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The optical, physical and chemical properties of the products of glyoxal uptake on ammonium sulfate seed aerosols

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Abstract

The heterogeneous reaction between gas phase glyoxal and ammonium sulfate (AS) aerosols, a proxy for inorganic atmospheric aerosol, was studied in terms of the dependence of the optical, physical and chemical properties of the product aerosols on initial particle size and ambient RH. The reactions were studied under different relative humidity (RH) conditions, varying from dry conditions ($\sim 20\%$ RH) and up to 90% RH, covering conditions prevalent in many atmospheric environments. At $\lambda = 355$ nm, the reacted aerosols demonstrate a substantial growth in optical extinction cross section, as well as in mobility diameter under a broad range of RH values ($35\text{--}90\%$). The ratio of the product aerosol to seed aerosol geometric cross section reached up to ~ 3.5 , and the optical extinction cross-section up to ~ 250 . The reactions show a trend of increasing physical and optical growth with decreasing seed aerosol size, from 100 nm to 300 nm, as well as with decreasing RH values from 90% to $\sim 40\%$. Optically inactive aerosols, at the limit of the Mie range (100 nm diameter) become optically active as they grow due to the reaction. AMS analyses of the reaction of 300 nm AS at RH values of 50% , 75% and 90% show that the main products of the reaction are glyoxal oligomers, formed by acetal formation in the presence of AS. In addition, imidazole formation, which is a minor channel, is observed for all reactions, yielding a product which absorbs at $\lambda = 290$ nm, with possible implications on the radiative properties of the product aerosols. The ratio of absorbing substances (C–N compounds, including imidazoles) increases with increasing RH value. A core/shell model used for the investigation of the optical properties of the reaction products of AS 300 nm with gas phase glyoxal, shows that the refractive index (RI) of the reaction products are in the range between $1.57\text{--}1.71$ for the real part and between $0\text{--}0.02$ for the imaginary part of the RI at 355 nm. The observed increase in the ratio of the investigated absorbing substances is slightly indicated in the RI values found by the model, as the imaginary part of the product RI increases from 0.01 to 0.02 with increasing RH. The imaginary part is expected to increase further at higher RH and become more substantial in cloud

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droplets. This study shows that the reaction of abundant substances present in atmospheric aerosols, such as AS, and gas phase glyoxal alters the aerosols' optical, physical and chemical properties and may have implications on the radiative effect of these aerosols.

1 Introduction

Aging of aerosol particles in the atmosphere is an important process, leading to products with new chemical and physical properties, higher molecular weight (MW), higher oxidation state, and therefore hygroscopicity, as well as altered optical properties. Therefore, investigating the effect of atmospheric condensed-phase reactions is vital for better understanding and modeling of aerosols' atmospheric role in climate and health. Aging occurs via oxidation reactions by abundant radicals such as OH, O₃, NO₃ and Cl, (Rudich, 2003; Rudich et al., 2007) as well as by gaseous organic oxidants, such as gas-phase glyoxal (Hallquist et al., 2009; Heald et al., 2008; Jimenez et al., 2009; Volkamer et al., 2007; Zhang et al., 2007), and may contribute to the formation of secondary organic aerosols (SOA).

SOA constitute an important fraction of tropospheric aerosol mass (Jimenez et al., 2009). They form via oxidation of VOCs and their consequent condensation onto aerosols, but also include products of condensed-phase chemical processes. An important path for SOA production is through cloud processing and reactions in aqueous phase aerosol (Carlton et al., 2007; Hallquist et al., 2009; Kroll and Seinfeld, 2008; Tan et al., 2009). SOA may contain oligomers and other high MW species, such as Humic Like Substances (HULIS) (Graber and Rudich, 2006; Kalberer et al., 2004). SOA physical, chemical and optical properties are a major source of uncertainty in aerosol sciences. Their formation is probably underestimated in current models, despite their large contribution to aerosol mass (Hallquist et al., 2009; Kanakidou et al., 2005).

Ammonium sulfate (AS) is a ubiquitous component of global anthropogenic aerosols, particularly in the accumulation mode, which renders it a highly efficient scatterer in

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the actinic light range. It scatters in both the visible and the near UV wavelengths, and is a hygroscopic aerosol component with a deliquescence point at approximately 80 % relative humidity (RH), while the dehydration process involves hysteresis, with an efflorescence point at ~35 % RH (Cruz and Pandis, 2000; Cziczo and Abbatt, 1999; Dinar et al., 2007).

Glyoxal is produced mainly from the photooxidation of biogenic and anthropogenic volatile organic compounds (VOCs) by OH radical-initiated oxidation of anthropogenic aromatic VOCs, alkenes and acetylene, reactions of O₃ with alkenes, and by the oxidation of biogenic molecules such as isoprene (Fu et al., 2008; Myriokefalitakis et al., 2008; Volkamer et al., 2009; Volkamer et al., 2007). Its typical atmospheric concentrations range from 0.01–5 ppb (Grosjean et al., 1996; Grossmann et al., 2003; Liggio et al., 2005b; Volkamer et al., 2007; Wittrock et al., 2006). Glyoxal is found in urban environments, such as mega-cities, over the oceans and in forests, including tropical rain forests, and in biomass burning regions, it is also found in marine and coastal sites, and countryside locations (Fu et al., 2008, 2009; Myriokefalitakis et al., 2008; Sinreich et al., 2007; Volkamer et al., 2007). Due to their high solubility and reactivity in the aqueous phase, glyoxal and other α -dicarbonyls contribute substantially to SOA formation, mainly via cloud processing, while the photolysis of glyoxal significantly contributes to HOx (OH+HO₂) chemistry in the gas phase (Carlton et al., 2007; Ervens et al., 2008; Fu et al., 2008; Hallquist et al., 2009). Glyoxal participates in reactive uptake when liquid water is present, and is therefore scavenged by hydrated aerosol, or fog and cloud droplets (Corrigan et al., 2008; Hastings et al., 2005; Kroll et al., 2005; Liggio et al., 2005b; Schweitzer et al., 1998).

