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Size-resolved aerosol emission factors and new particle formation/growth activity occurring in Mexico City during the MILAGRO 2006 Campaign

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Abstract

Measurements of the aerosol size distribution from 11 nm to 2.5 microns were made in Mexico City in March, 2006, during the MILAGRO field campaign. Observations at the T0 research site could often be characterized by morning conditions with high particle mass concentrations, low mixing heights, and highly correlated particle number and CO₂ concentrations, indicative that particle number is controlled by primary emissions. Average size-resolved and total number- and volume-based emission factors for combustion sources impacting T0 have been determined using a comparison of peak sizes in number and CO₂ concentration peaks. The number emission and volume emission factors for particles from 11 nm to 494 nm are 1.23×10^{15} particles, and 7.54×10^{11} cubic microns per kg of carbon, respectively. Uncertainty on the number emission factor is approximately a factor of 1.5. The mode of the number emission factor was between 25 and 32 nm, while the mode of the volume factor was between 0.25 and 0.32 microns. These emission factors are reported as log normal model parameters and are compared with multiple emission factors from the literature. In Mexico City in the afternoon, the CO₂ concentration drops during ventilation of the polluted layer, and the coupling between CO₂ and particle number breaks down, especially during new particle formation events when particle number is no longer controlled by primary emissions. Using measurements of particle number and CO₂ taken aboard the NASA DC-8, this emission factor was applied to the MCMA plume; the primary emission factor predicts less than 50% of the total particle number and the surplus particle count is not correlated with photochemical age. Primary particle volume and number in the size range 0.1–2 μm are similarly too low to predict the observed volume distribution. Contrary to the case for number, the apparent secondary volume increases with photochemical age. The size distribution of the apparent increase, with a mode at ~250 nm, is reported.

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1 Introduction

1.1 Motivation

Numerous studies have shown the adverse effects of particulate matter (PM) on human health, with increased interest placed on ultrafine particles which become more toxic per unit mass with decreasing size (Mills et al., 2009; Oberdorster et al., 2005; Osornio-Vargas et al., 2003; Delfino et al., 2005). Atmospheric aerosols are known to affect the climate system by altering cloud properties, often referred to as the indirect effect (Lohmann and Feichter, 2005). Currently, the indirect effect of aerosols on climate is highly uncertain and limits our understanding of climate sensitivity to anthropogenic perturbations (IPCC, 2007). Uncertainty of the indirect effect is in part attributed to the uncertainty in the number of cloud condensation nuclei (CCN), or the particles on which cloud droplets form. Whether a particle will act as a CCN or not depends on the particle size, composition, and water-vapor supersaturation. Particles are introduced into the atmosphere either by direct emissions (primary particles) or homogeneous nucleation of low volatility species (secondary particles). Primary particles are introduced into the atmosphere at generally larger sizes (10 nm or greater) when compared to secondary particles (Pierce and Adams, 2009). Some particles are emitted at CCN active sizes, while others must grow in size through coagulation and condensation. Nucleated particles begin as molecular clusters, and require substantial growth to become large enough to act as CCN. The dynamic evolution of the particle size distribution involves competition between coagulation, condensation, and nucleation. Primary and secondary particles vie for growth through condensation of low volatility gas phase species, which can also nucleate to form new particles (Wang and Penner, 2009). Therefore, any uncertainty in the size distributions or rates of primary emissions or of secondary particle formation may lead to large uncertainty in the predicted CCN concentration (Adams and Seinfeld, 2002).

Although accurate knowledge of the size distribution and number concentration of atmospheric particles has been determined to be critical for prediction of CCN

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concentration, global models representing aerosol number concentration assume the number and size of particles from mass emissions (Chang et al., 2009; Pierce and Adams, 2007; Yu et al., 2010), due to the fact that inventories of anthropogenic emissions are based on mass rather than number concentration. While size-resolved emission factors are increasingly available, they are often determined from source-based tests or from vehicle tunnel studies. Here, we recover a size resolved number-based emission factor representative of Mexico City urban emissions. Both the emission factor itself, and the method of its recovery (particularly if refined using fast number size distribution measurements), may be valuable for continued refinement of size resolved emission factors.

1.2 Background

Mexico City is the largest megacity in North America, with a population of over 20 million people. Nearly 40 million liters of fuel are consumed each day, producing thousands of tons of pollutants (Molina et al., 2008). The city is located in the basin of the central Mexican plateau (19.5° N) at an altitude of approximately 2200 m a.s.l., with mountains to the south, west, and east. The topography and meteorology of this area reduce ventilation of pollutants, especially in the morning before the boundary layer has increased to levels where coupling with the gradient regional wind occurs, thereby contributing to the persistent air quality problems in this area (Jauregui, 1988; de Foy et al., 2006).

The air pollution in Mexico City has been the topic of numerous studies. Raga and colleagues have provided an observational summary of studies conducted from 1960–2000, and concluded that the limited measurements of aerosols and their transport from major source areas was highly insufficient for understanding the evolution and environmental impacts of aerosols. This lack of information about the physical characteristics of aerosols motivated a study in 1997 by Baumgardner et al. (2000). Their measurements, taken from a mountain site in the southwest portion of the basin, concluded that a large fraction of the measured aerosols were from primary emissions

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based on a positive correlation with CO. High nighttime aerosol number concentrations were suggestive of recirculation within the basin that would further impact pollutant levels. In the spring of 2003, a large scale field campaign was conducted (MCMA-2003) which contributed to the understanding and improvement of air quality in Mexico City.

5 Dunn and colleagues (2004) conducted measurements of the aerosol size distribution from 3–48 nm in two locations in the Mexico City area; a rural site in the mountain pass in the southeast corner and another in the city's center. New particle formation events were observed on 3 of 10 days sampled. At the urban site, high concentrations of
10 15–25 nm particles correlated with high levels of NO_x and CO and sudden decreases in condensational surface area preceded the new particle formation events (Dunn et al., 2004). Related measurements showed the significance of secondary organic and inorganic aerosol production and its contribution to particulate matter concentrations (Volkamer et al., 2006). Salcedo et al. (2006) reported that the organic mass fraction of PM_{2.5} as measured at the CENICA site in southeast Mexico City is on average 56%,
15 with inorganics representing 28% (Salcedo et al., 2006). The size distribution was dominated by an internally mixed accumulation mode, and a smaller externally-mixed mode that was concluded to be related to traffic emissions (Molina et al., 2007).

Similar results were obtained during the MILAGRO field campaign in 2006. The size distribution and composition of particles from 10 to 33 nm was measured northwest
20 of the city in Tecamac by Smith and colleagues (Iida et al., 2008; Smith et al., 2008). Results from a case study of one new particle formation event suggest that the freshly nucleated particle are comprised of more organics than sulfates and that organic compounds play a dominant role in the growth of particles at this location (Smith et al., 2008). Using data collected onboard the NSF/NCAR C-130 with HR-ToF-AMS, De-
25 Carlo and colleagues determined that organic species dominated submicron aerosol near the pollution source and in the outflow and regional air (DeCarlo et al., 2008). Ground based measurements performed by Aiken et al. using HR-ToF-AMS also concluded that organic aerosol comprises half of the mass of fine particulate matter (Aiken et al., 2009). Source apportionment of fine organic aerosol determined that primary

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emissions from motor vehicles consistently account for half of the organic carbon in the urban area (Aiken et al., 2009).

The evolution of the size distribution and composition of aerosols over the Mexico City basin was measured by Kleinman and colleagues (Kleinman et al., 2008, 2009).

5 Measurements taken aboard the DOE aircraft sampled aerosols with photochemical ages ranging from fresh emissions to day old air masses. Findings included that a 5-fold increase in aerosol volume per CO as a result of 0.5–1 day of processing is due to more accumulation mode particles in aged plumes, rather than larger particles and that the increase in accumulation mode particles was due to condensational growth
10 from Aitken mode particles, rather than volume growth (Kleinman et al., 2009).

The air quality in the basin during the 2006 campaign was compared to 10 years of data collected by Mexico City's automatic monitoring network (RAMA), and it was concluded that PM_{2.5} and PM₁₀ levels were mostly within their usual range (deFoy et al., 2008). PM₁₀ maximum loadings occurred at either the end of the morning or
15 the end of the afternoon and PM_{2.5} concentrations peaked around noon. The average PM_{2.5} and PM₁₀ levels measured at T0 were 40 µg m⁻³ and 56 µg m⁻³ respectively, and were markedly impacted by traffic emissions during rush hours (Querol et al., 2008). Measurements of PM at multiple locations in parallel suggest that variability in PM
20 levels and composition may be determined largely by atmospheric mixing and mixed layer height rather than by emission sources (Querol et al., 2008).

The quantification of particulate emissions from combustion sources is important for multiple reasons including generation of emission inventories for the development of air management and control strategies, examination of specific sources on compliance with standards, and assessment of personal and environmental exposures (Zhang and
25 Morawska, 2002). Because many of the health and climate effects associated with atmospheric particles are influenced by size, particle emission factors should also be size resolved. Emission factors and inventories have been compiled for Mexico City; however no reports of a size resolved particle number-based emission factor exist to date.

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Size-resolved emission factors can be measured in various ways. Dynamometer tests are often performed in controlled settings (Ristovski et al., 1998), but may not be representative of fleet average emissions under real world dilution and engine load conditions. On-road techniques have the advantage of sampling a wide range of vehicles with varying loads and speeds under real-world dilution conditions and can be measured from a mobile laboratory or at a stationary location (Zavala et al., 2009; Kolb et al., 2004; Jiang et al., 2005). On-road emission factors are commonly determined using a stationary control volume, such as a roadway tunnel where emission factors can be calculated using a mass-balance approach (Jamriska and Morawska, 2001; Geller et al., 2005). This method has the advantage of providing fleet averaged emission factors over a large number of vehicles; however difficulties arise when sampling smaller particles. The reduced dilution of a control volume may change evolution of the size distribution and influence the applicability of the recovered size distribution (Kristensson et al., 2004).

