Atmos. Chem. Phys. Discuss., 10, 25959–25989, 2010 www.atmos-chem-phys-discuss.net/10/25959/2010/ doi:10.5194/acpd-10-25959-2010 © Author(s) 2010. CC Attribution 3.0 License.



This discussion paper is/has been under review for the journal Atmospheric Chemistry and Physics (ACP). Please refer to the corresponding final paper in ACP if available.

Homogenous nucleation of sulfuric acid and water at atmospherically relevant conditions

D. Brus^{1,2}, K. Neitola², T. Petäjä³, J. Vanhanen³, A.-P. Hyvärinen², M. Sipilä³, P. Paasonen³, H. Lihavainen², and M. Kulmala³

¹Laboratory of Aerosol Chemistry and Physics, Institute of Chemical Process Fundamentals Academy of Sciences of the Czech Republic, Rozvojová 135, 165 02 Prague 6, Czech Republic

²Finnish Meteorological Institute, Erik Palménin aukio 1, P.O. Box 503, 00100 Helsinki, Finland
³Department of Physics, University of Helsinki, P.O. Box 64, 00014 Helsinki, Finland

Received: 16 September 2010 - Accepted: 26 October 2010 - Published: 3 November 2010

Correspondence to: D. Brus (brus@icpf.cas.cz)

Published by Copernicus Publications on behalf of the European Geosciences Union.



Abstract

In this study the homogeneous nucleation rates of sulfuric acid and water were measured by using a flow tube technique. The goal was to directly compare particle formation rates obtained from atmospheric measurements with nucleation rates of freshly

- ⁵ nucleated particles measured with particle size magnifier (PSM) which has detection efficiency of unity for particles having mobility diameter of 1.5 nm. The gas phase sulfuric acid concentration in this study was measured with the chemical ionization mass spectrometer (CIMS), commonly used in field measurements. The wall losses of sulfuric acid were estimated from measured concentration profiles along the flow tube. The
- ¹⁰ initial concentrations of sulfuric acid estimated from loss measurements ranged from 10^8 to 3×10^9 molecules cm⁻³. The nucleation rates obtained in this study cover about three orders of magnitude from 10^{-1} to 10^2 cm³ s⁻¹ for commercial ultrafine condensation particle counter (UCPC) TSI model 3025A and from 10^1 to 10^4 cm³ s⁻¹ for PSM. The nucleation rates and the slopes (*d*In*J*/*d*In[H₂SO₄]) show satisfactory agreement when compared to empirical kinetic and estimates medals and the latest atmospheric
- ¹⁵ when compared to empirical kinetic and activation models and the latest atmospheric nucleation data.

1 Introduction

Atmospheric new particle formation consists of rather complicated sets of processes, the first of them is gas-to-particle nucleation which occurs naturally but might be also
easily influenced by anthropogenic emissions of gases such as SO₂. It is generally accepted that sulfuric acid is a robust source of new particles and plays a central role in atmospheric new particle formation (Weber et al., 1996, 1997; Kulmala, 2003). In number of field experiments (e.g. Weber et al., 1996; Sihto et al., 2006; Riipinen et al., 2007; Zhao et al., 2010) and also in some laboratory studies (e.g. Berndt et al., 2005, 2010; Young et al., 2008; Sipilä et al., 2010) the rate of particle formation is not adequately explained by binary classical homogenous nucleation (CNT), the theory



greatly under predicts the observed nucleation. According to Kashchiev (1982) the relationship (the slope, $d \ln J/d \ln [H_2 SO_4]$) between particle production rate and sulfuric acid concentration directly corresponds to number of molecules in critical cluster. In atmospheric measurements and also in laboratory studies (Sipilä et al., 2010; Berndt

- ⁵ et al., 2010) it was observed that particle number concentration followed a power-law dependency of about 1–2, compared to CNT prediction that suggest exponents from 4 to 9 (Vehkamäki et al., 2002). This discrepancy has been puzzling to atmospheric researchers for more than a decade. As a solution to this problem it has been suggested that other associate molecule as ammonia (Weber et al., 1996) or organic acids (Zhang
- et al., 2004, 2010; Metzger et al., 2010; Paasonen et al., 2010) may have a stabilizing effect on the clusters and allow nucleation to occur at much lower concentrations of sulfuric acid than needed by CNT. The first in-situ atmospheric measurements of sulfuric acid in troposphere by using chemical ionization mass spectrometer (CIMS) were reported by Eisele and Tanner in 1993. Since that time CIMS was used in many field
 studies in diverse locations around the world (e.g. Weber et al., 1996, 1997; Mauldin et
- al., 1998; Sihto et al., 2006; Riipinen et al., 2007; Petäjä et al., 2009, Zhao et al., 2010) and also in laboratory studies (e.g. Ball et al., 1999; Zhang et al., 2004; Benson et al., 2008; Young et al., 2008; Sipilä et al., 2010).

Compared to our previous study (Brus et al., 2010) we apply here a similar approach as is commonly used in atmospheric measurements. The gas phase sulfuric acid concentration was measured with the chemical ionization mass spectrometer (CIMS). The number concentration of freshly nucleated particles was measured in parallel with commercial UCPC TSI 3025A and particle size magnifier (PSM) with particle counting efficiency close to unity for particles of ~1.5 nm, (Vahnanen et al., 2010).

