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A Functional Group Oxidation Model (FGOM) for SOA formation and aging

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| DDr | A Functional Group Oxidation Model | | | | | | |
| Discussion | X. Zhang and J. H. Seinfeld | | | | | | |
| כמס | Title I | Page | | | | | |
| D | Abstract | Introduction | | | | | |
| _ | Conclusions | References | | | | | |
| | Tables | Figures | | | | | |
| in Dn D | 14 | ۶I | | | | | |
| DDDr | • | • | | | | | |
| | Back | Close | | | | | |
| | Full Screen / Esc | | | | | | |
| ssior | Printer-frien | dly Version | | | | | |
| עס | Interactive | Discussion | | | | | |
| D | | | | | | | |

Abstract

Secondary organic aerosol (SOA) formation from a volatile organic compound (VOC) involves multiple generations of oxidation that include functionalization and fragmentation of the parent carbon backbone and, likely, particle-phase oxidation and/or ac-

- ⁵ cretion reactions. Despite the typical complexity of the detailed molecular mechanism of SOA formation and aging, a relatively small number of functional groups characterize the oxidized molecules that constitute SOA. Given the carbon number and set of functional groups, the volatility of the molecule can be estimated. We present here a Functional Group Oxidation Model (FGOM) that represents the process of SOA for-
- ¹⁰ mation and aging. The FGOM contains a set of parameters that are to be determined by fitting of the model to laboratory chamber data: total organic aerosol concentration, and O:C and H:C atomic ratios. The sensitivity of the model prediction to variation of the adjustable parameters allows one to assess the relative importance of various pathways involved in SOA formation. An analysis of SOA formation from the high- and
- ¹⁵ Iow-NO_x photooxidation of four C₁₂ alkanes (n-dodecane, 2-methylundecane, hexylcyclohexane, and cyclododecane) using the FGOM is presented, and comparison with the Statistical Oxidation Model (SOM) of Cappa et al. (2012) is discussed.

1 Introduction

Development of a predictive model for secondary organic aerosol (SOA) formation
 and evolution represents one of the most challenging problems in atmospheric chemistry. In principle, the benchmark for describing SOA formation is a fully explicit chemical model, based on a complete multi-generation gas-phase oxidation mechanism for the parent VOC. A quintessential example is GECKO-A, in which an exhaustive gas-phase VOC oxidation mechanism is generated automatically based on chemical kinetic and mechanistic specifications (Aumont et al., 2005; Camredon et al., 2007; Lee-Taylor et al., 2011; Aumont et al., 2012). Vapor pressures are estimated (based



on group-contribution methods) for each of the thousands of products generated in the mechanism, from which gas-particle partitioning equilibrium constants are computed. The overall yield of SOA for the system as well as its complete chemical make-up are predicted. In such an explicit model, the predictions depend critically on the chemical rules and reaction rate constants imbedded in the mechanism. For the types of highly

- ⁵ rules and reaction rate constants imbedded in the mechanism. For the types of highly functionalized compounds involved in SOA formation, experimentally verified rate constants and mechanisms are generally available for only the first few steps of oxidation. Nonetheless, an explicit model like GECKO-A or MCM 3.1 (Jenkin et al., 1997; Saunders et al., 2003; Bloss et al., 2005) embodies the current state of understanding of
- atmospheric reaction mechanisms. The molecular vapor pressure prediction method used also plays an important role in the overall SOA yield predicted (Valorso et al., 2011), although this is the case for any SOA model. Whereas it is desirable to embody within a model as much basic understanding of the SOA formation and evolution process as possible, it is necessary to keep in mind that an important goal of SOA model development is a computational module that can be included in 3-dimensional
- atmospheric chemical transport models. Consequently, the challenge is to balance the desire for chemical fidelity with the need for computational feasibility.

Aerosol Mass Spectrometer (AMS) measurements of organic aerosols, now a routine component of atmospheric measurements and chamber experiments, enable deriva-

- tion of the atomic O:C and H:C ratios of SOA, from which one can infer its overall oxidation state (Kroll et al., 2011). The volatility of SOA is not generally measured routinely, although the volatility of the organic mixture is an essential property of SOA formation. Any SOA model requires a prediction of aerosol volatility, which is related to the molecular properties of the components comprising the SOA, as embodied, for ex-
- ²⁵ ample, in carbon number and oxidation state. The relationship between oxidation state and volatility is not unique; that is, molecular mixtures with the same overall oxidation state do not necessarily exhibit the same overall volatility.

The traditional approach to representing SOA formation, as embodied in the 2-product model (Odum et al., 1996), is based on fitting an empirical gas-particle



partitioning model to chamber SOA yield data as a function of the mass concentration of absorbing aerosol. The Volatility Basis Set (VBS) (Donahue et al., 2006) is an extension of the 2-product concept, in which VOC oxidation products are assigned to volatility "bins" spanning the range of ambient organic saturation mass concentrations

⁵ (*C*^{*}). As the oxidation products undergo multi-generation reaction (often referred to as "aging"), each generation of reaction products is assigned to a volatility class (Robinson et al., 2007). The extent of gas-particle partitioning associated with each volatility bin depends on both the *C*^{*} of that class and the total mass concentration of the absorbing medium (M_0). The stoichiometric assignments of oxidation products to volatility classes are determined through fitting the model to chamber data.

Compelling evidence exists from field observations that current SOA models consistently underpredict ambient organic aerosol concentration (Tsigaridis and Kanakidou, 2003; Heald et al., 2005, 2006; Volkamer et al., 2006; Jimenez et al., 2009). Such underprediction is almost certainly a result of a number of factors. First, the traditional

SOA models on which these earlier modeling studies are based generally do not account explicitly for degree of oxidation and aging, aspects that are key to prediction of SOA formation. Second, the laboratory chamber data on which current models are based generally do not exhibit the degree of oxidation observed in atmospheric organic aerosol. Finally, the VOC and POA emission inventories that drive atmospheric models in all likelihood understate the fluxes of key SOA-forming species.

Three recent studies have been directed at a next generation of SOA models. Pankow and Barsanti (2009) introduced the "carbon number – polarity grid" (CNPG) framework. They expanded the "2-product" concept to the "np + mP" approach (n products with m possible types of very low volatility compounds) to represent the lumped oxidation and accretion products (Barsanti et al., 2011). Three levels of reactions (the oxidation of parent hydrocarbon and n products, and the formation of m particle-phase compounds) are included to represent the complexity of the mix of condensable primary and secondary compounds as a function of time. Donahue et al. (2011, 2012a, b) developed the two-dimensional Volatility Basis Set (2D-VBS)



employing saturation mass concentration and the degree of oxidation (The average oxidation state of aerosol, OS_C (2×O:C – H:C), can be used interchangeably with the O:C ratio) to describe the coupled aging and phase partitioning of organic aerosol. A point in the 2-D space represents a suite of molecules; the presumption is that an ⁵ ensemble of molecules with the same C^* and OS_C behave similarly in the aggregate. In the 2-D VBS space, if the number of carbon atoms, $n_{\rm C}$, in the carbon backbone does not change during oxidation, products will form along lines of constant $n_{\rm C}$, i.e., diagonal lines extending from the lower right (high C^* vs. low OS_C) to the upper left (low C^* vs. high OS_C). When chamber data are available, the volatility distribution of the products is obtained by SOA mass yield fitting. The products are spread over the 10 OS_C coordinate to stitch together a distribution that matches bulk AMS data and any known products. Cappa and Wilson (2012) formulated the Statistical Oxidation Model (SOM) that describes SOA formation as a statistical evolution in the space of numbers of carbon and oxygen atoms, $n_{\rm C}$ and $n_{\rm O}$, with fitting parameters that govern the probability of fragmentation vs. functionalization, the number of oxygen atoms added per 15 functionalization reaction, and the decrease in vapor pressure accompanying addition of an oxygen atom.

The new generation of models, as exemplified by the 2D-VBS and the SOM, represents SOA formation and evolution in terms of the competition between functionalization and fragmentation, the extent of oxygen atom addition, and the change of volatility. As in the earlier generation of SOA models, predictions can be adjusted to fit laboratory chamber data by determination of the optimal values of a number of parameters in the model. (This is in contrast to the fully explicit chemical model that has, in principle, no adjustable parameters.) The 2D-VBS and SOM do not require an ex-

²⁵ plicit VOC oxidation mechanism, although chemical intuition is required in the choice of the adjustable parameters. We present here a new variation of the SOA model that is based on explicit chemical information in terms of the types of functional groups that result from the oxidation of a parent VOC. We term this the Functional Group Oxidation Model (FGOM). As in the 2D-VBS and the SOM, the FGOM is characterized by



a set of parameters that are determined by analysis of chamber data. Key issues that must be addressed in SOA model development include: (1) in fragmentation, does C-C bond scission occur at any of the C-C bonds in the molecule or does the model specify, for example, that fragmentation always produces C_1 and C_{n-1} ? (2) How important

- are particle-phase reactions to the evolution of SOA, and how is such chemistry represented? (3) To what extent does particle phase chemistry increase the oxidation state of the aerosol and what is its effect on particle volatility? Such issues have important implications for simulations of the overall mass concentration of SOA and its oxidation state.
- We first describe the development of the FGOM, and then apply it to SOA formation from the low- and high- NO_x photooxidation of four C₁₂ alkanes. An important aspect of the present work is a comparison of the FGOM to the SOM, both of which are built on a similar foundation. It turns out that, even with the number of adjustable parameters, complete matching to organic aerosol mass concentration and O:C and H:C ratios over the course of an oxidation is challenging. Assessing the sensitivity of predictions
- to assumptions made in the model development is crucial in interpreting results of its application.

