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NO_x emissions in France in 2019–2021 as estimated by the high-spatial-resolution assimilation of TROPOMI NO₂ observations

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Abstract. Since 2018, the TROPOspheric Monitoring Instrument (TROPOMI) on board the Sentinel-5 Precursor (S5P) has provided unprecedented images of nitrogen dioxide (NO₂) tropospheric columns at a relatively high spatial resolution with a daily revisit. This study aims at assessing the potential of TROPOMI–PAL data to estimate the national to urban NO_x emissions in France from 2019 to 2021, using the variational mode of the recent Community Inversion Framework (CIF) coupled to the CHIMERE regional transport model at a spatial resolution of $10 \text{ km} \times 10 \text{ km}$. The seasonal to inter-annual variations in the French NO_x emissions are analyzed. Special attention is paid to the current capability to quantify strong anomalies in the NO_x emissions at intraannual scales, such as the ones due to the COVID-19 pandemic, by using TROPOMI NO_2 observations.

At the annual scale, the inversions suggest a decrease in the average emissions over 2019-2021 of -3% compared to the national budget from the Copernicus Atmosphere Monitoring Service regional inventory (CAMS-REG) for the year 2016, which is used as a prior estimate of the national-scale emissions for each year by the Bayesian inversion framework. This is lower than the decrease of -14% from 2016 to the average over 2019–2021 in the estimates of the French Technical Reference Center for Air Pollution and Climate Change (CITEPA). The lower decrease in the inversion results may be linked in large part to the limited level of constraint brought by the TROPOMI data, due to the observation coverage and the ratio between the current level of errors in the observation and the chemistry-transport model, and to the NO₂ signal from the French anthropogenic sources.

Focusing on local analysis and selecting the days during which the TROPOMI coverage is good over a specific local source, we compute the reductions in the anthropogenic NO_x emission estimates by the inversions from spring 2019 to spring 2020. These reductions are particularly pronounced for the largest French urban areas with high emission levels (e.g., -26% from April 2019 to April 2020 in the Paris urban area), reflecting reductions in the intensity of vehicle traffic reported during the lockdown period. However, the system does not show large emission decreases for some of the largest cities in France (such as Bordeaux, Nice and Toulouse), even though they were also impacted by the lockdown measures.

Despite the current limitations for the monitoring of emissions at the national scale, or for some of the largest cities in France, these results open positive perspectives regarding the ability to support the validation or improvement of inventories with satellite observations, at least at the local level. This leads to discussions on the need for a stepwise improvement of the inversion configuration for a better extraction and extrapolation in space and time of the information from the satellite observations.

1 Introduction

In Europe, nitrogen dioxide (NO₂) is emitted mainly by road traffic, thermal power plants and industrial activities and is produced in the atmosphere by the oxidation of nitric oxide (NO), which is emitted by the same activities. NO₂ is of great interest due to its important role in many atmospheric processes with strong implications for air quality, health, climate change and ecosystems. It is one of the major air pollutants with adverse impact on health (Costa et al., 2014; EEA, 2020). Deposition of nitrogen compounds like nitrates, for which NO2 is a precursor, leads to eutrophication of ecosystems (Stevens et al., 2018). NO2 also indirectly affects the radiative forcing as a precursor of tropospheric ozone and particulate matter. NO2 is therefore one of the regulated air quality pollutants. Nevertheless, despite ongoing improvements in the overall air quality, levels of air pollutants above the standards of the European Union (EU) are still measured across Europe, and air pollution remains a major health concern for European citizens (EEA, 2023). For example, France was condemned in 2019 by the Court of Justice of the EU (CJEU) for non-compliance with Directive 2008/50/EC relating to ambient air quality and more specifically for systematically and persistently exceeding concentration limit values (CLVs; 40 μg.m⁻³ on an annual average) for NO₂, particularly in the Île-de-France area, close to traffic. According to Airparif (2022), the planned reductions in the emissions of nitrogen oxides $(NO_x = NO + NO_2)$ will still be insufficient by 2025 to respect the NO2 CLV, which should have been reached by January 2010.