25 2 Experimental setup

The same experimental setup as introduced in Brus et al. (2010) was used in this investigation. The experimental setup details and the principle of operation can be found



therein. Only brief description of apparatus, its principle of operation and differences associated with particle counting and determination of sulfuric acid concentration are discussed here. The experimental setup consists of five main parts: an atomizer, a furnace, a mixing unit, a nucleation chamber and a particle detector unit. A liquid solu-

- tion of known concentration and amount (0.22 ml/min) is introduced by the HPLC Pump (Waters 515) through a ruby micro-orifice (Bird Precision – 20 μm) together with particle free air (about 4 l/min) into the furnace. The dispersion is vaporized in a furnace (Pyrex glass tube) which is 60 cm long and has an internal diameter (I.D.) of 2.5 cm. The tube is wrapped with resistance heating wires. The temperature inside the furnace
- ¹⁰ is kept at approximately 470 K and controlled by a PID controller to within ±0.1 K (Digi-Trace, TCONTROL-CONT-02). After the furnace, the vapor is filtered with a Teflon filter (MITEXTM Millipore 5.0 μm LS) to remove any liquid residue or particulate impurities. The Teflon filter is placed on the perforated Teflon support pad just after the furnace, and before the entrance to the mixing unit. The filtered vapor is then introduced into
- the mixing unit, made of Teflon, and cooled by turbulent mixing with room temperature particle free air to about 320 K. The flow rate of the mixing air is about 8 l/min. The mixing unit dimensions are: O.D.=10 cm, I.D.=7 cm, height=6 cm. The mixing unit is kept at room temperature and it is not insulated. Both lines of particle free air are controlled by a flow rate controller to within ±3% (MKS type 250). The vapor gas mixture
- is then cooled to the desired nucleation temperature in a nucleation chamber, which is kept at a constant temperature with two liquid circulating baths (Lauda RK-20). The nucleation chamber is made of stainless steel, with an I.D. of 6 cm and an entire length of 200 cm. The concentration of water vapor is measured at the centre and far end of the nucleation chamber with two humidity and temperature probes (Vaisala HMP37E
- and humidity data processor Vaisala HMI38) to within ±3%. The aerosol number concentration is measured just after the nucleation chamber with an ultrafine condensation particle counter (UCPC) TSI model 3025A and simultaneously with particle size magnifier (PSM). The sulfuric acid concentration is measured also at the end of the nucleation chamber with chemical ionization mass spectrometer (CIMS).



The liquid samples of sulfuric acid and water mixture are prepared from a 0.01 M solution of H_2SO_4 (Reagecon, AVS purity) and ultrapure water (Millipore, TOC less than 10 ppb, resistivity 18.2 M Ω cm @ 25 °C). The desired solution concentration is prepared in two steps of dilution. First, 1 I of primary solution of concentration (1.96×10⁻⁴ mol/l) is made by adding 20 ml of 0.01 mol H_2SO_4 to 1 I of pure water. Then the desired final solution for a particular measurement is made. To cover RH's from 60% to 10% we

prepare 1 l of final solution from 0.5 ml to 70 ml of primary solution. The final solution concentration is always checked by Ion Chromatography with a lower detection limit of 0.02 mg/l of SO₄²⁻ in the analytical laboratory at the Finnish Meteorological Institute.

10 2.1 Chemical Ionization Mass Spectrometer, CIMS

5

Sulfuric acid was measured with a chemical ionization mass spectrometer, CIMS (Eisele and Tanner, 1993; Mauldin et al., 1998; Petäjä et al., 2009). The sulfuric acid in the sample flow is chemically ionized by (NO³⁻) ions. The reagent ions are generated by nitric acid and a 241Am alpha source and mixed in a controlled manner in a drift tube utilizing concentric sheath and sample flows together with electrostatic lenses.

Prior to entering the vacuum system, the chemically ionized sulfuric acid molecules pass through a layer of dry nitrogen flow in order to dehydrate the sulfuric acid. In the vacuum system the sulfuric acid clusters are dissociated to the core ions by collisions with the nitrogen gas seeping through the pinhole in the collision-dissociation cham-

²⁰ ber (Eisele and Tanner, 1993). The sample beam is collimated with a set of conical octopoles, mass filtered with a quadrupole and detected with a channeltron. The sulfuric acid concentration is determined by the ratio between the signals at mass 97 amu (HSO⁴⁻) and the reagent ion at mass 62 amu (NO³⁻) multiplied by the instrument and setup dependent calibration factor.

²⁵ The calibration factor is determined by photolyzing ambient water vapor with a mercury lamp to generate a known amount of OH radicals in front of the inlet. The produced OH radicals subsequently convert isotopically labeled ³⁴SO₂ into labeled sulfuric acid in a well defined reaction time yielding finally after ionization (H³⁴SO₄⁻). A nominal



detection limit of the CIMS instrument is 5×10^4 molecules cm⁻³ for a 5 min integration period. The error estimate in determined concentrations is estimated to be about factor of 2. CIMS was used also to detect sulfuric acid dimers. The calibration factor used for monomers was applied also in converting the dimer signal to concentration. Since the transmission for dimer (195 amu) can differ from monomer (97 amu), use of the single calibration factor causes error in the determined concentration. Therefore our results concerning the dimer concentrations are still somewhat qualitative.

2.2 Particle size magnifier

5

Particle size magnifier (PSM, Airmodus A09) used in this study is based on two recent major developments on the field of particle counting. First, on the work of Sgro and de la Mora (2004) (and the references therein) with the development of mixing type particle size magnifier for almost arbitrarily small particles, and second, on the study by lida et al. (2009) to find the most suitable working fluid to be used in a condensation particle counter (CPC). The critical dimensions and the geometry of the PSM are very

- ¹⁵ close to those given by Sgro and de la Mora (2004). Diethylene glygol was used as the working fluid. It has relatively high surface tension and low saturation vapor pressure. Because of these properties a high saturation ratio is acquired without homogeneous nucleation (lida et al., 2009). Diethylene glygol has also been experimentally tested in the ultrafine condensation nucleus counter (UFCNC) prototype (Stoltzenburg and
- McMurry, 1991) showing a superior performance in the sub-2 nm size range (lida et al., 2009). Due to low vapor pressure of diethylene-glygol the particles cannot easily grow to optical sizes (~1 µm in diameter). Therefore an external CPC (TSI 3010) is used for detecting the activated particles in this design. Calibration results (Vahnanen et al., 2010) have shown that PSM detects charged particles approaching efficiency of unity (practically diffusion loss limited) down to ~1.5 nm. Below that still ~25% of the
- smallest calibration ion (tetra-methyl-ammonium-ion) with mobility equivalent diameter of 1.05 nm, was activated in the PSM in comparison to reference electrometer (TSI 3068B). An assumption of unity detection efficiency in case of PSM is justified.