An effective method for determining a fuel-based, size-resolved, submicron particle emission factor uses the relationship between particle emissions to emission of a co-emitted gaseous pollutant (Janhall and Hallquist, 2005; Jiang et al., 2005). Carbon dioxide is the primary carbon-containing product of fuel combustion and can provide an estimate of the amount of fuel burned (McGaughey et al., 2004). Relating features (i.e. peaks) in the time series of a pollutant concentration (e.g. particle number) to corresponding peaks of CO₂ and CO, allows for the emission ratio to be expressed in units of particles per mole of the gaseous compound or related to the carbon content of fuel and expressed as amount of pollutant emitted per unit of fuel consumed. Alternatively, if the fleet average fuel economy is known or estimated, emission factors can be expressed per vehicle km. This method of determining fuel-based particle emission factors has previously been used in various locations (Kirchstetter et al., 1999; Westerdahl et al., 2009; Kittelson et al., 2004; Ning et al., 2008). In this study, the observed correlation between particle number and CO₂ concentrations was exploited in order to determine an average size resolved number based emission factor for Mexico

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City. Rather than a regression of total number or a size fractionated number upon CO₂ concentrations, a method was used to identify peaks of CO₂, and then to determine their corresponding paired peak in particle number. The baseline (against which the peaks were determined) was calculated using a successive moving average method as described by Watson and Chow (2001). This method should be robust even in the presence of slowly varying baseline concentrations of CO₂ or particle number. Key questions for this data analysis were the robustness of the method to new particle formation events, the uncertainty in the recovered emission factor, how the emission factor would compare with other vehicular emission factor studies, whether sources could be identified according to wind direction or source class, and how the recovered emission factor would compare with gas and size-resolved aerosol measurements in the megacity plume.

The technique as employed does not recover source-specific emission factors (e.g. on-road spark ignition, on-road diesel, coal combustion, etc.). However, it recovers the average emission factor over multiple combustion sources.

2 Instrumentation

Observations were made by the University of Iowa and by the Department of Energy Brookhaven National Laboratory in Mexico City during the MILAGRO field campaign in March, 2006, in order to characterize the T0 urban size distribution, to study the primary emissions around T0, and to characterize new particle formation and ultrafine particle growth. From 7–29 March, measurements were taken at the T0 research location, one of the three supersites selected for the campaign. The T0 location was located inside the Instituto Mexicano del Petróleo (IMP) in the northwestern part of the basin, which is an urban background site influenced by fresh roadway traffic emissions, residential emissions and at times local industrial emissions. Instrumentation used in this work is listed in Table 1.

Two datasets of particle size distribution are used in this work. The University of Iowa operated a Scanning Mobility Particle Sizer (TSI 3081) and an Aerodynamic Particle Sizer (TSI 3321), with relative humidity controlled inlets (Stanier et al., 2004), as shown in Fig. 1. The SMPS measured the size distribution from 10.9 nm to 478 nm at 5 min intervals, while the APS measured the distribution from 542 μm to 19.8 μm at 20 second intervals. Jian Wang of Brookhaven National Laboratory operated an SMPS system which measured the size distribution from 15 to 494 nm at 2 min intervals (Wang et al., 2003). The Iowa SMPS and APS switched each 5 min between a dried (typically <20% RH) configuration and a not dried configuration. The not dried configuration was intended to sample aerosols at ambient RH, however due to the increased temperature of the building housing the instruments relative to ambient conditions, the sample was dried before reaching the detector. The average ambient relative humidity was 43.3% ($\pm 23.2\%$) and the average relative humidity in the “not dried” sample line was 19.1% ($\pm 6.9\%$). In this work, the “not dried” channel is used for both the SMPS and APS analysis. The relative humidity (RH) of the BNL SMPS sample flow was always below 30% during MILAGRO, and was below 25% for a vast majority of the size distribution measurement, suggesting sampled aerosol particles were effectively dry. The BNL SMPS was calibrated using polystyrene latex standards. Data from the BNL SMPS were reduced using the data inversion procedure described by Collins et al. (2002).

CO₂ and meteorology measurements were made using a Licor LI-7500 CO₂ analyzer, which measured ambient CO₂ concentrations from an open field in the middle of the IMP, approximately 150 m away. The location of the CO₂ monitor with respect to the particle sizing equipment is shown in Fig. 2.

For evaluation of the particle number and particle volume time series, comparisons to nephelometer measurements of light scattering at 700 nm (Marley et al., 2009) was used. Comparisons to hourly PM_{2.5} readings from the Mexico City ambient air monitoring network were also used (RAMA, Red Automatica de Monitoreo Atmosferico, http://www.sma.df.gob.mx/simat/home_base.php).

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The measurement of wind speed and wind direction used in this analysis was obtained by the Department of Energy (N. Marley) at the T0 location using a Vaisala WXT 510 weather transmitter.

In order to apply the emission factor to the MCMA plume downwind of the city, measurements of aerosol size distribution and CO₂ obtained via aircraft were required. The measurements of the aerosol size distribution used in this work were obtained onboard the NASA DC-8 aircraft by the University of Hawaii group HiGEAR using a modified long differential mobility analyzer (10 nm–400 nm) (Roberts et al., 2010). Sampling was conducted through an NCAR solid diffuser inlet under dry instrument conditions (less than 30% RH). A modified LI-COR (6252) non-dispersive infrared gas analyzer, was used to determine the CO₂ mixing ratios and was operated under constant pressure (Vay et al., 2003). Measurements of benzene and toluene, used here to approximate plume age, were obtained using whole air samples collected by D. Blake (UC Irvine). Samples were collected in stainless steel canisters and analyzed at the UC Irvine laboratory within 7 days of collection (Baker et al., 2008).

3 Data analysis techniques

The Iowa SMPS number size distribution data used in this work were corrected for inlet losses, which occurred due to the tubing and valves required for the alternating drying of the sample (e.g. 32% transmission was estimated at 15 nm, and 93% transmission at 102 nm). The BNL SMPS was designed to measure the size distribution and concentration of aerosols at T0 without additional processing and therefore had a shorter inlet, less upstream valves and fittings, and a higher efficiency. The BNL inlet efficiency was calculated using the flow rates and tubing lengths of the sampling line as described in Wang et al. (2002) and applied to correct the SMPS data set. Then the transmission efficiencies for the Iowa SMPS were calculated to maximize hourly average agreement with the BNL SMPS. These are similar to transmission efficiencies measured using a reconstructed inlet after the study (Kalafut-Pettibone, 2009).

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To improve comparability between the APS and SMPS size distribution, the APS data (originally calibrated for aerodynamic diameter) was shifted to the smaller equivalent electrical mobility diameters using the approach of Khlystov et al. (2004). This requires an assumed density (1.43 g cm^{-3} was used based on the Mexico City density and composition as determined by DeCarlo et al. (2008)) and an assumption of spherical particles. APS transmission efficiencies were measured using a reconstructed inlet in the laboratory (Kalafut-Pettibone, 2009), and the inlet efficiency of the bypass (not-dried) inlet was 92% at 0.54 microns and 65% at 1.5 microns. However, using these efficiencies to correct the field data led to a poor overlap between the largest channel of the SMPS and the smallest channel of the APS, and it led to a low $\text{PM}_{2.5}$ mass estimated from the combined SMPS, volume, and aerosol density, when compared to nearby $\text{PM}_{2.5}$ monitoring sites operated by the RAMA network. Therefore an alternate correction was used as described below. The multiplier required for best agreement between the lower end of the APS size range (542 nm) and the upper end of the SMPS size distribution (which extended to 496 nm) was determined. The multiplier calculated in this manner was 2.86. This multiplier was assumed for all sizes. The reason for the difference in measured transmission and the required correction factor in the field (equivalent to a 35% transmission efficiency) is not understood. The relative transmission efficiency between the dried and bypass inlets was also calculated; however, only the results from the bypass (not dried) inlet is used in this work.

The decision to jointly analyze the CO_2 and number time series was based on an initial inspection of the dataset, where a Pearson correlation coefficient of 0.6 was noted between fine particles and CO_2 ; however, periods with much higher correlation existed. The early morning hours had the highest correlation, with 4 h correlations from 04:00–08:00 a.m. reaching 0.93 on some days. Figure 4 illustrates the correlation coefficient between particle number and CO_2 as a function of the time of day and also includes a scatterplot of number vs. CO_2 for the entire measurement period. Although the biosphere respiration contribution to the CO_2 concentration would also be at a maximum during the early morning, the biogenic sources of CO_2 were determined to

be insignificant (emissions were dominated by anthropogenic activity) according to flux measurements of CO₂ obtained during the campaign (Velasco et al., 2009).

The calculation of the number-based emission factor is based on the comparison of the size of the peaks in particle concentration to peaks in CO₂ concentration. The technique is an adaptation of previous emission factor calculations by Jiang et al. (2005) and utilized a simple peak identification and quantification scheme used by Watson and Chow (2001). In Jiang et al. (2005), a mobile laboratory with fast (1-s) CO₂ data, used the 5th lowest 1 second reading in a 3 min (180 s) window as the on-road background level. CO₂ peaks were then calculated by subtracting this background. Similar background subtraction was performed for all compounds of interest. CO₂ datapoints 42 ppm above background were included in a database of “peaks” and the ratios of pollutants (above their background concentrations) to the CO₂ peaks was compiled, and then averaged to determine the mean emission factor for each species in question.