2 Development of a Functional Group Oxidation Model (FGOM)

SOA formation from a parent VOC involves multiple generations of oxidation coupled to ²⁰ gas-particle partitioning. Note that a "generation" is simply the number of times that a molecule has reacted with oxidants (OH, O_3 or NO_3), starting from the parent VOC. In the general process depicted in Fig. 1, S_i^g represents the spectrum of gas-phase products at the *i*-th generation, each product having its own chemical identity and volatility. The multi-generation progression of the chemistry is characterized by a competition ²⁵ between functionalization and fragmentation. Each S_i^g may, in principle, partition to the aerosol phase, S_i^p . Once in the aerosol phase, S_i^p may itself undergo chemical reaction. In Fig. 1, these reactions are assumed to parallel mechanistically those in the



gas-phase, although that assumption need not be invoked. The particle-phase species may be converted to essentially non-volatile compounds, denoted N^p by reactions that can be either non-oxidative or oxidative in nature (Kroll and Seinfeld, 2008). The non-oxidative reactions are typified by the association of two condensed-phase organic ⁵ molecules.

The schematic in Fig. 1 defines the overall structure of the Functional Group Oxidation Model (FGOM). We term the model FGOM, in that a comprehensive chemical mechanism, such as MCM 3.1 or GECKO-A, is used to predict the types of functional groups that characterize the oxidation of the parent organic molecule, yet the detailed mechanism is not employed in the model itself. With the carbon number and the spectrum of functional groups specified, the vapor pressure of the compound defines the relationship between S_i^g and S_i^p . Because the full chemical mechanism of SOA formation is not generally known, predictive SOA models built upon chamber data are characterized by a set of parameters, the values of which are determined by optimal fitting to experimental data. Gas-phase oxidant (e.g. OH, O₃, and NO₃) concentrations need to be specified based on chamber conditions. The quantitative behavior of the FGOM is expressed by a set of kinetic equations for S_i^g , S_i^p , and N^p .

2.1 Functionalization

The essential nature of functionalization is the addition, for a particular parent VOC,

- of a certain number and type of functional groups at each generation of reaction. The mechanism underlying the addition of different functional groups is guided by the specific chemistry. More than one functional group can be added per generation, and the types of functional groups added vary by generation. Reaction rate constants for different functionalization steps are lumped according to the structure of the parent
- ²⁵ VOC molecule and the chemical transformation of different functional groups based on structure-activity relationships (Kwok and Atkinson, 1995). For purpose of discussion, one can consider OH reaction as the oxidative process in SOA formation, although, if appropriate, other oxidants, such as O_3 and NO_3 can be considered as



well. Consideration of the general mechanism of OH oxidation allows one to identify the types of functional groups involved. OH-initiated oxidation leads to a peroxy radical (RO₂), the fate of which controls the SOA-forming mechanism (Kroll and Seinfeld, 2008). At sufficiently low NO_x concentrations, RO₂ self-reactions will dominate the fate of RO₂:

 $RO_2 + HO_2 \rightarrow ROOH + O_2$

 $RO_2 + RO_2 \rightarrow 2RO + O_2$

 \rightarrow RCHO + ROH

 \rightarrow ROOR + O₂

In the presence of sufficient NO_x, RO₂ reacts preferentially with NO and NO₂, e.g.

$$RO_2 + NO \rightarrow RO + NO_2$$

15

10

 $\rightarrow \text{RONO}_2$

Alkoxy radicals (RO) formed either isomerize through a 1,5 H-atom shift to a δ-hydroxyalkyl radical, which introduces an extra OH group, react with O₂ to form carbonyls, or undergo fragmentation. Thus, functional groups introduced from peroxy radical reactions include hydroperoxide (–OOH), nitrate (–ONO₂), hydroxyl (–OH), and carbonyl (–C=O). If the parent hydrocarbon contains a non-aromatic ring, decomposition of the RO will not lead to a decrease of carbon number. The addition of OH introduces an extra hydroxyl group to the parent hydrocarbon. Addition of OH to an aromatic ring is mechanistically different from that to a C=C bond. The progressive photochemical reactions of oxidation products proceeds by: (1) H-atom abstraction from C-H bonds and, to a lesser extent, from the O-H bond of alcohols; (2) H-atom abstraction from C-H bonds of carbonyls; (3) oxidation of –OOH and –ONO₂ groups; and (4) Photolysis of –OOH and –ONO₂ groups.



The addition of various combinations of these four groups via oxidation accounts for a majority of the gas-phase reactions involving semi-volatile product aging. That is, a combination of these four groups is assumed to be a sufficient surrogate for all functional groups in terms of the contribution to vapor pressure, carbon number, and oxidation state. This assumption is based on two observations pertinent to the photooxidation of VOCs, IVOCs, and SVOCs (alkanes, alkenes, terpenes, and aromatics) in the troposphere. First, gas-phase mechanisms involving the addition and transformation of these four functional groups account for much of our current understanding of atmospheric photochemistry leading to the SOA formation. Second, other functional groups can be represented by a combination of these four surrogate functional groups with comparable chemical identities (volatility and oxidation state): the carboxylic acid group (–COOH) can be represented by the sum of a –C=O and a –OH group and the

ether group (-O-) can be replaced by a -C=O group.
Particle-phase oxidative reactions can occur via heterogeneous OH uptake to particles and OH-initiated oxidation in the bulk (e.g., Ravishankara, 1997; Knopf et al., 2005; Zahardis and Petrucci, 2007; Shiraiwa et al., 2011). As noted, we will assume that particle-phase oxidation reactions proceed in parallel with and via the same chemical mechanisms as in the gas phase. The reaction rate of OH oxidation of compound "X" in the particle phase can be expressed as:

$$\frac{d[X]_{p}}{dt} = k'_{p}[X]_{p}$$

20

where k'_p , the pseudo first-order particle-phase reaction rate constant, is related to the corresponding gas-phase reaction by $k'_p = r_p k_g [OH]_g$, and where k_g is the corresponding gas-phase bimolecular reaction rate constant, $[OH]_g$ is the gas-phase OH radical concentration, and r_p is the ratio of the OH reaction rate constant in the particle phase to that in the gas phase. r_p is one of the free parameters of the model. Since par-

to that in the gas phase. r_p is one of the free parameters of the model. Since particle size dependence of the kinetic parameters is not considered in the model, size dependence of r_p is not accounted for.



(1)

2.2 Fragmentation

More heavily oxidized molecules can fragment more easily (Kroll et al., 2009). Fragmentation thus tends to lead to a reduction in the ability to form organic aerosol. The probability of fragmentation f for a stable molecule has been proposed to be represented as a function of the O: C ratio of the molecule (Jimenez et al., 2009; Murphy et

sented as a function of the O: C ratio of the molecule (Jimenez et al., 2009; Murphy et al., 2011, 2012; Cappa and Wilson, 2012; Chacon-Madrid et al., 2012; Donahue et al., 2012b):

 $f = (O:C)^{f_v}$

- where f_v is a parameter to be determined by fitting of data. Fragmentation does not necessarily substantially change the average O:C ratio of SOA; fragmentation leads to products with fewer carbon and, generally, more oxygen atoms, although the ultimate contribution of these products to the SOA mass depends on their volatilities. Species with a higher O:C ratio have a greater probability of fragmentation. In this way, f_v is a key factor that influences the total organic aerosol mass concentration.
- In the SOA model structure, it is necessary to make an assumption concerning the mechanism of fragmentation. In the SOM, the distribution of fragments is computed in two alternative ways: (1) a random probability for the location of the C-C bond scission, thus affecting the resulting size of the scission product; and (2) fragmentation leads to production of only C_1 species, along with the appropriate co-product. Cappa
- and Wilson (2012) found that the "random fragmentation" and the "one carbon loss fragmentation" yield a very similar set of optimal parameters for chamber-generated α -pinene SOA. The 2-D VBS also assumes that the C-C bond scission is randomly distributed along the carbon backbone. The distribution of products over the OS_C – $n_{\rm C}$ space is computed in such a way that the less volatile bins to the left of the midpoint (nominally C) have the same O: C on the productor while the more volatile bins to
- ²⁵ (nominally $C_{n/2}$) have the same O:C as the precursor, while the more volatile bins to the right of the midpoint progress diagonally towards the bin with highest O:C (i.e., $C^* = 10^9 \,\mu g \,m^{-3}$, O:C = 1.0). We invoke here the assumption that fragmentation leads



to the formation of a one-carbon compound (C₁), which is formaldehyde, together with a co-product (C_{*n*-1}) that has the same collection of functional groups but one fewer carbon atom than the parent compound. The "random fragmentation" assumption in SOM acts to rapidly distribute mass over the entire $n_{\rm C}$ vs. $n_{\rm O}$ space. The initial products

- ⁵ upon fragmentation in the 2-D VBS form a "hockey stick" in the O:C vs. *C** space. The "one carbon loss fragmentation" assumption leads to a somewhat slower progression of prediction of smaller fragments. These three different assumptions concerning the mechanism of fragmentation lead eventually to a similar final state characterized by a distribution of intermediates of a more volatile and oxygenated nature.
- ¹⁰ The role of fragmentation is evident in the chemical mechanisms shown in Fig. 2 (see also the Supplement). It is assumed that any of the hydrogen atoms on a secondary carbon is more or less equally vulnerable to OH attack, so instead of the exact structure of a specific molecule, we focus on the type and number of functional groups associated with a carbon backbone. Under low-NO_x conditions, photolysis of –OOH
- ¹⁵ groups and peroxy-peroxy reactions are the primary pathways leading to scission of the carbon skeleton. In the domain of low-NO_x RO₂+ HO₂ dominated chemistry, the photolysis of an –OOH group generates an alkoxy radical, which decomposes to a carbonyl group (formaldehyde in the C₁+C_{*n*-1} case) and an alkyl radical (R). Reaction of R with O₂ followed by reaction with HO₂ regenerates the –OOH group; In RO₂ + RO₂
- ²⁰ dominated chemistry, the self-reaction of RO₂ produces –C=O, –OH, and RO. The decomposition of RO leads to –C=O and R'. R' reacts solely with O₂ to form R'O₂, which can initiate another RO₂ + RO₂ or RO₂ + HO₂ cycle. Under high-NO_x conditions, RO is produced primarily from the photolysis of –ONO₂ and RO₂ + NO reactions. RO decomposes to a carbonyl and an alkyl radical, which reacts with O₂ to reform RO₂. In
- ²⁵ the presence of a sufficiently high concentration of NO_x, RO₂ will preferentially react with NO to form an alkoxy radical or a $-ONO_2$ group again. We shall take $-ONO_2$ as the only functional group formed upon the photolysis of $-ONO_2$ because of its larger contribution to the decrease of vapor pressure than the fragments from the decomposition of an alkoxy radical. We also assume that the OH oxidation of an alcohol produces



solely a carbonyl, so the presence of –OH will not affect the type of functional groups added upon fragmentation.

