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**Artificial primary
marine aerosol
production**

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Artificial primary marine aerosol production: a laboratory study with varying water temperature, salinity and succinic acid concentration

J. Zábori¹, M. Matisāns¹, R. Krejci^{1,2}, E. D. Nilsson¹, and J. Ström¹

¹Department of Applied Environmental Science, Stockholm University, 114 18 Stockholm, Sweden

²Department of Physics, University of Helsinki, 00014 Helsinki, Finland

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Correspondence to: J. Zábori (julia.zabori@itm.su.se)

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Abstract

Primary marine aerosols are an important component of the climate system, especially in the remote marine environment. With diminishing sea-ice cover, better understanding of the role of sea spray aerosol on climate in the polar regions is required. As for Arctic Ocean water, laboratory experiments with NaCl water confirm that a few degrees change in the water temperature (T_w) gives a large change in the number of primary particles. Smaller particles with a dry diameter between 0.01 μm and 0.25 μm dominate the aerosol number density, but their relative dominance decreases with increasing water temperature from 0 °C where they represent 85–90 % of the total aerosol number to 60–70 % of the total aerosol number at 10 °C water temperature. This effect is most likely related to a change in physical properties and not to modification of sea water chemistry. A change of salinity between 15 g kg^{-1} and 35 g kg^{-1} showed no influence on the relative shape of a particle number size distribution, nor did a change in water temperature between 0 °C and 16 °C. An experiment where succinic acid was added to a NaCl water solution showed, that the number concentration of particles with $D_p < 0.312 \mu\text{m}$ decreased by 43 % when the succinic acid concentration in NaCl water at $T_w = 0 \text{ }^\circ\text{C}$ was increased from 0 $\mu\text{mol l}^{-1}$ to 2446 $\mu\text{mol l}^{-1}$. Different organic constituents and perhaps inorganic substances resulted in a particle number shift towards larger particle sizes, when comparing a size distribution resulting from pure NaCl water to size distributions resulting from Arctic Ocean water and resulting from NaCl water with a succinic acid concentration of 2446 $\mu\text{mol l}^{-1}$.

1 Introduction

Sea salt aerosols released by bubble bursting from the ocean into the atmosphere are the largest aerosol source globally with respect to aerosol mass (Textor et al., 2006) and play an important role for the radiative balance in remote marine areas. These aerosols influence the global top-of-the-atmosphere clear-sky radiative forcing (Hay-

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wood et al., 1999; Ma et al., 2008; Grini et al., 2002; Zhao et al., 2008) and are understood as an important source for cloud condensation nuclei (Pierce and Adams, 2006). Both, the direct and indirect effect of sea salt particles represent a negative radiative forcing. Due to sea ice retreat, as a consequence of global warming, the conditions and magnitude of Arctic primary marine aerosol source may rapidly change in the near future. During the last decades an accelerated decrease in sea ice extent in the Arctic Ocean has been observed (Comiso et al., 2008), resulting in an increase of the Arctic primary marine aerosol source area and altered energy and particle fluxes. Both the physical and chemical properties of the Arctic Ocean water are changing. Sea surface temperature has increased over the past few decades (Steele et al., 2008). Concurrently, salinity in surface waters is expected to decrease locally, given an expected enhancement in the negative mass balance of glaciers (Nuth et al., 2010). As a consequence of these altered environmental conditions the fauna and flora in the Arctic Ocean is also expected to change along with the sea water organic content (Wassmann and Reigstad, 2011; Tremblay et al., 2011).

The detailed process of primary marine aerosol emission as a result of a bubble bursting is complex. Even the simple parameter of water temperature influences the physical properties of water in various ways; an increase in water temperature lowers the surface tension, viscosity and the solubility of gases in the water, and increases the molecular diffusivity (Thorpe et al., 1992). All these parameters alter the air bubble behavior in the water, such as the time taken for bubble bursting, the bubble rise speed, and the exchange rate between the two phases in the water. Several parameterizations to describe marine aerosol emissions have been developed over the last few decades (Monahan et al., 1986; Gong, 2003; Mårtensson et al., 2003; Clarke et al., 2006). All parameterizations include the whitecap coverage (the result of a bubble plume reaching the surface), which has been shown to be a strong function of wind speed.

However, in this study we focus on the effect on NaCl particle generation from bubble bursting caused by changes in water temperature near freezing conditions, motivated by observations of Arctic Ocean water (Zábori et al., 2012) which revealed a strong

temperature dependence of particle emissions near freezing temperatures. We also investigate the influence of salinity and surface-active compounds represented by succinic acid on total aerosol numbers and number size distributions. The objective of this work, combined with laboratory experiments conducted with real Arctic sea water, is to quantify the relative importance of water temperature, salinity, and other dissolved species in influencing primary marine aerosol production in high latitude ocean waters. This might support parameter considerations for further aerosol emission parametrizations for this region.

2 Methods

2.1 Instrumentation

During the laboratory experiments we used an identical instrumental setup to that used during laboratory experiments with real Arctic sea water in 2010 (Zábori et al., 2012). Instruments providing information on total particle number concentrations, size distributions and parameters describing the state of the sampled air (air temperature and relative humidity) were contained in a compact rack. The sampled air was fed to the instruments through a 2 m long 1/4" diameter stainless steel tube. Due to the geometry of the aerosol sampling line and associated internal losses, the effective upper size limit was estimated to be $5\ \mu\text{m}$ (dry diameter, D_p). Relative humidity in the sampling line was measured by a Hygroclip SC04 (Rotronic) hygrometer and this was used to monitor whether hygroscopic growth could have impacted on particle number concentrations. However, since relative humidity was always lower than 45% during the experiments, all particle sizes are referred to as dry diameters.

Total aerosol number concentrations for D_p larger than $0.01\ \mu\text{m}$ were measured at 1 Hz by a TSI 3010 Condensation Particle Counter (CPC) and particles $0.25\ \mu\text{m} < D_p < 32\ \mu\text{m}$ were determined in 31 different size classes using a GRIMM 1.109 Optical Particle Counter (OPC) once every 6 s. Size distributions of smaller particles

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were measured by a custom-built Differential Mobility Analyzer connected to a TSI 3010 CPC. This Differential Mobility Particle Sizer (DMPS) system operated in the size range 0.01 μm to 0.398 μm , where particles were divided into 17 size classes. Water temperature (T_w) and salinity were measured by a Hach HQ40d conductivity and temperature sensor with a precision of $\pm 0.3^\circ\text{C}$ for temperature and $\pm 0.1\text{ g kg}^{-1}$ for salinity.

2.2 Aerosol generation

Primary marine aerosols are produced by bubble bursting on the water surface. To produce air bubbles in surface water numerous techniques can be applied: Pumping filtered air through a porous media situated in the water (e.g. Mårtensson et al., 2003), producing water waves in a simulation tank which are then allowed to break (e.g. Bowyer et al., 1990), or splashing water on a pre-existing water surface, for example by using a jet of water (e.g. Hultin et al., 2010). Fuentes et al. (2010a) evaluated different methods (except the large simulation tank) with respect to their ability to generate primary marine aerosol by analyzing the bubble spectrum, submicron size distribution, hygroscopicity and cloud condensation nuclei activity. The authors reported that different methods gave very different bubble spectra and as a result different aerosol properties. The use of oceanic water showed that the impinging water jet best simulated the marine primary particle production observed in the ocean and this is also the method employed in this study.

2.3 Surfactant

The surfactant chosen for the experiments was succinic acid, since it is a water-soluble organic compound and ubiquitous in different compartments of the environment (Mahiuddin et al., 2008; Hyvärinen et al., 2006). Succinic acid belongs to a group of low molecular weight dicarboxylic acids (DCA), which are amongst others synthesized and metabolized in different biochemical processes in sea water (Steinberg and Bada, 1984; Kester and Foster, 1963). Succinic acid is, with concentrations of less than

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$1 \mu\text{g l}^{-1}$ one of the lower concentrated DCA in sea water (Tedetti et al., 2006), but can still be considered to represent other DCA, such as oxalic acid, malonic acid and adipic acid.

2.4 Experimental setup

5 About 90 L of water per experiment was used to fill a stainless steel tank with a capacity of 190 L in total. The tank was used as a “buffer” tank from which water was pumped with a rate of 2.5 L min^{-1} using an EHEIM centrifugal pump into a carefully sealed polyethylene bottle (Nalgene Labware). A 5 mm inner diameter stainless steel nozzle in the PET bottle generated a constant water-jet that impinged on a pre-existing water surface. A simple overflow system within the PET bottle kept the water level stable and the water volume remained constant at 10 L. The distance between the nozzle
10 exit and the water surface was approximately 16 cm. Water that entered the overflow drain was fed back to the stainless steel tank by gravity. To avoid any contamination by room air, a filtered particle free air stream was pumped at 7 L min^{-1} into the PET bottle.
15 Repeated checks by switching off the water pump were made to ensure the integrity of the system and that particles in the room air did not enter the system. The aerosol sample was taken directly above the water surface in the PET bottle at a distance of about 18 cm (Fig. 1), with a sampling rate of 6 L min^{-1} . Water temperature and salinity were measured in the outflow of the PET bottle into the steel tank.

