

⁹Department of Biophysics, Graduate School of Science, Kyoto University, Kyoto, Japan

¹⁰Finnish Meteorological Institute, Helsinki, Finland

Received: 26 March 2012 – Accepted: 23 April 2012 – Published: 4 May 2012

Correspondence to: P. Paasonen (pauli.paasonen@helsinki.fi)

Published by Copernicus Publications on behalf of the European Geosciences Union.

ACPD

12, 11485–11537, 2012

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Abstract

Sulphuric acid is a key component in atmospheric new particle formation. However, sulphuric acid alone does not form stable enough clusters to initiate particle formation in atmospheric conditions. Strong bases, such as amines, have been suggested to stabilize sulphuric acid clusters and thus participate in particle formation. We modelled the formation rate of clusters with two sulphuric acid and two amine molecules ($J_{A_2B_2}$) at varying atmospherically relevant conditions with respect to concentrations of sulphuric acid ($[H_2SO_4]$), dimethylamine ([DMA]) and trimethylamine ([TMA]), temperature and relative humidity (RH). The modelled formation rates $J_{A_2B_2}$ were functions of sulphuric acid concentration with close to quadratic dependence, which is in good agreement with atmospheric observations of the connection between the particle formation rate and sulphuric acid concentration. The coefficients $K_{A_2B_2}$ connecting the cluster formation rate and sulphuric acid concentrations as $J_{A_2B_2} = K_{A_2B_2}[H_2SO_4]^2$ turned out to depend also on amine concentrations, temperature and relative humidity. We tested how the model results change if the clusters with two sulphuric acid and two amine molecules are assumed to act as seeds for heterogeneous nucleation of organic vapours (other than amines) with higher atmospheric concentrations than sulphuric acid. We also compared the modelled coefficients $K_{A_2B_2}$ with the corresponding coefficients calculated from the atmospheric observations (K_{obs}) from environments with varying temperatures and levels of anthropogenic influence. By taking into account the modelled behaviour of $J_{A_2B_2}$ as a function of $[H_2SO_4]$, temperature and RH, the atmospheric particle formation rate was reproduced more closely than with the traditional semi-empirical formulae based on sulphuric acid concentration only. The formation rates of clusters with two sulphuric acid and two amine molecules with different amine compositions (DMA or TMA or one of both) had different responses to varying meteorological conditions and concentrations of vapours participating to particle formation. The observed inverse proportionality of the coefficient K_{obs} with RH and temperature agreed best with the modelled coefficient $K_{A_2B_2}$ related to formation of

ACPD

12, 11485–11537, 2012

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

a cluster with two H_2SO_4 and one or two TMA molecules, assuming that these clusters can grow in collisions with abundant organic vapour molecules. In case this assumption is valid, our results suggest that the formation rate of clusters with at least two of both sulphuric acid and amine molecules might be the rate-limiting step for atmospheric particle formation. More generally, our analysis elucidates the sensitivity of the atmospheric particle formation rate to meteorological variables and concentrations of vapours participating in particle formation (also other than H_2SO_4).

1 Introduction

Formation of new aerosol particles and their growth has been observed to take place in a wide variety of environments (Kulmala et al., 2004; Kulmala and Kerminen, 2008; Mirme et al., 2010; Zhang et al., 2012). In terms of the total particle number concentration in the global atmosphere, particle formation is very likely to dominate over the primary aerosol particle sources (Spracklen et al., 2006; Yu et al., 2010). Atmospheric particle formation has been estimated to give a significant contribution to the global cloud condensation nuclei budget (Spracklen et al., 2008; Merikanto et al., 2009; Pierce and Adams, 2009; Yu and Luo, 2009), and to cause significant uncertainties in estimates of both current and future indirect climate forcing by aerosols (Wang and Penner, 2009; Kazil et al., 2010; Makkonen et al., 2012).

The most important vapour for atmospheric new particle formation is thought to be sulphuric acid, yet significant uncertainties related to the actual formation mechanism and role of other vapours in this process remain (Kerminen et al., 2010; Sipilä et al., 2010; Kirkby et al., 2011; Zhang et al., 2012). Field measurements indicate that the formation rate of new aerosol particles tend to be proportional to the ambient sulphuric acid concentration ($[\text{H}_2\text{SO}_4]$) to the power of 1–2 (Weber et al., 1996; Sihto et al., 2006; Riipinen et al., 2007; Kuang et al., 2008; Paasonen et al., 2010), but larger power values have also been presented (e.g., Wang et al., 2011). Theoretically, the power of one could be explained by an activation-type process (Kulmala et al., 2006) and the

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures



Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



power of two by a kinetic process (McMurry and Friedlander, 1979). The corresponding semi-empirical expressions for the particle formation rate, J , may be written as:

$$J = A [\text{H}_2\text{SO}_4], \quad (1)$$

$$J = K [\text{H}_2\text{SO}_4]^2, \quad (2)$$

where A and K are empirically determined coefficients (e.g., Riipinen et al., 2007). It should be noted that the power of two can be explained also by other than kinetically limited processes (Vehkamäki et al., 2012).

At present, there is no general consensus on whether Eq. (1) or (2) describes better the atmospheric particle formation, or should some more detailed parameterisation be used. This unfortunate situation stems from (i) the scarcity of atmospheric measurement data suitable for this purpose, (ii) challenges in analysing atmospheric particle formation events (Korhonen et al., 2011), (iii) the possibility of having more than one formation mechanism operating either simultaneously or under different atmospheric conditions, and (iv) the large variability of the empirically determined values of the coefficients A and K between different particle formation events and from site to site. In most of the previous studies the particle formation rate is compared only with the sulphuric acid concentration. The detailed role of other constituents and the impact of meteorological conditions are not known so far, even though the roles of organic vapours (e.g., Metzger et al., 2010) and relative humidity (Hamed et al., 2011) have been studied. In this manuscript we will investigate the variability of the coefficient K and the factors possibly responsible for this variability. We will focus on the kinetic-type process (Eq. 2), since this mechanism is tractable with the modelling framework that we will use as our main analysis tool. Furthermore, Paasonen et al. (2010) showed that even though the coupling between particle formation rate and sulphuric acid concentration was roughly linear (as in Eq. 1) when the data from several sites were investigated together, the analysis of single-site data sets revealed closer to quadratic couplings at each site (as in Eq. 2).

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

In order to approach the problem, we need to specify the vapours participating in particle formation. Petäjä et al. (2011) showed that under atmospherically relevant conditions, the formation rate of sulphuric acid dimers is close to the collision rate of sulphuric acid molecules. In order to explain this finding, a stabilizing agent (most probably a basic molecule) is required in a sulphuric acid-containing cluster. Several laboratory experiments have pointed out that amines, which are probably the strongest bases present in atmospheric with significant concentrations, have a marked enhancing impact on particle formation (Murphy et al., 2007; Berndt et al., 2010; Erupe et al., 2011; Bzdek et al., 2011; Zhao et al., 2011; Yu et al., 2012). These observations agree with quantum chemical calculations, e.g., Kurtén et al. (2008), Loukonen et al. (2010) and Ortega et al. (2012) concluded that amines are far more efficient than ammonia in stabilizing small H_2SO_4 -clusters in the atmosphere. However, Nadykto et al. (2011) estimated that the difference between amines and ammonia in their capacity for stabilizing H_2SO_4 -clusters is evened out by typically orders of magnitude higher atmospheric ammonia concentrations. The formation of stable H_2SO_4 dimers is the first step in acid-base cluster formation, and thereby gives the maximum particle formation rate in the kinetic-type process. Hypothetically, it might be that a cluster of two H_2SO_4 and amine molecules could be both large and (relatively) basic enough to grow further via collisions with acids other than H_2SO_4 as well – for example organic acids, which have much higher atmospheric concentrations than sulphuric acid. In this case, the formation rate of the clusters would be directly comparable to the observed particle formation rate.

Even though some atmospheric concentration measurements have been made during the last decades (Ge et al., 2011; Hanson et al., 2011), the knowledge of the concentrations of gas phase amine molecules in ambient air is still far from being complete. Even less is known about the spatial and temporal variation of their concentrations. In addition, the insufficient understanding of the role of other organic vapours in particle formation or in their initial growth increases the uncertainties in determining the mechanism behind atmospheric particle formation

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The particle formation rate is expected to be affected by the ambient temperature and humidity. The evaporation rate of molecules from a cluster increases exponentially with the temperature, so the stability of clusters decreases with an increasing temperature and a smaller fraction of the clusters will be able to grow into larger sizes. On the other hand, the presence of water molecules increases the stability of some sulphuric acid-amine-clusters (Loukonen et al., 2010), and thus the particle formation rate can also depend on relative humidity (RH).

The coefficient K has been observed to vary by roughly four orders of magnitude, from about 10^{-15} to 10^{-11} $\text{cm}^3 \text{s}^{-1}$ under different atmospheric conditions (Kuang et al., 2008; Paasonen et al., 2010). Our main goal in this paper is to shed light on the factors explaining this variability. More specifically, we aim to address the following questions: (i) how the formation rate of $(\text{H}_2\text{SO}_4)_2$ -amine clusters (i.e. stabilized sulphuric acid dimers) is limited by the concentrations of H_2SO_4 and amines, (ii) how sensitive this process is to the properties of different amines, (iii) what the roles of ambient temperature and humidity are, and (iv) how our results compare with atmospheric observations. We model the concentrations of clusters formed from sulphuric acid and two amines, di- and trimethylamine, in atmospherically relevant concentration and temperature ranges. The model results are presented in the way that they are comparable to atmospheric measurements. In addition to the model results, we compare the modelled formation rate of stable dimers to a variety of atmospheric data of particle formation rate and sulphuric acid concentration measured in very different atmospheric environments.

2 Materials and methods

In the following sections the abbreviations A, D, and T refer to sulphuric acid (H_2SO_4), dimethylamine ($(\text{CH}_3)_2\text{NH}$, DMA), and trimethylamine ($(\text{CH}_3)_3\text{N}$, TMA), respectively. The abbreviation B refers to base, here meaning either DMA or TMA. The number of molecules in a cluster is marked with a subscript and the concentration is marked with square brackets, e.g. $[\text{A}_2\text{D}_1\text{T}_1]$ stands for the concentration of clusters with two

5 sulphuric acid molecules, one DMA and one TMA molecule. Additionally we use $[A_{1,tot}]$ for the total sulphuric acid monomer concentration, i.e. the sum of concentrations $[A]$, $[A_1D_1]$ and $[A_1T_1]$. In the comparison between the model results and atmospheric observations we denote with K_{obs} the values of the coefficient K calculated from the measured particle formation rate and sulphuric acid concentration using Eq. (2).

2.1 Model system

10 The modelled cluster size space includes clusters consisting of 1–3 molecules of sulphuric acid and 0–2 amine molecules. However, the clusters A_3T_2 and $A_3D_1T_1$ were not included in the model, because the corresponding quantum chemical formation energies have not been calculated so far. Ortega et al. (2012) determined the cluster energies for clusters up to A_4D_4 , but we did not include these clusters into our system in order to keep the model system similar for both DMA and TMA. Including the larger clusters would also require an assumption that e.g. the organic vapours do not adhere to these clusters (due to their too high acidity/low basicity), which would complicate the system further. The effect of water molecules bound to A_2 , A_1D_1 and A_1T_1 -clusters on the concentration distribution was studied, but the hydrated clusters were not directly included in the dynamic model (see Sect. 2.1.3. for how hydration was implicitly accounted for). The model runs were performed in temperature range from 260 to 320 K and RH range from 0 to 100 %.

20 The amine concentrations were kept constant during the model runs. We made the calculations with $[DMA]$ to $[TMA]$ ratios of 1 : 1, 1 : 10 and 10 : 1. In order to make the comparison to measured sulphuric acid concentration possible (see Sect 2.1.1), we also kept the total sulphuric acid monomer concentration $[A_{1,tot}]$ constant by fixing the changes in $[A]$ to be inverse to changes in $[A_1B_1]$, i.e. $d[A]/dt = -d[A_1D_1]/dt - d[A_1T_1]/dt$.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



2.1.1 Kinetic model and rate coefficients

We used a kinetic model (DACM, Kulmala, 2010) for calculating the concentrations of sulphuric acid-amine-clusters. The steady state concentrations were determined by dynamical equations for each cluster. The time derivatives of the concentration of each cluster type included source terms from the collisions and evaporations resulting in that cluster. The sink terms originated from the evaporation rates from the cluster, its collisions with the other clusters within the model and losses due to the coagulation sink, which was set equal to 10^{-3} s^{-1} , a typical sink for molecular sized clusters in continental background areas (Dal Maso et al., 2007). We repeated all the model runs under the assumption that the A_2B_2 -clusters also grow out of the system by collisions with organic vapour molecules having concentration of 10^9 cm^{-3} . These collisions were included in the model by adding an additional loss term for the A_2B_2 -clusters.

The collision rates k_{ij} between clusters i and j were calculated as (Seinfeld and Pandis, 2006)

$$k_{ij} = \beta_{ij} c_i c_j, \quad (3)$$

where c_i stands for the concentration of cluster i , and β_{ij} is the collision coefficient between clusters i and j , obtained from kinetic gas theory as

$$\beta_{ij} = \pi(r_i + r_j)^2 \sqrt{\frac{8k_B T}{\pi \mu}}, \quad (4)$$

where k_B is the Boltzmann constant, T is the temperature, r_i is the radius of cluster i and $\mu = m_i m_j / (m_i + m_j)$ is the reduced mass.

The evaporation $e_{i+j \rightarrow i,j}$ rates were calculated as

$$e_{i+j \rightarrow i,j} = \gamma_{i+j \rightarrow i,j} c_{i+j}, \quad (5)$$

where $\gamma_{i+j \rightarrow i,j}$ is the evaporation coefficient of cluster $i + j$ evaporating into clusters i and j , and c_{i+j} is the concentration of cluster $i + j$. The evaporation coefficients were

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



acquired from the quantum chemical calculations described by Ortega et al. (2012). Assuming detailed balance, the evaporation coefficient can be calculated from the corresponding collision coefficient and the Gibbs free energies of formation of the evaporating cluster $i + j$ and the product clusters i and j :

$$\gamma_{i+j \rightarrow i,j} = \beta_{ij} \frac{p_{\text{ref}}}{k_{\text{B}}T} \exp\left(\frac{\Delta G_{i+j} - \Delta G_i - \Delta G_j}{k_{\text{B}}T}\right), \quad (6)$$

where β_{ij} is the collision coefficient, ΔG_i is the Gibbs free energy of formation from monomers of cluster i and p_{ref} is the reference pressure where the formation free energies have been calculated (which cancels because the values of ΔG_i are calculated in the same pressure).

