



Supplement of

Interpretable machine learning quantifies composition and size influences on aerosol spectral absorption

Wenfang Wang et al.

Correspondence to: Pengfei Tian (tianpf@lzu.edu.cn) and Min Chen (chenmin@lzu.edu.cn)

The copyright of individual parts of the supplement might differ from the article licence.

1 S1. Stokes Diameter Conversion Aerodynamic Diameter.

2 SMPS used 8.82-310.6 nm data and APS used 0.523-19.81 μm data. For particle
3 density of $1.7 \text{ g}\cdot\text{cm}^{-3}$ the particle diameter 310.6 nm corresponds to the aerodynamic
4 diameter $D_a = \sqrt{\rho_p/\rho_0} D_p \cong 405 \text{ nm}$, where $\rho_0 = 1 \text{ g}\cdot\text{cm}^{-3}$, and $\rho = 1.7 \text{ g}\cdot\text{cm}^{-3}$ is
5 constant over the whole size range (Shang et al., 2018).

6 S2. Offline chemical analysis of filter samples.

7 For WSII, one quarter of each filter was cut into small pieces and extracted with
8 10 mL ultrapure water ($>18 \text{ M}\Omega\cdot\text{cm}$) by ultrasonication for 30 min; the extraction was
9 repeated once, and the combined extracts were brought to 25 mL with ultrapure water.
10 The extract was filtered through a $0.45 \mu\text{m}$ membrane (GHP, model 4563, Pall Inc.,
11 USA) and stored at $4 \text{ }^\circ\text{C}$ prior to ion chromatography (881 Compact IC Pro, Metrohm,
12 Switzerland; and ICS-1500, Dionex Inc., USA). Cations (Na^+ , K^+ , NH_4^+ , Ca^{2+} , Mg^{2+})
13 were measured with ion chromatograph (881 Compact IC Pro, Metrohm, Switzerland)
14 equipped with a Metrosep C4 column. The eluent was $1.75 \text{ mM HNO}_3 + 0.75 \text{ mM}$
15 dipicolinic acid. Anions (Cl^- , NO_3^- , SO_4^{2-}) were measured with an ion chromatograph
16 (ICS-1500, Dionex Inc., USA) equipped with an IonPac AS22 column. The eluent was
17 $4.5 \text{ mM Na}_2\text{CO}_3 + 1.4 \text{ mM NaHCO}_3$. The method detection limits (MDLs) were 0.05,
18 0.10, 0.19, 0.04, 0.01, 0.02, and $0.02 \text{ mg}\cdot\text{L}^{-1}$ for K^+ , NH_4^+ , Ca^{2+} , Mg^{2+} , Cl^- , NO_3^- , and
19 SO_4^{2-} , respectively. Due to generally elevated Na^+ blanks associated with quartz-fiber
20 filters and glassware, Na^+ was excluded here.

21 One quarter of each quartz-fiber filter was used for elemental analysis. The filter
22 portion was cut into small pieces and transferred to a digestion bottle. Concentrated
23 nitric acid (HNO₃, 68%, 10 mL) was added, and the mixture was heated until digestion
24 was complete. After cooling, the digest was diluted to 20 mL with ultrapure water. The
25 resulting solutions were analyzed for Al, Ca, Mg, Fe, and Ti using inductively coupled
26 plasma-atomic emission spectrometry (ICP-AES; iCAP 7400, Thermo, USA). The
27 MDLs were 0.02, 0.19, 0.02, 0.07, and 0.00 mg·L⁻¹ for Al, Ca, Mg, Fe, and Ti,
28 respectively.

29 A circular area of each filter sample using a 0.504 cm² punch was cut for the
30 analysis of OC and EC by thermal/optical carbon analyzer (DRI Model 2015, USA)
31 following the thermal/optical reflectance (TOR) method (Chow et al., 2007). Briefly, a
32 filter punch was progressively heated to evolve carbon fractions under stepwise
33 temperature programs. OC fractions were evolved in a pure He atmosphere at 120 °C
34 (OC1), 250 °C (OC2), 450 °C (OC3), and 550 °C (OC4). Subsequently, EC fractions
35 were evolved in an oxidizing atmosphere of 2% O₂ and 98% He at 550 °C (EC1),
36 700 °C (EC2), and 800 °C (EC3). Evolved carbon was oxidized to CO₂ and then
37 reduced to CH₄ for detection using a flame ionization detector. Pyrolyzed organic
38 carbon (OPC) was monitored using the laser reflectance signal and was defined when
39 the reflected laser signal returned to its initial value after the introduction of O₂ into the
40 analysis atmosphere. OC was calculated as OC1 + OC2 + OC3 + OC4 + OPC, whereas
41 EC was calculated as EC1 + EC2 + EC3 – OPC. The MDLs were 0.18 and 0.04 μg·m⁻²

42 for OC and EC, respectively.

43 S3. Mie theory sensitivity experiments.

44 To provide a physical consistency check for the SHAP-based attribution, we
45 conducted a set of Mie-theory sensitivity experiments to isolate how (i) short-
46 wavelength absorptivity and (ii) particle-size structure can influence the absorption
47 Ångström exponent between 440 and 870 nm. The experiments use the AERONET-
48 retrieved column volume size distribution (VSD), expressed as $\frac{dV}{d\ln r}$ on a discrete
49 radius grid r spanning approximately 0.05–15 μm . For each radius bin, the number size
50 distribution is obtained by

$$51 \quad \frac{dN}{d\ln r} = \frac{1}{\frac{4}{3}\pi r^3} \times \frac{dV}{d\ln r} \quad (\text{S1})$$

52 Given the complex refractive index $m = n - ik$, where n is the real part and k is
53 the imaginary part, we compute the absorption efficiency $Q_{\text{abs}}(r, \lambda)$ using Mie theory
54 (Bohren and Huffman, 1983). The AAOD at wavelength λ is then evaluated by
55 integrating absorption cross section (up to a proportionality constant) weighted by the
56 number distribution,

$$57 \quad AAOD_{\lambda} \propto \int Q_{\text{abs}}(r, \lambda) \pi r^2 \frac{dN}{d\ln r} d\ln r \quad (\text{S2})$$

58 Because our goal is sensitivity and mechanistic consistency rather than an exact
59 retrieval of absolute AAOD, the proportionality constant is not required; all
60 comparisons are made within the same computational framework. The AAE between
61 440 and 870 nm is computed as

62
$$AAE = - \frac{\ln\left(\frac{AAOD_{440}}{AAOD_{870}}\right)}{\ln\left(\frac{440}{870}\right)} \quad (S3)$$

63 To isolate the influence of composition-related changes, we varied only the
 64 imaginary refractive index at 440 nm (k_{440}) while holding the size distribution fixed.
 65 Specifically, k_{440} was scaled by multiplicative factors of 0.6, 0.8, 1.0, 1.2, and 1.5, while
 66 k_{870} was kept unchanged. The real part n was held constant in these tests (1.5, the mean
 67 AERONET-retrieved value). This design emphasizes changes in absorption that
 68 preferentially affect shorter wavelengths, consistent with increased contributions from
 69 aerosol components that absorb more strongly in the near-UV (e.g., absorbing mineral
 70 dust and brown carbon), and avoids conflating this effect with broadband absorbers
 71 whose spectral dependence is weak.

