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Evidence of an elevated source of nucleation based on model simulations and data from the NIFTy experiment

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Abstract

New particle formation has been observed at a number of ground-based measurement sites. Prior research predominantly from Europe has provided evidence that this new particle formation, while observed in the near-surface layer, is actually occurring in atmospheric layers above the surface and appears to be focused in or close to the residual layer formed by the nocturnal inversion. Here, we present both observations and modeling for southern Indiana, which support this postulate. Based on simulations with a detailed aerosol dynamics model and the Weather Research and Forecasting model, along with data from ground-based remote sensing instruments and detailed gas and particle phase measurements, we show evidence that (i) the maximum rate change of ultrafine particle concentrations as observed close to the surface is always preceded by breakdown of the nocturnal inversion and enhancement of vertical mixing and (ii) simulated particle size distributions exhibit greatest accord with surface observations during and subsequent to nucleation only when initialized with a particle size distribution representative of clear atmospheric conditions, rather than the in situ (ground-level) particle size distribution.

1 Introduction and objectives

Particle nucleation has been observed with high frequency at a geographically diverse suite of ground-based measurement sites (Kulmala et al., 2011) and plays a key role in determining the ambient particle size distribution (Spracklen et al., 2006), but the controls and limitations on nucleation occurrence and growth remain uncertain.

Key uncertainties pertain, in part, to mechanistic deconvolution of the chemical and physics controls and include questions regarding variation of nucleation intensity and mechanisms with height, and specifically whether nucleation occurs principally within the atmospheric boundary layer or is initiated at or close to the residual layer (or free troposphere) (Boulon et al., 2011; Crumeyrolle et al., 2010; Pierce et al., 2011;

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Stratmann et al., 2003; Wehner et al., 2010). Indications of a link between the occurrence and intensity of nucleation and boundary layer dynamics is provided by observations that enhancement of turbulent kinetic energy, associated with entrainment and development of the boundary layer, is frequently observed just prior to the appearance of newly formed particles in the mixed layer (Nilsson et al., 2001; Pryor et al., 2011). Further balloon-borne observations during the SATURN experiment near Leipzig in Germany provided evidence that prior to the break-up of the nocturnal inversion nucleation was focused on the residual layer and, subsequent to erosion of the inversion and growth of the boundary layer, nucleation was observed throughout the planetary boundary layer (Stratmann et al., 2003; Wehner et al., 2007). Further evidence for an elevated source of nucleated particles at continental sites in Europe was provided by data collected near Cabauw in the Netherlands in the IMPACT field campaign (Wehner et al., 2010). In that research vertical profiles of particle size distributions were taken in a helicopter borne package and the evolution of the boundary layer was observed using a lidar system. As in the SATURN experiment, analysis of the IMPACT measurements led the authors to infer that “it is very likely that these particles observed at the ground were formed at higher altitudes and mixed downwards” (Wehner et al., 2010). The IMPACT authors postulate that the strong vertical variability of new particle formation is a result of enhanced turbulence and non-linear mixing in specific layers which resulted in supersaturation of precursor gases coupled with vertical heterogeneity of in situ aerosol concentrations and thus variations in the condensational sink leading to enhancement or suppression of nucleation. Here we examine this issue further using: (i) data from the Nucleation In Forests (NIFTy) experiment in southern Indiana in conjunction with simulations conducted using the Weather Research and Forecasting (WRF) model and (ii) a novel approach based on the application of a parameterized multi-component particle model (i.e. the University of Helsinki Multicomponent Aerosol (UHMA) model). Specifically we use the UHMA model to examine the sensitivity of nucleation and growth to the prevailing chemical and physical environmental conditions (i.e. abundance of nucleation precursors and condensational sink) in order to theorize

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the region of the atmosphere in which nucleation is initiated. On the basis of the sensitivity analysis conducted using as input to the model conditions observed close to the ground versus those associated with an elevated source of nucleation we infer that nucleation is initiated aloft. Supporting evidence for these assertions is drawn from

5 ground-based particle size distribution measurements, observed and modeled profiles of the meteorological state parameters and lidar backscatter measurements.

2 Methods

2.1 NIFTy

During the Nucleation In Forests (NIFTy) experiment, conducted between 5 and 31

10 May 2008, particle physical and chemical properties, gas phase concentrations and meteorological parameters of state were measured at three locations along an 80 km transect in southern Indiana from Indianapolis in the northeast to the small college town of Bloomington in the southwest (Pryor et al., 2011). The primary measurements used herein were collected during the later portion of the NIFTy experiment (14–27 May) at

15 the Morgan Monroe State Forest (MMSF) tower site, centrally located along this transect. A Scanning Mobility Particle Sizer (SMPS) system from TSI Inc. (SMPS 3936) and a Fast Mobility Particle Sizer (FMPS 3091) (TSI, Inc.) were deployed at MMSF to measure simultaneously particle size distributions from a height of 46 m (above a canopy of 28 m). Measured particle concentrations were corrected using the exper-

20 imentally derived tubing particle transmission efficiencies presented in (Pryor et al., 2010). Thirty-minute average concentrations of SO₂ and H₂SO₄ were measured below the forest canopy using a TECO (model 43S) monitor and a Chemical Ionization Mass Spectrometer (CIMS) (Eisele and Tanner, 1993; Mauldin et al., 2003; Petäjä et al., 2009). VOC concentrations were measured at the canopy top in six approximately 2-h

25 intervals starting at 09:00 (LST) on multi-sorbent cartridges and analyzed for isoprene, α -pinene, limonene, cumene (isopropylbenzene), benzene, toluene, ethyl benzene, o-,

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m-, and p-xylene. Meteorological conditions at the site were measured using micrometeorological equipment installed on this AmeriFlux tower, a scanning doppler lidar (Natural Power ZephIR lidar), a Vaisala tethersonde and a Vaisala ceilometer (CL31) (see Pryor et al. 2011 for a full description of the instrumentation deployed at MMSF). During NIFTy particle size distribution measurements were also made using an FMPS and SMPS, respectively at a site in Bloomington (18 km SW of MMSF) and in Indianapolis (60 km NE of MMSF).

Long-term measurements at the MMSF site indicate evidence of elevated concentrations of ultra-fine particles (D_p : 6–30 nm) on approximately 1 day in 5 (Pryor et al., 2010), with a highest frequency in May, consistent with observed high concentrations of ultra-fine particles on nearly half of all sampling days during NIFTy (Table 1) (Pryor et al., 2011).

2.2 UHMA model

The UHMA model is a box-model containing parameterizations of the dynamics of multicomponent particles including nucleation (based on parameterization of binary, ternary and kinetic nucleation), condensation, coagulation and dry deposition (Korhonen et al., 2004). In this study, the particle size distribution (PSD) is treated using a fixed sectional discretization of the size distribution in the modified version of the hybrid grid (Jacobson and Turco, 1995), which does not require any assumption regarding the PSD and allows us to describe accurately the variability associated with field measurements (Spracklen et al., 2005). Since only condensed core compounds need to be split among size sections, numerical diffusion is reduced (Jacobson and Turco, 1995). Herein we employ the UHMA model using the following assumptions:

- *The temporal variation of boundary layer depth* (and thus box model volume in which concentrations are assumed to be homogeneous) is simulated using a pseudo-sinusoidal profile evolving between 06:00 and 21:00 (LST) with maximum of 1000 m and minimum of 300 m (this minimum is maintained throughout the

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nighttime hours) based on data obtained from a tether sonde system deployed during NIFTy and confirmed based on simulations conducted using the WRF model. All other physical state parameters (e.g. temperature and humidity) are input hourly to the model based on observed conditions at 46 m.