The formation free energies of all the clusters were originally calculated at 298.15 K using a quantum chemical multi-step method that combines B3LYP/CBSB7 optimized geometry and frequency calculations with RI-CC2/aug-cc-pV(T + d)Z single point energy calculations (Ortega et al., 2012). The Gibbs free energy of formation in other temperatures was calculated as

$$\Delta G = \Delta H + T\Delta S, \quad (7)$$

where ΔH and ΔS are the formation enthalpy and entropy of formation of the cluster, respectively, obtained using harmonic oscillator and rigid rotor approximations for the clusters. The values of ΔH and ΔS were assumed to be constant over the studied temperature range.

Some assumptions of the evaporation rates were necessary to be made because quantum chemical free energies for all the relevant larger clusters have not been available. The rate coefficients for the following collisions were set to zero, because the corresponding evaporation from the resulting cluster occurs presumably very rapidly (Ortega et al., 2012): all collisions of the type $A_1B_2 + B$; $A_3 + A_n$; $A_3B_1 + A_n$; $A_2T_2 + T$ (as the computed evaporation rate of DMA from A_2D_3 is very fast, see Sect. 3.1). However, the collisions leading to clusters $A_2D_2T_1$ and $A_2D_1T_2$ were allowed. We have

11494

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



not computed the free energies for these clusters, but as it is obvious that they are very unstable, they were set to evaporate one amine molecule instantaneously when formed. Whether $A_2D_2T_1$ and $A_2D_1T_2$ -clusters evaporate a DMA or TMA molecule is determined by the relative stabilities of the resulting A_2D_2 , $A_2D_1T_1$ and A_2T_2 clusters, as can be seen by using Eq. (6) to calculate the ratio of two evaporation rates from the same initial cluster (i.e. with same ΔG_{i+j} , the differences between the collision rates β_{2A_2B+B} for different amines B are negligible, and the monomer energies ΔG_{DMA} and ΔG_{TMA} are equal to zero, by definition). With these boundary conditions, all the routes along which the clusters grow out of the system are included in the formation rate of clusters with two or more molecules of both A and B, $J_{A_2B_2}$ (see Sect. 2.1.3).

2.1.2 Effect of cluster formation on sulphuric acid concentration

The atmospheric sulphuric acid concentration is typically measured with Chemical Ionization Mass Spectrometers (CIMS; Berresheim et al., 2000; Petäjä et al., 2009; Sjöstedt et al., 2007). The measurement with the CIMS is based on selective ionization, during which the reagent nitrate ions (and/or its clusters with nitric acid) accept a proton from sulphuric acid molecule, converting it to a bisulphate ion. The signal of these bisulphate ions in the mass spectrometer is then converted to concentration via the calibration method described by Mauldin III et al. (1999). However, Kurten et al. (2011) estimated that the majority of the H_2SO_4 ·ammonia- and H_2SO_4 ·DMA-clusters gets charged in the CIMS charger. This charging leads to an immediate evaporation of the base molecule, and thus these clusters are detected as pure H_2SO_4 . In order to take this into account, we modelled the H_2SO_4 concentration measurement with the total concentration of sulphuric acid monomers, i.e. the sum of the pure sulphuric acid monomer concentration and the concentration of A_1B_1 -clusters

$$[A_{1,tot}] = [A] + [A_1T_1] + [A_1D_1]. \quad (8)$$

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

We ran the model with $[A_{1,\text{tot}}]$ in range of 10^5 to $3.16 \times 10^8 \text{ cm}^{-3}$, covering the typical observed ambient sulphuric acid concentrations (Petäjä et al., 2009; Mikkonen et al., 2011; Zheng et al., 2011).

2.1.3 Effect of water on the collision and evaporation rates

The concentration of water in the atmosphere is approximately ten orders of magnitude higher than those of sulphuric acid, ammonia and amines, and solving the birth-death equations for the full (three-dimensional) water-containing system is, therefore, computationally very challenging. To overcome this problem, we first solved the hydrate distributions for clusters with no more than two molecules (other than water) at different temperatures and relative humidities, and then calculated effective collision and evaporation rates of the other compounds as a weighted average over these hydrate distributions. We performed test calculations to justify this approach, and found that indeed the hydrate distributions stay in equilibrium during all the collision and evaporation processes relevant to this study.

2.1.4 Modelled cluster formation rate

The minimum requirement for a cluster to participate in particle formation is that it holds together long enough to collide with another cluster or molecule of nucleating vapour. We modelled the formation rate of a cluster as the difference between the sum of collision rates forming the cluster and the sum of evaporation rates from the cluster, i.e.

$$J_{ab} = \sum \beta_{a_i b_i} [a_i b_i] [a_j b_j] - \sum \gamma_{ab} [ab], \quad (9)$$

where the first term describes the formation of clusters with at least a acid molecules and b base molecules from the collisions of all the clusters $a_i b_i$ and $a_j b_j$ in which $a_i + a_j \geq a$ and $b_i + b_j \geq b$, and the second term describes the evaporation rate as a sum of the rates at which an acid or base molecule or an acid-base-cluster is evaporated

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



from cluster ab . This approach is analogous to the measured particle formation rates (e.g., Kulmala et al., 2007; Manninen et al., 2009), in which the external loss terms for a particle, coagulation and growth, are also taken into account.

In this study, we examine in detail $J_{A_2B_2}$, the formation rate of clusters including two or more sulphuric acid molecules and two or more amine molecules (for clarification, here are listed some of the collisions that are counted in $J_{A_2B_2}$: $A_2B_1 + B$; $A_1B_1 + A_1B_1$; $A_1B_2 + A_2$; $A_2B_1 + A_2B_1$, but e.g. $A_2B_2 + A$ is not, because the formation of A_2B_2 is already counted in $J_{A_2B_2}$). The conversions of A_2T_2 , $A_2D_1T_1$ and A_2D_2 due to their collisions with amine molecules, described in Sect. 2.1.1, were taken into account when calculating formation rates $J_{A_2B_2}$. Thus, we subtracted e.g. from the formation rate $J_{A_2T_2}$ calculated with Eq. (9) the proportion of collisions $A_2T_2 + D$ which results in $A_2D_1T_1$, and consistently added this proportion to $J_{A_2D_1T_1}$. The evaporations reckoned in the calculation of $J_{A_2B_2}$ are those from A_2B_2 -clusters only, because the evaporation rates of $A_{\geq 1}B_{\geq 1}$ from most of the $A_{> 2}B_{\geq 2}$ -clusters were not available (following consequences are discussed in Sect. 5.1). The reason for not analysing separately the formation rate of A_2B_1 clusters ($J_{A_2B_1}$) will be explained later.

The formation rate $J_{A_2B_2}$ is, as shown in the results, coupled with $[A_{1,tot}]$ to the power close to two. In order to study the effects of varying temperature, RH and concentrations of H_2SO_4 and amines on $J_{A_2B_2}$, we inspect the behaviour of the “observable” coefficient $K_{A_2B_2}$ in

$$J_{A_2B_2} = K_{A_2B_2} [A_{1,tot}]^2. \quad (10)$$

2.2 Atmospheric measurements

We reanalysed nine sets of measurements from around the Northern Hemisphere, all including data for aerosol particle number size distribution, sulphuric acid concentration, temperature and RH. The sulphuric acid concentrations at all the sites were measured with CIMS instruments. The measurements were conducted at the SMEAR II station in Hyytiälä (Finland) during QUEST II- and QUEST IV-campaigns in March–April 2003 and April–May 2005, respectively (Sihto et al., 2006; Riipinen et al., 2007),

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

and during the EUCAARI-campaign in April–May 2007 (Kulmala et al., 2011). The measurements at the German stations, Melpitz in May 2008 and Hohenpeissenberg from June 2007 to December 2008, were also conducted as a part of the EUCAARI campaign, as well as the measurements at San Pietro Capofiume (Italy) during June–July 2009. The data at Tecamac (municipality 40 km north from Mexico City, Mexico) were gathered in May 2006 (Iida et al., 2008), at Atlanta (US) (McMurry and Eisele 2005; Kuang et al., 2008) in August 2002 and at Beijing (China) between July and September 2008 (Wang et al., 2011).

The particle number size distribution measurements during the EUCAARI campaigns (Hyytiälä 2007, Melpitz, Hohenpeissenberg and S. P. Capofiume) were measured with a Neutral and Air Ion Spectrometer (NAIS, Kulmala et al., 2007). From these data the formation rate of particles with diameter of 2 nm was calculated directly as described in Paasonen et al. (2010). At the other sites the particle number size distribution was measured with a Differential Mobility Particle Sizer (DMPS). From these data we calculated first the formation rate of 4 nm particles using a similar procedure as with the NAIS data, and then back calculated the formation rate of 2 nm particles with the formula given by Kerminen and Kulmala (2002). From the QUEST campaigns in Hyytiälä we applied the data for the formation rate of 1.5 nm particles calculated with the methods described by Sihto et al. (2006) and Riipinen et al. (2007).

3 Results: simulated behaviour of the system

3.1 Cluster formation free energies and evaporation rates

The quantum chemical enthalpies ΔH , entropies ΔS and the related Gibbs free energies ΔG at $T = 280$ K of the clusters included in this study are presented in Table 1. The resulting evaporation rates are presented in Table 2, and those of A_1B_1 , A_2B_2 and A_2B_1 -clusters are depicted in Fig. 1. Electronic cluster formation energies calculated in this study (at the B3LYP/CBSB7//RI-CC2/aug-cc-pV(T+d)Z level) as well as the

coordinates of the most stable structure of each cluster (at the B3LYP/CBSB7 level) are presented in Supplement. The same information for the other clusters is presented by Ortega et al. (2012).

Evaporation from A_3B_1 -clusters was found to be very fast in comparison to A_2B_1 -clusters (see Table 2), and adding more acids while keeping the number of bases constant would increase the evaporation rate even more (Ortega et al., 2012). This implies that in order to form a stable cluster consisting of at least two H_2SO_4 molecules that is capable of growing further, the addition of a second base molecule is required. Thus, we focus on the formation of clusters consisting of at least two molecules of both sulphuric acid and amine molecules.

Trimethylamine is a stronger base than dimethylamine, making the A_1T_1 -cluster slightly more stable than the A_1D_1 -cluster. However, the nitrogen atom in the TMA molecule, already bound to three carbon atoms, is capable of forming only one hydrogen bond, whereas that in dimethylamine can form two. Thus, the A_2T_2 -cluster is much less stable than A_2D_2 cluster. Accordingly, the only crucial point in the formation of A_2D_2 is whether the A_1D_1 -cluster will hold together until it collides with another cluster or a sulphuric acid molecule. In the formation of A_2T_2 -clusters, also the competition between the evaporation of TMA from A_2T_2 cluster and the collision of A_2T_2 with the next condensing molecule is important. It should be noted that the evaporation of a TMA molecule from a A_2T_2 -cluster results in a very stable A_2T_1 -cluster which can then grow back by colliding either with an amine molecule or H_2SO_4 clustered with an amine molecule.

The relative stabilities of A_1T_1 - and A_1D_1 -clusters change when water molecules are introduced to the system. In typical atmospheric conditions ($T = 280$ K, $RH = 40\%$) the water molecule concentration is around 10^{17} cm^{-3} , leading to a collision frequency with water of about $5 \times 10^7 \text{ s}^{-1}$ for all molecules and clusters. As the evaporation rate of water from the $A_1D_1W_1$ -cluster is of the same order of magnitude (see right hand side in Table 2), a large fraction of $A_1D_1W_n$ -clusters is hydrated, decreasing the effective evaporation rate of DMA from $A_1D_1W_n$ -clusters. The evaporation rate of a water molecule

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

from $A_1T_1W_1$ clusters is two orders of magnitude higher, implying that roughly 1 % of all $A_1T_1W_n$ clusters are hydrated. However, even this minor fraction increases the effective evaporation rate of TMA from $A_1T_1W_n$ -clusters, because the evaporation of a TMA molecule is 100 times faster from the $A_1T_1W_1$ -cluster compared to that from A_1T_1 .
5 As a result, the effective evaporation rate of a TMA molecule from $A_1T_1W_n$ -clusters is higher than that of a DMA molecule from $A_1D_1W_n$ -clusters at RH of 20 % or higher.

Due to the large computational costs, we did not calculate the quantum chemical free energies of clusters consisting of water and more than two other molecules. For these clusters we assumed that they would evaporate water molecules so effectively that the
10 proportion of hydrated clusters is negligible, and does not affect the effective evaporation rate of a sulphuric acid or amine molecule. This has been shown to be a relevant assumption for the sulphuric acid A_2D_1 -clusters (Loukonen et al., 2010). Because the TMA molecule is capable of forming only one hydrogen bond, a $A_{>1}T_{\geq 1}$ -cluster (where water can make bonds only with the acid molecules) should bind with water even less
15 strongly than a $A_{>1}D_{\geq 1}$ -cluster. This is consistent with the far higher evaporation rates of water from the $A_1T_1W_1$ -cluster than from the $A_1D_1W_1$ -cluster (see Table 2). Thus, the presence of water molecules should not significantly affect the evaporation rates calculated for clusters with more than two sulphuric acid or amine molecules.

3.2 Sulphuric acid monomer distribution

20 We present first the results acquired with a [DMA] to [TMA] ratio of one, and describe in Sect. 3.3.4 how the results change if either concentration is higher.

As our goal in this study is to model the formation rate of H_2SO_4 -amine-clusters with respect to the sulphuric acid concentration measured with CIMS (approximated to equal the total sulphuric acid monomer concentration $[A_{1,tot}]$), we first inspect the
25 proportions of pure and clustered sulphuric acid monomers. In Fig. 2a we depict the fractions of pure H_2SO_4 molecules (solid lines with circles), A_1T_1 -clusters (solid lines) and A_1D_1 -clusters (dashed lines) of the total monomer concentration $[A_{1,tot}]$ under varying amine concentrations as a function of $[A_{1,tot}]$ at $T = 280$ K and in absence of water.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



When $[A_{1,tot}]$ is below 10^6 cm^{-3} (and the meteorological conditions are as above), the fraction of sulphuric acid monomers being clustered with an amine molecule depends only on the amine concentration, but when $[A_{1,tot}]$ surpasses 10^6 cm^{-3} the clustered fraction starts decreasing. If the amine concentration is 10^8 cm^{-3} or higher and an order of magnitude higher than $[A_{1,tot}]$, almost all of the monomers are clustered with amines, > 30 % with DMA and > 60 % with TMA.