As opposed to Jiang et al. (2005) where 1–10 s data were available for all species, and all measurements were collocated, in this study there were relatively slow measurements (e.g. 2 min for the BNL SMPS, and 5 min for the Iowa SMPS), and there was a separation distance of 150 m between the CO₂ measurement and the particle measurement. This required the adoption of an analysis using longer temporal averaging.

For isolation of number and CO₂ peaks to compare in this study, peaks were isolated relative to a baseline concentration calculated by the successive moving average technique of Watson and Chow (2001). Illustrated for a 24 h period in Fig. 3, the result of this technique is a 10 min time series of baseline and peak concentrations. Figure 3a shows 1-min CO₂ and 2-min number concentrations, the 10-min averages of both signals, and the baseline concentrations of each. Figure 3b shows the peaks above the baseline concentration for both signals (which will be denoted as ΔCO₂ and ΔNum) which are used to calculate the emission factor. The baseline and peak values were constructed as follows. Concentrations were averaged to a 10 min time basis. This put all the measurements on a common time basis regardless of their native time

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5 resolution, which ranged from 0.1 s (CO₂ concentration) to 5 min (University of Iowa SMPS). 6-h average concentrations were compared to the 10 min averaged data, and the lower of each pair was retained in a tentative baseline time series. Then the tentative baseline was smoothed using a 3-h averaging time, and the resulting values were
10 compared to the original 10 min time series, with retention of the lower values in the tentative baseline. A third round of averaging (1 h averaging time) and comparison was done to yield a final baseline time series which represents the regional background concentration plus the urban background. Watson and Chow (2001) approximate influence distances for short duration peaks using Pasquill Gifford curves (plot of distance from source versus plume dimension). The estimated horizontal plume dimensions are 70–1000 m at 2 km away from a continuous source. Plumes of these dimensions could cause short duration (<10 min) responses in the samplers. In other words, the baseline represents a regional background plus a well mixed urban background originating from distances likely farther than 0.2–2 km away. Peaks are caused by smaller plumes that
15 can traverse the sampling site in <10 min, which are likely from sources closer than 2 km.

The 50th percentile of the CO₂ peak population was 1.6 mg m⁻³ and the 95th percentile peak size was 25 mg m⁻³. To calculate emission factors, all CO₂ peaks above a threshold were selected, and the number/CO₂ ratio of each of these 10 min periods is calculated. A threshold of 6 mg m⁻³ of CO₂ was used and sensitivity to this threshold is discussed. The technique is applied to total number and also to the number concentration in specific diameter bins to give a size-resolved emission factor.
20

Emission factors determined at T0 are based on the ratio of the change in particle concentration to the change in mass concentration of CO₂, e.g. particle number cm⁻³ (mg m⁻³)⁻¹. Both concentrations are on an actual volume basis (not corrected to a standard pressure and temperature or for water vapor). This can be easily converted to particle number (kg C)⁻¹, which is invariant to changes in pressure, temperature, and water mixing ratio. Conversion to other ratio based metrics, such as particle number cm⁻³ (μmole CO₂ per mol)⁻¹ require specification of the relevant temperature and
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pressure, and whether the CO₂ concentration is on a dry or wet basis. The equation used to use measured emission factors at various pressure, temperature, and water vapor combinations is:

$$C_{n,\text{predict}} \Big|_{T,P,P_{\text{H}_2\text{O}}} = \left[C_{n,\text{measure}} / C_{\text{CO}_2,\text{measure}} \right] 1.807 \frac{(P - P_{\text{H}_2\text{O}})}{1\text{bar}} \frac{293\text{K}}{T} \left[\chi_{\text{CO}_2} \right]$$

where $C_{n,\text{predict}}$ is the predicted concentration in particle number cm⁻³ at the specified temperature (T in K), total pressure (P in bar), and water partial pressure ($P_{\text{H}_2\text{O}}$ in bar). $C_{n,\text{measure}} / C_{\text{CO}_2,\text{measure}}$ is the original measured emission factor (in the current work in particle number cm⁻³ and mg m⁻³, respectively, at local conditions). χ_{CO_2} is the CO₂ mixing ratio in $\mu\text{mole per mole dry air}$. 1.807 is a conversion factor specific to the selected units.

4 Results and discussion

4.1 Summary statistics of dataset

For the common period (shared among the ground-based instruments) of 10–25 March, the mean values and distributions of the key experimental variables are given in Table 2. The mean number concentration (as measured by SMPS) in the 15–494 nm size range was $2.1 \times 10^4 \text{ cm}^{-3}$. The mean CO₂ concentration measured at T0 was 559 mg m^{-3} ($392 \mu\text{mole mol}^{-1}$).

The time series of several of the key measurements in this work are shown in Fig. 5. Because of the uncertainty in the APS inlet transmission, several comparisons were made between metrics sensitive to aerosol volume and mass. Available at T0 was a 3-wavelength nephelometer, and the nearby (within 10 km) hourly PM_{2.5} monitoring sites, operated by the RAMA network. Figure 5a represents the time series of number and CO₂ concentration for the period from 10–25 March and illustrates the correlation

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between the two signals. Figure 5b provides a time series of particulate matter (PM) volume and aerosol scattering at 700 nm as measured by the nephelometer for the same time period. PM volume was then converted to PM mass using the assumed density of 1.43 g cm^{-3} as previously discussed and for comparison with local $\text{PM}_{2.5}$ measurements (Fig. 5c). The correlation coefficient (r) of PM mass and aerosol scattering (nephelometer, 550 nm) was determined to be 0.84 (not shown). The slope of the regression line, which represents the scattering per unit mass, is $4.1 \text{ m}^2 \text{ g}^{-1}$ which is in the range of expected values for urban pollution (Bates et al., 2006). Diurnal averages are provided in Fig. 5d and e. Aerosol number and CO_2 reached peak concentrations at 07:00–08:00 a.m. (all times Central Standard Time, which is UTC-6 h). CO_2 and particle number drop together due to boundary layer ventilation from 09:00–11:00 a.m., and then CO_2 continues to drop to a minimum at 04:00 p.m., while particle number (on average) increases during the afternoon, due to higher numbers at sizes below $\sim 30 \text{ nm}$. Aerosol volume peaks from 08:00–10:00 a.m., while extinction peaks at 10:00 a.m. as well. The wind rose representing wind speed and direction for the period from 10–25 March (Fig. 5f) shows winds predominantly from the west and east, with the high wind velocity periods from the east south east.

In order to further evaluate the APS data, a comparison with the 700 nm nephelometer data was conducted. The total aerosol volume as determined by the SMPS and APS was considered together with light scattering at 700 nm, for all days with over 6 h of data (11 days total). Hours during new particle formation were excluded. The average correlation (R^2) for the 11 days examined is 0.87. The same calculation was then repeated using only SMPS volume, rather than SMPS+APS volume and the correlation (R^2) decreased to 0.74.

4.2 Ultrafine particle growth events

Conditions at T0 could often be characterized by highly correlated particle number and carbon dioxide concentrations, indicative that particle number is controlled by primary emissions. During new particle formation events, particle number is no longer

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controlled by primary emissions and the correlation between particle number and CO₂ breaks down. This can be observed in the 17 March time series of number concentration and CO₂ shown in Fig. 3b, where peaks above baseline of CO₂ and number (ΔCO_2 and ΔNum) are correlated until the onset of the event shortly after 12:00 CST at which point number concentration increases and CO₂ levels decrease.

Ultrafine particle growth events, which were qualitatively identified by prominent and growing modes in the size range 10–15 nm, were observed on 6 of 16 (37.5%) days sampled at the T0 location. Events typically began between 10:30 and 13:00 local time. The growth rates of the ultrafine particles varied between 4.9 and 17.7 nm h⁻¹, which are somewhat higher than the average reported range for other urban areas (0.5–9 nm h⁻¹) (Kulmala et al., 2004). Table 3 provides a summary of the events and Fig. 6 plots the size-distribution from 15–500 nm for the 6 days where new particle formation occurred.

One commonality among all particle formation and growth events is lower than average particle number concentration just prior to the onset of the event. Favorable conditions for new particle formation often include low pre-existing aerosol concentrations, which may act as a sink for condensable vapors (Kulmala and Kerminen, 2008), and therefore this result is consistent with previous observations of new particle formation in Mexico City (Dunn et al., 2004) and various other locations (Clarke, 1993; Weber et al., 1997; Kulmala et al., 2004). It is known that the Mexico City basin experiences substantial ventilation during the afternoon (de Foy et al., 2006, 2009; Banta, 1985) which likely provides an explanation for rapid decreases in condensational surface area. Five of six of the particle formation/growth events sampled in this study have decreases in CO₂ and aerosol scattering leading to an increase in number concentration, consistent with previous observations.

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4.3 Recovered emission factor

The average number and volume emission factors as determined for Mexico City using the described method are 1.23×10^{15} particles (from 11 nm to 494 nm), and 7.54×10^{11} cubic microns (from 15 nm to 1800 nm) per kg of carbon, respectively. The mode of the number emission factor is between 25 and 32 nm, while the mode of the volume factor is between 0.25 and 0.32 microns as shown in Fig. 7.

An approximate estimate of the total uncertainty in the emission factor is done and includes contributions from both method and instrument uncertainty, and sampling error. A third source of error, representativeness error (e.g. how representative of the entire airshed is the recovered emission factor) is very difficult to quantify and is not considered in this work. However, its importance should not be minimized. The random uncertainty from the limited number of samples taken is the easiest to quantify, and we have taken advantage of the relatively large number of individual emission factors (~600) to calculate a 95% confidence interval on the mean. This contributes approximately $\pm 10\%$ to the grand average emission factor.

