**Atmospheric
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Stratospheric BrO abundance measured by a balloon-borne submillimeterwave radiometer

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Abstract

Measurements of mixing ratio profiles of stratospheric bromine monoxide (BrO) were made using observations of BrO otational line emission at 650.179 GHz by a balloon-borne SIS (superconductor-insulator-superconductor) submillimeterwave heterodyne receiver. The balloon was launched from Ft. Sumner, New Mexico (34° N) on 22 September 2011. Peak mid-day BrO abundance varied from 16 ± 2 ppt at 34 km to 6 ± 4 ppt at 16 km. Corresponding estimates of total inorganic bromine (Br_y), derived from BrO vmr (volume mixing ratio) using a photochemical box model, were 21 ± 3 ppt and 11 ± 5 ppt, respectively. Inferred Br_y abundance exceeds that attributable solely to decomposition of long-lived methyl bromide and other halons, and is consistent with a contribution from bromine-containing very short lived substances, Br_y^{VSLS}, of 4 ppt to 8 ppt. These results for BrO and Br_y were compared with, and found to be in good agreement with, those of other recent balloon-borne and satellite instruments.

1 Introduction

Stratospheric bromine contributes to ozone loss through halogen catalyzed cycles analogous to stratospheric chlorine chemistry, except that the bromine cycles have a significantly greater potential to deplete ozone (Daniel et al., 1999). In the daytime upper stratosphere, BrO is the most abundant of the inorganic bromine species comprising up to 70 % of Br_y, where

$$Br_y = BrO + HBr + Br + HOBr + BrONO_2 + BrCl \tag{1}$$

Stratospheric Br_y results from photodecomposition of naturally-occurring and anthropogenic bromine-containing source gases transported from the troposphere into the stratosphere. Although the total bromine loading of the stratosphere has declined by $1 \% \, yr^{-1}$ since 2001 (Dorf et al., 2006b; Hendrick et al., 2008; WMO, 2011) due to reductions in surface emissions of CH_3Br and other halons with long photochemical

lifetimes, some uncertainty remains in the quantitative closure of Br_y budget. Recent balloon-borne and satellite observations of BrO abundance indicate that Br_y levels in the mid- and lower stratosphere exceed that estimated by models based solely on long-lived organic bromine surface source gases (Wamsley et al., 1998; Salawitch et al., 2005).

The additional source of Br_y needed to reconcile the stratospheric bromine budget is generally understood to be from decomposition of naturally occurring bromine-containing source gases with very short photochemical lifetimes (VSLS) such as $CHBr_3$, CH_2BrCl, and CH_2BrCH_2Br, with current estimates of Br_y^{VSLS} ranging from 2 to 8 ppt (Salawitch et al., 2005; WMO, 2011; Brinckmann et al., 2012).

In situ measurements of BrO and total bromine in the stratosphere date from the early 1980's (Berg et al., 1980; Brune and Stimpfle, 1993). More recently, an extensive series of BrO and Br_y profiles, spanning several years and a range of season and latitude, were reported by (Dorf et al., 2006b) from balloon flights of the LPMA/DOAS (Limb Profile Monitor of the Atmosphere/Differential Optical Absorption Spectroscopy) instrument, the SAOZ-BrO (Systeme d'Analyse par Observation Zenithale) balloon-borne spectrometer (Pundt et al., 2002), and the in situ UV resonance fluorescence BrO instrument (Woyke et al., 1999). Both LPMA/DOAS and SAOZ-BrO measure BrO by UV absorption of direct sunlight.

Global-scale observations of stratospheric BrO have been made by four satellite instruments.

SCIAMACHY (Scanning Imaging Absorption Spectrometer for Atmospheric Chartography) (Bovensmann et al., 1999) on the sun-synchronous Envisat, in limb mode, provided altitude profiles of BrO from measurement of scattered sunlight. SCIAMACHY was operational from 2002 until the loss of Envisat in April 2012. BrO profiles applied in this study are from MPI for Chemistry, Mainz. The retrieval is performed in the wavelength range from 338–357.25 nm. For details, see Kühl et al. (2008) and Pukīte et al. (2010).

Daily zonal mean BrO profiles, covering 16 to 36 km altitude, are measured by the Optical Spectrograph and Infra Red Imager System (OSIRIS) instrument on the Odin spacecraft (McLinden et al., 2010). Constituent profiles are derived from observation of spectrally dispersed, limb scattered sunlight (274 nm to 810 nm wavelength range) from the upper troposphere into the lower mesosphere.

The Microwave Limb Sounder (MLS), a millimeter and submillimeterwave heterodyne receiver on the EOS Aura platform, has been in operation since August 2004 and derives vertical profiles of BrO and other gases from limb thermal emission spectra. Aura is in sun-synchronous 98° inclination orbit with ascending node equator at approximately 01:45 p.m. local time. MLS BrO profiles are derived from [81]BrO rotational line emission at 625 GHz and 650 GHz (Waters et al., 2006; Kovalenko et al., 2007b), as measured by the SLS in this study.

The Superconducting Submillimeter-wave Limb Emission Sounder (SMILES) (Masuko et al., 2002; Kikuchi et al., 2010) measured BrO abundance in the middle atmosphere from the Japanese Experiment Module (JEM) of the International Space Station (ISS) from 12 October 2009 to 21 April 2010. The SMILES instrument, like Aura MLS, Odin/SMR and the balloon SLS, is a submillimeterwave heterodyne receiver system. The SMILES optics include a single side band (SSB) filter to separate upper and lower sideband signals while MLS and SLS are double sideband (DSB) receivers. JEM/SMILES was the first space application of sensitive 4 K cooled SIS receiver technology mixers and HEMT (high electron mobility transistor) amplifiers for the Earth atmospheric observation. The ISS orbit is non-sun-synchronous at 58° inclination. SMILES observation frequency bands are 624.32–626.32 GHz and 640.12–650.32 GHz, which cover the same BrO emission lines as Aura MLS and SLS.

2 Observations and analysis

Data presented here are from a high altitude balloon flight of the Jet Propulsion Laboratory (JPL) Submillimeterwave Limb Sounder (SLS) instrument from the Columbia

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Scientific Balloon Facility site at Ft. Sumner, New Mexico, USA (34° N, 104° W). On-board the gondola, in addition to the SLS, were the JPL MkIV solar occultation Fourier transform infrared spectrometer (Toon, 1991) and in situ O_3 sensor (Proffitt and McLaughlin, 1983). The balloon was launched on 22 September 2011, reached

5 float altitude of approximately 39 km at 14:00 local time and remained above 37 km through the flight until flight termination the next day.

3 SLS instrument description

The Submillimeter Limb Sounder (SLS) is a heterodyne radiometer that measures molecular rotational thermal emission spectra from a limb observing geometry. The

10 SLS has been flown on numerous middle and high latitude flights since 1991 (Stach-nik et al., 1992, 1999). Measured gases include ClO, O_3, HCl, N_2O, HO_2, and HNO_3. In 2004, the instrument was upgraded with a superconductor-insulator-superconductor (SIS) mixer improving receiver sensitivity by 20 times compared to that with the Schot-tky diode mixer used on earlier flights. Spectrometers were also upgraded from analog

15 filterbanks to a single digital polyphase filterbank spectrometer that provides uniform channels with high spectral resolution and significantly improved gain stability. Figure 1 is a block diagram of the radiometer. Radiance from the atmospheric limb, sky cold reference and warm calibration reference are selected by a scanning 30 cm diameter primary antenna and scan plate. The measured field-of-view is approximately 0.6° full

20 width at half maximum. The double sideband down converter is a quasi-optic coupled SIS NbTiN junction mixer (Zmuidzinas, 1992). The tunable local oscillator (620 GHz to 680 GHz) is generated by a synthesizer, amplifier, multiplier and coupled to the mixer by a Mylar beamsplitter. The SIS junction and HEMT low noise amplifier (Ward et al., 2003; Ward, 2003) are maintained at 4.2 K in a liquid helium cryostat. Following down-

25 conversion in the SIS mixer, the output intermediate frequency (IF) band (4–8 GHz) is analyzed by the digital spectrometer.

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The digital spectrometer (Fig. 2) deployed on SLS was instrumental in obtaining excellent data quality. A digital polyphase spectrometer was implemented in a Reconfigurable Open Architecture Computing Hardware (ROACH) FPGA signal processing board developed by the Berkeley CASPER project (Werthimer, 2011). Two 3 Gsps (gigasamples per second) analog to digital converter cards, sampling the IF down converted to base band at an aggregate rate of 6 Gsps, provided 3 GHz of bandwidth with 8192 channels and frequency resolution of 375 kHz. To our knowledge this is the most advanced digital spectrometer deployed in a field experiment. For the measurements discussed here alternate channels were co-added to provide a reduced data rate due to the limited on-board data storage system. Unlike autocorrelator and FFT digital spectrometers, the polyphase implementation provides channels with steep and deep skirts. A Power PC interfaced to the Virtex 5 on the CASPER board performs additional data processing, telemetry down-link, and data storage on a USB flash drive.