The monomer concentration distribution is strongly dependent on the temperature (Fig. 2b). The fraction of H_2SO_4 clustered with amines starts decreasing at temperatures > 270 K, as the evaporation rate from A_1B_1 -clusters exceeds the coagulation sink. However, when the amine concentration increases and exceeds $[A_{1,tot}]$, the temperature at which this decrease becomes visible is elevated, because the increasing collision frequency between sulphuric acid and amine molecules fastens the reformation of the evaporated clusters. In addition to the prevailing concentrations and temperature, the monomer concentration distribution is affected by the relative humidity (RH, Fig. 2c). At temperatures below 270 K (or at high $[A_{1,tot}]$), the effect of RH is insignificant as evaporation is not the main sink for the clustered monomers. At higher temperatures, the ratio between $[A_1T_1]$ and $[A_1D_1]$ decreases from values larger than one to values below one at $\text{RH} \sim 20\%$. Above this RH, the proportion of A_1T_1 clusters continues decreasing with RH while the proportion of (hydrated) pure acid molecules increases.

At amine concentrations $> 10^7 \text{ cm}^{-3}$, the fraction of A_1D_1 -clusters is lower than that of A_1T_1 -clusters also at temperatures at which the evaporation rates of both are below the coagulation sink, and thus the loss rates are equal. This feature arises from the slight stability difference between the $A_1T_1W_0$ - and $A_1D_1W_0$ -clusters: under high amine concentrations the $A_1D_1(W_n)$ -clusters collide frequently with TMA-molecules, and due to the higher stability of the $A_1T_1W_0$ -cluster in comparison with $A_1D_1W_0$, the formed, very unstable $A_1D_1T_1$ -cluster evaporates more often DMA than TMA. Via this mechanism, a significant fraction of A_1D_1 -clusters are continuously converted to A_1T_1 -clusters.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

3.3 Modelled formation rates $J_{A_2B_2}$ and the corresponding coefficients $K_{A_2B_2}$

In the following we will consider only the formation rate of $A_{\geq 2}B_{\geq 2}$ -clusters, $J_{A_2B_2}$, calculated with Eq. (10). The formation rate $J_{A_2B_1}$ is not presented separately, because the high evaporation rate of A from A_3B_1 (see Table 2) does not allow the A_2B_1 -clusters to grow further before the addition of the second amine. The collisions of A_2B_1 with amine containing clusters are, instead, counted in $J_{A_2B_2}$. The modelled formation rates of clusters with at least two H_2SO_4 and amine molecules are presented in Fig. 3 as a function of the modelled total sulphuric acid monomer concentration $[A_{1,tot}]$ under varying amine concentration with the $[DMA]:[TMA]$ ratio equal to unity.

As an overall result, it is notable that all the modelled formation rates are coupled with the concentration $[A_{1,tot}]$ to the power close to two, as can be seen by comparing them with the $A_{1,tot}$ collision rate, $4.5 \times 10^{-10} \times [A_{1,tot}]^2$, indicated in Fig. 1 with a dashed black line. However, the power of $[A_{1,tot}]$ with which the formation rates $J_{A_2B_2}$ are connected varies slightly from two: e.g. $J_{A_2T_2}$ is steeper than the collision rate at low amine concentrations and gets shallower at high values of $[A_{1,tot}]$. A detailed analysis explaining the deviation of the powers from two and the decreasing formation rates $J_{A_2T_2}$ with increasing amine concentrations are given in the following sections.

3.3.1 Effect of temperature

The ambient temperature has an obvious impact on the fraction of sulphuric acid monomers being clustered with amines (Fig. 2b). The formation rate of A_2B_2 -clusters is strongly dependent on this fraction, because pure sulphuric acid molecules do not form stable dimers. Since the information of the clusterization of the sulphuric acid monomers is lost when the sulphuric acid concentration is measured with the CIMS (see Sect. 2.1.2), the formation rate of A_2B_2 -clusters can be expected to be affected by the varying temperature at a given value of $[A_{1,tot}]$, comparable to the measured concentrations. We studied this temperature effect by inspecting the temperature dependence of the coefficient $K_{A_2B_2}$, calculated from Eq. (10) as $K_{A_2B_2} = J_{A_2B_2}/[A_{1,tot}]^2$.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



To get an impression of the temperature and monomer concentration dependences of the coefficients $K_{A_2T_2}$ and $K_{A_2D_2}$, we first depict them as functions of temperature in the presence of only one amine (Fig. 4). The crucial factors affecting the value of $K_{A_2B_2}$ are the amine concentration and the values of the different sink terms, i.e. the evaporation rate, coagulation sink rate and collision rate of the A_2B_2 -cluster with sulphuric acid monomers and dimers (either clustered with amine or not), the latter determining the further growth of the A_2B_2 -clusters.

In principle $K_{A_2T_2}$ and $K_{A_2D_2}$ behave similarly. If we first inspect the lowest temperatures in Fig. 4, we can see that the collision rate of $A_{1,tot}$ ($\sim 4.5 \times 10^{-10} \text{ cm}^3 \text{ s}^{-1}$) is achieved when the amine concentration is roughly an order of magnitude higher than $[A_{1,tot}]$. The other main feature is the decrease of $K_{A_2B_2}$ as a function of temperature due to the decreasing proportion of clustered sulphuric acid monomers (see Fig. 2b). A third similarity between $K_{A_2T_2}$ and $K_{A_2D_2}$ is that at the lowest temperatures the values of both coefficients decrease with increasing $[A_{1,tot}]$. Consequently, at low temperatures

$$J_{A_2B_2} \sim [A_{1,tot}]^{<2}.$$

Despite the qualitatively similar behaviour of $K_{A_2T_2}$ and $K_{A_2D_2}$ in varying concentration and temperature space, there are some notable differences between Fig. 4a and b. First of all, the exponential decrease of $K_{A_2B_2}$ as a function of temperature is much steeper for TMA than for DMA: for a temperature increase of 30 K the coefficient $K_{A_2T_2}$ decreases roughly four orders of magnitude, but $K_{A_2D_2}$ only about two orders of magnitude. The steep decrease of $K_{A_2T_2}$ as a function of temperature results from the reversed temperature dependencies of the source and the sink of A_2T_2 -clusters: the source decreases with temperature, due to increased evaporation of A_1T_1 -clusters, whereas the main sink, the evaporation of TMA from A_2T_2 -clusters, increases. In case of A_2D_2 -clusters, the decrease in the source rate is similar, but the evaporation from A_2D_2 is smaller than the other sink terms over the whole studied temperature range, and the total sink does not change significantly with temperature.

Another difference between Fig. 4a and b is that $K_{A_2T_2}$ is strongly dependent on $[A_{1,tot}]$ also at high temperatures (lines corresponding to same [TMA] but different

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



$[A_{1,\text{tot}}]$ are separated also at $T > 300$ K), whereas the values of $K_{A_2D_2}$ corresponding to the same $[DMA]$ are not significantly different at high temperatures. This difference originates from different outcome of the competition between the growth and evaporation for different A_2B_2 -clusters. Due to the high stability of A_2D_2 , practically all of the formed A_2D_2 -clusters collide with $[A_{1,\text{tot}}]$ before evaporating even at $[A_{1,\text{tot}}] = 10^6 \text{ cm}^{-3}$. On the contrary, a TMA molecule is evaporated from most of the A_2T_2 clusters before collision growth at temperatures above 290 K at all studied concentrations $[A_{1,\text{tot}}]$. Consequently, an increase in $[A_{1,\text{tot}}]$ increases the probability of A_2T_2 -cluster growth before the evaporation of TMA, leading to over an order of magnitude increase in $K_{A_2T_2}$ due to an order of magnitude increase in $[A_{1,\text{tot}}]$. Thus, at high temperatures $J_{A_2T_2} \sim [A_{1,\text{tot}}]^{>3}$.

The values of $K_{A_2D_2}$ are for the most part independent of $[A_{1,\text{tot}}]$ at high temperatures, but when $[A_{1,\text{tot}}]$ reaches values over two orders of magnitude higher than $[DMA]$ (as in the blue solid lines with circles in Fig. 4b) the coefficient $K_{A_2D_2}$ starts increasing. With such a large difference in monomer concentrations, $[A_2D_1]$ becomes higher than $[DMA]$ and thus the collisions between two A_2D_1 clusters become a significant source term for $J_{A_2D_2}$. The same occurs in the case of TMA cluster formation as well (blue lines in Fig. 4a). This feature and its significance should be considered with caution, due to the opposite, presumably rapid, evaporations of A_2B_1 -clusters from the resulting A_4B_2 -cluster, which are not included in our model (Ortega et al., 2012; more detailed discussion in Sect. 5.1).

3.3.2 Effect of amine competition

The modelled coefficients $K_{A_2B_2}$ related to different A_2B_2 -clusters, when both the amines are present with a concentration ratio of 1 : 1, are depicted as a function of temperature in Fig. 5, corresponding to $[A_{1,\text{tot}}]$ equal to 10^6 (Fig. 5a), 10^7 (Fig. 5b) and 10^8 cm^{-3} (Fig. 5c).

Even though the concentration of A_1T_1 -clusters is larger than that of A_1D_1 -clusters at $RH = 0$ (see Sect. 3.2), the formation rates $J_{A_2B_2}$ are dominated by the formation rate

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

of clusters with two DMAs, $J_{A_2D_2}$. The values of $K_{A_2D_2}$ at an amine concentration ratio 1 : 1 are practically the same as those obtained in the absence of TMA (see Fig. 4b). At amine concentrations below or equal to $[A_{1,\text{tot}}]$, the formation rate $J_{A_2T_2}$ is also very similar to that computed in the absence of DMA, but when the amine concentrations exceed the sulphuric acid monomer concentration by over an order of magnitude the value of $K_{A_2T_2}$ starts to decrease with increasing amine concentrations at temperatures below ~ 300 K. This feature is caused by a similar conversion mechanism to that of converting A_1D_1 -clusters to more stable A_1T_1 -clusters via $A_1D_1T_1$. The cluster A_2D_2 is ~ 7 kcal mol⁻¹ more stable than A_2T_2 , whereas the $A_2D_1T_1$ -cluster is less than one kcal/mol more stable than A_2T_2 (see Table 1). Because the evaporation rate depends on the stability difference between the initial and final cluster(s), the evaporated amine from $A_2D_1T_2$ -clusters is DMA in ~ 22 % of evaporations, but from $A_2D_2T_1$ -clusters DMA is evaporated in only ~ 0.002 % of evaporations. With this mechanism A_2T_2 clusters are converted to $A_2D_1T_1$ and then to A_2D_2 by collisions with DMA. The higher the amine concentrations is, the more efficient the conversion.

The values and the temperature dependence of the coefficient $K_{A_2D_1T_1}$ are very similar to those of $K_{A_2T_2}$, the difference between the values of these two coefficients never exceeding a factor of two.

3.3.3 Effect of relative humidity

The effect of RH on the values of the modelled coefficients $K_{A_2B_2}$ at total sulphuric acid monomer concentration of 10^7 cm⁻³ and 300 K is presented in Fig. 6. The effect of RH arises from the changes in the effective evaporation rate from the smallest clusters. Thus, as stated in Sect. 3.2, the effect of RH is negligible at low temperatures or high $[A_{1,\text{tot}}]$, when evaporation is not the main sink for these clusters.

The differences between the formation rates of A_2D_2 , $A_2D_1T_1$ and A_2T_2 -clusters become more pronounced with an increasing RH, as the effective evaporation rate of A_1T_1 clusters increases and that of A_1D_1 clusters decreases.

3.3.4 Effect of varying DMA to TMA ratios

In order to inspect the effect of varying the [DMA] to [TMA] ratio we repeated the analysis by setting [DMA] an order of magnitude higher or lower than [TMA].

The behaviour of the modelled coefficient $K_{A_2D_2}$ in varying temperature and concentration space is affected only slightly by changes in the [DMA] to [TMA] ratio. Even if the concentration of DMA is an order of magnitude lower than the TMA concentration, the values of the coefficient $K_{A_2D_2}$ remain almost equal to those in the case of one-to-one amine concentration ratio at $[A_{1,tot}] = 10^6 \text{ cm}^{-3}$, and decrease only by a factor up to four when $[A_{1,tot}]$ increases to 10^8 cm^{-3} . If instead [TMA] is an order of magnitude lower than [DMA], the values of $K_{A_2D_2}$ remain the same as in the absence of TMA at all conditions. On the contrary, the formation rate of A_2T_2 clusters is strongly dependent on the amine concentration ratio. Similarly to the one-to-one amine concentration ratio, the coefficient $K_{A_2T_2}$ is practically equal to that in absence of DMA, as long as the DMA concentration is smaller than or equal to $[A_{1,tot}]$. However, when [DMA] exceeds $[A_{1,tot}]$, the increase of the value of $K_{A_2T_2}$ with increasing amine concentrations first stalls and then starts decreasing as [DMA] increases. This decrease at high DMA concentrations is caused by the higher probability of a A_2T_2 cluster to collide with a DMA molecule than with a sulphuric acid monomer, leading to a conversion chain $A_2T_2 \rightarrow A_2D_1T_1 \rightarrow A_2D_2$, as explained in Sect. 3.3.2. Thus, in comparison to a [DMA] to [TMA] ratio of one-to-one (Fig. 5a–c), the stalling of $K_{A_2T_2}$ occurs at an order of magnitude higher [TMA] for a ratio 1 : 10, and with the opposite concentration ratio at an order of magnitude smaller [TMA]. Consequently, at a 10 : 1 ratio the values of $K_{A_2T_2}$ are below $2 \times 10^{-12} \text{ cm}^{-3}$ at all the amine concentrations and temperatures. With this amine concentration ratio the value of the coefficient $K_{A_2D_1T_1}$ is a factor from 3 to 15 larger than the value of $K_{A_2T_2}$ at all conditions. With a [DMA] to [TMA] ratio of 1 : 10 $K_{A_2D_1T_1}$ is a factor from 4 to 12 smaller than $K_{A_2T_2}$. With both ratios the smallest values of $K_{A_2D_1T_1}$ correspond to the highest amine concentrations. The exponential decrease of $K_{A_2D_1T_1}$ with increasing temperature is still similar to that of $K_{A_2T_2}$ and thus clearly steeper than that of $K_{A_2D_2}$.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



3.3.5 Assuming organic vapours can adhere to A_2B_2 -clusters

As shown in Sect. 3.3.1 the formation rates $J_{A_2T_2}$, $J_{A_2D_1T_1}$ and $J_{A_2D_2}$ are coupled with sulphuric acid concentration to power between below two and slightly over three, depending on the temperature and the ratio of $[A_{1,tot}]$ to amine concentration. With the DACM model we can also investigate the outcomes of the assumption that two sulphuric acid molecules would be enough to form a cluster capable of growing by collisions with organic vapour molecules. If the concentration of this organic vapour is clearly higher than $[A_{1,tot}]$, the growth of the A_2B_2 -clusters is no longer strongly dependent on $[A_{1,tot}]$, but rather on the organic vapour concentration. We studied the changes in the behaviour of the coefficients $K_{A_2B_2}$ in case where A_2B_2 -clusters grow further (and do not evaporate back) via collisions with an organic vapour of concentration 10^9 cm^{-3} , which is well in the range of the typical total atmospheric organic acid concentrations (Chebbi and Carlier, 1996).

