



## Abstract

Temperature profiles from two satellite instruments – TIMED/SABER and Aura/MLS – have been used to calculate hydroxyl-layer equivalent temperatures for comparison with values measured from OH(6-2) emission lines observed by a ground-based spectrometer located at Davis Station, Antarctica (68° S, 78° E). The profile selection criteria – <500 km from the ground station and solar zenith angles > 97° – yielded a total of 2359 SABER profiles over 8 years (2002–2009) and 7407 MLS profiles over 5.5 years (2004–2009). The availability of simultaneous OH volume emission rate (VER) profiles from the SABER (OH-B channel) enabled an assessment of the impact of several different weighting functions in the calculation of OH-equivalent temperatures. The maximum difference between all derived hydroxyl layer equivalent temperatures was less than 3 K. Restricting the miss-distance and miss-time criteria showed little effect on the bias, suggesting that the OH layer is relatively uniform over the spatial and temporal scales considered. However, a significant trend was found in the bias between SABER and Davis OH of  $\sim 0.7$  K/year over the 8-year period with SABER becoming warmer compared with the Davis OH temperatures. In contrast, Aura MLS exhibited a cold bias of  $9.9 \pm 0.4$  K compared with Davis OH, but importantly, the bias remained constant over the 2004–2009 year period examined. The difference in bias behaviour of the two satellites has significant implications for multi-annual and long-term studies using their data.

## 1 Introduction

The SABER (Sounding of the Atmosphere by Broadband Emission Radiometry) instrument on board the NASA's TIMED (Thermosphere Ionosphere Mesosphere Electrodynamics) satellite has been providing temperature profiles as one of its level 2 products since early 2002 (Mertens et al., 2001). Several reports of comparisons between SABER temperatures and those measured by ground-based techniques and

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instruments have been published in the past few years (e.g., Mertens et al., 2004; Oberheide et al., 2006; Xu et al., 2006; López-González et al., 2007; Mulligan and Lowe, 2008; Remsberg et al., 2008; Smith et al., 2010). Each report provides an estimate of the offset or bias between SABER and the comparator dataset (see supplemental Table 2). All of these studies involve relatively short data runs, the longest of which was the three-year (2003–2006) study of Oberheide et al. (2006) and they all involve comparisons with data from Northern Hemisphere sites only.

This work differs from previous studies in the following respects: it uses ground-based temperature data from a Southern Hemisphere station – Davis Station, Antarctica (68° S, 78° E) and it includes the years 2002–2009 which enables us to examine the behaviour of the bias between SABER and the Davis dataset during this extended period. The instrument at Davis has been well documented (Greet et al., 1998; French et al., 2000) and has been providing well characterised temperature data in every Southern Hemisphere winter season since 1995 (Burns et al., 2003; French and Burns, 2004; French et al., 2005). We also compare our ground-based observations with temperature profiles from a second satellite instrument – the Microwave Limb Sounder (MLS) on board the EOS (Earth Observing System) Aura spacecraft (Schwartz et al., 2008) – over the period 2004–2009 to help contextualise our SABER-Davis OH results.

## 2 Instrumentation and data

### 2.1 Davis spectrometer

A 1.26 m f/9 Czerny-Turner scanning spectrometer has operated at Davis each year since 1995 recording the P-branch lines of the OH(6-2) band near 840 nm. Observations are made in the zenith with a 5.3° field-of-view and with an instrument resolution of ~ 0.16 nm. A cooled GaAs photomultiplier, operated in pulse counting mode, detects the sky emission. Acquisition time is of the order of 7 min per spectrum. Further details of the instrument are contained in Greet et al. (1998) and French et al. (2000).



orientation at ANMI. A measurable difference in the spectral radiance of the LBS was detected in 2007 as a result of the altered shape of the tungsten filament in the different orientations of the LBS. Correcting the ANMI calibrations to match the Davis vertical measurements results in cooler OH rotational temperatures by an average 0.9 K.

