



Supplement of

Microphysical evolution and column loading drive nonlinear regional contrast in black carbon top-of-atmosphere forcing

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39 **Section S1**

40 **Theoretical Concept of Santa Barbara DISORT Atmospheric Radiative Transfer Model (SBDART)**

41 In this study, the Santa Barbara DISORT Atmospheric Radiative Transfer (SBDART) model (Ricchiazzi et
42 al., 1998) (available at <https://github.com/paulricchiazzi/SBDART>) is utilized to calculate atmospheric
43 irradiance both including and excluding aerosol influences. Given that most solar radiation is concentrated
44 within the 0.25 to 4.0 μm wavelength range (Paulescu et al., 2013), the analysis is confined to this spectral
45 interval. To perform these irradiance computations, SBDART requires inputs such as columnar aerosol
46 optical depth (AOD), Single scattering albedo (SSA), and Angstrom exponent (ASY) as, per waveband
47 optical properties. In our case we have derived BCAOD and corresponding SSA and ASY at different
48 wavebands using a set of observational data and their uncertainties as explained in the main text.
49 Additionally, the column water vapor, surface spectral albedo, and surface ozone data obtained from MODIS
50 and TROPOMI satellites respectively are used to compute their respective local impacts on the radiative
51 stream. Monthly averages of model atmosphere were prepared using temperature, pressure, and ozone
52 vertical profiles sourced from ERA-5 reanalysis datasets
53 (<https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-pressure-levels?tab=form>), dynamically
54 updating the model atmosphere for each month.

55 The aerosol DRF is determined by evaluating the difference in net radiative fluxes, where the net flux (ΔF)
56 is defined as the difference between the downward irradiance (F_{\downarrow}) and upward irradiance (F_{\uparrow}):

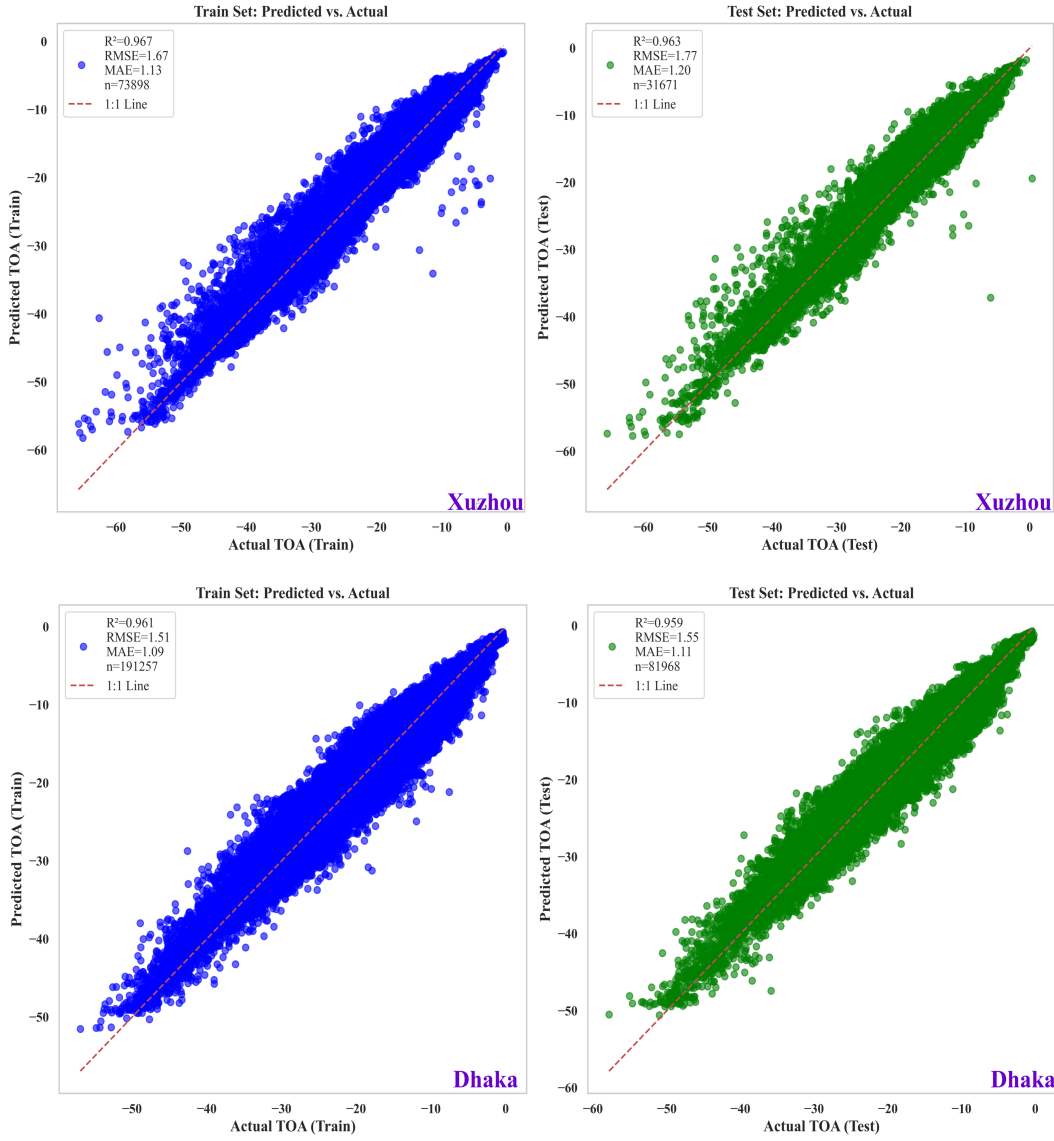
57
$$\Delta F = F_{\downarrow} - F_{\uparrow} \quad \text{Eq. 1}$$

