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Primary and secondary sources of formaldehyde in urban atmospheres: **Houston Texas region**

D. D. Parrish¹, T. B. Ryerson¹, J. Mellqvist², J. Johansson², A. Fried³,

D. Richter³, J. G. Walega³, R. A. Washenfelder^{1,4}, J. A. de Gouw^{1,4}, J. Peischl^{1,4}, K. C. Aikin^{1,4}, S. A. McKeen^{1,4}, G. J. Frost^{1,4}, F. C. Fehsenfeld^{1,4}, and

S. C. Herndon⁵

¹NOAA ESRL Chemical Sciences Division, 325 Broadway, Boulder, CO, USA

²Earth and Space Sciences, Chalmers University of Technology, Gothenburg, Sweden

³Earth Observing Laboratory, National Center for Atmospheric Research, Boulder, CO, USA ⁴CIRES, University of Colorado, Boulder, CO, USA

⁵Aerodyne Research, Inc., Billerica, Massachusetts, USA

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Correspondence to: D. D. Parrish (david.d.parrish@noaa.gov)

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Introduction

Formaldehyde (CH₂O) is an oxygenated volatile organic compound (VOC) that plays an important role in the formation of ozone pollution in urban areas. Both primary sources (i.e., direct emissions from anthropogenic sources) and secondary sources

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(i.e., production in the atmosphere during oxidation of other, directly emitted VOCs) contribute to atmospheric concentrations of CH₂O. Most secondary production of CH₂O is expected to occur during the atmospheric oxidation of ethene, propene and higher terminal alkenes, such as 1-butene, 1,3-butadiene and isoprene, but CH₂O is additionally formed more slowly from the oxidation of alkanes and aromatic compounds. CH₂O is lost from the atmosphere through photolysis, reaction with the hydroxyl radical (OH), and deposition.

Quantifying the relative contribution of primary and secondary CH_2O sources is crucial to developing effective ozone control strategies in urban areas. Photolysis of CH_2O is an important source of OH radicals, which are the species that initiate atmospheric photo-oxidation, and serves as a fuel for the photochemical cycles that produce ozone. Accumulation of CH_2O during nighttime hours from direct emissions could provide large CH_2O concentrations at dawn that could initiate photochemistry earlier in the diurnal cycle than would be the case in their absence. Thus, emissions from primary sources are an attractive target for regulatory efforts designed to reduce urban ozone concentrations.

Urban sources of atmospheric CH₂O have been investigated for decades. In Los Angeles in 1980 Grosjean (1982) measured concentrations as high as 48 ppbv, and reported measurements by others from the 1960s showing that CH₂O exceeded 100 ppbv in the worst photochemical episodes in that city. Based upon the observed diurnal cycle, Grosjean (1982) concluded that both direct anthropogenic emissions and photochemical production made substantial contributions to ambient CH₂O concentrations. A variety of statistical studies have attempted to quantify the relative amounts of ambient CH₂O contributed by primary and secondary sources in several cities, including Vancouver (Li et al., 1997), Houston (Friedfeld et al., 2002; Rappenglück et al., 2010; Buzcu Guven and Olaguer, 2011) and Mexico City (Garcia et al., 2006).

The quantification of primary and secondary formaldehyde sources is particularly important in Houston, Texas, which is characterized by strongly elevated atmospheric CH_2O concentrations (Wert et al., 2003; Ryerson et al., 2003; Martin et al., 2004).

Houston is home to a very large industrial sector associated with petrochemical and petroleum refining activity, and these industrial activities are associated with the elevated CH₂O concentrations. Given this industrial activity, the relative contributions from primary and secondary sources may be significantly different from most urban areas. Indeed, Olaguer et al. (2009) have argued that primary emissions from this industrial sector may make large contributions to ambient CH₂O, and thus should be identified, quantified and controlled.

In this work, we present analytical methods for quantifying both primary and secondary sources of CH2O. The major primary sources of CH2O that have been suggested to be important in Houston-Galveston-Brazoria (HGB) are motor vehicles and the area's industrial facilities. Primary emissions from the industrial facilities are derived from direct flux measurements, and those from the vehicle fleet are derived from measured ambient CH₂O to CO ratios under conditions dominated by vehicle emissions, combined with emission inventory estimates for vehicle CO emissions. The secondary sources of CH₂O in HGB are production from primary emissions of parent VOCs emitted from these same anthropogenic sources, as well as VOCs of biogenic origin. Photochemical oxidation initiated by OH during daytime is expected to dominate this secondary production, but nighttime oxidation initiated by ozone (O₃) or the nitrate radical (NO₃) reacting with those emitted VOCs also contributes. The amount of CH₂O produced by secondary sources is derived from the estimated yield of CH₂O from reacted VOCs combined with emission inventory estimates of industrial and vehicle VOC emissions. Although our primary goal is to provide a quantitative analysis of CH₂O emitted by primary sources and formed from secondary sources within the HGB ozone nonattainment area, the approach presented here is applicable to other urban areas and to other photochemical species.

The following section describes the data sets utilized in this paper, and Sects. 3 and 4 address emissions from petrochemical facilities and on-road vehicle emissions. Section 5 compares our results to other analyses and discusses the reasons for the divergent results, and Sect. 6 discusses the results and presents conclusions.

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The analysis presented here is based upon archived data sets that have been described elsewhere; only brief introductions and references to these descriptions are given here. NOAA conducted two airborne studies in the HGB region during the Tex-5 AQS 2000 (Ryerson et al., 2003; Wert et al., 2003) and TexAQS 2006 (Washenfelder et al., 2010; Peischl et al., 2010) field studies; those data are available at http://esrl.noaa.gov/csd/tropchem/. The aircraft platforms were the NCAR Electra in 2000 and the NOAA WP-3D in 2006. Airborne CH₂O concentrations were acquired by NCAR employing tunable infrared laser absorption spectroscopy. During the 2000 study a tunable diode laser absorption spectrometer described by Wert et al. (2003) was employed, while the 2006 study employed a tunable difference frequency generation laser absorption spectrometer, as described by Weibring et al. (2007). Both instruments provided 1-s to 10-s CH₂O measurements. Both aircraft campaigns included 1 Hz measurements of O₃, nitric oxide (NO), nitrogen dioxide (NO₂) total reactive nitrogen (NO_v), carbon monoxide (CO), sulfur dioxide (SO₂), and carbon dioxide (CO₂) (Ryerson et al., 1998, 1999, 2000; Holloway et al., 2000; Daube et al., 2002). Speciated VOCs were measured by gas chromatography (GC) of whole air samples acquired during each flight (Schauffler et al., 1999). Both aircraft campaigns included speciated VOC measurements by proton transfer reaction mass spectrometry (PTR-MS) (de Gouw and Warneke, 2007), and the 2006 field campaign included ethene (C₂H₄) measurements at 5 s resolution with laser photoacoustic spectroscopy (LPAS) (de Gouw et al., 2009). Parrish et al. (2009) give additional details of the 2006 measurements.

Chalmers University of Technology equipped a mobile van with Solar Occultation Flux (SOF) and mobile Differential Optical Absorption Spectroscopy (DOAS) instrumentation (Mellqvist, 1999; Rivera et al., 2010; Mellqvist et al., 2010a) to measure vertical columns of CH₂O, ethene, propene, and other VOCs in 2006 and in 2009. The SOF technique is based on open path Fourier Transform Infrared (FTIR) Spectroscopy

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using direct solar radiation as the light source, while the mobile DOAS is an open path system with scattered solar radiation as the light source. Installation in a mobile van allows continuous column concentration measurements to be performed while transecting an emission plume. These measurements, together with measured position and wind speed, make it possible to calculate emission fluxes in the plume. The accuracy of these flux determinations is estimated to be on the order of 30 %, primarily due to the uncertainty of the wind speed. The SOF results are available from Mellqvist et al. (2010b).