2.5 Experiments

Initially, measurements with clean de-ionized MilliQ water (MQ) were conducted to test the functionality of the experimental setup and to create a baseline case for further experiments with higher salinities. Water was collected from a MQ system (Millipore, MiliQ Plus 185) some days before the start of the experiment and was kept frozen
25 at temperatures of -20°C until the preparation for the experiment began. Before the experiment, the frozen MQ water was left to melt until it reached a water temperature

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of 7.4 °C. During the experiment the MQ water was then cooled down to 1.7 °C and again left to warm up to 7.4 °C. Cooling was facilitated using dry ice, kept in a stainless steel bucket in the buffer tank. By removing the cooling bucket the water temperature increased towards the ambient room temperature. Experiments were conducted for low water temperatures in order to emulate Arctic conditions and to investigate the large sensitivity of particle number concentration expected at low temperatures (Zábori et al., 2012).

A similar experiment was repeated for waters with different salinities and different cooling/warming temperature ranges. To minimize any influence of organic matter, salinities of 18 g kg⁻¹ and 35 g kg⁻¹ were achieved by adding NaCl (sodium chloride, puriss.p.a, Sigma-Aldrich) to fresh MQ water in the buffer tank. Once the respective salinity was stable in the outflow of the PET bubble, the experiment started with a cooling sequence, as described above. The experiment using water with a salinity of 18 g kg⁻¹ initially had a water temperature of 21.4 °C and was cooled down to 0.3 °C before the water was left to warm up to 12.6 °C. For the 35 g kg⁻¹ salinity water (representing typical ocean salinity) the corresponding temperatures were 16.8 °C at the start of the experiment, 0 °C was the lowest temperature and 14.2 °C at the end of the experiment.

To examine the influence of various brackish water conditions on the particle number size distribution, experiments with differing salinities, but a relatively stable water temperature were conducted. Since the setup did not allow for maintaining cold temperatures and a homogeneous salinity at the same time (stirring the water with the bucket in the tank was very cumbersome), it was decided to conduct the experiment with water at room temperature. After reaching a relatively stable water temperature the salinity was changed from clean MQ water to 35 g kg⁻¹ water in 15 incremental steps by adding NaCl.

In a final experiment the impact of an organic acid on the temperature dependent trend of particle number characteristics was examined. To achieve this water with a stable salinity of 35 g kg⁻¹ was cooled down from 20.4 °C to about 0 °C. While keeping

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the water temperature close to the freezing point, succinic acid (ACS reagent, >99.0 %, Sigma-Aldrich) was stepwise added to the water in the tank, starting with a concentration of $94 \mu\text{mol l}^{-1}$ which is within the same order of magnitude of the estimated dissolved organic carbon concentration in Arctic Ocean waters (Kivimäe et al., 2010).

5 After a succinic acid concentration of about $2450 \mu\text{mol l}^{-1}$ was reached, the water sample was left to warm up to a water temperature of about 13.3°C .

3 Data analysis

To test the functionality of the experimental setup and to create a baseline case two median size distributions of MQ water (MQ1 and MQ2) were compared to a median size distribution of Arctic Ocean water. The number of particle size distributions used to calculate the respective median is presented in Table 1.

To examine the influence of water temperature on particle number concentrations, data was organized into 1°C water temperature bins. Median particle number concentrations were calculated for each temperature bin. Due to a fluctuating water level in the first two experiments, data were interpolated for the MQ and 18 g kg^{-1} saline water before the median was estimated. On average, cooling water temperature down to the lowest temperature took 12 h 10 min whereas the following warming up time was on average 14 h 20 min. The cooling and warming rates for the different experiments are shown in Table 2.

20 Particle number size distributions of NaCl water with a salinity of 35 g kg^{-1} were averaged over four different water temperature ranges. The first range contained all size distributions resulting from waters having a temperature of 0°C , the second range included all size distributions between 1°C and 4°C , the third contained all size distributions resulting from water with T_w between 8°C and 11°C and the last temperature range included waters with a temperature between 13°C and 16°C . These temperature ranges were chosen to best represent the different steps of the experiment. The ranges 1°C to 4°C and 8°C to 11°C were covered twice, first when the water was cooled down

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and second when the water was left to warm up again. The size distribution resulting from water with a T_w between 13°C and 16°C is only based on water which was cooled down, since no data were available for the warming part. While the range 1°C to 4°C represents the lower part of the total temperature range, 8°C to 11°C represents the intermediate part and 13°C to 16°C represents the upper part. The size distribution of 0°C water represents the turning point between cooling and warming of the water. The number of size distributions used to calculate the median particle number size distribution is shown in Table 3.

To examine the influence of salinity on the particle number size distribution, different salinities from 0 g kg^{-1} NaCl to 35 g kg^{-1} NaCl were binned into intervals of 3 g kg^{-1} NaCl, starting from 0 g kg^{-1} NaCl. For the resulting salinity bins the median $dN/d\log D_p$ was calculated (based on at least five individual size distributions) to which lognormal functions were fitted. The fitting was made subjectively to bridge the overlapping region between the optical and electrical mobility size measurements around D_p 259 nm. The fitted size distributions were normalized to their corresponding integral number to simplify comparison.

The median particle number size distributions resulting from 35 g kg^{-1} NaCl water with different succinic acid concentrations ($0\ \mu\text{mol l}^{-1}$, $94\ \mu\text{mol l}^{-1}$, $565\ \mu\text{mol l}^{-1}$, $1505\ \mu\text{mol l}^{-1}$ and $2446\ \mu\text{mol l}^{-1}$) were compared to test the influence of different succinic acid concentrations in the water on particle number size distributions. Table 4 details the number of size distributions used to build the median for the different succinic acid concentrations.

To explore the effect of possible additional organics and other inorganic salts than NaCl on particle number size distribution, median number size distributions of unadulterated Arctic water, NaCl water with a succinic acid concentration of $2446\ \mu\text{mol l}^{-1}$ and of $0\ \mu\text{mol l}^{-1}$ were calculated. The water temperature for all waters was between -0.5°C and $+0.5^\circ\text{C}$ and salinity was kept between 35.0 g kg^{-1} and 35.6 g kg^{-1} . The number of distributions used to calculate the medians in different size ranges are summarized in Table 5. For comparison, the number size distributions were normalized to

their integral number based on the sum of four lognormal curve fits. This fitting was conducted for the same reasons as described above. Parameters for the fitted size distributions are listed in Table 6. For the overlapping size range of DMPS and OPC measurements the first two size bins of the OPC measurements of D_p 265 nm and D_p 290 nm were not included for fitting the lognormal functions, but are included for completeness in the other number size distribution figures. Data points were omitted based on comparisons of total number concentration with the integral of the number size distribution including the first two size channels and rejecting the first two channels. Including the first two channels led to an underestimation of the total particle number. This decision was supported by evidence that for these sizes the DMPS measurements provide higher quality data; measurements with monodisperse polystyrene latex particles showed a steeply decreasing counting efficiency of the GRIMM 1.109 OPC below 15 % for particles smaller than D_p 0.305 μm (Heim et al., 2008). In addition, measurements based on electrical mobility of the particles are not influenced by their color and likely less influenced by their shape compared to measurements based on optical properties.

4 Results

4.1 MQ water experiments

As a first step, an experiment with pure MQ water was performed to establish a basic test case for further experiments. Particle number concentration produced from MQ water increased slightly with increasing water temperature from 1.5 °C to 7.5 °C for particles with $D_p > 0.01 \mu\text{m}$ (Fig. 2). The largest Poisson error of 4 % was for the particle number concentration at $T_w = 1\text{--}2^\circ\text{C}$, corresponding to a deviation of about 0.5 particles cm^{-3} from the median concentration. Hence, the observed increasing trend is likely real. Nevertheless, the particle number concentration is relatively low (80 particles cm^{-3} for water temperatures between 7 °C and 8 °C) compared with waters

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of 18 g kg^{-1} and 35 g kg^{-1} salinity (at least $350 \text{ particles cm}^{-3}$ for water temperatures between 7°C and 8°C). For particles with $D_p > 0.25 \mu\text{m}$ the particle number concentration is negligible (Fig. 3).

Median aerosol number size distributions from experiments conducted with MQ water for two different water temperature ranges ($3.1^\circ\text{C} < T_w < 3.6^\circ\text{C}$ and $4.8^\circ\text{C} < T_w < 7.4^\circ\text{C}$) were measured about 4 h apart, and named MQ1 and MQ2, respectively (Fig. 4). As a reference, the two size distributions are compared to a number size distribution of Arctic Ocean water with a similar water temperature ($T_w = 3.8^\circ\text{C}$), seen in Fig. 4. Both MQ size distributions show a relatively low number concentration, which is decreasing from D_p $0.01 \mu\text{m}$ to about D_p $3 \mu\text{m}$, while the aerosol size distribution observed for real Arctic Ocean sea water shows an increasing concentration for particles with diameters between D_p $0.01 \mu\text{m}$ and D_p $0.16 \mu\text{m}$. MQ2 shows higher particle concentrations (in total $61 \text{ particles cm}^{-3}$ compared to $23 \text{ particles cm}^{-3}$ for MQ1), but matches the MQ1 number concentration for particle sizes larger than D_p $0.16 \mu\text{m}$.