## 2.2 SABER

The SABER instrument is a radiometer which measures Earth limb emission profiles over the spectral range 1.27–17  $\mu\text{m}$  from the TIMED satellite in circular orbit at 625 km inclined at 74° to the equator (Russell et al., 1999). The latitude coverage ranges from 54° S to 82° N or 82° S to 54° N depending on the yaw cycle. The satellite orbit precesses slowly to complete 24 h LT (local time) in 60 days. Temperature is retrieved from 15  $\mu\text{m}$  and 4.3  $\mu\text{m}$  CO<sub>2</sub> emissions over an altitude range of 10–105 km, with a vertical resolution of about 2 km, and along track resolution of 400 km (Mertens et al., 2002). A summary of the evolution of SABER data releases is provided by Remsberg et al. (2008). Retrievals employ iterative algorithms in the upper mesosphere lower thermosphere (UMLT) region to account for non-local thermodynamic equilibrium (non-LTE) radiative transfer effects. An assessment of these algorithms has been made by Mertens et al. (2001, 2002, 2004) and further detail of the v1.07 non-LTE retrieval algorithm is given by García-Comas et al. (2008).

Errors in the retrieved temperatures in the 80–100 km region are estimated to be in the range  $\pm 1.5$ –5 K if the kinetic temperature profile does not have pronounced vertical structure (García-Comas et al., 2008). Errors are greater in summer polar conditions, but these are not of concern here since our study concentrates on Southern Hemisphere polar winter. In addition to kinetic temperature, volume emission rate (VER) profiles are derived simultaneously from the OH-B channel, sensitive in the range 1.56–1.72  $\mu\text{m}$ , which includes mostly the OH(4-2) and OH(5-3) bands. While we compare temperatures from OH(6-2) band measurements at Davis in this study, altitude differences between all the vibrational levels are not expected to exceed 2 km (McDade, 1991). SABER calibration is maintained by reference to an internal blackbody

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(at 247 K) every 446 s. Between each up/down limb scan pair, views are also made of “cold space” and the instrument baffling as an estimate of internal stray light effects (Remsberg et al., 2008).

## 2.3 EOS Aura Microwave Limb Sounder

The EOS (Earth Observing System) Aura spacecraft was launched on 15 July 2004 into a near-polar, 705 km altitude, sun-synchronous orbit (Schwartz et al., 2008). It provides almost complete global coverage (82° S–82° N) with ~ 13 orbits per day. The MLS (Microwave Limb Sounder) instrument on board Aura observes thermal microwave emissions in different regions from 115 GHz to 2.5 THz. The MLS field-of-view is in the direction of orbital motion, and it scans the Earth’s limb vertically from ~ 5–100 km every 24.7 s. Temperature and geopotential height profiles are produced on a fixed vertical pressure grid by the MLS level 2 (version 2.2 data is used in this study) processing algorithm which is applied to the thermal microwave emissions near the spectral lines 118 GHz O<sub>2</sub> and 234 GHz O<sup>18</sup>O. It produces scientifically useful temperature profiles for geopotential heights corresponding to the range 316 hPa (~8 km) to 0.001 hPa (~97 km) (Schwartz et al., 2008). Recent studies utilizing MLS temperature measurements in the MLT region include an investigation of summer planetary-scale oscillations by Meek and Manson (2009) and of the quasi 5-day wave by von Savigny et al. (2007).

## 3 Data quality and analysis

### 3.1 Davis OH(6-2)

Hydroxyl observations at Davis are made when the sun is more than 7° below the horizon, which defines an observing season window between 8-February (day 049) and 23-October (day 296). This interval includes the warm winter mesopause period with partial coverage of the spring and autumn transition periods. No observations can be

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obtained during the cold summer mesopause period. Nightly averages are derived from all 7 min spectra that pass data quality selection criteria. Application of these criteria discards measurements with weighted standard deviation  $> 15$  K, counting errors  $> 10$  K, high backgrounds, large background slopes, and unusual changes of intensity between consecutive scans. Burns et al. (2003) have examined the influence of cloud, aurora and Fraunhofer absorption in moonlight in these data. A total of 1677 nightly averages are obtained for the 2002 to 2009 interval from nearly 125 000 individual spectra that pass the selection criteria. A minimum of 10, maximum of 163 (average of 74) samples contribute to each nightly average. Figure 1a shows these nightly averages with  $1\sigma$  error bars in comparison with the MSISE-90 model temperatures for 87 km at  $69^\circ$  S. The model shows good agreement with the observations at Davis, and illustrates the full seasonal cycle at this latitude.

