

**Atmospheric aerosol  
compositions in  
China**

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# Atmospheric aerosol compositions in China: spatial/temporal variability, chemical signature, regional haze distribution and comparisons with global aerosols

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## Abstract

During 2006 and 2007, the daily concentrations of major water-soluble constituents, mineral aerosol, organic carbon (OC) and elemental carbon (EC) in ambient PM<sub>10</sub> samples were investigated from 16 urban, rural and remote sites in various regions of China, and were compared with global aerosol. A large difference between urban and rural chemical species was found, normally with 1.5 to 2.5 factors higher in urban than in rural sites. Optically-scattering aerosols such as sulfate (~16%), OC (~15%), nitrate (~7%) and ammonium (~5%) consist of ~50% of the total aerosols with another ~35% from mineral aerosol also having a certain degree of scattering ability, indicating a dominant scattering feature of aerosols in China. Of the total OC, ~55%–60% can be attributed to the secondary organic carbon (SOC). The absorbing aerosol EC accounts for ~3.5% of the total PM<sub>10</sub>. Seasonally, maximum concentrations of most aerosol species are found in winter while mineral aerosol also peaks in spring. Second peaks were found for sulfate and ammonium in summer and for OC and EC in May and June. This can be considered as a typical seasonal pattern in various aerosol components in China. Aerosol acidity is normally neutral in most of urban areas, but becomes somewhat acidic in rural areas. Based on the surface visibility from 681 meteorological stations in China during 1957–2005, four major haze areas are also identified with similar visibility changes, namely, (1) Hua Bei Plain in N. China, plus the Guanzhong Plain; (2) E. China with the main body in the Yangtze River Delta area; (3) S. China with most areas of Guangdong and the Pearl River Delta area; (4) The Si Chuan Basin in S. W. China. The degradation of visibility in these areas is linked with the emission changes and high PM concentrations. Such quantitative chemical characterization of aerosols is essential in assessing their role in atmospheric chemistry and weather-climate effects, and in validating atmospheric models.

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## 1 Introduction

Tropospheric aerosols are highly variable in time and space from US (Malm and Schichtel, 2004; Tanner et al., 2004), Europe (Querol et al., 2008; Querol et al., 2004; Yin and Harrison, 2008), S. E. Asia (Kim et al., 2007; Oanha et al., 2006), S. Asia (Kumar et al., 2007; Rastogi and Sarin, 2005; Stone et al., 2010) to China (Zhang et al., 2008a; He et al., 2001; Hu et al., 2002; Yao et al., 2002; Ye et al., 2003; Zhang et al., 1993; Zhang et al., 2002), and contain optically-scattering sulfate, organic carbon (OC), nitrate, ammonium product, and mineral aerosol in most circumstance, as well as optically-absorbing materials, mainly BC with partial contributions from mineral aerosol (Yang et al., 2009; Zhang et al., 2008b). High loadings of these aerosols induce a forcing for climate change (Forster et al., 2007; Myhre, 2009) in addition to having an adverse effect on weather (Wang et al., 2010; Pérez et al., 2006) and air quality for human health (Delfino et al., 2003; Feng et al., 2007). The uncertainties of the impacts are recognized to be very large, especially when very limited understanding of the regional and global distributions of individual aerosol species is available. This is also further complicated by their mixing status and the influence of aerosols on cloud physics when CCN (cloud condensation nucleus) activation capacity of mineral aerosol is enhanced through heterogeneous chemical reaction with polluted gases (Kumar et al., 2010; Koehler et al., 2009; Levin et al., 1996), and water-soluble organic aerosol (WSOA) involves into the CCN activation (Decesari et al., 2003; Facchini et al., 1999; Pöschl et al., 2010).

Various aerosols and their precursors are released by human activities in every part of the world especially in recent decades, but are attracting special attentions in areas of rapid economic growth and high population density, such as in China. Because of China's vast territory, high aerosol loading and complex compositions, many new phenomena and issues occurred, e.g., the interaction between aerosol and clouds in addition to the traditional cloud physics. Comprehensive understanding of aerosol chemical compositions in China, distributions, variations and signatures is one of the

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important bases to characterize the aerosols in the world and reduce uncertainty of their weather-climate effects, and also is a foundation for further understanding the aerosols-cloud interaction. Unfortunately, although there were some investigations on various aerosol chemical compositions in China (Zhang et al., 2008a; He et al., 2001; Hu et al., 2002; Wang et al., 2005; Yao et al., 2002; Ye et al., 2003; Zhang et al., 1993; Zhang et al., 2009b), the earlier studies were limited in scope and time with very few having data exceeding one year and covering various urban, rural and remote regions of the country.

This paper presents a data set on major water-soluble constituents, mineral aerosol and carbonaceous species (OC and EC) in daily ambient PM<sub>10</sub> samples measured during 2006 and 2007 at 16 stations of the CMA Atmosphere Watch Network (CAWNET), covering seasonal variations on a longer basis, which are essential to accurately and objectively assess the impacts of various aerosols in China and their influences at regional and global scales. Visibility changes in recent decades in China and the exhibiting distributions of regional haze in China are also presented.

The purpose of this paper is to assess chemical concentration levels of the urban, regional background, and remote aerosol, and to obtain the seasonal variations at various locations in the haze areas. Through the analysis of aerosol acidities, OC/EC ratios and the secondary organic carbon (SOC) component, the chemical signatures are also identified. These observations and analyses provide a general picture of aerosols in China and can also be used in validating model results.

## 2 Sample collection and analyses

### 2.1 Site and sampling descriptions

24-hr aerosol filter samples were collected at 16 CAWNET stations on a one day in every three day basis during 2006 and 2007. These CAWNET stations are operated by the Chinese Meteorological Administration (CMA), as shown in Fig. 1 and described

in details in Table 1. The rural stations are located in representative areas with the sampling point distant from local sources and raised above the surrounding ground level. At the urban locations, the sampling points are usually 50–100 m above the city average elevation in order to receive mixed aerosols rather than those dominated by specific sources. Aerosol samples were collected by using MiniVol™ air sampler (Airmetrics, Oregon USA), operating at an ambient air flow rate of 5 L min<sup>-1</sup> for 24 h from 09:00 a.m. to 09:00 a.m. (BST) the following day. The filtration media were 47 mm Whatman quartz microfibre filters (QM/A) that were cleaned by heating at 800 °C for 3 h before using.

## 2.2 Aerosol chemical composition analyses

The elemental concentrations were analyzed directly by an X-Ray Fluorescence (XRF) method using X-Lab 2000 (SPECTRO, Germany) at the Key Laboratory for Atmospheric Chemistry (LAC), Chinese Academy of Meteorological Sciences (CAMS). Eighteen elements such as As, Br, Ca, Cr, Cu, Fe, K, Mg, Mn, Na, Ni, Pb, S, Se, Ti, V, Zn and Zr were analyzed by the Multi-Channel Analyzer that was calibrated using standard plastic sample to check the signal energy and the multi-channel twice a month. The standard calibration curve between the intensity of signal and the elemental concentration is determined by single-element Mylar filters with certain concentration. For different kind of aerosol filter, the data of blank filter is also used to form the standard calibration curve.

After XRF analysis, a portion of filter samples was allocated for the determinations of SO<sub>4</sub><sup>2-</sup>, NO<sub>3</sub><sup>-</sup>, NH<sub>4</sub><sup>+</sup> and other ionic species, including Na<sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, F<sup>-</sup> and Cl<sup>-</sup> by ion chromatography (Dionex 600 series). The details protocol can be found in reference (Zhang et al., 2002).

Thermal/optical reflectance carbon (TOR) analysis method for EC and OC content was performed following the Interagency Monitoring of Protected Visual Environments (IMPROVE) protocol (Chow et al., 1993, 2004). The detailed procedures at (LAC, CAMS) can be found in reference (Zhang et al., 2005b, 2008a).

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### 3 Results and discussions

All these investigations enable the analysis of the spatial and temporal distributions of aerosols in China. The analyses will also help to identify the source attributions and atmospheric processes that influence the air quality in China, especially the visibility, or haze problems. This session will present the general pictures of atmospheric aerosols in China, put them into the perspective of global aerosol distributions, and finally link them with the long term visibility observations in China.

#### 3.1 Aerosol chemical compositions in various areas of China and comparisons with global aerosol

The annual averaged daily chemical data sets from the 16 stations in China are listed in Table 2. Generally, mineral aerosol (including sand dust, urban fugitive dust and coal ash) is the largest aerosol component in China, accounting ~35 % of PM<sub>10</sub> (Fig. 2 and Table 2) and even reaching from 50 % to 60 % in northwestern China. The annual median concentration of ~55 μg m<sup>-3</sup> for mineral aerosol alone (Fig. 2) is almost equivalent to the sum of all kind of aerosols in urban Europe (discussed in the later section), showing the dynamical and substantial contributions of mineral aerosol to the relative high regional background of aerosol mass in China. The high mineral fraction is partially a result of Asian desert sources (Zhang et al., 1996, 2003) which are not far from most of northern part of China. On the other hands, some of high mineral concentrations are evidently caused by emissions from urban fugitive dust/fly ash sources, especially during non-spring period. For example, Ca is mainly associated with mineral aerosol in the PM<sub>10</sub> samples; but non-crustal Ca is found in the coarse particle fraction, most of which is expected to be from construction activities and coal-ash (Zhang et al., 2002).

Sulfate and OC are other two major components with light-scattering properties, with each accounting for ~15 % of total PM<sub>10</sub> while nitrate and ammonium is about 7 % and 5 %, respectively. But in the Qinghai-Xizang (Tibetan) Plateau and the desert in the northwestern areas, sulfate accounted for only about 3%–6 % of the PM<sub>10</sub>, while OC

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in the proportion of these regions did not drop. The absorbing EC accounts for only 3.5% (Fig. 2). The sum of concentrations of other water-soluble inorganic cations (called “other WSIC” in Table 2), including  $\text{Ca}^{2+}$ ,  $\text{K}^+$ ,  $\text{Mg}^{2+}$  and  $\text{Na}^+$ , are also listed. Trace elements (Ni, Cu, Zn, As, Se, Br, Sr, Zr and Pb) and other WSIC account for

~0.5% and ~5% of  $\text{PM}_{10}$ , respectively. Referring to Fig. 3, the large difference between urban and rural aerosols are found for mass concentrations of various major chemical species, especially for  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and EC with urban to rural ratios (called “urban impact” thereafter) of 2.1, 2.2, 1.9 and 2.6, respectively, exhibiting the major sources of these constituents are mainly located in urban areas. Probably due to the long-rang transport of the Asian dust, the urban to rural ratio for mineral aerosol is only ~ 1.5. The urban impact of OC, ~1.6, is also relatively low, probably due to more biomass burning sources in rural area, and the existence of large SOC that was favored in regionally-dispersed aerosol (discussed in details later). The mass concentrations of rural  $\text{Ca}^{2+}$ ,  $\text{Na}^+$ ,  $\text{Mg}^{2+}$  are not much lower than urban values, showing a considerable number of soluble cations in rural aerosols, although their mass concentrations are not high.