Several studies have investigated the processes by which glyoxal is introduced into aerosol and facilitates SOA formation (Carlton et al., 2007; Corrigan et al., 2008; De Haan et al., 2009; Galloway et al., 2009; Hastings et al., 2005; Kroll et al., 2005; Liggio et al., 2005a; Noziere et al., 2009; Schwier et al., 2010; Shapiro et al., 2009). The reactive uptake of glyoxal by ammonium ion-containing aerosol has been shown to be an important pathway for gas phase glyoxal depletion and can account for the observed

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5 excess of glyoxal in liquid phase aerosol compared to those expected based on its Henry's law values. The products of the reaction between inorganic salts, including AS, and glyoxal, as well as other carbonyl compounds, are mostly oligomeric species, formed by a reaction in which the salt acts as a catalyst for the oligomerization of carbonyl compounds by acetal formation (in the case of glyoxal) or by aldol condensation (Corrigan et al., 2008; Galloway et al., 2009; Kroll et al., 2005; Liggió et al., 2005b; Noziere et al., 2009, 2010). Noziere et al. (2009) have conducted a series of reactions between AS and glyoxal in the bulk phase, and have identified two main peaks in the UV-Vis absorption spectrum; a peak at $\lambda = 209$ nm, and a smaller peak at $\lambda = 290$ nm. 10 The first peak is attributed to the formation of glyoxal oligomers, produced via two suggested mechanisms, either oligomerization by acetal formation or nucleophilic attack of the glyoxal by the ammonium ion followed by water loss. The latter can also form heterocyclic products with C–N bonds, which can account for the observed peak at $\lambda = 290$ nm. According to Galloway et al. (2009) the oligomers absorb in the short UV (peak at 200 nm and less), which corresponds well with the UV-Vis spectrum peak at $\lambda = 209$ nm observed by Noziere et al. (2009), and are therefore optically irrelevant to the troposphere. In addition, Galloway et al. (2009) found that glyoxal oligomerization in the aqueous phase is reversible and concentration-dependent. The reaction of AS with glyoxal also produces imidazoles (the second mechanism suggested by Noziere et al. (2009)), which absorb UV radiation at $\lambda = 290$ nm, and as a result may influence atmospheric absorption at an important range of the actinic flux, where photolysis leads to OH formation (Cho, 2004; Galloway et al., 2009; Hewitt and Harrison, 1985; Noziere et al., 2009; Schulze, 1973; Sjøstedt et al., 2007). Therefore, aerosol absorption at this wavelength can have significant implications on radiative transfer and atmospheric chemistry. 25

The heterogeneous reaction between AS seed aerosols and gas phase glyoxal was also studied in terms of the chemical and physical properties of the product aerosols, as well as the kinetics and thermodynamics of the reaction (Galloway et al., 2009; Kroll et al., 2005; Liggió et al., 2005a, b; Shapiro et al., 2009). Galloway et al. (2009)

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and Kroll et al. (2005) performed their studies in the Caltech 28 m³ Teflon environmental chambers, using hydrated polydisperse aerosol populations under varying glyoxal gas phase concentrations and RH values. Under dark conditions, both studies have established the formation of oligomeric species of glyoxal and Galloway et al. (2009) have also observed the formation of C–N compounds, among which are imidazoles. Liggio et al. (2005a, b) used various hydration states of a monodisperse aerosol population and found similar chemical products. All studies show growth in size and mass under various hydration levels, and no growth under dry conditions, confirming that the mechanism of uptake and subsequent oligomerization of gas phase glyoxal on AS seed aerosols requires hydrated conditions. Galloway et al. (2009) demonstrate that the reaction of glyoxal with ammonia is induced by the hydration of the ammonium ion in the AS particle, and leads to imidazole formation (Galloway et al., 2009).

In this paper we study the changes in the optical properties of AS aerosols due to the heterogeneous reaction with gas phase glyoxal. We establish, for the first time, the dependence of the optical, chemical and physical properties of the reaction products on the initial particle size and on the RH conditions. The RH values chosen in this study are atmospherically relevant, ranging from low ambient RH values, found in arid regions up to values of 90 %, typical for tropical regions or in close vicinity of clouds. By following the change in optical, physical and chemical properties of a specific initial AS particle size while varying the RH values, we determine the effect of RH on the reaction. In the same manner, we inter-compare the effect of size on the reaction while keeping the RH constant. In addition, we determine the effect of the heterogeneous reaction on the optical properties of the aerosol particles using cavity ring down aerosol spectrometer (CRD-AS).

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2 Methodology

2.1 Aerosol generation and classification

A schematic of the laboratory setup used is shown in Fig. 1. Aerosols were generated by the method described in detail in our previous publications (Dinar et al., 2008; Lang-Yona et al., 2009; Riziq et al., 2007). Briefly, an aqueous solution of AS (0.04M–0.11M) is nebulized with dry nitrogen flow; the formed droplets are dried by passing through two silica gel diffusion dryer columns, and consequently charged by a neutralizer (TSI model 3012A). Size selection of the resulting dry polydisperse aerosol population is achieved with an electrostatic classifier (differential mobility analyzer (DMA), TSI model 3081), after which the nearly monodisperse aerosol flow is directed into a humidification stage consisting of a nafion humidifier with a temperature controlled bath. This setup allows the RH of the aerosol flow to be controlled by varying the bath temperature. The hydrated aerosol flow enters a 30 L reaction vessel into which gas phase glyoxal is introduced. The residence time of the particles with the gaseous glyoxal is approximately 1 h, after which the aerosol population is dried by two additional silica gel diffusion dryer columns and introduced into a scanning mobility particle sizer (SMPS, TSI model 3080) for determination of size distribution and concentration of the dried reacted particles. An $80 \text{ cm}^3 \text{ min}^{-1}$ flow is directed into the Aerodyne High Resolution Time of Flight Aerosol Mass Spectrometer (Hi-RES-TOF-AMS, Aerodyne) system for determination of the chemical properties, and $800 \text{ cm}^3 \text{ min}^{-1}$ are introduced into a pulsed cavity ring down (CRD) system for determination of their optical properties (see Fig. 1).

