The value of  $\gamma_{\text{ss}}$  remains unchanged upon addition of 1 M NaCl whereas it increased upon addition of 0.1 M  $\text{OH}^-$ . Analysis of uptake rates measured under gas-phase diffusion controlled conditions (0.5 M KOH film) gave the cited value for the accommodation coefficient on aqueous surfaces,  $\alpha = (9 \pm 4) \times 10^{-3}$ . Estimates of the diffusion coefficients  $D^{\text{ClNO}_2}$  in He,  $\text{N}_2$  and  $\text{H}_2\text{O}$  are reported as  $275 \pm 26$ ,  $75 \pm 6$  and  $100 \pm 20$  Torr  $\text{cm s}^{-1}$  respectively. Parameters determining reactive uptake rates into  $\text{Br}_2$  solutions were determined interdependently as  $H^2 D_1 k^{\text{II}} = (0.101 \pm 0.015) \text{ M cm}^2 \text{ s}^{-2}$ . The gas phase products were found to be  $\text{Br}_2$ ,  $\text{BrNO}_2$  and minor amounts of  $\text{BrCl}$ .  $\text{Br}_2$  was a secondary reaction product resulting from the condensed phase reaction of  $\text{BrNO}_2$  with  $\text{Br}^-$ .
- (e) Uptake study in a tubular flow reactor where  $\text{N}_2\text{O}_5$  was converted to  $\text{ClNO}_2$  by reaction on aqueous slurry of NaCl and the resultant  $\text{ClNO}_2$  reacted with NaCl/oxalic acid solution at  $\text{pH} = 1.8$ . Relatively fast uptake of  $\text{ClNO}_2$  attributed to acid catalysed reaction producing  $\text{Cl}_2$  as the main product. The reactant  $\text{N}_2\text{O}_5$  was monitored by CRDS after thermal-decomposition to  $\text{NO}_3$ . Products were monitored by  $\text{I}^-$ -CIMS. Analysis of uptake rates gave an estimate of  $>10^7 \text{ M}^{-1} \text{ s}^{-1}$  for the rate constant for the reaction of  $\text{ClNO}_2$  with  $\text{Cl}^-$ .

## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.01	274
$k^{\text{II}}$ ( $\text{M}^{-1} \text{s}^{-1}$ ) (Cl; $\text{pH} < 2$ )	$> 10^7$	
$H^2 D_1 k^{\text{II}}$ ( $\text{M cm}^2 \text{s}^{-2}$ ) (Br)	$0.101 \pm 0.015$	274
Reliability		
$\Delta \log(\alpha_b)$	$\pm 0.3$	

## Comments on preferred values

The results of the earlier studies are in broad agreement, and the recent work revealed an important change in mechanism at conditions of low pH leading to  $\text{Cl}_2$  production. The uptake of  $\text{ClNO}_2$  into the aqueous phase is slow into neutral solutions containing  $\text{Cl}^-$  but is significantly enhanced in acid solutions ( $\text{pH} < 2$ ).  $\text{Cl}_2$  is not formed from neutral solutions but is produced with high yield at  $\text{pH} = 1.8$  even at  $[\text{Cl}^-] = 0.05 \text{ M}$  (Roberts et al., 2008), probably due to an acid catalysed reaction. Uptake is also significantly enhanced on introduction of  $\text{Br}^-$  in the condensed phase, with formation of products  $\text{BrNO}_2/\text{Br}_2$  which will partition rapidly to the gas phase. Reactive uptake is controlled by chemical reaction in the bulk, i.e. in terms of the resistance model:

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1 k^{\text{I}})^{0.5}} \right\}^{-1} \quad \text{where } k^{\text{I}} = k^{\text{II}}[\text{X}^-]_{\text{aq}} \quad (\text{M, X} = \text{Cl}^- \text{ or } \text{Br}^-)$$

The preferred value of  $\alpha_b$  is taken from the work of Fickert et al. (1998). The preferred values for uptake on  $\text{Cl}^-$  are based on the work of Roberts et al. (2008), who used a value of  $H = 4 \times 10^{-2} \text{ Matm}^{-1}$  in their analysis. The preferred values for uptake on  $\text{Br}^-$  are based on the thorough analysis reported in the work of Fickert et

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



al. (1998). The liquid phase diffusion coefficient has not been determined but a value of  $\sim 1 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$  is normally used.

## References

- Behnke, W., George, C., Scheer, V., and Zetzsch, C.: J. Geophys. Res., 102, 3795–3804, 1997.
- 5 Fickert, S., Helleis, F., Adams, J., Moortgat, G. K., and Crowley, J. N.: J. Phys. Chem. A, 102, 10689–10696, 1998.
- Frenzel, A., Scheer, V., Sikorski, R., George, C., Behnke, W., and Zetzsch, C.: J. Phys. Chem. A, 102, 1329–1337, 1998.
- 10 George, C., Behnke, W., Scheer, V., Zetzsch, C., Magi, L., Ponche, J. L., and Mirabel, Ph.: Geophys. Res. Lett., 22, 1505–1508, 1995.
- Roberts, J. M., Osthoff, H. D., Brown, S. S., and Ravishankara, A. R.: Science, 321, 1059, 2008.
- Schweitzer, F., Mirabel, Ph., and George, C.: J. Phys. Chem., 102, 3942–3952, 1998.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

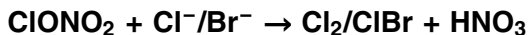
Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{\text{SS}}</math>, <math>\gamma_0</math></i>				
$\gamma = 0.0244 \pm 0.0023$	0.1 M NaCl	274.6	Deiber et al. (2004)	DFT-MS (a)
$\gamma = 0.041 \pm 0.0067$	0.01 M NaBr	274.1		
$\gamma = 0.046 \pm 0.0019$		277		
$\gamma = 0.047 \pm 0.0011$		280.8		
$\gamma = 0.044 \pm 0.0027$	0.025 M NaBr	274.3		
$\gamma = 0.041 \pm 0.0043$		276.2		
$\gamma = 0.055 \pm 0.0060$	0.05 M NaBr	274		
$\gamma = 0.054 \pm 0.0039$		277		
$\gamma = 0.053 \pm 0.0028$		281		
$\gamma = 0.057 \pm 0.0048$	0.10 M NaBr	274.9		
$\gamma = 0.049 \pm 0.0024$		276.2		
$\gamma = 0.056 \pm 0.0008$		279.9		
$\gamma = 0.0655 \pm 0.0015$	0.5 M NaBr	274.4		
$\gamma = 0.0663 \pm 0.0011$		276.5		
$\gamma = 0.073 \pm 0.0106$	1.0 M NaBr			

### 5 Comments

(a) Uptake rates measured onto 200  $\mu\text{m}$  droplets following loss of reactant in conventional droplet train apparatus. Droplet temperature controlled by evaporative cooling with adjustment of  $p(\text{H}_2\text{O})$ . Uptake coefficient determined with a simple correction for diffusion effects.  $\gamma$  measured as function of [NaCl] and [NaBr]. On NaCl  $\gamma$  was not significantly larger than on pure water, but  $\text{Cl}_2$  was observed as a unique gas-phase product. On NaBr droplets  $\gamma$  increased with increasing [Br<sup>-</sup>]

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


and was essentially independent of temperature over the small range investigated for all  $[\text{Br}^-]$ . Both  $\text{BrCl}$  and  $\text{Br}_2$  were detected as gas phase products.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.11	273–290
$H\sqrt{k^{\text{II}}} \text{ (M}^{1/2} \text{ atm}^{-1} \text{ s}^{-1/2}) \text{ (Br}^-)$	$1.0 \times 10^6$	273–290
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.2$	273–290
$\Delta \log(H\sqrt{k^{\text{II}}})$	$\pm 0.15$	273–290

### Comments on preferred values

The cited work is the only study of reactive uptake of  $\text{ClONO}_2$  on aqueous halide substrates  $\text{NaCl}$  and  $\text{NaBr}$ ; all the other reported studies used either solid substrates (ice) or sulphuric acid solutions. These studies showed that uptake led to  $\text{XCl}$  ( $\text{X} = \text{Cl}$  or  $\text{Br}$ ) and  $\text{HNO}_3$  formation. The dihalogens can partition to the gas phase, depending on their solubility. No gas phase products were observed from uptake on water droplets due to the high solubility of the products.

The measured uptake coefficient on  $\text{NaCl}$  was not significantly greater than on pure water droplets (see data sheet for  $\text{ClONO}_2 + \text{H}_2\text{O}$ , VI.A1.16). However it is expected that  $\text{Cl}_2$  rather than  $\text{HOCl}$  will be formed as products. When  $\text{Br}^-$  was present in solution,  $\gamma$  increased with increasing  $[\text{Br}^-]$  and was independent of temperature over the small range investigated for all  $[\text{Br}^-]$ . This was interpreted in terms of the resistance model with increasing liquid phase reaction rate of  $\text{ClONO}_2(\text{aq})$  due to reaction with

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Br<sub>2</sub> allowing accommodatopn controlled uptake at high [Br<sup>-</sup>];

$$\gamma = \left\{ \frac{1}{\alpha} + \frac{\bar{c}}{4HRT(D_1k^1)^{0.5}} \right\}^{-1} \text{ where } k^1(\text{s}^{-1})=k^{II} \times [\text{Br}^-]_{\text{aq}}(\text{M})$$

This allowed evaluation of the reactive uptake parameters for uptake on Br<sup>-</sup> containing solutions at 274.5 K:  $\alpha_b = (0.108 \pm 0.011)$  and the product  $H\sqrt{k^{II}} = 1.0 \times 10^6 \text{ M}^{1/2} \text{ atm}^{-1} \text{ s}^{-1/2}$  from a plot of uptake coefficients corrected for gas phase diffusion effects ( $1/\gamma - (1/\gamma_{\text{diff}})$ , vs.  $1/[\text{NaBr}]^{1/2}$ , according to the resistance model with  $D_1$  assumed to be  $5 \times 10^{-6} \text{ cm}^2 \text{ s}^{-1}$ . The recommended uptake parameters are based on this analysis. There are no reported values of  $H$  or  $D_1$  for ClONO<sub>2</sub>.

## References

10 Deiber, G., George, C., Le Calve, S., Schweitzer, F., and Mirabel, Ph.: Atmos. Chem. Phys., 4, 1291–1299, 2004.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

HOBr + Cl<sup>-</sup>/Br<sup>-</sup> (aq) → products

## Experimental data

Parameter	Aqueous Solution	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients, <math>\gamma</math></i>				
>0.2	NaCl (pH 0.3, 75 % RH)	295	Abbatt and Waschewsky (1998)	AFT-CIMS (a)
>0.2	NaCl (pH 7.2, 75 % RH)			
<1.5 × 10 <sup>-3</sup>	NaCl (unbuffered, 75 % RH)			
8 × 10 <sup>-3</sup>	NaCl and natural salt (pH ≈ -1 to -0.7, RH > 75 %)	296	Pratte and Rossi (2006)	AFT-MS (b)
<i>Accommodation coefficient, <math>\alpha</math></i>				
>0.01	NaCl/NaBr (pH 2–5)	274	Fickert et al. (1999)	WWFT-MS (c)
0.6 ± 0.2	NaBr	296 ± 2	Wachsmuth et al. (2002)	AFT (d)

## 5 Comments

(a) Flow tube at 933–1013 mBar N<sub>2</sub>. Number density (1–4 × 10<sup>4</sup> cm<sup>-3</sup>) and diameter (2–4 μm) of the NaCl particles were measured using an optical particle counter. HOBr (2–10 × 10<sup>12</sup> molecule cm<sup>-3</sup>), was made by passing damp N<sub>2</sub> with traces of Br<sub>2</sub> over HgO and was detected as SF<sub>5</sub>O<sup>-</sup> using SF<sub>6</sub><sup>-</sup> chemi-ions. No uptake observed to unbuffered and non-acidified chloride solutions. Uptake coefficient to acidified and buffered solutions is a lower limit, owing to diffusion limitations.

(b) Atmospheric pressure flow tube. Aerosols (≈200–300 nm diameter, 1.7–11 × 10<sup>-4</sup> cm<sup>2</sup> cm<sup>-3</sup> total surface area density) were made from acidified salt solutions (0.034 M Cl<sup>-</sup>) at pH of 1. The salts used were pure NaCl, re-crystallised sea-salt (RSS) and natural sea-salt (NSS). The composition of the aerosol (1.7–4.4 M Cl<sup>-</sup>, 2.5–6.4 M H<sub>2</sub>SO<sub>4</sub>) was varied by varying the relative humidity. Measured uptake coefficients were found to suddenly increase from ≈2 × 10<sup>-3</sup> to 8 × 10<sup>-3</sup> at RH ≈ 75–80 % for NaCl and RSS and remain high. For NSS a different behaviour

---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**


---

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)




**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[⏪](#)
[⏩](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


was observed with a peak in  $\gamma$  (also  $\approx 8 \times 10^{-3}$ ) at RH = 75 % and lower values (factor of 4 or more) at both higher and lower RH.

(c) Slow flowing aqueous film (50–100  $\mu\text{m}$  thick) on the internal surface of a flow tube operated at  $\approx 13$ –35 mbar He or N<sub>2</sub>. HOBr, Br<sub>2</sub> and BrCl were detected as positive ions using electron impact ionisation. HOBr ( $\approx 10^{12}$  molecule cm<sup>-3</sup>) was eluted into the flow tube from an aqueous solution. The lower limit to  $\alpha$  was obtained from the pressure dependence of the uptake coefficient to an acidified film and assuming a value for the diffusion coefficient of HOBr in H<sub>2</sub>O.

(d) Very low concentrations ( $\approx 300$  molecule cm<sup>-3</sup>) of HOBr detected using radioactive labelling. Uptake to 0.2 M NaBr aerosol particles (diameter 50–60 nm) at RH = 37 % and pH < 6. A small correction (4 %) for diffusion was applied.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.6	296
$k_{\text{ter}}$ (M <sup>-2</sup> s <sup>-1</sup> )	$> 5.6 \times 10^9$	296
$H$ (M atm <sup>-1</sup> )	$6.1 \times 10^3$	296
$D_1$ (cm <sup>2</sup> s <sup>-1</sup> )	$1.4 \times 10^{-5}$	296
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	296

## Comments on preferred values

The large uptake coefficients reported for HOBr reacting with acidified aqueous bromide and chloride indicate that the accommodation coefficient is large (Abbatt and

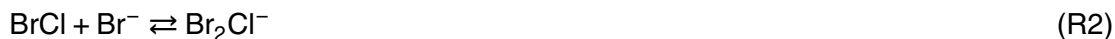
Waschewsky, 1998; Wachsmuth et al., 2002) and possibly unity. In terms of the resistance model, the uptake coefficient for HOBr reacting with sea-salt is given by:

$$\gamma = \left\{ \frac{1}{\alpha} + \frac{\bar{c}}{4HRT(D_1k')^{0.5}} \right\}^{-1}$$

where  $k' = k_{\text{ter}} [H^+][Cl^-]$  and the ionic concentrations are in M.  $R$  is  ~~$8.206 \times 10^{-2} \text{ LatmK}^{-1} \text{ mol}^{-1}$~~ . The termolecular rate constant,  $k_{\text{ter}}$ , was taken from Vogt et al. (1996), the solubility and diffusion coefficient for HOBr were taken from Frenzel et al. (1998).

This expression predicts uptake coefficients that are close to the accommodation coefficient for  $\text{pH} < 7$  when applied to fresh sea-salt with  $[Cl^-] \approx 5.3 \text{ M}$ . This is in accord with the large lower limit to the uptake coefficient measured by Abbatt and Waschewski, but is higher than the uptake coefficients (factor 20) determined by Rossi and Pratte (2006). The observation of much lower uptake coefficients when the droplets are not acidified and/or buffered (Abbatt and Waschewsky, 1998) is related to rapid depletion of available  $[H^+]$  when using large concentrations of HOBr.

The important aqueous phase reactions in sea-salt aerosol containing chloride and bromide ions may be briefly summarised as follows:



The relative efficiency of  $\text{Br}_2$  and  $\text{BrCl}$  release from acidified halide solutions ( $\text{pH} < 3$  or buffered to  $< 5.6$ ) with varying bromide/chloride ratios was found (Fickert et al., 1999) to be consistent with aqueous phase equilibrium constants for Reactions (R2) and (R3) (Wang et al., 1994). Following bulk accommodation to salt solutions with the approximate composition of sea-water ( $[Cl^-]/[Br^-] \approx 700$ ), HOBr reacts almost exclusively with

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



H<sup>+</sup>/Cl<sup>-</sup> to release mainly Br<sub>2</sub>. Only at reduced [Br<sup>-</sup>] are significant amounts of BrCl released (Fickert et al., 1999). Fickert et al. (1999) found that yields of Br<sub>2</sub> (or BrCl at low Br<sup>-</sup>) were close to 100% of the HOBr taken into solution as long as the pH was lower than ≈7.

5 Several experimental investigations of the uptake of HOBr to dry and frozen halide surfaces (both pure chloride or bromide or mixed) have revealed a similar chemistry to that outlined above (Kirchner et al., 1997; Mochida et al., 1998; Chu et al., 2002; Huff and Abbatt, 2002; Adams et al., 2002). On pure chloride samples, BrCl is released at high yield and on pure bromide samples, Br<sub>2</sub> is the sole product. On mixed chloro-  
10 ride/bromide samples both Br<sub>2</sub> and BrCl are observed, with a dependence on initial composition and temperature of the surface.

## References

- Abbatt, J. P. D. and Waschewsky, G. C. G.: J. Phys. Chem., 102, 3719–3725, 1998.
- Adams, J. W., Holmes, N. S., and Crowley, J. N.: Atmos. Chem. Phys., 2, 79–91, 2002,  
15 <http://www.atmos-chem-phys.net/2/79/2002/>.
- Chu, L. and Chu, L. T.: J. Phys. Chem. A, 103, 8640–8649, 1999.
- Fickert, S., Adams, J. W., and Crowley, J. N.: J. Geophys. Res., 104, 23719–23727, 1999.
- Frenzel, A., Scheer, V., Sikorski, R., George, C., Behnke, W., and Zetzsch, C.: J. Phys. Chem. A, 102, 1329–1337, 1998.  
20
- Huff, A. K. and Abbatt, J. P. D.: J. Phys. Chem. A, 106, 5279–5287, 2002.
- Kirchner, U., Benter, T., and Schindler, R. N.: Ber. Bunsen Ges. Phys. Chem., 101, 975–977, 1997.
- Mochida, M., Akimoto, H., van den Bergh, H., and Rossi, M. J.: J. Phys. Chem., 102,  
25 4819–4828, 1998.
- Pratte, P. and Rossi, M. J.: Phys. Chem. Chem. Phys., 8, 3988–4001, 2006.
- Vogt, R., Crutzen, P. J., and Sander, R.: Nature, 383, 327–330, 1996.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Wachsmuth, M., Gäggeler, H. W., von Glasow, R., and Ammann, M.: Atmos. Chem. Phys., 2, 121–131, 2002,  
<http://www.atmos-chem-phys.net/2/121/2002/>.  
Wang, T. X., Kelley, M. D., Cooper, J. N., Beckwith, C., and Margerum, D. W.: Inorg. Chem., 33, 5872–5878, 1994.

5

32276

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

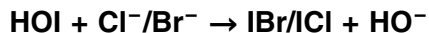
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.12



### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{ss}</math>, <math>\gamma_0</math></i>				
$\gamma = 2.2 \times 10^{-3}$	2 (Cl <sup>-</sup> )	274	Braban et al. (2007)	WWFT-EIMS (a)
$\gamma = 2.0 \times 10^{-3}$	pure water			

### 5 Comments

(a) The uptake of HOI ( $5\text{--}50 \times 10^{10}$  molecules  $\text{cm}^{-3}$ ) to aqueous surfaces was studied in the wetted-wall flow tube with MS detection at  $m/z = 144$ . Several combinations of dissolved concentrations of Cl<sup>-</sup> and Br<sup>-</sup> in solutions of different pH were used. Uptake on pure water films was also measured. Experiments were all conducted at 274 K and at flow tube total pressures of 13 and 20 Torr. HOI was prepared in situ via the reaction of O atoms (generated in a microwave discharge) with C<sub>2</sub>H<sub>5</sub>I or C<sub>3</sub>H<sub>7</sub>I in a flow of He at reduced pressure. The uptake coefficient values were independent of composition or pH of the solution, except for pure water or very low electrolyte concentration. Calculations using HOI diffusion coefficients of  $D_{\text{H}_2\text{O}}^{\text{HOI}} = 36.8 \text{ cm}^2 \text{ s}^{-1} \text{ Torr}^{-1}$ , and  $D_{\text{He}}^{\text{HOI}} = 318.2 \text{ cm}^2 \text{ s}^{-1} \text{ Torr}^{-1}$  at 274 K, obtained by extrapolation from the data of Holmes et al. (2000), showed that uptake is gas phase diffusion limited in all cases. Uptake of HOI produced dihalogens ICl and IBr which were partially released to the gas phase; IBr was preferentially released whenever Br<sup>-</sup> was present with Cl<sup>-</sup>.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.1	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	Not determined	

## Comments on preferred values

Uptake and reaction of HOI in aqueous films evidently occurs rapidly. However the results do not allow determination of the accommodation coefficient for HOI on aqueous surfaces as the uptake was diffusion limited, based on the diffusion coefficients reported by Holmes et al. (2001); only a lower limit of  $\alpha_{\text{HOI}} > 2.2 \times 10^{-3}$  can be deduced. Other studies using different surfaces suggest that  $\alpha$  is substantially higher. A lower limit of  $\alpha_{\text{HOI}} = 0.3$  was given by Holmes et al. (2001) for HOI on aqueous sulfuric acid surfaces, and  $\alpha_{\text{HOI}} > 0.12$  on dry salt particles at 253 K. Mossinger and Cox (2001) observed a reactive uptake coefficient of  $\gamma = 0.06$  on solid sea salt aerosol at 23 % RH, which leads to formation of IBr and ICl products. Thus it is highly likely that the accommodation coefficient of HOI on aqueous surfaces is at least 0.1 at ambient temperatures.

Rate coefficients for  $\text{HOI} + \text{X}^-$  in solution have not been reported. For HOBr the reaction is fast and uptake is accommodation limited under atmospheric conditions, and the same is likely for reactive uptake of HOI into salt droplets. The dihalogen products ICl and IBr formed by reaction with  $\text{Cl}^-$  or  $\text{Br}^-$  will partition rapidly to the gas phase. In mixed halide solutions IBr is preferentially released whenever  $\text{Br}^-$  is present with  $\text{Cl}^-$ , and ICl is only released when  $\text{Br}^-$  has been chemically depleted.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## References

Braban, C. F., Adams, J. W., Rodriguez, D., Cox, R. A., Crowley, J. N., and Schuster, G.: Phys. Chem. Chem. Phys., 9, 3136–3148, 2007.

5 Holmes, N. S., Adams, J. W., and Crowley, J. N.: Phys. Chem. Chem. Phys., 3, 1679–1687, 2001.

Mössinger, J. A. and Cox, R. A.: J. Phys. Chem. A, 105, 5165–5177, 2001.

Wachsmuth, M., Gaeggeler, H. W., Von Glasow, R., and Ammann, M.: Atmos. Chem. Phys., 2, 121–131, 2002,

<http://www.atmos-chem-phys.net/2/121/2002/>.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.13

### BrCl + Br<sup>-</sup>(aq) → Products

#### Experimental data

Parameter	Aqueous Solution	Temp./K	Reference	Technique/Comments
Accommodation coefficient, $\alpha$				
0.33 ± 0.18		270–285	Katrib et al. (2001)	DT-MS (a)

#### 5 Comments

(a) BrCl was eluted into the gas phase from an aqueous solution and detected by mass spectrometer at  $m/z = 116$ . Br<sub>2</sub> and Cl<sub>2</sub> were present as impurities. Droplet size was  $\approx 150 \mu\text{m}$ . The uptake of BrCl to bromide solutions did not follow first-order kinetics and no values of  $\gamma$  were reported. The accommodation coefficient was obtained from experiments using a different aqueous phase scavenger of BrCl. No uptake was observed for pure H<sub>2</sub>O at pH = 5.6 ( $\gamma < 10^{-3}$ ).

#### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.33	270–280
Reliability		
$\Delta \log(\alpha_b)$	±0.5	270–280

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

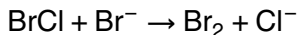
Interactive Discussion





## Comments on preferred values

The only experimental study of the uptake of BrCl to bromide containing aqueous solutions ( $5 \times 10^{-2}$  M) (Katrib et al., 2001) did not return values of  $\gamma$ . This was reported as being due to complications involving reaction of  $\text{Cl}_2$  impurity with the surface resulting in non first-order kinetics. BrCl is known to react in aqueous solutions with  $\text{Br}^-$  to form  $\text{Br}_2\text{Cl}$ , which is in equilibrium with  $\text{Br}_2$  and  $\text{Cl}^-$ . The net reaction can be written as:



Using the expression below and literature values for the aqueous-phase solubility of BrCl at 275 K ( $H = 4.6 \text{ M atm}^{-1}$ , Bartlett and Margerum, 1999), the rate constant for reaction of BrCl with  $\text{Br}^-$  ( $k_{\text{Br}} > 10^8 \text{ M s}^{-1}$ , Wang et al., 1994) an aqueous-phase diffusion coefficient of  $10^{-5} \text{ cm}^2 \text{ s}^{-1}$  and the value of  $\alpha$  reported above results in uptake coefficients between 0.02 and 0.2 when  $[\text{Br}^-]$  is varied between  $10^{-3}$  and 1 M. The rate of uptake is thus accommodation controlled at high  $[\text{Br}^-]$  ( $\gamma \rightarrow \alpha_b$ ) and chemically controlled at low  $[\text{Br}^-]$ . The calculated value for  $\gamma$  obtained at  $5 \times 10^{-2}$  M bromide ( $\gamma \approx 9 \times 10^{-2}$ ) appears to be consistent with the raw data presented in Fig. 2 of Katrib et al. (2001), from which (by comparison with uptake to NaI)  $\gamma$  is less than 0.1.

$$\gamma_b = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1 k')^{0.5}} \right\}^{-1} \quad \text{and} \quad k' = k_{\text{Br}}[\text{Br}^-] \quad \text{where} \quad [\text{Br}^-] \text{ is the bromide activity.}$$

This is however a simplification of the chemical interactions of molecular halogens in halide solutions, which are characterised by rapid equilibria involving several di- and tri-halogen species (Wang et al., 1999) and it is not obvious that a first-order analysis to obtain  $\gamma$  is appropriate. This may in part explain the observations of unusual kinetics of Katrib et al. (2001). Accurate modelling of the uptake and release of halogens from sea-water requires a more detailed treatment of the aqueous phase processes than possible here.

Studies of the uptake of BrCl to dry and frozen bromide containing surfaces indicate conversion to  $\text{Br}_2$ , as written above (Huff and Abbatt, 2000).

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

Bartlett, W. P. and Margerum, D. W.: *Env. Sci. Tech.*, 33, 3410, 1999.

Huff, A. K. and Abbatt, J. P. D.: *J. Phys. Chem.*, 104, 7284–7293, 2000.

5 Katrib, Y., Deiber, G., Schweitzer, F., Mirabel, P., and George, C.: *J. Aerosol Sci.*, 32, 893–911, 2001.

Wang, T. X., Kelley, M. D., Cooper J. N., and Margerum, D. W.: *Inorg. Chem.*, 33, 5872–5878, 1994.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

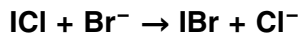
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.14



### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{\text{SS}}</math>, <math>\gamma_0</math></i>				
$\gamma = 2.6 \times 10^{-3}$	2 (Cl <sup>-</sup> )	274	Braban et al. (2007)	WWFT-EIMS (a)
$\gamma = 2.5 \times 10^{-3}$	2 (Br <sup>-</sup> )			
$\gamma = 3 \times 10^{-3}$	pure water 2 (Br <sup>-</sup> )			
$\gamma = (0.6 - 1.6) \times 10^{-3}$	15–209 Torr	278		
$\alpha = 0.01 - 1.0$	aq. NaBr aerosol			
$\gamma = 0.018 \pm 0.004$		274		AFT-CIMS (b)
$\gamma = (0.1 - 1) \times 10^{-3}$	Sea salt aerosol (5–87 % RH)	274		

### Comments

(a) The uptake of ICl ( $4\text{--}70 \times 10^{10}$  molecules  $\text{cm}^{-3}$ ) to aqueous surfaces was studied in the wetted-wall flow tube with MS detection at  $m/z = 144$ . Several combinations of dissolved concentrations of Cl<sup>-</sup> and Br<sup>-</sup> in solutions of different pH were used. Experiments were conducted at 274 and 293 K and at a range of flow tube pressures between 13 and 20 Torr, and at 278 K in a different system (see Holmes et al., 2001) over the pressure range 15–209 Torr. The uptake coefficient values were independent of composition or pH of the solution, but increased with decreasing pressure in the flow tube. Uptake is thus gas phase diffusion limited in all cases. Gas phase diffusion coefficients for ICl of  $D_{\text{H}_2\text{O}}^{\text{ICl}} = 57 \pm 6 \text{ cm}^2 \text{ s}^{-1} \text{ Torr}^{-1}$ , and  $D_{\text{He}}^{\text{ICl}} = 289 \pm 19 \text{ cm}^2 \text{ s}^{-1} \text{ Torr}^{-1}$  at 278 K, were obtained from the pressure de-

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



pendence of  $\gamma$ . Uptake of ICl into  $\text{Br}^-$  solutions produced IBr which was partially released to the gas phase.

(b) Atmospheric pressure aerosol flow tube with ICl ( $\approx 10^{13}$  molecule  $\text{cm}^{-3}$ ) measured by CIMS, together with products  $\text{Br}_2$  or IBr. Aerosols were generated from salt solutions in a constant output atomiser (TSI 3076), dried, and then equilibrated at controlled humidity by addition of “wet” carrier gas before entry into the flow tube. The size distribution was measured with a differential mobility analyser (DMA). The area weighted radius and surface area of the aerosol in the flow tube was 100–1000 nm, and  $(1-80) \times 10^{-3} \text{ cm}^2/\text{cm}^3$  respectively. Uptake coefficients were determined from the first order rate constants for ICl decay, corrected for wall loss, which were linearly dependent on surface area. Diffusion limitation was negligible for the size range used. Uptake of ICl onto NaBr aerosol ( $\text{pH} < 8$ ) led to IBr release with a yield of  $0.6 \pm 0.3$ . Lower uptake on sea salt attributed to depletion of  $\text{Br}^-$ . The uptake coefficient on NaBr aerosol was independent of RH in the range 10–70 % (and hence on  $[\text{Br}^-]$ ) and was independent of particle size. This indicates an accommodation controlled uptake process.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.016	298
$k^{\text{II}} (\text{M}^{-1} \text{s}^{-1}) (\text{Br})$	$>10^3$	298
$H/M \text{ atm}^{-1}$	110	298
$D_l/\text{cm}^2 \text{ s}^{-1}$	$5.0 \times 10^{-6}$	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298

32284

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred values

The recommended value of  $\alpha_b$  is based on the uptake coefficient into NaBr aerosol; the reaction of ICl with  $\text{Br}^-$  is fast and it is likely that reactive uptake of ICl into salt droplets is accommodation limited under atmospheric conditions. The dihalogen product IBr will partition rapidly to the gas phase. The uptake coefficient at lower  $[\text{Br}^-]$  is given in terms of the resistance model:

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1k^1)^{0.5}} \right\}^{-1} \text{ where } k^1(\text{s}^{-1}) = k^{II}[\text{Br}^-]_{\text{aq}}(\text{M})$$

At 298 K, a value of  $k^{II} = 5.0 \times 10^4$  can be used to estimate  $\gamma$ .

## References

- Braban, C. F., Adams, J. W., Rodriguez, D., Cox, R. A., Crowley, J. N., and Schuster, G.: Phys. Chem. Chem. Phys., 9, 3136–3148, 2007.
- Holmes, N. S., Adams, J. W., and Crowley, J. N.: Phys. Chem. Chem. Phys., 3, 1679–1687, 2001.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.15

### I<sub>2</sub> + Cl<sup>-</sup>/Br<sup>-</sup> → products

#### Experimental data

Parameter	[X]/M	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{ss}</math>, <math>\gamma_0</math></i>				
$\gamma = (2.0 \pm 0.3) \times 10^{-3}$	2 (Cl <sup>-</sup> ) 0.003 (Br <sup>-</sup> )	274	Braban et al. (2007)	WWFT-EIMS (a)
$\gamma_0 = (1.8 \pm 0.2) \times 10^{-3}$	pure water	274		

#### 5 Comments

(a) The uptake of I<sub>2</sub> at concentrations of  $(0.2-1) \times 10^{12}$  molecules  $\text{cm}^{-3}$  to aqueous surfaces was studied in the wetted-wall flow tube with MS detection at  $m/z = 144$ . Uptake was measured on pure water and with dissolved concentrations of Cl<sup>-</sup> (2 M) and Br<sup>-</sup> (0.003 M) to mimic sea water. Uptake was rapid and was gas phase diffusion limited in both cases, but on pure water the uptake coefficient reduced with increasing exposure length, indicating surface saturation. No gas phase products were observed.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.016	298
$H/M \text{ atm}^{-1}$	24	298
$D_1/\text{cm}^2 \text{ s}^{-1}$	$5.0 \times 10^{-6}$	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298

## Comments on preferred values

The uptake of IBr into aqueous halide is the same as into pure water, and is solubility limited, with no detectable reaction products in the gas phase. The recommendation is based on an analogy with ICl uptake on the same surfaces. The time dependent uptake coefficient is given by:

$$\gamma = \left\{ \frac{1}{\alpha} + \frac{\bar{c}\sqrt{\pi}}{8HRT} \sqrt{\frac{t}{D_1}} \right\}^{-1}$$

## References

Braban, C. F., Adams, J. W., Rodriguez, D., Cox, R. A., Crowley, J. N., and Schuster, G.: Phys. Chem. Chem. Phys., 9, 3136–3148, 2007.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A2.16



### Experimental data

Parameter	Aqueous Solution	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients, <math>\gamma</math></i> $\approx 4 \times 10^{-3} - 0.16$	$\text{Br}^- (1.25 \times 10^{-3} - 0.5 \text{M})$	263–293	Hu et al. (1995)	DT-MS (a)

### 5 Comments

(a) Flow tube at 8–27 mbar He (+H<sub>2</sub>O) with 120–250  $\mu\text{m}$  droplets. Cl<sub>2</sub> ( $5\text{--}100 \times 10^{12}$  molecule cm<sup>-3</sup>) was detected by mass spectrometer. The size of the uptake coefficient and its dependence on the bromide concentration could not be explained if only a bulk reaction was considered.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_b$	1	260–295
$\alpha_s$	1	260–295
$k_{(\text{Br}^-)}$ ( $\text{M}^{-1} \text{s}^{-1}$ )	$4.58 \times 10^{14} \exp(-2866/T)$	260–295
$H$ ( $\text{M atm}^{-1}$ )	$1.44 \times 10^{-6} \exp(-3181/T)$	260–295
$D_l$ ( $\text{cm}^2 \text{s}^{-1}$ )	$0.127 \exp(-2645/T)$	260–295
$R$ ( $\text{L atm mol}^{-1} \text{K}^{-1}$ )	$8.2057 \times 10^{-2}$	
$k_s \cdot K_{\text{LangC}} \cdot N_{\text{max}}$ ( $\text{L mol}^{-1} \text{cm}^{-1} \text{s}^{-1}$ )	$4.92 \times 10^{-4} \exp(4134/T)$	260–295
<i>Reliability</i>		
$\Delta \log(\gamma)$	$\pm 0.3$	260–295

## Comments on preferred values

The only data available for the uptake of  $\text{Cl}_2$  to bromide containing, aqueous solutions is that of Hu et al. (1995). They showed that the uptake of  $\text{Cl}_2$  could not be explained by bulk phase reaction alone (high values of measured  $\gamma$  and non-linear plot of inverse  $\gamma$  versus  $[\text{Br}^-]^{-0.5}$ ) and proposed that reaction within a surface-near volume was responsible. In this case, the measured uptake coefficient ( $\gamma$ ) is given by:

$$\frac{1}{\gamma} = \frac{1}{\alpha_s} + \frac{1}{\Gamma_s + \left(\frac{1}{\Gamma_{sb}} + \frac{1}{\Gamma_b}\right)^{-1}}$$

where:

$$\Gamma_s = \frac{4k_s a_{(\text{Br}^-)} K_{\text{LangC}} N_{\text{max}}}{\bar{c} (1 + K_{\text{LangC}} [\text{Cl}_2])} \text{ and}$$

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$$\Gamma_b = \frac{4HRT}{\bar{c}} \sqrt{D_1 k_{(\text{Br}^-)} a_{(\text{Br}^-)}}$$

and  $a_{(\text{Br}^-)}$  is the activity ( $\text{mol L}^{-1}$ ) of the dissolved bromide ion.

For efficient interfacial mass transport,  $\alpha_s = \alpha_b = 1$  the first equation simplifies to:

$$\frac{1}{\gamma} = 1 + \frac{1}{\Gamma_s + \Gamma_b}$$

When far from surface saturation (low  $K_{\text{LangC}}$  or low  $[\text{Cl}_2]$ ) the second expression also reduces to:

$$\Gamma_s = \frac{4k_s a_{(\text{Br}^-)} K_{\text{LangC}} N_{\text{max}}}{\bar{c}}$$

The units of the composite term  $k_s \cdot K_{\text{LangC}} \cdot N_{\text{max}}$  are  $\text{L mol}^{-1} \text{cm}^{-1} \text{s}^{-1}$  if the surface concentration of bromide ions is considered (as in this case) to be proportional to the bulk activity ( $\text{mol L}^{-1}$ ). The depth through which the surface reaction is considered to take place is integrated in the values of  $k_s \cdot K_{\text{LangC}} \cdot N_{\text{max}}$ , which were derived by fitting to experimental data.

The temperature dependence of  $\Gamma_s$  can arise from both the partition coefficient of  $\text{Cl}_2$  to the surface ( $K_{\text{LangC}}$ ) or via the surface rate coefficient ( $k_s$ ), though the former is expected to dominate. The parameterisation gives values of  $\gamma$  that are consistent with those presented by Hu et al. (1995) at all temperatures and bromide activities covered in their experiments. Diffusive effects were also taken into account when fitting to the experimental datasets by using

$$\frac{1}{\gamma} = 1 + \frac{1}{\Gamma_{\text{diff}}} + \frac{1}{\Gamma_s + \Gamma_b}$$

and effective gas-phase diffusion coefficients for  $\text{Cl}_2$  reported by Hu et al. (1995).

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The tabulated, temperature dependent expressions for aqueous phase rate coefficients, solubilities and aqueous phase diffusion coefficients listed for calculation of  $\Gamma_b$  were calculated from values given by Hu et al. (1995) at single temperatures. These values were taken in preference to other reports (e.g. of  $k_{(\text{Br}^-)}$  by Wang et al., 1994) in order to maintain an internally consistent set of parameters. Note that  $\gamma$  has almost no sensitivity to  $\alpha_s$  or  $\alpha_b$  as at high values of the uptake coefficient (when  $\alpha_s$  or  $\alpha_b$  may be rate limiting) the surface reaction dominates. At low uptake rates, the bulk-phase chemical reaction dominates and the effect of  $\alpha$  is diminished.

Studies of the uptake of  $\text{Cl}_2$  to dry and frozen bromide containing salts surfaces have identified  $\text{Br}_2$  and  $\text{BrCl}$  as the main products released to the gas-phase (Berko et al., 1991; Mochida et al., 1998; Huff and Abbatt, 2000; Adams et al., 2002; Santschi and Rossi, 2004).  $\text{BrCl}$  is formed initially and is converted to  $\text{Br}_2$  by reaction with  $\text{Br}^-$ . The same products are expected for the reaction on an aqueous surface, with the ratio of  $\text{BrCl}$  to  $\text{Br}_2$  defined by aqueous phase equilibria:



## References

- Adams, J. W., Holmes, N. S., and Crowley, J. N.: Atmos. Chem. Phys., 2, 79–91, 2002, <http://www.atmos-chem-phys.net/2/79/2002/>.
- Berko, H. N., McCaslin, P. C., and Finlayson-Pitts, B. J.: J. Phys. Chem., 95, 6951–6958, 1991.
- Hu, J. H., Shi, Q., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 99, 8768–8776, 1995.
- Huff, A. K. and Abbatt, J. P. D.: J. Phys. Chem., 104, 7284–7293, 2000.

Mochida, M., Hirokawa, J., Kajii, Y., and Akimoto, H.: Geophys. Res. Lett., 25, 3927–3930, 1998.

Santschi, C. and Rossi, M. J.: Phys. Chem. Chem. Phys., 6, 3447–3460, 2004.

Wang, T. X., Kelley, M. D., Cooper J. N., and Margerum, D. W.: Inorg. Chem., 33, 5872,  
5 1994.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Appendix A3

### Uptake on other aqueous electrolytes

#### VI.A3.5

#### $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$ (ammonium sulphate aerosols)

##### 5 Experimental data

Parameter	RH/%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{\text{SS}}</math>, <math>\gamma_0</math></i>				
0.043 ± 0.005	60	293	Mozurkewitch and Calvert (1988)	AFT-CL (a)
0.017 ± 0.002	93.5	297	Hu and Abbatt (1997)	AFT-CIMS/OPC (b)
0.023 ± 0.004	83.0			
0.053 ± 0.006	68.5			
0.044 ± 0.008	50.0			
0.00094 ± 0.00059	8	295 ± 1	Kane et al. (2001)	AFT-CIMS/SMPS (e)
0.0038 ± 0.0002	37			
0.012 ± 0.0003	55			
0.020 ± 0.0023	72			
0.033 ± 0.0013	86			
0.042 ± 0.0022	92			
0.0182 ± 0.0034	62.1	295 ± 2	Folkers et al. (2003)	FTIR/SMPS/APC (d)
0.0025–0.0063 (+ natural organic)	55–83			
0.0024 ± 0.0006	20	298	Hallquist et al. (2003)	AFT-CL/SMPS (e)
0.0047 ± 0.0013	35			
0.0141 ± 0.0040	50			
0.0149 ± 0.0052	70			
0.0164 ± 0.0048	80			
0.025 ± 0.0013	50	288		
0.0141 ± 0.0040	50	298		
0.0040 ± 0.008	50	308		
0.0057 ± 0.0013	25		Badger et al. (2006)	AFT-CL/SMPS (f)
0.0150 ± 0.0040	50			
0.0160 ± 0.0052	60			
0.0190 ± 0.0048	70			
0.00063 ± 0.00004 (6% humic acid)	25			
0.0040 ± 0.0010 (6% humic acid)	50			
0.0083 ± 0.0006 (6% humic acid)	70			
0.035 ± 0.002	50	263	Griffiths and Cox (2009)	AFT-CL/SMPS (g)
0.028 ± 0.002	50	275		
0.020 ± 0.005	50	283		
0.010 ± 0.002	50	293		
0.005 ± 0.002	50	303		

32293

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments

(a) Atmospheric pressure aerosol flow tube with  $\text{N}_2\text{O}_5$  ( $\approx 10^{13}$  molecule  $\text{cm}^{-3}$ ) measured by a modified chemiluminescence method, via thermal dissociation to  $\text{NO}_3$  and titration with  $\text{NO}$ , which was detected. Aerosols generated in a constant output atomiser, dried, and size selected with a differential mobility analyser (DMA) coupled to a condensation particle counter (CPC), to count the particles. The monodisperse aerosol was then equilibrated at controlled humidity before entry into the flow tube. The size and surface area of the aerosol in the flow tube was calculated from the deliquescence properties of the aerosol, determined in separate tandem DMA experiments. The typical diameter,  $d_{\text{mean}}$  was 0.08–0.2  $\mu\text{m}$ , with surface area density of  $1\text{--}5 \times 10^{-5}$   $\text{cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss, which were linearly dependent on surface area. Diffusion limitation was negligible for the size range used.

(b) Atmospheric pressure aerosol flow tube with detection of  $\text{N}_2\text{O}_5$  ( $7 \times 10^{12}$  molecule  $\text{cm}^{-3}$ ) by CIMS using  $\text{I}^-$  reagent ion. The aerosols were generated in an ultrasonic nebuliser and were equilibrated with the ambient humidity before entry into the flow tube. The size distribution, measured with an optical particle counter (OPC), was used to calculate the surface area of the aerosol in the flow tube. The peak in the surface area distribution,  $d_{\text{max}}$  was between 2 and 4  $\mu\text{m}$ . The counter was calibrated by collection of aerosol of known composition in a aqueous trap (assumed 100% efficient) and measurement of the electrical conductivity of the trapped electrolyte. Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss (Brown correction), and for diffusion limitation to the particle surface assuming an average diameter,  $d_{\text{max}}$  of the non-monodisperse aerosol. Cited errors on  $\gamma$  is  $\pm 1\sigma$  precision; the estimated potential systematic error arising mainly from measurement of the SA was  $\pm 25\%$ .

32294

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

(c) Atmospheric pressure aerosol flow tube with detection of  $\text{N}_2\text{O}_5$  ( $(5$  to  $200) \times 10^{12}$  molecule  $\text{cm}^{-3}$ ) by CIMS using  $\text{I}^-$  reagent ion. Aerosols generated in a constant output atomiser and were dried and equilibrated at controlled humidity before entry into the flow tube. The size distribution was measured with a scanning mobility particle sizer (SMPS) coupled to a condensation particle counter (CPC). The surface area of the aerosol in the flow tube was calculated from the observed size distribution assuming the particles are spherical. The mean diameter,  $d_{\text{mean}}$  was  $0.12 \mu\text{m}$ , with surface area density of  $0.0016 \text{ cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss, and for diffusion limitation to the particle surface using the size resolved Knudsen number for the non-monodisperse aerosol. Although the values of  $\gamma$  showed some scatter (only selected values from the 17 data points reported are cited), a large increase with RH throughout the range 8–92 % was observed, which was fitted by the expression:  $= 2.79 \times 10^{-4} + 1.3 \times 10^{-4} \times (\text{RH}) - 3.43 \times 10^{-6} \times (\text{RH})^2 + 7.52 \times 10^{-8} \times (\text{RH})^3$ . No dependence on the uptake coefficient on  $[\text{N}_2\text{O}_5]$  was observed.

(d) Static aerosol chamber with inert Teflon walls at ambient pressure and temperature.  $\text{N}_2\text{O}_5$  ( $\approx 10^{13}$  molecule  $\text{cm}^{-3}$ ) produced in situ by reaction of  $\text{NO}_2$  with ozone. Gas phase species measured by FTIR and UV spectroscopy. Polydisperse aerosol ( $d_{\text{a}} = 20 \text{ nm} - 5 \mu\text{m}$ ) generated by spraying dilute ammonium sulphate solutions, with size distribution measured by SMPS and aerodynamic particle size for particle diameters  $< 700 \text{ nm}$ . Uptake coefficients determined by fitting experimental time dependence of species with a numerical model of chemistry and integrated aerosol surface area.  ~~$(\text{NH}_4)_2\text{SO}_4$  coated by organic oxidation products of ambient VOC with added ozone.~~

(e) Atmospheric pressure aerosol flow tube with  $\text{N}_2\text{O}_5$  ( $(7.5$ – $125) \times 10^{13}$  molecule  $\text{cm}^{-3}$ ) measured via thermal dissociation to  $\text{NO}_3$  and titration with  $\text{NO}$ , which was detected by  $\text{O}_3$ -chemiluminescence. Deliquesced

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



aerosols were generated in a constant output atomiser, and conditioned by equilibration with an additional excess flow of controlled RH. Size distribution determined with a differential mobility analyser (DMA) coupled to Faraday cup electrometer to count the particles. The typical peak diameter,  $d_{\max}$  was 200 nm, with surface area density of  $(7 \text{ to } 37) \times 10^{-5} \text{ cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, after correction for wall loss and diffusion effects. The measured uptake coefficients were independent of  $[\text{N}_2\text{O}_5]$  or on RH above 50%, but  $\gamma$  declined with decreasing RH below 50%. This falloff was attributed to decreased liquid water amounts in the aerosol at low RH, supported by even lower  $\gamma$  values for “dry” aerosols below the efflorescence RH.