- *Particle removal.* Within the original UHMA model dry deposition processes are described by a semi-empirical parameterization validated for the boreal forest in Hyytiälä, Finland (Rannik et al., 2003). Deposition velocities over deciduous forests have been shown to be lower than those measured over boreal forests (Pryor et al., 2009), thus we implemented a physical parameterization of dry deposition (Slinn, 1982; Pryor and Binkowski, 2004) constrained to match observed size-resolved particle deposition velocities of sub-100 nm diameter particles as measured at the MMSF site (Pryor et al., 2009). No removal by wet deposition was parameterized.
- *A simplified gas phase chemical mechanism* was implemented following the work of (Boy et al., 2005) (Tables 2 and 3) and was solved using the Euler Backward Iterative approach. Observed concentrations of the inorganic gases (sulfuric dioxide (SO₂) and ammonia (NH₃)) and condensable organic gases were provided at hourly resolution to the model. Because we did not have direct observations of semi-volatile organic compounds, the concentration of the condensable organics was estimated from the observations of isoprene, α-pinene, cumene, limonene, benzene, toluene, ethyl-benzene, (m,p)-xylene, and (o)-xylene using Fractional Aerosol Coefficients (FAC) (Grosjean, 1992). It is acknowledged that this approach neglects the availability of oxidants and the relationship with condensed mass (Griffin et al., 2003), but is applied here to broadly represent the potential for the production of oxidation products with low volatility that might partition into the particle phase.
- Comprehensive evaluations of *nucleation parameterizations* (Zhang et al., 2010a, b; Sihto et al., 2006) have illustrated the challenge of selecting a theoretically

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appropriate nucleation parameterization and the huge range of variation of model skill when predictions based on differing parameterizations are compared to observed number concentrations (e.g. up to 3 orders of magnitude variations for the Aitken and accumulation mode). In order to select which of the nucleation parameterizations was optimal for the simulations in southern Indiana, we used the approach of Sihto et al. (2006) and Kuang et al. (2008) to examine the functional relationship between ultra-fine particle nucleation rate and observed sulfuric acid concentrations ($[\text{H}_2\text{SO}_4]$). If nucleation can be described using the activation approach, then the formation rate of new particles should scale linearly with $[\text{H}_2\text{SO}_4]$. If the nucleation process is best described by the kinetic approach then the formation rate will tend to scale with the square of $[\text{H}_2\text{SO}_4]$, while if it follows the ternary nucleation theory, with the critical cluster comprising NH_3 , H_2SO_4 and water molecules, the exponent should exceed 3 (Sihto et al., 2009). Thus:

$$J_{(1 \text{ nm})} = K \times [\text{H}_2\text{SO}_4]^n \quad (1)$$

where n represents the number of sulfuric acid molecules in each critical cluster. The best linear fit of the logarithm of the nucleation rate and the logarithm of sulfuric acid concentration for class A events during NIFTy provided a value of n of 1.78 and K of around 10^{-14} (Fig. 1). J_1 was computed based on (Kuang et al., 2008):

$$J_1(t) = J_{10}(t + \Delta t) \times \exp \left[\frac{1}{2} \times \frac{A_{\text{FUCHS}}}{GR_{1-10}} \times \sqrt{\frac{48k_{\text{BT}}}{\pi^2 \rho}} \left(\frac{1}{\sqrt{D_{\rho 1}}} - \frac{1}{\sqrt{D_{\rho 10}}} \right) \right] \quad (2)$$

Where J_{10} was derived from the measured PSD from the SMPS and the median Fuchs surface area of the pre-existing particles during $t + \Delta t$ was computed as:

$$A_{\text{FUCHS}} = \frac{4}{3} \pi \times \int_{D_{p6}}^{D_{p10}} D_p^2 \times \left(\frac{Kn + Kn^2}{1 + 1.71Kn + 1.33Kn^2} \right) n(D_p) dD_p \quad (3)$$

where $n(D_p)$ is the number concentration of particles of diameter D_p . The Knudsen number is defined as:

$$Kn = \frac{2\lambda}{D_p} \quad (4)$$

where λ is assumed to be 100 nm (Kuang et al., 2008), the particle density ρ is 1.3 kg m^{-3} and k_B is the Boltzmann constant. An average Δt of 0.5 h was estimated from the time shift required to match the measured sulfuric acid concentrations with the 6–10 nm diameter average particle number concentrations. Based on the results of the analysis of the nucleation rate as a function of sulfuric acid concentrations we chose to apply a kinetic approach to nucleation where the prefactor K is derived from the basic kinetic theory, assuming nucleation to be limited by the collision rate of sulfuric acid molecules (Seinfeld and Pandis, 2006), rather than estimated from observations (Sihto et al., 2006).

We need to acknowledge the regression fit between $\log[\text{H}_2\text{SO}_4]$ and $\log(J_1)$ is highly sensitive to the fitting time length since a fit conducted only on the duration of each nucleation event would provide higher slope values (Kuang et al., 2008). Because of the relative low temporal resolution of the available sulfuric acid measurements we fitted the model over a larger time span, than has been previously used.

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– *Initial particle size distribution.* For the base case simulations, the model was initialized using the PSD measured by the FMPS at midnight (LST) of each simulated day (14–27 May 2008). In order to reduce the sensitivity to measurement uncertainty in each size bin, the data from the FMPS were fitted to three log-normal modes (see the example in Fig. 2). The geometric mean diameter, standard deviation and number concentration from these modes were then used to derive the initial number concentration in each of the model size sections. The model was then run for 24 h and evaluated over the same time period.

The time-step of the simulations was set to 10 s. A sensitivity analysis was undertaken to assess model performance as a function of the number of sections used to describe the PSD, but for all other simulations, 250 sections, equally spaced in a logarithmic scale in the range of diameters between 3 nm and 10 μm , were used to match the discretization of the PSD for the sub 100-nm particles from the SMPS operated at MMSF in order to facilitate the evaluation of the model fit to the data.

Since the UHMA model has been previously theoretically validated (Korhonen et al., 2004), herein we use statistical metrics not to evaluate the model per se but as a diagnostic tool. Three primary metrics of model performance are used:

a. The presence, absence and “type” of nucleation event as defined using a subjective event classification (Boy and Kulmala, 2002):

class A: a new sudden particle mode appears in the diameter range below 25 nm and it persists and grows for more than 1 h

class B: a new particle mode is present but it is not visible at the smallest measured diameters. The computation of the growth rate may be uncertain because of high variations in the mode number concentration

class C: a new particle mode is present but does not show a clear growth

The day was defined as non-event if the aforementioned conditions are not met (see Table 1 for a summary of the observed data from the SMPS and FMPS in terms of the event classification).

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- b. The skill of the UHMA model in predicting the particle number size distribution is quantified by metrics such as nucleation intensity, the growth rate and the accuracy and timing of the peak number concentration. Nucleation intensity is quantified as the total number of particles with diameter between 10 and 30 nm during the two hours with the highest total number concentration:

$$\text{Intensity} = \sum_{h=1}^2 \sum_{D_p=10}^{30} N_{h,D_p} \quad (5)$$

The growth rate is defined as the rate at which the number geometric mean diameter ($\overline{D_{gN}}$) in the nucleation mode (6–30 nm) evolves:

$$\ln \overline{D_{gN}} = \frac{\sum_{j=6}^{30} n_j \times \ln D_j}{N_{6-30}} \quad (6)$$

where $\ln \overline{D_{gN}}$ is the value of $\ln D$ weighted according to the number of particles in that size interval, n_j is the number of particles in a group whose diameters are centered around D_j and N_{6-30} is the total number of particles in the diameter range of 6–30 nm. The growth rate is computed fitting a first order polynomial to the number geometric mean diameters occurring in the three hours subsequent to the minimum $\overline{D_{gN}}$.