Society is thus faced with a major environmental challenge: the need to rapidly reduce NO2 concentrations to levels that comply with the law (EU Directive 2016/2284) and do not impact human health or ecosystems and therefore to reduce anthropogenic NO_x emissions. An accurate account of NO_x emissions in space and time is needed to assess the effectiveness of policies aimed at reducing NO_x emissions. However, the quantification of anthropogenic NO_x emissions following a bottom-up (BU) approach, based on the statistics of activity sectors and fuel consumption and relying on emission factors per activity type, suffers from relatively large uncertainties. For example, at national and annual scales, these uncertainties reach 50 %-200 % depending on the activity sector in the European Monitoring and Evaluation Programme (EMEP) inventory (Kuenen and Dore, 2019), and these emission factors can be biased (e.g., with Dieselgate; Brand, 2016). Schindlbacher et al. (2021) reported uncertainty estimates for total national NO_x emissions ranging from 5% for Norway to 45% for Ireland. In addition, the use of proxies and typical temporal profiles inevitably introduces errors in the quantification of the spatiotemporal variability at high resolutions. In situ information could be used to analyze local emissions and their variations due to specific events or measures (Guevara et al., 2023). Finally, the BU inventories are often delivered with a 2-year lag. The assessment of air quality (AQ) policies (as mentioned above) would benefit from accurate emission inventories spatialized at a high resolution with a fast update capability, and integrating independent information could play a critical role in AQ analysis and policies. The use of atmospheric measurements to complement current BU approaches may then support the development of such inventories.

Since the 2000s, NO₂ atmospheric mixing ratios have been monitored around the world by space-borne instruments, such as the Global Ozone Monitoring Experiment (GOME; Burrows et al., 1999) and GOME-2 (Munro et al., 2016), the SCanning Imaging Absorption spectroMeter for Atmospheric CHartographY (SCIAMACHY; Burrows et al., 1995; Bovensmann et al., 1999) and the Ozone Monitoring Instrument (OMI; Levelt et al., 2018).

In this context, attempts have been made to develop socalled top-down (TD) methods, complementary to BU inventories, to deduce NO_x emissions from NO_2 satellite data. These methods are based on the principle that atmospheric levels and variations in NO_2 reflect the convolution of the amplitude and variations in NO_x emissions with atmospheric chemistry and physics. Through the statistical inverse modeling of the atmospheric chemistry and transport, one can derive estimates of the emissions based on the concentration fields. However, strong non-linear relationships exist between NO_x emissions and satellite NO_2 tropospheric vertical column densities (TVCDs) (Lamsal et al., 2011; Vinken et al., 2014; Miyazaki et al., 2017; Elguindi et al., 2020) due to the complex chemistry affecting NO_x in the atmosphere.

This complex atmospheric chemistry has been taken into account in various ways in more or less complex TD methods. Mass-balance approaches have been performed at the global (Lamsal et al., 2011; Vinken et al., 2014) and regional (Visser et al., 2019) scales, accounting for non-linear relationships between NO_x emission changes and NO₂ TVCDs via reactions with hydroxyl (OH) radicals but with simple scaling factors. However, Stavrakou et al. (2013) have shown that other direct or indirect NO_x sinks associated with other species (such as ozone (O₃) or the HO₂ radical) could significantly influence NO2 concentrations in the atmosphere and therefore TD estimates. A more detailed account of the complex NO_x chemistry should thus support more accurate derivations of NO_x emissions from NO_2 satellite data. Therefore, more elaborate approaches using chemistry-transport models (CTMs) with ensemble Kalman filter inverse modeling techniques or variational approaches have been used to infer NO_x emissions at the global (Müller and Stavrakou, 2005; Miyazaki et al., 2017) or regional scale (van der A et al., 2008; Mijling and van der A, 2012; Mijling et al., 2013; Lin, 2012; Ding et al., 2017; Savas et al., 2023). However, NO_x inversions from satellite NO_2 observations have resolution-dependent biases: coarse-resolution models are known to bear negative biases in NO_2 over large sources (Valin et al., 2011). It would therefore be essential to operate at finer spatial resolutions. In addition, monitoring NO_x emissions implies monitoring hotspots of emissions (large urban and industrial areas or strong point sources), where much of the global emissions are concentrated (57% of the global population lives in urban areas as of 2022; Ritchie and Roser, 2018). Individual monitoring of emission hotspots is required to assess the emission reduction policies, since the scales they target range from the regional scale (e.g., that of the EU) to the country scale or even to that of smaller territory units, e.g., cities. This in turn requires the mapping of NO_2 concentrations at high spatial and temporal resolutions.