The MQ water test experiments set a baseline for the instrumental setup used. The increase of smaller particles with water temperature (cf. Fig. 2) likely results from an accumulation of dissolved impurities over time rather than from an increase in water temperature. This was tested by tilting the bottle in which the aerosols were produced allowing for water in the bottle to mix. Instantaneously the particle number concentration of particles with a $D_p > 0.01 \mu\text{m}$ dropped down and increased slowly back to the previously concentration level (not shown here). This suggests that an accumulation of impurities on the water surface on which the water jet impinges, takes place. These substances might be either already present in the MQ water or are a consequence of flushing the system while conducting the experiments. Nevertheless the total number concentrations are negligible compared to the concentrations resulting from experiments with 18 g kg^{-1} and 35 g kg^{-1} saline water.

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4.2 Saline, “inorganic” water experiments

The second experimental stage was the inclusion of NaCl to MQ water to produce the simplest possible sea water proxy. Here, sodium chloride was added in several consecutive steps increasing salinity to a typical oceanic level around 35 g kg^{-1} . Similar to the previous test experiment with MQ water, changes in aerosol number concentration and size distribution were studied as function of T_w and also as a function of salinity. The particle number concentrations for waters with salinities of about 18 g kg^{-1} and 35 g kg^{-1} are shown as a function of water temperature (Fig. 2 and Fig. 3), and for comparison particle number concentrations resulting from Arctic Ocean water were added.

All median concentrations for particles with a $D_p > 0.01 \mu\text{m}$ and a salinity of 18 g kg^{-1} and 35 g kg^{-1} show a clear trend of decreasing aerosol concentrations of about one order of magnitude with increasing water temperature from the lowest temperature bin to a water temperature of about 10°C (Fig. 2). For water temperatures above about 10°C (varying between 8°C and 14°C depending on the experiment), the particle number concentration is rather stable. The particle production from the Arctic Ocean water has about the same temperature trend as the aerosols from the artificial waters. The absolute number concentrations cannot be compared directly due to different dilution factors of the aerosol sample with filtered air.

For particles with a $D_p > 0.25 \mu\text{m}$ a decrease of particle number concentration from the lowest water temperature to 10°C is clearly visible for 35 g kg^{-1} saline water, while the trend starts for 18 g kg^{-1} saline water at $T_w > 2^\circ\text{C}$ (Fig. 3). The particle number concentration for 35 g kg^{-1} saline water is stable for $T_w > 10^\circ\text{C}$ and for 18 g kg^{-1} saline water is stable for $T_w > 13.5^\circ\text{C}$. The total particle number concentration for particles with a $D_p > 0.25 \mu\text{m}$ decreased between three and four times from the lowest water temperatures around 0°C to above $10 - 13^\circ\text{C}$. The data resulting from Arctic Ocean water experiments exhibit the same trend for the particles $D_p > 0.25 \mu\text{m}$ albeit with a steeper slope.

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The ratio of the particle number concentration of particles with $D_p > 0.01 \mu\text{m}$ and $D_p > 0.25 \mu\text{m}$ for the salinities 18 g kg^{-1} and 35 g kg^{-1} are presented in Fig. 5. There is a decrease from the lowest water temperature to a temperature of about 10°C . For water temperatures larger than about 10°C no clear trend is observed. At low water temperatures close to 0°C , 85–90 % of the sea spray aerosol population consists of particles $0.01 \mu\text{m} < D_p < 0.25 \mu\text{m}$ in case of artificial sea water with only NaCl added and about 75–80 % for the Arctic Ocean water. At higher water temperatures the aerosol particles $0.01 \mu\text{m} < D_p < 0.25 \mu\text{m}$ represent between about 60–70 % of aerosol population for both, artificial and real sea waters.

Overall, laboratory experiments using MQ NaCl mixture show consistent number density dominated by sea spray aerosol by particles $0.01 \mu\text{m} < D_p < 0.25 \mu\text{m}$. This is also consistent with previous studies, e.g. on Atlantic Ocean sea water (Hultin et al., 2010) or in tropical Pacific Ocean (Clarke et al., 2006).

Comparing particle number size distributions of 35 g kg^{-1} NaCl water of the four water temperature ranges 0°C , 1°C to 4°C , 8°C to 11°C and 13°C to 16°C show that the number size distribution resulting from the two lower temperature ranges have significant higher particle number concentrations than the size distributions based on the two higher temperature ranges for particle sizes $0.012 \mu\text{m} < D_p < 1.5 \mu\text{m}$ (see Fig. 6). All four distributions overlap for dry diameters outside this range. The two number size distributions resulting from higher water temperatures ($T_w = 8\text{--}11^\circ\text{C}$ and $T_w = 13\text{--}16^\circ\text{C}$) overlap over the whole size range. A local maximum between D_p $0.2 \mu\text{m}$ and $0.25 \mu\text{m}$ can be identified for all four size distributions. To demonstrate the influence of salinity on particle number size distributions for the salinity range $0\text{--}35 \text{ g kg}^{-1}$, log-normally fitted size distributions normalized to the integral number density for the different salinity bins were compared (Fig. 7). Relative particle number concentrations for particles between $0.10 \mu\text{m} < D_p < 0.23 \mu\text{m}$ are higher for the lower salinities and decrease with increasing salinity until they remain stable at a salinity of about 15 g kg^{-1} . The dominant mode seen in Fig. 7 also has a slight shift from about 0.142 to $0.225 \mu\text{m}$ D_p when the salinity increases from $0\text{--}3 \text{ g kg}^{-1}$ to $12\text{--}15 \text{ g kg}^{-1}$, above which the mode diameter is stable.

Relative particle number concentrations for smaller and larger sizes differ as well for different salinities, but not in an identifiable direction. One exception is an indication of a weak tail of smaller particles from about 15 g kg^{-1} . Generally, it can be said that there is a local maxima for all salinity bins at around $0.225 \mu\text{m}$ and the number size distributions do not vary considerably in relative shape for salinities higher than 15 g kg^{-1} .

4.3 Succinic acid experiment

Succinic acid experiments were conducted to test the influence of different water temperatures on water containing a representative organic acid in addition to the NaCl. Also, number size distributions resulting from NaCl waters with different succinic acid concentrations for a stable water temperature were compared to each other. Additionally, a size distribution of aerosols resulting from NaCl water was compared to a size distribution of aerosols resulting from NaCl water with added succinic acid and a size distribution of particles generated from Arctic Ocean water to demonstrate the combined influence of organics and other inorganic salts on the number size distribution.

Comparable to waters having a salinity of 18 g kg^{-1} and 35 g kg^{-1} without added succinic acid, the 35 g kg^{-1} saline water with a succinic acid concentration of $2446 \mu\text{mol l}^{-1}$ caused a similar decreasing trend in particle number concentration with increasing water temperature (Fig. 2 and Fig. 3). Nevertheless, the particle number concentration for $D_p > 0.01 \mu\text{m}$ for water with succinic acid added is on an about 1/3 lower level compared to the other saline waters. As the temperature approaches about 10°C the curves converge and the aerosol production from the succinic acid spiked water matches production during the inorganic experiments. For higher water temperatures the particle number concentrations level out. While an about tenfold decrease of particle number concentration was observed for the inorganic waters (for $D_p > 0.01 \mu\text{m}$), the particle number concentration decreased only six times from the lowest temperature to above 10°C for the water spiked with succinic acid. For particles $D_p > 0.25 \mu\text{m}$ there is no clear difference between inorganic experiments and experiments with succinic acid added to

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the water, see Fig. 3. The decreasing rates with increasing temperature are about three to four times and approximately similar for all saline experiments.

A more detailed analysis shows that the change of particle number concentration between aerosols produced with and without adding succinic acid is not constant over the full temperature range (Fig. 8). For median temperatures between 1.5°C and 13.5°C the difference was always larger for the particles with $D_p > 0.01 \mu\text{m}$ than for the particles with $D_p > 0.25 \mu\text{m}$. While the change is increasing with increasing water temperature for particles with $D_p > 0.25 \mu\text{m}$ (from about 15 to 35%), a smaller increase is observed for the particles with a $D_p > 0.01 \mu\text{m}$ (from about 27% to about 37%).

As for all other saline waters, the ratio between $D_p > 0.01 \mu\text{m}$ and $D_p > 0.25 \mu\text{m}$ shows a decreasing trend with a corresponding water temperature increase from 0°C to 7°C. At low T_w about 0°C 75–80% of all particles are within the size range $0.01 \mu\text{m} < D_p < 0.25 \mu\text{m}$ (Fig. 5).

Figure 9 shows particle number size distributions resulting from NaCl waters to which different amounts of succinic acid were added, resulting in succinic acid concentrations of $94 \mu\text{mol l}^{-1}$, $565 \mu\text{mol l}^{-1}$, $1505 \mu\text{mol l}^{-1}$ and $2446 \mu\text{mol l}^{-1}$ in the water. As a reference, a size distribution produced from pure NaCl water is shown in addition. All median water temperatures were in the range between -0.2 and +0.2°C and medians of salinity were in the range between 35.4 and 35.7 g kg⁻¹. The succinic acid concentration of $94 \mu\text{mol l}^{-1}$ is the same order of magnitude as the estimated dissolved organic carbon concentration in the Arctic Ocean (Kivimäe et al., 2010) and four orders of magnitude higher than the observed concentrations of succinic acid in the ocean (Tedetti et al., 2006). A decrease of particle number concentration with increasing succinic acid concentration is observed for particles $0.025 \mu\text{m} < D_p < 0.300 \mu\text{m}$. Compared to the number size distribution based on pure NaCl water, the size distributions resulting from water with a $94 \mu\text{mol l}^{-1}$ and $2446 \mu\text{mol l}^{-1}$ succinic acid concentration show a 9% and 43% reduced particle number concentration for $D_p < 0.312 \mu\text{m}$.