In each experiment, we set an initial dry AS particle size and conducted the reaction with gas phase glyoxal in the ~ 1 h residence time reaction vessel, under various RH values (Fig. 1). For the measurement of the optical properties we used the CRD-AS system at $\lambda = 355 \text{ nm}$. By choosing this wavelength we were hoping to observe the increase in absorption due to the formation of imidazoles. Although their peak absorption is at $\lambda = 290 \text{ nm}$, there is a tail of absorption extending to $\lambda = 355 \text{ nm}$ as well.

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Due to increased absorption by the mirrors, shorter wavelengths could not be studied. The wavelength at $\lambda = 355$ nm is atmospherically important because it is within the actinic light reaching the earth's atmosphere and the absorption by organic substances is enhanced as we move from the visible towards the UV part of the spectrum. For each RH value, we measured prior, during and after the reaction, the extinction coefficient (α_{ext}) of the dry particles using the CRD-AS system at $\lambda = 355$ nm and calculated the extinction cross section (σ_{ext}) from Eqs. (1)–(2), while simultaneously measuring the dry mean diameter of the particles with the SMPS.

2.2 CRD-AS system

The CRD-AS system was described in details elsewhere (Dinar et al., 2008; Lang-Yona et al., 2009; Riziq et al., 2007, 2008). Briefly, a pulsed laser light of ~ 1 mJ at $\lambda = 355$ nm is injected into a cavity containing two highly reflective mirrors (reflectivity, $R = 99.9$) where it undergoes multiple passes. The light intensity decays exponentially with a characteristic time constant which is a function of the mirror reflectivity and the losses within the cavity. When scatterers and/or absorbers are introduced into the cavity, the light decays faster and the extinction coefficient can be established (Eq. 1).

$$\alpha_{\text{ext}} = \frac{L}{dc} \left[\frac{1}{\tau} - \frac{1}{\tau_0} \right] \quad (1)$$

where α_{ext} is the measured extinction coefficient, τ_0 is the characteristic decay time for aerosol-free cavity, τ is the cavity ring-down time in the presence of particles, L/d is the ratio of the total cavity length to the cavity length occupied by particles, and c is the speed of light. Using the particle number concentration, the extinction cross section (σ_{ext}) at a specific wavelength can be calculated (Eq. 2). (Lang-Yona et al., 2009; Pettersson, 2004; Riziq et al., 2007, 2008).

$$\sigma_{\text{ext}} = \frac{\alpha_{\text{ext}}}{N} \quad (2)$$

Where N is the particle number concentration, determined by a condensation particle counter (CPC, TSI model 3022A). For spherical particles, the extinction efficiency (Q_{ext}) can be calculated using Eq. (3).

$$Q_{\text{ext}} = 4 \frac{\sigma_{\text{ext}}}{\pi D^2} \quad (3)$$

5 where D is the particle mean diameter. By fitting the curve of Q_{ext} versus the size parameter χ (the ratio of the particle circumference (πD) to the laser's wavelength (λ), given by $\chi = \pi D/\lambda$) to a Mie theory calculation, the complex refractive index (RI) can be retrieved by varying independently the real and the imaginary parts of the RI.

2.3 Glyoxal gas generation and concentration measurement

10 Gas phase glyoxal was added to the reaction vessel by flowing $50 \text{ cm}^3 \text{ min}^{-1}$ of pure dry N_2 through a bulb containing solid phase glyoxal trimer dihydrate (GTD) (97%, Sigma Aldrich) heated to 110°C , based on the method used by Corrigan et al. (2008); Galloway et al. (2009); Hastings et al. (2009); Kroll et al. (2005); Liggio et al. (2005a, b).

15 The concentration of the glyoxal gas was measured with the CRD-AS system at a wavelength of $\lambda = 405 \text{ nm}$, using equation 4 for measurement of gaseous species.

$$\text{conc} = \frac{\alpha_{\text{abs}}}{\sigma_{\text{abs}}} \quad (4)$$

20 Where α_{abs} is the absorption coefficient measured by the CRD-AS system [cm^{-1}], σ_{abs} is the absorption cross-section [$\text{cm}^2 \text{ molecule}^{-1}$], and conc is the gas concentration [molecule cm^{-3}]. The value of σ_{abs} at $\lambda = 405 \text{ nm}$ is $4.491 \times 10^{-20} \text{ cm}^2 \text{ molecule}^{-1}$ (Volkamer et al., 2005). This measurement yields a glyoxal concentration in the reaction vessel of $1.6 \pm 0.5 \text{ ppb}$.

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3 Results and discussion

Following 1 h of mixing between a selected size of AS aerosol and gas phase glyoxal, the resulting aerosol size distribution, optical and chemical properties were measured with an SMPS, a CRD-AS system and an AMS, respectively. The reaction time is within the timescale previously observed for the heterogeneous reaction between AS particles and gas phase glyoxal (Galloway et al., 2009; Kroll et al., 2005; Liggio et al., 2005a; b).

3.1 Optical and physical properties

In this section we establish the dependence of the optical and physical properties of the dry product aerosols on initial particle size and on the RH (and hence, hydration state of the aerosol) for the reaction between AS particles and gas phase glyoxal.