4   Radiometric calibration

Radiometric calibration was performed as described for earlier versions of the SLS instrument (Stachnik et al., 1992, 1999). The spectrometer response at each frequency i, C^i, is proportional to atmospheric radiance plus receiver noise power. As part of each limb scan sequence, the scan plate is rotated to view cold sky radiance (+45° elevation angle) and an on-board ambient temperature calibration black-body.

For each spectrometer channel,

$$C^i_{\mathrm{limb|sky|cal}} = g^i \times (P^i_{\mathrm{limb|sky|cal}} + P^i_{\mathrm{receiver}}) \qquad (2)$$

where $C_{\mathrm{limb|sky|cal}}$ is the spectrometer response with scan plate viewing limb, sky, or cal and P^i_{receiver} is the noise power contributed internally by the receiver system from the SIS mixer, amplifiers, second heterodyne stage and spectrometer. Limb-sky difference

radiance, $P_{\text{limb-sky}}^i$ for channel i, is then

$$P_{\text{limb-sky}}^i = (C_{\text{ant}}^i - C_{\text{sky}}^i)/g^i(t) \qquad (3)$$

Gain coefficients, $g^i(t)$ (counts/Kelvin), drift with time due to instrument temperature changes, local oscillator drive power and ambient pressure and are determined by linear interpolation of the instantaneous gain, $g^i(\text{cal}_n)$, measured during the ambient temperature calibration target scene at $t = \text{cal}_n$ for limb scan n.

$$g^i(\text{cal}_n) = \frac{C_{\text{cal}}^i - C_{\text{sky}}^i}{P_{\text{cal}} - P_{\text{sky}}^i} \qquad (4)$$

The calibration target black-body radiant power, P_{cal}, in Kelvins, is the Planck temperature evaluated at 650 GHz. The "sky" reference temperature, P_{sky}^i, is essentially the 3 K cosmic background brightness, except near the center of strong atmospheric emission lines where the residual air mass above the balloon altitude adds non-negligible atmospheric radiance of, maximally, 5 % of P_{cal} for balloon altitude of 38 km. To reduce systematic gain error, an atmospheric radiative transfer model (discussed in Sect. 5) with climatological trace gas concentrations is used to estimate P_{sky}^i.

5   Profile retrieval

Mixing ratio profiles of atmospheric trace gases are derived from the SLS calibrated radiances, $P_{\text{l-s}}^i$, through an instrument radiance "forward" model that calculates the expected radiances from an input atmospheric state model. Pre-flight measurements of channel shape, sideband ratio, antenna field-of-view are used, together with balloon height and radiosonde temperature profiles, in the forward model to calculate and convolve ray radiances. The 1-dimensional atmospheric state model parameterizes constituent concentration with altitude as a set of triangular basis functions on 2 km centers

from the surface to 100 km. The atmospheric model currently includes approximately 180 of the likely "most significant" gases and isotopologues for which spectroscopic data is available. Pressure, temperature and water vapor profiles in the model are from local radiosondes launched within the duration of the balloon flight. Ray radiance is calculated at each frequency by line by line summation of Voigt lineshape contributions along a path from balloon through the model atmosphere for each channel frequency element $r^i(v)$ and for each antenna scan plate elevation angle. Spectroscopic and partition function data are from the JPL catalog (Pickett et al., 1998). Expressions for N_2 and O_2 collision-induced absorption and water vapor continuum are from (Goyette and DeLucia, 1990; Rosenkranz, 1998). Instrumental field-of-view smearing is incorporated by convolving the ray radiances with a Gaussian beam shape function matched to the measured beam parameters.

Mixing ratio profiles are estimated from measured radiances by the linear optimal estimation retrieval scheme (Houghton et al., 1984; Rodgers, 2000). Briefly, optimal estimation retrieval is essentially a least-squares fit of observedradiances, weighted by observation uncertainty (covariance), to a linearized atmospheric radiance model, $y - y_0 = \mathbf{K}(x - x_0)$, where the matrix $\mathbf{K}$ of partial derivatives relates the vector of radiances y to the vector of parameters in atmospheric model state vector, x about linearization point (x_0). y_0 is the calculated radiance for parameter vector (x_0). The partial derivative matrix is calculated by finite difference in the forward model along with the radiances. The inversion is stabilized by addition of 'virtual' measurements derived from atmospheric model initial, a priori, state. The retrieval result is bound to the a priori where measurements are either absent or excessively noisy. The minimum variance estimate of the atmospheric state vector is

$$\Delta x = x - x_0 = \mathbf{S}_a \mathbf{K}^T (\mathbf{K} \mathbf{S}_a \mathbf{K}^T + \mathbf{S}_e)^{-1} \Delta x \tag{5}$$

where $y - y_0 = \Delta y$ is the vector of measurement residuals, $\mathbf{S}_a$ is the covariance matrix of the a priori, $\mathbf{S}_e$ is the diagonal measurement covariance and x_0 is the a priori atmospheric state.

6 Retrieval precision and accuracy

The overall uncertainty in retrieved BrO profiles has contributions from radiometric precision from thermal noise and systematic errors from radiance calibration error, standing wave spectral artifacts, antenna pointing, spectroscopic parameters used in the forward radiance model and retrieval error. Radiometric precision is limited by thermal noise

$$\Delta T = \frac{T_{sys}}{\sqrt{B\tau}} \tag{6}$$

where T_{sys} is the sum of atmospheric radiance, $T_{cal|limb|sky}$, and total receiver noise, $T_{receiver}$, expressed in Kelvins; B is the effective measurement noise bandwidth in s^{-1} and τ is the integration time in seconds. $T_{receiver}$ for the SIS receiver is approximately 280 K averaged across the IF band pass. The noise contribution from background atmospheric radiance varies with tangent height from a few K at 38 km to 250 K when scanning down to the lowermost stratosphere. The precision in retrieved parameters, as limited by radiometric noise, is given by

$$\mathbf{S} = (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1} \tag{7}$$

which relates measurement covariance, $\mathbf{S}_e$ to uncertainty in the retrieved parameter covariance. Diagonal elements of $\mathbf{S}$ are the estimated uncertainties in the retrieved atmospheric state vector elements. Systematic uncertainty in the BrO vmr retrieval due to standing waves and other spectral artifacts, remaining after day-night differencing, are estimated from Eq. (5) with the post-retrieval residual substituted for $\mathbf{\Delta y}$. Overall radiance calibration scaling error, due to uncertainty in the calibration target effective black body temperature, is approximately 2 % as estimated by comparison of measured and calculated radiance at the center of optically-thick O_3 lines in the receiver band pass. The SLS instrument pointing is derived from an onboard solid-state gyroscope, inclinometer, and magnetometer package accurate to approximately 0.1° which implies

a height uncertainty of 0.6 km at 30 km tangent height. In post-flight processing, the pointing uncertainty was reduced to approximately 0.3 km by fitting a pointing offset to match the sharp onset of measured and calculated radiance at tangent heights below the tropopause.

⁵ ## 7 Measurements

7.1 BrO spectroscopy

The rotational emission spectrum of [81]BrO in its ground vibrational and electronic state consists of Λ-doubled and hyperfine split lines spaced at ~ 25 GHz (Pickett et al., 1998). Two lines, the $J = \frac{49}{2} - \frac{47}{2}$ at 624.768 and the $J = \frac{51}{2} - \frac{49}{2}$ at 650.179 GHz are within the SLS local oscillator frequency tuning range. Although the BrO line near 624 GHz has fewer spectroscopic interferences, the BrO $J = \frac{51}{2} - \frac{49}{2}$ feature near 650.179 GHz enables simultaneous measurement of the [35]ClO lines at 649.4 GHz, 3 HO_2 lines and 2 O_3 emission lines within the IF band pass of the SLS instrument (see Fig. 3). Additionally, the BrO feature at 650 GHz is also observed by the Aura EOS MLS and ISS JEM-SMILES instruments. This provides the opportunity for direct spectral comparison. Line parameters, pressure line width and temperature dependencies for these transitions are from recent studies by Yamada et al. (2003) and Drouin et al. (2001).

BrO profiles are retrieved from the average of daytime (solar zenith angle less than $\sim 85°$) limb scan spectra recorded between 13:30 and and 18:00 local time. Spectra cover the altitude range, in 2 km steps, from balloon float at 39.6 km to below the tropopause height where the limb path becomes opaque due to water vapor absorption. Figure 3 shows a sample of the flight-average full 3 GHz bandwidth spectra corresponding to limb tangent heights of 38, 36, 34, 30 km. Prominent, near 4500 MHz IF, in these spectra are two optically thick features corresponding to O_3 lines at 638.903 GHz and at 647.839 GHz in the receiver lower and upper sidebands, respectively. These near equivalent strength lines provide a convenient in-flight high precision measure of the

receiver sideband response ratio. As shown in Fig. 4, the [81]BrO emission line octet at 650.178 GHz (6730 MHz IF) appears on the shoulder of stronger OO[17]O lines in the same sideband and adjacent to another O_3 feature.

