



*Supplement of*

## **Lagrangian particle–based simulation of aerosol-dependent vertical variation of cloud microphysics in a laboratory convection cloud chamber**

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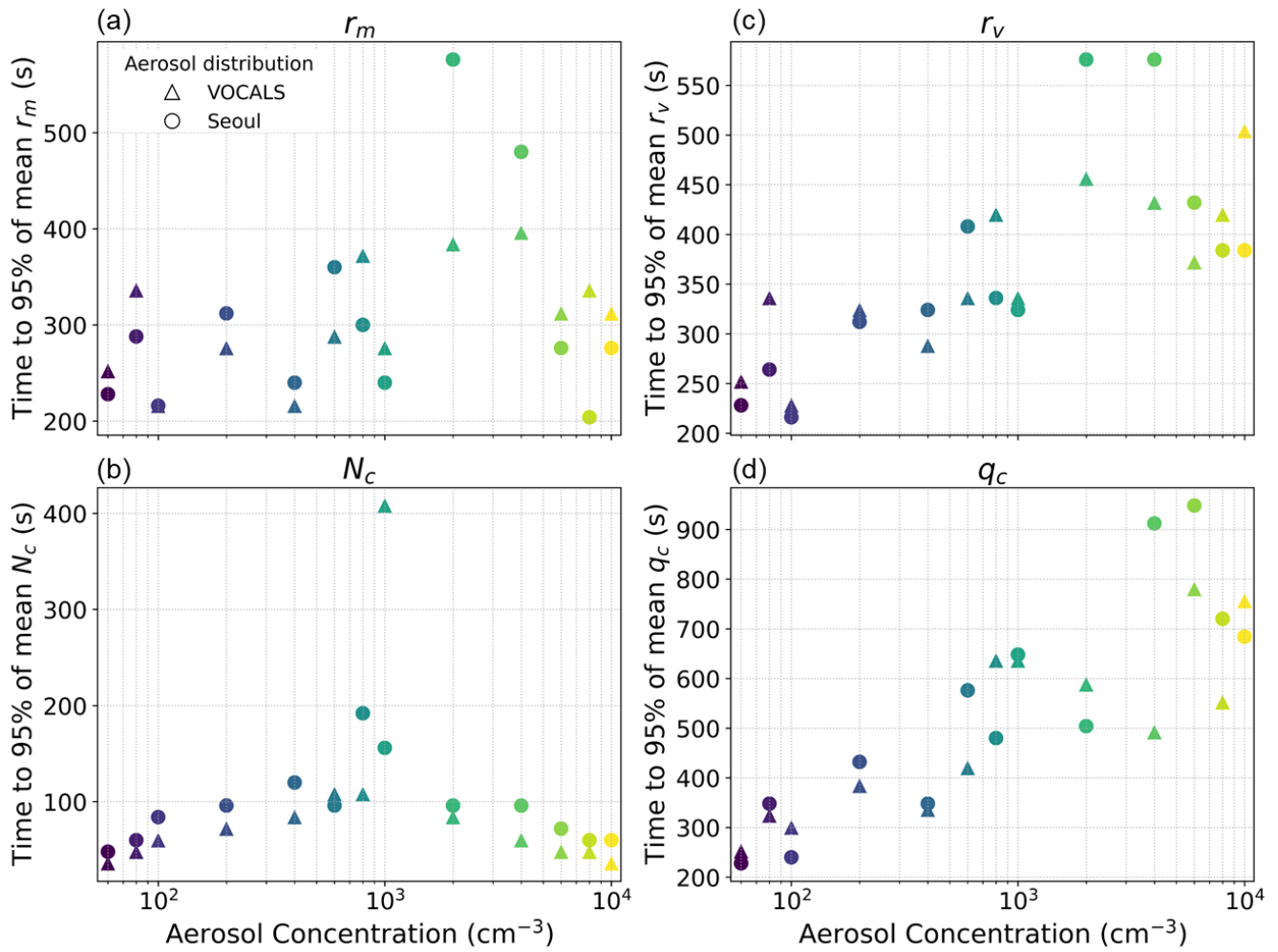
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## Analysis of Microphysical Equilibration Timescales and Their Dependence on Aerosol Concentration

