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Ship-based detection of glyoxal over the remote tropical Pacific Ocean

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Abstract

We present the first detection of glyoxal (CHOCHO) over the remote tropical Pacific Ocean in the Marine Boundary Layer (MBL). The measurements were conducted by means of the University of Colorado Ship Multi-Axis Differential Optical Spectroscopy (CU SMAX-DOAS) instrument aboard the research vessel Ronald H. Brown. The research vessel was on a cruise in the framework of the VAMOS Ocean-Cloud-Atmosphere-Land Study – Regional Experiment (VOCALS-REx) and the Tropical Atmosphere Ocean (TAO) projects lasting from October 2008 through January 2009 (74 days at sea). The CU SMAX-DOAS instrument features a motion compensation system to characterize the pitch and roll of the ship and to compensate for ship movements in real time. We found elevated mixing ratios of up to 170 ppt CHOCHO located inside the MBL up to 3000 km from the continental coast over biologically active upwelling regions of the tropical Eastern Pacific Ocean. This is surprising since CHOCHO is very short lived (atmospheric life time ~ 2 h) and highly water soluble (Henry's Law constant $H=4.2 \times 10^5$ M/atm). This CHOCHO cannot be explained by transport of it or its precursors from continental sources. Rather, the open ocean is a source for CHOCHO to the atmosphere. Dissolved Organic Matter (DOM) photochemistry in surface waters is a source for Volatile Organic Compounds (VOCs) to the atmosphere, e.g. acetaldehyde. The extension of this mechanism to very soluble gases, like CHOCHO, is not straightforward since the air-sea flux is directed from the atmosphere into the ocean. For CHOCHO, the dissolved concentrations would need to be extremely high in order to explain our gas-phase observations by this mechanism (40–70 μM CHOCHO, compared to ~ 0.01 μM acetaldehyde and 60–70 μM DOM). Further, while there is as yet no direct measurement of VOCs in our study area, measurements of the CHOCHO precursors isoprene, and/or acetylene over phytoplankton bloom areas in other parts of the oceans are too low (by a factor of 10–100) to explain the observed CHOCHO amounts. We conclude that our CHOCHO data cannot be explained by currently understood processes. Yet, it supports first global source estimates of 20 Tg/year CHOCHO from the

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oceans, which likely is a significant source of secondary organic aerosol (SOA). This chemistry is currently not considered by atmospheric models.

1 Introduction

Glyoxal (CHOCHO) is the smallest α -dicarbonyl compound. It is primarily formed by oxidation of numerous Volatile Organic Compounds (VOCs) such as aromatics and alkenes, but is also emitted directly into the atmosphere (Calvert et al., 2000, 2002; Volkamer et al., 2001). VOC oxidation, and thus the CHOCHO production, is driven by photochemical activity in the atmosphere. Although high peak concentrations could be observed in urban areas mainly due to anthropogenic activity (Volkamer et al., 2005a; Sinreich et al., 2007), the largest global source of CHOCHO is from the oxidation of biogenic VOC emissions, primarily isoprene (Fu et al., 2008; Stavrou et al., 2009). Major sources also are biomass burning and biofuel usage (Fu et al., 2008). The total continental source of CHOCHO to the atmosphere is constrained by satellite measurements, and ranges between 94 and 108 Tg/year (Stavrou et al., 2009), of which about 50% are currently explained in terms of known sources (Myriokefalitakis et al., 2008; Fu et al., 2008); the remaining 50% of the CHOCHO source is currently unaccounted for. Satellite data (Wittrock et al., 2006) in combination with model assumptions regarding the sinks of CHOCHO (TM4) made a first quantification of this source, and indicate that tropical oceans could add an additional 20 Tg/year to the global CHOCHO source (Myriokefalitakis et al., 2008). CHOCHO is rapidly lost by photolysis and reactions with OH radicals. It can also be taken up in cloud droplets (Ervens et al., 2004), and/or aqueous particles (Liggio et al., 2005; Volkamer et al., 2009a) and form secondary organic aerosols (SOA). The global SOA source from CHOCHO over land is currently estimated to range from 6.4 to 21 Tg/year (Fu et al., 2008, Stavrou et al., 2009). The atmospheric life time of CHOCHO is about 1.3 h for overhead sun conditions (Volkamer et al., 2005a), and the global mean life time is between 2.5 and 3 h (Myriokefalitakis et al., 2008; Fu et al., 2008; Stavrou et al., 2009).

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Although CHOCHO has been measured for over two decades (Kieber et al., 1990), CHOCHO measurements continue to be scarce. Also, almost all observations were performed over land and in the Northern Hemisphere (Volkamer et al., 2010). Information over the remote ocean is only available from several satellite instruments. Wittrock et al. (2006) and Vrekoussis et al. (2009) published global maps of CHOCHO using the SCanning Imaging Absorption SpectroMeter for Atmospheric Cartography (SCIAMACHY). Worldwide CHOCHO distribution maps also were derived from the Ozone Monitoring Instrument (OMI, available at <http://www.cfa.harvard.edu/~tkurosui/SatelliteInstruments/OMI/SampleImages/CHOCHO/index.html>) and from the Global Ozone Monitoring Experiment 2 (GOME-2). However, they vary strongly in the amount of CHOCHO, which reflects the fact that measurements from space are inherently difficult to conduct due to low albedo over oceans and cloud coverage (specifically over the tropics).

Attempts to calculate the global CHOCHO distribution have been made using atmospheric models, including TM4 (Myriokefalitakis et al., 2008), GEOS-Chem (Fu et al., 2008) and IMAGES (Stavrakou et al., 2009). However, a comparison with satellite measurements shows a large mismatch in remote tropical ocean regions: While satellites observe CHOCHO over the remote tropical ocean the models do not predict any CHOCHO there. This triggered the question, whether chemical models are missing sources, or whether the satellite retrievals have been affected by a retrieval artifact. Since satellites point to the Earth's surface the latter could be due to absorption interference with chlorophyll in the ocean which is found in parts of the ocean where CHOCHO was retrieved from satellite (see http://earthobservatory.nasa.gov/GlobalMaps/view.php?d1=MY1DMM_CHLORA).

In order to investigate this mismatch we have developed the CU SMAX-DOAS device, which is optimized for the deployment on ships (see Sect. 3 and Volkamer et al., 2009b), and performed MAX-DOAS measurements of CHOCHO on the research vessel Ronald H. Brown during its cruise through the tropical Pacific Ocean between October 2008 and January 2009.