The excellent dynamic range of the digital spectrometer is evident from Fig. 3 that show spectra covering the full 280 K brightness range of atmospheric signals, Fig. 4 that shows a zoomed in region covering BrO, and Fig. 5 that shows the day minus night difference spectra with a full scale range of 0.3 K. An additional property of the digital spectrometer is its ability to "tie together" contiguous channels due to the use of a common digitizer, which explains the excellent small signal performance of the measurement system, and which also allows a reduced chopping (Dicke switching) rate to be used with minimal impact to the quality of spectral contrast data.

Since BrO abundance exhibits a strong diurnal behavior forming $BrONO_2$ rapidly following sunset, and O_3 does not diurnally vary in the middle stratosphere, interfering O_3 lines, residual standing-wave artifacts, other interferences and offsets were removed by subtracting, from the daytime spectra at each tangent height, corresponding averaged nighttime spectra. The nighttime average consists of spectra recorded after 23:00 local solar time at the position of the balloon. This procedure is particularly effective for balloon measurements since, in the absence of strong wind shear, the balloon platform, drifting with the local circulation, permits continuous (approximate) observation of a single air parcel.

The resulting day-night spectra, shown in Fig. 5, for limb-path tangent heights of 38, 34, 32 and 28 km reveal a 200 mK amplitude [81]BrO line and two prominent HO_2 features also visible due to the diurnal variation of HO_2 mid-stratosphere abundance. BrO and HO_2 profiles were retrieved on a 2 km grid from the day-night spectra using the optimal estimation procedure. Synthetic day-night spectra, calculated using the retrieved BrO and HO_2 profiles, are shown (red) in Fig. 5.

The retrieved BrO profile and estimated uncertainty are shown in Fig. 6 in comparison with a BrO profile from EOS MLS. BrO profile uncertainty, as described in Sect. 6, is the combined error from radiometric precision, calibration bias, pointing offset, and

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retrieval residuals. The variation of profile uncertainty with altitude can be explained as follows. At heights near to and above the balloon float altitude (~ 38 km), uncertainty increases as the instrumental observation path length decreases from a limb path length of ~ 300 km to a much shorter upward slant path observing geometry (~ 50 km). In the middle altitude range (37 km to 25 km), best sensitivity is achieved due to the low level of background radiance, minimal pointing error, and narrow tangent point beamwidth. Below 25 km, uncertainty increases due to increasing thermal noise from water vapor continuum emission, and effect of the propagation of gain error with increasing radiance background.

7.2 Inferred Br_y

BrO, as the most abundant of inorganic bromine gases in the daytime stratosphere, provides the best indication of total Br_y. Since BrO is in rapid photochemical equilibrium with $BrONO_2$, Br, and HOBr, the BrO comprises a variable fraction of Br_y depending on local solar time, latitude, altitude, atmospheric composition and season.

Following previous approaches (Kovalenko et al., 2007b; McLinden et al., 2010), we estimate Br_y profile from measured BrO using a photochemical model. In this study, the University of California, Irvine photochemical box model (Prather, 1992; McLinden et al., 2000; Brohede et al., 2008) was used. The model was constrained to climatological profiles for $O_3, NO_y, Cl_y, N_2O, CH_4$ and H_2O. For the model Br_y profile, the correlation from (Wamsley et al., 1998), increased by an assumed VSLS bromine contribution of 6 ppt, was used. Br_y was estimated using a simple linear scaling relationship between model and measured BrO and Br_y. At each altitude,

$$[Br_y^{measured}] = \frac{[Br_y^{model}]}{[BrO^{model}]}[BrO^{measured}] \tag{8}$$

Since BrO measurements are derived from day-night spectra, BrO^{model} is calculated from corresponding model daytime minus nighttime averaged differences. After sunset,

BrO rapidly forms $BrONO_2$ in the mid- and lower stratosphere so this diurnal correction to the model is small (less than 0.5 ppt) below 38 km. Applying Eq. (8) at each altitude to measured BrO, gives an estimate of the Br_y altitude profile which is shown plotted in Fig. 7. Uncertainty in the Br_y estimate is simply scaled from that estimated for BrO and neglects biases that might be introduced by input parameters to the photochemical model.

From the average Br_y from 28 km to 38 km, where uncertainty in measured BrO and sensitivity to photochemical model inputs are minimal, we derive estimated value for total inorganic bromine, Br_y, of 21 ± 3 ppt. The uncertainty in total Br_y is from the extreme values plus estimated uncertainty of Br_y within the 28 km to 38 km altitude range. We estimate of the implied contribution to Br_y from very short lived bromine organics by subtracting Br_y attributed to CH_3Br and other long-lived bromine source gases using the Br_y-N_2O correlation expressions in (Wamsley et al., 1998) scaled to 16 ppt of CH_3Br plus halon peak abundance (WMO, 2011) assuming age-of-air of 5 yr and 2006 as the year of stratospheric entry. Comparison in Fig. 7 of total Br_y to that from long-lived bromine organics is consistent with a VSLS contribution of 6 ± 4 ppt above 28 km. Below 28 km, a slight increasing trend in estimated Br_y^{VSLS} to 8 ppt at 16 km is evident but within the larger estimated uncertainty bounds for that altitude.

7.3 Comparison with other measurements

The Microwave Limb Sounder (MLS) instrument (Waters et al., 2006) on the EOS Aura satellite measures BrO abundance profiles also monitoring the $J = \frac{51}{2} - \frac{49}{2}$ rotational emission lines at 650.179 GHz. However, MLS uses a different local oscillator frequency (642 870.0 MHz) than the SLS balloon instrument (643 444.0 MHz) and therefore has different spectral interferences from the lines in the receiver opposite sideband. The MLS radiometric noise level necessitates significant spatial and temporal averaging to obtain comparable BrO mixing ratio precision (Kovalenko et al., 2007b). Since the EOS-Aura MLS is in sun-synchronous orbit measuring at 13:45 and 01:45 local time for the ascending and descending parts of the orbit respectively, spectral

interference due to ozone emission and other bias can be effectively removed by taking day-night (ascending-descending) differences of the retrieved BrO profiles. In Fig. 6, we compare the MLS BrO zonal (25° N to 35° N) bi-monthly mean (blue line) of day-night retrieved profiles using the new algorithm described in Millán et al. (2012). The date range of MLS data is centered on the balloon flight date, 22 September 2011. The MLS retrieval profile precision bound for the bi-monthly mean is shown as blue-dashed lines. In Fig. 6, we also show the MLS BrO profile for September 2005 from Kovalenko et al. (2007b) which was derived using a slightly different retrieval procedure. For both MLS retrieval cases, the pressure range over which the data are useful is 10 hPa to 4.6 hPa (approximately 30 km to 37 km) which is indicated on the plot by closed circles, and open circles outside this range. Agreement between both MLS and SLS profiles is generally good and within the MLS precision bounds at all altitudes. In the region between 10 hPa to 4.6 hPa, profiles agree to within 3 ppt.

Figure 8 shows the Br_y profile (black line) inferred from the SLS BrO profile mixing ratio using the U.C. Irvine photochemical box compared to Br_y inferred from the MLS BrO average for 2005 using SLIMCAT and WACCM photochemical models (Millán et al., 2012). The range of scientifically useful MLS measurements (10 to 4.6 hPa) is indicated by solid circles. MLS and SLS Br_y results are consistent in the upper stratosphere (above 6 hPa) finding total inorganic stratospheric bromine at approximately 21 ppt. Uncertainties in inferred Br_y are indicated by dashed black and dashed red lines for SLS and MLS, respectively. For reference, the Br_y estimated from long-lived organic gases from (Wamsley et al., 1998) scaled to 16 ppt for an age-of-air of 5 yr (2006) is also plotted (cyan curve).

Closest available JEM-SMILES BrO data, matching season and local time of the SLS balloon flight, are from October 2009. Data presented here are from the JEM/SMILES operational products version 300.

Figure 9 compares SLS results with daily averaged SMILES BrO profiles (blue lines) for October 12 (11 profiles), 13 (5 profiles), and 15 (14 profiles) in 2009 with local time between 11:00 and 16:00 and with latitude between 25° N and 40° N. Local solar times

for the selected SMILES overpasses were sufficiently close to the SLS measurement time to make the diurnal correction negligible compared to the combined measurement uncertainty. Retrieval error in the daily averaged SMILES BrO profiles, which includes only spectral noise and null space error, is shown as blue-dashed lines in Fig. 9. Data where the SMILES measurement response is low (less than 0.67) are indicated by open circles in the figure. We find good agreement between SMILES and SLS profiles near 25 km but at other altitudes agreement is at the limits of combined uncertainties.

