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We thank both reviewers for their time and effort and their helpful and constructive comments. The original comments by the reviewers are in color. Our reply is in black.

There was apparently a problem with the page and line numbering in the review by Michael Prather. Now, both the page and line numbers indicated in the original review and the page and line numbers corresponding to the numbering of the published manuscript in ACPD are given. We thank the reviewer for revising his page and line numbers.

In addition to the changes requested by the reviewers, we have made some further changes. Two new tables were added. Table 7 shows the response in time-integrated global mean surface air temperature to an emission pulse. Table 5 provides the coefficient to represent the multi-model mean responses by sums of exponentials. The responses as simulated by individual models were also fitted. The coefficients are tabulated in the supplementary information for all models and for the responses in CO<sub>2</sub>. Results of the fits are compared with the model output in a complementary figure provided in the supplementary. Figure 5 (new figure 6) was adjusted; the results for the 5000 GtC pulse were scaled by a factor 1/50. This allows the reader to better compare the results for the 100 GtC versus the 5000 GtC pulse. A few groups slightly revised their output; all numerical values in the MS were checked again and revised where necessary. Revisions in numerical values are minor and do not affect our conclusions.

The revised manuscript is attached (see supplement). All changes in the text are highlighted in color using the “track change mode” of the text editor program.

**Review by M. Prather**

Review of ACP-2012-580  
Joos et al., "Carbon dioxide and climate impulse response functions for the computation of greenhouse gas metrics: A multi-model analysis"

This manuscript is a well written, important and new synthesis of the GWP and related metrics that has implications ranging from the scientists to the negotiators. It really must be published, but I hope that the authors will consider some revisions that correct from a misreading and that correct some seemingly longstanding misconceptions about time scales. In this, I admit that the references suggested here are primarily papers I have worked on, but feel that some of these published works are relevant to the discussion.

Michael Prather

Cross-cutting problems:

1-Airborne fraction is discussed, but it is not clearly noted that the value is specific to a given time, and that it depends of course (as the authors have shown) on the historical record of emissions to date. Thus it needs to have a year assigned to it when used.

Done. Clarified to read: “This last condition is expressed in terms of the 20<sup>th</sup> century cumulative airborne fraction,  $\gamma$ .”

45 2-Non-linearity is casually thrown about here, but the examples to demonstrate that the results  
46 are not linear are ridiculous, viz, of course the difference between 100 Gt-C and 5000 Gt-C  
47 does not scale as the latter puts us into a totally different carbon world. What is needed here  
48 is a recognition that the perturbations can and should be linearized (and will scale) if we deal  
49 with C pulses that are relevant to the government decisions on mitigation being proposed (1 to  
50 10 Gt-C). The paper presents inadvertently the misleading result that results are always non-  
51 linear. I know this is not true because we have the same issues with atmospheric chemistry  
52 which is inherently non-linear (bi-molecular reactions), but it is linear for CH<sub>4</sub> perturbations  
53 over a wide range (but not pulses of 1000 ppb).

54

55 In fact, if the AGWP is as non-linear as stated, then the AGWP reported here for 100  
56 Gt-C is not that for pulses of <1 Gt-C. Did you somehow scale the non-linearity out to  
57 get to the linear regime. Since no one would calculate a GWP for CH<sub>4</sub> or N<sub>2</sub>O with a  
58 100 Gt pulse (the chemistry of the atmosphere would be unrecognizable! as would the  
59 residence time of that pulse.) Thus what is needed is the AGWP for 0.1 Gt-C. Do you  
60 believe that the tabulated values in the paper are appropriate for this.

61

62 **Done.** A new figure is inserted where AGWP is plotted versus the magnitude of the initial  
63 pulse input. For this, we performed an additional set of 200 impulse response calculations  
64 varying the magnitude of the pulse between one GtC and 5000 GtC and for a constant  
65 background concentration of 389 ppm and for a varying background following RCP6. The  
66 results allow one to extrapolate the AGWP to an infinitely small pulse size. The difference in  
67 AGWP for an infinitely small pulse compared to a pulse of 100 GtC is about 1 percent and  
68 much less than the overall uncertainty in AGWP. Thus, the tabulated values given in the paper  
69 are appropriate.

70

71 The results also demonstrate that AGWP values start to scatter for pulses smaller than about  
72 50 GtC in the Bern3D-LPJ model. The reason for this scatter is that the signal-to-noise ratio is  
73 decreasing as variability in simulated CO<sub>2</sub> becomes more relevant.