2 Measurement technique

For the measurement of CHOCHO the well established Differential Optical Absorption Spectroscopy (DOAS) technique was applied (Platt, 1994). DOAS uses the position and optical density of narrow band absorption features (<5 nm width) to selectively detect and quantify trace gases by applying the Lambert–Beer law. It is a remote sensing technique which is inherently calibrated and which unequivocally identifies the trace gases of interest by their characteristic narrow band absorption structures (Platt, 1999). In particular, the Multi-Axis (MAX-)DOAS technique employed here uses scattered sunlight observed from multiple viewing directions. This increases the sensitivity to trace gases close to the surface due to differences in the respective light path distributions, i.e. a lower “elevation angle” leads to longer light paths through a trace gas layer near the instrument altitude and thus to a stronger absorption signal (elevation angle is defined as the angle between the horizontal and the pointing direction of the telescope) (e.g., Hönninger and Platt, 2002; van Roozendaal et al., 2003; Hönninger et al., 2004; Wagner et al., 2004). Additionally, measurements at different elevation angles provide information about the vertical profile of a trace gas. When assuming a well mixed trace gas layer (e.g. within the atmospheric boundary layer), its vertical extent and concentration can be determined with good accuracy (e.g., Wittrock et al., 2004; Sinreich et al., 2005; Heckel et al., 2005; Frieß et al., 2006). To determine this, typically the trace gas O_4 is used. O_4 is the oxygen dimer molecule whose concentration is proportional to the concentration of O_2 squared. Thus, the concentration profile shape is well known and is quantitatively dependent only on the air pressure. For this reason O_4 can be used as a valid tracer of the light path distribution in the atmosphere (e.g., Wagner et al., 2004; Sinreich et al., 2005; Frieß et al., 2006).

MAX-DOAS measurements from ships are rare so far and go along with the challenge of the stabilization of the viewing angle (Kreher et al., 1995; Leser et al., 2003; Wagner et al., 2007; Martin et al., 2009). For this cruise, we actively accounted for the pitch and roll of the ship, as described in Sect. 3.

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3 Instrumentation

The University of Colorado Ship MAX-DOAS (CU SMAX-DOAS) instrument consists of a telescope, spectrograph/detector units, and a laptop PC. The telescope collects scattered sunlight, and was mounted on “deck 03 forward” close to the bow of the ship. The telescope is comprised of a quartz prism and a quartz lens (cone angle of 0.3°) which focuses incoming light on a quartz fiber fixed in the focal plane. The prism is mounted to rotate vertically by a stepper motor (0.09° motor precision) allowing a rotation of 360° . In zero position (view to horizon), the telescope pointed to 256° (clockwise) relative to the bow tip of the ship. Solar stray light spectra were acquired sequentially from the following elevation angles: 90° , 0.25° , 0.6° , 1.5° , 3.8° , 10° , 25° , 90° , 179.75° , 179.4° , 178.5° , 176.2° , 170° , 155° , 205° , and -45° . An internal motor encoder reads the position of the motor axis (encoder accuracy 0.18°), and software assures that the motor positions are reached. Two inclinometers arranged perpendicularly to each other measured roll and pitch of the ship during measurement with a precision of 0.1° each. The information from these sensors is processed by software into the coordinate system along the line of sight of the telescope, and is used to drive the stepper motor in such way that it compensates for the ship’s movement and keeps the elevation angle on a software determined target level. A third inclinometer was periodically mounted on top of the prism housing to assure the accuracy of our angle stabilization: At a swell of up to 6° roll (peak to peak) and up to 1.3° pitch (peak to peak) the deviations were $\pm 0.56^\circ$ standard deviation. The integration time of each spectrum was set to 30 s; one sequence of elevation angles was completed after about 8 min. The light is focused by the telescope on a single fiber which is coupled to 3 spectrograph/detector units “QE65000” of Ocean Optics Inc. equipped with two-dimensional CCD-detectors with 1044 wavelength channels (model Hamamatsu S7031-1006). The spectrograph/detector unit used for the CHOCHO measurements projected a wavelength range from about 415 to 524 nm. The optical spectrometer bench was heated to a temperature of 30°C to minimize changes in optical projecting

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properties while the detector itself was cooled to 0°C to reduce dark current. In leg 1, the spectrometer temperature was stable to ±0.1°C except from 21 to 23 October when problems with the heating foil of the spectrometer limited the stability to ±3°C. In leg 2 and 3, the spectrometer temperature was stable to ±0.03°C. The optical resolution is 0.5 nm FWHM (full width at half maximum), which was determined by measuring the second order of the 253.65 nm mercury emission line at 507.30 nm. This emission line was also used to determine the instrumental slit function and later to convolve high resolution literature cross sections to the optical resolution of our instrument. The measurement routine was controlled by a laptop PC and a CompactRio computer (model CompactRIO-9072, National Instruments Corp.). Both were operated fully automated by a custom Labview code, and worked reliably during the whole measurement period. The target CCD saturation level for the spectra was set to 75%. Spectra which reached the maximum CCD capacity of at least one channel of the detector because of light changes during acquisition were automatically rejected.

4 Ship cruise

The MAX-DOAS measurements over the tropical Pacific Ocean were conducted in the framework of two projects, (1) the VAMOS Ocean-Cloud-Atmosphere-Land Study – Regional Experiment (VOCALS-REx, see <http://www.eol.ucar.edu/projects/vocals/>) and (2) the Tropical Atmosphere Ocean (TAO, see <http://www.pmel.noaa.gov/tao/>). The cruise of the RV Ronald H. Brown during VOCALS-REx was split into two legs: Leg 1 (20 October–3 November 2008) started upon entering the Ecuadorian waters coming from Panama. The vessel headed south to 20° S latitude and 85° W longitude where it stopped for a buoy replacement and then cruised eastwards to Arica, Chile. Leg 2 (9 November–3 December 2008) started in Arica going west to this buoy and back, however not straight, but with some deviations to North and South. The TAO leg (9 December 2008–11 January 2009) started from Arica and headed straight to 8° S and 95° W to service a line of ocean buoys going north. At 0° latitude, the RV Ron Brown

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sailed to the 110° W buoy line continuing with buoys located from 8° S through 8° N. Lastly, the buoys between 8° N and 2° N at the 95° W line were serviced, before the vessel went straight to Panama City.