The potential for error in determination of number by the SMPS is ~15% due primarily to uncertain flowrates, charging efficiency, and inlet efficiencies (Khlystov et al., 2004). The uncertainty in the IOWA SMPS is likely higher due to the more severe inlet losses described in the experimental section. However, the BNL size distribution is used preferentially over the IOWA SMPS size distribution where available. The potential uncertainty due to the threshold selection in the calculation of the emission factor is thought to be within 20% (see below).

Combining these three sources of potential error, we find that the overall emission factor has potential error of approximately a factor of 1.5. The uncertainty in the APS size range and in the volume emission factor are larger. Also, the uncertainty on emission factors for shorter periods (e.g. daily or sub-daily periods) is higher.

As stated in the experimental section, the CO₂ measurement location and the number measurement location were separated by 150 m, and true collocation would likely

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yield a significant improvement in the power of this technique, possibly improving the error characteristics and definitely allowing use of shorter averaging times, if desirable.

The emission factor was fit to establish the corresponding lognormal mode parameters using constrained nonlinear optimization of an objective function that weighted the normalized number distribution of the emission factor by 0.8 and the normalized volume distribution of the emission factor by 0.2. A constraint was put on the width of the modes such that $\log \sigma$ was required to be 0.15 or higher; this avoids overly narrow peaks. A single mode fit achieved an R^2 of 0.94 with respect to the number distribution and an R^2 of 0.85 with respect to the volume distribution. This fit had parameters of N 1.65×10^{15} particles per kg C, D_p of 30 nm, and $\log \sigma$ of 0.41 (shown in Figs. 7 and 9). Significant improvement in the fit could be achieved but only by moving to 3 and 4 mode fits. The parameters of these fits are shown in Table 4. The single mode fit parameters were very insensitive to the relative weighting of number and volume. Fitting based on volume only gave single mode (number size distribution) parameters of D_p of 30 nm and $\log \sigma$ of 0.39.

Sensitivity of the emission factors to the CO_2 threshold used in the peak identification process was performed. Higher emission factors resulted from the use of lower thresholds. 6 mg m^{-3} (the 72nd percentile of the CO_2 peaks) was selected as the best threshold because it represented a good compromise between the tendency to have more emission factor values for small threshold values and the tendency of the variability in emission factors to increase substantially with small peaks. Using the threshold of 25 mg m^{-3} , the fewest number of peaks were isolated for analysis. For example, the BNL instrument time series yields 110 peaks during the 16 days analyzed when a threshold of 25 mg m^{-3} is used; this increases to 505, 691, 731, and 1044 peaks using the 8, 6, 4, and 2 mg m^{-3} thresholds, respectively. The confidence interval of the mean number emission factor was smallest at 6 mg m^{-3} relative to those using small thresholds (more values to average, but with a higher variability), and higher thresholds (similar variability but fewer values informing the average). Jiang et al. (2005) considered changes to the threshold of $\pm 24\%$ and showed that the CO emission factor

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only changed by 1–4% for this amount of change in the threshold. A similar perturbation to our threshold (e.g. 4.56, 6, and 7.44 mg m⁻³) shows a larger variability in the EF (14% increase at 4.56 and 5% decrease at 7.44). Based on comparison of mean emission factors and their confidence intervals at thresholds of 4, 6, 8 and 25 mg m⁻³, we estimate threshold selection as causing uncertainty of approximately 20%. Using 25 mg m⁻³, the recovered emission factor decreased by 47% relative to the value at 6 mg m⁻³, while the confidence interval widened. It is unclear whether the observed decrease is due to differences in sources, differences in plume processing, or to correlated errors between ΔNum and ΔCO_2 . However, the number of these plumes (with ~ 25 mg m⁻³ above baseline CO_2) is small and basing the overall emission factor on a larger sample is advisable. No time periods were excluded in the recovery of the emission factors. The spectral shape of the recovered size-resolved emission factor is insensitive to the threshold.

4.4 Dependence of emission factors on time of day, day of study, and wind direction

The average number emission factor as a function of time of day, and day of the month were examined in order to determine the impact that meteorological changes, biomass burning, or new particle formation may have on the overall emission factor.

A regional scale overview of the meteorological conditions during the MILAGRO campaign is provided by Fast et al. (2007). Based on the meteorology, the campaign was split into three regimes. The first regime occurred prior to 14 March and was characterized by dry, sunny conditions. The second regime began with a “Norte” or cold-surge event, which is characterized by northerly near-surface winds and increased humidity on 14 March with a gradual drying of the atmosphere over the next few days. A second norte event, lasting only a few hours, passed through on 22 March bringing light rain. The third regime began with the most significant cold-surge on 23 March. Substantial precipitation was seen on both the 24th and 25th. The impact of biomass burning organic aerosol at T0 was characterized by Aiken and coauthors (2010). Two

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high fire periods are identified from March 11–15 and March 17.5–23.5; a low fire period coincident with increased precipitation is identified from March 24–29. Details of the ultrafine particle formation and growth events observed at T0 are detailed in Table 3.

The average emission factor for each day of the study was calculated. Day to day variation showed no statistical significance with the exception of 2 days (14 March and 24 March) which were slightly higher. This could not be explained by local fire activity or unique meteorological conditions. It is concluded that a time period longer than the 16 days analyzed here would be required to determine the correlation between air mass origin and emission factor with statistical significance. The retrieval method works by selecting peak areas above a baseline concentration, and slow increases in particle concentration (or CO₂), such as those that might occur from changes in regional biomass burning activity may not influence the recovered emission factor. Days with new particle formation and growth activity have similar recovered emission factors as days without new particle formation and growth, and thus the method is not influenced by particle formation events.

The diurnal pattern of the emission factor is observed using 3- averages as a function of time across all 16 days (Fig. 8). The number on each bar represents the total number of CO₂ peaks that were identified and used to calculate an emission factor for each time period. As expected, the largest number of CO₂ peaks was identified from 06:00 to 09:00, coincident with the morning rush hour traffic. Relatively very few peaks are detected from 12:00 to 18:00, during the time which the basin experiences ventilation and dilution. Stated another way, 86% of the peaks used to calculate the daily average emission factors are detected in the hours from 00:00–12:00 and 21:00–24:00, causing the afternoons to have less of an effect on the overall emission factor.

An analysis of the wind directional dependence was performed, including plotting of pollution roses for number and volume, pollution roses for emission factors, and conditional probability functions for CO₂ and emission factors. Normalizing for frequency of different wind directions, combined “number and CO₂” plumes were about 2 to 4 times more likely to originate from specific directions (the west, from the south, and

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from the northeast) than from other directions. These directions correspond to an industrial area with many SO₂ point sources (to the west), Vallejo Ave. (to the west), Eje 4 Ave. (to the south) and Eje Central Ave. (to the northeast). While CO₂ had a directional signature, the emission factors themselves were much less dependent on direction. The 10 min periods with highest number to CO₂ peak size ratios preferentially occurred under winds from the west. But more typical emission factors showed no directional dependence. Incidentally, the most common morning wind direction was from the west, possibly contributing to the directionality of the highest emission factors. The fact that emission factors are largely non-directional eliminates the possibility of recovering source-specific emission factors without additional source-specific markers. However, at the same time, it makes the recovered emission factor potentially more representative of a wide geographical area.

4.5 Comparison to emission factors from literature

The size-resolved number emission factors determined in this work are compared to those of three different experimental studies, and the result is shown in Fig. 9. Additionally, a representative size distribution used as input to a global aerosol model (Spracklen et al., 2010) is included in the figure. The first study of interest was conducted in the Minneapolis metropolitan area by Kittelson and colleagues in summer of 2002 (Kittelson et al., 2006; Johnson et al., 2005). On-road mobile measurements including the size distribution from 10–300 nm, CO and CO₂, were collected under typical highway driving conditions. Then, the apportionment between SI and diesel vehicles was calculated using a differential traffic volume correlation method. CO₂ and CO contributions were apportioned using the same method, and assuming these species represent the total carbon emitted, a conversion to fuel specific emission factors was made.

The other two studies of interest were conducted in the Caldecott Tunnel in Northern California. The tunnel has two separate traffic bores; one which allows both heavy- and light-duty traffic, and the second which allows light-duty traffic only. In each of

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median radius of 30 nm (Spracklen et al., 2010; Stier et al., 2005). Then emission rates (Tg yr^{-1}) compiled from emission inventories are distributed across the representative log-normal distribution functions (Spracklen et al., 2010; Dentener et al., 2006). For this reason, comparison on an absolute scale with the emission factor presented here is difficult. However, the shape of model input emission distribution is compared to the emission factor distribution recovered in this work in Fig. 9. While the mode of the two distributions is similar, the model assumed distribution is narrower, and therefore may potentially underestimate the number of particles below 30 nm and above 40 nm relative to the mode.