In Fig. 4, the annual or seasonal mean mass concentrations ( $\mu\text{g m}^{-3}$ ) of six major types of aerosols in  $\text{PM}_{10}$  are synthesized with at least one entire year data from various rural and urban sites in 16 regions of the world, to compare the aerosol chemical compositions of China with those of global aerosols.

*Mineral Aerosol* generally is the largest component in most areas of the world (Fig. 4). The highest mass concentration levels are found in urban areas of S. Asia, HBP (Hua Bei Plain) in N. China and SCB (Si Chuan Basin) in S. W. China with mean value about 80–85  $\mu\text{g m}^{-3}$ . The rural area in N. W. China also has relative higher mineral concentrations of around 85  $\mu\text{g m}^{-3}$ . Mineral concentrations decrease in the urban areas of S. China, YRD (Yangzi River Delta), and rural areas of E. and N. E. China, High Asian Area (higher than 3500 m a.s.l) where concentrations are spatially uniform ranging from 20 to 40  $\mu\text{g m}^{-3}$ , exhibiting the relative higher regional background of mineral mass in whole Asian regions. Urban areas of S. E. Asia, US and Europe have

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lower mean mineral concentration of  $\sim 10\text{--}20\ \mu\text{g m}^{-3}$ . In rural Europe and US the lowest concentrations, generally below  $10\ \mu\text{g m}^{-3}$ , are found with few deserts and good vegetation cover.

*Sulfate*, as the second largest fraction of aerosol mass in some areas of the world (Fig. 4), finds the highest concentrations in China, generally vary at two levels,  $34\text{--}40\ \mu\text{g m}^{-3}$  at hot-spots from SCB to HBP, and  $15\text{--}25\ \mu\text{g m}^{-3}$  at other areas of China, which is similar to the level of  $\sim 20\ \mu\text{g m}^{-3}$  in urban of S. Asia, slightly higher than the S. E. Asian level of  $\sim 10\ \mu\text{g m}^{-3}$ . Large amount of coal consumptions in China, not only from the industrial activities but also from residential heating in many relatively small cities and towns, is the main reason for the higher sulfate mass observed there, which is consistent with the large  $\text{SO}_2$  emissions, e.g. north and east China (Cao et al., 2010; Lu et al., 2010). Of the total global  $\text{SO}_2$  emissions,  $\sim 100\text{--}120\ \text{Tg}$ , in 2000–2010, emission from coal-power plants account for  $\sim 47\%$ . The major sources of current global emissions are also considered to be in Asia, especially in China, India and other emerging industry countries, where the total emission is around  $47.0\ \text{Tg}$  in 2006 (Zhang et al., 2009a). Europe and America have somewhat lower emissions. This is why the relative lower sulfate concentrations are found between  $1.0\text{--}4.0\ \mu\text{g m}^{-3}$  from Europe to most of areas of US.

*Organic Carbon* has its highest mass concentrations in urban S. Asia with mean value of  $\sim 66\ \mu\text{g m}^{-3}$ . In China, the OC mass concentrations are at level of  $33\text{--}36\ \mu\text{g m}^{-3}$  from urban of SCB and HBP to  $\sim 15\text{--}20\ \mu\text{g m}^{-3}$  in the other regions. Unusually high OC,  $\sim 25\ \mu\text{g m}^{-3}$ , is found in rural northwest China, urban S. Asia and High Asian Area, showing the possible influences of open biomass burning. Of the total global OC emissions, about 74% is attributable to the biomass burning (Bond et al., 2004). On the basis of 2003 data, Asian OC emissions is estimated to be  $9.2\ \text{Tg C}$ , of which China and India contribute  $\sim 2.6$  and  $\sim 2.2\ \text{Tg C}$ , respectively (Ohara et al., 2007). In addition, secondary organic carbon (SOC) also contributes to the total OC (Miyazaki et al., 2006; Sullivan et al., 2006; Weber et al., 2007; Zhang et al., 2005b, 2008a). The estimated uncertainties of global annual SOC production are large, with a production rate ranging

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from 2.5 to 44.5 Tg OM per year from biogenic volatile organic compounds (VOC) and 0.05 to 2.6 Tg OM per year from anthropogenic VOC (Tsigaridis and Kanakidou, 2003). OC concentrations are relatively lower in US and Europe decreasing from urban mean  $\sim 6.0 \mu\text{g m}^{-3}$  to rural level of  $\sim 1.0\text{--}2.5 \mu\text{g m}^{-3}$ .

*Nitrate* usually is the fourth largest aerosol species in most regions of the world (Fig. 4). Not surprisingly, nitrate tends to be highest around urban HBP, SCB in China, plus urban S. Asia with mean concentrations from  $13\text{--}18 \mu\text{g m}^{-3}$ , where ammonium and nitrogen oxide emissions are greatest (Cao et al., 2010; Ohara et al., 2007; Streets et al., 2003; Zhang et al., 2009a). The global emission inventory suggests the major sources for nitrogen oxide are from fossil fuel combustion, in which the emissions from coal- power plant and motor-vehicle are  $\sim 22\%$  and  $41\%$ , respectively. The total global emissions of 2000–2010 is estimated to be  $\sim 80\text{--}110$  Tg, a slight increasing over the previous century<sup>11</sup>. The world's largest emissions are reported from China,  $\sim 20.8$  Tg in 2006 (Zhang et al., 2009), followed by the United States,  $\sim 14.9$  Tg in 2005. Outside hot-spots area, nitrate mean concentrations are between  $\sim 6.0\text{--}10 \mu\text{g m}^{-3}$  in other parts of China. Urban US, Europe and S. E. Asia also have similar or lower level of  $\sim 5.0 \mu\text{g m}^{-3}$ , while areas in High Asian Area, rural Europe and US have the averaged nitrate concentrations of around  $1.0\text{--}2.0 \mu\text{g m}^{-3}$ .

*Ammonium* is still mainly from the emission of coal-combustion both from industry and domestic consumptions, especially in China. While domestic animals, use of synthetic N fertilizers, biomass burning, human activities and pets are other major contributors to ammonium (Cao et al., 2010). Like sulfate, the more coal consumptions from a region, the higher the ammonium are found. The mean concentrations generally vary at a level,  $12\text{--}14 \mu\text{g m}^{-3}$  in urban S. Asia, HBP, SCB of China, and at another level of  $4.0\text{--}7.0 \mu\text{g m}^{-3}$  in other areas of China and urban S. E. Asia, which is bit higher than the levels of  $2\text{--}3 \mu\text{g m}^{-3}$  in urban US and Europe,  $0.2\text{--}1.3$  in rural US, Europe and High Asian Area (Fig. 4).

<sup>11</sup> Emission Database for Global Atmospheric Research, [www.mnp.nl/edgar](http://www.mnp.nl/edgar); The Greenhouse Gas and Air Pollution Interactions and Synergies, <http://gains.iiasa.ac.at>.

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*Elemental Carbon* is also called as EC, black carbon (BC) or light absorbing carbon (LAC) is the smallest proportion of major aerosol components in most parts of the world. But in urban S. Asia, High Asian Area and urban Europe, EC concentrations are higher than ammonium ion that is usually the second smallest aerosol component.

The EC has its highest level of mass concentrations,  $\sim 10 \mu\text{g m}^{-3}$ , in urban S. Asia, HBP and SCB of China, decrease to  $\sim 3.0\text{--}6.0 \mu\text{g m}^{-3}$  in other parts of China, urban S. E. Asia, and to  $\sim 3.0 \mu\text{g m}^{-3}$  in urban US and Europe, and decline further to  $\sim 0.2\text{--}2.0 \mu\text{g m}^{-3}$  from High Asian Area to rural Europe and US. Currently, Asia and Africa are considered to be two major sources of EC emissions accounting for  $\sim 39\%$  and  $\sim 25\%$  of total emission, respectively (Bond et al., 2004).

### 3.2 Seasonal patterns of aerosol chemical species in China

The spatial variability in the seasonal patterns of each of the major aerosol types is explored (Fig. 5). In general, the bulk  $\text{PM}_{10}$  mass concentrations tend to be higher during the winter half year from the autumn to winter (14 of the 16 CAWNET sites) and lower in summer, having a second peak in spring (11 of the 16). In wintertime large amount of coal consumptions, especially from residential heating (Cao et al., 2010; Lu et al., 2010), is the main reason for the highest sulfate and other chemical species observed in winter. Spring sand and dust storm (SDS) contributes a lot to tropospheric aerosol. In China the emissions of OC, VOC, BC, CO,  $\text{NH}_3$ ,  $\text{NO}_x$  and PM have two peaks in May to June and October after wheat and food crops harvests, respectively, associated with the increase of biomass burning, fertilizer and other agricultural activities (Fig. 6). This is the reason why the bulk aerosol mass show a general seasonal variation pattern with a peak in winter, dropping and then increasing in spring, reaching a minimum in summer and then increasing in autumn. The following sites had winter maxima: Longfengshan, Zhengzhou, XiAn, Jinsha, Nanning, Akdala, Gaolanshan and Gucheng. Autumn maxima found in Jinsha, Panyu and Dunhuang. Spring maxima occurred in LinAn, Taiyangshan, Lasha, Dunhuang. But in the SCB (Chengdu) there is no obvious seasonal variation found. This is also true for Lasha in High Asian Area

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(Fig. 5). The seasonal variation for each component of Chinese aerosols is depicted as follows:

*Mineral Aerosol:* Monthly mean concentrations of mineral aerosol over various CAWNET sites (Fig. 5) and its changes averaged for urban and rural site data (Fig. 6) illustrate that the relative higher concentration levels are found in March, April and May of Spring, showing the influences of Asian SDS (Zhang et al., 2003; Gong et al., 2006). The mineral concentrations then decline from June to August, and rise from September to December to the level similar to Spring's, especially for urban mineral aerosol, showing the contribution from the urban coal-ash and other fugitive dust during heating season from mid-autumn to winter. The rural mean concentrations of mineral aerosol change moderately, but remain similar to the urban ones.

*Organic Carbon:* The monthly mean OC at urban and rural sites varies similarly (Fig. 6), with the highest level in winter, continuously declining in the thereafter months, increasing their mean concentration in May to June, dropping again and remaining the minimum level during summer, and then gradually rising the concentration from Oct. to winter. This monthly changing pattern not only reflects the maximum winter emissions from fossil fuel combustions but also reveals the peak emissions associated with the agricultural activities from May to June and October in China. It can be seen that the seasonal variation of OC is most similar to the total emission pattern over China (Fig. 6) (Cao et al., 2006).