Figure 2a shows the final optical extinction cross-sections (σ_{ext}), and Fig. 2b the final mean diameters (nm) for the dry aerosol products of the reaction of 250 nm AS particles with gas phase glyoxal under varying RH values. The RH range is from 20 % (dry aerosol) up to 90 %. Up until a threshold of 35 % RH, the reaction does not occur at all, there is no change in optical properties and size. At the RH value of 35 %, a reaction is observed. Table 1 show that the maximum change occurs at 40 % RH, where the dry mean diameter of the particles measured by the SMPS grows in ~ 1 h from 250 nm to 420 nm, and σ_{ext} reaches $6.83 \times 10^{-9} \text{ cm}^2$, a 6 fold increase from the initial value of σ_{ext} of $1.13 \times 10^{-9} \text{ cm}^2$ for AS aerosol. The magnitude of change decreases in both size and σ_{ext} with increasing RH value from 40 % to 90 %, however it remains significant at all RH values up to 90 %, where the size increases to 290 nm and σ_{ext} reaches $2.99 \times 10^{-9} \text{ cm}^2$ (values and their corresponding error are given in Table 1). Therefore, at all RH values in which the reaction occurs, i.e. 35 %–90 %, the reaction leads to more optically active aerosols. The final mean diameters (nm) and σ_{ext} of the products of the ~ 1 h heterogeneous reactions of AS 100 nm, 200 nm and 300 nm particles with glyoxal gas at RH values varying from 50 % to 90 % have also been studied, and exhibit

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the same trend as the AS 250 nm (Table 1). As expected, no reaction occurs in dry AS particles, since the reaction requires hydrated conditions (Galloway et al., 2009; Kroll et al., 2005; Liggió et al., 2005a, b). The hygroscopic growth of AS has been extensively studied (Cruz and Pandis, 2000; Cziczo and Abbatt, 1999; Dinar et al., 2007; Gysel et al., 2002). Its deliquescence RH (DRH) was found at $\sim 80\%$ with hysteresis upon dehydration, leading to an efflorescence point at $\sim 35\%$ RH. Galloway et al. (2009), Kroll et al. (2005), and Liggió et al. (2005a, b) also show that a reaction occurs at low relative humidities of RH 50 % and less. A recent study by Mikhailov et al. (2009) demonstrates that water monolayers are formed on the surface of AS particle at low RH values from approximately $\sim 20\%$ RH. Therefore, we suggest that the reaction we observe under low RH conditions may proceed through water adsorption onto the AS particles by forming a thin aqueous layer on the particle. The aqueous conditions allow for the dissolution of the AS and the glyoxal gas into the outer layer of the particle, leading to a highly concentrated aqueous component around the particle. This may also explain why the change in the optical and physical cross sections of the product aerosol is enhanced with decreasing RH values in the reaction. This phenomenon is elaborately explored in the next section (see Sect. 3.2).

Figure 3 shows the comparison of the growth ratios of the optical extinction cross-sections, for all measured AS sizes. Figure 3 and Table 1 demonstrate the dependence of the growth in physical and optical cross sections on the initial AS particle size. The growth in the optical and physical cross sections of the product aerosol is enhanced with decreasing initial AS particle size. Additionally, the optical extinction growth ratios exceed the geometric growth (Table 1). This is especially evident in the case of AS 100 nm particles. In this case, while the growth in geometric cross section due to the reaction is in the same order as for all other AS particle sizes, ranging from 1.81 for RH 90 % to 3.28 for RH 50 % (Table 1), the optical cross sections increase by two orders of magnitude in all measured RH values. At 50 % RH, σ_{ext} increases 244 times from 2.34×10^{-12} of AS 100 nm particle to $5.71 \times 10^{-10} \text{ cm}^2$ after the reaction, at 75 % RH it increases 119 fold to 2.79×10^{-10} and at 90 % RH 59 times reaching σ_{ext}

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of 1.39×10^{-10} . Therefore, the effect of this reaction can be significant since under most atmospheric conditions it can alter a major fraction of the aerosol population from having negligible radiative effect into optically active aerosols. Furthermore, as expected from Fig. 2, the reaction, followed by optical and physical growth, occurs already at 35 % RH and is most significant at 40 % RH, which has great implications on the radiative properties of the reacted aerosol under these environmental conditions.

To investigate the optical properties of the chemical products due to the heterogeneous uptake of glyoxal by AS aerosol, we calculated the Q_{ext} of the product particles using Eq. (3). All reacted aerosols have Q_{ext} values that are different from the values of pure AS.

We implemented the obtained Q_{ext} values in a core/shell model to obtain the RI of the reaction products which is assumed to be in a reacted shell over an unreacted core of ammonium sulfate (modified from the code of Liu et al. (2007)). The core/shell model assumes that the structure of the product aerosol is that of an AS core and the shell in this model is assumed to be the added material from the substances formed by the reaction. The model was used to investigate the reaction of gas phase glyoxal with AS 300 nm seed aerosols at 50 %, and 75 % RH. It could not be used for 90 % RH because this value is above the AS particles' deliquescence point and hence a core+shell structure of the final product particles is not applicable, but rather complete mixing of the AS and the reaction products is expected. Since there is no deliquescence at 50 % and 75 % RH, we assume a small layer of water around the AS seed, and therefore can say that the change in core diameter due to dissolution is negligible. The RI retrieval of a mixed particle with a single known Q_{ext} value can result in a very wide range of RIs. The input for the core/shell model is the AS RI, the total Q_{ext} of the particle, the particle total size, and the sizes of the core and shell, presented as the mean diameters obtained by the SMPS. The model is based on Mie theory and therefore assumes spherical particles. It retrieves the RI of the shell by first varying the possible RI values of the shell and calculating the corresponding Q_{ext} and then finding the minimum square deviation between the measured and the modeled Q_{ext} values. The core of the

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particles is the AS 300nm seed, with an RI of $n = 1.553 \pm 0.005 + i0.002 \pm 0.013$. The final sizes of the product particles are 436 nm, and 378 nm, for RH values of 50 %, and 75 %, respectively (Table 1), and the shell size is calculated as the difference between the total final particle size and the AS 300 nm seed. Table 2 shows the results of the model calculations for the RIs of the shells for the two experiments we investigated. The obtained RI value at 355 nm is $n = 1.67 + i0.01$ with a range between $n = 1.7 + i0$ and $n = 1.59 + i0$ for the 436 nm particle with Q_{ext} of 4.47 ± 0.34 , and $n = 1.71 + i0.02$ with a range between $n = 1.68 + i0$ and $n = 1.57 + i0.02$ for the 378 nm particle with Q_{ext} of 4.60 ± 0.35 . Since we only have one value of measured Q_{ext} as an input for the model, we expect that the output RI range will have large uncertainty associated with the limited data provided to the model. Nonetheless, this model provides a first estimation of the RI of the reaction products at $\lambda = 355$ nm. This may significantly contribute to the scarcely available information of optical properties of organic SOA from such reactions. To validate our results, we compared these values to the values found by Dinar et al. (2008) for HULIS of biogenic sources (K-pusztá). We expect the values to have some similarity since the RI we obtained is that of the main reaction products, which are glyoxal oligomers (see next section), similar to the organic oligomers which constitute the HULIS compounds. The RI value of $n = 1.616 + 0.023i$ at $\lambda = 390$ nm obtained by Dinar et al. (2008) is in qualitative good agreement with the RI values we obtained.