72 To isolate the role of particle-size structure, we held the refractive index fixed and
 73 perturbed the size distribution by shifting the characteristic radii of the fine and coarse
 74 modes. We separated the VSD into a fine subset and a coarse subset using a radius
 75 threshold of $r = 0.6 \mu\text{m}$. For each subset, we performed a log-radius shift,

76
$$\left(\frac{dV}{d\ln r}\right)_{new}(r) = \left(\frac{dV}{d\ln r}\right)_{old}\left(\frac{r}{s}\right) \quad (S4)$$

77 where s is the shift factor. We used $s = 0.8, 0.9, 1.0, 1.1,$ and 1.2 , corresponding to
 78 moving the characteristic radius to 80%, 90%, 100%, 110%, and 120% of the baseline
 79 value. The shift is implemented by linear interpolation in $\ln r$, with values outside the
 80 original subset radius range set to zero. To ensure that this perturbation represents a
 81 change in size location rather than loading, we renormalize the shifted subset so that its

82 mode-integrated volume is conserved,

$$83 \quad \int \left(\frac{dV}{d\ln r} \right)_{new} d\ln r = \int \left(\frac{dV}{d\ln r} \right)_{old} d\ln r \quad (S5)$$

84 The unperturbed subset is kept unchanged, so that “fine-mode shift” modifies only
 85 the fine subset while the coarse subset remains fixed, and vice versa for “coarse-mode
 86 shift”. The resulting sensitivity curves are summarized in Fig. S11.

87 Table S1. Correlation Matrix of Aerosol Chemical Species

	NH ₄ ⁺	K ⁺	Mg ²⁺	Ca ²⁺	Cl ⁻	SO ₄ ²⁻	NO ₃ ⁻	OC	EC	Al	Ca	Fe	Mg	Ti
NH ₄ ⁺	1.00	0.90	0.25	-0.07	0.50	0.95	0.99	0.83	0.78	0.08	0.07	0.34	0.23	0.20
K ⁺		1.00	0.36	-0.07	0.52	0.85	0.89	0.84	0.82	0.03	-0.03	0.24	0.20	0.09
Mg ²⁺			1.00	0.76	0.74	0.24	0.23	0.35	0.33	0.30	0.31	0.03	0.34	0.08
Ca ²⁺				1.00	0.33	-0.05	-0.07	-0.06	-0.05	0.11	0.11	-0.15	0.11	-0.09
Cl ⁻					1.00	0.45	0.48	0.57	0.54	0.35	0.39	0.13	0.40	0.18
SO ₄ ²⁻						1.00	0.95	0.76	0.72	0.09	0.08	0.27	0.19	0.21
NO ₃ ⁻							1.00	0.82	0.77	0.07	0.06	0.34	0.23	0.19
OC								1.00	0.82	0.10	0.10	0.35	0.22	0.08
EC									1.00	0.00	-0.02	0.14	0.12	0.00
Al										1.00	0.88	0.07	0.91	0.11
Ca											1.00	0.15	0.86	0.20
Fe												1.00	0.11	0.60
Mg													1.00	0.07
Ti														1.00

88

89

90 Table S2. Summary statistics of standardized MLR coefficients and model performance
 91 from 1000 bootstrap resamples for the extended model specifications.

coef	mean	std	p2.5	p97.5
a	0.006633	0.069351	-0.12142	0.148516
EC	-7.5E+11	9.37E+11	-2.7E+12	9.93E+11
OM	-2.1E+12	2.57E+12	-7.3E+12	2.73E+12
FMD	-7.3E+12	9.12E+12	-2.6E+13	9.67E+12
nd-WSII	-7.1E+12	8.9E+12	-2.5E+13	9.43E+12
D _{SMPS}	0.032792	0.125148	-0.18773	0.299769
D _{APS}	0.446719	0.086028	0.300897	0.630525
R ²	0.74	0.03	0.67	0.77
RMSE	0.13	0.01	0.13	0.15
MAE	0.10	0.01	0.09	0.16

92 The extended model refers to the standardized MLR specification including EC and
 93 OM fractions. coef denotes the regression coefficient (including the intercept term, “a”).
 94 mean and std are the bootstrap mean and standard deviation of each coefficient across
 95 1000 resamples. p2.5 and p97.5 are the 2.5th and 97.5th percentiles of the bootstrap
 96 distribution, respectively, forming the percentile-based 95% bootstrap confidence
 97 interval.

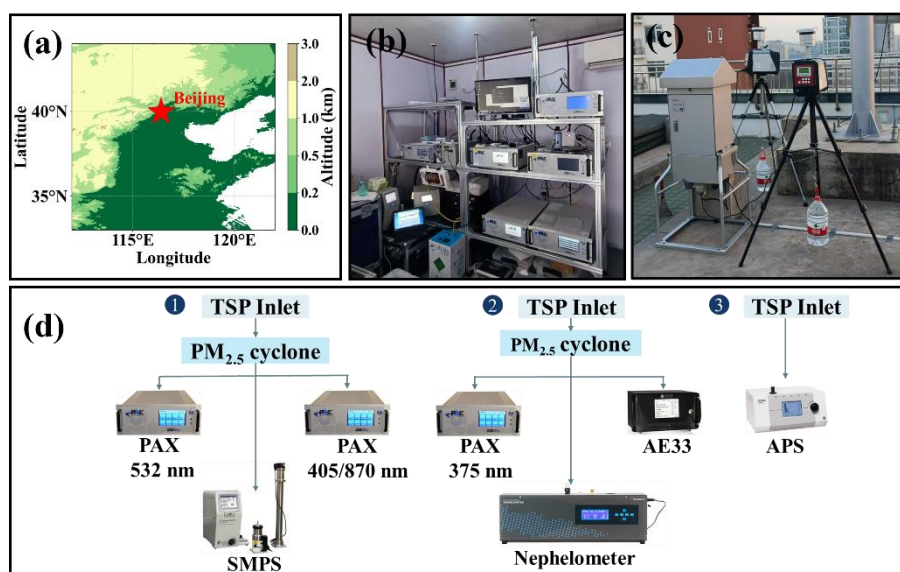
98
 99

100 Table S3. Summary statistics of standardized MLR coefficients and model performance
 101 from 1000 bootstrap resamples for the reduced model specifications.