- c. We also analyzed the accuracy of prediction of the peak number concentration and its timing, defined as the ratio between the maximum simulated ($\max N_{6-10}^{\text{sim}}$) and observed ($\max N_{6-10}^{\text{obs}}$) number concentration of particles of size 6–10 nm and the time difference (Δt) at which they occur respectively:

$$\text{accuracy} = \frac{\max N_{6-10}^{\text{sim}}}{\max N_{6-10}^{\text{obs}}} \quad (7)$$

$$\Delta t = t \left(\max N_{6-10}^{\text{obs}} \right) - t \left(\max N_{6-10}^{\text{sim}} \right) \quad (8)$$

2.3 WRF simulations

The Weather Research and Forecasting model Version 3 (WRFV3), applied using the physics schemes listed in Table 4, was used to simulate meteorological conditions during 11–26 August 2008 over a parent domain (324×274 grid cells with a spatial resolution of 9 km) and a nested domain (310×259 grid cells with a spatial resolution of 3 km) (Fig. 3). The model was run using the Advanced Research WRF (ARW) solver in the non-hydrostatic mode with 35 vertical levels and initialized with lateral boundary conditions as simulated by the North American Mesoscale Model (NAM). The land cover data were specified from the USGS 24-category data at a resolution of 3.7 km and 0.9 km for the parent and nested domain respectively. The WRF simulations were conducted principally to examine the vertical evolution of the planetary boundary layer and thus the primary physics scheme of interest to the current work is the PBL scheme. The Mellor-Yamada-Janjic scheme (a TKE-based 2.5-level closure) was selected because it is relatively computationally efficient, has been widely used in prior research and has been demonstrated to generate relatively representative temperature profiles and PBL heights in applications elsewhere (Hu et al., 2010).

3 Results

3.1 Evidence for an elevated source of nucleation from observations

As described below, observations of meteorological parameters at the MMSF site support the hypothesis of nucleation initiation above the surface with subsequent entrainment of the freshly nucleated particles into the mixed layer. In order to make composites of multiple event days in this analysis, following the approach used in Pryor et al. (2011),

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all times were converted to a normalized scale in which the time is presented relative to the maximum gradient (i.e. rate change) in 10 nm particle number concentrations.

Turbulence intensity as measured with the ZephIR lidar indicates a strong link between the occurrence of high concentrations of ultra-fine particles at the MMSF site and boundary layer dynamics. The greatest increase in the number concentration of 6–10 nm particles occurs approximately one hour prior to the peak of nucleation intensity at 46 m during class A events (Fig. 4a). Nearly simultaneously turbulence intensity rapidly increases from low values throughout the lowest 300 m of the atmosphere (representative of a stable nocturnal atmosphere) to more turbulent conditions associated with erosion of the nocturnal inversion and development of a fully mixed planetary boundary layer (Fig. 4b). Thus erosion of the nocturnal inversion and increased vertical mixing always preceded the maximum concentration of ultra-fine particles and was observed to precede the maximum rate change of ultrafine particle concentrations by approximately 0.5 h.

Further evidence supporting the hypothesis of an elevated source of nucleation is provided by the analysis of the condensational sink (CS) (Kulmala et al., 2001; Dal Maso et al., 2002) measured in the surface layer (at 46 m) during event class A days. The CS is a metric quantifying the role of preexisting particles in removing condensable vapors from the atmosphere, thus a high CS is expected to suppress nucleation. As in research conducted elsewhere, there is a lack of correspondence between the observed CS in near-surface measurements just prior to the largest rate change of ultrafine particles and the concentration of ultrafine particles (Fig. 5) (Boulon et al., 2011). This also supports our speculation of an elevated initiation of nucleation.

The Vaisala ceilometers has a wavelength of 910 nm and thus shows strongest response to the accumulation mode particles that also tend to dominate the PSD of sulfuric acid and thus the condensational sink. The ceilometer backscatter measurements were processed following the approach of Stratmann et al. (2003), to examine evidence for elevated atmospheric layers with relatively low aerosol loading (i.e. low backscatter) and thus more favorable for nucleation occurrence (i.e. lower CS). In this

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approach the backscatter signals from each 20 m increment in each 5 min period from each day is normalized to the average backscatter at that height over the period 06:00–12:00 (LST). The backscatter profiles were integrated with vertical temperature profiles derived from simulations with the WRF model (Fig. 6) to examine the presence of a clear inversion and residual layer and to link the erosion of this layer to the appearance of elevated concentrations of ultrafine particles. A couple of hours prior to nucleation appearance in the surface layer, event class A days are always characterized by a strong temperature inversion with a base located at approximately 400 m (Fig. 6g). This inversion is much weaker during event class B and class C days (Fig. 6h) and completely absent during non-event days which are usually characterized by greater cloud cover (Fig. 6i). During event days, at the time newly formed particles are detected close to the surface, the nocturnal inversion is almost completely eroded. Moreover the relatively low elevation of the base of the residual layer (i.e. 600–700 m), coupled with intense mixing phenomena due to boundary layer dynamics, may favor the advection of new particles formed aloft into the mixed layer. Combining the vertical profiles of temperature with the relative backscatter signal derived from ceilometer data provides direct evidence of boundary layer dynamics and mixing processes. The ceilometers data indicate that on event class A days there was a clear evidence for a low backscatter layer associated with the inversion and residual layer (Fig. 6a). Just prior to the maximum rate change of 10 nm particles in the surface observations the inversion is eroded (Fig. 6), during the transition to a fully mixed layer with high particle number concentration. This transition often occurs simultaneously with or earlier than detection of substantial numbers of sub 10-nm particles detected close to the ground, which strongly supports our hypothesis of particle entrainment from aloft.

Indirect evidence of nucleation being initiated aloft may be derived by examining the concentrations of the smallest detected particles during event class A days at the MMSF site. The ratio between the number concentration of 3 nm and 6 nm particles in the hour of maximum 6 nm particle concentrations (computed after applying the empirical transmission efficiency correction presented in Pryor et al. (2010)) is on average

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60 % and thus is not high enough to explain the observed 6 nm particle concentrations. The low concentrations of 3 nm particles might be related to limitations of the deployed instrument and the experimental setup but may also reflect early aging (i.e. growth by condensation and coagulation) of freshly nucleated particles formed elsewhere from the sampling site as also speculated by Pierce et al. (2011). Following the logic presented in Birmili and Wiedensohler (2000), the presence of a closed contour in the particle number size distribution profile (see Fig. 7) is also indicative of non-local nucleation (i.e. that the particles being observed were formed earlier – in this case in an elevated layer).