5 The DOAS analysis

5 The primary data product of the DOAS spectral analysis is the retrieval of the “differential Slant Column Density” (dSCD, difference between the actual SCD and the SCD of the Fraunhofer reference spectrum – FRS) for the trace gas of interest. The SCD is the integrated trace gas concentration along the light path. Thus the SCD depends on several parameters, namely the trace gas concentration profile, the measurement geometry, the sun’s position and the optical properties of the atmosphere, which is mainly determined by clouds and aerosol load. To account for the complexity of the light path distribution radiative transfer calculations were performed. This allows for interpretation of the measurements quantitatively by means of calculated “Air Mass Factors” (AMFs). The AMF expresses the light path ratio between the slant and the vertical path through the atmosphere (Solomon et al., 1987; Marquard et al., 2000).

15 We used the WinDOAS software package (Fayt and van Roozendael, 2001) from IASB (Belgium Institute for Space Aeronomy) to retrieve the CHOCHO dSCDs with the DOAS analysis of the measured spectra. Several trace gas cross sections as well as the logarithm of a FRS, a Ring spectrum (Grainger and Ring, 1962) and a polynomial of degree 3 were fitted to the logarithm of the acquired spectrum based on a non-linear least squares fitting algorithm. The Ring spectrum (calculated by the Software package DOASIS; Kraus, 2006) fulfills the purpose of correcting for Raman scattering and the polynomial for Rayleigh and aerosol extinction. Also included in the fitting routine was an intensity offset (linear polynomial) to account for instrumental stray light.

20 The wavelength-pixel-mapping of the analyzed spectrum was allowed to shift against that of the FRS, the Ring spectrum and the cross sections (Stutz and Platt, 1996).

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In our retrieval the FRS and Ring spectra were continuously updated, and measurement spectra were analyzed relative to the closest FRS measured prior to measurement spectrum. This procedure minimizes the influence of instrumental instabilities, and largely eliminates absorptions of trace gases located in the stratosphere. The

5 Ring spectrum was calculated from the appropriate FRS and was updated accordingly. Using WinDOAS, the wavelength-pixel-mapping was determined by fitting the FRS to a high resolution solar spectrum (Kurucz et al., 1984) convolved with the instrument's slit function. For the CHOCHO evaluation, a spectral range from 433 to 458 nm was chosen, encompassing two major CHOCHO absorption bands including the relatively
10 strong band at around 455 nm (see Fig. 1).

The following high resolution absorption cross sections were included in the retrieval: CHOCHO (Volkamer et al., 2005b), NO₂ at 294 K (Vandaele et al., 1997), ozone at 223 K (Bogumil et al., 2003) and O₄ (Hermans et al., 1999). These cross sections were convolved with the instrumental slit function to match the spectral resolution of
15 the instruments. In addition, we used a water vapor reference spectrum derived from our own atmospheric measurements on 11 November 2008, at 17:29 UT due to the poor quality of literature water vapor cross sections in the CHOCHO wavelength range (see Sinreich et al., 2007; and the references therein for details on the procedure). At this time, the ship was at 19.9° S and 75.2° W where CHOCHO was mostly below the
20 detection limit. Since the water vapor reference spectrum was derived by the instrument itself, its specific optical properties inherently lead to a better spectral retrieval with less systematic structures. A comparison of two otherwise identical CHOCHO retrievals, one using our measured water vapor cross section and one using a water vapor cross section from literature (Rothman et al., 2005) led to indistinguishable
25 results for the CHOCHO dSCDs within the spectral analysis uncertainty, but showed significantly larger residual structures when using the cross section from literature.

The O₄ dSCDs were retrieved in the wavelength range between 465 and 486 nm, from the strong absorption band at 477 nm. Cross sections of O₄ (Hermans et al., 1999), NO₂ at 294 K (Vandaele et al., 1997), H₂O (Rothman et al., 2005) and of O₃

(Bogumil et al., 2003) were fitted. Furthermore, a Ring spectrum, a polynomial of degree 2 and a linear intensity offset were included in the fit.

6 Quality assurance and data filtering

The data set underwent several filters to eliminate a possible impact of ship exhaust to the data, to ensure a specific signal-to-noise level, and to provide a subset of data with cloud-free conditions for retrieving volume mixing ratios (see Sect. 7.3).

- Ship exhaust: During the cruise and while on station to service buoys, the exhaust from the ship was likely to be blown into the line of sight of the MAX-DOAS device at times. In order to prevent the ship exhaust from contributing to the CHOCHO data set the NO₂ dSCDs (taken from the CHOCHO analysis) of 1.5°, 90° and 178.5° were compared. In the open sea, in general, no significant tropospheric NO₂ values are expected. If a significant tropospheric NO₂ signal was detected in one of the shallow viewing directions but not in the vertical or the other shallow direction, the results were excluded for the whole measurement sequence of the respective pointing direction. However, close to the coast NO₂ plumes from land could be measured. Then, a significant tropospheric NO₂ signal was found in all three directions, and no filtering was applied.
- Signal-to-noise: Only CHOCHO dSCDs above a conservative detection limit were considered (except for Fig. 2 where this filter is not applied). The dSCD detection limit was set to 2 times the root mean square of the fit residual divided by the strongest convolved (differential) cross section peak in the analysis wavelength window ($5 \times 10^{-19} \text{ cm}^2$; Volkamer et al., 2005b).
- Cloud screening: The data set was divided into subsets of data according to whether clouds were present or not. In order to find a measure for cloudless conditions, a cloud index was developed based on the concept of a color index introduced by Wagner et al. (1998) and recently applied by Takashima et

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al. (2009). The ratio $R(\text{elevation angle})$ of the spectral intensities from the edges of the spectrometer's wavelength range (420 nm and 520 nm) was calculated for each spectrum, as a measure how “blue” the sky is. A higher value means a sky with fewer and thinner clouds and a lower value consequently more and thicker clouds. It turned out that the ratios are inversely proportional to elevation angle, reflecting the effect of scattering from sea spray aerosols, that are more abundant close to the surface. However, they are not separated by elevation angle under cloudy conditions. Thus the ratio of the color index at the 3.8° elevation angle – $R(3.8)$ – and the one at the 25° elevation angle – $R(25)$ – provide a sharp threshold between cloudy and sunny conditions. Both ratios cover a significant area of the sky, and particularly almost all of the area of measurement viewing directions. For elevation angles larger than 90° , the corresponding ratio was applied: $R(176.2)/R(155)$. The calculated ratios resulted in two significant maxima at about 1 (cloudy conditions) and 1.5 (cloud-free conditions), so that the value 1.3 was chosen as a suitable threshold between cloudy and cloud-free conditions. This classification was also found to be consistent with campaign notes on cloudiness. The cloud index is plotted in Fig. 2a,b as a two-value index where the high value indicates cloudy conditions and the low value no clouds.