### 3.2 SABER

SABER version 1.07 data (available at <http://saber.gats-inc.com>) were used in this study. A total of 2547 profiles (typically 2–4 profiles/day) satisfied the selection criteria of tangent point within a 500 km radius of Davis and solar zenith angle  $> 97^\circ$  (night-time profiles corresponding to the spectrometer measurements) over the 2002–2009 interval. The satellite 60-day yaw cycle results in SABER observations over Davis in the same three time intervals each year (days 75–140, 196–262 and 323–014). However, the solar zenith angle criterion ( $> 97^\circ$ ) completely rejects the summer interval (days 323–014). Temperature profiles were rejected if the VER was unusual, viz, the FWHM of the VER profile was  $> 2\sigma$  from the mean profile width ( $4.8 \text{ km} < \text{FWHM} < 11.5 \text{ km}$ ) or if the pressure interpolated to 87 km was  $> 2\sigma$  from mean pressure ( $0.0012 < \text{pressure} < 0.005 \text{ mb}$ ) at that altitude (188 profiles rejected, leaving 2359). Profiles were also rejected if their VER profiles departed significantly from a Gaussian shape (measured by the  $\chi^2$  fit value) mostly eliminating double peaked VER profiles in a similar fashion to Winick et al. (2009). Following the

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application of these criteria, only those profiles that were within 8 h of the closest OH(6-2) nightly average were retained giving a total of 2060 profiles.

In order to compare satellite temperatures with those derived from the ground-based instrument, we derive hydroxyl layer equivalent temperatures from the temperature profiles using a weighting function which is representative of the hydroxyl layer (e.g., Oberheide et al., 2006; López-González et al., 2007; Mulligan and Lowe, 2008). In the first part of this study, we concentrated on SABER profiles because of the availability of simultaneous OH 1.6  $\mu\text{m}$  VER profiles. Five different weighting profiles were considered as follows:

- W\_SL: The weighting function from She and Lowe (1998) of the form.

$$W_{\text{SL}} = (a / (1 + \exp(-(z - b + c/2)/d))) (1 - 1 / (1 + \exp(-(z - b - c/2)/e)))$$

where the coefficients  $a = 27$ ;  $b = 87.133$ ;  $c = 3.8682$ ;  $d = 1.8235$ ;  $e = 2.9306$ , calculated over the tangent point altitude ( $Z$ ) range  $70 < Z < 110$  km.

- W\_G87: A Gaussian centred at 87 km with FWHM 8 km calculated over the range  $70 < Z < 104$  km.
- W\_VER: the SABER (OH-B) VER profile, evaluated from the altitude of the peak OH VER  $\pm 20$  km.
- W\_GFIT: A Gaussian fitted to the altitude of the OH VER peak  $\pm 20$  km, but then weighted to  $\pm 17$  km.
- W\_VERm: A profile derived from the mean of the 2359 VER profiles that comprise the initial selection set.

These weighting profiles are shown in Fig. 2 with the mean (and standard error) of the 2359 SABER temperature profile retrievals across the MLT region. Multiplying each temperature profile by the corresponding weighting profile yields  $T_{\text{SL}}$ ,  $T_{\text{G87}}$ ,  $T_{\text{VER}}$ ,  $T_{\text{GFIT}}$  and  $T_{\text{VERm}}$ . The SABER temperature interpolated to 87 km ( $T_{\text{Alt}}$ ) was also calculated.  $T_{\text{GFIT}}$  temperatures, derived in this way are shown in Fig. 1b.

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## 4 Comparison of SABER with Davis OH(6-2)