3.2 Evidence for an elevated source of atmospheric particles derived from model analyses

A key consideration in evaluating any particle dynamics model is the sensitivity of the model to the description of the particle size distribution (Zhang et al., 1999). Indeed, reconciling the need for computational efficiency and adequate representation of the PSD is a key factor for implementing particle dynamics within regional and global climate models (Spracklen et al., 2006). A model sensitivity analysis to the size distribution discretization was conducted to define the optimum number of sections as a function of model performance and computational time (Table 5). The results show little sensitivity to the number of sections for higher number of sections (i.e. for simulations conducted using 300 sections rather than 250), hence 250 sections (over the whole simulated range of diameters: i.e. between 3 nm and 10 μ m) were selected to match the size resolution of the SMPS operated at MMSF and thus has 80 sections in the size range 3–105 nm). However, it is worthy of note that based on our sensitivity analyses it appears that model skill improves only slightly setting more than 80 bins, thus given the computational time required by 250 sections is around nine times longer than assuming 80 sections, fewer than 250 sections may be used for longer term simulations.

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When driven by measured sulfur dioxide concentrations and initialized using the observed PSD at midnight (LST) on each individual day, the UHMA model correctly simulates the occurrence or absence of nucleation (Table 1). However, these simulations indicate a clear systematic under-prediction of the growth rate (Table 5 and Fig. 8) and typically an underestimation of nucleation intensity (Table 1). Further there is a marked offset in the timing of the maximum concentration of particle numbers for $D_p = 6-10$ nm (Table 5). There are a number of possible causes for these systematic biases since the simulated particle number concentrations are determined by the balance between formation processes (the nucleation rate scales with the square of sulfuric acid concentrations) and removal processes (dry deposition and condensational sink). The kinetic nucleation approach was identified to be the dominant nucleation mechanism and since the dry deposition algorithm has been reformulated according to a physically based data-constrained description, thus we examined three other possible causal mechanisms of these biases using model sensitivity analyses of:

- *The initial particle loading.* In order to investigate the model sensitivity to the initial PSD, we ran a simulation wherein UHMA was initialized using the measured PSD on May 14 and then left to run for the entire 14 day period. A consistent underestimation of growth rates remained (cf. Fig. 9a and b). An additional set of simulations was performed initializing the model with a PSD representative of “clear” atmospheric conditions as described by Seigneur et al. (1986) (Table 6). This idealized PSD is applied as being representative of the lower backscatter associated with the residual layer. When this idealized PSD is used to initialize the UHMA model, the simulated nucleation intensity shows better accord with the observations (Table 1) and event class A growth rates are closer to the observed values, although a small negative bias in GR remains (Fig. 10 and Table 5). The intensity of nucleation is over-estimated relative to the ground-based measurements in simulations with the clear case PSD but this is rather consistent with nucleation occurring in a relatively shallow atmospheric layer and then being mixed, diluted and reduced by coagulation during transport to the ground-

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based observing systems. However, initializing the model with clear atmospheric conditions also appears to overestimate formation rates during non-event days. Therefore we suggest that factors other than the initial particle loading may also play a major role in controlling nucleation occurrence.

- *The availability of nucleation precursors.* Since the kinetic approach assumes the nucleation rate is dependent on the square of sulfuric acid concentrations, part of the underestimation of simulated growth rates may be attributed to the negative bias in simulated sulfuric acid concentrations (Fig. 11). In the simulations, $[\text{H}_2\text{SO}_4]$ results from the reaction between sulfur dioxide measured close to the ground and photochemically produced OH. The chemical mechanism shows good skill in simulating OH concentrations relative to observations. The simulated average peak concentration during event class A days is $4.21 \times 10^6 \text{ cm}^{-3}$ compared to the observed value of $4 \times 10^6 \text{ cm}^{-3}$, whereas the average simulated maximum $[\text{OH}]$ during non-event days is $2.72 \times 10^6 \text{ cm}^{-3}$, compared to measured value of $8 \times 10^5 \text{ cm}^{-3}$ (Pryor et al., 2011). Thus we speculate that the negative bias in simulated $[\text{H}_2\text{SO}_4]$ is due to a negative bias in SO_2 concentrations due to their measurement below the forest canopy. To investigate if the systematic bias in $[\text{H}_2\text{SO}_4]$ was responsible for a negative bias in simulated nucleation an empirical correction factor (derived by fitting modeled $[\text{H}_2\text{SO}_4]$ to the observations, Fig. 11) was applied to the modeled values. Thus a simulation was conducted in which the model was initialized each day using the observed PSD and the concentration of $[\text{H}_2\text{SO}_4]$ was enhanced. Simulated growth rates are increased by an average factor of 1.2 and the agreement with observed values improved, but they remain consistently lower than those computed from both the clear case simulations and observations (Table 5). This underestimation of growth rates even with enhanced $[\text{H}_2\text{SO}_4]$ is also observed when the model is initialized on May and left to run for the entire 14 day period (Fig. 12a).

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- Another factor regulating the *availability of condensable vapors* and the growth rate of ultrafine particles is the mass accommodation coefficient which describes the probability of a gas molecule will stick to the pre-existing particles. Mass accommodation coefficient values lower than 1 may thus enhance the concentration of condensable compounds available for nucleation hence reducing the influence of the CS. When the mass accommodation coefficient was set to a value of 0.43 (based on work by Pöschl et al. (1998)), UHMA simulations indicate only a weak sensitivity to this parameter (Fig. 12b and Table 5).

4 Discussion and concluding remarks

In this study we provide independent evidence, both from observations and modeling results, supporting the hypothesis of an elevated source of nucleation in the residual layer, as also advanced at other, mostly European, sites (Boulon et al., 2011; Pierce et al., 2011; Stratmann et al., 2003). Evidence in support of this postulate derived from observations at the MMSF site located in southern Indiana and thus within the polluted Ohio River Valley include:

- There is a clear relationship between the occurrence of elevated ultrafine particle concentrations and enhancement of turbulence intensity consistent with erosion of the nocturnal boundary layer and vertical transfer of recently nucleated particles (Fig. 4).
- The absence of a relationship between the near-surface condensational sink one hour prior to the appearance of significant ultrafine particle concentrations and concentrations of sub 10-nm particles (Fig. 5). Based on backscatter data from a ceilometer, atmospheric conditions more favorable to new particle formation (i.e. lower CS) are found in the residual layer (Fig. 6).
- The primary distinction between event and non-event days, in terms of boundary layer dynamics, is the consistent presence of a capping inversion on event class A

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days which is always eroded by the time of nucleation observation in the surface layer (Fig. 6g). This modeling evidence strongly supports the hypothesis of new particles formed in the residual layer and then mixed down to the surface layer with the erosion of the nocturnal inversion.

- The low ratio between 3 nm and 6 nm particle number concentrations measured in the near-surface layer at the MMSF site may be the result of an early aging of freshly nucleated particles before detection during their advection from higher atmospheric levels to the mixed layer.