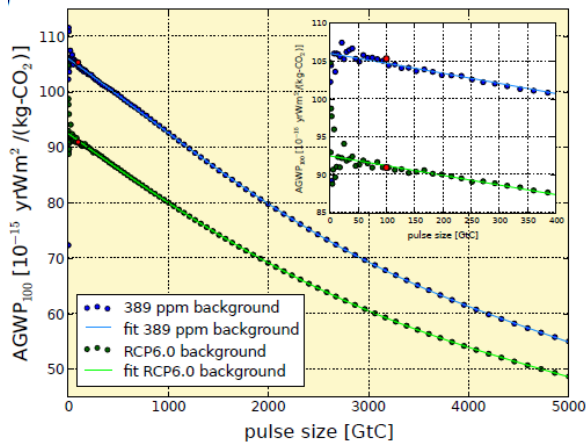
74

75 The following text is included in the revised MS in section 4.4.2 Pulse size:

76 “This smaller climate response per unit emission is a consequence of the smaller time-integrated  
77 forcing per unit emissions for larger pulses. The time-integrated radiative forcing at year 100 is  
78 smaller by 39 % for PI5000 than for PI100. The decrease in radiative efficiency (Equation 3) more  
79 than compensates for the larger time-integrated  $IRF_{CO_2}$  in PI5000 than PI100.

80 Next, the influence of the pulse size on the Absolute Global Warming Potential of CO<sub>2</sub> at year 100 is  
81 investigated in more detail (Figure 5b). Specifically, we ask how representative is the AGWP<sub>CO<sub>2</sub></sub> as  
82 determined with a pulse input of 100 GtC in our standard setup for the limiting case of an infinitely  
83 small pulse. The pulse size was varied between 1 GtC and 5000 GtC in the Bern3D-LPJ both for  
84 constant background conditions of 389 ppm as well as for the RCP6.0 case. AGWP<sub>CO<sub>2</sub></sub>( $t=100$  yr) is  
85 plotted versus pulse size in Figure 5b. A polynomial fit through the data points yields a continuous  
86 relationship between pulse size and AGWP over the range from 0 to 5000 GtC. The results show that  
87 AGWP<sub>CO<sub>2</sub></sub>( $t=100$  yr) for an infinitely small pulse is only about 1.2% higher than for a pulse size of 100  
88 GtC. Results also show that internal climate variability affect the computed AGWP<sub>CO<sub>2</sub></sub> significantly for  
89 small pulses of a few GtC only in the Bern3D-LPX. This is evidenced by the scatter in results for small  
90 pulses. In conclusion, the AGWP<sub>CO<sub>2</sub></sub> values tabulated in Table 4 are a good approximation for the  
91 limiting case of infinitely small carbon additions or removals to the atmosphere. “

92



93  
 94 Figure 5 (b)  $AGWP_{CO_2}$  versus pulse size for two different background conditions. Circles represent  
 95 results from individual simulations and the lines fits through the results. The standard pulse size of  
 96 100 GtC is indicated by red circles.

97  
 98 In addition, we have also determined the influence of varying background conditions. This is  
 99 described in section 4.4.1 Background conditions:

100 “The response for temporally varying background conditions is in addition explored with one  
 101 model (Bern3D-LPJ) for illustrative purposes. Emissions of  $CO_2$  and non- $CO_2$  agents are  
 102 prescribed to follow those from the Representative Concentration Pathways RCP2.6, RCP4.5,  
 103 RCP6.0 and RCP8.5 in the control setup. The same procedure was applied to determine the  
 104 IRF as in the standard setup. However, forcing ( $CO_2$ , non- $CO_2$ , aerosoles, landuse area) was  
 105 extended based on the RCPs until year 2300 as described in (Zickfeld et al., 2012). After year  
 106 2300, the forcing is extended until year 3010 by using 2300 values. The pulse was released in  
 107 year 2010 instead of 2015 as in the 389ppm background scenario. The evolution of  $IRF_{CO_2}$   
 108 (Figure 5a) is relatively similar between the standard case (389 ppm background) and  
 109 RCP2.6, but very different for the three other RCP cases.  $IRF_{CO_2}$  decreases in all cases to  
 110 about 70% in the first two decades after the pulse. Then, it continues to decrease for the  
 111 standard and the RCP2.6 cases, whereas  $IRF_{CO_2}$  increases again in the other cases as  
 112 atmospheric  $CO_2$  and global warming continues to rise in these scenarios. For RCP8.5, the  
 113 pulse fraction remaining airborne is still well above 80% at year 1000. The time-integrated  
 114  $IRF_{CO_2}$  evaluated at year 100 is 62 years for the 389 ppm background and 66, 68, 69 and 75  
 115 years for RCP2.6, RCP4.5, RCP6.0, and RCP8.5, respectively. The resulting perturbation in  
 116 radiative forcing is evaluated as difference in forcing between the control without pulse and  
 117 the corresponding pulse run and using the non-linear equation (3). AGWP range between 105  
 118 and  $85 \times 10^{-15} \text{ yr W m}^{-2} \text{ kg-CO}_2^{-1}$  for the five cases and at year 100. The RCP8.5 case,  
 119 although featuring the largest time-integrated  $IRF_{CO_2}$ , has the smallest AGWP of the five  
 120 cases as the radiative efficiency decreases with higher  $CO_2$  concentration.”