Log-normally fitted and normalized number size distributions resulting from 35 g kg⁻¹ NaCl water, 35 g kg⁻¹ NaCl water with a succinic acid concentration of $2446 \mu\text{mol l}^{-1}$

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and a normalized number size distribution resulting from Arctic Ocean water were compared to study the influence of different organic contents and salts other than NaCl on the particle number size distribution (Fig. 10). The Arctic Ocean water size distribution includes the contribution of unknown organic substances and inorganic salts other than NaCl, while the other size distributions are caused only by NaCl with or without succinic acid. Between particles with D_p 0.07 to 0.14 μm and particles with D_p 0.16 to 0.3 μm the synthetic water size distributions show relatively higher values than the Arctic Ocean size distribution with a distinct narrow peak close to D_p 0.25 μm . For particles $D_p > 0.316 \mu\text{m}$ the normalized particle number concentration is higher for the ocean water up to a D_p of about 1.8 μm . For particle diameters larger than about D_p 1.8 μm the normalized concentrations converge. The shapes of the normalized fitted size distributions of the two synthetic waters do not differ significantly, except for the magnitude between D_p 0.07 μm and 0.16 μm and between D_p 0.3 μm and 1.0 μm (Fig. 10). For the particles between D_p 0.07 μm and 0.16 μm the relative number is higher for the inorganic NaCl water than for the NaCl water with succinic acid. For the size range between D_p 0.3 μm and 1.0 μm it is reversed.

5 Discussion and conclusion

When comparing particle number size distributions from different studies, it is important to consider that the mechanism used for generating the aerosols might alter the size distribution. Fuentes et al. (2010a) observed four aerosol modes at geometric diameters of 0.014 μm , 0.048 μm , 0.124 μm and 0.334 μm produced by a multiple plunging water jet system. From a higher significance of the fourth mode in the jet system compared to an atomizer, a glass filter and an aquarium diffuser as aerosol production mechanisms, it was concluded that the fourth mode is linked to a mechanism particularly occurring in the jet system. A water jet hitting the water surface can itself produce aerosols and is influencing the air bubbles that rise towards the water surface. Sellegri et al. (2006) also observed a higher significance of a mode $D_p \geq 0.300 \mu\text{m}$ for aerosols

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produced by a weir compared to aerosols produced by spargers. Most probably the particle number size distributions in the current study are also influenced by the aerosol production mechanism. However, since all size distributions were generated using a similar setup, a comparison between the different size distributions is reasonable. That said, comparison with other results presented in the literature has to be treated with caution due to different experimental setup and aerosol generation mechanisms, and different types of water and/or surfactants.

5.1 Total particle number concentration, number size distribution and water temperature

All tested NaCl waters with salinities of 18 g kg^{-1} and 35 g kg^{-1} show a steep decrease of particle number concentration from temperatures of about 0°C up to a temperature of about $8\text{--}14^\circ\text{C}$. The ratio between particles with a $D_p > 0.01 \mu\text{m}$ and particles with a $D_p > 0.25 \mu\text{m}$ is decreasing with increasing T_w . This is an effect due to a more pronounced decrease of particles with $D_p > 0.01 \mu\text{m}$ compared to the decrease of particles with $D_p > 0.25 \mu\text{m}$. The same patterns were observed during experiments with Arctic Ocean waters.

Further, it was demonstrated that the size distributions resulting from 35 g kg^{-1} NaCl water, for different water temperature ranges, all show the same local maxima between D_p $0.20 \mu\text{m}$ and 0.25 . For the sizes $0.012 \mu\text{m} < D_p < 1.5 \mu\text{m}$ the particle number size distributions resulting from the lower water temperature ranges (0°C , 1°C to 4°C) show higher particle number concentrations than the size distribution generated from warmer water (8°C to 11°C and 13°C to 16°C). Particle number concentrations resulting from the two highest water temperature ranges overlapped for the whole size spectrum. Zábori et al. (2012) used Arctic Ocean water and a setup similar to that presented in this paper for the experiments. Within the water temperature range between -2°C and 5°C it was observed that particle number concentrations for all sizes between D_p $0.012 \mu\text{m}$ to about D_p $1.8 \mu\text{m}$ exhibited higher particle number concentrations for lower water temperatures compared to higher water temperatures. For sizes larger than D_p

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1.8 μm the concentrations converge. Local maxima were found to be at 0.180 μm and 0.57 μm . It was concluded that the shape of the size distribution is not influenced by the water temperature. Nevertheless, the relation between the observed local maxima varied slightly with water temperature. The experiments of Hultin et al. (2011) with ocean water from the Baltic Sea presented a significant decrease of particle number concentration for particles $0.02 \mu\text{m} < D_p < 1.8 \mu\text{m}$ from about 750 particles cm^{-3} at 4 $^{\circ}\text{C}$ to about 170 particles cm^{-3} at 14 $^{\circ}\text{C}$. The aerosols were generated by a nearly identical setup to the one used in the current study. The studies are also comparable in that the temperature was gradually changed for the same water.

Bowyer et al. (1990) used a ROYCO Model 225/519 particle counter with six size channels from about $r > 0.25$ to $r > 5 \mu\text{m}$ (80 % RH radius (r)). An increasing particle number production with increasing T_w (between 0 $^{\circ}\text{C}$ and 30 $^{\circ}\text{C}$) from the $r > 1.5 \mu\text{m}$ channel to the largest size channel was observed. The effect increased with size and for the largest channel ($r > 5 \mu\text{m}$) an increase of particle number of the factor 2-3 over the whole temperature range occurs. For the two smallest size channels (that is for particles $r < 1.5 \mu\text{m}$) and T_w below about 15 $^{\circ}\text{C}$ the particle concentration decreased with increasing T_w , up to a factor 4. For temperatures higher than about 13 $^{\circ}\text{C}$ the particle number concentration remained constant. The data Bowyer et al. (1990) used was from the same experiments as Monahan et al. (1982), using a 3 m long white cap simulation tank in which two waves of collected coastal water one on each side were produced and the bubble generation when the waves collided was used to emulate ocean conditions. Bowyer et al. (1990) tried several different regimes to change the water temperature. For most of their data the aerosol is counted and sized continuously as the water temperature changes with only small temperature increments between each reading, much like in the current study. With large increases in temperature between sets of two measurements, the particle production continued to decrease with an increase in T_w through the entire temperature range, well above 15 $^{\circ}\text{C}$.

For particles $D_p < 0.1 \mu\text{m}$, Mårtensson et al. (2003) observed increasing particle production with decreasing temperature using synthetic sea water (with constant water

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temperatures at -2°C , 5°C , 15°C and 23°C). The trend was the opposite for sizes above D_p $0.35\ \mu\text{m}$. Between D_p 0.1 and $0.35\ \mu\text{m}$ there was no clear trend. In the experiments of Mårtensson et al. (2003) air bubbles were generated by pressing clean air through a glass frit, which was shown to be a less representative method to mimic sea spray formation (Fuentes et al., 2010a).

It can be concluded that the observed decrease of particle number concentration with increasing water temperature up to about 8 – 14°C is likely independent of the mechanism with which the aerosols are produced, at least for small particles. Mårtensson et al. (2003) and Bowyer et al. (1990) agree that larger particles have the opposite temperature trend, but they disagree very much at which size the positive temperature trend begins. It is difficult to speculate on why Bowyer et al. (1990) and Mårtensson et al. (2003) observe this trend for the larger particles, given the huge differences in experimental setup and instrumentation.

Mårtensson et al. (2003) observed that between T_w 15°C and 23°C the production of the smallest particles decreased in line with the higher temperature ranges. In this it differs from the current study, with Hultin et al. (2011) and with most of the Bowyer et al. (1990) experiments. In Mårtensson et al. (2003) experiments, aerosol production was allowed to equilibrate at a constant water temperature, while T_w was changed continuously within our experiments and the study of Hultin et al. (2011). Interestingly, Bowyer et al. (1990) also observed a decrease of particle production with increasing T_w up to 30°C , when considering experiments conducted over large temperature increments. Hence, it may be that the existence of a break in the temperature trend, and its exact location on the temperature scale, affects particle production.

Besides the setup, the observed trend of decreasing particle number concentration with increasing water temperature seems to be qualitatively independent of the three different water types. NaCl water (this study), artificial sea water (Mårtensson et al., 2003), and real ocean water (Bowyer et al., 1990) show a similar water temperature dependent trend. Therefore, it is more likely that the observed trend is driven by a physical change rather than by a chemical (or biological) change of the sea water,

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though details in the results may still depend on chemistry or experimental setup. This is supported by the reproducibility of the results when conducting the experiments with temperatures going from high water temperatures down to lower temperatures and up again to higher water temperatures.