3.2 AMS analysis results

The changes in chemical composition were monitored with the high-resolution time of flight aerosol mass spectrometer (Hi-RES-TOF AMS). Figure 4 shows the ratio of the mass of the main glyoxal oligomer fragments, of fragments of C–N compounds and of imidazole (m/z 68) to sulfate mass, before the reaction of AS 300 nm with gaseous phase glyoxal, and after the reaction at RH values of 50 %, 75 % and 90 %. The mass peaks considered in the AMS analysis are taken from the studies of Galloway et al. (2009) and Liggio et al. (2005a, b). The fragments of glyoxal oligomerization

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products are composed of m/z 29, 30, 47, 58, 60, 77, 88, 105, 117, 135, 145, 175, 192 and 193. The C–N compounds are the sum of m/z 41, 46, 52, 53, 57, 68, 69, 70 and 96. We use m/z 68 for the imidazole peak from the analysis of Galloway et al. (2009). All masses are normalized to sulfate mass to account for losses during the reaction, according to the method used by Galloway et al. (2009) and by Liggiio et al. (2005a, b). The glyoxal oligomerization products increase after the reaction onset by two orders of magnitude compared to the background values present before the reaction, and decrease with increasing RH, in accordance with the observed trend in geometric cross section and σ_{ext} shown in Figs. 2–3 and in Table 1. This trend can be attributed to an increase in the dilution of the AS aerosol with increasing RH. As suggested by Liggiio et al. (2005b), the dilution in higher RH values slows the oligomerization process. Noziere et al. (2009) have shown that the NH_4^+ ion acts as a catalyst for the reaction of glyoxal oligomerization by acetal formation, and hence the dilution of the AS due to higher water content in the particle is likely to decrease the reaction rate.

The glyoxal oligomerization products are the main products of this reaction, with total mass/sulfate mass ratios of 4.5 ± 0.1 after the reaction at RH 50 %, 1.8 ± 0.1 at RH 75 %, and 1.2 ± 0.1 for RH 90 %. The normalized mass values of C–N compounds are between 1–2 orders of magnitude smaller and the normalized imidazole mass is 2–3 orders of magnitude less than the glyoxal oligomer fragments. The normalized C–N compounds and the imidazole products have increased by up to two orders of magnitude compared with pre-reaction background values, exhibiting the same trend of decreasing normalized mass with increasing RH values.

Figure 5 shows the ratios (normalized to sulfate mass) of C–N compounds to glyoxal oligomerization products and the ratio of imidazoles only, to the glyoxal oligomerization products versus RH values. The C–N compounds' ratio are 30×10^{-3} , 38×10^{-3} , and 46×10^{-3} for the reactions conducted at RH values of 50 %, 75 % and 90 %, respectively. The ratio of the imidazoles are 3×10^{-3} , 5×10^{-3} , 8×10^{-3} for the RH values of 50 %, 75 % and 90 %, respectively. The increase in both ratios with increasing RH values indicates an increase in the formation of absorbing substances with increasing

water content of the reacting particles. This suggests that higher water content enhances the imidazole formation pathway shown in Galloway et al. (2009) (Fig. 9 of their paper) and we can assume that in cloud droplets containing AS, where the water content is higher, this reaction will be further enhanced, leading to a higher ratio of absorbing material, and potentially altering the optical properties by an addition of an absorbing component to their RI.

The RI values derived by the core/shell model are in the range between 1.57–1.71 for the real part and 0–0.02 for the imaginary part of the RI (Table 2). This is attributed to the main products measured by the AMS, i.e., fragments of glyoxal oligomeric species. This is the first time, to our knowledge, that an estimation of the RI of these oligomers was performed. The observed increase in the ratio of the investigated absorbing substances (C-N compounds, including imidazoles) to the glyoxal oligomers with increasing RH values is slightly indicated in the RI values found by the model, as the imaginary part of the product RI increases from 0.01 to 0.02 with increasing RH. Although the increase is within the model error, we cannot rule out that it indicates a trend of growth in the imaginary part of the RI with increasing RH. The fact that this is not a significant increase is expected from the difference of three orders of magnitude between the main oligomeric products, which are non-absorbing, and the absorbing products (see Table 2 and Figs. 4–5).

A comparison of the results of the AS 300 nm particles mass growth in our study to the values measured in Galloway et al. (2009) and Liggio et al. (2005a, b) is presented in Table 3. The ratio of final mass to the seed mass (normalized by sulfate mass) is in the same order of magnitude as found by Liggio et al (2005b). Our value of 13.30 ± 0.79 is slightly smaller than their value of 16, and our calculated mass growth rate is one order of magnitude smaller. Since we found in this study that the reaction is enhanced with decreasing initial particle size, and our initial seed aerosol size is approximately twice the size used by Liggio et al. (2005b) this difference can be expected. Moreover, the glyoxal gas concentration in our study is ~ 2 ppb compared to ~ 5 ppb used by Liggio et al. (2005b) (see Table 3). Both ours and the value of Liggio et al. (2005b) for the

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ratio of final mass (after four hours) to seed mass are two orders of magnitude larger than that found by Galloway et al. (2009). However, the growth rates cannot be directly compared, since this study and Liggio et al. (2005b) use a monodisperse distribution of aerosols reacting with gas phase glyoxal, and both Galloway et al. (2009) and Kroll et al. (2005) use a polydisperse distribution for the reaction. As shown in this study, the extent of particle growth depends on the initial particle size, therefore not all particles in the population used by Galloway et al. (2009) and Kroll et al. (2005) grow to the same extent, while in this study and in Liggio et al. (2005b) the mass growth per particle can be determined and compared.