121  
 122  
 123

124 19801/11  
 125 Abstract p.2 / line 18 Please give the percent for the ‘remainder’ as we are unsure if the 60%  
 126 is absolute or just the fraction of the 76% taken up.  
 127 Done.

128  
 129 801/16  
 130 2/22 Do you not want to note that this depends on the future scenario. Both here and in the  
 131 text, the mention that it depends very much on which scenario we follow is subdued.

- 132 a) Text clarified in the abstract by pointing out that the pulse is added to a constant  
133 background concentration: “For a 100 Gt-C emission pulse added to a constant CO<sub>2</sub>  
134 concentration of 389 ppm, ...”  
135 b) Text added on p2 line 28 of original MS: “Uncertainties related to the size of the  
136 emission pulse, the atmospheric and climatic background conditions or the choice of  
137 the future scenario, and the carbon cycle-climate feedback are assessed in sensitivity  
138 simulations.”  
139

140 801/19 (re the agreement to 15% is over what time scales)  
141 2/25-26 Over what time scales? At the very long it seems like the different IPCC models were  
142 very different.  
143

144 Done. Text clarified to read: “Estimates for time-integrated response in CO<sub>2</sub> published in the  
145 IPCC First, Second, and Fourth Assessment and our multi-model best estimate all agree  
146 within 15% during the first 100 years.  
147

148 801/21  
149 2/28 Is this stating the obvious that larger pulses have larger integrated effects, or that it is  
150 non-linear in relative response. In which case this conclusion really needs to be modified to  
151 address the problem noted above.

152 Done. Text clarified to read: “The integrated CO<sub>2</sub> response, normalized to the pulse size, is  
153 lower for pre-industrial conditions, compared to present day, and lower for smaller pulses  
154 than larger pulses.”  
155

156 803/6  
157 4/1 “to change.” With what? Scenarios? New knowledge?

158 Done. Text clarified to read: “...the atmospheric and climatic conditions influencing the CO<sub>2</sub>  
159 response continue to change with time.”  
160

161 803/26-28  
162 4/23 Explain what this (different years, different CO<sub>2</sub>) means to first-order in the value of the  
163 AGWP for example.

164 Done. Text clarified to read: “Such changes in the background concentration cause both the  
165 radiative forcing and the response function to change, but the changes partially cancel leading  
166 to smaller changes in the AGWP (Caldeira and Kasting, 1993;Reisinger et al., 2011).  
167 .  
168

169 804/7  
170 4/30 Why not give some obvious references here, e.g., [Prather and Hsu, 2010]

171 Done. Reference added.  
172

173 804/22-26  
174 5/13 One of the few papers following this “causal chain” that several of us worked on is  
175 [Prather et al., 2009]

176 Done. Reference added.  
177

178 805/5 (although this is my personal use, since lifetime is primarily a ‘budget’ number)  
179 5/22 Lifetimes should more accurately be ‘time scales’ see later notes and discussion of such  
180 terms [Prather, 2007]

181 Done. Text changed to “The different perturbation timescales imply ..”  
182

183 806/7ff

184 The IRF function discussion is great, and indeed it is a Green's fn, but that is a complex mix  
185 of e-folds, not a single time scale for most all gases. The problem begins with line 13  
186 Done. The text has been revised as suggested and as further explained in reply to the next few  
187 comments.