The SCIAMACHY (Bovensmann et al., 1999) instrument on sun-synchronous ENVISAT measures BrO by ultraviolet absorption spectroscopy of limb scattered sunlight. Figure 10 compares SCIAMACHY BrO (red line) from the closest coincident overpass (23 September 2011, 32° N 100° W) against the SLS BrO profile (black solid circles). SCIAMACHY BrO uncertainty is larger below 18 km and above 33 km due to low measurement response (below 0.65) at those altitudes, and those data are indicated in the plot by open circles. Also shown is the SLS profile adjusted (black open circles) to the local solar time (10:15 a.m.) of the SCIAMACHY overpass. The comparison is shown both in concentration (left panel) and volume mixing ratio (right panel) units. Local solar time correction was approximated using the previously described photochemical box model and a simple linear scaling relation

$$BrO^{meas}_{t_{SCIA}} = \frac{BrO^{model}_{t_{SCIA}}}{BrO^{model}_{t_{SLS}}} BrO^{meas}_{t_{SLS}} \qquad (9)$$

where $BrO^{meas}_{t_{SLS}}$ is the profile measured by SLS, $BrO^{model}_{t_{SLS}}$ is the model profile at the mean local time of SLS measurements, and $BrO^{meas}_{t_{SCIA}}$ is the resulting SLS profile adjusted to the local solar time of the SCIAMACHY daytime overpass. This procedure enables meaningful comparison of measurements made at different local times and solar zenith angles, however, additional uncertainty may enter from uncertainty in the photochemical parameters used in the model as discussed by Sioris et al. (2006). Profiles show excellent agreement across the entire altitude range of valid measurements.

For comparison with recent BrO measurements from balloon, we include in Fig. 11 the BrO profile from a flight of the LPMA/DOAS (Limb Profile Monitor of the Atmosphere/Differential Optical Absorption Spectrometer) instrument (Pfeilsticker et al., 2000; Dorf et al., 2006a) in October 2003. The balloon was launched from Aire sur

⁵ l'Adour, France (43.7° N, 0.3° W), at 10° higher latitude than that of Ft. Sumner, but at similar season, local time and solar zenith angle.

In Fig. 12, the SLS BrO profile is compared with BrO profiles from two flights of the SAOZ-BrO balloon instrument. SAOZ-BrO data are from flights launched at Aire sur l'Adour, France (43.7° N) on 1 October 2002 (blue dots line) and at Vanscoy, Canada

¹⁰ (52.0° N) on 25 August 2004 (green dots lines). These SAOZ-BrO data were previously compared with BrO profiles from MLS (Kovalenko et al., 2007a), SCIAMACHY Rozanov et al. (2011) and OSIRIS (McLinden et al., 2010). Since SAOZ-BrO makes observations during balloon ascent, solar zenith angles vary with height. During the 1 October 2002, the solar zenith angle was approximately 81° for measurements at

¹⁵ 15 km to 87° at 29 km. For the 25 August 2004 flight, solar zenith angles were 79° at 15 km to 89° at 35 km. The SLS BrO profile scaled, using the photochemical model procedure as described above, to local solar time corresponding to the mean solar zenith angle during the SAOZ observations is shown plotted in Fig. 12 as black open circles and line.

²⁰ LPMA/DOAS and SLS profiles are in good agreement within mutual error extents over the full coincident altitude range. The SLS profile shows excellent agreement with SAOZ-BrO profile from the October 2003 Aire Sur l'Adour balloon flight. SLS BrO is in good agreement with the SAOZ-BrO Vanscoy profile below about 26 km. We note that these four profiles were made over a range of latitude (34° to 52° N), season (summer

²⁵ to autumn), and year (2002–2011) and differences due to BrO variability are expected. Furthermore, the SAOZ-BrO and LPMA/DOAS observations were made close to the peak stratospheric bromine abundance but the SLS measurement after several years of bromine decline (WMO, 2011).

Figure 13 shows SLS estimated Br_y (black line) mixing ratio plotted against N_2O as proxy for age of air. Plots of Br_y against a dynamical tracer provide a more meaningful inter-comparison of measurements made at various latitudes and during various seasons.

The Br_y-N_2O tracer correlation from the Odin/OSIRIS instrument (McLinden et al., 2010) is shown in Fig. 13. N_2O data is from the Submillimeter Millimeterwave Radiometer (SMR) on the Odin platform. Odin/OSIRIS curve (purple line) is the monthly mean of AM and PM observations for latitudes between 40° S to 40° N. LPMA/DOAS Br_y correlation (green dots-line) from a balloon flight in March 2003 from Kiruna, Sweden (68° N) is also shown (Kovalenko et al., 2007a).

SLS Br_y uncertainty range (black-dashed) is the BrO uncertainty scaled by the Br_y/BrO ratio from the photochemical box model. Abundance of Br_y (cyan line) attributed to long-lived organic sources is from (Wamsley et al., 1998) scaled to the estimated age of air (16 ppt peak abundance) for these observations. Estimated Br_y^{VSLS} abundance (red, uncertainty as red-dashed) is Br_y minus the estimate of inorganic bromine attributed to long-lived bromine source gases.

8 Conclusions

Balloon-borne remote sensing submillimeterwave emission measurements of BrO indicate mid-day abundances ranging from 16 ± 2 ppt at 34 km to 6.5 ± 4 at 16 km. Total Br_y abundance, inferred from BrO and a photochemical model, is 21 ± 3 ppt. The inferred Br_y abundance profile was found to exceed the levels of Br_y attributable solely to photodecomposition of CH_3Br and other long-lived source gases throughout the stratosphere. These results are consistent with the view that decomposition of short-lived bromine-containing source gases deposits at least 4 ppt inorganic bromine in the lowermost stratosphere. Figure 14, adapted from WMO (2011) and Millán et al. (2012), compares SLS Br_y^{VSLS} with recent estimations of the VSLS contribution to

stratospheric inorganic bromine (Br_y^{VSLS}) from satellite, ground-based and balloon-borne instruments. Vertical dashed line marks 6 ppt, the central values for Br_y attributed to VSLS. Results are in good general agreement with other recent measurements and support the conclusions (WMO, 2011).

References

Berg, W. W., Crutzen, P. J., Grahek, F. E., Gitlin, S. N., and Sedlacek, W. A.: 1st measurements of total chlorine and bromine in the lower stratosphere, Geophys. Res. Lett., 7, 937–940, 1980. 28893

Bovensmann, H., Burrows, J. P., Buchwitz, M., Frerick, J., Noel, S., Rozanov, V. V., Chance, K. V., and Goede, A. P. H.: SCIAMACHY: Mission objectives and measurement modes, J. Atmos. Sci., 56, 127–150, 1999. 28893, 28905

Brinckmann, S., Engel, A., Bönisch, H., Quack, B., and Atlas, E.: Short-lived brominated hydrocarbons – observations in the source regions and the tropical tropopause layer, Atmos. Chem. Phys., 12, 1213–1228, doi:10.5194/acp-12-1213-2012, 2012. 28893

Brohede, S., McLinden, C. A., Urban, J., Haley, C. S., Jonsson, A. I., and Murtagh, D.: Odin stratospheric proxy NO_y measurements and climatology, Atmos. Chem. Phys., 8, 5731–5754, doi:10.5194/acp-8-5731-2008, 2008. 28902

Brune, W. H. and Stimpfle, R. M.: In situ measurements of stratospheric reactive trace gases, Chapter 5, 133–184, doi:10.1021/ba-1993-0232.ch005, Advances in Chemistry, Vol. 232, ISBN13: 9780841224704eISBN: 9780841224148, 1993. 28893

Daniel, J. S., Solomon, S., Portmann, R. W., and Garcia, R. R.: Stratospheric ozone destruction: The importance of bromine relative to chlorine, J. Geophys. Res.-Atmos., 104, 23871–23880, 1999. 28892

Dorf, M., Bösch, H., Butz, A., Camy-Peyret, C., Chipperfield, M. P., Engel, A., Goutail, F., Grunow, K., Hendrick, F., Hrechanyy, S., Naujokat, B., Pommereau, J.-P., Van Roozendael, M., Sioris, C., Stroh, F., Weidner, F., and Pfeilsticker, K.: Balloon-borne stratospheric BrO measurements: comparison with Envisat/SCIAMACHY BrO limb profiles, Atmos. Chem. Phys., 6, 2483–2501, doi:10.5194/acp-6-2483-2006, 2006a. 28906

Dorf, M., Butler, J. H., Butz, A., Camy-Peyret, C., Chipperfield, M. P., Kritten, L., Montzka, S. A., Simmes, B., Weidner, F., and Pfeilsticker, K.: Long-term observations of stratospheric bromine reveal slow down in growth, Geophys. Res. Lett., 33, L24803, doi:10.1029/2006GL027714, 2006b. 28892, 28893, 28924

Drouin, B. J., Miller, C. E., Muller, H. S. P., and Cohen, E. A.: The rotational spectra, isotopically independent parameters, and interatomic potentials for the X-1 (2)Pi(3/2) and X-2 (2)Pi(1/2) states of BrO, Journal of Molecular Spectroscopy, 205, 128–138, 2001. 28900

Goyette, T. M. and DeLucia, F. C.: The Pressure Broadening of the $3_{12} \rightarrow 2_{20}$ Transition of Water Between 80 K and 600 K, Journal of Molecular Spectroscopy, 143, 346–358, 1990. 28898