Model sensitivity analyses indicate nucleation rates appear to be highly dependent on the PSD used to initialize the simulations with generally lower sensitivity to H₂SO₄ availability, the mass accommodation coefficient and the number of sections used to represent the PSD. Thus results from the UHMA simulations that also support the postulate that nucleation is not occurring in the near-surface layer include the following:

- The base case simulations with UHMA, driven by the PSD measured at 46 m and sulfur dioxide concentrations measured close to the surface, indicate consistent underestimation of nucleation intensities and growth rates (Figs. 8, 9 and 10, and Tables 1 and 5). Higher agreement between observed and simulated PSD is obtained when the initial PSD is set to conditions representative of clear atmospheric conditions (Fig. 10b) and when higher sulfuric acid concentrations are applied (Fig. 12a). Both of these conditions may be found at higher levels in the atmosphere and thus lead us to infer that nucleation is initiated aloft and the newly formed particles are then mixed downward with the breaking of the residual layer and formation of the convective boundary layer (Pierce et al., 2011; Stratmann et al., 2003; Wehner et al., 2010).
- Simulated peak number concentrations in the size range of 6-10 nm exhibit a positive bias compared to observations both in the case of the initialization with the observed PSD and with clear atmospheric conditions. However, the time at which

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5 this peak occurs is on average shifted towards an earlier nucleation start when assuming clear atmospheric conditions. This result supports our hypothesis of new particles formed aloft (e.g. in the residual layer), then mixed down and diluted with the formation of the mixed layer (thus also explaining the slightly lower observed nucleation intensities compared to simulated values). The greatest temporal gradient (i.e. rate change) of 6–10 nm particle number concentrations occurs in the hour prior to the maximum in near-surface observations which is also consistent with the postulate that nucleation occurs aloft and also implies an essential role for boundary layer dynamics in observations of elevated ultrafine particle concentrations in the surface layer.

15 Although the postulates and findings advanced here cannot be considered conclusive, they are consistent with prior research that has indicated that new particle formation likely occurs principally at or close to the residual layer (or free troposphere) (Boulon et al., 2011; Lauros et al., 2011; Pierce et al., 2011; Stratmann et al., 2003; Wehner et al., 2010). Given the key role played by boundary layer dynamics and the location of the capping inversion in dictating the occurrence of high particle concentrations near the surface, findings from this study may provide insights for the vertical resolution required by regional models to adequately represent aerosol dynamics.

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Table 1. Classification of event days at the MMSF site during the NIFTy experiment for 14–27 May based on the subjective classification of (Boy and Kulmala, 2002). The hour of maximum rate change of 10 nm particle concentrations (start hour) is provided in local standard time (LST). Observed nucleation intensities [cm^{-3}], computed as the number of particles with diameter between 10 and 30 nm in the two hours of highest concentration, are shown in the 4th column. The final three columns show model simulation results setting 250 sections and initializing the model with the measured particle size distribution (measured PSD) and clear atmospheric conditions (clear case) wherein the PSD is as described by (Seigneur et al., 1986).

	Observations SMPS (and FMPS)			Simulations		
Day	Event	Start hour (LST)	Nucleation Intensity [cm ⁻³] from the SMPS	Initialized with measured PSD	Nucleation Intensity [cm ⁻³] (measured PSD)	Nucleation Intensity [cm ⁻³] (clear case)
14	Non-event (Non-event)	Non-event	Non-event	Non-event	Non-event	Non-event
15	Non-event (Non-event)	Non-event	Non-event	Non-event	Non-event	Non-event
16	C (C)	9	3.58×10^4	A	1.52×10^5	1.89×10^5
17	A (A)	9	1.06×10^5	A	3.82×10^4	1.54×10^5
18	C (C)	10	5.17×10^4	C	4.56×10^1	7.45×10^4
19	A (A)	9	6.95×10^4	C	8.31×10^1	1.58×10^5
20	B (B)	13	1.18×10^4	A	5.59×10^2	9.16×10^4
21	A (A)	8	2.47×10^4	A	1.88×10^2	8.65×10^4
22	A (A)	9	4.90×10^4	A	5.32×10^4	1.43×10^5
23	Non-event (Non-event)	Non-event	Non-event	Non-event	Non-event	Non-event
24	A (A)	10	2.22×10^4	A	2.30×10^4	1.90×10^5
25	A (C)	9	5.37×10^4	A	3.90×10^3	1.86×10^5
26	Non-event (Non-event)	Non-event	Non-event	Non-event	Non-event	Non-event
27	Non-event (Non-event)	Non-event	Non-event	Non-event	Non-event	Non-event

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Table 2. Chemical reactions implemented within the UHMA model (Reference 1 = Hertel et al. (1993), 2 = Boy et al. (2006), 3 = Seinfeld and Pandis (2006), 4 = Simpson et al. (1997)). Second order rate constants are expressed in [ppb⁻¹ s⁻¹] for a temperature of 298K. Photolysis rate constants are expressed in s⁻¹ and computed as $J = I \times (\cos\Theta)^m \times e^{-\eta \times \sec\Theta}$ wherein Θ is the solar zenith angle. Typical peak values of photolysis rate constants for the MMSF site during a clear sky day in May are reported.

Reaction	Rate constant	Reference
SO ₂ + OH = H ₂ SO ₄ + HO ₂	2.50×10^{-2}	1
NO ₂ + OH = HNO ₃	2.72×10^{-1}	1
OH + CO = HO ₂	6.67×10^{-3}	1
O ₃ + OH = HO ₂	1.67×10^{-3}	1
HCHO + OH = HO ₂ + CO	2.50×10^{-1}	1
HO ₂ + HO ₂ = H ₂ O ₂ + O ₂	6.91×10^{-2}	1
organics + OH = products	1.53×10^{-1}	1
H ₂ + OH = HO ₂ + H ₂ O	1.65×10^{-1}	2
HO ₂ + NO = OH + NO ₂	2.00×10^2	3
HO ₂ + O ₃ = OH + 2O ₂	4.76×10^{-2}	3
O ₃ = 0.2OH + 0.8O ₃ (P)	3.05×10^{-4}	4
HCHO = 2HO ₂ + CO	3.21×10^{-4}	4

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Table 3. Input concentrations [ppb] to the chemical mechanism based on Boy and Kulmala (2002) and typical spring values measured in Southern Indiana by the Indiana Department of Environmental Management.

Compound	Concentration [ppb]
O ₃	30
NO ₂	5
NO	2.5
CO	800
HCHO	10
CH ₄	1900
H ₂	500
HO ₂	3.7×10^{-3}

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Table 4. Physics schemes (Skamarock et al., 2008) adopted for the WRF simulations.

Physics option	Adopted scheme
Microphysics	WRF Single-Moment 6-class
Longwave Radiation	Rapid Radiative Transfer Model (RRTM)
Shortwave Radiation	Dudhia
Surface layer	Eta
Land Surface	Noah Land Surface Model
Planetary boundary layer	Mellor-Yamada-Janjic
Cumulus parameterizations	Kain-Fritsch

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Table 5. Measured and simulated growth rates [nm h^{-1}] for class A events. Column 1 indicates the day in May 2008, column 2 the observed growth rates (GR) and column 3 shows simulated GR computed initializing the model with clear atmospheric conditions and setting 250 sections. Columns 4 and 5 refer to GR simulated initializing the model with measured PSD and setting 250 sections when the sulfuric acid (H_2SO_4) and mass accommodation coefficient (α) correction are applied respectively, as described in Sect. 3.2. Columns 6–10 summarize the results from the sensitivity analysis of simulated growth rates on the number of sections (20, 80, 150, 250 and 300 sections) when the model is initialized with measured PSD. The ratio between simulated and measured maximum 6–10 nm particle number concentrations and their timing ($\Delta t = t_{\text{obs}} - t_{\text{sim}}$), computed initializing the model each day with the measured PSD and with clear atmospheric conditions, are also reported (columns 11–14).