Results and discussion 3

Two separate experiments were conducted in the Finnish Meteorological Institute (FMI) flow tube. The nucleation rates of sulfuric acid and water were measured as a function of initial sulfuric acid concentration at three different relative humidities (16, 32

and 57%). Also the nucleation rate temperature dependency was investigated; the ex-5 periments were conducted at three temperatures (25, 15 and 5°C) while keeping the relative humidity close to 50%. To obtain the initial sulfuric acid concentration, the sulfuric acid losses were estimated separately for all experimental conditions. The main reason why we focused on obtaining initial sulfuric acid concentration in our flow tube was that the concentration of prepared solution of sulfuric acid and water is known for 10 each particular experiment and thus the initial sulfuric acid concentration determined with CIMS and IC method can be mutually compared.

3.1 H₂SO₄ losses

15

- Sulfuric acid wall losses were determined experimentally by measuring the losses of sulfuric acid concentration along the nucleation chamber. Two sets of experiments were conducted. First, relative humidity was changed (16, 32 and 57%) and nucleation temperature (25°C) was kept constant. Second, nucleation temperature was changed (25, 15 and 5°C) and relative humidity was kept constant (~50%). The nucleation chamber consists of two 1 m long interchangeable parts; one of them is equipped with
- 4 holes in equal distance of 20 cm from beginning and from each other. In the first set 20 of measurements the holes were in upper position so we measured sulfuric acid losses for relative humidities 16, 32, 57% in distances of 20, 40, 60, and 80 cm from the beginning and then at the end (200 cm) of the nucleation chamber. The slopes obtained from the fits to experimental data $ln([H_2SO_4])$ vs. distance in the nucleation chamber
- stand for the loss rate coefficient, k_{obs} (cm⁻¹). To be able to measure along whole 25 nucleation chamber an additional CIMS inlet sampling tube had to be used, which is a stainless steel tube with I.D. 10 mm and its whole length was 122 cm (100 cm straight



+ 22 cm elbow-pipe). The sulfuric acid losses in the CIMS' inlet sampling tube itself were estimated in separate experiment by using two different lengths of straight sampling tube, 50 and 100 cm. In Table 1 can be found the observed loss rate coefficients, k_{obs} , mean flow velocity, v, residence time, t, the wall loss factor, WLF, and the wall loss factor estimated for CIMS' inlet sampling tube, WLF_{inlet}. The overall sulfuric acid 5 losses in the nucleation tube and the inlet sampling tube are then WLF×WLF_{inlet}. The initial sulfuric acid concentration [H₂SO₄]_{init} is then obtained as [H₂SO₄]_{measured}×WLF ×WLF_{inlet}. The wall loss factor (WLF) increases with relative humidity, which is due to presence of temperature gradient in first 50 cm in the nucleation chamber, see Brus et al. (2010) for details. The presence of temperature gradient (both axial and radial) im-10 poses thermophoretic force towards the cooled nucleation chamber wall (set to 25°C) and also increases the value of diffusion coefficient, thus increasing WLF in first 50 cm. There is no temperature gradient in the CIMS' sampling tube and the WLF_{inlet} is behaving as expected, the WLF_{inlet} is increasing with decreasing RH, (Hanson and Eisele, 2000). 15

In the second set of loss measurements the positions of nucleation chamber parts were exchanged, so the holes were in lower part of nucleation chamber. This was done to ensure the reproducibility of experiment at relative humidity ~50%, and also find out how big role plays the axial temperature gradient (thermophoresis and higher diffusion

- ²⁰ coefficient) in first 50 cm of the nucleation chamber. The sulfuric acid losses were measured at three nucleation temperatures (25, 15 and 5 °C) and relative humidity ~50% at distances of 120, 140, 160, 180, and 200 cm. Table 2 contains the obtained loss rate coefficients together with accompanied parameters. The WLF is generally smaller then in the first set of experiment due to smaller and constant diffusion coefficient; i.e. there
- is no axial temperature gradient present in lower part of nucleation chamber. The WLF is increasing with decreasing nucleation temperature again due to increased radial temperature gradient. The WLF_{inlet} is even pronounced because the CIMS' sampling inlet tube was not temperature controlled, but only well insulated.



3.2 Nucleation rates

The number concentrations of freshly nucleated particles were measured as a function of initial sulfuric acid concentration at several levels of relative humidities. In all experiments two different counting systems were used in parallel. An UCPC TSI 3025A

- ⁵ which was calibrated with silver particles to a mobility diameter d_{50} cut-off of 2.28 nm. As a second counting system a mixing type particle size magnifier (PSM) with close to unity detection efficiency for mobility equivalent diameter of 1.5 nm was used. The initial concentration of sulfuric acid was estimated from loss measurements using CIMS and it ranged from 10⁸ to 3×10⁹ molecules cm⁻³. The onset of nucleation for UCPC TSI 3025A particle counter was observed at sulfuric acid initial concentrations about
- ¹⁰ TSI 3025A particle counter was observed at sulfuric acid initial concentrations about 10⁸ and for PSM at about 10⁷ molecules cm⁻³. The different counting efficiency of both counters lead to different slopes in plot of nucleation rate vs. sulfuric acid concentration. The biggest difference in counting between UCPC TSI 3025A and PSM is at lowest nucleation rates, about a factor of 200, and the smallest difference is at high-
- est nucleation rates, about a factor of 3. From the obtained slopes it is obvious that both lines will merge at certain point, where the particle diameter of grown particles for UCPC TSI 3025A will also reach the counting efficiency of unity. The detailed comparison and explanation of differences among several counting systems can be found in Sipilä et al. (2010). The linear fits to experimental data for both particle counters are
- ²⁰ presented in Table 3. Nucleation time in our experiment is defined as time from the nucleation zone maxima to the end of the flow tube; which is half of the total residence time. Nucleation zone was determined experimentally (Brus et al., 2010) and also with Fluent CFD model (Herrmann et al., 2010), the maxima of nucleation zone was found at distance of about 1 m up from the nucleation chamber end. The nucleation rate is then
- defined as particle number concentration divided by nucleation (or half of residence) time. The highest uncertainty in nucleation rate is estimated to be factor of ±2 when considering unlikely shift in position of nucleation zone maximum from 50 cm to 150 cm in the nucleation chamber. The resulting nucleation rates at relative humidities of 16,



32, 47 and 57% cover about three orders of magnitude from 10^{-1} to 10^2 cm³ s⁻¹ and from 10^1 to 10^4 cm³ s⁻¹ for UCPC TSI 3025A and PSM, respectively, see Fig. 1a and b. It has to be pointed out that experimental data at RH=32% were already published in Sipilä et al. (2010) and experimental data at RH=47%, are taken from temperature dependency measurements (see Sect. 3.3) to show the experiment reproducibility.