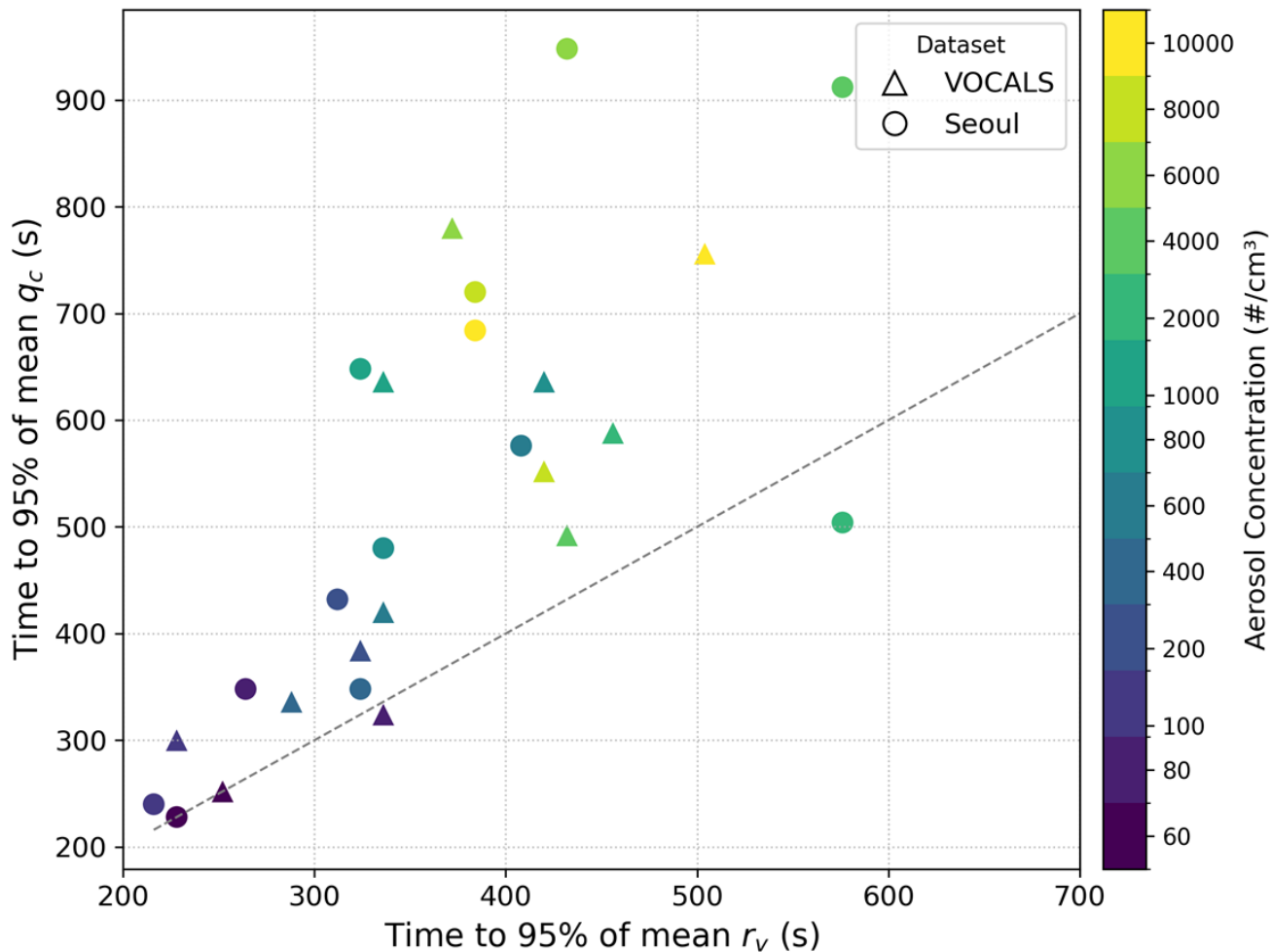
10

Figure 2 in the text shows that each microphysical variable—cloud droplet number concentration ( $N_c$ ), mean droplet radius ( $r_m$ ), and cloud water mixing ratio ( $q_c$ )—exhibits distinct timescales to reach quasi-equilibrium conditions, with these timescales varying systematically according to aerosol concentration. Figures S1 and S2 further elucidate how aerosol-induced changes in supersaturation impact these equilibration timescales. Figure S1 presents the equilibration times, defined as the time to reach 95% of their 15-30 minute average values, for  $q_c$ ,  $N_c$ ,  $r_m$ , and volume-mean droplet radius ( $r_v$ ) across a range of aerosol concentrations. These equilibration times differ systematically:  $N_c$  equilibrates fastest because most aerosol particles rapidly exceed the cloud droplet threshold radius of 1  $\mu\text{m}$  used in this study. Droplet radius ( $r_m$ ,  $r_v$ ) equilibrate next, whereas  $q_c$  reaches equilibrium slowest, particularly under high aerosol loading, due to the continued slow growth of larger droplets. Figure S2 explicitly compares the equilibration times of  $q_c$  and  $r_v$ , highlighting how their difference increases with aerosol concentration. This result indicates that  $q_c$  continues to evolve significantly after droplet sizes have equilibrated. Physically, this might occur because  $q_c$ , being weighted more heavily toward larger droplets (via  $r^3$ ), continues to increase as those large droplets slowly grow. Under high aerosol concentrations, reduced supersaturation prolongs the growth of these larger droplets, further delaying the equilibration of  $q_c$  relative to  $r_v$ . Consequently, the gap between the two equilibration timescales becomes increasingly pronounced as aerosol loading might increase.

30



**Figure S1.** Time required for each microphysical variable to reach 95% of its time-averaged value over the quasi-equilibrium period (15–30 minutes), as a function of aerosol concentration, is shown for VOCALS (triangles) and Seoul (circles) cases. Panels show the timescales associated with the following variables: (a) mean droplet radius ( $r_m$ ), (b) cloud droplet number concentration ( $N_c$ ), (c) volume-mean droplet radius ( $r_v$ ), and (d) cloud liquid water mixing ratio ( $q_c$ ). The plots highlight distinct timescales for each variable to approach steady-state behavior.



40

**Figure S2.** Relationship between the time to reach 95% of the mean quasi-equilibrium (15–30 min) volume-mean droplet radius ( $r_v$ ) and the mean quasi-equilibrium cloud water mixing ratio ( $q_c$ ), for VOCALS (triangles) and Seoul (circles) cases. Marker color represents aerosol concentration. The dashed grey 1:1 line highlights that  $q_c$  generally equilibrates more slowly than  $r_v$ , especially under high aerosol loading.

45