



# Meteorological influence on surface ozone trends in China: assessing uncertainties caused by multi-dataset and multi-method

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Received: 20 April 2025 – Discussion started: 19 May 2025

Revised: 26 August 2025 – Accepted: 9 September 2025 – Published: 27 October 2025

**Abstract.** China has witnessed notable increases in surface ozone ( $\text{O}_3$ ) concentrations since 2013, with meteorology identified as a critical driver. However, meteorological contributions vary with different meteorological datasets and analytical methods, and their uncertainties remain unassessed. This study leveraged decadal observational maximum daily 8-hour average  $\text{O}_3$  records (2013–2022) across China, revealing intensified nationwide  $\text{O}_3$  pollution with increasing  $\text{O}_3$  trends of 0.79–1.31 ppb  $\text{yr}^{-1}$  during four seasons. We gave special focus on uncertainties of meteorology-driven  $\text{O}_3$  trends by using diverse meteorological datasets (ERA5, MERRA2, FNL) and diverse analytical methods (Multiple Linear Regression, Random Forest, GEOS-Chem model). A useful statistic (coefficient of variation, CV) was adopted as an uncertainty quantification metric. For multi-dataset analysis, models driven by different meteorological datasets exhibited the maximum meteorology-driven  $\text{O}_3$  trend (+0.55 ppb  $\text{yr}^{-1}$ , multi-dataset mean) with the highest consistency (CV = 0.25) in spring. The FNL-driven model always obtained larger trends compared to ERA5 and MERRA2, which could be attributed to inability to accurately evaluate planetary boundary layer height in FNL dataset. For multi-method analysis, three methods demonstrated optimal consistency in winter (CV = 0.40) and the worst consistency in summer (CV = 2.00). The meteorology-driven  $\text{O}_3$  trends obtained from GEOS-Chem model were almost smaller than those obtained by other two methods, partly resulting from higher simulated  $\text{O}_3$  values before 2018. Overall, all analyses driven by diverse meteorological datasets and analytical methods drew a robust conclusion that meteorological conditions almost boosted  $\text{O}_3$  increases during all seasons; the uncertainties caused by different analytical methods were larger than those caused by diverse meteorological datasets.

## 1 Introduction

Since 2013, the Chinese government has implemented a series of policies to mitigate air pollution resulting from the rapid industrial and urban expansion, such as the “Air Pollution Prevention and Control Action Plan” (Wang, 2021). Several criteria air pollutants exhibited decreases due to the emission control efforts, but not ozone ( $\text{O}_3$ ) (Qi et al., 2023; Shen et al., 2020). In China,  $\text{O}_3$  concentrations were increased by  $50\text{--}124\text{ }\mu\text{g m}^{-3}$  from 2015–2022 (Yao et al., 2024). The formation of surface  $\text{O}_3$  depends nonlinearly on its precursors and is strongly influenced by meteorological conditions and anthropogenic emissions (Wang et al., 2017). The impact of emission-related factors on  $\text{O}_3$  increases in China over the past decade has been extensively debated, including the ineffective control of volatile organic compounds (VOCs) emissions, the heightened  $\text{O}_3$  photochemical production due to the rapid decrease in  $\text{PM}_{2.5}$ , and the reduced nitric oxide (NO) titration effect (Li et al., 2019, 2022; Lin et al., 2021a; Liu and Wang, 2020b; Ren et al., 2022).

Meteorological conditions also play a crucial role in shaping surface  $\text{O}_3$  trends (Liu et al., 2023; Lu et al., 2019b), resulting in increased  $\text{O}_3$  concentrations during warm seasons over most of the United States, the European Union, and China from 2014–2019 (Lyu et al., 2023). In China, the meteorological impacts on  $\text{O}_3$  levels may be comparable to the anthropogenic contributions (Li et al., 2020; Liu and Wang, 2020a). From 2013 to 2018, meteorology could account for 43 % of the daily variability in summer surface  $\text{O}_3$  concentrations in eastern China (Han et al., 2020). Adverse meteorological conditions were identified as the cause of the worsening  $\text{O}_3$  trends during 2015–2020 in Beijing–Tianjin–Hebei (BTH), Yangtze River Delta (YRD), and Pearl River Delta (PRD) regions (Hu et al., 2024b). In YRD, Dang et al. (2021) found that meteorological factors contributed 84 % of the  $\text{O}_3$  increase during the summers of 2012–2017. In PRD, meteorological conditions contributed 83 % of the increasing  $\text{O}_3$  trends during the summers of 2015–2019 (Mousavinezhad et al., 2021). After 2019, meteorological conditions tended to improve  $\text{O}_3$  air quality (Liu et al., 2023; Wang et al., 2023). Compared to 2019, the wetter and cooler meteorological conditions in 2020 reduced  $\text{O}_3$  concentrations by 2.9 ppb in eastern China (Yin et al., 2021). However, during 2022’s summer, a notable rebound in  $\text{O}_3$  levels occurred with  $\text{O}_3$  concentrations rising by 12–15 ppb in China compared to 2021, which was attributable to the extreme heatwave events (Qiao et al., 2024). With climate change, the frequency of extreme  $\text{O}_3$  pollution events is expected to increase (Gao et al., 2023; Ji et al., 2024). Given the shifted meteorological effects on  $\text{O}_3$  and climate change, it is imperative to conduct  $\text{O}_3$ –Meteorology research focusing on longer time frames to gain deeper insights into the long-term changes in  $\text{O}_3$  concentrations (Wang et al., 2024a).

Studies conducted over the past six years to determine the meteorological influence on the surface  $\text{O}_3$  trend have

been systematically reviewed, as documented in Table S1. The meteorological influence on surface  $\text{O}_3$  concentrations is commonly assessed by using the traditional statistical model (TSM), machine learning model (MLM), and chemical transport model (CTM), driven by reanalysis meteorological products such as the fifth-generation European Centre for Medium-Range Weather Forecasts atmospheric reanalysis of the global climate (ERA5), the Modern-Era Retrospective Analysis for Research and Applications, version 2 (MERRA2), and the National Centres for Environmental Prediction (NCEP) Final Operational Global Analysis data (FNL). Although several studies have demonstrated that meteorological impacts derived from CTM results can corroborate the results of TSM (Liu et al., 2023; Yan et al., 2024) or MLM (Ni et al., 2024; Yin et al., 2021), uncertainties in the determination of meteorological effects on surface  $\text{O}_3$  concentrations cannot be neglected. For example, Pan et al. (2023) reported that the meteorological impact on  $\text{O}_3$  trends in Beijing during 2013–2020 was  $+0.52\text{ ppb yr}^{-1}$ , which is only half of the value estimated by Gong et al. (2022).

Uncertainties in quantifying the drivers of  $\text{O}_3$  trends can be ascribed to the discrepancies between different meteorological datasets and between different methods (Guo et al., 2021; Weng et al., 2022; Yao et al., 2024). Regarding the uncertainty caused by different meteorological datasets, the meteorologically driven annual variations of  $\text{O}_3$  concentrations from 2017–2019 identified by the MERRA2-driven TSM are not consistent with the ERA5-driven TSM during the summer of YRD (Hu et al., 2024a; Qian et al., 2022). During the summers of 2013–2019 in YRD, Li et al. (2019) reported a trend of  $+0.7\text{ ppb yr}^{-1}$  in meteorology-driven  $\text{O}_3$  concentrations using the MERRA2-driven TSM, while the trend of Yan et al. (2024) was  $-0.3\text{ }\mu\text{g m}^{-3}\text{ yr}^{-1}$  using the ERA5-driven TSM. Regarding the uncertainty caused by different methods, the meteorology-driven  $\text{O}_3$  trend identified by MLM for 2019–2021 was 2.4 times larger than that identified by CTM based on the same meteorological dataset input (MERRA2) in the North China Plain (NCP) during summer (Wang et al., 2024a). In BTH, from 2021 to 2022, Luo et al. (2024) identified a negative meteorological contribution based on the ERA5-driven MLM, while Yan et al. (2024) suggested a positive contribution ( $+4.3\text{ }\mu\text{g m}^{-3}$ ) based on the ERA5-driven TSM during summer.