3.3 Temperature dependency

The effect of temperature on nucleation rate was studied and the experimental results are presented for both counters (UCPC TSI 3025A and PSM) separately in Fig. 2a and b. The nucleation rates were measured as a function of sulfuric acid initial concen-¹⁰ tration at three temperatures 25, 15 and 5°C, the relative humidity was kept close to 50%. The experiment was conducted in a way that for one prepared solution of sulfuric acid and water, first all flow tube parameters (RH, T, total flow) were adjusted to measure nucleation temperature of 25°C. After the experiment was finished the parameters were readjusted to measure nucleation temperature of 15°C by changing temperature

of nucleation chamber wall and flows to keep relative humidity close to 50%. Finally nucleation temperature of 5 °C was measured in the same way and the sulfuric acid – water solution was changed afterwards.

The nucleation rate shows an enhancement of more than one order of magnitude per decreasing the nucleation temperature by 20°C at sulfuric acid concentration of 10⁹ molecules cm⁻³ for both particle counters. At sulfuric acid concentration of 10⁸ molecules cm⁻³ the measured data show no enhancement of nucleation rate because of different slopes of isotherms. The steepest slope was observed at temperature of 5°C which is in disagreement with prediction of CNT (Vehkamäki et al., 2002). This might be due to undercounting of both particle counters at lower sulfuric acid concentrations, CNT predicts about 30% smaller critical cluster size at 5°C than at 25°C. The experimental data for 15 and 25°C lie almost on top of each other; this is probably due to experimental difficulties we observed at lower temperatures 15 and 5°C. The resulting slopes of fits to experimental data are collected in Table 4.



3.4 Dimer formation

The formation of sulfuric acid dimer in both its hydrated and unhydrated form is the first step in sulfuric acid and water nucleation process (Hanson and Lovejoy, 2006). The working amu range of the CIMS used in this study was from 40 to 250, which allowed us

to observe individual sulfuric molecules as HSO_4^- at 97 amu and also sulfuric acid dimer 5 cluster as HSO₄·H₂SO₄ at amu 195. It must be pointed here that results concerning the dimer concentrations are still only qualitative. Generally, the observed concentration of dimer was in units of percent of the monomer concentration, which agrees with earlier studies (Eisele and Hanson, 2000). A slight RH dependency in monomer to dimer relation was observed. For the nucleation temperature of 25 °C the increasing trend 10 in the ratio (M/D) of monomer (97 amu) to dimer (195 amu) from ~100 to ~200 was observed with increasing relative humidity from 16 to 56%, see Fig. 3. The slope of the fit to all data in Fig. 3 is 1.08 and $R^2 = 0.76$.

The monomer to dimer ratio as a function of nucleation temperature can be seen in Fig. 4. The data are averages over whole isotherm with corresponding standard devia-15 tions as error bars. The M/D ratio is about factor of 3 larger for nucleation temperature of 25°C (M/D=224) than for 5°C (M/D=85), Eisele and Hanson (2000) reported M/D value ~40 at ~240 K. The trend of M/D ratio is decreasing with decreasing temperature in our study. Similar trend was also observed in Eisele and Hanson (2000) but only for cluster bigger than trimer. 20

25

3.5 Comparison to our previous data

A discrepancy was found in the results of this study compared to data published earlier (Brus et al., 2010). In our earlier study the method of bubblers was used to estimate concentration and the losses of sulfuric acid along the flow tube as the total sulfate (SO_4^{2-}) concentration obtained via ion chromatography (IC) analysis. In this study the initial sulfuric acid concentration measured with CIMS method reaches about 20% at RH~50% and only about 1% at RH~16% of total sulfate concentration obtained via



ion chromatography (IC) analysis of the prepared liquid samples and consequent initial sulfuric acid concentration calculated by mass balance, see Fig. 5.

There might be several reasons for such observations. The CIMS measures only monomer (97 amu) in a gas phase, the dimer concentration (195 amu) was usually ⁵ less then 1% of monomer concentration. This might indicate that the rest of sulfuric acid is in another form. The losses of sulfuric acid into particles is marginal, it was in the range of few per mille to maximum of 3% for sulfuric acid concentration range from 10⁸ to 10⁹ molecule cm⁻³.

What has to be also considered is shielding of sulfuric acid with water molecules. The hydration of sulfuric acid takes always place whenever traces of water are involved in the process of nucleation. According to classical theory of hydration made by Jaecker-Voirol and Mirabel (1987, 1988) and validated by Kulmala et al. (1991), only about 10% of sulfuric acid is in unhydrated form at relative humidity of 50%. Salcedo et al. (2004) studied the effect of relative humidity on the detection of sulfur dioxide and sulfuric acid

- and found negative effect on the sensitivity of the CIMS to SO₂ and H₂SO₄ because water molecules form clusters with reactant and product ions thus shielding the molecules from being ionized. They claim that the effect can be avoided by increasing the CIMS' inlet flow tube temperature to 150 °C. On the other hand e.g. Eisele and Tanner (1993) in their study claim that the CIMS measurements are sensitive to total sulfuric acid
- without discrimination between free acid and monoacid hydrates, or even between free and higher-order acid clusters and their hydrates. Water is far more volatile than sulfuric acid and any water associated with an ion may be driven off as the ion is sampled through the collisional-dissociation chamber (CDC) of the CIMS, (Eisele and Hanson, 2000). Our results are contra intuitive in the case of water molecule shielding. The
- ²⁵ sulfuric acid concentration measured with CIMS is decreasing with decreasing relative humidity. If the shielding would be due to water molecules then sulfuric acid concentration would be increasing towards the lower relative humidity. Second possibility is involvement of ammonia in shielding of sulfuric acid molecules by creation of stable clusters of ammonium sulfate or ammonium bisulfate. Even though the concentration



of ammonia was always below the detection limit of ion chromatography (IC) analysis (0.02 mg/l), we have no doubts that there is always certain level of ammonia present in our experiment even though the ultrapure water and particle free clean air is used. The IC ammonia detection limit (0.02 mg/l) for our experimental setup corresponds to

mixing ratio of 0.5 ppb of ammonia, this corresponds to concentration about one order higher than sulfuric acid concentration measured with CIMS and close to ratio of unity to total sulfate concentration obtained from IC analysis and subsequent mass balance calculation.