7 Results and discussion

7.1 Spectral retrieval of CHOCHO

Spectral proof for the detection of CHOCHO is given in Fig. 1. The spectrum shows an example CHOCHO retrieval from a 30 s spectrum that was taken on 28 December 2008, at 03:29 p.m. UT at an elevation angle of 179.4° (0.6° above the horizon). At this time the RV Ron Brown was cruising at about 1.5° N and 110.3° W at the most western part of the whole cruise, hence about 2000 km from continental land. The fitted optical

density of CHOCHO versus wavelength is shown in red, whereas the black line shows the sum of the scaled cross section and the residual. The root mean square (RMS) variation of the optical density of the residual in this evaluation is 3.7×10^{-4} and the optical density of CHOCHO is about 1.5×10^{-3} (dSCD of 2.69×10^{15} molec/cm²) which is about 4 times the residual and about 9 times the fit error (3.07×10^{14} molec/cm²). A spectral artifact caused by chlorophyll in the water is highly unlikely for our MAX-DOAS measurements since the measurements shown in this paper are all from sky spectra and only exceptionally hit the ocean (with its generally low albedo).

Figure 2a and b show one week of CHOCHO and O₄ dSCDs (in units of molec/cm² and in arbitrary units of molec²/cm⁵, respectively) retrieved from multiple elevation angles at the beginning of the TAO leg from 11–17 December 2008. Since solar stray light MAX-DOAS requires sunlight no data is collected at night, which explains the gaps in the time series. For lucidity reasons, only the elevation angles pointing to one side of the ship (i.e. 0.6°, 1.5°, 3.8°, 10°, and 25°) are shown and color coded, and we note that the dSCDs measured to the other side (i.e. 179.4° to 155°) exhibited similar values. Due to the strong swell the 0.25° (and 179.75°) elevation angle often pointed below the horizon at the sea surface, and this data is therefore omitted. Clearly, the CHOCHO dSCDs are separated by elevation angles and are higher in lower elevation angles indicating that CHOCHO is enhanced near the instrument altitude, i.e. inside the Marine Boundary Layer (MBL). Notably, the dSCD values of 0.6° and 1.5° (as well as 179.4° and 178.5°) lie mostly on top of each other, and 3.8° (and 176.2°) on 17 December behaved the same as well. Such overlapping was observed during the whole measurement period. We have demonstrated elsewhere that this overlapping is typically caused by aerosol extinction (Volkamer et al., 2009b); the collapse of dSCDs of the lowest elevation angles onto an almost single value reflects the fact that the bulk of the scattering events are located inside the boundary layer. Also, the O₄ dSCD values of the lowest elevation angles in Fig. 2b lie on top of each other whereby values of higher elevation angles are clearly separated, which confirms the existence of elevated aerosol loads. Furthermore, in both plots the two value cloud index introduced

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in Sect. 6 is shown. Figures 1 and 2a demonstrate for the first time that CHOCHO is located in the MBL over the tropical Pacific Ocean.

CHOCHO is highly water soluble and partitions readily to water droplets in clouds (Henry's law constant $H_{\text{eff}}=4.2\times 10^5$ M/atm at 298 K; Ip et al., 2009). At a typical liquid water content of 1 g/m^3 , about 90% of CHOCHO is dissolved in the cloud water and 10% is present in the gas phase (see Volkamer et al., 2009a). Indeed, the CHOCHO dSCDs of the 25° elevation angle (red dots) often showed slightly elevated values when it was cloudy (higher cloud index value). This indicates that even in clouds, the gas phase CHOCHO is still observable. This effect should also be seen for the other elevation angles. However, the relative contribution of the light paths inside clouds is smaller, and noise and other scattering effects are stronger or even dominant. Since O_4 is not water soluble and the vertical profile is well known, even in the presence of clouds, O_4 dSCDs provide a meaningful tracer of the photon path distribution. The effect of clouds is typically to decrease differences in the O_4 dSCDs of different elevation angles, most visible on 17 December. Under cloud free conditions the O_4 dSCDs of the different elevation angles are clearly separated and are more consistent.

Figure 3a shows the CHOCHO dSCDs of 1.5° and 178.5° elevation angle along the cruise track. Consistently higher values were retrieved during the TAO leg. The values typically vary between 1 and 3×10^{15} molec/cm² (individual values up to 6×10^{15} molec/cm²), and higher values tend to be observed on the way from Arica to the first buoy and during the Western and Northern part of the cruise. However, the CHOCHO dSCD values reflect the overlap of photon light paths and CHOCHO distribution. The following section describes how the O_4 dSCDs can be used to isolate light path effects to derive CHOCHO volume mixing ratios.

7.2 Conversion of CHOCHO dSCDs into volume mixing ratios

We used the radiative transfer model McArtim (Deutschmann, 2009) to convert CHOCHO dSCDs into volume mixing ratios and exploited the special feature of overlapping dSCDs at low elevation angles (up to 3.8°) to simplify this conversion, following

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principles introduced by Volkamer et al. (2009b). This approach is based on the observation that pointing below a certain angle yields dSCD values that are essentially indistinguishable within the analysis error. Consequently, the light path distributions of these lower angles must be very similar. This effect is attributed to aerosols: Assuming a well mixed MBL the aerosol load is high enough so that the differences in the path lengths which usually arise within different elevation angles are no longer observable. As shown in Fig. 2, the collapse of dSCDs is observed for CHOCHO (absorbing strongest at 455 nm) as well as for O₄ absorbing at 477 nm.