*Elemental Carbon:* The time series of monthly concentrations for EC are generally similar with OC both for urban and rural data (Fig. 6), exhibiting a typical monthly and seasonal pattern like OC, showing a common source for these two components.

*Sulfate:* Generally, the changing pattern of monthly sulfate is similar to that of carbonaceous aerosols (Fig. 5), but the biggest difference appears in summer with a second peak. The monthly sulfate, especially at urban sites, increases its concentration from June and reaches its second peak in July (Fig. 6). This is probably due to plenty of  $\text{SO}_4^{2-}$  and gas-phase  $\text{NH}_3$  in urban area, enhanced photochemistry, low air mass ventilation on a regional scale. The less varied change of rural sulfate shows again

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that the urban impact does not greatly extend over the rural areas in China.

Comparing urban and rural aerosol species, overall, the relative higher (Fig. 4) monthly mean concentrations of various aerosol species at urban sites varied similarly with those at rural sites (Fig. 6). But the differences are enlarged normally during winter for most of the aerosol species, exhibiting urban emission impact in winter. As for sulfate and ammonium, however, the significantly increased gap between urban and rural mean are found in July, showing the summer influence on these secondary aerosol formation, especially for sulfate and ammonium product (Fig. 6).

*Ammonium:* The monthly changes of ammonium for urban and rural average are similar to sulfate but not to nitrate (Fig. 6), showing the evidences to support the neutralization of acidic sulfate is favored over the formation of ammonium nitrate.

*Nitrate:* Unlike sulfate and ammonium which go up in July, nitrate and OC share a similar monthly changing pattern (Figs. 5 and 6), especially in urban area, indicating a common motor-vehicle sources influences for them. After ~300,000 heavy-emission vehicles were banned during 1 to 19 July of 2008 before the Beijing 08 Olympic Game (8–24 August of 2008), organics and nitrate concentrations were reduced by 25–40%, while the sulfate and ammonium basically changed very little. After further cutting the traffic by half from 20 July to 7 August at Beijing area, the concentration of organics and nitrate particles further decreased ~6–20% with little changes in sulfate and ammonium (Zhang et al., 2009b). This suggests that the motor vehicle-related fossil fuel combustion is still the dominant source of nitrate and OC at Chinese urban cities. Because the neutralization of acidic sulfate is favored over the formation of ammonium nitrate, less ammonium nitrate can be formed due to insufficient  $\text{NH}_3$  left in summer. That is why no second peak was observed for nitrate in summer as for sulfate and ammonium. Generally, ammonium nitrate is formed in areas characterized by high  $\text{NH}_3$  and  $\text{HNO}_3$  concentrations and low temperatures (Pathak et al., 2004).

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### 3.3 Aerosol chemical signatures

#### 3.3.1 Aerosol acidities

The calculated ion balance for urban and rural aerosols show that in most cases the imbalance between cations and anions (denoted  $\Delta C$ ) are less than several percent of the ionic budget (Fig. 3). The ion balance calculation is based on the following equations for anions (1) and for cations (2):

$$\text{Anion equivalence}(\mu\text{mole m}^{-3}) = [F^{-}] + [Cl^{-}] + [NO_3^{-}] + 2[SO_4^{2-}] \quad (1)$$

$$\text{Cation equivalence}(\mu\text{mole m}^{-3}) = [Na^{+}] + [NH_4^{+}] + [K^{+}] + 2[Ca^{2+}] + 2[Mg^{2+}] \quad (2)$$

After taking into account experimental errors, the imbalance between cations and anions ( $\Delta C$ ) and the ionic budget are expressed in micro-equivalents per cubic meter in Eqs. (3) and (4), respectively.

$$\Delta C = [Na^{+}] + [NH_4^{+}] + [K^{+}] + 2[Ca^{2+}] + 2[Mg^{2+}] - [F^{-}] - [Cl^{-}] - [NO_3^{-}] - 2[SO_4^{2-}] \quad (3)$$

$$S = [Na^{+}] + [NH_4^{+}] + [K^{+}] + 2[Ca^{2+}] + 2[Mg^{2+}] + [F^{-}] + [Cl^{-}] + [NO_3^{-}] + 2[SO_4^{2-}] \quad (4)$$

Among the cations, the most abundant mass is found to be  $NH_4^{+}$  with annual mean concentration of 11.9 and 5.6  $\mu\text{g m}^{-3}$  in urban and rural areas, respectively (Fig. 3), followed by  $Ca^{2+}$ , with a few urban and rural difference. But the mass concentrations of  $K^{+}$ ,  $Mg^{2+}$  and  $Na^{+}$  are not high. Of the anions,  $SO_4^{2-}$  and  $NO_3^{-}$  are the dominant species contributing to the aerosol acidities, both of which correlate highly with  $NH_4^{+}$  and poorly with  $Ca^{2+}$  and  $K^{+}$ , especially in urban area (Fig. 7) while the role of  $Cl^{-}$  is minor, with less correlation with  $NH_4^{+}$ . Furthermore, the relatively higher correlations are found between  $SO_4^{2-}$ ,  $NO_3^{-}$  and  $NH_4^{+}$  in urban than in rural areas. The correlations between  $SO_4^{2-}$  and  $NH_4^{+}$  are higher than between  $NO_3^{-}$  and  $NH_4^{+}$  (Fig. 7), indicating again that ammonium sulfate is probably the predominant ammonium product in aerosols both for urban and rural areas in China.

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$\text{NH}_4^+$  mass concentration are much higher in urban than in rural areas, which implies that the major sources of  $\text{NH}_4^+$  are located in urban areas. In China's three major economic development areas, Beijing-Tianjin Area, the Yangzi River Delta Area, and the Pearl River Delta Area, for instance, waste and coal consumption accounts for ~67 % to 85 % total  $\text{NH}_3$  emissions, of which people and poultry wastes, and coal-combustion from industry, commercial and residential usage are the major contributors mainly in urban areas (Cao et al., 2010). In the rural area, the major source of ammonia emissions is from the agricultural activities, which accounts for ~14 % to 28 % total  $\text{NH}_3$  emissions, with the nitrogen fertilizer as the major contributor (Cao et al., 2010). This is why one observed relative higher ammonium in urban relative to rural, and why the relatively higher correlations are found between  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and  $\text{NH}_4^+$  in urban than in rural areas.

Although the poor correlations between  $\text{SO}_4^{2-}$  and  $\text{Ca}^{2+}$  and  $\text{K}^+$  are found in urban areas, a relative higher correlation is found between and  $\text{K}^+$  in rural area with a very small slope (0.062) (Fig. 8), showing somewhat degree combination between  $\text{K}^+$  and  $\text{SO}_4^{2-}$  under the condition of low  $\text{NH}_3$  levels. Because of the very low concentration of  $\text{K}^+$  ( $\sim 1 \mu\text{g m}^{-3}$ ) in rural area, the impact to aerosol acidities resulted from depleting  $\text{K}^+$  by  $\text{SO}_4^{2-}$  can be negligible. As mentioned previously, the correlations between  $\text{SO}_4^{2-}$  and  $\text{Ca}^{2+}$  are poor both in urban and rural areas. However in rural area, especially in remote Duhuang (DH) and Lasha (LS) (Fig. 8), the relationship between  $\text{SO}_4^{2-}$  and  $\text{Ca}^{2+}$  is significant different from other rural sites, exhibiting relative stronger neutralization of acidic sulfate by  $\text{Ca}^{2+}$  in these remote sites with very low  $\text{SO}_4^{2-}$ ,  $\text{NH}_4^+$  mass and relative higher mass of  $\text{Ca}^{2+}$ . We take ionic specie ( $\text{CaSO}_4$ ) into aerosol acidity estimation (Nenes et al., 2011) for Dunhuang and Lasha in Table 2.

The comparisons between observed  $\text{NH}_4^+$  mass concentration (named " $\text{NH}_4^+_{\text{obs}}$ " in the right panel of Fig. 7) and the  $\text{NH}_4^+$  amount needed to fully neutralize the measured  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and  $\text{Cl}^-$  (called " $\text{NH}_4^+_{\text{neu}}$ ") are also presented in Fig. 7. The annual median values of these data at each site are also presented in Table 2. The aerosol acidity is

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evaluated by comparing the ratio between  $\text{NH}_4^+_{\text{obs}}$  and  $\text{NH}_4^+_{\text{neu}}$ . Particles are “more acidic” if the observed ammonium is significantly lower (by 25 % and beyond, Zhang et al., 2005a) than the predicted values, and are neutralized if the two values are close. Normally most of the larger than “0.75” median values of aerosol acidity are found at urban sites, expect for XiAn and Dalian, and the lower values found at rural sites, except for Changde and Jinsha. The observed  $\text{NH}_4^+$  is generally about 84 % and 74 % of the neutralized  $\text{NH}_4^+$  mass, for urban and rural aerosol, respectively, suggesting somewhat acidic for rural aerosol and the existence of the extra  $\text{SO}_4^{2-}$  and other anion molecules in the particles in the form of bisulfate ( $\text{HSO}_4^{2-}$ ) or others in rural area. Referring to Table 2, the ratio of  $\text{NH}_4^+_{\text{obs}}$  and  $\text{NH}_4^+_{\text{neu}}$  at Xian is relative lower than the “normal” urban value, probably due to the relative higher  $\text{SO}_4^{2-}$  found in the Guanzhong Plain (XiAn) where the high coal-combustion emission and higher sulfate concentration were observed (Zhang et al., 2002). The  $\text{NH}_4^+$  concentration finds non-proportional low at Dalian, probably resulting in abnormal lower ratio of  $\text{NH}_4^+_{\text{obs}}$  and  $\text{NH}_4^+_{\text{neu}}$  found in Dalian. The extraordinary low ratio of  $\text{NH}_4^+_{\text{obs}}$  and  $\text{NH}_4^+_{\text{neu}}$  found in Dunhuang and Lhasa such relative remote sites with less anthropogenic activities, probably due to the very low  $\text{NH}_4^+$  concentrations measured around the detection limit of instrument.

In order to assess the overall spatial and temporal proportions of SOC in OC for better parameterizing the SOC in models, the monthly averaged SOC in total OC are presented in Table 3, associated with OC to EC ratios. According to Castro (Castro et al., 1999), the contribution of SOC is estimated from the following equation:

$$\text{SOC} = \text{OC}_{\text{total}} - \text{EC} \times (\text{OC}/\text{EC})_{\text{min}} \quad (5)$$

where SOC is the secondary OC,  $\text{OC}_{\text{total}}$  denotes the total OC, and  $(\text{OC}/\text{EC})_{\text{min}}$  is the observed minimum OC/EC ratio, which has a unique value identified for each site.