Water temperature influences several physical properties involved in the shaping of the air bubble plume. Since sea spray is mainly produced by bursting bubbles, the water temperature is possibly important for the aerosol production. An increase in water temperature reduces viscosity of the water, decreases the solubility of gases and increases the molecular diffusivity. The lower viscosity of water results in higher rise speeds of the bubbles, influencing the gas exchange between the bubble and the water. A lower solubility of the gases increases the number of bubbles, but is compensated by an increase of the diffusivity (Thorpe et al., 1992). Also surface tension at the water and air interfaces (ocean-atmosphere as well as ocean-bubble air) is expected to change. A higher water temperature decreases the surface tension and therefore might alter the bubble spectra as well as the manner in which bubbles burst and droplets form. Thorpe et al. (1992) stated that the net effect of water temperature on bubble production is an exponential decrease of mean bubble concentrations with a halving for every 10 °C. Even if the net effect of water temperature on bubble concentration seems to be known, a net effect on the particle number production is not. However, the laboratory experiments together with measurements performed using real Arctic sea water and a comparison with the literature strongly point towards physically based temperature dependence of the bubble bursting process and sea spray formation. The chemical composition or the role of organic matter seems to play a less important role in these processes.

5.2 Impact of organic content on particle number and number size distribution

For testing the influence of organics on particle number concentration and number size distribution, experiments with NaCl waters having succinic acid concentrations of 94 $\mu\text{mol l}^{-1}$, 565 $\mu\text{mol l}^{-1}$, 1505 $\mu\text{mol l}^{-1}$ and 2446 $\mu\text{mol l}^{-1}$ were conducted. Salinity was about 35.5 g kg^{-1} and water temperature about 0 °C. When the highest suc-

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cinic acid concentration of $2446 \mu\text{mol l}^{-1}$ was reached, water temperature was left to warm up from 0°C to about 13°C . It was observed that higher succinic acid concentrations led to lower particle number concentrations for $D_p < 312 \mu\text{m}$ at a water temperature of 0°C . The artificial NaCl water mixture with a succinic acid concentration of $2446 \mu\text{mol l}^{-1}$ resulted in a lower total particle number concentration over a wide temperature range (from about 0°C to $10\text{--}13^\circ\text{C}$, depending on the salinity), especially for $0.01 \mu\text{m} < D_p < 0.25 \mu\text{m}$, compared to NaCl water mixtures to which no succinic acid was added. The difference in number concentration between water with and water without succinic acid for particles with $D_p > 0.25 \mu\text{m}$ was smaller. For median water temperatures higher than 8.5°C the difference was less than 5%.

A comparison between the log-normally fitted and normalized size distributions from the two synthetic waters showed that the aerosol number concentration resulting from pure NaCl water is relatively higher within $0.07 \mu\text{m}$ to $0.16 \mu\text{m}$ D_p than the number concentration of the size distribution resulting from the NaCl water plus succinic acid. For the size range $0.3 \mu\text{m} < D_p < 1 \mu\text{m}$ it is reversed. Between $0.316 \mu\text{m} < D_p < 1.8 \mu\text{m}$ the the Arctic Ocean water caused relatively higher number concentrations than the synthetic waters.

Sellegrì et al. (2006) generated aerosols by a weir and compared normalized size distributions resulting from synthetic water (including Cl^- , Na^+ , SO_4^{2-} , Mg^{2+} , Ca^{2+} and K^+) and normalized size distributions resulting from synthetic water plus 5 mg l^{-1} sodium dodecyl sulphate (SDS), which is probably the most commonly used detergent surfactant. For sizes about $0.06 \mu\text{m} < D_p < 0.28 \mu\text{m}$ the inorganic water caused higher relative number concentrations than the organic. The number concentration in the size distribution resulting from organic water exceeded that of the inorganic water for sizes from $0.28 \mu\text{m} < D_p < 0.55 \mu\text{m}$. These observations agree quite well with ours when comparing the size distributions derived from synthetic water to the size distributions produced by using Arctic Ocean water. For particle sizes lower than $D_p 0.06 \mu\text{m}$ Sellegrì et al. (2006) reported that the size distribution derived from organic water was characterized by a higher particle number concentration than the one derived from inorganic water. This is

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not observed in our measurements. For all three waters the size distributions differ relatively little for sizes smaller than D_p $0.07 \mu\text{m}$. Consequently, the influence of organics is negligible for the production of aerosols within these sizes, at least for our experimental setup and succinic acid. Sellegri et al. (2006) concluded that a size distribution shift to smaller sizes appeared to happen when adding SDS. Within our study, when adding succinic acid to the water or when using Arctic Ocean water, a shift to larger sizes was observed instead. Differences to the study of Sellegri et al. (2006) might be explained by the differences of the experimental setup, the difference of the synthetic water used including the difference in organic compound and concentration which was used. It appears that the results found when spiking synthetic water with one organic molecule cannot be extrapolated to other organic compounds. It should however be noted that the effect of succinic acid qualitatively agree better with what we found for the complex mixture of organics in the real ocean water, while the effect described by Sellegri et al. (2006) for SDS was the opposite.

Tyree et al. (2007) used a diffuser to produce air bubbles in synthetic sea water (Cl^- , Na^+ , SO_4^{2-} and Mg^{2+}) and compared resulting aerosol number size distributions to number size distributions resulting from artificial sea water plus 0.1, 1 and 10 mg l^{-1} oleic acid. In contrast to succinic acid is oleic acid insoluble in water. Tyree et al. (2007) observed a uni-modal size distribution at around D_p $0.1 \mu\text{m}$ for all experiments. Water with oleic acid resulted in higher total number concentrations compared to the synthetic water without any organic compound. The same was observed by Garrett (1968). A glass frit was used to produce air bubbles and the water was synthesized by adding NaCl until a salinity of 30‰ was reached. After adding oleic acid the particle number concentration increased. Garrett (1968) did not observe an increase of particles if the synthesized NaCl water was filtered twice before the experiment. It was concluded that foam stabilizing substances were present both in real sea water and in the NaCl water he synthesized. These substances delay the bubble bursting process. Adding oleic acid reduces the foam and thus the stability of the bubbles and result in higher particle production due to a higher bubble bursting rate.

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Fuentes et al. (2010b) conducted experiments with a sea water proxy (artificial sea water enriched with diatomaceous bioexudates). For aerosol generation a multiple plunging water jet system was used. Compared to artificial (inorganic) seawater without any exudates the resulting number size distributions were shifted within the submicron size range to smaller sizes. This is qualitatively in agreement with Sellegri et al. (2006), but opposite to our results. An increase of particle number of particles with a $D_p < 0.1 \mu\text{m}$ was observed for an increase of diatomaceous bio-exudates concentration between $175 \mu\text{M OC} < 0.2 \mu\text{m}$ and $300 \mu\text{M OC} < 0.2 \mu\text{m}$ (where $\text{OC} < 0.2 \mu\text{m}$ stands for "organic carbon concentration" of water filtered through a filter with a pore size of $0.2 \mu\text{m}$). This is also qualitatively in agreement with Sellegri et al. (2006) (the increase in number of particles below $0.06 \mu\text{m } D_p$ when SDS was added), but not seen in our results. Even if the aerosol generation is similar in the current study, the organics added to the artificial seawater are quite different. Fuentes et al. (2010b) account not only for dissolved organic carbon, but also for colloidal organic matter attributed to the water insoluble part of primary marine aerosols. The insoluble organic carbon is not considered in our experiment when using NaCl water with succinic acid (since succinic acid is soluble), but probably is present in the Arctic Ocean water sample. Fuentes et al. (2010b) used the exudates of algae present in high biological active regions on the west coast of Africa, which is presumably different to the late winter Arctic Ocean water sample used in our experiments.

If and how much the adding of succinic acid during our experiment might have influenced possible stabilizing substances present in the NaCl water is not clear. There are at least two mechanisms through which the succinic acid might influence the particle number production. On the one hand, it might stabilize or destabilize bubbles on the water surface due to reactions with other substances and on the other hand succinic acid might have influenced the bubble bursting through changes in the surface tension of the water and bubble surfaces.

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5.3 Impact of salinity on particle number and size distributions

Log-normally fitted and normalized size distributions were compared for different salinities from 0 g kg^{-1} up to 35 g kg^{-1} while the water temperature was stable. It was observed that relative particle number concentrations for particles between $0.10\text{ }\mu\text{m} < D_p < 0.23\text{ }\mu\text{m}$ are higher for the lower salinities and decrease with increasing salinity until they remain stable at a salinity of about 15 g kg^{-1} or higher. Scott (1975) examined the coalescence ability and air bubble stability on the water surface for different salinities (0 g l^{-1} , 80 g l^{-1} and 350 g l^{-1}). The coalescence of small air bubbles was more efficiently inhibited for higher salinities resulting in a lower rising velocity and therefore a delayed arrival to the water surface. Bubbles with a higher stability on the water surface were observed as a result of higher salinities. Blanchard (1963,1988) stated that small bubbles produce rather jet drops and large bubbles rather film drops. If we assume that the 35 g kg^{-1} saline water produces more small bubbles which result in jet drop production, we would expect an increase of particles in the super micrometer aerosol range.

Mårtensson et al. (2003) compared number size distributions with salinities of 0 g kg^{-1} , 9.3 g kg^{-1} and 33 g kg^{-1} . At higher salinity, higher number concentrations for particles with $D_p > 0.2\text{ }\mu\text{m}$ but a conserved shape of the distribution was observed. Mårtensson et al. (2003) stated that the change was consistent with a total aerosol volume linearly proportional to the salinity. That would be the logical result if simply every droplet left a smaller dry aerosol particle due to lower salinity, but it would require the fresh sea spray size spectra to stay unchanged despite that the bubble spectra change.