On the basis of the above-mentioned, large uncertainties caused by multi-dataset or multi-method exist in  $\text{O}_3$ –Meteorology analyses. However, available intercomparisons of  $\text{O}_3$  analyses mainly focused on predicting the  $\text{O}_3$  concentrations. For example, Wang et al. (2024b) and Weng et al. (2023) compared the differences in  $\text{O}_3$  concentration prediction caused by different datasets and models, respectively. The uncertainties in quantifying the meteorological contributions to  $\text{O}_3$  trends caused by multi-dataset and multi-method remain unassessed. In addition, previous studies have predominantly focused on summer  $\text{O}_3$  pollution, although re-

ports indicate an extension of the O<sub>3</sub> pollution season to winter and spring across major clusters in China (Cao et al., 2024a; Li et al., 2021) and an unfavourable meteorological impact on O<sub>3</sub> air quality in spring and winter in BTH (Luo et al., 2024). It is essential to conduct an intercomparison of meteorology-driven O<sub>3</sub> quantification using multi-dataset and multi-method across all seasons.

This study utilised 10-year (2013–2022) surface O<sub>3</sub> observations in China to investigate long-term O<sub>3</sub> trends and quantify the meteorological influence on O<sub>3</sub> trends using diverse meteorological datasets and analytical methods. Figure 1 shows the framework, and the main objectives were: (1) to assess uncertainties in identifying the meteorological influences caused by multi-dataset. This was achieved by employing the TSM (i.e. multiple linear regression, MLR) driven by different reanalysis meteorological products (i.e. ERA5, MERRA2, and FNL); (2) to assess uncertainties in identifying meteorological effects caused by multi-method. This was achieved by establishing three models corresponding to TSM (i.e. MLR), MLM (i.e. random forest, RF), and CTM (i.e. GEOS-Chem, GC), each driven by the MERRA2 product; (3) to calculate the mean of meteorology-driven O<sub>3</sub> trends driven by three datasets (multi-dataset mean) and three methods (multi-method mean), so as to derive relatively robust results.

Our paper is structured as follows: Sect. 2 briefly introduces the details of surface O<sub>3</sub> observations and different meteorological datasets, as well as the framework of three methods, namely MLR, RF, and GC. The quantification of the uncertainties in meteorology-driven O<sub>3</sub> trends caused by multi-dataset and multi-method is presented in Sec. 3. Section 4 concludes the paper. The findings of this study provide a scientific foundation for developing regional and seasonal strategies to mitigate and manage O<sub>3</sub> pollution in China.

## 2 Data and Methods

### 2.1 Surface O<sub>3</sub> and meteorological data sources

Hourly surface O<sub>3</sub> observations from over 1000 state-controlled stations operated by the China National Environmental Monitoring Centre from 2013–2022 were used to analyse the long-term O<sub>3</sub> trends across all seasons: spring (March–April–May), summer (June–July–August), autumn (September–October–November), and winter (December–January–February). The maximum daily 8 h average (MDA8) O<sub>3</sub> was calculated as an air quality indicator after filtering out abnormal data using the *z*-scores method. For detailed information on data quality control, refer to He et al. (2017).

In this study, we selected three widely used reanalysis products to assess the uncertainties caused by different meteorological datasets. Variables during 2013–2022 from ERA5, MERRA2, and FNL, as detailed in Table S2, were selected as meteorological inputs for building MLR models.

These reanalyses have spatial resolutions of 0.25° × 0.25°, 0.625° × 0.5°, and 1° × 1° on a global latitude-longitude grid, respectively. In Sect. 3.2, we also incorporate the NCEP FNL reanalysis product with a spatial resolution of 0.25° × 0.25° (FNL025) for the period 2016–2022 to explore the effect of spatial resolution on the analysis of uncertainties caused by multi-dataset.

### 2.2 Methods for obtaining long-term series and meteorological influence

#### 2.2.1 Kolmogorov–Zurbenko (KZ) filter

The KZ filter, known for its ability to extract low frequency signals from time series data and handle missing values, has been extensively applied to analysing air quality variations (Eskridge et al., 1997; Rao and Zurbenko, 1994; Wise and Comrie, 2005). This filter is particularly useful in studying variations in air quality over time. The original time series of air pollutants or meteorological variables ( $X(t)$ ) can be decomposed by the KZ filter into the following form:

$$X(t) = X_{ST}(t) + X_{SN}(t) + X_{LT}(t) \quad (1)$$

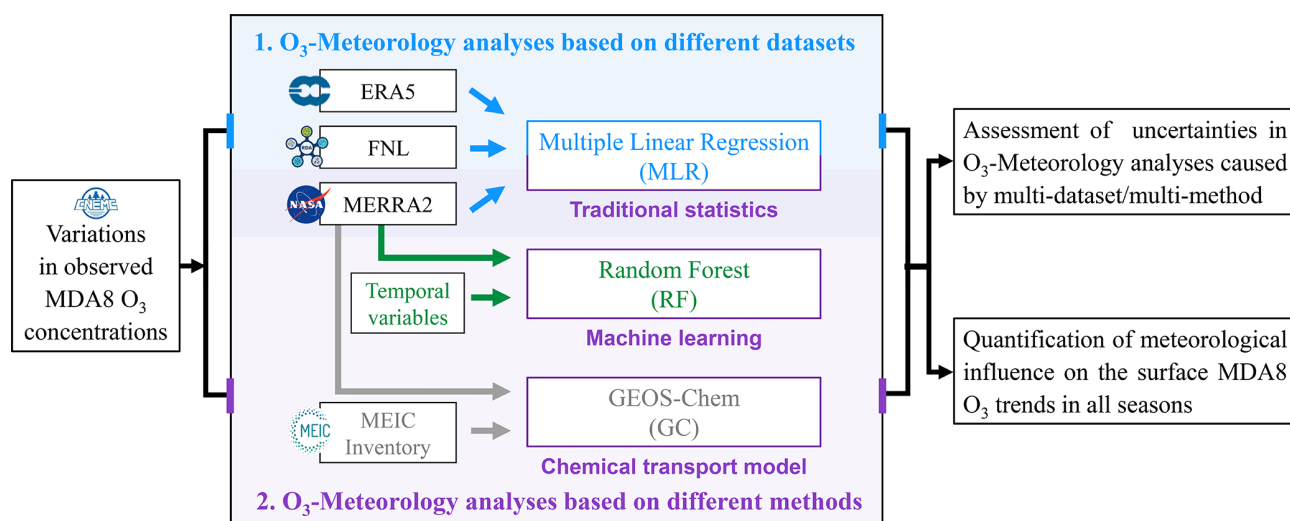
$$X_{LT}(t) = KZ_{(365,3)}X(t) \quad (2)$$

$$X_{BL}(t) = KZ_{(15,5)}X(t) \quad (3)$$

$$X_{ST}(t) = X(t) - X_{BL}(t) \quad (4)$$

In the decomposition process,  $X(t)$  represents the original daily time series, while  $X_{ST}(t)$ ,  $X_{SN}(t)$ , and  $X_{LT}(t)$  denote the short-term, seasonal, and long-term components, respectively. The baseline component,  $X_{BL}(t)$ , is defined as the sum of  $X_{SN}(t)$  and  $X_{LT}(t)$ . The  $KZ_{(p,q)}$  filter executes  $q$  iterations with  $p$  as the moving average window length of  $X(t)$  series. Specially,  $X_{LT}(t)$  is derived using the  $KZ_{(365,3)}$  filter, capturing long-term changes with periods exceeding 1.7 years.  $X_{BL}(t)$  is obtained through the  $KZ_{(15,5)}$  filter, encompassing both seasonal and long-term components.  $X_{ST}(t)$  represents short-term fluctuations with period less than 33 days in the original time series.  $X_{SN}(t)$  is derived as the difference between  $X_{BL}(t)$  and  $X_{LT}(t)$ , corresponding to seasonal variation on a timescale of months. The KZ filter can fill in missing values by using iterated moving average technique. Although not all of the ozone measurement sites were active over the entire period 2013–2022, missing value problems can be handled for most stations after we conduct three iterations with.

In this study, all statistical analyses were performed at the seasonal scale (spring: March–April–May; summer: June–July–August; autumn: September–October–November; winter: December–January–February). For each season, the  $KZ_{(365,3)}$  filter was applied to extract the long-term trends in observed, meteorology-driven, and emission-driven MDA8 O<sub>3</sub> concentrations (see details in Fig. S1) during 2013–2022, as detailed in Sect. 2.2.2, 2.2.3, and 2.2.4.