58 The net flux is calculated separately at the top of the atmosphere (TOA) and the Earth's surface (SFC). To
59 isolate the effect of aerosols, the net flux with aerosols present (F_{aero}) is compared against the net flux without
60 aerosols ($F_{\text{non-aero}}$). The forcing at TOA and SFC are computed by subtracting the non aerosol net flux from
61 the aerosol net flux:

62
$$\Delta F_{\text{TOA}} = \Delta F_{\text{aero,TOA}} - \Delta F_{\text{non-aero,TOA}} \quad \text{Eq. 2}$$

63
64 In this study, SBDART used its default near-surface-dominant aerosol vertical distribution, with aerosol
65 loading concentrated within the lower atmosphere/boundary layer and decreasing with altitude. SBDART is
66 a widely used model to analyze the aerosol DRF (Fu et al., 2017; Patel et al., 2017).

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Figure S1: Comparison of predicted TOA values from the ML model against actual TOA values for both training and test datasets. The top row represents results for Xuzhou, while the bottom row represents Dhaka. The left panels show predictions on the training set, while the right panels show predictions on the test set. Each scatter plot includes key performance metrics: the coefficient of determination (R^2), Root Mean Square Error (RMSE), and Mean Absolute Error (MAE). The red dashed line represents the 1:1 reference line, indicating perfect agreement between predicted and actual TOA values. Model hyperparameters were set as follows: `n_estimators=100`, `max_depth=10`, `min_samples_split=5`, `min_samples_leaf=3`, `max_features='sqrt'` for Random Forest regressors.

81 **Table S1: Train and Test Performance from 5-Fold Cross-Validation of the developed ML TOA**
 82 **Models in Xuzhou and Dhaka**

Folds	Xuzhou (Train)			Dhaka (Train)		
	R ²	RMSE	MAE	R ²	RMSE	MAE
1	0.96	1.7	1.2	0.96	1.5	1.1
2	0.96	1.8	1.2	0.96	1.5	1.1
3	0.96	1.7	1.2	0.96	1.6	1.1
4	0.96	1.8	1.2	0.96	1.6	1.1
5	0.96	1.8	1.2	0.96	1.5	1.1
	Xuzhou (Test)			Dhaka (Test)		
1	0.96	1.8	1.2	0.96	1.6	1.1
2	0.96	1.8	1.2	0.96	1.6	1.1
3	0.96	1.9	1.2	0.96	1.6	1.1
4	0.96	1.7	1.2	0.96	1.5	1.1
5	0.96	1.8	1.2	0.96	1.6	1.1

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84 **Section S2**

85 **AERONET-based consistency assessment of COSMO-RTM BC TOA forcing efficiency**

86 To evaluate whether the COSMO-RTM BC TOA forcing fields used as the machine-learning target are
 87 physically consistent with independent aerosol radiative constraints, COSMO-RTM BC TOA forcing
 88 efficiency was compared with AERONET Version 3 Level 2.0 total-aerosol forcing efficiency over Dhaka.
 89 Because AERONET provides total-aerosol radiative forcing rather than BC-isolated forcing, this analysis is
 90 not intended as a strict one-to-one validation of BC TOA DRF. Instead, it evaluates whether the COSMO-
 91 RTM BC TOA forcing efficiency shows physically consistent regime-dependent, and proportional behavior
 92 relative to AERONET total-aerosol TOA forcing efficiency under absorbing aerosol conditions.

93 For Dhaka, AERONET Version 3 Level 2.0 inversion retrievals of TOA DRF were matched with COSMO-
 94 RTM BC TOA DRF estimates by retaining model pixels within a 20 km radius of the AERONET station.
 95 Model grid values were aggregated to daily means to match the temporal resolution of the AERONET
 96 inversion product. AERONET AAOD at 550 nm was derived from measured absorption AAOD (440-870)
 97 and AAE, and the FE was calculated as:

98
$$FE_{\text{AERONET}} = \frac{\text{TOA DRF}_{\text{AERONET}}}{\text{AAOD}_{550}}$$

99 representing total aerosol forcing efficiency from AERONET. For the COSMO-RTM side, BC FE was
 100 calculated as:

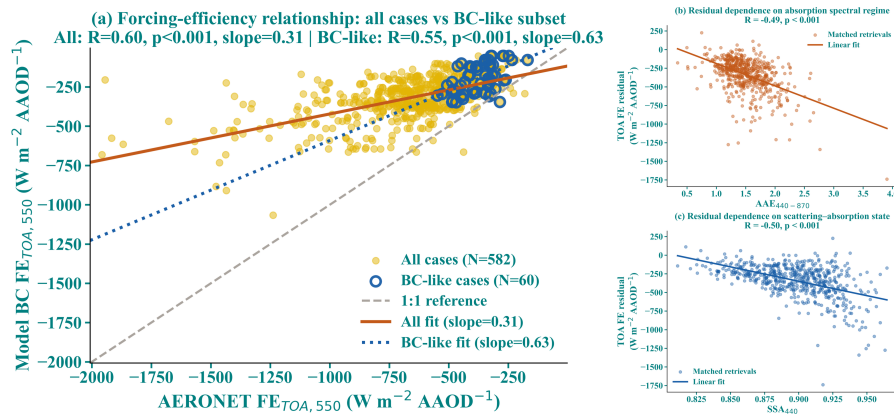
102
$$FE_{\text{BC}} = \frac{\text{TOA DRF}_{\text{BC}}}{\text{BCAAOD}_{550}}$$

101 using COSMO-SBDART-derived BC TOA forcing and the corresponding BC absorption optical depth.

103 The study used AAOD rather than AOD because the purpose of this comparison is to evaluate absorbing-
 104 aerosol radiative response. For BC-specific forcing, absorption optical depth is the relevant optical driver,
 105 whereas normalization by total AOD would mix BC absorption with scattering aerosol and other non-BC
 106 absorbing components.

107 **Regime-dependent strengthening of the COSMO-RTM and AERONET TOA FE relationship**

108 The study examined whether the COSMO-RTM BC TOA FE and AERONET TOA FE relationship changes
 109 systematically with aerosol absorption regime (Figure S2a). Across all matched retrievals, the ordinary
 110 regression slope between FE_{modelBC} and FE_{AERONET} was 0.31, with a significant correlation ($R = 0.6$; $p < 0.001$).
 111 When the analysis was restricted to the more absorbing, spectrally BC-like subset, defined by cases where
 112 SSA_{440} was below the 30th percentile and $0.8 < AAE_{440-870} < 1.4$ (Liu et al., 2018; Luo et al., 2023), the slope
 113 increased to 0.63 while the relationship remained significant ($R = 0.55$; $p < 0.001$). This increase indicates
 114 that the proportional scaling between COSMO-RTM BC FE and AERONET total-aerosol FE becomes closer
 115 under more BC-like absorption conditions.



116 **Figure S2: Regime-dependent consistency between COSMO-RTM BC TOA FE and AERONET total-**
 117 **aerosol TOA FE. (a) Scatterplot of COSMO-RTM BC-only TOA FE against AERONET total-aerosol**
 118 **TOA FE for all matched retrievals and for the spectrally BC-like subset. (b) Residual dependence on**
 119 **AAE440-870. (c) Residual dependence on SSA440.**

121 The residual analysis (Figure S2(b and c) further supports this regime dependence. The forcing-efficiency
 122 residual, is not randomly distributed with respect to aerosol optical state. Instead, it shows significant
 123 dependence on both $AAE_{440-870}$ and SSA_{440} . Residuals decrease systematically with increasing AAE ($R = -$
 124 0.49 ; $p < 0.001$), indicating that the separation between AERONET total-aerosol FE and model BC-only FE
 125 changes as the absorption spectrum becomes less BC-like (Figure S2b). A similar but slightly stronger
 126 dependence is observed with SSA_{440} ($R = -0.50$; $p < 0.001$), showing that the residual structure is also
 127 controlled by the column scattering-absorption balance (Figure S2c).

128 The regime-dependent analysis shows that the COSMO-RTM BC TOA FE and AERONET TOA FE
 129 relationship strengthens under low SSA and spectrally BC like conditions and the residuals are systematically
 130 structured by AAE and SSA. These results suggest that the COSMO-RTM BC and AERONET differences
 131 are systematically organized by aerosol optical regime rather than being structureless residual scatter.

132 **Relative-scaling diagnostic**

133 While the regime-dependent analysis establishes that the COSMO-RTM BC TOA and AERONET TOA FE
134 relationship strengthens under BC-like optical conditions (more absorbing conditions), it does not by itself
135 determine whether the magnitude of COSMO-RTM BC-only FE is reasonable relative to AERONET total-
136 aerosol FE. To address this, a relative-scaling diagnostic was used to define the expected BC-to-total aerosol
137 benchmark.