coef	mean	std	p2.5	p97.5
a	0.00	0.07	-0.13	0.14
FMD	0.35	0.17	0.04	0.71
nd-WSII	-0.16	0.17	-0.50	0.17
D _{SMPS}	-0.02	0.12	-0.24	0.26
D _{APS}	0.44	0.09	0.30	0.64
R ²	0.74	0.02	0.68	0.76
RMSE	0.13	0.01	0.13	0.15
MAE	0.10	0.01	0.09	0.11

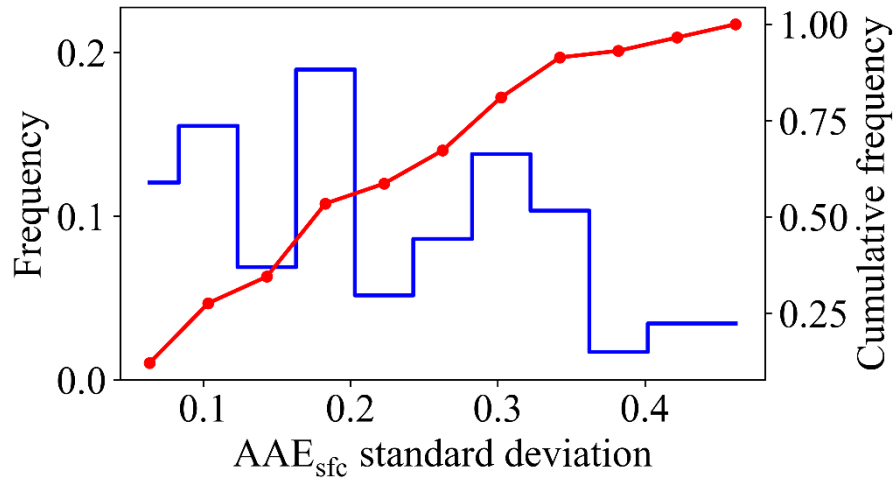
102 The reduced model refers to the standardized MLR specification excluding EC and OM
 103 fractions.

104



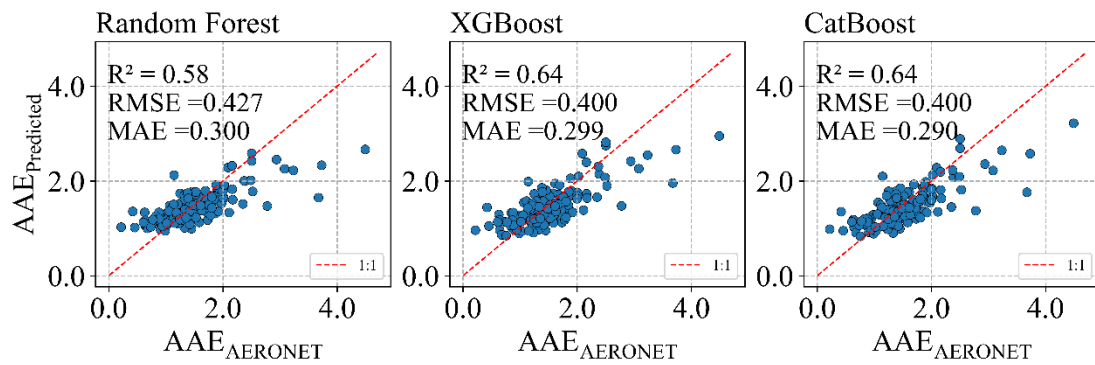
105

106 **Figure S1.** Geographic location of observation site and observation instruments. (a)
 107 Geographical map, location of the field campaign. (b, c) Photographs during the field
 108 campaign. (d) Airflow settings of the online instruments.