**Figure 1.** The framework of the uncertainty assessment in this study.

### 2.2.2 Stepwise MLR for separating meteorological influence

As vividly illustrated in Fig. S1, a data-based TSM (i.e., MLR integrating the KZ filter) was employed to separate the observed MDA8 O<sub>3</sub> concentrations into meteorology-driven and emission-driven concentrations (Sadeghi et al., 2022; Shang et al., 2023; Zhang et al., 2022a). We initially applied the KZ filter to disassemble the MDA8 O<sub>3</sub> time series and all meteorological variables listed in Table S2 into short-term, baseline, and long-term components at individual state-controlled stations for each season. Subsequently, a series of screening processes aligned with our previous research (Wang et al., 2024c) were executed to perform stepwise MLR on the short-term/baseline MDA8 O<sub>3</sub> concentrations and a group of meteorological variables series, respectively. The established MLR model is presented herein:

$$C_{s,r}(t) = b_{0,s,r} + \sum_{i=1}^k b_{i,s,r} \times \text{Met}_i(t) + \varepsilon \quad (5)$$

Here,  $C_{s,r}(t)$  represents the MDA8 O<sub>3</sub> concentration for season  $s$  and monitoring station  $r$ , while  $\text{Met}_i(t)$  signifies the  $i$ th meteorological variable out of a total of  $k$ , and  $b_{i,s,r}$  is the corresponding regression coefficient.  $b_{0,s,r}$  denotes the intercept term, and  $\varepsilon$  is the residual term. After establishing MLR models for the short-term and baseline components in each season, we obtain their respective residual terms. The total residuals, which represent the sum of residuals from baseline variables and short-term variables, primarily reflect anthropogenic influences. We then applied a KZ<sub>(365,3)</sub> filter to these aggregated residuals to derive long-term emission-driven and meteorology-driven O<sub>3</sub> variations. Finally, the meteorology-driven O<sub>3</sub> trends and emission-driven O<sub>3</sub> trends were obtained through Least Square Method.

The constructed MLR models driven by meteorological variables from ERA5, MERRA2, or FNL in each season will

allow a comprehensive analysis of multi-dataset uncertainties. The meteorological impact on O<sub>3</sub> trends derived from the MERRA2-driven MLR model will also be integrated into the analysis of multi-method uncertainties to improve the comparability of results.

### 2.2.3 Random forest (RF) for deriving meteorological influence

The application of MLM in O<sub>3</sub> air quality research is becoming increasingly prevalent due to its superior accuracy, user-friendly nature, and capability to capture nonlinear relationships (Ni et al., 2024; Yao et al., 2024; Zhang et al., 2022b). Considering the limited influence of discrepancy in O<sub>3</sub>-Meteorology analyses stemming from different machine learning algorithms (Wang et al., 2024a), we opted to build a representative MLM known as the meteorological normalisation model based on the RF algorithm (Ding et al., 2023; Ji et al., 2024; Zhang et al., 2023), to delineate meteorology- and emission-driven O<sub>3</sub> concentrations.

RF stands out as a tree-based ensemble learning algorithm adept at handling nonlinear issues and reducing overfitting (Breiman, 2001). An RF model was developed for each state-controlled station in each season to predict the MDA8 O<sub>3</sub> concentration using the Python package “Sklearn-RandomForestRegressor”. The predictors included six temporal variables (year, month of a year, day of a week, day of a month, day of a year, Unix time), serving as proxies for anthropogenic emission intensity (Grange et al., 2018), alongside six MERRA2 meteorological variables as listed in Table S2 (i.e. SLP, T2max, U10, V10, RH2, PBLHday). The training dataset comprised 70 % of the data, while the remaining 30 % was reserved for model evaluation. A statistical cross-validation technique was employed to determine optimal hyperparameters for enhancing RF prediction per-



formance (Weng et al., 2022). Coefficient of determination ( $R^2$ ) values were utilised to assess model performance for each station. Over 70 % of state-controlled stations showed  $R^2 \geq 0.5$  in all seasons (Fig. S2b), which is consistent with the 0.4–0.6 range reported in comparable studies (Weng et al., 2022; Lu et al., 2024). Stations with  $R^2 < 0.5$  were excluded to avoid significant attribution uncertainty that could be introduced by the RF performance. To evaluate the robustness of the  $R^2 \geq 0.5$  criterion, we performed sensitivity analyses using thresholds of  $R^2 \geq 0.6$  and  $R^2 \geq 0.4$ , to ensure that our conclusions are not an artifact of an arbitrary cutoff (Table S3).

After establishing the RF model, both the original time variables and resampled meteorological variables were utilised as input data. For meteorological normalisation, we implemented the protocol of Vu et al. (2019). Meteorological variables were resampled by randomly selecting data from the two weeks before and after the specified date, while temporal proxies remained fixed. To derive the de-weathered MDA8  $O_3$  concentration for a given day (e.g. 1 March 2013), the random resampling process was iterated 1000 times. The mean predicted  $O_3$  under average meteorological conditions, which refers to de-weathered  $O_3$ , corresponds to the emission-driven  $O_3$  concentration. The meteorology-driven MDA8  $O_3$  concentrations for each season were computed as the difference between observed concentrations and de-weathered concentrations. Detailed processes are shown in Fig. S2a. The  $KZ_{(365,3)}$  filter was then applied to obtain long-term components, and meteorology-driven  $O_3$  trends were derived using Least Square Method.

#### 2.2.4 GEOS-Chem (GC) simulation for quantifying meteorological influence

The numerical analysis of surface  $O_3$  in China was performed with the GC classic version 13.3.3 (<https://github.com/geoschem/GCClassic/releases/tag/13.3.3>, last access: 22 October 2025). Developed as a global 3-D model, GC incorporates a fully coupled  $O_3$ – $NO_x$ –VOCs–aerosol–halogen chemical mechanism, driven by the MERRA2 meteorological input. Numerous studies have leveraged GC to simulate  $O_3$  air quality in China, demonstrating alignment between observational data and model outcomes (Dai et al., 2024; Dang et al., 2021; Li et al., 2019; Lu et al., 2019a). We employed the nested-grid GC to simulate the long-term surface  $O_3$  concentrations and to quantify the meteorology-driven MDA8  $O_3$  trends over China. The nested-grid domain was set over China's mainland (15–55° N, 70–140° E) with a horizontal resolution of 0.5° latitude by 0.625° longitude and 47 vertical layers extending up to an altitude of 0.01 hPa. A global simulation with a horizontal resolution of  $2^\circ \times 2.5^\circ$  provided the chemical boundary conditions for the nested-grid simulation every 3 h. To ensure model stability and accuracy, a 6-month spin-up simulation was con-

ducted before the commencement of the targeted 10-year period from March 2013 to February 2023.

Emissions management within GC is facilitated by the Harmonized Emissions Component, a system introduced by Lin et al. (2021b). Anthropogenic emissions are sourced from the Community Emissions Data System (CEDS) inventory globally, with specific overwriting by the Multi-resolution Emission Inventory for China (MEIC) within the Chinese region. The simulations for 2021–2022 adopt a similar approach to Zhai et al. (2021), using 2019 MEIC emissions with  $NO_x$  emissions reduced by 8 %–13 % and 2017 MEIC with VOCs emissions reduced by 10 %–14 %, based on the policy released by Ministry of Ecology and Environment of the People's Republic of China. For natural emissions, biogenic VOCs, soil, and lightning  $NO_x$  were calculated online in the model. Emissions from biomass burning, ships, and aircraft are sourced from the Global Fire Emissions Database, the CEDS inventory, and the 2019 Aircraft Emissions Inventory Code, respectively.

In order to assess the model's performance and to get a quantification of meteorology-driven  $O_3$  trends during the period of 2013–2022, two sets of simulations were conducted: (1) BASE: the standard simulation of  $O_3$  concentrations from 2013 to 2022, where both meteorological fields and emissions (including anthropogenic, natural, and biomass emissions) vary year by year from 2013 to 2022; (2) FixE2013: a “fixed-emission simulation” where meteorological conditions vary from 2013 to 2022 while anthropogenic emissions remain constant at 2013 levels. The FixE2013 simulation is designed to quantify the meteorological influence on  $O_3$  variations. The FixE2013 simulation is designed to obtain the MDA8  $O_3$  concentrations driven solely by meteorological changes and further quantify the meteorological influence on  $O_3$  variations in four seasons. After applying the  $KZ_{(365,3)}$  filter to derive the long-term meteorology-driven series, trends were calculated through Least Square Method. Figure S3 evaluates the performance of the GC simulation for 2013–2022. The GC model generally captures the monthly variability in MDA8  $O_3$  over China and three megacity clusters, with the correlation coefficients greater than 0.80, although it always shows a high bias of surface  $O_3$  in warm seasons (Dai et al., 2024), which can be attributed to its inability to capture the complex terrain, local pollution sources and meteorological conditions, or overestimates of the correlations between the surface  $O_3$  concentration and temperature (Shen et al., 2022; Sun et al., 2021).