4.6 Application of emission factors to Mexico City plume

The average emission factor of $385.5 \text{ particles cm}^{-3} (\text{mg m}^{-3} \text{CO}_2)^{-1}$ was used to predict number concentration in the aged Mexico City plume based on CO_2 measurements taken aboard the NASA DC-8. The volume emission factor (through 2 microns) was $0.35 \mu\text{m}^3 \text{ cm}^{-3} (\text{mg m}^{-3} \text{CO}_2)^{-1}$. On a molar basis at the temperature, pressure, and water partial pressure of the aircraft sampling (296 K, 0.75 bar, 2.3 mbar $P_{\text{H}_2\text{O}}$), the emission factors are $516 \text{ particles cm}^{-3} \mu\text{mol CO}_2 \text{ mol}^{-1}$ and $0.47 \mu\text{mol CO}_2 \text{ mol}^{-1}$. At 1.013 bar and 293 K, the emission factors are somewhat higher (706 and 0.64, respectively). This predicted number concentration was compared to number concentrations measured on the aircraft. Because of extensive secondary particle formation and growth in the plume, agreement between the prediction based on the urban primary size distribution and the in situ measurement is not expected; rather the amount of disagreement may indicate the extent of plume processing. The data used for this analysis was taken on 11 March, when the NASA DC-8 conducted a flight in order to sample the aged Mexico City plume as well as the near-source plume. A pass over the city (T0) was made (19.45 lat. -99.07 lon. at 2530 m altitude), and the Mexico City plume is easily identified by the CO_2 mixing ratio gradient where concentrations up to $400.5 \mu\text{mol mol}^{-1}$ are seen over the city, with a gradual decrease to background levels ($382 \mu\text{mol mol}^{-1}$) traveling outward.

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The number size distribution evolution at T0 during this day had particle volume peaking at 08:00 CST, a minimum in aerosol volume and aerosol scattering at noon, aerosol growth and mass increases during the early afternoon, and a significant drop in aerosol volume and aerosol scattering from an air mass change at ~16:30 CST. The DC-8 flew over Mexico City at high altitude (~9 km a.s.l.) and then flew north over T0, T1 and T2 at an altitude of ~2.5 km a.s.l., or <1 km above ground). In this analysis, we used five measurements of the particle size distribution beginning at 14:19 CST and covering a 10-min duration. The first measurement was obtained over the city, followed by measurements at 15.7, 30.7, 46.6 and 61.8 km (where CO₂ has decreased to within 0.5 μmol mol⁻¹ of the background concentration). Measurements of CO₂ were obtained at 1-s resolution throughout the duration (Vay et al., 2009). The photochemical age at the time of each number concentration measurement was calculated using the measured ratio of toluene to benzene as described by Warneke (2007).

For this comparison, the baseline CO₂ and particle number concentrations aloft were determined using the same successive moving average subtraction method that was applied to the ground based measurements. During this period (consisting of five datapoints), the baseline of number concentration from the successive subtraction technique averages 3997 cm⁻³ and ranges from 3512 cm⁻³ to 4388 cm⁻³. The results from this analysis are provided in Fig. 10. The ΔCO₂ (green), which is the increase in CO₂ mixing ratio above a baseline as measured by the aircraft, is plotted on the right axis versus photochemical age. The corresponding predicted increase in particle number concentration above a baseline (ΔNum) is calculated by multiplying ΔCO₂ by the average emission factor and is represented by the dashed line. The actual measured ΔNum is represented by the solid red line.

As shown in Fig. 10, the measured increase in particle number concentration is significantly higher than predicted using the emission factor determined at the ground.

Our speculation is that there is a source of particles during plume aging that significantly increases particle number while leaving CO₂ unchanged. New particle formation is a possible explanation for the elevated particle number concentrations. The

specific appearance of Fig. 10 is sensitive to the number baseline used to determine the increase relative to background. The increase in the baseline by 3400 cm^{-3} (on average) would bring the predicted ΔNum and the measured ΔNum into much closer agreement. The interpretation of this new background (which would be on average 7400 cm^{-3}) is that it would require a particle source not associated with CO_2 – and this could be from new particle formation, or from other non-combustion sources. For comparison, the average number baseline calculated for all samples below 3 km a.s.l. was 1800 cm^{-3} and the 90th percentile of the low altitude (<3 km a.s.l.) datapoints was 3600 cm^{-3} , so 7400 cm^{-3} would be considered a high background value. The average total number concentration at altitudes <3 km a.s.l. is 3400 cm^{-3} .

Figure 11 shows the analysis using the size resolved measurements (and the size resolved emission factor, represented by a single lognormal mode as discussed above) from the DC-8 rather than the total number. Figure 11a shows the measured number distribution (symbols) and the predicted number size distribution based on the primary emission factor of this work. The prediction is equal to a background distribution plus the emission factor times the increment in CO_2 above background (ΔCO_2). The shape of the background distribution is calculated as the average of the measured number distributions from 20:16–20:27 UTC, and the number is fixed at the 3997 cm^{-3} (the background total number calculated above). Both this shape and absolute number are uncertain. But even allowing for a higher background, the measured particle concentrations is well in excess of the primary prediction at less than 40 nm for three of the four distributions. At sizes from 40–100 nm, the prediction is low for two of the three distributions.

Figure 11b and c focus on the aerosol volume distribution. Figure 11b graphs actual and predicted volume distributions based on the primary emission factor of this work. The background distribution function is calculated as the minimum of the volume distributions from 20:16–20:27 UTC. The integral of this background volume distribution through 2 microns is $3.59\text{ }\mu\text{m}^3\text{ cm}^{-3}$. The integral of the volume distributions in Fig. 11b range from $5.6\text{--}18.8\text{ }\mu\text{m}^3\text{ cm}^{-3}$.

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Figure 11c shows the volume distributions normalized to the CO₂ increment above baseline. The photochemical ages, calculated by the measured benzene/toluene ratio, are shown in the figure and range from 2–15 h (referenced to 0 h set using the benzene toluene ratio at 20:19). Consistent with the conclusions of Kleinman et al. (2009) the volume integral (and also the number concentration from 0.1–1 microns, which is well correlated with volume) increase significantly with photochemical age; furthermore the change in shape (increase in volume without significant increase in mean size) is consistent with condensational growth rather than volume growth. While the number of measurements in this work is insufficient to show if this increase is linear with photochemical age as shown by Kleinman et al., the apparent secondary volume can be compared to the primary emitted volume from the emission factor of this work. In the 20:19 distribution, the volume above calculated baseline (through 2 microns) $0.91 \pm 0.14 \mu\text{m}^3 \text{cm}^{-1}$ per $\mu\text{mol mol}^{-1}$, while the primary emission factor calculated at T0 (to within a factor 1.5) is 0.47. Thus the primary emissions account for only ~34% of the volume below 2 microns, even in this relatively fresh sample. In the later samples with VOC ages of 13–15 h, the fraction explained by the primary emissions drops to ~20%.

Using the CO to CO₂ molar ratio of 0.045 determined in from aircraft measurements of the Mexico City plume (Vay et al., 2009) one can convert the CO₂ based emission factors recovered in this work to CO-based emission factors. The emission factor of 706 particles cm^{-3} ($\mu\text{mol mol}^{-1}$)⁻¹ used above converts to 15 690 particles cm^{-3} ($\mu\text{mol mol}^{-1} \text{CO}$)⁻¹ at 1.013 bar and 293 K. The volume emission factor of $0.64 \mu\text{m}^3 \text{cm}^{-3} \mu\text{mol CO}_2 \text{mol}^{-1}$ (through 2 microns) converts to $14.2 \mu\text{m}^3 \text{cm}^{-3}$ per $\mu\text{mol mol}^{-1}$ of CO. The volume emission factor through 0.44 microns (which matches the upper size limit of the DMA in Kleinman et al., 2009) is then $7.7 \mu\text{m}^3 \text{cm}^{-3}$ per $\mu\text{mol mol}^{-1}$ of CO. From Fig. 2 in Kleinman et al. (2009), the y intercept of the DMA volume versus photochemical age is $\sim 11 \mu\text{m}^3 \text{cm}^{-3}$ per $\mu\text{mol mol}^{-1}$ of CO.

This preliminary analysis of a small number of distributions from the DC-8, normalized to CO₂ and referenced to a primary emission factor, is consistent with the more

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comprehensive analysis of aerosol aging on the the G-1 aircraft. The primary emission factor of this work predicts a fraction of the observed volume aloft close to Mexico City ranging from ~34–70%. The increase in aerosol volume relative to a marker of dilution (CO₂ increment above baseline) is similar to that documented by Kleinman et al. (2009) with a larger dataset from the G-1. The emission factor also cannot account for the observed aerosol number distribution, particularly below 100 nm. The mismatch is not systematic with photochemical age (as it appears to be for aerosol volume).

5 Conclusions

Measurements of the aerosol size distribution from 11 nm to 2.5 microns were made in Mexico City at the T0 site in March, 2006, during the MILAGRO field campaign. Morning conditions included high particle mass concentrations, low mixing heights, and highly correlated particle number and CO₂ concentrations. Mean concentrations of particle number and volume are: 20 800 cm⁻³ and 18 600 cm⁻³ (from the BNL and Iowa SMPS, respectively), and 15.4, 12.0, and 10.1 μm³ cm⁻³ (from the BNL SMPS, Iowa SMPS, and APS respectively). Size ranges for these summary values run from 15–494 nm, 11–478 nm, and 0.56–2.5 μm, respectively.

Ultrafine particle growth events, defined as a prominent growing mode at 10–15 nm, were observed on 6 of 16 (37.5%) days sampled at the T0 location. Events typically began between 10:30 and 13:00 local time. The growth rates of the ultrafine particles varied between 4.9 and 17.7 nm h⁻¹.

Average size-resolved and total number- and volume-based emission factors for combustion sources impacting T0 have been determined using comparison of the areas of particle number and CO₂ concentration peaks. The analysis is most sensitive to peaks smaller than 1 h in time duration, and is insensitive to slow changes in the CO₂ and particle number baselines. The technique employed in the current work is somewhat sensitive to the CO₂ peak threshold selection used, but is not

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sensitive to afternoon nucleation events, which happen during a time when CO₂ does not peak above the threshold. The number emission and volume emission factors are 1.23×10^{15} particles, and 7.54×10^{11} cubic microns per kg of carbon, respectively for the particle size distribution from 11 nm to 494 nm. The uncertainty on the number emission factor is approximately a factor of 1.5. The mode of the number emission factor was between 25 and 32 nm, while the mode of the volume factor was between 0.25 and 0.32 microns.