Hendrick, F., Johnston, P. V., De Maziere, M., Fayt, C., Hermans, C., Kreher, K., Theys, N., Thomas, A., and Van Roozendael, M.: One-decade trend analysis of stratospheric BrO over Harestua (60 degrees N) and Lauder (45 degrees S) reveals a decline, Geophys. Res. Lett., 35, L14801, doi:10.1029/2008GL034154, 2008. 28892

Houghton, J. T., Taylor, F., and Rodgers, C.: Remote Sounding of atmospheres, Cambridge University Press, New York, NY 10022, 1984. 28898

Kikuchi, K., Nishibori, T., Ochiai, S., Ozeki, H., Irimajiri, Y., Kasai, Y., Koike, M., Manabe, T., Mizukoshi, K., Murayama, Y., Nagahama, T., Sano, T., Sato, R., Seta, M., Takahashi, C., Takayanagi, M., Masuko, H., Inatani, J., Suzuki, M., and Shiotani, M.: Overview and early results of the Superconducting Submillimeter-Wave Limb-Emission Sounder (SMILES), J. Geophys. Res.-Atmos., 115, D23306, doi:10.1029/2010JD014379, 2010. 28894

Kovalenko, L. J., Jucks, K. W., Salawitch, R. J., Toon, G. C., Blavier, J. F., Johnson, D. G., Kleinbohl, A., Livesey, N. J., Margitan, J. J., Pickett, H. M., Santee, M. L., Sen, B., Stachnik, R. A., and Waters, J. W.: Observed and modeled HOCl profiles in the midlatitude stratosphere: Implication for ozone loss, Geophys. Res. Lett., 34, L19801, doi:10.1029/2007GL031100, 2007a. 28906, 28907, 28926

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Kovalenko, L. J., Livesey, N. L., Salawitch, R. J., Camy-Peyret, C., Chipperfield, M. P., Cofield, R. E., Dorf, M., Drouin, B. J., Froidevaux, L., Fuller, R. A., Goutail, F., Jarnot, R. F., Jucks, K., Knosp, B. W., Lambert, A., MacKenzie, I. A., Pfeilsticker, K., Pommereau, J. P., Read, W. G., Santee, M. L., Schwartz, M. J., Snyder, W. V., Stachnik, R., Stek, P. C., Wagner, P. A., and Waters, J. W.: Validation of Aura Microwave Limb Sounder BrO observations in the stratosphere, J. Geophys. Res.-Atmos., 112, D24S41, doi:10.1029/2007JD008817, 2007b. 28894, 28902, 28903, 28904, 28919

Kühl, S., Pukite, J., Deutschmann, T., Platt, U., and Wagner, T.: SCIAMACHY limb measurements of NO_2, BrO and OClO. Retrieval of vertical profiles: Algorithm, first results, sensitivity and comparison studies, Adv. Space Res., 42, 1747–1764, 2008. 28893

Masuko, H., Shiotani, M., and the SMILES mission team: JEM/SMILES Mission Plan, version 2.1., Tech. Rep. NASDA/CRL, National Space Development Agency of Japan, and Communications Research Laboratory, http://smiles.nict.go.jp/Mission_Plan/, 2002. 28894

McLinden, C. A., Olsen, S. C., Hannegan, B., Wild, O., Prather, M. J., and Sundet, J.: Stratospheric ozone in 3-D models: A simple chemistry and the cross-tropopause flux, J. Geophys. Res.-Atmos., 105, 14653–14665, 2000. 28902

McLinden, C. A., Haley, C. S., Lloyd, N. D., Hendrick, F., Rozanov, A., Sinnhuber, B. M., Goutail, F., Degenstein, D. A., Llewellyn, E. J., Sioris, C. E., Van Roozendael, M., Pommereau, J. P., Lotz, W., and Burrows, J. P.: Odin/OSIRIS observations of stratospheric BrO: Retrieval methodology, climatology, and inferred Br(y), J. Geophys. Res.-Atmos., 115, D15308, doi:10.1029/2009JD012488, 2010. 28894, 28902, 28906, 28907, 28926

Millán, L., Livesey, N., Read, W., Froidevaux, L., Kinnison, D., Harwood, R., MacKenzie, I. A., and Chipperfield, M. P.: New Aura Microwave Limb Sounder observations of BrO and implications for Br_y, Atmos. Meas. Tech., 5, 1741–1751, doi:10.5194/amt-5-1741-2012, 2012. 28904, 28907, 28921, 28927

Pfeilsticker, K., Sturges, W. T., Bosch, H., Camy-Peyret, C., Chipperfield, M. P., Engel, A., Fitzenberger, R., Muller, M., Payan, S., and Sinnhuber, B. M.: Lower stratospheric organic and inorganic bromine budget for the Arctic winter 1998/99, Geophys. Res. Lett., 27, 3305–3308, 2000. 28906

Pickett, H. M., Poynter, R. L., Cohen, E. A., Delitsky, M. L., Pearson, J. C., and Muller, H. S. P.: Submillimeter, millimeter, and microwave spectral line catalog, J. Quant. Spectrosc. Ra., 60, 883–890, 1998. 28898, 28900

Prather, M.: Catastrophic loss of stratospheric ozone in dense volcanic clouds, J. Geophys. Res.-Atmos., 97, 10187–10191, doi:10.1029/92JD00845, 1992. 28902

Proffitt, M. H. and McLaughlin, R. J.: Fast-response dual-beam uv-absorption ozone photometer suitable for use on stratospheric balloons, Review of Scientific Instruments, 54, 1719–1728, 1983. 28895

Puķīte, J., Kühl, S., Deutschmann, T., Platt, U., and Wagner, T.: Extending differential optical absorption spectroscopy for limb measurements in the UV, Atmos. Meas. Tech., 3, 631–653, doi:10.5194/amt-3-631-2010, 2010. 28893

Pundt, I., Pommereau, J. P., Chipperfield, M. P., Van Roozendael, M., and Goutail, F.: Climatology of the stratospheric BrO vertical distribution by balloon-borne UV-visible spectrometry, J. Geophys. Res.-Atmos., 107, 4806, doi:10.1029/2002JD002230, 2002. 28893

Rodgers, C.: Inverse Methods for Atmospheric Sounding: Theory and Practice, World Scientific, Singapore, 2000. 28898

Rosenkranz, P. W.: Water Vapor Microwave Continuum Absorption: a Comparison of Measurements and Results, Radio Science, 33, 919–928, 1998. 28898

Rozanov, A., Kühl, S., Doicu, A., McLinden, C., Puķīte, J., Bovensmann, H., Burrows, J. P., Deutschmann, T., Dorf, M., Goutail, F., Grunow, K., Hendrick, F., von Hobe, M., Hrechanyy, S., Lichtenberg, G., Pfeilsticker, K., Pommereau, J. P., Van Roozendael, M., Stroh, F., and Wagner, T.: BrO vertical distributions from SCIAMACHY limb measurements: comparison of algorithms and retrieval results, Atmos. Meas. Tech., 4, 1319–1359, doi:10.5194/amt-4-1319-2011, 2011. 28906

Salawitch, R. J., Weisenstein, D. K., Kovalenko, L. J., Sioris, C. E., Wennberg, P. O., Chance, K., Ko, M. K. W., and McLinden, C. A.: Sensitivity of ozone to bromine in the lower stratosphere, Geophys. Res. Lett., 32, L05811, doi:10.1029/2004GL021504, 2005. 28893

Sioris, C. E., Kovalenko, L. J., McLinden, C. A., Salawitch, R. J., Van Roozendael, M., Goutail, F., Dorf, M., Pfeilsticker, K., Chance, K., von Savigny, C., Liu, X., Kurosu, T. P., Pommereau, J. P., Bosch, H., and Frerick, J.: Latitudinal and vertical distribution of bromine monoxide in the lower stratosphere from Scanning Imaging Absorption Spectrometer for Atmospheric Chartography limb scattering measurements, J. Geophys. Res.-Atmos., 111, D14301, doi:10.1029/2005JD006479, 2006. 28905

Stachnik, R. A., Hardy, J. C., Tarsala, J. A., Waters, J. W., and Erickson, N. R.: Submillimeter wave heterodyne measurements of stratospheric ClO, HCl, O_3, and HO_2 – 1st results, Geophys. Res. Lett., 19, 1931–1934, 1992. 28895, 28896

Stachnik, R. A., Salawitch, R., Engel, A., and Schmidt, U.: Measurements of chlorine partitioning in the winter Arctic stratosphere, Geophys. Res. Lett., 26, 3093–3096, 1999. 28895, 28896

Toon, G. C.: The JPL MkIV Interferometer, Opt.Photonic News, 2, 19–21, 1991. 28895

5 Wamsley, P. R., Elkins, J. W., Fahey, D. W., Dutton, G. S., Volk, C. M., Myers, R. C., Montzka, S. A., Butler, J. H., Clarke, A. D., Fraser, P. J., Steele, L. P., Lucarelli, M. P., Atlas, E. L., Schauffler, S. M., Blake, D. R., Rowland, F. S., Sturges, W. T., Lee, J. M., Penkett, S. A., Engel, A., Stimpfle, R. M., Chan, K. R., Weisenstein, D. K., Ko, M. K. W., and Salawitch, R. J.: Distribution of halon-1211 in the upper troposphere and lower stratosphere and the
10 1994 total bromine budget, J. Geophys. Res.-Atmos., 103, 1513–1526, 1998. 28893, 28902, 28903, 28904, 28907, 28920, 28921, 28926