Volatility estimation 2.3

In fragmentation, if sufficient addition of functional groups takes place, each of the products may still have a sufficiently low vapor pressure to be considered as semi-volatile. 5 A critical feature of the model is a relationship between saturation mass concentration, C^* , and the carbon number of a compound.

From the point of view of predicting physical properties, description in terms of functional groups tends to be more useful than in terms of elemental composition. But, by the same token, the experimentally measureable quantities tend to be best related to 10 elemental composition. In order to reconcile these two approaches to molecular characterization in terms of SOA, it is useful to develop a correlation that relates vapor pressure (or C^*) to both functional groups and elemental amounts. The theory underlying this approach will be developed generally and subsequently applied in analysis of SOA formation from C_{12} alkanes. 15

Frequently used vapor pressure estimation methods are based on expressing the logarithm of vapor pressure as a linear combination of contributions from individual functional groups (Capouet and Müller, 2006; Pankow and Asher, 2008; Camponelle et al., 2011). On the other hand, an approximate linear relationship between C^* and elemental composition, as developed by Donahue et al. (2011), proves to be advanta-20 geous. Here we merge the "group contribution" and "element contribution" methods in order to relate the carbon number to a given saturation concentration of a compound *i*. The resulting expression is

$$\log_{10} \frac{10^6 p_{\mathrm{L},i}^0 (12n_{\mathrm{C}}^i + hn_{\mathrm{C}}^i + 16n_{\mathrm{O}}^i + 14n_{\mathrm{N}}^i)}{RT} = \left(n_{\mathrm{C}}^0 - n_{\mathrm{C}}^i\right) b_{\mathrm{C}} - n_{\mathrm{O}}^i b_{\mathrm{O}} - 2\frac{n_{\mathrm{C}}^i n_{\mathrm{O}}^i}{n_{\mathrm{C}}^i + n_{\mathrm{O}}^i} b_{\mathrm{CO}} - \sum n_{\mathrm{func}}^i b_{\mathrm{func}}$$
(2)

where $b_{\rm C}$ is the carbon-carbon interaction term, $b_{\rm O}$ is the oxygen-oxygen interaction 25 term, b_{CO} is the carbon-oxygen interaction term, b_{func} is the interaction term for different 32576

Discussion Paper ACPD 12, 32565-32611, 2012 **A Functional Group Oxidation Model Discussion** Paper X. Zhang and J. H. Seinfeld **Title Page** Introduction Abstract Conclusions References **Discussion** Paper **Tables Figures** Back **Discussion** Paper Full Screen / Esc **Printer-friendly Version** Interactive Discussion

functional groups, $n_{\rm C}^0$ is the reference carbon number, $n_{\rm C}^i$ is the carbon number of compound *i*, $n_{\rm O}^i$ is the oxygen number of compound *i*, n_N^i is the nitrogen number of compound *i*, n_N^i is the number of different functional groups in compound *i*, and *h* is the ratio of hydrogen number to carbon number of compound *i*. For the application to

- the alkane system to be presented, we have determined $n_{\rm C}^0$, $b_{\rm C}$, $b_{\rm O}$, $b_{\rm CO}$, and $b_{\rm func}$ by fitting the above equation to the estimated vapor pressures of 1080 standard alkane compounds (see Supplement). After optimal fitting of Eq. (2) to the standard compound vapor pressures, the predicted carbon number of a compound lies within the range of 10% uncertainty of its actual value.
- ¹⁰ To illustrate the application of the volatility-composition correlation, consider the mechanism of the OH-oxidation of dodecane ($C_{12}H_{26}$). For a typical ambient organic aerosol loading of several μ g m⁻³, let us assume for purpose of illustration that any compound with $C^* < 10^3 \mu$ g m⁻³ can partition appreciably to the particle phase. The grey shaded region in Fig. 3 defines the particle phase "boundary" for compounds having each an appreciable to appreciable the section of the particle phase.
- ¹⁵ ing carbon numbers 12 and fewer. The lines define the carbon number range and types of functional groups for oxidation products of dodecane under low-NO_x conditions. The magenta line in the lower left hand corner of the grey shaded region, for example, indicates that a product with two –OOH groups and 12 to 7 carbon atoms has sufficiently low volatility to lie in the semi-volatile region defined by $C^* < 10^3 \,\mu g \,m^{-3}$. However, the
- ²⁰ cyan line in the upper right hand corner of the "condensable region" indicates that a molecule with only one -C=O group would need at least 17 carbon atoms to fall into the semi-volatile region. This would necessitate some types of accretion reaction to occur given the C₁₂ parent molecule. The trade-off between functionalization and fragmentation in the volatility of the products becomes evident in such a graphic representation.

25 2.4 Gas-particle partitioning

Gas-particle equilibrium partitioning of semi-volatile products is assumed. From Raoult's law, $p_i = p_{L,i}^0 \gamma_i \chi_i$, where $p_{L,i}^0$ (atm) is the vapor pressure of semi-volatile



compound *i* at the given temperature, γ_i is a molality-based activity coefficient, and γ_i is the mole fraction of compound *i* in the particle phase. The gas-particle partition coefficient $K_{\text{p},i}$ (m³ mol⁻¹) is expressed as:

$$K_{\mathrm{p},i} = \frac{S_i^{\mathrm{p}}/S_T^{\mathrm{p}}}{S_i^{\mathrm{g}}} = \frac{RT}{p_{\mathrm{L},i}^{\mathrm{0}}\gamma_i}$$

⁵ where S_i^g is the molar concentration (mol m⁻³) of *i* in the gas phase; S_i^p is the molar concentration (mol m⁻³) of *i* in the particle phase; S_T^p is the molar concentration per unit volume of air (mol m⁻³) of all the species in the particle phase. *R* (m³ atm mol⁻¹ K⁻¹) is the gas constant, and *T* (K) is temperature. The saturation concentration C_i^* (g m⁻³) is expressed as:

10
$$C_{i}^{*} = \frac{C_{i}^{g}C_{T}^{p}}{C_{i}^{p}} = \frac{p_{L,i}^{0}\gamma_{i}M_{wi}}{RT}$$

20

where C_i^g is the mass concentration (gm^{-3}) of *i* in the gas phase; C_i^p is the mass concentration (gm^{-3}) of *i* in the particle phase; C_T^p is the total organic aerosol concentration (gm^{-3}) . M_{wi} is the molecular weight of *i*. Thus the fraction of *i* in the particle phase, ϕ_i^p is:

15
$$\phi_{i}^{p} = \frac{1}{1 + 1/(K_{p,i}S_{T}^{P})} = \frac{1}{1 + C_{i}^{*}/C_{T}^{P}}$$

Note that the molality-based activity coefficient is assumed to be unity here. Non-ideal interactions between organic and inorganic species in the aerosol phase influence the gas/particle partitioning of semi-volatiles. Thermodynamic models show that the saturation concentration of species *i*, C_i^p , is usually underpredicted by assuming ideal condensed phase behavior, and the extent of underprediction differs by functional groups

(3)

(4)

(5)

and is affected by aerosol water content, as well as temperature and SOA loading (Zuend et al., 2010, 2012). Inclusion of particle-phase nonideality and phase separation is beyond the scope of the current study.

2.5 Particle-phase accretion reactions

- ⁵ Recent modeling studies have suggested that dehydration reactions in the particle phase can play an important role in determining the characteristics of the oxidation state (Chen et al., 2011; Cappa and Wilson, 2012), although the detailed mechanisms are uncertain. In the present FGOM framework, particle-phase accretion reactions (non-oxidative) are allowed to occur and assumed to generate essentially non-volatile
 ¹⁰ products. These reactions are treated as bimolecular, with a reaction rate constant k_a (m³ mol⁻¹ s⁻¹), which is a fitting parameter of the model. We assume that all of these non-volatile compounds can be represented as a single generalized species, which is
- characterized by its time-dependent molar amount and the set $[C_x, H_y, O_z]$, where the parameters *x*, *y*, and *z* represent the numbers of carbon, hydrogen, and oxygen atoms ¹⁵ characterizing the non-volatile particle-phase composition. The parameters *x*, *y*, *z* are tunable parameters of the model. Note that certain limits are set for mathematically optimizing the *x* value in order to conserve the carbon mass: the upper limit of the *x* value is the sum of carbon numbers of two molecules with the parent carbon skeleton and the *x* value can not be lower than the sum of carbon numbers of two molecules that ²⁰ have the lowest carbon numbers in the category of semi-volatile compounds defined by Eq. (2). The oxygen and hydrogen masses are not conserved in the specification
 - of y and z, in view of the formation of water molecules during the oligomerization processes.