188

189 806/13-20 (IRFs are truly a mix of exponentials unless the Jacobians change during the  
190 decay)

191 6/22-25 This simple-minded approximation should not continue to be propagated in this  
192 important paper. Please look at [Prather, 2007] for discussion of the different time scales. We  
193 have approximated the impact of long-lived gases as a single e-fold. It is not correct. For  
194 example, N2O emissions have no effective decay for 2-3 years until they get into the  
195 stratosphere. For all chemically reactive species there are indeed several times scales as there  
196 are for CO2. Unlike what is said in the CO2 discussion, these are not parametric fits, but are  
197 the fundamental modes of that species in the atmosphere. Further the steady-state lifetime of a  
198 perturbation (does not work for CO2 except over >106 yr) is the integrated impact of a pulse  
199 (see P2007). It can be approximated as a single e-fold but that is only approximate. Thus in  
200 truth all trace gas emissions follow the eqn 5.

201 Done. Text changed to read:

202 "It is convenient to describe the  $IRF_x$  by exponential functions (Prather, 2007;Maier-Reimer and  
203 Hasselmann, 1987).

$$204 \quad IRF_x(t) = \sum_{i=0}^n a_{x,i} \cdot \exp\left(\frac{-t}{\tau_{x,i}}\right) \quad \text{for } t \geq 0. \quad (5)$$

205 The unitless coefficients  $a_{x,i}$  represent a fraction that is associated with a certain nominal time  
206 scale  $\tau_{x,i}$  and their sum equals 1. In turn the AGWP for gas x is:

$$207 \quad AGWP_x(t) = A_x \sum_{i=0}^n a_{x,i} \cdot \tau_{x,i} \left(1 - \exp\left(\frac{-TH}{\tau_{x,i}}\right)\right) \quad (6)$$

208

209 and further down:

210 "Most GHGs are involved in complex chemical reactions in the atmosphere and are  
211 transported within the atmosphere. A local perturbation in one species invokes perturbations  
212 elsewhere on a range of timescales and often involving many other species. The chemistry-  
213 transport system can be linearized and represented with the help of eigenvalue decomposition  
214 following Prather, 2007. Then, it becomes clear that the perturbation time scales  $\tau_{x,i}$  represent  
215 the (negative inverse) eigenvalues of the leading chemical modes of gas x. "

216

217

218 806/23

219 7/6 The TAU\_x above is a time scale (should be a sum over different decay rates), the TAU\_x  
220 used in this line for chemical gases should be the steady-state lifetime of a perturbation as this  
221 is the weighted sum of the different modes excited by the pulse.

222 Done. Text changed to read:

223 AGWP increases with increasing time horizon TH to finally approach a constant value for TH several  
224 times larger than the largest perturbation time scale of gas x. The AGWP becomes the product of the  
225 "steady-state" life time of a perturbation,  $\tau_{x,ss}$ , (Prather, 2007) and the radiative efficiency, i.e.,  
226  $AGWP_x = A_x \tau_{x,ss}$ . The steady-state perturbation lifetime is the weighted sum over all time scales  
227 ( $\tau_{x,ss} = \sum a_{x,i} \tau_{x,i}$ )..