The University of Houston conducted extensive measurements at Moody Tower, a site on the top of a 65 m building in Houston, Texas during the TexAQS-II radical and aerosol measurement project (TRAMP) (Lefer and Rappenglück, 2010), which was a component of the second Texas Air Quality Study (TexAQS II) (Parrish et al., 2009). Lefer and Rappenglück (2010) and references therein describe the measurements including CH_2O , CO, O_3 , NO_y , and the photolysis rate of NO_2 (j NO_2). The analysis in the present paper utilizes the CH_2O (measured by Hantzsch reaction fluorescence) and CO (measured by Gas Filter Correlation) data. The measurements were conducted from 13 August to 2 October 2006. The results reported here are based on 10-s averaged data that were provided to us by the TRAMP measurement team on 23 May 2008.

Baylor University deployed a Piper Aztec aircraft in the HGB region during the summer of 2006 (Baylor University, 2009; Olaguer et al., 2009). Measurements included CH₂O (measured by Hantzsch reaction fluorescence), O₃, NO, NO₂, NO_y, CO, and VOCs (measured by canister sampling with gas chromatograph/flame ionization detection analysis). The data are available from the Texas Environmental Research Consortium (TERC) website: http://projects.tercairquality.org/AQR/H063.

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3.1 Quantification of formaldehyde formed from oxidation of petrochemical HRVOC emissions

Analysis of observations made during the TexAQS 2000 study (Ryerson et al., 2003; Wert et al., 2003; Kleinman et al., 2002, 2003; Daum et al., 2003) established that the petrochemical industrial facilities in Houston consistently emit large amounts of VOCs and oxides of nitrogen ($NO_x = NO + NO_2$) to the atmosphere. The VOCs characteristically include especially large concentrations of highly reactive volatile organic compounds (HRVOCs), in particular the alkenes ethene and propene. During day-time, these emissions produce plumes of elevated O_3 concentrations downwind from the sources, and analysis confirmed that the initial hydrocarbon reactivity in the petrochemical source plumes is primarily due to the alkenes. These plumes also contain high (as much as $> 30 \, \text{ppbv}$) concentrations of CH_2O formed as a secondary product of the HRVOC oxidation (Wert et al., 2003). Figure 1 shows one example of such a plume observed downwind of the Houston Ship Channel (HSC).

The evolution of the relationship between O₃ and CH₂O measured aboard the NCAR Electra in the 27 August 2000 plume is illustrated in Fig. 1 and quantitatively examined in Fig. 2. The flight involved multiple, crosswind transects flown upwind and downwind from HSC. The molar enhancement ratio of CH₂O to O₃ produced in the plume at a particular downwind transect is given by the slope of the linear correlation between the measurements made during that transect. In Fig. 2 all linear correlations are

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required to pass through the estimated background concentrations of CH₂O and O₃ appropriate for that day: 0.5 ppbv CH₂O (the concentration in background air over the Central Gulf of Mexico, Gilman et al., 2009) and 31.7 ppbv O₃ (the O₃ concentration at CH₂O = 0.5 ppbv calculated from the CH₂O-O₃ correlation for the farthest upwind transect at 29.0° N). Downwind of HSC the concentrations of both species increased rapidly, and by the second transect at ~ 24 km downwind (30.0° N) CH₂O reached its maximum concentration and the two species were well correlated ($r^2 = 0.88$). On subsequent transects, O₃ reached its maximum concentration, but the ratio of CH₂O to O₃ continually decreased through the farthest downwind plume transect while the correlation continued to increase to a maximum of $r^2 = 0.94$. Figure 3 summarizes the CH₂O to O₃ ratios at the downwind transects and compares the 27 August flight to a second flight conducted under similar conditions on 28 August (see Fig. 8 of Ryerson et al., 2003 and Figs. 5 and 6 of Wert et al., 2003).

The photochemical evolution of CH₂O in the plume illustrated in Figs. 2 and 3 suggests a useful approach for calculating the flux of secondary CH₂O formed in plumes downwind of petrochemical facilities. The peak CH₂O concentration is reached early in the plume transport since the daytime lifetimes of its HRVOC precursors are short (3-8 h for ethene and 1–2.5 h for propene, Wert et al., 2003). This slows CH₂O production as transport proceeds. In addition, the lifetime of CH₂O is also short (3 to 4 h in the sunlit lower troposphere, Seinfeld and Pandis, 1998). This leads to a rapid decrease of the CH₂O concentration when production slows.

Given these constraints, the total quantity of secondary CH₂O formed from primary HRVOC emissions can be calculated from the product of the total emissions times the yield of CH₂O produced during the atmospheric oxidation of these alkenes. The total HRVOC emissions in HGB are available from emission inventories and from direct ambient measurements of HRVOC fluxes in the downwind plumes. However, since the available inventories generally underestimate the alkene emissions from these facilities by large factors, we cannot use the 2005 National Emission Inventory (NEI) (Ryerson et al., 2003; de Gouw et al., 2009; Mellqvist et al., 2010a). Instead we use an inventory

(Brioude et al., 2011; Kim et al., 2011) that has been modified on a facility-by-facility basis to agree with the measured fluxes of ethene and propene (Mellqvist et al., 2010a). Since the lifetime of the alkenes are generally shorter than the time for transport of air masses out of HGB, this calculation will provide a realistic estimate of the secondary source of CH₂O from the petrochemical facilities.

On this basis the results of the quantification of the secondary CH₂O flux from specific petrochemical facilities and the total HGB area are given in Table 1. Assuming that OH is the primary oxidant of the alkenes, Seinfeld and Pandis (1998) give the product yields of 1.44 molecules CH₂O per molecule ethene and 0.86 molecules CH₂O per molecule propene. The product of the emission flux of each alkene times the product yield of CH₂O from that alkene yields an estimate of the secondary CH₂O formed from that alkene. A sum over the emitted alkenes gives an estimate of the total secondary CH₂O. Table 1 gives the alkene fluxes directly measured from specific facilities, as well as the integration over the entire HGB region (latitude 28.9 to 30.6° N; longitude 94.4 to 96.2° W) from the emission inventory. Table 1 also gives the flux of secondary CH₂O that would result from the atmospheric oxidation of those primary alkene emission fluxes.

The quantification of secondary CH₂O formation derived above has uncertainties that are difficult to quantify. The result may be an overestimate, as a fraction of ethene and propene may not react before leaving HGB, but instead may be an underestimate because the contribution from oxidation of emissions of heavier alkenes, alkanes and aromatics is neglected. Wert et al. (2003) present an analysis of the "CH₂O production potential" of the individual VOCs measured in specific atmospheric samples. This CH₂O production potential gives the total rate at which CH₂O is formed from all measured VOCs during oxidation by OH radicals. For the eight most concentrated (i.e., least photochemically processed) VOC samples collected over industrial regions in HGB, the terminal alkenes, largely ethene and propene, on average, composed 95 % of total CH₂O production potential. However, this percentage gives an instantaneous picture of CH₂O formation early in the oxidation of the plume, while the total CH₂O

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production derived above is an integration over the time that the emissions remain in the HGB region. It must be noted that NO₃ and O₃ also are important oxidants of alkenes heavier than ethene (Brown et al., 2011); however these oxidation pathways are less important than OH, and they also produce CH₂O in significant yields. In 5 summary, it appears likely that the above quantification of secondary production is an underestimate for daytime emissions, when contributions from heavier alkenes, alkanes and aromatics are neglected, but an overestimate for nighttime emissions when chemical processing is slower, and some fraction of the emissions can be transported out of HGB before reacting. There are also uncertainties in the CH₂O yield from the oxidation of the alkenes. Overall, since some of the uncertainties likely compensate for others, a conservative estimate for the uncertainty of the quantification of the rate of secondary CH₂O formation is judged to be ±40%. This value is reflected in the uncertainties indicated in Table 1.