#### 2.3 Assessment of uncertainties caused by multi-dataset and multi-method

In this study, the coefficient of variation (CV) is applied to assess the uncertainties in  $O_3$ –Meteorology analyses caused by different meteorological datasets or methods. The CV, calculated as the ratio of the standard deviation (SD) to the mean, serves as a statistical metric commonly utilised to

measure the diversity within datasets or models (Bedeian and Mossholder, 2000; Chen et al., 2019). Compared to other comparators (e.g. range, inter-quartile range, and SD), the CV is a unit-free measure that quantifies percentage variation relative to the mean and is less sensitive to outliers and heavy-tailed distributions (Högel et al., 1994; Chattamvelli and Shanmugam, 2023). In this study, higher CVs indicate lower consistency of meteorologically driven O<sub>3</sub> trends derived from different datasets or methods. To give a more quantitative assessment, consistency levels were classified as strong and weak with  $CV < 0.5$  and  $CV > 1.0$ , respectively (Wang et al., 2022a). Given the possibility of disparate meteorology-driven O<sub>3</sub> trends detected by different datasets or methods, we consider the absolute value of the CV as a quantitative indicator of the uncertainties. For each season, when examining the uncertainties arising from different datasets, the CV represents the SD of trends derived from the ERA5, MERRA2, and FNL-driven MLR models divided by the mean. Similarly, in the context of multi-method uncertainties, the CV is the SD of trends identified by the MLR, RF, and GC models divided by the mean.

## 3 Results

### 3.1 Observed trends in surface O<sub>3</sub> concentration

Figure 2 shows the trends in observed MDA8 O<sub>3</sub> concentrations over a 10-year period during four seasons. Noteworthy increases in O<sub>3</sub> concentrations were observed at 78 %–93 % of state-controlled stations over the years, with the national trend being  $+1.31$ ,  $+0.93$ ,  $+0.79$ , and  $+0.80$  ppb yr<sup>−1</sup> in spring, summer, autumn, and winter, respectively.

The major eastern megacity clusters in China also displayed their highest MDA8 O<sub>3</sub> increase trends in spring, with trends of  $+1.16$  ppb yr<sup>−1</sup> in BTH,  $+1.61$  ppb yr<sup>−1</sup> in YRD, and  $+1.48$  ppb yr<sup>−1</sup> in PRD, which has been reported in previous studies (Cao et al., 2024b; Chen et al., 2020; Wang et al., 2022b). During summer, BTH and YRD faced more severe challenges in O<sub>3</sub> prevention and control compared to PRD, with rising MDA8 O<sub>3</sub> trends in the former two regions being about three times higher than that in PRD (Fig. 2b).

In terms of O<sub>3</sub> growth rates, Shanxi province and Anhui province ranked the top two provinces in China over the past decade in all seasons except for winter, consistent with Zhao et al. (2020). In spring and winter, O<sub>3</sub> concentrations increased in all provinces, with trends of  $+0.39$  to  $+2.75$  ppb yr<sup>−1</sup> and  $+0.42$  to  $+1.30$  ppb yr<sup>−1</sup>, respectively. Notably, Jilin province experienced an obvious improvement in O<sub>3</sub> air quality during summer and autumn, with decreasing trends of  $−0.74$  and  $−0.38$  ppb yr<sup>−1</sup>, respectively, which was also confirmed by Gong et al. (2022). As mentioned in Section 1, variations in O<sub>3</sub> concentrations are fundamentally modulated by emissions and meteorology. This section mainly documents observed O<sub>3</sub> trends, and the quantitative

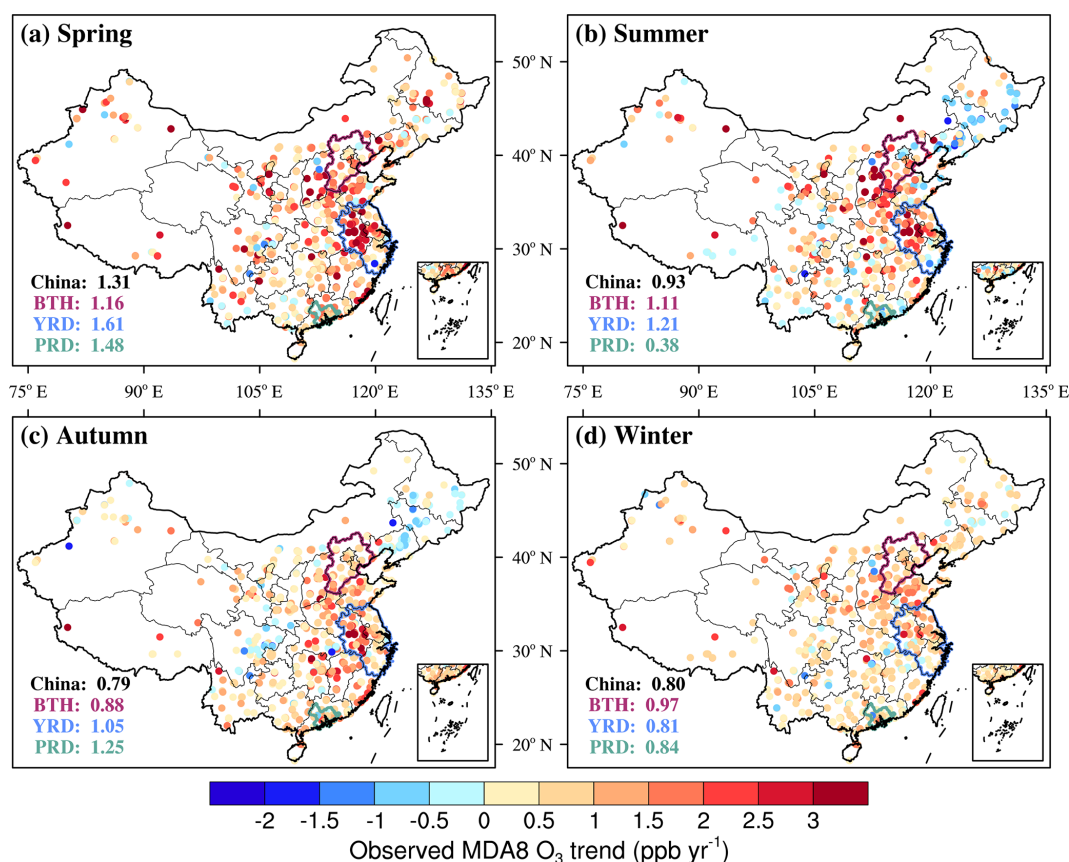
contributions of emissions and meteorology to MDA8 O<sub>3</sub> variations will be discussed in Sect. 3.2.

The annual and seasonal mean MDA8 O<sub>3</sub> concentrations across China are detailed in Figs. S4 and S5, providing a holistic depiction of the persisting spread of O<sub>3</sub> pollution since 2013. On a national average, the O<sub>3</sub> air quality was worst in summer, with the average O<sub>3</sub> levels exceeding the air quality standard Grade I limit of 50 ppb almost every year. Notably, the summer of 2019 marked a peak period for O<sub>3</sub> pollution, with an average concentration of 59.7 ppb (Fig. S5b).

### 3.2 Uncertainty in meteorology-driven O<sub>3</sub> trends caused by multi-dataset

The traditional statistical method (the MLR model), which has a relatively low computational cost but can provide valuable insights into the quantification of meteorological contributions to O<sub>3</sub> trends, was used to investigate the uncertainties in O<sub>3</sub>–Meteorology analyses caused by different meteorological datasets. As shown in Fig. 3a, meteorological conditions contribute to an increase in MDA8 O<sub>3</sub> concentrations across all seasons in China, with the multi-dataset mean trends ranging from  $+0.19$  ( $\pm 0.47$ ) ppb yr<sup>−1</sup> to  $+0.55$  ( $\pm 0.45$ ) ppb yr<sup>−1</sup>. All three dataset-driven MLR models indicate that meteorology leads to the most rapid increase in MDA8 O<sub>3</sub> concentrations in spring, with trends ranging from  $+0.47$  ( $\pm 0.47$ ) ppb yr<sup>−1</sup> to  $+0.71$  ( $\pm 0.59$ ) ppb yr<sup>−1</sup>, and a low CV of 0.25. This suggests a high consistency among the three datasets in assessing the meteorological influence on surface O<sub>3</sub> concentrations. During summer and autumn, meteorological influences on O<sub>3</sub> show greater spatial heterogeneity (with higher SD) and larger variability among multi-datasets (with higher CV). Specifically in autumn, the meteorology-driven O<sub>3</sub> trend derived from the FNL-driven MLR model is 4.1 times larger than that derived from the ERA5-driven MLR model. Lu et al. (2024) compared meteorology-driven O<sub>3</sub> trends derived from ERA5- and MERRA2-driven MLR models during the summers of 2013–2019. Their findings revealed that ERA5-derived trends were lower than those from MERRA2 in YRD and PRD, whereas trends derived from ERA5 were comparable to those from MERRA2 in BTH. This inter-study consensus further validates the robustness of our methodological framework.