In this study the liquid solutions of sulfuric acid and water were prepared in the same way as in previous study (Brus et al., 2010) also the same range of total sulfate concentrations when calculated by mass balance was observed, and the similar nucleation rates when compared to UCPC 3025A were obtained for the same range of total sulfate concentration.

3.6 Comparison to atmospheric nucleation data

Many scientific groups found and confirmed that the vapor concentration of sulfuric acid in atmosphere is often strongly connected with new particle formation. The correlation of sulfuric acid vapor concentrations and formation rate of neutral aerosol particles can be generally expressed with two models, the kinetic model of McMurry (1980) and the activation model of Kulmala et al. (2006). Parameters of both models are determined
 empirically from atmospheric data. Both models are dependent on the sulfuric acid concentration, kinetic model quadratically and activation model linearly:

 $J = (K \text{ or } A) \times [H_2 SO_4]^E$,

where *K* is a kinetic coefficient ranging from 10^{-14} to 10^{-11} cm³ s⁻¹ and *A* activation coefficient ranging from 10^{-7} to 10^{-5} s⁻¹ (e.g. Weber et al., 1996; Sihto et al., 2006; Riipinen et al., 2007; Paasonen et al., 2010), *E* is an exponent associated with number of sulfuric acid molecules in critical cluster (Kashchiev, 1982), it is usually found to be in between values 1 and 2 when applied to atmospheric data.

Discussion Paper ACPD 10, 25959-25989, 2010 Homogenous nucleation of sulfuric acid and water **Discussion** Paper D. Brus et al. **Title Page** Introduction Abstract Conclusions References **Discussion** Paper **Figures** Back **Discussion** Paper Full Screen / Esc **Printer-friendly Version** Interactive Discussion

(1)

In this study we compare our experimental data to latest atmospheric data analysis made by Paasonen et al. (2010) where the CIMS was used at four measurement sites – Hyytiälä (Finland), Hohenpeissenberg and Melpitz (Germany), and San Pietro Capofiume (Italy). However, as Paasonen et al. (2010) concluded, the nucleation rates

- ⁵ in Hohnepeissenberg were not closely connected to sulfuric acid concentration, and thus our comparison is made only to the data from the other three sites, see Fig. 6. The formation rates of 2 nm neutral particles (J_2) were obtained at all stations from particle size distributions recorded on nucleation event days. Such dataset can be directly compared to nucleation rates obtained with PSM in our study. The exponents
- from linear fits to our experimental data range from 1.2 to 2.2, depending on relative humidity and nucleation temperature. The worse agreement between our experiment and atmospheric data was found for the highest nucleation temperature (25 °C) and the lowest relative humidity (RH 16%), see Fig. 6. The kinetic and activation coefficients obtained from our experimental data are in close agreement to atmospheric ones even
- ¹⁵ though the range of relative humidities and temperatures of atmospheric data is quite wide, see Tables 5 and 6. The median kinetic and activation coefficients of whole dataset presented in Paasonen et al. (2010) (Table 4 therein) are $K = 26 \times 10^{-14} \text{ cm}^{-3} \text{s}^{-1}$ and $A = 9.7 \times 10^{-7} \text{ s}^{-1}$. In our study we found median coefficients for whole dataset to be $K = 0.1 \times 10^{-14} \text{ cm}^{-3} \text{ s}^{-1}$ and $A = 7.85 \times 10^{-7} \text{ s}^{-1}$, thus favouring the activation mechanism in nucleation process. However, this kind of interpretation has to be considered with cautiousness, because the nucleation coefficients may be strongly dependent on some other quantities on a low-volatility organic vapor concentration as suggested by
- some other quantities, e.g. low-volatility organic vapor concentration as suggested by Paasonen et al. (2010).

4 Conclusions

In this study the homogeneous nucleation rates were measured in two separate sets of experiment. In first one we tested the influence of relative humidity in the range from 16 to 57% and in second one the influence of temperature on nucleation rate



at three different nucleation temperatures 25, 15 and 5 °C. Two condensation particle counters (UCPC TSI 3025A and PSM with CPC TSI 3010) with different *d*₅₀ detection efficiency were used in parallel to count freshly nucleated particles. The gas phase sulfuric acid concetration was measured with CIMS. The initial concentration of sulfuric acid was estimated from loss measurements using CIMS and it ranged from 10⁸ to 3×10^9 molecules cm⁻³. The losses of sulfuric acid along the flow tube were estimated for each particular set of experimental conditions. The onset of nucleation for UCPC TSI 3025A was observed at sulfuric acid initial concentrations at about 10⁸ and for PSM at about 10⁷ molecules cm⁻³. The resulting nucleation rates at relative humidities of 16, 32, 47 and 57% cover about three orders of magnitude from 10⁻¹ to 10² and from 10¹ to 10⁴ for UCPC TSI 3025A and PSM, respectively. The nucleation rate shows an enhancement of more than one order of magnitude per decreasing the nucleation temperature by 20 °C at sulfuric acid concentration of 10⁹ molecules cm⁻³ the measured data

- ¹⁵ show no enhancement of nucleation rate because of different slopes of isotherms. The concentration of dimers was found to be less than one percent of monomer concentration. A slight RH dependency in monomer to dimer relation was observed. For the nucleation temperature of 25 °C the increasing trend in the monomer (97 amu) to dimer (195 amu) ratio (M/D) from ~100 to ~200 was observed in relative humidity range from
- 16 to 56%. The trend of M/D ratio is decreasing with decreasing temperature in our study. The M/D ratio was estimated to be about factor of 3 larger for nucleation temperature 25°C (M/D=224) than for 5°C (M/D=85). Obtained experimental nucleation rate data were also compared to two empirical (kinetic and activation) models. The obtained median activation coefficients are close to the atmospheric ones, whereas
- ²⁵ the kinetic coefficients were from one to three orders of magnitude smaller. However it has to be pointed out that these coefficients may be strongly dependent on some other quantities like low-volatility organic vapor concentration. The exponents obtained from fits to our data are in the range of 1.2 to 2.2, depending on relative humidity and nucleation temperature. Even though the sulfuric acid concentration measured with



CIMS is only 10% of the total sulfate concentration obtained via Ion Chromatography analysis and subsequent mass balance (Brus et al., 2010), the slopes in figures *J* vs. [H₂SO₄] and *J* vs. total sulfate are the same. This probably means that the form and amount of active sulfuric acid involved in nucleation process itself is ambiguous and it
⁵ is limited from left side by free [H₂SO₄] and from right side by total sulfate concentration. The participation of ammonia can not be disproved in our nucleation experiment, even though the concentration of ammonia was always below the detection limit of IC

analysis, the ammonia mixing ratio of 0.5 ppb and less is probably present during the process of new particle formation.