Consistent with the previous evaluation of OC and EC in 2006 from CAWNENT (Zhang et al., 2008a), the two year data from 2006 to 2007 also show a relative higher OC/EC value of  $\sim 5.5$  in rural aerosols, and a lower urban one of 3.5 that is similar to the ratios of emission inventories in China. The OC/EC ratios is 2.0 for fossil fuel combustion,

1.4 for gasoline and 3.0 for diesel oil, and 3.3 for biomass fuel combustion on the basis of the China emission inventory (Cao et al., 2006) . This reflects the domination of primary emissions to the urban carbonaceous aerosol, especially fossil and biomass fuel combustions.

5 The high OC/EC ratios of ~5.5 for rural aerosols are probably contributed from open biomass burning in fields, which has a calculated OC/EC ratio of 7.1 in emission inventory (Cao et al., 2006). Referring to the monthly changes of OC/EC in Fig. 10, the OC/EC ratios peak in June and October, especially for rural aerosols due to the increasing influence of primary emission and precursors in June and October as described in previous section (Cao et al., 2010). The high OC/EC ratios can also be contributed from SOC formation in rural areas (Zhang et al., 2008a) as a noticeable rise of the SOC was estimated in June and October (Fig. 6). The annual averaged fractions of SOC estimation show that the contributions from SOC to OC concentrations are ~56 % (52–64 %) for the urban aerosols and ~58 % (35–82 %) for the regionally representative aerosols (Table 3). Despite its importance, the formation processes of SOC are highly uncertain, and the lack of a quantitative understanding of secondary organic formation processes in the atmosphere is one of the most important weaknesses in the present understanding of atmospheric aerosol and its climate effect (Hallquist et al., 2009; Tsigaridis and Kanakidou, 2007). Although here is a very rough estimate for SOC with substantial uncertainty, it still can provide some insight into the SOC contributions to the total OC.

### 3.4 Regional haze distributions and classifications in China

As a result of the increasing aerosol concentrations in China, the air quality has been deteriorating for a couple of decades, associated especially with the increasing hazy weathers. However, there was no systematic monitoring network for aerosols in China to quantify the trend. As the visibility is a good indicator of aerosol loadings in the atmosphere, the analysis of the visibility can provide some estimates of the aerosol trends in China.

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Based on the 14:00 (Beijing Local Time) surface visibility data from 681 meteorological stations in China during 1957–2005, excluding the data for rainy, snowy, foggy, misty, and sand and dust storm days, the regional haze distributions and its trend every five years are illustrated in Fig. 9. With less anthropogenic influences before 1980's, few areas have mean surface visibility less than 15 km. Eastern China was the low visibility region with ~25–30 km annual mean visibility.

Based upon above visibility data, the temporal and spatial distribution of visibility are used to find the regions with similar visibility trends by rotated empirical orthogonal function (REOF) method (Fig. 10). The REOF can resolve the principal component of the visibility data on the basis of their temporal and spatial variations, and identify the distributions of high correlating area of each principal (called mode in Fig. 10). Nine regions with similar visibility trends were found with some sub-regions (Fig. 11a).

*Region I:* i.e. N. E. China, with two sub-regions: one is in the major grain producing areas (I<sub>(1)</sub>) mainly in Heilongjiang Province; another is in old industrial base (I<sub>(2)</sub>) mainly in Liaoning province.

*Region II:* i.e. Hua Bei Plain (HBP) in N. China, including one of the China's rapid economic development regions, Beijing-Tianjin area, as well as the provinces of Hebei, Shangdong and Henan. The Guanzhong Plain in the south of the Loess Plateau (II<sub>(2)</sub>) is also assigned into this region, mainly because of the similar visibility change.

*Region III:* i.e. E. China, including another rapid economy development region, the Yangtze River Delta area and the provinces of Hubei, Anhui, Jiangsu, Shanghai, Zhejiang, Hunan and Jiangxi.

*Region IV:* mainly including Fujian Province and southeast hilly area of China.

*Region V:* i.e. S. China, including the third major rapid economic development region, the Pearl River Delta area, and the provinces of Guangdong and Guangxi.

*Region VI:* i.e. S.W. China, including the Si Chuan Basin (SCB).

*Region VII:* i.e. southwest of the Yunnan-Guizhou Plateau.

*Region VIII:* i.e. High Asian Area (with averaged height ~3500 m a.s.l., including the main area of the Qinghai-Xizang (Tibetan) Plateau).

*Region IX*: i.e. N.W. desert areas, including Gobi and arid areas with several sub-regions.

With the changes in every five year mean visibilities from various regions in China (Fig. 9), and inferring from the spatial difference between the first five year since 2000's and the early of 1960's (Fig. 11b), the visibility in almost all regions in eastern China at recently years had ~7 to 15 km losses relative to these in the early of 1960's. The most significant visibility loss areas were found in Regions II and III (Fig. 11b), showing the more serious problems of regional haze in these regions. Region V and some small areas in northeast and northwest China also had the similar visibility loss rate (Fig. 11a). The total loss of visibility was about 10 km in eastern China since 1957 with a mean rate of decrease of about  $0.2 \text{ km yr}^{-1}$ . In western China, only half of the loss rate was found, probably due to less population density and relative lower economic growth (Fig. 11c). Because the visibility is highly attributable to the fine aerosol loading, today's haze is not a completely natural phenomenon. The regional atmosphere often shows turbidity gray, which was called the "Grey Haze" phenomenon in China as well.

Still referring to Fig. 11, relative higher population density areas in China (red series color regions in Fig. 11c) often correspond to large emissions of aerosol and its precursors, such as Region II, III, V and VI (Fig. 11d) where both relative higher  $\text{PM}_{10}$  levels (Table 2) and poor visibility were found (Figs. 9 and 11b). These four regions can be considered as four major regional haze areas in China (shadow areas marked in Fig. 11a) with high GDP growth and high aerosol emissions (Fig. 11d) mainly from industrial activities, residential and commercial sectors. The high emissions also superimposed with the contributions of topography and meteorological features form the haze problems there. Although the total emissions in Region I are higher than those in Regions V and VI (Fig. 11d), the overall visibility loss is much smaller than in Regions V and VI, which is probably due to its larger area and favorable meteorological conditions for pollutants to diffuse. Consequently, this region is not regarded as one of the major haze areas in China even though some local poor visibility was found such as in Liaoning province with high emissions.

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## 4 Summary

Two years (2006–2007) of aerosol chemical compositions at 16 urban, rural and remote sites as well as the regional haze distributions are investigated for a better understanding of the characterization of regional haze and atmospheric aerosols, their distributions, variations and chemical signatures in China.

The major findings of this study can be summarized as follows:

- Aerosol composition data show that  $PM_{10}$  in almost all CAWNET sites, except at the two remote sites of Akdala and Zhuzhang, is characterized by high levels of mineral aerosol, which is partially a result of Asian desert dust transport, and contributions from urban fugitive dust/fly ash sources. The annual mean concentrations of  $\sim 75$  and  $55 \mu g m^{-3}$  for mineral aerosol in urban and rural areas, respectively, are almost equivalent to the sum of all kind of aerosols in urban Europe, showing the dynamical and substantial contributions of mineral aerosol to the relative high regional background of aerosol mass in China. This character is also true for aerosol in other Asian areas, which is the most significant difference between Asian and Europe, US aerosol compositions. Due to the high level of mineral aerosol in Asia and especially in China, in the future it is worth to pay particular concern on CCN activation capacity enhancement of mineral aerosol through heterogeneous chemical reaction with polluted gases in the region.
- In China, the total PM levels follow a clear seasonal pattern with a winter peak and summer minimum. A broader pattern is observed, especially at northwestern and northern China, with uplift in spring due to the higher frequency of Asian sand and dust episodes. No obvious seasonal variation was found for SCB, the large base in S. W. China, and Lhasa in High Asian Area, exhibiting relative independent variations and less influences from other areas.
- Similar to bulk PM, the relative higher concentration is found for mineral aerosol in spring that is a signal of Asian desert dust. A typical summer minimum and

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autumn to winter increasing pattern is also observed for urban mineral, showing the large influences of urban fugitive dust on the seasonal variation of this aerosol specie.

- As the second largest aerosol component in most areas of China, the relative higher sulfate concentrations are found with annual mean values of  $\sim 34$  and  $16 \mu\text{g m}^{-3}$  at urban and rural sites, respectively. The sulfate levels at urban sites increase progressively from June to July to exhibit an obvious uplift in mid-summer, due to plenty of  $\text{SO}_4^{2-}$  and gas-phase  $\text{NH}_3$  in urban area, enhanced photochemistry, low air mass ventilation on a regional scale. As a result, less ammonium nitrate can be formed due to insufficient  $\text{NH}_3$  left at urban area. That is why no second peak for nitrate in summer is found as for sulfate and ammonium product, showing the strong ability of sulfate to associate with ammonium. The annual mean nitrate levels are  $\sim 15$  and  $7.9 \mu\text{g m}^{-3}$  for urban and rural aerosol, respectively. Because of the increase of winter coal-combustion emission in both urban and rural areas as the largest source for nitrate precursors in the country, it is not difficult to understand why nitrate peaks are observed in winter associated with low temperature.
- Because of much higher  $\text{NH}_4^+$  mass concentration found in urban ( $\sim 12 \mu\text{g m}^{-3}$ ) than in rural areas ( $5.6 \mu\text{g m}^{-3}$ ), and the important role of  $\text{NH}_3$  in aerosol acidity, the aerosols in urban areas are nearly neutralized by having relative higher  $\text{NH}_4^+$ , leaving the aerosols in the rural area somewhat acidic, even where some  $\text{Ca}^{2-}$  and  $\text{K}^+$  can participate in the neutralization processes.
- Almost with the same level of sulfate, OC aerosol finds the annual mean concentrations of  $\sim 30$  and  $18 \mu\text{g m}^{-3}$  at urban and rural sites, respectively. EC generally vary at two levels,  $\sim 8.6 \mu\text{g m}^{-3}$  for urban hotspot, and  $\sim 3.4 \mu\text{g m}^{-3}$  for regional background. The OC/EC ratio is close to 8.7 at remote sites and 5.5 at rural ones, which can be explained by the biomass burning and the enhanced formation of SOC from both anthropogenic and biogenic precursors. The OC/EC ratio

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is around 3.5 for urban aerosol, showing the strong influence of local emission source. The SOC can contribute an annually averaged ~55 % and 60 % to the total urban and rural OC aerosols, respectively, and this source of OC enhanced around May to June and October, especially in regional background aerosol with increasing contribution from biomass burning emission in China. Accurate determination and better parameterization of SOC in numerical models is still subject to a great uncertainty, and is an important factor to accurately estimate the impact of aerosol scattering effects in modeling community.