2.6 Summary of the Functional Group Oxidation Model (FGOM)

²⁵ The structure of the FGOM is summarized as follows. The type and number of functional groups added to the parent molecule over multi-generation oxidation are



specified from consideration of an explicit chemical mechanism, such as GECKO-A or MCM (MCM is employed in the current FGOM model structure). As functionalization proceeds, a spectrum of products that are characterized by various combinations of functional groups together with the parent carbon backbone is generated. The accom-

- ⁵ panying fragmentation leads to products with the same functional groups as the parent product but with fewer C atoms. The automatically generated fragments are subject to the same oxidation mechanisms as the parent products. Vapor pressures of oxidation products are estimated by the vapor prediction model "EVAPORATION" (Camponelle et al., 2011). Products formed partition to the particle phase, where they may either
- ¹⁰ be oxidized following the same gas-phase mechanism, or undergo accretion reactions. The model contains six tunable parameters: r_p , the ratio of the particle-phase oxidative reaction rate constant to that in the gas phase; f_v , the parameter characterizing the probability of fragmentation; k_a , the accretion reaction rate constant in the particle phase; and a matrix $[C_x, H_y, O_z]$ of the carbon, hydrogen, and oxygen numbers of the per valetile products.
- the non-volatile products. These six parameters are incorporated in a set of differential equations governing the kinetics of all the species in the mechanism. The differential equations are solved numerically, with instantaneous equilibrium partitioning calculated at each time step. The number of differential equations for a certain VOC system depends on the chemical nature of the parent hydrocarbon (e.g., structure, volatility,
- etc.) and the complexity of the functionalization mechanism. Evolution of these six parameters describes the gas- and particle-phase chemistry, together with the instantaneous equilibrium partitioning process. Their values are determined by optimal fitting of the model simulations of organic mass growth and elemental compositions with laboratory chamber data (organic aerosol mass, O:C, and H:C). These three proper-
- ties are equally weighted to constitute the overall optimization objective function. The Levenberg-Marquardt method implemented in MATLAB's "System Identification Toolbox" (MathWorks, 2002) is used for the nonlinear minimization. The essential feature of the FGOM model that distinguishes it from explicit chemical models is the simplified treatment of fragmentation processes: the fragmentation operator (in the form of



 f_v) is applied to determine the probability that a given stable molecule will fragment, instead of accounting for the detailed free radical chemistry. Owing to the consideration of carbon-number conserved oxidation reactions based on the explicit chemistry, the FGOM model differs from the 2-D VBS or SOM models in terms of the treatment of particle-phase oxidative and non-oxidative reactions through the parameters r_p , k_a , and $[C_x, H_v, O_z]$.

3 Application to SOA formation for C₁₂ alkanes

The Functional Group Oxidation Model is applied to predict SOA formation from the OH-initiated oxidation of four C₁₂ alkanes (dodecane, methylundecane, hexylcyclohex ane, and cyclododecane) under both high- and low-NO_x conditions. Experiments were carried out in the dual 28 m³ Teflon reactors in the Caltech Environmental Chamber. Detailed descriptions of experimental protocols can be found elsewhere (Yee et al., 2012; Craven et al., 2012). Experimental conditions are summarized in Table 1. Different OH radical sources were employed to achieve high- and low-NO_x levels. For low-NO_x conditions, experiments were designed to capture 36 h of total photooxida-

- ¹⁵ low-NO_x conditions, experiments were designed to capture 36 h of total photooxidation during which the OH concentration produced from the photolysis of H_2O_2 remains essentially constant. Under high-NO_x conditions, the OH source is HONO. The photolysis rate of HONO is about an order of magnitude higher than that of H_2O_2 , leading to a depletion of OH within ~ 3 h. For both cases, the time-dependent OH concentration
- ²⁰ can be obtained by optimal fitting of the predicted parent hydrocarbon decay with its measured temporal profile. Three measured SOA properties, i.e., organic mass growth, O:C ratio, and H:C ratio, are employed to constrain the performance of FGOM. The best-fit parameters obtained for the four C₁₂ alkanes for both high-NO_x and low-NO_x conditions are given in Table 2.
- Questions associated with the application of FGOM are: (1) How well does the model reproduce the measured SOA properties? (2) Are all six parameters equally influential in terms of representing SOA formation and evolution? and (3) How well do these



mathematically best-fit parameters express the actual chemistry thought to be occurring in the system? Some general findings include: f_{y} values are generally lower under high-NO_x conditions, indicating more intense fragmentation. The abundance of RO is the most important determinant of the degree of fragmentation. In the presence of

- NO_{v} , the primary fate of RO_{2} is reaction with NO, producing RO. All of the low- NO_{v} 5 experiments fall into the HO₂ + RO₂ dominant regimes, as > 95 % RO₂ will react with HO₂. Although the photolysis of the -OOH group formed from HO₂ + RO₂ reaction leads to the formation of RO, the yield of RO via this path is still an order of magnitude lower than under high-NO_x conditions. In addition, f_{i} values for cycloalkanes are higher than those for straight-chain alkanes, which is consistent with the mechanism of the decomposition of RO in a non-aromatic ring.
- 10

Optimal $r_{\rm p}$ values for the eight alkane systems range from 10^{-3} to 10^{-1} , that is, the particle-phase oxidation reaction rates are predicted to be at least one order of magnitude slower than those in the gas phase. The role of particle-phase oxidative reactions in the SOA formation from C₁₂ alkanes is, therefore, not strongly influential 15 during the initial period of time. However, for species with very low vapor pressure (e.g., 1×10^{-11} atm), 99.8% of which will end up being in the particle phase at a moderate organic aerosol loading, e.g., 40 µg m⁻³, particle-phase oxidative reactions are as in-

- fluential as those in the gas phase even if r_p is of the order of 10^{-3} . Most semi-volatile products generated from the functionalization channel of the FGOM have vapor pres-20 sure higher than 10⁻¹⁰ atm, and as a result, their particle-phase oxidative reactions are predicted to be a minor path when $r_{\rm p}$ is less than 10^{-1} , a prediction that is consistent with field and experimental evidence (Murphy et al., 2007; DeCarlo et al., 2008; George et al., 2008). Optimal k_a values for the eight alkane systems range from 10^3
- to $10^4 \text{ m}^3 \text{ mol}^{-1} \text{ s}^{-1}$, indicating that accretion reactions occurring in the particle phase 25 are predicted to play an important role in SOA properties, given the high local concentrations of the species in the condensed phase. Take the low-NO_v photooxidation of dodecane as an example; the predicted non-volatile species eventually account for ~ 30 % of the total organic aerosol, see Fig. 9. Optimal values of the matrix $[C_x, H_v, O_z]$



for the eight systems indicate a highly dehydrated nature of the non-volatile species. Since the explicit particle-phase reaction mechanisms in these systems are not yet clear, one cannot assess unambiguously the effect of particle-phase reactions on the SOA aging. We will return to the importance of the five particle-phase parameters in Sect. 3.3.

3.1 Organic aerosol growth

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The predicted (FGOM) temporal profiles of organic aerosol mass (C_{OA}), together with the wall-loss-corrected chamber data, for the four C₁₂ alkane systems based on the optimal fit parameters in Table 2 are shown in Fig. 4. SOA yields under high-NO_x conditions are lower than those under low-NO_x conditions because of both the shorter OH radical exposure time and the more intense fragmentation in the presence of NO_x. The SOA growth curves under various NO_x levels show different patterns. Particles start to grow immediately after the initiation of photochemistry under high-NO_x conditions, whereas a short lag occurs before organic aerosol mass accumulates under low-NO_x. The initial OH concentration under high-NO_x conditions is an order of magnitude higher than that in low-NO_x, which substantially accelerates the progression of multi-generation oxidation and consequently the early stage formation of low-volatility compounds. In addition, the RO chemistry occurring under high-NO_x conditions leads to products spanning a wider range of volatilities after only one generation of oxidation.

- The level off of C_{OA} at the end of the experiments is generally not captured for all the eight alkane systems. The level off could have resulted from re-evaporation of particles under low-NO_x conditions and the cessation of OH production under high-NO_x conditions. Reproduction of C_{OA} requires accurate representation of the trade-off between the fragmentation of semi-volatile products and the formation of particle-phase nonvolatile products, which results in an accumulation of total particle organic mass. The
- C_{OA} growth curves for cyclododecane under high-NO_x conditions are not as closely simulated as for the other alkanes. For cyclododecane, particles are observed to grow fairly rapidly after the inception of photooxidation.



An essential feature of the FGOM is the prediction of the emergence of a spectrum of functional groups as the oxidation proceeds. For the alkanes investigated here, the C₁₂ backbone is not predicted to break until 3–5 functional groups have been added to the parent chain. As a result, 7-8 generations of oxidation are predicted to occur until the last remaining C_{12} backbone is fragmented. In the FGOM model, every possi-5 ble oxidation step leading to functionalization is implicitly considered in the generalized mechanism. The extent to which aging proceeds then depends on the availability of oxidants. Figure 5 shows the predicted contribution of different generations of products to the total organic aerosol mass for the eight alkane systems. The OH exposure time under low-NO_v conditions is, on average, ~ 3 times higher than that under high-NO_v 10 conditions. As a result, the 2nd and 3rd generation products are dominant components at high-NOx, whereas 4th and later generation products account for over 50 % of the total aerosol mass at the end of the low-NO_x experiments. The temporal distribution of multi-generational products reflects different chemical mechanisms governing the photochemistry under high- vs. low-NO_x conditions. Due to the importance of RO 15 chemistry in the presence of NO_x , a spectrum of products spanning a large range of volatility and oxidation states are formed upon the OH initiated photooxidation. Under low-NO_x conditions, products containing the –OOH group dominate the particle phase due to their low volatilities, although most of these are consumed rapidly in the gas

²⁰ phase. In general, the extent of agreement between predicted and observed chamber data for the initial SOA growth depends on the matrix of functional groups generated as oxidation proceeds. For example, GECKO-A generates 2% trifunctional species in the first-generation products from C₁₆ alkanes (Aumont et al., 2012), products that are not included in MCM 3.1. Because of their low volatility, a small amount of trifunctional species could contribute substantially to the particle mass.