Since 2017, the TROPOspheric Monitoring Instrument (TROPOMI; Veefkind et al., 2012) on board the Copernicus Sentinel-5 Precursor (S5P) has monitored atmospheric NO₂ with high-resolution imaging (pixel size of about 5.6 km × 3.6 km since August 2019), which should support the quantification of anthropogenic emissions at national to local scales. With a swath as wide as approximately 2600 km on ground, TROPOMI also provides an unprecedented daily coverage. It theoretically covers any point of the Earth 1 to 2 times a day. The need for surface solar irradiance, the cloud cover and the quality filtering limit the number of pixels in numerous images locally (e.g., at high latitudes), but the possibility of having a follow-up week by week (even day by day) remains for large portions of the globe.

To fully exploit these TROPOMI satellite images, variational inversion systems seem particularly well adapted, since they allow high-dimensional problems to be solved (Elbern et al., 2000; Quélo et al., 2005; Pison et al., 2009; Henze et al., 2009; Cao et al., 2022), typically addressing the emission fluxes at high spatial and temporal resolutions and assimilating a large number of data, such as those provided in TROPOMI images. The non-linearities of the chemistry of NO_x can be dealt with in a variational inversion framework which drives a regional CTM using a manageable chemistry scheme to simulate NO_2 concentrations and whose adjoint code is available.

In this context, this study assesses the potential of the TROPOMI observations to inform us about NO_x emissions in France from 2019 to 2021 at national to urban scales. The primary target of the inversions and analysis is the anthropogenic emissions, i.e., mainly those due to the combustion of fossil- and biofuels. The soil emissions (including the impact of agricultural practices), which represent a much smaller share of the total national emissions and which are more diffuse, are assumed to be more difficult to diagnose and are kept as a secondary target of the inversions.

We use the high-dimensional variational inversion drivers of the recent Community Inversion Framework (CIF; Berchet et al., 2021). The CIF drives a configuration of the CHIMERE regional CTM (Menut et al., 2013; Mailler et al., 2017) covering France at $10 \,\mathrm{km} \times 10 \,\mathrm{km}$ spatial resolution,

including a chemistry module taking into account the complex NO_x chemistry in gas phase, its non-linearities, and its adjoint (Fortems-Cheiney et al., 2021). This relatively fine spatial resolution makes it possible to focus on the largest French urban areas. The period 2019–2021 covers the phase of the COVID-19 crisis in spring 2020 during which NO₂ concentrations and anthropogenic NO_x emissions were expected to have significantly dropped over Europe (Bauwens et al., 2020; Menut et al., 2020; Diamond and Wood, 2020; Ordóñez et al., 2020; Petetin et al., 2020; Barré et al., 2021; Gaubert et al., 2021; Deroubaix et al., 2021; Lee et al., 2021; Souri et al., 2021; Levelt et al., 2022; Guevara et al., 2021, 2022, 2023). In France, the population was confined from 17 March to 10 May, and all public spaces deemed nonessential to daily life in the country were shut down. Then, from 11 May to 1 June, lockdown restrictions were progressively lifted. The population was confined again from 30 October to 15 December. The analysis of the emissions from 2019–2021 should thus provide insights on the current capability to quantify strong anomalies in the NO_x emissions at intra-annual scales by using satellite NO₂ observations.

Our configuration of the CHIMERE CTM, the NO_2 TROPOMI satellite observations, and the variational inversion framework and set up are described in Sect. 2. Section 3 presents the results of our study, including comparisons between TROPOMI–PAL NO_2 TVCDs and their CHIMERE-simulated equivalents and the analysis of the spatiotemporal variability in the French NO_x emissions. Our conclusions are given in Sect. 4.

2 Data and methods

2.1 Prior estimates of the emission maps

The principle of the inversion is to correct a priori emission maps, referred to later on as "prior emissions". In this study, the inversion controls NO and NO_2 emissions based on the Bayesian update of a prior estimate of these emissions. Therefore, there is a need for independent maps of the NO_x emissions to derive this prior estimate. There is also a need for estimates of the anthropogenic emissions of 15 species, including non-methane volatile organic compounds (NMVOCs) and carbon monoxide (CO), that are used in the chemistry scheme of the atmospheric chemistry-transport model, even though they are not controlled by the inversions (these emissions remain fixed during the inversion process; see Sect. 2.5).