Representing the emission factor as log normal modes yields parameters that may facilitate intercomparison of the emission factor and its use in modeling studies. For example, a single lognormal mode fit achieved an R^2 of 0.94 with respect to the number distribution and an R^2 of 0.85 with respect to the volume distribution. This fit had parameters of N 1.65×10^{15} particles per kg C, D_p of 30 nm, and $\log \sigma$ of 0.41 (shown in Figs. 7 and 9). While CO₂ and particle number exhibited preference for specific directions as determined by conditional probability of peaks versus wind direction, emission factors were largely insensitive to wind direction. While this limits potential to isolate specific sources, it does perhaps increase the ability to generalize the emission factor to wider parts of the Mexico City airshed.

The recovered emission factor resembles those used in some global modeling studies to represent carbonaceous aerosols. It also has a large number of common features with tunnel and chase studies aimed at isolating vehicle emission factors. The size-resolved emission factor from Mexico City is between several of the light-duty (lower emission factor) and heavy-duty (higher emission factor) emission factors from the United States. At sizes smaller than 30 nm, the emission factor is lower than many vehicular emission factors, possibly due to the greater degree of scavenging of ultrafine particles between the point of emission and the point of sampling in our study versus the tunnel and chase studies. At sizes larger than 30 nm, the emission factors from Mexico City are larger than several published vehicular emission factors.

Better collocation of the CO₂, meteorology, and number size distribution measurement, and the use of faster instrumentation for the size distribution would increase

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the power of this technique to isolate shorter duration peaks and eliminate sources of uncertainty in the emission factor (although much of the variability is real and would remain).

Comparison of the size-resolved primary emission factor to the number distribution in the Mexico City plume aloft, as measured by the NASA DC-8 on 11 March 2006, was conducted. A plume with CO₂ peaking at 11 μmol mol⁻¹ above background was isolated. The emission factor of this work explains (at 11 μmol mol⁻¹ enhancement) only about 5700 cm⁻³ of the number enhancement (which totaled 16 000 cm⁻³ above background). This relative underprediction continued in subsequent samples with lower CO₂ enhancements but was not correlated with photochemical age. The recovered emission factor was consistent with recent work showing that the ratio of total to primary aerosol in urban plumes can quickly grow to values of 2 to 5.

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Table 1. Measurements used in the data analysis for this work.

Instrument	Group (PI)	Measurement	Temporal Resolution
Aerodynamic Particle Sizer (TSI Inc., 3321)	University of Iowa (C. Stanier)	Size distribution from 0.542 μm to 19.8 μm	20 s
SMPS (TSI Inc., 3081)	University of Iowa (C. Stanier)	Size distribution from 10.9 nm to 478 nm	5 min
SMPS (TSI Inc., 3081)	Brookhaven	Size distribution from 15 nm to 494 nm	2 min
LI-COR LI-7500	University of Iowa (W. Eichinger)	CO ₂	1 s
Integrating Nephelometer (TSI Inc., 3563)	Department of Energy (N. Marley)	Aerosol scattering at 450, 550, and 700 nm	
Vaisala WXT510	Department of Energy (N. Marley)	Wind speed and direction	1 min 40 s
Thermo Andersen PM _{2.5} Beta Attenuation Monitor (CAM, SJA), TEOM 1400a-FDMS 8500 PM _{2.5} (MER)	Mexico City Ambient Air Monitoring Network (RAMA)	PM _{2.5}	Hourly
Modified LI-COR 6252 analyzer	NASA (S. Vay)	CO ₂ (DC-8 aircraft)	1 s
Modified DMA	NASA (A. Clarke)	Particle number (DC-8 aircraft)	
Canister samples of trace gases	UC Irvine (D. Blake)	Benzene, Toluene (DC-8 aircraft)	1–2 min 4–5 min

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Table 2. Mean values and distributions of ground-based measurements at T0 from 10–25 March 2006.

Measurement	Mean	Maximum	Minimum
CO ₂	559 mg m ⁻³	659 mg m ⁻³	492 mg m ⁻³
	392 μmol mol ⁻¹	454 μmol mol ⁻¹	359 μmol mol ⁻¹
Particle Number (BNL)	2.1 × 10 ⁴ cm ⁻³	5.7 × 10 ⁴ cm ⁻³	7.1 × 10 ³ cm ⁻³
Particle Number (Iowa, not dried)	1.9 × 10 ⁴ cm ⁻³	4.1 × 10 ⁴ cm ⁻³	1.9 × 10 ³ cm ⁻³
Temperature	17.6 °C	29.4 °C	8.8 °C
Relative Humidity	39.4 %	86.6%	6.3 %

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Table 4. Parameters for lognormal modal fits of the emission factor.

Fit	Mode	N # (kg C) ⁻¹	D _p (nm)	log sigma	Notes
Single mode fit, 80% weighted to number distribution, 20% weighted to volume distribution	1	1.65 × 10 ¹⁵	29.6	0.41	0.94 R ² toward N; 0.85 R ² relative to V
Three mode fit, 100% weighted toward number distribution	1	1.12 × 10 ¹⁵	32.7	0.30	0.94 R ² toward N;
	2	2.20 × 10 ¹⁴	16.9	0.15	0.46 R ² relative to V
	3	1.80 × 10 ¹⁴	151	0.21	
Three mode fit, 100% weighted toward volume distribution	1	2.03 × 10 ¹⁵	32.7	0.37	0.73 R ² toward N;
	2	6.56 × 10 ¹¹	1342	0.59	0.98 R ² relative to V
	3	1.70 × 10 ¹³	229	0.15	
Four mode fit, 80% weighted to number distribution, 20% weighted to volume distribution	1	1.03 × 10 ¹⁵	34.3	0.29	0.99 R ² toward N;
	2	2.83 × 10 ¹⁴	17.6	0.15	0.94 R ² relative to V
	3	1.94 × 10 ¹⁴	144.5	0.22	
	4	1.79 × 10 ¹²	1278	1.09	

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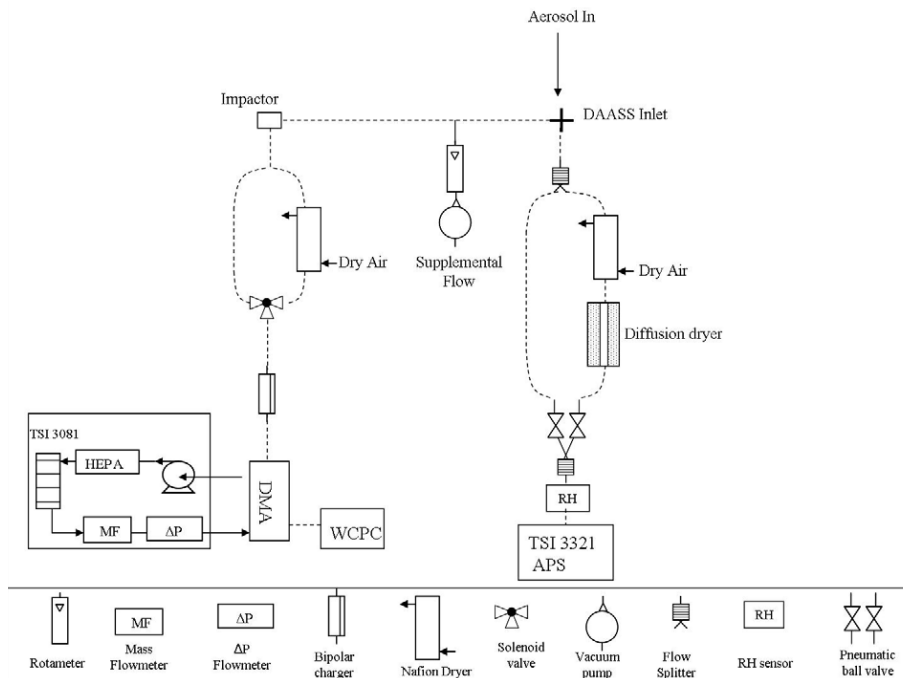


Fig. 1. University of Iowa RH-controlled inlets and sizing instrumentation.

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Fig. 2. Aerial view of the T0 research site and surrounding land use. The locations of the CO₂ monitor and SMPS instrumentation are denoted by an “x”.

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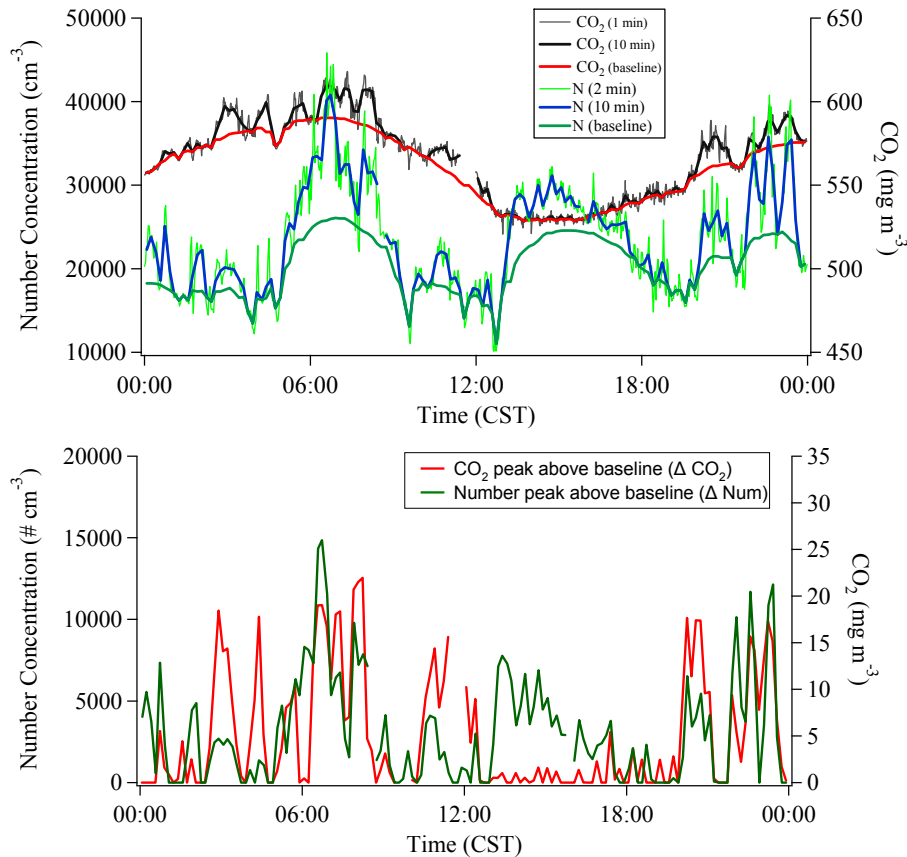


Fig. 3. Time series data for 17 March, illustrating the method used to retrieve the number-based emission factors based on the correlation between CO₂ and particle number for Mexico City.