3.2 Direct measurement of the primary formaldehyde flux from petrochemical facilities

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The most direct measurement of the primary flux of CH₂O from industrial facilities in HGB is that reported by Mellqvist et al. (2010b) and Johansson et al. (2010), who deployed a mobile van just downwind of specific industrial areas to measure emission fluxes in the plumes from the facilities. Table 1 presents a summary of measurements conducted in 2009, which found relatively small fluxes of CH₂O immediately downwind of the industrial facilities. Mellqvist et al. (2010b) argue that these CH₂O fluxes represent mostly primary emissions, because the measurements were made so close to the facility that transport times were short enough that secondary formation was assumed to contribute little to the observed CH₂O fluxes.

Mellqvist et al. (2010b) and Johansson et al. (2010) present one flux measurement that allows our determination of the quantity of secondary CH₂O formation to be tested. On 20 May 2009 under easterly winds they measured the flux of alkenes and CH₂O in the coalesced plume from the HSC and Mont Belvieu areas during a transect on

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the west side of the HSC (see Fig. 58 of Mellqvist et al., 2010b). The transport time was sufficient (\sim 2–3 h from Mont Belvieu) for substantial photochemical production of CH $_2$ O to have proceeded. The measured CH $_2$ O flux was about 1200 kg h $^{-1}$, and the plume still had a significant flux of unreacted alkenes (e.g., 490 kg h $^{-1}$ ethene). When these unreacted alkenes do react, the ultimate total flux of CH $_2$ O is expected to be at least 1960 kg h $^{-1}$, which agrees to within 4 % with the combined 2040 kg h $^{-1}$ secondary source calculated by summing the separate contributions from HSC and Mont Belvieu in Table 1.

A comparison of primary and secondary CH_2O fluxes from the petrochemical facilities is included in Table 1. Summing over the three petrochemical industrial areas, $4\pm2\%$ of the CH_2O flux is of primary origin and $96\pm2\%$ is of secondary origin, produced during photochemical oxidation of primary alkene emissions. We take this relative primary-secondary partitioning to be characteristic of the entire petrochemical sources of CH_2O in HGB.

3.3 Sporadic formaldehyde emission events from petrochemical facilities

Olaguer et al. (2009) have focused attention on sporadic episodes in the HGB area characterized by very high reported concentrations of CH_2O up to 52 ppbv (Eom et al., 2008). They argue that direct primary emissions can possibly explain these high concentrations. Here we briefly discuss the expected signature of concentrations of trace species within plumes of primary CH_2O emissions, and then examine two episodes that have received particular attention (Olaguer et al., 2009). The goal is to determine if secondary formation alone is adequate to explain the observed CH_2O concentrations, or if there is substantial evidence for significant sporadic episodes of primary CH_2O emissions.

A unique signature is expected for measurements made within a fresh plume of primary CH₂O emissions. Initially upon emission of primary CH₂O the enhanced CH₂O concentrations would not be accompanied by enhanced O₃ concentrations. In contrast,

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secondary production of CH₂O is generally accompanied by production of O₃. Plumes with significantly enhanced CH₂O concentrations without correlated O₃ concentration enhancements were not encountered in either of the two NOAA airborne field campaigns conducted during TexAQS 2000 (Wert et al., 2003; Ryerson et al., 2003) and TexAQS 2006 (Washenfelder, et al., 2010). Figure 2 shows the relationship between CH₂O and O₃ found on 27 August 2000, which was typical of that found in all the research flights conducted by NOAA during the two TexAQS studies. The number of coincident CH₂O and O₃ data points (14 031 10-s averages and 146 624 1-s averages in 2000 and 2006, respectively) represent over 14 000 km flight distance in each study from 14 days in 2000 and 12 days in 2006. Many individual plumes were examined during the analysis performed for publications based on these data (Wert et al., 2003; Ryerson et al., 2003; Washenfelder, et al., 2010). The TexAQS 2006 study included nighttime flights (Brown et al., 2009), when primary emissions of CH₂O would be particularly obvious, but evident plumes of primary CH₂O emissions were not encountered. If concentrated plumes (i.e., several ppbv enhancements) of fresh CH₂O primary emissions are present in the HGB region, they were not encountered in either of these aircraft studies.

It is, of course, impossible to prove that primary emissions never play a significant or even a dominant role in some isolated episodes. A plume of primary CH_2O emissions released in daytime would be expected to produce significant amounts of O_3 from the photochemical processing of CH_2O as long as sufficient NO_x is also present, so a plume of primary CH_2O emissions would soon lose its unique signature. However, it is possible to investigate if secondary formation alone is adequate to explain specific observed episodes. Here we examine two episodes that have received particular attention.

During a morning flight on 31 August 2006, the Baylor Aztec aircraft repeatedly sampled a plume over and downwind of the HSC. This plume contained CH_2O concentrations higher than the instrument could quantify ($\sim 9\,\mathrm{ppbv}$), as well as high concentrations of a variety of primary species and ozone (see Supplement and Fig. 8 of

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Olaguer et al., 2009). Examination of the original data set (Baylor University, 2009) demonstrates that this plume represented a very complicated air mass with separate parts of the plume showing markedly different ratios of the primary pollutants NO_v, CO and SO₂. It is also evident that relatively fresh emissions (i.e., those with a large fraction of NO_v still present as NO_x) were mixing with aged pollution, as indicated by high O₃ concentrations approaching 200 ppbv, which is the highest O₃ observed by the Baylor Aztec during 2006. The time resolution of the CH₂O instrument (~ 1 min) was not adequate to resolve the rapid concentration changes encountered by the aircraft. Hence, it is undetermined whether the high observed CH₂O concentrations were associated with the fresh emissions or the aged pollution. It is apparent however, that the observed high O₃ concentrations are consistent with very high concentrations of secondary CH₂O; for example Wert et al. (2003) report CH₂O > 30 ppbv in a plume with O₃ ~ 150 ppbv. Thus, the measurements reported by the Baylor Aztec in the 31 August 2006 plume do not provide strong evidence for primary emissions of CH₂O as the main source of this plume. Rivera et al. (2010) report the flux of CH₂O from the HSC on this same day, and conclude that its source was predominately secondary production from VOC emissions within HSC.

Eom et al. (2008) report the observation of a CH₂O plume during the morning of 27 September 2006 at the Lynchburg Ferry USEPA site in Baytown, TX. This plume reached a maximum concentration of 52 ppbv, which is reportedly the maximum ambient concentration of CH₂O ever observed in the HGB region. There was no conclusive evidence for the source of this CH₂O. Based upon poor correlation with O₃ and other arguments, the authors argue that primary CH₂O emissions may have played a role. A definitive examination of the sources of CH₂O in this (or any other) plume requires consideration of the recent transport of the sampled air parcel. Meteorological analyses (see Supplement) indicate that the air from the HGB region on 26 September was transported south over Galveston Bay and returned to the HGB area at the time that the 27 September plume was observed. The stagnation and recirculation transport pattern of this plume is ideal for accumulation of high CH₂O concentrations from

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secondary processing of the HRVOC emissions from the HSC. Until the transport and chemical processing that occurred in this plume are understood in detail, no definitive assignment of the source of CH₂O in this plume is possible. In summary, no strong evidence has been presented for episodes of sporadic CH₂O primary emissions from the petrochemical facilities in the HGB region.