The organic vapour concentration has a negligible influence on the values of $K_{A_2D_2}$, because even without organic vapours the evaporation of A_2D_2 -clusters does not significantly decrease the formation rate $J_{A_2D_2}$. However, the behaviour of $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ changes if the growth of the related A_2B_2 -clusters can also occur by collisions with abundant organic vapour molecules. In Fig. 7 we have depicted the values of $K_{A_2T_2}$ at 1 : 1 amine concentration ratio resulting from the assumption that the organic vapour can attach to the A_2T_2 clusters and to promote further growth.

The high concentration of an organic vapour attaching to the A_2B_2 -clusters increases the values of the coefficients $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ by two mechanisms. First, in Fig. 7 the value of $K_{A_2T_2}$ does not change when $[A_{1,tot}]$ increases from 10^6 to 10^7 cm^{-3} , whereas in the absence of the organic vapour a similar change in $[A_{1,tot}]$ increased the value of $K_{A_2T_2}$ by over an order of magnitude (blue solid lines in Fig. 5a–c). As explained in Sect. 3.3.1, the reason for the increase in the values of $K_{A_2T_2}$ with $[A_{1,tot}]$ was the increasing probability of A_2B_2 -clusters to grow before evaporation. In the case where abundant organic vapour is mainly responsible for the growth of the A_2B_2 -clusters, this

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



probability and thus also the values of the related coefficients $K_{A_2T_2}$ become independent of $[A_{1,tot}]$ (elevated values of $K_{A_2T_2}$ at high temperatures and $[A_{1,tot}] = 10^8 \text{ cm}^{-3}$ seem to result, at least partly, from the limitations of our model, see Sects. 3.3.1 and 5.1) A second mechanism affecting the values of $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ is the diminished impact of the conversion of A_2T_2 -clusters to $A_2D_1T_1$ - and further to A_2D_2 -clusters. This effect becomes less significant as the A_2T_2 - and $A_2D_1T_1$ -clusters collide more often with an organic vapour molecule than with a DMA molecule. At amine concentrations equal to the applied organic vapour concentration, 10^9 cm^{-3} , this effect decreases the value of $K_{A_2T_2}$. The impact of organic vapour collisions on the values of the coefficient $K_{A_2D_1T_1}$ is similar to the above described effect on $K_{A_2T_2}$.

4 Comparison of model results with atmospheric observations

When interpreting the comparison of model results and ambient data, it should be kept in mind that the formation rate of 2 nm particles (or 1.5 nm particles in case of Hyytiälä 2003 and 2005 data) is somewhat lower than the initial particle formation rate. This is due to the coagulation losses during the growth of particles up to 2 nm (e.g., Kerminen and Kulmala, 2002). The diameter of the A_2B_2 -clusters comparable to the electrical mobility diameters applied in the ambient data would be roughly 1.2 nm (Ehn et al., 2011).

The measured atmospheric formation rate of 2 nm particles (J_{2nm}) is depicted in Fig. 8a as a function of the sulphuric acid monomer collision rate. In Fig. 8b we show the values of K_{obs} as a function of temperature. In addition to the observed values of the coefficient K_{obs} , we show in Fig. 8b the values of the modelled coefficients $K_{A_2D_2}$ and $K_{A_2T_2}$ at low amine concentrations for comparison. We do not present separately the comparison between K_{obs} and the sum of all coefficients $K_{A_2B_2}$, because this sum differs from $K_{A_2D_2}$ only in case of 1 : 10 ratio of [DMA] and [TMA], in which case it is practically equal to $K_{A_2T_2}$.

are associated with a cubed sulphuric acid concentration at $T > 290$ K, which is not observed in the atmospheric data. The smallest value for the variability, 130, resulted from the modelled $K_{A_2D_2}$ with a [DMA] to [TMA] ratio 1 : 10 in the absence of organic vapours. The smallest variabilities related to the modelled coefficients $K_{A_2T_2}$ or $K_{A_2D_1T_1}$ (360 and 280, respectively) were obtained with a [DMA] to [TMA]-ratio of 10 : 1 in the presence of organic vapours growing the A_2B_2 -clusters. The smaller variabilities between the observed and modelled formation rates being related to $K_{A_2D_2}$ than those related to coefficients $K_{A_2T_2}$ or $K_{A_2D_1T_1}$ in presence of organic vapours may result from the assumed constant amine concentration: as the overall temperature dependence of K_{obs} is closer to that of $K_{A_2D_2}$ (Fig. 8b), the smallest overall variability is also achieved with this coefficient. However, this analysis shows that the formation rates of A_2T_2 - or $A_2D_1T_1$ -clusters fail to reproduce the measured particle formation rates if organic vapours are not let to adhere to these clusters.

Even though the strongest impact on the modelled coefficients $K_{A_2B_2}$ is that of temperature, the different RH dependencies of $K_{A_2D_2}$, $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ (see Fig. 6) can provide a clue to which mechanism is closest to the atmospheric observations. To study the RH dependence of the observed coefficient K_{obs} , we first need to take into account the marked impact of the temperature on both K_{obs} and RH. This was made by scaling the values of the coefficient K_{obs} to the same temperature with the temperature dependence of the modelled coefficients $K_{A_2B_2}$. This scaling is functionally equivalent to moving all the data points in Fig. 8b to the same temperature by following the model lines, but now including the impact of varying concentrations $[A_{1,tot}]$ on the temperature dependencies of $K_{A_2B_2}$ (see Fig. 4). The numerical comparison of the correlations between RH and the values of the observed coefficients K_{obs} at the measured temperature, as well as those with the temperature scaled to 300 K (K_{300}) are presented in Table 3 for all the sites separately. The logarithmic value of K_{300} correlates negatively with RH at all the measurement sites except Hohenpeissenberg and Atlanta. However, at Atlanta, as well as at Melpitz, the correlation is not significant (p -value $\gg 10^{-5}$). The values of β in $K_{300} = \alpha \times \exp(\beta \times RH)$ from the linear regression fits to the K_{300} values

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

interpolated with $K_{A_2T_2}$ with organic vapours were in the range of -12 to -2.7 at all sites with a significant correlation except for Hohenpeissenberg. Similar values were achieved with other $K_{A_2B_2}$ -coefficients used for interpolation, as well. The linear regression fits to the RH-dependencies of $K_{A_2D_2}$, $K_{A_2T_2}$ and $K_{2A_1T_1T}$ (see Fig. 6) result in β -values of 0.033, -1.6 and -0.68 , respectively. Thus, the observed RH dependence suggests that the atmospheric particle formation rate is better associated with the formation rate of A_2T_2 -clusters than with the formation rate of A_2D_2 -clusters.

5 Discussion

5.1 On the stability of clusters with three or more acids

We investigated the formation rate of A_2B_2 -clusters with the assumption that the A_2B_2 -clusters with DMA or/and TMA are capable of growing further when colliding with sulphuric acid monomers. Ortega et al. (2012) have presented theoretical evidence that this is the case concerning the DMA-clusters, but the growth ability of A_2T_2 -clusters is not self-evident. This doubt arises from the fact that the three methyl groups in the TMA-molecule leave room for the nitrogen atom to make only one hydrogen bond, whereas the nitrogen in DMA can make two hydrogen bonds. Thus, the TMA molecules themselves are not capable of binding the third sulphuric acid (or other vapour molecule) to the A_2T_2 -cluster, but it could attach to the cluster by bonding to the sulphuric acid molecules. However, the high stability of the A_2T_1 -cluster, in which one other sulphuric acid molecule is not bound directly to the TMA, shows that also the bond between two sulphuric acids in a base-stabilized cluster can be very strong. In addition, some experimental studies have shown that increasing [TMA] increases the particle formation rate (Erupe et al., 2011; Yu et al., 2012).

In the discussion related to Figs. 4 and 7 we pointed out that the values of all coefficients $K_{A_2B_2}$ increase with the total sulphuric acid monomer concentration, $[A_{1,tot}]$, at high temperatures when $[A_{1,tot}]$ is orders of magnitude larger than the amine

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

concentration. This feature originates from the increasing concentrations of A_2B_1 -clusters, which under the given conditions exceeds the concentrations of gas phase amines, and thus the collisions between two A_2B_1 -clusters contributes significantly to the positive terms in $J_{A_2B_2}$ (see Eq. 9). Since the applied cluster energy calculations do not cover the resulting A_4B_2 -cluster, the negative terms describing the reverse evaporation are not affected. Ortega et al. (2012) showed that A_4D_2 -clusters evaporate A_2D_1 -cluster (thus forming two A_2D_1 -clusters) at relatively high rate (3 s^{-1} at 298.16 K). Thus, it seems that the above discussed increase in the values of $K_{A_2B_2}$ would be less pronounced than that seen in Figs. 4, 5 and 7, if the evaporation rates for larger clusters were taken into account. The evaporations of other $A_{\leq 4}D_{\leq 4}$ clusters resulting in clusters smaller than A_2D_2 are slow ($< 5 \times 10^{-3}\text{ s}^{-1}$ at 298.16 K, Ortega et al., 2012). Since A_4B_2 -clusters are formed efficiently only when $[A_{1,\text{tot}}]$ is orders of magnitude higher than $[B]$, we may assume that at other conditions the evaporations of the $A_{>2}D_{\geq 2}$ -clusters do not have marked effect on the formation rate $J_{A_2D_2}$. However, the effects of the corresponding evaporations on $J_{A_2T_2}$ and $J_{A_2D_1T_1}$ remain to be evaluated by further studies.

5.2 On the comparison to ambient data

We showed that the formation rates $J_{A_2B_2}$ are not always coupled with the sulphuric acid concentration to the power of exactly two, but the exponent varies from below two to greater than three. Values below two are related to the situation in which the amine and sulphuric acid monomer concentrations are of the same order of magnitude. In this case e.g. increasing the sulphuric acid concentration with constant amine concentration decreases the proportion of sulphuric acid monomers being clustered with amines, and consequently the formation rate of A_2B_2 -clusters per sulphuric acid monomer collisions decreases. The exponent values below two were found only for temperatures below 300 K. Exponent values larger than two are possible for the formation rate $J_{A_2D_2}$ only when $[A_{1,\text{tot}}]$ exceeds the amine concentration by more than

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

two orders of magnitude and the temperature is over 290 K. The formation rates $J_{A_2T_2}$ and $J_{A_2D_1T_1}$, instead, are associated with close to cubed sulphuric acid concentrations when $A_{1,tot}$ are the most abundant or only vapours/clusters able to promote further growth of A_2T_2 - and $A_2D_1T_1$ -clusters. If an organic vapour is present in equal or higher concentrations than that of sulphuric acid and can attach to the A_2B_2 -clusters, the exponent values of three are not encountered. Intriguingly, the range in which the exponent varies agrees with the atmospheric observations of the association between particle formation rate and sulphuric acid concentration, as the exponents basing on observations are typically from one to two, and in some cases even three. However, the exponents of the initial and observed particle formation rates are not necessarily the same, but they may change during the growth to the size of the observed particles (Korhonen et al., 2011).

Our model runs were made with three different [DMA] to [TMA] ratios, 1 : 1, 10 : 1 and 1 : 10. These ratios cover mostly the measured ratios at sites varying from rural to agricultural and urban areas (Ge et al., 2011). However, it should be noted that the measured DMA and TMA concentrations do not necessarily describe the concentrations of the free molecular concentrations, but perhaps also e.g. the amine molecules clustered with acids might be detected. This is a similar problem to the consideration of whether the CIMS detects the A_1B_1 -clusters or not, discussed in Sect. 2.1.2. Furthermore, in the atmosphere a wide variety of different amines exists, and other amines can also play a role in the cluster formation process. In particular, monomethylamine and some ethylamines have been reported to have concentrations comparable to those of DMA and TMA (Ge et al., 2011). They could have a significant effect on particle formation – those with more alkyl-groups due to their strong basicity, and e.g. monomethylamine due to its ability to form up to three hydrogen bonds and thus possibly facilitate the formation of larger sulphuric acid-amine-clusters together with other more basic amines.

The steepness of the decreasing values of $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ with temperature results from two cluster types, A_1B_1 and A_2B_2 , with evaporation rates that strongly affect

complexity of determining the impact of temperature on atmospheric particle formation. There are several processes which may elevate the observed particle formation rate when temperature increases: in addition to decreasing humidity and presumably increasing amine concentration, increasing temperature elevates exponentially the emissions of volatile organic compounds (Gunther et al., 1995) which after oxidation grow the freshly formed particles (e.g., Riipinen et al., 2011). Thus, even if the particle formation rate did not increase with temperature, at least the proportion of formed particles surviving to 2 nm size should. With all these factors potentially increasing the observed particle formation rate at higher temperatures, the suppressing effect of temperature itself must be quite strong to result in negative correlation between the temperature and the coefficient K_{obs} .

6 Summary and conclusions

We modelled the formation rate of clusters consisting of two sulphuric acid and two amine molecules with a kinetic cluster model (DACM, Kulmala, 2010). These clusters are necessary intermediates in the formation of larger sulphuric acid-amine-clusters. If atmospheric particle formation is a phenomenon related to formation of acid-base-clusters, it seems evident that the atmospheric particle formation rate should not exceed the formation rate of these clusters. We applied quantum chemical calculations for determining the evaporation rates for the clusters of sulphuric acid (A), trimethylamine (T) and dimethylamine (D) in a temperature range from 260 K to 320 K, and studied the formation rates $J_{A_2B_2}$ of $A_2T_2^-$, $A_2D_2^-$ and $A_2D_1T_1^-$ -clusters.