Tyree et al. (2007), who used a diffuser and artificial water of different salinities (from 1 to 70‰), showed that the geometric mean diameter was almost constant with salinity, centered near 100 nm dry diameter, except for the lowest salinities where it changed towards slightly smaller sizes. This is in good agreement with our results. For the aerosol number concentration they found that it increased with increasing salinity, which can-

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not be confirmed from our measurements. Hultin et al. (2011) observed higher particle number concentrations over the whole measurement range for higher salinity ranges (examining real ocean water with salinities about 6 g kg^{-1} , 7 g kg^{-1} and 35 g kg^{-1} salinity), which is at least qualitatively in agreement with Tyree et al. (2007).

5 Russell and Singh (2006) experimented with three different sodium salts. They found that the aerosol mass was linear to the aqueous salt concentration, in agreement with Mårtensson et al. (2003). However, both number concentration and mean diameter depended on the one-sixth power of the aqueous concentration. This suggests that ionic forces retard the fluid motion and limit the bubble bursting process. They found 10 that diameter, number and mass concentrations related to the non-dimensional fluid-mechanical Eötvös number (E_o), which indicates that the forces controlling the bubble bursting mechanism are described effectively by the ratio of buoyant and surface forces. Since they worked with 2 mm sized bubbles, their study concerns the few large bubbles and the film drops presumably produced from them. The large majority of bubbles in 15 a full bubble spectrum (more representative of real ocean conditions) will be around $100 \mu\text{m}$ diameter, and their E_o number will be far below the critical value found by Russell and Singh (2006).

The inorganic system with only salt and water may appear simple. Yet, when comparing studies in the literature and the current study, the results on how salinity influences 20 the particle number size distribution differ considerably. These differences might arise from different combinations of salts in the waters used for the experiments, the presence of some organic compounds in the commercial aquarium salt mixtures used in some experiments, or the experimental setup. It can at least be concluded that the difference in salinity between 15 g kg^{-1} and 35 g kg^{-1} NaCl concentration in our experiments have only little effect on the shape of the particle number size distribution, a 25 conclusion supported by both Hultin et al. (2011) and Tyree et al. (2007).

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6 Future Implications

The effect on sea spray formation by changes in water temperature may have implications for the future climate in the Arctic. The decrease of particle number concentration with increasing water temperature might decrease the direct aerosol effect over the Arctic Ocean. On the one hand the decreased direct effect might result in a decrease of the negative radiative forcing at the top of the atmosphere, since less sea spray particles are emitted over the relatively dark sea surface. However, on the other hand a decreased direct effect may result in a decrease of the positive radiative forcing caused by a reduced sea spray concentration consisting of partly absorbing particles above a bright surface such as sea ice (Forster et al., 2007). At the same time the indirect effect of aerosols should decrease as well. Less small particles imply a reduction in cloud condensation nuclei and therefore might decrease cloud cover and with it lower albedo over the Arctic Ocean. This applies mainly to open Arctic waters, as Nilsson et al. (2001) showed that sea spray emissions within the central Arctic Ocean fields of sea ice are 1–2 orders of magnitude smaller than from open sea. During the last decades, Arctic Sea ice has retreated faster than climate models predict (Stroeve et al., 2007). Nilsson et al. (2001) was the first to suggest that in the Arctic, climate change would cause changes in sea spray emission due both to changes in local meteorology (e.g. higher wind speed) and the retreat of sea ice, which would also generate larger areas with higher sea spray emissions, a factor which may combine or compete with the ice albedo feed back. To this we can add the effects of changing water temperature. Recently, Struthers et al. (2011) used a GCM to quantify the increase in climate forcing caused by sea spray in a climate scenario with largely ice free summers in 2100, and reported these to be between -0.2 and -0.4 Wm^{-2} ; a negative feedback on the climate system. The model predicted that the change in the first indirect aerosol effect (cloud albedo effect) is approximately ten times greater than the change in direct aerosol forcing. Struthers et al. (2011) used the temperature dependency found by Mårtensson et al. (2003). However, numerous factors are expected to change in

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the Arctic regions and influence its climate (increased water vapor concentration and changing general circulation patterns etc.). The warming trend of the Arctic region is clear, but precise predictions are difficult. Nevertheless, implementing the boosted sea spray particle number emissions at low water temperatures into climate models may be important. A sensitivity test including also the effects of wind speed and sea ice on primary marine emissions might give valuable information about the relevance of the new water temperature dependence of particle number production and contribute to improving future climate predictions. As shown previously, a comparison between laboratory experiments is difficult due to different aerosol generation mechanisms, different waters and different organic substances used. More studies are needed for examining the influence of water temperature, salinity, soluble and insoluble organic substances present in different marine environments on particle number production by using only one experimental setup.

7 Summary

Earlier experiments conducted with Arctic Ocean water showed a dependence of primary marine particle number production on low water temperatures suggesting the development of a sea spray parameterization for high latitude oceans. Laboratory experiments with aerosol production by an impinging water jet on a water surface and subsequent bubble bursting were conducted to clarify the temperature dependence observed with ocean waters and potential influence of other parameters (salinity, organic content). Initially, experiments ran with pure NaCl water to exclude any influence on particle production by organic substances contained in the water. Later on succinic acid was added to the pure NaCl water to test the influence of a representative organic surfactant in the ocean. The total particle number concentration was examined as a function of water temperature (in the range from about 0°C up to 21°C) for salinities of 0 g kg⁻¹, 18 g kg⁻¹, and 35 g kg⁻¹ and for water with a salinity of 35 g kg⁻¹ and a succinic acid concentration of 2446 μmol l⁻¹. Particle number size distributions result-

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ing from NaCl water as a function of water temperature (examined $T_w = 0^\circ\text{C}$, $T_w = 1-4^\circ\text{C}$, $T_w = 8-11^\circ\text{C}$ and $T_w = 13-16^\circ\text{C}$) and as a function of succinic acid concentration ($0\ \mu\text{mol l}^{-1}$, $94\ \mu\text{mol l}^{-1}$, $565\ \mu\text{mol l}^{-1}$, $1505\ \mu\text{mol l}^{-1}$ and $2446\ \mu\text{mol l}^{-1}$) were analysed. Additionally, normalized particle number concentrations for 12 different salinity ranges between 0 and $35\ \text{g kg}^{-1}$ were compared to test the influence of salinity on relative particle number size distributions at a stable water temperature. The key findings of this study can be summed up to:

- Changes in water temperatures close to the freezing point strongly influence the total particle number concentration, especially for particles with $0.01\ \mu\text{m} < D_p < 0.25\ \mu\text{m}$. This is likely a consequence of physical rather chemical processes.
- Particle number concentrations resulting from NaCl water with a succinic acid concentration of $2446\ \mu\text{mol l}^{-1}$ show the same water temperature dependent trend as pure NaCl waters, but on a lower particle concentration level.
- A water temperature change between 0°C and 16°C does not change the shape of a particle number size distribution resulting from NaCl water ($35\ \text{g kg}^{-1}$ salinity).
- An increase of succinic acid concentration from $0\ \mu\text{mol l}^{-1}$ to $2446\ \mu\text{mol l}^{-1}$ in four steps at $T_w = 0^\circ\text{C}$ tends to decrease the particle number concentration for particles with $0.01\ \mu\text{m} < D_p < 0.312\ \mu\text{m}$ by about 43%.
- High concentrations of succinic acid and other unknown organics and salts present in Arctic Ocean water alter the size distribution resulting of pure NaCl water. A shift to larger sizes is observed from pure NaCl water to Arctic Ocean water.
- Salinity changes between $15\ \text{g kg}^{-1}$ and $35\ \text{g kg}^{-1}$ do not influence the shape of the particle number size distribution between $0.01\ \mu\text{m} < D_p < 5\ \mu\text{m}$ when using pure NaCl water.

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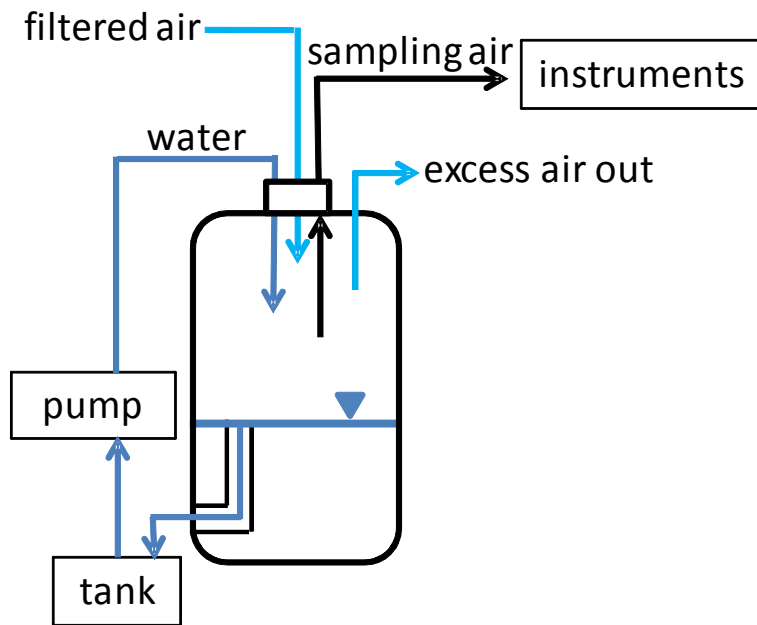


Fig. 1. Schematic picture of bubble bursting experimental setup. The tank was used as a buffer to recirculate the sea water sample through the PET bottle, where primary marine aerosol was produced by an impinging water jet. Darker blue lines represent water, and the triangle symbol indicates water surface in the bottle.