Acknowledgements. Authors would like to acknowledge KONE foundation, EUCAARI project (The European Integrated project on Aerosol Cloud Climate and Air Quality Interactions) and FCoE (Finnish Center of Excellence of Academy of Finland, project no. 1118615) for their financial support.

References

20

25

¹⁵ Ball S. M., Hanson D. R., and Eisele F. L., and McMurry P. H.: Laboratory studies of particle nucleation: Initial results for H₂SO₄, H₂O, and NH₃ vapors., J. Geophys. Res., 104(D19), 23709–23718, 1999.

Benson, D. R., Young, L. H., Kameel, F. R., and Lee, S.-H.: Laboratory-measured nucleation rates of sulfuric acid and water binary homogeneous nucleation from the SO₂ + OH reaction, Geophys. Res. Lett., 35, L11801, doi:10.1029/2008GL033387, 2008.

Berndt, T., Böge, O., Stratmann, F., Heintzenberg, J., and Kulmala, M.: Rapid formation of sulfuric acid particles at near atmospheric conditions, Science, 307, 698–700, 2005.

Berndt, T., Stratmann, F., Sipilä, M., Vanhanen, J., Petäjä, T., Mikkilä, J., Grüner, A., Spindler, G., Lee Mauldin III, R., Curtius, J., Kulmala, M., and Heintzenberg, J.: Laboratory study on new particle formation from the reaction OH + SO₂: influence of experimental conditions,

 H_2O vapour, NH_3 and the amine tert-butylamine on the overall process, Atmos. Chem. Phys., 10, 7101–7116, doi:10.5194/acp-10-7101-2010, 2010.

Brus, D., Hyvärinen, A.-P., Viisanen, Y., Kulmala, M., and Lihavainen, H.: Homogeneous nucle-



ation of sulfuric acid and water mixture: experimental setup and first results, Atmos. Chem. Phys., 10, 2631–2641, doi:10.5194/acp-10-2631-2010, 2010.

- Eisele, F. L. and Hanson, D. R: First measurements of prenucleation molecular clusters, J. Phys. Chem. A, 104, 830–836, 2000.
- ⁵ Eisele, F. L. and Tanner, D.: Measurement of the gas phase concentration of H_2SO_4 and methane sulfonic acid and estimates of H_2SO_4 production and loss in the atmosphere, J. Geophys. Res., 98, 9001–9010, 1993.
 - Hanson, D. R. and Eisele, F. L.: Diffusion of H₂SO₄ in humidified nitrogen: Hydrated H₂SO₄, J. Phys. Chem. A, 104, 1715–1719, 2000.
- ¹⁰ Hanson D. R. and Lovejoy E. R.: Measurements of the thermodynamics of the hydrated dimmer and trimer of sulfuric acid, J. Phys. Chem. A, 110(31), 9525–9527, 2006.
 - Herrmann, E., Brus, D., Hyvärinen, A.-P., Stratmann, F., Wilck, M., Lihavainen, H. and Kulmala,
 M.: A computational fluid dynamics approach to nucleation in the water-sulfuric acid-system,
 J. Phys. Chem. A, 114(31), 8033–8042, 2010.
- ¹⁵ Iida, K., Stolzenburg, M., and McMurry, P. H.: Effect of Working Fluid on Sub-2 nm Particle Detection with a Laminar Flow Ultrafine Condensation Particle Counter, Aerosol Sci. Technol., 43, 81–96, 2009.
 - Jaecker-Voirol, A., Mirabel, P., and Reiss, H.: Hydrates in supersaturated binary sulphuric acidwater vapor: A reexamination, J. Chem. Phys., 87, 4849–4852, 1987.
- ²⁰ Jaecker-Voirol A. and Mirabel P.: Nucleation rate in a binary mixture of sulfuric acid and water vapor, J. Phys. Chem., 92(12), 3518–3521, 1988.
 - Kashchiev, D: On the relation between nucleation work, nucleus size and nucleation rate, J. Chem. Phys., 76, 5098–5102, 1982.
 - Kulmala, M., Lazaridis, M., Laaksonen, A., and Vesala, T.: Extended hydrates interaction
- ²⁵ model: Hydrate formation and the energetics of binary homogeneous nucleation, J. Chem. Phys., 94(11), 7411–7413, 1991.

Kulmala, M.: How particles nucleate and grow, Science, 302, 1000–1001, 2003.

- Kulmala, M., Lehtinen, K. E. J., and Laaksonen, A.: Cluster activation theory as an explanation of the linear dependence between formation rate of 3nm particles and sulphuric acid concentration, Atmos. Chem. Phys., 6, 787–793, doi:10.5194/acp-6-787-2006, 2006.
- concentration, Atmos. Chem. Phys., 6, 787–793, doi:10.5194/acp-6-787-2006, 2006.
 Mauldin III, R. L., Frost, G., Chen, G., Tanner, D., Prevot, A., Davis, D., and Eisele, F.: OH measurements during the First Aerosol Characterization Experiment (ACE 1): Observations and model comparisons, J. Geophys. Res., 103, 16713–16729, 1998.