### 4.1 The effect of selection criteria on bias and correlation

SABER weighted temperatures, for each profile shape, are compared to the closest OH(6-2) rotational temperature, as individual samples (2060 coincidences) and as nightly averages (847 coincidences). The differences (SABER – Davis OH) are shown in Fig. 1c for the T\_GFIT evaluation. The mean difference ( $T_{\text{GFIT}} - T_{\text{OH}}$ ) in this example is 1.25 K with a standard error of 0.27 K. Table 1 summarizes the mean differences for the 2060 individual profiles, and 847 nightly average comparisons. Overall, SABER and OH temperatures lie within 2 K of each other, no matter which weighting profile is chosen, and whether the comparison is made on an individual profile or nightly average basis. Profiles based on the SABER VER ( $T_{\text{VER}}$ ,  $T_{\text{GFIT}}$  and  $T_{\text{VERm}}$ ) generally lead to a warm bias while weighting with the She and Lowe (1998) profile leads to a cool bias. The smallest standard deviation in the bias is obtained with the  $W_{\text{VERm}}$  profile. Spatial and temporal variability may contribute to the differences between Davis OH and SABER temperatures since we have accepted all overpasses within 500 km in range and 8 h in time. The effect on bias has been examined with more restrictive spatio-temporal coincidence criteria; restricting the range of miss-distance acceptance in 100 km steps and miss-time acceptance from  $\pm 8$  h to  $\pm 15$  min. No systematic trend in the bias is observed with tighter restrictions on either criterion for any of the weighting functions considered (see supplemental Tables 3 and 4 available on the web). The consistency of biases over the span of spatial and temporal ranges considered indicate that the OH layer is largely uniform over these scales, or that differences are averaged out in the wide field-of-view of SABER. Burns et al. (2003) found good correlation between Davis OH and sodium lidar temperatures at Syowa station (69° S, 39° E) at a distance of 1500 km. The majority of the day-to-day variability is in response to large scale planetary waves with spatial extents of several thousand km (French and Burns, 2004; French et al., 2005). Tidal amplitudes are small at this latitude (generally less

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than 2 K) which contributes to the lack of bias deviation over timescales up to 8 h (Hagan et al. 1999).

## 4.2 SABER-Davis OH bias trend

Our analysis thus far has treated the 2002–2009 comparison dataset as a whole. Returning to Fig. 1c we now examine the trend in bias over the 8-year time span. Means for each year, and for each yaw cycle “campaign”, are plotted in Fig. 3a. A linear fit shows an apparent trend in the bias increasing at  $\sim 0.7$  K/year (implying that SABER temperatures are getting warmer compared to the OH measurements) and is independent of the weighting function used. To our knowledge such a trend has not been reported in previous comparisons.

A calibration drift in either instrument could produce such a trend. The total change in response correction for the Davis instrument over 2002–2009 equates to about 1.1% or 1.1 K in temperature from the  $P_1(2)/P_1(5)$  ratio (which contributes  $\sim 57\%$  on average to the weighted mean) and less for the other ratios. The uncertainty in the response calibration is  $\sim 0.3$  K. Thus in terms of a calibration drift for the Davis instrument a change in bias of  $\sim 0.7$  K/year (a total of 5.6 K over 8 years) is conservatively at least 5 times the maximum change in response correction and at least 19 times the uncertainty.

Another possibility for the source of a bias trend could be an inaccurate representation of long term changes in  $\text{CO}_2$  or O concentrations in the SABER non-LTE retrieval algorithm, perhaps as a consequence of the decrease in solar flux over the life of the SABER mission (J.-H. Yee, personal communication, 2010). If SABER is the source of the drift, it has important implications for the use of its temperatures in the assessment of long term change. Further information on the bias trend may be gained over the next few years as solar activity increases, but for now the source of the trend in bias remains unresolved.

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### 4.3 Aura MLS-Davis OH bias and trend

As a check on the bias trend we have performed the same comparison exercise using available data from the Microwave Limb Sounder (MLS) on NASA's Aura satellite. A total of 7407 MLS Level 2 (version 2.2) profiles between July 2004 and December 2009 satisfied the same acceptance range ( $< 500$  km) and solar zenith angle ( $> 97^\circ$ ) criteria used for the SABER analysis. Since our SABER analysis demonstrated that OH-equivalent temperatures are relatively insensitive to the shape of the weighting function, we used only a simple Gaussian weighting profile centred on an altitude of 87 km and a FWHM of 8.7 km to calculate hydroxyl layer equivalent temperatures. A significant cold bias of  $9.9 \pm 0.4$  K (MLS colder than Davis OH) was obtained as the mean of 1034 nightly averages over the 5.5 years. Importantly, however, biases derived for each year do not show an increasing drift between 2004 and 2009 as shown in Fig. 3b.