Our experiments were conducted for the hydration process of AS aerosols (i.e. while increasing the RH). The dehydration process such as in the case of a drying cloud droplet or saturated aerosol particles, involves efflorescence and hysteresis. While the hydration process leads to a deliquescence point for AS particles at approximately 80 % RH, dehydration, their efflorescence point is at ~35 % RH (Dinar et al., 2007). This suggests that under the same RH conditions (in the 40–80 % range), the reaction of an evaporating AS droplet will be slower than for the reaction under hydration conditions, due to the higher water content of the dehydrating particles compared to the hydrating ones.

4 Atmospheric implications

The heterogeneous reaction of AS particles with glyoxal under hydrated conditions occurs at a very wide RH range, from 35 % to 90 % and is therefore relevant under atmospheric conditions varying from arid and semi arid areas to tropical regions. Due to reaction on hydrated AS particles there is formation of secondary organic material composed of glyoxal oligomers and C–N containing products that result in particle growth. The enhancement in the optical cross section due to the chemical reactions and the subsequent growth in size affect the potential radiative properties of the reacted particles significantly. This phenomenon is especially important for small aerosols that

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are not optically active prior to the reaction. Following the reaction, the change in size renders them optically active (up to 2 orders of magnitude enhancement in optical extinction cross section).

The increase in particle size due to the reaction affects the particles' microphysical properties in addition to their radiative properties. Larger particles act as better cloud condensation nuclei (CCN). However, in order to fully establish the microphysical properties of the product aerosols, the hygroscopicity of the glyoxal oligomerization product compounds needs to be studied.

The glyoxal gas concentration used in this study is ~ 2 ppb, in the range of atmospheric glyoxal concentrations, found to be between 0.01–5 ppb (Grosjean et al., 1996; Grossmann et al., 2003; Liggio et al., 2005b; Volkamer et al., 2007; Wittrock et al., 2006), and the experiments were conducted for 1 h. This suggests that the observed changes due to the reaction can occur within 1 h in the ambient atmosphere. Although Galloway et al. (2009) have shown that the glyoxal oligomerization in the particle phase is reversible, the temporary change in particle size and optical properties due to the reaction can cause a long term effect on both radiative properties and cloud formation in the region it occurs.

This study shows that the heterogeneous reaction of glyoxal and AS aerosols can affect both the SOA mass and the radiative effect of the aerosol. The extent of change depends strongly on the size, and the RH as well as the direction of RH change of the reacting particles.

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Table 1. The final sizes and optical cross sections compared to the initial values, for AS particles with mean diameters of 101 nm, 201 nm, 252 nm and 304 nm, measured by the SMPS, and the growth ratios of the geometric and optical cross sections for all measured RH values.

Initial size (nm)	% RH ($\pm 3\%$)	Final size (nm)	Ratio of final to initial geometric cross section growth	Initial σ_{ext} (cm ²)	Final σ_{ext} (cm ²)	σ_{ext} growth ratio
101 \pm 6	50	186 \pm 26	3.38 \pm 0.03	2.34 $\times 10^{-12}$ $\pm 9.76 \times 10^{-13}$	5.71 $\times 10^{-10}$ $\pm 1.02 \times 10^{-12}$	244
	75	156 \pm 20	2.40 \pm 0.02		2.79 $\times 10^{-10}$ $\pm 9.37 \times 10^{-13}$	119
	90	136 \pm 10	1.81 \pm 0.01		1.39 $\times 10^{-10}$ $\pm 8.27 \times 10^{-13}$	59
201 \pm 15	50	357 \pm 39	3.15 \pm 0.03	4.09 $\times 10^{-10}$ $\pm 1.00 \times 10^{-12}$	3.71 $\times 10^{-9}$ $\pm 1.99 \times 10^{-12}$	9
	75	312 \pm 33	2.42 \pm 0.02		2.67 $\times 10^{-9}$ $\pm 2.16 \times 10^{-12}$	7
	91	272 \pm 26	1.83 \pm 0.01		1.65 $\times 10^{-9}$ $\pm 1.90 \times 10^{-12}$	4
252 \pm 19	19	248 \pm 18	0.97 \pm 0.01	1.13 $\times 10^{-9}$ $\pm 5.46 \times 10^{-11}$	1.25 $\times 10^{-9}$ $\pm 3.60 \times 10^{-12}$	1
	30	246 \pm 20	0.95 \pm 0.01		1.21 $\times 10^{-9}$ $\pm 8.44 \times 10^{-13}$	1
	35	363 \pm 50	2.07 \pm 0.02		4.31 $\times 10^{-9}$ $\pm 4.89 \times 10^{-12}$	4
	40	422 \pm 42	2.80 \pm 0.02		6.83 $\times 10^{-9}$ $\pm 6.23 \times 10^{-12}$	6
	50	398 \pm 45	2.50 \pm 0.02		5.57 $\times 10^{-9}$ $\pm 6.65 \times 10^{-12}$	5
	75	366 \pm 38	2.10 \pm 0.02		4.72 $\times 10^{-9}$ $\pm 5.55 \times 10^{-12}$	4
	91	290 \pm 29	1.32 \pm 0.01		2.99 $\times 10^{-9}$ $\pm 3.39 \times 10^{-12}$	3
304 \pm 22	50	436 \pm 51	2.06 \pm 0.02	2.06 $\times 10^{-9}$ $\pm 1.85 \times 10^{-11}$	6.64 $\times 10^{-9}$ $\pm 9.82 \times 10^{-12}$	3
	75	378 \pm 43	1.55 \pm 0.01		5.09 $\times 10^{-9}$ $\pm 1.68 \times 10^{-11}$	2
	91	339 \pm 33	1.24 \pm 0.01		4.37 $\times 10^{-9}$ $\pm 7.97 \times 10^{-12}$	2

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Table 3. Comparison to literature values of the mass growth rate and the ratio of final mass to seed mass (normalized by sulfate mass) assuming linear reaction rate, from AMS data, after 4 h of reaction of AS 300 nm with glyoxal gas.