The light path can be calculated from the O₄ dSCDs and applied to the CHOCHO dSCDs. A correction factor calculated from radiative transfer modeling accounts for the difference in the absorption wavelengths and vertical profile shapes of the two trace gases. Volkamer et al. (2009b) demonstrated this for two cases in which the trace gas absorbs at the same wavelengths as O₄ (360 nm and 477 nm). Here we apply the same approach using O₄ at 477 nm for estimating the effective light path of a CHOCHO layer absorbing at 455 nm. The radiative transfer calculations give the correction factors, which, by means of the measured O₄ dSCDs, enable to convert CHOCHO dSCDs into volume mixing ratios. The correction factors were calculated for different aerosol loads and different SZAs. For the modeling, an albedo of 0.05, a single scattering albedo of 0.97 and an aerosol asymmetry parameter of 0.75 were chosen. The CHOCHO layer is assumed to be located close to sea level with a constant mixing ratio up to a certain height (box profile), as well as the aerosol load being constant (relative to the air pressure) over this same height. These heights were set to a typical value of 1 km (Sara Tucker, personal communication). Sensitivity studies varying MBL height showed little sensitivity for the low elevation angles used here, i.e. with 500 m layer height correction factors were <30% different to those with 1 km height; comparisons with 2 km layer height yielded even smaller differences (<18%). For deriving the correction factors, the aerosol extinction profile was chosen to reflect the typical behavior of the measured dSCDs (i.e. collapsing of the lowest elevation angles), and we derived mean correction factors for aerosol extinction values ranging from 0.2 to 0.6 km⁻¹ ground aerosol

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extinction. For values below 0.2 km^{-1} the values of the small elevation angles typically separate, and above 0.6 km^{-1} the 3.8° elevation angle clearly overlaps with the lower elevation angle values, which was not often observed in the measurements (see Volkamer et al., 2009b). In this interval the mean value was taken each and the standard deviation based on the variability within this interval was calculated for different solar zenith angles (SZA). The correction factor was expressed as a function of the SZA (ranging between 0.5 and 0.78), and was applied to the dSCDs of CHOCHO for 1.5° and 178.5° , using the light path calculated from the respective O_4 dSCD values (assuming a O_4 vertical column density, VCD, of $1.33 \times 10^{43} \text{ molec}^2/\text{cm}^5$ and a O_4 mean concentration of the lowest kilometer of $2.60 \times 10^{37} \text{ molec}^2/\text{cm}^6$, in arbitrary units), presuming 1 part per trillion (ppt) is $2.5 \times 10^7 \text{ molec}/\text{cm}^3$. Above 80° SZA, the correction factor increases significantly while the signal-to-noise decreases strongly, and no values above 80° SZA were chosen for the conversion into mixing ratios.

The radiative transfer modeling showed that a cloud cover would make a significant difference compared to a cloud-free scenario, on which the calculations were based. Thus, the cloud filter described in Sect. 6 was used to filter out all 1.5° and 178.5° dSCD values which were measured under cloudy conditions. In order to ensure that the assumptions of the overlapping of dSCD values are met, only 1.5° and 178.5° values which do not deviate from the respective 0.6° and 179.4° dSCD values by more than 20% were taken into account. This was valid for 86% of the values after the filters described in Sect. 6 were applied. The 14% which did not meet this criterion were not used for the mixing ratio conversion.

The overall error of the mixing ratio conversion is dominated by the assumptions of the approach especially of the mixing layer height and is at most 30%, which also includes the DOAS fit error.

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Elevated CHOCHO mixing ratios of up to 170 ppt were observed inside the MBL up to 3000 km from the continental coast. Figure 3b shows the CHOCHO mixing ratio along the cruise track. The color coding has been capped at 120 ppt to reflect most values and to emphasize their variability. During the TAO leg CHOCHO mixing ratios varied between 25 ppt and 170 ppt (see Fig. 3d), and were above the detection limit of few 10 ppt almost everywhere in the probed area. The lowest CHOCHO mixing ratios were found near the buoy at 20° S and 85° W, at the western tip of the most southern latitudes probed during the cruises. These mixing ratios correspond to a boundary layer vertical column of 0.6 to 4.0×10^{14} molec/cm², assuming a homogeneous box profile over a MBL height of 1 km. The mean value amounts to $(1.7 \pm 0.6) \times 10^{14}$ molec/cm². Also, the effective path length L_{eff} in the MBL for our MAX-DOAS instrument was calculated from the observed O₄ dSCDs ($L_{\text{eff}} = \text{dSCD} / \text{mixing ratio}$) and are plotted in Fig. 3c. The lowest L_{eff} values were observed during the Northwest part of the TAO leg, where the highest CHOCHO mixing ratios were found. Similarly, the highest values of L_{eff} were found in the Southern part of the cruise, and correspond to the lowest CHOCHO mixing ratios on the cruise. The effective path length is related to the aerosol extinction; a shorter effective path length is primarily caused by higher aerosol extinction. Indeed, we observe an anti-correlation of the CHOCHO mixing ratio and the effective path length, as shown in Fig. 3d. CHOCHO uptake to aerosols is most efficient on aqueous aerosols (Liggio et al., 2005; Volkamer et al., 2009a), such as those of the tropical Pacific Ocean, and aerosol loss can affect the atmospheric life time of CHOCHO (Volkamer et al., 2007). The anti-correlation in Fig. 3d does not show a 1:1 correspondence, but rather reflects a complex coupling of CHOCHO and aerosols. Heald et al. (2006) stated that the GEOS-Chem global model underpredicts the organic aerosol concentrations observed in the MBL for different regions of the world, and invoked a missing organic aerosol source of 27 Tg C/yr from the oceans that was correlated with photochemical activity. CHOCHO uptake to aerosols could potentially explain the observed

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organic aerosol concentrations in the MBL (Fu et al., 2008). However, Fig. 3d could equally be interpreted in terms of CHOCHO related aerosol sources (SOA formation), or be indicative of similar mechanisms forming aerosols and CHOCHO.

CHOCHO is very short lived, and any elevated CHOCHO indicates local sources.

The atmospheric life time of CHOCHO is estimated to be about 2 h as a typical daytime value, and shorter at solar noon. It is determined primarily by the abundance of solar radiation around the Equator, and to a lesser extent OH radical initiated CHOCHO loss. The 24 h average OH concentration during VOCALS-REx was estimated to be 1.4×10^6 molec/cm³ (Yang et al., 2009), corresponding to a peak daytime OH of about 5×10^6 molec/cm³. This is in reasonable agreement with observations over the tropical Atlantic Ocean (on average several 10^6 molec/cm³, though midday peaks reached almost 10^7 molec/cm³; Whalley et al., 2010). We note that a primary OH radical source is from ozone photolysis, and ozone concentrations can be very low over the tropical Pacific Ocean (Kley et al., 1996). Thus, extrapolating measurements from the Atlantic to the Pacific Ocean may not be straightforward. For the further discussion we assume $[\text{OH}] = 3 \times 10^6$ molec/cm³ as reasonable during daytime. Further, dry deposition of CHOCHO also can play an important role in a shallow MBL, since CHOCHO is highly water soluble ($H_{\text{eff}} = 4.2 \times 10^5$ M/atm; Ip et al., 2009). With a typical deposition velocity of 3–4 cm/s, the CHOCHO loss to the ocean surface can add 20 to 30% to the daytime loss rate from photolysis and OH reaction; dry deposition is relatively more important at night. Also, CHOCHO uptake by aerosols to form SOA can compete with rapid gas-phase losses and will further lower the atmospheric lifetime of CHOCHO (Volkamer et al., 2007; Fu et al., 2008; Stavrou et al., 2009; Ervens and Volkamer, 2010). The CHOCHO life time in the tropical Pacific Ocean is thus shorter than the global mean life time of CHOCHO (2.5 to 3 h; Myriokefalitakis et al., 2008; Fu et al., 2008; Stavrou et al., 2009), and about comparable to that in urban photochemical hot-spots (Volkamer et al., 2005a, 2007). Given the short life time of 1–2 h during the day, we can rule out transport and oxidation of precursors emitted from terrestrial sources as CHOCHO source. Typical precursors emitted over land are compounds like isoprene, toluene,

