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# Acetylene C<sub>2</sub>H<sub>2</sub> retrievals from MIPAS data and regions of enhanced upper tropospheric concentrations in August 2003

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# Abstract

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Acetylene ( $C_2H_2$ ) volume mixing ratios (VMRs) have been successfully retrieved from MIPAS Level 1B radiances during August 2003. The data presented here contain most information between 300 hPa and 100 hPa based on the averaging kernels, with information also at lower altitude levels (up to 500 hPa) albeit with some influence from the 5 300 hPa level. In our  $C_2H_2$  retrievals, data at altitude levels above 100 hPa must be treated with caution. Systematic errors are less than 10% at the upper levels but can reach higher levels at 300 hPa in the tropics due to water vapour influences. Random errors per point are less than 15% at lower pressure levels and are closer to 30% at 100 hPa.

Global distributions of both the absolute  $C_2H_2$  and ratios to MOPITT 150 hPa retrievals of carbon monoxide (CO) confirm some significant features for this important hydrocarbon in a characteristic summer month (August 2003), showing tight correlations regionally but globally emphasising the differences between sources and lifetimes

- of CO and  $C_2H_2$ . The ratios to CO are estimated to be accurate to approximately 10%. 15 A strong isolation of C<sub>2</sub>H<sub>2</sub> within the Asian monsoon anticyclone is observed, evidencing convective transport into the upper troposphere, horizontal advection within the anticyclone at 200 hPa, distinct but measurable gradients at the westward edge of the vortex and formation of a secondary dynamical feature over the Asian Pacific. The data
- for  $C_2H_2$  strongly support evidence for a strong isolated core to the anticyclone with 20 distinct gradients surrounding this core. Within this region, there is a relatively lower correlation of C<sub>2</sub>H<sub>2</sub> and CO suggesting difference in injection ratios or more likely due to expected chemical processing.

A second strong feature to the global distributions is observed in the enhancement and outflow of biomass burning from Africa at 200 hPa, both north-westward and east-25 ward from 10° S. The easterly flow shows high C<sub>2</sub>H<sub>2</sub> ratios to CO which have significantly decayed before reaching Australia. In the biomass burning regions,  $C_2H_2$  and CO are relatively tightly correlated. C<sub>2</sub>H<sub>2</sub> enhancements are observed to penetrate to



lower altitudes in the African biomass outflow in this month compared to uplift observed in the Asian monsoon anticyclone region.

Overall, the data show the distinctive nature of C<sub>2</sub>H<sub>2</sub> distributions, confirm in greater detail than previously possible features of hydrocarbon enhancements in the upper troposphere and highlight the future use of MIPAS hydrocarbon data for testing model transport and OH decay regimes in the middle to upper troposphere.

# 1 Introduction

The burning of vegetation, both living and dead, can release large quantities of gases into the atmosphere. Biomass burning is therefore a major source for the injection of trace gases into the atmosphere. Biomass burning and combustion were thought to be the major sources of acetylene ( $C_2H_2$ ) (Hegg et al., 1990; Blake et al., 1996; Whitby and Altwicker, 1978). However, recent work also suggests that biofuel emissions may contribute as the dominant source of  $C_2H_2$  with Xiao et al. (2007) estimating that as much as half of the global  $C_2H_2$  source is due to biofuel (3.3 Tg yr<sup>-1</sup>) with the remainder

- <sup>15</sup> being due to fossil fuel (1.7 Tg yr<sup>-1</sup>) and biomass burning (1.6 Tg yr<sup>-1</sup>). Streets et al. (2003) found from their emission study that 45% of the C<sub>2</sub>H<sub>2</sub> emissions in Asia were due to biofuel. It should be noted that Streets et al. (2003) make a clear distinction between what they term open biomass burning (i.e. forest fires) and the combustion of biofuels (wood, crop residue, dung, etc.) in domestic cooking and heating. The importance of biofuel as a C H, source has also hear identified in Africa, where Battachi
- <sup>20</sup> tance of biofuel as a  $C_2H_2$  source has also been identified in Africa, where Bertschi et al. (2003) state that in their study of Zambia, the emission of  $C_2H_2$  due to biofuel was significantly greater than that due to savanna biomass burning.

As it is estimated that Asia accounts for around 70% of global biofuel emissions as well as being a significant source for biomass burning and industrial emissions, it follows that Asia is the major source of  $C_2H_2$  emissions. Africa is also an important source of  $C_2H_2$  emissions, whether through biomass burning, fossil fuel or biofuel emissions. The work presented here is intended to confirm the large  $C_2H_2$  concentrations emanating from these regions.



It has been shown that the concentration of  $C_2H_2$  is strongly correlated to CO and the ratio of the two can be used as a robust tracer for the photochemical evolution of the air mass since it last encountered a combustion source (Xiao et al., 2007) due to the relatively long life-times involved and the fact that the correlation remains strong over this period. This allows the relative photochemical age of biomass plumes to be estimated as well as the amount of photochemical processing that the plume has undergone.

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In addition to the interest in  $C_2H_2$  as a tracer for the transport of biomass/biofuel burning,  $C_2H_2$  in itself plays an important role in the formation of glyoxal (CHOCHO) and has implications for the production of secondary organic aerosol (SOA) as discussed by Volkamer et al. (2009). For this reason, the ability to retrieve global distributions of  $C_2H_2$  may also prove important for future air quality and climate simulations, hence this study.

