



Supplement of

Enhanced atmospheric oxidation and particle reductions driving changes to nitrate formation mechanisms across coastal and inland regions of north China

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Text S1

Our isotope blank measurements followed the same procedure as the sample isotope analysis. Specifically, in the sample measurements, after purging with high-purity nitrogen, 20 nmol of nitrogen was added to the headspace vial containing the *Pseudomonas aureofaciens* (ATCC13985) strain. For the blank measurements, no sample was added, and after 24 hours, 10 M NaOH was directly injected to quench the reaction before the analysis. The peak area in the chromatogram represents the absolute amount of N₂O reduced by the strain, and the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values correspond to the $\delta^{15}\text{N}$ and $\delta^{18}\text{O}$ values of the sample. The peak area for the samples was around 10, while the peak areas for the two blank measurements were only 0.371 and 0.336, indicating an influence on the isotope values of less than 5%, which is negligible and thus not considered.

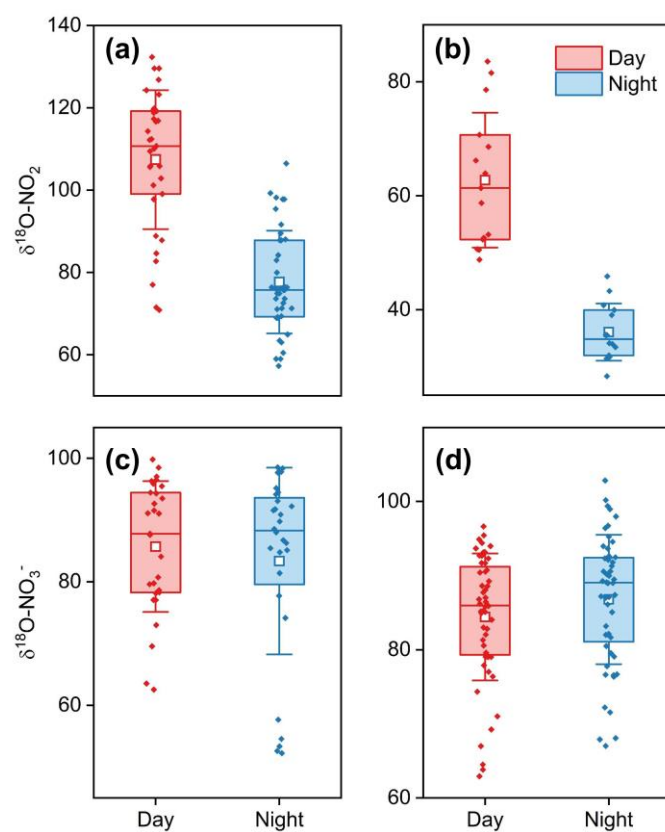


Figure S1. The diurnal values of atmospheric $\delta^{18}\text{O-NO}_2$ in Hefei winter (Zhang et al., 2025) (a) and in Nanchang summer (Cao, 2022) (b), and $\delta^{18}\text{O-NO}_3^-$ in Tianjin winter (Feng et al., 2020) (c) and in Nanjing winter (Zhang et al., 2022) (d).

Text S2

In most studies, the tropospheric $\delta^{15}\text{N}$ -NO_x was often assumed as 0‰ following Walters and Michalski (2016), Luo et al. (2023) and Deng et al. (2024). In addition, the tropospheric $\delta^{18}\text{O}$ -H₂O_(g) in Beijing in winter was determined as -27.9‰ in Wen et al. (2010), and in Qingdao, it was determined as -18.6‰ in Wang et al. (2022). The tropospheric $\delta^{18}\text{O}$ -NO_x ranged from 112‰ to 122‰ (Michalski et al., 2014; Walters and Michalski, 2016). The f_{NO_2} values in Beijing and Qingdao were 0.655 (Luo et al., 2023) and 0.786 (Lian et al., 2022) in winter, respectively.