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and acetylene, which at an OH concentration of 3×10^6 molec/cm³ have atmospheric life times of about 1 h, 15 h, and 5 days, respectively. Atmospheric models that represent land sources show enhanced CHOCHO only within few 100 km from the coast, but do not predict any CHOCHO over the remote tropical Pacific Ocean (Myriokefalitakis et al., 2008; Fu et al., 2008). This is consistent with the fact that typical transport times over a 3000 km distance over the Pacific Ocean are much longer than the atmospheric life times of terrestrial precursors. An air mass coming directly from the continent would take about 1 week to reach the ship (at a wind speed of 5 m/s, about 10 knots). Figure 4 shows 48 h HYSPLIT (HYbrid Single-Particle Lagrangian Integrated Trajectory) backward trajectory calculations for corner points of the cruise (Draxler and Rolph, 2010; Rolph, 2010). The trajectories are calculated for days when the ship was passing through these points. They are shown for the morning (green), midday (blue) and evening (red), and for three different air mass heights (100 m, 500 m and 1000 m). The prevailing wind direction was southeast rather than coming directly from the continent, which is in good agreement with the ship's meteorology data. This prolongs the transport times from land sources considerably. Since a significant CHOCHO source due to continental outflow can be excluded, the observed elevated CHOCHO must originate from the ocean.

Kieber et al. (1990) showed that compounds like formaldehyde, acetaldehyde or glyoxylate are produced from the photochemical oxidation of Dissolved Organic Matter (DOM) in natural waters when exposed to sunlight. Millet et al. (2010) recently parameterized data from Kieber et al. (1990) and showed that DOM photochemistry in the ocean is the second largest global source for acetaldehyde (57 Tg/year). Relying on the satellite retrievals of CHOCHO, the spatial distribution of the acetaldehyde source indeed appears to somewhat resemble that of CHOCHO at tropical latitudes (Wittrock et al., 2006; Vrekoussis et al., 2009). However, we argue that this spatial correspondence is at least partly coincidental, and should not be taken as evidence that the CHOCHO source from the ocean follows the same oxidation mechanism as had been identified for acetaldehyde by Kieber et al. (1990). This is because CHOCHO is

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much more soluble than acetaldehyde (30 000 times higher Henry's law constant). In order to explain 100 ppt of boundary layer CHOCHO in equilibrium with ocean surface waters by means of a bulk ocean water photo oxidation mechanism, the ocean would need to be a 4×10^{-5} M concentrated solution of CHOCHO. For comparison, the ocean concentration of acetaldehyde needed to explain atmospheric observations in Millet et al. (2010) (typically 20–200 ppt) is 1.3–20 nM. In fact, for CHOCHO the dissolved concentration would need to be even higher, since the atmospheric concentration of CHOCHO is not necessarily in steady state with the ocean due to rapid gas-phase losses and atmospheric dynamics. Typical concentrations of DOM in the tropical Eastern Pacific Ocean are $6-7 \times 10^{-5}$ M (Hansell et al., 2009), and a good fraction of this DOM has an essentially different molecular identity as CHOCHO. Typical CHOCHO concentrations in the air over open Caribbean Sea and Sargasso Sea were on the order of 80 ppt, and the corresponding seawater CHOCHO concentration was measured to be 0.5–5 nM (Zhou and Mopper, 1990). CHOCHO concentrations measured 100 km east of the Bahamas were 0.3 nM in the bulk ocean, and about 6 nM in the surface microlayer (Zhou and Mopper, 1997). Notably, CHOCHO is enriched by a factor of 4–20 in the surface microlayer compared to the bulk ocean, reflecting a higher production rate from more abundant precursors (Zhou and Mopper, 1997), and/or radical sources (Reeser et al., 2009). However, Zhou and Mopper (1990, 1997) measured 4 orders of magnitude less CHOCHO in the surface microlayer than is compatible with equilibrium air-sea partitioning and atmospheric CHOCHO concentrations. Based on these measurements the net air-sea flux of CHOCHO is directed from the atmosphere into the ocean. This implies that the CHOCHO source over the Caribbean and Sargasso Sea was airborne. We find comparable CHOCHO concentrations over the tropical remote Pacific Ocean, where continental transport can be ruled out. If the above ocean measurements extrapolate to the tropical Pacific Ocean, it appears that a bulk water photochemical process can only explain a very minor portion ($\sim 10^{-4}$ – 10^{-5}) of the CHOCHO source over the remote ocean.