Here, the prior estimates of anthropogenic NO_x emissions are based on a combination of the Copernicus Atmosphere Monitoring Service regional inventory (CAMS-REG) (Kuenen et al., 2022) for the year 2016 and the Inventaire National Spatialisé (INS; Ministère de la Transition écologique et solidaire, 2012) for the year 2012. Hereafter, the term "anthropogenic emissions" mainly corresponds to emissions due to the combustion of fossil and biofuels, and it excludes the

emissions due to soil fertilization by agriculture, such as CAMS-REG and INS.

CAMS-REG is an inventory of the anthropogenic emissions of pollutants in Europe which is spatialized at a 0.05° longitude × 0.1° latitude resolution. Annual and national budgets in this inventory are based on the officially reported emission data by European countries to the Convention on Long-Range Transboundary Air Pollution and the EU National Emission Ceilings Directive. These budgets are then disaggregated based on proxies of the different sectors, separating point sources and area sources, described in Kuenen et al. (2022). Default profiles for typical emission height by source type (Kuenen et al., 2022), which account for the average effective emission height (including plume rise), are based on Bieser et al. (2011). Temporal disaggregation is based on temporal profiles provided per sector with typical month-to-month, weekday-to-weekend and diurnal variations (Ebel et al., 1997; Menut et al., 2012).

The prior estimates of the anthropogenic emissions are completely derived from CAMS-REG outside France. In France, they are derived from the annual budgets and the temporal and vertical profiles of CAMS-REG, but the horizontal spatialization is based on the INS, i.e., using proxies at municipal scale from this inventory. The inventory product used for the prior estimate of the emissions is therefore called "CAMS-REG/INS" in the following.

Following the GENEMIS recommendations (Kurtenbach et al., 2001; Aumont et al., 2003), the anthropogenic NO_x emissions have been speciated as 90 % of NO, 9.2 % of NO_2 and 0.8 % of nitrous acid (HONO) emissions.

The NO biogenic soil emissions are prescribed using simulations from the Model of Emissions of Gases and Aerosols from Nature (MEGAN) model (Guenther et al., 2006), with an $\sim 1 \,\mathrm{km} \times 1 \,\mathrm{km}$ spatial resolution, which, in principle, does not take the impact of agricultural practices into account, even though it covers both natural and agricultural areas. Recent studies have indeed shown that MEGAN significantly underestimates soil emissions in agricultural areas (Oikawa et al., 2015; Almaraz et al., 2018; Sha et al., 2021; Zhu et al., 2023). However, there are large uncertainties in the NO_x emissions due to agriculture, and, in principle, there could be some overlapping between the agricultural and purely natural soil NO_x emission estimates. This explains why these emissions are not provided by CAMS-REG (Kuenen et al., 2022). Therefore, we do not include a specific agricultural soil NO_x emissions component in our prior estimation of the NO_x emissions.

Lightning NO_x fluxes, whose impact on NO_2 concentrations is very small in Europe even in summer (Menut et al., 2020), are not accounted for. Fire emissions are also ignored, as we assume that they only slightly contribute to the total NO_x emissions.

These emission maps have been aggregated to the grid of CHIMERE for the years 2019 to 2021 (Sect. 2.2 and Fig. 1).

2.2 Configuration of the CHIMERE CTM

We use the CHIMERE v2013 model to simulate fields of concentrations of gaseous chemical species in a domain that covers France and its vicinity (39.5–54.5° N, 11° W–12° E; see Fig. 1). The model horizontal grid is zoomed (Siour et al., 2013), with a 10 km resolution regular sub-grid in the center of the domain covering the whole of France and a 50 km resolution in the corners of the domain (Fig. 1). It corresponds to 166 (longitude) × 122 (latitude) horizontal grid cells. The model has 20 vertical layers from the surface to 200 hPa, with 8 layers within the first 2km.

The model is driven by the European Centre for Medium-Range Weather Forecasts (ECMWF) global meteorological fields (Owens and Hewson, 2018). Both CHIMERE (Menut et al., 2013) and its adjoint code operate the MELCHIOR-2 chemical scheme with more than 100 reactions (Lattuati, 1997; Derognat et al., 2003), including 24 for inorganic chemistry. Considering the short lifetime of NO₂, we do not consider its import from outside the domain: its boundary conditions are set to zero. Nevertheless, the lateral and top boundaries for other species, such as ozone (O₃), nitric acid (HNO₃), peroxyacetyl nitrate (PAN) and formaldehyde (HCHO), participating in the NO_x chemistry are considered. Initial and boundary conditions when relevant are specified using a nested run of CHIMERE (Siour et al., 2013) over a European domain (31.75-74.25° N, 15.25° W-35.75° E) with a spatial resolution of 0.5°, which itself uses boundary and initial conditions from climatological values from the LMDz-INCA global model (Szopa et al., 2009).