Ward, J. S.: Observations of Carbon Monoxide in the Starburst Galaxy M82 with a 690 GHz Wide Spectral Bandwidth Receiver, Ph.D. thesis, California Institute of Technology, 2003. 28895

15 Ward, J. S., Zmuidzinas, J., Harris, A. I., and Isaak, K. G.: A (CO)-C-12 J=6-5 map of M82: The significance of warm molecular gas, Astrophysical Journal, 587, 171–185, 28 UNIV CHICAGO PRESS Part 1, 2003. 28895

Waters, J. W., Froidevaux, L., Harwood, R. S., Jarnot, R. F., Pickett, H. M., Read, W. G., Siegel, P. H., Cofield, R. E., Filipiak, M. J., Flower, D. A., Holden, J. R., Lau, G. K., Livesey, N. J.,
20 Manney, G. L., Pumphrey, H. C., Santee, M. L., Wu, D. L., Cuddy, D. T., Lay, R. R., Loo, M. S., Perun, V. S., Schwartz, M. J., Stek, P. C., Thurstans, R. P., Boyles, M. A., Chandra, K. M., Chavez, M. C., Chen, G. S., Chudasama, B. V., Dodge, R., Fuller, R. A., Girard, M. A., Jiang, J. H., Jiang, Y. B., Knosp, B. W., LaBelle, R. C., Lam, J. C., Lee, K. A., Miller, D., Oswald, J. E., Patel, N. C., Pukala, D. M., Quintero, O., Scaff, D. M., Van Snyder, W., Tope, M. C., Wagner,
25 P. A., and Walch, M. J.: The Earth Observing System Microwave Limb Sounder (EOS MLS) on the Aura satellite, Ieee Transactions on Geoscience and Remote Sensing, 44, 1075–1092, 61 IEEE-INST ELECTRICAL ELECTRONICS ENGINEERS INC, 2006. 28894, 28903

Werthimer, D.: The CASPER collaboration for high-performance open source digital radio astronomy instrumentation, in: General Assembly and Scientific Symposium, 2011 XXXth
30 URSI, 1–4, doi:10.1109/URSIGASS.2011.6051282, 2011. 28896

WMO: Scientific Assessment of Ozone Depletion: 2010, Global Ozone Research and Monitoring Project-Report No. 52, World Metrological Organization, Geneva, Switzerland, 2011. 28892, 28893, 28903, 28906, 28907, 28908, 28927

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5 Yamada, M. M., Kobayashi, M., Habara, H., Amano, T., and Drouin, B. J.: Submillimeter-wave measurements of the pressure broadening of BrO, J. Quant. Spectrosc. Ra., 82, 391–399, 2003. 28900

Zmuidzinas, J.: Quasi-optical slot antenna SIS Mixers, IEEE Transactions on Microwave Theory and Techniques, 40, 1992. 28895

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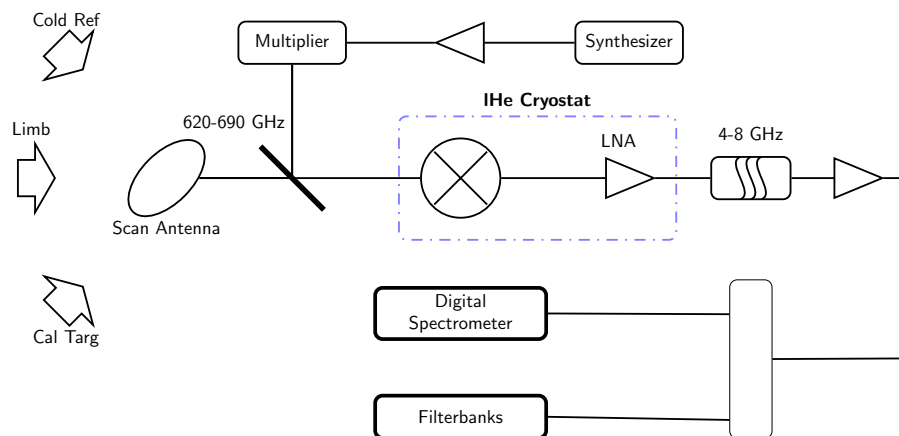
Fig. 1. Diagram of the SLS instrument. Atmospheric radiance is collected by a scanning antenna system, combined by a Mylar beam splitter with a frequency tunable local oscillator, downconverted in an SIS mixer. The IF signal is amplified, filtered, downconverted in a second heterodyne stage to the input frequency range of the digital and analog spectrometers.
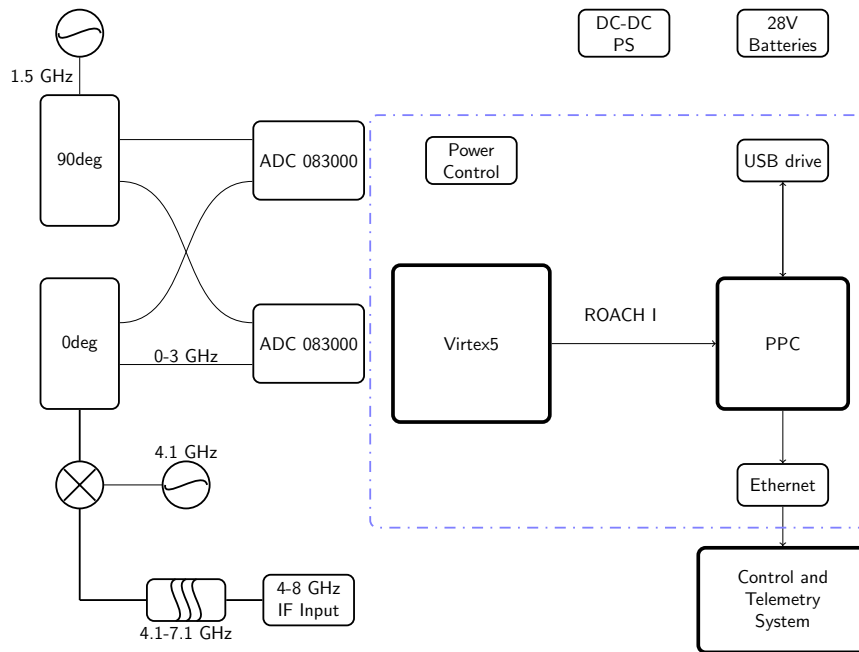
Fig. 2. Diagram of the digital spectrometer sub-system. Two interleaved National ADC083000 digitizers sample the input signal at 6 G samples/sec. Signal bitstream is converted to a power spectrum by an algorithm implemented in the Xilinx Virtex 5 FPGA. Code in a PowerPC collects spectra data , stores data onboard in a flashdrive, and transmits a real-time compressed sample of spectra to the ground.

Fig. 3. Measured double-sideband emission spectra with LO frequency 643.444 GHz. Sample spectra from the September 2011 balloon flight showing the full spectrometer bandwidth corresponding to limb tangent heights of 38, 36, 34 and 30 km. Region containing the [81]BrO lines is at the far right. Abscissa scales show frequency in as intermediate frequency (IF), and upper/lower sideband frequency prior to heterodyne downconversion.