138 The theoretically expected TOA FE scaling was derived from the the ratio of two separately computed
139 fractions, independent of the fitted FE regression. The BC absorption fraction quantifies BC's contribution
140 to total column absorption:

141
$$R_{\text{abs}} = \frac{\text{BC AAOD}_{550}}{\text{AAOD}_{550}^{\text{AERONET}}}$$

142 The BC forcing fraction quantifies the magnitude of BC TOA forcing from COSMO-RTM relative to the
143 AERONET total-aerosol TOA forcing:

144
$$R_{\text{DRF}} = \frac{|\text{DRF}_{\text{TOA,BC}}|}{|\text{DRF}_{\text{TOA,AERONET}}|}$$

145 If the COSMO-RTM BC microphysics and radiative-transfer calculation are physically consistent with the
146 independent AERONET total-aerosol forcing constraint, then the observed scaling between FE_{BC} and
147 $\text{FE}_{\text{AERONET}}$ should be consistent with:

148
$$\text{Expected TOA FE scaling} = \frac{R_{\text{DRF}}}{R_{\text{abs}}}$$

149 This expectation arises because the model BC FE represents radiative effect per unit BC absorption, while
150 AERONET total FE represents radiative effect per unit total aerosol absorption. Their ratio therefore scales
151 with the relative contribution of BC to both absorption and TOA forcing. Importantly, this expected scaling
152 is not fitted from the FE regression itself; rather, it provides a physically constrained benchmark for
153 evaluating whether the BC-only FE magnitude is reasonable relative to the AERONET total-aerosol FE. It is
154 therefore not a direct validation of isolated BC forcing, but a relative-scaling consistency test designed for
155 the case where the observational product represents total aerosol forcing and the model isolates the BC
156 component. Since this benchmark is a proportional ratio, it was compared with the origin-constrained
157 COSMO-RTM TOA FE - AERONET total aerosol TOA FE slope, $\text{FE}_{\text{BC}} = m\text{FE}_{\text{AERONET}}$, rather than with
158 the ordinary OLS slope that includes an intercept. The ordinary OLS slope was used above only to diagnose
159 regime-dependent strengthening of the relationship.

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167 **Table S2. Relative-scaling diagnostic between COSMO-RTM BC TOA FE and AERONET total-**
 168 **aerosol TOA FE.**

Regime	R_{abs} [95% CI]	R_{DRF} [95% CI]	Expected TOA FE scaling [95% CI]	Observed origin- constrained slope [95% CI]	Median FE_{BC}/FE_{AERONET}
All cases	0.66[0.62- 0.69]	0.32[0.29- 0.35]	0.48[0.44-0.53]	0.46[0.43- 0.48]	0.48
BC-like subset (SSA₄₄₀ < 30th percentile and AAE 0.8- 1.4)	0.51[0.45- 0.57]	0.25[0.17- 0.32]	0.50[0.33-0.62]	0.53[0.46- 0.61]	0.55

169 Table S2 shows that the observed proportional COSMO-RTM BC TOA FE - AERONET total aerosol TOA
 170 FE scaling is consistent with the expected BC-to-total aerosol benchmark. For all matched cases, the observed
 171 origin-constrained slope was 0.46 [0.43-0.48], closely overlapping the expected scaling of 0.48 [0.44-0.53].
 172 For the BC-like subset, the observed slope increased to 0.53 [0.46-0.61], while the expected scaling was 0.50
 173 [0.33-0.62]. The median FE_{BC}/FE_{AERONET} ratio also increased from 0.48 to 0.55 under more absorbing BC-
 174 like conditions. Thus, the magnitude of COSMO-RTM BC TOA FE while is not expected to match
 175 AERONET total-aerosol TOA FE one-to-one, but it falls within the physically interpretable BC-to-total
 176 aerosol scaling range. The decrease from 0.66 to 0.51 should not be interpreted as a weakening of the FE
 177 consistency result, because R_{abs} and R_{DRF} are both recomputed for a different optically filtered subset. The
 178 relevant diagnostic is not whether R_{abs} increases or decreases, but whether the ratio R_{DRF}/R_{abs} agrees with the
 179 observed origin-constrained FE slope. This agreement is maintained for both all cases and the more absorbing
 180 BC-like subset.

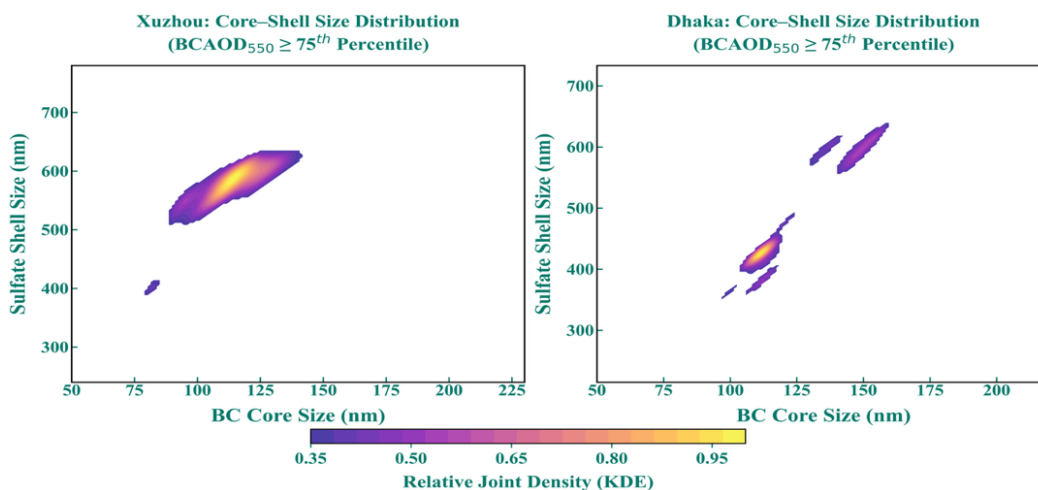
181 Overall, this analysis supports the physical coherence of the observation-constrained COSMO-RTM BC
 182 forcing target. Although it does not constitute direct validation of isolated BC DRF, the COSMO-RTM BC
 183 TOA FE shows regime-dependent and proportional consistency with independent AERONET total-aerosol
 184 radiative forcing constraints.