The aerosol module of CHIMERE is not considered in the simulations and inversions, as the adjoint of this module is not available (Fortems-Cheiney et al., 2021; Savas et al., 2023).

2.3 TROPOMI-PAL observations

TROPOMI, launched on board Sentinel-5 Precursor (S5P) in October 2017, is in a near-polar sun-synchronous orbit (altitude of $\sim824\,\mathrm{km}$) with an ascending node equatorial crossing at $\sim13:40$ Mean Local Solar Time. With an orbital cycle of 16 d and 14 orbits a day, the satellite passes over the same geographical area every cycle of 227 orbits. With a swath as wide as 108° – approximately 2600 km on ground – TROPOMI provides daily coverage for NO2. Observations over our domain span from around 09:30 to 14:30 local time with data from 2 to 3 orbits per day, with the major coverage around noon (11:30 to 13:30).

An evaluation of the TROPOMI v1.3 product with surface remote sensing observations indicated a systematic low bias of TROPOMI NO₂ tropospheric vertical column densities (TVCDs) of typically $-23\,\%$ to $-37\,\%$ in clean to slightly polluted conditions and as high as $-51\,\%$ over highly polluted areas (Verhoelst et al., 2021) compared to ground-based measurements. This negative bias has been mainly attributed

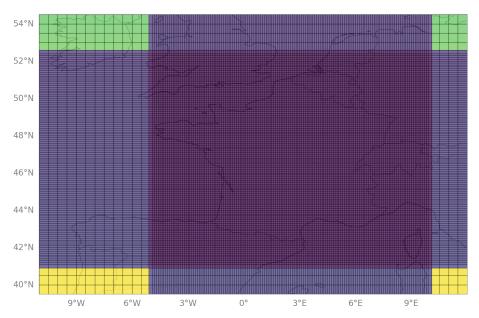


Figure 1. Domain of our CHIMERE configuration: $10 \, \text{km}$ resolution regular sub-grid in purple, $50 \, \text{km}$ resolution in green and yellow, and $10 \, \text{km} \times 50 \, \text{km}$ resolution in blue.

to a negative cloud height bias in the Fast Retrieval Scheme for Clouds from Oxygen absorption band (FRESCO) implementation (van Geffen et al., 2022b), and efforts have been made to correct it in the TROPOMI–PAL product (Eskes et al., 2021).

Here, we use the TROPOMI-PAL reprocessed data (Eskes et al., 2021), available from 2019 to 11 November 2021. We use a recent reprocessing of the TROPOMI data, called RPRO version 2.4, to cover the end of the year 2021. This latest reprocessing uses a new higher-resolution directional Lambertian-equivalent reflectivity derived from TROPOMI observations (van Geffen et al., 2022a). Nevertheless, the last evaluation of the TROPOMI RPRO v2.4 product with surface remote sensing observations still indicates significant biases of TROPOMI, with NO₂ TVCDs of typically +13% over clean areas to -40% over highly polluted areas (Lambert et al., 2023). These biases could partly be explained by the relatively coarse horizontal resolution of the prior global TM5-MP profiles $(1^{\circ} \times 1^{\circ})$ used in the retrieval process (which can be neglected when applying the retrieval averaging kernels to the model) (Douros et al., 2023). However, these biases are also largely attributed to systematic errors in the retrieved cloud pressure, surface albedo used, etc. (Boersma et al., 2016; Douros et al., 2023).

We select the data with a quality assurance (QA) value of 0.75 for both products, following the criteria of van Geffen et al. (2022b). The latest version of the TROPOMI NO₂ Algorithm Theoretical Basis Document (ATBD) is now for product version 2.6 (TROPOMI ATBD of the total and tropospheric NO₂ data products, KNMI and S5P-KNMI-L2-0005-RP, issue 2.4.0; van Geffen et al., 2022a).