4 Formaldehyde fluxes from on-road vehicles in HGB

In this section, we quantify the fluxes of primary CH₂O emissions from on-road vehicles in the HGB region, and estimate the rate of secondary formation of CH₂O during the atmospheric oxidation of the alkenes emitted by these vehicles.

4.1 Determination of the primary emission flux from on-road vehicles

To estimate the flux of primary CH₂O from on-road vehicle emissions, we multiply the CH₂O to CO emission ratio deduced from field observations in Houston by the total CO emission rate from on-road vehicles in HGB. This latter quantity is available from emission inventories constrained by ambient measurements. The CH₂O to CO emission ratio is quantified from the relationship between the concentrations of these two species observed during the morning traffic peak. This time period is selected because traffic related sources can dominate the ambient CH₂O concentrations, and the loss of CH₂O from the atmosphere is minimized because OH levels are suppressed by high NO_x concentrations and photolysis is still slow. The predominant source of CO in HGB is on-road vehicle emissions, so the ambient enhancement ratio of CH₂O to CO is not affected by dilution. In the following, all emission ratios are expressed as molar ratios, not mass ratios.

A preliminary analysis prepared for the TexAQS II Rapid Science Synthesis (Cowling et al., 2007) estimated that the primary emissions of $\rm CH_2O$ from mobile sources were, as an upper limit, 0.18 to 0.30 % of the CO emissions. This estimate was based upon nighttime measurements made on the NOAA research vessel Ronald H. Brown

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and WP-3D aircraft (see Fig. E2 of Cowling et al., 2007). This estimate was deemed an upper limit, due to the possibility that the sampled air had been photochemically processed to at least some extent during the preceding daytime period, or that some fraction of the observed formaldehyde had been produced from nighttime secondary production through O_3 or NO_3 reaction with primary VOCs. These findings are broadly consistent with previous determinations of the CH_2O to CO emission ratios of ~ 0.2 to 0.3% in Los Angeles (Grosjean, 1982), 0.10 to 0.14% in Denver, Colorado (Anderson et al., 1996), and 0.24% in Rome (Possanzini et al., 1996).

Rappenglück et al. (2010) report CH₂O and CO measured at Moody Tower in Houston, Texas as part of the TRAMP study (Lefer and Rappenglück, 2010). As shown in Fig. 4, the relationship between the concentrations of these two species measured at all times of day is not well represented by a single linear correlation. Thus, sources other than direct emissions from the on-road vehicle fleet must be important. The large open circle and dotted black line in Fig. 4 show the CH₂O–CO relationship expected if background air from the Central Gulf of Mexico with 80 ppbv CO and 0.5 ppbv CH₂O (Gilman et al., 2009) were transported into HGB and impacted only by on-road vehicle emissions with a CH₂O to CO emission ratio of 0.3%. Virtually none of the Moody Tower data lies on this reference line, but it does define the lower envelope of the observed CH₂O as a function of CO.

To obtain the best estimate for the CH₂O to CO emission ratio for on-road vehicles from the Moody Tower data set, we examine the correlation between these two species in the period before and during the morning traffic peak on individual days. The time window on each day is generally selected to include a pre-sunrise CO minimum, which represents the background air on that specific day to which the traffic emissions are added, and extend to the morning CO maximum. Only days with substantial CO enhancements (selected as peak CO exceeding 480 ppbv) are included in this evaluation. The color-coded points in Fig. 4 identify the 13 days during the TRAMP measurements when both CH₂O and CO data were collected during the morning traffic peak, and the peak CO exceeded 480 ppbv.

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Only one (18 September) of the 13 days with strong morning CO enhancements closely approximates the reference line in Fig. 4. That day was nearly ideal for evaluating the on-road vehicle emission ratio. During the entire preceding day (a Sunday) the wind remained southerly (171 ± 19°; average ± standard deviation) and brisk $(4.5 \pm 1.3 \,\mathrm{m\,s}^{-1})$. These winds brought relatively clean marine air to the Moody Tower site; for example, between midnight and 1:00 a.m. local standard time on 18 September, $O_3 = 9.6 \pm 0.2$ ppbv, $CO = 93 \pm 2$ ppbv, $NO_v = 2.4 \pm 0.2$ ppbv, and $CH_2O =$ 0.84 ± 0.04 ppbv. Between midnight and 6:00 a.m. the wind decreased in speed and rotated through westerly to northerly. By 6:00 a.m., winds were nearly calm allowing traffic emissions to accumulate in the resulting stagnant air. Since the petrochemical facilities lie generally east of the Moody Tower, no industrial emissions are expected to have impacted the measurements under such wind conditions (see Rappenglück et al., 2010). This expectation is supported by the measured SO₂, which remained below 0.6 ppbv during the predawn period. Figure 5 shows the gradual increase in CO, NO_v and CH₂O during this time. (The Supplement gives similar plots for all 13 days.) From the predawn CO minimum to the morning maximum, CH₂O was well correlated with CO ($r^2 = 0.92$) with a linear regression slope of 0.0026 ± 0.0003 (average $\pm 95\%$ confidence limit). This linear fit is included in Fig. 4. Since little day-to-day variability is expected in the HGB on-road vehicle fleet (at least for weekdays), the best estimate for the CH₂O to CO emission ratio is 0.26±0.03%, which agrees with the 0.18–0.30% upper limit estimate of Cowling et al. (2007). The 0.26 ± 0.03 % estimate is also an upper limit, since secondary production of CH₂O from the VOCs co-emitted with CH₂O by on-road vehicles are mixed with the primary emissions, even though the meteorological conditions on 18 September limit the time that the vehicle emissions remained in the atmosphere before measurement.

The slopes derived from the linear regressions for all 13 days with strong morning CO enhancements vary widely, which reflects variability of the influence of other sources rather than variability in the vehicle fleet emissions themselves. Figure 4 shows the linear fits and Table 2 summarizes the slopes derived from those fits for all 13 days.

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Except for 18 September, the rush hour data all lie well above the reference line. This is attributed to transport of CH₂O to Moody Tower from other sources within the HGB area. The variability of the slopes is attributed to the degree of correlation or anticorrelation of transported plumes with the morning traffic. Figure 6 illustrates two days that exemplify high correlation and high anti-correlation. On 15 September strong correlation ($r^2 = 0.85$) between a transported plume with high CH₂O concentrations and the morning CO maximum resulted in a relatively large slope (0.0066 ± 0.0017) due to the transport enhanced CH₂O concentrations (compare upper panel of Fig. 6 with Fig. 5, which use the same concentration scales). In contrast, on 20 September, transported air with high CH₂O concentrations reached Moody Tower throughout the early morning period, with the peak arriving before the CO traffic peak, which resulted in a negative correlation with CO (r = -0.41) and a negative slope (-0.0035 ± 0.0030).