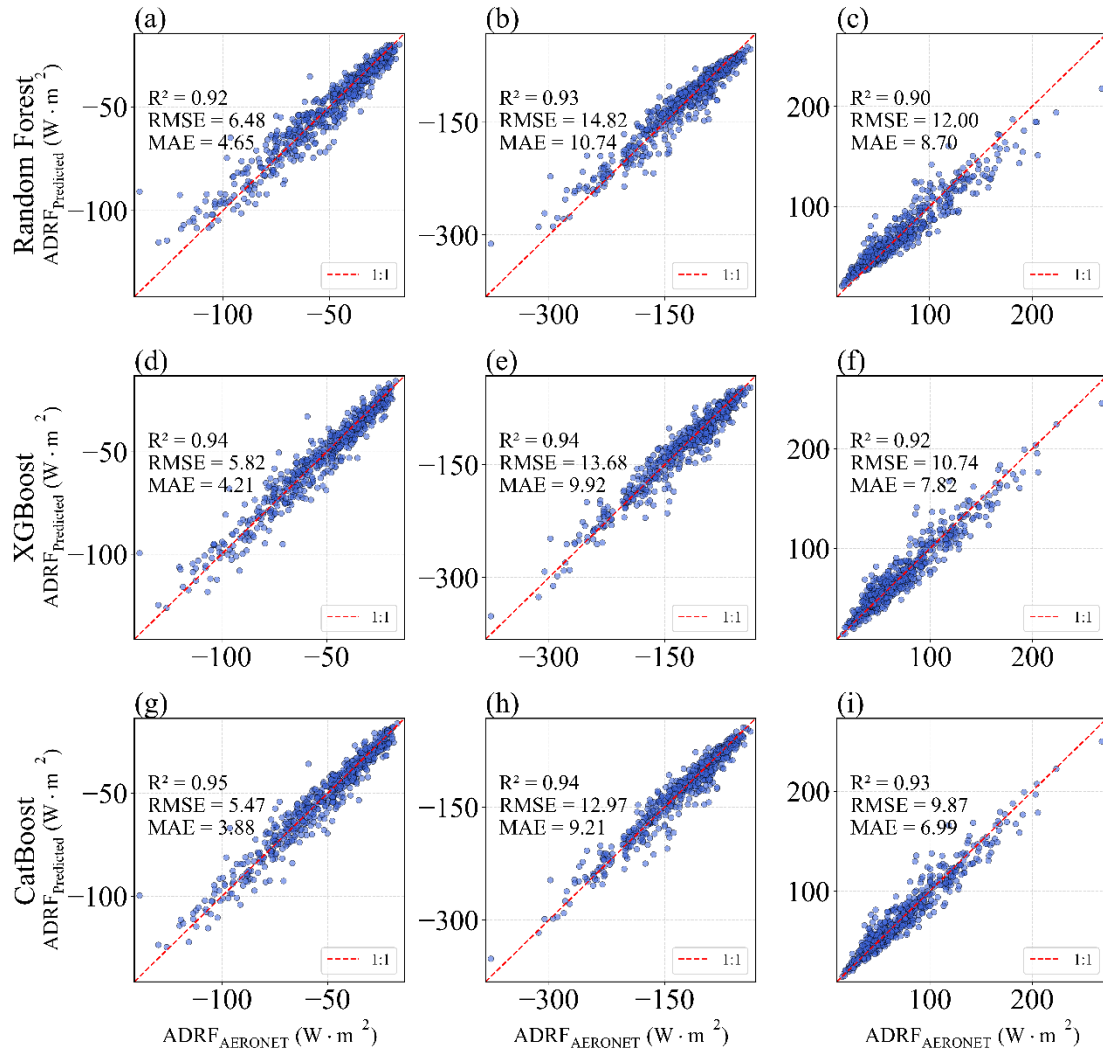
109

110 **Figure S2.** Frequency (blue) and cumulative frequency (red) distributions of within-
 111 window AAE_{sfc} standard deviation (11.5 h).



112

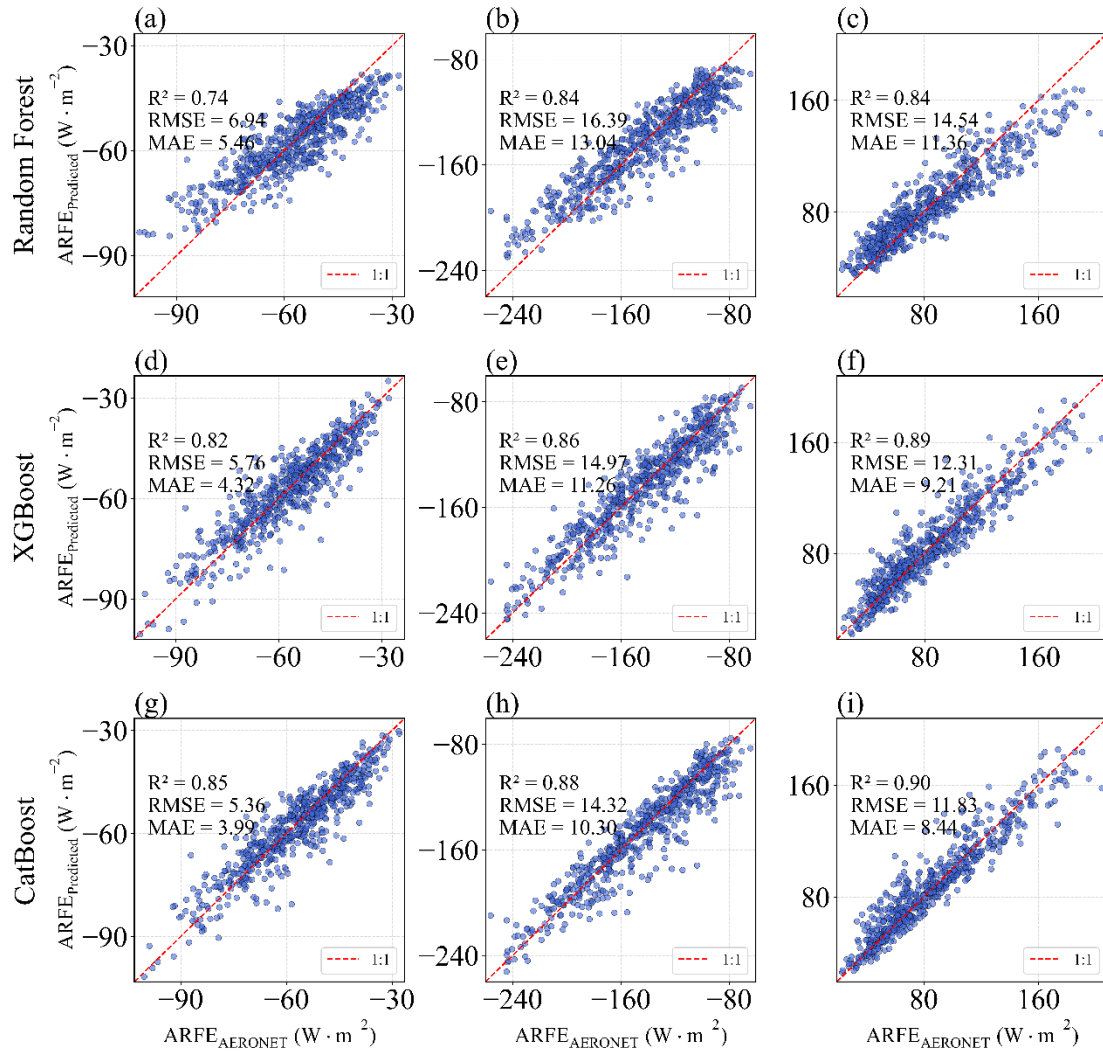
113 **Figure S3.** Comparison of aerosol absorption Ångström exponent (AAE) predictions
 114 and AERONET observations by different models.



115

116 **Figure S4.** Comparison of aerosol direct radiative forcing (ADRF) predictions and