### 5 Conclusions

A comparison of 8 years (2002–2009) of SABER temperature profiles with Davis hydroxyl rotational temperatures derived with Langhoff et al. (1986) transition probabilities shows:

- the SABER-Davis OH bias was within 2 K averaged over 2060 measurement pairs, and within 1.25 K averaged over 847 nightly averages, for all weighting profiles considered.
- Profiles based on the SABER VER (T\_VER, T\_GFIT and T\_VERm) generally lead to a warmer SABER-OH bias while the She and Lowe (1998) and Gaussian at 87 km weighting profiles generally lead to a cooler bias.
- More restrictive spatio-temporal coincidence criteria have little effect on the biases obtained indicating that the hydroxyl layer is largely uniform over the scales considered here.

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- The SABER-OH bias appears to be increasing at  $\sim 0.7$  K per year over the comparison interval. To our knowledge such a trend has not been previously reported.
- This trend is not apparent in equivalent comparisons with Aura/MLS measurements.

5 Recently, Remsberg et al. (2008) have reported an assessment of the quality of the version 1.07 temperature-pressure profiles of the middle atmosphere from TIMED/SABER. They conclude that the SABER v1.07 temperature distributions can be used to generate the near-global, seasonal and interannual variations of  $T_k$  in the UMLT for the period 2002 to 2008. We agree that the SABER temperatures are a unique  
10 resource for researchers in the UMLT region. Our detailed comparison of SABER temperatures with OH temperatures recorded at Davis over an 8-year period (2002–2009) suggest that some caution may be needed when using SABER temperatures over an extended time-period.

**Supplementary material related to this article is available online at:**

15 **[http://www.atmos-chem-phys-discuss.net/10/21547/2010/  
acpd-10-21547-2010-supplement.pdf](http://www.atmos-chem-phys-discuss.net/10/21547/2010/acpd-10-21547-2010-supplement.pdf)**

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**Table 1.** A comparison of the mean temperatures and SABER-OH biases for each weighting profile for 2060 individual profiles and 847 nightly averages compared to the closest OH temperature measurement ( $T_{\text{OH}}$ ) and closest OH nightly average ( $T_{\text{OHNA}}$ ).

SABER Profiles	$T_{\text{OH}}$	$T_{\text{OHNA}}$	$T_{\text{Alt}}(\text{K})$	$T_{\text{SL}}(\text{K})$	$T_{\text{G87}}$	$T_{\text{VER}}$	$T_{\text{GFIT}}$	$T_{\text{VERm}}$
2060 Profiles								
Weighted Temp (K)	202.00	203.39	203.76	201.30	202.38	203.81	203.80	203.19
Standard Deviation (K)	12.28	7.43	13.60	9.91	9.85	9.77	10.28	8.74
Standard Error (K)	0.27	0.16	0.30	0.22	0.22	0.22	0.23	0.19
Difference from $T_{\text{OH}}$			1.76	-0.70	0.38	1.81	1.80	1.19
SD of difference			14.72	12.75	12.00	11.90	11.90	11.54
Difference from $T_{\text{OHNA}}$			0.37	-2.09	-1.01	0.42	0.41	-0.20
SD of difference			12.44	9.33	8.59	8.39	8.51	7.63
Linear correlation $R^2$			0.179	0.203	0.287	0.305	0.336	0.319
Nightly Averages								
		$T_{\text{OHNA}}$	$T_{\text{Alt}}$	$T_{\text{SL}}$	$T_{\text{G87}}$	$T_{\text{VER}}$	$T_{\text{GFIT}}$	$T_{\text{VERm}}$
847 Profiles								
Weighted Temp (K)		203.11	204.26	201.97	202.91	204.34	204.35	203.65
Standard Deviation (K)		7.34	11.77	9.42	9.18	9.03	9.58	8.18
Standard Error (K)		0.3	0.4	0.3	0.3	0.3	0.3	0.3
Difference from $T_{\text{OHNA}}$			1.16	-1.14	-0.19	1.23	1.25	0.54
SD of the difference			10.56	8.98	7.99	7.67	7.85	7.20