2.4 Comparison between simulated and observed NO₂ TVCDs

To make relevant comparisons between simulations and satellite observations, the averaging kernels (AKs), which are associated with each observed TVCD and represent the vertical sensitivity of the satellite retrieval (Eskes and Boersma, 2003), are applied to the simulated concentration field. Due to the over-sampling resulting from the model's horizontal resolution being coarser than the TROPOMI data, we aggregate spatially and temporally the TROPOMI observations at the CHIMERE resolution into so-called super-observations. Within a given grid cell and time step, the super-observation is the observation (TVCD and AKs) which is the closest to the mean of the TROPOMI TVCDs. The error associated with each super-observation is also derived from the observation closest to the mean value and is subsequently included in the total so-called "observation error" (see Sect. 2.5). Our derivation of the error associated with each super-observation is thus conservative compared to other studies (Boersma et al., 2016), where the super-observation uncertainty is reduced compared to that of individual observations.

The reduction in uncertainty when combining several observations accounts for the fact that the retrieval errors include random noise (in particular, instrumental noise) without spatial correlation, i.e., errors which are independent from one observation to the other. However, as discussed above, the TROPOMI NO₂ observations bear large systematic errors from the retrieval process, which can exhibit significant spatial correlations. This explains our conservative attribution of observation errors to the super-observations.

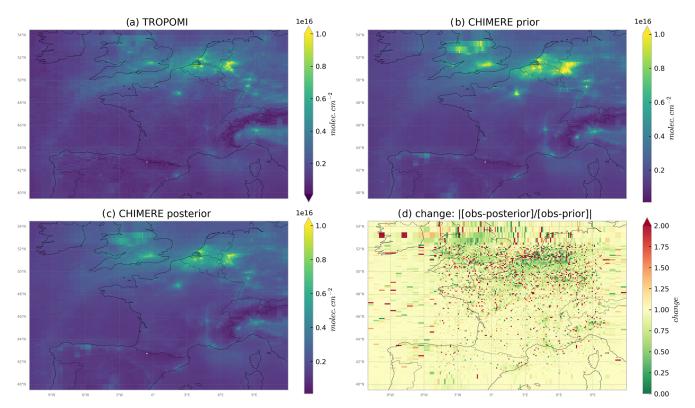


Figure 2. Averages of NO₂ TVCDs by (**a**) the TROPOMI–PAL data; (**b**) the CHIMERE simulation using prior emissions from CAMS-REG/INS and MEGAN, described in Sect. 2.1; and (**c**) the CHIMERE simulation using the posterior emissions from the inversion, for April 2020, in molec. cm⁻². (**d**) Ratio of the posterior and prior biases between NO₂ TVCDs simulated with CHIMERE and the TROPOMI–PAL observations. All ratios lower than 1, in green, demonstrate that posterior emission estimates improve the simulation compared to the prior ones.

The corresponding column of NO_2 in CHIMERE is vertically interpolated (at TROPOMI's super-observation location) on the vertical levels of the super-observation retrieval, and vertically integrated with the AKs of the super-observation, to yield the NO_2 simulated TVCD to be compared to the super-observation TVCD, as illustrated in Fig. 2a and b for the month of April 2020.

2.5 Variational inversion of NO_X emissions

The inversion of NO_x emissions consists of correcting the prior estimate of the emissions (presented in Sect. 2.1) to improve the fit between NO_2 TROPOMI–PAL satellite data and their simulated equivalents, using a Bayesian variational inversion framework similar to that of Fortems-Cheiney et al. (2021).

Series of 7 d inversion windows – independent from each other – are run and then combined to provide a corrected ("posterior") estimate of NO_x emissions over the whole period of analysis (2019–2021). For each inversion window, the posterior estimate of the emissions is found by iteratively minimizing the cost function J(x):

$$J(\mathbf{x}) = \frac{1}{2} (\mathbf{x} - \mathbf{x}^{b})^{\mathrm{T}} \mathbf{B}^{-1} (\mathbf{x} - \mathbf{x}^{b})$$
$$+ \frac{1}{2} (\mathcal{H}(\mathbf{x}) - \mathbf{y})^{\mathrm{T}} \mathbf{R}^{-1} (\mathcal{H}(\mathbf{x}) - \mathbf{y}), \tag{1}$$

where x, \mathcal{H} , y, \mathbf{B} and \mathbf{R} are, respectively, the control vector, the observation operator, the satellite observations, the prior error covariance matrix and the observation error covariance matrix. The control vector \mathbf{x} gathers variables for the correction of the surface NO and NO₂ emissions, with \mathbf{x}^b corresponding to the prior estimate of the control vector.