- McMurry, P. H.: Photochemical Aerosol Formation from SO₂: A Theoretical Analysis of Smog Chamber Data, J. Colloid Interface Sci., 78, 513–527, 1980.
- Metzger, A., Verheggen, B., Dommen, J., Duplissy, J., Prevot, A. S., Weingartner, E., Riipinen, I., Kulmala, M., Spracklen, D. V., Carslaw, K. S., and Baltensperger, U.: Evidence for the role
- of organics in aerosol particle formation under atmospheric conditions, P. Natl. Acad. Sci., 107, 6646–6651, doi:10.1073/pnas.0911330107, 2010.
 - Paasonen, P., Nieminen, T., Asmi, E., Manninen, H. E., Petäjä, T., Plass-Dülmer, C., Flentje, H.,
 Birmili, W., Wiedensohler, A., Hõrrak, U., Metzger, A., Hamed, A., Laaksonen, A., Facchini,
 M. C., Kerminen, V.-M., and Kulmala, M.: On the roles of sulphuric acid and low-volatility
 arrania vanaure in the initial store of atmospheric new particle formation. Atmos. Cham.
- ¹⁰ organic vapours in the initial steps of atmospheric new particle formation, Atmos. Chem. Phys. Discuss., 10, 11795–11850, doi:10.5194/acpd-10-11795-2010, 2010.
 - Petäjä, T., Mauldin, III, R. L., Kosciuch, E., McGrath, J., Nieminen, T., Paasonen, P., Boy, M., Adamov, A., Kotiaho, T., and Kulmala, M.: Sulfuric acid and OH concentrations in a boreal forest site, Atmos. Chem. Phys., 9, 7435–7448, doi:10.5194/acp-9-7435-2009, 2009.
- Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K., Teinilä, K., Kerminen, V.-M., Laaksonen, A., and Lehtinen, K. E. J.: Connections between atmospheric sulphuric acid and new particle formation during QUEST III-IV campaigns in Heidelberg and Hyytil, Atmos. Chem. Phys., 7, 1899–1914, doi:10.5194/acp-7-1899-2007, 2007. Salcedo, D., Villalta, P. W., Varutbangkul, V. Wormhoundt, J. C., Miake-Lye, R. C., Worsnop, D.
- ²⁰ R., Ballenthin, J. O., Thorn, W. F., Viggiano, A. A., Miller, T. M., Flagan, R. C., and Seinfeld, J. H., Int. J. Mass. Spectrom. 231, 17–30, 2004.
 - Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Dal Maso, M., Petäjä, T., Riipinen, I., Korhonen, H., Arnold, F., Janson, R., Boy, M., Laaksonen, A., and Lehtinen, K. E. J.: Atmospheric sulphuric acid and aerosol formation: implications from atmospheric measurements for nucleation and early growth mechanisms. Atmos. Chem. Phys., 6, 4079–4091, doi:10.5194/acp-6-4079-
- early growth mechanisms, Atmos. Chem. Phys., 6, 4079–4091, doi:10.5194/acp-6-4079-2006, 2006.
 - Sipilä, M. Berndt, T., Petäjä, T., Brus, D., Vanhanen, J., Stratmann, F., Patokoski, J., Mauldin III, R. L., Hyvärinen, A.-P., Lihavainen, H., and Kulmala, M.: The role of sulfuric acid in atmospheric nucleation, Science, 327, 5970, 1243–1246, doi:10.1126/science.1180315, 2010.
- ³⁰ Sgro, L. A. and de la Mora, J. F.: A simple turbulent mixing CNC for charged particle detection down to 1.2 nm, Aerosol Sci. Technol., 38, 1–11, 2004.
 - Stoltzenburg, M. R. and McMurry, P. H.: An ultrafine aerosol condensation nucleus counter, Aerosol Sci. Technol., 14, 48–65, 1991.

iscussion Pa	ACPD 10, 25959–25989, 2010			
aper I Discussi	Homogenous nucleation of sulfuric acid and water D. Brus et al.			
ion Paper	Title I	Page		
	Abstract Conclusions	Introduction References		
scussion	Tables	Figures		
Paper	4	>1 >		
Dis	Back Full Scre	Close en / Esc		
cussion	Printer-friendly Version			
Pape	Interactive Discussion			



- Vanhanen, J., Mikkilä, J., Sipilä, M., Manninen, H. E., Lehtipalo, K., Siivola, E., Petäjä, T., and Kulmala, M.: Particle size magnifier for nano-CN detection, Aerosol Sci. Technol., in press, 2010.
- Vehkamäki, H., Kulmala, M., Napari, I., Lehtinen, K. E. J., Timmreck, C., Noppel M.,
- and Laaksonen, A.: An improved parameterization for sulfuric acid-water nucleation rates for tropospheric and stratospheric conditions, J. Geophys. Res., 107(D22), 4622, doi:10.1029/2002JD002184, 2002.
 - Weber, R. J., Marti, J., McMurry, P. H., Eisele, F., Tanner, D. J., and Jefferson, A.: Measured atmospheric new particle formation rates: implications for nucleation mechanisms, Chem.
- ¹⁰ Eng. Commun., 151, 53–64, 1996.

20

Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measurements of new particle formation and ultrafine particle growth rates at a clean continental site, J. Geophys. Res., 102, 4373–4385, 1997.

Young, L. H., Benson, D. R., Kameel, F. R., Pierce, J. R., Junninen, H., Kulmala, M., and Lee,

- S.-H.: Laboratory studies of H₂SO₄/H₂O binary homogeneous nucleation from the SO₂+OH reaction: evaluation of the experimental setup and preliminary results, Atmos. Chem. Phys., 8, 4997–5016, doi:10.5194/acp-8-4997-2008, 2008.
 - Zhang, R., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X., Molina, L. T., and Molina, M. J.: Atmospheric new particle formation enhanced by organic acids, Science, 304 (5676),

1487-1490, doi:10.1126/science.1095139, 2004.

- Zhang, R.: Getting to the Critical Nucleus of Aerosol Formation, Science, 328(5984), 1366– 1367, doi:10.1126/science.1189732, 2010.
- Zhao, J., Eisele, F. L., Titcombe, M., Kuang, C., and McMurry, P. H.: Chemical ionization mass spectrometric measurements of atmospheric neutral clusters using the cluster-CIMS,
- ²⁵ J. Geophys. Res., 115, D08205, doi:10.1029/2009JD012606, 2010.