## Bias trends between Hydroxyl, SABER and MLS temperatures

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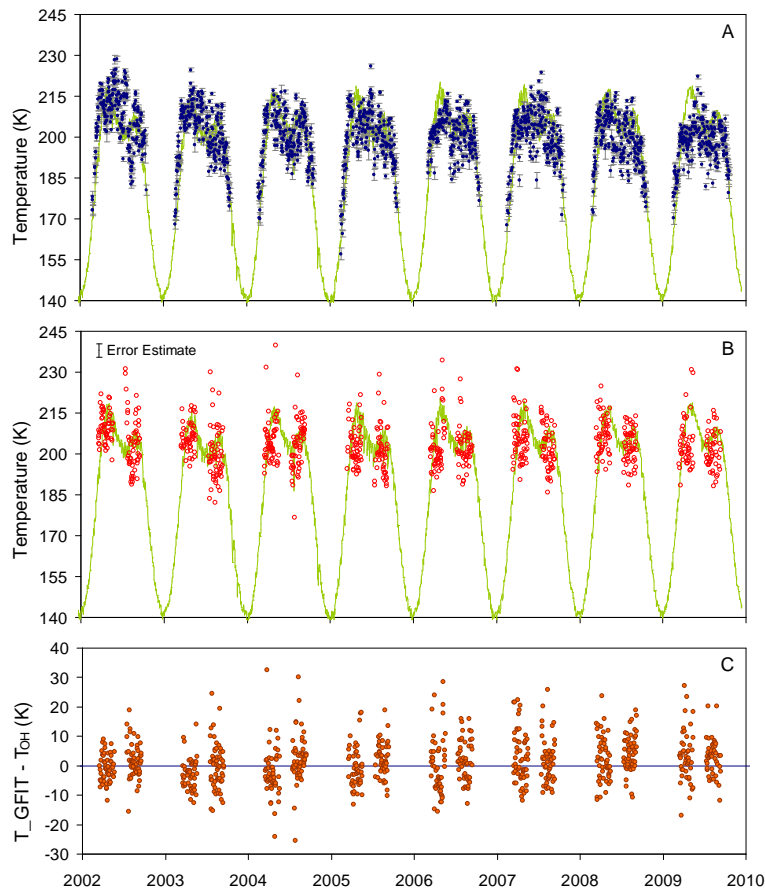
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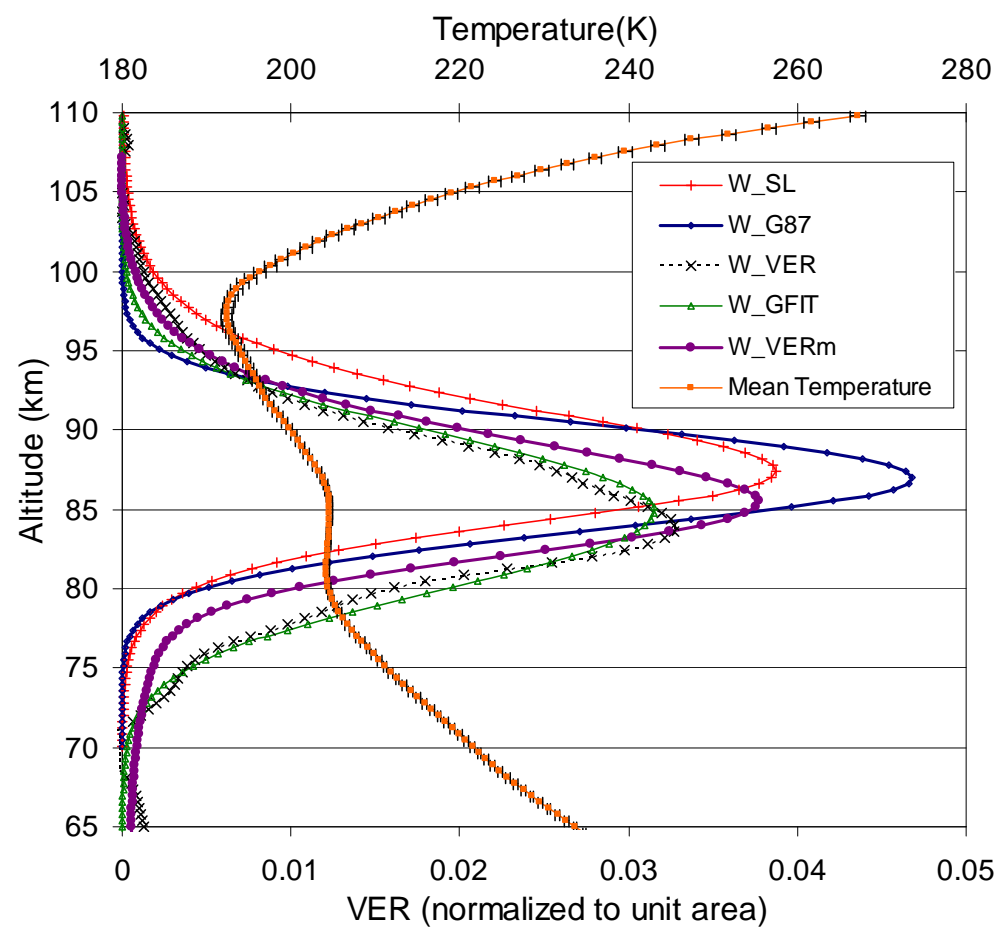
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**Fig. 1.** (A): Davis hydroxyl nightly average temperatures (1677 nights) overplotted on MSISE-90 model temperatures for 69° S. Error bars for the Davis OH series are 1 standard error-in-the-mean. Panel (B): SABER  $T_{\text{GFIT}}$  nightly averages (847 nights) with an estimate of the SABER error (5 K) plotted for comparison with the OH temperatures. Panel (C): the SABER-Davis OH difference for coincident nights.