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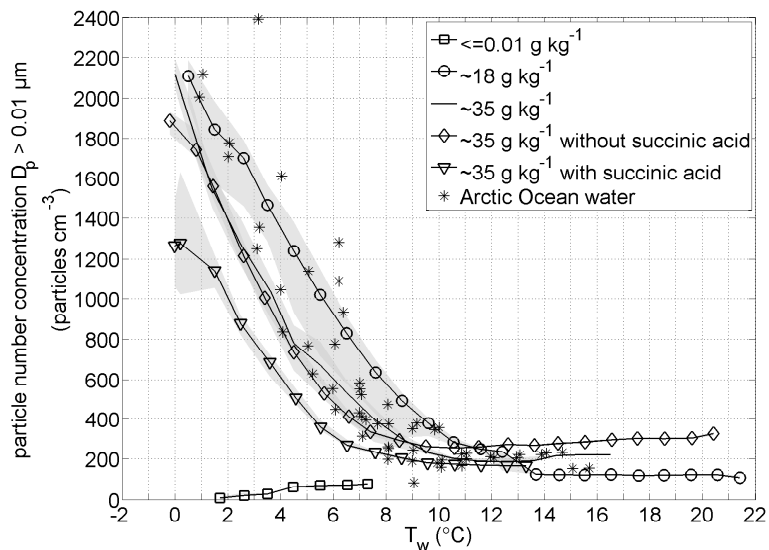


Fig. 2. Median particle number concentrations ($D_p > 0.01 \mu\text{m}$) for waters with different salinities and organic content as a function of water temperature (T_w). Grey shaded areas indicate the 25th and 75th percentiles.

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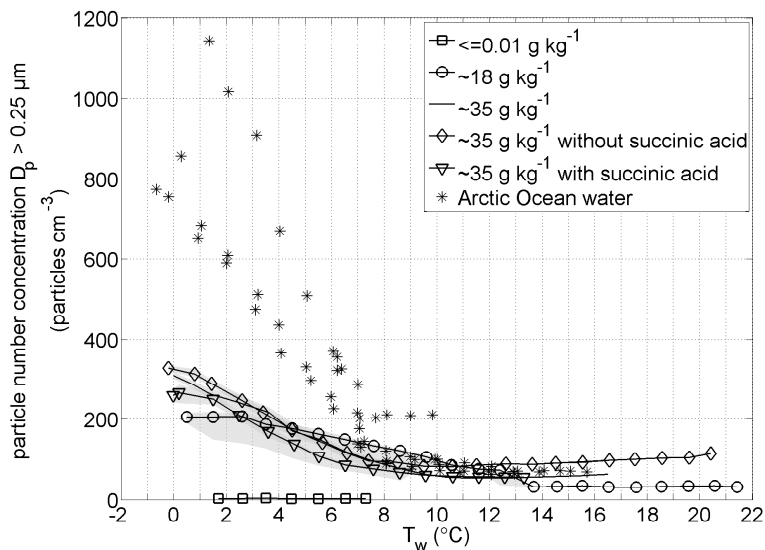


Fig. 3. Median particle number concentrations ($D_p > 0.25 \mu\text{m}$) for waters with different salinities and organic content as a function of water temperature (T_w). Grey shaded areas indicate the 25th and 75th percentiles.

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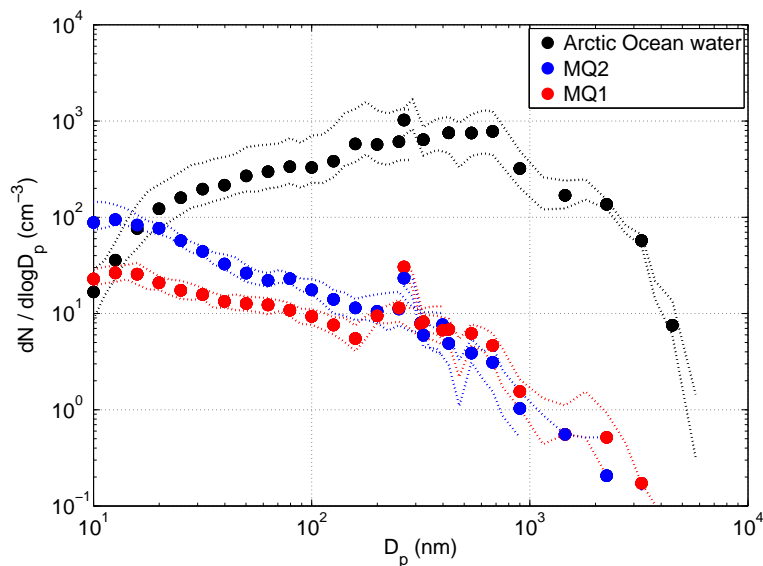


Fig. 4. Number size distribution of Milli-Q water (MQ1 with a median T_w of 3.4°C and MQ2 with a median T_w of 5.6°C) compared to the number size distribution of Arctic Ocean water (median T_w is 3.8°C). The number size distribution of MQ2 was measured about 4 h later than the size distribution of MQ1. Dotted lines represent the 25th and 75th percentiles. Only every second data point is shown for clarity.

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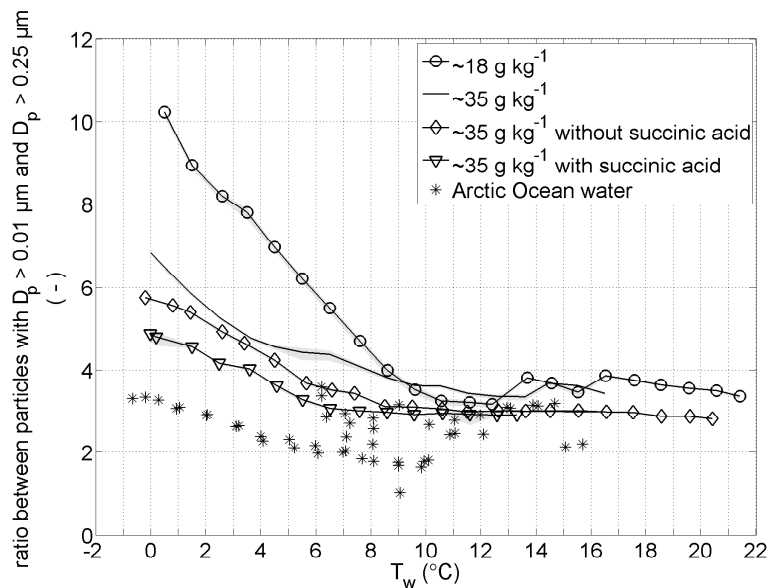


Fig. 5. Ratio between particles $D_p > 0.01 \mu\text{m}$ and $D_p > 0.25 \mu\text{m}$ for waters with different salinities and organic content as a function of water temperature (T_w). For comparison data from Arctic Ocean water is shown in addition. Grey shaded areas indicate the 25th and 75th percentiles.

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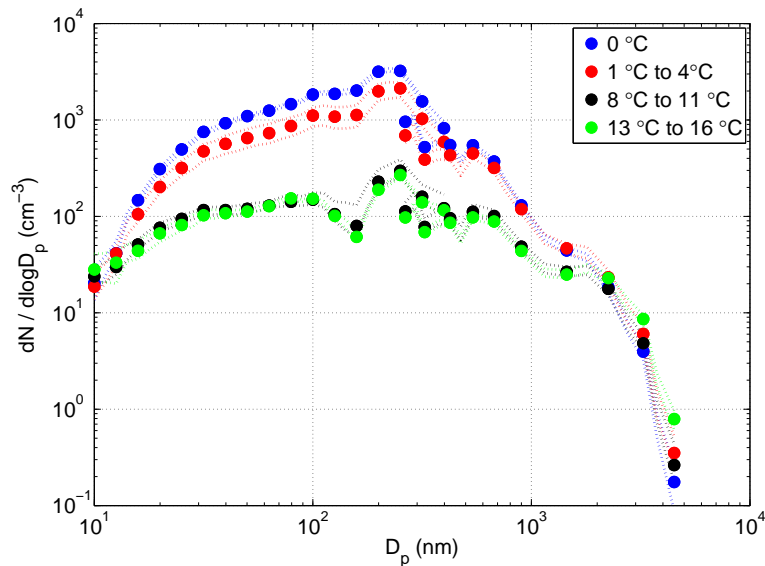


Fig. 6. Particle number size distributions from 35 g kg^{-1} NaCl water for different water temperature ranges. Median size distributions were calculated for temperatures at 0°C , between 1°C and 4°C , between 8°C and 11°C and between 13°C and 16°C . Dotted lines represent the 25th and 75th percentiles. Only every second data point is shown for clarity.