(f) Same experimental details as (e); study focussed on effect of mixed AS-humic acid aerosols on  $\gamma$ , the aerosol surface area in range  $0.12\text{--}0.50 \text{ m}^2 \text{ m}^{-3}$  with an area weighted particle diameter of  $\sim 300 \text{ nm}$ .

(g) Same experimental method as in (e), except that the halocarbon-wax-coated flow tube and the DMA were contained in a modified low temperature chamber so that aerosol  $S_a$  could be measured at the same temperature as the gas uptake region, leading to improved accuracy of uptake measurements.  $\text{N}_2\text{O}_5$  mixing ratios were around 300 ppbv and aerosol surface area density varied in the range  $0.02 \text{ to } 0.5 \text{ cm}^2 \text{ cm}^{-3}$ . The aerosol volume to surface ratio varied between  $3 \text{ and } 4 \times 10^{-9} \text{ m}$ .



## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.03	298
$\gamma$	$0.244 - 7.9 \times 10^{-4} T(K)$	260–305
$k^I$ ( $s^{-1}$ )	$k^{II} \times [H_2O]_{aq}$ (M)	
$k^{II}$ ( $M^{-1} s^{-1}$ )	$1.0 \times 10^5$	298
<i>Reliability</i>		
$\Delta \log(\alpha)$	$\pm 0.3$	298
$\Delta \log(\gamma)$	$\pm 0.3$ at 50 % RH	260–305

## Comments on preferred values

Uptake studies conducted on aqueous  $(NH_4)_2SO_4$  aerosols as a function of relative humidity show broad agreement in uptake coefficient values, but there are differences in detail. Only data for deliquesced aerosols are presented; at RH below the efflorescence RH for  $(NH_4)_2SO_4$  aerosol ( $\sim 37\%$ ) the particles may crystallise and under these conditions the uptake coefficient reduces dramatically to  $\gamma < 0.001$  due to reduced liquid volume. In most studies  $\gamma$  shows little dependence on humidity above  $\sim 50\%$  RH with  $\gamma = 0.015 - 0.02$  at 298 K, but below 50 % RH  $\gamma$  declines significantly as RH reduces. Exceptions are the study of Hu and Abbatt (1997) in which  $\gamma$  increased as RH reduces below 95 % and Kane et al. (2001) in which  $\gamma$  continued to increase monotonically above 50 % RH. These opposite trends have not been reproduced in any of the other studies and suggest some source of systematic error. Hu and Abbatt used larger particles and for size calibration a trapping method was used, which could have underestimated the aerosol surface area of smaller particles at lower RH. Kane et al. give insufficient experimental detail of the sizing method using SMPS to evaluate the reliability of SA measurements at high RH. We therefore recommend an expression giving an approximately constant value of  $\gamma$  above 50 % RH at 298 K.

32297

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

The value of  $\gamma$  above 50 % RH is close to that on pure water drops ( $\sim 0.015$  at 298 K), and only slightly lower than that on  $\text{H}_2\text{SO}_4$  droplets ( $\gamma = 0.036 \pm 0.008$ , Hallquist et al., 2000). The uptake leads to hydrolysis of  $\text{N}_2\text{O}_5$  and formation of  $\text{HNO}_3$  which partially transfers to the gas phase. The lack of a dependence on the water content of deliquesced droplets at relative humidity  $>50\%$ , suggests that uptake is controlled by a surface process, either mass accommodation or surface reaction. At lower RH,  $\gamma$  declines, but the data show large scatter, which may be due to the presence of a mixed size population of solid and liquid particles, as a result of efflorescence. These data suggest that uptake at lower RH is limited by the rate of hydrolysis of  $\text{N}_2\text{O}_5$  in bulk liquid phase, hence depends on  $[\text{H}_2\text{O}]_{\text{aq}}$  and, for small particles, on particle volume.

The recommended expression for RH dependence uses a size dependent resistance-model formulation:

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1 k^{\text{I}})^{0.5}} \left[ \coth\left(\frac{r}{l}\right) - \left(\frac{l}{r}\right) \right] \right\}^{-1}$$

The recommended value of  $\alpha_b = 0.03$  is based on  $\gamma$  observed for uptake on water, malonic acid (Thornton et al., 2004) and  $\text{H}_2\text{SO}_4$  droplets. The recommended liquid phase rate constant  $k^{\text{I}} (= k^{\text{II}} \times [\text{H}_2\text{O}]_{\text{aq}})$ , calculated using  $k^{\text{II}} = 1 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ , which is intermediate between the values derived by Thornton et al. (2003) from uptake of  $\text{N}_2\text{O}_5$  on malonic acid aerosol ( $2.5 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ ) and by Mentel et al. (1999) from uptake on  $\text{NaNO}_3$  aerosols ( $1.5 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ ). The RH dependence of  $\gamma$  can be calculated using water mass fractions taken from the AIM database and using  $D_1 = 1 \times 10^{-5} \text{ cm}^2 \text{ s}^{-1}$  and  $2 \text{ M atm}^{-1}$ , values used in the analyses for  $k^{\text{II}}$ . The expression fits the  $>50\%$  RH data well but overestimates uptake rates at low RH. The reacto-diffusive parameter ( $l = [D_1/k^{\text{II}}[\text{H}_2\text{O}]]^{0.5}$ ) predicts a significant size dependence of  $\gamma$  for  $r < 100 \text{ nm}$ .

The temperature dependence of  $\text{N}_2\text{O}_5$  uptake reported earlier by Hallquist et al. (2003) has recently been confirmed by Griffiths and Cox (2009), using improved determination of aerosol size and  $S_a$ . There is a substantial negative temperature dependence of  $\gamma$  at constant RH of 50 %. This is consistent with accommodation con-

trolled uptake but the effect is much larger than the  $T$  dependence of  $\alpha$  deduced from the results obtained for  $\text{H}_2\text{SO}_4$  droplets at 50 % RH. The recommended expression is an empirical linear fit to the data over the stated temperature range and should be used with caution outside this range.

## 5 References

Badger, C. L., Griffiths, P. T., George, I., Abbatt, J. P. D., and Cox, R. A.: J. Phys. Chem. A, 110, 6986, 2006.

Griffiths, P. T. and Cox, R. A.: Temperature dependence of heterogeneous uptake of  $\text{N}_2\text{O}_5$  by ammonium sulfate aerosol, Atmos. Sci. Lett., 10, 159–163, 2009.

10 Hallquist, M., Stewart, D. J., Baker, J., and Cox, R. A.: J. Phys. Chem., 104A, 3984, 2000.

Hallquist, M., Stewart, D. J., Stephenson, S. K., and Cox, R. A.: Phys. Chem. Chem. Phys., 5, 3453, 2003.

15 Folkers, M., Mentel, Th. F., and Wahner, A.: Geophys. Res. Lett., 30, art. no. 1644, doi:10.1029/2003GL017168, 2003.

Hu, J. H. and Abbatt, J. P. D.: J. Phys. Chem., A101, 871, 1997.

Kane, S. M., Caloz, F., and Leu, M.-T.: J. Phys. Chem., 105A, 6465, 2001.

Mentel, T. F., Sohn, M., and Wahner, A.: Phys. Chem. Chem. Phys., 1, 5451, 1999.

Mozurkewich, M. and Calvert, J. G.: J. Geophys. Res., 93, 15889, 1988.

20 Thornton, J. A., Braban, C. F. and Abbatt, J. P. D.: Phys. Chem. Chem. Phys., 5, 4593–4603, 2003.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Experimental data

Parameter	RH/%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{ss}</math>, <math>\gamma_0</math></i>				
0.020 ± 0.0036	13	293	Mozurkewitch and Calvert (1988)	AFT-CL (a)
0.0218 ± 0.0055	35			
0.050 ± 0.0055	45			
0.0482 ± 0.0055	62.5			
0.0382 ± 0.0055	77.5			
0.0015 ± 0.00042	13	295 ± 1	Kane et al. (2001)	AFT- CIMS/SMPS
0.0031 ± 0.0023	32			(b)
0.012 ± 0.0032	38			
0.018 ± 0.0027	48			
0.022 ± 0.0019	63			
0.035 ± 0.011	74			
0.046 ± 0.012	89			
0.069 ± 0.0095	99			
0.020 ± 0.0020	60	295 ± 2	Folkers et al. (2003)	FTIR/SMPS/APC (c)
0.0190 ± 0.0040	67			
0.0187 ± 0.0043	79.7			
0.0025–0.0063 (+ natural organic)	55–83			
0.0031 ± 0.0017	20	298	Hallquist et al. (2003)	AFT-CL/SMPS (d)
0.0041 ± 0.0020	35			
0.018 ± 0.008	50			
0.023 ± 0.0013	70			
0.015 ± 0.006	80			
0.030 ± 0.0022	50	263		
0.0140 ± 0.0062	50	268		
0.0129 ± 0.0045	50	273		
0.036 ± 0.0018	50	283		
0.0037 ± 0.008	50	308		
0.0133 ± 0.0009	30	298	Badger (2006)	AFT-CL/SMPS (e)
0.015 ± 0.002	50			
0.0196 ± 0.0009	60			
0.024 ± 0.001	70			
0.036 ± 0.009	50	263	Griffiths and Cox (2009)	AFT-CL /SMPS (f)
0.005 ± 0.002	50	268		
0.014 ± 0.006	50	278		
0.012 ± 0.001	50	283		
0.011 ± 0.001	50	288		
0.016 ± 0.003	50	293		
0.003 ± 0.001	50	303		
0.028 ± 0.006	60	298	Bertram and Thornton (2009)	AFT-CIMS/SMPS (g)

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Parameter	RH/%	Temp./K	Reference	Technique/Comments
0.005 ± 0.001	5			
0.013 ± 0.002	10			
0.016 ± 0.003	15			
0.024 ± 0.004	25			
0.028 ± 0.005	40			
0.025 ± 0.004	50			
0.026 ± 0.005	50			
0.033 ± 0.007	65			
0.029 ± 0.007	75			



## Comments

(a) Atmospheric pressure aerosol flow tube with  $\text{N}_2\text{O}_5$  ( $\approx 10^{13}$  molecule  $\text{cm}^{-3}$ ) measured by a modified chemiluminescence method, via thermal dissociation to  $\text{NO}_3$  and titration with  $\text{NO}$ , which was detected. Aerosols generated in a constant output atomiser, dried, and size selected with a differential mobility analyser (DMA) coupled to a condensation particle counter (CPC), to count the particles. The monodisperse aerosol was then equilibrated at controlled humidity before entry into the flow tube. The size and surface area of the aerosol in the flow tube was calculated from the deliquescence properties of the aerosol, determined in separate tandem DMA experiments. The typical diameter,  $d_{\text{mean}}$  was 0.08–0.2  $\mu\text{m}$ , with surface area density of  $1\text{--}5 \times 10^{-5}$   $\text{cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss, which were linearly dependent on surface area. Diffusion limitation was negligible for the size range used.

(b) Atmospheric pressure aerosol flow tube with detection of  $\text{N}_2\text{O}_5$  [(5 to 200)  $\times 10^{12}$  molecule  $\text{cm}^{-3}$ ] by CIMS using  $\text{I}^-$  reagent ion. Aerosols generated in a constant output atomiser and were dried and equilibrated at controlled humidity before entry into the flow tube. The size distribution was measured with a scanning mobility particle sizer (SMPS) coupled to a condensation particle counter (CPC). The surface area of the aerosol in the flow tube was calculated from the observed size distribution assuming the particles are spher-

32301

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



ical. The mean diameter,  $d_{\text{mean}}$  was  $0.12 \mu\text{m}$ , with surface area density of  $0.0016 \text{ cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss, and for diffusion limitation to the particle surface using the size resolved Knudsen number for the non-monodisperse aerosol. Although the values of  $\gamma$  showed some scatter (only selected values from the 17 data points reported are cited), a large increase with RH throughout the range 13–99% was observed, which was fitted by the expression:  $=2.79 \times 10^{-4} + 1.3 \times 10^{-4} \times (\text{RH}) - 3.43 \times 10^{-6} \times (\text{RH})^2 + 7.52 \times 10^{-8} \times (\text{RH})^3$ . No dependence on the uptake coefficient on  $[\text{N}_2\text{O}_5]$  was observed.

(c) Static aerosol chamber with inert teflon walls at ambient pressure and temperature.  $\text{N}_2\text{O}_5$  ( $\approx 10^{13} \text{ molecule cm}^{-3}$ ) produced in situ by reaction of  $\text{NO}_2$  with ozone. Gas phase species measured by FTIR and UV spectroscopy. Polydisperse aerosol ( $\text{dia.} = 20 \text{ nm} - 5 \mu\text{m}$ ) generated by spraying dilute ammonium bisulphate solutions, with size distribution measured by SMPS for particle diameters  $< 700 \text{ nm}$  and for larger diameters by aerodynamic particle sizer. Uptake coefficients determined by fitting experimental time dependence of species with a numerical model of chemistry and integrated aerosol surface area. ~~To investigate effect of organic film on uptake,  $\text{NH}_4\text{HSO}_4$  coated by organic oxidation products of ambient VOC with added ozone. A noticeable reduction in  $\gamma$  was observed in this case.~~

(d) Atmospheric pressure aerosol flow tube with  $\text{N}_2\text{O}_5$  ( $7.5 - 125 \times 10^{13} \text{ molecule cm}^{-3}$ ) measured by ~~via~~ thermal dissociation to  $\text{NO}_3$  and titration with  $\text{NO}$ , which was detected by  $\text{O}_3$ -chemiluminescence. Deliquesced aerosols were generated in a constant output atomiser, and conditioned by equilibration with an additional excess flow of controlled RH. Size distribution determined with a differential mobility analyser (DMA) coupled to Faraday cup electrometer to count the particles. The typical peak diameter,  $d_{\text{max}}$  was  $200 \text{ nm}$ , with surface area density of  $(7 - 37) \times 10^{-5} \text{ cm}^2/\text{cm}^3$ . Uptake coefficients were

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, after correction for wall loss and diffusion effects. The measured uptake coefficients were independent of  $[\text{N}_2\text{O}_5]$  or on RH above 50 %, but  $\gamma$  declined with decreasing RH below 50 %. This falloff was attributed to decreased liquid water amounts in the deliquesced aerosol at low RH.

- (e) Same experimental details as (d); the aerosol surface area in range  $0.06\text{--}0.27\text{ m}^2\text{ m}^{-3}$  with an area weighted particle diameter of 260 nm.
- (f) Same experimental method as in (d), except that the halocarbon-wax-coated flow tube and the DMA were contained in a modified low temperature chamber so that aerosol  $S_a$  could be measured at the same temperature as the gas uptake region, leading to improved accuracy of uptake measurements.  $\text{N}_2\text{O}_5$  mixing ratios were around 300 ppbv and aerosol surface area density varied in the range  $0.02$  to  $0.5\text{ cm}^2\text{ cm}^{-3}$ . The aerosol volume to surface ratio varied between 3 and  $4 \times 10^{-9}\text{ m}$ .
- (g) Measurements of the reactive uptake coefficient for  $\text{N}_2\text{O}_5$  on  $\text{NH}_4\text{HSO}_4$  aerosols in an aerosol flow tube at 1 bar and at room temperature, using both time dependence and surface area dependence (at fixed time) of  $[\text{N}_2\text{O}_5]$  decay.  $\gamma$  was determined as a function of relative humidity (RH 5 to 75 %), and aerosol  $[\text{NO}_3^-]$ . The low detection threshold for  $\text{N}_2\text{O}_5$  (<5 ppt) afforded by CIMS minimised the impact of accumulated  $[\text{NO}_3^-]$  which can reduce uptake rates. A strong increase in  $\gamma$  with RH up to 50 % was observed, above which it remained constant.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	0.04	298
$k^I$ (s <sup>-1</sup> )	$k^{IIx} [\text{H}_2\text{O}]_{\text{aq}}$ (M)	
$k^{II}$ (M <sup>-1</sup> s <sup>-1</sup> )	$1.5 \times 10^5$	298
$D_I/\text{cm}^2 \text{ s}^{-1}$	$1 \times 10^{-5}$	298
$H/M \text{ atm}^{-1}$	2	298
<i>Reliability</i>		
$\Delta \log(\alpha)$	$\pm 0.3$	298
$\Delta \log(\gamma)$ at 50 % RH	$\pm 0.3$	

## Comments on preferred values

All studies used deliquesced or supersaturated aqueous  $\text{NH}_4\text{HSO}_4$  aerosols, with uptake coefficients measured as a function of RH.  $\gamma$  was generally seen to increase with RH over the range 5–90 % but there is poor agreement between the different studies. The early study of Mozurkewitch and Calvert (1988) reported much larger uptake coefficients than from more recent work even after the 15 % downward revision of their reported  $\gamma$  values suggested by Fried et al. (1995). The results of Kane et al. (2001) are very scattered at RH < 50 % and show a large increase in  $\gamma$  above 70 % RH which has not been reproduced in any of the other studies. The former authors suggest that evaporation of  $\text{NH}_3$  could lead to acidification of the particles, forming  $\text{H}_2\text{SO}_4$  which is more reactive. Kane et al. give insufficient experimental detail of the sizing method using SMPS to evaluate the reliability of SA measurements at high RH. The uptake leads to hydrolysis of  $\text{N}_2\text{O}_5$  and formation of  $\text{HNO}_3$  which partitions between the gas phase and the  $\text{NH}_4\text{HSO}_4$  aerosols as aqueous nitrate.

All other studies show little dependence of  $\gamma$  on RH above ~50 % RH. Most of the studies give a value of  $\gamma \approx 0.02$ , but the results of Bertram and Thornton (2009) are

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





distinctly higher throughout the range. These authors used much lower  $[\text{N}_2\text{O}_5]$  than all other studies and, based on their quantitative analysis of their data for uptake in the presence of  $\text{NaNO}_3$ , they believe their uptake coefficients to be free of nitrate suppression.

The  $\gamma$  value at high RH and  $\sim 295$  K is similar to that on pure water drops ( $\sim 0.015$  at 298 K), and only slightly lower than that on  $\text{H}_2\text{SO}_4$  droplets ( $\gamma = 0.036 \pm 0.008$ , Hallquist et al., 2000). The lack of a dependence on the water content of deliquesced droplets at  $\text{RH} > 50\%$ , suggests that uptake is controlled by a surface process, either mass accommodation or surface reaction. At lower RH,  $\gamma$  declines. Since  $\text{NH}_4\text{HSO}_4$  aerosols do not effloresce easily this result suggests that uptake at lower RH is limited by the rate of hydrolysis of  $\text{N}_2\text{O}_5$  in the bulk liquid phase, and hence on  $[\text{H}_2\text{O}]_{\text{aq}}$ . At low RH the hydrolysis rate slows and in addition the size of the particles is smaller, leading to volume-limited uptake coefficients. Bertram and Thornton favour this model over the whole RH range, and attribute the constancy of  $\gamma$  above  $\sim 50\%$  RH to a reduction in  $[\text{H}_2\text{O}]_{\text{aq}}$  dependence of the dissociation of  $[\text{N}_2\text{O}_5]_{\text{aq}}$ . It is probable that the inhibition by nitrate is responsible for the large mismatch of the data of Bertram and Thornton at low  $[\text{N}_2\text{O}_5]$  and the other studies at  $\text{RH} < 40\%$ . The differences above  $50\%$  RH are less and may still be accounted for by estimates which have been made for the accumulated  $[\text{NO}_3^-]$  under the experimental conditions, and volume-limited uptake assumed by Bertram and Thornton is not indicated over the whole range of studies (see Griffiths et al., 2009).

We prefer to recommend an expression for the uptake on pure  $\text{NH}_4\text{HSO}_4$  aerosols using a resistance-model formulation which gives an approximately constant value of  $\gamma$  above  $70\%$  RH at 298 K.

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1k^1)^{0.5}} \right\}^{-1}$$

The recommended value of  $\alpha_b = 0.04$  at 298 K is based on  $\gamma$  observed for uptake on malonic acid (Thornton et al., 2004) and  $\text{H}_2\text{SO}_4$  droplets (Hallquist et al., 2003).  $k^1$  is

32305

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



calculated using the recommended liquid phase rate constant for  $\text{N}_2\text{O}_5 + [\text{H}_2\text{O}]_{\text{aq}}$ :  $k^{\text{II}} = 1.5 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ , which is based on the values derived Mentel et al. (1999) from uptake on  $\text{NaNO}_3$  aerosols ( $1.5 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ ), and on mixed  $\text{NH}_4\text{HSO}_4$ /dicarboxylic acid aerosols by Griffiths et al. (2009). The RH dependence of  $\gamma$  can be calculated using water mass fractions taken from the AIM database. At >50 % RH this expression provides a good compromise between the Bertram and Thornton data and the main body of the experimental data. Observed uptake rates at low RH are underestimated with this formula, due to the nitrate effect and volume limited uptake, as discussed above. The size dependence of  $\gamma$  can be evaluated by Eq. (23) in the introduction for heterogeneous datasheets using the reacto-diffusive parameter  $q (= \sqrt{D_1/k^{\text{II}}[\text{H}_2\text{O}]_{\text{aq}}})^{0.5}$ ; the recommended values predict a significant size dependence of  $\gamma$  for  $r \leq 100 \text{ nm}$ .

Hallquist et al. (2003) reported a temperature dependence of  $\text{N}_2\text{O}_5$  uptake on  $\text{NH}_4\text{HSO}_4$  aerosols at constant RH of 50 %. Their data showed considerable scatter but generally  $\gamma$  increased as  $T$  decreased in the range 263–308 K, but with a region of apparently zero temperature dependence between 270–290 K. The recent results of Griffiths and Cox (2009), which had reduced uncertainties in determination of the actual  $S_a$  of the aerosol, confirm this complex trend. However the origin of this behaviour is not consistent with our general parameterisation for uptake on aqueous liquids and we do not offer a recommendation at this stage.

## References

- Badger, C. L.: PhD thesis, University of Cambridge, 2006.  
Bertram, T. H. and Thornton, J. A.: Atmos. Chem. Phys., 9, 8351–8363, 2009, <http://www.atmos-chem-phys.net/9/8351/2009/>.  
Hallquist, M., Stewart, D. J., Baker, J., and Cox, R. A.: J. Phys. Chem., 104A, 3984, 2000.  
Hallquist, M., Stewart, D. J., Stephenson, S. K., and Cox, R. A.: Phys. Chem. Chem. Phys., 5, 3453, 2003.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Fried, A., Henry, B. E., Calvert, J. G., and Mozurkewich, M.: J. Geophys. Res., 99, 3517, 1994.
- Folkers, M., Mentel, Th. F., and Wahner, A.: Geophys. Res. Lett., 30, art. no. 1644, doi:10.1029/2003GL017168, 2003.
- 5 Griffiths, P. T. and Cox, R. A.: Atmos. Sci. Lett., 10, 159–163, 2009.
- Griffiths, P. T., Badger, C. L., Cox, R. A., Folkers, M., Henk, H. H., and Mentel, T. F.: J. Phys. Chem. A, 113, 5082–5090, 2009.
- Kane, S. M., Caloz, F., and Leu, M.-T.: J. Phys. Chem., 105A, 6465, 2001.
- Mentel, T. F., Sohn, M., and Wahner, A.: Phys. Chem. Chem. Phys., 1, 5451, 1999.
- 10 Mozurkewich, M. and Calvert, J. G.: J. Geophys. Res., 93, 15889–15896, 1988.
- Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: Phys. Chem. Chem. Phys., 5, 4593–4603, 2003.

## VI.A3.7

### $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$ ( $\text{Na}^+/\text{NH}_4^+$ nitrate aerosols)

#### Experimental data

Parameter	$[\text{NO}_3]/[\text{NH}_4\text{HSO}_4]$ M/wt%	RH/%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{\text{SS}}</math>, <math>\gamma_0</math></i>					
0.0018 ± 0.0020		48	295 ± 2	Wahner et al. (1999)	FTIR/SMPS/APC (a)
0.0032 ± 0.0006		62			
0.023 ± 0.012		88			
0.0040 ( $\text{NH}_4\text{NO}_3$ )		54			
0.0065 ( $\text{NH}_4\text{NO}_3$ )		60			
0.0090 ( $\text{NH}_4\text{NO}_3$ )		70			
0.016 ( $\text{NH}_4\text{NO}_3$ )		79			
0.00048 ± 0.00011		20	298	Hallquist et al. (2003)	AFT-CL/SMPS (b)
0.00049 ± 0.00023		30			
0.00065 ± 0.00025		50			
0.0010 ± 0.00022		70			
0.00063 ± 0.00064		50	263		
0.00151 ± 0.00096		50	278		
0.00031 ± 0.00096		50	288		
0.025 ± 0.005	1.1 (0.075)	60	298	Bertram and Thornton (2009)	AFT-CIMS/SMPS (c)
0.018 ± 0.004	2.0 (0.065)	60			
0.012 ± 0.003	3.3 (0.051)	60			
0.009 ± 0.003	4.4 (0.041)	60			
0.006 ± 0.003	6.6 (0.028)	60			
0.001 ± 0.003	9.2 (0.0)	55			

#### 5 Comments

(a) Static aerosol chamber with inert teflon walls at ambient pressure and temperature.  $\text{N}_2\text{O}_5$  ( $\approx 10^{13}$  molecule  $\text{cm}^{-3}$ ) produced in situ by reaction of  $\text{NO}_2$  with ozone. Gas phase species measured by FTIR and UV spectroscopy. Polydisperse aerosol ( $d_{\text{a}} = 20 \text{ nm} - 5 \mu\text{m}$ ) generated by spraying dilute ammonium bisulfate solutions, with size distribution measured by SMPS for particle diameters

10

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



<700 nm and aerodynamic particle sizer for larger diameters. The mean volume/surface ratio of the aerosol ( $=r_{av}/3$ ) was typically  $1.1 \times 10^{-5} \text{ cm}^3/\text{cm}^2$ . Uptake coefficients determined by fitting experimental time dependence of species with a numerical model of chemistry and integrated aerosol surface area. At high RH  $\gamma$  for uptake on  $\text{NaNO}_3$  aerosols was similar to that observed for sulphates and NaCl, but as RH declined  $\gamma$  declined to values a factor of 10 lower at  $\sim 50\%$  RH. A similar trend was observed for uptake on deliquesced  $\text{NH}_4\text{NO}_3$  aerosols.

(b) Atmospheric pressure aerosol flow tube with  $\text{N}_2\text{O}_5$  ( $7.5\text{--}125 \times 10^{13} \text{ molecule cm}^{-3}$ ) measured by via thermal dissociation to  $\text{NO}_3$  and titration with  $\text{NO}$ , which was detected by  $\text{O}_3$ -chemiluminescence. Deliquesced aerosols were generated in a constant output atomiser, and conditioned by equilibration with an additional excess flow of controlled RH. Size distribution determined with a differential mobility analyser (DMA) coupled to Faraday cup electrometer to count the particles. The typical peak diameter,  $d_{max}$  was 200 nm, with surface area density of  $(1.9\text{--}7.01) \times 10^{-3} \text{ cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, after correction for wall loss and diffusion effects. The measured uptake coefficients were approximately a factor of 10–20 lower than on sulphate aerosols at comparable RH and showed a weak increase with RH over the range 20–80%. No significant temperature dependence of  $\gamma$  was observed over the range 263–298 K.

(c) Measurements of the reactive uptake coefficient for  $\text{N}_2\text{O}_5$  on  $\text{NH}_4\text{HSO}_4$  aerosols containing nitrate in an aerosol flow tube at 1 bar and at room temperature, using both time dependence and surface area dependence (at fixed time) of  $[\text{N}_2\text{O}_5]$  decay.  $\gamma$  was determined as a function of  $[\text{NaNO}_3]$  at relative humidity of 60% and aerosol containing  $[\text{NH}_4\text{HSO}_4]$  of (0.028 to 0.086) wt%. At 55% RH result for aerosol containing zero  $\text{NH}_4\text{HSO}_4$ . The low initial  $[\text{N}_2\text{O}_5]$  (<5 ppt) afforded by CIMS minimised the impact of  $[\text{NO}_3^-]$  accumulated as a result of hydrolysis. A strong decrease in  $\gamma$  with increasing  $[\text{NO}_3^-]$  was observed.

## Preferred values

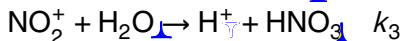
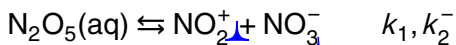
Parameter	Value	T/K
$\alpha_b$	0.04	298
$k^I$ ( $s^{-1}$ )	$k^{II \times} [H_2O]_{aq}$ (M)	
$k^{II}$ ( $M^{-1} s^{-1}$ )		298
$k_2/k_3$	0.054	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298
$\Delta \log(\gamma)$	$\pm 0.3$ at 50 % RH	

$[NO_3^-]$  and  $[H_2O]$  are molar concentrations of nitrate ion and water in aqueous aerosol

## Comments on preferred values

All studies show that the uptake of  $N_2O_5$  on aerosols containing nitrate is much reduced except at RH near 100 %. The so called “nitrate effect” was first described by Wahner et al. (1996) and was accounted for by a liquid phase hydrolysis mechanism involving reversible dissociation of solvated  $N_2O_5$ , according to the reaction mechanism:

$N_2O_5(g) \rightarrow N_2O_5(aq)$  accommodation,  $\alpha_b$



Thus the hydrolysis rate in solution was inhibited by the presence of nitrate ions which reduced uptake rate, making it chemically controlled as opposed to accommodation controlled. The inhibiting effect on  $\gamma(N_2O_5)$  of  $NO_3^-$  in aerosols containing sulphates and organics has also been investigated by Bertram and Thornton (2009) and by Griffiths et al. (2009). These studies gave values of  $(6.0 \pm 1.0) \times 10^{-2}$  and  $3.3 \times 10^{-2}$  for the ratio  $k_2/k_3$ , compared to a value of  $\sim 0.10$  reported by Wahner et al. (1998).

32310

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The  $\gamma$  values observed by Hallquist et al. (2003) were substantially smaller than observed by Mentel et al. (1999), who used aerosols with particles of larger mean size. This difference can be qualitatively accounted for by the effect of increased reacto-diffusive length as a result of the slower liquid phase hydrolysis, leading to volume-controlled uptake in small particles.

The recommended expression for RH dependence uses a size dependent resistance-model formulation:

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1 k^1)^{0.5}} \left[ \coth\left(\frac{r}{l}\right) - \left(\frac{l}{r}\right) \right] \right\}^{-1}$$

The recommended value of  $\alpha_b = 0.04$  is based on  $\gamma$  observed for uptake on malonic acid (Thornton et al., 2004) and  $\text{H}_2\text{SO}_4$  droplets and  $k^1 (= k^{\text{II}} \times [\text{H}_2\text{O}]_{\text{aq}})$ . The rate first order rate coefficient  $k^1$  is calculated using an expression for  $k^{\text{II}}$  modified for the presence of  $\text{NO}_3^-$  based on a hydrolysis mechanism steady-state treatment of the  $\text{H}_2\text{ONO}_2^+$  reactive intermediate:

$$k^1 = k^{\text{II}} \times [\text{H}_2\text{O}] \text{ where } k^{\text{II}} = k_0^{\text{II}} \left\{ 1 - \frac{k_2[\text{NO}_3^-]}{k_2[\text{NO}_3^-] + k_3[\text{H}_2\text{O}]} \right\}$$

$k_0^{\text{II}}$  the value in the absence of  $\text{NO}_3^-$ , was set at  $1.5 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ , which is the same as recommended for  $\text{N}_2\text{O}_5$  hydrolysis in other aqueous aerosols. The ratio  $k_2/k_3 = 0.053$  is the mean of the values deduced in the studies of Bertram and Thornton (2009), Griffiths et al. (2009) and Wahner et al. (1998). The RH dependence of  $\gamma$  can be described using water and nitrate concentrations taken from the AIM database, and with  $H(D_1)^{0.5} (\text{M atm}^{-1} \text{ cm s}^{-0.5}) = 2.0 \times 10^{-5}$ .

The reacto-diffusive parameter  $(D_1/k^1)^{0.5}$  predicts a significant size dependence of  $\gamma$  for  $r < 100 \text{ nm}$ . The expression for 100 nm particles fits the Mentel data quite well but overestimates uptake rates observed by Hallquist et al. (2003) at higher RH.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Hallquist et al. (2003) offer the only reported temperature dependence of  $\text{N}_2\text{O}_5$  uptake. They observed no significant temperature dependence of  $\gamma$  at constant RH of 50%. This is consistent with aqueous chemistry control of uptake rate.

## References

- 5 Hallquist, M., Stewart, D. J., Stephenson, S. K., and Cox, R. A.: Phys. Chem. Chem. Phys., 5, 3453–3463, 2003.
- Mentel, T. F., Sohn, M., and Wahner, A.: Phys. Chem. Chem. Phys., 1, 5451–5457, 1999.
- Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: Phys. Chem. Chem. Phys, 5, 4593–4603, 2003.
- 10 Wahner, A., Mentel, T. F., Sohn, M., and Stier, J.: J. Geophys. Res., 103 (D23), 31103–31112, 1998.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## VI.A3.8

 $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$  (aqueous organic aerosols)

## Experimental data

Parameter	RH/%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{ss}</math>, <math>\gamma_0</math></i>				
<i>Malonic Acid</i>				
$0.0020 \pm 0.0005$	10	298	Thornton et al. (2003)	AFT-CIMS (a)
$0.009 \pm 0.001$	20			
$0.0165 \pm 0.002$	30			
$0.022 \pm 0.0035$	40			
$0.025 \pm 0.003$	50			
$0.031 \pm 0.004$	70			
<i>Humic acid</i>				
$0.0001 \pm 0.0001$	25	298	Badger (2006)	AFT-CL/SMPS (b)
$0.0003 \pm 0.0001$	50			
$0.0010 \pm 0.0004$	75			
<i>Oxalic Acid (solid)</i>				
$0.0031 \pm 0.001$	74	$295 \pm 2$	Griffiths et al. (2008)	FTIR/SMPS/AFT/CL (c)
<i>Malonic Acid</i>				
$0.008 \pm 0.003$	30			
$0.012 \pm 0.002$	50			
$0.0169 \pm 0.0042$	64			
$0.016 \pm 0.003$	65			
<i>Succinic Acid</i>				
$<0.0006$	30			
$<0.0003$	50			
$0.0052 \pm 0.0003$	55			
$0.0051 \pm 0.0008$	64			
$0.006 \pm 0.001$	65			
$0.009 \pm 0.003$	70			
<i>Glutaric acid</i>				
$0.0006 \pm 0.0003$	30			
$0.0040 \pm 0.0009$	50			
$0.0081 \pm 0.0004$	65			
$0.0113 \pm 0.0029$	67			
$0.008 \pm 0.002$	75			

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments

(a) Measurements of the reactive uptake coefficient for  $\text{N}_2\text{O}_5$  hydrolysis,  $\gamma$ , on sub-micron organic aerosols in an aerosol flow tube at 1 bar and at room temperature, as a function of relative humidity (RH), aerosol phase,  $\text{N}_2\text{O}_5$  partial pressure, and mean aerosol size. Aerosol phase and relative humidity were determined simultaneously, and chemical ionization mass spectrometry was used to detect the decay rate of  $\text{N}_2\text{O}_5$  in the presence of malonic acid or azelaic acid aerosol. The  $\gamma$  on solid malonic acid was determined to be less than 0.001 (RH = 10–50 %), and on solid azelaic acid,  $\gamma$  was  $0.0005 \pm 0.0003$ . The values cited are for aqueous malonic acid aerosol. Evidence presented for an inverse dependence of  $\gamma$  on the initial concentration of  $\text{N}_2\text{O}_5$  between  $(1.5\text{--}7) \times 10^{11}$  molecule  $\text{cm}^{-3}$ , and a dependence on particle size for aerosol with surface area-weighted radii less than 100 nm at 50 % RH. Super-saturated malonic acid aerosol results are consistent with  $\text{N}_2\text{O}_5$  hydrolysis being both aerosol volume-limited where, for RH < 50 %, water is the limiting reagent, and also with a surface-specific process. The rate coefficient for reaction of  $\text{N}_2\text{O}_5$  with  $\text{H}_2\text{O}$  in solution retrieved for from the volume-limited regime was  $2.4 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ .

(b) Atmospheric pressure aerosol flow tube with  $\text{N}_2\text{O}_5$  ( $\sim 5 \times 10^{12}$  molecule  $\text{cm}^{-3}$ ) measured via thermal dissociation to  $\text{NO}_3$  and titration with  $\text{NO}$ , which was detected by  $\text{O}_3$ -chemiluminescence. Aqueous humic acid aerosols were generated from filtered solutions of its Na salt in a constant output atomiser, and conditioned by equilibration with an additional excess flow of controlled RH. Size distribution determined with a differential mobility analyser (DMA). The typical peak diameter,  $d_{\text{max}}$  was 170 nm, with surface area density of  $1.2$  to  $5.0 \times 10^{-2} \text{ cm}^2/\text{cm}^3$ . Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, after correction for wall loss and diffusion effects. The measured  $\gamma$  on humic acid aerosol increased linearly with RH in the range 25–75 % RH and was typically a factor of 50 lower than on sulphate. Uptake on aerosols of a range of

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



internally mixed composition of humic acid and  $(\text{NH}_4)_2\text{SO}_4$  were also measured. Even 6 wt % of humic acid caused a factor of  $>2$  reduction in uptake coefficient on  $(\text{NH}_4)_2\text{SO}_4$  at 70 % RH. This falloff was attributed to a decrease in bulk accommodation coefficient of  $\text{N}_2\text{O}_5$  at the surface due to surfactant properties of humic acid.

- (c) Uptake measurements on aqueous dicarboxylic acids (malonic, succinic, glutaric, oxalic) made using two different systems at 1 bar pressure: an aerosol flow tube system as described in comment (b); and a large static aerosol chamber, with  $\text{N}_2\text{O}_5$  ( $\approx 10^{13}$  molecule  $\text{cm}^{-3}$ ) produced in situ by reaction of  $\text{NO}_2$  with ozone, and gas phase species measured by FTIR and UV spectroscopy. Poly-dispersed aerosol (diam. range = 20 nm–5  $\mu\text{m}$ ) generated by spraying dilute solutions of carboxylic acids, with size distribution measured by SMPS for particle diameters  $<700$  nm and for larger diameters by aerodynamic particle sizer. Uptake coefficients determined by fitting experimental time dependence of species with a numerical model of chemistry and integrated aerosol surface area. The results for uptake on dicarboxylic acid aerosols at different humidities and phase were in good agreement using the two methods. The  $\gamma$  values all increased linearly with RH up to 70 % and for succinic and glutaric acid were distinctly lower than for malonic acid. For the latter,  $\gamma$  values were lower than observed by Thornton et al. (2003), and were not observed to level off above 50 % RH. This was attributed to a nitrate effect caused by accumulation of  $\text{HNO}_3$  product in the aqueous particles. Rate constants for liquid phase reaction of  $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$  were obtained in this study for malonic acid:  $1.8 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ ; for succinic acid:  $0.79 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ ; for glutaric acid:  $1.2 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$  from analysis of the volume limited uptake data.

## Preferred values

Parameter	Value	T/K
$\alpha_b$ (dicarboxylic acids)	0.035	298
$\gamma$ (humic acid)	$3.1 \times 10^{-5} \exp(0.046 \times \text{RH}(\%))$	298
$k^I$ ( $\text{s}^{-1}$ )	$k^{II} [\text{H}_2\text{O}]_{\text{aq}} \text{ (M)}$	
$k^{II}$ ( $\text{M}^{-1} \text{s}^{-1}$ ) malonic acid	$1.0 \times 10^5$	298
$k^{II}$ ( $\text{M}^{-1} \text{s}^{-1}$ ) succinic acid	$3.0 \times 10^4$	298
$k^{II}$ ( $\text{M}^{-1} \text{s}^{-1}$ ) glutaric acid	$5.0 \times 10^4$	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298
$\Delta \log(\gamma)$ humic acid	$\pm 0.3$ at 50 % RH	298

## Comments on preferred values

All studies used deliquesced or supersaturated aqueous aerosols, with uptake coefficients measured as a function of RH.  $\gamma$  was generally seen to increase with RH over the range 20–90 %, although for malonic acid the  $\gamma$  value at RH > 50% Thornton et al. (2004) is constant and close to that on  $\text{H}_2\text{SO}_4$  droplets ( $\gamma = 0.036 \pm 0.008$ ). The lack of a dependence on the water content of deliquesced droplets at RH > 50 %, suggests that uptake is controlled by surface accommodation. At lower RH for malonic and for the C4 and C5 acids,  $\gamma$  declines, which suggests that uptake is limited by the rate of hydrolysis of  $\text{N}_2\text{O}_5$  in the bulk liquid phase, and hence on  $[\text{H}_2\text{O}]_{\text{aq}}$ . For very small particles at low RH uptake rate can become dependent on particle volume. The uptake rates observed by Griffiths et al. (2009) are lower than observed by Thornton et al. (2003). This was attributed to a nitrate effect caused by accumulation of  $\text{HNO}_3$  product in the aqueous particles, also indicated in the decline of  $\gamma$  with  $[\text{N}_2\text{O}_5]$  observed by Thornton et al. (2003). A numerical model was developed by Griffiths et al. involving explicit treatment of the coupled processes involved in uptake, including the nitrate effect. The

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

rate coefficient for reaction of  $\text{N}_2\text{O}_5$  with  $\text{H}_2\text{O}$  in solution obtained using this model to analyse malonic acid data for 70 % RH ( $[\text{H}_2\text{O}] = 23 \text{ M}$ ) was  $k^{\text{II}} = 2.1 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ . This is larger than that deduced by Thornton et al. (2003),  $k^{\text{II}} = 2.5 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ , leading to the conclusion that rate constants derived assuming volume limited uptake are too low. Thus the rate constants derived in their study from analysis of the linear region; for succinic acid:  $0.79 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ ; for malonic acid:  $1.8 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ ; and for glutaric acid:  $1.2 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$  are also likely to be underestimated.

The recommended expression for RH dependence uses a size dependent resistance-model formulation:

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{\bar{c}}{4HRT(D_1k^{\text{I}})^{0.5}} \left[ \coth\left(\frac{r}{l}\right) - \left(\frac{l}{r}\right) \right] \right\}^{-1}$$

The recommended value of  $\alpha_b = 0.035$  is based on  $\gamma$  observed for uptake on malonic acid.  $k^{\text{I}}$  is calculated using the recommended liquid phase rate constant for  $\text{N}_2\text{O}_5 + [\text{H}_2\text{O}]_{\text{aq}}$ :  $k^{\text{II}} = 1 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ , which is intermediate between the values derived by Thornton et al. (2003) from hydrolysis of  $\text{N}_2\text{O}_5$  on malonic acid aerosol ( $2.5 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ ) and by Mentel et al. (1999) from uptake on  $\text{NaNO}_3$  aerosols ( $1.5 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ ). The RH dependence of  $\gamma$  can be calculated using water mass fractions taken from the AIM database, and using  $H(D_1)^{0.5} (\text{M atm}^{-1} \text{ cm s}^{-0.5}) = 6.3 \times 10^{-3}$ . The expression fits the average values reported for malonic acid in the range 30–60 % RH reasonably well, but the RH dependence is weaker than observed over the range. The reacto-diffusive parameter ( $\{D_1/k^{\text{II}}[\text{H}_2\text{O}]\}^{0.5}$ ) predicts a significant size dependence of  $\gamma$  for  $r < 100 \text{ nm}$ .

The lower uptake rates observed on succinic and glutaric acid particles cannot be entirely rationalised by reduced water mass fractions or acidity/nitrate effects (Badger, 2006), suggesting that the effective liquid phase rate constants ( $k^{\text{II}}$ ) is lower for these acids by ~70 % and 50 % compared to the recommended value for malonic acid.

The recommended uptake coefficient for  $\text{N}_2\text{O}_5$  on humic acid as a function of RH is based on the results of Badger et al. (2006).

32317

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

## References

Badger, C. L., Griffiths, P. T., George, I., Abbatt, J. P. D., and Cox, R. A.: J. Phys. Chem., 110, 6986–6994, 2006.

Badger, C. L.: PhD thesis, University of Cambridge, 2006.

5 Griffiths, P. T., Badger, C. L., Cox, R. A., Folkers, M., Henk, H. H., and Mentel, T. F.: J. Phys. Chem. A, 113, 5082–5090, doi:10.1021/jp8096814, 2009.

Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: Phys. Chem. Chem. Phys., 5, 4593–4603, 2003.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A3.9

### HO<sub>2</sub> + NH<sub>4</sub>HSO<sub>4</sub> (aq) → products

#### Experimental data

	Temp./K	Reference	Technique/Comments
Accommodation coefficients: $\alpha$ >0.2	293	Mozurkewich et al. (1987)	AFT (a)



#### 5 Comments

(a) Uptake of HO<sub>2</sub> (10<sup>8</sup>-10<sup>9</sup> molecule cm<sup>-3</sup>) to deliquescent particles (0.05 to 0.1 μm) containing CuSO<sub>4</sub>. HO<sub>2</sub> was formed by passing a mixture of H<sub>2</sub> and H<sub>2</sub>O over a hot Nichrome wire and detected by chemically amplified conversion to NO<sub>2</sub>. Levels of HO<sub>2</sub> were sufficiently low to neglect loss due to gas-phase self reaction. The average experimental uptake coefficient was  $\gamma = 0.4 \pm 0.08$  with no measurable dependence on the CuSO<sub>4</sub> molality between 0.006–0.07 M. The uptake coefficient dropped sharply when the CuSO<sub>4</sub> solution was less than  $\sim 10^{-3}$  M and was zero when less than  $10^{-4}$  M. The authors make a conservative estimate of 0.2 for the lower limit to the accommodation coefficient,  $\alpha_b$ .

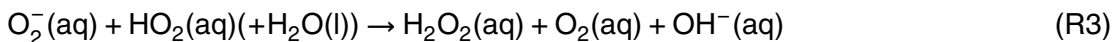
#### 15 Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.5	293 K
Reliability		
$\Delta (\log \alpha_b)$	undetermined	

## Comments on preferred values

The single experimental investigation of HO<sub>2</sub> uptake to NH<sub>4</sub>HSO<sub>4</sub> derived a lower limit to the accommodation coefficient of 0.2, which is consistent with results on other aqueous surfaces where uptake coefficients as large as 0.5 have been measured (e.g. VI.A3.10).

The uptake of HO<sub>2</sub> in aqueous solution with pH > 5, is presently believed to be driven by self-reaction and acid-base dissociation of HO<sub>2</sub> (pK<sub>a</sub> ~ 4.7) with formation of H<sub>2</sub>O<sub>2</sub> (Reactions R2, R3). In the presence of transition metal ions (TMI) the reaction of HO<sub>2</sub> and especially O<sub>2</sub><sup>-</sup> (Reaction R4) can be important:



If a first-order loss process for HO<sub>2</sub> or O<sub>2</sub><sup>-</sup> in the aqueous phase dominates (e.g. reaction with TMI such as Cu(II)), and assuming equal rates of reaction throughout the particle, the uptake coefficient can be calculated from the expression below:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{\bar{c}}{4H^{\text{eff}}RT\sqrt{k_{\text{TMI}}[\text{TMI}]D_1}}$$

$H^{\text{eff}} = H^{\text{HO}_2} (1 + K_{\text{eq}}/[\text{H}^+], K_{\text{eq}} = 2.1 \times 10^{-5} \text{ M at } 298 \text{ K (Jacob, 2000)}, H^{\text{HO}_2} = 9.5 \times 10^{-6} \exp(5910/T) \text{ (Hanson et al., 1992) and } D_1 = [1 \times 10^{-5}(T/298)]/(1.09 \times 10^8 \exp(-0.068T) + 0.873) \text{ cm}^2 \text{ s}^{-1} \text{ (Schwartz, 1984; Thornton et al., 2008) where the denominator in the } D_1 \text{ term was derived from a fit to the water viscosity data of Hallett (1963).}$

According to the reaction scheme above, in the absence of TMI, the rates of loss of aqueous-phase HO<sub>2</sub> are quadratically dependent on [HO<sub>2</sub>]<sub>aq</sub> and [O<sub>2</sub><sup>-</sup>]<sub>aq</sub> and are thus



strongly dependent on the gas-phase concentration of HO<sub>2</sub>. At low HO<sub>2</sub> concentrations (e.g. as found in the atmosphere) the liquid phase reactions become rate limiting and  $\gamma$  is expected to be much smaller as observed in dilute solutions by Mozurkewich et al. (1987) and the simple formalism above breaks down. Thornton and Abbatt (2005) suggest that the rate of loss of HO<sub>2</sub> from the gas-phase (in molecule cm<sup>-3</sup> s<sup>-1</sup>) is best described by a system in thermodynamic (Henry's law) equilibrium so that (Thornton et al., 2008):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{3\bar{c}N_A}{8000 (H^{\text{eff}}RT)^2 k_{\text{aq}}[\text{HO}_2]r}$$

$k_{\text{aq}}$  can be calculated from the rate coefficients for Reaction (R2) ( $k_2$ ) and (R3) ( $k_3$ ) (Bielski et al., 1985) and the pH:

$$k_{\text{aq}} = \frac{k_2 + \left(\frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}}\right) k_3}{\left(1 + \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}}\right)^2}$$

This formalism predicts that the loss of HO<sub>2</sub> to particles is favoured by high HO<sub>2</sub> mixing ratios, low temperatures (higher solubility) and low pH. At low concentrations of HO<sub>2</sub> (whereby the self reaction and reaction with O<sub>2</sub><sup>-</sup> are inefficient), values of  $\gamma$  of <0.005 are calculated, which are however much less than the uptake coefficients of Taketani et al. (2008) who investigated the uptake of low concentrations of HO<sub>2</sub> to (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> particles. In addition, as discussed by Hanson et al. (1992) and Thornton and Abbatt (2005), there is considerable uncertainty (factor of 2.5) associated with the solubility of HO<sub>2</sub> ( $H^{\text{HO}_2}$ ) and its temperature dependence. The above schemes also do not account for the RH dependence of uptake observed for (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> particles (see datasheet VI.A3.10).