Whilst vertical profiles of  $C_2H_2$  have been observed from aircraft measurements (Smyth et al., 1996; Talbot et al., 2003) and more recently from satellite observations (Rinsland et al., 2005), there is still some uncertainty as to the distribution of  $C_2H_2$  due to the poor understanding of the emissions from the large number of varying sources (Streets et al., 2003). Recent work by Park et al. (2008) has observed enhancements of  $C_2H_2$  inside the Asian monsoon anticyclone from space with the Atmospheric Chem-

- istry Experiment (ACE) FTIR instrument (Bernath et al., 2005). Due to the solar occultation technique used there are relatively few observations over the tropical region, with the majority of ACE observations occuring at polar latitudes. This sparse spatial sampling makes it necessary, particularly in the tropics, to average occultation data over multiple months and years.
- In this paper, a description is provided of global retrievals of  $C_2H_2$  performed from MIPAS infrared limb emission spectra for the upper troposphere. These provide greater temporal and spatial resolution than the ACE instrument albeit with a much lower signal to noise ratio (SNR) compared to the solar occultation method employed by ACE. Despite the higher noise, high spatial resolution distributions of  $C_2H_2$  are successfully



determined on sub-monthly timescales. The results of this paper focus on the distribution for August 2003, a characteristic summer month where the influence of the Asian monsoon anticyclone is observed in conjunction with considerable fire activity in Africa.

# 2 The MIPAS instrument

- The Michelson Interferometer for Passive Atmospheric Sounding (MIPAS) is a core in-5 strument onboard the ENVIronmental SATellite (ENVISAT) Satellite launched in March 2002. Measurements are performed by mid-infrared limb emission sounding of the atmosphere with a nominal mode (March 2002–March 2004) of 17 tangent altitudes at 68 km, 60 km, 52 km, 47 km, 42 km and then every 3 km until the final mid-tropospheric measurement at 6 km. MIPAS is a Fourier transform infrared spectrometer with a 0.025 unapodised spectral resolution and measures over a large spectral range, from 685 cm<sup>-1</sup> to 2410 cm<sup>-1</sup>, observing the region of the atmospheric spectrum where there are a variety of molecules with vibration-rotation bands with well-defined absorption lines (Fischer et al., 2008). Due to this, MIPAS is able to detect a wide range of trace species and is operationally used to retrieve pressure, temperature,  $O_3$ ,  $H_2O$ ,  $CH_4$ , HNO<sub>3</sub>, N<sub>2</sub>O and NO<sub>2</sub> from 6 to 68 km (Raspollini et al., 2006). In addition, various non-operational trace species have been detected and retrieved from MIPAS spectra including PAN (Moore and Remedios, 2010; Glatthor et al., 2007), HCFC-22 (Moore and Remedios, 2008) and  $C_2H_6$  (von Clarmann et al., 2007).
- <sup>20</sup> This work utilises the full resolution L1B MIPAS spectra from August 2003 where the instrument was measuring at 17 vertical levels with an unapodised spectral resolution of 0.025 cm<sup>-1</sup> before technical issues resulted in a reduction in resolution to 0.06 cm<sup>-1</sup> after August 2004.

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## 3 Determination of C<sub>2</sub>H<sub>2</sub> in the upper troposphere

The objective of this work was to examine the  $C_2H_2$  spectral signatures in MIPAS L1B spectra with a full optimal estimation retrieval algorithm in order to identify regions of high  $C_2H_2$  volume mixing ratios (VMRs). These regions will be discussed in the context of their production and transport due to biomass burning.

The  $C_2H_2$  infrared signature in the microwindow used for this work is outlined in Fig. 1 where the contributions from each species to the total radiance (black) are shown. The Oxford Reference Model (RFM) (Dudhia, 2005b) was used to model the spectra as observed by MIPAS. The RFM is a line-by-line radiative transfer model based on the GENLN2 model (Edwards, 1992) and can simulate the MIPAS infrared spectra taking into account the instrument lineshape and field of view. The spectroscopic information is taken from the HITRAN2004 spectral database (Rothman et al., 2005; Jacquemart et al., 2003) and in this case the simulation uses a standard atmospheric climatology (Remedios et al., 2007) with an enhancement of the  $C_2H_2$  profile typical of biomass 15 burning taken from Rinsland et al. (2005).

In Fig. 1, one line of the  $v_5$  band of  $C_2H_2$  (purple line) can be seen to be the dominant feature in the 776.0 cm<sup>-1</sup> to 776.15 cm<sup>-1</sup> spectral range and unlike other  $C_2H_2$  spectral lines in this region at 755 cm<sup>-1</sup>, 762 cm<sup>-1</sup> and 766.7 cm<sup>-1</sup> the line is not masked by a strong ozone feature (Rinsland et al., 1998). This clear absorption line allows the unambiguous identification of enhanced  $C_2H_2$  in MIPAS L1B spectra in this microwindow.

## 3.1 Retrieval method

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A full optimal estimation approach was taken to retrieve the  $C_2H_2$  VMR. The MIPAS Orbital Retrieval using Sequential Estimation or MORSE (Dudhia, 2005a) scheme is an optimal estimation scheme developed by the University Of Oxford based on the approach taken by Rodgers (2000) and has recently been used to successfully retrieve peroxyacetyl nitrate (PAN) from MIPAS observations by Moore and Remedios (2010) where further details of the retrieval scheme are provided.