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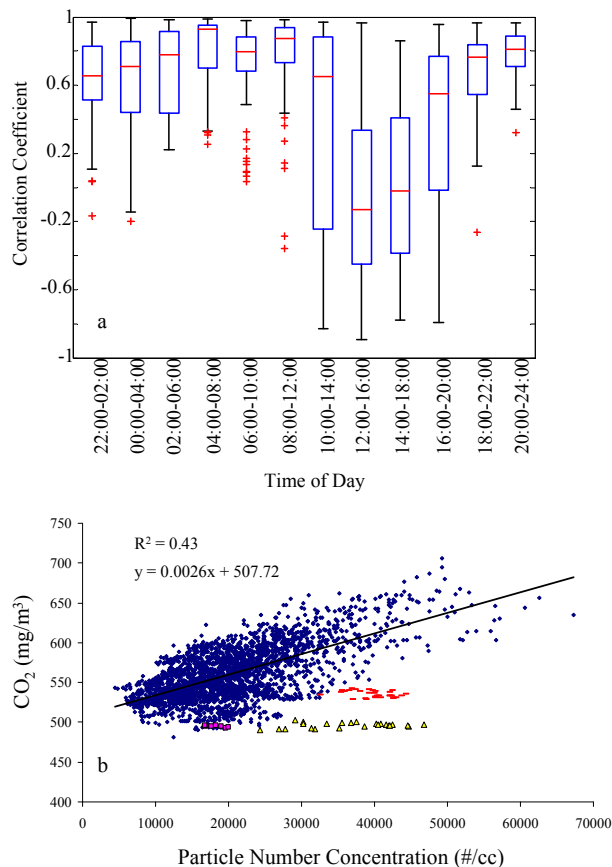
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Fig. 4. (a) Correlation coefficient of 10-min averaged CO₂ and total number over 4-h periods and (b) 10-min averages of CO₂ concentration versus particle number concentration for the entire period from 9 March to 29 March. Squares, triangles and dashed lines highlight measurements obtained during new particle formation events.

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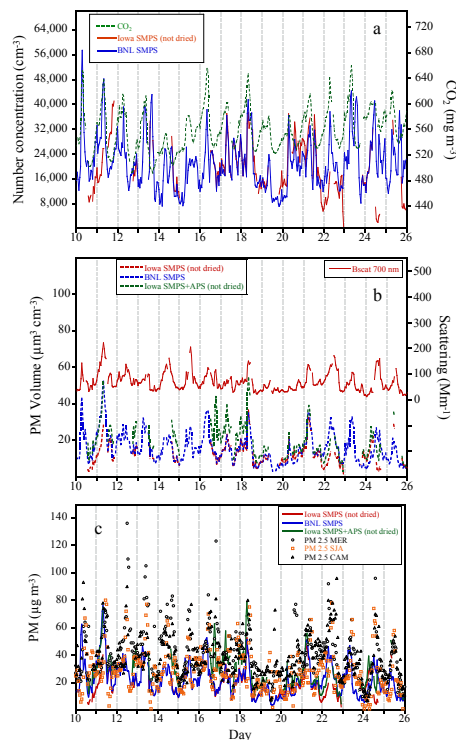


Fig. 5. Time series measurements made at the T0 research site during March 2006. **(a)** Particle number concentration (SMPS) and CO₂, **(b)** Particle volume (SMPS and APS) and total aerosol scattering (nephelometer), **(c)** estimated PM_{2.5} (SMPS and APS) and nearby (within 10 km) PM_{2.5} measurements from 3 different RAMA monitoring locations, **(d)** diurnal average for CO₂ and particle number concentration **(e)** diurnal average for particle volume (SMPS, APS), extinction (nephelometer), and PM mass (RAMA), **(f)** grand average wind rose for wind speed and direction at T0, in meters per second.

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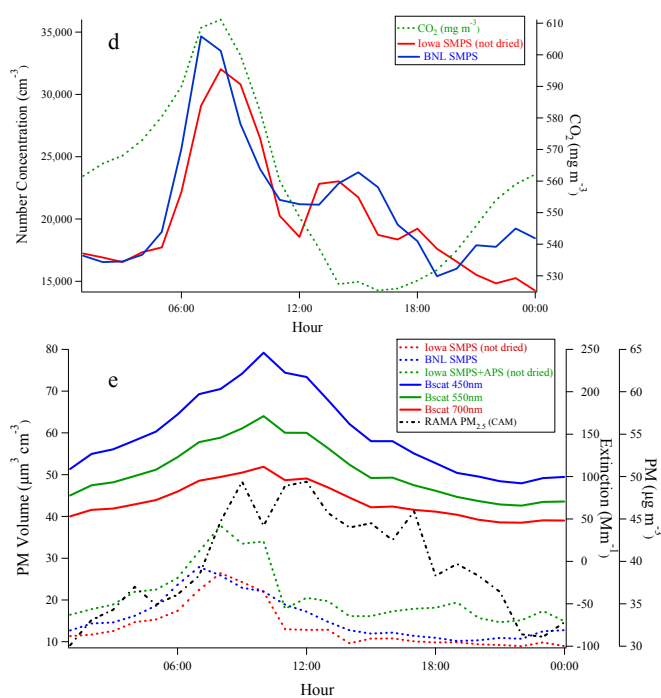


Fig. 5. Continued.

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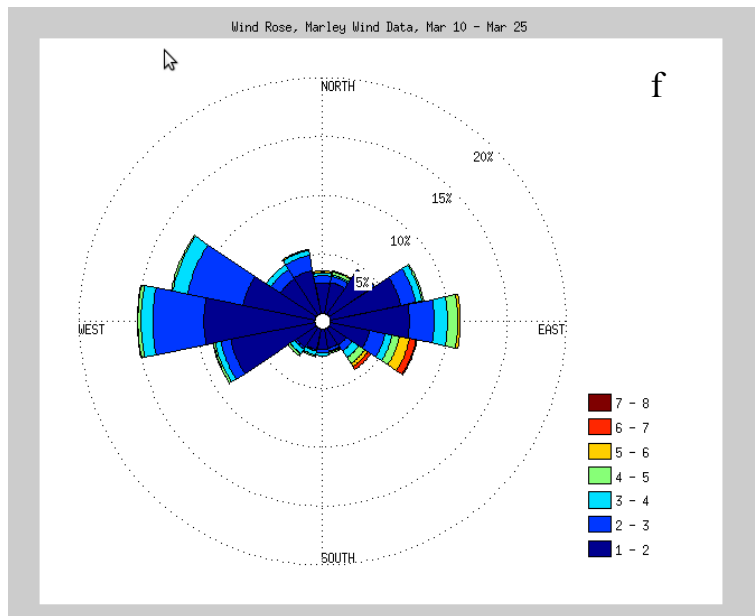


Fig. 5. Continued.

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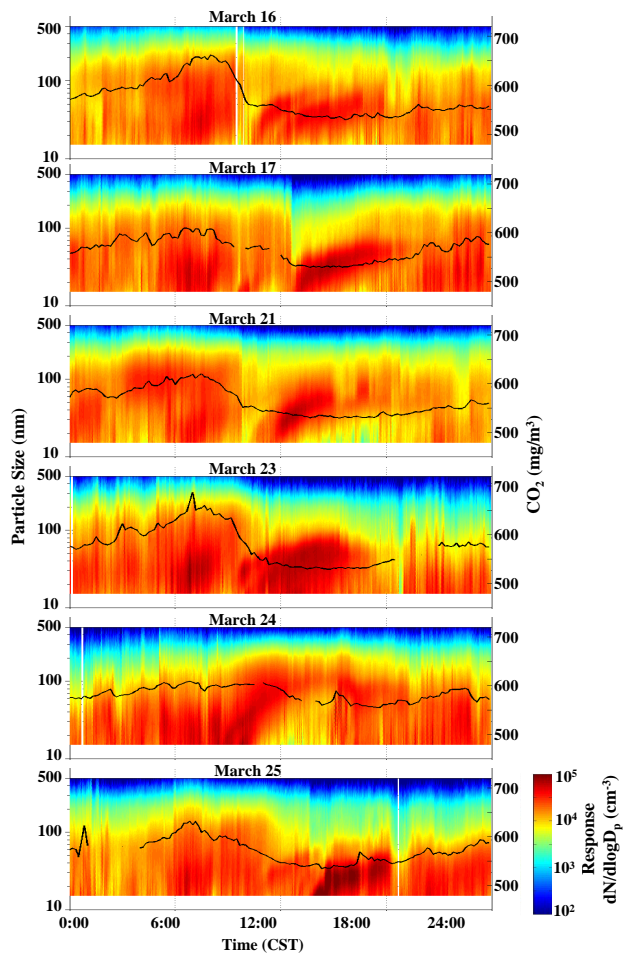


Fig. 6. Colorplots of ultrafine growth events.