228

229 807/1  
230 7/9 CO<sub>2</sub> and N<sub>2</sub>O(?) - add another gas here.  
231 Done. Text reads now: “..such as CO<sub>2</sub>, N<sub>2</sub>O, or SF<sub>6</sub>.”  
232  
233 807/4ff  
234 7/12 Add a restriction on your increasing time horizon (e.g., from 20 to 500 yr) This statement  
235 fails for very short times (<3 yr) when the CO<sub>2</sub> drops rapidly, but N<sub>2</sub>O and CH<sub>4</sub> decay only  
236 slowly! The statement fails for very short-term GWPs.  
237 Done. Text reads now: “For instance, the GWP values for CH<sub>4</sub>, which has an adjustment time of  
238 approximately 12 years, decrease with increasing time horizon (except for time horizons of a few  
239 years only), since GWP is defined with the integrated *RF* of CO<sub>2</sub> in the denominator.”  
240  
241 807/8  
242 7/16 SF<sub>6</sub> is an extreme example  
243 Done. Text reads now: “For long-lived gases (e.g. N<sub>2</sub>O, SF<sub>6</sub>) the development in GWP is  
244 controlled by both the increasing integrals of the radiative forcing by the long-lived gas and  
245 CO<sub>2</sub>.”  
246  
247 807/15  
248 7/22 “remains airborne AND in all reservoirs.”  
249  
250 Done. Text reads now: “A substantial fraction of the initial perturbation by the emission pulse  
251 remains in the atmosphere and the ocean for millennia.”  
252  
253 807/20  
254 7/25 This structure applies to all gases, not just CO<sub>2</sub> In this section there is no talk about  
255 emission of fossil C that will be converted to CO<sub>2</sub> in the atmosphere (CO, CH<sub>4</sub>, C<sub>2</sub>H<sub>6</sub>). It  
256 should be mentioned somewhere.  
257 Done. Equations (3) and (4) of the original manuscript are now replaced by equations (5) and  
258 (6) (see responses above).  
259  
260 We do not discuss oxidation of non-CO<sub>2</sub> carbon compounds as this distracts from the text  
261 flow and it is not of practical relevance for this study. The vast majority of anthropogenic CO<sub>2</sub>  
262 is emitted in the form of CO<sub>2</sub>, but CO<sub>2</sub> is also perturbed through the oxidation of  
263 anthropogenic CO, methane and other carbon compounds in the atmosphere. This carbon flux  
264 from oxidation within the atmosphere is two orders of magnitude smaller than the total  
265 anthropogenic carbon emissions. For example, assume that the current 1 ppm excess in  
266 atmospheric methane decays with a mean steady state perturbation life time of 12 years yields  
267 a flux of ~0.1 ppm/yr=0.2 GtC/yr only. The carbon emissions of non-CO<sub>2</sub> compounds is  
268 included in the carbon flux from fossil fuel burning and land use change that currently total  
269 about 10 GtC/yr.  
270  
271 807/24-25  
272 8/4 Disagree, these are the eigenvalues of the system, not just some fitting. You sort of say  
273 this below in line 14  
274 We disagree in the case of CO<sub>2</sub>. Text has been clarified to read: “The continuum of timescales  
275 involved in the redistribution of CO<sub>2</sub> can be approximated in practice by a few timescales only. It is  
276 usually sufficient to consider three to four terms in the sum in equation (5). Then, the coefficients  
277  $a_{\text{CO}_2,i}$  and  $\tau_{\text{CO}_2,i}$  have no direct process-based meaning, but are fitting parameters chosen to represent  
278 a given model-based *IRF*<sub>CO<sub>2</sub></sub>.”



279 Often equally good fits of the IRF of CO<sub>2</sub> can be obtained by different sets/values of  $a_{CO_2,i}$  and  $\tau_{CO_2,i}$   
 280 preventing a direct interpretation of the time scales. This is somewhat different to the eigenvalue  
 281 decomposition that reveals (leading) timescales that are directly linked to reaction rates.  
 282 (Eigenvalue decomposition can be applied to ocean box-diffusion models or box models and for such  
 283 simplified systems the timescales are linked to the diffusion or transport coefficients of the model)

284  
 285  
 286

287 808/19

288 8/24 I remain unconvinced that this statement is true for small pulses. IF you agree, please fix  
 289 here and throughout.

290

291 Please see the discussion in Joos et al., Tellus, 1996. Figure 2 of this paper included below  
 292 shows that the simple representation of the carbon cycle with an atmospheric response

293 function,  $CO_2(t) = \int_{t_0}^t e(t') \cdot IRF_{CO_2}(t-t') dt' + CO_2(t_0)$ , yields large deviation in the air-sea  
 294 flux and thus in atmospheric CO<sub>2</sub>.