- As to the global aerosols, the mineral, sulfate and OC fractions are also found to be responsible for most of the aerosol mass (less than  $PM_{10}$ ) at the majority of areas thought the world, followed by nitrate, ammonium and EC. Globally, the mass concentrations of these major aerosols in North America and Europe are relatively lower than those in S. E. Asia, S. Asia and China.
- Based on the similar variation trend of observed visibility, China can be classified into nine typical regions, four of which have seen the largest loss of visibility in recent decades, associated with higher  $PM_{10}$ , called major haze regions. These high PM and hazy regions are attributed to the proximity of emission sources, associated topography and meteorological features. Because the visibility is highly attributable to the fine aerosol loading, today's haze is not a completely natural phenomenon. The regional atmosphere often shows scattering and turbidity gray, which are called the “Grey Haze” phenomenon in China as well.

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**Table 1.** CMA Atmosphere Watch Network (CAWNET) site and sampling descriptions.

Station Name	CAWNET station description
Akdala (AKD) (47°06' N, 87°58' E)	Remote station, 562 m a.s.l., 55 km west of Fuhai county (Altai area, Xinjiang province, China), and ~250-300 km southeast of Kazakhstan. Aerosol samples collected from a 10 m-tall building.
Chengdu (CD) (30°39' N, 104°2.4' E)	Urban station, 496 m a.s.l., in the center of Chengdu city, Sichuan province. Aerosol samples collected from a 91 m-tall building.
Dalian (DL) (38°54' N, 121°37.8' E)	Urban station, 91.5 m a.s.l., southeast of Dalian center at city margin. Aerosol samples collected from a 5 m-tall building.
Dunhuang (DH) (40°9' N, 94°40.8' E)	Rural station, 1139 m a.s.l., 1.5 km northeast of Dunhuang city, Gansu province. Aerosol samples collected from a 5 m-tall sampling container.
Gaolanshan (GLS) (36°0' N, 105°51' E)	Rural station 5 km north from Lanzhou city (1531 m a.s.l.), Gansu province. Aerosol samples collected from a 544 m-tall hill.
Gucheng (GC) (39°7.8' N, 115°48' E)	Rural station but within area of rapid urbanization: 15.2 m a.s.l., 38 km southwest of Baoding city, Hebei province. Aerosol samples collected from an 8 m-tall building.
Jinsha (JS) (29°37.8' N, 114°12' E)	Rural station, 416 m a.s.l., 105 km north of Wuhan city, Hubei province. Aerosol samples collected from an 8 m-tall building.
Lhasa (LS) (29°40.2' N, 91°7.8' E)	Rural station in remote area, 3663 m a.s.l., within Lhasa city, Qinghai-Xizang Plateau. Aerosol samples collected from a several meter tall building.
LinAn (LA) (30°18' N, 119°44' E)	Rural station, 139 m a.s.l., 150 km northeast of Shanghai, and 50 km west of Hangzhou city (Zhejiang province, China). Aerosol samples collected from a 10 m-tall building.
Longfengshan (LFS) (44°43.8' N, 127°36' E)	Rural station, 331 m a.s.l., 175 km northeast of Harbin city. Aerosol samples collected from a 6 m-tall building.
Nanning (NJ) (22°49.2' N, 108°21' E)	Urban station, 84 m a.s.l., in Nanning city, Guangxi province. Aerosol samples collected from a 97 m-tall building.
Panyu (PY) (23°0' N, 113°21' E)	Urban station in Panyu district of Guangzhou city (5 m a.s.l.), Guangdong province. Aerosol samples collected from a 140 m-tall hill.
Shangri-La, Zhuzhang (SRL) (28°00' N, 99° 43' E)	Remote station, 3583 m a.s.l., 12 km northeast of Shangri-La county (Diqing area, Yunnan province, China)
Taiyangshan, Changde (TYS) (29°10.2' N, 111°42.6' E)	Rural station, 563 m a.s.l., 18 km northwest from Changde city, Hunan province. Aerosol samples collected from an 8 m-tall building.
Xian (XA) (34°25.8' N, 108°58.2' E)	Urban station in northern margin of Xian city, but within Jin River Industrial District in the Guanzhong Plain, 363 m a.s.l., 20 km north of center of Xian city, Shaanxi province. Aerosol samples collected from a 4 m-tall sampling container.
Zhengzhou (ZZ) (34°46.8' N, 113°40.8' E)	Urban station, 99 m a.s.l., in the center of Zhengzhou city, Henan province. Aerosol samples collected from a 56 m-tall building. Aerosol samples collected from a ~5 m-tall building. Aerosol samples collected from a 10 m-tall building.

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**Table 2.** Annual averaged daily concentrations ( $\mu\text{g m}^{-3}$ ) for  $\text{PM}_{10}$ , various aerosol compositions in various regions of China during 2006 and 2007.

Stations	Type	Region	$\text{PM}_{10}$	Mineral	$\text{SO}_4^{2-}$	OC	$\text{NO}_3^-$	$\text{NH}_4^+$	EC	$\text{Cl}^-$	Other WSIC <sup>a</sup>	Trace element <sup>b</sup>	Aerosol acidity <sup>c</sup>	No.
<i>Urban</i>														
Panyu (PY)	urban	Pearl River Delta economic region in S. China	147.4	45.0	26.8	22.3	11.7	8.6	7.9	1.9	7.8	1.0	0.89	139
Zhengzhou (ZZ)	urban	Province south of Beijing	271.3	104.1	45.0	29.2	22.6	16.5	9.2	4.6	15.8	1.1	0.85	157
Chengdu (CD)	urban	SI Chuan Basin	235.7	85.5	40.5	36.3	15.1	14.0	10.8	4.7	12.6	1.2	0.93	203
Gucheng (GC)	semi-urban	Province south of Beijing	253.8	83.0	35.5	38.5	20.0	15.4	11.0	7.2	10.1	2.6	0.80	202
XiAn (XA)	urban	Guanzhong Plain in central China	328.5	118.2	46.7	42.6	20.7	14.4	12.7	4.8	18.8	1.4	0.70	159
Nanning (NJ)	urban	City in S. China	98.9	23.9	21.6	17.9	5.1	5.8	4.0	1.1	6.9	0.46	0.90	189
Dalian (DL)	urban	Southern tip of Liaodong peninsula in N. E. China	156.8	68.0	23.3	20.2	13.5	7.7	5.3	2.6	9.2	0.66	0.67	195
<i>Regional</i>														
Taiyangshan, Changde (TYS)	rural	Dongting and Buoyan lakes basin in E. China	128.2	33.8	28.8	13.8	8.5	7.9	2.7	0.4	7.5	0.42	0.94	132
Longfengshan (LFS)	rural	Northeast China	81.8	19.3	10.0	15.9	4.9	2.5	2.3	0.7	5.6	0.27	0.56	195
Dunhuang (DH)	rural	North margin of Kumutage Desert in N. W. China	158.5	92.6	6.6	26.7	2.3	0.4	3.6	2.5	8.7	0.65	0.53 <sup>d</sup>	127
LinAn (LA)	rural	Yangzi River Delta economic regions in E. China	114.2	31.6	21.7	15.1	8.6	6.8	4.3	1.0	5.4	0.60	0.58	171
Jinsha (JS)	rural	Lower reach of Yangzi River in E. China	107.4	31.7	26.6	15.3	7.2	7.6	3.0	0.6	6.2	0.42	0.96	110
Lhasa (LS)	rural	Qinghai-Tibetan Plateau	77.4	39.0	2.9	21.7	2.3	0.2	3.8	1.3	6.7	0.42	0.56 <sup>e</sup>	162
Gaolanshan (GLS)	rural	Gansu province in N. W. China	187.1	97.3	16.7	19.1	18.4	6.5	3.8	3.2	11.0	0.81	0.52	199
<i>Background</i>														
Shangri-La, Zhuzhang (SRL) <sup>e</sup>	remote	Southwest tip of S. E. China	nd	0.85	1.6	3.1	0.45	0.15	0.34	nd	nd	nd	nd	74
Akdala (AKD) <sup>e</sup>	remote	Northwest tip of N. W. China	nd	4.7	3.3	2.9	0.38	0.60	0.35	nd	nd	nd	nd	56

<sup>a</sup> Other WSIC: other water-soluble inorganic cation, including  $\text{Ca}^{2+}$ ,  $\text{K}^+$ ,  $\text{Mg}^{2+}$  and  $\text{Na}^+$ ; <sup>b</sup> Trace element, including Ni, Cu, Zn, As, Se, Br, Sr, Zr, Pb; <sup>c</sup> Aerosol acidity: median value of the ratio between  $\text{NH}_4^+_{\text{obs}}$  (observed  $\text{NH}_4^+$  in  $\mu\text{mole m}^{-3}$ ) and  $\text{NH}_4^+_{\text{neu}}$  (the  $\text{NH}_4^+$  amount needed to fully neutralize the measured  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and  $\text{Cl}^-$ ); <sup>d</sup> Ionic specie ( $\text{CaSO}_4$ ) is also included in aerosol acidity estimation; <sup>e</sup> SRL and AKD data were during summer, 2004 to spring, 2005 from the reference (Qu et al., 2009)





**Table 3.** Monthly averages of SOC of total OC and ratio of OC and EC in various regions of China on the basis of 2006 and 2007 data.