Fig. 4. Daytime emission spectra observed corresponding to limb tangent heights of 38, 34, and 32 km showing the region of ^{81}BrO and overlapping O_3 lines. Abscissa scales show frequency in as intermediate frequency (IF), and upper/lower sideband frequency prior to heterodyne downconversion.
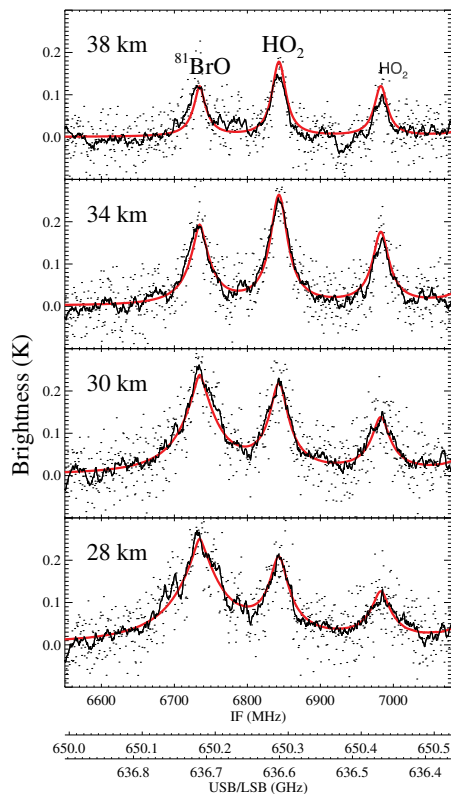
Fig. 5. Day-night difference emission spectra of the same spectral region shown in Fig. 4. Spectrometer data are shown at 750 kHz (dots) resolution and boxcar smoothed (thick solid black line). Synthetic spectra calculated using retrieved BrO profiles are also plotted (red-solid). Abscissa scales show frequency in as intermediate frequency (IF), and upper/lower sideband frequency prior to heterodyne downconversion. HO_2 line occurring in lower sideband is indicated by the san-serif font.
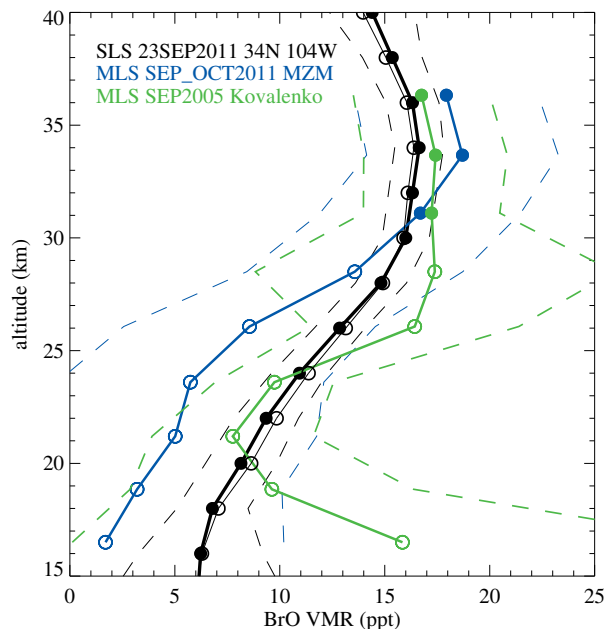
Fig. 6. SLS BrO profile (black) from the 09/23/2011 balloon flight compared with MLS. The MLS BrO profile (blue-solid) is zonal mean (25° N to 35° N) for the period 9/1/2011 through 10/31/2011. The MLS 2 month zonal mean averaged precision is indicated by the blue-dashed lines. The range of scientifically useful MLS measurements is 10 to 4.6 hPa corresponding to approximately 30 km to 37 km. Also shown (green dots-lines) is the MLS BrO profile for September 2005 from (Kovalenko et al., 2007b) derived using different retrieval procedure. Uncertainty range is indicated by the green dashed lines. The range of useful MLS measurements is indicated by solid circles for both data sets.

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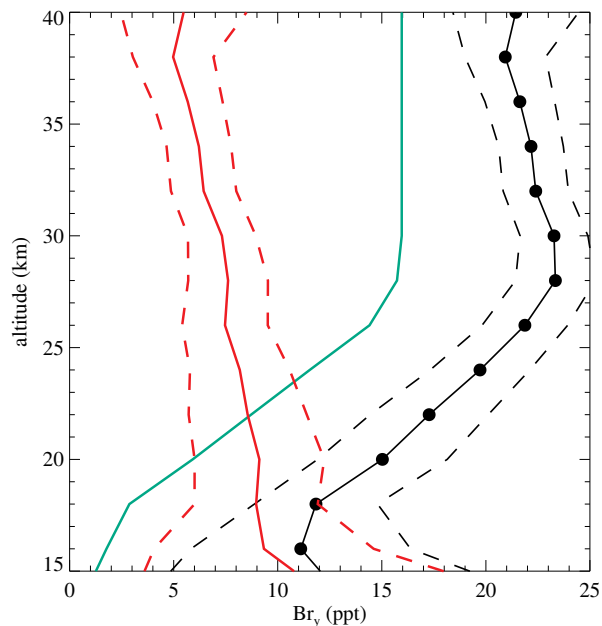
Fig. 7. Br_y mixing ratio inferred from SLS measured BrO (black) using a photochemical box model. Br_y uncertainty range (black-dashed) is the BrO uncertainty scaled by Br_y/BrO from the photochemical box model. Abundance of Br_y from long-lived organic sources (cyan) is from (Wamsley et al., 1998) scaled to 16 ppt at peak. Estimated VSLS abundance (red, uncertainty as red-dashed) is Br_y minus the estimate of inorganic bromine attributed to long-lived bromine source gases.
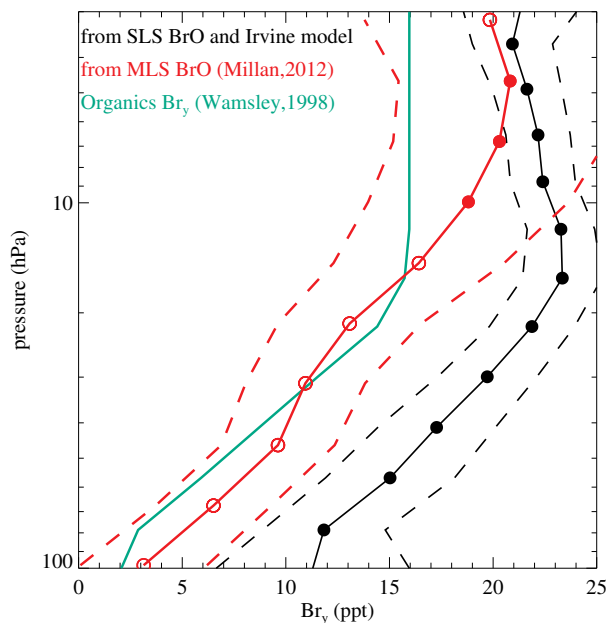
Fig. 8. Comparison of Br$_y$ mixing ratio scaled from SLS measured BrO (black) using the U.C. Irvine photochemical box model against Br$_y$ inferred from MLS BrO for 2005 (Millán et al., 2012). The range of scientifically useful MLS measurements (10 to 4.6 hPa) is indicated by solid circles. Uncertainties in Br$_y$ are indicated by dashed black and dashed red lines for SLS and MLS, respectively. Br$_y$ from long-lived organic gases from (Wamsley et al., 1998) scaled to 16 ppt for an age-of-air of 5 yr (2006) is also plotted (cyan curve).
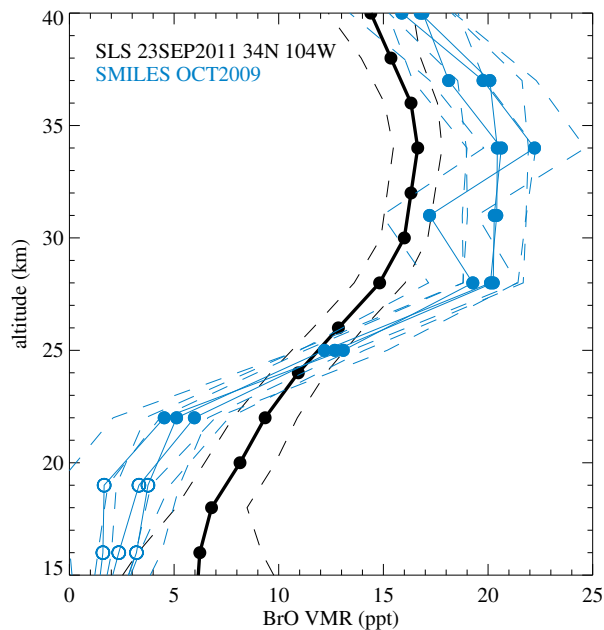
Fig. 9. SLS BrO profile (black dots lines) from the 23 September 2011 balloon flight compared with measurements by the ISS JEM/SMILES instrument in October 2009 (blue-dots-lines). Uncertainty in the SLS BrO profile is shown by the black-dashed lines. SMILES profiles shown are daily averages for 12, 13, 15 October 2009 with local time between 11:00 and 16:00 and with latitude between 25° N and 40° N. Retrieval error on each SMILES BrO daily average, spectral noise and null space error only, is shown as the blue-dashed lines. SMILES data with low measurement response (< 0.67) are indicated by open circles.
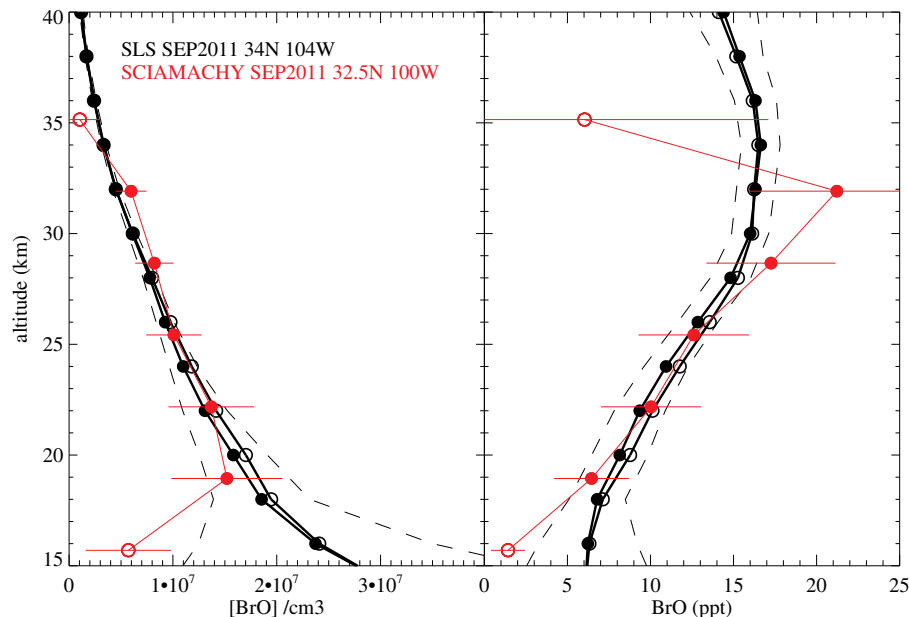
Fig. 10. SLS BrO profile (black-solid circles) from the 23 September 2011 balloon flight compared with SCIAMACHY BrO. Left panel shows the comparison in number density; right panel is the same data plotted as volume mixing ratio. The SCIAMACHY BrO profile (red dots-line) is from the closest coincident overpass (23 September 2011, 32° N 100° W). Also shown (black open circles) is SLS profile scaled to the local time (10:15 a.m.) of the SCIAMACHY overpass using the photochemical model described in the text. Uncertainty in the SLS BrO profile is shown by the black-dashed lines. Larger SCIAMACHY BrO uncertainty below 18 km and above 33 km is due to low measurement response (below 0.65) at those altitudes and indicated by open circles.
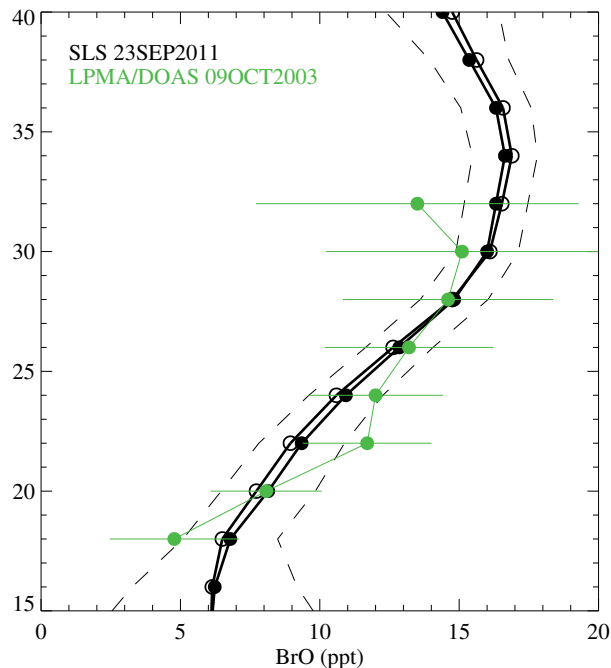
Fig. 11. SLS BrO profile (black-solid circles) from the 23 September 2011 balloon flight compared with a BrO profile (green dots-line) from the LPMA/DOAS instrument (Dorf et al., 2006b) . The LPMA/DOAS profile is from a balloon flight launched from Aire sur l'Adour, France (43.7° N, 0.3° W) on 9 October 2003. The SLS profile adjusted to the mean local solar time corresponding to the DOAS measurement is also plotted (black open circles).
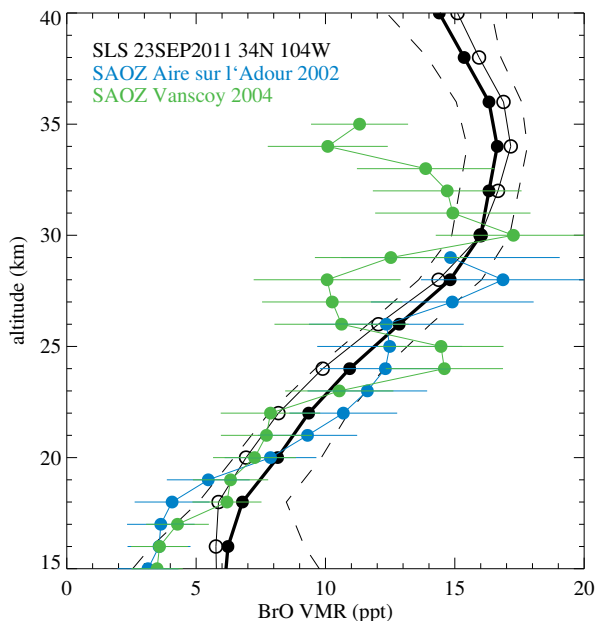
Fig. 12. SLS BrO profile from the 23 September 2011 balloon flight compared with BrO data from two balloon flights of the SAOZ-BrO instrument. SAOZ data are from flights launched at Aire sur l'Adour, France (43.7° N) (blue dots lines) on 1 October 2002 and at Vanscoy, Canada (52.0° N) on 25 August 2004 (green dots-lines). SAOZ measurements were made at solar zenith angles from 81° (15 km) to 87° (29 km) and from 79° (15 km) to 89° (35 km) for the October 2002 and August 2004 flights, respectively. The SLS BrO profile scaled to local time corresponding to the mean solar zenith angle during the SAOZ-BrO observations is shown plotted as black open circles-lines.