Day	Observations	Simulations											
	GR (nm h ⁻¹)	GR (nm h ⁻¹):	GR (nm h ⁻¹): initialization with measured PSD							Initialization with measured PSD		Initialization with clear case	
		initialization	Number of sections							$\frac{\max[N_{6-10 \text{ nm}}]_{\text{sim}}}{\max[N_{6-10 \text{ nm}}]_{\text{meas}}}$	Δt (h)	$\frac{\max[N_{6-10 \text{ nm}}]_{\text{sim}}}{\max[N_{6-10 \text{ nm}}]_{\text{meas}}}$	Δt (h)
			with clear case	H ₂ SO ₄ correction	$\alpha = 0.43$	20	80	150	250	300			
17	1.43	1.50	1.28	0.64	0.32	0.58	0.64	0.64	0.64	1.94	-2	5.41	+1
19	2.13	1.68	0.90	0.45	0.26	0.96	1.08	0.96	0.93	0.01	-1	3.00	+1
21	3.16	1.86	1.65	0.61	0.33	0.51	0.52	0.52	0.52	2.14	-2	8.39	0
22	2.47	1.19	1.53	0.63	0.38	0.56	0.61	0.64	0.63	2.19	-1	4.44	+1
24	3.73	1.93	1.92	0.80	0.43	0.73	0.79	0.78	0.79	14.6	-2	30.6	+1
25	3.40	2.22	2.56	1.02	0.65	0.94	0.94	0.95	0.94	2.29	-1	7.85	0
Mean	2.72	1.73	1.64	0.69	0.40	0.71	0.76	0.75	0.74	3.86	-1.50	9.95	0.67

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Table 6. Initial conditions of geometric mean diameter [μm], standard deviation and number concentration [cm^{-3}] for each mode of the PSD associated with clear atmospheric conditions and the average of observed PSD from the FMPS measurements.

Property	Clear Case			Measurements		
	Aitken	Accumulation	Coarse	Aitken	Accumulation	Coarse
Geometric mean diameter [μm]	0.021	0.103	0.929	0.058	0.026	1.107
Standard Deviation	1.80	1.60	2.20	1.67	1.48	2.20
Number Concentration [cm^{-3}]	1.256×10^3	6.451×10^2	7.252×10^{-1}	7.664×10^3	1.965×10^3	2.920×10^2

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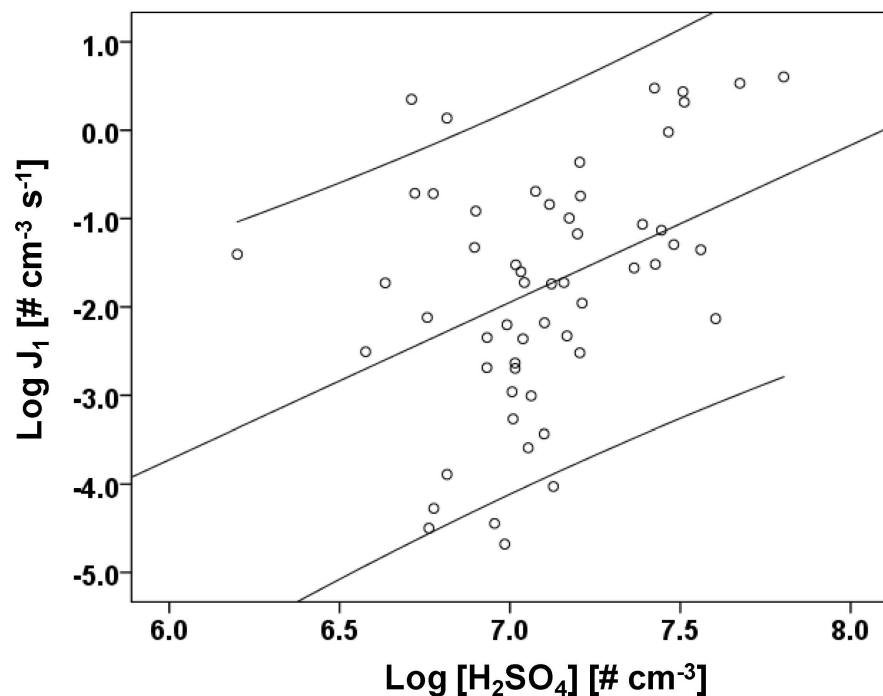


Fig. 1. Relationship between the base 10 logarithm of sulfuric acid concentrations [cm^{-3}] and the base 10 logarithm of the nucleation rate of 1 nm particles [$\text{cm}^{-3} \text{s}^{-1}$] during nucleation hours of class A events. The estimated regression coefficients are statistically significant, giving a regression line of $\text{Log}_{10} J_1 = -14.39 \times \text{Log}_{10} J_1 [\text{H}_2\text{SO}_4]^{1.78}$. The 90 % confidence intervals for K ($[-21.28, -7.51]$) and n ($[0.81, 2.75]$) are also shown.

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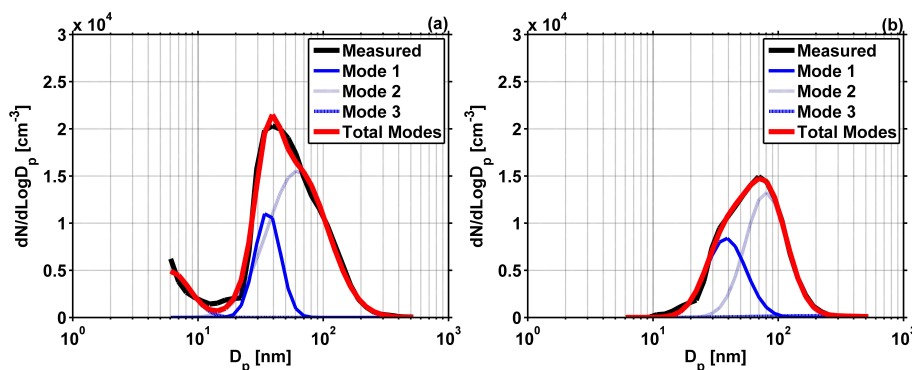


Fig. 2. Fit of the initial PSD to a three lognormal function using data measured by the FMPS at 00:00 LST for **(a)** an example A-event day (19 May) and **(b)** a non-event day (26 May).

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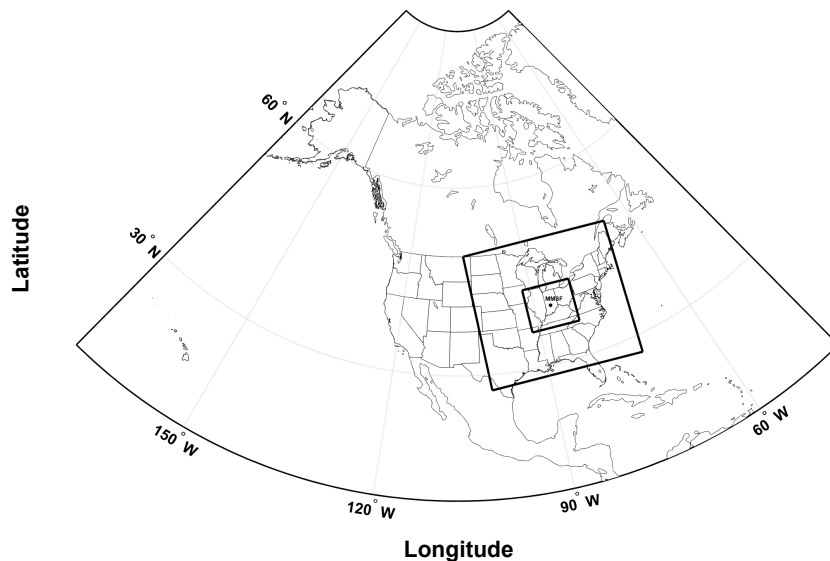


Fig. 3. Parent and nested domain simulated by WRF imposing a spatial resolution of 9 km and 3 km respectively. The location of the MMSF site is also shown.