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**Fig. 2.** SABER weighting profiles considered in this study (see text for descriptions), compared with the mean (winter) mesopause temperature profile derived from the 2359 SABER profiles within 500 km range.

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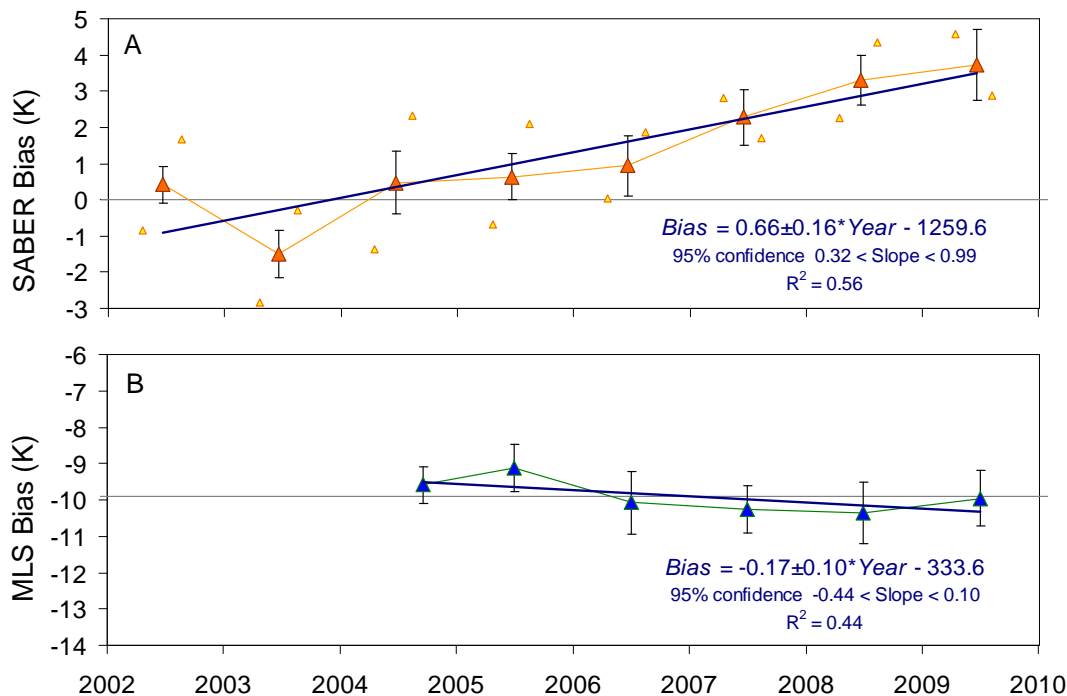
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**Fig. 3.** Panel (A): The mean SABER-DavisOH bias for each year, and for each yaw-cycle campaign, from Fig. 1c. A straight line fit to the latter shows a trend in the bias of  $\sim 0.7$  K/annum. Panel (B): MLS-DavisOH bias for each year. A straight-line fit to this data is not significantly different from zero trend at the 95% confidence level.

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