## **Supplementary Material**

"Isoprene suppression of new particle formation in a mixed deciduous forest" by Kanawade, et al.

## **Condensation sink calculations**

The condensation sink (*CS*), defined as the rate at which condensable vapors condense on existing particles, was calculated based on the following equations (Kulmala et al., 2001;Erupe et al., 2010);

$$CS = 2\pi D \cdot 10^{-7} \int_{0}^{\infty} d_{p} \beta_{m}(d_{p}) n(d_{p}) dd_{p} = 2\pi D \cdot 10^{-7} \sum_{i} \beta_{m} d_{p_{i}} N_{i}$$
(S1)

$$\beta_{m} = \frac{1 + Kn}{1 + 0.337Kn + \frac{4Kn}{3\alpha} + \frac{4Kn^{2}}{3\alpha}}$$
(S2)

$$Kn = \frac{2\lambda}{d_p}$$
(S3)

where  $d_p$  is particle diameter (nm),  $\beta_m$  is transition correctional factor (dimensionless) (Fuchs and Sutugin, 1971), Kn is the Knudsen number (dimensionless), D is diffusion coefficient [D=0.104 cm<sup>2</sup> s<sup>-1</sup> under our typical RH, (Hanson and Eisele, 2000)],  $\lambda_v$  is the mean free path of the vapor molecules (123 nm),  $N_i$  is particle number concentration in size bin *i*,  $\alpha$  is the mass accommodation (taken to be unity) and  $n(d_p)$  is particle size distribution function. Aerosol sizes measured at UMBS were predominantly smaller than <100 nm and so we used NDMA data (3-109 nm) to calculate *CS*. We found that this *CS* values were only slightly different from those calculated sizes from NDMA and LDMA together (3-800 nm), with a 30% difference at maximum.

## References

Erupe, M. E., Benson, D. R., Li, J., Young, L.-H., Verheggen, B., Al-Refai, M., Tahboub, O., Cunningham, V., Frimpong, F., Viggiano, A. A., and Lee, S.-H.: Correlation of aerosol nucleation rate with sulfuric acid and ammonia in Kent ohio: An atmospheric observation, 2010, 115, D23216, doi:10.1029/2010JD013942, 2010.

Fuchs, N. A., and Sutugin, A. G.: Highly dispersed aerosol, in: Topics in current aerosol research, edited by: Hidy, G. M., and Brock, J. R., Pergamon, New York, 1971.

Greenberg, J., Guenther, A., Petron, G., Wiedinmyer, C., Vega, O., Gatti, L. V., Tota, J., and Fisch, G.: Biogenic VOC emissions from forested Amazonian landscapes, Global Biogeochem. Cycles, 10, 651-662, 2004.

Hakola, H., Tarvainen, V., Laurila, T., Hiltunen, V., Hellen, H., and Keronen, P.: Seasonal variation of VOC concentrations above a boreal coniferous forest, Atmos. Env., 37, 1623-1634, 2003.

Hanson, D. R., and Eisele, F. L.: First measurement of pre-nucleation molecular clusters, J. Phys. Chem., 104, 2000.

Kesselmeier, J., Kuhn, U., Rottenberger, S., Biesenthal, T., Wolf, A., Schebeske, G., Andreae, M. O., Ciccioli, P., Brancaleoni, E., Frattoni, M., Oliva, S. T., Botelho, M. L., Silva, C. M. A., and Tavares, T. M.: Concentrations and species composition of atmospheric volatile organic compounds (VOCs) as observed during the wet and dry season in Rondonia (Amazonia), J. Geophys. Res., 107, D20, doi:10.1029/2000JD000267, 2002.

Kim, S., Karl, T., Helmig, D., Daly, R., Rasmussen, R., and Guenther, G.: Measurement of atmospheric sesquiterpenes by proton transfer reation-mass spectrometry (PTR-MS), Atmos. Meas. Tech., 2, 99-112, 2009.

Kuhn, U., Andreae, M. O., Ammann, C., Araujo, A. C., Brancaleoni, E., Ciccioli, P., Dindorf, T., Frattoni, M., Gatti, L. V., Ganzeveld, L., Kruijt, B., Lelieveld, J., Lloyd, J., Meixner, F. X., Nobre, A. D., Poschl, U., Spirig, C., Stefani, P., Thielmann, A., Valentini, R., and Kesselmeier, J.: Isoprene and monoterpene fluxes from Central Amazonian rainforest inferred from towerbased and airborne measurements, and implications on the atmospheric chemistry and the local carbon budget, Atmos. Chem. and Phys., 7, 2855-2879, 2007.

Kulmala, M., dal Maso, M., Makela, J. M., Pirjola, L., Vakeva, M., Aalto, P., Miikkulainen, P., Hameri, K., and O'Dowd, C. D.: On the formation, growth, and composition of nucleation mode particles, Tellus B., 53, 479-490, 2001.