Our purpose was to perform the modelling with input parameters as close to atmospheric experimental circumstances as possible. As recent experimental (Petäjä et al., 2011) and theoretical (Kurtén et al., 2011) studies have suggested, the state-of-art instrument for measuring sulphuric acid concentration, Chemical Ionization Mass Spectrometer (CIMS), detects the main fraction of $A_1D_1^-$ -clusters as pure sulphuric acid. We assumed that the same holds for $A_1T_1^-$ -clusters, and thus compared the formation rates

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

$J_{A_2B_2}$ with the total concentration of the sulphuric acid monomers, $[A_{1,tot}]$, calculated as a sum of the concentrations of A, A_1D_1 and A_1T_1 . We ran the model with $A_{1,tot}$ and amine concentrations from 10^5 to 10^9 cm^{-3} , which likely covers the typical atmospheric conditions, even though direct measurements of the gas phase amine concentrations are currently not available. The effect of meteorological conditions and vapours participating to particle formation was studied by examining the behaviour of coefficients $K_{A_2B_2}$ in $J_{A_2B_2} = K_{A_2B_2} \times [A_{1,tot}]^2$ as functions of temperature, RH and vapour concentrations.

We found that the coefficients $K_{A_2B_2}$ are dependent on the amine and sulphuric acid concentration, on the ratios between these concentrations and of the amine concentrations with each other, as well as on temperature and RH. In the case where the amine concentration is an order of magnitude higher than $[A_{1,tot}]$ and the temperature is below 270 K, a large enough fraction of sulphuric acid molecules is clustered with the amines to result in a formation rate close to the collision rate of $A_{1,tot}$. Close to collision rate values for $J_{A_2B_2}$ were also found at higher temperatures, up to 300 K, under high amine concentrations ($> 10^8$ cm^{-3}).

The temperature dependence of the coefficients $K_{A_2B_2}$ arising from the exponential increase of the cluster evaporation rates as functions of temperature, was found out to be very strong. The value of the coefficient $K_{A_2D_2}$ decreased roughly by an order of magnitude for an increase of 15 K in temperature, because of the increasing evaporation rate from the A_1D_1 -cluster. A similar temperature change decreased the values of the coefficients $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ by two orders of magnitude, because in addition to evaporation of A_1D_1 - or A_1T_1 -clusters the evaporation rates from A_2T_2 and $A_2D_1T_1$ -clusters limited the formation rate. The exponential decrease of the coefficients $K_{A_2B_2}$ started in temperatures between 260 and 310 K. The lowest on-set temperatures for the decrease were related to low atmospheric sulphuric acid concentrations ($\sim 10^6$ cm^{-3}) together with similar or lower amine concentrations, and the on-set temperature increased with increasing $[A_{1,tot}]$ and when the amine concentration was orders of magnitude higher than $[A_{1,tot}]$. In addition to temperature, the relative humidity had an im-

pact on the formation rates $J_{A_2B_2}$. The values of the coefficients $K_{A_2T_2}$ and $K_{A_2D_1T_1}$ decreased with an increasing RH, whereas $K_{A_2D_2}$ was not significantly affected by RH. This effect was less pronounced than that of temperature: the steepest decrease was obtained for $K_{A_2T_2}$ which decreased roughly by a factor two when RH increased by 20 %-units.

In addition to the influence of meteorological parameters and amine concentration (or the proportion of amine concentration to sulphuric acid concentration), we found that the values of $K_{A_2B_2}$ are affected directly by the concentration of sulphuric acid. Thus the formation rates $J_{A_2B_2}$ are not always associated with sulphuric acid concentration to power of two exactly, but at low atmospherically relevant temperatures the value of the exponent decreases from two. On the other hand, at high atmospherically relevant temperatures the formation rates $J_{A_2T_2}$ and $J_{A_2D_1T_1}$ turned out to be associated with sulphuric acid to the power of three, if the sulphuric acid is the only or the main compound growing the A_2B_2 -clusters further. Such power law dependencies with exponents higher than two have been reported, e.g. during the hot season in Beijing (Wang et al., 2012).

Finally we made a comparison between the modelled coefficients $K_{A_2B_2}$ and the coefficients K_{obs} determined from the measurements of particle formation rates and sulphuric acid concentrations conducted in diverse atmospheric conditions. When K_{obs} is plotted against temperature, the overall figure has a similar decreasing trend with temperature than that of $K_{A_2D_2}$. However, at several sites the coefficient K_{obs} exerted steeper decrease with increasing temperature, similar to those of the modelled coefficients $K_{A_2T_2}$ and $K_{A_2D_1T_1}$. Furthermore, the coefficient K_{obs} showed a similar decreasing trend with increasing RH to the coefficient $K_{A_2T_2}$. Even if the amines mainly determining the atmospheric particle formation rates were different than those studied here, our results give some indications of the process. The similarities between atmospheric particle formation rates and the modelled formation rates $J_{A_2T_2}$ and $J_{A_2D_1T_1}$ suggest that atmospheric particle formation from gas phase acids and bases occurs via exactly two cluster sizes for which evaporation is the main sink. Because the inverse

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

proportionality between coefficients $K_{A_2B_2}$ related to TMA and RH arises eventually from the complete surrounding of the nitrogen atom in TMA by alkyl groups (thus leaving no room for hydrogen bonds with water), the observed decrease in particle formation rates with increasing RH suggest that atmospheric particle formation is related, if not particularly to TMA, then to some other tertiary amines.

The results of our study show that both evolution of measurement techniques for determining the atmospheric amine concentrations and further quantum chemical calculations on the sulphuric acid-amine-clusters are necessary for determining the main pathways of the atmospheric particle formation. Elucidating the stabilities of clusters with other amines in the size range studied here would substantially increase our understanding of the distribution of atmospheric sulphuric acid clusters and their capability to grow. Furthermore, the stability of clusters consisting of TMA and more than two sulphuric acid molecules, as well as the possibility of some oxidized organic vapours to adhere to the A_2B_2 -clusters, are of much interest. From the experimental point of view, the effects of temperature and the prevailing sulphuric acid concentration on the power of sulphuric acid concentration associated with the observed particle formation rates should be investigated in more detail, both within and between particle formation event days.

Supplementary material related to this article is available online at:

<http://www.atmos-chem-phys-discuss.net/12/11485/2012/acpd-12-11485-2012-supplement.pdf>.

Acknowledgements. This work was funded by European Research Council (27463-ATMNUCLE and 257360-MOCAPAF), Academy of Finland (project number 1127372, Center of Excellence program project number 1118615, LASTU program project number 135054), and Maj and Tor Nessling Foundation (project numbers 2010212 and 2011200). We thank CSC – IT Center for Science in Espoo, Finland, for computing time, and Sanna-Liisa Sihto for providing the particle formation rate data related to QUEST II and IV campaigns.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



References

- Berndt, T., Stratmann, F., Sipilä, M., Vanhanen, J., Petäjä, T., Mikkilä, J., Grüner, A., Spindler, G., Lee Mauldin III, R., Curtius, J., Kulmala, M., and Heintzenberg, J.: Laboratory study on new particle formation from the reaction OH + SO₂: influence of experimental conditions, H₂O vapour, NH₃ and the amine tert-butylamine on the overall process, *Atmos. Chem. Phys.*, 10, 7101–7116, doi:10.5194/acp-10-7101-2010, 2010.
- Berresheim, H., Elste, T., Plass-Dülmer, C., Eisele, F. L., and Tanner, D. J.: Chemical ionization mass spectrometer for long-term measurements of atmospheric OH and H₂SO₄, *Int. J. Mass Spectrom.*, 202, 91–109, 2000.
- Bzdek, B. R., Ridge, D. P., and Johnston, M. V.: Amine reactivity with charged sulfuric acid clusters, *Atmos. Chem. Phys.*, 11, 8735–8743, doi:10.5194/acp-11-8735-2011, 2011.
- Dal Maso, M., Sogacheva, L., Aalto, P. P., Riipinen, I., Komppula, M., Tunved, P., Korhonen, L., Suur-Uski, V., Hirsikko, A., Kurten, T., Kerminen, V.-M., Lihavainen, H., Viisanen, Y., Hansson, H.-C., and Kulmala, M.: Aerosol size distribution measurements at four Nordic field stations: identification, analysis and trajectory analysis of new particle formation events, *Tellus B*, 59, 350–361, 2007.
- Ehn, M., Junninen, H., Schobesberger, S., Manninen, H. E., Franchin, A., Sipilä, M., Petäjä, T., Kerminen, V.-M., Tammet, H., Mirme, A., Mirme, S., Hörrak, U., Kulmala, M., and Worsnop, D. R.: An instrumental comparison of mobility and mass measurements of atmospheric small ions, *Aerosol Sci. Tech.*, 45, 522–532, 2011.
- Erupe, M. E., Viggiano, A. A., and Lee, S.-H.: The effect of trimethylamine on atmospheric nucleation involving H₂SO₄, *Atmos. Chem. Phys.*, 11, 4767–4775, doi:10.5194/acp-11-4767-2011, 2011.
- Ge, X., Wexler, A. S., and Clegg, S. L.: Atmospheric amines – Part I. A review, *Atmos. Environ.*, 45, 524–546, 2011.
- Guenther, A., Hewitt, C. N., Erickson, D., Fall, R., Geron, C., Graedel, T., Harley, P., Klinger, L., Lerdau, M., McKay, W. A., Pierce, T., Scholes, B., Steinbrecher, R., Tallamraju, R., Taylor, J., and Zimmermann, P.: A global model of natural volatile organic compound emissions, *J. Geophys. Res.*, 100, 8873–8892, doi:10.1029/94JD02950, 1995.
- Hamed, A., Korhonen, H., Sihto, S.-L., Joutsensaari, J., Järvinen, H., Petäjä, T., Arnold, F., Nieminen, T., Kulmala, M., Smith, J. N., Lehtinen, K. E. J., and Laaksonen, A.: The role

ACPD

12, 11485–11537, 2012

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

of relative humidity in continental new particle formation, *J. Geophys. Res.*, 116, D03202, doi:10.1029/2010JD014186, 2011.

Hanson, D. R., McMurry, P. H., Jiang, J., Tanner, D., and Huey, L. G.: Ambient pressure proton transfer mass spectrometry: detection of amines and ammonia, *Environ. Sci. Technol.*, 45, 8881–8888, 2011.

Iida, K., Stolzenburg, M. R., McMurry, P. H., and Smith, J. N.: Estimating nanoparticle growth rates from size-dependent charged fractions: analysis of new particle formation events in Mexico City, *J. Geophys. Res.*, 113, D05207, doi:10.1029/2007JD009260, 2008.

Kazil, J., Stier, P., Zhang, K., Quaas, J., Kinne, S., O'Donnell, D., Rast, S., Esch, M., Ferrachat, S., Lohmann, U., and Feichter, J.: Aerosol nucleation and its role for clouds and Earth's radiative forcing in the aerosol-climate model ECHAM5-HAM, *Atmos. Chem. Phys.*, 10, 10733–10752, doi:10.5194/acp-10-10733-2010, 2010.

Kerminen, V.-M. and Kulmala, M.: Analytical formulae connecting the “real” and the “apparent” nucleation rate and the nuclei number concentration for atmospheric nucleation events, *J. Aerosol Sci.*, 33, 609–622, 2002.

Kerminen, V.-M., Petäjä, T., Manninen, H. E., Paasonen, P., Nieminen, T., Sipilä, M., Junninen, H., Ehn, M., Gagné, S., Laakso, L., Riipinen, I., Vehkamäki, H., Kurten, T., Ortega, I. K., Dal Maso, M., Brus, D., Hyvärinen, A., Lihavainen, H., Leppä, J., Lehtinen, K. E. J., Mirme, A., Mirme, S., Hörrak, U., Berndt, T., Stratmann, F., Birmili, W., Wiedensohler, A., Metzger, A., Dommen, J., Baltensperger, U., Kiendler-Scharr, A., Mentel, T. F., Wildt, J., Winkler, P. M., Wagner, P. E., Petzold, A., Minikin, A., Plass-Dülmer, C., Pöschl, U., Laaksonen, A., and Kulmala, M.: Atmospheric nucleation: highlights of the EUCAARI project and future directions, *Atmos. Chem. Phys.*, 10, 10829–10848, doi:10.5194/acp-10-10829-2010, 2010.

Kirkby, J., Curtius, J., Almeida, J., Dunne, E., Duplissy, J., Ehrhart, S., Franchin, A., Gagné, S., Ickes, L., Kürten, A., Kupc, A., Metzger, A., Riccobono, F., Rondo, L., Schobesberger, S., Tsagkogeorgas, G., Wimmer, D., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Downard, A., Ehn, M., Flagan, R. C., Haider, S., Hansel, A., Hauser, D., Jud, W., Junninen, H., Kreissl, F., Kvashin, A., Laaksonen, A., Lehtipalo, K., Lima, J., Lovejoy, E. R., Makhutov, V., Mathot, S., Mikkilä, J., Minginette, P., Mogo, S., Nieminen, T., Onnela, A., Pereira, A., Petäjä, T., Schnitzhofer, R., Seinfeld, J. H., Sipilä, M., Stozhkov, Y., Stratmann, F., Tome, A., Vanhanen, J., Viisanen, Y., Vrtala, A., Wagner, P. E., Walther, H., Weingartner, E., Wex, H., Winkler, P. M., Carslaw, K. S., Worsnop, D. R., Baltensperger, U., and Kulmala, M.:

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



The role of sulfuric acid, ammonia and galactic cosmic rays in atmospheric aerosol nucleation, *Nature*, 476, 429–433, 2011.

Korhonen, H., Sihto, S.-L., Kerminen, V.-M., and Lehtinen, K. E. J.: Evaluation of the accuracy of analysis tools for atmospheric new particle formation, *Atmos. Chem. Phys.*, 11, 3051–3066, doi:10.5194/acp-11-3051-2011, 2011.

Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F., L.: Dependence of nucleation rates on sulfuric acid vapor concentration in diverse atmospheric locations, *J. Geophys. Res.*, 113, D10209, doi:10.1029/2007JD009253, 2008.

Kulmala, M.: Dynamical atmospheric cluster model, *Atmos. Res.*, 98, 201–206, 2010.

Kulmala, M. and Kerminen, V.-M.: On the formation and growth of atmospheric nanoparticles, *Atmos. Res.*, 90, 132–150, 2008.

Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V.-M., Birmili, W., and McMurry, P. H.: Formation and growth rates of ultrafine atmospheric particles: a review of observations, *J. Aerosol Sci.*, 35, 143–176, 2004.