If we assume that, on average, CH₂O from other (non-vehicle) sources transported to Moody Tower is uncorrelated with the morning CO traffic peak, then the linear regression slopes derived for the morning traffic peaks averaged over a large number of days should provide a measure of the CH₂O to CO emission ratio for on-road vehicles alone. The weighted average (i.e., each day's slope weighted by the inverse of the square of its confidence limit, Bevington, 1969) of the regression slopes for all 13 days is 0.30±0.02%, which is in excellent agreement with the result above for 18 September and the estimate of Cowling et al. (2007).

A recent tunnel study (Ban-Weiss et al., 2008) suggests significantly lower CH₂O to CO emission ratio for on-road vehicles. Using 2006 measurements made in a San Francisco Bay Area highway tunnel, these workers derive molar ratios of 0.062 % and 0.149 % for light duty, gasoline fueled vehicles and medium duty/heavy duty diesel fueled trucks, respectively. Both of these results are significantly lower than the result from the 2006 ambient measurements presented here. The reason for the differences between the two studies is not well established, but it may reflect the specific driving conditions, the vehicle mix and the relative absence of cold starts in the tunnel. However, the tunnel study does suggest that the result from the present work likely

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overestimates rather than underestimates the CH₂O to CO emission ratio for on-road vehicles in HGB.

Quantification of formaldehyde formed from oxidation of on-road VOC emissions

Following a procedure similar to that of Sect. 3.1, the total amount of secondary CH₂O that can form within HGB from on-road vehicle emissions can be estimated from the product of the total ethene and propene emissions from vehicles times the product vield of CH₂O from these alkenes. Rather than relying upon emission inventories to provide total ethene and propene emissions, we use the measured alkene to CO emission ratios multiplied by total CO emissions. This latter quantity will be taken from emission inventories, since this aspect of inventories has been more extensively tested. The primary CH₂O emission flux determined in the preceding section is also based upon the total CO emissions, so any uncertainty in this quantity will not affect the determination of the relative amount of primary versus secondary CH₂O associated with vehicle emissions. In this section, we again neglect any unreacted ethene or propene and CH₂O produced from oxidation of alkane, aromatics, and heavier alkenes.

Warneke et al. (2007) have derived the emission ratios of ethene and propene to CO characteristic of urban emissions using ambient measurements near the US east coast. They find good agreement with the results of Baker et al. (2008), who analyzed measurements from 28 US cities. Both of these studies generally quantified the ratios from on-road vehicle emissions, since that is the primary source of alkenes and CO in most of these cities. Since the vehicle fleet and the hydrocarbon gasoline composition does not vary markedly among different regions of the US, the Warneke et al. (2009) results are taken to be representative of the HGB vehicle fleet. Table 3 gives these alkene to CO ratios, as well as the secondary CH₂O to CO ratio implied by these ratios combined with the product yields of CH₂O from these alkenes (Seinfeld and Pandis, 1998) discussed earlier.

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Table 3 includes the integration of the on-road vehicle emissions of CO, ethene, propene and CH₂O in the HGB region, which is defined here as latitude 28.9 to 30.6° N and longitude 94.4 to 96.2° W. The integration is performed on the NEI 2005 inventory provided by EPA. However, CO emissions in the NEI 2005 inventory, which is based upon the MOBILE6 emission model, exceeds measured CO concentrations by about a factor of 2 (Parrish, 2006; Brioude et al., 2011). Consequently, to obtain an accurate estimate we reduce the integrated CO emission estimate by half. The alkene and CH₂O to CO emission ratios then allow total emissions of the alkenes and CH₂O to be derived, which are included in Table 3 in the row labeled "best estimate". For all species except CO these "best estimate" emissions are in good agreement (±25%) with the integrated NEI 2005 emissions.

Here again, the estimate of the secondary CH₂O may be an overestimate, since some of the ethene and propene may be transported out of the HGB region before reacting to form CH₂O, but may be an underestimate as CH₂O produced from oxidation of alkane, aromatics, and heavier alkenes is not included. The emission ratio of the alkenes to CO are estimated as accurate to ±30 % (Warneke et al., 2009), which are taken as the uncertainties for the primary emissions of the alkenes, while the estimate for the uncertainty of the secondary CH₂O formation rate is taken as ±40 % for reasons similar to the arguments given in Sect. 3.1.

Table 3 summarizes the estimated primary CH₂O emitted and secondary CH₂O formed from the on-road vehicle fleet. The primary emission estimate is based upon the ambient CH₂O to CO ratio measured during the morning traffic peak, and hence is an upper limit. These results indicate that no more than 28±8% of the CH₂O from the on-road vehicle fleet in HGB is of primary origin, with the remainder, at least $72 \pm 8\%$, of secondary origin, produced from oxidation of alkenes also emitted by the on-road vehicles. This estimated apportionment is expected to approximately apply to all US urban areas.

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The first problem is that the correlation-based studies explicitly or implicitly addressed source contributions to measured ambient CH_2O concentrations at particular sites, while the present analysis addresses the total mass of CH_2O emitted and formed within the entire HGB region. It is the emission fluxes and production rates (expressed as mass or moles per unit time) that quantify the amount of CH_2O emitted or produced within HGB, and it is these quantities that determine the importance of CH_2O to the photochemical production of O_3 within HGB. It is critical to note that measured ambient concentrations at any particular location are affected not only by emission fluxes and production rates, but also by transport (including dilution) processes and loss rates. The relative contributions to measured ambient concentrations are directly related to the relative emission fluxes and production rates only if the loss rates and the effects of transport and dilution are identical for each of the sources. In the case of CH_2O , this

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direct relationship does not apply, because secondary sources are at a maximum rate

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during the daytime when dilution and photochemical loss rates are also at a maximum. The diurnal cycle of CH₂O in HGB provides an example of the potentially confounding effects of dilution and loss rates. Observed surface concentrations of CH₂O (Fig. 7a) exhibit a relatively modest daytime maximum, but those daytime concentrations are present throughout a deep mixed convective boundary layer (CBL). Nighttime concentrations average only a factor of 2 lower than the daytime maxima, but represent a much shallower mixed layer. After normalizing those observed concentrations for mixing height (Fig. 7b), the average daytime maximum is more than a factor of 10 higher than the average nighttime concentrations. In addition to the greater dilution of formaldehyde during the day, the lifetime of CH₂O (3 to 4 h in full sun, Seinfeld and Pandis, 1998) is relatively short during the day, but much longer at night. Thus, CH₂O from any particular source would accumulate to higher concentrations at night than

The preceding discussion indicates that CH₂O from different sources is expected to experience a wide spectrum of loss rates and transport effects depending upon the diurnal dependence of the source strength. Hence, any analysis that aims to determine the relative importance of different sources must account for these confounding effects. In the present work, careful attention is given to ensure comparison between sources on the basis of total mass of formaldehyde emitted or produced, not directly on observed concentrations. Figure 1 of Buzcu Guven and Olaguer (2011) shows that the source factors derived from correlation analyses can have very strong diurnal variation. Such analyses based solely upon concentrations without accounting for varying transport and loss rates are expected to err substantially.

during the day, even if the emission rates and dilution effects remained constant.