Until these apparent discrepancies have been resolved by further experiments, we make no recommendation for parameterising HO<sub>2</sub> uptake to aqueous aerosol. We refer

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



to recent publications for a more detailed description of the effect of different parameterisation schemes (Thornton et al., 2008; Macintyre and Evans, 2011).

## References

- 5 Bielski, B. H. J., Cabelli, D. E., Arudi, R. L., and Ross, A. B.: J. Phys. Chem. Ref. Data, 14, 1041–1100, 1985.
- Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.
- Jacob, D. J.: Atmos. Env., 34, 2131–2159, 2000.
- 10 Macintyre, H. L. and Evans, M. J.: Atmos. Chem. Phys., 10, 7409–7414, 2010, <http://www.atmos-chem-phys.net/10/7409/2010/>.
- Morita, A., Kanaya, Y., and Francisco, J. S.: J. Geophys. Res.-Atmos., 109, art. no. D09201, doi:10.1029/2003JD004240, 2004.
- Mozurkewich, M., McMurry, P. H., Gupta, A., and Calvert, J. G.: J. Geophys. Res.-Atmos., 92, 4163–4170, 1987.
- 15 Schwartz, S. E.: J. Geophys. Res.-Atmos., 89, 1589–1598, 1984.
- Thornton, J. and Abbatt, J. P. D.: J. Geophys. Res.-Atmos., 110, D08309, doi:10.1029/2004JD005402, 2005.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A3.10

### HO<sub>2</sub> + (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> (aq) → products

#### Experimental data



	Temp./K	Reference	Technique/Comments
Accommodation coefficients: $\alpha_b$			
>0.5	298	Thornton and Abbatt (2005)	AFT (a)
>0.5	296 ± 2	Taketani et al. (2008)	AFT (b)

#### 5 Comments

(a) Uptake of HO<sub>2</sub> ( $2.5\text{--}5.0 \times 10^{10}$  molecule cm<sup>-3</sup>) to deliquescent particles (mean, surface area weighted radius of 125 nm) at RH = 40–45 % containing 0.01–0.1 M CuSO<sub>4</sub>, and buffered to a pH of 5.1. HO<sub>2</sub> was formed in the reaction of H atoms with O<sub>2</sub> (the former made in a microwave discharge of H<sub>2</sub>). Detection of HO<sub>2</sub> was by CIMS as O<sub>2</sub><sup>-</sup> using F<sup>-</sup> reagent ions, or conversion of HO<sub>2</sub> to H<sub>2</sub>SO<sub>4</sub> and detection of the product (as HSO<sub>4</sub><sup>-</sup>) using I<sup>-</sup> ions.

(b) Uptake of HO<sub>2</sub> ( $\sim 10^8$  molecule cm<sup>-3</sup>) to (NH<sub>4</sub>)<sub>2</sub>SO<sub>4</sub> particles (mean surface area weighted diameter of 80–110 nm) at RH between 45 and 75 %. HO<sub>2</sub> was generated by the photolysis of H<sub>2</sub>O in air and detected as OH (by LIF) following conversion in reaction with NO. The particles contained CuSO<sub>4</sub> ( $\sim 0.5$  M) to scavenge HO<sub>2</sub> in order to determine  $\alpha_b$ .

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

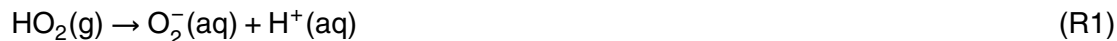
## Preferred values

Parameter	Value	T/K
$\alpha$	>0.5	290–300
Reliability		
$\Delta(\log \alpha_b)$	undetermined	

## Comments on preferred values

Thornton and Abbatt (2005) derived an average experimental uptake coefficient of  $\gamma = 0.5 \pm 0.1$  on particles containing  $\text{CuSO}_4$  and suggest that this is a lower limit to the accommodation coefficient. In the absence of  $\text{CuSO}_4$  non-exponential loss of  $\text{HO}_2$  was observed, which the authors attributed to the aqueous phase self reaction of  $\text{HO}_2$  and derived an aqueous-phase rate coefficient of  $1 \pm 0.25 \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$  for this process, consistent with kinetic data on the reaction of  $\text{HO}_2$  with itself and with  $\text{O}_2^-$ . Using significantly lower  $\text{HO}_2$  concentrations, Taketani et al. (2008) observed exponential  $\text{HO}_2$  decays with  $\gamma$  equal to  $\sim 0.5$  when  $\text{CuSO}_4$  was present, defining the lower limit to  $\alpha_b$ . In the absence of  $\text{CuSO}_4$   $\gamma$  increased with RH with values of 0.11 (45%), 0.15 (55%), 0.17 (65%) and 0.19 (75%). Taketani et al. speculate that the RH dependence stems from the effect of particle size changes with RH and the reaction occurring in the transition regime between surface and volume limited uptake. Lower uptake coefficients have been reported for  $\text{HO}_2$  interaction with dry  $(\text{NH}_4)_2\text{SO}_4$  particles with Taketani et al. (2008) reporting values of 0.04–0.05 and Gershenson et al. (1995) reporting 0.011.

The uptake of  $\text{HO}_2$  in aqueous solution with  $\text{pH} > 5$ , is presently believed to be driven by self-reaction and acid-base dissociation of  $\text{HO}_2$  ( $\text{pK}_a \sim 4.7$ ) with formation of  $\text{H}_2\text{O}_2$  (Reactions R2, R3). In the presence of transition metal ions (TMI) the reaction of  $\text{HO}_2$  and especially  $\text{O}_2^-$  (Reaction R4) can be important:



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

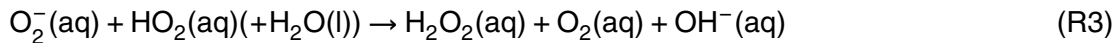
Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





If a first-order loss process for  $\text{HO}_2$  or  $\text{O}_2^-$  in the aqueous phase dominates (e.g. reaction with TMI such as  $\text{Cu}(\text{II})$ ), and assuming equal rates of reaction throughout the particle, the uptake coefficient can be calculated from the expression below:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{\bar{c}}{4H^{\text{eff}}RT\sqrt{k_{\text{TMI}}[\text{TMI}]D_1}}$$

$H^{\text{eff}} = H^{\text{HO}_2} (1 + K_{\text{eq}}/[\text{H}^+])$ ,  $K_{\text{eq}} = 2.1 \times 10^{-5} \text{ M}$  at 298 K (Jacob, 2000),  $H^{\text{HO}_2} = 9.5 \times 10^{-6} \exp(5910/T)$  (Hanson et al., 1992) and  $D_1 = [1 \times 10^{-5}(T/298)] / (1.09 \times 10^8 \exp(-0.068T) + 0.873) \text{ cm}^2 \text{ s}^{-1}$  (Schwartz, 1984; Thornton et al., 2008) where the denominator in the  $D_1$  term was derived from a fit to the water viscosity data of Hallett (1963).

According to the reaction scheme above, in the absence of TMI, the rates of loss of aqueous-phase  $\text{HO}_2$  are quadratically dependent on  $[\text{HO}_2]_{\text{aq}}$  and  $[\text{O}_2^-]_{\text{aq}}$  and are thus strongly dependent on the gas-phase concentration of  $\text{HO}_2$ . At low  $\text{HO}_2$  concentrations (e.g. as found in the atmosphere) the liquid phase reactions become rate limiting and  $\gamma$  is expected to be much smaller as observed in dilute solutions by Mozurkewich et al. (1987) and the simple formalism above breaks down. Thornton and Abbatt (2005) suggest that the rate of loss of  $\text{HO}_2$  from the gas-phase (in molecule  $\text{cm}^{-3} \text{ s}^{-1}$ ) is best described by a system in thermodynamic (Henry's law) equilibrium so that (Thornton et al., 2008):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{3\bar{c}N_A}{8000 (H^{\text{eff}}RT)^2 k_{\text{aq}}[\text{HO}_2]r}$$

32325

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$k_{\text{aq}}$  can be calculated from the rate coefficients for Reactions (R2) ( $k_2$ ) and (R3) ( $k_3$ ) (Bielski et al., 1985) and the pH:

$$k_{\text{aq}} = \frac{k_2 + \left( \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}} \right) k_3}{\left( 1 + \frac{K_{\text{eq}}}{[\text{H}^+]_{\text{aq}}} \right)^2}$$

This formalism predicts that the loss of  $\text{HO}_2$  to particles is favoured by high  $\text{HO}_2$  mixing ratios, low temperatures (higher solubility) and low pH. At low concentrations of  $\text{HO}_2$  (whereby the self reaction and reaction with  $\text{O}_2^-$  are inefficient), values of  $\gamma$  of  $<0.005$  are calculated, which are however much less than the uptake coefficients of Taketani et al. (2008) who investigated the uptake of low concentrations of  $\text{HO}_2$  to  $(\text{NH}_4)_2\text{SO}_4$  particles. In addition, as discussed by Hanson et al. (1992) and Thornton and Abbatt (2005), there is considerable uncertainty (factor of 2.5) associated with the solubility of  $\text{HO}_2$  ( $H^{\text{HO}_2}$ ) and its temperature dependence. The above schemes also do not account for the RH dependence of uptake observed by Taketani et al. (2008).

Until these apparent discrepancies have been resolved by further experiments, we make no recommendation for parameterising  $\text{HO}_2$  uptake to aqueous aerosol. The hypothesis of a fast disproportionation reaction  $2\text{HO}_2 \rightarrow \text{H}_2\text{O}_2 + \text{O}_2$  in aqueous solution advanced by Thornton and Abbatt (2005) is inconsistent with the measured value of  $\gamma$  for  $\text{HO}_2$  loss on both NaCl and  $(\text{NH}_4)_2\text{SO}_4$  aerosol when extrapolated to the  $\text{HO}_2$  concentrations lower by a factor of 500 used in the work of Taketani et al. (2008). We refer to recent publications for a more detailed description of the effect of different parameterisation schemes (Thornton et al., 2008; Macintyre and Evans, 2011).

## References

Bielski, B. H. J., Cabelli, D. E., Arudi, R. L., and Ross, A. B.: J. Phys. Chem. Ref. Data, 14, 1041–1100, 1985.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Gershenzon, Y. M., Grigorieva, V. M., Ivanov, A. V., and Remorov, R. G.: Faraday Discuss., 100, 83–100, 1995.

Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.

5 Jacob, D. J.: Atmos. Env., 34, 2131–2159, 2000.

Macintyre, H. L. and Evans, M. J.: Atmos. Chem. Phys., 10, 7409–7414, 2010, <http://www.atmos-chem-phys.net/10/7409/2010/>.

Morita, A., Kanaya, Y., and Francisco, J. S.: J. Geophys. Res.-Atmos., 109, 2004.

10 Mozurkewich, M., McMurry, P. H., Gupta, A., and Calvert, J. G.: J. Geophys. Res.-Atmos., 92, 4163–4170, 1987.

Schwartz, S. E.: J. Geophys. Res.-Atmos., 89, 1589–1598, 1984.

Taketani, F., Kanaya, Y., and Akimoto, H.: J. Phys. Chem. A, 112, 2370–2377, 2008.

Thornton, J. and Abbatt, J. P. D.: J. Geophys. Res.-Atmos., 110, D08309, doi:10.1029/2004JD005402, 2005.

ACPD

12, 32109–32472, 2012

---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

---

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Appendix A4

### Uptake on liquid sulfuric acid

#### VI.A4.0

#### $O_3 + H_2SO_4 \rightarrow$ products

#### 5 Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$<1 \times 10^{-6}$	300–270	Baldwin and Golden (1979)	Knudsen-MS (a)
$(4.79 \pm 1.38) \times 10^{-11}$ (75 wt % $H_2SO_4$ )	223	Olszyna et al. (1979)	(b)
$<1 \times 10^{-6}$	195	Dlugokencky and Ravishankara (1992)	CWFT-CLD (c)
$(1.2 \pm 0.08) \times 10^{-6}$ (98 wt % $H_2SO_4$ )	239	Il'in et al. (1992)	UV (d)
$(1.6 \pm 0.08) \times 10^{-6}$ (98 wt % $H_2SO_4$ )	258		
$(1.75 \pm 0.07) \times 10^{-6}$ (98 wt % $H_2SO_4$ )	273		

#### Comments

(a) No evidence for uptake of  $O_3$  to  $H_2SO_4$  surface which contained less than 5 % water.

10 (b) Static experiment with  $H_2SO_4$  coated on glass beads in a round bottom flask.  $O_3$  analysed ex-situ. Slow loss of  $O_3$  was observed over several hours to derive the values of  $\gamma$  listed.

15 (c)  $\approx 2$  mm thick (frozen)  $H_2SO_4$  films were prepared from a bulk solution of 50 or 97 wt %.  $[O_3]$  was either  $5 \times 10^8$  or  $2 \times 10^9$  molecule  $cm^{-3}$ . The geometric surface area was used to calculate  $\gamma$ .

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(d) Static reactor with O<sub>3</sub> measured using optical absorption at 254 nm. Initial [O<sub>3</sub>] was  $1-6 \times 10^{15}$  molecule cm<sup>-3</sup>.

### Preferred values

Parameter	Value	T/K
$\gamma$	$<10^{-6}$	200–220 K
Reliability		
$\Delta \log(\gamma)$	undetermined	

### Comments on preferred value

The studies of the uptake of O<sub>3</sub> to various H<sub>2</sub>SO<sub>4</sub> surfaces confirm that the interaction is very weak with uptake coefficients lower than  $10^{-6}$ . Olszyna et al. (1979) measured comparable uptake coefficients on pure H<sub>2</sub>SO<sub>4</sub> and H<sub>2</sub>SO<sub>4</sub> doped with various cations (Ni<sup>2+</sup>, Cu<sup>2+</sup>, Cr<sup>3+</sup>, Al<sup>3+</sup>, Fe<sup>3+</sup>, NH<sub>4</sub><sup>+</sup> and Mn<sup>2+</sup>). Doping with Fe<sup>2+</sup> resulted in  $\gamma$  values close to  $10^{-9}$ . Uptake coefficients of the order of  $10^{-8}$  to  $10^{-9}$  were obtained on frozen H<sub>2</sub>SO<sub>4</sub> surfaces at 217–263 (Harker and Ho, 1979). The preferred value is based on the uptake coefficients obtained in the flow tube experiments of Dlugokencky and Ravishankara (1992) which were carried out at stratospheric temperatures and using low [O<sub>3</sub>].

### References

- Baldwin, A. C. and Golden, D. M.: Science, 206, 562–563, 1979.  
Dlugokencky, E. J. and Ravishankara, A. R.: Geophys. Res. Lett., 19, 41–44, 1992.  
Harker, A. B. and Ho, W. W.: Atmos. Environ., 13, 1005–1010, 1979.  
Il'in, S. D., Selikhanovich, V. V., Gershenson, Y. M., and Rozenshtein, V. B.: Sov. J. Chem. Phys., 8, 1858–1880, 1991.  
Olszyna, K., Cadle, R. D., and Depena, R. G.: J. Geophys. Res., 84, 1771–1775, 1979.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.1

### OH (g) + H<sub>2</sub>SO<sub>4</sub> (aq) → products

#### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/ wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma_{ss} = (4.5 \pm 0.5) \times 10^{-4}$	>96	298	Baldwin and Golden (1980)	Kn-MS (a)
1	96	298	Gershenson et al. (1986)	(b)
$\gamma_{ss} > 8 \times 10^{-2}$	28	249	Hanson et al. (1992)	WWFT-LIF(c)
<i>Accommodation coefficients: <math>\alpha_b</math></i>				
>0.2	45–96	220–298	Cooper and Abbatt (1996)	CWFT-RF (d)

#### 5 Comments

(a) Hydroxyl radicals were generated by reaction of hydrogen atoms from microwave discharge of hydrogen with excess NO<sub>2</sub>. The uptake coefficient as listed in the table was not found to depend on OH pressure. The sulfuric acid solution was >95 wt %.

(b) Hydroxyl radicals were generated by reaction of hydrogen atoms from microwave discharge of hydrogen with NO<sub>2</sub>, leading to OH concentrations of  $2 \times 10^{12}$  molecule cm<sup>-3</sup> and lower. OH was detected by EPR. H<sub>2</sub>SO<sub>4</sub> was coated onto a quartz rod that was inserted into a Teflon coated quartz flow tube at about 6 mbar. No error is given to account for uncertainty due to the effects of gas phase diffusion.

(c) OH [ $(0.5\text{--}3) \times 10^{11}$  molecule cm<sup>-3</sup>] was generated through the reaction H + NO<sub>2</sub> and was monitored using LIF. A fluid 28 % H<sub>2</sub>SO<sub>4</sub>-H<sub>2</sub>O mixture was used at 249.5 K (2.1 mbar total pressure). The value listed in the table was corrected for

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



gas phase diffusion. The diffusion coefficient for OH in H<sub>2</sub>O vapor was taken to be the value of the self-diffusion coefficient for H<sub>2</sub>O,  $116 \pm 20 \text{ Torr cm}^2 \text{ s}^{-1}$ .

(d) Coated-wall flow tube with resonance fluorescence detection of OH. First-order loss rate of OH measured in the presence of various types of surfaces. OH ( $< 5 \times 10^{10} \text{ molecule cm}^{-3}$ ) was generated using both F + H<sub>2</sub>O and H + NO<sub>2</sub>. Total pressure in the flow tube was 1.33 mbar. Measured uptake was corrected for gas phase diffusion; the following binary diffusion coefficients were used:  $0.035 T^{1.75} \text{ Torr cm}^2/\text{s}$  for OH in He,  $0.0063 T^{1.75} \text{ Torr cm}^2/\text{s}$  for OH in H<sub>2</sub>O.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	1	220–298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.7$	220–298

### Comments on preferred values

The experiment by Hanson et al. (1992) was close to the diffusion limit, and the corrected uptake coefficient was still roughly consistent with bulk reaction limited uptake due to OH reacting with HSO<sub>4</sub><sup>-</sup> (Buxton et al., 1988), indicating that the bulk accommodation coefficient is significantly larger than this. The very high H<sub>2</sub>SO<sub>4</sub> concentration (due to pumping on it) and thus lower HSO<sub>4</sub><sup>-</sup> concentration may have led to the lower uptake coefficient reported by Baldwin and Golden (1980). Gershenson et al. (1986) estimated the uptake coefficient to be 1, also at room temperature, for 96% sulfuric acid not further concentrated. The experiment by Cooper and Abbatt (1996) was performed at a lower total pressure in the flow tube to reduce gas phase diffusion limitation and at high enough H<sub>2</sub>SO<sub>4</sub> concentration to provide a strong enough bulk liquid sink for OH, leading to 0.2 as lower limit, without indication of temperature dependence. We

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



provide relatively large error limits to the recommended value for  $\alpha_b$  to take into account the uncertainty associated with diffusion limitations in Gershenzon's et al. (1986) experiment.

## References

- 5 Baldwin, A. C. and Golden, D. M.: J. Geophys. Res., 85, 2888–2889, 1980.  
Buxton, G. V., Greenstock, C. L., Helman, W. P., and Ross, A. B.: J. Phys. Chem. Ref. Data, 17, 513–886, 1988.  
Cooper, P. L. and Abbatt, J. P. D.: J. Phys. Chem., 100, 2249–2254, 1996.  
Gershenzon, Y. M., Ivanov, A. V., Kucheryavyi, S. I., and Rozenshtein, V. B.: Kin. Catal., 27, 923–927, 1986.
- 10 Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.2

### HO<sub>2</sub> + H<sub>2</sub>SO<sub>4</sub> → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
>0.05	249.5	Hanson et al. (1992)	WWFT-LIF (a)
>0.1	243	Gershenson et al. (1995)	CRFT-EPR (b)
0.055 ± 0.02	223	Cooper and Abbatt (1996)	CWFT-RF (c)
<0.01	295	Thornton et al. (2005)	AFT-CIMS (d)
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
>0.2	223	Cooper and Abbatt (1996)	CWFT-RF (c)
0.8 ± 0.3	295	Thornton et al. (2005)	AFT-CIMS (d)

#### 5 Comments

(a) Uptake of HO<sub>2</sub> ( $5\text{--}30 \times 10^{10}$  molecule cm<sup>-3</sup>) to 28 wt% H<sub>2</sub>SO<sub>4</sub> films ≈0.3 mm thick. HO<sub>2</sub> was formed in the reaction of F with H<sub>2</sub>O<sub>2</sub> and detected as OH after reaction with NO. HO<sub>2</sub> uptake was limited by diffusion through the 1 Torr of He bath gas.

10 (b) HO<sub>2</sub> was detected either directly ( $[\text{HO}_2] = 3\text{--}5 \times 10^9$  molecule cm<sup>-3</sup>) or as OH following reaction with NO ( $[\text{HO}_2] = 1\text{--}3 \times 10^{11}$  molecule cm<sup>-3</sup>). H<sub>2</sub>SO<sub>4</sub> films were either 80 or 96 wt%.

15 (c) Uptake of HO<sub>2</sub> to 55 wt% H<sub>2</sub>SO<sub>4</sub> to determine  $\gamma$ .  $\alpha$  was determined by doping the H<sub>2</sub>SO<sub>4</sub> with 0.1 M CuSO<sub>4</sub>. HO<sub>2</sub> was formed in the reaction of F with H<sub>2</sub>O<sub>2</sub> and detected as OH after reaction with NO.

32333

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(d)  $\text{H}_2\text{SO}_4$  aerosol (diameter  $\approx 100$  nm) at 35–40 % RH.  $\alpha$  was determined by doping the  $\text{H}_2\text{SO}_4$  aerosol (made from 0.0005–0.005 M aqueous solutions) with 0.1 M  $\text{CuSO}_4$ . No dependence on RH was observed over the small range covered.  $\text{HO}_2$  (at concentrations of  $\approx 4 \times 10^{10}$  molecule  $\text{cm}^{-3}$ ) was formed in the reaction of H with  $\text{O}_2$  and detected as  $\text{O}_2^-$  using  $\text{F}^-$  as chemi-ion or as  $\text{HSO}_4^-$  following conversion to  $\text{H}_2\text{SO}_4$  (addition of NO and  $\text{SO}_2$ ) and ionisation with  $\text{I}^-$ . Loss of  $\text{HO}_2$  to 55 wt %  $\text{H}_2\text{SO}_4$  was indistinguishable from loss due to the reactor walls, hence the upper limit to  $\gamma$ .

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.8	220–300
<i>Reliability</i>		
$\Delta\alpha_b$	+0.2 -0.4	220–300

### Comments on preferred value

Two studies (Cooper and Abbatt, 1996; Thornton and Abbatt, 2005) show that bulk accommodation of  $\text{HO}_2$  to sulphate aerosol is very efficient with values of  $\alpha_b$  consistent with unity. The results from the four studies of the net uptake coefficient are however rather divergent, with values of  $\gamma$  obtained that vary from  $>0.1$  to  $<0.01$ . A probable explanation is related to the different experimental temperatures, with low temperatures favouring large uptake coefficients. Thornton and Abbatt (2005) suggest that the results from these studies are internally consistent and the increase in uptake coefficient at low temperature is due to enhanced solubility of  $\text{HO}_2$ . Moreover, Thornton and Abbatt (2005) suggest that the net uptake depends on the gas-phase  $\text{HO}_2$  concentration which impacts strongly on aqueous phase loss rates (due to self-reaction) of  $\text{HO}_2$ . As

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



all experiments were carried out using HO<sub>2</sub> concentrations orders of magnitude greater than found in the atmosphere, no expression for  $\gamma$  is given.

## References

Cooper, P. L. and Abbatt, J. P. D.: J. Phys. Chem., 100, 2249–2254, 1996.

5 Gershenzon, Y. M., Grigorjeva, V. M., Ivanov, A. V., and Remorov, R. G.: Faraday Discuss., 83–100, 1995.

Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A. R.: J. Phys. Chem., 96, 4979–4985, 1992.

10 Thornton, J. and Abbatt, J. P. D.: J. Geophys. Res., 110, art. no. D08309, doi:10.1029/2004jd005402, 2005.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

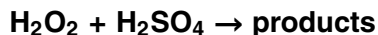
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.3



### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$7.8 \times 10^{-4}$	270–300	Baldwin and Golden (1979)	Knudsen-MS (a)

### 5 Comments

(a) After evacuating the Knudsen reactor containing the concentrated  $\text{H}_2\text{SO}_4$  sample, the remaining substrate was suggested to contain less than 5 % water.

### Preferred values

Parameter	Value	T/K
$\gamma$	$7.8 \times 10^{-4}$	270–300
<i>Reliability</i>		
$\Delta \log(\gamma)$	0.5	

### 10 Comments on preferred value

The single published study of the uptake of  $\text{H}_2\text{O}_2$  to sulphuric acid agrees with unpublished data (obtained using a Knudsen reactor) carried out using 60 wt %  $\text{H}_2\text{SO}_4$  at 203–223 K (Myhre and Nielsen, 1998) who derived values of  $\gamma$  in the range  $(0.69\text{--}3.8) \times 10^{-4}$  with the lower values obtained at the lowest temperature. There appears to be no information regarding the products of the interaction. We therefore recommend

15 the uptake coefficient of Baldwin and Golden (1979) with expanded error limits.

32336

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## References

- Baldwin, A. C. and Golden, D. M.: Science, 206, 562–563, 1979.
- Myhre, C. E. L. and Nielsen, C. J.: Polar Stratospheric Ozone 1997, Proceedings of the Fourth European Ozone Symposium, Schliersee, Bavaria, Germany, edited by: Harris, N. R. P., Kilbane-Dawe, T., and Amantidis, G. T., p. 589, European Communities, 1998.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.4

### NO (g) + H<sub>2</sub>SO<sub>4</sub> (aq) → products

#### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma < 5.0 \times 10^{-6}$	70	193–243	Saastad et al. (1993)	static (a)

#### 5 Comments

(a) Measurement of the total pressure drop in a static system over 70% H<sub>2</sub>SO<sub>4</sub>-H<sub>2</sub>O monitored by MS. The solution was believed to be a supercooled solution from its visual appearance. Total pressure was about 10<sup>-2</sup> mbar, NO pressures 10<sup>-5</sup>–10<sup>-2</sup> mbar.

#### 10 Preferred values

Parameter	Value	T/K
$\gamma$	$< 5 \times 10^{-6}$ (70 % H <sub>2</sub> SO <sub>4</sub> )	200–298
<i>Reliability</i>		
$\Delta \log(\gamma)$	undetermined	

#### Comments on preferred values

The upper limit to  $\gamma$  from the single study of Saastad et al. (1993) was adopted for the recommendation.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## References

Saastad, O. W., Ellermann, T., and Nielsen, C. J.: Geophys. Res. Lett., 20, 1191–1193, 1993.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.5

### NO<sub>2</sub> (g) + H<sub>2</sub>SO<sub>4</sub> (aq) → products

#### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: γ</i>				
$\gamma < 5.0 \times 10^{-6}$		193–243	Saastad et al. (1993)	(a)
$(4.1 \pm 1.5) \times 10^{-7}$	45	298	Kleffmann et al. (1998)	Bubbler-IC/IR (b)
$(2.5 \pm 0.5) \times 10^{-7}$	70			
$(5.2 \pm 0.5) \times 10^{-7}$	96			
$\gamma < 7.0 \times 10^{-7}$	39–68	203–243	Langenberg et al. (1998)	(c)
<i>Solubility: H</i>				
$0.5 \pm 0.1$	39	243	Langenberg et al. (1998)	(c)
$1.4 \pm 0.5$	59			
$0.5 \pm 0.2$	68			
$50 \pm 7$	39	203		
$65 \pm 4$	59			
$37 \pm 6$	68			
$3.45 \times 10^{-9} \exp(4800/T)$	59			
$4.39 \times 10^{-9} \exp(4600/T)$	68			

#### 5 Comments

- (a) Measurement of the total pressure drop in a static system over 70 % H<sub>2</sub>SO<sub>4</sub>-H<sub>2</sub>O monitored by MS. From its appearance, the solution was believed to be super-cooled. NO<sub>2</sub> pressures were 10<sup>-5</sup>-10<sup>-2</sup> mbar.
- (b) NO<sub>2</sub> (10<sup>14</sup> molecule cm<sup>-3</sup>) in 740 Torr of N<sub>2</sub> was circulated through a thermostated bubbler containing 125 to 500 ml of water. NO<sub>2</sub> and HNO<sub>2</sub> were measured using tunable diode laser absorption spectrometry. Nitrate and nitrite in the liquid was measured using ion chromatography, showing that they were initially formed at

32340

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



equimolar amounts. The first order rate constant for  $\text{NO}_2$  decay did not depend on the  $\text{NO}_2$  concentration, indicating overall first-order behaviour, and was proportional to the estimated bubble surface area. This led the authors to conclude that  $\text{NO}_2$  uptake would be limited by a surface process. The increasing reactivity with increasing  $\text{H}_2\text{SO}_4$  concentration was explained by formation of  $\text{HNO}_2^+$  as intermediate.

- (c) Measurement of chromatographic retention of  $\text{NO}_2$  in a  $\text{H}_2\text{SO}_4$  coated, thermostated quartz capillary using a chemiluminescence detector. Equilibrium with  $\text{N}_2\text{O}_4$  was accounted for in the analysis. This experiment showed reactive loss of  $\text{NO}_2$  of up to 40 %.

### Preferred values

Parameter	Value	T/K
$\gamma$	$<10^{-6}$ ( $>45\%$ $\text{H}_2\text{SO}_4$ )	200–298
$H$	$3.45 \times 10^{-9} \exp(4800/T)$ (59 % $\text{H}_2\text{SO}_4$ ) $4.39 \times 10^{-9} \exp(4600/T)$ (68 % $\text{H}_2\text{SO}_4$ )	200–250
<i>Reliability</i>		
$\Delta \log(\gamma)$	undetermined	
$\Delta \log(H)$	$\pm 0.5$	200–250

### Comments on preferred values

In view of the uncertainties related to the mass transfer characteristics pertinent to bubbler experiments (Lee and Schwartz, 1981; Cheung et al., 2000) it is not clear to what degree the experiment by Kleffmann et al. (1998) was solubility limited. Langenberg et al. (1998) provide an upper bound of the reactive uptake coefficient consistent with these values. We therefore used these to recommend an upper limit to the up-

take coefficient and to extend the temperature range over that provided by Saastad et al. (1993).

No recommendation for the bulk accommodation coefficient is given. We adopt the temperature dependent solubility reported by Langenberg et al. for the two higher acid concentrations. The temperature dependence corresponds to an enthalpy of solvation of  $-39.9$  and  $-38.5$  kJ mol<sup>-1</sup> for the 59 and 68 wt % solutions, respectively.

## References

Cheung, J. L., Li, Y. Q., Boniface, J., Shi, Q., Davidovits, P., Worsnop, D. R., Jayne, J. T., and Kolb, C. E.: J. Phys. Chem. A, 104, 2655–2662, 2000.

Kleffmann, J., Becker, K. H., and Wiesen, P.: Atmos. Environ., 32, 2721–2729, 1998.

Lee, Y.-N. and Schwartz, S. E.: J. Phys. Chem., 85, 840–848, 1981.

Langenberg, S., Proksch, V., and Schurath, U.: Atmos. Environ., 32, 3129–3137, 1998.

Saastad, O. W., Ellermann, T., and Nielsen, C. J.: Geophys. Res. Lett., 20, 1191–1193, 1993.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.6

### NO<sub>3</sub> + H<sub>2</sub>SO<sub>4</sub> → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$<1 \times 10^{-3}$ (60–95 wt % H <sub>2</sub> SO <sub>4</sub> )	170–200	Fenter and Rossi (1997)	Knudsen-LIF (a)

#### 5 Comments

(a) Uptake of NO<sub>3</sub> ( $10^{10}$ – $10^{13}$  molecule cm<sup>-3</sup>) to 0–95 wt % H<sub>2</sub>SO<sub>4</sub>. A flow of H<sub>2</sub>O was added to maintain the H<sub>2</sub>SO<sub>4</sub> composition over the course of an experiment. NO<sub>3</sub> was formed in the thermal decomposition of N<sub>2</sub>O<sub>5</sub>. NO<sub>3</sub> losses in the Knudsen reactor were attributed to reaction with a Halocarbon wax coating, so that a value of  $\gamma < (6.0 \pm 6.8) \times 10^{-4}$  was derived for interaction with H<sub>2</sub>SO<sub>4</sub>. The authors quote a final upper limit of  $1 \times 10^{-3}$ .

#### Preferred values

Parameter	Value	T/K
$\gamma$	$<1 \times 10^{-3}$ (60–95 wt % H <sub>2</sub> SO <sub>4</sub> )	200–270
<i>Reliability</i>		
$\Delta \log(\gamma)$	undetermined	

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred value

The single study (Fenter and Rossi, 1997) of the interaction of  $\text{NO}_3$  with  $\text{H}_2\text{SO}_4$  at various concentrations and temperatures derived an upper limit to the uptake coefficient, which we adopt as the preferred value.

## 5 References

Fenter, F. F. and Rossi, M. J.: J. Phys. Chem., 101, 4110–4117, 1997.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





**HNO<sub>2</sub> + H<sub>2</sub>SO<sub>4</sub> (aq) → products****Experimental data**

<i>Parameter</i>	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma_{SS} = (1.60 \pm 0.09) \times 10^{-2}$	65.3	213.5	Zhang et al. (1996)	CWFT-CIMS (a)
$\gamma_{SS} = (3.2 \pm 0.5) \times 10^{-2}$	68.8	218.6		
$\gamma_{SS} = (6.6 \pm 0.6) \times 10^{-2}$	71.3	222.6		
$\gamma_{SS} = (9.1 \pm 1.6) \times 10^{-2}$	73.0	226.1		
$\gamma_{SS} = (2.0 \pm 0.5) \times 10^{-4}$	55	220	Fenter and Rossi (1996)	Kn-MS (b)
$\gamma_{SS} = (2.0 \pm 0.5) \times 10^{-3}$	60	215–235		
$\gamma_{SS} = (5.5 \pm 1.0) \times 10^{-2}$	80	250–265		
$\gamma_{SS} = 0.31 \pm 0.02$	95	220–273		
$\gamma = (2.8 \pm 1.4) \times 10^{-6}$	60	298	Baker et al. (1999)	AFT-CLD (c)
$\gamma = (8.1 \pm 2.7) \times 10^{-6}$	60			
<i>Solubility: <math>H^*</math> (M atm<sup>-1</sup>)</i>				
$(4.2 \pm 0.5) \times 10^2$	31.5	250	Becker et al. (1996)	Static-TDLAS/IC (d)
$(1.1 \pm 0.2) \times 10^2$	31.5	267		
$(3.1 \pm 0.5) \times 10^1$	31.5	283		
$(2.0 \pm 0.2) \times 10^1$	31.5	298		
$(3.1 \pm 0.5) \times 10^2$	51.2	249		
$(6.4 \pm 0.7) \times 10^1$	51.2	267		
$(1.7 \pm 0.3) \times 10^1$	51.2	283		
$7.3 \pm 0.1$	51.2	298		
$(4.3 \pm 0.2) \times 10^4$	67.3	251		
$(1.2 \pm 0.5) \times 10^4$	67.3	268		
$(4.1 \pm 1.0) \times 10^3$	67.3	283		
$(1.6 \pm 0.2) \times 10^3$	67.3	298		
$1.3 \times 10^7$	83	295	Longfellow et al. (1998)	RWFT-CIMS (e)
$9.8 \times 10^3$	71	295		
$1.2 \times 10^3$	66	296		
$1.5 \times 10^6$	79	295		
$2.8 \times 10^8$	79	250		
$4.6 \times 10^9$	79	230		
$2.6 \times 10^3$	65	270		
$9.0 \times 10^3$	65	249		

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
1.6 × 10 <sup>5</sup>	65	220		
6.1 ± 1.1	50.5	300.5	Baker et al. (1999)	Bubbler-CLD (c)
10.2 ± 1.1	57.2	302		
14 ± 2	58.2	303		
15 ± 3	59.1	299		
25 ± 4	59.3	300		
38 ± 6	61.0	300		
<i>rate constants k<sub>II</sub> (M<sup>-1</sup> s<sup>-1</sup>)</i>				
300 ± 40	50.5	300.5	Baker et al. (1999)	Bubbler-CLD (c)
185 ± 40	56.5	301.5		
122 ± 15	57.2	302		
81 ± 10	58.2	303		
39 ± 5	59.1	299		
32 ± 4	59.3	300		
7.8 ± 0.8	61.0	300		
4.5 ± 0.5	61.4	298.5		
0.13 ± 0.01	65.4	300		

## Comments

- (a) CIMS detection of HNO<sub>2</sub> after reaction with SF<sub>6</sub><sup>-</sup>. 0.1 mm thick liquid H<sub>2</sub>SO<sub>4</sub>. The partial pressure of HNO<sub>2</sub> was around 6 × 10<sup>-7</sup> mbar. The sulphuric acid composition was controlled by maintaining a fixed temperature and H<sub>2</sub>O partial pressure of 6 × 10<sup>-4</sup> mbar in the gas flows. Persistent uptake over more than 1 h was interpreted as formation of NO<sup>+</sup>HSO<sub>4</sub><sup>-</sup> in the film. Its precipitation as a solid was not observed.
- (b) H<sub>2</sub>SO<sub>4</sub> solutions prepared by dilution of 95 wt %. Saturation of the HNO<sub>2</sub> uptake takes place after deposition of 3 % monolayer. The uptake appears to be reversible.
- (c) Typically 10<sup>7</sup> particles cm<sup>-3</sup> of 120 nm diameter were entrained in an aerosol laminar flow tube and interacted with HNO<sub>2</sub> in the concentration range 305 to 610 ppb. Aerosol diagnostics included a counter and an optical extinction measurement in the 200 to 390 nm range whereas the gas phase was monitored using a NO<sub>x</sub>

32346

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



chemiluminescence analyzer. The observed uptake coefficient was found to be roughly proportional to the partial pressure of  $\text{HNO}_2$  in these experiments. Separate bubbler experiments were performed to measure the bimolecular rate constant for the reaction  $2\text{HNO}_2 \Rightarrow \text{NO} + \text{NO}_2 + \text{H}_2\text{O}$ . The rate constant was found to fall off sharply beyond 56 wt % sulphuric acid. The rate constant was not sufficient to explain the observed uptake of  $\text{HNO}_2$  to aerosol particles by reaction in the bulk, so that a surface reaction mechanism was invoked.

(d) Solubility of  $\text{HNO}_2$  in a 11-L Pyrex glass reactor was measured directly by monitoring both gas phase composition by tunable diode laser spectrometry and liquid phase by ion chromatography. Formation of  $\text{NO}_2$  due to  $2\text{HNO}_2 = \text{NO} + \text{NO}_2 + \text{H}_2\text{O}$  was observed. Since the  $\text{NO}_2$  concentration equilibrated with time, this reaction was included into calculating the effective solubility. A further, very slow overall loss of  $\text{HNO}_2$  and concomitant formation of  $\text{N}_2\text{O}$  with time was also observed.

(e) Determination of solubility was performed in a 2.2 cm i.d. coated wall flow tube and a 1.84 cm i.d. rotating wetted wall flow tube with CIMS detection of  $\text{HNO}_2$ . Depending on the sulphuric acid concentration range, solubility could be determined in relative or absolute modes. No kinetics of uptake of  $\text{HNO}_2$  alone was reported.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.05	220–300
$A$ (M atm <sup>-1</sup> )	$4.2 \times 10^{-6} \exp(4873/T)$	
$B$ ((wt %) <sup>-1</sup> )	$13.16/T - 0.0856$	
$C$ (M atm <sup>-1</sup> )	$2.0 \times 10^8 \exp(-14000/T)$	
$D$ ((wt %) <sup>-1</sup> )	$297.3/T - 0.474$	
$k_b$ (M <sup>-1</sup> s <sup>-1</sup> )	320	298
$\gamma_{gs}$	$<10^{-5}$	298
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	undetermined	
$\Delta \log(H^*)$	$\pm 0.3$	220–300
$\Delta \log(k_b)$	$\pm 0.1$	298
$\Delta \log(\gamma_{gs})$	undetermined	

## Comments on preferred values

Solubility of HNO<sub>2</sub> in H<sub>2</sub>SO<sub>4</sub> solutions depends strongly on composition. The exponential decrease with sulphuric acid content up to about 50 wt % is due to a salting out effect at decreased water activity (Becker et al., 1996). Dissociation of HNO<sub>2</sub> into H<sup>+</sup> and NO<sub>2</sub><sup>-</sup> is only important below 1 wt % (see datasheet V.A1.7; Park and Lee, 1988). Above about 55 wt %, solubility of HNO<sub>2</sub> increases strongly due to the formation of either of NO<sup>+</sup>, H<sub>2</sub>ONO<sup>+</sup>, NO<sup>+</sup>HSO<sub>4</sub><sup>-</sup>, or a combination thereof (Burley and Johnston, 1992, and references therein; Riordan et al., 2005). The available studies determining the effective solubility of HNO<sub>2</sub> directly by Becker et al. (1996), Longfellow et al. (1998) and Baker et al. (1999) agree very well as far as they overlap in composition and temperature. We recommend using an expression based on a parameterisation proposed

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

by Becker et al. (1996):

$$H^*(\text{Matm}^{-1}) = A \exp(B(wt)) + C \exp(D(wt))$$

The preferred constants  $A$ ,  $B$ ,  $C$  and  $D$  listed in the table were determined by further taking into account the data at lower temperature and higher wt% by Longfellow et al. (1998). While the low wt% part represented by the first term in the expression for  $H^*$  can be traced back to the salting coefficient based on the Setschenow equation (see Becker et al. 1996 for details), the other parameters were adjusted to broadly fit the available data without explicitly accounting for equilibria with  $\text{NO}^+$ ,  $\text{H}_2\text{ONO}^+$  and  $\text{NO}^+\text{HSO}_4^-$ .

Uptake of  $\text{HNO}_2$  into sulphuric acid solutions can be driven by the effective solubility encompassing dissociation, protonation and the disproportionation into  $\text{NO}_2$  and  $\text{NO}$ , assuming that the time scale for equilibration is given by diffusion in the liquid phase:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{sol}}} \quad \text{with } \Gamma_{\text{sol}} = \frac{4 \cdot H_{\text{HNO}_2}^* RT}{\bar{c}_{\text{HNO}_2} \cdot \sqrt{\pi}} \cdot \sqrt{\frac{D_{\text{l,HNO}_2}}{t}} \quad (1)$$

The diffusion coefficient for  $\text{HNO}_2$  is parameterized by  $D_{\text{l,HNO}_2} = c_{\text{HNO}_2} T / \eta$ ; with  $c_{\text{HNO}_2} = 6.90 \times 10^{-8} \text{ cm}^2 \text{ cP K}^{-1} \text{ s}^{-1}$ , estimated as suggested by Klassen et al. (1998) using a molar volume of  $36 \text{ cm}^3/\text{mol}$  (da Silva et al., 2006). For the viscosity, we suggest using the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into tropospherically more relevant dilute solutions at high  $T$ :

$$\eta = aT^{-1.43} \exp(448\text{K}/(T - T_0)), \quad (2)$$

with  $a = 169.5 + 5.18 wt - 0.0825 wt^2 + 3.27 \times 10^{-3} wt^3$ ,

and  $T_0 = 144.11 + 0.166 wt - 0.015 wt^2 + 2.18 \times 10^{-4} wt^3$

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The kinetic experiments by Zhang et al. (1996) and Fenter and Rossi (1996) were likely limited by solubility and bulk accommodation. With  $\alpha_b > 0.05$  and the recommended values for the effective solubility, the above equation yields uptake coefficients consistent with those reported for times between 10 and 100 s.

The bulk reaction rate constant  $k_b$  for the disproportionation reaction was adopted from Baker et al. (1999) by assuming that the reduction of the effective bimolecular rate constant was due to the decreasing concentration of unprotonated  $\text{HNO}_2$  in solution. For the calculation of the concentration of unprotonated  $\text{HNO}_2$ , the first term of the recommended expression for the effective solubility can be used:

$$[\text{HNO}_2] = \rho_{\text{HONO}} A \exp(Bwt) \quad (3)$$

This leads to effective rates consistent with those measured by Baker et al. (1999). This rate constant is far too small to affect uptake of  $\text{HNO}_2$  to sulfuric acid appreciably over the short time scales of the kinetic experiments. The reaction of  $\text{HNO}_2$  with  $\text{NO}^+\text{HSO}_4^-$  to yield  $\text{N}_2\text{O}$  was not quantified in the study by Wiesen et al. (1995) in a way that would allow extracting a rate constant.

For the conditions of the aerosol flow tube experiments by Baker et al. (1999), gas phase  $\text{HNO}_2$  rapidly equilibrates with the solution of the submicron particles, so that the uptake coefficient was due to an irreversible loss process in the particle phase. Since the bulk reaction rate constant  $k_b$  was several orders of magnitude too low to explain the measured uptake coefficients, uptake was attributed to a surface process. Baker et al. suggest an Eley-Rideal type reaction, which was adopted here for a recommendation of an upper limit for  $\gamma_{\text{gs}}$  in the atmospherically relevant  $\text{HNO}_2$  concentration range.

## References

- Baker, J., Ashbourn, S. F. M., and Cox, R. A.: Phys. Chem. Chem. Phys., 1, 683–690, 1999.
- Becker, K. H., Kleffmann, J., Kurtenbach, R., and Wiesen, P.: J. Phys. Chem., 100, 14984–14990, 1996.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



da Silva, G., Dlugogorski, B. Z., and Kennedy, E. M.: AICHE Journal, 52, 1558–1565, 2006.

Fenter, F. F. and Rossi, M. J.: J. Phys. Chem., 100, 13765–13775, 1996.

Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.