The MORSE approach involves the inversion of the measured spectral radiances (y) in order to obtain the best solution for the retrieved parameters (the state vector x) with the associated random error, e. The relationship between the retrieved parameters and the measured radiances is defined by a forward model F(x), in this case the RFM, which performs the radiative transfer calculations to calculate the expected radiance given the input parameters.

As described by Rodgers (2000), the solution to the above is constrained with respect to the a priori information (in this case standard climatologies) by the a priori covariance and the solution is found by minimising the cost function  $\chi^2$  via an iterative procedure where the state vector (x) is updated after each iteration.

This iterative procedure for the state vector is

$$\boldsymbol{x}_{i+1} = \boldsymbol{x}_i - (\mathbf{S}_{\mathbf{a}}^{-1} + \mathbf{K}_i^T \mathbf{S}_{\mathbf{y}}^{-1} \mathbf{K}_i + \gamma \mathbf{D})^{-1} (-\mathbf{K}_i^T \mathbf{S}_{\mathbf{y}}^{-1} [\boldsymbol{y} - \mathbf{F}(\boldsymbol{x}_i)] + \mathbf{S}_{\mathbf{a}}^{-1} [\boldsymbol{x}_i - \boldsymbol{x}_{\mathbf{a}}])$$
(1)

with the cost function taking the form

$$\chi^{2} = (\boldsymbol{y} - \boldsymbol{K}\boldsymbol{x})^{T} \boldsymbol{S}_{\boldsymbol{y}}^{-1} (\boldsymbol{y} - \boldsymbol{K}\boldsymbol{x}) + (\boldsymbol{x} - \boldsymbol{x}_{a})^{T} \boldsymbol{S}_{\boldsymbol{a}}^{-1} (\boldsymbol{x} - \boldsymbol{x}_{a})$$

<sup>15</sup> **S**<sub>a</sub> is the a priori covariance matrix where diagonal values equivalent to a 1000% covariance of the C<sub>2</sub>H<sub>2</sub> VMR along with a correlation length of 3 km were necessary to account for the large variations in VMR over biomass burning regions. **S**<sub>y</sub> is the measurement error covariance and was taken to be equivalent to a conservative estimate of the MIPAS spectral noise in Band A of 40 nW/(cm<sup>2</sup> sr cm<sup>-1</sup>) (Fischer et al., 2008); note that the noise in apodised spectra can reduce to lower than 20 nW/(cm<sup>2</sup> sr cm<sup>-1</sup>).

The reduction of the cost function at each iteration is used to test for convergence and the final value is useful in determining whether the obtained solution is a sensible value. Ideally, the value for  $\chi^2$  should be equal to the number of degrees of freedom in the retrieval.

In addition, in order to accurately retrieve  $C_2H_2$  it was necessary to obtain accurate VMRs of the other interfering species in the spectral range. Therefore, prior to retrieving  $C_2H_2$ , MORSE was used to retrieve profiles of pressure/temperature,  $H_2O$ ,  $O_3$  and

(2)

HNO<sub>3</sub> respectively which allowed an improvement to the climatology provided as a priori knowledge for subsequent retrievals.

# 3.2 Retrieval setup

 $C_2H_2$  VMRs were retrieved for August 2003 for the 9 km to 30 km nominal MIPAS altitudes. The results presented in this paper are primarily for the upper-tropospheric limb measurements at a 12 km nominal altitude with a corresponding average pressure of 200 hpa. The retrieval used MORSE L1C files which were created by extracting the relevant radiance information for the  $C_2H_2$  microwindow from the MIPAS L1B spectra and applying a Norton-Beer Medium apodisation. The retrieved data was filtered to remove points where the retrieval quality was not deemed to be satisfactory using the  $\chi^2$  parameter and a cloud index (Spang et al., 2004) value as indications of the quality of the retrieval. Only retrieved data where the cloud index was greater than 4.0 (corresponding to clear-sky conditions) and the  $\chi^2$  value was less than 2.0 were used. As a further indication of the retrieval quality, comparisons were performed to confirm that there were only small differences between the measured and simulated spectra in

the retrieval process. These residuals are shown in Fig. 2 where average residuals of less than 40 nW/(cm<sup>2</sup> sr cm<sup>-1</sup>) were found, consistent with MIPAS measurement noise. Encouragingly the mean bias in the residuals (dashed line in Fig. 2) is also small.

# 3.3 Error analysis

- Figure 3 shows the final errors calculated for the retrieval, showing the total calculated error (solid line), along with the random component (dotted line) and the systematic component (dashed line). The systematic model parameter errors shown for each variable are  $1\sigma$  values calculated from measured biases in the MIPAS data and were taken from Fischer et al. (2008). Uncertainties of 2% were assumed for pressures and
- <sup>25</sup> 1 K for temperatures. Water vapour was assumed to have an uncertainty of 20%, with 10% for  $O_3$  and HNO<sub>3</sub> and 5% for CCl<sub>4</sub>. The errors associated with the Gain, Shift



and Spread were calculated using perturbations to the instrument line shape of 4%, 2% and 4% respectively and an offset error of  $2 \text{ nW/(cm}^2 \text{ sr cm}^{-1})$  was assumed, with these figures taken from Fischer et al. (2008). The error relating to the C<sub>2</sub>H<sub>2</sub> spectroscopy was conservatively assumed to be 5% with Jacquemart et al. (2003) giving an error of between 2% and 5%.