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191 **Table S3: Categorizing the Predictors based on their percentiles for conditional analysis**

Percentile Conditions	BCAOD ₅₅₀	BC Size	Mixing-state
<25 th	Clean	small	Heavily coated
>=25 th & <50 th	fairly polluted	medium	Moderately coated
>=50 th & <75 th	moderately polluted	Big	Minimally coated
>=75 th	polluted	Very Big	Barely coated

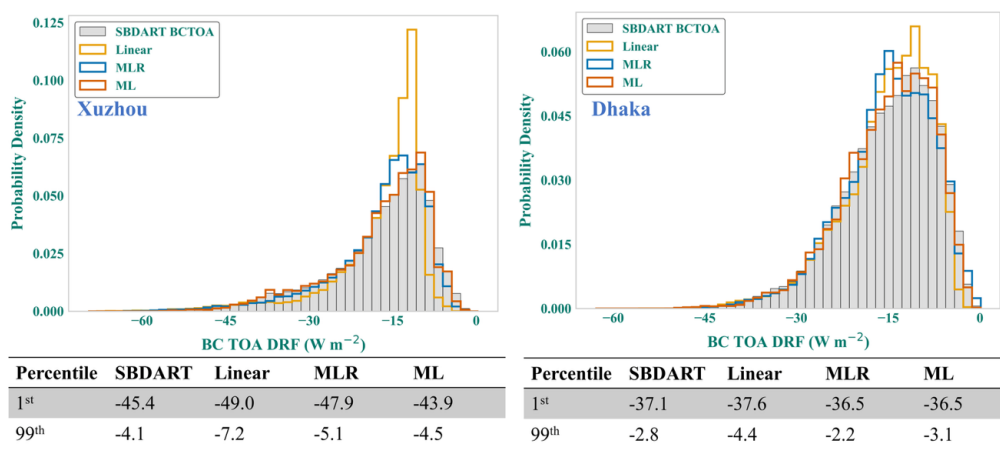
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194 **Figure S3: Dominant BC core-sulfate shell size distributions under polluted conditions (BCAOD₅₅₀**
 195 **>75th percentile in Xuzhou and Dhaka. Color shading shows the relative joint KDE density normalized**
 196 **by the regional maximum, indicating the most frequent core-shell microphysical states during high-**
 197 **loading conditions.**

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200 **Figure S4: Probability density functions (PDFs) of BC TOA values from SBDART and the three**
 201 **surrogate models (Linear, MLR, and ML) for Xuzhou (left) and Dhaka (right). Tabulated percentile**

202 values at the 1st and 99th percentiles are provided below each panel to compare the distributional
203 extremes.