5 Longfellow, C. A., Imamura, T., Ravishankara, A. R., and Hanson, D. R.: J. Phys. Chem. A, 102, 3323–3332, 1998.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

10 Wiesen, P., Kleffmann, J., Kurtenbach, R., and Becker, K. H.: Faraday Discuss., 100, 121–127, 1995.

Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

Zhang, R., Leu, M.-T., and Keyser, L. F.: J. Phys. Chem., 100, 339–345, 1996.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A4.8

 $\text{HNO}_3 + \text{H}_2\text{SO}_4 \rightarrow \text{products}$ 

## Experimental data

Parameter	$[\text{H}_2\text{SO}_4]/\text{wt } \%$	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma_{\text{SS}} = 2.0 \times 10^{-3}$	75	230	Tolbert et al. (1988)	Kn-MS (a)
$\gamma_{\text{SS}} = 4.0 \times 10^{-3}$	96.5	295		
$\gamma = 0.28$	5	188	Reihs et al. (1990)	Kn-MS (b)
$\gamma = 0.25$	40	188		
$\gamma = 0.12$	58	188		
$\gamma = 0.07$	74	188		
$\gamma = 0.05$	83	188		
$\gamma = 0.001$	96	188		
$\gamma = 0.07 \pm 0.01$	15–66	223		
$\gamma = 0.015$	74	223		
$\gamma = 0.003$	87	223		
$\gamma = 0.11 \pm 0.01$	73	283	Van Doren et al. (1991)	DT-IR (c)
<i>Accommodation coefficients: <math>\alpha</math></i>				
$\alpha_{\text{s}} > 0.8$ (70 wt % $\text{D}_2\text{SO}_4$ )		213	Morris et al. (2000)	MB (d)
$\alpha_{\text{b}} > 0.8$ (70 wt % $\text{D}_2\text{SO}_4$ )				
<i>Solubility: <math>H</math></i>				
$H^* = 3.56 \times 10^{-3} \exp(3320/T)$	87	188–240	Reihs et al. (1990)	(b)
$H^* = 8.54 \times 10^{-3} \exp(3550/T)$	74			
$H^* = 2.02 \times 10^{-1} \exp(3190/T)$	66			
$H^* = 7.47 \times 10^{-8} \exp(7160/T)$	58			
$(4 \pm 1) \times 10^3 \text{ M atm}^{-1}$	73	283	Van Doren et al. (1991)	(c)
$H^* = 9.03 \times 10^{-8} \exp(7750/T)$	40	190–230	Zhang et al. (1993)	(e)
$H^* = 4.85 \times 10^{-8} \exp(7470/T)$	50			
$H^* = 2.44 \times 10^{-8} \exp(7300/T)$	60			
$H^* = 6.32 \times 10^{-9} \exp(7240/T)$	70			
$H^* = 6.13 \times 10^{-9} \exp(7030/T)$	75			

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Comments

- (a) The bulk substrate was cooled slowly to temperatures of 210 K to 230 K. The recovery yield of  $\text{HNO}_3$  that was condensed at 230 K on 75 %  $\text{H}_2\text{SO}_4$  was approximately 20 % in contrast to 100 % for pure ice.
- (b) The sulphuric acid substrate was prepared from bulk solutions and cooled to the temperature range 188 K to 240 K.  $\gamma$  was time dependent, the uptake coefficients given in the table are those measured after 100 s (the experiments spanning a total time of 2000 s). The time dependence was used to estimate the effective solubility assuming solubility limited uptake. The diffusion coefficient in the liquid phase was extrapolated from a room temperature value of  $10^{-5} \text{ cm}^2 \text{ s}^{-1}$  in 60 %  $\text{H}_2\text{SO}_4$  using temperature dependent viscosity data from the literature. The error in  $H^*$  is estimated to be a factor of 3. The expressions given in the table were obtained from fits to a Van't Hoff plot.
- (c) Fast droplet train (73 %  $\text{H}_2\text{SO}_4$ - $\text{H}_2\text{O}$  droplets) traversing a flow tube.  $\text{HNO}_3$  was detected by TDLAS. The  $\gamma$  values were corrected for gas-phase diffusion. The solubility was obtained by assuming that the measured uptake coefficient is due to solubility limited uptake. The bulk accommodation coefficient was assumed to be independent of  $\text{H}_2\text{SO}_4$  concentration and similar to that of water (0.17). The liquid phase diffusion coefficient was estimated to be  $9 \times 10^{-7} \text{ cm}^2 \text{ s}^{-1}$  at 283 K based on the estimated temperature dependence of viscosity data.
- (d) Molecular beam experiment with  $\text{HNO}_3$  beam produced from the expansion of a 1 %  $\text{HNO}_3$  in  $\text{H}_2$  mixture, leading to an incident  $\text{HNO}_3$  energy  $150 \text{ kJ/mol}$  hitting the 70.5 %  $\text{D}_2\text{SO}_4$  at 213 K. Fluxes were estimated to be below  $10^{15} \text{ cm}^{-2} \text{ s}^{-1}$ . Scattered or desorbing molecules were detected with a mass spectrometer at an exit angle of  $45^\circ$ . A large fraction of incident  $\text{HNO}_3$  molecules transfers their kinetic energy to the substrate; the inelastically scattered  $\text{HNO}_3$  molecules lost 85 % of their original energy on average. We use this as a lower bound to  $\alpha_s$  reported in

32353

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



the table, even though the authors do not report the trapping probability for HNO<sub>3</sub>. They argue that it is close to 1 at low, atmospherically relevant incident energies. 95 % of those molecules trapped on the surface undergo proton exchange in the bulk of D<sub>2</sub>SO<sub>4</sub>. Therefore  $\alpha_b$  is listed in the table with the same lower bounds.

- 5 (e) Static vapour pressure measurement of HNO<sub>3</sub> (MS) over a stirred H<sub>2</sub>SO<sub>4</sub>/HNO<sub>3</sub>/H<sub>2</sub>O solution.

### Preferred values

Parameter	Value	T/K
$\alpha_s$	>0.8	188–223
$\alpha_b$	>0.8	188–223
$m_1$ (wt % <sup>-2</sup> K)	0.14	190–300
$m_2$ (wt % <sup>-1</sup> K)	-36	190–300
$m_3$ (K)	8980	190–300
$b_1$ (wt % <sup>-2</sup> )	0.00063	190–300
$b_2$ (wt % <sup>-1</sup> )	0.012	190–300
$b_3$	14.7	190–300
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	undetermined	
$\Delta \log(\alpha_s)$	undetermined	
$\Delta \log(H^*)$	±0.5	190–300

### Comments on preferred values

- 10 The molecular beam experiment by Morris et al. (2000) provides the most direct picture of HNO<sub>3</sub> interacting with the surface of sulphuric acid. Most of the HNO<sub>3</sub> molecules colliding with the surface are trapped, and nearly all of these undergo rapid proton exchange due to efficient solvation in the bulk. HNO<sub>3</sub> forms very stable complexes

32354

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



with  $\text{H}_2\text{O}$  and  $\text{H}_2\text{SO}_4$  that may also be important for the stabilization of  $\text{HNO}_3$  in the interfacial region (Yang and Finlayson-Pitts, 2001; Fairbrother and Somorjai, 2000). We therefore follow Morris et al. (2000) for recommending a lower limit to  $\alpha_s$  and  $\alpha_b$ .

In Knudsen cell and droplet train experiments, uptake of  $\text{HNO}_3$  into  $\text{H}_2\text{SO}_4$  solutions was time dependent and followed the behaviour expected for solubility limited uptake in the studies by Reihs et al. (1990) and Van Doren et al. (1991). The slow evaporation observed by Tolbert et al. (1988) with thermal desorption spectrometry appears to be consistent with this. Bulk accommodation was therefore not rate limiting the uptake. The decreasing trend with increasing sulphuric acid concentration reflects the increasing degree of solubility limitation at short interaction times. Initial uptake coefficients extracted from the time dependent traces are not reported that could be compared to the Morris et al. (2000) study.

Estimating Henry's Law constants from the time dependent uptake data carries a large uncertainty, as high experimental stability is required to allow reliable  $t^{1/2}$  fits to the data and also due to the uncertainty associated with estimating the temperature dependent diffusion coefficients. The solubility data extracted from the kinetic experiments by Reihs et al. (1990) and extrapolated to higher temperatures are consistent with vapour pressure measurements for the ternary  $\text{HNO}_3$ - $\text{H}_2\text{SO}_4$ - $\text{H}_2\text{O}$  system reported by Vandoni (1944) at 273 K. We use the Zhang et al. (1993) solubility data based on  $\text{HNO}_3$  vapour pressure measurements, which covered the temperature range from 230 to 190 K. We adopt a parameterization of their data in our table of preferred values in which:

$\log H^* = m/T - b$  where

$$m = m_1[\text{H}_2\text{SO}_4]^2 + m_2[\text{H}_2\text{SO}_4] + m_3 \text{ and } b = b_1[\text{H}_2\text{SO}_4]^2 + b_2[\text{H}_2\text{SO}_4] + b_3$$

and the  $\text{H}_2\text{SO}_4$  concentration  $[\text{H}_2\text{SO}_4]$  is in wt %

As pointed out by Taleb et al. (1996), at high temperature, Zhang's data are consistent with Vandoni's data. Taleb et al. (1996) also provide a more comprehensive model

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



to predict vapour pressure and activity of H<sub>2</sub>O, HNO<sub>3</sub> and H<sub>2</sub>SO<sub>4</sub>, which reproduces nicely the maximum in the HNO<sub>3</sub> vapor pressure of about 1:1 (mole fraction) HNO<sub>3</sub>-H<sub>2</sub>SO<sub>4</sub> mixtures reported by Vandoni (1944) and which is in closer agreement with Zhang's data than the work of Luo et al. (1995).

## 5 References

Carslaw, K. S., Clegg, S. L., and Brimblecombe, P.: J. Phys. Chem., 99, 11557–11574, 1995.

Fairbrother, D. H. and Somorjai, G. A.: J. Phys. Chem. B, 104, 4649–4652, 2000.

10 Luo, B. P., Carslaw, K. S., Peter, T., and Clegg, S. L.: Geophys. Res. Lett., 22, 247–250, 1995.

Morris, J. R., Behr, P., Antman, M. D., Ringeisen, B. R., Splan, J., and Nathanson, G. M.: J. Phys. Chem. A, 104, 6738–6751, 2000.

Reihs, C. M., Golden, D. M., and Tolbert, M. A.: J. Geophys. Res., 95, 16545–16550, 1990.

15 Taleb, D. E., Ponche, J. L., and Mirabel, P.: J. Geophys. Res., 101, 25967–25977, 1996.  
Tolbert, M. A., Rossi, M. J., and Golden, D. M.: Geophys. Res. Lett., 15, 847–850, 1988.

Van Doren, J. M., Watson, L. R., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 95, 1684–1689, 1991.

20 Vandoni, R.: Mem. Serv. Chim. Etat, 31, 87–111, 1944.

Yang, H. S. and Finlayson-Pitts, B. J.: J. Phys. Chem. A, 105, 1890–1896, 2001.

Zhang, R. Y., Wooldridge, P. J., and Molina, M. J.: J. Phys. Chem., 97, 8541–8548, 1993.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A4.9

### HO<sub>2</sub>NO<sub>2</sub> + H<sub>2</sub>SO<sub>4</sub> → products

#### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficient: <math>\gamma_0</math></i>				
0.23	58.3	207.9	Zhang et al. (1997)	CWFT-CIMS (a)
0.20	58.3	213.5		
0.13	58.3	218.9		
0.07	58.3	226.8		
<i>H* (D<sub>1</sub>)<sup>0.5</sup></i>				
2.95 × 10 <sup>-10</sup> exp (5940/T)	52.9	209–229	Zhang et al. (1997)	CWFT-CIMS (a)
1.37 × 10 <sup>-8</sup> exp (4980/T)	58.3	208–227		
1.31 × 10 <sup>-5</sup> exp (3320/T)	66.4	201–215		
1.75 × 10 <sup>-3</sup> exp (2030/T)	73.8	204–224		

#### Comments

(a) Coated-wall laminar flow reactor using CIMS detection of PNA using both SF<sub>6</sub><sup>-</sup> and I<sup>-</sup> as source ions. PNA is either detected as HO<sub>2</sub>NO<sub>2</sub>F<sup>-</sup> or HO<sub>2</sub>NO<sub>2</sub><sup>-</sup>. The first-order rate coefficient measured at a typical PNA concentration of 5 × 10<sup>7</sup> molecule cm<sup>-3</sup> was corrected for gas phase diffusion (10 % for the smallest and a factor of four for the largest  $\gamma$  values) using a gas phase diffusion constant of PNA in He of  $\rho D_1 = 319 \text{ mbar cm}^2 \text{ s}^{-1}$ . Liquid sulphuric acid films were approximately 0.1 mm thick. The uptake of PNA was reversible when the flow was halted. The time dependence of  $\gamma$  was interpreted as solubility limited uptake and used to extract  $H^*(D_1)^{0.5}$  values. Representative expressions for these are listed in the table.

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.2	205–225
$H^*$ (Matm <sup>-1</sup> )	$5.5 \times 10^{-23} \exp(wt(0.52 - 1.05 \times 10^{-4}(wt - 53)^2)) \times \exp((20683 - 346.29wt + 2.044wt^2)/T)$	205–225
$c$ (cm <sup>2</sup> cP K <sup>-1</sup> s <sup>-1</sup> )	$6.02 \times 10^{-8}$	205–225
$\Delta H_{\text{soln}}^0$ (kJ Mol <sup>-1</sup> )	-58.6	
$S_{\text{soln}}^0$ (J Mol <sup>-1</sup> K <sup>-1</sup> )	-159.0	
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	undetermined	
$\Delta \log(H)$	±0.15	205–225
$\Delta \Delta H_{\text{soln}}^0$ kJ Mol <sup>-1</sup>	±4.5	
$\Delta S_{\text{soln}}^0$ (J Mol <sup>-1</sup> K <sup>-1</sup> )	±15	

## Comments on preferred values

The only study available presents a careful analysis of time dependent uptake of PNA into liquid sulphuric acid films. The largest initial uptake coefficient reported forms the basis for a lower limit to  $\alpha_b$ . To obtain the recommended values for  $H^*$  from the  $H^*(D_1)^{0.5}$  values reported by Zhang et al. (1997), we estimate the diffusion coefficient based on the Wilke and Chang (1955) method, as suggested by Klassen et al. (1998) for a range of other species:

$$c = \frac{7.4 \times 10^{-8} \sqrt{K_{\text{solvent}}}}{V_A^{0.6}}$$

Klassen et al. (1998) found  $\gamma_{\text{solvent}} = 64$  to well represent H<sub>2</sub>SO<sub>4</sub> solutions in this concentration range. With the partial molar volume  $V_A$  of 45.183 cm<sup>3</sup> mol<sup>-1</sup>, we obtain the value for  $c$  listed above that can be used to calculate the diffusion coefficient via

32358

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$D_{\text{I,HO}_2\text{NO}_2} = cT/\eta$ . For the viscosity, we suggest to use the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into tropospherically more relevant dilute solutions at high  $T$ :

$$\eta = aT^{-1.43} \exp(448\text{K}/(T - T_0)),$$

with  $a = 169.5 + 5.18 wt - 0.0825 wt^2 + 3.27 \times 10^{-3} wt^3$ ,  
and  $T_0 = 144.11 + 0.166 wt - 0.015 wt^2 + 2.18 \times 10^{-4} wt^3$

The  $c$  values (and thus also the diffusion coefficients) are about a factor of 2 higher than those estimated by Zhang et al. (1997) based on another approach.

The effective solubilities of PNA in  $\text{H}_2\text{SO}_4$  are strongly  $T$ -dependent but only slightly dependent upon  $\text{H}_2\text{SO}_4$  concentration owing to the fact that PNA is a weak acid whose degree of dissociation is vanishingly small at relevant  $[\text{H}_2\text{SO}_4]$ .

The resulting  $T$ -dependence of  $H^*$  allowed the evaluation of the dissolution parameters of PNA according to the following reaction: (1)  $\text{HO}_2\text{NO}_2(\text{g}) \rightleftharpoons \text{HO}_2\text{NO}_2(\text{aq})$ ; (2)  $\text{HO}_2\text{NO}_2(\text{aq}) \rightleftharpoons \text{O}_2\text{NO}_2^- + \text{H}^+$ . The values of  $\Delta H_{\text{soln}}^0$  and  $S_{\text{soln}}^0$  varied from  $-62.8$  to  $-54.4 \text{ kJ mol}^{-1}$  and  $-174$  to  $-143.9 \text{ J mol}^{-1} \text{ K}^{-1}$  in the given  $\text{H}_2\text{SO}_4$  concentration range.

No reaction products were detected in agreement with the identical shape of the uptake and desorption trace of PNA.

## References

- Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.  
Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.  
Wilke, C. R. and Chang, P.: Aiche Journal, 1, 264–270, doi:10.1002/aic.690010222, 1955.  
Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.  
Zhang, R., Leu, M.-T., and Keyser, L. F.: J. Phys. Chem. A, 101, 3324–3330, 1997.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**NH<sub>3</sub> + H<sub>2</sub>SO<sub>4</sub> (aqueous sulphuric acid-aerosol) → products****Experimental data**

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{SS}</math>, <math>\gamma_0</math></i>			
0.3 (20 % H <sub>2</sub> SO <sub>4</sub> )	248–288	Swartz et al. (1999)	DT-UV/Vis (a)
0.8 (40 % H <sub>2</sub> SO <sub>4</sub> )			
1.0 (55 % H <sub>2</sub> SO <sub>4</sub> )			
1.0 (75 % H <sub>2</sub> SO <sub>4</sub> )			
0.98 ± 0.15 (40–65 % H <sub>2</sub> SO <sub>4</sub> )	293–297	Hanson and Kosciuch (2003)	AFT-CIMS (b)
1.04 ± 0.14 (15–40 % H <sub>2</sub> SO <sub>4</sub> )	287–292		

**Comments**

(a) Uptake experiment of gas phase NH<sub>3</sub> on train of droplets whose size was in the range 150–300 μm entrained in a flowing mixture of helium and water vapour (between 2.9 and 23.3 mbar). The reaction time was between 2 and 15 ms, the ammonia concentration was in the range 10<sup>13</sup> to 2 × 10<sup>14</sup> cm<sup>-3</sup> and was monitored using a VUV lamp emitting at λ = 121.6 nm. The cited uptake coefficients were corrected for gas diffusion effects. The uptake coefficient increases as a function of acid concentration and reaches unity at about 55 wt % H<sub>2</sub>SO<sub>4</sub>. The distinct negative temperature dependence of γ observed in aqueous solutions (Shi et al., 1999) decreases with increasing concentration in the range 20–70 % H<sub>2</sub>SO<sub>4</sub>. The increased NH<sub>3</sub> uptake in acid solution is apparently due to reaction between NH<sub>3</sub><sup>+</sup> and H<sup>+</sup> at the gas-liquid interface. The results yielded parameters required to model the reaction of NH<sub>3</sub> with H<sup>+</sup> at the gas-liquid interface. These uptake experiments were expanded to include a detailed study of gas transport to a moving train of droplets.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





(b) Uptake of 1–12 ppb NH<sub>3</sub> in N<sub>2</sub> on 15–65 % H<sub>2</sub>SO<sub>4</sub> aerosol flowing in an atmospheric pressure flow tube (810 mbar) which was characterised using a DMA/CNC combination. The particle size distribution peaked close to 100 nm at a particle number concentration in the range 4 × 10<sup>4</sup> to 1.5 × 10<sup>5</sup> particle cm<sup>-3</sup> to afford a NH<sub>3</sub>-H<sub>2</sub>SO<sub>4</sub> ratio of typically 0.10. NH<sub>3</sub> was monitored using SIMS of protonated water clusters of the type NH<sub>4</sub><sup>+</sup>(H<sub>2</sub>O)<sub>m</sub> with m between 1 and 5. Corrections to the data have been made owing to an important NH<sub>4</sub><sup>+</sup> and a smaller NH<sub>3</sub> background signal.

### Preferred values

Parameter	Value	T/K
$\gamma$ ( $\geq 50$ % H <sub>2</sub> SO <sub>4</sub> )	1.0	265–300
<i>Reliability</i>		
$\Delta \log(\gamma)$	±0.3	265–300 ?

### Comments on preferred values

Uptake of ammonia on sulphuric acid surfaces results in reaction to form NH<sub>4</sub><sup>+</sup> and HSO<sub>4</sub><sup>-</sup>. The available data for [H<sub>2</sub>SO<sub>4</sub>] above 50 wt % are in agreement, showing an uptake coefficient close to unity. However for more dilute solutions, the results of uptake studies on fine aerosols (Hanson and Kosciuch, 2003) disagree with those from the droplet train experiments, which show a decline in  $\gamma$  with increasing water content below 45 wt %. In view of these apparent inconsistencies we confine our recommendation to concentrated acid ( $\geq 50$  wt %). The possible causes for this discrepancy at lower acid concentrations have been the subject of literature debate (Worsnop et al., 2004; Hanson and Kosciuch, 2004; Hanson et al., 2004; Morita et al., 2004a). One possible cause lies in the use of an inadequate gas resistance model in the interpretation of results from the droplet train experiments, which is discussed by Morita et al. (2004b)

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



in their study of the uptake of H<sub>2</sub>O on water surfaces. A clear resolution will require further investigation.

The uptake coefficients all exceed the bulk accommodation coefficient,  $\alpha_b$ , for uptake on pure water, and may indicate a direct chemical interaction of NH<sub>3</sub> at the surface, e.g. by protonation.

## References

Hanson, D. and Kosciuch, E.: J. Phys. Chem. A, 107, 2199–2208, 2003.

Hanson, D. R. and Kosciuch, E.: J. Phys. Chem. A, 108, 8549–8551, 2004.

Hanson, D. R., Sugiyama, M., and Morita, A.: J. Phys. Chem. A, 108, 3739–3744, 2004.

Morita, A., Sugiyama, M., Kameda, H., Koda, S., and Hanson, D. R.: J. Phys. Chem. B, 108, 9111–9120, 2004a.

Morita, A., Sugiyama, M., and Koda, S.: J. Phys. Chem. A, 108, 8544–8545, 2004b.

Shi, Q., Davidovits, P., Jayne, J. T., Worsnop, D. R., and Kolb, C. E.: J. Phys. Chem. A, 103, 8812–8823, 1999.

Swartz, E., Shi, Q., Davidovits, P., Jayne, J. T., Worsnop, D. R., and Kolb, C. E.: J. Phys. Chem. A, 103, 8824–8833, 1999.

Worsnop, D. R., Williams, L. R., Kolb, C. E., Mozurkewich, M., Gershenson, M., and Davidovits, P.: J. Phys. Chem. A, 108, 8546–8548, 2004.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A4.11

 $\text{N}_2\text{O}_5 + \text{H}_2\text{O}$  (aqueous sulphuric acid aerosol)

## Experimental data

Parameter	RH/%	H <sub>2</sub> SO <sub>4</sub> /wt%	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math>, <math>\gamma_{SS}</math>, <math>\gamma_0</math></i>					
0.10 ± 0.01	1–10		293	Mozurkewitch and Calvert (1988)	AFT-CL (a)
0.139 ± 0.009	1–10		274		
0.12 ± 0.03		40	226	Hanson and Ravishankara (1991a)	WWFT-CIMS (b)
0.14 ± 0.03		60			
0.10 ± 0.02		70			
0.10 ± 0.02		75			
0.058 ± 0.006		73	283	Van Doren et al. (1991)	DT-TDLAS (c)
0.055 ± 0.010		74–96	220 ± 5	Williams et al. (1994)	Kn-MS (d)
0.06 ± 0.04		79–84	293	Fried et al. (1994)	AFT-CL (e)
0.103 ± 0.006		64–81	273		
0.12 ± 0.016		57–76	260		
0.148 ± 0.011		63–76	247		
0.086 ± 0.009		61–63	247		
0.146 ± 0.036		69–71	230–234		
0.102 ± 0.011		65–69	231–234		
0.077 ± 0.019		54–64	225–231		
0.076 ± 0.018	60		270	Hanson and Lovejoy (1994)	AFT-CIMS (f)
0.081 ± 0.032	60		230		
0.061 ± 0.013	70		296		
0.109 ± 0.024	70		230		
0.077 ± 0.016	80		270		
0.090 ± 0.020	80		230		
0.11		53	200	Zhang et al. (1995)	WWFT-CIMS (g)
0.034		29, 16.4 % HNO <sub>3</sub>	195		
0.021		5, 41 % HNO <sub>3</sub>	200		
0.045		40, 10% HNO <sub>3</sub>	200		
0.10		5, 41 % HNO <sub>3</sub>	220		
0.075 ± 0.047		65	210–230	Beichert and Finlayson-Pitts (1996)	Kn-MS (h)
0.023 ± 0.004	90	17	297	Hu and Abbatt (1997)	AFT-CIMS/OPC (i)
0.038 ± 0.012	58.5	40			
0.060 ± 0.020	19.5	59			

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Parameter	RH/%	H <sub>2</sub> SO <sub>4</sub> /wt%	Temp./K	Reference	Technique/Comments
0.050 ± 0.008	8.5	66			
0.187 ± 0.028		69	240	Robinson et al. (1997)	DT-TDLAS (j)
0.154 ± 0.023		69	260		
0.159 ± 0.024		39	231		
0.086 ± 0.0129		39	255		
0.11 ± 0.02		60	210–230	Hanson (1997)	WWFT and AFT-CIMS (k)
0.092		55, 5% HNO <sub>3</sub>	202		
0.021		45, 15% HNO <sub>3</sub>	230		
0.045		59, 0.08% HNO <sub>3</sub>			
0.10		58, 1–2% HNO <sub>3</sub>			
0.033 ± 0.004	8–80	26.3–55	298	Hallquist et al. (2000)	AFT-CL/SMPS (l)
0.028 ± 0.008	40	49	298		
0.036 ± 0.004	40	48	288		
0.049 ± 0.009	40	44	278		
0.091 ± 0.019	40	42	268		
0.092 ± 0.013	40	42	263		
0.045	0.7	81	295 ± 1	Kane et al. (2001)	AFT-CIMS/SMPS (m)
0.045 ± 0.013	2.9	76			
0.030	7.7	70			
0.075	16	63			
0.061	28.4	55			
0.050		38 ± 2	193.6	Wagner et al. (2005)	SR-FTIR (n)
0.044		32, 10% HNO <sub>3</sub>			
0.033		17, 34% HNO <sub>3</sub>			
0.016		9, 48% HNO <sub>3</sub>			
		58, 1–2% HNO <sub>3</sub>			

## Comments

(a) Atmospheric pressure aerosol flow tube with N<sub>2</sub>O<sub>5</sub> ( $\approx 10^{13}$  molecule cm<sup>-3</sup>) measured by a modified chemiluminescence method, via thermal dissociation to NO<sub>3</sub> and titration with NO, which was detected. Aerosols generated in a constant output atomiser, dried, and size selected with a differential mobility analyser (DMA) coupled to a condensation particle counter (CPC). The monodisperse aerosol was then equilibrated at controlled humidity before entry into the flow tube. The typical diameter,  $d_{\text{mean}}$  was 0.08–0.2 μm, with surface area density of  $1\text{--}5 \times 10^{-5}$  cm<sup>2</sup>/cm<sup>3</sup>. Uptake coefficients were determined from the first order rate

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss, which were linearly dependent on surface area. Diffusion limitation was negligible for the size range used. The uptake of  $\text{N}_2\text{O}_5$  on  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  was independent of RH in the range 1–10 %.

- (b) Wetted wall flow tube coupled to CIMS detection. Aqueous  $\text{H}_2\text{SO}_4$  film residence time of 20–30 s.  $p(\text{H}_2\text{O})$  was  $\sim 1.3 \times 10^{-3}$  mbar. The temperature dependence of  $\gamma$  as measured for the 60 % and 70 solution: none was found within the reported error limits.
- (c) Fast train of 200  $\mu\text{m}$  73 %/ $\text{H}_2\text{O}$  droplets traversing a flow tube with TDLAS detection. Pressures = 13.3 mbar and droplet-gas interaction times of 1–2 ms.
- (d) Knudsen cell technique using MS detection.  $p(\text{N}_2\text{O}_5)$  was 0.4 to  $10 \times 10^{11}$  molecule  $\text{cm}^{-3}$ .  $\gamma$  value of 0.02–0.03 – independent of temperature, concentration (74 to 96 %  $\text{H}_2\text{O}$ ).  $\gamma$  value corrected for saturation effects.
- (e) High pressure (0.3 to 0.8 atm) flow tube using slow flow conditions and submicron  $\text{H}_2\text{SO}_4$  aerosol generated by homogeneous nucleation from the reaction of  $\text{SO}_3 + \text{H}_2\text{O}$ . Particle size 60–250 nm diameter at 225 to 293 K, with  $\text{H}_2\text{SO}_4$  54 to 82 wt %  $\text{N}_2\text{O}_5$  (30 to 200 ppb) detected by titrating with NO in a heated quartz tube, using TDLAS to monitor NO and also  $\text{H}_2\text{O}$  vapor.
- (f) Kinetics of  $\text{N}_2\text{O}_5$  hydrolysis on sulfuric acid aerosol measured in laminar flow reactor at 825 mbar of  $\text{N}_2$ , using CIMS detection. The aerosol surface area was determined by UV extinction using Mie theory. The  $\text{H}_2\text{SO}_4$  particle concentration were in the range 5–100  $\times 10^4$  particles  $\text{cm}^{-3}$  with diameters between 0.2 to 0.4  $\mu\text{m}$ . Initial  $\text{N}_2\text{O}_5$  concentration was of the order of 1 to  $5 \times 10^{13}$  molecule  $\text{cm}^{-3}$ . The uptake measurements on 60 %  $\text{H}_2\text{SO}_4$  at 230 K may have been influenced to some extent by  $\text{HNO}_3$  uptake and hence partial saturation.
- (g) Wetted wall flow reactor – CIMS. Liquid sulfuric acid films were applied to the cooled wall of a horizontally mounted tube. Total pressure = 0.53 mbar of He

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



with partial pressure of  $\text{N}_2\text{O}_5$  of  $6.7 \times 10^{-7}$  mbar. The  $\gamma$  values decrease with increasing  $\text{HNO}_3$  content hence decreasing temperature at constant  $\text{H}_2\text{O}$  and  $\text{HNO}_3$  partial pressure. On the binary  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  surface  $\gamma = 0.1$  independent of temperature. In the temperature range 195 to 220 K  $\gamma$  may be expressed as  $\gamma = -0.379 + 0.00213T$  at  $p_{\text{H}_2\text{O}} = 5 \times 10^{-7}$  mbar and  $p_{\text{HNO}_3} = 6.7 \times 10^{-7}$  mbar.

- (h) Knudsen flow reactor using MS detection.  $\text{N}_2\text{O}_5$  concentration was in the range 3 to  $90 \times 10^{11}$  molecule  $\text{cm}^{-3}$ .
- (i) Atmospheric pressure aerosol flow tube with detection of  $\text{N}_2\text{O}_5$  ( $7 \times 10^{12}$  molecule  $\text{cm}^{-3}$ ) by CIMS using  $\text{I}^-$  reagent ion. The aerosols were generated in an ultrasonic nebuliser and were equilibrated with the ambient humidity before entry into the flow tube. The size distribution, measured with an optical particle counter (OPC), was used to calculate the surface area of the aerosol in the flow tube. The peak in the surface area distribution,  $d_{\text{max}}$  was between 2 and 4  $\mu\text{m}$ . The counter was calibrated by collection of aerosol of known composition in an aqueous trap (assumed 100% efficient) and measurement of the electrical conductivity of the trapped electrolyte. Uptake coefficients were determined from the first order rate constants for  $\text{N}_2\text{O}_5$  decay, corrected for wall loss (Brown correction), and for diffusion limitation to the particle surface assuming an average diameter,  $d_{\text{max}}$  of the non-monodisperse aerosol. The first order rate constant for the title reaction scaled linearly with the aerosol surface area. Cited errors on  $\gamma$  is  $\pm 1\sigma$  precision; the estimated potential systematic error arising mainly from measurement of the SA was  $\pm 25\%$ .  $\text{HNO}_3$  accumulation in the aerosol was insignificant.
- (j) Fast train of 200  $\mu\text{m}$   $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  droplets traversing a flow tube with TDLAS detection. Pressures = 13.3 mbar and droplet-gas interaction times of 2 to 20 ms. The temperature of the droplets was inferred from the water partial pressure measured by TDL absorption. A negative temperature dependence of  $\gamma_{\text{ss}}$  was

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



observed for  $T \geq 230$  K and  $\gamma_{ss}$  slightly increased with increasing concentration of  $H_2SO_4$  measured at 39, 54 and 69%. Model involving neutral and acid catalysed mechanism for  $N_2O_5$  hydrolysis fitted data from several laboratories.

- (k) Flow reactors with sulfuric acid wall film (0.2 mm thickness) and sub-micron aerosol (particle size 0.1  $\mu m$ ) – CIMS detection. The pressure was 0.5 mbar He in the wall coated tube and 240 mbar of  $N_2$  in the aerosol flow tube. Initial  $[N_2O_5]$  was  $5 \times 10^{11}$  molecule  $cm^{-3}$  for the aerosol and  $5 \times 10^{11}$  molecule  $cm^{-3}$  for the bulk liquid experiment.
- (l)  $N_2O_5$  uptake experiment in an aerosol flow tube at characteristic  $N_2O_5$  concentrations of 500 ppb coupled to  $NO_2$  detection using a chemiluminescence detector after thermal decomposition of  $N_2O_5$  and following  $NO_3$  titration with  $NO$ . Typical aerosol number densities were  $(6-10) \times 10^5$  particles  $cm^{-3}$ . The  $\gamma$  values were independent of RH in the range 8–80 %, and hence on aerosol  $H_2O$  content in the range 24.3–64.5 wt %  $H_2SO_4$ .  $\gamma$  showed a negative temperature dependence at RH = 40 %.
- (m) Atmospheric pressure flow tube with contact times between 1.5 and 7 s once the 100 nm diameter aerosol was well mixed with  $N_2O_5$  that was detected using CIMS. The reaction probability for  $N_2O_5$  on  $H_2SO_4$  aerosols shows only a weak dependence upon the relative humidity over the range 1–90 %. The  $\gamma$  value over this range obtained from a fit to all available data is given by the following expression:  $\gamma[H_2SO_4] = 0.052 - 2.79 \times 10^{-4} \times (RH)$ .
- (n) Uptake of  $N_2O_5$  prepared in situ in a large aerosol chamber (AIDA) equipped with aerosol metrology instrumentation, gas-and particular phase  $H_2O$  detection using hygrometers and FTIR extinction measurement capabilities. Saturated ternary solutions (STS) consisting of  $H_2SO_4/HNO_3/H_2O$  were formed by the interaction of  $N_2O_5$  between the initial 38 wt %  $H_2SO_4/H_2O$  background aerosol with  $H_2O$

vapour from ice-coated chamber walls. The decrease of  $\gamma$  with increasing mixing ratio of  $\text{HNO}_3$  is attributed to the nitrate effect at low temperatures.

### Preferred values

Parameter	Value	$T/\text{K}$
$\alpha_b$	0.042	298
$\alpha_b$	$2.3 \times 10^{-5} \exp(2240/T)$	240–300
$\gamma_r$	$[(7353/T) - 24.83]^{-1}$	210–300
$\Delta \log(\alpha)$	$\pm 0.1$	298
$\Delta \log(\gamma)$	$\pm 0.3$ at 50 % RH	260–305

### 5 Comments on preferred values

There is a large body of experimental data on the uptake of  $\text{N}_2\text{O}_5$  on  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  surfaces covering the wide range of temperature and humidity relevant for the atmosphere between the surface and the lower stratosphere. The results are generally consistent between the different studies which used both bulk and dispersed (aerosol) surfaces. At room temperature  $\gamma$  shows little dependence on RH in the range 8–80 %, and hence on aerosol  $\text{H}_2\text{O}$  content in the corresponding range 20–70 wt %  $\text{H}_2\text{SO}_4$ . However at low temperatures there is a fall off in  $\gamma$  with decreasing water content.  $\gamma$  shows a distinct negative temperature dependence at  $T > 230 \text{ K}$ , but at lower temperature the temperature dependence becomes positive. The uptake leads to hydrolysis of  $\text{N}_2\text{O}_5$  and formation of  $\text{HNO}_3$  which transfers to the gas phase. However at low temperatures the increased solubility of  $\text{HNO}_3$  in the  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  particles leads to a reduction in the uptake rate.

The lack of a dependence on the water content of  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  surfaces and the negative temperature dependence at  $T > 230 \text{ K}$  suggests that uptake is controlled by a surface process, either mass accommodation or surface reaction. It is suggested by

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





Robinson et al. (1997) that the change in behaviour at low temperature is due to uptake under these conditions becoming limited by the rate of hydrolysis of  $\text{N}_2\text{O}_5$  in bulk liquid phase. They presented a phenomenological model which accounts for the observed dependence of the uptake coefficients on concentration and temperature. Two hydrolysis pathways are proposed, a direct reaction with  $\text{H}_2\text{O}$  and an acid-catalysed reaction involving dissociation of  $\text{N}_2\text{O}_5$  promoted by  $\text{H}^+$  ions.

The recommended expression for  $\gamma$  uses a resistance-model formulation but with an empirical representation of the temperature dependent liquid phase resistance due to chemical reaction,  $\gamma_r$ :

$$\gamma = \left\{ \frac{1}{\alpha_b} + \frac{1}{\gamma_r} \right\}^{-1}$$

The recommended value of  $\alpha_b$  is based on a linear least squares fit to all data at  $T > 240\text{ K}$ , plotted in the form  $\ln[\alpha/(1-\alpha)]$  vs.  $1/T$ . The temperature dependence of the  $\gamma_r$  term was determined by plotting the difference term  $(1/\gamma_{\text{obs}} - 1/\alpha)$  against  $1/T$  and using linear regression to evaluate the empirical relationship between  $1/\gamma_r$  and temperature.

## References

- Badger, C. L., Griffiths, P. T., George, I., Abbatt, J. P. D., and Cox, R. A.: J. Phys. Chem. A, 110, 6986–6994, 2006.
- Beichert, P. and Finlayson-Pitts, B. J.: J. Phys. Chem., 100, 15218–15228, 1996.
- Griffiths, P. T. and Cox, R. A.: Atmos. Sci. Lett., 10, 159–163, 2009.
- Hallquist, M., Stewart, D. J., Baker, J., and Cox, R. A.: J. Phys. Chem. A, 104, 3984–3990, 2000.
- Hallquist, M., Stewart, D. J., Stephenson, S. K., and Cox, R. A.: Phys. Chem. Chem. Phys., 5, 3453–3463, 2003.
- Fried, A., Henry, B. E., Calvert, J. G., and Mozurkewich, M.: J. Geophys. Res., 99, 3517–3532, 1994.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Folkers, M., Mentel, T. F., and Wahner, A.: Geophys. Res. Lett., 30, art. no. 1644, doi:10.1029/2003GL017168, 2003.

Golden, D. M., Manion, J. A., Reihls, C. M., and Tolbert, M. A.: in: CHEMRAWN VII: Chemistry of the Atmosphere: The Impact of Global Change, edited by: Calvert, J. G., Oxford University Press, 1992.

Hanson, D. R.: Geophys. Res. Lett., 24, 1087–1090, 1997.

Hanson, D. R. and Ravishankara, A. R.: J. Geophys. Res., 96, 17307–17314, 1991a.

Hu, J. H. and Abbatt, J. P. D.: J. Phys. Chem. A, 101, 871–878, 1997.

Kane, S. M., Caloz, F., and Leu, M.-T.: J. Phys. Chem. A, 105, 6465–6470, 2001.

Mentel, T. F., Sohn, M., and Wahner, A.: Phys. Chem. Chem. Phys., 1, 5451–5454, 1999.

Mozurkewich, M. and Calvert, J. G.: J. Geophys. Res., 93, 15889–15896, 1988.

Robinson, G. N., Worsnop, D. R., Kolb, C. E., and Davidovits, P.: J. Geophys. Res., 102, 3583–3601, 1997.

Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: Phys. Chem. Chem. Phys., 5, 4593–4603, 2003.

Van Doren, J. M., Watson, L. R., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 95, 1684–1689, 1991.

Wagner, R., Naumann, K.-H., Mangold, A., Möhler, O., Saathoff, H., and Schurath, U.: J. Phys. Chem. A, 109, 8140–8148, 2005.

Williams, L. R., Manion, J. A., Golden, D. M., and Tolbert, M. A.: J. Appl. Meteorol., 33, 785–790, 1994.

Zhang, R., Leu, M.-T., and Keyser, L. F.: Geophys. Res. Lett., 22, 1493–1496, 1995.

Zhang, R., Leu, M.-T., and Keyser, L. F.: J. Geophys. Res., 100, 18845–18854, 1995a.

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.12

### H<sub>2</sub>CO(g) + H<sub>2</sub>SO<sub>4</sub> (aq) → products

#### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
0.01	67	223	Tolbert et al. (1993)	Kn-MS (a)
0.08	75			
$(2.7 \pm 0.4) \times 10^{-3}$	10	267	Jayne et al. (1996)	DT-TDL (b)
$(6.7 \pm 0.9) \times 10^{-3}$	55	241		
$(1.14 \pm 0.09) \times 10^{-2}$	85	300		
$2 \times 10^{-3}$	48–95	197–215	Iraci and Tolbert (1997)	SF-MS (c)
<i>Solubility, <math>H^*</math> (M atm<sup>-1</sup>)</i>				
$2.0 \times 10^6$	48	206	Iraci and Tolbert (1997)	SF-MS (c)
$4.0 \times 10^6$	64	207		
$4.0 \times 10^6$	76	201		
$6.3 \times 10^6$	86	209		
$1.7 \times 10^6$	95	207		

#### 5 Comments

(a) Uptake experiments of CH<sub>2</sub>O on quiescent and stirred sulfuric acid solutions were performed in the range 60 to 75 %wt H<sub>2</sub>SO<sub>4</sub> at low temperature. The concentration of CH<sub>2</sub>O ranged from  $2 \times 10^{10}$  to  $2 \times 10^{11}$  molecule cm<sup>-3</sup> and no new gas phase products were detected during uptake. Gentle stirring of the sulfuric acid solution increased the rate of uptake by CH<sub>2</sub>O on the order of a factor of four or so owing to surface saturation at low temperatures, slow diffusion in the liquid and large gas concentrations.

10

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

- (b) Uptake to 10–85 wt% droplets at temperatures between 241 and 300 K. Time-dependent (reactive) uptake of  $\text{CH}_2\text{O}$  was observed and interpreted in terms of formation of  $\text{CH}_2(\text{OH})_2$  and protonated formaldehyde,  $\text{CH}_3\text{O}^+$ , at high acidities. The formation of a chemisorbed surface complex of  $\text{CH}_2\text{O}$  at the gas-liquid interface has been invoked under mildly acidic conditions (<20% acid).
- (c) Stirred flow reactor with  $\text{H}_2\text{CO}$  partial pressures of  $4 \times 10^{-6}$  to  $2.7 \times 10^{-3}$  mbar. Thin  $\text{H}_2\text{SO}_4$  films were made from the reaction of  $\text{SO}_3$  with  $\text{H}_2\text{O}$  and supported on an optical window monitored by transmission FTIR spectroscopy.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.04	200–300
$c$	$8 \times 10^{-8} \exp(-0.0272 \text{ wt})$	200–300
$H_{\text{CH}_2\text{O}}$ ( $\text{M atm}^{-1}$ )	$\exp(-(-0.0456 + 55.5/T) \times 0.46 m_{\text{H}_2\text{SO}_4})$	200–300
$H_{\text{CH}_2(\text{OH})_2}^*$ ( $\text{M atm}^{-1}$ )	$H_{\text{CH}_2\text{O}} (1 + K_2 a_{\text{H}_2\text{O}})$	200–300
$H_{\text{CH}_3\text{O}^+}^*$ ( $\text{M atm}^{-1}$ )	$H_{\text{CH}_2\text{O}} (1 + K_3 [\text{H}^+])$	200–300
$K_2$ ( $\text{M}^{-1}$ )	$\exp(4020/T - 5.83)$	200–300
$K_3$ ( $\text{M}^{-1}$ )	$0.56 \exp(8.84 (T - 260)/T)$	200–300
$k_{\text{H}_2\text{O}}$ ( $\text{M}^{-1} \text{ s}^{-1}$ )	$7800 \exp(-1910/T) (1 + 870[\text{H}^+])$	200–300
$\Gamma_{b,\text{poly}}$	0.001	197–215
<i>Reliability</i>		
$\Delta \log(\alpha)$	$\pm 0.3$	200–300
$\Delta \log(H)$	$\pm 0.3$	200–300
$\Delta \log(K)$	$\pm 0.3$	200–300
$\Delta \log(\gamma)$	$\pm 0.7$	200–300

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

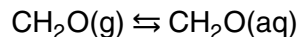
Interactive Discussion



## Comments on preferred values

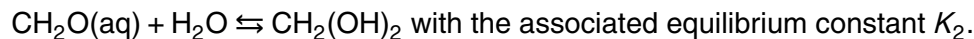
By analogy to known condensed phase kinetics CH<sub>2</sub>O is both solvated to methylene-glycol (CH<sub>2</sub>(OH)<sub>2</sub>) ~~as well as existing~~ in protonated form (CH<sub>3</sub>O<sup>+</sup>) in H<sub>2</sub>SO<sub>4</sub>. The uptake of HCHO to H<sub>2</sub>SO<sub>4</sub> shows a negative temperature dependence at constant [H<sub>2</sub>SO<sub>4</sub>] and a pronounced H<sub>2</sub>SO<sub>4</sub> concentration dependence:  $\gamma$  increases with increasing concentration starting from a concentration where formaldehyde protonation is thought to be the dominant solvation mechanism. Jayne et al. (1996) present a set of parameters that reproduced the time dependent uptake of CH<sub>2</sub>O for the short time scale (20 ms) of their experiment, from which we adopt the value for  $\alpha_b$ . In absence of static measurements we adopt their parameters relevant for the calculation of the solubility as a recommendation, since they appropriately link to more dilute solution as well as known hydrolysis rate constants (Jayne et al., 1992).

$H_{\text{CH}_2\text{O}}$  describes the physical Henry's Law for formaldehyde:

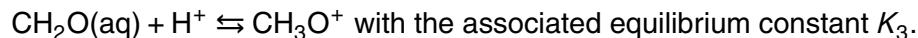


It includes a Setchenow coefficient referenced to that for HOCl.

$H_{\text{CH}_2(\text{OH})_2}^*$  includes the ~~further equilibrium due to~~ hydration to the more soluble gem-diol form:



$H_{\text{CH}_3\text{O}^+}$  (Matm<sup>-1</sup>) can be used to describe the partitioning into the protonated form of formaldehyde:



We suggest using the parameterisation of Shi et al. (2001) for the proton concentration and the water activity:

$$(\text{H}^+) = \exp \left( 60.51 - 0.095 wt + 0.0077 wt^2 - 1.61 \times 10^{-5} wt^3 \right. \\ \left. - (1.76 + 2.52 \times 10^{-4} wt^2) T^{0.5} + (-805.89 + 253.05 wt^{0.076}) / T^{0.5} \right)$$

32373

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$$a_w = \exp[(-69.775X - 18253.7X^2 + 31072.2X^3 - 25668.8X^4)(1/T - 26.9033/T^2)]$$

where the mole fractions can be calculated from weight percent concentration by:  $X = wt/(wt + (100 - wt)98/18)$

The three Henry's Law constants can be used to calculate the overall equilibrium uptake and partitioning into  $\text{CH}_2\text{O}$  (aq),  $\text{CH}_2(\text{OH})_2$ , and  $\text{CH}_3\text{O}^+$ . Jayne et al. (1996) also invoked a surface process, parameterized by a time dependent direct gas-surface reaction probability to fully explain the observed uptake coefficient at short gas-droplet interaction times. Since the surface density of the corresponding surface complex remains low it does not affect the overall equilibrium of the other species. Iraci and Tolbert (1997) did not see evidence for the presence of  $\text{CH}_2(\text{OH})_2$  in IR spectra taken after long exposure time at 210 K and 80%  $\text{H}_2\text{SO}_4$ , but rather suspected formation of a polymer, e.g., through dehydration of  $\text{CH}_2(\text{OH})_2$ , which would lead to an apparently higher effective solubility than predicted from the equilibria above alone. The absence of  $\text{CH}_2(\text{OH})_2$  may have been due to the fall off of  $K_2$  at high  $\text{H}_2\text{SO}_4$  concentration. Since the long term kinetics has not been measured at higher  $T$  and lower  $[\text{H}_2\text{SO}_4]$ , where  $\text{CH}_2(\text{OH})_2$  concentrations would be higher, we refrain from giving detailed kinetic parameters for this potential sink of formaldehyde, but simply recommend a constant value of  $10^{-3}$  for  $\Gamma_{b,poly}$  with large error limits, so that the overall  $\gamma$  becomes consistent with the steady state uptake observed by Iraci and Tolbert after long exposure times.