Ortega, J., Helmig, D., Guenther, A., Harley, P., Pressley, S., and Vogel, C.: Flux estimates and oh reaction potential of reactive biogenic volatile organic compounds (BVOCs) from a mixed northern hardwood forest, Atmos. Environ., 41, 5479-5495, 2007.

Sellegri, K., Hanke, M., Umann, B., Arnold, F., and Kulmala, M.: Measurements of organic gases during aerosol formation events in the boreal forest atmosphere during QUEST, Atmos. Chem. Phys., 5, 373-384, 2005.

Spirig, C., Guenther, A., Greenberg, J. P., Calanca, P., and Tarvainen, V.: Tethered balloon measurements of biogenic volatile organic compounds at a boreal forest site, Atmos. Chem. Phys., 4, 215-229, 2004.

Table S1. Comparison of R (ratio of isoprene to MT carbon) calculated from emission flux and concentrations. This table intends to show that R values calculated from emission flux and concentrations show the same trend as depicted in Table 1.

	Locations	Michigan format	Amazon forestb	Einland forest
Parameters		whenigan torest	Amazon lorest	riniand iorest
Isoprene	emission flux $(mgC m^{-2} h^{-1})$	1.32±0.7 <sup>a</sup> 3.9±0.8 (Kim et al., 2009) <sup>c</sup> 2.5 (Ortega et al., 2007)	3.4±3.6 (Kuhn et al., 2007) 1.9-8.8 (Greenberg et al., 2004) <sup>c</sup>	0.03 (Spirig et al., 2004) <sup>c</sup>
	concentration (ppbv)	4.4±0.36 <sup>a</sup> 4.8 (Kim et al., 2009)	9.0 (Kesselmeier et al., 2002) 2.2-9.8 (Greenberg et al., 2004)	0.05-0.25 (Hakola et al., 2003) 0.1-0.2 (Sellegri et al., 2005)
MT	$(mgC m^{-2} h^{-1})$	0.05±0.02 <sup>a</sup> 0.18±0.07 (Kim et al., 2009) <sup>c</sup> 0.11 (Ortega et al., 2007)	0.36±0.58 (Kuhn et al., 2007) 0.22-0.76 (Greenberg et al., 2004) <sup>c</sup>	0.16 (Spirig et al., 2004) <sup>c</sup>
	(ppbv)	0.35±0.24 <sup>a</sup> 0.2 (Kim et al., 2009)	0.6-0.7 (Kesselmeier et al., 2002) 0.27-0.69 (Greenberg et al., 2004)	0.1-0.3 (Hakola et al., 2003) 0.2-0.4 (Sellegri et al., 2005)
<i>R</i> (from emission flux)		26.4±4.5 <sup>a</sup> 21.4±16 (Kim et al., 2009) 22.0 (Ortega et al., 2007)	9.4±6.2 (Kuhn et al., 2007) 15.2 (Greenberg et al., 2004)	0.18 (Spirig et al., 2004)
R (from concentration) <sup>d</sup>		6.3±0.8 <sup>a</sup> 12.0 (Kim et al., 2009)	7.0 (Kesselmeier et al., 2002) 6.3 (Greenberg et al., 2004)	0.37 (Hakola et al., 2003) 0.25 (Sellegri et al 2005)

<sup>a</sup>this study.

<sup>b</sup>data from wet season only.

<sup>c</sup> for unit conversion, the ratio of compound to carbon of 1.13 was used.

<sup>d</sup>the ratio calculated from isoprene and MT concentrations (ppbv) was  $([Isoprene] \times 5)/([MT] \times 10)$ , where 5 and 10 indicate the number of carbon atoms in a isoprene and MT molecule, respectively.



Figure S1. The meaured (a) temperature (red symboled line) and relative humidity (blue symboled ine), and (b) wind direction (grey symboled line) and wind speed (dark cyan symboled line) shown as a function of time on July 16 (solid circles) and August 2 (open diamonds).



Figure S2. Time evolution of the key atmospheric parameters prescribed or simulated in the IIN box model for July 16 event. (a) Temperature (black asterisks) RH (black open triangles). (b) OH (blue open diamond), SO<sub>2</sub> (black plus signs), and H<sub>2</sub>SO<sub>4</sub> (red asterisks). (c) The particle size distributions simulated from the IIN microphysical box model along the backward air trajectory for July 16 event, using prescribed key atmospheric parameters. The nucleation simulations were performed forward, whereas HYSPLIT backward trajectory calculations were backward; thus the time t = 0 corresponds to the time the model simulations start (near the power plant) and t = 35 hours corresponds to the time at the forest measurement site. Time evolution of the key atmospheric parameters prescribed or simulated in the IIN box model for August 2 event were also show similar results.