A second major problem with the three earlier studies is that they are based on multivariate correlation approaches, and interpretation of the results required assumptions regarding the cause of the correlations; however, the hypothesized causes are incorrect in important respects. First, all three studies take CO and two of the studies (Rappenglück et al., 2010; Buzcu Guven and Olaguer, 2011) take SO₂ as markers for

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primary emissions of CH₂O. They also assume that O₃ (Friedfeld et al., 2002) or PAN (Rappenglück et al., 2010; Buzcu Guven and Olaguer, 2011) is a reliable marker for secondary production of CH₂O. They then further assume that any correlation of CH₂O with CO or SO₂ indicates primary emission, and that only correlation of CH₂O with O₃ 5 or PAN can indicate secondary production. However, none of the studies presents analysis to support these assumptions; in effect they assume that correlation proves cause. They neglect to consider that ambient CH₂O concentrations may well correlate with ambient concentrations of CO from mobile source emissions and SO₂ from industrial emissions because those same sources also emit large quantities of reactive VOCs that form secondary formaldehyde. None of the three studies presents any evidence regarding the actual source of the formaldehyde that correlates with the primary emission tracers.

The TexAQS 2000 aircraft data discussed above in Sect. 3 (Figs. 1–3) can illuminate the dominant cause of the correlation of CH₂O with SO₂. The 27 and 28 August flights sampled the plume from HSC under similar meteorological conditions. Figure 8 shows the CH₂O vs. SO₂ correlation for those two flights with the measurements divided into relatively fresh emissions (grey points) and the more aged plume (red points). The fresh emissions have a weak correlation, a small CH₂O to SO₂ ratio and small CH₂O concentrations compared to more aged emissions sampled further downwind. The stronger correlation of CH₂O with SO₂ and increasing CH₂O concentrations that appear during transport is the signature of secondary formation of CH₂O in a plume initially rich in SO₂. The correlation grows into the plume through photochemical processing during plume transport. Co-located primary emissions of CH₂O and SO₂ would have the strongest correlation and highest concentrations of both species closest to the source. Hence, correlation of CH₂O with SO₂ without a detailed analysis of the cause of the correlation cannot be taken as indicative of primary emissions of CH₂O.

Stutz et al. (2011) utilized a mobile laboratory during the 2009 field study to investigate CH₂O plumes downwind from industrial facilities. They investigated the spatial extent of the plumes and evaluated them in the context of the prevailing wind to identify

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the sources of observed plumes. A single primary CH_2O source was identified, which was in the Texas City area with an emission rate of $\sim 25\,\mathrm{kg}\,\mathrm{h}^{-1}$, corresponding to the total Texas City primary CH_2O emissions measured by Mellqvist et al. (2010b) (Table 1). In this plume the CH_2O/SO_2 ratio was 0.07–0.12, much smaller than the 0.4–1.3 ratio found downwind of the HSC (Fig. 8), which again indicates that the CH_2O downwind of HSC is of secondary origin.

Similar considerations apply to correlations of CH₂O with CO. Vehicle emissions of CO and VOCs, including alkenes, accumulate together in urban air masses. Photochemical processing produces CH₂O, which leads to significant correlations of ambient concentrations of CH₂O and CO. Figure 9 illustrates the development of this correlation observed in the 27 and 28 August flights. As the air moves downwind, increased concentrations of both CO (from accumulation of emissions) and CH₂O (from accumulation of photochemical production) are observed. There is significant correlation of CH₂O with CO (r = 0.76 for all data in Fig. 9), with higher correlations and different slopes observed downwind of HSC (r = 0.87, red to orange points in Fig. 9) and downwind of the central urban area (r = 0.83, green to purple points in Fig. 9). Importantly, nearly all of the observed CH₂O is due to secondary production, as the ratio of CH₂O to CO in primary emissions from vehicles (black dotted line in Fig. 9; see Sect. 4.1) is a factor of 15 to 30 smaller than the observed CH₂O vs. CO slopes. In summary, it is incorrect to assume that correlations of CH₂O with either SO₂ or CO necessarily indicate primary emissions of CH₂O.

Similarly, neither O_3 nor PAN can necessarily be taken as a tracer for secondary CH_2O formation without firm analysis to justify that assumption. Further, the correlation coefficient and slope between these species and CH_2O vary significantly depending upon the precursor mix and degree of processing. The formation of both O_3 and PAN requires both VOCs and NO_x to be present. The photochemical processing of an emitted plume with large amounts of reactive VOCs without NO_x would be expected to form copious amounts of secondary CH_2O , but little or no O_3 or PAN. Alternatively, the photochemical processing of a plume with large primary emissions of both CH_2O

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and NO_x would be expected to form large amounts of O₃, but any remaining unreacted CH2O that correlated with that O3 would be considered secondary. Figure 2 shows an example of the variability of the CH₂O correlation with O₃ within the HSC plume (east of -95.5° longitude). Downwind of the Houston central urban area (west ₅ of -95.5° longitude), the CH₂O correlation with O₃ is significant (r = 0.72) but with a much smaller slope (0.07 ppbv CH₂O/ppbv O₃) than observed downwind in the HSC plume (as large as 0.15 ppbv CH₂O/ppbv O₃). The coincident CH₂O and PAN data from the 27 and 28 August flights are much more limited, but variability in correlation coefficient and slope between these two species is also apparent. For example, downwind of the Houston central urban area, the CH₂O vs. PAN correlation coefficient is 0.91 with a slope of 2.7 ppbv CH₂O/ppbv PAN; the corresponding values downwind of HSC are 0.77 with a slope of 4.4 ppbv CH₂O/ppbv PAN.

In summary, the correlations between ambient concentrations CH₂O and those of primary pollutants (e.g., SO₂ and CO) and other secondary products (e.g., O₃ and PAN) arise from complex atmospheric interactions, vary substantially depending upon the mix of precursors in and air mass, and are strongly affected by transport and loss processes. Consequently, source apportionment analyses based solely on correlations cannot be expected to be reliable. The problems with such approaches are expected to be particularly severe when attempting source apportionment analyses of secondary species such as CH₂O, since such a large number of processes are involved in determining the correlations between the atmospheric concentrations of various secondary and primary species.

Discussion and conclusions

We have evaluated the rates of secondary production and primary emission of CH₂O from petrochemical industrial facilities and on-road vehicles in Houston Texas region based upon ambient measurements made in the 2000-2009 period and a measurement constrained emission inventory based upon the EPA NEI 2005. This evaluation

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(summarized in Table 4) shows that by far the predominant source of CH₂O (92±4% of total) is secondary production formed during the atmospheric oxidation of the alkenes emitted from the petrochemical facilities that characterize the industrial activity in HGB. These same facilities also emit much smaller amounts of primary CH₂O (4±2% of total); these primary emissions (in contrast to the alkene emissions) are well predicted by current emission inventories. CH₂O from the on-road vehicle fleet (4±2% of total) is also dominated by the secondary CH₂O formed from the alkenes directly emitted by the vehicles. We quantified an upper limit for the amount of primary CH₂O emitted by this fleet; that amount is relatively small (28±8% of the vehicle total), and is well predicted by current emission inventories.

This evaluation indicates that there is no strong observational evidence for large primary CH₂O emissions beyond those presently included in emission inventories. There is also no need to hypothesize such emissions for models to adequately reproduce observed CH₂O or O₃ concentrations within HGB. Several studies (Wert et al., 2003; Jiang and Fast, 2004; Byun et al., 2007; Kim et al., 2011) have shown reasonable agreement with observations when the ethene and propene emissions are increased according to the results of measured emissions from the petrochemical facilities.

Since CH₂O is dominated by secondary production, there is no large fraction of CH₂O sources in HGB that can respond to direct, emission control efforts focused on primary CH₂O emissions, although the Texas City source (Stutz et al., 2011) discussed above could be controlled by a focused effort. Ongoing efforts to control HRVOC emissions from the petrochemical facilities and VOC emission controls on the motor vehicle fleet will effectively control secondary CH₂O formation in HGB.