The total error on the retrieved VMR was estimated to be 15.6% at 200 hpa, with the majority being related to the random retrieval noise (14%) and only 6.9% related to the systematic errors. Of those systematic errors, the uncertainty in the  $C_2H_2$  spectroscopy and instrument line-shape effects along with the uncertainty in the pressure and temperature were the dominant contributors; there is only a minimal contribution from the uncertainty in the VMRs of other species in this spectral window. As the  $C_2H_2$ profile decreases rapidly with altitude, above 120 hpa there is a greatly reduced signal and hence the random error increases substantially. At these high altitudes the  $O_3$ and HNO<sub>3</sub> uncertainties add to the overall error whilst at lower altitudes (300 hpa and above) the uncertainty in water vapour VMP has a much larger effect.

above), the uncertainty in water vapour VMR has a much larger affect.

# 4 C<sub>2</sub>H<sub>2</sub> retrieval results

## 4.1 Global behaviour

Initially in the results that follow we concentrate on the 200 hpa level, at which the errors are smallest and random error dominates single profile retrievals.

Figure 4a shows all of the data at 200 hpa for the MORSE retrievals which pass the cloud and quality filters during August 2003. The data are averaged onto a regular grid with a 5° resolution using a distance-weighted approach to calculate the mean; a scaling distance r of 10° from the grid box centre is adopted with the weighting contribution following a  $1 - x^2/r^2$  relationship, with the contribution reaching zero at 10° from the box centre (i.e. at x = r). This method maintained the sharper gradients reasonably well as well as producing a robust average of the individual observation



values. The data are not interpolated vertically to avoid interpolation errors and the average pressure of each set of retrieval levels is given instead, following the procedure of Moore and Remedios (2010).

The plot shows some key features of the C<sub>2</sub>H<sub>2</sub> distributions observed in the August timeframe. Strong signals of source regions and transport of higher C<sub>2</sub>H<sub>2</sub> mixing ratios are seen over the African biomass burning region and the Pacific outflow from Asia. In addition, a strong enhancement of C<sub>2</sub>H<sub>2</sub> is observed over the Middle-Eastern region and clearly shows the enhanced C<sub>2</sub>H<sub>2</sub> VMRs due to the Asian monsoon anticyclone as also reported by Park et al. (2008). Only a small signature is seen in the vicinity of the Amazon. It is also noticeable that the distributions in the most northerly latitudes, whilst greater than those in the corresponding southerly latitudes, are much lower in mixing ratio than in the extended African and Asian outflows.

# 4.2 C<sub>2</sub>H<sub>2</sub> – CO relationships

The VMR of carbon monoxide is expected to be heavily correlated to that of  $C_2H_2$ (Wang et al., 2004) with the ratio between the two acting as an indicator for the relative age of the air mass and the extent of atmospheric processing it has undergone (Xiao et al., 2007; Smyth et al., 1996, 1999). Hence in order to examine the  $C_2H_2$  – CO relationship further the 150 hpa MOPITT (Measurement Of Pollution In The Troposphere) Level 2 Version 3 data for the same time period has also been investigated. Maps of the concentrations of the two gases are shown in Fig. 4.

The comparison of the two maps largely supports the interpretation that there is much in common between the two gases with respect to major inputs to the upper troposphere. However, there are also interesting differences. The C<sub>2</sub>H<sub>2</sub> mixing ratios do not show strong features over Northern high latitudes and the Amazon unlike the

<sup>25</sup> CO data. In the Pacific region, the C<sub>2</sub>H<sub>2</sub> appears to provide a tighter definition of the local transport, close to Asia, rather than the CO which appears more even in VMR, indicating that C<sub>2</sub>H<sub>2</sub> distributions may prove useful in further constraining particular transport pathways. These aspects indicate that C<sub>2</sub>H<sub>2</sub> data could provide some significant and complementary information to those contained in CO distributions.



Some care should be taken in interpreting detailed comparisons of the two data sets as whilst MIPAS is observing at a 12 km nominal altitude with an approximate 3 km field of view and a mean pressure of 200 hpa, MOPITT is a nadir-sounding instrument with the retrieved vertical levels a result of its broad averaging kernels in the troposphere.

- <sup>5</sup> Typically the MOPITT CO 150 hpa averaging kernels have sensitivity between 400–100 hpa. We have estimated, by applying averaging kernels to typical profiles, that the effect of vertical resolution may lead to a systematic underestimation of the  $C_2H_2/CO$  ratio of approximately 10% but that a strong correlation would still be expected between the relevant trace gas data from these two instruments.
- <sup>10</sup> The strong African biomass burning signature is clearly observed in both datasets, indicating that a  $C_2H_2$  enhancement due to biomass burning is present. In addition, there are strong CO and  $C_2H_2$  enhancements observed over Asia, particularly the Asian outflow into the Northern Pacific Ocean. Although the data in the South-East Asia region is more sparse due to the cloud associated with the Asian monsoon, en-<sup>15</sup> hancements are observed over this region where data is present.

It is the strong enhancement of both CO and  $C_2H_2$  in the Middle East region which is of particular interest. Rather than local production, the probable cause of this feature is long-range transport from the convective region over South-East Asia via the Easterly Jet associated with the Asian monsoon anticyclone into the Middle East.