RH value	Glyoxal gas concentration (ppb)			Mass Growth rate ($\mu\text{g/p/min}$)		Final mass/ seed mass after 4 h		
	This study	Liggio et al. (2005a)	Galloway et al. (2009)	This study	Liggio et al. (2005b)	This study (assuming linear reaction rate)	Liggio et al. (2005a)	Galloway et al. (2009)
50 %	~2	~5	~67	8.85×10^{-12} $\pm 3.29 \times 10^{-13}$	6×10^{-11}	13.30 ± 0.79	16	0.35
75 %	~2			2.83×10^{-12} $\pm 3.68 \times 10^{-13}$	–	6.98 ± 0.50	–	–
90 %	~2			1.59×10^{-12} $\pm 3.13 \times 10^{-13}$	–	5.68 ± 0.40	–	–

* RH values for Liggio et al. (2005a, b) are 49 % and seed aerosol size of 127.8 nm and for Galloway et al. (2009) the RH is 55 %.

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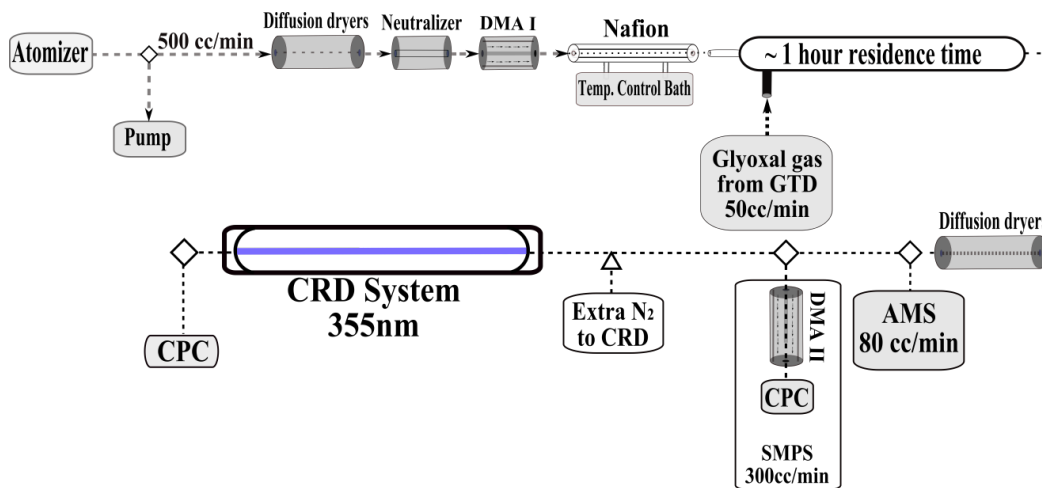


Fig. 1. The experimental setup for the heterogeneous reaction experiments of AS particles with glyoxal. * Acronyms:

- DMA – Differential Mobility Analyzer
- AMS – Aerosol Mass Spectrometer
- CPC – Condensation Particle Counter
- SMPS – Scanning Mobility Particle Sizer
- CRD – Cavity Ring Down (spectrometer)
- Hepa – High efficiency particulate air

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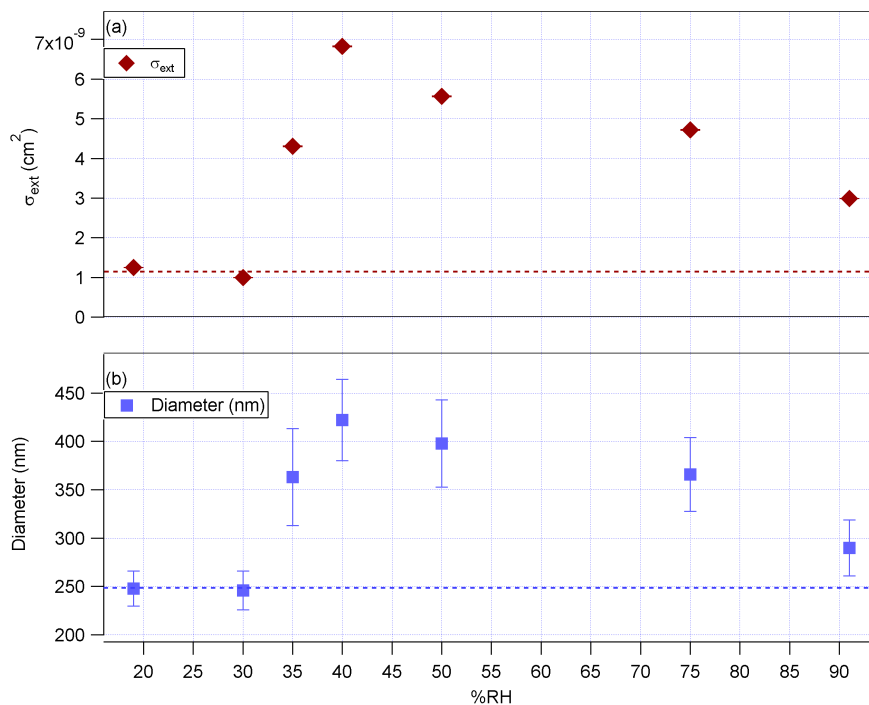


Fig. 2. The size and extinction cross section (σ) vs. RH for AS 250 nm. The results of the ~ 1 h heterogeneous reaction of AS 250 nm particles with gas phase glyoxal at RH values varying from 20% to 90% (while hydrating). **(a):** The extinction cross sections, σ_{ext} (cm²) (red diamonds). **(b):** the final mean diameters (nm) (blue cubes).