Stations		1	2	3	4	5	6	7	8	9	10	11	12
<i>Urban</i>													
Panyu (PY)	SOC of OC	0.52 (±0.14)	0.50 (±0.15)	0.48 (±0.18)	0.53 (±0.12)	0.42 (±0.13)	0.52 (±0.07)	0.62 (±0.06)	0.53 (±0.13)	0.59 (±0.13)	0.55 (±0.12)	0.59 (±0.08)	0.62 (±0.09)
	OC/EC	4.1	4.0	3.7	3.9	3.7	4.8	2.4	2.9	2.7	3.0	3.3	4.0
Zhengzhou (ZZ)	SOC of total OC	0.49 (±0.11)	0.55 (±0.12)	0.56 (±0.12)	0.51 (±0.08)	0.60 (±0.12)	0.54 (±0.10)	0.42 (±0.10)	0.53 (±0.08)	0.45 (±0.12)	0.57 (±0.09)	0.58 (±0.12)	0.54 (±0.10)
	OC/EC	4.1	4.0	3.7	3.9	3.7	4.8	2.4	2.9	2.7	3.0	3.3	4.0
Chengdu (CD)	SOC of total OC	0.48 (±0.15)	0.57 (±0.13)	0.50 (±0.13)	0.54 (±0.10)	0.66 (±0.10)	0.52 (±0.12)	0.48 (±0.05)	0.51 (±0.12)	0.50 (±0.12)	0.54 (±0.09)	0.49 (±0.12)	0.49 (±0.10)
	OC/EC	4.1	4.0	3.7	3.9	3.7	4.8	2.4	2.9	2.7	3.0	3.3	4.0
Gucheng (GC)	SOC of total OC	0.65 (±0.10)	0.69 (±0.09)	0.68 (±0.13)	0.69 (±0.06)	0.68 (±0.08)	0.69 (±0.10)	0.58 (±0.11)	0.49 (±0.12)	0.50 (±0.08)	0.65 (±0.14)	0.69 (±0.07)	0.70 (±0.05)
	OC/EC	4.1	4.0	3.7	3.9	3.7	4.8	2.4	2.9	2.7	3.0	3.3	4.0
XiAn (XA)	SOC of total OC	0.61 (±0.10)	0.64 (±0.07)	0.56 (±0.14)	0.55 (±0.08)	0.54 (±0.09)	0.56 (±0.09)	0.41 (±0.10)	0.48 (±0.09)	0.47 (±0.12)	0.43 (±0.13)	0.54 (±0.10)	0.64 (±0.04)
	OC/EC	4.1	4.0	3.7	3.9	3.7	4.8	2.4	2.9	2.7	3.0	3.3	4.0
Nanning (NJ)	SOC of total OC	0.60 (±0.06)	0.63 (±0.09)	0.61 (±0.08)	0.61 (±0.09)	0.68 (±0.06)	0.67 (±0.06)	0.71 (±0.09)	0.56 (±0.10)	0.56 (±0.11)	0.58 (±0.09)	0.62 (±0.07)	0.59 (±0.05)
	OC/EC	4.3	4.7	4.4	4.6	5.8	5.5	5.9	4.0	4.0	4.3	4.3	4.4
Dalian (DL)	SOC of total OC	0.42 (±0.13)	0.49 (±0.15)	0.61 (±0.10)	0.68 (±0.07)	0.67 (±0.10)	0.64 (±0.13)	0.68 (±0.11)	0.53 (±0.14)	0.47 (±0.08)	0.50 (±0.09)	0.48 (±0.11)	0.44 (±0.06)
	OC/EC	2.9	3.3	4.4	6.1	5.0	4.4	5.5	3.4	3.3	3.4	3.2	3.1
<i>Regional</i>													
Taiyangshan, Changde (TYS)	SOC of total OC	0.50 (±0.13)	0.69 (±0.11)	0.51 (±0.11)	0.52 (±0.08)	0.58 (±0.14)	0.60 (±0.12)	0.57 (±0.11)	0.50 (±0.18)	0.49 (±0.13)	0.41 (±0.10)	0.46 (±0.10)	0.47 (±0.07)
	OC/EC	5.0	6.2	5.0	5.5	6.3	6.0	5.5	5.1	5.0	4.3	4.6	4.6
Longfengshan (LFS)	SOC of total OC	0.69 (±0.04)	0.69 (±0.07)	0.72 (±0.06)	0.79 (±0.05)	0.78 (±0.06)	0.82 (±0.07)	0.78 (±0.13)	0.67 (±0.06)	0.70 (±0.09)	0.74 (±0.08)	0.69 (±0.15)	0.61 (±0.08)
	OC/EC	6.3	6.0	7.2	9.3	8.8	10.7	10.3	6.1	6.8	8.4	6.8	5.0
Dunhuang (DH)	SOC of total OC	0.61 (±0.06)	0.68 (±0.13)	0.78 (±0.15)	0.71 (±0.15)	0.55 (±0.20)	0.63 (±0.15)	0.52 (±0.23)	0.43 (±0.23)	0.40 (±0.25)	0.55 (±0.17)	0.67 (±0.06)	0.54 (±0.20)
	OC/EC	7.7	8.6	10.3	8.3	6.7	7.1	4.8	4.1	4.5	6.6	9.5	7.9
LinAn (LA)	SOC of total OC	0.56 (±0.14)	0.61 (±0.08)	0.46 (±0.13)	0.59 (±0.07)	0.49 (±0.09)	0.54 (±0.15)	0.55 (±0.16)	0.48 (±0.11)	0.46 (±0.10)	0.45 (±0.10)	0.53 (±0.10)	0.47 (±0.09)
	OC/EC	3.6	3.9	3.3	4.0	3.7	3.7	3.7	3.4	3.7	3.2	3.8	3.5
Jinsha (JS)	SOC of total OC	0.39 (±0.11)	0.35 (±0.21)	0.48 (±0.15)	0.46 (±0.13)	0.57 (±0.16)	0.63 (±0.13)	0.55 (±0.11)	0.45 (±0.18)	0.41 (±0.17)	0.51 (±0.23)	0.44 (±0.09)	0.38 (±0.06)
	OC/EC	4.6	5.0	5.8	5.3	6.2	5.6	6.1	4.9	4.8	4.9	4.9	4.4
Lhasa (LS)	SOC of total OC	0.64 (±0.06)	0.69 (±0.07)	0.73 (±0.06)	0.73 (±0.07)	0.76 (±0.07)	0.70 (±0.06)	0.64 (±0.15)	0.63 (±0.10)	0.53 (±0.07)	0.57 (±0.07)	0.65 (±0.06)	0.67 (±0.08)
	OC/EC	5.5	5.6	6.9	5.9	7.3	5.4	5.3	5.7	5.6	5.2	5.7	5.3
Gaolanshan (GLS)	SOC of total OC	0.45 (±0.10)	0.53 (±0.13)	0.61 (±0.11)	0.71 (±0.12)	0.67 (±0.09)	0.67 (±0.09)	0.57 (±0.10)	0.50 (±0.09)	0.50 (±0.13)	0.50 (±0.10)	0.47 (±0.11)	0.46 (±0.09)
	OC/EC	4.6	5.0	5.8	5.3	6.2	5.6	6.1	4.9	4.8	4.9	4.9	4.4

\* Annual arithmetic mean value with standard deviation in parentheses.

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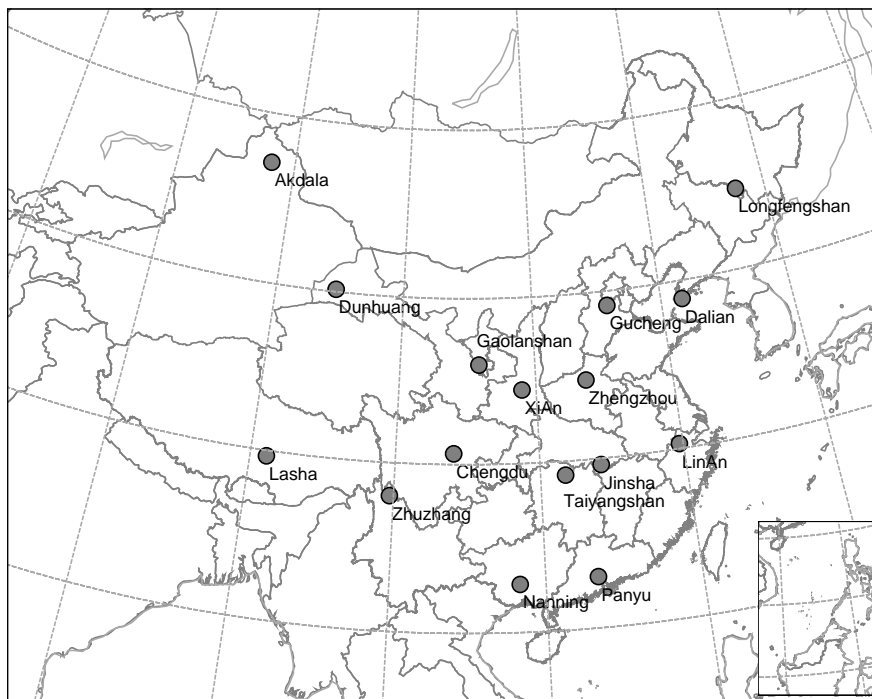
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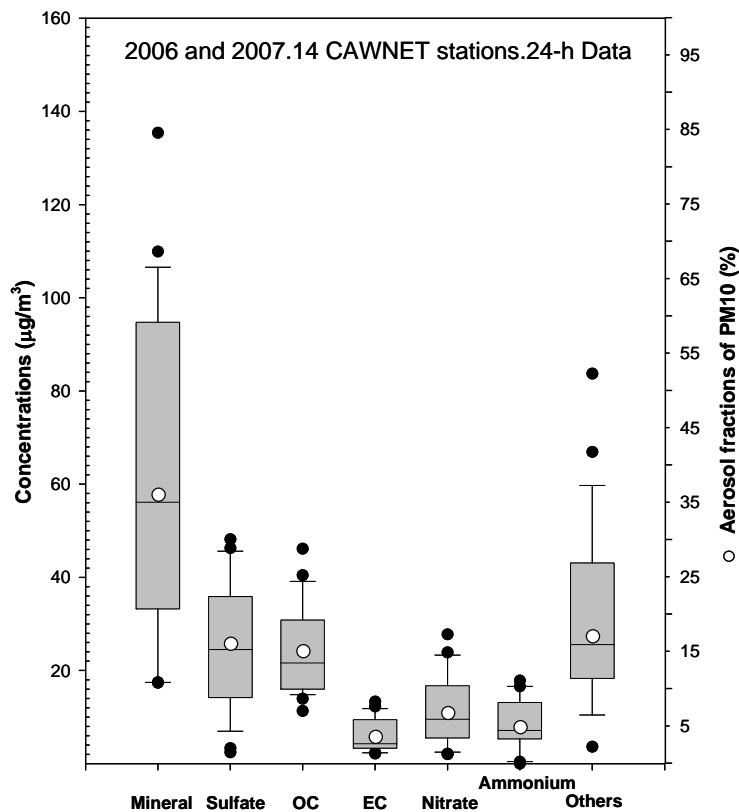
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**Fig. 1.** Locations of sixteen CAWNET stations.[Title Page](#)[Abstract](#)[Introduction](#)[Conclusions](#)[References](#)[Tables](#)[Figures](#)[◀](#)[▶](#)[◀](#)[▶](#)[Back](#)[Close](#)[Full Screen / Esc](#)[Printer-friendly Version](#)[Interactive Discussion](#)

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**Fig. 2.** Annual averaged mass concentrations of aerosol chemical species and its percentage of PM<sub>10</sub> in China.