The overall uptake coefficient is given by the expression:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{sol,HCHO}} + \frac{1}{\frac{1}{\Gamma_{\text{sol,CH}_2(\text{OH})_2} + \frac{1}{\Gamma_{b,\text{CH}_2(\text{OH})_2}}} + \Gamma_{\text{sol,CH}_3\text{O}^+} + \Gamma_{b,poly}}$$

The time dependent solubility limited uptake terms are given by:

$$\Gamma_{\text{sol,X}} = \frac{4H_X RT \sqrt{\frac{4D_i}{\pi t}}}{\bar{C}_{\text{HCHO}}}$$

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The hydration reaction is parameterized by:

$$\frac{1}{\Gamma_{b,CH_2(OH)_2}} = \frac{4H_{HCHO}RT\sqrt{D_1a_wk_{H_2O}}}{\bar{c}_{HCHO}}$$

For the hydration rate constant  $k_{H_2O}$  we adopt the expression provided by Jayne et al. (1996) that is consistent with more dilute aqueous solutions.

For the diffusion coefficient, we suggest using the expression for the  $c$  parameter given above, based on Jayne et al. (1996), via  $D_1 = cT/\eta$ . For the viscosity, we suggest using the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into the tropospherically more relevant dilute solutions at high  $T$ :

$$\eta = aT^{-1.43} \exp(448 K / (T - T_0)),$$

$$\text{with } a = 169.5 + 5.18 wt - 0.0825 wt^2 + 3.27 \times 10^{-3} wt^3,$$

$$\text{and } T_0 = 144.11 + 0.166 wt - 0.015 wt^2 + 2.18 \times 10^{-4} wt^3$$

The combination of all three experiments shows that the rate of uptake of  $CH_2O$  by  $H_2SO_4$  solutions increases with acidity of  $H_2SO_4$  at 228 K for  $[H_2SO_4] \geq 70$  wt % to values close to  $\gamma = 0.10$ . At higher temperatures in the range of 241–300 K the solubility of dissolved  $CH_2O$  and methyleneglycol ( $CH_2(OH)_2$ ) increases with increasing temperature and decreases with increasing acid activity. For protonated formaldehyde ( $CH_3O^+$ ) the solubility increases with increasing acid activity as well as decreasing temperature.

## References

- Iraci, L. T. and Tolbert, M. A.: J. Geophys. Res., 102, 16099–16107, 1997.  
Jayne, J. T., Duan, S. X., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Phys. Chem., 96, 5452–5460, 1992.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Jayne, J. T., Worsnop, D. R., Kolb, C. E., Swartz, E., and Davidovits, P.: J. Phys. Chem., 100, 8015–8022, 1996.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

5 Tolbert, M. A., Pfaff, J., Jayaweera, I., and Prather, M. J.: J. Geophys. Res., 98, 2957–2962, 1993.

Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





$\text{CH}_3\text{C}(\text{O})\text{OONO}_2$  (PAN) +  $\text{H}_2\text{SO}_4$  (aq) → products

## Experimental data

Parameter	H <sub>2</sub> SO <sub>4</sub> wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma_0 = 0.008$ (85 wt % H <sub>2</sub> SO <sub>4</sub> )		200	Zhang and Leu (1997)	WWFT-CIMS (a)
<i>Solubility: <math>H^*</math> (M atm<sup>-1</sup>)</i>				
$6.47 \times 10^4$	46	199	Zhang and Leu (1997)	WWFT-CIMS (a)
$5.54 \times 10^4$	46	216		
$2.0 \times 10^4$	54	208		
$8.75 \times 10^2$	54	226		
$2.42 \times 10^4$	59	208		
$4.01 \times 10^3$	59	216		
$5.58 \times 10^4$	72	208		
$3.31 \times 10^3$	72	222		

## 5 Comments

(a) PAN was observed to be reversibly taken up into sulfuric acid owing to its solubility. The quantity  $H^*(D_1)^{0.5}$  was determined from the time-dependent uptake of PAN on sulfuric acid. The solubility constant  $H^*$  was extracted using calculated diffusion coefficients based on the model developed by Luo et al. (1994) using the measured viscosities of H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O solution from Williams and Long (1995). The  $H^*$  values were found to depend strongly on temperature and less strongly on acid composition (46–72 wt %). For example, the effective  $H^*$  at 208–222 K on 72 wt % H<sub>2</sub>SO<sub>4</sub> was approximately a factor of 2–3 higher than extrapolated from measurements at higher temperatures for more dilute solutions.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_b$	>0.008	190–230
$\Delta H/\text{kJ mol}^{-1}$ (46 wt H <sub>2</sub> SO <sub>4</sub> %)	50.2	195–220
$\Delta/k\text{J mol}^{-1}$ (72 wt % H <sub>2</sub> SO <sub>4</sub> )	78.2	208–220
$\Delta S/\text{J mol}^{-1} \text{K}^{-1}$ (46 wt % H <sub>2</sub> SO <sub>4</sub> )	160	195–220
$\Delta S/\text{J mol}^{-1} \text{K}^{-1}$ (72 wt % H <sub>2</sub> SO <sub>4</sub> )	284	195–220
<i>Reliability</i>		
$\Delta \log \alpha_b$	undetermined	
$\Delta \log H^*$	±0.3	15–220

## Comments on preferred values

At H<sub>2</sub>SO<sub>4</sub> concentrations lower than 72 wt % PAN uptake is reversible and solubility controls the amount taken up to the solution. The preferred values for the effective Henry's law solubility,  $H^*$ , are taken from the study by Zhang and Leu (1997), which are consistent with studies on water surfaces at higher temperatures (Kames and Schurath, 1995). The temperature dependence of  $H^*$  for a given acid composition can be expressed as:  $\ln H^* = -\Delta H/RT + \Delta S/R$ . The values of the PAN solvation enthalpy and entropy  $\Delta H$  and  $\Delta S$  vary with acid composition in the range 46–72 wt %. The values given in the table are for the lower and upper end of composition range and cover the range 200–222 K.

Measurement of accommodation coefficients in the time dependent regime leading to saturation is difficult as adsorption and desorption are not separated in time. The recommendation is based on the initial uptake coefficient reported by Zhang and Leu (1997) and is given a large uncertainty.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## References

Kames, J. and Schurath, U.: J. Atmos. Chem., 21, 151–164, 1995.

Luo, B. P., Clegg, S. L., Peter, T., Muller, R., and Crutzen, P. J.: Geophys. Res. Lett., 21, 49–52, doi:10.1029/93gl02956, 1994.

5 Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

Zhang, R. and Leu, M. T.: J. Geophys. Res., 102, 8837–8843, 1997.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
0.42	72	223		
0.44	72	233		
0.47	72	238		
<i>Solubility:</i> H (Matm <sup>-1</sup> ); <i>Diffusion:</i> D <sub>1</sub> (cm <sup>2</sup> s <sup>-1</sup> )				
H* = 9.9 × 10 <sup>-13</sup> exp(9030/T)	50	220–230	Williams and Golden (1993)	Kn-MS (i)
H* = 3.0 × 10 <sup>-4</sup> exp(4150/T)	55	215–230		
H* = 1.3 × 10 <sup>-5</sup> exp(4640/T)	60	200–230		
HD <sub>1</sub> <sup>1/2</sup> = 3.6 × 10 <sup>-4</sup> exp(2060/T)	59.6	200–225	Hanson and Ravishankara (1993)	CWFT-CIMS (j)
HD <sub>1</sub> <sup>1/2</sup> = 1.1 × 10 <sup>-3</sup> exp(2070/T)	55.6	205–225		
HD <sub>1</sub> <sup>1/2</sup> = 9.2 × 10 <sup>-4</sup> exp(2440/T)	50.5	200–230		
H* = 2.7 × 10 <sup>-4</sup> exp(4910/T)	45	195–220		
H* = 1.0 × 10 <sup>-4</sup> exp(5110/T)	45.3	195–220		
H* = 2.6 × 10 <sup>-4</sup> exp(4650/T)	50	200–220		
H* = 6.8 × 10 <sup>-6</sup> exp(5310/T)	51.1	200–220		
H* = 1.2 × 10 <sup>-5</sup> exp(6100/T)	35	210–230	Zhang et al. (1993)	Static (k)
H* = 6.5 × 10 <sup>-6</sup> exp(5880/T)	40	200–230		
H* = 1.0 × 10 <sup>-6</sup> exp(5530/T)	50	195–230		
H* = 4.2 × 10 <sup>-7</sup> exp(4920/T)	60	195–210		
H* = 5.0 × 10 <sup>-4</sup> exp(5060/T)	43	205–225	Elrod et al. (1995)	Static (l)
H* = 1.2 × 10 <sup>-4</sup> exp(4940/T)	50	205–225		
H* = 1.0 × 10 <sup>3</sup>	59.5	251	Hanson and Lovejoy (1996)	RWFT-CIMS (c)
H* = 3.3 × 10 <sup>-5</sup> exp(5460/T)	45	205–225	Hanson (1998)	RWFT-CIMS (m)
H* = 6.9 × 10 <sup>-6</sup> exp(5450/T)	50			
HD <sub>1</sub> <sup>1/2</sup> = 1.4 × 10 <sup>-6</sup> exp(4230/T)	49	230–260	Robinson et al. (1998)	DT-TDLAS (d)
HD <sub>1</sub> <sup>1/2</sup> = 9.8 × 10 <sup>-7</sup> exp(3900/T)	54	230–260		
HD <sub>1</sub> <sup>1/2</sup> = 4.8 × 10 <sup>-5</sup> exp(2500/T)	59	235–265		
HD <sub>1</sub> <sup>1/2</sup> = 7.9 × 10 <sup>-2</sup>	69	240		

## Comments

(a) The bulk substrate (1 mL to 5 mL) was cooled slowly to temperatures of 210 K to 230 K. The water vapour pressure was 5.3 × 10<sup>-4</sup> mbar. Measured uptake coefficients did not depend on the appearance of the sample (glassy or crystalline). At 230 K, the measured uptake coefficient was diffusion limited; the larger value is therefore listed as lower limit. The recovery yield of HCl was at maximum 5% when the solution was warmed to 253 K.

32381

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



(b) Droplet diameters were between 110 and 280  $\mu\text{m}$ . The HCl concentration (circa  $10^{12}$  molecule  $\text{cm}^{-3}$ ) was measured downstream of the flow tube by tunable diode laser absorption. The uptake coefficients listed in the table were corrected by gas phase diffusion using  $D_{\text{HCl-He}} = 515 \text{ torr cm}^2 \text{ s}^{-1}$  and  $D_{\text{HCl-H}_2\text{O}} = 120 \text{ torr cm}^2 \text{ s}^{-1}$ . The time dependent uptake coefficient was evaluated in terms of solubility driven uptake.  $H^*$  was estimated by  $H^* = HK_a/[\text{H}^+]$  at low  $\text{H}_2\text{SO}_4$  concentrations and using the Hammett acidity concept to estimate the degree of dissociation of HCl in high wt %  $\text{H}_2\text{SO}_4$  solution. For the liquid phase diffusion coefficient,  $10^{-5} \text{ cm}^2 \text{ s}^{-1}$  was used.

(c) Uptake into a known amount of  $\text{H}_2\text{SO}_4$  solution (around  $1 \text{ cm}^3$ ).

(d) Time dependent uptake coefficients were measured using droplet diameters between 120 and 250  $\mu\text{m}$ . The HCl concentration of about  $10^{12}$  molecule  $\text{cm}^{-3}$  was measured downstream of the flow tube by tunable diode laser absorption. The time dependent uptake coefficients were corrected for gas phase diffusion using  $D_{\text{HCl-He}} = 0.701 \text{ atm cm}^2 \text{ s}^{-1}$ ,  $D_{\text{HCl-Ar}} = 0.157 \text{ atm cm}^2 \text{ s}^{-1}$  and  $D_{\text{HCl-H}_2\text{O}} = 0.166 \text{ atm cm}^2 \text{ s}^{-1}$ . By varying the droplet size, the  $\text{Kn}$  numbers could be raised to 1 allowing determination of uptake coefficients in the range of 0.01 to 1. The corrected uptake coefficients were evaluated in terms of solubility driven uptake to determine of  $H^*D_1^{1/2}$  and  $\alpha_b$ .

(e) Uptake of HCl was monitored through the size changes of 30–70  $\mu\text{m}$  diameter droplets held in an electrodynamic trap with HCl pressures at 0.01 to 0.1 mbar. The time dependent uptake was evaluated to obtain an independent measurement of the diffusion coefficient. For data in the phase transfer limited regime,  $\alpha_b$  was estimated from the slope of a plot of the rate of uptake versus the HCl partial pressure. The authors indicated that the HCl pressure was not homogeneous throughout the chamber, so that the  $\alpha_b$  value reported is probably too low.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

(f) Molecular beams produced from the expansion of 2 % HCl in H<sub>2</sub>, 5 % HCl in He or 10 % HCl in N<sub>2</sub>. The substrate was 70 % D<sub>2</sub>SO<sub>4</sub> at 213 K. Scattered or desorbing molecules were detected with a mass spectrometer at an exit angle of 45°. At the lowest energy, the fraction of in-elastically scattered molecules was less than 10 %, from which we provide a lower bound to  $\alpha_s$  reported in the table. Only 11 % of those molecules trapped on the surface undergo proton exchange in the bulk of D<sub>2</sub>SO<sub>4</sub> leading to the value of  $\alpha_b$  listed in the table, if the proton exchange rate is limited by the solvation rate.

(g) Molecular beam experiment using the same setup as in Morris et al. (2000), but covering a broader range of D<sub>2</sub>SO<sub>4</sub> concentration. Again, the measured fraction of HCl undergoing proton exchange to DCl is interpreted as  $\alpha_b$  assuming that  $\alpha_s$  is close to 1, since inelastically scattered HCl molecules were not observed at low incident beam energies. The values given in the table were interpolated from the data provided as a plot only. A few experiments were also presented with DCl on H<sub>2</sub>SO<sub>4</sub>, showing about 20 % higher values, tentatively interpreted as isotopic effects. Behr et al. (2001) also provide values for the solubility based on HCl residence time measurement that are in line with Carslaw et al. (1995) at low acid concentration but are higher above 65 wt %.

(h) Molecular beam experiment comparable to the setup used by Morris et al. (2000).

(i) Time dependent HCl uptake coefficients measured in a Knudsen cell to derive values of  $H^*D_1^{1/2}$ . Values of  $H^*$  were obtained by estimating the temperature dependent diffusion coefficient from the measured temperature dependence of the viscosity. Initial uptake coefficients were not reported. The expressions given in the table were obtained by fitting to the reported data.

(j) Time dependent HCl uptake coefficients measured in a cylindrical flow tube with H<sub>2</sub>SO<sub>4</sub> contained in a boat were analysed to derive values of  $H^*D_1^{1/2}$ . Direct measurements of  $H^*$  were done by static vapour pressure measurements. The param-

32383

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



eters listed in the table were obtained by fitting to the tabulated datasets given by Hanson and Ravishankara (1993). Initial uptake coefficients were not reported.

- (k) Static vapour pressure measurement using a mass spectrometer over a stirred  $\text{H}_2\text{SO}_4/\text{HCl}/\text{H}_2\text{O}$  solution (0.01–5 wt % HCl). The expression given in the table was obtained by fitting to the reported data.
- (l) Vapour pressure measurement over ternary and quaternary solutions in a wetted wall flow tube. The expression given in the table was obtained by fitting to the reported data.
- (m) Vapour pressure measurement over  $\text{H}_2\text{SO}_4$  solutions in a rotating wetted wall flow tube coupled to a CIMS. HCl was added to the solution at  $10^{-2}$  to  $10^{-3}$  M. Linearity between HCl concentration and flow rate was checked to assure equilibrated conditions. The expression in the table is obtained from a fit to the data as reported. The value of  $H^*$  reported in Table 1 of Hanson (1998) for the 50 wt % solution ( $1.36 \times 10^7$ ) was corrected to  $1.36 \times 10^6$  to be consistent with the value plotted in the figure. Hanson (1998) also reanalysed data from Hanson and Ravishankara (1993) to take into account proper handling of the  $\text{SF}_6^-$  signal during CIMS calibration and a HCl background signal from memory effects on tubing downstream of the flow tube, which resulted in fair agreement with the newer measurements.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_s$	>0.9	213–238
$k_{\text{sol}}$ ( $\text{s}^{-1}$ )	$7.84 \times 10^{10}/\eta$ (cP)	190–300
$k_{\text{des}}$ ( $\text{s}^{-1}$ )	$8.0 \times 10^{17} \exp(-5000/T)$	190–300
$H^*$ ( $\text{Matm}^{-1}$ )	$(0.094 - 0.61X + 1.2X^2) \exp(-8.68 + (8515 - 10718X^{0.7})/T)$	190–300
$c$ ( $\text{cm}^2 \text{cPK}^{-1} \text{s}^{-1}$ )	$7.8 \times 10^{-8}$	190–300
$A$ ( $\text{Matm}^{-1} \text{K}^{1.43}$ )	$169.5 + 5.18 (\text{wt} \%) - 0.0825 (\text{wt} \%)^2 + 3.27 \times 10^{-3} (\text{wt} \%)^3$	190–300
$T_0$ (K)	$144.11 + 0.166 (\text{wt} \%) - 0.015 (\text{wt} \%)^2 + 2.18 \times 10^{-4} (\text{wt} \%)^3$	190–300
<i>Reliability</i>		
$\Delta \log(\alpha_s)$	$\pm 0.3$	213–238
$\Delta \log(k_{\text{sol}})$	$\pm 0.3$	190–300
$\Delta \log(k_{\text{des}})$	$\pm 0.3$	190–300
$\Delta \log(H^*)$	$\pm 0.3$	190–300

## Comments on preferred values

Uptake of HCl into  $\text{H}_2\text{SO}_4$  solutions is fast. The molecular beam experiments provide the most direct measurement of the initial steps of HCl uptake. At low (close to thermal) incident beam energies, nearly all HCl molecules are thermally accommodated on the surface, evident from the absence of inelastically scattered molecules, leading to the recommended high value for  $\alpha_s$ . The fraction of HCl molecules that actually undergo exchange with the bulk liquid phase can be tracked by observing proton exchanged DCI molecules desorbing from the surface. Assuming that solvation of HCl is the rate limiting step, this fraction is used to obtain the bulk accommodation coefficient  $\alpha_b$  (Behr et al., 2009).

The kinetic data by Watson et al. (1990) did not allow to clearly differentiate the contributions by bulk accommodation and solubility. Robinson et al. (1998) provided a more reliable data set at lower temperature using the droplet train technique that involved sufficiently high Knudsen number conditions to extract large bulk accommodation coefficients. The Robinson et al. and Behr et al. data sets agree in that  $\alpha_b$  is large

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

( $>0.6$ ) up to moderate sulphuric acid contents. However, the temperature dependence observed is negative for the 39 and 48 wt% solutions above 230 K for the Robinson et al. data and positive for the 72 wt% solution below 240 K as obtained by Behr et al. Robinson et al. interpret their negative  $T$  dependence as driven by the Gibbs free energy of the transition state at the surface. This would be inconsistent with the high single value measured by Hanson and Lovejoy (1996) at 274 K on a 34 wt% solution. Behr et al. (2009) use capillary wave theory to explain the strong negative correlation of  $\alpha_b$  with viscosity (determining  $k_{sol}$ ). Our recommendation follows the suggestion by Behr et al. (2009) for  $k_{sol}$ . We use an Arrhenius expression for  $k_{des}$  with an activation energy in between that of Behr et al. and Robinson et al. Using

$$\alpha_b = k_{sol} / (k_{des} + k_{sol})$$

this leads to reasonable agreement with the data sets considered. It follows the strong negative dependence on sulphuric acid content at low  $T$ , whereby we prefer the newer Behr et al. (2009) data with respect to the absolute value while still aligning it parallel to the concentration dependence of the earlier data. The recommended parameterization resolves the apparent inconsistency of positive temperature dependence at low temperature and high sulphuric acid concentration and negative temperature dependence for the more dilute solutions.

In the absence of a reactive sink in the aqueous phase, uptake is driven by solubility, characterized by the effective Henry's Law constant  $H^*$  and the diffusion coefficient,  $D_1$ .

The more reliable solubility values are derived from experiments, in which the solubility was directly measured: Elrod et al. (1995), Hanson (1998) and the one data point by Hanson and Lovejoy (1996). The Zhang et al. (1993) data are likely too low due to HCl calibration issues as noted by Elrod et al. Many kinetic experiments were analysed in terms of solubility limited uptake to obtain  $H^* D^{1/2}$ . This requires estimating diffusion coefficients to obtain  $H^*$ , which in turn are parameterized in terms of the viscosity. We recommend using the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into tropospherically more relevant

dilute solutions at high  $T$ :

$$\eta = AT^{-1.43} \exp(448 K / (T - T_0))$$

The temperature and concentration dependent diffusion coefficient can then be obtained by the Stokes-Einstein equation

5  $D_1 = cT/\eta$

The diffusion coefficients of HCl thus obtained are in reasonable agreement with data obtained by Klassen et al. (1998), but in some disagreement with measurements at  $T < 190$  K by Schwell et al. (2000), though at only two acid concentrations. Schwell et al. argue that the Stokes Einstein approach used by Klassen et al. may underestimate the diffusion coefficient at temperatures below 200 K.

10 This procedure to obtain  $H^*$  from the measured  $HD^{1/2}$  values leads to fair agreement with the directly measured  $H^*$  values and aligns the higher temperature droplet train experiments by Robinson et al. (1998) to those at lower  $T$ .

15 We recommend the mole fraction based expression by Shi et al. (2001) for  $H^*$  over 190–300 K for up to 70 wt % solutions ( $X = wt / (wt + (100 - wt)98/18)$ ) noting the potential formation of  $\text{HSO}_3\text{Cl}$  as suggested by Robinson et al. (1998) to explain the higher  $H^*$  at >70 wt % solution. Shi's parameterization builds upon the thermodynamic models of Carslaw et al. (1995) and Luo et al. (1995) and uses an expression modified from that used by Hanson (1998).

## 20 References

Behr, P., Morris, J. R., Antman, M. D., Ringeisen, B. R., Splan, J. R., and Nathanson, G. M.: Geophys. Res. Lett., 28, 1961–1964, 2001.

Behr, P., Scharfenort, U., Ataya, K., and Zellner, R.: Phys. Chem. Chem. Phys., 11, 8048–8055, doi:10.1039/b904629a, 2009.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Carlsaw, K. S., Clegg, S. L., and Brimblecombe, P.: J. Phys. Chem., 99, 11557–11574, 1995.

Eicher, L. D. and Zwolinski, B. J.: J. Phys. Chem., 75, 2016–2024, 1971.

Elrod, M. J., Koch, R. E., Kim, J. E., and Molina, M. J.: Faraday Discuss., 100, 269–278, 1995.

Hanson, D. R. and Lovejoy, E. R.: J. Phys. Chem., 100, 6397–6405, 1996.

Hanson, D. R. and Ravishankara, A. R.: J. Phys. Chem., 97, 12309–12319, 1993.

Hanson, D. R.: J. Phys. Chem. A, 102, 4794–4807, 1998.

Hayduk, W. and Minhas, B. S.: Can. J. Chem. Eng., 60, 295–299, 1982.

Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.

Luo, B. P., Carlsaw, K. S., Peter, T., and Clegg, S. L.: Geophys. Res. Lett., 22, 247–250, 1995.

Morris, J. R., Behr, P., Antman, M. D., Ringeisen, B. R., Splan, J., and Nathanson, G. M.: J. Phys. Chem. A, 104, 6738–6751, 2000.

Robinson, G. N., Worsnop, D. R., Jayne, J. T., Kolb, C. E., Swartz, E., and Davidovits, P.: J. Geophys. Res., 103, 25371–25381, 1998.

Schwell, M., Baumgärtel, H., Weidinger, I., Kramer, B., Vortisch, H., Wöste, L., Leisner, T., and Rühl, E.: J. Phys. Chem. A, 104, 6726–6732, 2000.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

Tolbert, M. A., Rossi, M. J., and Golden, D. M.: Geophys. Res. Lett., 15, 847–850, 1988.

Watson, L. R., Vandoren, J. M., Davidovits, P., Worsnop, D. R., Zahniser, M. S., and Kolb, C. E.: J. Geophys. Res., 95, 5631–5638, 1990.

Williams, L. R. and Golden, D. M.: Geophys. Res. Lett., 20, 2227–2230, 1993.

Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

Zhang, R. Y., Wooldridge, P. J., and Molina, M. J.: J. Phys. Chem., 97, 8541–8548, 1993.

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



ClO + H<sub>2</sub>SO<sub>4</sub> (aq) → products

## Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$(2.0 \pm 1) \times 10^{-5}$	90	295	Martin et al. (1980)	CWFT-EPR (a)
$(1.5 \pm 1) \times 10^{-4}$	85	260		
$(2.2 \pm 1) \times 10^{-4}$	80	240		
$<1.0 \times 10^{-4}$	60–70	213	Abbatt (1996)	CWFT-RF (b)

5 **Comments**

(a) Measurement of the uptake kinetics in a fast flow tube with EPR detection of ClO. Pure Cl<sub>2</sub> at 0.67 mbar was discharged in a  $\mu$ -wave cavity and combined downstream with an excess of O<sub>3</sub> in order to generate ClO. The quartz flow tube was coated with halocarbon wax and the discharge tube with solid B<sub>2</sub>O<sub>3</sub> in order to minimize wall losses of ClO. The inside of the flow tube was coated with H<sub>2</sub>SO<sub>4</sub> and the H<sub>2</sub>O vapor pressure was held constant throughout the temperature range at  $6.7 \times 10^{-4}$  mbar leading to a changing composition of the H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O mixture as the temperature was changed (96 % at 300 K, 80 % at 240 K).

(b) Uptake rates in a coated wall flow tube at 1.3 mbar total pressure of He coupled to a resonance fluorescence (RF) detector after chemical conversion. ClO was generated from the reaction Cl + O<sub>3</sub> → ClO + O<sub>2</sub> by discharging Cl<sub>2</sub>/He in a microwave cavity and adding O<sub>3</sub> downstream at typical [O<sub>3</sub>] in the range from 1 to  $3 \times 10^{13}$  molecule cm<sup>-3</sup>. After ClO interacted with the active surface it was converted to Cl in the reaction ClO + NO → Cl + NO<sub>2</sub> which was detected using

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



a RF-excited Cl<sub>2</sub>/He lamp emitting in the VUV at 119 nm in conjunction with a flowing O<sub>2</sub> filter. The sensitivity of the RF lamp was measured using the reaction H + Cl<sub>2</sub> → HCl + Cl. The ClO concentrations used in this work were on the order of 1 to 2 × 10<sup>11</sup> molecule cm<sup>-3</sup>.

## 5 Preferred values

No recommendation.

## Comments on preferred values

In the work of Martin et al. (1980) no separation between the effects of changing composition of the H<sub>2</sub>SO<sub>4</sub> solution (considered to be negligible) and the temperature dependence of  $\gamma$  was performed. All kinetic results were obtained from the full numerical integration of the parabolic flow system including axial diffusion and resulted in a pronounced negative temperature dependence of the uptake coefficient  $\gamma_{ss}$ . Martin et al. observed HCl as product of the heterogeneous reaction. No effect of irradiation by simulated sunlight on  $\gamma$  was found within a  $\pm 20\%$  variation of  $\gamma$ . Martin et al. (1980) report a higher sensitivity for ClO-detection compared to Abbatt (1996). In the light of the disagreement between the two datasets we make no recommendations, but note that uptake of ClO to H<sub>2</sub>SO<sub>4</sub> at stratospheric temperatures and composition is inefficient.

## References

- Martin, L. R., Judeikis, H. S., and Wun, M.: J. Geophys. Res., 85, 5511–5518, 1980.  
20 Abbatt, J. P. D.: Geophys. Res. Lett., 23, 1681–1684, 1996.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## VI.A4.16

### HOBr + H<sub>2</sub>SO<sub>4</sub> (aq) → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
Uptake coefficients: $\gamma$ >0.06 (60 wt % H <sub>2</sub> SO <sub>4</sub> )	228	Abbatt (1994)	CWFT-MS (a)
Solubility: $H$ (M atm <sup>-1</sup> )			
$3.8 \times 10^5$ (69.8 wt %)	212	Abbatt (1995)	CWFT-MS (b)
$9.0 \times 10^4$ (69.8 wt %)	228		
$\approx 1 \times 10^6$ (60 wt % H <sub>2</sub> SO <sub>4</sub> )	210	Hanson and Ravishankara (1995)	CWFT-CIMS (c)
$4.6 \times 10^{-4} \exp[(4500 \pm 480)/T]$	213–238	Waschewsky and Abbatt (1999)	CWFT-MS (d)
$1.4 \times 10^5$ (58–69 wt %)	250	Hanson (2003)	CWFT-CIMS (e)
$\approx 4 \times 10^4$	260		
$\approx 2.5 \times 10^4$	270		

#### 5 Comments

(a) HOBr ( $0.5\text{--}5 \times 10^{12}$  molecule cm<sup>-3</sup>) was generated by the reaction sequence: H + NO<sub>2</sub> → OH + NO and OH + Br<sub>2</sub> → HOBr, ionised by electron impact and detected as HOBr<sup>+</sup>. The carrier gas flow was humidified to maintain the H<sub>2</sub>SO<sub>4</sub> concentration.

10 (b) HOBr (<1 × 10<sup>12</sup> molecule cm<sup>-3</sup>) was generated by the reaction sequence: H + NO<sub>2</sub> → OH + NO and OH + Br<sub>2</sub> → HOBr, ionised by electron impact and detected as HOBr<sup>+</sup>. The carrier gas flow was humidified to maintain the H<sub>2</sub>SO<sub>4</sub> concentration (69.8 wt %). Some evidence was found for a second-order component of HOBr loss due to self reaction. Values of  $HD_1^{1/2} = (30 \pm 15)$  and  
 15  $(20 \pm 10)$  M atm<sup>-1</sup> cm s<sup>-1/2</sup> were obtained at 212 and 228, respectively and con-

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



verted to the values listed in the table using diffusion coefficients calculated as described by Klassen et al. (1998).

- (c) HOBr detected using  $\text{SF}_6^-$  chemi-ions. The evolution of the HOBr signal following uptake of  $\text{BrONO}_2$  ( $1-3 \times 10^{-7}$  Torr) to the  $\text{H}_2\text{SO}_4$  surface was analysed to derive time dependent uptake coefficients and solubility. A value of  $HD_1^{1/2}$  of 110 ( $\pm 40$ )  $\text{M atm}^{-1} \text{cm s}^{-1/2}$  was converted to the solubility listed in the table using  $D_1 \approx 10^{-8} \text{cm}^2 \text{s}^{-1}$ .
- (d) HOBr ( $\approx 5 \times 10^{-10}$  atm, measured by UV absorption at 254 nm) was generated ex-situ by passing a humidified flow of  $\text{Br}_2$  in He over  $\text{HgO}$ , ionised by electron impact and detected as  $\text{HOBr}^+$ . The carrier gas flow was humidified to maintain the  $\text{H}_2\text{SO}_4$  concentration. The time dependent HOBr signal was used to derive values of  $HD_1^{1/2}$  at different temperatures and  $\text{H}_2\text{SO}_4$  concentrations (59.7–70.1 wt %). The uptake to fresh  $\text{H}_2\text{SO}_4$  surfaces was found to be as much as a factor of two larger than to surfaces previously exposed to HOBr. Only data from fresh surfaces is reported. Values of  $HD_1^{1/2}$  were converted to  $H_{\text{HOBr}}$  using diffusion coefficients calculated as described by Klassen et al. (1998). The temperature dependence of the solubility yields a heat of dissolution of  $-38 (\pm 4) \text{kJ mol}^{-1}$ .
- (e) Rotating CWFT with stirring of the  $\text{H}_2\text{SO}_4$  film (58–70 wt %). HOBr (typically  $10^{-10}$  atm) was generated by reacting  $\text{BrONO}_2$  with water and detected using  $\text{SF}_6^-$  chemi-ions.  $H_{\text{HOBr}}$  was determined by measuring the equilibrium amount of HOBr taken up to a known volume of  $\text{H}_2\text{SO}_4$ . Correction was applied to take into account the uptake of e.g.  $\text{Br}_2\text{O}$  with the total error assigned as 50 %. The temperature dependence of the solubility yields a heat of dissolution of  $-52 (\pm 15) \text{kJ mol}^{-1}$ . Values of  $H_{\text{HOBr}}$  at temperatures other than 250 K were not listed and no parameterisation was given, hence the values in the table at 260 and 270 K were taken from a plot and are approximate.



## Preferred values

Parameter	Value	T/K
$H_{\text{HOBr}}$	$5.22 \times 10^{-5} \exp(5427/T)$	210–270
Reliability		
$\Delta \log H_{\text{HOBr}}$	$\pm 0.5$	210–270

## Comments on preferred value

At low HOBr concentrations its uptake is time dependent and reversible (Waschewsky and Abbatt, 1999; Hanson, 2003) so that a net uptake coefficient is insufficient to describe the interaction, which is driven by solubility and liquid phase diffusion. Waschewsky and Abbatt (1999) suggest that the previous observation (Abbatt, 1995) of irreversible loss of HOBr to a  $\text{H}_2\text{SO}_4$  surface was due to self-reaction caused by high [HOBr]. Waschewsky and Abbatt (1999) and Hanson (2003) found no significant variation of  $H_{\text{HOBr}}$  with  $\text{H}_2\text{SO}_4$  concentration at any given temperature, potentially a result of protonation which may increase the effective solubility at high  $\text{H}_2\text{SO}_4$  wt % and compensate the expected decrease in physical solubility.

The data of Waschewsky and Abbatt (1999) and Hanson (2003) were obtained in different temperature regimes and extrapolated data are in poor agreement with much larger values (at some temperatures by a factor of >10) obtained by Hanson (2003). Hanson (2003) presents arguments that larger values of  $H_{\text{HOBr}}$  are more compatible with measurements of the rate coefficient for reaction of HOBr with HCl in  $\text{H}_2\text{SO}_4$ , which otherwise would exceed the diffusion limit. In addition, he argues that the use of time dependent uptake coefficients to derive  $H_{\text{HOBr}}$  is more prone to systematic error than measurements in an equilibrated solution that also do not require estimates of  $D_1$ . Following Hanson (2003), our preferred value for  $H_{\text{HOBr}}$  uses an average heat of solvation from Waschewsky and Abbatt (1999) and Hanson (2003), and the value of  $H_{\text{HOBr}}$  at 250 K derived by Hanson (2003).

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

Abbatt, J. P. D.: Geophys. Res. Lett., 21, 665–668, 1994.

Abbatt, J. P. D.: J. Geophys. Res., 100, 14009–14017, 1995.

Hanson, D. R.: J. Geophys. Res., 108, art. no. 4239, doi:10.1029/2002JD002519, 2003.

5 Hanson, D. R. and Ravishankara, A. R.: Geophys. Res. Lett., 22, 385–388, 1995.

Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.

Waschewsky, G. C. G. and Abbatt, J. P. D.: J. Phys. Chem. A, 103, 5312–5320, 1999.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



HBr + H<sub>2</sub>SO<sub>4</sub> → products

## Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Solubility: H<sup>+</sup> (M atm<sup>-1</sup>)</i>				
log( <i>H<sup>+</sup></i> ) = 2680/ <i>T</i> - 5.00	54	200–240	Williams et al. (1995)	Knudsen-MS (a)
log( <i>H<sup>+</sup></i> ) = 2480/ <i>T</i> - 4.93	60			
log( <i>H<sup>+</sup></i> ) = 2680/ <i>T</i> - 4.87	66			
log( <i>H<sup>+</sup></i> ) = 2680/ <i>T</i> - 4.80	72			
5.95 × 10 <sup>6</sup>	59.6	210	Abbatt (1995)	CWFT-MS (b)
log( <i>H<sup>+</sup></i> ) = 2900/ <i>T</i> - 3.8	40.3	218–298	Abbatt and Nowak (1997)	CWFT-MS (c)
log( <i>H<sup>+</sup></i> ) = 2800/ <i>T</i> - 4.9	48.8			
log( <i>H<sup>+</sup></i> ) = 2400/ <i>T</i> - 4.5	59.6			
log( <i>H<sup>+</sup></i> ) = 2400/ <i>T</i> - 4.9	50.5			
log( <i>H<sup>+</sup></i> ) = 2500/ <i>T</i> - 5.8	64.4			
log( <i>H<sup>+</sup></i> ) = 2200/ <i>T</i> - 5.8	69.8			
see comment and preferred values		195–250	Kleffmann et al. (2000)	(d)
<i>Diffusion coefficient: D<sub>1</sub> (cm<sup>2</sup> s<sup>-1</sup>)</i>				
7.9 × 10 <sup>-8</sup> / <i>T</i> / <i>η</i>	30–72	220–300	Klassen et al. (1998)	(e)

## 5 Comments

(a) Values of  $HD^{0.5}$  were extracted from time dependent HBr uptake to H<sub>2</sub>SO<sub>4</sub> films of 59.8, 66.1 and 71.9 wt % and at temperatures between 200 and 240 K. Values of  $HD^{0.5}$  were converted to solubilities using diffusion coefficients calculated using  $D = cT/\eta$  where  $\eta$  is the viscosity of H<sub>2</sub>SO<sub>4</sub> (taken from Williams and Long, 1995) and  $c$  was estimated to be  $1 \times 10^{-7} \text{ cm}^2 \text{ cP/sK}$ . This value of  $c$  compared well to later measured values (Klassen et al., 1998). Experiments were also conducted by measuring the vapour pressure of HBr above its solution (at known concentration) in H<sub>2</sub>SO<sub>4</sub> at 54, 60 and 66 wt %.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- (b) Values of  $HD^{0.5}$  were extracted from time dependent HBr uptake to  $H_2SO_4$  films of 59.6, 64.4 and 69.8 wt % and at temperatures between 233 and 200 K.  $H$  was extracted using diffusion coefficients for HBr in  $H_2SO_4$  as described by Williams et al. (1995) (see comment (a)). Only the result for  $HD^{0.5}$  at 210 K (59.6 wt %) was listed though data were obtained over the temperature/concentration range indicated. The solubility listed was calculated using temperature and concentration dependent diffusion coefficients of Klassen et al. (1998).
- (c) Equilibrium concentrations of HBr were measured above cold, HBr-containing  $H_2SO_4$  samples of known HBr concentration composition and wt % (40.3–60.5 %). Solubility data at room temperature were obtained by bubbling  $N_2$  through a HBr/ $H_2SO_4$  solution of known composition and analysing the HBr mole fraction in the gas by re-dissolution in  $H_2O$  with electrical conductivity measurement of ion concentrations. The data listed also contain a parameterisation of measurements (at 59.6, 64.4 and 69.8 wt %) reported in Abbatt (1995), and described in comment (b).
- (d) Head space above stirred, bulk solutions of known composition was analysed by tuneable diode laser absorption spectroscopy of HBr.  $H_2SO_4$  solutions of 53–75 wt % were used.
- (e) Modified diaphragm cell method. The value of the parameter  $\gamma$  as independent of the  $H_2SO_4$  concentration and agrees well (within 5%) with calculations using the Wilke and Chang (1955) expression.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$m_1$ (wt % <sup>-2</sup> K)	$-1.977 \times 10^{-4}$	
$m_2$ (wt % <sup>-1</sup> K)	$-2.096 \times 10^{-2}$	
$m_3$ (K)	4.445	
$b_1$ (wt % <sup>-2</sup> )	$-8.979 \times 10^{-5}$	
$b_2$ (wt % <sup>-1</sup> )	$2.141 \times 10^{-2}$	
$b_3$	-6.067	
$c$ (cm <sup>2</sup> cP K <sup>-1</sup> s <sup>-1</sup> )	$7.9 \times 10^{-8}$	220–300
$A$	$169.5 + 5.18 wt - 0.0825 wt^2$ $+ 3.27 \times 10^{-3} wt^3$	
$T_0$	$144.11 + 0.166 wt - 0.015 wt^2$ $+ 2.18 \times 10^{-4} wt^3$	
<i>Reliability</i>		
$\Delta \log H^*$	$\pm 0.3$	

## Comments on preferred value

The available data on the interaction of HBr with H<sub>2</sub>SO<sub>4</sub> solutions indicates very high solubility, especially in dilute solutions and at low temperatures. Efficient dissociation of HBr in H<sub>2</sub>SO<sub>4</sub> implies that the solubilities are effective ( $H^*$ ). Data have been obtained using static or stirred samples with head-space analysis or using time dependent uptake coefficients in combination with diffusion coefficients of HBr in H<sub>2</sub>SO<sub>4</sub>. Most authors estimate the overall uncertainty as being better than a factor of two, which is borne out by the reasonable agreement from group to group. Kleffmann et al. (2000) have analysed the experimental datasets to produce a H<sub>2</sub>SO<sub>4</sub> concentration and temperature dependent parameterisation of HBr effective solubility. This is adopted in our

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

table of preferred values in which:

$$\log_{10} H^* = 1000 m/T + b \text{ where}$$

$$m = m_1[\text{H}_2\text{SO}_4]^2 + m_2[\text{H}_2\text{SO}_4] + m_3 \text{ and } b = b_1[\text{H}_2\text{SO}_4]^2 + b_2[\text{H}_2\text{SO}_4] + b_3$$

5 and the  $\text{H}_2\text{SO}_4$  concentration  $[\text{H}_2\text{SO}_4]$  is in wt %

Note that the experimentally derived solubilities are factors of 2–6 lower than those calculated using activity coefficients (Luo et al., 1995).

Klassen et al. (1998) have parameterised the diffusion coefficient for HBr in  $\text{H}_2\text{SO}_4$  as:

$$10 D_1 = cT/\eta$$

where  $c$  is a constant ( $\text{cm}^2 \text{cPK}^{-1} \text{s}^{-1}$ ) and  $\eta$  is the viscosity of  $\text{H}_2\text{SO}_4$  at a given wt % and temperature. Shi et al. (2001), have taken viscosity data from Williams and Long (1995) and for pure water to derive an extended formulation to cover  $\text{H}_2\text{SO}_4$  viscosities from 0 to 80 wt %, which can be used with the  $c$  constant above to derive  $D_1$  over the same concentration range using:

$$\eta = AT^{-1.43} \exp(448/(T - T_0))$$

This parameterisation results in values of  $\eta$  which agree to better than 10 % (for 40–70 wt %  $\text{H}_2\text{SO}_4$  solutions) with those of Klassen et al. (1998).

## References

- 20 Abbatt, J. P. D.: J. Geophys. Res.-Atmos., 100, 14009–14017, 1995.  
Abbatt, J. P. D. and Nowak, J. B.: J. Phys. Chem. A, 101, 2131–2137, 1997.  
Klassen, J. K., Hu, Z. J., and Williams, L. R.: J. Geophys. Res.-Atmos., 103, 16197–16202, 1998.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



- Kleffmann, J., Becker, K. H., Broske, R., Rothe, D., and Wiesen, P.: J. Phys. Chem. A, 104, 8489–8495, 2000. (Erratum published in vol. 105, page 11593).
- Luo, B., Carslaw, K. S., Peter, T., and Clegg, S.: Geophys. Res. Lett., 22, 247–250, 1995.
- 5 Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res.-Atmos., 106, 24259–24274, 2001.
- Wilke, C. R. and Chang, P.: AIChE J., 1, 264–270, 1955.
- Williams, L. R., Golden, D. M., and Huestis, D. L.: J. Geophys. Res.-Atmos., 100, 7329–7335, 1995.
- 10 Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

## VI.A4.18

### HOI + H<sub>2</sub>SO<sub>4</sub> → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficients: <math>\alpha_b</math></i>			
>0.3 (>58 wt % H <sub>2</sub> SO <sub>4</sub> )	253	Holmes et al. (2001)	CWFT-MS (a)
<i>Uptake coefficients: <math>\gamma</math></i>			
$1 \times 10^{-5} \exp\{(1700 \pm 200)/T\}$	195–252	Allanic and Rossi (1999)	Knudsen-MS (b)
$>4 \times 10^{-2}$ (>58 wt % H <sub>2</sub> SO <sub>4</sub> )	253	Holmes et al. (2001)	CWFT-MS (a)

#### 5 Comments

(a) HOI (at concentrations  $<10^{10}$  molecule cm<sup>-3</sup>) was formed in the reaction of O(<sup>3</sup>P) with C<sub>3</sub>H<sub>7</sub>I. The uptake to >58 wt % H<sub>2</sub>SO<sub>4</sub> was continuous and its size independent of the HOI concentration. No evidence for formation of iodine containing products was observed. The surface concentration of H<sub>2</sub>SO<sub>4</sub> could have been much greater than the initial value of 58 wt % as the bulk gas-flow was not humidified.

(b) HOI (at approximate concentrations of  $\sim 10^{10}$  molecule cm<sup>-3</sup>) was formed in the reaction of O(<sup>3</sup>P) with C<sub>2</sub>H<sub>5</sub>I. 40–70 wt % H<sub>2</sub>SO<sub>4</sub> solutions were kept at temperatures between 195 and 252 K. The gas-phase H<sub>2</sub>O flow was adjusted to prevent surface concentration changes. The uptake coefficients displayed a negative temperature dependence, with the parameters in the Table derived from an unweighted fit to the datapoints reported at each temperature.



## Preferred values

None

## Comments on preferred values

Both Allanic and Rossi (1999) and Holmes et al. (2001) found the uptake of HOI to  $\text{H}_2\text{SO}_4$  solutions to be continuous (i.e. non-saturating), suggestive of very high solubility, possibly enhanced by protonation to form  $\text{H}_2\text{OI}^+$ . The continuous uptake could also be caused by the presence of impurities in the HOI source reacting with HOI in/on  $\text{H}_2\text{SO}_4$ .

Allanic and Rossi observed the immediate formation of  $\text{I}_2$  as gas-phase product, albeit at a yield which was estimated to account for less than 50% of the HOI taken up. In contrast, Holmes et al. report no gas-phase iodine containing products. It is conceivable that the products observed by Allanic and Rossi were formed on surfaces other than  $\text{H}_2\text{SO}_4$  in their Knudsen reactor. The single datapoint of Holmes et al. at 253 K does not fit with the trend in  $\gamma$  with temperature observed by Allanic and Rossi, but is too high. This may reflect enhanced uptake to very concentrated  $\text{H}_2\text{SO}_4$  surfaces which may have been present in the Holmes et al. experiments. On the other hand, Allanic and Rossi (1999) did not observe a change in HOI uptake coefficient when the  $\text{H}_2\text{O}$  partial pressure above an initially 70 wt %  $\text{H}_2\text{SO}_4$  solution was increased by a factor of 4.

Upon adding NaCl to the  $\text{H}_2\text{SO}_4$  solution, Holmes et al. (2001) observed ICl formation with 100% yield, suggesting that dissolved HOI (or  $\text{H}_2\text{OI}^+$ ) reacts efficiently with halide ions.

The accommodation coefficient presented by Holmes et al. (2001) was derived from observation of pressure dependent values of  $\gamma$  on  $\text{H}_2\text{SO}_4$  solutions of largely unknown composition and thus no recommendation is made.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

Allanic, A. and Rossi, M. J.: J. Geophys. Res., 104, 18689–18696, 1999.