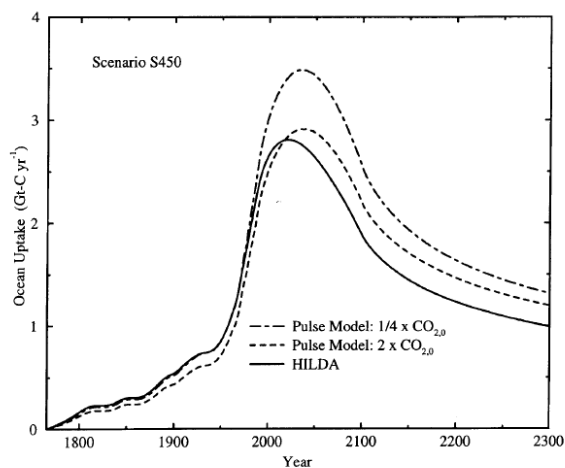


Fig. 2. Oceanic carbon uptake for IPCC scenario S450 as simulated with the HILDA model (solid line) and the atmospheric pulse response model using the response functions shown in Fig. 1. For the atmospheric pulse model, eq. (1) is solved simultaneously with the following mass balance equation:  $e = dN_a/dt + dN_{oc}/dt$  which expresses that the emissions,  $e$ , are equal to the change in atmospheric,  $dN_a/dt$ , and oceanic inventory,  $dN_{oc}/dt$ . Atmospheric CO<sub>2</sub> concentration and the change in atmospheric inventory is prescribed in order to stabilize concentration at 450 ppm (Enting et al., 1994).

295  
 296

297 809/4

298 9/4 By “setup” do you mean future scenarios? Please expand a bit

299 Done. Text reads now: “.. depends on the details of the experimental setup (background  
 300 concentration, pulse size)..”

301

302 810/3-15 // I am not sure what I was concerned with here, I think it was about the consistency,  
 303 but skip it unless it evokes some interest on your part.

304 9/23-27 This argument only works if you have evidence that the two results co-vary in a self-  
 305 consistent model, otherwise there is nothing wrong with what has been done. Please make  
 306 clear is this is so and you find it from the results here.

307 Skipped as suggested.

308

309 812/1ff // this comment is a bit pedantic, but having just worked hard on a paper to consider  
 310 all these effects, it seems like the 4-place accuracy is not right. The stratospheric fill is a  
 311 more important bias (affects the CH<sub>3</sub>CCl<sub>3</sub> to CH<sub>4</sub> OH-lifetime scaling.)

312 11/14 The number quoted 2.123 GtC/ppmv is wrong. For one, the units are ppm in mole  
313 fraction, they are not by “volume”. No one measures by volume, and then the virial  
314 corrections for CO<sub>2</sub> will change the mole fraction anyway – so no ‘v’. Two, the coefficient  
315 appears to be in error, see the recent publication [Prather et al., 2012] that clarifies the  
316 mistakes that have been made in converting thus. The dry-air mass of the atmosphere must be  
317 different, and one cannot assume that current CO<sub>2</sub> perturbations are well mixed into the  
318 stratosphere. The CO<sub>2</sub> number (if one takes 98.5% fill) is 2.086 Gt-C/ppm. Thus there is a 2%  
319 bias. The 98.5% is about right given the 2 ppm/yr increase and a 2.5-year lag to fill most of  
320 the stratosphere. If 100% it is 2.118.

321 Done. Thank you for pointing this out.

- 322 - ppmv replaced by ppm
- 323 - Text added: “Recently, the factor to convert ppm into mol was slightly revised to  
324 0.1765 (±5%) Pmol/ppm (Prather et al., 2012); this yields a conversion factor of  
325 2.120 GtC/ppm (0.1765 Pmol/ppm x 12.01 gC/mol) when assuming that the CO<sub>2</sub> is  
326 distributed evenly in the atmosphere as done here. For current emissions, the increase  
327 in the stratosphere lags the tropospheric increase and a 1 ppm change in the  
328 troposphere may corresponds to a mean atmospheric change that is about 1 to 2%  
329 lower. In the following these uncertainties of order 2% are neglected.”

330

331 812/18

332 11/31 It would be good to know what range of non-CO<sub>2</sub> forcing was included? e.g., CH<sub>4</sub>  
333 lifetime?

334 Sensitivity simulations (see SI) show that results of this study do hardly depend on the  
335 treatment of non-CO<sub>2</sub> forcings.

336

337 817/5-9

338 15/25 Could the reason here be that the simple models just do not have the degrees of freedom  
339 that the physical models do?

340 MAGGICC is a box-type model with highly simplified assumptions on ocean and land  
341 dynamics. GENIE and Bern3D-LPJ include both a 3-d dynamic ocean circulation model and a  
342 dynamic global vegetation model. The latter two models have by design a large degree of  
343 freedom in ocean and land dynamics, albeit they lack variability in the atmosphere.

344

345 817/26 (& less so 12)

346 16/13 Please give the year for the airborne fraction, it is not constant in general.

347

348 Done. Text clarified to read: “This last condition is expressed in terms of the 20<sup>th</sup> century  
349 cumulative airborne fraction  $\gamma$ ”

350

351 819/26

352 18/3-9 Why give the equilibrium response values, they are not relevant?