- <sup>20</sup> It is expected that  $C_2H_2$  is highly correlated to CO and the  $C_2H_2/CO$  ratio can provide important information relating to the age of the biomass plume and the amount of atmospheric processing it has undergone (Xiao et al., 2007). For this purpose, the 200 hpa MORSE  $C_2H_2$  VMRs were correlated against the 150 hpa MOPITT CO VMRs for the globally averaged monthly data and the result is shown in Fig. 5a with the red points indicating the mean values of the data along with their standard deviation,
- separated into 2 ppbv bins. This figure clearly shows a strong correlation between the two VMRs (with r = 0.76) but at the same time, there appear to be two distinct domains within this correlation with distinctly different gradients.



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For the region containing the eastward Asian outflow into the Pacific Ocean (10° N to 40° N, 130° E to 110° W), a correlation of 0.91 is observed (Fig. 6c) with a steep gradient (5.25). This strong correlation and steep gradient suggest that there is considerable 25 transport between Asia and North America and that this occurs on a relatively short time-scale before the  $C_2H_2$  has time to photochemically age.

American gradient is the smallest of the four regions studied here.

- very different. The highest  $C_2H_2$  VMRs of 150 pptv are observed directly over the African source region, whereas the Amazon region shows much smaller C<sub>2</sub>H<sub>2</sub> VMRS for the same CO mixing ratios. In the African sector, where significant transport into the Atlantic is observed, the gradient of the correlation is significantly higher (3.17) than in South America (1.41) indicating either very different source ratios of the two gases 20 or else a very different photochemical regime. It is interesting to note that the South
- 0.81 (Fig. 6a) and 0.80 (Fig. 6b) respectively. However, the regions are otherwise
- The major biomass burning regions of Southern Africa (30° S to 10° N, 30° W to 80° E) and South America (30° S to 10° N, 180° W to 30° W) both show strong correlations, 15
- 0.67. The tail of low  $C_2H_2$  values compared to high CO values in the Northern Hemisphere plot stems from the high latitude data where CO is enhanced in a number of features whereas C<sub>2</sub>H<sub>2</sub> is not.

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- to the lack of significant sources of C<sub>2</sub>H<sub>2</sub> or long-range transport south of the equator whereas not only does the Northern Hemisphere contain the strong Asian  $C_2H_2$ sources but there is a considerably higher background of CO in the Northern Hemisphere when compared to the Southern Hemisphere resulting in a lower correlation of 10
- To examine this further, the correlation was performed separately for the Northern Hemisphere (Fig. 5b) and the Southern Hemisphere (Fig. 5c). From the C<sub>2</sub>H<sub>2</sub>-CO correlations it becomes apparent that the two domains seen in the global correlation are due to the combination of the differences in C<sub>2</sub>H<sub>2</sub> sources and transport between the two hemispheres. The strong correlation of 0.92 for the Southern Hemisphere is due



Finally, the correlation for the Asian monsoon anticyclone region  $(10^{\circ} \text{ N to } 40^{\circ} \text{ N}, 30^{\circ} \text{ E to } 110^{\circ} \text{ E})$  has a value of 0.59 (Fig. 6d). Unlike the previous examples, the region here is largely isolated with convective injection toward the east (see Sect. 4.4). Although the correlation is reasonable, it appears that there is considerable variability in <sup>5</sup> either source ratios of C<sub>2</sub>H<sub>2</sub> to CO or in the photochemical age of the C<sub>2</sub>H<sub>2</sub> observed.

Having illustrated that the  $C_2H_2$  and CO relationship maintains a largely strong correlation and remains preserved through the processes of convection and mixing as suggested by Smyth et al. (1996), the  $C_2H_2/CO$  ratio can be used to further analyse the transport dynamics from the combustion source regions and to explore further the

- <sup>10</sup> transport out of the injection regions into the upper troposphere. The gridded  $C_2H_2$ and CO values (as shown in Fig. 4) are used to calculate the  $C_2H_2$ /CO ratio (Fig. 7). The ratio is expected to be highest over source regions where  $C_2H_2$  and CO are both produced from the combustion process and diminish with time as the photochemical age of air increases with atmospheric processing. Due to the difference in lifetimes of
- $_{15}$  C<sub>2</sub>H<sub>2</sub> (approximately 2 weeks) and CO (approximately 2 months), for a high C<sub>2</sub>H<sub>2</sub>/CO ratio to exist, the C<sub>2</sub>H<sub>2</sub> must be relatively young. Hence, if a high C<sub>2</sub>H<sub>2</sub>/CO ratio is observed (i.e. greater than 2 pptv/ppbv, Smyth et al., 1996) then either the observation is over a region where combustion is taking place or the observed plume has been transported from such a source region in a relatively short space of time.
- It is therefore possible to identify clear transport mechanisms from the C<sub>2</sub>H<sub>2</sub>/CO ratio in Fig. 7. The high ratio located over southern Africa is maintained as the plume travels north-west out into the Atlantic Ocean. In contrast, the eastward transport of this southern African plume towards the Indian Ocean and Australia would appear to be relatively slower, allowing the plume time to undergo a significant amount of atmospheric processing. A further transport pathway is observed from Asia out into the Pacific Ocean and is related to the tongue of the Asian Monsoon anticyclone that extends out over the Pacific. This process is known to transport air relatively guickly
- and this is confirmed by the fact that a high  $C_2H_2/CO$  ratio is maintained out over the ocean where there are no source regions present.