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A simplified calculation using the SCIAMACHY CHOCHO VCDs, and the TM4 global model has quantified that to explain CHOCHO over the oceans requires a global source of 20 Tg/year (Myriokefalitakis et al., 2008). This source could be related to oceanic emission of VOC precursors that form CHOCHO as the result of their oxidation in the atmosphere. The most important natural precursors are isoprene (CHOCHO yield is 8%; Bey et al., 2001), toluene (30% yield; Volkamer et al., 2005b), and acetylene (65% yield; Hatakayama et al., 1986). Assuming an OH concentration of 3×10^6 molec/cm³, and an atmospheric life time of CHOCHO of 2 h (conservative upper limit), the VOC concentration needed to explain 100 ppt of CHOCHO is either 600 ppt of isoprene, 2.8 ppb toluene, or 8.9 ppb acetylene, or lower concentrations if multiple compounds were present. There are few observations of VOCs over the remote open ocean, and most of them focus on isoprene (Bonsang et al., 1992; Broadgate et al., 1997, 2004; Baker et al., 2000; Matsunaga et al., 2002; Meskhidze and Nenes, 2006; Yassaa et al., 2008). The currently measured isoprene concentrations do not exceed 40 ppt and mostly are on the order of 10 ppt. All observations were made at higher latitudes than our probed area, but global estimates of isoprene concentration maps confirm these values for the probed area (Palmer and Shaw, 2005; Arnold et al., 2009; Gantt et al., 2009). Plass-Dülmer et al. (1995) derived a global upper limit for acetylene from ethane emissions over seawater of 0.16–0.54 Tg/year, which on average corresponds to less than 20 ppt acetylene. In lack of any measurements of toluene over the open sea, it seems unlikely to explain the observed CHOCHO production from secondary VOC chemistry over the open ocean. Thus neither continental outflow of VOC precursors from land sources, nor VOC emissions from the open ocean, nor a direct CHOCHO source from DOM photochemistry in the bulk ocean or surface microlayer are likely explanations for our observations. The high solubility of CHOCHO requires the source to be airborne, either in form of a process at the air-sea interface, or by a bubble bursting process. Further research is needed to determine the formation mechanism of CHOCHO over the remote tropical ocean.

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The MAX-DOAS measurements onboard the research vessel Ronald H. Brown showed elevated CHOCHO differential Slant Column Densities (dSCDs) over the remote tropical Pacific Ocean, up to 3000 km from continental land sources. Ours are the first direct CHOCHO measurements in the remote marine boundary layer (MBL). Previous satellite measurements had also retrieved CHOCHO over the remote ocean, though different satellites give inconsistent values. The TM4 and GEOS-Chem models do not predict any CHOCHO over the open ocean (Myriokefalitakis et al., 2008; Fu et al., 2008), and currently cannot explain these observations.

By means of O_4 dSCDs and radiative transfer modelling the CHOCHO dSCDs of the 1.5° and 178.5° elevation angles were converted into mixing ratios, and the effective light paths in the MBL were estimated. Mixing ratios were typically 120 ± 36 ppt, and higher in the north-western part of the probed area (closer to the Equator); lower values were observed further south (typically around 50–60 ppt). An anti-correlation is observed between the CHOCHO mixing ratios, and the effective path length (an indicator for aerosol extinction). This is compatible with aerosol uptake of CHOCHO to form secondary organic aerosol (SOA), or could indicate similar mechanisms forming aerosols and CHOCHO.

The source of CHOCHO located up to 3000 km from continental land is unclear. Transport of CHOCHO or its precursors from terrestrial sources can be ruled out, due to the long travel time (>1 week) of the air mass from the continental coast to reach the ship. Also, the production of CHOCHO by photo degradation of marine dissolved organic matter (DOM) in the ocean is unlikely to be an efficient source for atmospheric CHOCHO due to the high solubility of CHOCHO. However, an indirect link with DOM cannot be excluded. Lastly, the production of CHOCHO from VOC emissions would require VOC precursor amounts that are much higher than have been observed to date in the open ocean. The tropical Eastern Pacific Ocean is a particularly productive upwelling region and is a poorly probed atmospheric environment. More research is

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needed to understand the formation mechanism, and to quantify the impact of the reactive gas release from the open ocean on atmospheric chemistry, and possibly climate.

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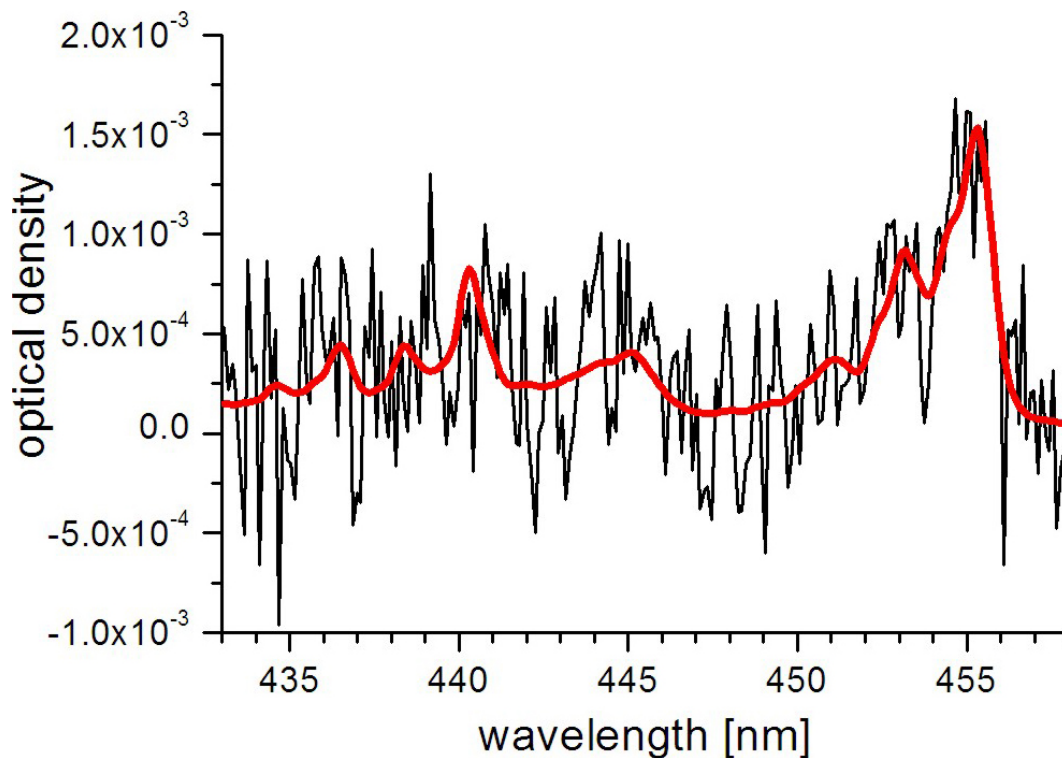


Fig. 1. Spectral retrieval for elevated CHOCHO (dSCD of 2.69×10^{15} molec/cm² and RMS of 0.37×10^{-3}). The example spectrum was recorded on 28 December 2008 at 03:29 p.m. UT at an elevation angle of 179.4°. The optical density of CHOCHO versus wavelength is shown in red, whereas the black line shows the sum of the scaled cross section and the residual.