$$\begin{aligned}\delta^{15}\text{N} - \text{NO}_3^- &= \gamma \times [\delta^{15}\text{N} - \text{NO}_3^-]_{\text{OH}} + (1 - \gamma) \times [\delta^{15}\text{N} - \text{NO}_3^-]_{\text{N}_2\text{O}_5} \\ &= \gamma \times [\delta^{15}\text{N} - \text{HNO}_3]_{\text{OH}} + (1 - \gamma) \times [\delta^{15}\text{N} - \text{HNO}_3]_{\text{N}_2\text{O}_5}\end{aligned}\quad (\text{S1})$$

$$\begin{aligned}\delta^{18}\text{O} - \text{NO}_3^- &= \gamma \times [\delta^{18}\text{O} - \text{NO}_3^-]_{\text{OH}} + (1 - \gamma) \times [\delta^{18}\text{O} - \text{NO}_3^-]_{\text{N}_2\text{O}_5} \\ &= \gamma \times [\delta^{18}\text{O} - \text{HNO}_3]_{\text{OH}} + (1 - \gamma) \times [\delta^{18}\text{O} - \text{HNO}_3]_{\text{N}_2\text{O}_5}\end{aligned}\quad (\text{S2})$$

$$\begin{aligned}[\delta^{15}\text{N} - \text{HNO}_3]_{\text{OH}} &= \delta^{15}\text{N} - \text{NO}_2 \\ &= 1000 \times \left[\frac{(^{15}\alpha_{\text{NO}_2/\text{NO}} - 1)(1 - f_{\text{NO}_2})}{(1 - f_{\text{NO}_2}) + (^{15}\alpha_{\text{NO}_2/\text{NO}} \times f_{\text{NO}_2})} \right] + \delta^{15}\text{N} - \text{NO}_x\end{aligned}\quad (\text{S3})$$

$$[\delta^{15}\text{N} - \text{HNO}_3]_{\text{N}_2\text{O}_5} = 1000 \times (^{15}\alpha_{\text{N}_2\text{O}_5/\text{NO}_2} - 1) + \delta^{15}\text{N} - \text{NO}_x \quad (\text{S4})$$

$$\begin{aligned}[\delta^{18}\text{O} - \text{HNO}_3]_{\text{OH}} &= \frac{2}{3} \times [\delta^{18}\text{O} - \text{NO}_2]_{\text{OH}} + \frac{1}{3} \times [\delta^{18}\text{O} - \text{OH}]_{\text{OH}} \\ &= \frac{2}{3} \times \left[\frac{1000 \times (^{18}\alpha_{\text{NO}_2/\text{NO}} - 1) \times (1 - f_{\text{NO}_2})}{(1 - f_{\text{NO}_2}) + (^{18}\alpha_{\text{NO}_2/\text{NO}} \times f_{\text{NO}_2})} + [\delta^{18}\text{O} - \text{NO}_x] \right] \\ &\quad + \frac{1}{3} \times \left[(\delta^{18}\text{O} - \text{H}_2\text{O}_{(g)}) + 1000 \times (^{18}\alpha_{\text{OH}/\text{H}_2\text{O}_{(g)}} - 1) \right]\end{aligned}\quad (\text{S5})$$

$$[\delta^{18}\text{O} - \text{HNO}_3]_{\text{N}_2\text{O}_5} = \left(\delta^{18}\text{O} - \text{NO}_2 + 1000 \times (^{18}\alpha_{\text{N}_2\text{O}_5/\text{NO}_2} - 1) \right) \times \frac{5}{6} + \delta^{15}\text{N} - \text{H}_2\text{O} \times \frac{1}{6} \quad (\text{S6})$$

$$1000(^m\alpha_{x/y} - 1) = \frac{A}{T^4} \times 10^{10} + \frac{B}{T^3} \times 10^8 + \frac{C}{T^2} \times 10^6 + \frac{D}{T} \times 10^4 \quad (\text{S7})$$

Table S1. Values of $\delta^{18}\text{O}$ from atmospheric components.

Components	Values (‰)	References
O_3	From 80 to 130	Michalski et al., 2011
O_2	23.5	Kroopnick and Craig; 1972
H_2O (g) in Beijing winter	-27.9	Wen et al., 2010
H_2O (g) in Qingdao winter	-18.6	Wang et al., 2022
$\cdot\text{OH}$ in Beijing winter	From -72.4 to -64.9	$\delta^{18}\text{O}-\text{OH} = \delta^{18}\text{O}-\text{H}_2\text{O}_{(\text{g})} + 1000(^{18}\alpha_{\text{X/Y}} - 1)$ (Walters and Michalski, 2016)
$\cdot\text{OH}$ in Qingdao winter	From -61.2 to -57.8	

Table S2. $^{15}\alpha_{A/B}$ and $^{18}\alpha_{A/B}$ regression coefficients as a function of the temperature ($150\text{ K} \leq T \leq 450\text{ K}$) (Walters and Michalski, 2015, 2016).