The fact that the  $C_2H_2/CO$  ratio agrees so closely with the known transport mechanisms in these regions provides substantial confidence in the approach taken. Furthermore, it allows the Easterly Jet transport from Asia into the Middle East to be discussed in the context of the  $C_2H_2/CO$  ratio. The high  $C_2H_2/CO$  ratio extends over the whole of the region between Asia and the Middle East and again clearly shows the strong 5 chemical isolation of the monsoon. As this region is entirely over land some care must be taken to distinguish between a maintained high ratio due to fast transport or the alternative that the ratio remains high due to there being various combustion sources distributed over the whole region. There is no CO enhancement in the MOPITT surface data in this region and the correlation for this region (Fig. 6d) remains relatively 10 strong. Along with this, the fact that a clear gradient exists in the distribution of the high  $C_2H_2$  and CO values (Fig. 4) between the Middle East and Northern Africa provides confidence that it is the Easterly Jet transport that is being observed. The anticyclone continues to transport newly formed  $C_2H_2$  and CO into the region from the biomass and biofuel sources in Asia, leaving no time for substantial photochemical processing 15 to occur before it is circulated back towards the source region. It is this circulation due to the monsoon anticyclone which causes the persistence of the high  $C_2H_2/CO$  which

To summarise this section, the  $C_2H_2$ -CO relationship has been examined for a va-<sup>20</sup> riety of regions and it was found that strong correlations persist between the two species following transport into the upper troposphere. This in turn justified the use of a  $C_2H_2/CO$  ratio in order to examine characteristics of the observed transport pathways as well as identifying the chemical isolation due to the Asian monsoon anticyclone. The observation of ratios greater than 2 pptv/ppbv clearly suggests both localised injection

is observed.

and fast photochemical transport. In the following sections we examine in more detail the C<sub>2</sub>H<sub>2</sub> distributions relating to African biomass burning and the Asian monsoon anticyclone.



# 4.3 African biomass burning

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In order to examine some of the features of the  $C_2H_2$  distributions in more detail, crosssections have been plotted for both the meridional and zonal distributions. The NCEP mean tropopause pressure for August 2003 is shown as a red dashed line in order to give an indication of the location of the tropopause and to allow some preliminary discussion on whether any tropopause penetration into the stratosphere is being observed.

The C<sub>2</sub>H<sub>2</sub> distribution relating to the African biomass burning is investigated through the use of zonal cross-sections taken along the 5° N and 10° S lines of latitude, passing through the African biomass burning C<sub>2</sub>H<sub>2</sub> enhancement (Fig. 8). The cross-section taken along 5° N shows how the large biomass burning signal is largely confined to below 250 hpa, well below the tropopause, with the outflow of C<sub>2</sub>H<sub>2</sub> generally less than 100 pptv. The cross-section taken further south at 10° S shows a similar distribution but this time with the strongest enhancement of C<sub>2</sub>H<sub>2</sub> much more confined and ex-

tending somewhat higher into the atmosphere suggesting more intense biomass burning in this region associated with a greater uplift of  $C_2H_2$ . Evidence for outflow of this biomass burning both towards the east and west is also apparent with  $C_2H_2$  values above 100 pptv observed over both the Indian and Atlantic Oceans.

To examine this further, meridional cross-sections of the retrieved  $C_2H_2$  distribution were taken along the 20° E (Fig. 9a), 40° E (Fig. 9b) and 75° E (Fig. 9c) longitude lines.

The cross-section at 20° E (Fig. 9a) passes through the western edge of the anticyclone in the Northern Hemisphere and through the large African biomass burning region in the Southern Hemisphere. The vertical transport of the  $C_2H_2$  enhancement related to the African biomass burning is found to be relatively weak with the major enhancement ( $C_2H_2 > 150$  pptv) only reaching 250 hpa. It is worth noting that in the stratospheric data, there are apparently weak enhancements of  $C_2H_2$  above high values in the troposphere due to the African biomass region. Correlations suggests that these are a retrieval artefact and no other evidence of enhanced stratospheric values could be observed. This subject is returned to in the next section.



This section has shown how the  $C_2H_2$  retrieval from MIPAS is capable of providing information not only on the African biomass burning sources of  $C_2H_2$  but also on the behaviour in terms of uplift and transport away from these sources.

#### 4.4 The Asian monsoon anticyclone

- In contrast to Africa, the strong C<sub>2</sub>H<sub>2</sub> enhancement is observed as high as 150 hpa, reaching the tropopause in the meridional cross-section inside the core of the anticyclone at 75° E (Fig. 9c), This suggests that the observed uplift from the monsoon convection is considerably stronger or occurs on a faster timescale compared to the weaker convection observed over Africa. The meridional distributions and outgoing
  longwave radiation (OLR) (Fig. 12b) show that the main monsoon convection is occurring between 0 and 20° N, with some convection extending further north. The data show very clearly that the C<sub>2</sub>H<sub>2</sub> mixing ratios are very strongly related to the convective systems (see also Fig. 11). Following injection, the enhancement moves northwards and eastwards from the convective region into the centre of the anticyclone as the air is
- <sup>15</sup> convected upwards. This also is consistent with the work of Park et al. (2007) who note a clear distinction between the location of the anticyclone circulation and the convective region over India and South-East Asia. The 150 hpa geopotential height anomaly calculated from NCEP meteorological data (Fig. 12a) shows the position of the Asian monsoon anti-cyclone and the associated wind field for this month.
- <sup>20</sup> Zonal cross-sections of  $C_2H_2$  were taken along the 20° N and 30° N latitude lines passing through the strong Asian monsoon anticyclone isolation. The cross-section taken across the centre of the anticyclone core at 30° N (Fig. 10a) shows the extent of the chemical isolation with a strong gradient of  $C_2H_2$  both longitudinally between 20° E and 50° E at the western edge of the anticyclone and vertically as the  $C_2H_2$  is con-
- strained beneath the tropopause. The dramatic change in the C<sub>2</sub>H<sub>2</sub> profiles across the monsoon boundary shows the strength of the chemical isolation inside the monsoon as previously identified by Park et al. (2008) but with much higher spatial and temporal resolution. It also clearly demonstrates the distinct gradient regions surrounding the