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205 **Section S3.**

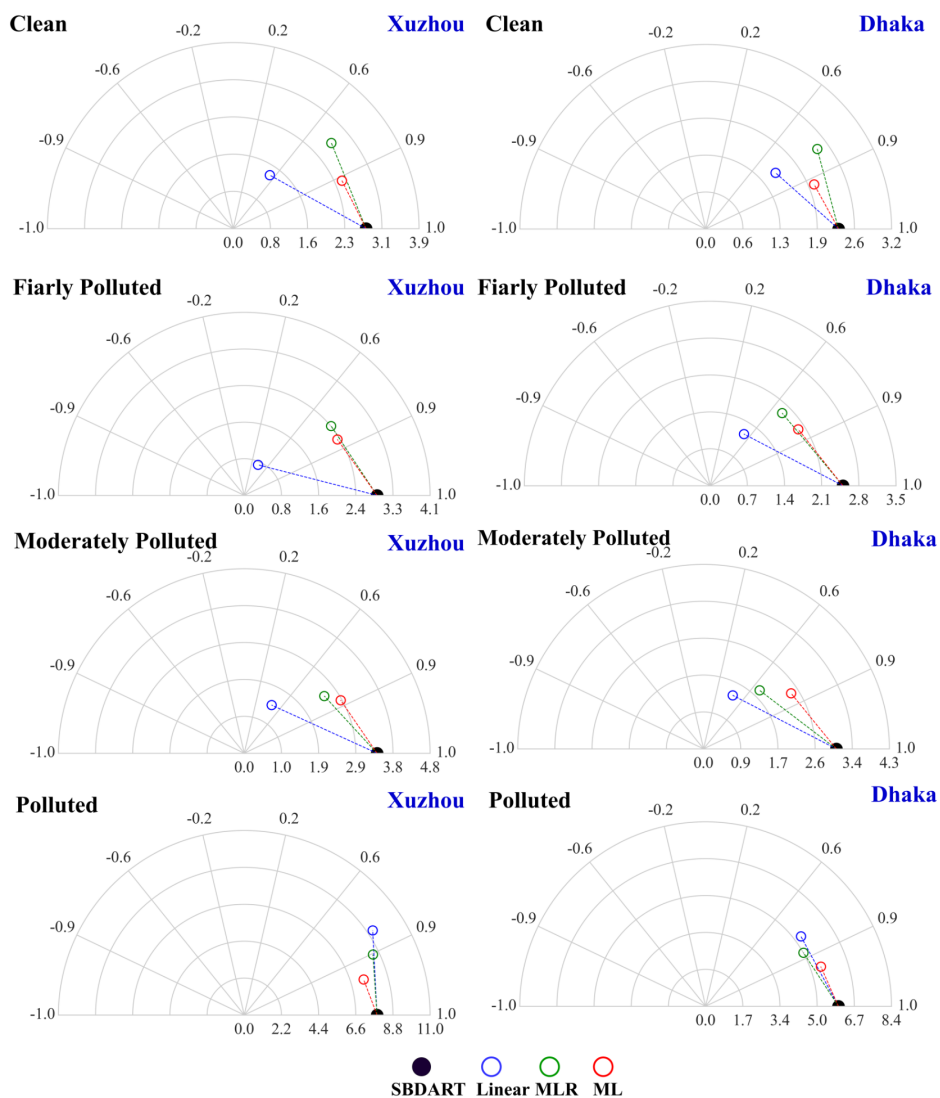
206 **Model performance comparison across different Pollution conditions.**

207 The performance of the model was also analyzed under different pollution conditions, and has been illustrated
208 using the Taylor Diagram (Figure S5) over Xuzhou and Dhaka. All the models demonstrate variable efficacy
209 across pollution conditions in Xuzhou and Dhaka, with ML consistently outperforming linear and MLR
210 approaches.

211 In Xuzhou, ML achieves superior correlations (0.85 to 0.96) (Figure S5), lower RMSE (1.1 to 2.3 Wm⁻²),
212 and enhanced variance capture especially in moderate to high pollution condition, compared to MLR and
213 linear models. Its near-zero mean bias (MBE: -0.07 to 0.05 Wm⁻²) across different pollution conditions and
214 tighter 95% confidence demonstrates minimal systematic error and high precision. Under clean conditions,
215 ML achieves a lower RMSE (1.1 Wm⁻²), reducing errors by 48% and 59% compared to MLR and linear
216 model. Similarly, over Dhaka, ML further excels, with bias corrected RMSE (0.9 Wm⁻²) approximately 36%
217 lower than MLR (1.4 Wm⁻²) and 40% lower than linear (1.5 Wm⁻²), underscoring ML's reliability in clean
218 scenarios over both regions. This improvement is noteworthy because AOD and AAOD retrievals/inversions
219 from multiple platforms, typically exhibit higher accuracy in moderate and polluted conditions, while clean
220 conditions pose challenges due to weaker aerosol signals (Dubovik et al., 2000, 2006; Dubovik and King,
221 2000). The challenges in retrieving AOD and AAOD directly affect the quantification of direct radiative
222 forcing (DRF), particularly for black carbon (BC). Our model benefits from a more size- and mixing state-
223 flexible aerosol representation that captures aerosol physical and optical properties both more precisely and
224 in a more physically accurate manner. This in turn results in improved radiative transfer model (RTM)
225 simulations of BC DRF, which the ML model then reproduces with high fidelity. Under fairly and moderately
226 polluted conditions of Xuzhou, the ML model outperforms competing methods, achieving a bias-corrected
227 RMSE approximately 18% lower than MLR and 44% lower than the linear model. The ML model's mean
228 bias error (MBE) is near zero; 0.01 Wm⁻² for fairly polluted and 0.02 Wm⁻² for moderately polluted
229 conditions-indicating minimal bias. This contrasts with MLR, which shows a slight underprediction reflected
230 by negative MBE (-0.2 Wm⁻²) in fairly polluted conditions and a positive bias (0.5 Wm⁻²) in moderately
231 polluted conditions. Similarly, the linear model exhibits larger biases of -0.5 Wm⁻² and 0.9 Wm⁻²
232 respectively for these pollution levels. The ML model's low bias paired with lower error metrics highlights
233 its improved accuracy and reliability across these conditions.