Here the definition of x ensures that the inversion solves separately for the two main types of NO emissions, the anthropogenic and the biogenic emissions (without any further sectorization or decomposition into more detailed emission components), and for the anthropogenic NO₂ emissions at the 1 d and model grid cell (i.e., 50 to 10 km) resolution temporally and horizontally and over three vertical levels for the anthropogenic emissions (accounting for the range of injection heights discriminating between the anthropogenic sources). With such a control vector, the prior NO/NO₂ anthropogenic emission ratio speciation from the GENEMIS recommendations (see Sect. 2.1) is not conserved in the in-

version, but the analysis focuses on the NO_x emissions as the sum of the NO and NO_2 emissions.

Furthermore, contrary to Fortems-Cheiney et al. (2021), we implicitly aim to characterize the prior uncertainty in both the anthropogenic and biogenic NO_x emissions with log-normal distributions. This allows the inversion system to apply high variations in NO_x emissions while ensuring positivity, unlike the classic corrections of the emission with scaling factors with a Gaussian distribution of prior uncertainty. However, the uncertainty in the control vector must follow a Gaussian distribution, as required by the use of Eq. (1). Therefore, here, the control vector x is defined as the logarithm of the scaling factors to be applied to the prior estimate of the emissions, and the posterior anthropogenic or biogenic emission estimate at a given grid cell of the model and for a given day f_i is derived from the corresponding control parameter x_i as $f_i = \exp(x_i)$. f_i^{prior} .

Our control vector \mathbf{x} is therefore characterized as follows:

- It contains the logarithm of the scaling coefficients for NO anthropogenic emissions at a 1 d temporal resolution, for the 166 (longitude) × 122 (latitude) horizontal grid cells of the model, and over three vertical bands of injection heights for the emissions (from 0 to 25 m, from 25 to 1900 m and from 1900 to 12 000 m); this is done essentially to reduce the dimensionality of the problem along this axis, as most emissions are concentrated in the first two layers. The corrections applied are the same for all layers within a band.
- It contains the logarithm of the scaling coefficients for NO₂ anthropogenic emissions at the same temporal and spatial resolutions as for NO.
- It contains the logarithm of the scaling coefficients for NO biogenic emissions at a 1 d temporal resolution, for the 166 (longitude) × 122 (latitude) model grid cells at the surface level.

NO and NO₂ 3D initial conditions (specified using a nested run of CHIMERE, as described in Sect. 2.2) are not controlled and are set once initially for all 7 d windows at 00:00 UTC on the first day of these windows. We therefore do not account for the potential update of the concentrations during a previous 7 d window due to the inversions. They have a low impact on the inversion, since the first satellite observations over the domain are around noon, while the NO_{χ} lifetime is short (on the order of a few hours; Hakkarainen et al., 2024).

The uncertainties in the observations y, together with those in the observation operator \mathcal{H} , and the uncertainties in the prior estimate of the control vector x^b are assumed to have a Gaussian distribution. Therefore, they are fully characterized by error covariance matrices.

The prior uncertainty is defined at the resolution of the control vector. Thus, the terms of the prior error covariance matrix **B** reflect the uncertainties in the logarithm of the anthropogenic and biogenic emissions of NO_x at 1 d and 50 to 10 km resolution. This matrix is set block diagonal (see Eq. 2), with one block corresponding to the anthropogenic emissions and the other one corresponding to the biogenic emissions, assuming that there is no correlation between the respective uncertainties in the prior estimates for these two types of emissions.

Another assumption is that, at the 1 d and 50 to 10 km resolution, there are no spatial or temporal correlations in the prior uncertainties in the anthropogenic emissions, due to the heterogeneity of these emissions. Therefore, the first block of **B** corresponding to the logarithms of the anthropogenic emissions is set diagonally. Each diagonal element is set at $(0.3)^2$: the range associated to this σ value in the log-space corresponds to a factor ranging between 74 %–135 % in the emission space at the 1 d resolution and at the model's grid scale.

The second block of **B** corresponding to the biogenic fluxes accounts for space correlations in the uncertainties in these emissions, which are assumed to be more homogeneous in space and time than the anthropogenic emissions and to decrease exponentially with distance. They are set with a $\lambda_0 = 30$ km decorrelation length. On the diagonal, uncertainties are set to a value of $(0.6)^2$. The range associated to this value σ in the log-space corresponds to a factor ranging between 55 %–182 % in the emission space at the 1 d resolution and at the model's grid scale.

$$\mathbf{B} = \begin{pmatrix} \begin{pmatrix} \sigma_{ii}^{\text{ant}} & \dots & 0 \\ \vdots & \ddots & \vdots \\ 0 & \dots & \sigma_{jj}^{\text{ant}} \end{pmatrix} & \mathbf{0} \\ & \mathbf{0} & \begin{pmatrix} \sigma_{j+1,j+1}^{\text{bio}} & \dots & \nu_{k,j+1}^{\text{bio}} \\ \vdots & \ddots & \vdots \\ \nu_{j+1,k}^{\text{bio}} & \dots & \sigma_{kk}^{\text{bio}} \end{pmatrix} \end{pmatrix}, \quad (2)$$

with ν in the form of $\nu_{nm} = \sigma_n \sigma_m e^{\frac{-\Delta x}{\lambda_0}}$.