We find no evidence that sporadic episodes of primary CH₂O emissions from the petrochemical industrial facilities make a significant contribution to CH₂O in HGB. Although we do not quantify other possible sources of primary emissions, such as off-road mobile sources, these are not expected to constitute major CH₂O emission sources in HGB. Secondary formation of CH₂O from biogenic VOCs, especially isoprene, has not been addressed, and air coming into the Houston area from forested

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regions to the north and east may contain a significant amount of secondary formaldehyde formed from isoprene. This biogenic secondary CH₂O could play a role in initiating the photochemical processing of the ozone precursors emitted in Houston.

The correlation-based analyses of Friedfeld et al. (2002), Rappenglück et al. (2010) 5 and Buzcu Guven and Olaguer (2011) reached conclusions in conflict with those presented here. However, those studies are flawed because (1) they analyze ambient concentrations, not the total quantity of CH₂O emitted or formed and do not account for differential dilution and loss processes between sources, and (2) they rely only on correlations without firmly establishing the causes of the correlations. Analyses presented here indicate that the assumed causes were in fact incorrect. Similar problems must be suspected in any correlation-based analyses of CH2O sources conducted in other urban areas (e.g., Li et al., 1997; Garcia et al., 2006). Indeed, all correlationbased source apportionment analyses of secondary species must be investigated for similar problems before their conclusions can be confidently accepted.

Supplementary material related to this article is available online at: http://www.atmos-chem-phys-discuss.net/11/32601/2011/ acpd-11-32601-2011-supplement.pdf.

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Table 1. Summary of the measured and inventory average primary emission fluxes and estimated secondary formation rate of CH₂O from petrochemical facilities in the HGB given as 24-h averages. Units are kg h⁻¹ except as noted.

Area	Primary Ethene	Primary Propene	Secondary CH ₂ O ¹	Primary CH ₂ O
HSC	500 ± 54^2	642 ± 108^2	1165 ± 490	45 ± 21^2
Mont Belvieu	444 ± 174^2	303 ± 189^2	871 ± 520	17 ± 7^2
Texas City	122 ± 41^2	54 ± 22^2	221 ± 120	22 ± 5^2
Total HGB	2550 ³	4250 ³	6550 ± 2620	317 ³
Total HGB ⁴	91	101	220 ± 90	10.6

¹ Estimated from product of the fluxes of ethene and propene multiplied by the CH₂O product yield of the respective alkene.

² Measured emissions (Mellqvist et al., 2010b).

³ Emission inventory (Kim et al., 2011) integrated over the HGB.

⁴ Units in kmole h⁻¹.

Table 2. Slopes derived from linear regressions of CH_2O vs. CO for the selected morning vehicle traffic peak periods during 2006. Data were collected at the Moody Tower site.

Date	Local Time	Slope ± C.L.*	r^2
20 Aug	04:20-07:20	0.0073 ± 0.0027	0.68
21 Aug	04:20-06:50	0.0017 ± 0.0005	0.77
23 Aug	03:30-07:40	0.0063 ± 0.0010	0.87
24 Aug	05:00-07:00	0.0020 ± 0.0027	0.21
2 Sep	04:10-07:10	0.0055 ± 0.0014	0.83
6 Sep	03:00-07:40	0.0025 ± 0.0007	0.68
7 Sep	03:50-06:50	0.0048 ± 0.0022	0.57
15 Sep	05:50-08:20	0.0066 ± 0.0017	0.85
18 Sep	04:10-08:10	0.0026 ± 0.0003	0.92
20 Sep	02:10-07:10	-0.0035 ± 0.0030	0.17
26 Sep	04:40-07:30	0.0053 ± 0.0009	0.92
27 Sep	04:20-07:50	0.0033 ± 0.0017	0.48
28 Sep	06:00-08:20	0.0036 ± 0.0015	0.71
Average		0.0030 ± 0.0002	

 $^{^{*}}$ C.L. = 95 % confidence limit of the slope.

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Table 3. Summary of emission fluxes of CO, ethene, propene and formaldehyde estimated for the HGB on-road vehicle fleet, given as 24-h averages. The indicated uncertainties are estimated 95 % confidence limits.

	Primary CO	Primary Ethene	Primary Propene	Secondary CH ₂ O	Primary CH ₂ O
X _i /CO ¹ HGB NEI 2005 ⁵ HGB best estimate ⁶	- 1684 842 ⁷	4.6 ± 1.4^{2} 4.0 3.8 ± 1.2^{8}	1.4 ± 0.4^{2} 1.1 1.15 ± 0.34^{8}	7.7 ± 2.3^{3} - 6.5 ± 2.6^{8}	3.0 ± 0.2^4 2.0 $2.5^{8.9}$

¹ Units: 10⁻³ mole/mole CO.

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² Emission ratio of alkene to CO derived from ambient measurements (Warneke et al., 2007).

³ Estimated from the sum of two terms, one for ethene and one for propene; each term is the product of the emission ratio of the alkene to CO, and the CH₂O yields of the respective alkene.

⁴ Emission ratio of CH₂O to CO from ambient measurements at Moody Tower in 2006 – see discussion in text.

⁵ NEI 2005 inventory integrated over HGB. Units: kmole h⁻¹ on average summer weekday.

⁶ Units: kmole h⁻¹ on average summer weekday.

⁷ Taken as 50 % of NEI 2005 integration – see discussion in text.

⁸ Product of X_i/CO and primary CO emissions.

⁹ As discussed in the text, this is an upper limit; no confidence limit is indicated.

Table 4. Summary of the rates of secondary production and primary emission of CH_2O in HGB given as 24-h averages. Units of absolute rates are kmol h⁻¹ and uncertainties of primary emissions are estimated as $\pm 30\%$.

Source	Secondary	Primary
Point sources	220 ± 90 (92 %)	10.6 (4 %)
On-road vehicles	$6.5\pm2.6\ (3\%)$	2.5 (1%)
Total	227 ± 90	13.1
Percent total	$95\pm3\%$	$5\pm3\%$

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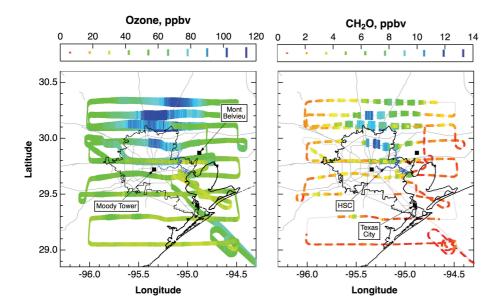


Fig. 1. Distributions of ozone (left) and formaldehyde (right) downwind of the HSC measured by the Electra aircraft during TexAQS 2000. The data were collected between 12:00 and 18:00 local standard time, and are plotted on the 27 August 2000 flight track, with the symbols sized and color-coded according to the measured mixing ratios of the respective species as indicated by the keys above each plot. During this flight, measured winds were southerly (wind direction = $162 \pm 17^{\circ}$) and steady (wind speed = $5.4 \pm 1.5 \,\mathrm{m \, s}^{-1}$), where standard deviations of the respective quantities are indicated. Text boxes with arrows indicate approximate locations of specific petrochemical complexes and a measurement site referred to in the text.