		A	B	C	D
$^{15}\alpha_{A/B}$	$\text{N}_2\text{O}_5/\text{NO}_2$	0.69398	-1.9859	2.3876	0.16308
	NO_2/NO	3.8834	-7.7299	6.0101	-0.17928
$^{18}\alpha_{A/B}$	NO/NO_2	-0.04129	1.1605	-1.8829	0.74723
	$\cdot\text{OH}/\text{H}_2\text{O}_{(\text{g})}$	2.1137	-3.8026	2.5653	0.5941
	$\text{N}_2\text{O}_5/\text{NO}_2$	-0.54136	0.13073	1.2477	-0.1272

Table S3. Equations for calculating the statistical evaluation indices.

Statistical index	Formula
1. Mean Bias	$\text{MB} = \frac{1}{N} \sum_1^N (\text{Sim} - \text{Obs})$
2. Root Mean Square Error	$\text{RMSE} = \sqrt{\frac{1}{N} \sum_1^N (\text{Sim} - \text{Obs})^2}$
3. Index of agreement, IOA	$\text{IOA} = 1 - \frac{\sum_{i=1}^N (\text{Sim} - \text{Obs})^2}{\sum_{i=1}^N (\text{Sim} - \text{Obs} + \text{Obs} - \overline{\text{Obs}})^2}$
4. Normalized Mean Bias	$\text{NMB} = \frac{1}{N} \sum_1^N \left(\frac{\text{Sim} - \text{Obs}}{\text{Obs}} \right)$
5. Normalized Mean Error	$\text{NME} = \frac{1}{N} \sum_1^N \left \frac{\text{Sim} - \text{Obs}}{\text{Obs}} \right $
6. Correlation coefficient (R)	$\text{R} = \frac{1}{N} \sum_{i=1}^N \left[\frac{(\text{Sim} - \overline{\text{Sim}})(\text{Obs} - \overline{\text{Obs}})}{S_p S_o} \right]$ $S_p = \left[\frac{1}{N} \sum_{i=1}^N (\text{Sim} - \overline{\text{Sim}})^2 \right]^{\frac{1}{2}}$ $S_o = \left[\frac{1}{N} \sum_{i=1}^N (\text{Obs} - \overline{\text{Obs}})^2 \right]^{\frac{1}{2}}$

Table S4. Sources of nitrate observation data for the winter of 2013 and the winter of 2018 in the NCP.

City	Winter, 2013	Winter, 2018
Beijing	Song et al. (2019)	Fan et al. (2020)
Tianjin	Yao et al. (2020)	Observation
Shijiazhuang	Wang et al. (2016)	Zhou et al. (2020)
Jinan	Cheng et al. (2021)	Observation
Zhengzhou	Wei et al. (2019)	Dong et al. (2020)
Qingdao	Observation	Observation
Yantai	/	/

The NO_3^- observation data collected during the winter of 2018 for Tianjin were sourced from direct observations by the group of Li Xiaodong at Tianjin University (sampling site: Building 19 rooftop, Tianjin University; coordinates: 39.11°N, 117.16°E). The NO_3^- observation data collected during the winter of 2018 in Jinan were sourced from observations by the group of Xue Likun at Shandong University (sampling site: Jinan City Environmental Monitoring Station; coordinates: 36.66°N, 117.05°E). For Qingdao, NO_3^- observation data for both the winter of 2018 and the winter of 2013 were derived from our own observations.