Figure 3b–d depicts the meteorological impact on the MDA8 O<sub>3</sub> trends in the three megacity clusters (BTH, YRD, and PRD). Meteorology caused the MDA8 O<sub>3</sub> increase in most of the megacity clusters and seasons, except for BTH during autumn. In seasons where the meteorological effects derived from the three MLR models are all positive, the multi-dataset mean trends ranged from  $+0.09$  ( $\pm 0.38$ ) to  $+0.33$  ( $\pm 0.13$ ) ppb yr<sup>−1</sup> in BTH,  $+0.18$  ( $\pm 0.20$ ) to  $+0.68$  ( $\pm 0.56$ ) ppb yr<sup>−1</sup> in YRD, and  $+0.73$  ( $\pm 0.36$ ) to  $+1.13$  ( $\pm 0.45$ ) ppb yr<sup>−1</sup> in PRD. Consistent with Fig. 3a, meteo-



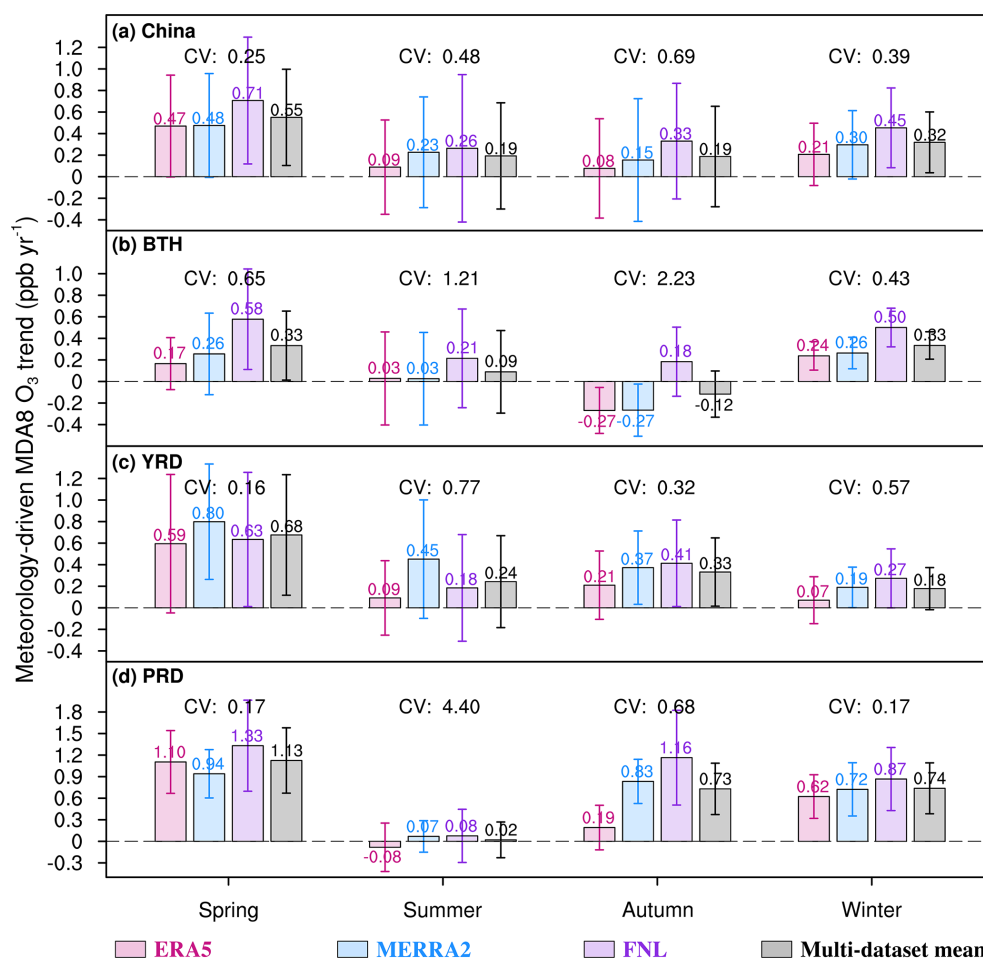
**Figure 2.** Trends in observed MDA8 O<sub>3</sub> concentrations in China from 2013–2022 during (a) spring, (b) summer, (c) autumn, and (d) winter. Values in black, purple, blue, and green represent the mean trends for the whole China, BTH, YRD, and PRD, respectively.

rology triggered the most rapid increase in MDA8 O<sub>3</sub> concentrations in spring across the three megacity clusters. The largest meteorological impact in BTH during spring was also revealed by Luo et al. (2024). Large CVs ( $> 1.0$ ) were observed in BTH during summer and autumn. Notably, the meteorological influence calculated by the three dataset-driven MLR models even showed opposite trends in BTH during autumn, indicating challenges in assessing the meteorological impacts on surface O<sub>3</sub> concentrations. In contrast, in YRD and PRD, the three MLR models demonstrated high consistency across almost all seasons. Although the largest CV reached 4.40 in PRD during summer, it was considered acceptable because the three MLR models indicated that meteorology had a minor influence (less than  $+0.1$  ppb yr<sup>-1</sup>) on O<sub>3</sub> trends.

From a provincial perspective in Fig. S6, we can also see that the meteorological contributions to O<sub>3</sub> trends are positive during spring and winter. Large uncertainties in O<sub>3</sub>-Meteorology analyses were identified during summer and autumn. There were 7 and 12 provinces with controversial meteorological contributions identified by the three dataset-driven MLR models in summer and autumn, respectively.

Figure 4 displays the spatial distribution of the CV values from the perspective of state-controlled stations in four seasons. Consistent with the national and provincial perspectives, the least uncertainties in O<sub>3</sub>-Meteorology analyses were observed in spring, with CVs less than 0.5 at 45 % of stations. Obvious discrepancies in meteorology-driven O<sub>3</sub> trends are found in summer and autumn, particularly in the NCP and northwestern China, with CVs greater than 1.0 at 33 %–40 % of the stations. In autumn, it is noteworthy that the uncertainties caused by multi-dataset are lower in the south than in the north. Previous studies that employed MLR models to predict O<sub>3</sub> concentration also revealed that the MLR had better performance in the south than in the north (Han et al., 2020; Lu et al., 2024).

Based on the three dataset-driven MLR models, the meteorological and anthropogenic contributions to the MDA8 O<sub>3</sub> trends in China during 2013–2022 were further examined. As presented in Fig. 5, both meteorological conditions and anthropogenic emissions lead to O<sub>3</sub> increases. According to the ERA5- and MERRA2-driven MLR models, variations in anthropogenic emissions were identified as the dominant factor driving the increase in MDA8 O<sub>3</sub> concentrations across all seasons, with anthropogenic contributions ranging from



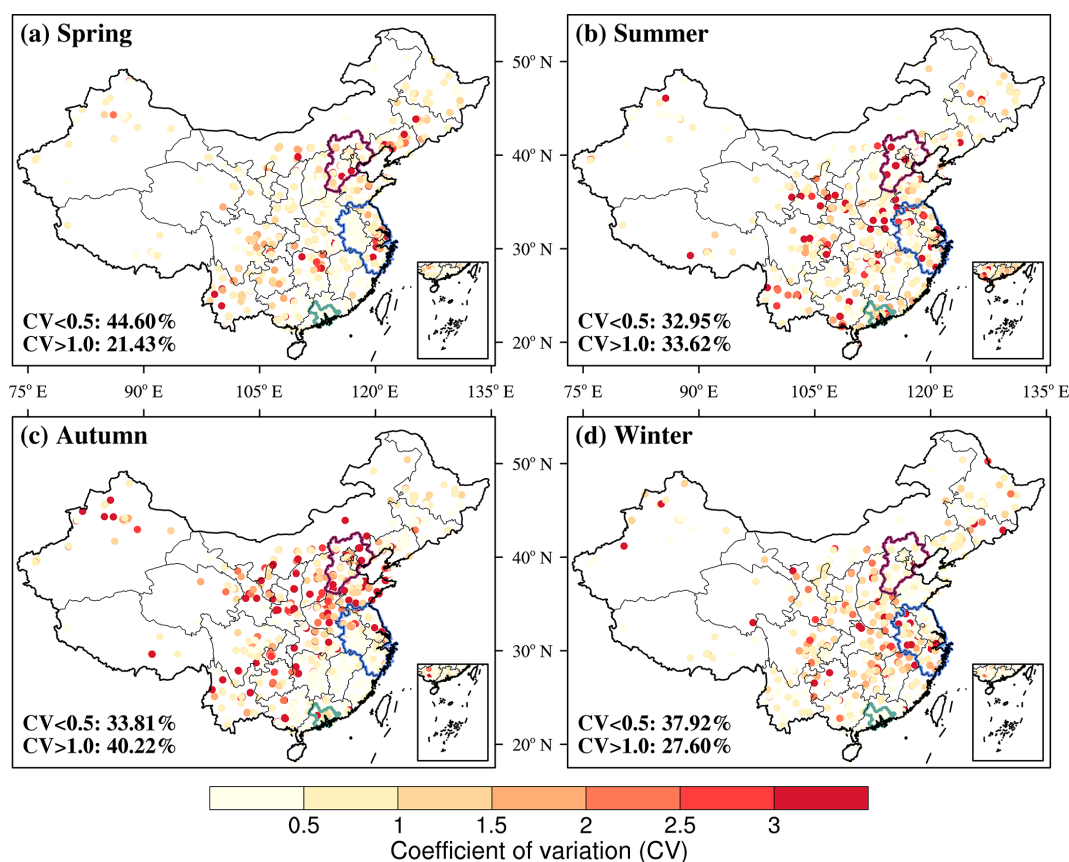
**Figure 3.** Meteorology-driven MDA8 O<sub>3</sub> trends in (a) the whole China, (b) BTH, (c) YRD, and (d) PRD during four seasons. Values in red, blue, and purple represent trends calculated by ERA5-, MERRA2-, and FNL-driven multiple linear regression (MLR) models, respectively. The fourth black bar represents the multi-dataset mean trend. Error bars indicate  $\pm 1$  standard deviation (SD) of site-level trends calculated from all available monitoring stations within each region. The absolute value of the coefficient of variation (CV) for each season is also shown.