Kulmala, M., Lehtinen, K. E. J., and Laaksonen, A.: Cluster activation theory as an explanation of the linear dependence between formation rate of 3nm particles and sulphuric acid concentration, *Atmos. Chem. Phys.*, 6, 787–793, doi:10.5194/acp-6-787-2006, 2006.

Kulmala, M., Riipinen, I., Sipilä, M., Manninen, H. E., Petäjä, T., Junninen, H., Dal Maso, M., Mordas, G., Mirme, A., Vana, M., Hirsikko, A., Laakso, L., Harrison, R. M., Hanson, I., Leung, C., Lehtinen, K. E. J., and Kerminen, V.-M.: Toward direct measurement of atmospheric nucleation, *Science*, 318, 89–92, 2007.

Kulmala, M., Asmi, A., Lappalainen, H. K., Baltensperger, U., Brenguier, J.-L., Facchini, M. C., Hansson, H.-C., Hov, Ø., O'Dowd, C. D., Pöschl, U., Wiedensohler, A., Boers, R., Boucher, O., de Leeuw, G., Denier van der Gon, H. A. C., Feichter, J., Krejci, R., Laj, P., Lihavainen, H., Lohmann, U., McFiggans, G., Mentel, T., Pilinis, C., Riipinen, I., Schulz, M., Stohl, A., Swietlicki, E., Vignati, E., Alves, C., Amann, M., Ammann, M., Arabas, S., Artaxo, P., Baars, H., Beddows, D. C. S., Bergström, R., Beukes, J. P., Bilde, M., Burkhardt, J. F., Canonaco, F., Clegg, S. L., Coe, H., Crumeyrolle, S., D'Anna, B., Decesari, S., Gilaroni, S., Fischer, M., Fjaeraa, A. M., Fountoukis, C., George, C., Gomes, L., Hälloran, P., Hamburger, T., Harrison, R. M., Herrmann, H., Hoffmann, T., Hoose, C., Hu, M., Hyvärinen, A., Hörrak, U., Iinuma, Y., Iversen, T., Josipovic, M., Kanakidou, M., Kiendler-Scharr, A., Kirkevåg, A., Kiss, G., Klimont, Z., Kolmonen, P., Komppula, M., Kristjánsson, J.-E., Laakso, L., Laaksonen, A., Labonnote, L., Lanz, V. A., Lehtinen, K. E. J., Rizzo, L. V.,

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Makkonen, R., Manninen, H. E., McMeeking, G., Merikanto, J., Minikin, A., Mirme, S., Morgan, W. T., Nemitz, E., O'Donnell, D., Panwar, T. S., Pawlowska, H., Petzold, A., Pienaar, J. J., Pio, C., Plass-Duelmer, C., Prévôt, A. S. H., Pryor, S., Reddington, C. L., Roberts, G., Rosenfeld, D., Schwarz, J., Seland, Ø., Sellegri, K., Shen, X. J., Shiraiwa, M., Siebert, H., Sierau, B., Simpson, D., Sun, J. Y., Topping, D., Tunved, P., Vaattovaara, P., Vakkari, V., Veeffkind, J. P., Visschedijk, A., Vuollekoski, H., Vuolo, R., Wehner, B., Wildt, J., Woodward, S., Worsnop, D. R., van Zadelhoff, G.-J., Zardini, A. A., Zhang, K., van Zyl, P. G., Kerminen, V.-M., S Carslaw, K., and Pandis, S. N.: General overview: European Integrated project on Aerosol Cloud Climate and Air Quality interactions (EUCAARI) – integrating aerosol research from nano to global scales, *Atmos. Chem. Phys.*, 11, 13061–13143, doi:10.5194/acp-11-13061-2011, 2011.

Kurtén, T., Loukonen, V., Vehkamäki, H., and Kulmala, M.: Amines are likely to enhance neutral and ion-induced sulfuric acid-water nucleation in the atmosphere more effectively than ammonia, *Atmos. Chem. Phys.*, 8, 4095–4103, doi:10.5194/acp-8-4095-2008, 2008.

Kurtén, T., Petäjä, T., Smith, J., Ortega, I. K., Sipilä, M., Junninen, H., Ehn, M., Vehkamäki, H., Mauldin, L., Worsnop, D. R., and Kulmala, M.: The effect of H₂SO₄ – amine clustering on chemical ionization mass spectrometry (CIMS) measurements of gas-phase sulfuric acid, *Atmos. Chem. Phys.*, 11, 3007–3019, doi:10.5194/acp-11-3007-2011, 2011.

Loukonen, V., Kurtén, T., Ortega, I. K., Vehkamäki, H., Pádua, A. A. H., Sellegri, K., and Kulmala, M.: Enhancing effect of dimethylamine in sulfuric acid nucleation in the presence of water – a computational study, *Atmos. Chem. Phys.*, 10, 4961–4974, doi:10.5194/acp-10-4961-2010, 2010.

Makkonen, R., Asmi, A., Kerminen, V.-M., Boy, M., Arneth, A., Hari, P., and Kulmala, M.: Air pollution control and decreasing new particle formation lead to strong climate warming, *Atmos. Chem. Phys.*, 12, 1515–1524, doi:10.5194/acp-12-1515-2012, 2012.

Manninen, H. E., Nieminen, T., Riipinen, I., Yli-Juuti, T., Gagné, S., Asmi, E., Aalto, P. P., Petäjä, T., Kerminen, V.-M., and Kulmala, M.: Charged and total particle formation and growth rates during EUCAARI 2007 campaign in Hyytiälä, *Atmos. Chem. Phys.*, 9, 4077–4089, doi:10.5194/acp-9-4077-2009, 2009.

Mauldin III, R., Tanner, D., Heath, J., Huebert, B., and Eisele, F.: Observations of H₂SO₄ and MSA during PEM-Tropics-A, *J. Geophys. Res.*, 104, 5801–5816, 1999.

McMurry, P. H. and Eisele, F. L.: Preface to topical collection on nucleation and growth in urban Atlanta, *J. Geophys. Res.*, 110, D22S01, doi:10.1029/2005JD006644, 2005.

- McMurry, P. H. and Friedlander, S. K.: New particle formation in the presence of aerosol, *Atmos. Environ.*, 13, 1635–1651, 1979.
- Meng, Z. Y., Lin, W. L., Jiang, X. M., Yan, P., Wang, Y., Zhang, Y. M., Jia, X. F., and Yu, X. L.: Characteristics of atmospheric ammonia over Beijing, China, *Atmos. Chem. Phys.*, 11, 6139–6151, doi:10.5194/acp-11-6139-2011, 2011.
- Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J., and Carslaw, K. S.: Impact of nucleation on global CCN, *Atmos. Chem. Phys.*, 9, 8601–8616, doi:10.5194/acp-9-8601-2009, 2009.
- Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S., Weingartner, E., Riipinen, I., Kulmala, M., Spracklen, D. V., Carslaw, K. S., and Baltensperger, U.: Evidence for the role of organics in aerosol particle formation under atmospheric conditions, *P. Natl. Acad. Sci. USA*, 107, 6646–6651, doi:10.1073/pnas.0911330107, 2010.
- Mikkonen, S., Romakkaniemi, S., Smith, J. N., Korhonen, H., Petäjä, T., Plass-Duelmer, C., Boy, M., McMurry, P. H., Lehtinen, K. E. J., Joutsensaari, J., Hamed, A., Mauldin III, R. L., Birmili, W., Spindler, G., Arnold, F., Kulmala, M., and Laaksonen, A.: A statistical proxy for sulphuric acid concentration, *Atmos. Chem. Phys.*, 11, 11319–11334, doi:10.5194/acp-11-11319-2011, 2011.
- Mirme, S., Mirme, A., Minikin, A., Petzold, A., Hörrak, U., Kerminen, V. -M., and Kulmala, M.: Atmospheric sub-3 nm particles at high altitudes, *Atmos. Chem. Phys.*, 10, 437–451, doi:10.5194/acp-10-437-2010, 2010.
- Murphy, S. M., Sorooshian, A., Kroll, J. H., Ng, N. L., Chhabra, P., Tong, C., Surratt, J. D., Knipping, E., Flagan, R. C., and Seinfeld, J. H.: Secondary aerosol formation from atmospheric reactions of aliphatic amines, *Atmos. Chem. Phys.*, 7, 2313–2337, doi:10.5194/acp-7-2313-2007, 2007.
- Nadykto, A. B., Yu, F., Jakovleva, M. V., Herb, J., and Xu, Y.: Amines in the Earth's atmosphere: a density functional theory study of the thermochemistry of pre-nucleation clusters, *Entropy*, 13, 554–569, 2011.
- Ortega, I. K., Kupiainen, O., Kurtén, T., Olenius, T., Wilkman, O., McGrath, M. J., Loukonen, V., and Vehkamäki, H.: From quantum chemical formation free energies to evaporation rates, *Atmos. Chem. Phys.*, 12, 225–235, doi:10.5194/acp-12-225-2012, 2012.
- Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flenje, H., Birmili, W., Wiedensohler, A., Hörrak, U., Metzger, A., Hamed, A., Laaksonen, A., Facchini, M. C., Kerminen, V.-M., and Kulmala, M.: On the roles of sulphuric acid and

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



low-volatility organic vapours in the initial steps of atmospheric new particle formation, *Atmos. Chem. Phys.*, 10, 11223–11242, doi:10.5194/acp-10-11223-2010, 2010.

Petäjä, T., Mauldin, III, R. L., Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy, M., Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a boreal forest site, *Atmos. Chem. Phys.*, 9, 7435–7448, doi:10.5194/acp-9-7435-2009, 2009.

Petäjä, T., Sipilä, M., Paasonen, P., Nieminen, T., Kurtén, T., Ortega, I. K., Stratmann, F., Vehkamäki, H., Berndt, T., and Kulmala, M.: Experimental observation of strongly bound dimers of sulphuric acid nucleating in the atmosphere, *Phys. Rev. Lett.*, 106, 228302, doi:10.1103/PhysRevLett.106.228302, 2011.

Pierce, J. R. and Adams, P. J.: Uncertainty in global CCN concentrations from uncertain aerosol nucleation and primary emission rates, *Atmos. Chem. Phys.*, 9, 1339–1356, doi:10.5194/acp-9-1339-2009, 2009.

Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K., Teinilä, K., Kerminen, V.-M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between atmospheric sulphuric acid and new particle formation during QUEST III–IV campaigns in Heidelberg and Hyytiälä, *Atmos. Chem. Phys.*, 7, 1899–1914, doi:10.5194/acp-7-1899-2007, 2007.

Riipinen, I., Pierce, J. R., Yli-Juuti, T., Nieminen, T., Häkkinen, S., Ehn, M., Junninen, H., Lehtipalo, K., Petäjä, T., Slowik, J., Chang, R., Shantz, N. C., Abbatt, J., Leaitch, W. R., Kerminen, V.-M., Worsnop, D. R., Pandis, S. N., Donahue, N. M., and Kulmala, M.: Organic condensation: a vital link connecting aerosol formation to cloud condensation nuclei (CCN) concentrations, *Atmos. Chem. Phys.*, 11, 3865–3878, doi:10.5194/acp-11-3865-2011, 2011.

Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Dal Maso, M., Petäjä, T., Riipinen, I., Korhonen, H., Arnold, F., Janson, R., Boy, M., Laaksonen, A., and Lehtinen, K. E. J.: Atmospheric sulphuric acid and aerosol formation: implications from atmospheric measurements for nucleation and early growth mechanisms, *Atmos. Chem. Phys.*, 6, 4079–4091, doi:10.5194/acp-6-4079-2006, 2006.

Sipilä, M., Berndt, T., Petäjä, T., Brus, D., Vanhanen, J., Stramann, F., Patokoski, J., Mauldin III, R. L., Hyvärinen, A.-P., Lihavainen, H., and Kulmala, M.: The role of sulfuric acid in atmospheric nucleation, *Science*, 327, 1243–1246, doi:10.1126/science.1180315, 2010.

Sjöstedt, S. J., Huey, L. G., Tanner, D. J., Peischl, J., Chen, G., Dibb, J. E., Lefer, B., Hutterli, M. A., Beyersdorf, A. J., Blake, N. J., Blake, D. R., Sueper, D., Ryerson, T., Burkhardt, J.,

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

- and Stohl, A.: Observations of hydroxyl and the sum of peroxy radicals at Summit, Greenland during summer 2003, *Atmos. Environ.*, 41(24), 5122–5137, 2007.
- Spindler, G., Teichmann, U., and Sutton, M. A.: Ammonia dry deposition over grassland-micrometeorological flux-gradient measurements and bidirectional flux calculations using an inferential model, *Q. J. Roy. Meteor. Soc.*, 127, 795–814, 2001.
- 5 Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V.-M., Mann, G. W., and Sihto, S.-L.: The contribution of boundary layer nucleation events to total particle concentrations on regional and global scales, *Atmos. Chem. Phys.*, 6, 5631–5648, doi:10.5194/acp-6-5631-2006, 2006.
- 10 Spracklen, D. V., Carslaw, K. S., Kulmala, M., Kerminen, V.-M., Sihto, S.-L., Riipinen, I., Merikanto, J., Mann, G. W., Chipperfield, M. P., Wiedensohler, A., Birmili, W., and Lihavainen, H.: Contribution of particle formation to global cloud condensation nuclei concentrations, *Geophys. Res. Lett.*, 35, L06808, doi:10.1029/2007GL033038, 2008.
- Vehkamäki, H., McGrath, M. J., Kurtén, T., Julin, J., Lehtinen, K. E. J., and Kulmala, M.: Re-thinking the application of the first nucleation theorem to particle formation, *J. Chem. Phys.*, 15 136, 094107, doi:10.1063/1.3689227, 2012.
- Wang, M. and Penner, J. E.: Aerosol indirect forcing in a global model with particle nucleation, *Atmos. Chem. Phys.*, 9, 239–260, doi:10.5194/acp-9-239-2009, 2009.
- Wang, Z. B., Hu, M., Yue, D. L., Zheng, J., Zhang, R. Y., Wiedensohler, A., Wu, Z. J., Nieminen, T., and Boy, M.: Evaluation on the role of sulfuric acid in the mechanisms of new particle formation for Beijing case, *Atmos. Chem. Phys.*, 11, 12663–12671, doi:10.5194/acp-11-12663-2011, 2011.
- 20 Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measured atmospheric new particle formation rates: implications for nucleation mechanisms, *Chem. Eng. Commun.*, 151, 53–64, 1996.
- Yu, F. and Luo, G.: Simulation of particle size distribution with a global aerosol model: contribution of nucleation to aerosol and CCN number concentrations, *Atmos. Chem. Phys.*, 9, 7691–7710, doi:10.5194/acp-9-7691-2009, 2009.
- Yu, F., Luo, G., Bates, T. S., Anderson, B., Clarke, A., Kapustin, V., Yantosca, R. M., Wang, Y., and Wu, S.: Spatial distributions of particle number concentrations in the global troposphere: Simulations, observations, and implications for nucleation mechanisms, *J. Geophys. Res.*, 30 115, D17205, doi:10.1029/2009JD013473, 2010.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[⏪](#)[⏩](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

Yu, H., McGraw, R., and Lee, S. H.: Effects of amines on formation of sub-3 nm particles and their subsequent growth, *Geophys. Res. Lett.*, 39, L02807, doi:10.1029/2011GL050099, 2012.