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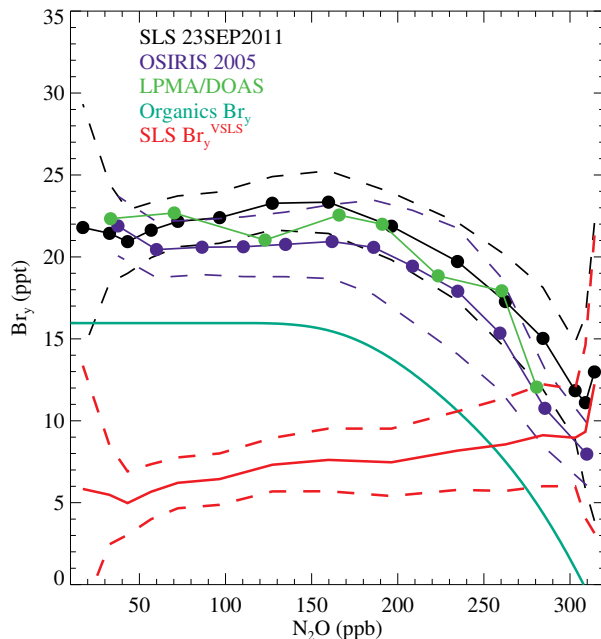Printer-friendly Version

Interactive Discussion

Fig. 13. Correlation of Br_y mixing ratio inferred from SLS measured BrO against N_2O mixing ratio (black). The Br_y tracer correlation from the Odin-OSIRIS and the Odin/Submillimeter Millimeter Radiometer instruments (McLinden et al., 2010) is shown as purple dots-lines. Odin-OSIRIS data is the monthly mean of a.m. and p.m. observations for the 40° S to 40° N latitude band. LPMA/DOAS Br_y correlation (green dots-line) is from a balloon flight in March 2003 from Kiruna, Sweden (Kovalenko et al., 2007a). SLS Br_y uncertainty range (black-dashed) is the BrO uncertainty scaled by Br_y/BrO from the photochemical box model. Abundance of inorganic bromine (cyan) from long-lived source gases is from (Wamsley et al., 1998) scaled as in Fig. 7. Estimated Br_y^{VSLS} abundance (red, uncertainty as red-dashed) is Br_y minus the estimate of inorganic bromine attributed to long-lived bromine source gases.
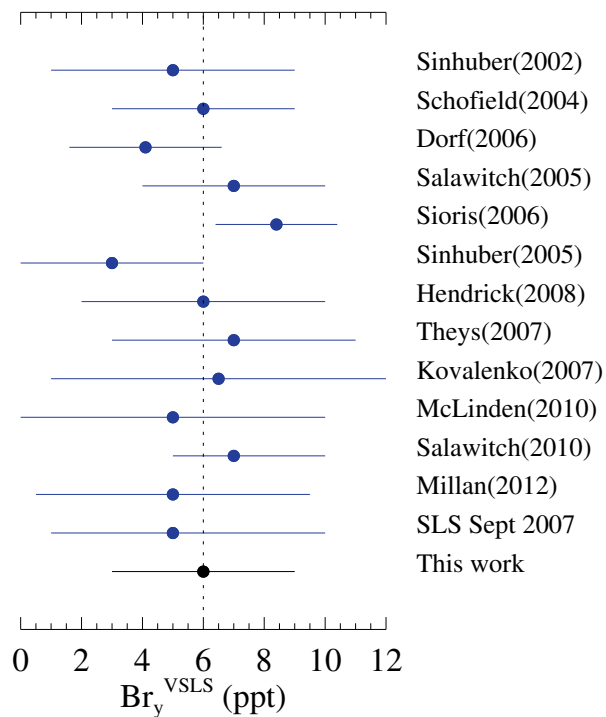
Fig. 14. Adapted from WMO (2011) Tables 1–14 and Millán et al. (2012), summary of recent estimations of the VSLS contribution to stratospheric inorganic bromine (Br_y^{VSLS}) from satellite, ground-based and balloon-borne instruments. Vertical dashed line marks 6 ppt, the central values for Br_y attributed to VSLS.