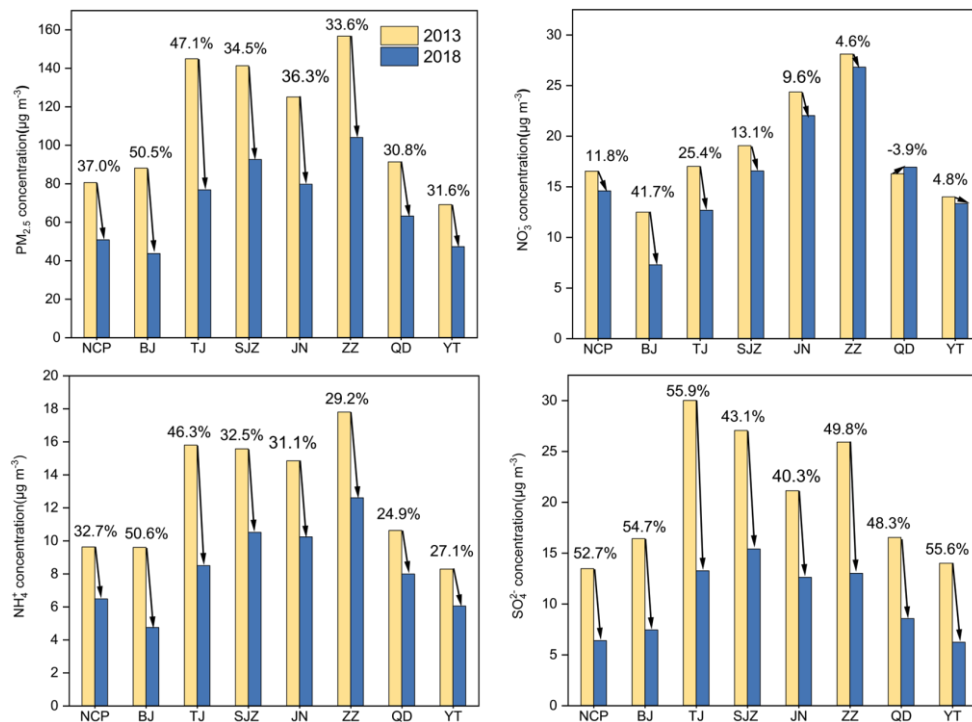


Figure S2. Concentrations of PM_{2.5} and its components in seven major cities in the NCP region during the winters of 2013 and 2018.

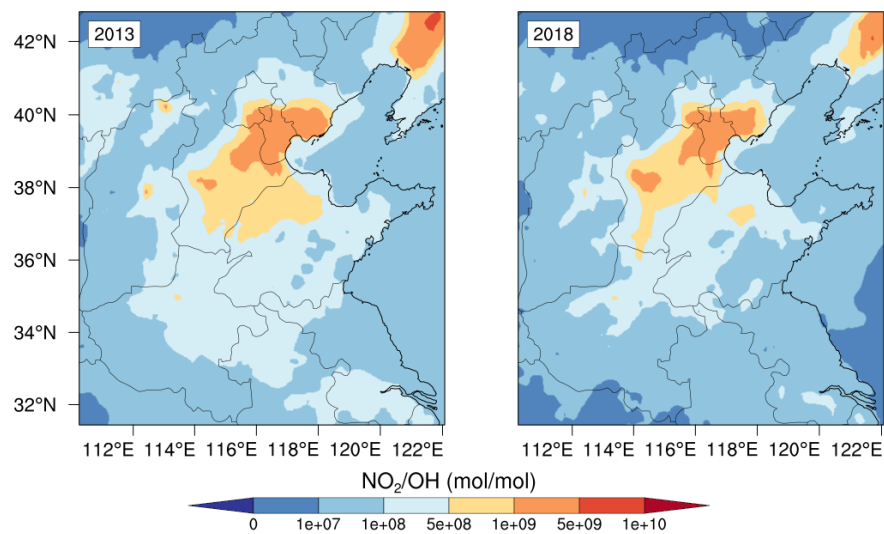


Figure S3. Spatial distribution of the NO₂/OH molar ratio in the NCP region during the winters of 2013 and 2018.

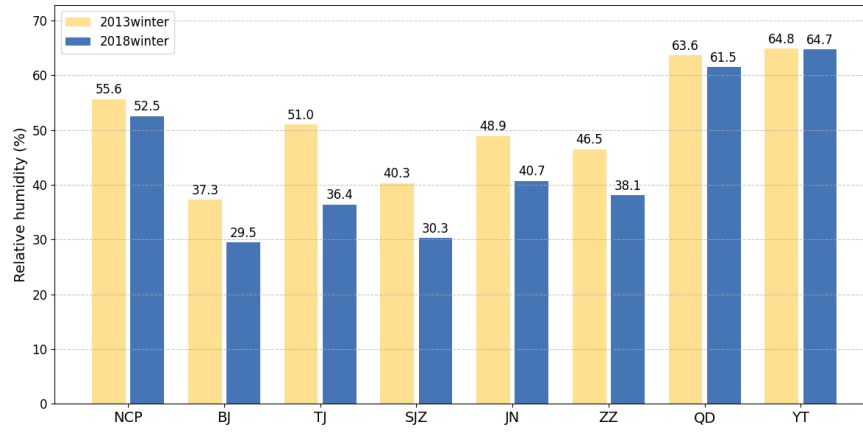


Figure S4. Relative humidity in the NCP and seven major cities (2013, 2018).

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