Zhang, R., Khalizov, A., Wang, L., Hu, M., and Xu, W.: Nucleation and growth of nanoparticles in the atmosphere, *Chem. Rev.*, 114, 5839–5849, doi:10.1021/cr2001756, 2012.

Zhao, J., Smith, J. N., Eisele, F. L., Chen, M., Kuang, C., and McMurry, P. H.: Observation of neutral sulfuric acid-amine containing clusters in laboratory and ambient measurements, *Atmos. Chem. Phys.*, 11, 10823–10836, doi:10.5194/acp-11-10823-2011, 2011.

Zheng, J., Hu, M., Zhang, R., Yue, D., Wang, Z., Guo, S., Li, X., Bohn, B., Shao, M., He, L., Huang, X., Wiedensohler, A., and Zhu, T.: Measurements of gaseous H₂SO₄ by AP-ID-CIMS during CAREBeijing 2008 Campaign, *Atmos. Chem. Phys.*, 11, 7755–7765, doi:10.5194/acp-11-7755-2011, 2011.

Table 1. Formation enthalpies ΔH and entropies ΔS , calculated at 298 K, and the corresponding Gibbs free energies ΔG at $T = 280$ K. The values of ΔH and ΔG are given in [kcal mol^{-1}], and ΔS in [$\text{cal mol}^{-1} \text{K}^{-1}$].

Cluster	ΔH	ΔS	ΔG	Cluster	ΔH	ΔS	ΔG
A ₂	-17.85	-33.42	-8.49	A ₁ W ₁	-11.34	-29.70	-3.02
A ₃	-35.82	-72.19	-15.61	A ₁ W ₂	-22.49	-60.69	-5.50
A ₁ D ₁	-24.65	-31.01	-15.97	A ₁ W ₃	-33.89	-100.19	-5.84
A ₂ D ₁	-57.06	-74.40	-36.23	A ₁ W ₄	-45.97	-131.91	-9.04
A ₃ D ₁	-81.46	-112.90	-49.85	A ₁ W ₅	-54.74	-160.16	-9.89
A ₁ D ₂	-40.08	-66.37	-21.5	A ₂ W ₁	-31.64	-68.74	-12.39
A ₂ D ₂	-87.52	-105.16	-58.08	A ₂ W ₂	-45.67	-106.43	-15.87
A ₃ D ₂	-113.58	-155.37	-70.08	A ₂ W ₃	-57.35	-136.68	-19.08
A ₂ D ₃	-108.24	-152.75	-65.47	A ₂ W ₄	-68.27	-168.62	-21.06
A ₁ T ₁	-26.04	-34.54	-16.37	A ₂ W ₅	-82.95	-208.44	-24.59
A ₂ T ₁	-57.16	-75.30	-36.08	D ₁ W ₁	-6.65	-28.43	1.31
A ₃ T ₁	-75.32	-112.92	-43.70	D ₁ W ₂	-15.80	-66.40	2.79
A ₁ T ₂	-42.23	-71.66	-22.17	A ₁ D ₁ W ₁	-37.89	-65.76	-19.48
A ₂ T ₂	-84.48	-118.41	-51.33	A ₁ D ₁ W ₂	-48.79	-99.37	-20.97
A ₁ D ₁ T ₁	-41.01	-68.15	-21.93	A ₁ D ₁ W ₃	-58.81	-134.40	-21.18
A ₂ D ₁ T ₁	-83.86	-113.64	-52.04	A ₁ D ₁ W ₄	-69.46	-164.94	-23.28
				A ₁ D ₁ W ₅	-80.10	-199.35	-24.28
				A ₁ T ₁ W ₁	-35.52	-66.46	-16.91
				A ₁ T ₁ W ₂	-45.20	-95.26	-18.53

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

⏪

⏩

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Table 2. Evaporation rates from the modelled clusters at $T = 280$ K in units $[\text{s}^{-1}]$. The fastest evaporation rate from each cluster type is highlighted with bold font.

Initial cluster	Evaporated molecule/cluster					Initial cluster	Evaporated molecule			
	A	D	A ₁ D ₁	T	A ₁ T ₁		W	A ₁ W _n	D	T
A ₂	2 × 10³	–	–	–	–	A ₁ W ₁	5 × 10 ⁷	5 × 10 ⁷	–	–
A ₃	3 × 10⁴	–	–	–	–	A ₁ W ₂	2 × 10 ⁸	2 × 10 ⁸	–	–
A ₁ D ₁	4 × 10⁻³	4 × 10⁻³	–	–	–	A ₂	–	2 × 10 ³	–	–
A ₂ D ₁	2 × 10⁻⁶	3 × 10 ⁻¹²	2 × 10⁻⁶	–	–	A ₂ W ₁	1 × 10 ⁷	5 × 10 ²	–	–
A ₃ D ₁	3 × 10⁻¹	3 × 10 ⁻¹⁷	2 × 10 ⁻¹⁰	–	–	A ₂ W ₂	3 × 10 ⁷	2 × 10 ²	–	–
A ₁ D ₂	7 × 10 ⁻¹⁰	7 × 10⁵	7 × 10⁵	–	–	A ₂ W ₃	5 × 10 ⁷	6 × 10 ¹	–	–
A ₂ D ₂	3 × 10 ⁻¹⁹	1 × 10⁻⁷	5 × 10 ⁻¹¹	–	–	A ₂ W ₄	5 × 10 ⁸	2 × 10 ²	–	–
A ₃ D ₂	5 × 10⁰	3 × 10 ⁻⁶	1 × 10 ⁻⁴	–	–	A ₁ D ₁	–	4 × 10 ⁻³	4 × 10 ⁻³	–
A ₂ D ₃	1 × 10 ⁻²⁰	3 × 10⁴	2 × 10 ⁻¹²	–	–	A ₁ D ₁ W ₁	3 × 10 ⁷	2 × 10 ⁻³	2 × 10 ⁻³	–
A ₁ T ₁	2 × 10⁻³	–	–	2 × 10⁻³	–	A ₁ D ₁ W ₂	1 × 10 ⁹	1 × 10 ⁻²	1 × 10 ⁻²	–
A ₂ T ₁	4 × 10⁻⁶	–	–	4 × 10 ⁻¹²	4 × 10⁻⁶	A ₁ T ₁	–	2 × 10 ⁻³	–	2 × 10 ⁻³
A ₃ T ₁	1 × 10⁴	–	–	2 × 10 ⁻¹²	2 × 10 ⁻⁵	A ₁ T ₁ W ₁	6 × 10 ⁹	2 × 10 ⁻¹	–	2 × 10 ⁻¹
A ₁ T ₂	–	–	–	4 × 10⁵	4 × 10⁵	A ₁ T ₁ W ₂	9 × 10 ⁸	1 × 10 ⁰	–	1 × 10 ⁰
A ₂ T ₂	2 × 10 ⁻¹³	–	–	2 × 10⁻²	4 × 10 ⁻⁵					
A ₁ D ₁ T ₁	–	7 × 10⁵	3 × 10 ⁵	3 × 10 ⁵	7 × 10⁵					
A ₂ D ₁ T ₁	4 × 10 ⁻¹⁴	6 × 10⁻³	5 × 10 ⁻⁶	7 × 10 ⁻³	5 × 10 ⁻⁶					

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

Table 3. Numerical comparison of the RH dependencies of the observed coefficients K at the measured temperature ($K = K_{\text{obs}}$) and when K_{obs} is scaled to $T = 300$ K ($K = K_{300}$). Presented values are the correlation coefficient R , the p-value describing the probability of coincidental correlation of this strength, and the value of β in $K = \alpha \times \exp(\beta \times \text{RH})$ from linear regression fits. The compared values of K_{300} are scaled to 300 K with the temperature dependence of coefficient K_{A2T2} from the model run with organic vapours, but the correlations were similar when scaling was done with the temperature dependence of any of the coefficients K_{A2D2} , K_{A2T2} or K_{A2D1T1} .

Site/Campaign	$K = K_{\text{obs}}$			$K = K_{300}$		
	$R(\text{RH}, \log(K))$	$p(\text{RH}, \log(K))$	Fitted β in $K \sim e^{\beta \times \text{RH}/100}$	$R(\text{RH}, \log(K))$	$p(\text{RH}, \log(K))$	Fitted β in $K \sim e^{\beta \times \text{RH}/100}$
Hyytiälä 2003	-0.16	3×10^{-2}	-1.4	-0.31	2×10^{-5}	-2.7
Hyytiälä 2005	-0.24	3×10^{-3}	-3.2	-0.36	3×10^{-6}	-4.7
Hyytiälä 2007	-0.53	3×10^{-21}	-6.7	-0.74	6×10^{-48}	-12
Melpitz	0.27	1×10^{-4}	4.3	-0.04	6×10^{-1}	-0.44
Hohenpeissenberg	0.49	2×10^{-18}	5.8	0.40	9×10^{-13}	4.5
S. P. Capofiume	-0.14	3×10^{-1}	-1.2	-0.49	5×10^{-5}	-4.4
Mexico City	-0.42	1×10^{-5}	-1.7	-0.69	1×10^{-15}	-4.1
Atlanta	0.34	2×10^{-2}	6.8	0.20	2×10^{-1}	4.0
Beijing	0.03	7×10^{-1}	0.15	-0.39	5×10^{-7}	-3.3
ALL SITES	-0.16	2×10^{-10}	-2.5	-0.30	1×10^{-31}	-4.4

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)


Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

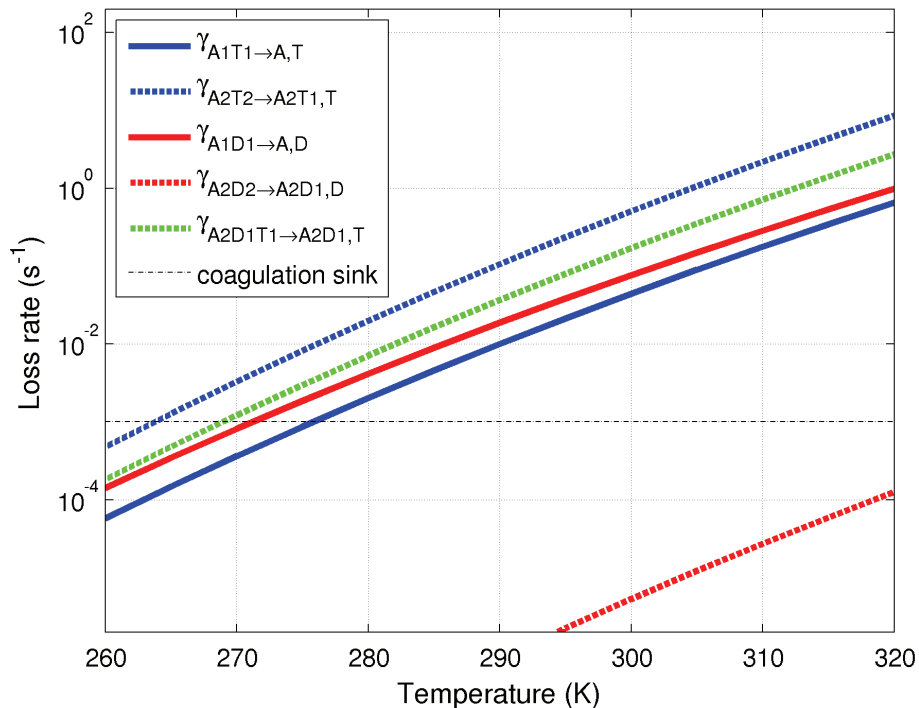


Fig. 1. Evaporation rates from A_1B_1 , A_2B_2 and A_2B_1 -clusters as a function of temperature. The presented evaporation rates are the highest rates corresponding to each cluster (see Table 2). The coagulation rate used in the model is depicted for comparison.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

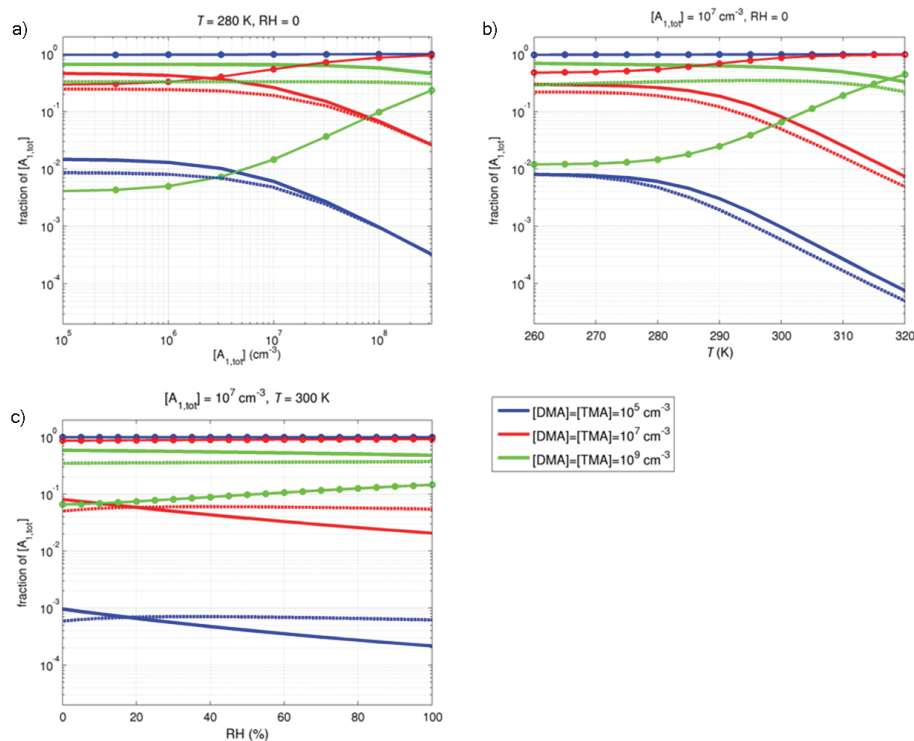


Fig. 2. Contributions of pure A (solid lines with circles), A_1T_1 -cluster (solid lines) and A_1D_1 -cluster (dashed lines) to the total H_2SO_4 monomer concentration $[A_{1,\text{tot}}]$ at varying amine concentrations (a) as a function of the total H_2SO_4 monomer concentration, (b) as a function of temperature and (c) as a function of relative humidity (RH). Temperatures, RHs and total sulphuric acid monomer concentrations in the model runs are shown in the panels.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