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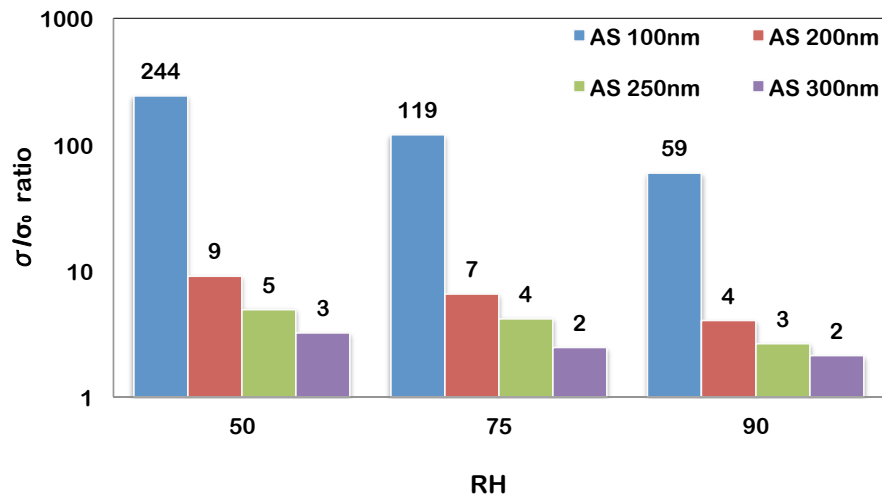


Fig. 3. The ratio of the extinction cross section after reaction (σ ; cm^2) to the initial extinction cross section (σ_0) of the dry AS particles with initial sizes of 100 nm (blue), 200 nm (red), 250 nm (green), and 300 nm (purple) after a ~ 1 h heterogeneous reaction with glyoxal gas at RH values varying from 50 % to 90 %.

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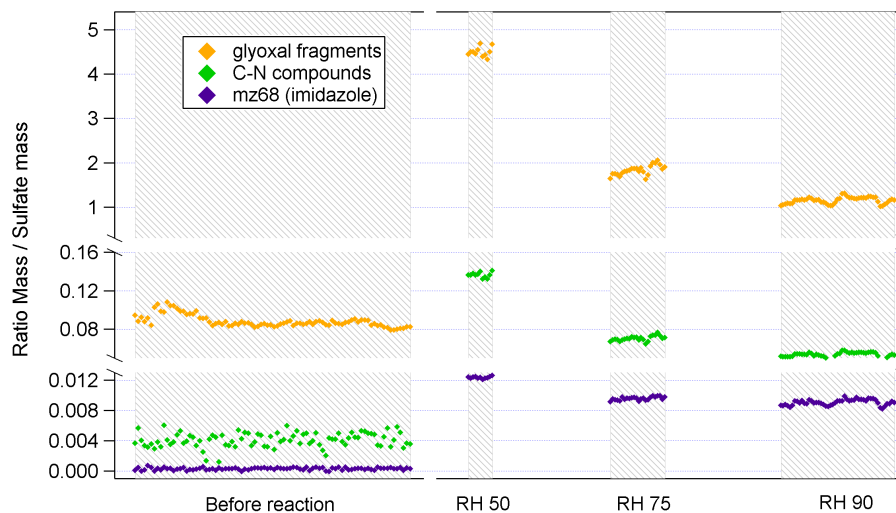


Fig. 4. The total glyoxal fragments (orange diamonds), total C–N compounds (green diamonds) and m/z 68 (imidazole) (purple diamonds) normalized to sulfate (SO_4) mass for the reaction of AS 300 nm particles with glyoxal are presented for 4 periods: **(a)** before the reaction (pure AS particles), **(b)** after ~ 1 h reaction at RH 50 %, **(c)** after ~ 1 h reaction at RH 75 %, and **(d)** after ~ 1 h reaction at RH 90 %.

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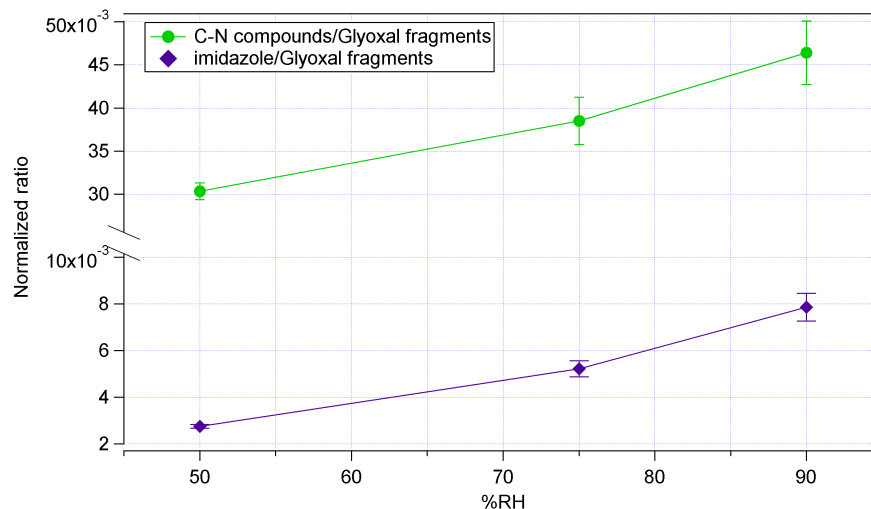


Fig. 5. The ratio of the total C–N compounds to glyoxal fragments (green diamonds) and the imidazoles to glyoxal fragments (purple diamonds) for the reaction of AS 300 nm particles with glyoxal gas at the measured RH values is presented. Both ratios show an increase with increasing RH value and decreasing total mass, since the mass of the glyoxal fragments decreases more drastically with increasing RH than the total C–N compounds and imidazoles. All masses are normalized to sulfate mass.

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