3.2 Elemental aerosol composition

Figure 6 shows the measured and simulated particle average O:C ratios for the OH oxidation of dodecane, methylundecane, hexylcyclohexane, and cyclododecane. The



experiments simulated here result in moderate levels of oxidation, i.e. O: C ~ 0.3 in the semi-volatile oxygenated organic aerosol (SV-OOA) range (Ng et al., 2010, 2011). For most experiments, O: C spans the range from 0 to 0.5. Note that in analyzing chamber results, O: C begins at zero and then rapidly increases as the early low volatility compounds condense. In this region, the total organic aerosol amount (C_{OA}) is small and in a range where AMS measurements are less reliable. As C_{OA} increases, species with

- higher volatilities are capable of partitioning to the particle phase, causing O: C to decrease. The O: C trend thus represents a competition between the decrease due to the partitioning of higher volatility compounds with fewer O-atoms and the increase that re-
- ¹⁰ sults from partitioning of more oxygenated compounds as multiple generations unfold. Owing to less data reliable at the onset of AMS sampling, we consider the data points in the middle of the span, i.e. $O: C \sim 0.2-0.3$, as the "standard" for the comparison with simulations from "sim.1". Predictions at high-NO_x conditions are generally consistent with the measured ratio. For low-NO_x conditions, predictions are in agreement
- ¹⁵ with measured O: C within 30 % uncertainty at the beginning and match the measured O: C after ~ 20 h reaction. In general, the O: C ratios under high-NO_x conditions are predicted to be higher than under low-NO_x conditions because the formation of $-ONO_2$ groups introduces more oxygen to the particles than the -OOH group. No increase in O: C is observed under high-NO_x conditions because of the experimental depletion of
- OH radicals after ~ 3 h reaction. As noted, the FGOM does not reproduce the observed O:C data for the cyclododecane case. The initial organic mass growth observed from the OH oxidation of cyclododecane is consistent with rapid functionalization. Based on gas-phase chemical mechanisms, cycloalkanes should become more oxygenated as the breakage of the carbon ring introduces more than one extra O atom, compared
- with straight-chain alkanes. However, the measured O:C for the cyclododecane low-NO_x case is the lowest among the four alkanes. The model can replicate the measured ratio only if the average oxygen number of non-volatile compounds is decreased to 3. The model prediction suggests a more effective dehydration process occurring in cyclododecane than other systems, as long as the AMS measurements are comparable



for the 8 experiments and the matrix of functional groups is applicable for the four C_{12} alkanes.

- Measured and predicted H: C ratios of the SOA are shown in Fig. 7. The simulated H:C ratios match the observed after several hours of reaction for each experiment. However, the starting points varied by 15%: predicted H:C ratio starts from 2, de-5 creasing to ~ 1.7 along with the oxidation, whereas AMS measurements show that the H: C ratio is roughly constant, with an average ranging from 1.6 to 1.7, over the course of each experiment. The functional group contributing the highest decrease in H:C ratio is -C=0. On the basis of the gas-phase chemistry, at least 3 -C=0 additions to the parent hydrocarbon are needed to reach the observed H: C ratio. One generation 10 of oxidation can introduce at most one -C=O to the SOA, as the functional groups that are most likely to induce partitioning into the particle phase are -OOH and -ONO₂,
- resulting in a higher contribution to the decrease of vapor pressure than -C=O. This means that > 3 generations of oxidation are needed to produce the "right" molecule.
- Note that certain combinations of functional groups, like –OOH and –OH, that have 15 sufficiently low vapor pressure to be incorporated in the particles, do not change the H: C ratio. Therefore, products with H: C ratio even lower than 1.7 generated via either gas- or particle- phase chemistry could exist. While we have not explicitly considered the -C(O)OH group, this group can be represented by the sum of -C=O and -OH.
- In addition, -COOH groups can be produced only via fragmentation, which might be 20 accompanied by an increase of vapor pressure. Thus, the contribution of -COOH to the chemical composition of alkane SOA is deemed less important as those of other functional groups. Dehydration processes in the particle phase can play an important role in obtaining the desired H: C ratio, but our simulations suggest that this process is likely not fast enough to drive the H : C ratio immediately to \sim 1.7.
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3.3 Role of particle-phase chemistry

Experimental evidence indicates the potential importance of oligomerization processes in the SOA formation (Gao et al., 2004; Kalberer et al., 2004; Surratt et al., 2007; Chan



et al., 2011). To assess the role of non-oxidative particle-phase reactions on SOA formation from C_{12} alkanes, one may shut off the accretion reaction channel ($k_a = 0$) and fit only f_v and r_p to the SOA data. Although under this condition the model is capable of reproducing the observed organic aerosol growth curves, O:C and H:C ratios are overpredicted by 35% and 30%, respectively. Progressive oxidation chemistry is not capable of explaining the measured O:C and H:C ratios. Thus, gas-phase oxidation alone, coupled to gas-particle partitioning, does not lead to condensed phase OA with sufficiently low volatility and high oxidation state. One concludes that the nonoxidative chemistry afforded by the accretion reaction channel is required to simulate the observed oxidation state of the alkane SOA by producing essentially non-volatile

compounds with a highly dehydrated nature.

Particle-phase oxidative reactions constitute an additional channel in the "VOC to SOA pump". As noted, it is assumed that these reactions parallel those in the gas phase in terms of the functionalization and fragmentation pathways available, but at a rate

- ¹⁵ that is a factor of r_p (< 1) of the gas-phase rate. Fragmentation at each generation both in the gas and particle phases leads to a systematic cleavage of the original carbon backbone. If this cleavage is not compensated for by additional functionalization, the molecules produced will have higher volatility than their precursors, and a decrease in overall organic aerosol mass occurs regardless of whether the pathway occurs in the
- 20 gas or particle phase. The existence of the particle-phase oxidation channel acts to accelerate the overall SOA evolution and therefore can lead to a more rapid aging than in its absence.

Figure 8 shows the simulated organic aerosol growth from the photooxidation of dodecane under low-NO_x conditions when $r_p = 0$ or 1, with other parameters held at their ²⁵ best-fit values. When $r_p = 0$, the predicted organic aerosol growth is able to match the measured C_{OA} because the r_p value for low-NO_x dodecane obtained from the optimization is 10^{-3} , implying a negligible particle-phase oxidation pathway. When $r_p = 1$, the predicted C_{OA} is less than 20% of that measured in the chamber after 35 h of reaction. Although the incorporation of the accretion channel increases the predicted



 C_{OA} , this increase is insignificant in the presence of highly effective particle-phase oxidation reactions ($r_p = 1$). Obviously, the " $r_p = 1$ " simulation is an unrealistic case such that two parallel oxidation pathways with the same set of reaction rate constants exist in the gas and particle phases during SOA aging. The overall gas- and particle- phase rates of functionalization and fragmentation processes are accelerated in such a way that the decrease of volatility from the addition of functional groups cannot compensate for the increase of volatility due to the cleavage of the carbon backbone. As a result, the semi-volatile products are mostly consumed during the rapid aging process to produce intermediates with a higher volatility circumventing the accumulation of the organic aerosol.

4 Comparison of FGOM with SOM

As noted earlier, the two SOA models, 2-D VBS and SOM, employ distinct, but related, frameworks to characterize SOA properties, that is, saturation concentration vs. average oxidation state for 2-D VBS and carbon number vs. oxygen number for SOM. Both ¹⁵ models represent functionalization and fragmentation, and each treats the fragmentation process with a tunable parameter "*P*_f", that represents the probability of fragmentation. The molar yields of smaller molecules from fragmentation processes are assumed to be evenly distributed in 2-D VBS but to be randomly assigned or with only one carbon atom loss in SOM. For functionalization, 2-D VBS assumes that each generation of ²⁰ oxidation reduces the saturation concentration between one and six decades and adds between one and three oxygen atoms. In the SOM, the extent of decrease of vapor pressure and the average number of oxygen atoms added per generation of oxidation are tunable parameters (see Table 3).

The 2-D VBS is built on the matrix of C^* vs. OS_C , which reflects the average car-²⁵ bon and oxygen numbers based on the relationship between C^* and the elemental composition developed by Donahue et al. (2011). The evolution of a certain "bin" on the 2D-VBS matrix upon photooxidation is governed by the two competing kernels: Discussion Paper ACPD 12, 32565-32611, 2012 **A Functional Group Oxidation Model Discussion** Paper X. Zhang and J. H. Seinfeld **Title Page** Introduction Abstract Conclusions References Discussion Paper **Figures** Back **Discussion** Paper Full Screen / Esc **Printer-friendly Version** Interactive Discussion

functionalization and fragmentation. An important constraint on these two kernels is the conservation of carbon mass: functionalization leads to no change of carbon mass but increase of oxygen mass; and fragmentation leads to products that have the same total carbon mass as the parent compound. In this manner, one needs to define two

- ⁵ generic functionalization and fragmentation kernels based on best understanding of the chemistry. For example, if a reaction adds 3 -OH groups to a carbon backbone, C^* will decrease substantially and OS_C rises modestly, so the products end up in a bin that has a nominal n_C a bit larger than the average n_C of the reactant bin. As a result, a number of free parameters representing the stoichiometric amount of product bins
- ¹⁰ need to be assigned manually after each generation of oxidation. The only actual fitting parameter in 2-D VBS is thus the functionalization vs. fragmentation branching, which is represented as " $P_{\rm f}$ ". Since the 2-D VBS and SOM share similar principles in describing the evolution of SOA, we employ SOM here to compare with the performance of FGOM for the low-NO_x dodecane case.
- ¹⁵ Both FGOM and SOM attempt to reproduce the observed chamber generated SOA properties (organic aerosol mass, O:C and H:C ratio) through optimal fitting of the empirical parameters in each model, e.g. Figs. 4, 6, and 7 here and Figs. 1 and 3 in Cappa et al. (2012). Note that in SOM, the van Krevelen slope is used to determine the H:C that corresponds to the simulated O:C, as opposed to trying to match the
- ²⁰ absolute H: C values. Necessarily, the aerosol chemical composition generated by the two models is different. In the SOM, the individual species are characterized in a statistical manner: by the average carbon and oxygen number ($n_{\rm O}/n_{\rm C}$ pairs), pairs that do not necessarily correspond to actual compounds since functional group information is not explicitly simulated. The FGOM automatically generates all possible compounds
- with the combination of four different functional groups and certain carbon numbers, as suggested by the generalized gas-phase chemistry. One can map the simulation results of these two models onto the $OS_C n_C$ space (Fig. 9). Atmospheric aging leads to increased functionality on the C_{12} skeleton and thus higher average oxidation state, which is reflected by both models. The SOM predicts $OS_C (-1.1 \sim -0.6)$ vs. FGOM



 $(-1.8 \sim -1.0)$. The AMS measured OS_C (= 2 × O:C - H:C) of dodecane SOA is bounded by these two predictions, ranging from -1.2 to -1. The underestimation of OS_C by FGOM at the early stage is largely due to the starting point of the H:C ratio, as discussed in Sect. 3.2. The SOM-generated OS_C is shaped primarily by the simulated

⁵ O:C ratio, since the adjustable parameter "H-atom loss per O-atom added" is determined as a constant. The overprediction of OS_C thus results from the overprediction of the O:C ratio at the end of the simulation.