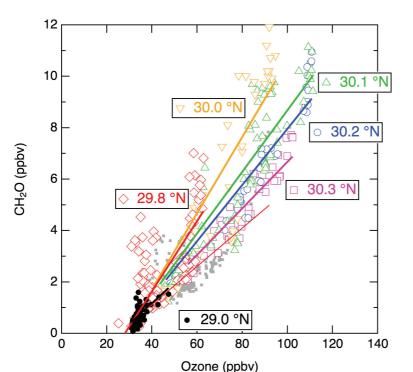


Fig. 2. Relationship of formaldehyde versus ozone mixing ratios measured during the 27 August 2000 flight. The data collected at one upwind (29.0° N latitude) and five downwind transects from HSC (east of -95.5° longitude) are shown by different symbols color-coded according to latitude as indicated in the annotations. All other data are shown as grey dots. Linear least squares fits to the data from each transect are shown also color-coded. These fits all pass through the background mixing ratios of O_3 (31.7 ppbv) and CH_2O (0.5 ppbv) as explained in text.

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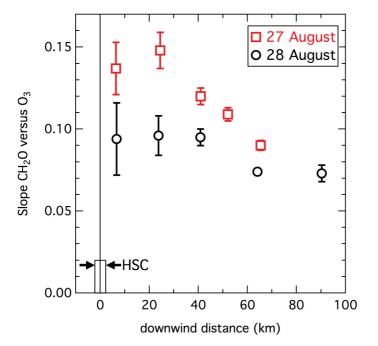


Fig. 3. Dependence of the slope with 95 % confidence limits of the ${\rm CH_2O}$ versus ${\rm O_3}$ relationship as a function of downwind distance from HSC. The 27 August 2000 data are from the linear regressions illustrated in Fig. 2; the 28 August 2000 data are from a similar analysis of a second flight conducted under similar meteorological conditions. The bar with arrows indicates the location and approximate width of the HSC industrial region. The farthest downwind transect corresponds to about 6 h transport time.

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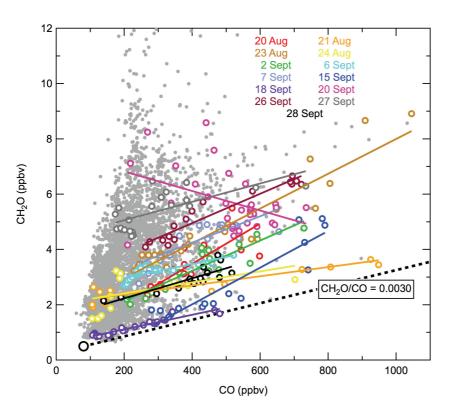


Fig. 4. Relationship between CH_2O and CO observed at Moody Tower during TRAMP. Gray points include all data with CH_2O mixing ratios ≤ 12 ppbv. Small circles color-coded by date indicate the morning traffic peak data discussed further in the text. The solid colored lines indicate the linear, least-squares fit to the respective color-coded data. The large black circle indicates the Central Gulf of Mexico mixing ratios reported by Gilman et al. (2009), and the heavy, dotted black line indicates the expected mixing ratio enhancements from primary emissions of CH_2O and CO in a ratio of 0.3%.

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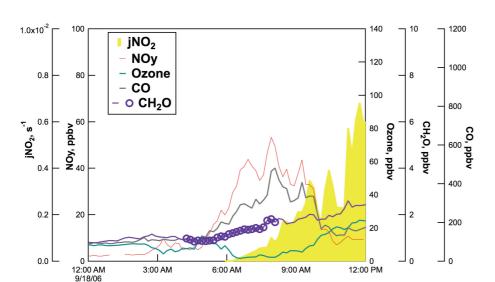


Fig. 5. Time series of the photolysis rate of NO_2 and the mixing ratios of NO_y , O_3 , CH_2O and CO observed during the morning of 18 September 2006 at Moody Tower. Small circles indicate the CH_2O data for that day included in the linear regression illustrated in Fig. 4. Time is given as local standard time (CST).

Time, CST

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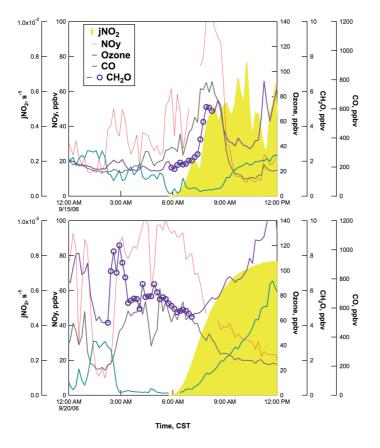


Fig. 6. Time series observed during the mornings of 15 and 20 September 2006 at Moody Tower in the same format as Fig. 5. Small circles indicate the CH₂O data for those days included in the linear regressions illustrated in Fig. 4.

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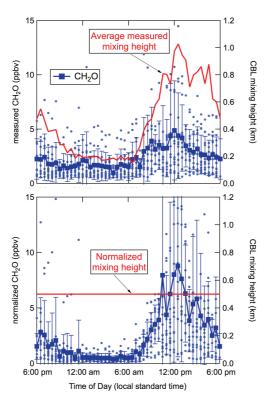


Fig. 7. CH₂O concentrations and mixing heights measured aboard the NOAA research vessel Ronald H. Brown during TexAQS 2006 within the HGB area. **(a)** The light blue points include all 30-min averages recorded during the study, and the dark blue symbols indicate averages and standard deviations for 30 min diurnal periods. The red line indicates average mixing height (i.e., CBL depth) **(b)** The calculated CH₂O concentrations expected if the integrated column concentration in **(a)** were uniformly mixed to a constant mixing height of 500 m (after Gilman et al., 2009).

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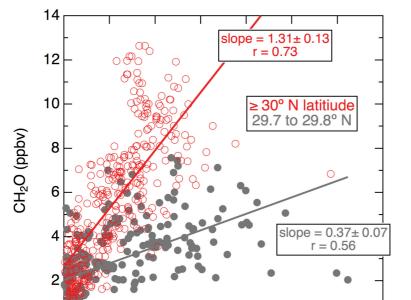


Fig. 8. Relationships between CH₂O and SO₂ measured by the Electra on 27 and 28 August 2000 within the plume from HSC (taken as east of –95.5° longitude to avoid plume from Parish power plant that moves over the western part of the city). The track for the first flight is shown in Fig. 1, and the second flight track was similar. Data are color-coded according to whether they were collected directly over HSC and immediately downwind (grey points, 29.7–29.8° N) or further downwind (red points, 30.0–30.3° N). The lines and annotations of the respective colors indicate the linear regressions to the data sets.

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SO₂ (ppbv)

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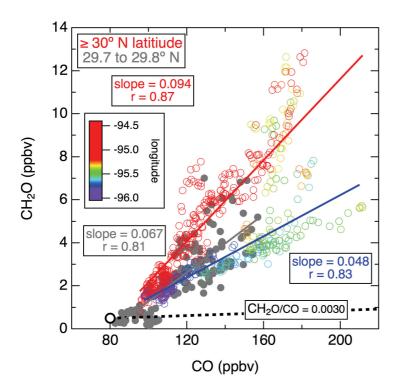


Fig. 9. Relationships between CH₂O and CO measured by the Electra on 27 and 28 August 2000 in the same format as Fig. 8. The track for the first flight is shown in Fig. 1, and the second flight track was similar. Data from within the entire plume downwind from the Houston area are included. Data are color-coding according latitude range (grey points, 29.7-29.8° N; colored points, 30.0-30.3° N) and longitude according to color-scale in plot. The lines and annotations of the respective colors indicate the linear regressions to the data sets divided by latitude range and longitude (red east and blue west of -95.4° longitude). The dotted black line indicates the expected mixing ratio enhancements from primary emissions of CH₂O and CO from the on-road vehicle fleet with a ratio of 0.3%.

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