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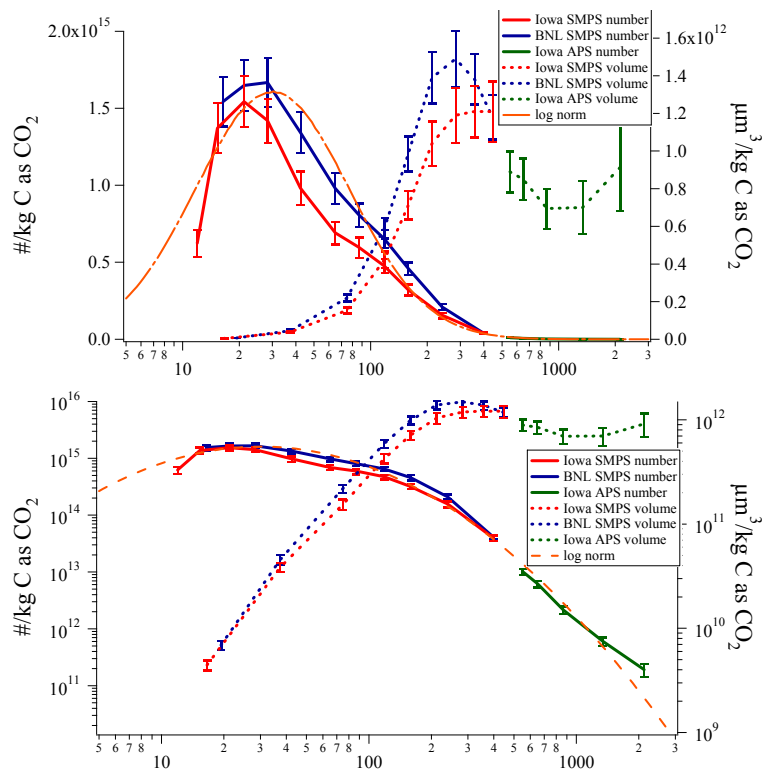


Fig. 7. Grand average number and volume emission factor distributions shown on both a log and linear scale. Error bars represent the 95% confidence interval based on variability in sampled values (see text for discussion of additional sources of uncertainty).

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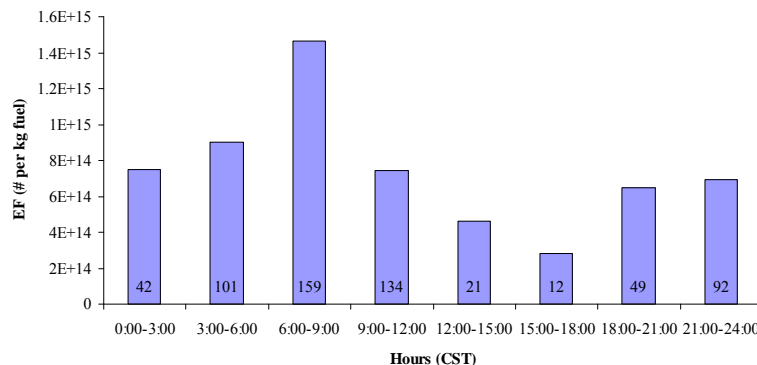
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Fig. 8. Average emission factor versus time of day. The numbers on the graph represent the total number of CO₂ peaks that went into the calculation of the emission factor for each 3 h period.

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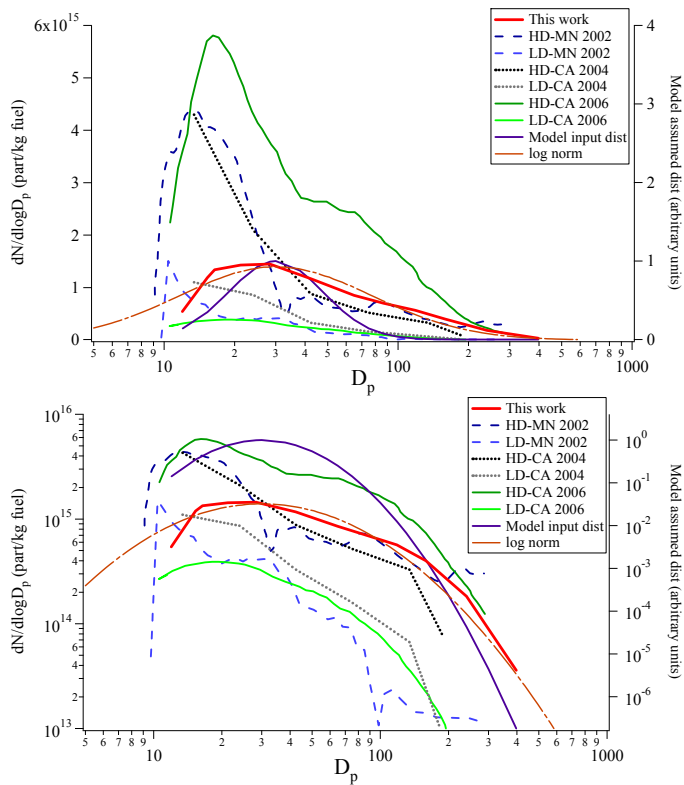


Fig. 9. Comparison of measured emission factor distribution from Mexico City to experimentally determined, fleet averaged emission factors and a model assumed carbonaceous emission distribution. The model assumed distribution is listed in arbitrary units to allow for comparison of the distribution shape.

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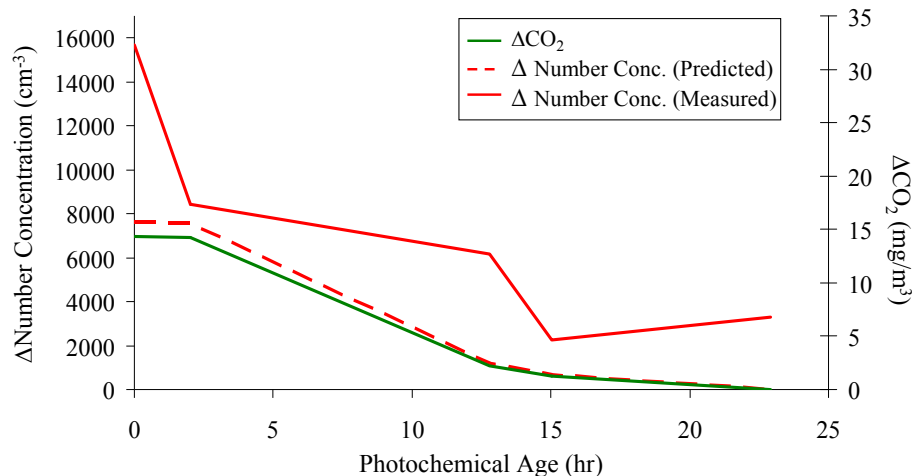


Fig. 10. Comparison of the predicted increase in particle number above the baseline concentration (dashed red) associated with an increase in CO₂ above baseline (green), with the measured increase in particle number (red). Measurements of CO₂ and particle number concentration were obtained on 11 March 2006 aboard the NASA DC-8 flight which transected the Mexico City plume (14:15–14:30 CST). The emission factor used to predict the increase in particle number is 687.9 #/ppm CO₂.

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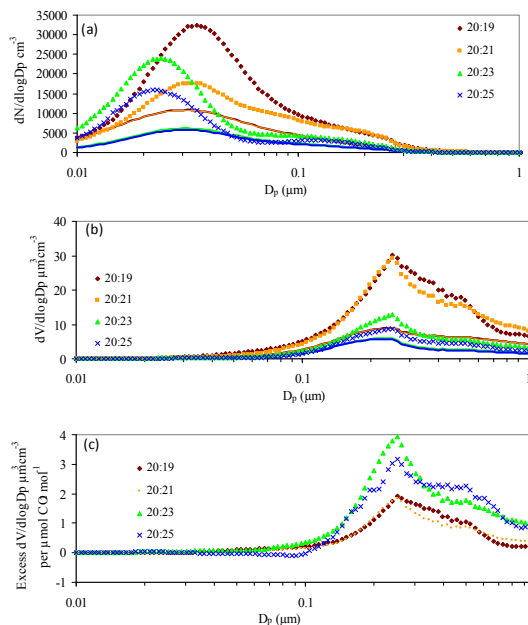


Fig. 11. Measured size and volume distributions in the DC-8 plume near Mexico City during the 11 March, 2006 flight at average altitude of 342 m above ground at 14:20 local time. Panel (a) shows measured number size distributions (symbols) with predicted size distributions (lines) based on a regional background plus the primary emission factor (this work) times the CO_2 increment above background. Panel (b) shows the same as (a) except number size distributions have been converted to volume distribution. Panel (c) shows the difference between measurement and background + primary, normalized to the CO_2 increment in the plume. VOC ages using the benzene/toluene ratio relative to the ratio at 20:19 are shown by each distribution. Integrals of the volume distributions through 2 microns in (c) are 0.86 ± 0.15 , 0.88 ± 0.15 , 2.0 ± 1.0 , and $1.7 \pm 2.0 \mu\text{m}^3 \text{cm}^{-3}$ per ppm. For reference, the primary emission factor (through 2 microns) is $0.51 \mu\text{m}^3 \text{cm}^{-3}$ per ppm.

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