Holmes, N. S., Adams, J. W., and Crowley, J. N.: Phys. Chem. Chem. Phys., 3, 1679–1786, 2001.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

32402



Acetone + H<sub>2</sub>SO<sub>4</sub> (aq) → products

## Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
$\gamma_0 = 0.7$	85	200	Duncan et al. (1998)	(a)
$\gamma_{ss} = 0.4$	85	200		
$\gamma_{ss} = 0.25$	75	201	Duncan et al. (1999)	(a)
$\gamma_{ss} = 0.48$	85	221		
$\gamma_{ss} = 0.57$	85	221		
$\gamma_{ss} = 5.2 \times 10^{-5}$	89.4	298	Esteve and Noziere (2005)	(b)
$\gamma_{ss} = 5.6 \times 10^{-6}$	85.0			
$\gamma_{ss} = 4.7 \times 10^{-6}$	80.8			
$\gamma_{ss} = 1.6 \times 10^{-6}$	73.9			
<i>Solubility: <math>H^*</math> (Matm<sup>-1</sup>)</i>				
$H^* = 7.1 \times 10^7$	70	180	Duncan et al. (1999)	(a)
$H^* = 1.2 \times 10^7$	70	187		
$H^* = 2.4 \times 10^6$	70	195		
$\log H^* = -7.50 + 2920/T$	67.7	220–239	Klassen et al. (1999)	Knud-MS (c)
$\log H^* = -7.43 + 2750/T$	60.7	223–234		
$\log H^* = -7.39 + 2650/T$	54.4	213–226		
$\log H^* = -7.35 + 2580/T$	48.3	212–222		
$\ln H^* = 3.0 - [H_2SO_4]_f - 4850$ ( $1/298.15 - 1/T$ )	40–75	198–292	Kane et al. (1999)	WWFT-MS (d)
$H^* = 2.1 \times 10^4$	50	230	Imamura and Akiyoshi (2000)	WWFT-PIMS (e)
$H^* = 1.3 \times 10^3$	50	250		
$H^* = 8.2 \times 10^4$	60	230		
$H^* = 4.4 \times 10^3$	60	250		
$H^* = 2.5 \times 10^4$	69	250		
$H^* = 2.2 \times 10^5$	76	250		
$H^* = 2.2 \times 10^4$	76	270		
$H^* = 9.5 \times 10^4$	79	270		

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments

- (a) Low pressure, molecular flow reactor with FTIR and MS. Reflection-Absorption infrared spectroscopic analysis of  $\text{H}_2\text{SO}_4$  films (0.05–0.1  $\mu\text{m}$  thick, 180–260 K, 35–96 wt %  $\text{H}_2\text{SO}_4$ ) exposed to acetone ( $10^{-7}$ – $10^{-4}$  mbar) reveal presence of mesityl oxide and trimethylbenzene. The reaction proceeds via formation of protonated acetone, which, in  $\text{H}_2\text{SO}_4$  solutions of >70 wt %, can undergo self-dehydration to form mesityl oxide. Trimethylbenzene is formed from mesityl oxide in solutions of >85 wt %. At concentrations lower than 70 wt % acetone uptake is reversible and the effective solubility,  $H^*$ , was determined by integrating the uptake until the thin film was saturated. Errors on  $H^*$  are estimated as 25 %. Uptake coefficients obtained by gas-analysis using MS. The enthalpy and entropy of solution were determined as  $-66 \text{ kJ mol}^{-1}$  and  $-249 \text{ J mol}^{-1} \text{ K}^{-1}$ , respectively. The solubility data (180–195 K) can be described by:  $H^* = \exp\{(T * 216.7 - 66000)/8.314 * T\}$  for 70 wt %  $\text{H}_2\text{SO}_4$ .
- (b) Rotated WWFT generally at  $\approx 100$  mbar with MS detection of acetone at high concentrations ( $1$ – $5 \times 10^{-5}$  atm). The experimental uptake coefficient,  $\gamma_{\text{ss}}$ , was observed to increase with [acetone]. Reactive uptake observed at  $[\text{H}_2\text{SO}_4] > 70$  wt % and reversible uptake at lower concentrations. Uptake coefficients were corrected for gas diffusion.
- (c) Knudsen reactor using 48–68 wt %  $\text{H}_2\text{SO}_4$  at 210–240 K. Solubility determined by measuring the time dependence of the acetone uptake and using liquid-phase diffusion coefficients calculated from viscosity data. Uncertainty estimated as  $\pm 33$  %. Addition of ammonium sulphate reduced the solubility slightly.
- (d) WWFT reactor using 40–87 wt %  $\text{H}_2\text{SO}_4$  at 198–300 K. Uptake in dilute solutions (<75 wt %) was found to be reversible. Solubility determined by measuring the time dependence of the acetone uptake and using liquid-phase diffusion coefficients calculated from viscosity data. The dependence of the solubility on the

32404

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$\text{H}_2\text{SO}_4$  concentration and temperature is given by the equation in the table where  $f = 0.23 + 5/T$  and  $[\text{H}_2\text{SO}_4]$  is the molality of the  $\text{H}_2\text{SO}_4$  solution. Note that the original expression given by the authors contains a typographical error, which has been subsequently corrected (Leu, 2003). The error limits on  $H^*$  are reported to be 50%. In  $\text{H}_2\text{SO}_4$  solutions of greater than 80 wt %, both mesityl oxide and trimethylbenzene were observed by MS when liquid acetone was mixed with  $\text{H}_2\text{SO}_4$ .

(e) WWFT with PIMS detection of reversible acetone uptake to thin  $\text{H}_2\text{SO}_4$  films (50–79 wt %) at 230–270 K. Solubility determined by integrating the acetone uptake until the film was saturated. Uncertainties in  $H^*$  reported to be  $\pm 30\%$ .

### Preferred values

Parameter	Value	$T/\text{K}$
$\ln H^*$	$3.0 - [\text{H}_2\text{SO}_4]f - 4850 (1/298.15 - 1/T)$ (40–75 wt % $\text{H}_2\text{SO}_4$ )	195–295 K
Reliability		
$\Delta \log H$	$\pm 0.5$	

### Comments on preferred values

The experimental data show that acetone is protonated to some extent in  $\text{H}_2\text{SO}_4$  solutions and, at  $\text{H}_2\text{SO}_4$  concentrations greater than 75 wt %, dehydration reactions result in the formation of mesityl oxide and trimethylbenzene. At concentrations lower than 70 wt % acetone uptake is reversible and solubility controls the amount taken up to the solution. The preferred values for the effective Henry's law solubility are taken from the most extensive study by Kane et al. (1999), which covered the largest temperature and  $\text{H}_2\text{SO}_4$  concentration range.  $f = 0.23 + 5/T$  and  $[\text{H}_2\text{SO}_4]$  is the molality of the  $\text{H}_2\text{SO}_4$  solution. Where comparison is possible (overlap in  $T$  and  $\text{H}_2\text{SO}_4$  wt %) the agreement between the studies is generally within a factor of 2–4. For this reason we present error limits significantly greater than the 25–50 % uncertainties reported by the various

authors. The uptake coefficients presented by Duncan et al. (1998, 1999) and Esteve and Noziere (2005) diverge by several orders of magnitude, most likely due to the different temperature regimes investigated, and no recommendation is given.

## References

- 5 Duncan, J. L., Schindler, L. R., and Roberts, J. T.: Geophys. Res. Lett., 25, 631–634, 1998.
- Duncan, J. L., Schindler, L. R., and Roberts, J. T.: J. Phys. Chem. B, 103, 7247–7259, 1999.
- Esteve, W. and Noziere, B.: J. Phys. Chem. A, 109, 10920–10928, 2005.
- 10 Imamura, T. and Akiyoshi, H.: Geophys. Res. Lett., 27, 1419–1422, 2000.
- Kane, S. M., Timonen, R. S., and Leu, M. T.: J. Phys. Chem. A, 103, 9259–9265, 1999.
- Klassen, J. K., Lynton, J., Golden, D. M., and Williams, L. R.: J. Geophys. Res., 104, 26355–26361, 1999.
- Leu, M. T.: Int. Rev. Phys. Chem., 22, 341–376, 2003.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.20

### CHBr<sub>3</sub> + H<sub>2</sub>SO<sub>4</sub> → products

#### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>			
$3 \times 10^{-3}$	200	Hanson and Ravishankara (1993)	CWFT-CIMS (a)

#### 5 Comments

(a) Uptake study using a coated wall laminar flow tube coupled to CIMS detection. On liquid 58 % H<sub>2</sub>SO<sub>4</sub> a reversible physical adsorption was observed with an initial uptake coefficient  $\gamma_0 = 3 \times 10^{-3}$  (displayed) and 0.5 s exposure time within the wall-coated laminar flow tube. For longer exposure times (~20 s)  $\gamma$  decreases to  $2 \times 10^{-4}$ , most likely owing to partial saturation of the uptake and a rate-limiting acid-catalyzed hydrolysis.

#### Preferred values

No recommendation.

In view of the single determination with very few experimental data at one temperature and one H<sub>2</sub>SO<sub>4</sub> concentration no recommendation is given.

#### References

Hanson, D. R. and Ravishankara, A. R.: NATO ASI Series (17), 281, in: The Tropospheric Chemistry of Ozone in the Polar Regions, edited by: Niki, H. and Becker, K. H., Springer Verlag, 1993.

32407

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/ wt %	Temp./K	Reference	Technique/Comments
<i>Accommodation coefficient: <math>\alpha_b</math></i>				
0.02–0.09	43	298	Van Dingenen and Raes (1991)	(a)
>0.4	73–98	303	Pöschl et al. (1998)	CWFT-CIMS (b)
0.96 ± 0.14	7–40	295	Hanson (2005)	AFT-CIMS (c)
0.76 ± 0.05	47–65			

### Comments

(a) Uptake of H<sub>2</sub>SO<sub>4</sub> was derived from particle growth measurements in a 230 L spherical batch reactor. Particle size distributions were measured with a differential mobility analyser. Seed H<sub>2</sub>SO<sub>4</sub> particles were first produced by reaction of SO<sub>2</sub> with OH. OH was produced by photolysis of a NO<sub>x</sub> mixture, thus likely from photolysis of HNO<sub>2</sub> formed on the chamber walls. Loss of these particles within the reactor was then measured under dark conditions. Growth of the particles was observed in a third stage by again inducing H<sub>2</sub>SO<sub>4</sub> formation as before, but with higher light intensity, resulting in growth of H<sub>2</sub>SO<sub>4</sub> particles due to condensation of H<sub>2</sub>SO<sub>4</sub>. OH concentrations were estimated from the decay of the ratio of the concentrations of propene and propane also added to the reactor. The estimated OH level was used to infer the H<sub>2</sub>SO<sub>4</sub> concentrations. When deriving the bulk accommodation coefficient from the growth rates, the Fuchs-Sutugin correction factor was used to account for gas phase diffusion.

(b) Coated wall flow tube operated at 0.5–10 Torr with CIMS detection of H<sub>2</sub>SO<sub>4</sub> after reaction with SF<sub>6</sub><sup>-</sup>. H<sub>2</sub>SO<sub>4</sub> was delivered from a thermostated reservoir kept at

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)




## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



348 K. The reactor wall was coated by a thin  $\text{H}_2\text{SO}_4$  film, the composition of which was determined by the  $\text{H}_2\text{O}$  pressure. The uptake coefficient (interpreted as bulk accommodation coefficient) was determined by fitting the first order loss rates taking into account gas phase diffusion. The best fit was obtained for a value of 0.65. Taking into account the uncertainties led to the lower limit listed in the table.

- (c) Uptake to sulphuric acid aerosol was studied in a laminar flow reactor coupled to CIMS detection using  $\text{HNO}_3$  as source of primary ions. Sulphuric acid particles were generated by homogeneous nucleation from supersaturated vapour leading to a lognormal particle size distribution within 50–120 nm, with a few  $10^4$  particles per  $\text{cm}^3$ , characterised by a differential mobility analyzer. Gas-phase sulphuric acid was generated by reaction of OH with  $\text{SO}_2$  leading to concentrations of  $3 \times 10^9 \text{ cm}^{-3}$  in the flow tube. This concentration is at least two orders of magnitude above the vapour pressure of sulphuric acid. Uptake to the particles did not lead to significant growth. The measured uptake coefficients were corrected for gas phase diffusion using the Fuchs-Sutugin correction factor. The diffusion coefficient was directly measured based on the observed wall loss rates in absence of aerosol particles. Its negative humidity dependence was in line with previous measurements by Hanson and Eisele (2000).

### Preferred values

Parameter	Value	T/K
$\alpha_b, [\text{H}_2\text{SO}_4] < 50 \text{ wt } \%$	1	200–300
$\alpha_b, [\text{H}_2\text{SO}_4] > 50 \text{ wt } \%$	>0.7	200–300
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	200–300

## Comments on preferred values

Of the two more recent studies, the aerosol flow tube study by Hanson (2005) is much less affected by gas phase diffusion and can thus more reliably return high uptake coefficients. The slight trend for a lower bulk accommodation coefficient apparent in Hanson's data is not significant enough to provide a parameterized expression for the composition dependence. It is thus reflected by providing a lower limit of 0.7 for solutions >50 wt %. The coated wall flow tube study by Pöschl et al. (1998) is in agreement with these high values for  $\alpha_b$ , keeping in mind that the wall loss rates were practically in the diffusion limit of that experiment. The earlier estimate of  $\alpha_b$  based on particle growth measurements certainly suffered from the fact that neither the concentration of  $\text{H}_2\text{SO}_4$  nor that of its precursor  $\text{SO}_2$  was directly measured. In addition, the condensation rate derived was very sensitive to the wall loss rate of particles, which had to be determined in a separate experiment.

## References

- Hanson, D. R.: J. Phys. Chem. A, 109, 6919–6927, 2005.  
Pöschl, U., Canagaratna, M., Jayne, J. T., Molina, L. T., Worsnop, D. R., Kolb, C. E., and Molina, M. J.: J. Phys. Chem. A, 102, 10082–10089, 1998.  
Van Dingenen, R. and Raes, F.: Aerosol Sci. Technol., 15, 93–106, 1991.  
Hanson, D. R. and Eisele, F.: J. Phys. Chem. A, 104, 1715–1719, 2000.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

HOCl + HCl (aqueous sulphuric acid aerosol) → H<sub>2</sub>O + Cl<sub>2</sub>

## Experimental data

Parameter	$p_{\text{HOCl}}/\text{mbar}$	$p_{\text{HCl}}/\text{mbar}$	$[\text{H}_2\text{SO}_4]/\text{wt}\%$	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>						
$\gamma_{\text{HOCl}} = 1 \times 10^{-3}$		$7 \times 10^{-6}$	60	263	Hanson and Ravishankara (1991)	WWFT-CIMS (a)
$\gamma_{\text{HCl}} = 6 \times 10^{-3}$	$1.5 \times 10^{-5}$	$9 \times 10^{-6}$	60	215		
$\gamma_{\text{HOCl}} = 0.02 - 0.15$		$10^{-8} - 10^{-6}$	59.6	200–205	Hanson and Ravishankara (1993)	WWFT-CIMS (b)
$\gamma_{\text{HOCl}} = 0.45$		$4 \times 10^{-7}$	50	198	Zhang et al. (1994)	WWFT-MS (c)
$\gamma_{\text{HOCl}} = 7 \times 10^{-2}$		$4 \times 10^{-7}$	60	205		
$\gamma_{\text{HOCl}} = 4 \times 10^{-3}$		$4 \times 10^{-7}$	68	214		
$\gamma_{\text{HOCl}} = 0.1$		$4 \times 10^{-7}$	57	202		
$\gamma_{\text{HOCl}} = 0.2$		$8 \times 10^{-7}$	50	198		
$\gamma_{\text{HOCl}} = 0.5$		$2 \times 10^{-6}$	60	205		
$\gamma_{\text{HOCl}} = 4.5 \times 10^{-3}$		$2 \times 10^{-6}$	59.6	251	Hanson and Lovejoy (1996)	RWFT-CIMS (d)
$\gamma_{\text{HOCl}} = 1.3 \times 10^{-2}$		$2 \times 10^{-5}$				
$\gamma_{\text{HOCl}} = 2.7 \times 10^{-2}$		$7 \times 10^{-5}$				
$\gamma_{\text{HOCl}} = 2.5 \times 10^{-2}$ (70 nm)		$3 \times 10^{-3}$	59.5	251	Hanson and Lovejoy (1996)	AFT-CIMS (d)
$\gamma_{\text{HOCl}} = 7.5 \times 10^{-2}$ (200 nm)		$3 \times 10^{-3}$		251		
$\gamma_{\text{HOCl}} = 1.0 \times 10^{-1}$ (400 nm)		$3 \times 10^{-3}$				
$\gamma_{\text{HOCl}} = 0.07$ ( $a_{\text{H}_2\text{O}} = 0.10$ )		$3 \times 10^{-3}$				
$\gamma_{\text{HOCl}} = 0.4$ ( $a_{\text{H}_2\text{O}} = 0.28$ )		$3 \times 10^{-3}$				
$\gamma_{\text{HOCl}} = 0.86$ ( $a_{\text{H}_2\text{O}} = 0.56$ )		$3 \times 10^{-3}$				
$\gamma_{\text{HCl}} = 0.1$ (38 nm)	$1.5 \times 10^{-3}$		34	274		
$\gamma_{\text{HCl}} = 0.32$ (160 nm)	$1.5 \times 10^{-3}$		34			
$\gamma_{\text{HCl}} = 0.58$ (420 nm)	$1.5 \times 10^{-3}$		34			
$\gamma_{\text{HCl}} = 0.08$ ( $a_{\text{H}_2\text{O}} = 0.58$ )	$1.5 \times 10^{-3}$			272–276		
$\gamma_{\text{HCl}} = 0.44$ ( $a_{\text{H}_2\text{O}} = 0.69$ )	$1.5 \times 10^{-3}$					
$\gamma_{\text{HCl}} = 0.75$ ( $a_{\text{H}_2\text{O}} = 0.79$ )	$1.5 \times 10^{-3}$					
$\gamma_{\text{HOCl}} = 0.013$		$5 \times 10^{-7}$	52	250	Donaldson et al. (1997)	RWFT-CIMS (e)
$\gamma_{\text{HOCl}} = 0.09$		$2 \times 10^{-5}$	52			
$\gamma_{\text{HOCl}} = 0.011$		$2 \times 10^{-5}$	62.5	250		
$\gamma_{\text{HOCl}} = 0.045$		$2 \times 10^{-4}$	62.5			
$\gamma_{\text{HOCl}} = 0.0048$		$4 \times 10^{-5}$	67	250		
$\gamma_{\text{HOCl}} = 0.018$		$1 \times 10^{-3}$	67			
$\gamma_{\text{HOCl}} = 0.043$		$5 \times 10^{-7}$	58	220		
$\gamma_{\text{HOCl}} = 0.15$		$1 \times 10^{-5}$	58			

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Parameter	$p_{\text{HOCl}}/$ mbar	$p_{\text{HCl}}/$ mbar	$[\text{H}_2\text{SO}_4]/$ wt %	Temp./ K	Reference	Technique/Comments
$Y_{\text{HOCl}} = 0.05$		$2 \times 10^{-6}$	60	220		
$Y_{\text{HOCl}} = 0.12$		$4 \times 10^{-5}$	60			
$Y_{\text{HOCl}} = 0.023$		$1 \times 10^{-6}$	63	220		
$Y_{\text{HOCl}} = 0.070$		$2 \times 10^{-5}$	63			
$Y_{\text{HOCl}} = 2.0 \times 10^{-3}$ (70 nm)		$10^{-3}$	67	250		AFT-CIMS (e)
$Y_{\text{HOCl}} = 5.5 \times 10^{-3}$ (180 nm)		$10^{-3}$				
$Y_{\text{HOCl}} = 1.3 \times 10^{-2}$ (400 nm)		$10^{-3}$				
$Y_{\text{HOCl}} = 1.3 \times 10^{-2}$ (680 nm)		$10^{-3}$				
<i>Reacto-diffusive parameters</i>						
$H_{\text{HOCl}}(D_b k_b H_{\text{HCl}}^*)^{1/2}$ $= 4.0 \times 10^4 \text{ M cm s}^{-1} \text{ atm}^{-3/2}$			59.6	251	Hanson and Lovejoy (1996)	RWFT-CIMS (d)
$l_{\text{HOCl}} = 1.2 \times 10^{-5} \text{ cm}$			59.5	251	Hanson and Lovejoy (1996)	AFT-CIMS (d)
$l_{\text{HCl}} = 1.2 \times 10^{-4} \text{ cm}$			34	274		
$(k_b H_{\text{HOCl}})^{1/2} = 1.1 \times 10^4 (\text{atm s})^{-1/2}$			34	274		

### Comments

- (a) The cold aqueous solution of  $\text{H}_2\text{SO}_4$  flowed down the walls of the vertical flow tube.  $p(\text{H}_2\text{O})$  was generally around  $1.3 \times 10^{-3}$  mbar.  $[\text{HCl}]$  for the experiments on the 40 % solution was  $2 \times 10^{11}$ , for the 60 to 75 % solution it was in the range  $1$  to  $3 \times 10^{11}$  molecule  $\text{cm}^{-3}$ .
- (b) First order loss of HOCl was observed in a cylindrical flow tube with  $\text{H}_2\text{SO}_4$  applied as a coating for the 200–205 K measurements or contained in a boat for the measurement at 219 K. HOCl was generated by reacting HF with  $\text{Ca}(\text{OCl})_2$  powder. HCl was added via uptake from the gas phase at pressures  $>3 \times 10^{-7}$  mbar, leading to  $[\text{Cl}^-] > 10^{-5}$  M, in excess of HOCl. The value for  $k_b$  reported and listed in the table was obtained by estimates of the diffusion coefficients and solubilities as reported from the same study.
- (c) ~~Wetted wall flow tube~~ with water vapour pressure maintained constant at  $1.5 \times 10^{-3}$  Torr, to give 45–70 wt %  $\text{H}_2\text{SO}_4$  solutions for  $195 \text{ K} < T < 220 \text{ K}$ . Detection limits were  $5 \times 10^{-8}$  Torr for HOCl  $5 \times 10^{-7}$  Torr for HCl. Production of  $\text{Cl}_2$

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



matched the loss of HOCl or HCl. Experiments with both excess of HCl and HOCl were performed. Uptake coefficients were consistent between analysis of reactant loss and  $\text{Cl}_2$  formation. Kinetics were not affected by the presence of  $\text{HNO}_3$  at  $6.5 \times 10^{-7}$  mbar.

- (d) Bulk measurements of first-order loss of either HOCl or HCl under conditions of mutual excess of the other, respectively. HOCl, HCl and  $\text{Cl}_2$  were detected using  $\text{SF}_6^-$  reactant ion. The parameter listed in the table is the slope of a plot of  $1/\gamma$  vs. the square root of the HCl pressure.

Aerosol flow tube experiment operated at 180 Torr of  $\text{N}_2$ ,  $10 \text{ STP cm}^3 \text{ s}^{-1}$  and 251–276 K. The aerosol mode diameter was 200 nm with a geometric standard deviation of 1.2. Under the conditions of the experiment, for both HOCl and HCl diffusion controlled loss to the flow tube walls occurred.

- (e) ~~Rotating wetted wall and aerosol flow tube~~ experiments identical to those done by Hanson and Lovejoy (1996). The total pressure in the ~~wetted wall flow tube~~ was at maximum a few mbar (mostly He), with a flow rate of  $3\text{--}4 \text{ STP cm}^3 \text{ s}^{-1}$ .

### Preferred values

Parameter	Value	T/K
$\alpha_{\text{b,HCl}}$	see datasheet VI.A4.14	
$\alpha_{\text{b,HOCl}}$	1	
$k_{\text{b}} (\text{M}^{-1} \text{s}^{-1})$	$80 [\text{H}^+] T/\eta$	190–273
<i>Reliability</i>		
$\Delta \log(k_{\text{b}})$	$\pm 0.3$	

## Comments on preferred values

HOCl reacts efficiently with HCl under acidic conditions (Eigen and Kustin, 1962). Most of the kinetic data covering tropospheric and lower stratospheric conditions were obtained by one group using coated wall and aerosol flow tubes. The measured uptake coefficients clearly established that kinetics is driven by reaction in the bulk (Lovejoy and Hanson, 1996). Donaldson et al. (1997) suggested protonated HOCl,  $H_2OCl^+$ , being the reactive species, based on the strong sensitivity of the first order loss rate constant in the bulk on acidity and on theoretical evidence (Koch et al., 1997; Francisco and Sander, 1995; and references therein). Shi et al. (2001) suggested the expression adopted here for the second order rate constant to explicitly include the proton molarity and assumed that the rate constant is limited by diffusion of HOCl (leading to a dependence of the apparent rate constant on viscosity). As already pointed out by Shi et al., taking the revised parameterizations for  $H_{HOCl}$  (datasheet VI.A4.13) and  $H_{HOCl}^*$  (datasheet VI.A4.14), this leads to good agreement with all available data of  $\gamma_{HOCl}$  for a wide range of conditions, including those by Zhang et al. (1994), another flow tube experiment at low temperature.

The recommended parameterisation reproduces the available data of  $\gamma_{HOCl}$  and  $\gamma_{HCl}$  reasonably well and notably also the size dependence for both, as measured by Hanson and Lovejoy (1996) and Donaldson et al. (1997). The reacto-diffusive length is in the range of 0.1 to 1  $\mu m$ . The parameterization also reproduces the dependence of the uptake as a function of water activity.

The uptake coefficients of HOCl and HCl on sulphuric acid in presence of HCl or HOCl, respectively, can be obtained from:

$$\frac{1}{\gamma_{HOCl}} = \frac{1}{\alpha_{b,HOCl}} + \frac{\bar{c}_{HOCl}}{4H_{HOCl}RT\sqrt{D_{l,HOCl}\cdot k_b\rho_{HCl}H_{HOCl}^*}\left[\coth\left(\frac{r_p}{l_{HOCl}}\right) - \left(\frac{l_{HOCl}}{r_p}\right)\right]}$$

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$$\frac{1}{Y_{\text{HCl}}} = \frac{1}{\alpha_{\text{b,HCl}}} + \frac{\bar{c}_{\text{HCl}}}{4H_{\text{HCl}}^* RT \sqrt{D_{\text{l,HCl}} \cdot k_{\text{b}} \rho_{\text{HOCl}} H_{\text{HOCl}}} \left[ \coth\left(\frac{r_{\text{p}}}{l_{\text{HCl}}}\right) - \left(\frac{l_{\text{HCl}}}{r_{\text{p}}}\right) \right]$$

$\alpha_{\text{b,HCl}}$  is assumed to be one (see datasheet VI.A4.13). For  $\alpha_{\text{b,HCl}}$  we recommend the parameterization given on datasheet VI.A4.14, which leads to slightly better agreement with data than assuming  $\alpha_{\text{b,HCl}} = 1$ :  $\alpha_{\text{b,HCl}} = k_{\text{sol}} / (k_{\text{des}} + k_{\text{sol}})$  with  $k_{\text{sol}} = 7.84 \times 10^{10} / \eta$  and  $k_{\text{des}} = 8.0 \times 10^{17} \exp(-5000/T)$ .

We suggest using the parameterisation of Shi et al. for the proton concentration:

$$[\text{H}^+] = \exp \left[ 60.51 - 0.095 wt + 0.0077 wt^2 - 1.61 \times 10^{-5} wt^3 - (1.76 + 2.52 \times 10^{-4} wt^2) T^{0.5} + (-805.89 + 253.05 wt^{0.076}) / T^{0.5} \right]$$

We take the preferred expression for the solubility of HOCl from datasheet VI.A4.13,

$$H_{\text{HOCl}} = 1.91 \times 10^{-6} \exp(5862.4/T) \exp(-(0.0776 + 59.18/T)[\text{H}_2\text{SO}_4])$$

and that for HCl from datasheet VI.A4.14:

$$H_{\text{HCl}}^* = (0.094 - 0.61X + 1.2X^2) \exp(-8.68 + (8515 - 10718X^{0.7})/T)$$

The mole fraction,  $X$ , of sulfuric acid is given by  $X = wt / (wt + (100 - wt)98/18)$  and  $[\text{H}_2\text{SO}_4]$  in mol L<sup>-1</sup> can be calculated from the sulfuric acid content using density data provided by Shi et al.

The diffusion coefficients for HOCl and HCl are parameterized by  $D_{\text{l},X} = c_X \times 10^{-8} T / \eta$ ; with  $c_{\text{HOCl}} = \text{cm}^2 \text{ cP K}^{-1} \text{ s}^{-1}$  (see datasheet VI.A4.13) and  $c_{\text{HCl}} = 7.8 \times 10^{-8} \text{ cm}^2 \text{ cP K}^{-1} \text{ s}^{-1}$  (see datasheet VI.A4.14). For the viscosity, we suggest using the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into tropospherically more relevant dilute solutions at high  $T$ :

$$\eta = AT^{-1.43} \exp(448K/(T - T_0)),$$

32415

ACPD

12, 32109–32472, 2012

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



with  $A = 169.5 + 5.18wt - 0.0825wt^2 + 3.27 \times 10^{-3}wt^3$ ,  
and  $T_0 = 144.11 + 0.166wt - 0.015wt^2 + 2.18 \times 10^{-4}wt^3$

5 The reacto-diffusive lengths needed to account for finite particle sizes are given by:

$$l_{\text{HOCl}} = (D_{\text{I,HOCl}} / (k_{\text{b}} \rho_{\text{HOCl}} H_{\text{HOCl}}^*))^{0.5}$$

$$l_{\text{HCl}} = (D_{\text{I,HCl}} / (k_{\text{b}} \rho_{\text{HCl}} H_{\text{HCl}}))^{0.5}$$

## References

- 10 Eigen, M. and Kustin, K. J.: J. Am. Chem. Soc., 84, 1355–1361, 1962.  
Hanson, D. R. and Lovejoy, E. R.: J. Phys. Chem., 100, 6397–6405, 1996.  
Hanson, D. R. and Ravishankara, A. R.: J. Geophys. Res., 96, 17307–17314, 1991.  
Hanson, D. R. and Ravishankara, A. R.: J. Phys. Chem., 97, 12309–12319, 1993.  
Zhang, R. Y., Leu, M. T., and Keyser, L. F.: J. Phys. Chem., 98, 13563–13574, 1994.  
15 Donaldson, D. J., Ravishankara, A. R., and Hanson, D. R.: J. Phys. Chem. A, 101, 4717–4725, 1997.  
Koch, T. G., Banham, S. F., Sodeau, J. R., Horn, A. B., McCoustra, M. R. S., and Chesters, M. A.: J. Geophys. Res., 102, 1513–1522, 1997.  
Francisco, J. S. and Sander, S. P.: J. Chem. Phys., 102, 9615–9618, 1995.  
20 Shi, Q., Jayne, J. T., Kolb, C. E. Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.  
Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## VI.A4.23



### Experimental data

Parameter	Temp./K	Reference	Technique/Comments
<i>Uptake coefficient, <math>\gamma</math></i> 0.02–0.235 (HBr = $2-96 \times 10^{-9}$ atm)	228	Abbatt and Nowak (1997)	CWFT-MS (a)

### 5 Comments

(a)  $(5-20) \times 10^{10}$  molecule  $\text{cm}^{-3}$  HOCl admitted into the flow tube via the movable injector. Experimental uptake coefficients for HOCl were obtained at 228 K and 69.3 wt%  $\text{H}_2\text{SO}_4$  with various amounts of gas-phase HBr providing an excess concentration in the  $\text{H}_2\text{SO}_4$  film.

### 10 Preferred values

Parameter	Value	T/K
$k_{\text{HOCl}+\text{HBr}}$ ( $\text{M}^{-1} \text{s}^{-1}$ )	$1 \times 10^7$	228
$\alpha_b$	1	228
<i>Reliability</i>		
$\Delta \log(k_{\text{HOCl}+\text{HBr}})$	$\pm 0.3$	

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Comments on preferred value

The preferred values are based on the dataset of Abbatt and Nowak (1997). The expression:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_b} \text{ with } \Gamma_b = \frac{4HRT \sqrt{D_1 k_{\text{HOCl+HBr}}} [\text{HBr}]}{\bar{c}}$$

- 5 was constrained using preferred values of  $H$  and  $D_1$  for HOCl and  $H$  for HBr (IUPAC, 2012, datasheets VI.A4.17 and VI.A4.13) with  $k_{\text{HOCl+HBr}}$  varied to adjust  $\gamma$  to match the experimental data.  $\alpha_b$  was assumed to be unity. The preferred value of  $k_{\text{HOCl+HBr}}$  is larger than reported by Abbatt and Nowak owing to use of lower values of  $H\sqrt{D_1}$ . As noted by Abbatt and Nowak, the value of  $k_{\text{HOCl+HBr}}$  is orders of magnitude larger than
- 10 that reported for low-acidity aqueous solutions (Kumar and Margerum, 1987).

## References

- Abbatt, J. P. D. and Nowak, J. B.: J. Phys. Chem. A, 101, 2131–2137, 1997.  
Kumar, K. and Margerum, D. W.: Inorg. Chem., 26, 2706–2711, 1987.

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

ClONO<sub>2</sub> + HCl (aqueous sulphuric acid aerosol) → HONO<sub>2</sub> + Cl<sub>2</sub>

## Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ] (wt %)	[HCl]/[HNO <sub>3</sub> ] M	Temp./K	Reference	Technique/Comments
γ (ClONO <sub>2</sub> )					
3 × 10 <sup>-3</sup>	65		210	Tolbert et al. (1988)	Kn-MS (a)
2 × 10 <sup>-3</sup>	40		218	Hanson and Ravishankara (1991)	WWFT-CIMS (b)
2 × 10 <sup>-4</sup>	60		215		
<1 × 10 <sup>-4</sup>	65		215		
<2 × 10 <sup>-5</sup>	70		220		
<2 × 10 <sup>-5</sup>	75		230		
0.038	46.6	<10 <sup>-4</sup>	202	Hanson and Ravishankara (1994)	WWFT-CIMS (c)
0.10	46.6	10 <sup>-3</sup>			
0.30	46.6	6 × 10 <sup>-3</sup>			
0.16 ± 0.06	49	5.8 × 10 <sup>-3</sup>	197	Zhang et al. (1994)	WWFT-CIMS (d)
0.10 ± 0.05	51	2.2 × 10 <sup>-3</sup>	198.5		
0.04 ± 0.013	55	3.3 × 10 <sup>-4</sup>	201		
(9.5 ± 3.2) × 10 <sup>-3</sup>	60	2.5 × 10 <sup>5</sup>	205.5		
(1.8 ± 1.6) × 10 <sup>-3</sup>	65	2.5 × 10 <sup>-6</sup>	211		
0.3	43	8.2 × 10 <sup>-4</sup>	194	Elrod et al. (1995)	WWFT-MS (e)
0.083	47	1.4 × 10 <sup>-4</sup>	196		
0.038	51	2.4 × 10 <sup>-5</sup>	198		
0.01	55.2	3 × 10 <sup>-6</sup>	201		
0.11	20.2	4.88 × 10 <sup>-4</sup> (28.35 % HNO <sub>3</sub> )	195		
0.058	44	6.18 × 10 <sup>-5</sup> (6.1 % HNO <sub>3</sub> )	197		
0.021	50.3	1.38 × 10 <sup>-5</sup> (2.26 % HNO <sub>3</sub> )	199		
0.07	45	10 <sup>-8</sup> mbar	203–205	Hanson (1998)	WWFT-CIMS (f)
0.6	45	10 <sup>-6</sup> mbar			
0.03	51	10 <sup>-8</sup> mbar			
0.019	51	10 <sup>-6</sup> mbar			
0.011	55	10 <sup>-8</sup> mbar			
0.043	55	10 <sup>-6</sup> mbar			
0.18 ± 0.10	44	10 <sup>-6</sup> mbar, 4.4 % HNO <sub>3</sub>	205		
0.25 ± 0.15	20.3	10 <sup>-6</sup> mbar, 25.6 % HNO <sub>3</sub>			
0.38 ± 0.1	49 ± 1	1.1 × 10 <sup>-7</sup> mbar	240	Hanson (1998)	AFT-CIMS (g)
0.51 ± 0.14	49 ± 1	2 × 10 <sup>-7</sup> mbar			

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments

- (a) The H<sub>2</sub>O vapor pressure in the flow reactor was  $5.3 \times 10^{-4}$  mbar ( $\pm 30\%$ ). Before ClONO<sub>2</sub> uptake at typical pressures of  $6.4 \times 10^{-3}$  mbar, the 65 % H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O solution was exposed to  $1.7 \times 10^{-3}$  mbar HCl for an hour to result in an estimated mole fraction ratio HCl:H<sub>2</sub>O =  $2 \times 10^{-4}$ . More than 85 % of the HCl is recoverable as Cl<sub>2</sub> from the 65 % H<sub>2</sub>SO<sub>4</sub> solution at 210 K. Resulting  $\gamma$  values are very similar to experiments without added HCl.
- (b) The cold aqueous solution of H<sub>2</sub>SO<sub>4</sub> was flowed down the walls of the vertical flow tube with a liquid film residence time of 20–30 s.  $p(\text{H}_2\text{O})$  was generally around  $1.3 \times 10^{-3}$  mbar, [HCl] for the experiments on the 40% solution was  $2 \times 10^{11}$  for the 60 to 75% solution was in the range 1 to  $3 \times 10^{11}$  molecule cm<sup>-3</sup>. The error limits are reported to be (+100 %, -50 %) except for the 40 and the 60 % solution. The uptake of ClONO<sub>2</sub> in the presence of HCl is not accelerated compared to experiments without HCl. A number of experiments were performed at high HCl concentrations with the result that  $\gamma$  is significantly enhanced at comparable ClONO<sub>2</sub> concentrations: for a 60 % solution  $\gamma$  was  $2 \times 10^{-4}$  (215 K), 0.006 (215 K) and  $\sim 0.01$  (238 K) for gas phase HCl concentrations of  $2 \times 10^{10}$ ,  $7 \times 10^{12}$  and  $\sim 5 \times 10^{12}$  molecule cm<sup>-3</sup>.
- (c) Liquid H<sub>2</sub>SO<sub>4</sub> (46.6 % to 65 %) was applied as a cold liquid to the inner wall of the flow tube. The HCl was either added to the liquid (46.6 and 51 %) or taken up from the gas phase (55.6, 57.5, 58.5, 59.8 and 65 %). No significant temperature dependence of the uptake coefficient in the temperature range 192 to 208 K was observed. The uptake coefficient increased with increasing  $p(\text{HCl})$  and increasing H<sub>2</sub>O activity. Significant uptake was observed at [HCl]  $\rightarrow 0$  which was attributed to the ClONO<sub>2</sub> hydrolysis reaction. It was found from fitting experimental data over a wide range that a surface-mediated reaction between Cl<sup>-</sup> and ClONO<sub>2</sub> scaling

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



with  $p(\text{HCl})$  was required in addition to the competitive bulk reaction to account for the observed uptake rates.

- (d) Total pressure of He 0.67 mbar and linear flow velocity  $890 \text{ cm s}^{-1}$ . The partial pressure of  $\text{H}_2\text{O}$  (5 ppm at 100 mbar) was held constant throughout the range of 195 to 220 K in order to control  $[\text{H}_2\text{SO}_4]$  between 45 and 70% as a function of temperature.  $p(\text{ClONO}_2)$  and  $p(\text{HCl})$  were varied between  $(1-2.5) \times 10^{-7}$  and  $(2.7-27) \times 10^{-7}$  mbar, respectively. The value of  $\gamma$  increases by a factor of ten with increasing  $[\text{HCl}]$  in the given range. Identical values of  $\gamma$  were measured on liquid sulfuric acid containing up to 5% of  $\text{HNO}_3$  and 5 ppm  $\text{H}_2\text{O}$ .
- (e) Reactive uptake experiment using a vertically-mounted ~~wetted wall flow reactor~~ coupled to a differentially-pumped beam-sampling quadrupole MS. The substrates were bulk ternary or quaternary solutions of  $\text{H}_2\text{SO}_4\text{-H}_2\text{O-HCl-HNO}_3$ . The uptake measurements were performed at 203 K, and the composition of the solutions selected to correspond to various atmospheric equilibrium temperatures in the range 194–201 K.  $p(\text{HCl})$  over the solutions was measured at temperatures in the range 208–233 K to determine  $H^*(\text{HCl})$  and  $p(\text{ClONO}_2)$  was of the order of  $1.3 \times 10^{-6}$  mbar. It was established that  $\gamma_{\text{ss}}$  was essentially independent of temperature at a given composition and that it was controlled by the HCl solubility in the condensed phase.
- (f) Uptake experiment in rotating ~~wetted wall flow tube~~ using CIMS detection for  $\text{ClONO}_2$ ,  $\text{Cl}_2$ ,  $\text{HOCl}$  and water vapor. HCl was doped into the 49.5, 51, 53 and 55%  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  solutions from the gas phase. For 45% sulfuric acid solutions HCl was added by mixing pure with sulfuric acid containing known amounts of HCl. The measured uptake is the sum of the reactions of  $\text{ClONO}_2$  with HCl and the hydrolysis with  $\text{H}_2\text{O}$ . The parametrization of the total uptake contains an interfacial term for  $\text{ClONO}_2 + \text{HCl}$  which is important for  $[\text{HCl}] > 10^{-3} \text{ M}$ . For liquid ternary solutions containing  $\text{HNO}_3$  the uptake was significantly lower compared

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

to pure sulfuric acid solutions: for partial pressures corresponding to 5 ppm H<sub>2</sub>O and 5 ppb HNO<sub>3</sub> at 19 km altitude  $\gamma_{ss}$  was typically lower by a factor of two.

- (g) Uptake on sub-micron aerosol particles of H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O (49 ± 1 wt %) doped with HCl, measured at 240 K and 120 mbar total pressure to determine surface accommodation coefficient,  $\alpha_s$ . The surface-area-weighted mean diameter was 160 nm and the number density  $(2-5) \times 10^5 \text{ cm}^{-3}$ . The particles were doped with HCl ( $p(\text{HCl}) = 1.1$  and  $2.0 \times 10^4 \text{ mbar}$ ), resulting in an HCl content of  $9.6 \times 10^{-3}$  and  $1.7 \times 10^{-2} \text{ M}$ , respectively. The weak dependence of  $\gamma$  on  $p(\text{HCl})$  suggests that reaction rate is not quite fast enough to force the measured  $\gamma$  to the value of  $\alpha_s$ . A best estimate for  $\alpha_s$  based on a crude analysis gave  $0.72 \pm 0.1$  but a value of  $\alpha_s = 1$  is consistent with the measurements.

### Preferred values

Parameter	Value	T/K
$\alpha_s$	1.0	298
$b_0/\text{M}^{-2} \text{ cm}^{-2}$	$7.9 \times 10^{-11}$	
$k_{\text{H}_2\text{O}}/\text{s}^{-1}$	$1.95 \times 10^{10} \exp(2800/T)$	190–280
$k_{\text{H}^+}/\text{M}^{-1} \text{ s}^{-1}$	$1.22 \times 10^{12} \exp(6200/T)$	190–280
$\Gamma_s$	$66.12 \exp(-1374/T) \times H_{\text{ClONO}_2} \times M_{\text{HCl}}$	
$H_{\text{ClONO}_2}/\text{Mbar}^{-1}$	$1.6 \times 10^{-6} \exp(4710/T) \exp[(-0.306 + 24.0/T) M_{\text{H}_2\text{SO}_4}]$	200–280
$D_l/\text{cm}^2 \text{ s}^{-1}$	$5 \times 10^{-8} T/\eta$	180–240
<i>Reliability</i>		
$\Delta \log(\alpha)$	±0.1	298
$\Delta \log(\gamma)$	±0.3	200–280

$M_{\text{H}_2\text{SO}_4}$  = molarity H<sub>2</sub>SO<sub>4</sub> (Mol dm<sup>-3</sup>);  $M_{\text{HCl}}$  = molarity HCl (Mol dm<sup>-3</sup>) =  $p(\text{HCl}) \times H_{\text{HCl}}^*$ .

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

## Comments on preferred values

There is a large body of experimental data on the reactive uptake of ClONO<sub>2</sub> in the presence of HCl on H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O surfaces, mainly covering the range of temperature and humidity, and acid composition relevant for the upper troposphere and lower stratosphere. The uptake leads to efficient formation of HNO<sub>3</sub> and Cl<sub>2</sub> products both of which transfer rapidly to the gas phase. At low  $p(\text{HCl})$ , when hydrolysis of ClONO<sub>2</sub> contributes, HOCl is also observed as a product.

The kinetic results are generally consistent between the different studies which used both bulk and dispersed (aerosol) surfaces, although the uptake coefficients measured on bulk surfaces often required large correction for gas phase diffusion effects, because of fast reactive uptake of ClONO<sub>2</sub> (or HCl) at the surface. The reactive uptake coefficient increases strongly with H<sub>2</sub>O content of H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O solutions in the corresponding range 40–70 wt % at constant  $p(\text{HCl})$ , which reflects the increasing solubility of HCl in H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O as relative humidity increases. The uptake coefficient shows only a very weak dependence on temperature.

The uptake coefficient also shows a complex dependence on  $p(\text{HCl})$ , or on  $[\text{HCl}]_l$  in experiments where the solutions were specifically prepared. At intermediate  $[\text{HCl}]_l$  Elrod et al. (1997) found  $\gamma$  proportional to  $[\text{HCl}]_l^{1/2}$ , whilst Zhang et al. (1994) who used higher  $p(\text{HCl})$ , gave close to linear dependence on  $[\text{HCl}]_l$ . Hanson and Ravishankara (1994) covered a wider range of  $p(\text{HCl})$ , and found that for intermediate and high HCl amounts,  $\gamma$  values were proportional to  $p(\text{HCl})^{1/2}$  and  $p(\text{HCl})$  respectively, attaining a  $\gamma$  value of 0.3 at  $1.3 \times 10^{-4}$  mbar HCl at 202 K. At low  $p(\text{HCl})$ ,  $\gamma$  ends to a constant value of  $\sim 0.01$ , which is due to the additional contribution of hydrolysis of ClONO<sub>2</sub> to the overall  $\gamma$ . The observed effects of temperature can also be rationalised by its effect on HCl solubility in H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O solution of varying mole fraction, together with compensating effects on the rate coefficient of the liquid phase reaction. At low temperatures the presence of HNO<sub>3</sub> in the H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O solutions leads to significantly lower  $\gamma$  relative to solutions

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



without HNO<sub>3</sub> (Hanson, 1998). The effect increases with [HNO<sub>3</sub>]/[H<sub>2</sub>SO<sub>4</sub>] and at equal wt% amounts to a 50 % reduction in  $\gamma$  at 205 K.

Hanson and Ravishankara (1994) suggested that two pathways control the rate of heterogenous reaction of ClONO<sub>2</sub> with HCl: a direct surface reaction; and a bulk phase reaction involving reaction of solvated ClONO<sub>2</sub> with Cl<sup>-</sup>, leading to the distinct kinetic dependencies on  $p(\text{HCl})$ . Hanson (1998) has reanalysed their earlier data, and those of Elrod et al. (1997) and Zhang et al. (1994) in the light of improved measurements of HCl solubility, and has formulated a parameterisation for overall  $\gamma$  taking into account the different reactive processes, using a resistance-model formulation, Eq. (1):

$$\frac{1}{\gamma} = \frac{1}{\alpha_s} + \frac{1}{\frac{1}{\Gamma_b} + \frac{1}{\alpha_s \frac{k_{\text{sol}}}{k_{\text{des}}}}} + \Gamma_s \quad (1)$$

Hanson's parameterisation gives overall  $\gamma$  for uptake of chlorine nitrate as a function of mole fraction of H<sub>2</sub>SO<sub>4</sub> and  $p(\text{HCl})$  for stratospheric conditions near 200 K, using an empirical representation of the composition-dependent liquid phase resistance due to chemical reaction.

Shi et al. (2001) have adopted the Hanson model and extended the parameterisation to include temperature dependence, using re-evaluated H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O composition and temperature dependent expressions for HCl solubility,  $H_{\text{HCl}}^*$  (see VI.A4.14), chlorine nitrate solubility,  $H_{\text{ClONO}_2}$ , and resistance term  $\Gamma_b$  for the ClONO<sub>2</sub>+H<sub>2</sub>O reaction (see VI.A4.25). Shi et al. (2001) assumed that the ClONO<sub>2</sub> + HCl reaction is H<sup>+</sup> catalysed and its rate is diffusion controlled, as deduced by Hanson (1998). They also assume that the direct surface reaction of ClONO<sub>2</sub> with HCl depends on bulk-liquid [ClONO<sub>2</sub>] and [HCl] through their respective  $T$ - and composition-dependent  $H$  values.