Table 1. Sulfuric acid losses in the first half of the nucleation chamber at three relative hu-
midities, RH (16, 32 and 57%), $T = 25^{\circ}$ C, where k_{obs} is the observed loss rate coefficients, v is
mean flow velocity, t is residence time, WLF is the wall loss factor, and WLF _{inlet} is the wall loss
factor estimated for CIMS' inlet sampling tube.

RH [%]	$k_{\rm obs} [{\rm cm}^{-1}]$	<i>v</i> [cm s ⁻¹]	<i>t</i> [s]	WLF	WLF_{inlet}
57	-0.01	7.9	25.2	6.6	3.3
32	-0.0092	7.4	27.1	6.5	4.4
16	-0.0092	7.3	27.5	6.0	4.6



Table 2. Sulfuric acid losses in the second half of the nucleation chamber at three tempera-
tures, T (25, 15 and 5 °C), RH~50%, where k_{obs} is the observed loss rate coefficients, v is mean
flow velocity, t is residence time, WLF is the wall loss factor, and WLF _{inlet} is the wall loss factor
estimated for CIMS' inlet sampling tube.

<i>T</i> [°C]	$k_{\rm obs} [{\rm cm}^{-1}]$	$v [{\rm cms^{-1}}]$	<i>t</i> [s]	WLF	WLF_{inlet}
25	-0.0063	7.7	25.3	3.6	3.9
15	-0.0069	7.2	27.6	4.0	4.4
5	-0.0072	7.2	27.6	4.2	6.8



Discussion Pa	AC 10, 25959–2	PD 25989, 2010
per Discussion	Homog nucleation acid an D. Brua	jenous of sulfuric d water s et al.
Pape	Title	Page
P	Abstract	Introduction
	Conclusions	References
iscussi	Tables	Figures
on Pa	14	►I.
aper	•	► E
_	Back	Close
Discuss	Full Scre	en / Esc
sion I	Printer-iner	
Paper		

Table 3. The slopes $(d \ln J/d \ln [H_2 SO_4])$ of fits to data in Fig. 1. at three different relative humidities (57, 32 and 16%), and nucleation temperature 25 °C.

RH [%]	slope (PSM)	R^2 (PSM)	slope (3025A)	R ² (3025A)
57	1.7	0.96	2.9	0.94
32	1.3	0.75	2.4	0.83
16	1.5	0.93	2.0	0.97

Discussion I	AC 10, 25959–	ACPD 10, 25959–25989, 2010		
Paper Discussio	Homog nucleation acid an D. Bru	genous of sulfuric od water s et al.		
n Pape	Title	Page		
_	Abstract	Introduction		
	Conclusions	References		
)iscussi	Tables	Figures		
ion F	I.€	►I.		
Daper	•	Þ		
	Back	Close		
Discuss	Full Scre	Full Screen / Esc		
ion	Printer-Intendiy Versión			
Paper		Discussion		

Table 4. The slopes $(d \ln J/d \ln [H_2SO_4])$ of fits to data in Fig. 2. at three different temperatures (5, 15 and 25 °C) and relative humidity ~50%.

<i>T</i> [°C]	slope (PSM)	R^2 (PSM)	slope (3025A)	R ² (3025A)
5	2.2	0.99	2.9	0.99
15	1.5	0.99	1.9	0.99
25	1.2	0.98	2.1	0.99

Discussion Pa	ACPD 10, 25959–25989, 2010			
per Discussior	Homogenous nucleation of sulfuric acid and water D. Brus et al.			
n Pap	Title	Page		
er	Abstract	Introduction		
	Conclusions	References		
iscuss	Tables	Figures		
ion P	14	۶I		
aper	•	•		
_	Back	Close		
Discu	Full Screen / Esc			
Printer-friendly Version				
1 Pap	Interactive Discussion			
)er	•			

Table 5. Calculated median kinetic (K) and activation (A) coefficients at different levels of relative humidity, exponent E is taken from the linear fit to PSM data of this study.

RH [%]	$K [{\rm cm}^3{\rm s}^{-1}]$	<i>A</i> [s ⁻¹]	E (PSM)
57	1.4×10 ⁻¹⁵	7.2×10 ⁻⁷	1.7
32	6.9×10 ⁻¹⁶	4.9×10 ⁻⁷	1.3
16	1.3×10 ⁻¹⁶	3.9×10 ⁻⁸	1.5

Discussion Pa	AC 10, 25959–2	ACPD 10, 25959–25989, 2010		
per Discussior	Homog nucleation acid an D. Brus	jenous of sulfuric d water s et al.		
1 Pap	Title	Page		
θŗ	Abstract	Introduction		
	Conclusions	References		
iscuss	Tables	Figures		
tion P	14	۶I		
aper	•	•		
_	Back	Close		
Discu	Full Scre	en / Esc		
ssion	Printer-friendly Version			
1 Pap	Interactive	Interactive Discussion		
)er	@ ()			

Table 6. Calculated median kinetic (K) and activation (A) coefficients at different nucleation temperatures, exponent E is taken from the linear fit to PSM data of this study.

<i>T</i> [°C]	<i>K</i> [cm ³ s ⁻¹]	A [s ⁻¹]	E (PSM)
5	1.0×10^{-14}	1.0×10^{-6}	2.2
15	3.6×10 ⁻¹⁵	8.5×10 ⁻⁷	1.5
25	8.0×10 ⁻¹⁶	2.6×10 ⁻⁶	1.2



Fig. 1. Nucleation rates of sulfuric acid and water as a function of sulfuric acid initial concentration, nucleation temperature T=25 °C. Particle number concentration measured with UCPC TSI 3025 A (A) and particle size magnifier (PSM) (B).











Fig. 3. Sulfuric acid dimer concentration as a function of monomer concentration at three different relative humidities and nucleation temperature T = 25 °C. The slope of the fit to all data is 1.08 and $R^2 = 0.76$. The M/D ratio is increasing from ~100 to ~200 with increasing relative humidity from 16 to 57%.





Fig. 4. Monomer to dimer ratio as a function of three different nucleation temperatures T=5, 15, 25 °C.

Full Screen / Esc

Printer-friendly Version

Interactive Discussion

Discussion Paper



Fig. 5. The initial sulfuric acid monomer concentration measured by CIMS as a function of initial total sulfate concentration determined by ion chromatography analyses and subsequent mass balance calculations.