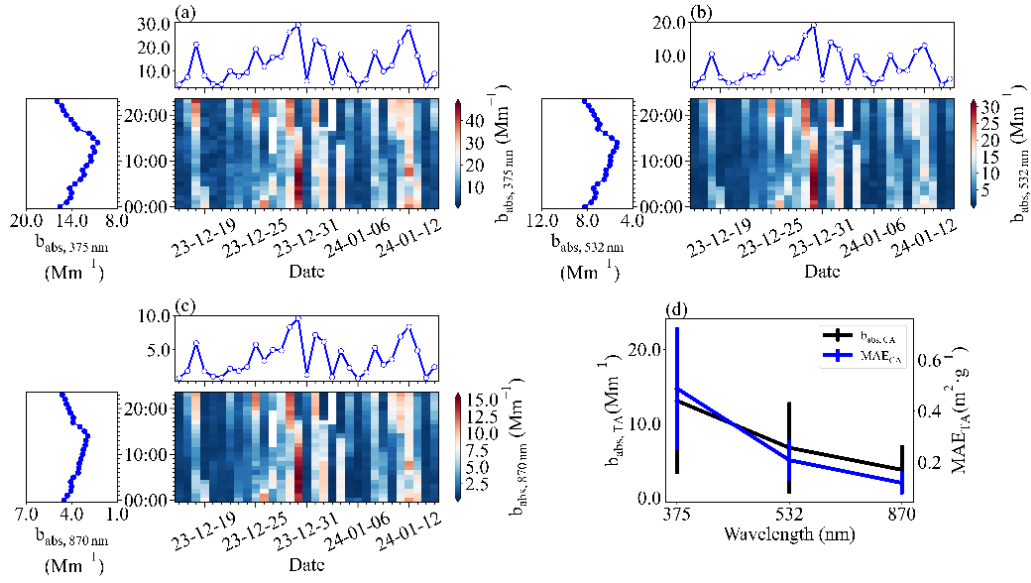
117 AERONET observations by different models.



118

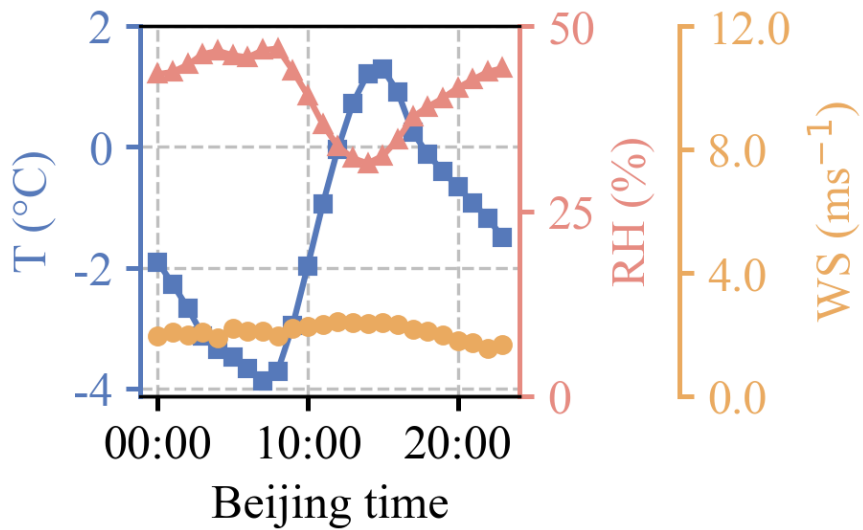
119 **Figure S5.** Comparison of aerosol radiative forcing efficiency (ARFE) predictions and
 120 AERONET observations by different models.

121



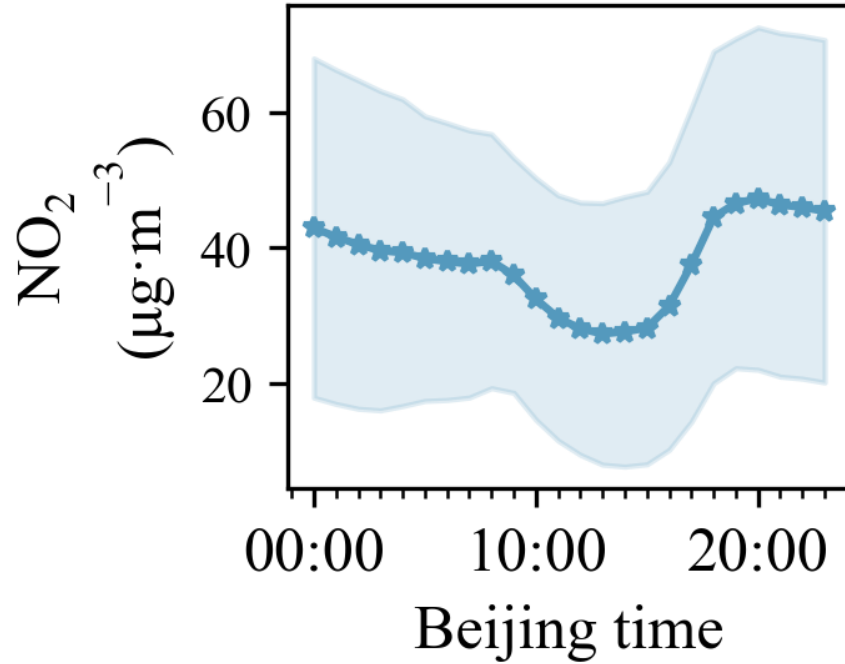
122

123 **Figure S6.** (a, b, c) The time series of mass concentrations of absorption coefficients at
 124 375, 405, 870 nm, mass absorption efficiency at 375, 405, 870 nm. (d). Wavelength
 125 dependence of absorption coefficients and mass absorption efficiency.



126

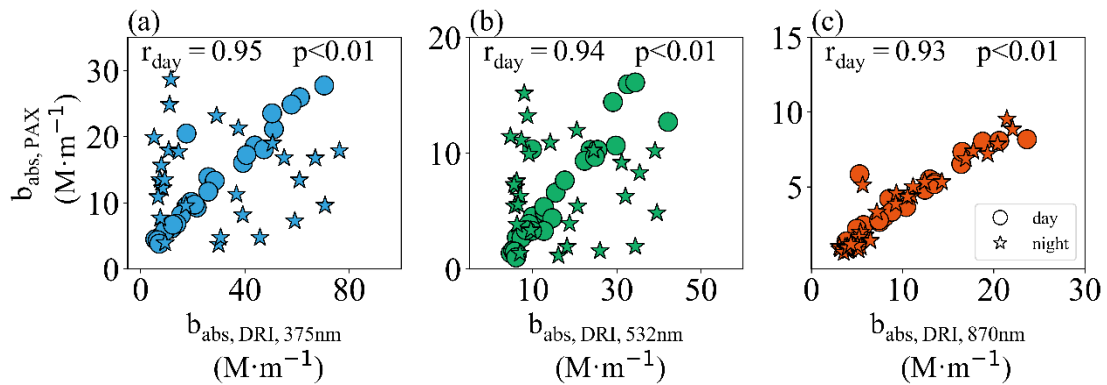
127 **Figure S7.** Daily variation of Temperature, Relative Humidity, Wind Speed.



128

129 **Figure S8.** Daily variation of nitrogen dioxide (NO_2). Hourly NO_2 mass concentrations
 130 were obtained from the China National Environmental Monitoring Network for the
 131 Beijing urban site.