The preferred value  $\alpha_s$  is based on the measured uptake coefficients on HCl-doped aerosols reported by Hanson (1998), which support a surface accommodation coefficient near unity ( $\alpha_s = 1$ , see comment (g) above). It follows that the term  $\alpha_s k_{\text{sol}}/k_{\text{des}}$  which is equivalent to the surface to bulk resistance term,  $\Gamma_{\text{sb}}$ , must also be close to

32424

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





unity, and Eq. (1) reduces to:

$$\frac{1}{\gamma} = 1 + \frac{1}{(\Gamma_b + \Gamma_s)} \quad (2)$$

The resistance term  $\Gamma_b$  for the bulk reaction probability is given by:  $\Gamma_b = 4RTH(D_1k^1)^{1/2}/c$  where  $D_1$  and  $H$  are respectively the diffusion coefficient and Henry's constant for ClONO<sub>2</sub>.  $k^1$  is the sum of the first order rate coefficients for reaction of ClONO<sub>2</sub> with Cl<sup>-</sup> ( $k_{HCl}^1$ ) and H<sub>2</sub>O ( $k_{H_2O}^1$ ) in the bulk. Combining the resistance terms for these two processes acting in parallel gives:

$$\Gamma_b = \Gamma_{H_2O} \sqrt{1 + \frac{k_{HCl}^1}{k_{H_2O}^1}} \quad (3)$$

Here  $\Gamma_{H_2O} = 4RTH(D_1k_{H_2O}^1)^{1/2}/c$  and the recommended temperature and composition dependence of the first order rate coefficient,  $k_{H_2O}^1$ , and solubility of ClONO<sub>2</sub>,  $H$ , are given in the functional forms described in the data sheet for the ClONO<sub>2</sub> + H<sub>2</sub>O reaction (see VI.A4.25). The overall hydrolysis rate constant is given by the expression:

$$k_{H_2O}^1 (s^{-1}) = (k_{H_2O} + k_{H^+} a_{H^+}) a_w$$

where  $a_w$  = water activity;  $a_{H^+}$  = acid activity (Mol dm<sup>-3</sup>).  $a_w$  is calculated from:

$$a_w = \exp[(-69.775X - 18253.7X^2 + 31072.2X^3 - 25668.8X^4)(1/T - 26.9033/T^2)].$$

The diffusion controlled rate coefficient for the ClONO<sub>2</sub> + HCl reaction is given by the expression:

$$k_{HCl}^1 (s^{-1}) = b_0 D_1 a_{H^+} H_{HCl}^* \rho(HCl) \quad (4)$$

The recommended value for the constant  $b_0$  ( $=7.9 \times 10^{-11} \text{ M}^{-2} \text{ cm}^{-2}$ ), was determined from the experimental measurement of the reacto-diffusive length at 250 K at high [HCl] (Hanson, 1998). The value of  $H_{\text{HCl}}^*$  is given as a function of  $\text{H}_2\text{SO}_4$  mole fraction,  $X$ , by the expression:

$$5 \quad H_{\text{HCl}}^* = (0.094 - 0.61X + 1.2X^2) \exp(-8.68 + (8515 - 10718X^{0.7})/T) \quad (5)$$

as given in data sheet for HCl solubility VI.A4.14. The acid activity is derived from the data from the thermodynamic model of Carslaw et al. (1997). For the expression for  $a_{\text{H}^+}$  in the table the Carslaw model was extended to include the acidity of pure water, to provide values of  $a_{\text{H}^+}$  as a function of acid wt % (wt) in units of  $\text{M dm}^{-3}$  extending to dilute solution (wt  $\sim 0\%$ ):

$$10 \quad a_{\text{H}^+} = \exp \left[ 60.51 - 0.095 \text{ wt} + 0.0077 \text{ wt}^2 - 1.61 \times 10^{-5} \text{ wt}^3 - (1.76 + 2.52 \times 10^{-4} \text{ wt}^2) T^{0.5} \right. \\ \left. + (-805.89 + 2.53.05 \text{ wt}^{0.076})/T^{0.5} \right]$$

Diffusion coefficients,  $D_1$ , are calculated using the expression:  $D_1 = CT/\eta$ , with  $C = 5 \times 10^{-8} \text{ cm}^2 \text{ cPK}^{-1} \text{ s}^{-1}$  (taken from Klassen et al., 1998). Viscosity data for  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  obtained by Williams and Long (1995) was reparameterized to give a more general formulation for  $D_1$  covering the temperature range 200–300 K and 0–80 wt %  $\text{H}_2\text{SO}_4$ :

$$20 \quad \eta = AT^{-1.43} \exp(448/T - T_0)$$

$$A = 169.5 + 5.18 \text{ wt} - 0.0825 \text{ wt}^2 + 3.27 \times 10^{-3} \text{ wt}^3$$

$$T_0 = 144.11 - 0.166 \text{ wt} + 0.015 \text{ wt}^2 - 2.18 \times 10^{-4} \text{ wt}^3$$

The resistance term  $\Gamma_s$  for the surface reaction probability was parameterised by Shi et al. (2001) as follows:

$$25 \quad \Gamma_s = 66.12 \exp(-1374/T) H_{\text{ClHONO}_2} H_{\text{HCl}}^* \rho_{\text{HCl}} \quad (6)$$

32426

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

which assumes a dependence of surface reaction rate on liquid phase concentrations of reactants, which is determined by their solubilities. An Arrhenius temperature dependence is assumed and the recommended constant term was evaluated by fitting to the ClONO<sub>2</sub> uptake rate data at high [HCl]. This expression gave a better representation of the temperature dependence of the surface reaction component than that used by Hanson (1998).

The model gives a good fit to the data over the range of conditions employed and will give reasonably accurate values of  $\gamma$  (ClONO<sub>2</sub>) for UTLS conditions. Overall the deviation reported by Shi et al. (2001) from experimental data used for the fit was 40 %, compared with 46 % for the parameters reported by Hanson (1998). These form the basis of the recommended uncertainty in  $\gamma$  over the temperature range 200–280 K. The temperature dependence of the  $\gamma_r$  is in fact rather small due to compensating effects of the temperature dependence of  $D_1$  and  $H_{\text{HCl}}^*$ . For detailed discussion of the model/measurement comparison the paper by Shi et al. (2001) should be consulted.

## References

- Donaldson, D. J., Ravishankara, A. R., and Hanson, D. R.: J. Phys. Chem. A, 101, 4717–4725, 1997.
- Carslaw, K. S., Clegg, S. L., and Brimblecombe, P.: J. Phys. Chem., 99, 11557–11574, 1995.
- Elrod, M. J., Koch, R. E., Kim, J. E., and Molina, M. J.: Faraday Discuss., 100, 269–278, 1995.
- Hanson, D. R. and Ravishankara, A. R.: J. Geophys. Res., 96, 17307–17314, 1991.
- Hanson, D. R. and Ravishankara, A. R.: J. Phys. Chem., 98, 5728–5735, 1994.
- Hanson, D. R.: J. Phys. Chem. A, 102, 4794–4807, 1998.
- Huthwelker, T., Peter, Th., Luo, B. P., Clegg, S. L., Carslaw, K., and Brimblecombe, P.: J. Atm. Chem., 21, 81–95, 1995.
- Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

Tolbert, M. A., Rossi, M. J., and Golden, D. M.: Geophys. Res. Lett., 15, 847–850, 1988.

5 Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

Zhang, R., Leu, M.-T., and Keyser, L. F.: J. Phys. Chem., 98, 13563–13574, 1994.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

$\text{ClONO}_2 + \text{H}_2\text{O} \text{ (aqueous sulphuric acid aerosol)} \rightarrow \text{HONO}_2 + \text{HOCl}$ 

## Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ] wt %	Temp./K	Reference	Technique/Comments
$\gamma, \gamma_{\text{ss}}, \gamma_0 \text{ (ClONO}_2\text{)}$				
$(3.2 \pm 0.8) \times 10^{-3}$	95.6		Rossi et al. (1987)	Kn-MS (a)
$2.6 \times 10^{-3}$	65	210	Tolbert et al. (1988)	Kn-MS (b)
$3 \times 10^{-4}$	75	230		
$0.064 \pm 0.010$	40	218	Hanson and Ravishankara (1991a)	WWFT-CIMS (c)
$(3.1 \pm 0.5) \times 10^{-3}$	60	215		
$(1.2 \pm 0.2) \times 10^{-3}$	65	215		
$(3.9 \pm 0.7) \times 10^{-4}$	70	220		
$(1.9 \pm 0.3) \times 10^{-4}$	75	230		
$(8.1 \pm 1.3) \times 10^{-2}$	40	220 ± 5	Williams et al. (1994)	Kn-MS (d)
$(4.1 \pm 0.7) \times 10^{-3}$	57.5			
$(2.1 \pm 0.35) \times 10^{-4}$	75			
$(3.8 \pm 1.2) \times 10^{-2}$	46.6	202	Hanson and Ravishankara (1994)	WWFT-CIMS (e)
$(6.1 \pm 2.0) \times 10^{-3}$	57.5			
$(9.3 \pm 3.0) \times 10^{-4}$	65			
$(2.0 \pm 0.6) \times 10^{-2}$	50	197.5	Zhang et al. (1994)	WWFT-CIMS (f)
$(2.5 \pm 0.7) \times 10^{-3}$	60	205		
$(7.5 \pm 2.5) \times 10^{-4}$	65	211		
$(3.0 \pm 1.0) \times 10^{-4}$	70	218		
$2 \times 10^{-3}$	65	250	Hanson and Lovejoy (1995)	AFT-CIMS (g)
0.11	53	200	Zhang et al. (1995)	WWFT-CIMS (h)
0.034	29, 16.4% HNO <sub>3</sub>	195		
0.021	5, 41% HNO <sub>3</sub>	200		
0.045	40, 10% HNO <sub>3</sub>	200		
0.10	5, 41% HNO <sub>3</sub>	220		
$(3.7 \pm 0.6) \times 10^{-3}$	59	241	Robinson et al. (1997)	DT-TDLAS (i)

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Parameter	[H <sub>2</sub> SO <sub>4</sub> ] wt %	Temp./K	Reference	Technique/Comments
$(5.60 \pm 0.84) \times 10^{-2}$	39	229		
$(2.62 \pm 0.39) \times 10^{-2}$	39	259		
$0.071 \pm 0.025$	36	245	Ball et al. (1998)	AFT-CL (j)
$0.038 \pm 0.009$	43	248		
$0.0094 \pm 0.003$	54	253		
$0.113 \pm 0.023$	36.5	203	Hanson (1998)	WWFT CIMS (k)
$0.086 \pm 0.017$	36.5	230		
$0.053 \pm 0.011$	45	203		
$0.038 \pm 0.008$	45	230		
$0.011 \pm 0.002$	55	203		
$(1.6 \pm 0.3) \times 10^{-4}$	75	270		
$(1.4 \pm 0.3) \times 10^{-4}$	75	250		
$(1.1 \pm 0.2) \times 10^{-4}$	75	230		
$(6.4 \pm 1.3) \times 10^{-5}$	75	217		
$(2.5 \pm 1.0) \times 10^{-5}$	75	200		
$0.019 \pm 0.004$	44, 4.6% HNO <sub>3</sub>	203		
$0.042 \pm 0.016$	20, 26% HNO <sub>3</sub>	205		

## Comments

- (a) Simultaneous flows of H<sub>2</sub>O ( $1.5 \times 10^{13}$  molecule cm<sup>-3</sup>) and ClONO<sub>2</sub> ( $9.4 \times 10^{11}$  –  $4.5 \times 10^{13}$  molecule cm<sup>-3</sup>) were exposed to bulk 95.6 % H<sub>2</sub>SO<sub>4</sub> at ambient temperature;  $\gamma$  is independent of [ClONO<sub>2</sub>] and declines with exposure time.
- (b) The H<sub>2</sub>O vapor pressure in the flow reactor was  $5.3 \times 10^{-4}$  mbar ( $\pm 30\%$ ). Uptake of ClONO<sub>2</sub> at typical pressures of  $6.4 \times 10^{-3}$  mbar on 65 % H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O.
- (c) Flowing aqueous H<sub>2</sub>SO<sub>4</sub> film.  $p(\text{H}_2\text{O})$  was  $\sim 1.3 \times 10^{-3}$  mbar and [ClONO<sub>2</sub>] was approx.  $10^{10}$ – $10^{11}$  molecule cm<sup>-3</sup>. The temperature dependence of  $\gamma$  as measured for the 60% and 70% H<sub>2</sub>SO<sub>4</sub> solution: none was found within the reported error limits.

(d)  $[\text{ClONO}_2]$  was approx.  $10^{12}$ – $10^{13}$  molecule  $\text{cm}^{-3}$ . The following analytical expression was obtained for the uptake/reaction probability ( $\gamma$ ) of  $\text{ClONO}_2$ :  $\log \gamma = 1.87 - 0.074 \times [\text{H}_2\text{SO}_4]$  (wt %) ( $\pm 15\%$ ) valid for the range 40 to 80%  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  mixtures at 220 K.

(e) Quiescent liquid  $\text{H}_2\text{SO}_4$  surface with CIMS detection.  $\text{H}_2\text{SO}_4$  (46.6% to 65%) was applied as a cold liquid to the inner wall of the flow tube.

(f) Walls coated with a 70%  $\text{H}_2\text{SO}_4$  solution of approximately 0.1 mm thickness. The partial pressure of  $\text{H}_2\text{O}$  was held constant (5 ppm at 100 mbar pressure) throughout the temperature range 195–220 K in order to control the acid concentration with temperature leading to  $[\text{H}_2\text{SO}_4]$  in the range 45 to 70 % by weight.

(g) High pressure (0.3 to 0.8 bar) flow tube using slow-flow conditions and sub-micron  $\text{H}_2\text{SO}_4$  aerosol generated by homogeneous nucleation from the reaction of  $\text{SO}_3 + \text{H}_2\text{O}$ . Particle size 60–250 nm diameter. The uptake coefficient for the reaction of  $\text{ClONO}_2$  with 60 wt % sulfuric acid aerosol increases monotonically with particle size at 250 K. The reacto-diffusive length ( $l$ , the effective  $l_{\text{eff}}$  depth over which reaction occurs) derived from these experiments is  $0.037 \pm 0.007 \mu\text{m}$ .

(h) Two methods of film preparation of defined ternary composition  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}/\text{HNO}_3$  at low temperatures resulted in identical values of  $\gamma$  which does not change in the presence of  $\text{HNO}_3$ , even at a limiting composition of 15%  $\text{HNO}_3/30\%$   $\text{H}_2\text{SO}_4$  at 195 K.

(i) Fast train of 200  $\mu\text{m}$   $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  droplets traversing a flow tube with TDLAS detection. Pressures = 13.3 mbar and droplet-gas interaction times of 2 to 20 ms. The temperature of the droplets was inferred from the water partial pressure measured by TDL absorption. A negative temperature dependence of  $\gamma_{\text{ss}}$  was observed for  $T \geq 230$  K and  $\gamma_{\text{ss}}$  slightly increased with increasing concentration of  $\text{H}_2\text{SO}_4$  measured at 39, 54 and 69%. A model involving neutral and acid catalysed mechanism for  $\text{ClONO}_2$  hydrolysis was used to extract values of  $\alpha_{\text{p}}$  and liquid

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



phase rate constants by fitting data from several laboratories over a range of temperature and composition. The temperature dependence of  $\alpha_b$  was given in the form  $\alpha_b/(1 - \alpha_b) = \exp(-17.89 + 4515/T)$ .

- (j) High pressure (0.3 to 0.8 atm) flow tube using slow flow conditions and sub micron  $\text{H}_2\text{SO}_4$  aerosol generated by homogeneous nucleation from the reaction of  $\text{SO}_3 + \text{H}_2\text{O}$ . Particles of a narrow size range were selected with a DMA for each  $\text{H}_2\text{SO}_4$  concentration used: mean size 95 nm, 104 nm and 63 nm radius for 36, 43 and 54 wt % respectively.  $\text{ClONO}_2$  (30 to 200 ppb) detected by titrating with NO in a heated quartz tube, using TDLAS to monitor NO and also  $\text{H}_2\text{O}$  vapor. The average value of  $\gamma$  was used to derive a value of  $l = 26$  nm for the reacto-diffusive length at  $\sim 43\%$   $\text{H}_2\text{SO}_4$
- (k) Uptake of  $\text{ClONO}_2$  measured in two flow reactors. *a*: Rotating wetted wall flow tube with sulfuric acid wall film (0.2 mm thickness). The pressure was 0.5 mbar He in the coated wall flow tube and 240 mbar of  $\text{N}_2$  in the aerosol flow tube. The range of  $[\text{H}_2\text{SO}_4]$  investigated was 36.5 to 55%  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  at 203 K, 36.5 to 45.0% at 230 K and 75% in the temperature range 200 to 270 K. *b*: sub-micron aerosol (particle size 0.1  $\mu\text{m}$ )  $49 \pm 1\%$   $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$ . CIMS detection for  $\text{ClONO}_2$ , HOCl and water vapor. Initial concentration was  $5 \times 10^{11}$  molecule  $\text{cm}^{-3}$  in both studies. The results for small sulfuric acid particles resulted in a lower limit of  $\alpha_s > 0.5$ .  $\text{ClONO}_2$  reaction probabilities were also measured on  $\text{H}_2\text{SO}_4$  solutions containing significant amounts of  $\text{HNO}_3$ . In contrast to previous reports,  $\text{HNO}_3$  was found to have a significant reducing effect on  $\gamma$  for  $\text{ClONO}_2$ .

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_b$	1	298
$k_{\text{H}_2\text{O}}/\text{s}^{-1}$	$1.95 \times 10^{10} \exp(2800/T)$	190–280
$k_{\text{H}^+}/\text{M}^{-1} \text{s}^{-1}$	$1.22 \times 10^{12} \exp(6200/T)$	190–280
$H_{\text{ClONO}_2}/\text{M atm}^{-1}$	$1.6 \times 10^{-6} \exp(4710/T) \times \exp[(-0.306 + 24.0/T) \times M_{\text{H}_2\text{SO}_4}]$	200–280
$D_l/\text{cm}^2 \text{s}^{-1}$	$5 \times 10^{-8} T/\eta$	180–240
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.1$	298
$\Delta \log(\gamma)$	$\pm 0.2$	200–280

$M_{\text{H}_2\text{SO}_4}$  = molarity  $\text{H}_2\text{SO}_4$  ( $\text{M}$ ); wt = weight %  $\text{H}_2\text{SO}_4$ ; X = mole fraction  $\text{H}_2\text{SO}_4$

## Comments on preferred values

- 5 There is a large body of experimental data on the uptake of  $\text{ClONO}_2$  on  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  surfaces covering relevant temperatures, humidity and sulphuric acid aerosol composition for the atmosphere between the surface and the lower stratosphere. The results are generally consistent between the different studies, which used both bulk and dispersed (aerosol) surfaces. The uptake leads to hydrolysis of  $\text{ClONO}_2$  and formation of  $\text{HOCl}$  and  $\text{HNO}_3$ , which both transfer to the gas phase. The presence of  $\text{HNO}_3$  in the  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  solutions leads to a reduction in the uptake rate (Hanson, 1998; Zhang et al., 1995). Based on uptake measurements of Hanson (1998) on  $\text{HNO}_3/\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  solutions corresponding to conditions in the polar lower stratosphere ( $\sim 5$  ppb  $\text{HNO}_3$  in the gas phase), the  $\gamma$  values would be approximately a factor of 2 lower than for  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  solutions.

Uptake rates show strong dependence on  $\text{H}_2\text{O}$  content of the sulphuric acid solution/aerosol,  $\gamma_{\text{ClONO}_2}$  decreasing in the range 20–70 wt %  $\text{H}_2\text{SO}_4$ .  $\text{H}_2\text{O}$  content depends on both temperature and relative humidity and consequently there is a com-

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



plex variation of  $\gamma$  for atmospheric conditions: generally low temperature and high RH favour rapid reactive uptake with  $\gamma$  values  $\sim 0.1$  in 40%  $\text{H}_2\text{SO}_4$  falling off to  $\sim 10^{-4}$  in 75%  $\text{H}_2\text{SO}_4$ . At low temperature the effect of composition dominates leading to weak negative  $T$  dependence of  $\gamma$ .

5 These characteristics indicate that uptake is determined by chemical reaction in the liquid droplet and can be interpreted in terms of the resistance model. Thus the overall uptake coefficient is given by Eq. (1)

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_b} \quad (1)$$

$$10 \quad \Gamma_b = 4RTH(D_1 k^1)^{1/2} / \bar{c} \quad (2)$$

where  $H$  is the Henry's ~~solubility~~ constant of  $\text{ClONO}_2$ ,  $D_1$  is the diffusion coefficient of  $\text{ClONO}_2$ , and  $k^1$  the first order rate constant for hydrolysis of  $\text{ClONO}_2$  in  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  solutions.

15 Robinson et al. (1977) presented a phenomenological model addressing solubility, diffusion and chemical reactivity of  $\text{ClONO}_2$ , which accounts for the observed dependence of the uptake coefficients on concentration and temperature. Two hydrolysis pathways are proposed, a direct reaction with  $\text{H}_2\text{O}$  and an acid-catalysed reaction involving  $\text{H}^+$  ions to promote the dissociation:  $\text{ClONO}_2 + \text{H}^+ \rightarrow \text{HOCl} + \text{NO}_2^+$ . The hydrolysis rate coefficient representing the two pathways was given in the form:  
20  $k^1 = k_{\text{H}_2\text{O}} a_w + k_{\text{H}^+} a_{\text{H}^+}$ ; here  $a_w$  and  $a_{\text{H}^+}$  are the activities of  $\text{H}_2\text{O}$  and  $\text{H}^+$  in solution respectively. They extracted temperature and composition dependencies for the individual parameters,  $\alpha_b$  and  $k^1$  from a fit of their experimental  $\gamma$  values using this model, together with a parameterisation of  $H$  based on solubility data given by Huthwelker et al. (1995) for  $\text{HOCl}$  as a proxy for  $\text{ClONO}_2$ .  $D_1$  was parameterized using the expres-  
25 sion:  $D_1 = cT/\eta$ , with  $c$  evaluated from viscosity data for  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  solutions obtained by Williams and Long (1995).

---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**M. Ammann et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

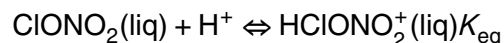
Printer-friendly Version

Interactive Discussion



Hanson (1998) has used a similar model, which gives  $\gamma$  values as a function of mole fraction of  $\text{H}_2\text{SO}_4$ ,  $X$ , which is defined for a given temperature and water activity (RH). Only a single direct hydrolysis reaction was used for  $k^1$ . Composition dependent values of  $\Gamma_b$  were obtained from a fit of Eq. (1) to experimental  $X$ -dependent  $\gamma$  values for 200–205 K, assuming  $\alpha_b = 1.0$ :  $\Gamma_b = \exp(-0.392 - 13.13X - 50.914X^2)$  (Eq. 3). This simple expression can be used to calculate  $\gamma$  from Eq. (1) with reasonable accuracy for specified values of  $\text{H}_2\text{SO}_4$  wt% in the range 40–65% in and for temperatures near 200 K; at  $\geq 70$  wt% the observed  $\gamma$  values are seriously underpredicted. Thus this formula only applies for only for a limited range of lower stratospheric conditions. Hanson also derived individual parameterisations for variation of  $H$ ,  $D_1$ ,  $\text{H}_2\text{SO}_4$  acidity ( $a_{\text{H}^+}$ ,  $a_{\text{H}_2\text{O}}$ ) as a function of mole fraction of  $\text{H}_2\text{SO}_4$  at temperatures near 200 K.

Shi et al. (2001) have reported further analysis using these models for representation of uptake coefficients for  $\text{ClONO}_2$  hydrolysis and reaction with  $\text{HCl}$  in  $\text{H}_2\text{SO}_4$  solutions. They adopt a general mechanism involving both a direct and an acid-catalyzed channel for  $\text{ClONO}_2$  reaction with  $\text{H}_2\text{O}$  and  $\text{HCl}$ , with the reactive species controlled by following equilibrium:



The hydrolysis rate constant is given by:

$$k^1 = k_{\text{H}_2\text{O}}a_w + k_{\text{H}^+}a_{\text{H}^+}a_w, \quad (3)$$

where the acid catalysed rate constant,  $k_{\text{H}^+}$ , includes the term  $K_{\text{eq}}$ , arising from the assumption that protonated species are in equilibrium. The values of  $D_1$ ,  $H$ ,  $\text{H}_2\text{SO}_4$  acidity and water activity ( $a_{\text{H}^+}$ ,  $a_w$ ) were parameterized independently of the experimental uptake measurements.

The thermodynamic model of Carslaw et al. (1995) was used to derive a parameterization of  $a_w$  in terms of the mole fraction ( $X$ ) of  $\text{H}_2\text{SO}_4$ , ( $X = \text{wt}/(\text{wt} + (100 - \text{wt})98/18)$ ):

$$a_w = \exp[(-69.775X - 18253.7X^2 + 31072.2X^3 - 25668.8X^4)(1/T - 26.9033/T^2)].$$

Note however that for atmospheric modelling  $a_w$  is usually calculated from the local  $\text{H}_2\text{O}$  mixing ratio and temperature. For the acid activity the Carslaw model was extended to include the acidity of pure water, to provide values of  $a_{\text{H}^+}$  as a function of acid  $wt$  in units of  $\text{M dm}^{-3}$  extending to dilute solution ( $wt \sim 0\%$ ):

$$5 \quad a_{\text{H}^+} = \exp \left[ 60.51 - 0.095 wt + 0.0077 wt^2 - 1.61 \times 10^{-5} wt^3 - (1.76 + 2.52 \times 10^{-4} wt^2) T^{0.5} + (-805.89 + 253.05 wt^{0.076}) / T^{0.5} \right].$$

Diffusion coefficients,  $D_1$  were estimated using the expression:  $D_1 = cT/\eta$ , with  $c = 5 \times 10^{-8} \text{ cm}^2 \text{ cPK}^{-1} \text{ s}^{-1}$  (taken from Klassen et al., 1998). Viscosity data for  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  obtained by Williams and Long (1995) was re-parameterized to give a more general formulation covering the temperature range 200–300 K and 0–80  $wt$   $\text{H}_2\text{SO}_4$ :

$$\eta = AT^{-1.43} \exp(448/T - T_0)$$

$$15 \quad A = 169.5 + 5.18 wt - 0.0825 wt^2 + 3.27 \times 10^{-3} wt^3$$

$$T_0 = 144.11 - 0.166 wt + 0.015 wt^2 - 2.18 \times 10^{-4} wt^3$$

The variations of  $H$  and  $k_1$  with acid strength are constrained by the experimental values of  $H(k_1)^{1/2}$ , which is determined from experimental uptake coefficients using Eq. (2) for defined  $D_1$ . Following Robinson et al. (1997), the solubility of  $\text{HOCl}$  was used as proxy for  $\text{ClONO}_2$ , and is given in the functional form:

$$20 \quad H = H_0 \exp(B/T) \exp(-S_t M_{\text{H}_2\text{SO}_4}) \quad (4)$$

$$S_t = c_s + d/T \quad (5)$$

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**


---

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[⏪](#)
[⏩](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


$H(\text{HOCl})$  was re-parameterized taking into account the higher temperature and wt%  $\text{H}_2\text{SO}_4$  data of Donaldson et al. (1997) which extended and improved that of Huthwelker et al. (1995). The Setchenow coefficient,  $S_t$ , was assumed to depend on the *molarity* (rather than molality) of the  $\text{H}_2\text{SO}_4$  solution.  $H$  was derived from indirect analysis of aerosol kinetics measurements of Hanson and Lovejoy (1995), which provided uptake coefficient,  $\gamma$ , and reacto-diffusive length,  $l$ , for hydrolysis of  $\text{ClONO}_2$  on 60%  $\text{H}_2\text{SO}_4$  at 250 K. This provides a value for  $H_{\text{ClONO}_2}$  at a single temperature and to obtain the temperature dependence of  $H$  (and also of  $k^1$ ), use was made of the kinetic data for the  $\text{ClONO}_2 + \text{HCl}$  reaction, for which the temperature dependence includes that of  $H_{\text{ClONO}_2}$  and  $k_{\text{HCl}}^1$ . Assuming the latter is diffusion limited (following Hanson, 1998) its temperature dependence is fixed, allowing extraction of the following parameters for calculation of the solubility of  $\text{ClONO}_2$ :  $H^0 = 1.6 \times 10^{-6}$ ,  $B = 4710$ ,  $c_s = 0.306$ , and  $d = 24.0$  in Eqs. (4) and (5).

The rate coefficients for the direct and acid catalysed hydrolysis were expressed in Arrhenius form:

$$k_{\text{H}_2\text{O}} = A_{\text{H}_2\text{O}} \cdot \exp(-E_{\text{H}_2\text{O}}/T)$$

$$k_{\text{H}^+} = A_{\text{H}^+} \cdot \exp(-E_{\text{H}^+}/T)$$

The values of  $k^1$  for experimental conditions were calculated from Eq. (3), using the parameterisations for  $a_{\text{H}^+}$ ,  $a_{\text{H}_2\text{O}}$  given above. The  $A$  and  $E$  values for the temperature dependence of  $k_{\text{H}}$  and  $k_{\text{H}_2\text{O}}$  were determined from a global fit to experimental uptake coefficients for both the  $\text{ClONO}_2$  hydrolysis and  $\text{HCl}$  reaction data, assuming the reaction  $\text{ClONO}_2 + \text{HCl}$  is diffusion limited, in the same procedure as the determination of  $H_{\text{ClONO}_2}$  over a range of temperature. The values obtained from fitting of the data from (Hanson, 1998; Ball et al., 1998; Robinson et al., 1997; Hanson and Lovejoy, 1995; Zhang et al., 1994; Manion et al., 1994; Hanson and Ravishankara, 1991, 1994) were:

$$A_{\text{H}_2\text{O}} = 1.95 \times 10^{10} \text{ s}^{-1} E_{\text{H}_2\text{O}} = 2800 \text{ K}$$

$$A_{\text{H}^+} = 1.22 \times 10^{12} \text{ M}^{-1} \text{ s}^{-1} E_{\text{H}^+} = 6200 \text{ K}$$

All three models give a good description of the experimental  $\gamma$  values for the range of conditions used to derive the parameters. However the parameterisations for  $H$ ,  $D_1$ ,  $k_1$  and the other data on which they are based, differ considerably. The simpler parameterisation for  $\Gamma_b$  (Eq. 3) obtained by Hanson gives a reasonable representation of observed  $\gamma$  values as a function  $[\text{H}_2\text{SO}_4]$ , but only up to 65 wt%, and at temperatures around 200 K. The more complex parameterisation of Shi et al. (2001) gives an excellent description of the experimentally observed data over a range of temperature 190–260 K at  $[\text{H}_2\text{SO}_4]$  in the range 35–75 wt%. The parameters of Shi et al., given in the table above, are recommended as the preferred values for hydrolysis of  $\text{ClONO}_2$  covering most atmospheric conditions. The standard deviation of uptake coefficients calculated using the Shi et al. model with respect to all experimental data used in the global fit was 32% ( $1\sigma$ ). This forms the basis of the recommended uncertainty in  $\gamma$  over the temperature range 200–280 K.

## References

- Ball, S. M., Fried, A., Henry, B. E., and Mozurkewich, M.: *Geophys. Res. Lett.*, 25, 3339–3342, 1998.
- Carlsaw, K. S., Clegg, S. L., and Brimblecombe, P.: *J. Phys. Chem.*, 99, 11557–11574, 1995.
- Donaldson, D. J., Ravishankara, A. R., and Hanson, D. R.: *J. Phys. Chem. A*, 101, 4717–4725, 1997.
- Golden, D. M., Manion, J. A., Reihls, C. M., and Tolbert, M. A.: in: *CHEMRAWN VII: Chemistry of the Atmosphere: The Impact of Global Change*, edited by: Calvert, J. G., Oxford University Press, 1992.
- Hanson, D. R. and Ravishankara, A. R.: *J. Geophys. Res.*, 96, 17307–17314, 1991a.
- Hanson, D. R. and Ravishankara, A. R.: *J. Phys. Chem.*, 98, 5728–5734, 1994.
- Hanson, D. R. and Lovejoy, E. R.: *Science*, 267, 1326–1328, 1995.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**M. Ammann et al.

---

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- Hanson, D. R.: J. Phys. Chem. A, 102, 4794–4807, 1998.
- Huthwelker, T., Peter, Th., Luo, B. P., Clegg, S. L., Carslaw, K., and Brimblecombe, P.: J. Atm. Chem., 21, 81–95, 1995.
- Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.
- 5 Manion, J. A., Fittschen, C. M., Golden, D. M., Williams, L. R., and Tolbert, M. A.: Isr. J. Chem., 34, 355–363, 1994.
- Robinson, G. N., Worsnop, D. R., Kolb, C. E., and Davidovits, P.: J. Geophys. Res., 102, 3583–3601, 1997.
- Rossi, M. J., Malhotra, R., and Golden, D. M.: Geophys. Res. Lett., 14, 127–130, 1987.
- 10 Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.
- Tolbert, M. A., Rossi, M. J., Malhotra, R., and Golden, D. M.: Science, 238, 1258–1260, 1987.
- Williams, L. R., Manion, J. A., Golden, D. M., and Tolbert, M. A.: J. Appl. Meteorol., 33, 785–790, 1994.
- 15 Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.
- Zhang, R., Leu, M.-T., and Keyser, L. F.: Geophys. Res. Lett., 22, 1493–1496, 1995.
- Zhang, R., Leu, M.-T., and Keyser, L. F.: J. Phys. Chem., 98, 13563–13574, 1994.

HOBr + HCl (aqueous sulphuric acid aerosol) → BrCl + H<sub>2</sub>O

## Experimental data

Parameter	H <sub>2</sub> SO <sub>4</sub> wt %	[HCl]/ $\rho$ (HOBr)	Temp./K	Reference	Technique/Comments
$\gamma$ (HOBr)					
0.023	69.8	HCl = $0.22 \times 10^{12} \text{ cm}^{-3}$	228	Abbatt (1995)	CWFT-MS (a)
0.060	69.8	HCl = $0.96 \times 10^{12} \text{ cm}^{-3}$			
0.12	69.8	HCl = $2.5 \times 10^{12} \text{ cm}^{-3}$			
0.20	69.8	HCl = $6.0 \times 10^{12} \text{ cm}^{-3}$			
0.2	60	$\rho$ (HCl) = $2.7 \times 10^{-7} \text{ mbar}$	210	Hanson and Ravishankara (1995)	CWFT-MS (b)
$\gamma$ (HCl)					
$0.11 \pm 0.04$	59.7	$\rho$ (HOBr) = $8 \times 10^{-7} \text{ mbar}$	228	Waschewsky and Abbatt (1999)	CWFT-MS (c)
$0.032 \pm 0.008$	65.6	$\rho$ (HOBr) = $1 \times 10^{-6} \text{ mbar}$	228		
$0.060 \pm 0.015$	70.1	$\rho$ (HOBr) = $3.4 \times 10^{-5} \text{ mbar}$	228		
0.19	58	$\rho$ (HOBr) = $1 \times 10^{-5} \text{ mbar}$	250	Hanson (2003)	CWFT-CIMS (d)
0.07	62	$\rho$ (HOBr) = $1 \times 10^{-5} \text{ mbar}$			
0.011	65	$\rho$ (HOBr) = $1 \times 10^{-5} \text{ mbar}$			
0.002	69.5	$\rho$ (HOBr) = $1 \times 10^{-5} \text{ mbar}$			

## 5 Comments

(a) HOBr ( $<1 \times 10^{12} \text{ molecule cm}^{-3}$ ) was generated by the reaction sequence:  $\text{H} + \text{NO}_2 \rightarrow \text{OH} + \text{NO}$  and  $\text{OH} + \text{Br}_2 \rightarrow \text{HOBr}$ , ionised by electron impact and detected as  $\text{HOBr}^+$ . The carrier gas flow was humidified to maintain the H<sub>2</sub>SO<sub>4</sub> concentration (69.8 wt %). Excess HCl (gas phase concentration  $(0.2\text{--}7.0) \times 10^{12} \text{ molecule cm}^{-3}$ ) added simultaneously with HOBr ( $>1 \times 10^{12} \text{ molecule cm}^{-3}$ ). HOBr loss was first order, and no systematic dependence of  $\gamma$  on [HOBr] was observed. Using measured values of  $\text{HD}_1^{1/2}$  for HOBr

10

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





---

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**


---

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)




[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


( $20 \pm 10 \text{ M atm}^{-1} \text{ cm s}^{-1/2}$ ) and for  $H^*$  for HCl the cited  $\gamma$  value constrained the bimolecular rate constant for the HOBr + HCl reaction to be  $1.4 \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ .

(b) HOBr was generated from the hydrolysis reaction of  $\text{BrONO}_2$  on 60 %  $\text{H}_2\text{SO}_4$  and detected using  $\text{SF}_6^-$  chemi-ions. BrCl was observed as a gas phase product indicating a fast effective bimolecular reaction on  $\text{H}_2\text{SO}_4$ . A bimolecular rate constant for the HOBr + HCl reaction of  $(0.3\text{--}1) \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$  was derived from the uptake data.

(c) HOBr ( $\approx 5 \times 10^{-10} \text{ bar}$ , measured by UV absorption at 254 nm) was generated ex-situ by passing a humidified flow of  $\text{Br}_2$  in He over  $\text{HgO}$ , ionised by electron impact and detected as  $\text{HOBr}^+$ . Decay of HCl and formation of BrCl was measured in presence of excess HOBr. Uptake coefficients determined from both kinetic curves were within 10 %.  $\gamma$  increased with  $\rho(\text{HOBr})^{1/2}$  at 65–70 %. Bimolecular rate constants for the HOBr + HCl reaction derived from the uptake data using  $H^*(\text{HOBr})$  determined in the same study were in the range of  $(1.6\text{--}270) \times 10^5 \text{ M}^{-1} \text{ s}^{-1}$ , increasing strongly with wt %  $\text{H}_2\text{SO}_4$  and less so with temperature (213–228 K). The  $[\text{H}_2\text{SO}_4]$  dependence was attributed to an acid catalysed mechanism and the temperature dependence to diffusion limitation.

(d) Rotating CWFT with stirring of the  $\text{H}_2\text{SO}_4$  film (58–70 wt %). HOBr (typically  $10^{-10} \text{ atm}$ ) was generated by reacting  $\text{BrONO}_2$  with water and detected using  $\text{SF}_6^-$  chemi-ions. First order loss of HCl was measured in the presence of excess HOBr  $(0.2\text{--}2.0) \times 10^{-10} \text{ bar}$ .  $\Gamma_b$  was extracted assuming  $\alpha_b = 1.0$ , and was  $\propto \rho(\text{HOBr})^{1/2}$ . Using  $H_{\text{HCl}}$  from Carslaw et al. (1995) and  $H_{\text{HOBr}}$  determined in the same study (see reaction VI.A4.16) values of  $k^{\text{II}}$  were derived. The values ( $k^{\text{II}} = (2.5\text{--}6.0) \times 10^7 \text{ M}^{-1} \text{ s}^{-1}$  over the range 58–69.5 %  $\text{H}_2\text{SO}_4$  at 250 K) showed a lower dependence on acid strength compared to results of Waschewsky and Abbatt (1999), obtained at lower temperatures.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	1.0	190–250
$k^{II}$ ( $M^{-1} s^{-1}$ )	$\exp(154-1.63 W) \exp(-(38500-478 W)/T)$	200–230
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	298
$\Delta \log(k^{II})$	$\pm 0.2$	210–230

## Comments on preferred values

The experimental studies show that the uptake coefficient of HOBr or HCl on  $H_2SO_4/H_2O$  containing HCl or HOBr increases with the concentration of the co-reactant in solution. The  $\gamma_{HOBr}$  values from the earlier study of Abbatt (1995) are uncertain because the assumed relative solubilities of HOCl and HCl were incorrect. The measured  $\gamma_{HCl}$  in the studies of Waschewsky and Abbatt (1999) and Hanson (2003) agree well and the  $\gamma$  values show a strong increase with the water content of the  $H_2SO_4$  and a rather weak temperature dependence. The kinetics are consistent with the resistance model with a large accommodation coefficient ( $\alpha_b = 1$ ), and rate of uptake controlled by bulk phase chemical reaction  $HOBr + HCl$ :

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_b} \quad (1)$$

$$\Gamma_b = \frac{4RT H_{HOBr}}{\bar{c}} \sqrt{D_I k^{II} H_{HCl}^* \rho_{HCl}} \text{ for } \gamma_{HOBr} \quad (2)$$

$$\text{or } \Gamma_b = \frac{4RT H_{HCl}^*}{\bar{c}} \sqrt{D_I k^{II} H_{HOBr}^* \rho_{HOBr}} \text{ for } \gamma_{HCl} \quad (3)$$

32442

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



where  $H^*$  are the Henry's solubility constants and  $D_1$  are the liquid diffusion coefficients for HOBr and HCl which are a function of mole fraction of  $\text{H}_2\text{SO}_4$ ,  $X$ .

There are significant differences in the heterogeneous reaction rate constants derived from the experimental uptake coefficients in the two studies. Thus the values of  $k^{\text{II}}$  derived by Waschewsky and Abbatt (1999) from their data at 213–238 K was a factor of  $\sim 8$  lower than the values expected from extrapolation from the 250 K value of  $k^{\text{II}}$  derived by Hanson (2003), based on an estimated activation energy of  $15 \text{ kJ mol}^{-1}$ . This discrepancy can be largely attributed to the use of different values for  $H_{\text{HOBr}}$  in the derivation of  $k^{\text{II}}$ . The preferred value for  $H_{\text{HOBr}} = 5.22 \times 10^{-5} \exp(5427/T)$  is based on Hanson (2003) (see IUPAC evaluation VI.4.16), which showed only a weak dependence on acid strength in the range 60–70 %  $\text{H}_2\text{SO}_4$ . Using this choice of  $H_{\text{HOBr}}$  Hanson obtained coherent set of  $k^{\text{II}}$  as a function of  $T$  and wt %  $\text{H}_2\text{SO}_4$ , from the results of all studies.  $k^{\text{II}}$  showed a very strong  $T$ -dependence at 60 wt % acid and a much weaker one at 70 %. Hanson gives an expression for  $k^{\text{II}}$  for stratospheric conditions ( $wt = \text{wt} \% \text{H}_2\text{SO}_4$ ):

$$k^{\text{II}}(\text{M}^{-1} \text{s}^{-1}) = \exp(154 - 1.63 \cdot wt) \exp(-(38500 - 478 \cdot wt)/T)$$

This equation reproduces the experimentally derived  $k^{\text{II}}$  values at 210–228 K but the predicted values at 238 and 250 K are over-estimated by a factor of 3 and 6 respectively. Nevertheless this is the recommended expression to calculate the values of the uptake coefficients for stratospheric conditions, using Eqs. (1)–(3). The  $X$ -dependent expressions used for  $H_{\text{HCl}}^*$  (based on Shi et al., 2001; see this evaluation: VI.A4.14) and for  $H_{\text{HOBr}}^*$  (based on Hanson, 2003; see this evaluation: VI.A4.16), are given below:

$H_{\text{HOBr}}^*/\text{M} \cdot \text{bar}^{-1}$	$5.22 \times 10^{-5} \exp(5427/T)$	210–270
$H_{\text{HCl}}^*/\text{M} \cdot \text{bar}^{-1}$	$(0.094 - 0.61X + 1.2X^2) \exp(-8.68 + (8515 - 10718X^{0.7})/T)$	203
$D_{\text{HCl}}/\text{cm}^2 \text{s}^{-1}$	$7.8 \times 10^{-8} T/\eta$	190–240

32443

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



( $X$  = mole fraction  $\text{H}_2\text{SO}_4$ ,  $\eta$  = viscosity  $\text{H}_2\text{SO}_4$ )

The expression for  $D_1(\text{HCl})$  from Klassen et al. (1998), based on viscosity data given by Shi et al. (2001) (see this evaluation: VI.A4.25).

## 5 References

Abbatt, J. P. D.: J. Geophys. Res., 100, 14009–14017, 1995.

Hanson, D. R. and Ravishankara, A. R.: Geophys. Res. Lett., 22, 385–388, 1995.

Hanson, D. R.: J. Phys. Chem. A, 102, 4794–4807, 1998.

Hanson, D. R.: J. Geophys. Res., 108, art. no. 4239, doi:10.1029/2002JD002519, 2003.

10 Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

Waschewsky, G. C. G. and Abbatt, J. P. D.: J. Phys. Chem. A, 103, 5312–5320, 1999.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



### Experimental data

Parameter	wt % H <sub>2</sub> SO <sub>4</sub>	Temp./K	Reference	Technique/Comments
$\gamma(\text{HOBr})$ >0.25	69.8	228	Abbatt (1995)	CWFT-MS (a)

### 5 Comments

(a) HOBr ( $<1 \times 10^{12}$  molecule cm<sup>-3</sup>) was generated by the reaction sequence:  $\text{H} + \text{NO}_2 \rightarrow \text{OH} + \text{NO}$  and  $\text{OH} + \text{Br}_2 \rightarrow \text{HOBr}$ , ionised by electron impact and detected as  $\text{HOBr}^+$ . The carrier gas flow was humidified to maintain the H<sub>2</sub>SO<sub>4</sub> concentration (69.8 wt %). HBr (concentration  $\geq 1 \times 10^{12}$  molecule cm<sup>-3</sup>) was added simultaneously with HOBr ( $<1 \times 10^{12}$  molecule cm<sup>-3</sup>). HOBr loss was first order.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.3	190–250
$k^{\text{II}}/\text{M}^{-1} \text{s}^{-1}$	$5 \times 10^4$	228
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.5$	298

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


## Comments on preferred values

The single study of Abbatt (1995) at 228 K reported a rapid uptake coefficient of HOBr on a 69.8% H<sub>2</sub>SO<sub>4</sub> surface in the presence of 1 × 10<sup>12</sup> molecule cm<sup>3</sup> HBr. Experimental constraints did not allow measurements at lower concentrations. The authors derived a lower limit for the second order rate constant of  $k^{\text{II}} > 5 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$  for the liquid phase reaction: HOBr + HBr → H<sub>2</sub>O + Br<sub>2</sub> at 228 K, using the resistance model with an assumed accommodation coefficient of  $\alpha_b = 1$ , and using values of  $HD^{1/2}$  for HOBr and  $H_{\text{HBr}}^*$  from their own study. However if the IUPAC recommended expressions for the solubility terms ( $H_{\text{HOBr}} = 1.14 \times 10^6 \text{ M bar}^{-1}$  from VI.A4.16 and  $H_{\text{HBr}}^* = 7 \times 10^3 \text{ M bar}^{-1}$  (from VI.A4.17), and for the diffusion coefficients in 70% wt H<sub>2</sub>SO<sub>4</sub> at 228 K, the Abbatt  $\gamma$  value gives  $k^{\text{II}} > 6.8 \times 10^2 \text{ M}^{-1} \text{ s}^{-1}$ . Abbatt presents an argument for value of  $k^{\text{II}}$  closer to the diffusion controlled limit ( $10^7 \text{ M}^{-1} \text{ s}^{-1}$ ) so the assumptions that  $\alpha_b = 1$  and that the experimental uptake coefficient was controlled by HBr+HOBr liquid phase reaction rates, may be incorrect. In view of this we are unable to recommend a value for  $k^{\text{II}}$  consistent with  $\alpha_b = 1$ .

However, an estimate of the reactive uptake coefficient for HOBr + HBr for stratospheric conditions can be obtained using Eqs. (1) + (2), assuming  $\alpha_b = 0.3$  and a value of  $k^{\text{II}} = 5 \times 10^4 \text{ M}^{-1} \text{ s}^{-1}$ , as deduced by Abbatt (1995):

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_b} \quad (1)$$

$$\Gamma_b = \frac{4RT H_{\text{HOBr}}}{\bar{c}} \sqrt{D_1 k^{\text{II}} H_{\text{HBr}}^* \rho_{\text{HBr}}} \quad (2)$$

The values  $H_{\text{HOBr}}$  and  $H_{\text{HBr}}^*$ , are taken from the IUPAC evaluation and are given below:

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$H^*(\text{HOBr})/\text{M bar}^{-1}$	$5.22 \times 10^{-5} \exp(5427/T)$	210–270
$H^*(\text{HBr})/\text{M bar}^{-1}$	$10^{\{1000 \times (-1.977 \times 10^{-4} wt^2 - 2.096 \times 10^{-2} wt + 4.445)/T + (-8.979 \times 10^{-5} wt^2 + 2.141 \times 10^{-2} wt - 6.067)\}}$	195–250
$D_1(\text{HOBr})$	$6.4 \times 10^{-8} T/\eta$	200–300

In the expression for  $H^*(\text{HBr})$ ,  $wt$  is the  $\text{H}_2\text{SO}_4$  concentration in wt%.  $D_1(\text{HOBr})$  is taken from the work of Klassen et al. (1998), with viscosity calculated from the parameterisation given by Shi et al. (2001).