monsoon anticyclone core. The zonal cross-section at 30° N (Fig. 10b) transects the convective region of the monsoon located over India and South-East Asia as indicated by OLR values less than  $205 \text{ W/m}^2$  (Fig. 12b). Although the centre of the convective region contains no cloud-free data, the enhancements observed on the edge of this <sup>5</sup> region support the suggestion that the C<sub>2</sub>H<sub>2</sub> is uplifted from within this convective region. Finally, the zonal cross-section at 30° N (Fig. 10a) shows very nicely the eastward transport of enhanced VMRs at 200 hPa.

Figure 11 shows the  $C_2H_2$  distributions at the four MIPAS levels with average pressures of 313, 200, 120 and 72 hPa (the 9, 12, 15 and 18 km nominal altitudes). Figure 11b at an average pressure of 200 hPa illustrates the convection-related uplift, the

- <sup>10</sup> ure 11b at an average pressure of 200 hPa illustrates the convection-related uplift, the distinct gradients maintained around the core of the anti-cyclone and the enhancement of the VMRs in the core. In addition, one can observe a westward extension of the anticyclone which results in high mixing ratios in an apparent maximum near 160 degrees east. Such a feature could clearly be mis-interpreted in nadir sounding data emphasising the complementarity of good limb sounding observations.
  - In order to define the extent of the anticyclone system, the 150 hpa geopotential height anomaly is calculated from NCEP meteorological data (Fig. 12a) and confirms that the  $C_2H_2$  is strongly isolated within the anticyclone core. The ability to capture the chemical isolation caused by the monsoon boundary so well in satellite data provides valuable information on the extent of its influence.

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A set of typical averaging kernels for these  $C_2H_2$  MIPAS retrievals over the monsoon anticyclone region are shown in Fig. 13 and clearly show strong peaks at approximately 300 hpa, 200 hpa and 120 hPa; for example, the 200 hpa averaging kernel peaks with a value of 0.72, indicating that there is a relatively small dependence on the a priori.

<sup>25</sup> The C<sub>2</sub>H<sub>2</sub> response function above 120 hpa shows that these data are highly reflective of the 120 hPa level itself and so should be highly correlated with it. Our data reflect this in both global and local correlations (not shown). Therefore, although there is an interesting enhanced feature at 72 hPa in our data sets, we cannot assert that this feature is unambiguously a stratospheric enhancement since it overlies the enhanced



VMRs in the troposphere below it. These vertical correlations in our retrievals suggest that interpretation of this feature in our MIPAS data set requires considerable care.

In a recent paper, Randel et al. (2010) identify the transport of air masses from the surface deep into the stratosphere through the use of HCN retrieved from ACE. They

<sup>5</sup> conclude that the monsoon circulation provides an effective pathway for pollution from Asia, India and Indonesia to enter the global stratosphere. It is not clear whether our retrievals are entirely sensitive to the stratosphere but they do show good promise for verifying model transport pathways in a given year.

This section has focused on the chemical isolation of the  $C_2H_2$  distribution relating to the Asian monsoon anticyclone. The vertical and horizontal extent of this isolation has been mapped to a much higher spatial resolution than has previously been possible. The data compares well against other indicators of the location of the anticyclone such as the geopotential height, and shows both the strong eastward transport in the core of the vortex and the strong gradients at the edge of the vortex. The convective impact on  $C_2H_2$  is very clear from our data using both the regions of cloudiness and the OLR. Enhancements of the VMR could not clearly be distinguished in the stratosphere due

#### to vertical correlations in the retrievals.

#### 5 Conclusions

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 $C_2H_2$  has been successfully retrieved from MIPAS for August 2003 with strong signatures associated with the Asian monsoon anticyclone chemical isolation and African biomass burning clearly evident.

Once retrieved, the  $C_2H_2$  VMR has been used in conjunction with the 150 hpa MO-PITT CO VMR to examine the correlation between  $C_2H_2$  and CO. When performed globally, these calculations showed a strong correlation between the  $C_2H_2$  and CO of

0.76 but with two distinct domains due to the Northern and Southern Hemispheres. When correlated individual, the Northern Hemisphere had a correlation of 0.67 showing the much greater variability in the transport and sources compared to the Southern



Hemisphere whose correlation was 0.918 but with significantly lower  $C_2H_2$  values. Defining the geographical regions to limit the influence of the variability from different sources allowed stronger correlations to be revealed. Africa, South America, the Asian monsoon anticyclone and the Asian Pacific outflow regions all correlated strongly with

<sup>5</sup> the difference in gradients giving some indication of the amount of transport occurring in the different regions. Even for the case over the Middle East/Asia where these is no clear ocean background influence to distinguish the enhancements from, a reasonable correlation between  $C_2H_2$  and CO still existed.