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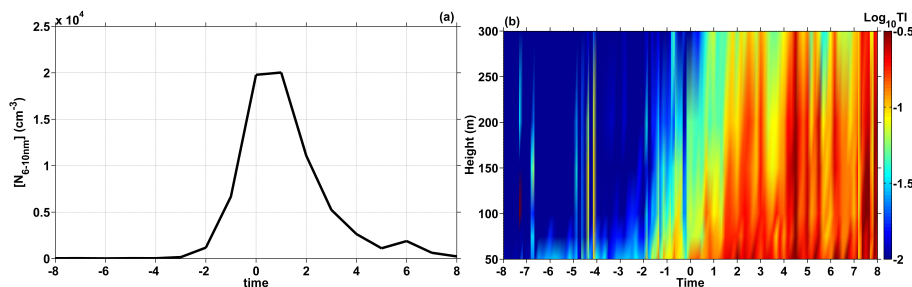


Fig. 4. (a) Mean 6–10 nm particle number concentrations (cm^{-3}) and (b) \log_{10} of turbulence intensity estimated from the 10 min average ZephIR lidar measurements during event class A days. The time coordinate is expressed relative to the hour the maximum rate change of 10 nm particle concentrations was observed.

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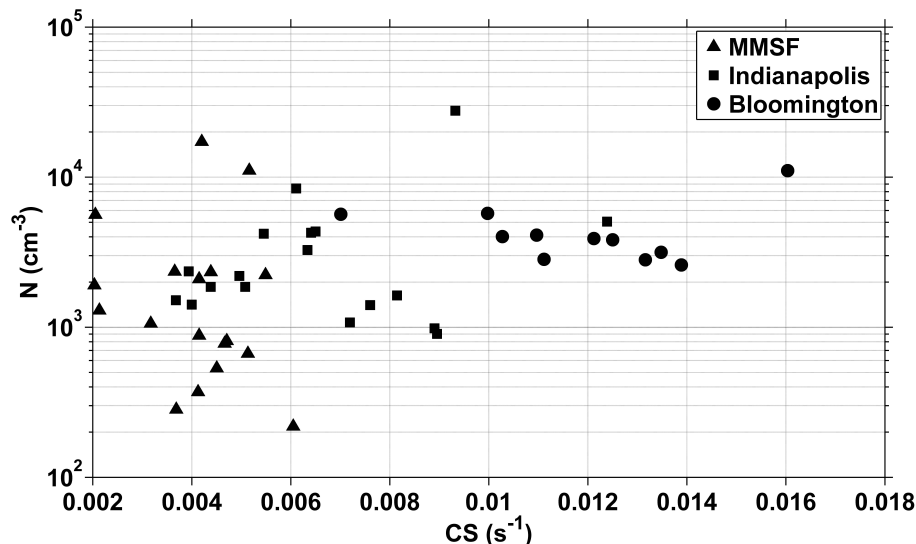


Fig. 5. Scatterplot of condensational sink (CS) (s^{-1}) and ultrafine particle number concentrations (cm^{-3}). CS is computed for event class A days for the period 1 h prior to the maximum rate change of 10 nm particles. Particle number concentrations are for 6–10 nm diameters in data from MMSF and for $D_p = 10\text{--}30$ nm particle number concentration for the Bloomington and Indianapolis sites.

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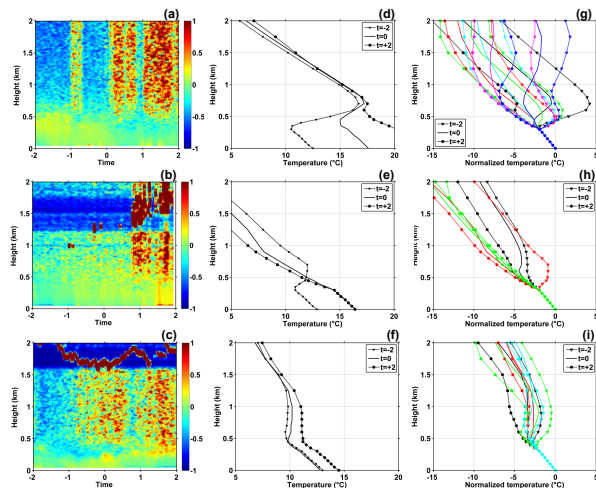


Fig. 6. Examples of relative backscatter signal computed from ceilometers data and vertical temperature profiles (°C) simulated by WRF during a **(a–b)** class A event (17 May), **(c–d)** class C event (18 May) and **(e–f)** non-event (23 May) for times t normalized relative to the hour of maximum rate change of 10 nm particle number concentrations. The relative backscatter b_{rel} is defined according to (Stratmann et al., 2003): $b_{rel}(z) = \frac{b(z) - \overline{b(z)}}{\overline{b(z)}}$, where $b(z)$ is the backscatter

measured at the height z and $\overline{b(z)}$ is the average backscatter at a specific height z during the analyzed time period (06:00 a.m.–12:00 p.m. LST). A moving average filter of 5 min is applied and the vertical resolution is 20 m. Panels in column 3 summarize normalized vertical temperature profiles (°C) simulated by WRF during **(g)** event class A, **(h)** event class B and C, **(i)** non-event days between 14 and 26 May for times t normalized relative to the hour of maximum rate change of 10 nm particle number concentrations. Temperatures in frames g, h and i have been normalized relative to the ground value and different colors refer to different days.

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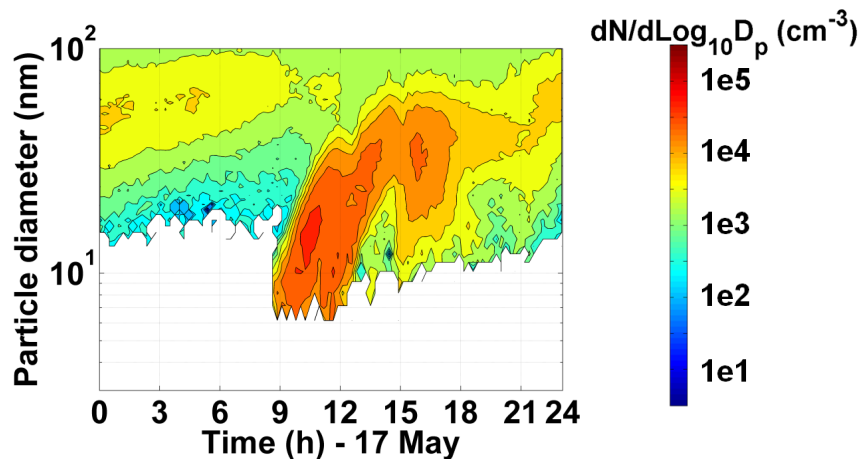


Fig. 7. Contour plot showing the particle number size distribution [cm^{-3}], expressed as $dN/d\text{Log}D_p$, measured on 17 May. The presence of a closed contour line around the highest particle number concentrations may imply non-local nucleation (Birmili and Wiedensohler, 2000).