## References

- Abbatt, J. P. D.: J. Geophys. Res., 100, 14009–14017, 1995.  
 Klassen, J. K., Hu, Z., and Williams, L. R.: J. Geophys. Res., 103, 16197–16202, 1998.  
 Shi, Q., Jayne, J. T., Kolb, C. E. Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



BrONO<sub>2</sub> + H<sub>2</sub>SO<sub>4</sub> (l) → products

## Experimental data

Parameter	H <sub>2</sub> SO <sub>4</sub> /wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>				
0.5 (+0.5, -0.2)	45	210	Hanson and Ravishankara (1995)	CWFT-CIMS (a)
0.4 (+0.6, -0.2)	60	210		
0.3 (+0.7, -0.1)	70	220		
1.0 (0.1 $\mu$ m radius)	48	229	Hanson et al. (1996)	AFT-CIMS (b)
0.81 (0.1 $\mu$ m radius)	47	249		
0.75 (0.1 $\mu$ m radius)	66	272		
0.26 (0.1 $\mu$ m radius)	76.5	294		
0.8	66	250	Hanson (2003)	RWFT and AFT-CIMS (c)
0.95	72	250		
0.5	75	250–293		
0.2	79	250–293		
0.15	82	250–293		

## 5 Comments

(a) Uptake measurement on a ~~wetted wall flow tube~~ coupled to CIMS detection. The source ion was SF<sub>6</sub><sup>-</sup> and BrONO<sub>2</sub> and HOBr were monitored as BrONO<sub>2</sub>F<sup>-</sup> and SF<sub>5</sub>O<sup>-</sup>, respectively. Large corrections (factor 5) to the observed loss rates of BrONO<sub>2</sub> were made to take diffusive effects into account. HOBr was detected as the primary hydrolysis product of the title reaction at a typical partial pressure of BrONO<sub>2</sub> of (1–3) × 10<sup>7</sup> molecule cm<sup>-3</sup>.

(b) Aerosol flow reactor operated at total pressures (N<sub>2</sub>) ranging from 239 ± 40 (majority of experiments) to 825 ± 13 mbar between 249 and 298 K. BrONO<sub>2</sub> (10<sup>12</sup> molecule cm<sup>-3</sup>), HOBr and HNO<sub>3</sub> were detected as BrONO<sub>2</sub>F<sup>-</sup>, SF<sub>5</sub>O<sup>-</sup> and

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

$\text{NO}_3^-$ , respectively, using  $\text{SF}_6^-$  as reagent ion. The sulfuric acid aerosol particle size distribution was log-normal with a typical surface area-weighted peak radius of 0.1  $\mu\text{m}$ . Uptake coefficients  $\gamma$  were independent of the particle radius over a range of 0.05 to 0.17  $\mu\text{m}$  for 78% wt  $\text{H}_2\text{SO}_4$  particles at 298 K. In some experiments, HCl was added to the  $\text{H}_2\text{SO}_4$  in order to trap HOBr (as BrCl) in order to prevent potential regeneration of  $\text{BrONO}_2$  from reaction of  $\text{HNO}_3$  with HOBr.

(c) Rotating CWFT with stirring of the  $\text{H}_2\text{SO}_4$  film (58–85 wt %, 240–293 K) at 0.5–2.6 mbar total pressure. The aerosol flow tube was operated at room temperature and 160 Torr total pressure. Initial  $\text{BrONO}_2$  concentration was  $10^{11}$ – $10^{12}$   $\text{cm}^{-3}$  in the RWFT and  $10^{12}$ – $10^{13}$   $\text{cm}^{-3}$  in the AFT. Uptake coefficients showed a strong fall off towards high  $\text{H}_2\text{SO}_4$  wt %, a few representative values are listed in the table.

### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.8	210–300
$\Gamma_b$	$0.11 + \exp(29.2 - 0.40 \text{ wt})$	210–230
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.10$	210–300
$\Delta \log(\gamma)$	$\pm 0.15$	210–300

### Comments on preferred values

The three studies yield consistent results, with values of  $\gamma$  close to unity in diluted (45 wt %)  $\text{H}_2\text{SO}_4$ , and decreasing to 0.15 at  $\sim 80$  wt %. Owing to the large rate constant for  $\text{BrONO}_2$  hydrolysis in the condensed phase (the authors estimate a diffusion-limited rate constant  $k^{\text{II}}$  on the order of  $10^9 \text{ M}^{-1} \text{ s}^{-1}$ ) no size-dependence of  $\gamma$  is expected, because the reacto-diffusive length,  $(D_1/k^{\text{II}})^{0.5}$  is short. HOBr (and  $\text{HNO}_3$ ) is seen as

the major product of BrONO<sub>2</sub> uptake to H<sub>2</sub>SO<sub>4</sub>. Some Br<sub>2</sub> was also observed, which likely stems from reactions of HOBr and not from an additional BrONO<sub>2</sub> loss process.

Since the BrONO<sub>2</sub> solubility is not known, a detailed parameterisation of the uptake coefficient is not possible. We therefore adopt the expression suggested by Hanson (2003) that reproduces the uptake coefficients measured on both bulk and aerosol from all three studies:

$$\frac{1}{\gamma} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_b}$$

## References

Hanson, D. R. and Ravishankara, A. R.: Geophys. Res. Lett., 22, 385–388, 1995.

Hanson, D. R., Ravishankara, A. R., and Lovejoy, E. R.: J. Geophys. Res., 101, 9063–9069, 1996.

Hanson, D. R.: J. Geophys. Res., 108, art. no. 4239, doi:10.1029/2002JD002519, 2003.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**HNO<sub>2</sub> + HCl (aqueous sulphuric acid aerosol) → ClNO + H<sub>2</sub>O****Experimental data**

Parameter	$\rho_{\text{HCl}}/\text{mbar}$	$\rho_{\text{HNO}_2}/\text{mbar}$	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>						
$Y_{\text{HCl}} = (2.0 \pm 0.5) \times 10^{-2}$	$4.0 \times 10^{-7}$	$6.7 \times 10^{-7}$	60.8	208	Zhang et al. (1996)	CWFT-CIMS (a)
$Y_{\text{HCl}} = (1.1 \pm 0.4) \times 10^{-2}$			65.1	213		
$Y_{\text{HCl}} = (1.6 \pm 0.1) \times 10^{-2}$			68.0	217		
$Y_{\text{HCl}} = (2.0 \pm 0.2) \times 10^{-2}$			71.3	223		
$Y_{\text{HCl}} = (0.5 - 1.0) \times 10^{-4}$	$1.5 \times 10^{-5}$	$1.3 \times 10^{-5}$	62.5–65	225–230	Fenter and Rossi (1996)	Kn-MS (b)
$Y_{\text{HCl}} = (2.0 \pm 0.7) \times 10^{-3}$			60	225–230		
$Y_{\text{HNO}_2} = 1.1 \times 10^{-3}$	$4.8 \times 10^{-6}$		50	250	Longfellow et al. (1998)	RWFT-CIMS (c)
$Y_{\text{HNO}_2} = 1.5 \times 10^{-2}$	$1.6 \times 10^{-4}$		50	250		
$Y_{\text{HNO}_2} = 3.7 \times 10^{-2}$	$8.5 \times 10^{-7}$		50	205		
$Y_{\text{HNO}_2} = 1.5 \times 10^{-1}$	$1.2 \times 10^{-5}$		50	205		
$Y_{\text{HNO}_2} = 3.8 \times 10^{-2}$	$4.4 \times 10^{-6}$		60	219		
$Y_{\text{HNO}_2} = 9.2 \times 10^{-2}$	$2.8 \times 10^{-5}$		60	219		
$Y_{\text{HNO}_2} = 9.4 \times 10^{-3}$	$2.8 \times 10^{-6}$		60	230		
$Y_{\text{HNO}_2} = 5.9 \times 10^{-2}$	$5.4 \times 10^{-5}$		60	230		
$Y_{\text{HNO}_2} = 2.9 \times 10^{-3}$	$5.6 \times 10^{-6}$		60	250		
$Y_{\text{HNO}_2} = 2.4 \times 10^{-2}$	$9.6 \times 10^{-5}$		60	250		
$Y_{\text{HNO}_2} = 6.8 \times 10^{-2}$ ( $r_p = 64$ nm)	$4.8 \times 10^{-4}$		60	250		
$Y_{\text{HNO}_2} = 5.5 \times 10^{-2}$ ( $r_p = 232$ nm)	$4.8 \times 10^{-4}$		60	250		
$Y_{\text{HNO}_2} = 1.5 \times 10^{-2}$ ( $r_p = 83$ nm)	$1.2 \times 10^{-4}$		60	250		
$Y_{\text{HNO}_2} = 1.5 \times 10^{-2}$ ( $r_p = 306$ nm)	$1.2 \times 10^{-4}$		60	250		
$Y_{\text{HCl}} = 5.1 \times 10^{-3}$		$1.1 \times 10^{-7}$	70	215		
$Y_{\text{HCl}} = 1.1 \times 10^{-2}$		$1.6 \times 10^{-6}$	70	215		
$Y_{\text{HCl}} = 1.0 \times 10^{-3}$		$1.6 \times 10^{-6}$	67	250		
$Y_{\text{HCl}} = 5.2 \times 10^{-3}$		$1.1 \times 10^{-5}$	67	250		
$Y_{\text{HCl}} = 3.3 \times 10^{-4}$		$3.1 \times 10^{-6}$	70	269		
$Y_{\text{HCl}} = 1.3 \times 10^{-3}$		$1.6 \times 10^{-5}$	70	269		

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments

- (a) ~~Coated wall flow tube~~ with CIMS detection of  $\text{HNO}_2$  after reaction with  $\text{SF}_6^-$ . 0.1 mm thick liquid  $\text{H}_2\text{SO}_4$ . The total pressure was 0.5 mbar He; the partial pressure of  $\text{HNO}_2$  and HCl was around  $6 \times 10^{-7}$  mbar and  $6 \times 10^{-7}$  mbar, respectively. The sulphuric acid composition was controlled by maintaining fixed temperature and  $\text{H}_2\text{O}$  partial pressure of  $6 \times 10^{-4}$  mbar in the gas flows. CINO was observed as product with approximately unity yield. Separation of solubility driven HCl uptake from reaction driven uptake was achieved by allowing enough time for establishing steady state at each injector position.
- (b) Knudsen cell reactor with MS detection.  $\text{H}_2\text{SO}_4$  solutions were prepared by dilution of a 95 wt% solution.  $\text{HNO}_2$  and HCl were prepared by adding  $\text{H}_2\text{SO}_4$  to  $\text{NaNO}_2$  and  $\text{NaCl}$ , respectively. MS traces were consistent with full conversion of HCl into CINO in presence of  $\text{HNO}_2$ . Reaction probabilities (as listed in the table) were obtained from the difference of the HCl loss in presence and absence of  $\text{HNO}_2$  to account for HCl uptake due to solubility of HCl alone.
- (c) Most uptake experiments were performed in a 2.2 cm i.d. ~~coated wall flow tube~~ and a 1.84 cm i.d. rotating ~~wetted wall flow tube~~ with CIMS detection of  $\text{HNO}_2$ . The reaction of  $\text{HNO}_2$  with  $\text{SF}_6^-$  was revisited to come up with an improved estimate of the rate constant and the branching ratio for the formation of  $\text{SF}_5^-$  vs.  $\text{HFNO}_2^-$ . The observed loss of  $\text{HNO}_2$  in the flow tube was corrected for the hydrolysis reaction of the product CINO yielding back  $\text{HNO}_2$  by taking into account the reaction probability of CINO of  $2.5 \times 10^{-3}$  on 60 wt% and  $3.0 \times 10^{-3}$  on 50 wt% solutions at 250 K determined in separate experiments.  $\text{HNO}_2$  uptake to deliquesced NaCl at 268 K did not lead to detectable amounts of  $\text{HNO}_2$ . A few experiments were performed with aerosol particles also listed in the table.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$\alpha_s$	1	200–300
$\Gamma_{s,\text{HNO}_2}$	$9 \times 10^4 \exp(-3000/T) [\text{HCl}] H_{\text{NO}_+}^*$	200–300
$H_{\text{NO}_+}^*$ (M atm <sup>-1</sup> )	$C \exp(D \text{ wt})$	200–300
$C$ (M atm <sup>-1</sup> )	$2.0 \times 10^8 \exp(-14000/T)$	200–300
$D$ ((wt%) <sup>-1</sup> )	$297.3/T - 0.474$	200–300
$k_b$ (M <sup>-1</sup> s <sup>-1</sup> )	$25[\text{H}^+] T/\eta$	200–300
$\Gamma_{s,\text{HCl}}$	$4 \times 10^{-3} [\text{NO}^+] H_{\text{HCl}}^*$	200–300
<i>Reliability</i>		
$\Delta \log(\alpha_s)$	±0.3	200–300
$\Delta \log(H^*)$	±0.3	200–300
$\Delta \log(k_b)$	±0.3	200–300

## Comments on preferred values

The available studies agree that efficient reaction of HNO<sub>2</sub> with HCl occurs in sulphuric acid solutions to yield ClNO. Most likely, the reaction proceeds through one of the protonated forms of HNO<sub>2</sub>, e.g., NO<sup>+</sup>HSO<sub>4</sub><sup>-</sup>. The data by Longfellow et al. (1998) cover a sufficiently wide parameter range to allow constraining the kinetics. The pressure dependence observed for the uptake of HNO<sub>2</sub> in presence of HCl deviates from the expectation based on bulk reaction alone and thus is indicative of a surface process.

While Longfellow et al. reported individual contributions for the surface vs. bulk terms to overall uptake for each individual solution composition and temperature, we adopt here a more general parameterisation. Following a similar procedure as for HOCl and HCl, the bulk reaction is parameterised with the acid concentration and the viscosity

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

to represent diffusion limited kinetics of an acid catalysed reaction that largely controls temperature and composition dependence.

In absence of limitations by surface accommodation or surface to bulk transfer, for the uptake of HNO<sub>2</sub>, the uptake coefficient simplifies to:

$$\frac{1}{\gamma_{\text{HONO}}} = \frac{1}{\alpha_s} + \frac{1}{\Gamma_{s,\text{HONO}} + \frac{1}{\frac{1}{\Gamma_{sb,\text{HONO}}} + \frac{1}{\Gamma_{b,\text{HONO}}}}} \approx \frac{1}{\Gamma_{s,\text{HONO}} + \Gamma_{b,\text{HONO}}}$$

The surface reaction is assumed to be a direct reaction between NO<sup>+</sup> and HCl. In absence of data constraining their surface concentrations, we simply assume that the latter is proportional to the bulk concentration, which is tied to the gas phase pressures through the Henry's Law constant. The equilibrium constant describing the concentration of NO<sup>+</sup> is given by the second term of the expression for the solubility of HNO<sub>2</sub> given in datasheet VI.A4.7, as given in the preferred value table. The concentration of HCl is calculated using the expression for the solubility recommended in datasheet VI.A4.14:

$$H_{\text{HCl}}^* = (0.094 - 0.61X + 1.2X^2) \exp(-8.68 + (8515 - 10718X^{0.7})/T)$$

The mole fraction,  $X$ , of sulfuric acid is given by  $X = wt/(wt + (100 - wt)98/18)$ .

The bulk reaction rate constant requires knowing the proton concentration. We suggest using the parameterisation of Shi et al. (2001):

$$[\text{H}^+] = \exp \left[ 60.51 - 0.095 wt + 0.0077 wt^2 - 1.61 \times 10^{-5} wt^3 - (1.76 + 2.52 \times 10^{-4} wt^2) T^{0.5} + (-805.89 + 253.05 wt^{0.076})/T^{0.5} \right]$$

For the viscosity, we suggest using the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into tropospherically more relevant dilute solutions at high  $T$ :

$$\eta = AT^{-1.43} \exp(448\text{K}/(T - T_0)),$$

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Discussion Paper | Discussion Paper | Discussion Paper | Discussion Paper

with  $A = 169.5 + 5.18(wt) - 0.0825(wt)^2 + 3.27 \times 10^{-3}(wt)^3$ ,  
and  $T_0 = 144.11 + 0.166(wt) - 0.015(wt)^2 + 2.18 \times 10^{-4}(wt)^3$

5 This allows calculating the contribution of the bulk reaction according to:

$$\Gamma_{b,HNO_2} = 4H_{HNO_2}^* RT \sqrt{D_{l,HONO} \cdot k_b^{\parallel} \rho_{HCl} H_{HCl}^*} \left[ \coth \left( \frac{r_p}{l_{HNO_2}} \right) - \left( \frac{l_{HNO_2}}{r_p} \right) \right] / \bar{c}_{HNO_2}$$

For the effective solubility of  $HNO_2$ , we suggest using the expression recommended on datasheet VI.A4.7:  $H^* = A \exp(B(wt)) + C \exp(D(wt))$ , with  $A = 4.2 \times 10^{-6} \exp(4873/T)$ ;  $B = 13.16/T - 0.0856$ ;  $C = 2.0 \times 10^8 \exp(-14000/T)$ ;  $D = 297.3/T - 0.474$ .

10 The diffusion coefficient for  $HNO_2$  is parameterized by  $D_{l,HNO_2} = C_{HNO_2} T / \eta$ ; with  $C_{HNO_2} = 6.90 \times 10^{-8} \text{ cm}^2 \text{ cP K}^{-1} \text{ s}^{-1}$ , estimated as suggested by Klassen et al. (1998) using a molar volume of  $36 \text{ cm}^3/\text{mol}$  (da Silva et al., 2006). The reactivity-diffusive length needed to account for finite particle sizes is given by  $l_{HNO_2} = (D_{l,HNO_2} / (k_b \rho_{HCl} H_{HCl}^*))^{0.5}$

15 Similarly, for the uptake coefficient of HCl, the expression for the combined contribution of surface and bulk processes, is given by:

$$\frac{1}{\gamma_{HCl}} = \frac{1}{\alpha_s} + \frac{1}{\Gamma_{s,HCl} + \frac{1}{\frac{1}{\Gamma_{sb,HCl}} + \frac{1}{\Gamma_{b,HCl}}}}$$

The surface accommodation coefficient is assumed to be one. For calculating  $\Gamma_{s,HCl}$ , the tendency of HCl to adsorb to the surface is again expressed as proportional to the Henry's Law constant, while the concentration of  $NO^+$  is given by  $\rho_{HNO_2} H_{NO^+}^* \cdot \Gamma_{sb,HCl} =$   
20  $k_{sol} / k_{des}$  is obtained using  $k_{sol} = 7.84 \times 10^{10} / \eta$  (cP) and  $k_{des} = 8.0 \times 10^{17} \exp(-5000/T)$  from datasheet VI.A4.14. The expression for the resistance due to the bulk reaction is given by:

$$\Gamma_{b,HCl} = 4H_{HCl}^* RT \sqrt{D_{l,HCl} \cdot k_b^{\parallel} \rho_{HNO_2} H_{HNO_2}^*} \left[ \coth \left( \frac{r_p}{l_{HCl}} \right) - \left( \frac{l_{HCl}}{r_p} \right) \right] / \bar{c}_{HCl}$$

32455

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The diffusion coefficient of HCl can be expressed as  $D_{\text{HCl}} = 7.8 \times 10^{-8} T / \eta$  as explained on datasheet VI.A4.14. The reactivity-diffusive length needed to account for finite particle sizes is given by  $l_{\text{HCl}} = (D_{\text{HCl}} / (k_b \rho_{\text{HNO}_2} H_{\text{HNO}_2}^*))^{0.5}$ .

This surface and bulk reaction model describes the available data relatively well. The strongly non-linear dependencies of the solubility of  $\text{HNO}_2$  and HCl impose strong constraints on adjusting the model to the data, only three parameters in the two surface reaction terms and one for the bulk rate coefficient expression are adjusted. Assuming the bulk reaction only does not allow bringing the  $\text{HNO}_2$  uptake coefficients into consistency with the HCl uptake coefficients, and it would lead to quite strong disagreement with the pressure dependence for both, 50 and 60 wt% solutions. The only temperature dependence in the formalism required to fit the data was for the surface reaction of  $\text{HNO}_2$ , likely to represent the temperature dependence of the surface coverage. The case for the surface process for HCl is relatively poor. Longfellow et al. did not include it given the scatter in the data, even though the slope of the pressure dependence appears to be steeper than expected for bulk only reaction. We include this surface reaction to also improve agreement with the Zhang et al. (1996) data, even though it leads to some overestimate of the Longfellow data at 219 K, but we refrain from including a temperature dependence there in absence of more data to constrain it. The HCl uptake coefficients reported by Fenter and Rossi (1996) appear to be at least an order of magnitude lower than expected based on the relatively large  $\text{HNO}_2$  pressures used.

## References

- Zhang, R., Leu, M.-T., and Keyser, L. F.: J. Phys. Chem., 100, 339–345, 1996.
- Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.
- Longfellow, C. A., Imamura, T., Ravishankara, A. R., and Hanson, D. R.: J. Phys. Chem. A, 102, 3323–3332, 1998.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





da Silva, G., Dlugogorski, B. Z., and Kennedy, E. M.: *AIChE Journal*, 52, 1558–1565, 2006.

Klassen, J. K., Hu, Z., and Williams, L. R.: *J. Geophys. Res.*, 103, 16197–16202, 1998.

Fenter, F. F. and Rossi, M. J.: *J. Phys. Chem.*, 100, 13765–13775, 1996.

5 Williams, L. R. and Long, F. S.: *J. Phys. Chem.*, 99, 3748–3751, 1995.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



**HNO<sub>2</sub> + HBr (aqueous sulphuric acid aerosol) → BrNO + H<sub>2</sub>O****Experimental data**

Parameter	$\rho_{\text{HBr}}/\mu\text{bar}$	$\rho_{\text{HNO}_2}/\mu\text{bar}$	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficients: <math>\gamma</math></i>						
$\gamma_{\text{HBr}} = 4.0 \times 10^{-2}$	$10^{-5}$ – $10^{-4}$	$10^{-5}$ – $10^{-4}$	40	210	Seisel and Rossi (1997)	Kn-MS (a)
$\gamma_{\text{HBr}} = 4 \times 10^{-4}$	$10^{-5}$ – $10^{-4}$	$10^{-5}$ – $10^{-4}$	95	270		
$\gamma_{\text{HNO}_2} = 5.5 \times 10^{-3}$	$10^{-5}$ – $10^{-4}$	$10^{-5}$ – $10^{-4}$	40	210		
$\gamma_{\text{HNO}_2} = 2 \times 10^{-3}$	$10^{-5}$ – $10^{-4}$	$10^{-5}$ – $10^{-4}$	52–69	210–230		
$\gamma_{\text{HNO}_2} = 2.2 \times 10^{-2}$	$10^{-5}$ – $10^{-4}$	$10^{-5}$ – $10^{-4}$	95	270		

5 **Comments**

(a) H<sub>2</sub>SO<sub>4</sub> solutions were prepared by dilution of a 95 wt % solution. HNO<sub>2</sub> was prepared by adding H<sub>2</sub>SO<sub>4</sub> to NaNO<sub>2</sub>. Formation of BrNO was observed for all solution compositions. Uptake coefficients as listed in the table were due to the total uptake of HBr and not corrected for solubility limited uptake of HBr alone.

10 **Preferred values**

Parameter	Value	T/K
$\alpha_b$	1	200–300
$k_b$ (M <sup>-1</sup> s <sup>-1</sup> )	0.1	200–300
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	±0.7	200–300
$\Delta \log(k_b)$	±1	200–300

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Comments on preferred values

Seisel and Rossi (1997) observed conversion of  $\text{HNO}_2$  to  $\text{BrNO}$ . The interpretation of the kinetic data is somewhat complicated by the fact that solubility limited uptake of  $\text{HBr}$  or  $\text{HNO}_2$  was not always clearly separated from reaction limited uptake. E.g., at low wt %, the uptake of  $\text{HBr}$  was not sensitive to the presence of  $\text{HNO}_2$ ; similarly, at high wt %, the uptake of  $\text{HNO}_2$  was not sensitive to the presence of  $\text{HBr}$ . At intermediate compositions, the uptake coefficients seem to represent reaction limited conditions. Taking into account the composition dependent solubilities of  $\text{HBr}$  and  $\text{HNO}_2$  and assuming a simple bulk reaction to drive uptake leads to reasonable agreement with data especially for the  $\text{HNO}_2$  uptake coefficient, and reproduces the decreasing trend of the uptake coefficient of  $\text{HBr}$  with increasing wt % and the slightly increasing uptake coefficient of  $\text{HNO}_2$  over the same composition range. In absence of a more extensive data set over a larger range of pressures, we refrain from invoking a surface reaction to explain the relatively high uptake coefficients at high wt %, which are clearly underestimated by the model.

The uptake coefficient of  $\text{HNO}_2$  is given by:

$$\frac{1}{\gamma_{\text{HONO}}} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{b,\text{HONO}}}$$

$$\Gamma_{b,\text{HONO}} = 4H_{\text{HNO}_2}^* RT \sqrt{D_{i,\text{HNO}_2} \cdot k_b^{\text{II}} \rho_{\text{HBr}} H_{\text{HBr}}^*} \left[ \coth \left( \frac{r_p}{l_{\text{HNO}_2}} \right) - \left( \frac{l_{\text{HNO}_2}}{r_p} \right) \right] / \bar{c}_{\text{HNO}_2}$$

For the effective solubility of  $\text{HNO}_2$ , we suggest using the expression recommended on datasheet VI.A4.7:  $H^* = A \exp(B(wt)) + C \exp(D(wt))$ , with  $A = 4.2 \times 10^{-6} \exp(4873/T)$ ;  $B = 13.16/T - 0.0856$ ;  $C = 2.0 \times 10^8 \exp(-14000/T)$ ;  $D = 297.3/T - 0.474$ .

The diffusion coefficient for  $\text{HNO}_2$  is parameterized by  $D_{i,\text{HNO}_2} = C_{\text{HNO}_2} T / \eta$ ; with  $C_{\text{HNO}_2} = 6.90 \times 10^{-8} \text{ cm}^2 \text{ cPK}^{-1} \text{ s}^{-1}$ , estimated as suggested by Klassen et al. (1998), but using a molar volume of  $36 \text{ cm}^3 / \text{mol}$  (da Silva et al., 2006).

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



For the viscosity, we suggest using the parameterization presented by Shi et al. (2001), which fits well to data by Williams and Long (1995) but extends into tropospherically more relevant dilute solutions at high  $T$ :

$$\eta = AT^{-1.43} \exp(448K/(T - T_0)),$$

with  $A = 169.5 + 5.18wt - 0.0825wt^2 + 3.27 \times 10^{-3}wt^3$ ,

and  $T_0 = 144.11 + 0.166wt - 0.015wt^2 + 2.18 \times 10^{-4}wt^3$

According to data sheet VI.A4.17, the solubility of HBr can be parameterised by:

$$\log_{10} H_{\text{HBr}}^* = 1000m/T + b,$$

where  $m = m_1[\text{H}_2\text{SO}_4]^2 + m_2[\text{H}_2\text{SO}_4] + m_3$  and  $b = b_1[\text{H}_2\text{SO}_4]^2 + b_2[\text{H}_2\text{SO}_4] + b_3$  and the  $\text{H}_2\text{SO}_4$  concentration  $[\text{H}_2\text{SO}_4]$  is in wt%.

$$m_1(\text{wt}\%^{-2}K) = -1.977 \times 10^{-4}; m_2(\text{wt}\%^{-1}K) = -2.096 \times 10^{-2}; m_3(K) = 4.445;$$

$$b_1(\text{wt}\%^{-2}) = -8.979 \times 10^{-5}; b_2(\text{wt}\%^{-1}) = 2.141 \times 10^{-2}.$$

The reactivity-diffusive length needed to account for finite particle sizes is given by  $l_{\text{HNO}_2} = (D_{\text{l,HNO}_2}/(k_b \rho_{\text{HBr}} H_{\text{HBr}}^*))^{0.5}$ .

Similarly, the uptake coefficient of HBr is given by:

$$\frac{1}{\gamma_{\text{HBr}}} = \frac{1}{\alpha_b} + \frac{1}{\Gamma_{\text{b,HBr}}}$$

$$\Gamma_{\text{b,HBr}} = 4H_{\text{HBr}}^* RT \sqrt{D_{\text{l,HBr}} \cdot k_b^{\text{ll}} \rho_{\text{HNO}_2} H_{\text{HNO}_2}^*} \left[ \coth\left(\frac{r_p}{l_{\text{HBr}}}\right) - \left(\frac{l_{\text{HBr}}}{r_p}\right) \right] / \bar{c}_{\text{HBr}}$$

The diffusion coefficient of HBr can be expressed as  $D_{\text{l,HBr}} = 7.9 \times 10^{-8} T / \eta$  as explained on datasheet VI.A4.17. The reactivity-diffusive length needed to account for finite particle sizes is given by  $l_{\text{HBr}} = (D_{\text{l,HBr}} / (k_b \rho_{\text{HNO}_2} H_{\text{HNO}_2}^*))^{0.5}$ .

32460

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## References

- da Silva, G., Dlugogorski, B. Z., and Kennedy, E. M.: *AIChE J.*, 52, 1558–1565, 2006.
- Klassen, J. K., Hu, Z., and Williams, L. R.: *J. Geophys. Res.*, 103, 16197–16202, 1998.
- Seisel, S. and Rossi, M. J.: *Ber. Bunsen Ges. Phys. Chem. Chem. Phys.*, 101, 943–955, 1997.
- Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: *J. Geophys. Res.*, 106, 24259–24274, 2001.
- Williams, L. R. and Long, F. S.: *J. Phys. Chem.*, 99, 3748–3751, 1995.

ACPD

12, 32109–32472, 2012

### Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



HOCl + H<sub>2</sub>SO<sub>4</sub> → products

## Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Solubility: H (M atm<sup>-1</sup>), Diffusion: D<sub>1</sub> (cm<sup>2</sup> s<sup>-1</sup>)</i>				
$HD_1^{1/2} = 3.4 \times 10^{-5} \exp(2640/T)$ 58.5		200–230	Hanson and Ravishankara (1993)	CWFT-CIMS (a)
$HD_1^{1/2} = 3.9 \times 10^{-5} \exp(2810/T)$ 55.6		205–250		
$HD_1^{1/2} = 1.8 \times 10^{-5} \exp(3070/T)$ 50.5		200–230		
$HD_1^{1/2} = 2.6 \times 10^{-6} \exp(3590/T)$ 46		200–230		
$HD_1^{1/2} = 41$	57	204	Zhang et al. (1994)	CWFT-MS (b)
$H = 1.4 \times 10^3$	59.5	251	Hanson and Lovejoy (1996)	CWFT-CIMS (c)
$D_1 = (6 \pm 2) \times 10^{-7}$				
$HD_1^{1/2} = 9$	58	220	Donaldson et al. (1997)	CWFT-CIMS (d)
$HD_1^{1/2} = 7.1$	60			
$HD_1^{1/2} = 4.6$	63			
$HD_1^{1/2} = 2.7$	67			
$HD_1^{1/2} = 0.9$	75			
$H = 2.5 \times 10^3$	52	250		
$H = 1.1 \times 10^3$	62.5			
$H = 1.1 \times 10^3$	65			
$H = 0.9 \times 10^3$	67			
$H = 1.0 \times 10^3$	70			

## 5 Comments

(a) HOCl was made in-situ from hydrolysis of ClONO<sub>2</sub> in H<sub>2</sub>SO<sub>4</sub> or ex-situ by the reaction of Ca(OCl)<sub>2</sub> with HCl. Time dependent HOCl uptake coefficients were analysed to derive values of  $HD_1^{1/2}$  for 59.6, 55.6, 50.5 and 46 wt % H<sub>2</sub>SO<sub>4</sub> solu-

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



tions. The parameters listed in the table were obtained by fitting to the tabulated datasets given by Hanson and Ravishankara (1993).

- (b) Time dependent HOCl uptake coefficients were analysed to derive values of  $HD_1^{1/2}$ . Only a single value for  $HD_1^{1/2}$  was reported. The experiment was conducted at 204 K with the  $H_2O$  vapour pressure adjusted to  $3.8 \times 10^{-4}$  Torr. This was converted to a  $H_2SO_4$  concentration using the measurements of Zhang et al. (1993). Other datasets obtained were stated to be in good agreement with Hanson and Ravishankara (1993).
- (c) Rotated CWFT. HOCl (usually present at  $\sim 3 \times 10^{10}$  molecule  $cm^{-3}$ ) was synthesised by flowing HF over  $Ca(OCl)_2$ . The solubility was measured directly, the diffusion coefficient was derived from the dependence of the uptake coefficient on the HCl concentration in solution (i.e. measurement of  $H(D_1k)^{1/2}$  where  $k$  is the first order constant for reaction of HOCl with dissolved HCl).
- (d) Rotated CWFT. The  $H_2SO_4$  concentration (49–75 wt %) was adjusted by variation of the  $H_2O$  partial pressure. The HOCl solubility at 250 K was measured directly. Values for  $HD_1^{1/2}$  at 220 K were derived from time dependent uptake coefficients.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## Preferred values

Parameter	Value	T/K
$c$ (cm <sup>2</sup> cP K <sup>-1</sup> s <sup>-1</sup> )	$6.4 \times 10^{-8}$	200–250
$A$	$169.5 + 5.18 wt - 0.0825 wt^2 + 3.27 \times 10^{-3} wt^3$	
$T_0$	$144.11 + 0.166 wt - 0.015 wt^2 + 2.18 \times 10^{-4} wt^3$	
$H^0$ (M atm <sup>-1</sup> )	$1.91 \times 10^{-6}$	
$B$ (K)	5862.4	
$c_s$	0.0776	
$d$	59.18	
Reliability		
$\Delta \log(H)$	$\pm 0.3$	200–250

## Comments on preferred value

The majority of the experimental data on HOCl interaction with H<sub>2</sub>SO<sub>4</sub> solutions has come from one research group (Hanson and Ravishankara, 1993; Hanson and Lovejoy, 1996; Donaldson et al., 1997). Solubilities have been derived directly from stirred solutions or indirectly (via measurement of  $HD_1^{1/2}$ ) using static solutions.

The diffusion of HOCl in H<sub>2</sub>SO<sub>4</sub> has not been investigated directly. Klassen et al. (1998) have parameterised the diffusion coefficient for HOCl in H<sub>2</sub>SO<sub>4</sub> as:

$$D_1 = cT/\eta$$

where  $c$  is a constant (cm<sup>2</sup> cP K<sup>-1</sup> s<sup>-1</sup>) and  $\eta$  is the viscosity of H<sub>2</sub>SO<sub>4</sub> at a given wt% and temperature. Shi et al. (2001), have taken viscosity data from Williams and Long (1995) and for pure water to derive an extended formulation to cover H<sub>2</sub>SO<sub>4</sub> viscosities from 0 to 80 wt%, which can be used with the  $C$  constant above to derive  $D_1$  over the same concentration range using:

$$\eta = AT^{-1.43} \exp(448/(T - T_0))$$

32464

ACPD

12, 32109–32472, 2012

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





This parameterisation results in values of  $\eta$  which agree to better than 10% (for 40–70 wt% H<sub>2</sub>SO<sub>4</sub> solutions) with those of Klassen et al. (1998). Solubilities derived from measurements of  $HD_1^{1/2}$  and values of  $D_1$  calculated as described above are in good agreement.

Based on the experimental data of Hanson and Ravishankara (1993) and solubility data for pure H<sub>2</sub>O, Huthwelker et al. (1995) developed a semi-empirical expression for the solubility of HOCl in H<sub>2</sub>SO<sub>4</sub> solutions. Experimental data which appeared later (Donaldson et al., 1996) indicates an HOCl solubility which is substantially larger than that calculated at high H<sub>2</sub>SO<sub>4</sub> concentrations (>65 wt%). Donaldson et al. (1996) suggested application of a temperature independent, empirical correction factor ( $f$ ) to the calculated solubilities of Huthwelker et al.:  $f = 1 + 1.052 \exp(0.273x \text{ wt}\% - 65.66)$ . Use of this factor and the diffusion coefficients of Huthwelker et al. (1995) also aligns the direct measurements of  $HD_1^{1/2}$  presented in Donaldson et al. In addition, Shi et al. (2001) have reanalysed the available data for HOCl solubility in H<sub>2</sub>SO<sub>4</sub> and H<sub>2</sub>O and derived a further expression, which reproduces the data well. We adopt their formulation, which can be used for a wide range of temperatures (190–250 K) and H<sub>2</sub>SO<sub>4</sub> concentrations (0–75 wt%):

$$H = H^0 \exp(B/T) \exp(-SM_{\text{H}_2\text{SO}_4})$$

where  $S$  is the Setchenow coefficient ( $S = c_s + d/T$ ) and  $M_{\text{H}_2\text{SO}_4}$  is the molarity of the H<sub>2</sub>SO<sub>4</sub> solution, which can be calculated from temperature and concentration dependent H<sub>2</sub>SO<sub>4</sub> densities.

## References

- Donaldson, D. J., Ravishankara, A. R., and Hanson, D. R.: J. Phys. Chem. A, 101, 4717–4725, 1997.
- Hanson, D. R. and Lovejoy, E. R.: J. Phys. Chem., 100, 6397–6405, 1996.
- Hanson, D. R. and Ravishankara, A. R.: J. Phys. Chem., 97, 12309–12319, 1993.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Huthwelker, T., Clegg, S. L., Peter, T., Carslaw, K., and Luo, B. P.: J. Atmos. Chem., 21, 81–95, 1995.

Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res.-Atmos., 106, 24259–24274, 2001.

5 Williams, L. R. and Long, F. S.: J. Phys. Chem., 99, 3748–3751, 1995.

Zhang, R., Wooldridge, P. J., Abbatt, J. P. D., and Molina, M. J.: J. Phys. Chem., 97, 7351–7358, 1993.

Zhang, R. Y., Leu, M. T., and Keyser, L. F.: J. Phys. Chem., 98, 13563–13574, 1994.

ACPD

12, 32109–32472, 2012

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



## VI.A4.32

### CH<sub>3</sub>SO<sub>3</sub>H + H<sub>2</sub>SO<sub>4</sub> (aqueous) → products

#### Experimental data

Parameter	[H <sub>2</sub> SO <sub>4</sub> ]/wt %	Temp./K	Reference	Technique/Comments
<i>Uptake coefficient: <math>\gamma</math></i>				
0.86 ± 0.08	47–90 < 50 %	296	Hanson (2005)	AFT-CIMS (c)
0.64 ± 0.2	50	296		
0.25 ± 0.05	65	296		

#### 5 Comments

(c) Uptake to sulphuric acid aerosol was studied in a laminar flow reactor coupled to CIMS detection using HNO<sub>3</sub> as source of primary ions. Sulphuric acid particles were generated by homogeneous nucleation from supersaturated vapour leading to a lognormal particle size distribution within 50–120 nm, with a few 10<sup>4</sup> particles per cm<sup>3</sup>, characterised by a differential mobility analyzer. Concentrations of CH<sub>3</sub>SO<sub>3</sub>H were 3 × 10<sup>10</sup> cm<sup>-3</sup> in the flow tube. The measured uptake coefficients were corrected for gas phase diffusion using the Fuchs-Sutugin correction factor. The diffusion coefficient was directly measured based on the observed wall loss rates in absence of aerosol particles. Its average value was 0.0786 atm cm<sup>2</sup> s<sup>-1</sup> over the full range of humidity.

## Preferred values

Parameter	Value	T/K
$\alpha_b$	1	296
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.3$	296

## Comments on preferred values

5 The aerosol flow tube study by Hanson (2005) leads to uptake coefficients of  $\text{CH}_3\text{SO}_3\text{H}$  not significantly different from unity for solution compositions below 50 wt %, from which we adopt a value of 1 for  $\alpha_b$ . Hanson argues that the solubility of  $\text{CH}_3\text{SO}_3\text{H}$  may be lower at higher  $\text{H}_2\text{SO}_4$  concentration, so that uptake ran into solubility equilibrium during his experiments, which would explain the low uptake coefficient for high wt % solutions. No data of  $\text{CH}_3\text{SO}_3\text{H}$  solubility in  $\text{H}_2\text{SO}_4$  solutions are available to our knowledge  
10 to assess this in more detail.

## References

Hanson, D. R.: J. Phys. Chem. A, 109, 6919–6927, 2005.

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion





## Experimental data

Parameter		[H <sub>2</sub> SO <sub>4</sub> ] (wt %)	[HCl] M	Temp./K	Reference	Technique/Comments
$\gamma$ (N <sub>2</sub> O <sub>5</sub> )	$\Phi$ (ClNO <sub>2</sub> )					
0.11 ± 0.02	0.02 ± 0.002	50	0.0001	220	Talukdar et al. (2012)	RWFT-CIMS (a)
0.11 ± 0.01	0.24 ± 0.03		0.001	210		
0.11 ± 0.01	0.42 ± 0.03		0.002	210		
0.10 ± 0.01	0.92 ± 0.09		0.01	205		
0.11 ± 0.01	1.06 ± 0.11		0.1	215		
0.11 ± 0.01	0.047 ± 0.006	60	0.0001	220		
0.087 ± 0.01	0.282 ± 0.03		0.001	210		
0.11 ± 0.01	0.35 ± 0.03		0.0019	210		
0.124 ± 0.005	0.74 ± 0.01		0.005	205		
0.087 ± 0.005	0.92 ± 0.09		0.05	223		
0.14 ± 0.01	1.00 ± 0.11		0.1	215		

5 **Comments**

(a) Uptake coefficients of N<sub>2</sub>O<sub>5</sub> and yields of ClNO<sub>2</sub> and Cl<sub>2</sub> (relative to measured N<sub>2</sub>O<sub>5</sub> loss) in the interaction on neat and HCl-doped aqueous H<sub>2</sub>SO<sub>4</sub> (50% and 60%) were measured at 205–225 K. All gaseous species were measured using calibrated CIMS with I<sup>-</sup> as reagent ion. The acids were applied as a cold liquid to a rotating inner wall of the flow tube. Standard HCl solution was mixed with the aqueous H<sub>2</sub>SO<sub>4</sub> to produce the desired [HCl]<sub>liq</sub>. Gaseous HCl/He mixtures were added to the flow tube to maintain the correct [HCl]<sub>liq</sub>. There was no significant dependence of the uptake coefficient on temperature, [HCl]<sub>liq</sub> or [H<sub>2</sub>SO<sub>4</sub>] in the

10

## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



range covered. The reaction yields of ClNO<sub>2</sub> showed a strong function of [HCl] in the range (0.00006–0.1 M HCl); yields of Cl<sub>2</sub> were <1 % in all experiments.

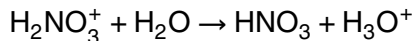
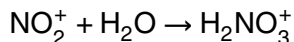
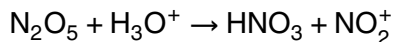
### Preferred values

Parameter	Value	T/K
$\alpha_b$	0.11	200–220
$k_{\text{Cl}}/k_{\text{H}_2\text{O}}$ (50 wt % H <sub>2</sub> SO <sub>4</sub> )	$(1.0 \pm 0.3) \times 10^{-4}$	190–280
$k_{\text{Cl}}/k_{\text{H}_2\text{O}}$ (60 wt % H <sub>2</sub> SO <sub>4</sub> )	$(1.6 \pm 0.4) \times 10^{-4}$	190–280
<i>Reliability</i>		
$\Delta \log(\alpha_b)$	$\pm 0.1$	200–230

### 5 Comments on preferred values

The cited study provides the only experimental data on the reactive uptake of N<sub>2</sub>O<sub>5</sub> in the presence of HCl on H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O surfaces. The uptake leads to formation of HNO<sub>3</sub> and ClNO<sub>2</sub> products. No formation of Cl<sub>2</sub>, which is produced from reaction of N<sub>2</sub>O<sub>5</sub> in aqueous salt solutions containing Cl<sup>-</sup> at higher temperatures and low pH, was observed on H<sub>2</sub>SO<sub>4</sub>/H<sub>2</sub>O surfaces (IUPAC datasheets VI.A2.06 and VI.A2.09).

These observations can be explained by a mechanism involving an acid-catalysed solvation of N<sub>2</sub>O<sub>5</sub> to form H<sub>2</sub>NO<sub>3</sub><sup>+</sup> followed by competition between its reaction with Cl<sup>-</sup> and with H<sub>2</sub>O. In the absence of HCl, uptake of N<sub>2</sub>O<sub>5</sub> on aqueous H<sub>2</sub>SO<sub>4</sub> leads solely to hydrolysis, producing HNO<sub>3</sub>. There is quite compelling evidence that under these conditions this occurs by an acid catalysed mechanism (Robinson et al., 1997):



## Evaluated kinetic and photochemical data for atmospheric chemistry

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

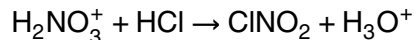
Full Screen / Esc

Printer-friendly Version

Interactive Discussion



In the presence of HCl the following reaction occurs:



As  $p(\text{HCl})$  increases the observed yield of  $\text{ClNO}_2$  product increases due to competition between reaction of  $\text{H}_2\text{NO}_3^+$  with  $\text{H}_2\text{O}$  and  $\text{Cl}^-$ , although the overall reactive uptake coefficient remains unchanged, and is equal to the recommended value for uptake on  $\text{H}_2\text{SO}_4/\text{H}_2\text{O}$  surfaces without HCl present (IUPAC, 2012). The yield of  $\text{ClNO}_2$  is given by the expression:

$$\Phi(\text{ClNO}_2) = \frac{k_{\text{R}}[\text{Cl}^-]}{k_{\text{R}}[\text{Cl}^-] + [\text{H}_2\text{O}]_l} \text{ where } k_{\text{R}} = \frac{k_{\text{Cl}}}{k_{\text{H}_2\text{O}}}$$

The recommended values for  $\alpha_b$  and  $k_{\text{Cl}}/k_{\text{H}_2\text{O}}$  are based on the work of Talukdar et al. (2012) together with consideration of the data for the uptake of  $\text{N}_2\text{O}_5$  into aqueous  $\text{H}_2\text{SO}_4$  over a wider range of temperature and compositions. The molarity of  $\text{Cl}^-$  in aqueous  $\text{H}_2\text{SO}_4$  can be calculated using the recommended expression for  $H_{\text{HCl}}^*$  from Shi et al. (2001):  $M_{\text{HCl}} = \text{molarity HCl (Mol dm}^{-3}\text{)} = p(\text{HCl}) \times H_{\text{HCl}}^*$ . Note that the recommended value of  $k_{\text{Cl}}/k_{\text{H}_2\text{O}}$  was derived using  $[\text{H}_2\text{O}]$  expressed as molarity.

It is generally believed (Thornton et al., 2003; Griffiths et al., 2009) that in neutral or mildly acidic solution, free  $\text{NO}_2^+$  is not formed and  $\text{HNO}_3$  and other products (e.g.  $\text{ClNO}_2$  in the presence of  $\text{Cl}^-$ ) are formed from competitive bulk liquid reactions of  $\text{H}_2\text{NO}_3^+$  formed via (slower) heterolytic dissociation of solvated  $\text{N}_2\text{O}_5$  (aq) to  $\text{H}_2\text{NO}_3^+$  (aq) and  $\text{NO}_3^-$  (aq).

## References

- Griffiths, P. T., Badger C. L., Cox, R. A., Folkers, M., Henke, H. H., and Mentel, T. F.: J. Phys. Chem. A, 113, 5082–5090, 2008.
- Robinson, G. N., Worsnop, D. R., Kolb, C. E., and Davidovits, P.: J. Geophys. Res., 102, 3583–3601, 1997.

**Evaluated kinetic and photochemical data for atmospheric chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Shi, Q., Jayne, J. T., Kolb, C. E., Worsnop, D. R., and Davidovits, P.: J. Geophys. Res., 106, 24259–24274, 2001.

Talukdar, R. K., Burkholder, J. B., Roberts, J. M., Portmann, R. W., and Ravishankara, A. R.: J. Phys. Chem. A, 116, 6003–6014, 2012.

- 5 Thornton, J. A., Braban, C. F., and Abbatt, J. P. D.: Phys. Chem. Chem. Phys., 5, 4593–4603, 2003.

**Evaluated kinetic  
and photochemical  
data for atmospheric  
chemistry**

M. Ammann et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