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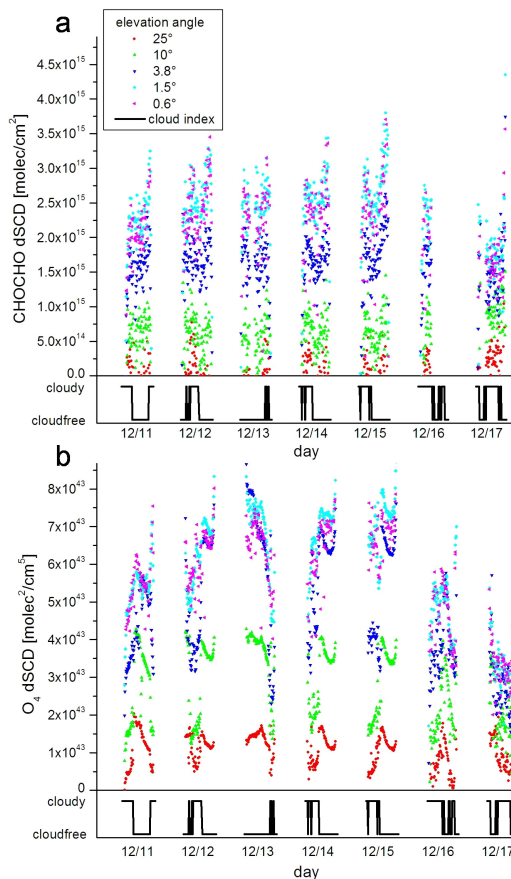


Fig. 2. (a) One week of the retrieved CHOCHO dSCDs from 11–17 December 2008. The elevation angles (color coded) are well separated except for the lowest elevation angles and on 17 December. A two level cloud index is added whereby the high value means cloudy conditions and the low value no clouds. (b) The same as (a) for O_4 dSCDs.

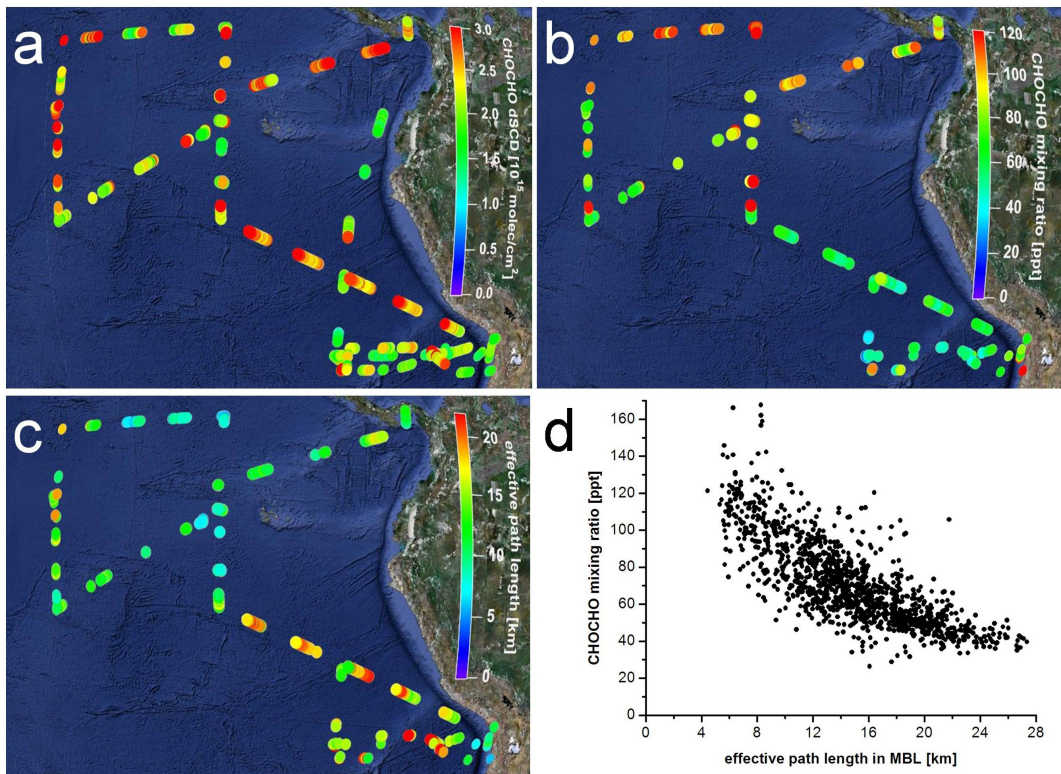


Fig. 3. Cruise plots (westerly of South America) of **(a)** CHOCHO dSCDs at 1.5° and 178.5° elevation angles, **(b)** CHOCHO volume mixing ratios (given in ppt) derived as described in Sect. 7.3, and **(c)** effective path length in the MBL (satellite underlay pictures: Google Earth). **(d)** Correlation plot of CHOCHO mixing ratio versus the effective path length, which shows an anti-correlation.

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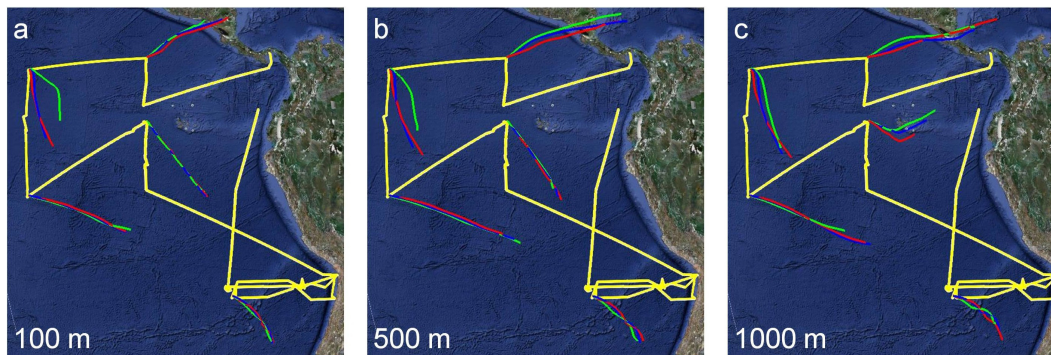


Fig. 4. 48 h HYSPLIT (HYbrid Single-Particle Lagrangian Integrated Trajectory) backward trajectory calculations for corner points of the cruise (yellow line) at the days the ship was there (Draxler and Rolph, 2010; Rolph, 2010). The trajectories are calculated for the morning (green), midday (blue) and evening (red) each for air mass heights at the starting points of **(a)** 100 m, **(b)** 500 m and **(c)** 1000 m.

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