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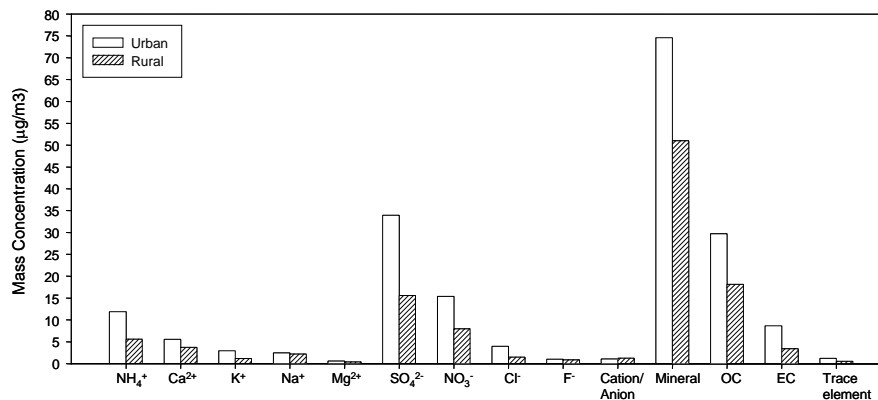
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**Fig. 3.** Annual averaged mass concentrations of various chemical species in urban and rural aerosols in China. The balance between cations and anions (cation/anion) are expressed in micro-equivalents per cubic meter in this diagram.

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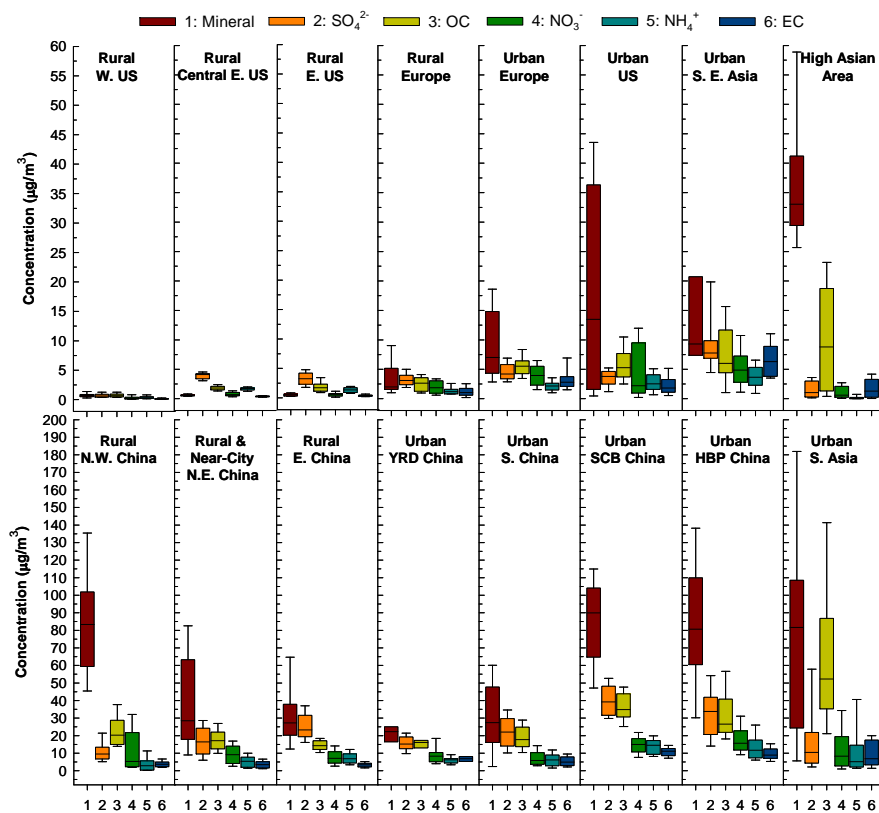


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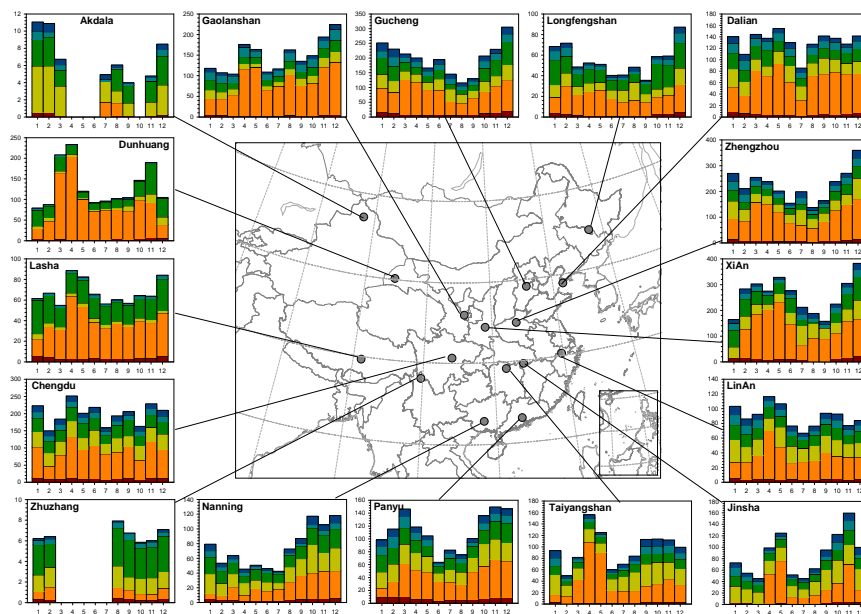
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**Fig. 4.** Box plot summarized the annual, seasonal or monthly mean mass concentration ( $\mu\text{g m}^{-3}$ ) of six major types of aerosol particles in diameter smaller than 10 micron with at least an entire year data from various rural and urban sites in 16 areas of the world continental area. These include: (1) *Rural W. US* (Malm and Schichtel, 2004); (2) *Rural Central E. US* ((Malm and Schichtel, 2004); (3) *Rural E. US* (Liu et al., 2005; Malm and Schichtel, 2004; Malm et al., 1994; Tanner et al., 2004); (4) *Rural Europe* (Gullu et al., 2000; Hueglin et al., 2005; Kocak et al., 2007; Putaud et al., 2004; Puxbaum et al., 2004; Querol et al., 2001, 2009; Rodriguez et al., 2004; Theodosi et al., 2010; Viana et al., 2008; Yin and Harrison, 2008; Yttri, 2007); (5) *Urban Europe* (Hueglin et al., 2005; Lenschow et al., 2001; Lonati et al., 2005; Perez et al., 2008; Putaud et al., 2004; Querol et al., 2008; Querol et al., 2001; Rodriguez et al., 2002; Rodriguez et al., 2004; Roosli et al., 2001; Viana et al., 2006; Viana et al., 2007; Yin and Harrison, 2008); (6) *Urban US* (Chow et al., 1993a; Ito et al., 2004; Kim et al., 2000; Liu et al., 2005; Malm and Schichtel, 2004; Sawant et al., 2004; Tanner et al., 2004); (7) *Urban S. E. Asia* (Kim et al., 2007b; Lee and Kang, 2001; Oanha et al., 2006); (8) *High Asian Area*, with mean height higher than 3500 m a.s.l (Decesari et al., 2009; Qu et al., 2009; Ram et al., 2010; Rastogi and Sarin, 2005; Shresth et al., 2000; This work); (9) *Rural N. W. China* .(Zhang et al., 2008a; This work); (10) *Rural & Near-City N. E. China* (Zhang et al., 2008a; This work); (11) *Rural E. China* (Zhang et al., 2008a; This work); (12) *Urban YRD (Yangzi River Delta) in E. China* (Wang et al., 2003; Wang et al., 2006; Yao et al., 2002; Ye et al., 2003); (13) *Urban S. China* (Hagler et al., 2006; Zhang et al., 2008a; This work); (14) *Urban SCB (Si Chuan Basin) in S. W. China* (Zhang et al., 2008a; This work); (15) *Urban HBP (Hua Bei Plain) in N. China* (Hu et al., 2002; Oanha et al., 2006; Roosli et al., 2001; Wang et al., 2005; Yao et al., 2002; Zhang et al., 2008a; This work); (16) *Urban S. Asia* (Chakraborty and Gupta, 2010; Khare and Baruah, 2010; Kumar et al., 2007; Lodhi et al., 2009; Raman et al., 2010; Rastogi and Sarin, 2005; Safai et al., 2010; Stone et al., 2010).

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**Fig. 5.** Monthly averaged daily concentrations changes of aerosol chemical species in various regions of the CAWNET network.

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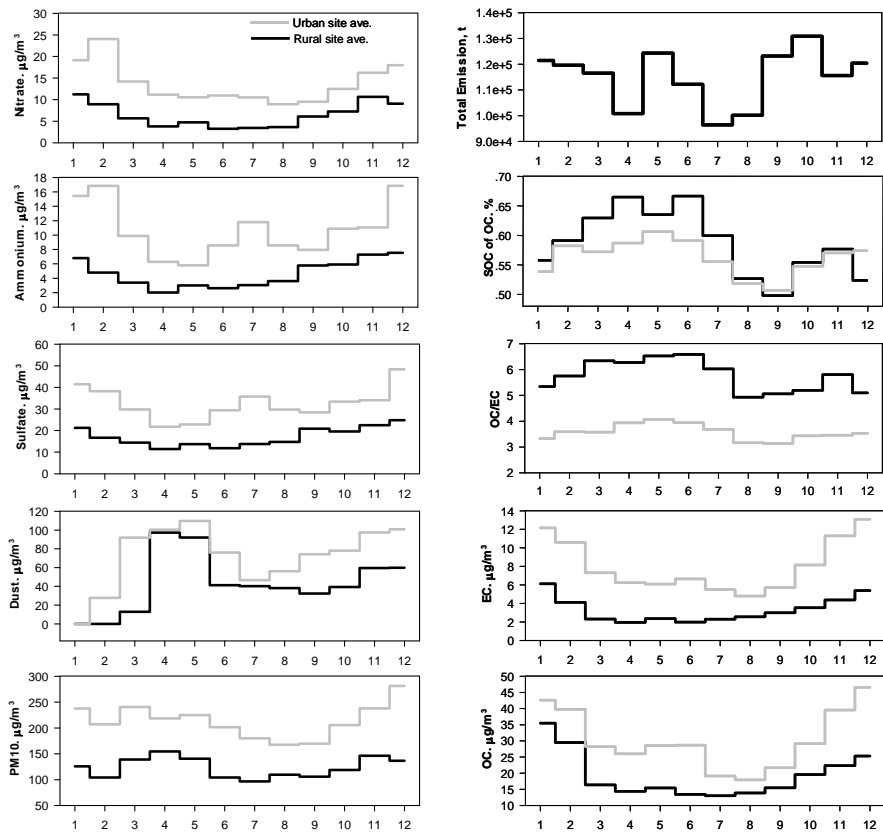
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**Fig. 6.** Monthly changes of urban and rural mean concentrations of various aerosol species in China, also including the monthly variations of total aerosol and its precursor emission data.