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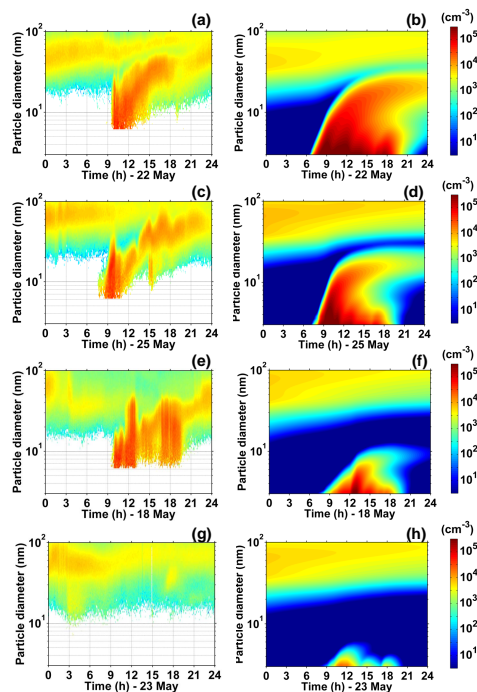


Fig. 8. Example particle number size distribution [cm^{-3}], expressed as $\text{dN}/\text{dLog}D_p$ from measurements (panels **a**, **c**, **e**, **g**) and as simulated by the UHMA (panels **b**, **d**, **f**, **h**) during the NIFTy experiment. The reported days are representative of (**a**)–(**b**) an event class A and (**g**)–(**h**) a non-event day showing good model skill, (**c**)–(**d**) an event class A and (**e**)–(**f**) an event class C day in which particle growth is not captured by the model. Model simulations have been performed setting 250 sections in order to match the size resolution of the SMPS.

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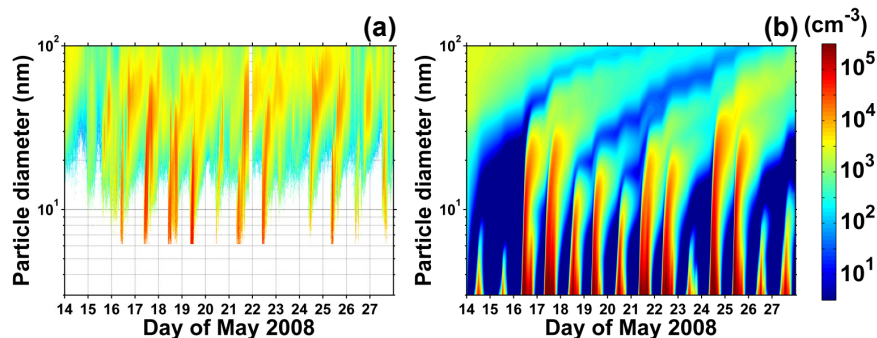


Fig. 9. Particle number size distribution [cm^{-3}], expressed as $\text{dN}/\text{dLog}D_p$, **(a)** measured by the SMPS during the NIFTy experiment and **(b)** simulated. Model simulations have been performed setting 250 sections in order to match the size resolution of the SMPS and was initialized with the PSD observed on the 14th of May at 00:00 LST.

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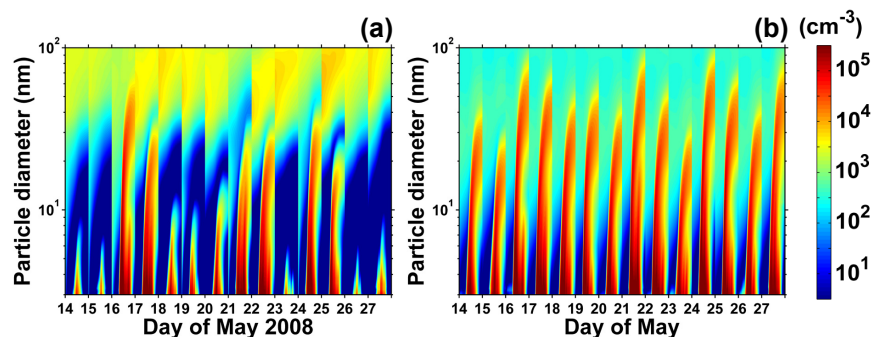


Fig. 10. Simulated particle number size distribution [cm^{-3}], expressed as $\text{dN}/\text{dLog}D_p$. Model simulations were performed using 250 sections to match the size resolution of the SMPS. The model was initialized every day at 00:00 LST using (a) the measured PSD and (b) the PSD representative of clear atmospheric conditions as described in Seigneur et al. (1986).

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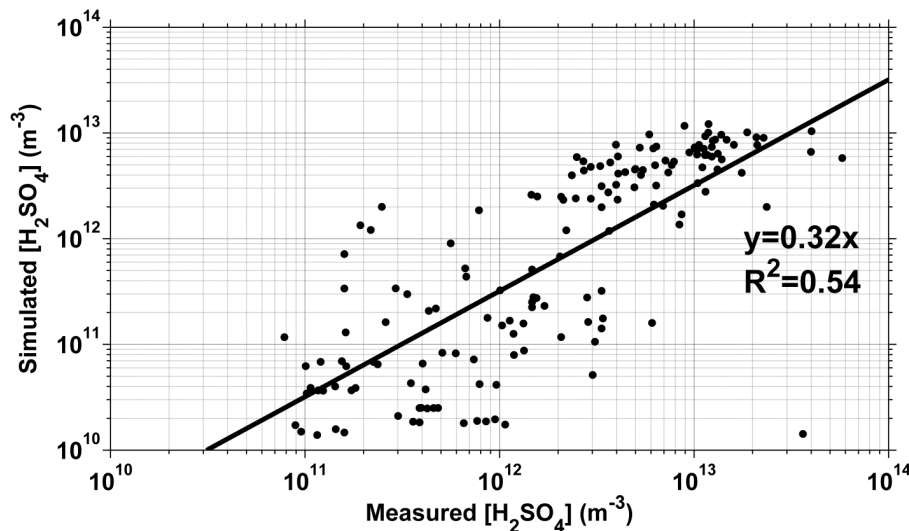


Fig. 11. Regression analysis between hourly averages of measured and simulated sulfuric acid concentrations $[\text{H}_2\text{SO}_4]$ (m^{-3}) for event days during the NIFTy experiment. As shown, simulated values are negatively biased relative to the in situ measurements.

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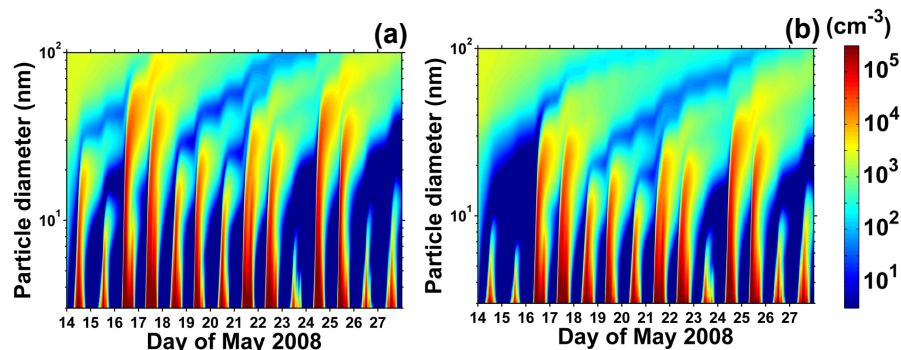


Fig. 12. Simulated particle number size distribution [cm^{-3}], expressed as $dN/d\text{Log}D_p$ showing model sensitivity on **(a)** sulfuric acid concentrations which have been enhanced by a factor 3.125 in order to match the observed values and **(b)** on the mass accommodation coefficient, set to 0.43 for each condensable compound (Pöschl et al., 1998). Model simulations have been performed setting 250 sections in order to match the size resolution of the SMPS and initialized with the PSD observed on 14 May at 00:00 LST.

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