353

354 Text clarified to read: “What magnitude in the SAT response is to be expected from this  
355 forcing? The equilibrium response in global mean surface air temperature (SAT) to these  
356 forcing values are 0.49°C (year 0), 0.21°C (year 100) and 0.13°C (year 1000) when assuming  
357 for illustrative purposes a typical mid-range climate sensitivity of 3°C for a nominal doubling  
358 of CO<sub>2</sub>.”

359

360

361 820/24ff

362 19/6 No, it could be done with more ensembles.



363  
364 Done. We do not say that it is not possible with enough ensembles. Text clarified to read:  
365 “This internal variability of the more comprehensive models makes the extraction of the  
366 response in SAT challenging for these models and a well-defined separation of the forced  
367 response from the models’ internal variability is not achieved when relying on single  
368 simulations.”

369  
370 821/5-20// was looking for clarity on which perturbation was being present with all the  
371 numbers.

372 19/29 Recommend that you repeat the 100 Gt-C number here as a reminder of what the 20 Gt-  
373 C is with respect to.

374  
375 Done.

376  
377 823/14  
378 21/8 Is not this ‘historical’, rather than background, since the history matters.

379  
380 Done. In our setup, the exact history of the CO<sub>2</sub> evolution up to year 2010 does hardly  
381 influence the IRF as the IRF is obtained by difference from an experiment with the pulse  
382 emission and a corresponding simulation without pulse emissions but the same atmospheric  
383 history. The effect of different historical emissions (within the range of historical emissions  
384 datasets) will be much smaller than the effects of other issues considered in this article. The  
385 IRF do depend on how emissions evolve in the future. This is now explicitly demonstrated in  
386 a new figure (Figure 6a). This figure compares the IRFs obtained under a constant CO<sub>2</sub>  
387 concentration background of 389 ppm and for time-varying backgrounds corresponding to the  
388 four RCPs.

389  
390  
391 825/2ff  
392 22/23 Basically the ocean uptake for 5000 Tg-C is flux limited.

393 The ocean uptake rate is influenced by the carbonate chemistry (Revelle factor) and the rates  
394 of surface-to-deep transport.

395  
396 825/17ff  
397 23/13 Again, I think this is incorrect for useful ranges of C pulses  
398 Please see our response to point 1.

399  
400 826/21 This was directed at conclusion here and present the paragraphs above that appear to  
401 be only using one model. 24/2 Is this correct? only one model (Bern3D-LPJ) ? Then most of  
402 the discussion on this page is weak and maybe should be dropped.  
403 We believe it is important to discuss these issues in an illustrative way.

404  
405 827/20-21  
406 24/24 The uncertainty you are talking about cannot be compared directly to the negotiating  
407 positions of governments (Time Horizon uncertainty is not scientific, is it?).

408  
409 Done. Text clarified to read: “The subjective choice of the time horizon has a much larger  
410 influence on the range in absolute global warming potential of CO<sub>2</sub> and in the global warming  
411 potential of most other agents than uncertainties associated with the computation of these

412 values for a given time horizon. The uncertainty in AGWP (in units of  $\text{yr W m}^{-2}$  per  $\text{kg-CO}_2$ )  
413 can be mapped to a range in the time horizon (in units of year).”

414

415 827/28: ...hexafluoride calculated with a single e-fold decay of 12 yr....” since these are not  
416 necessarily lifetimes.

417 25/2 “with a single e-fold perturbation: : :

418 Done. Text modified as requested.

419

420 829/3-4

421 26/3 There may be other recent papers that followed this causal chain and quantified it,  
422 consider referencing [Prather et al., 2009]

423 Done. Here we reference Steinacher et al., Nature, in revision, 2012. The Prather 2009  
424 reference has been given in the introduction

425

426 829/13 “anthropogenic” is clumsy but better than this.

427 26/11 So something other than “man-made” – it does not translate well into American  
428 English.

429 Replaced with anthropogenic that is also used in the following lines.

430

431

432 829/16-17

433 26/14 What do you mean by all greenhouse gases – this result is very dependent on what is  
434 included – aerosols, O3, CH4, CFCs: : :.

435

436 Aerosols are not GHGs. The others are included as detailed by Strassmann et al., CD, 2009.

437

438

439 I did not review the Appendix.

440

441