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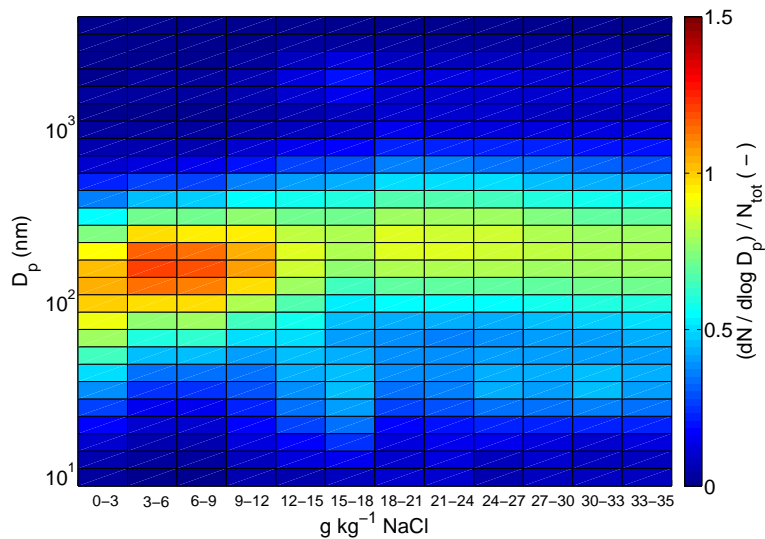


Fig. 7. Normalized fitted particle number size distributions for different salinity ranges.

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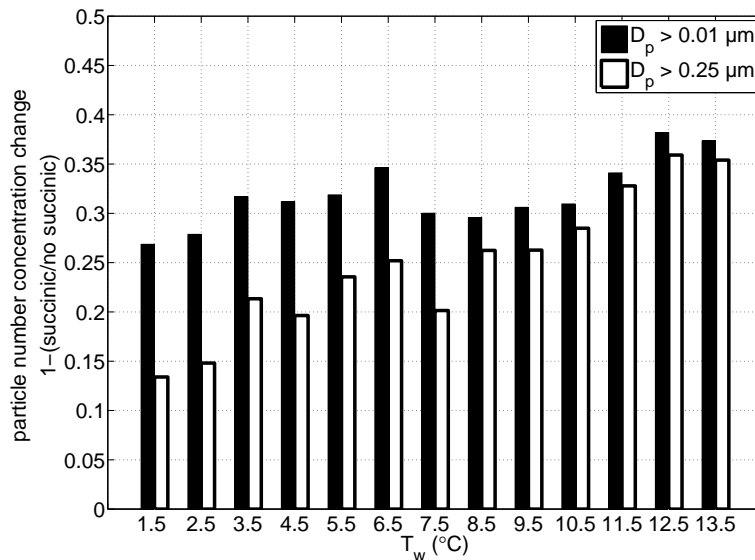


Fig. 8. Change of median particle number concentrations between particle number concentrations after adding succinic acid and before adding succinic acid as a function of water temperature (T_w).

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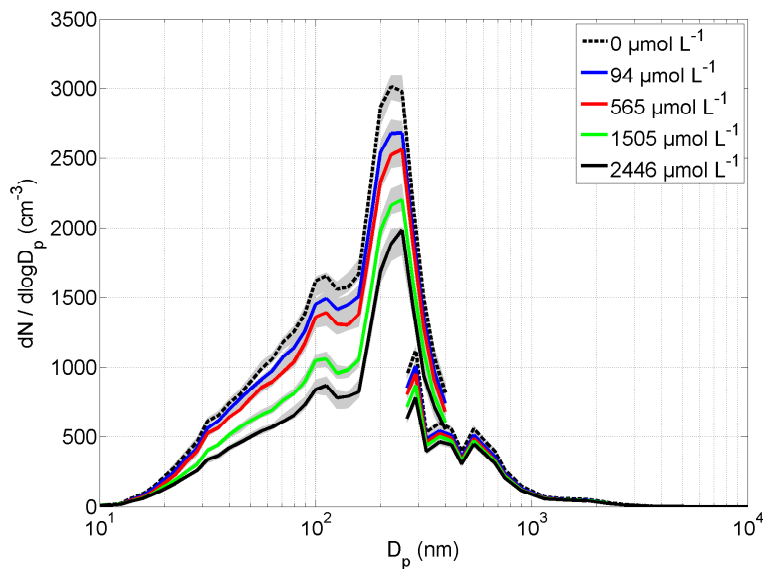


Fig. 9. Particle number size distributions for different succinic acid concentrations. Water temperature medians are between 0.2 and -0.2°C and medians of salinity between 35.4 and 35.7 g kg^{-1} . Measurement time per concentration between 30 and 70 min. Grey shaded areas indicate the area between the 75th and 25th percentiles.

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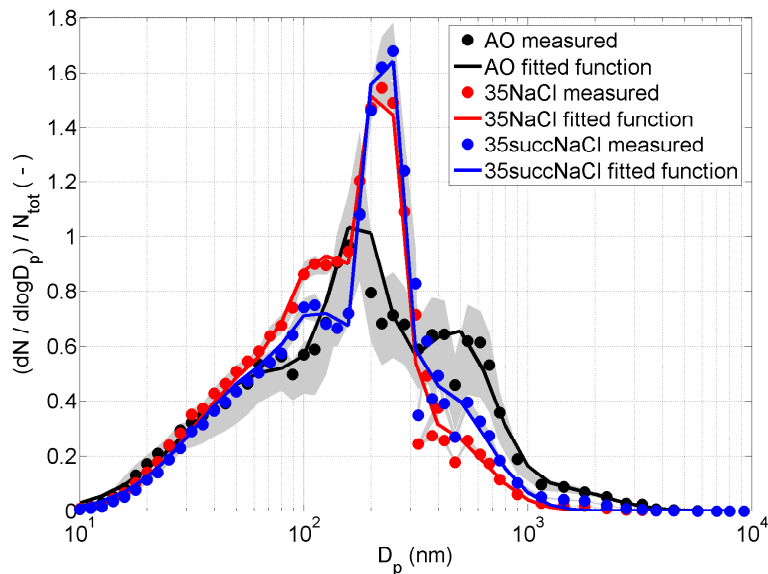


Fig. 10. Comparison between the normalized fitted particle number size distribution of Arctic Ocean water (water temperature median -0.3°C , salinity median 35.0 g kg^{-1}), sodium chloride water without succinic acid (water temperature median is 0.1°C and median salinity is 35.3 g kg^{-1}) and sodium chloride water with succinic acid (water temperature median is -0.1°C and median salinity is 35.6 g kg^{-1}). The normalized size distributions measured and the 25th and 75th percentiles (grey shaded areas) are shown in addition.

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Table 1. Number of number size distributions used to calculate a median size distribution for the Arctic Ocean versus MQ number size distribution comparison.

	Arctic Ocean Water	MQ1	MQ2
$0.01 \mu\text{m} < D_p < 0.389 \mu\text{m}$	-	50	32
$0.01 \mu\text{m} < D_p < 0.281 \mu\text{m}$	138	-	-
$0.265 \mu\text{m} < D_p < 32 \mu\text{m}$	6578	1740	1112

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Table 2. Cooling and warming rates for the different experiments.

	MQ	18 g kg ⁻¹	35 g kg ⁻¹	35 g kg ⁻¹ without succinic acid	35 g kg ⁻¹ with succinic acid
Cooling rate °C h ⁻¹	1.1	1.2	1.4	1.5	-
Warming rate °C h ⁻¹	0.4	0.8	1.1	-	1.1

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Table 3. Number of number size distributions used to calculate the median size distribution for the 35 g kg^{-1} saline water within different temperature ranges.

	0 °C	1 °C to 4 °C	8 °C to 11 °C	13 °C to 16 °C
$0.01 \mu\text{m} < D_p < 0.389 \mu\text{m}$	23	103	59	32
$0.265 \mu\text{m} < D_p < 32 \mu\text{m}$	880	3620	2130	1120

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Table 4. Number of size distributions used to calculate a median size distribution for a comparison of the influence of different succinic acid concentrations on the particle number size distribution.

Succinic acid concentration ($\mu\text{mol l}^{-1}$)	$0.01 \mu\text{m} < D_p < 0.389 \mu\text{m}$	$0.265 \mu\text{m} < D_p < 32 \mu\text{m}$
0	7	490
94	14	980
565	6	430
1505	10	700
2446	9	630

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Table 5. Number of size distributions used to calculate a median size distribution for a comparison of the influence of different constituents on the particle number size distribution.

	Arctic Ocean Water	NaCl 35 g kg ⁻¹	NaCl 35 g kg ⁻¹ with succinic acid
0.01 μm < D _p < 0.389 μm	–	17	21
0.01 μm < D _p < 0.281 μm	8	–	–
0.265 μm < D _p < 32 μm	259	600	770

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Table 6. Parameters for fitted lognormal aerosol number size distributions with N_x as the number concentration (cm^{-3}), $D_{g,x}$ as the geometric diameter (μm) and σ_x as the standard deviation for the different modes and waters: Arctic Ocean water (AO), NaCl water with a salinity of 35 g kg^{-1} and a succinic acid concentration of $0 \mu\text{mol l}^{-1}$ (35 NaCl) and NaCl water with a salinity of 35 g kg^{-1} and a succinic acid concentration of $2446 \mu\text{mol l}^{-1}$ (35succNaCl).

	AO	35NaCl	35succNaCl
N1	430	880	465
Dg1	0.073	0.07	0.07
σ_1	2.3	2.08	2.1
N2	220	340	115
Dg2	0.18	0.124	0.123
σ_2	1.33	1.335	1.33
N3	245	565	307
Dg3	0.48	0.225	0.226
σ_3	1.48	1.20	1.17
N4	35	280	240
Dg4	1.46	0.41	0.41
σ_4	1.59	1.58	1.59

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