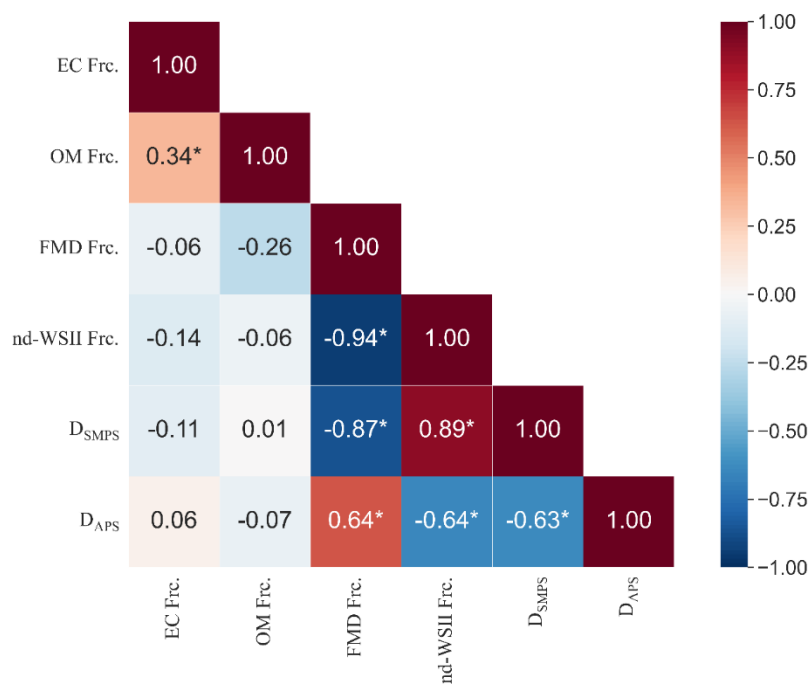
132



133

134 **Figure S9.** Scatter plots of the aerosol absorption coefficients between PAX and DRI.
 135 (a) 375 nm. (b) 532 nm. (c) 870nm.