The average carbon numbers predicted by these two models cover distinctly different ranges. The temporal profile of aerosol $n_{\rm C}$ predicted by FGOM decreases initially

- to ~ 10.6 and then increases to ~ 15.2 at the end of the experiment. In the FGOM, two principal pathways govern the dynamics of carbon number: C-C bond cleavage and particle-phase accretion reactions. The degree of C-C bond cleavage determines the lower limit of the average SOA carbon number, whereas the extent of accretion reactions sets the carbon number upper limit. The optimal FGOM fitting results suggest
- ¹⁵ that oligomerization in the particle phase is required to predict the observed oxidative aging of dodecane SOA, in particular the evolution of the O:C and H:C. In the timedependent trend of $n_{\rm C}$ generated by SOM, the progressive addition of oxygen atoms along with gas-phase aging allow the formation of small highly oxygenated fragments that have sufficiently low volatilities to contribute appreciably to the particle phase.
- To visualize the differences between these two approaches that more or less equally capture the chamber-generated SOA properties, the simulated SOA chemical compositions are shown in the log C^* vs. OS_C space in Fig. 9. The SOM simulated composition exhibits saturation concentrations of monomers that are, on average, one order of magnitude lower than those in FGOM. Values of the best-fit parameters in SOM
- for the low-NO_x dodecane case can be found in Table 2 in Cappa et al. (2012). The decrease of vapor pressure upon the addition of one oxygen on a log scale (Δ LVP), 2.20, corresponds to the addition of an –OH group, which leads to the largest decrease of vapor pressure per oxygen atom according to the vapor pressure prediction models "SIMPOL.1" and "EVAPORATION". This Δ LVP value suggests that in order



to reproduce the observed concentration of chamber-generated organic aerosol, the gas-phase chemistry should proceed such that the addition of any functional group on the carbon chain results in a decrease of the logarithm of vapor pressure by 2.20 per oxygen atom. The vapor pressure estimation module developed for use in FGOM for the alkanes constrains the Δ LVP values per O-atom addition to range from 0.79 to 1.97 upon the addition of the array of functional groups (the predicted decreases of the logarithm of vapor pressure (atm) upon the addition of –OOH, –ONO₂, –OH, and –C=O are 2.83, 2.36, 1.97, and 1.20, respectively). This is confirmed as the probability of fragmentation in SOM ($m_{\text{frag}} = 0.077$) is higher than that in FGOM ($f_v = 0.61$ with accretion channel and $f_v = 0.77$ without accretion channel).

The average addition of O-atoms per generation of oxidation in SOM is 1.22, which, together with the Δ LVP value, is consistent with the addition of one –OH group. The addition of an –OH group corresponds to a van Krevelen slope of 0, as compared with either the observed H:C vs. O:C slope, –0.82, or the best-fit value of H-atom loss per

- O-atom added, 1.77 (see Fig. 10). Neither the best-fit parameters in the SOM nor those in the FGOM can reproduce the observed O:C or H:C ratios without a dehydration reaction channel. This addresses a fundamental question: to what extent can gasphase chemistry alone account for the SOA formation? Several products generated by the gas-phase mechanism in FGOM have been detected in the chamber dodecane low-
- NO_x experiments (Yee et al., 2012) and these products have sufficiently low volatility to form a substantial amount of organic aerosol mass. Nonetheless, the inescapable conclusion is that the occurrence of accretion reactions is necessary to reconcile both the disagreement between the indications of the best-fit parameters in SOM and the inability of the FGOM to reproduce the measured elemental composition based solely
- ²⁵ on gas-phase chemistry. Evaluating the extent to which oligomer formation is important for virtually all SOA parent VOCs awaits a more complete understanding of gas-phase chemistry and measured SOA chemical properties.



5 Conclusions

A Functional Group Oxidation Model (FGOM) is developed that is a compact representation of key processes, such as functionalization, fragmentation, and oligomerization, in the formation and evolution of SOA. A central component of the FGOM is a statistical

- fit of a large number of estimated compound vapor pressure relating the carbon number to the saturation concentration of a compound. The temporal profiles of organic aerosol growth and elemental composition (O: C and H: C) generated from OH-initiated chamber oxidation of four C₁₂ alkanes have been employed to evaluate the performance of this model. In general, the FGOM satisfactorily reproduces characteristic features of
- ¹⁰ SOA formation through optimal fitting of 6 tunable parameters (r_p , f_v , k_a , and [C_x , H_y , O_z]). Analysis of model output also reveals some general trends important in chamber SOA studies. Organic aerosol yield depends most strongly on three variables: the probability of fragmentation to produce volatile compounds, oxidation reactions involving semi-volatile compounds in the particle phase, and irreversible particle-phase accretion
- reactions. The probability of fragmentation largely determines the overall concentration of organic aerosol mass formed, whereas the latter two parameters govern the shape of the particle growth curve. The elemental composition of particles, O:C and H:C, is dependent not only on the functionalization in the gas phase, but also, probably to a larger extent, on potential particle-phase chemistry. Comparison between the FGOM
- and SOM models reveals the potential importance of dehydration processes in shaping the aerosol chemical composition. The statistical fit of compound vapor pressure is developed here using 1080 alkane molecules. This correlation can be extended to other classes of parent VOCs in SOA formation. The FGOM model can be applied to other VOC systems, such as alkenes, aromatics and terpenes, in such a way that the number
- of semi-volatile products that are characterized by the combination of relevant carbon numbers and different functional groups are included in this correlation. The role of gas-phase vs. particle-phase chemistry can be evaluated based on the indications of best-fit parameters obtained by the optimization computation.



Supplementary material related to this article is available online at: http://www.atmos-chem-phys-discuss.net/12/32565/2012/ acpd-12-32565-2012-supplement.pdf.

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References

15

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- Aumont, B., Szopa, S., and Madronich, S.: Modelling the evolution of organic carbon during its gas-phase tropospheric oxidation: development of an explicit model based on a self generating approach, Atmos. Chem. Phys., 5, 2497–2517, doi:10.5194/acp-5-2497-2005, 2005.
- Aumont, B., Valorso, R., Mouchel-Vallon, C., Camredon, M., Lee-Taylor, J., and Madronich, S.: Modeling SOA formation from the oxidation of intermediate volatility n-alkanes, Atmos. Chem. Phys., 12, 7577–7589, doi:10.5194/acp-12-7577-2012, 2012.
 - Barsanti, K. C., Smith, J. N., and Pankow, J. F.: Application of the np+mP modeling approach for simulating secondary organic particulate matter formation from *α*-pinene oxidation, Atmos. Environ., 45, 6812–6819, 2011.
 - Bloss, C., Wagner, V., Jenkin, M. E., Volkamer, R., Bloss, W. J., Lee, J. D., Heard, D. E., Wirtz, K., Martin-Reviejo, M., Rea, G., Wenger, J. C., and Pilling, M. J.: Development of a detailed chemical mechanism (MCMv3.1) for the atmospheric oxidation of aromatic hydrocarbons, Atmos. Chem. Phys., 5, 641–664, doi:10.5194/acp-5-641-2005, 2005.
- ²⁰ Camredon, M., Aumont, B., Lee-Taylor, J., and Madronich, S.: The SOA/VOC/NOx system: an explicit model of secondary organic aerosol formation, Atmos. Chem. Phys., 7, 5599–5610, doi:10.5194/acp-7-5599-2007, 2007.
 - Capouet, M. and Müller, J.-F.: A group contribution method for estimating the vapour pressures of a-pinene oxidation products, Atmos. Chem. Phys., 6, 1455–1467, doi:10.5194/acp-6-1455-2006, 2006.
 - Cappa, C. D. and Wilson, K. R.: Multi-generation gas-phase oxidation, equilibrium partitioning, and the formation and evolution of secondary organic aerosol, Atmos. Chem. Phys., 12, 9505–9528, doi:10.5194/acp-12-9505-2012, 2012.



- Cappa, C. D., Zhang, X., Loza, C. L., Craven, J. S., Yee, L. D., and Seinfeld, J. H.: Application of the Statistical Oxidation Model (SOM) to secondary organic aerosol formation from photooxidation of C₁₂ Alkanes, Atmos. Chem. Phys. Discuss., 12, 27077–27109, doi:10.5194/acpd-12-27077-2012, 2012.
- ⁵ Chacon-Madrid, H. J., Murphy, B. N., Pandis, S. N., and Donahue, N. M.: Simulations of smogchamber experiments using the two-dimensional volatility basis set: linear oxygenated precursors, Environ. Sci. Technol., 46, 11179–11186, 2012.
 - Chan, M. N., Surratt, J. D., Chan, A. W. H., Schilling, K., Offenberg, J. H., Lewandowski, M., Edney, E. O., Kleindienst, T. E., Jaoui, M., Edgerton, E. S., Tanner, R. L., Shaw, S. L., Zheng, M.,
- ¹⁰ Knipping, E. M., and Seinfeld, J. H.: Influence of aerosol acidity on the chemical composition of secondary organic aerosol from β -caryophyllene, Atmos. Chem. Phys., 11, 1735–1751, doi:10.5194/acp-11-1735-2011, 2011.
 - Chen, Q., Liu, Y. J., Donahue, N. M., Shilling, J. E., and Martin, S. T.: Particle-phase chemistry of secondary organic material: modeled compared to measured O:C and H:C elemental ratios provide constraints. Environ. Sci. Technol., 45, 4763–4770, 2011.
- ratios provide constraints, Environ. Sci. Technol., 45, 4763–4770, 2011.
 Compernolle, S., Ceulemans, K., and Müller, J.-F.: Technical Note: Vapor pressure estimation methods applied to secondary organic aerosol constituents from a-pinene oxidation: an intercomparison study, Atmos. Chem. Phys., 10, 6271–6282, doi:10.5194/acp-10-6271-2010, 2010.
- ²⁰ Craven, J. S., Yee, L. D., Ng, N. L., Canagaratna, M. R., Loza, C. L., Schilling, K. A., Yatavelli, R. L. N., Thornton, J. A., Ziemann, P. J., Flagan, R. C., and Seinfeld, J. H.: Analysis of secondary organic aerosol formation and aging using positive matrix factorization of high-resolution aerosol mass spectra: application to the dodecane low-NO_x system, Atmos. Chem. Phys. Discuss., 12, 16647–16699, doi:10.5194/acpd-12-16647-2012, 2012.
- DeCarlo, P. F., Dunlea, E. J., Kimmel, J. R., Aiken, A. C., Sueper, D., Crounse, J., Wennberg, P. O., Emmons, L., Shinozuka, Y., Clarke, A., Zhou, J., Tomlinson, J., Collins, D. R., Knapp, D., Weinheimer, A. J., Montzka, D. D., Campos, T., and Jimenez, J. L.: Fast airborne aerosol size and chemistry measurements above Mexico City and Central Mexico during the MILAGRO campaign, Atmos. Chem. Phys., 8, 4027–4048, doi:10.5194/acp-8-4027-2008, 2008.
- ³⁰ Donahue, N. M., Robinson, C. O., and Pandis, S. N.: Coupled partitioning, dilution, and chemical aging of semivolatile organics, Environ. Sci. Technol., 40, 2635–2643, 2006.



Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set: 1. organic-aerosol mixing thermodynamics, Atmos. Chem. Phys., 11, 3303–3318, doi:10.5194/acp-11-3303-2011, 2011.

Donahue, N. M., Kroll, J. H., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility

- basis set Part 2: Diagnostics of organic-aerosol evolution, Atmos. Chem. Phys., 12, 615– 634, doi:10.5194/acp-12-615-2012, 2012a.
 - Donahue, N. M., Henry, K. M., Mentel, T. F., Kiendler-Scharr, A., Spindler, C., Bohn, B., Brauers, T., Dorn, H. P., Fuchs, H., Tillmann, R., Wahner, A., Saathoff, H., Naumann, K.-H., Möhler, O., Leisner, T., Müller, L., Reinning, M.-C., Hoffman, T., Salo, K., Hallquist, M., Frosch, M.,
- Bilde, M., Tritscher, T., Barmet, P., Praplan, A. P., DeCarlo, P. F., Domenn, J., Prévôt, A. S. H., and Baltensperger, U.: Aging of biogenic secondary organic aerosol via gas-phase OH radical reactions, P. Natl. Acad. Sci., 109, 13503–13508, 2012b.

Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J., Yoo, K. Y., Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle phase acidity

- and oligomer formation in secondary organic aerosol, Environ. Sci. Technol., 38, 6582–6589, 2004.
 - George, I. J., Slowik, J., and Abbatt, J. P. D.: Chemical aging of ambient organic aerosol from heterogeneous reaction with hydroxyl radicals, Geophys. Res. Lett., 35, L13811, doi:10.1029/2008GL033884, 2008.
- Heald, C. L., Jacob, D. J., Park, R. J., Russell, L. M., Huebert, B. J., Seinfeld, J. H., Lia, H., and Weber, R. J.: A large organic aerosol source in the free troposphere missing from current models, Geophys. Res. Lett., 32, L18809, doi:10.1029/2005GL023831, 2005.
 - Heald, C. L., Jacob, D. J., Turquety, S., Hudman, R. C., Weber, R. J., Sullivan, A. P., Peltier, R. E., Atlas, E. L., deGouw, J. A., Warneke, C., Holloway, J. S., Newman, J. A., Flocke, F. M.,
- and Seinfeld, J. H.: Concentrations and sources of organic carbon aerosols in the free troposphere over North American, J. Geophys. Res., 111, D23S47, doi:10.1029/2006JD007705, 2006.
 - Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M.,
- Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin,
 C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn,
 M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J.
 A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick,



F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, Science, 326, 1525–1529, 2009.

Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile organic compounds: a protocol for mechanism development, Atmos. Environ., 31, 81–104, 1997.

Kalberer, M., Paulsen, D., Sax, M., Steinbacher, M., Dommen, J., Prevot, A. S. H., Fisseha, R.,

Weingartner, E., Frankevich, V., Zenobi, R., and Baltensperger, U.: Identification of polymers as major components of atmospheric organic aerosols, Science, 303, 1659–1662, 2004.
 Knopf, D. A., Anthony, L. M., and Bertram, A. K.: Reactive uptake of O₃ by multicomponent and multiphase mixtures containing oleic acid, J. Phys. Chem. A, 109, 5579–5589, 2005.

Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of low-volatility organics in the atmosphere, Atmos. Environ., 42, 3593–3624, 2008.

Kroll, J. H., Smith, J. D., Che, D. L., Kessler, S. H., Worsnop, D. R., and Wilson, K. R.: Measurement of fragmentation and functionalization pathways in the heterogeneous oxidation of oxidized organic aerosol, Phys. Chem. Chem. Phys., 11, 8005–8014, 2009.

Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Charles,

- Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Charles, E. K., and Worsnop, D. R.: Carbon oxidation state as a metric for describing the chemistry of atmospheric organic aerosol, Nature Chem., 3, 133–139, 2011.
 - Kwok, E. S. C. and Atkinson, R.: Estimation of hydroxyl radical reaction rate constants for gas-phase organic compounds using a structure-reactivity relationship: An update, Atmos. Enviton., 29, 1685–1695, 1995.
 - Lee-Taylor, J., Madronich, S., Aumont, B., Baker, A., Camredon, M., Hodzic, A., Tyndall, G. S., Apel, E., and Zaveri, R. A.: Explicit modeling of organic chemistry and secondary organic aerosol partitioning for Mexico City and its outflow plume, Atmos. Chem. Phys., 11, 13219– 13241, doi:10.5194/acp-11-13219-2011, 2011.

30 MathWorks, MATLAB, 2002.

5

15

25

Murphy, D. M., Cziczo, D. J., Hudson, P. K., and Thomson, D. S.: Carbonaceous material in aerosol particles in the lower stratosphere and tropopause region, J. Geophys. Res., 112, D04203, doi:10.1029/2006JD007297, 2007.

| | 12, 32565–32611, 2012 | | | | | | | |
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| | Title Page | | | | | | | |
| 5 | Abstract | Introduction | | | | | | |
| , | Conclusions | References | | | | | | |
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32597

- Murphy, B. N., Donahue, N. M., Fountoukis, C., and Pandis, S. N.: Simulating the oxygen content of ambient organic aerosol with the 2D volatility basis set, Atmos. Chem. Phys., 11, 7859–7873, doi:10.5194/acp-11-7859-2011, 2011.
- Murphy, B. N., Donahue, N. M., Fountoukis, C., Dall'Osto, M., O'Dowd, C., Kiendler-Scharr, A., and Pandis, S. N.: Functionalization and fragmentation during ambient organic aerosol
- A., and Pandis, S. N.: Functionalization and fragmentation during ambient organic aerosol aging: application of the 2-D volatility basis set to field studies, Atmos. Chem. Phys., 12, 10797–10816, doi:10.5194/acp-12-10797-2012, 2012.
 - Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt,
- L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H., Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets from Aerosol Mass Spectrometry, Atmos. Chem. Phys., 10, 4625–4641, doi:10.5194/acp-10-4625-2010, 2010.

Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H., and Worsnop, D.

- R.: Changes in organic aerosol composition with aging inferred from aerosol mass spectra, Atmos. Chem. Phys., 11, 6465–6474, doi:10.5194/acp-11-6465-2011, 2011.
 - Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle partitioning and secondary organic aerosol yields, Environ. Sci. Technol., 30, 2580–2585, 1996.
- Pankow, J. F. and Asher, W. E.: SIMPOL.1: a simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, Atmos. Chem. Phys., 8, 2773–2796, doi:10.5194/acp-8-2773-2008, 2008.
 - Pankow, J. F. and Barsanti, K. C.: The carbon number-polarity grid: A means to manage the complexity of the mix of organic compounds when modeling atmospheric organic particulate matter, Atmos. Environ., 43, 2829–2835, 2009.
 - Ravishankara, A. R.: Heterogeneous and multiphase chemistry in the troposphere, Science, 276, 1058–1065, 1997.

25

- Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop,
 A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking organic aerosols: Semivolatile
 emissions and photochemical aging, Science, 315, 1259–1262, 2007.
- emissions and photochemical aging, Science, 315, 1259–1262, 2007.
 Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-

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aromatic volatile organic compounds, Atmos. Chem. Phys., 3, 161–180, doi:10.5194/acp-3-161-2003, 2003.

Shiraiwa, M., Ammann, M., Koop, T., and Pöschl, U.: Gas uptake and chemical aging of semisolid organic aerosol particles, P. Natl. Acad. Sci., 108, 11003–11008, 2011.

⁵ Surratt, J. D., Kroll, J. H., Kleindienst, T. E., Edney, E. O., Claeys, M., Sorooshian, A., Ng, N. L., Offenberg, J. H., Lewandowski, M., Jaoui, M., Flagan, R. C., and Seinfeld, J. H.: Evidence for organosulfates in secondary organic aerosol., Environ. Sci. Technol., 41, 517–527, 2007.

Tsigaridis, K. and Kanakidou, M.: Global modelling of secondary organic aerosol in the troposphere: a sensitivity analysis, Atmos. Chem. Phys., 3, 1849–1869, doi:10.5194/acp-3-1849-2003, 2003.

Valorso, R., Aumont, B., Camredon, M., Raventos-Duran, T., Mouchel-Vallon, C., Ng, N. L., Seinfeld, J. H., Lee-Taylor, J., and Madronich, S.: Explicit modelling of SOA formation from a-pinene photooxidation: sensitivity to vapour pressure estimation, Atmos. Chem. Phys., 11, 6895–6910, doi:10.5194/acp-11-6895-2011, 2011.

10

25

¹⁵ Volkamer, R., Jimenez, J. L., sanMartini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T., Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from anthropogenic air pollution: Rapid and higher than expected, Geophys. Res. Lett., 33, L17811, doi:10.1029/2006GL026899, 2006.