As the  $C_2H_2$  and CO remain well correlated through transport and mixing, this allowed the  $C_2H_2/CO$  ratio to be calculated and analysed in the context of the photochemical age of air and the time since the air mass had last encountered a combustion source. From this analysis it was found that known transport mechanisms from Southern Africa into the Atlantic/Indian Oceans and the transport from Asia into the Pacific Ocean were well-defined with the relative speed of each transport system indicated by

- <sup>15</sup> the C<sub>2</sub>H<sub>2</sub>/CO ratio. It was also noted that the C<sub>2</sub>H<sub>2</sub> appeared to be more tightly confined along the transport pathways compared to the CO which due its longer lifetime does become slightly more mixed. Furthermore, a strong C<sub>2</sub>H<sub>2</sub> signal was retrieved over the Middle East region and the C<sub>2</sub>H<sub>2</sub>/CO ratio allowed us to observe enhanced C<sub>2</sub>H<sub>2</sub> concentrations resulting from the fast outflow from Asia associated with the monsoon
- <sup>20</sup> anticyclone. This enhancement was maintained by the influence of the anticyclone associated with the monsoon, which acts as a barrier to further transport and leads to the chemical isolation in the region. This was verified further through the use of trajectory modelling (not shown) which provided further confidence in both the  $C_2H_2$  retrieval and in the use of the  $C_2H_2/CO$  ratio to act as an indicator for the photochemical age of air <sup>25</sup> in future work.

By examining cross-sections through the anticyclone, the horizontal and vertical extent of the isolation was shown. The location of the convective region over India and South-East Asia as identified by OLR data was also evident from these cross-sections, suggesting that it is from this region where the  $C_2H_2$  is convected into the



upper troposphere. Analysis of the geopotential height anomaly and 150 hpa wind vectors were also used to verify the location of the anticyclone which was found to be in very good agreement with the  $C_2H_2$  distribution, identifying  $C_2H_2$  as an appropriate dynamical tracer.

- <sup>5</sup> A similar analysis was performed for the African biomass burning enhancements. This showed the behaviour of the uplift, which was observed not to penetrate as high into the troposphere as for the Asian monsoon anticyclone. This, along with the observed outflow from this region, has allowed the areas of more intense biomass burning activity to be identified.
- In summary, we have shown that is is possible to use MIPAS data to explore the upper-tropospheric distributions of  $C_2H_2$  with a much higher temporal and spatial resolution than has previously been possible. The ratio of  $C_2H_2$ -CO was shown to provide information on the transport of photochemically aged plumes and the amount of atmospheric processing they have undergone. In addition, we have shown that  $C_2H_2$  can act as a tracer not only for investigating effects such as transport and convection but
- also for identifying dynamical effects such as the monsoon's chemical isolation. Due to the high spatial sampling, this data would prove suitable for further exploitation by comparisons to modelled data in order to study these effects further.

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**Fig. 3.** Contributions to the  $C_2H_2$  retrieval error. The total error (solid line), random error (dotted line) and the systematic error (dashed line) are all shown as well as the various components to the systematic error such as the Instrument Line Shape and spectroscopic errors.











**Fig. 5.** Correlation of the 200 hpa MORSE  $C_2H_2$  VMRs against 150 hpa MOPITT CO VMRs averaged for August 2003 for Globally, the Northern Hemisphere and the Southern Hemisphere. The two distinct domains observed in the global correlation are due to the differences in sources and transport for the Northern Hemisphere and Southern Hemisphere. The dashed lines indicate the line of best fit through the correlations. The red points show the average values in 2 ppbv bins, with the associated error bars indicating the standard deviation within each bin.





**Fig. 6.** 200 hpa MORSE  $C_2H_2$  VMRs against 150 hpa MOPITT CO VMRs averaged for August 2003 for: Africa [30° S to 10° N, 30° W to 80° E], South America [30° S to 10° N, 180° W to 30° W], Asian Pacific Outflow [10° N to 40° N, 130° E to 110° W] and the Asian monsoon anticyclone region [10° N to 40° N, 30° E to 110° E]. Again, the dashed lines indicate the line of best fit through the correlations. The red points show the average values in 2 ppbv bins, with the associated error bars indicating the standard deviation within each bin.





























(d) 72 hpa (18 km nominal altitude)

Fig. 11. The MORSE retrieved  $C_2H_2$  distributions for August 2003 located over the Asian monsoon anticyclone region for the 9 km, 12 km, 15 km and 18 km nominal MIPAS tangent altitudes. The strong Asian monsoon anticyclone isolation is clearly evident as well as significant transport from Asia towards North America, particulary at lower altitudes.





**Fig. 12.** The Geopotential Height Anomaly and Outgoing Longwave Radiation for August 2003. The geopotential height anomaly is used as an indication for the anticyclone location and is shown to be in strong agreement with the enhanced  $C_2H_2$  chemically isolated by the monsoon anticyclone. The OLR is used to indicate areas of deep convection (i.e. OLR <  $205 \text{ W/m}^2$ ) which are outlined by the white contour lines. These show the large area of deep convection over India and South-East Asia, on the South-Eastern edge of the Asian monsoon anticyclone.