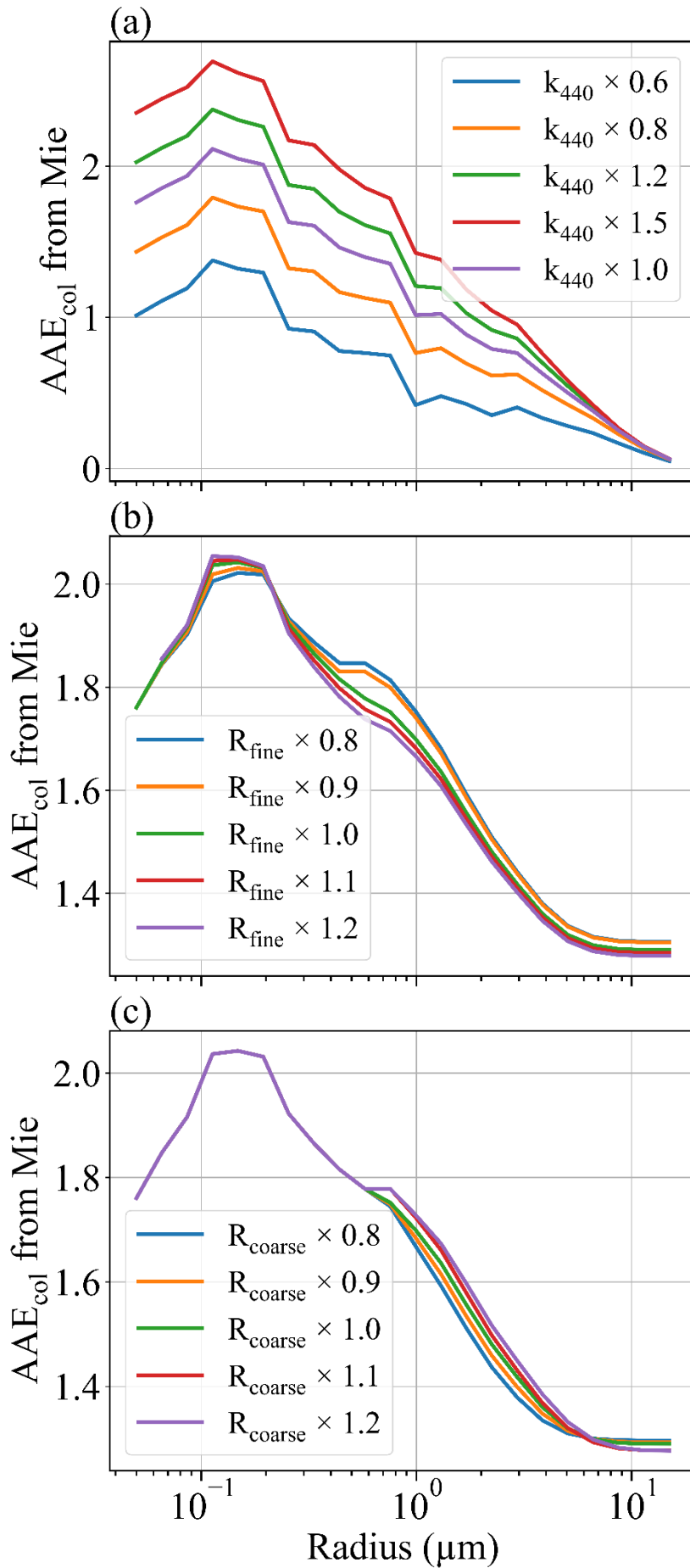
136



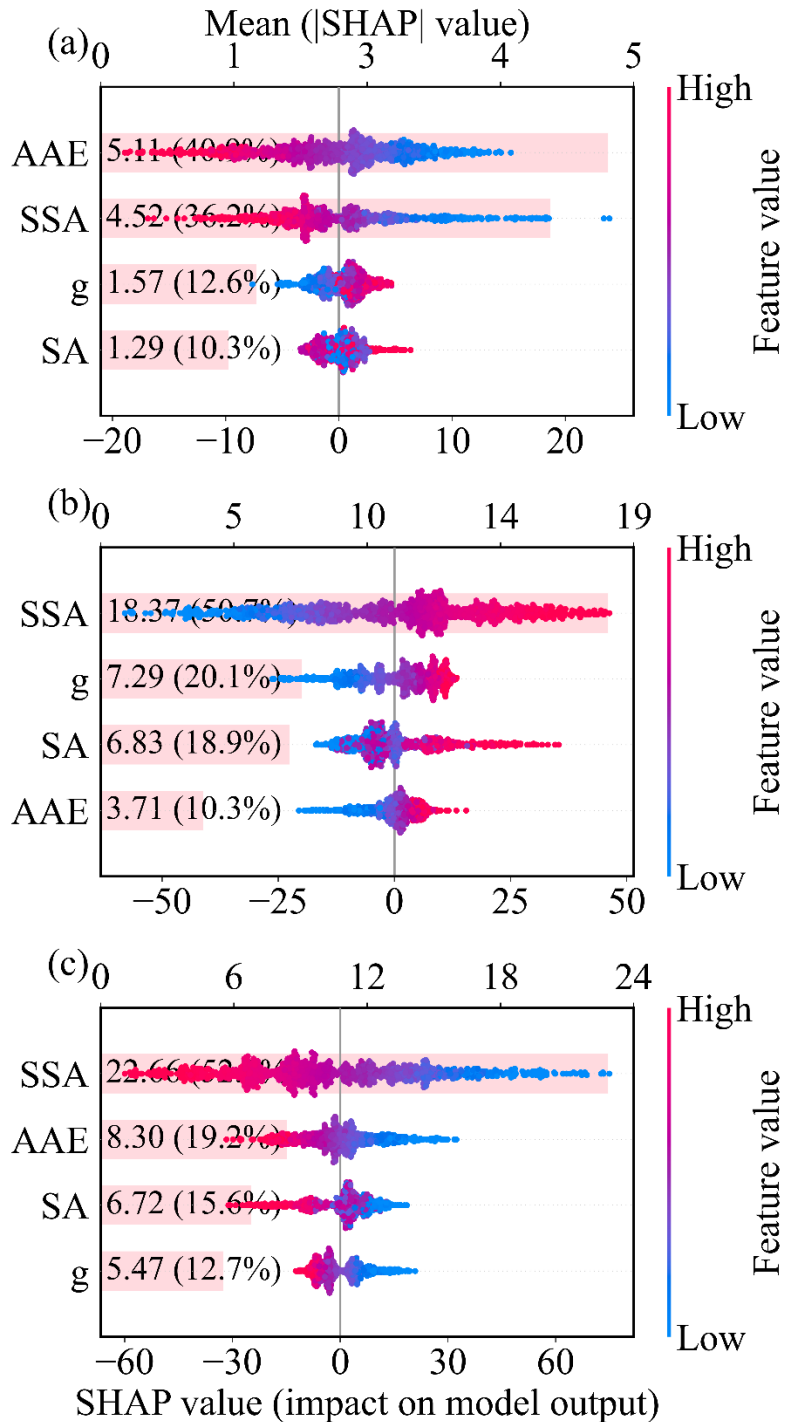
137

138 **Figure S10.** Correlation analysis among chemical composition (non-dust water-soluble
 139 ions (nd-WSII), fine mineral dust (FMD), elemental carbon (EC), organic matter (OM))
 140 fraction and size parameter (fine-mode mean diameter from SMPS (D_{SMPS}), coarse-
 141 mode mean diameter from APS (D_{APS})).

142



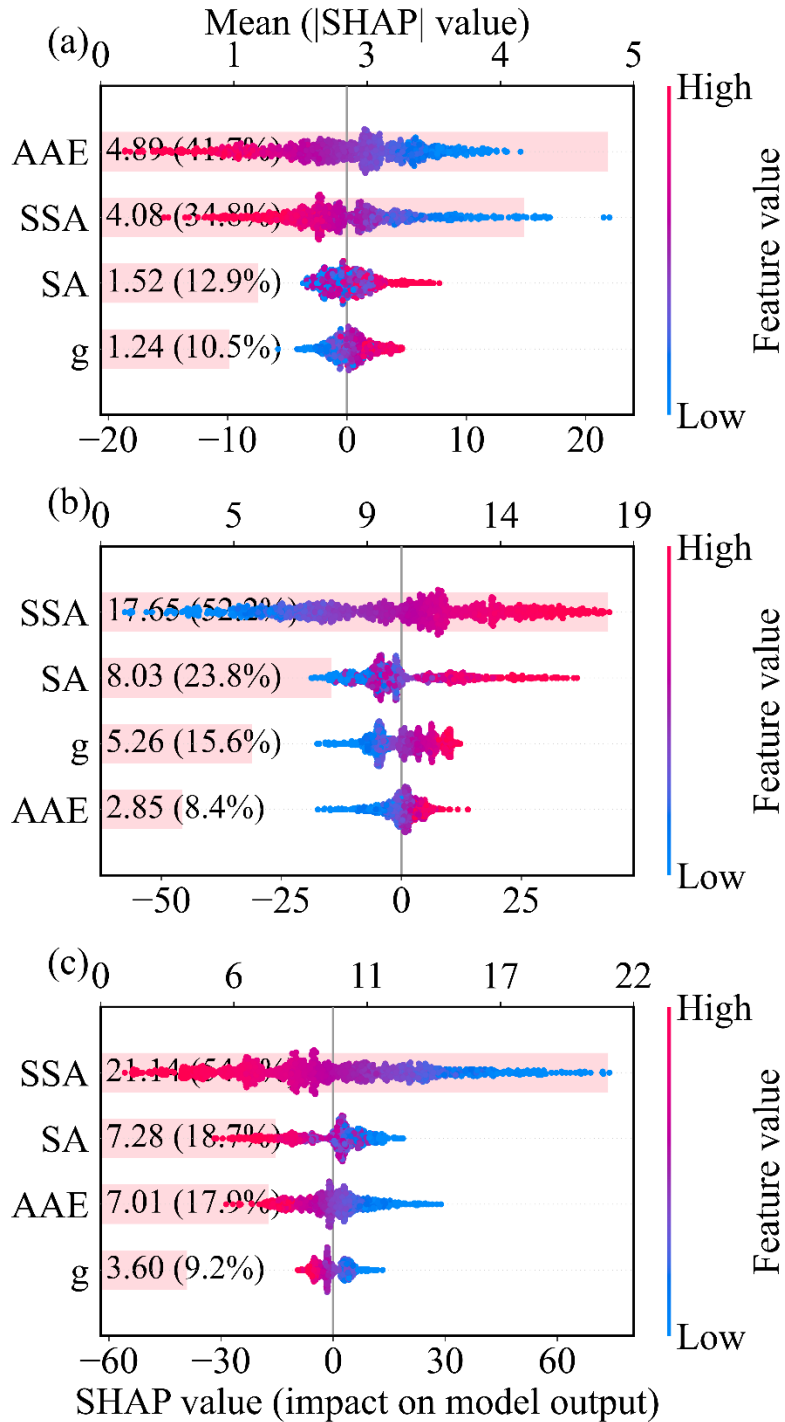
144 Figure S11. Mie-theory sensitivity experiments linking composition-related
145 absorptivity and size structure to AAE. (a) Sensitivity to shortwave absorptivity: the
146 imaginary refractive index at 440 nm (k_{440}) is scaled by factors of 0.6, 0.8, 1.0, 1.2, and
147 1.5 while the size distribution and k_{870} are held fixed (the real part is fixed at $n=1.5$, the
148 AERONET mean). (b) Sensitivity to fine-mode size: the fine-mode characteristic radius
149 is shifted to 80%, 90%, 100%, 110%, and 120% of its baseline value while conserving
150 the fine-mode integrated volume; the coarse mode is unchanged. (c) Sensitivity to
151 coarse-mode size: the coarse-mode characteristic radius is shifted to the same factors
152 while conserving the coarse-mode integrated volume; the fine mode is unchanged.
153 Curves are shown as a function of particle radius; changes in AAE reflect the differing
154 impacts of absorptivity and size perturbations under otherwise fixed conditions.