63.2 % to 90.4 %. The results suggest that more stringent emission control policies should be implemented to counteract the adverse effects of meteorological influences on O<sub>3</sub> concentrations.

It is interesting to note that the FNL-driven model almost always gave relatively larger predictions of meteorologically driven O<sub>3</sub> trends compared to the models driven by ERA5 and MERRA2. To investigate whether this discrepancy was due to the coarser spatial resolution of the FNL dataset, a comparison was made between the FNL025-driven MLR model ( $0.25^\circ \times 0.25^\circ$ ) and the FNL-driven MLR model ( $1.0^\circ \times 1.0^\circ$ ). As depicted in Fig. S7, the deviation of the meteorology-driven trends calculated by the two MLR models was less than 0.1 in China and three megacity clusters across four seasons, indicating that different spatial resolutions have little effect on O<sub>3</sub>-Meteorology analyses. Further examination was conducted to assess the influence of meteorological variables on O<sub>3</sub>-Meteorology analyses. Table S4

lists the 10-year trends in each meteorological factor and shows a great discrepancy in the variable “PBLHday”. Zuo et al. (2023) also reported that FNL exhibited the highest uncertainty for the evaluation of PBLH compared to ERA5 and MERRA2, and that its performance may be constrained by complex underlying terrain and static instability (Guo et al., 2021). As Fig. S8 shows, constructing the FNL-driven MLR models using six meteorological variables without “PBLHday” can reduce the estimated meteorological impact by 0.08 to 0.20 ppb yr<sup>-1</sup>. To obtain a more reliable estimate, it is recommended to use MERRA2 reanalysis dataset due to its eclectic result (Fig. 3) and avoid using FNL because of the uncertainty brought by PBLH when separating meteorological and anthropogenic influences on O<sub>3</sub> concentrations in China.



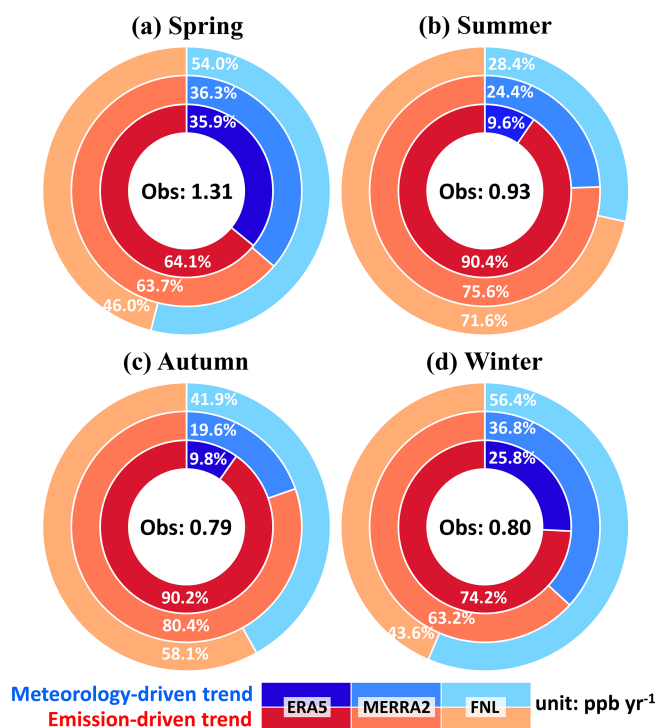


**Figure 4.** The absolute value of the coefficient of variation (CV) for each state-controlled monitoring station in China during (a) spring, (b) summer, (c) autumn, and (d) winter. The CV is calculated by the standard deviation (SD) of the trends derived from ERA5-, MERRA2-, and FNL-driven MLR models divided by the mean. The darker colour means the larger uncertainty in quantifying the meteorological impact on observed O<sub>3</sub> trends. The proportion of state-control stations with CV less than 0.5 and greater than 1.0 is also shown. The outline marked in purple, blue, and green represents the region of BTH, YRD, and PRD, respectively.

### 3.3 Uncertainty in meteorology-driven O<sub>3</sub> trends caused by multi-method

This section discusses the uncertainties caused by multi-method (i.e. MLR, RF, GC), all of which are driven by the MERRA2 dataset. Figure 6 illustrates the meteorology-driven MDA8 O<sub>3</sub> trends calculated by the MLR, RF, and GC models. For the whole of China, the large uncertainties are evident during summer, when the meteorology-driven O<sub>3</sub> trends derived from the MLR model are notably larger than those from the RF and GC models, with a CV of 2.00 (Fig. 6a). In the other three seasons, the multi-method mean trends, ranging from  $+0.17 (\pm 0.37)$  to  $+0.26 (\pm 0.27)$  ppb yr<sup>-1</sup>, are 1.1 to 2.1 times lower than those computed by the three dataset-driven MLR models (Fig. 3a), all models converge on the conclusion that meteorological conditions contribute to the deterioration of O<sub>3</sub> air quality. Meteorology-driven MDA8 O<sub>3</sub> trends exhibited minor variations across different  $R^2$  thresholds (Table S3), indicating that the trends are not an artifact of an arbitrary cutoff.

In YRD and PRD, the three models exhibit strong agreement across all seasons, with a maximum CV of 0.61. The low uncertainties are further corroborated by consistent CV estimates derived under different RF's  $R^2$  thresholds (Table S3). Across these regions, where meteorology leads to an increase in O<sub>3</sub> concentrations with multi-method mean trends of  $+0.17 (\pm 0.08)$  to  $+0.47 (\pm 0.22)$  ppb yr<sup>-1</sup> in YRD and  $+0.10 (\pm 0.12)$  to  $+0.83 (\pm 0.19)$  ppb yr<sup>-1</sup> in PRD. Notably, the most rapid meteorology-driven O<sub>3</sub> increase is also observed in spring (Fig. 6c and d), which is consistent with Fig. 3c and d. Lu et al. (2024) also demonstrated a high degree of consistency among the MLR, ML, and GC models in PRD during summer. Specifically, all three models indicated that meteorology contributed approximately 25 % of O<sub>3</sub> variability over the period 2013–2019. In BTH, the three models perform consistently well only in winter, with meteorology-driven O<sub>3</sub> trends ranging from  $+0.09 (\pm 0.07)$  to  $+0.26 (\pm 0.15)$  ppb yr<sup>-1</sup> and a CV of 0.55. It is also observed that in summer and autumn, meteorology plays a relatively small role in influencing O<sub>3</sub> air quality despite the



**Figure 5.** Percentage contributions of meteorological conditions (blue) and anthropogenic emissions (red) to the trends in observed MDA8 O<sub>3</sub> concentrations calculated by ERA5-, MERRA2-, and FNL-driven multiple linear regression (MLR) models in China during (a) spring, (b) summer, (c) autumn, and (d) winter. Values in black represent the observed MDA8 O<sub>3</sub> trends averaged over the whole China.

controversial results obtained by the three models (Fig. 6b). This finding aligns with a study focusing on the O<sub>3</sub> air quality in BTH from 2015 to 2022 (Luo et al., 2024), which suggested that meteorological conditions tend to increase MDA8 O<sub>3</sub> concentration by only 0.01 µg m<sup>-3</sup> in summer and decrease MDA8 O<sub>3</sub> concentration by 0.3 µg m<sup>-3</sup> in autumn from 2015–2022.