Yee, L. D., Craven, J. S., Loza, C. L., Schilling, K. A., Ng, N. L., Canagaratna, M. R., Ziemann,

- P. J., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from low-NO_x photooxidation of dodecane: evolution of multigeneration gas-phase chemistry and aerosol formation, J. Phys. Chem. A, 116, 6211–6230, 2012.
 - Zahardis, J. and Petrucci, G. A.: The oleic acid-ozone heterogeneous reaction system: products, kinetics, secondary chemistry, and atmospheric implications of a model system – a review, Atmos. Chem. Phys., 7, 1237–1274, doi:10.5194/acp-7-1237-2007, 2007.
 - Zuend, A. and Seinfeld, J. H.: Modeling the gas-particle partitioning of secondary organic aerosol: the importance of liquid-liquid phase separation, Atmos. Chem. Phys., 12, 3857-3882, doi:10.5194/acp-12-3857-2012, 2012.

Zuend, A., Marcolli, C., Peter, T., and Seinfeld, J. H.: Computation of liquid-liquid equilibria and

³⁰ phase stabilities: implications for RH-dependent gas/particle partitioning of organic-inorganic aerosols, Atmos. Chem. Phys., 10, 7795–7820, doi:10.5194/acp-10-7795-2010, 2010.

| iscussion | ACPD 12, 32565–32611, 2012 | | | | | | |
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Table 1. Secondary organic aerosol yields and initial experimental conditions for OH oxidation of C_{12} alkanes.

| | Compound | OH source | $\begin{array}{c} \text{Seed Conc.} \\ (\text{NH}_4)_2 \text{SO}_4 \\ \mu\text{m}^3 \text{cm}^{-3} \end{array}$ | HC ₀ (ppb) | NO ₀ (ppb) | NO ₂₀ (ppb) | OH_0 (mole cm ⁻³) | Yield [*] (%) |
|----------------------|------------------|-----------|---|--------------------------|--------------------------|---------------------------|---------------------------------|---------------------------|
| High NO _x | dodecane | H_2O_2 | 8.6 | 32.2 | 343 | 429 | 4.20×10^{7} | 5.8 |
| | methylundecane | H_2O_2 | 15.4 | 72.4 | 366 | 424 | 3.90×10^{7} | 5.6 |
| | hexylcyclohexane | H_2O_2 | 12.1 | 22.1 | 320 | 399 | 3.19×10^{7} | 17.7 |
| | cyclododecane | H_2O_2 | 7.5 | 13.8 | 289 | 444 | 2.40×10^{7} | 41.3 |
| Low NO _x | dodecane | HONO | 10.8 | 34.0 | <2 | < 4 | 2.59×10^{6} | 16.0 |
| | methylundecane | HONO | 15.2 | 28.0 | <2 | < 4 | 2.29 × 10 ⁶ | 15.1 |
| | hexylcyclohexane | HONO | 10.5 | 15.6 | <2 | < 4 | 2.81 × 10 ⁶ | 34.8 |
| | cyclododecane | HONO | 13.8 | 10.4 | <2 | < 4 | 3.04 × 10 ⁶ | 47.3 |

 * The values determined here are final yields, i.e., the mass of aerosol formed per mass of alkane reacted after 3 h and 35 h, respectively, in high- vs. low-NO_x experiments.

Table 2. Best-fit parameters to chamber data for OH-initiated oxidation of C_{12} alkanes (For each alkane, the top panel is full set nonlinear minimization and the bottom panel is just to fit f_v and r_p to the total organic mass concentration).

| - | | _ | L. | | -1 |
|----------------------|-------------------|------------------|-------------------------|---|---------------------|
| | | f _v a | $r_{\rm p}^{\rm D}$ | k _a ^c | $[C_x H_y O_z]^{a}$ |
| | | | | (m ³ mol ⁻¹ s ⁻¹) | |
| High NO _x | | 0.57 | 1.12×10^{-2} | 9.99 × 10 ³ | [21.21 21.83 3.23] |
| | dodecane | 0.65 | 9.90 × 10 ⁻³ | - | _ |
| | mothylundooppo | 0.66 | 1.07×10^{-2} | 1.02×10^{4} | [24.00 30.21 1.16] |
| | methylundecane | 0.81 | 1.00 × 10 ⁻² | - | _ |
| | hovulovolohovono | 0.76 | 1.35×10^{-1} | 5.28×10^4 | [22.12 31.23 5.84] |
| | nexylcyclonexarie | 1.10 | 1.00 × 10 ⁻³ | - | _ |
| | avaladadaaana | 0.89 | 1.00×10^{-3} | 1.05×10^4 | [23.49 27.42 3.30] |
| | cyclododecane | 1.42 | 1.00 × 10 ⁻³ | _ | _ |
| Low NO _x | | 0.61 | 1.00×10^{-3} | 5.95 × 10 ³ | [24.00 39.38 5.59] |
| | dodecane | 0.77 | 1.00 × 10 ⁻³ | _ | _ |
| | mathulundaaana | 0.67 | 1.00×10^{-3} | 5.70 × 10 ³ | [24.00 37.78 5.32] |
| | methylundecane | 0.91 | 1.00 × 10 ⁻³ | - | _ |
| | havelavalahavana | 0.82 | 8.99×10^{-2} | 6.01 × 10 ³ | [18.03 26.44 6.00] |
| | nexylcyclonexane | 1.00 | 1.00 × 10 ⁻³ | - | _ |
| | avaladadaaana | 0.89 | 1.00×10^{-3} | 1.86 × 10 ³ | [23.36 31.42 3.29] |
| | cyclouodecalle | 2.00 | 1.00 × 10 ⁻³ | - | _ |

^a The probability of fragmentation ($P_{\rm f}$) should be within the range of 0 to 1.

^b The upper and lower limits for r_{p} are set to be 1 and 10^{-3} , respectively.

^c No upper limit is applied to k_a .

^d The upper limit of *x* value should be 24 for C_{12} alkanes studied here and the *x* value can not be lower than the sum of two molecules with the lowest carbon numbers in the category of semi-volatile compounds defined by Eq. (2); the upper limit of *y* value should be the hydrogen number of the C_{24} alkane; and *z* value ranges from 0 to the sum of two molecules that have the largest oxygen number generated from the functionalization channel.



Table 3. Comparison of the treatment of gas-phase chemistry, including functionalization and fragmentation, particle-phase chemistry, including oxidation and accretion reactions, and gas-particle partitioning among three models: (1) two-dimensional Volatility Basis set (2D-VBS), (2) Statistical Oxidation Model (SOM), and (3) Functional Group Oxidation Model (FGOM).

| | | 2D-VBS ^a | SOM ^b | FGOM |
|---------------------------------|-----------|--|---|--|
| Empirical parameters | | - Probability of fragmentation $P = (O:C)^{f}$ | - Probability of fragmentation $P = c_{\text{frag}} \times \text{O or } P = (\text{O}:\text{C})^{\text{mfrag}}$ | - Probability of fragmentation $P = (O:C)^{fv}$ |
| | | Decrease in saturation concentration per generation | Decrease in vapor pressure per O-atom added | Particle phase oxidation rates |
| | | Number of O-atoms added per generation | Number of O-atoms added per generation | Particle phase accretion reaction rates |
| | | | Number of H-atoms lost per O-atom added | Chemical properties of the non-volatile species |
| Gas-phase chemistry | Func. | Change in the O : C ratio Decrease in C* | Addition of oxygen Decrease in vapor pressure | Addition of functional groups Decrease in vapor pressure |
| | Frag. | – A function of O : C ratio | A function of oxygen number or O : C ratio | – A function of O: C ratio |
| Particle- phase chemistry | Oxidation | OH radical uptake coefficient is unity Follows gas-phase oxidation mechanisms | OH radical uptake coefficient is unity One extra oxygen is added per generation of oxidation Fragmentation is treated the same as gas-phase | Reaction rates are best-fit parameters Follows gas-phase oxidation mechanisms Fragmentation is treated the same as gas-phase |
| | Accretion | - | - | Bimolecular reaction with an adjustable reaction rate constant The elemental composition of non-volatile species is defined by empirical parameters |

^a Assumed that C-C bond scission is randomly distributed along the carbon backbone, resulting in an even molar distribution for all carbon numbers between C_{n-1} and C_1 .

^b The distribution of fragments upon fragmentation is computed in two ways: (1) a random probability for the location of C-C bond scission, and (2) fragmentation leads to C_{n-1} and C_1 and species.

Discussion Paper **ACPD** 12, 32565-32611, 2012 **A Functional Group Oxidation Model** X. Zhang and **Discussion Paper** J. H. Seinfeld **Title Page** Introduction Abstract Conclusions References **Discussion** Paper **Tables Figures** Back **Discussion** Paper Full Screen / Esc **Printer-friendly Version** Interactive Discussion





Fig. 1. Scheme for SOA formation and evolution, showing multi-generational gas- and particlephase reactions, in the Functional Group Oxidation Model. "S" and "N" denote semi-volatile and non-volatile compounds, respectively. Note that the single product for each generation shown here represents a spectrum of products spanning a range of volatilities and oxidation states. The oxidation of semi-volatile compounds comprises two processes: functionalization (red arrows) and fragmentation (yellow arrows). Instantaneous equilibrium partitioning is assumed (green arrows). Non-oxidative (accretion) reactions are considered as bimolecular irreversible reactions (blue arrows).





Fig. 2. Chemical mechanisms of alkane photooxidation underlying fragmentation processes.



32603











Fig. 4. Simulated (colors) and observed (black) time-dependent organic aerosol growth from the photooxidation of four C_{12} alkanes under high- (red) and low- (green) NO_x conditions. Note that "sim.1" represents the full fitting of the six empirical parameters in the FGOM to the chamber data and "sim.2" refers to fitting by only two parameters, r_p and f_v , to the organic mass concentration.



Fig. 5. Time evolution of predicted gas- and particle- phase multi-generational products from OH oxidation of C_{12} alkanes under high- and low-NO_x conditions (upper and lower panel, respectively).





























Fig. 10. van Krevelen diagram. FGOM and SOM simulated H:C vs. O:C temporal profile, together with chamber measured data (black asterisks) for low-NO_x dodecane system.