155

156 **Figure S12.** SHAP analysis with AOD fixed at its 25th percentile quantifies the relative
 157 contributions of single scattering albedo (SSA), asymmetry parameter (g), surface
 158 albedo (SA) and columnar AAE in driving aerosol radiative forcing efficiency (ARFE)
 159 variations at the top of the atmosphere (a), bottom (b), and in the atmosphere (c). The

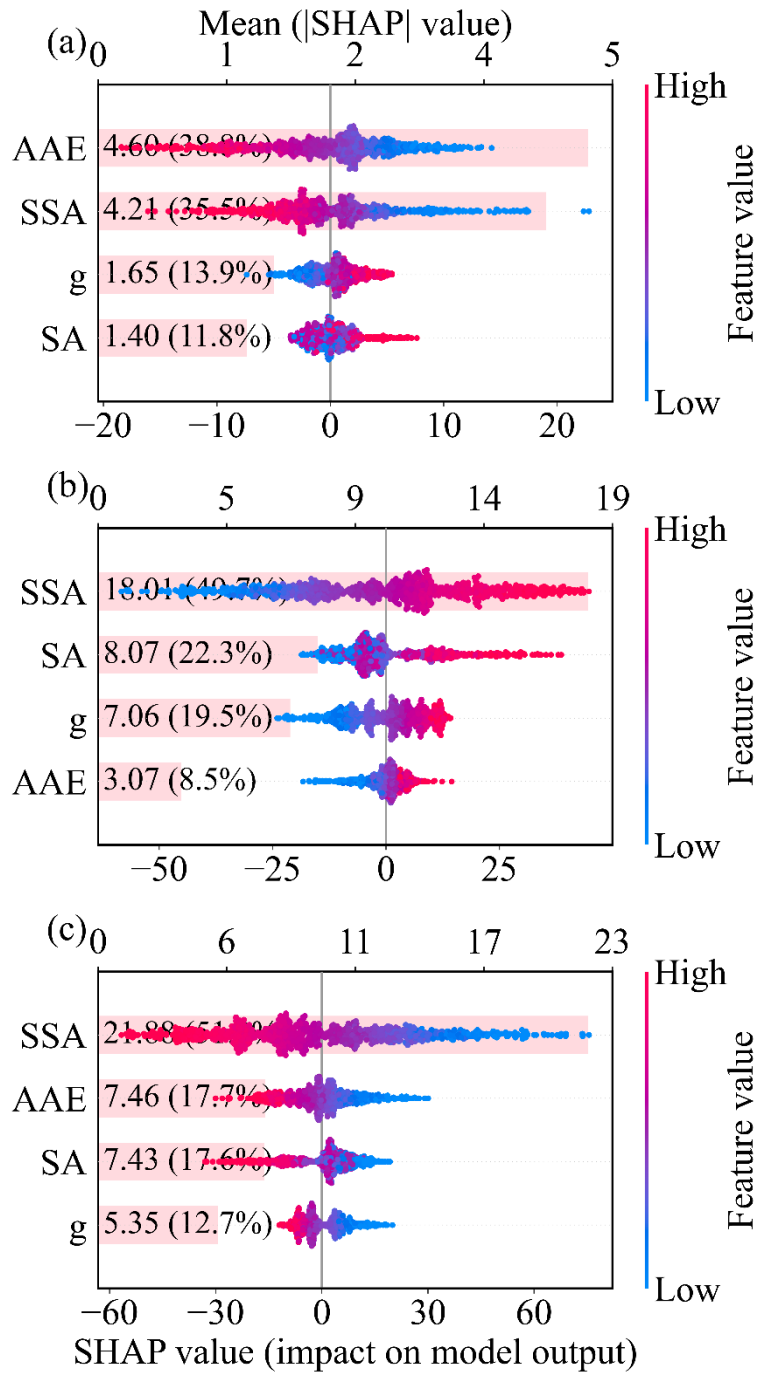
160 mean absolute SHAP values (numbers in parentheses) indicate the relative contribution
 161 of each predictor to the model output.



162

163 **Figure S13.** Same as Figure S12, except aerosol optical depth (AOD) is fixed at its 75th

164 percentile.



166

167 **Figure S14.** Same as Figure S12, except aerosol optical depth (AOD) is fixed at its

168 mean value.

169

170 **References**

171 Bohren, C. F. and Huffman, D. R.: *Absorption and Scattering of Light by Small*
172 *Particles*, John Wiley & Sons, Inc., New York, 1983.

173 Chow, J. C., Watson, J. G., Chen, L.-W. A., Chang, M. C. O., Robinson, N. F., Trimble,
174 D., and Kohl, S.: The IMPROVE_A Temperature Protocol for Thermal/Optical
175 Carbon Analysis: Maintaining Consistency with a Long-Term Database, *Journal*
176 *of the Air & Waste Management Association*, 57, 1014–1023,
177 <https://doi.org/10.3155/1047-3289.57.9.1014>, 2007.

178 Shang, D., Hu, M., Zheng, J., Qin, Y., Du, Z., Li, M., Fang, J., Peng, J., Wu, Y., Lu, S.,
179 and Guo, S.: Particle number size distribution and new particle formation under
180 the influence of biomass burning at a high-altitude background site at Mt. Yulong
181 (3410 m), China, *Atmos. Chem. Phys.*, 18, 15687–15703,
182 <https://doi.org/10.5194/acp-18-15687-2018>, 2018.