In addition, Fig. 6 illustrates that the meteorology-driven O<sub>3</sub> trends obtained from GC are relatively smaller. As shown in Fig. S3 and Table S5, this difference could partly be attributed to the higher O<sub>3</sub> levels and lower O<sub>3</sub> increases simulated by the GC model before 2018. The GC's systematic overestimation of O<sub>3</sub> concentrations, as well as underestimation of O<sub>3</sub> increases, was also reported by Lu et al. (2024), in which the GC captured 13.6 %–81.1 % of the observed O<sub>3</sub> increases in China during the summer of 2000–2019. It is crucial to take into account the overestimation of low-level O<sub>3</sub> observations, as noted in previous studies (Hu et al., 2024c; Mao et al., 2024). To validate this hypothesis, we compared the meteorology-driven O<sub>3</sub> trends calculated by MLR with those calculated by GC from 2018–2022, and a higher agreement was found over 2018–2022 compared to the 2013–2022 period in Fig. S9. The trends driven by RF model are eclectic

in more cases (Fig. 6) and recommended to isolate meteorological and anthropogenic drivers.

From a provincial perspective, as depicted in Fig. S10, the three models together indicate that meteorology causes an O<sub>3</sub> increase in winter across almost all provinces except for Guizhou and Sichuan. In summer and autumn, meteorology leads to a decrease in 5 provinces, mainly in northeastern China, with trends ranging from −0.42 to −0.11 ppb yr<sup>-1</sup>. Interestingly, across all seasons, the three models introduce less uncertainty in the developed east coast regions such as Jiangsu, Fujian, and Guangdong compared to other provinces. This suggests that quantifying meteorological impact on O<sub>3</sub> levels in these developed regions along the east coast of China is relatively reliable.

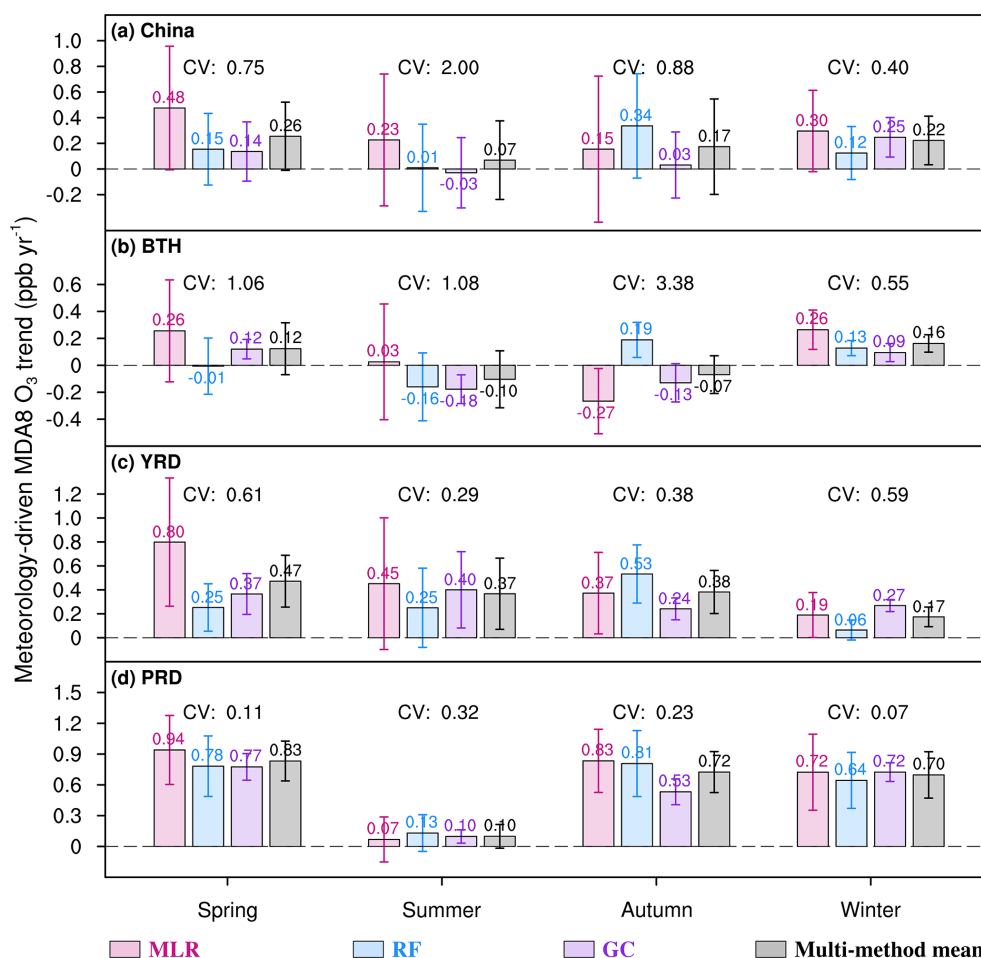
From the perspective of state-controlled stations, Fig. 7 shows the spatial distribution of the CV during four seasons. The lowest disparities in the meteorology-driven MDA8 O<sub>3</sub> trends persist in winter, with CVs of less than 0.5 recorded at 29 % of the stations. In the other three seasons, however, significant discrepancies in meteorology-driven O<sub>3</sub> trends are prominent, with CVs greater than 1.0 at least 48 % of the stations. Similar to Fig. 4, it is noteworthy that in autumn, the uncertainties caused by multi-method are more pronounced in the northern regions compared to the southern regions.

## 4 Limitations

While this study advances understanding of meteorological contributions to O<sub>3</sub> trends, several limitations warrant attention in future work. Though the reanalysis meteorological dataset is generated observationally, inherent constraints exist, including parameterization uncertainties affecting O<sub>3</sub>-relevant physical processes (Janjić et al., 2018; Davidson and Millstein, 2022) and resolution constraints.

Regarding analytical approaches, machine learning efficiently captures nonlinear O<sub>3</sub>-meteorology relationships without requiring explicit physicochemical parameterizations, enabling scalable multi-site analysis. However, its inability to resolve chemical mechanisms and sensitivity to predictor selection remain key constraints. Conversely, while GC mechanistically resolves chemistry-transport interactions and enables source attribution, it propagates uncertainties from emission inventories and chemical mechanisms into trend estimates.

Future studies could be improved in the following ways: First, more meteorological datasets and methods should be used to provide more robust uncertainty quantification in O<sub>3</sub>-meteorology analyses. Second, implementing clustering techniques (e.g. K-means algorithm) could identify sub-regional drivers at ecotones, enhancing spatial resolution beyond our regional framework. Finally, the Lindeman-Merenda-Gold indices can be employed to quantitatively resolve the contributions of specific meteorological variables. The mechanistic understanding of O<sub>3</sub> drivers would be im-



**Figure 6.** Meteorology-driven MDA8 O<sub>3</sub> trends in (a) the whole China, (b) BTH, (c) YRD, and (d) PRD during four seasons. Values in red, blue, and purple represent trends calculated by multiple linear regression (MLR), random forest (RF), and GEOS-Chem (GC) models, respectively. The fourth black bar represents the multi-method mean trend. Error bars indicate  $\pm 1$  standard deviation (SD) of site-level trends calculated from all available monitoring stations within each region. The absolute value of the coefficient of variation (CV) for each season is also shown.

proved by integrating additional variables, such as solar radiation, soil moisture, and climate indices (e.g. El Niño–Southern Oscillation). Clustering techniques would be valuable to augment the region-based approach and would provide better understanding of the similarity between stations.