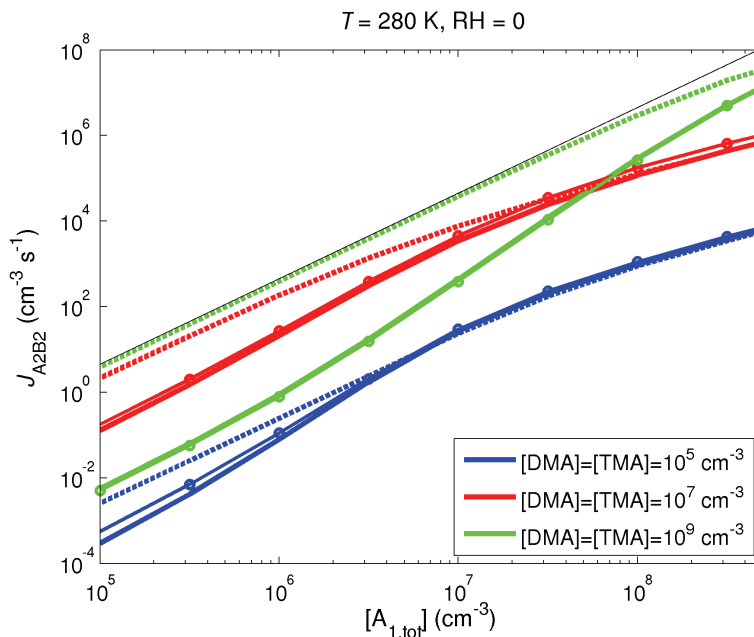


Fig. 3. The modelled formation rates J_{A2T2} (solid lines), J_{A2D2} (dashed lines) and J_{A2D1T1} (solid lines with circles) as a function of the modelled total sulphuric acid monomer concentration $[A_{1,tot}]$ at $T = 280 \text{ K}$ under varying amine concentration with a 1 : 1 [DMA] : [TMA]-ratio. The solid black line represents the collision rate $4.5 \times 10^{-10} \text{ cm}^3 \text{ s}^{-1} \times [A_{1,tot}]^2$ calculated for all the H_2SO_4 monomers.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

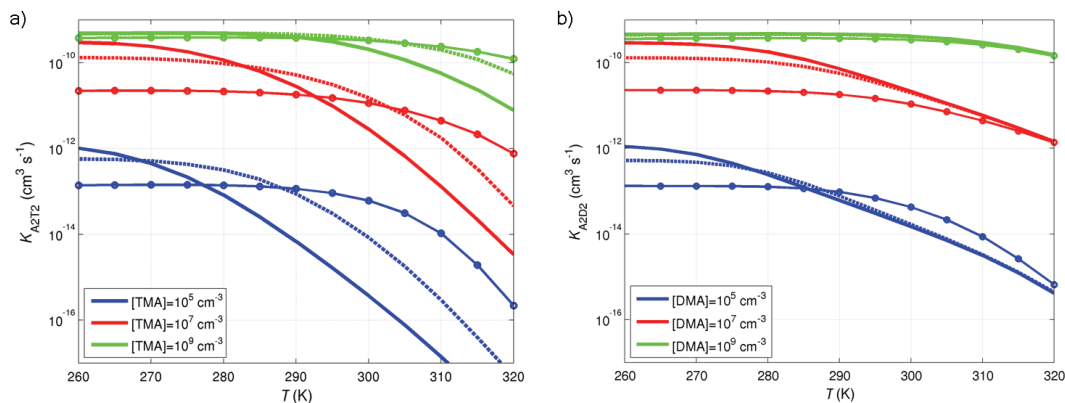


Fig. 4. The modelled coefficients K_{22} , the ratio of J_{22} and squared total sulphuric acid monomer concentration, as a function of temperature for (a) A_2T_2 -clusters and (b) A_2D_2 -clusters in cases when only one amine is present. The solid lines correspond to total H_2SO_4 monomer concentration $[A_{1,tot}] = 10^6 \text{ cm}^{-3}$, the dashed to $[A_{1,tot}] = 10^7 \text{ cm}^{-3}$ and the solid lines with circles to $[A_{1,tot}] = 10^8 \text{ cm}^{-3}$.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

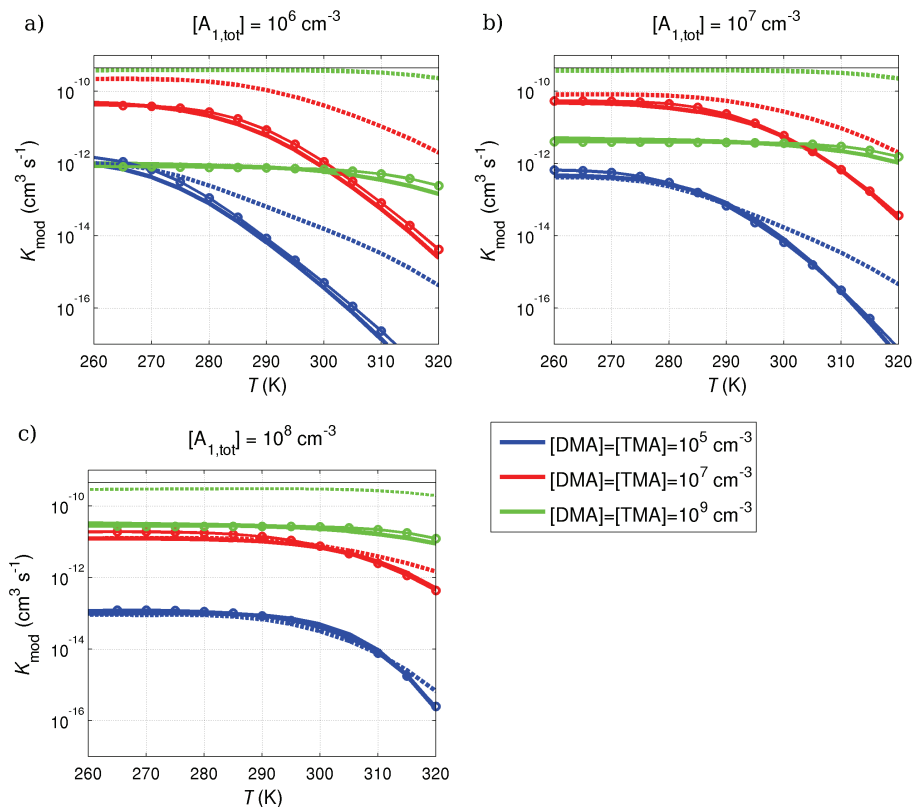


Fig. 5. Modelled coefficients K_{22} , the ratio of J_{22} and squared total sulphuric acid monomer concentration, as a function of temperature under DMA to TMA ratio 1 : 1 (concentrations shown in the legend), when the total sulphuric acid monomer concentration $[A_{1,\text{tot}}]$ equals to **(a)** $10^6 \text{ cm}^{-3} \text{ s}^{-1}$, **(b)** $10^7 \text{ cm}^{-3} \text{ s}^{-1}$ and **(c)** $10^8 \text{ cm}^{-3} \text{ s}^{-1}$. Solid lines correspond to $K_{A_2T_2}$, the dashed lines to $K_{A_2D_2}$ and the solid lines with circles to $K_{A_2D_1T_1}$. The black line shows the collision rate of all H_2SO_4 monomers.

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

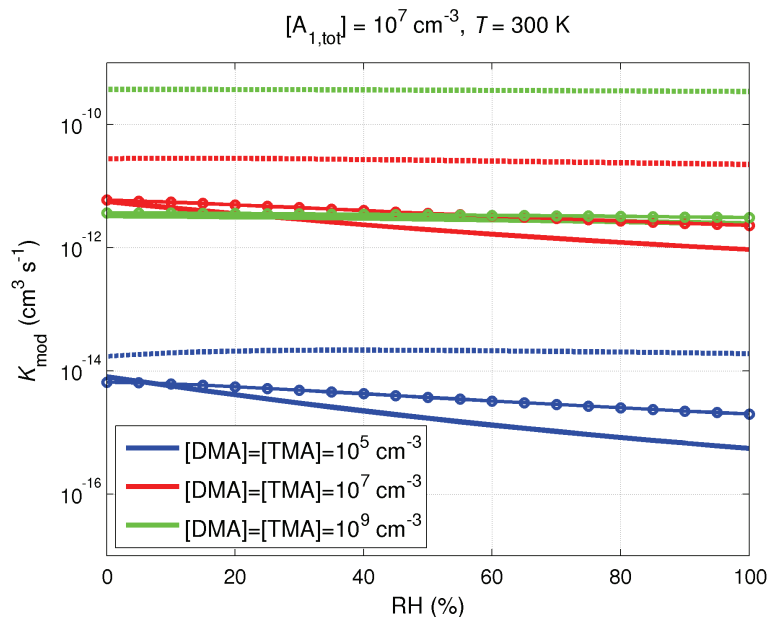


Fig. 6. The modelled coefficients K_{22} as functions of relative humidity (RH). Solid lines correspond to $K_{A_2T_2}$, the dashed lines to $K_{A_2D_2}$ and the solid lines with circles to $K_{A_2D_1T_1}$.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion



Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

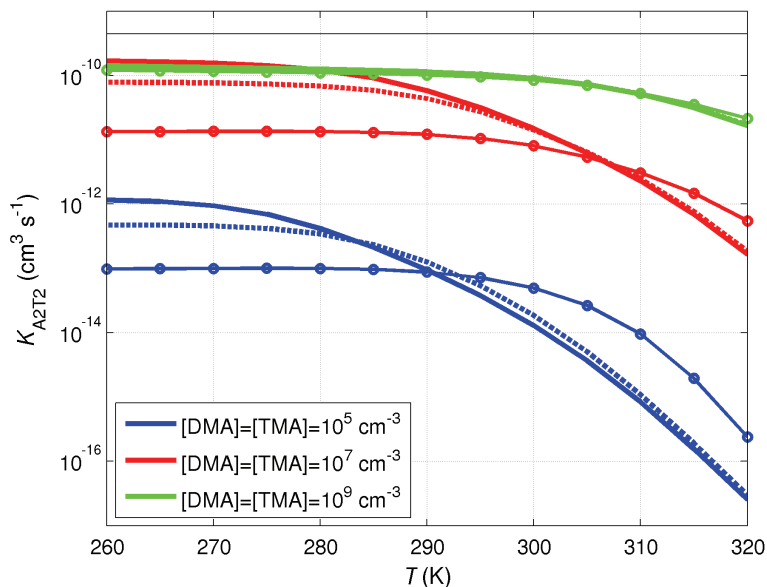


Fig. 7. The modelled coefficient $K_{A_2T_2}$ as a function of temperature assuming that the A_2T_2 and $A_2D_1T_1$ -clusters are grown out of the model when colliding with organic vapour molecules at concentration of 10^9 cm^{-3} . The solid lines correspond to total H_2SO_4 monomer concentration $[A_{1,\text{tot}}] = 10^6 \text{ cm}^{-3}$, the dashed to $[A_{1,\text{tot}}] = 10^7 \text{ cm}^{-3}$ and the solid lines with circles to $[A_{1,\text{tot}}] = 10^8 \text{ cm}^{-3}$.

[Title Page](#)
[Abstract](#)
[Introduction](#)
[Conclusions](#)
[References](#)
[Tables](#)
[Figures](#)
[◀](#)
[▶](#)
[◀](#)
[▶](#)
[Back](#)
[Close](#)
[Full Screen / Esc](#)
[Printer-friendly Version](#)
[Interactive Discussion](#)

Sulphuric acid-amine clusters in varying atmospheric conditions

P. Paasonen et al.

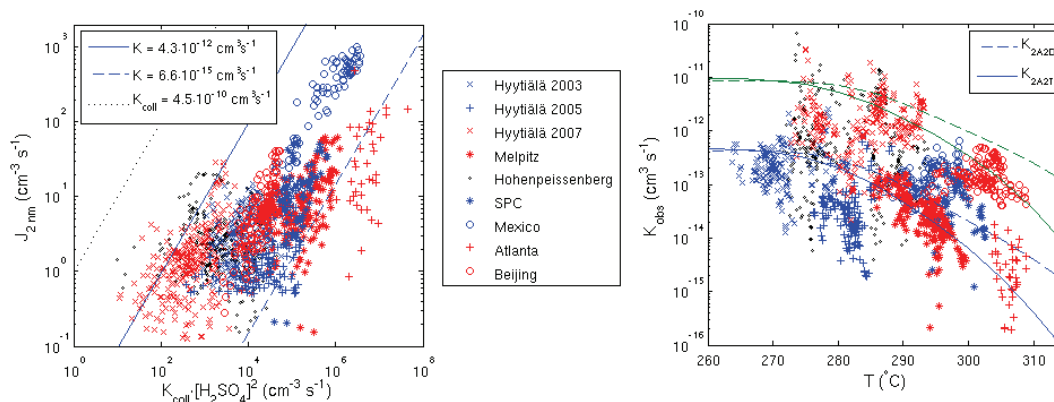


Fig. 8. (a) The measured formation rate of 2 nm particles ($J_{2\text{nm}}$) as a function of sulphuric acid monomer collision rate. The solid and dashed lines correspond to the 95th and 5th percentile of the data point specific values of $J/(K_{\text{coll}} \times [\text{H}_2\text{SO}_4]^2)$, respectively, leaving 90 % of the data points between the lines. The collision rate between sulphuric acid monomers is depicted with a dotted line. Panel **(b)** shows the observed coefficient K_{obs} as a function of temperature. The solid and dashed lines correspond to the values of the modelled coefficients K_{A2D2} and K_{A2T2} , respectively, at amine concentrations of 10^5 cm^{-3} (blue lines) and 10^6 cm^{-3} (green lines), $\text{RH} = 50\%$ and $[\text{A}_{1,\text{tot}}] = 10^7 \text{ cm}^{-3}$.

Title Page

Abstract

Introduction

Conclusions

References

Tables

Figures

◀

▶

◀

▶

Back

Close

Full Screen / Esc

Printer-friendly Version

Interactive Discussion