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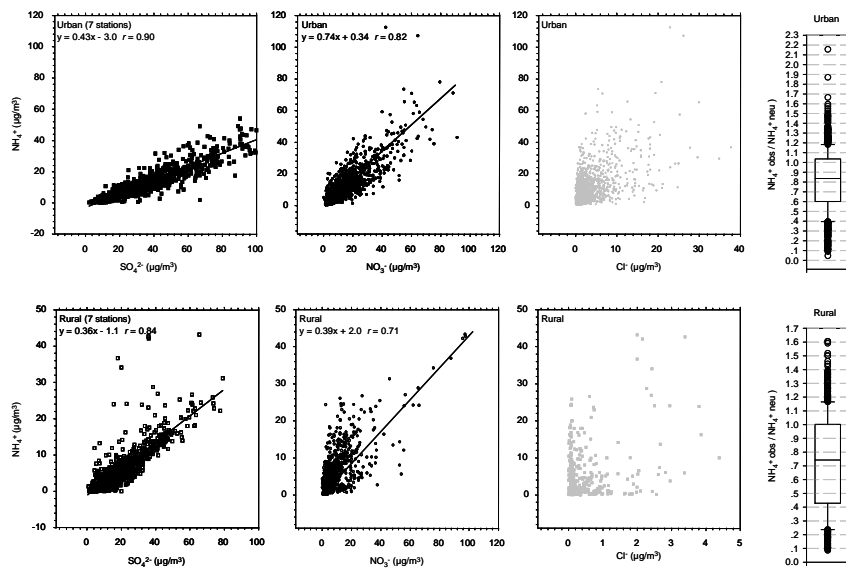
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**Fig. 7.** Correlations among  $\text{NH}_4^+$ ,  $\text{SO}_4^{2-}$ ,  $\text{NO}_3^-$  and  $\text{Cl}^-$ , as well as aerosol acidities estimation.

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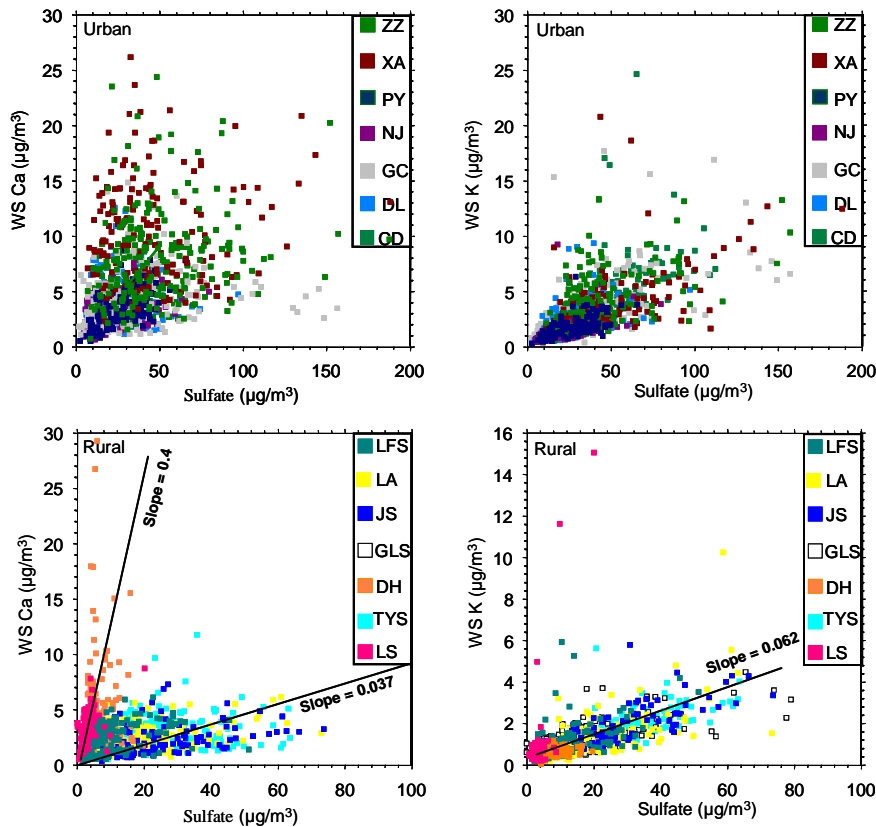
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**Fig. 8.** Correlations among sulfate, water soluble Ca and K for urban and rural aerosols in China.

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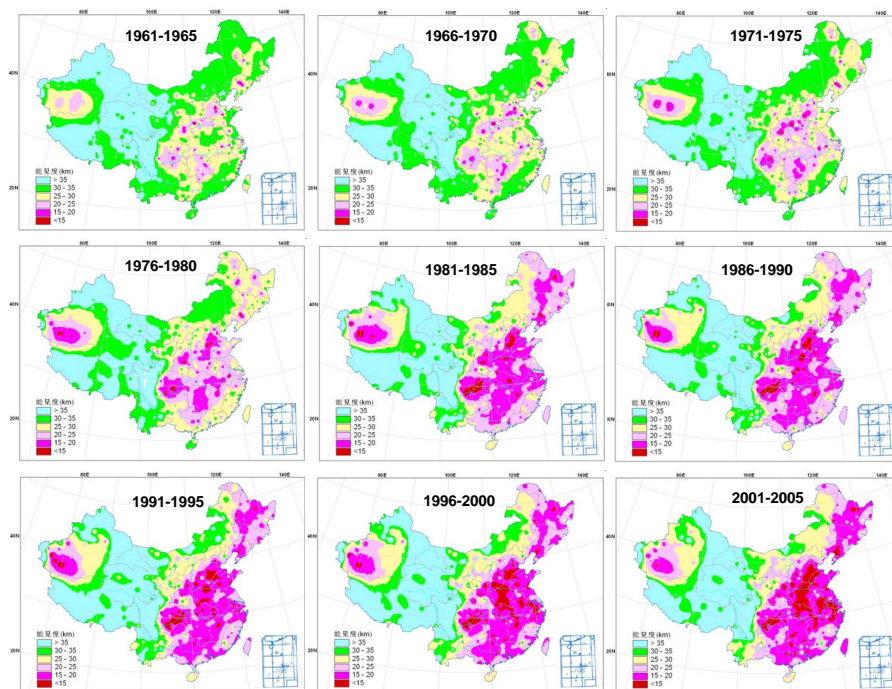


Fig. 9. Every five year mean visibility changes since 1960.

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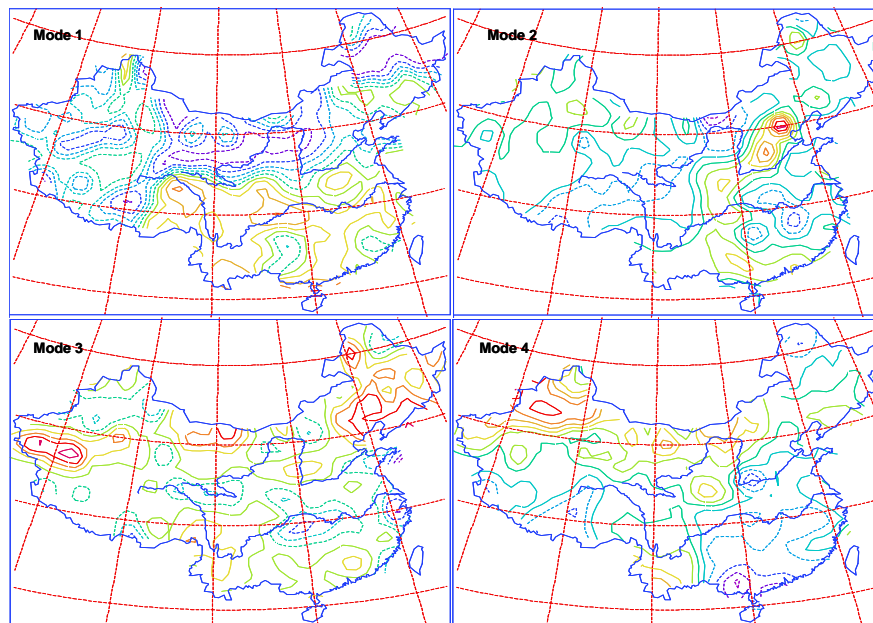
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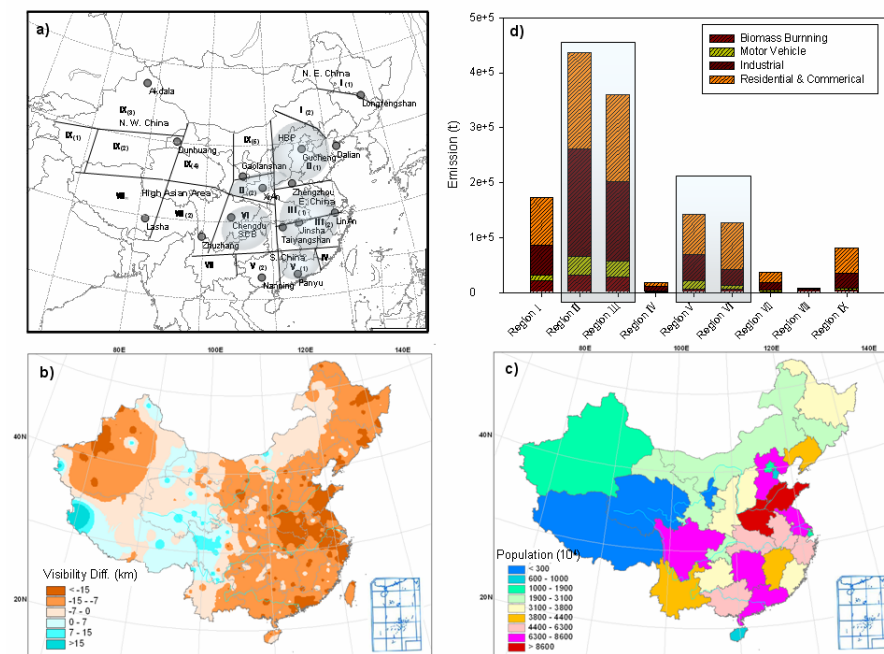


**Fig. 10.** Visibility mode (principal component) coefficient distributions that resolved by REOF method on the basis of 14:00 surface visibility data from 681 meteorological stations in China during 1957–2005. Contours denote the correlation among visibility variations in different regions.

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**Fig. 11.** (a) Nine similar visibility changing regions in China. The four major haze regions are marked with shadow; (b) Differences in five year mean visibilities between 2001–2005 and 1961–1965. Minus value with gradually darker colors represent the loss of visibility; (c) Population during 2001–2005 (NBSC, 2008); (d) Total emission of aerosol and its precious gas from nine regions of China. Original emission data from the reference (Cao et al., 2010).

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