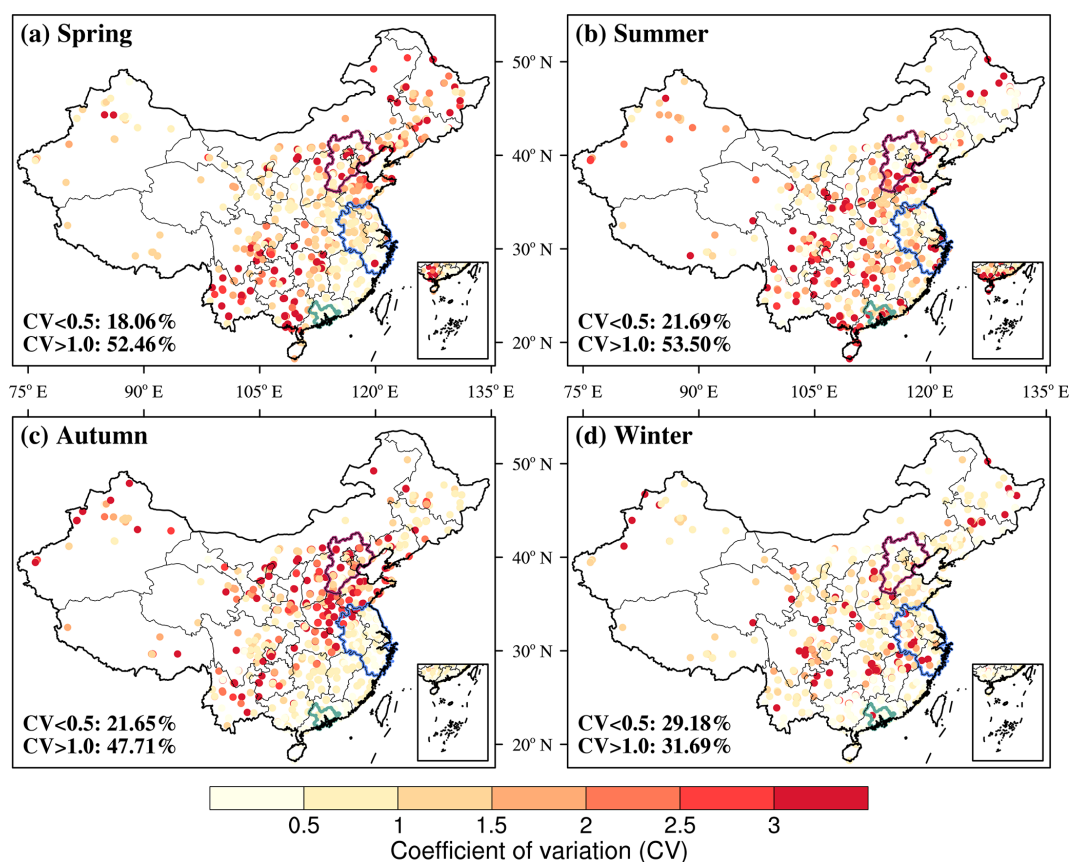
## 5 Conclusions and discussions

This study used the 10-year (2013–2022) surface O<sub>3</sub> observations to clarify O<sub>3</sub> variations during four seasons in China, and quantify the meteorological impacts on O<sub>3</sub> trends, with a special focus on the uncertainties of meteorology-driven O<sub>3</sub> trends. Diverse meteorological datasets (ERA5, MERRA2, FNL) and analytical methods (MLR, RF, GEOS-Chem) were employed to systematically analyse the uncertainties in meteorology-driven O<sub>3</sub> trends caused by multi-dataset and multi-method which have not been assessed before. The coefficient of variation (CV) was adopted as a met-

ric to assess the uncertainty. The main conclusions are as follows:

Over the past decade, increasing trends in MDA8 O<sub>3</sub> were observed at over 78 % of state-controlled stations across all seasons, with the national trend of +1.31, +0.93, +0.79, and +0.80 ppb yr<sup>-1</sup> in spring, summer, autumn, and winter, respectively.

We first applied the MLR model (driven by ERA5, MERRA2, and FNL, respectively), which has proven its usefulness and reliability in O<sub>3</sub>–Meteorology analyses, to assess uncertainties caused by multi-dataset. For the whole China, all three dataset-driven MLR models indicate that meteorological conditions have led to an increase in MDA8 O<sub>3</sub> concentrations in four seasons, with multi-dataset mean trends ranging from +0.19 to +0.55 ppb yr<sup>-1</sup>. The models driven by different meteorological datasets showed a maximum meteorology-driven O<sub>3</sub> trend of +0.55 ppb yr<sup>-1</sup> with the highest consistency (CV = 0.25) in spring. The FNL-driven



**Figure 7.** The absolute value of the coefficient of variation (CV) for each state-controlled monitoring station in China during (a) spring, (b) summer, (c) autumn, and (d) winter. The CV is calculated by the standard deviation (SD) of the trends derived from multiple linear regression (MLR), random forest (RF), and GEOS-Chem (GC) models divided by the mean. The darker colour means the larger uncertainty in quantifying the meteorological impact on observed  $O_3$  trends. The proportion of state-control stations with CV less than 0.5 and greater than 1.0 is also presented. The outline marked in purple, blue, and green represents the region of BTH, YRD, and PRD, respectively.

model always obtained larger meteorology-driven  $O_3$  trends compared to the models driven by ERA5 and MERRA2, which could be attributed to the inability to accurately evaluate PBLH in the FNL dataset. The dominant influence of anthropogenic emissions on  $O_3$  increase was also identified, highlighting the need for more stringent emission control policies to mitigate the adverse effects of meteorological conditions.

We further applied the MLR, RF, and GEOS-Chem models to obtain the meteorological influence on  $O_3$  trends to explore the uncertainties caused by multi-method. In China and three megacity clusters, the three methods consistently indicated positive meteorological contributions to  $O_3$  increases during spring and winter, with multi-method mean trends ranging from  $+0.12$  to  $+0.83$  ppb  $yr^{-1}$  and  $+0.17$  to  $+0.70$  ppb  $yr^{-1}$ , respectively. In summer and autumn, especially in BTH, where the meteorological influence was relatively lower, three methods gave conflicting predictions of meteorological influence on  $O_3$  with CVs greater than 1.08. For the whole China, three different methods demon-

strated optimal consistency in winter with a CV of 0.40 and the worst consistency in summer with a CV of 2.00. The meteorology-driven  $O_3$  trends obtained from GEOS-Chem model were almost relatively smaller than those obtained by other two methods, which could partly be attributed to the higher  $O_3$  values simulated by the GEOS-Chem model before 2018.

All analyses driven by diverse meteorological datasets and analytical methods drew a consistent finding: meteorological conditions almost contribute to  $O_3$  increase across all seasons. The uncertainties of meteorology-driven  $O_3$  trends caused by different analytical methods were larger than those caused by diverse meteorological datasets. Considering that the favourable effects of meteorology on  $O_3$  pollution tend to be weaker after 2019 and the effects of COVID-19, it is necessary to conduct research over different periods and longer periods. In addition, further research is needed to focus on the meteorological contributions to  $O_3$  trends in northern China due to larger uncertainties.



**Data availability.** The surface O<sub>3</sub> observations are obtained from <https://softnet.net/air/> (last access: 22 October 2025). The ERA5, MERRA2, FNL, and FNL025 reanalysis meteorological data are taken from <https://cds.climate.copernicus.eu/datasets> (last access: 22 October 2025), [http://geoschemdata.wustl.edu/ExtData/GEOS\\_0.5x0.625/MERRA2/](http://geoschemdata.wustl.edu/ExtData/GEOS_0.5x0.625/MERRA2/) (last access: 22 October 2025), <https://rda.ucar.edu/datasets/d083002/> (last access: 22 October 2025), and <https://rda.ucar.edu/datasets/ds083-3/> (last access: 22 October 2025), respectively. The code of the GEOS-Chem (version 13.3.3) model is available at <https://zenodo.org/records/5748260> (The International GEOS-Chem User Community, 2021). The MDA8 O<sub>3</sub> observations and analytical results derived from MLR, RF, and GEOS-Chem can be obtained from <https://doi.org/10.5281/zenodo.15859027> (Wang and Zhu, 2025).

**Supplement.** The supplement related to this article is available online at <https://doi.org/10.5194/acp-25-13863-2025-supplement>.

**Author contributions.** JZ designed the research, and XW performed simulations and analysed the results. GJ, XC, and ZY helped with the model simulations. LC, XJ, and HL provided useful comments on the paper. XW and JZ wrote the paper with contributions from all co-authors.

**Competing interests.** The contact author has declared that none of the authors has any competing interests.

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**Acknowledgements.** The author Xueqing Wang would like to thank Zhenze Liu from Nanjing University of Information Science and Technology, and Yongqi Zhao from Institute of Atmospheric Physics, Chinese Academy of Sciences, for their help in analysing the research results. We also thank the anonymous reviewers for their constructive suggestions that helped improve the paper.

**Financial support.** This research has been supported by National Key Research and Development Program of China (no. 2022YFE0136100), National Nature Science Foundation of China (nos. 42293323, 42007195, 42305121), and State Environmental Protection Key Laboratory of Sources and Control of Air Pollution Complex (no. SCAPC202114).

**Review statement.** This paper was edited by Jason West and reviewed by two anonymous referees.

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