234 Similar performance was observed in Dhaka, where under fairly polluted conditions, ML model achieves a
235 bias-corrected RMSE reduction of approximately 22% compared to MLR decreasing from 1.8 to 1.4 Wm⁻²,
236 and even more pronounced reduction of 33% compared to linear model. Similarly, under moderate pollution
237 condition, ML reduces bias corrected RMSE by 23% over MLR (1.7 vs. 2.2 Wm⁻²) and by approximately
238 37% compared to linear model. MLR exhibits slightly enhanced prediction of the spread in Dhaka compared

239 to linear, however it fails to align its predictions with the SBDART trend or relationships in the observations.
 240 This indicates, while MLR might predict larger deviations from the mean at the right time (matching the
 241 variability), these deviations may be occurring in the wrong direction or phase relative to observations, and
 242 likely are driven by different underlying physics than the cleaner cases observed in Xuzhou, possibly
 243 necessitating improvements to the underlying physical model to account for this non-linearity.
 244 Under polluted conditions in Xuzhou, ML's bias corrected RMSE is 36% and 55% lower than MLR and
 245 linear models respectively. While over Dhaka, the variance captured by ML (SBDART: 5.9 vs. ML: 5.5) also
 246 high correlation (0.94) and bias corrected RMSE reduced by more than 30%. Overall, ML offered superior
 247 reproduction ability of the underlying physical model in all conditions, with the performance gap widening
 248 as pollution magnitude strengthens.
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 251 **Figure S5: Taylor diagram comparing the performance of Linear, MLR, and ML models in predicting**
 252 **TOA radiative forcing for Xuzhou (left) and Dhaka (right) under different pollution conditions (Clean,**

253 **Fairly Polluted, Moderately Polluted, and Polluted). The diagrams illustrate the standard deviation**
254 **(radial) and correlation coefficient (angular) of each model relative to SBDART (black dot). The**
255 **proximity of model points to the reference (SBDART) indicates prediction accuracy, with colors**
256 **representing different models.**

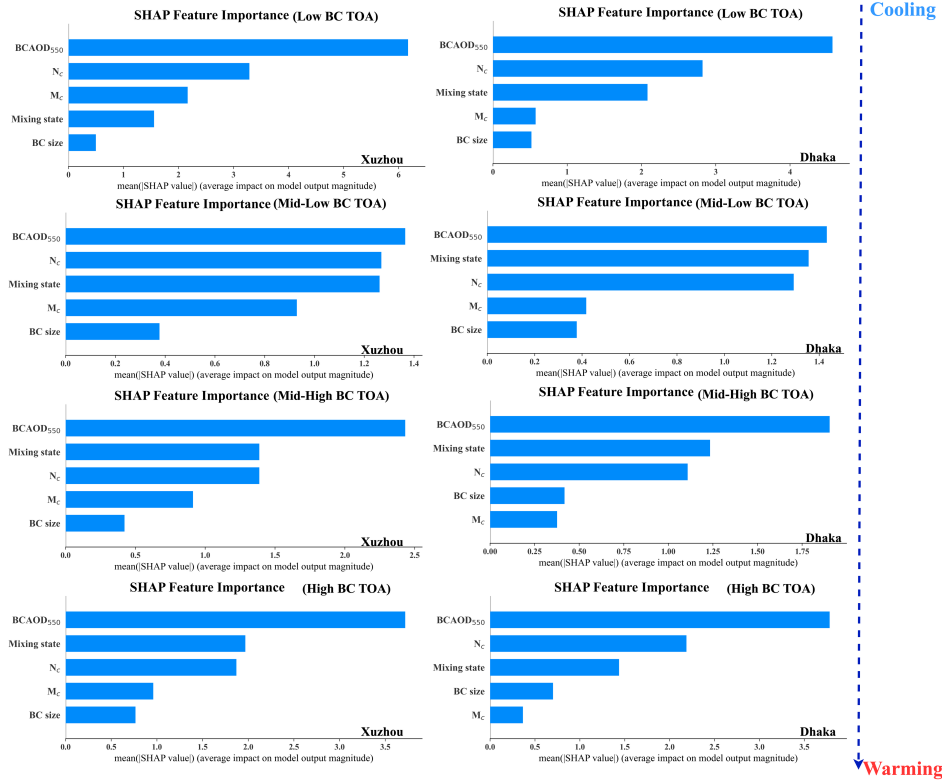
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258 **Section S4**

259 **SHAP Analysis across different TOA regimes.**

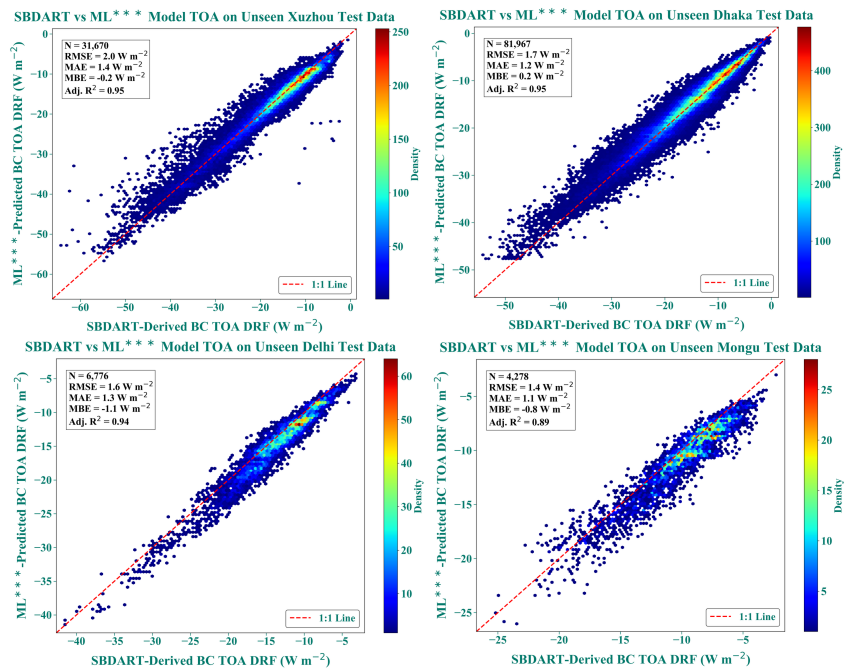
260 SHAP analysis performed across different BC TOA ranges (Figure S6) further elucidates how predictor
261 importance evolves as predicted aerosol radiative effects transition from cooling to warming. In both regions,
262 BCAOD₅₅₀ dominates the ML predicted radiative forcing, transitioning from scatter-driven cooling at high
263 values (Xuzhou: 0.4-2.2; Dhaka: 0.41-1.9) under low TOA to absorption-driven warming at lower values
264 (Xuzhou: 0.01-0.45; Dhaka: 0.01-0.43) under high BC TOA. Notably, overlapping BCAOD₅₅₀ ranges (0.4-
265 0.45) exhibit both positive and negative SHAP contributions (warming and cooling tendencies), consistent
266 with competing absorption and scattering influences in the underlying RTM-derived training data, and
267 suggesting that the model learns sensitivity to small changes in aerosol microphysics that shift the predicted
268 overall effect. This shift is accompanied by a systematic narrowing of the range of aerosol properties
269 associated with cooling as predicted BC TOA forcing intensifies. In Xuzhou, the upper bound of BCAOD₅₅₀
270 associated with cooling decreases from 2.2 (Low BC TOA) to 2.0 (High BC TOA), while in Dhaka, it narrows
271 from 1.9 to 0.8. Concurrently, the lower bound of BCAOD₅₅₀ associated with warming shifts towards smaller
272 values, indicating although total attenuation decreases, the plausible transition from scattering-to absorption-
273 dominated regimes may drive the BC TOA towards more warming. These patterns emphasize that BC TOA
274 transition is governed by coupled evolution of aerosol loading and microphysical changes.

275 A key regional contrast emerges in the role of mixing state, whose relative importance increases with TOA
276 forcing, particularly from mid-high to high BC TOA conditions in Xuzhou. This growth exceeds that of N_c,
277 despite N_c retaining a higher overall SHAP magnitude. In Dhaka, mixing-state SHAP effects peak under
278 intermediate BC TOA conditions and weaken at high and low BC TOA, indicative of a different coupling
279 between coating evolution and column loading. BC size plays a marginal role in both regions but gains
280 modest importance in Dhaka's High BC TOA (75.6% positive SHAP for >220 nm cores), consistent with
281 environments where larger, weakly coated BC cores contribute preferentially to absorption. M_c exerts
282 negligible influence in Dhaka but is associated with cooling negative SHAP contribution in Xuzhou only at
283 high loading (>450 mg m⁻²), consistent with enhanced scattering associated with heavily polluted industrial
284 environments. Overall, while BCAOD₅₅₀ remains the most influential predictor, mixing state emerges as a
285 strong modulator of predictor BC TOA transition regimes under intermediate forcing conditions, with N_c
286 exerting a strong impact under high BC TOA regimes in the learned mapping.



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Figure S6: Shapley Additive Explanations (SHAP) analysis showing the feature importance and feature ranking through mean SHAP values for predicting TOA over Xuzhou and Dhaka for different transitional ranges of BC TOA forcing



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Figure S7: Performance of the revised combined model (ML^{***}) on held-out test data across the four regions (30% from each region, 124,691 total samples, completely unseen during training). Panels show SBDART-derived vs ML^{***}-predicted BC TOA DRF

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Figure S8: Photographs of Xuzhou under high aerosol loading, taken within minutes of each other using identical camera settings (12 mm, f/1.8, ISO 50), with only a slight shutter-speed adjustment to avoid sensor saturation when facing the Sun. Panel (a) shows the view toward the Sun (not directly at it), where strong forward scattering produces dense haze and reduced visibility (2024-10-15 07:42 local time). Panel (b) shows the view in the opposite direction, where apparent scattering and haze are noticeably weaker and visibility is enhanced (2024-10-15 07:44 local time). The pronounced directional contrast qualitatively illustrates anisotropic scattering under heavy aerosol loading, providing visual context for the microphysics-driven absorption–scattering competition discussed in the main text.

Photo courtesy: Prof. Jason Blake Cohen.

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