 ${\cal H}$ is the observation operator, linking the control variables in the log-space to the simulated equivalents of the super-observations. It includes the exponential operator and scaling factor converting the maps of the logarithm of the coefficient for the emissions into emission maps at the spatial and temporal resolutions of CHIMERE, the atmospheric chemistry and transport model CHIMERE itself, and the extraction of the TVCDs from CHIMERE where and when we have TROPOMI super-observations.

The uncertainties on the observations y and on the observation operator are characterized by the so-called observation error covariance matrix \mathbf{R} , set up here as a diagonal matrix based on the assumption that these errors are not correlated in space or time when aggregated at the model 50 to $10 \, \mathrm{km}$ and $1 \, \mathrm{h}$ resolution. The variance in the observation errors corresponding to individual observations in the diagonal of \mathbf{R} is the quadratic sum of the error we have assigned to the TROPOMI–PAL super-observations (see Sect. 2.4) and of an

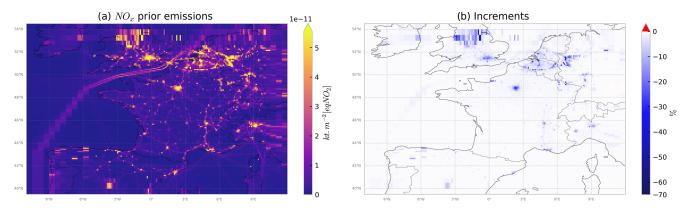


Figure 3. (a) Total prior NO_x fluxes (anthropogenic emissions from CAMS-REG/INS and biogenic emissions from MEGAN; see Sect. 2.1 for details) and (b) relative increments to the total prior emissions from the inversion in % for April 2020.

estimate of the errors from the observation operator. We assume that the observation operator error is dominated by the chemistry-transport modeling errors and by the errors associated with the discrepancies between the spatial representativity of the super-observations and of the model corresponding column: it is set at 30 % of the retrieval value. It was set at 20 % by Fortems-Cheiney et al. (2021) at a coarser resolution and is increased here to take into account the mismatch between the shape and location of the real and simulated atmospheric patterns at our finer resolution (see Sect. 2.4).

The minimum of the cost function J is searched for with the iterative M1QN3 limited-memory quasi-Newton minimization algorithm (Gilbert and Lemaréchal, 1989). At each iteration, the computation of the gradient of J relies on the adjoint of the observation operator and, in particular, on the adjoint of CHIMERE. In the results presented in Sect. 3, as a compromise between computational time and the level of convergence of the iterative minimization of J in the inversions, the minimization is regarded as satisfactory when the norm of the gradient of J is reduced by 80%.

The calculation of the uncertainty in the posterior estimates of emissions is challenging when using a variational inverse system (Kadygrov et al., 2015; Rayner et al., 2019; Fortems-Cheiney et al., 2021), and it is not done here. It would require a large ensemble of time-consuming inversions (especially due to the handling of chemistry) to enable a proper sampling of the uncertainties and thus a proper derivation of the statistics of uncertainty.

3 Results

3.1 Seasonal and inter-annual estimates of total French NO_x emissions

3.1.1 Fit between TROPOMI–PAL super-observations and their simulated equivalents

Before analyzing the results in terms of emissions, we check the behavior of the inversion by comparing the per-

formances of the prior and posterior simulations in reproducing the spatial and temporal variations of the observations. As an illustration, TROPOMI-PAL and the corresponding CHIMERE NO2 TVCDs are shown in Fig. 2a and b, respectively, for April 2020. The TROPOMI-PAL observations and their NO₂ simulated equivalents present similar spatial patterns, with hotspots (TVCDs higher than 1×10^{16} molec. cm⁻²) over urban areas and low values over rural areas during the whole simulated period from 2019 to 2021 (illustrated in Fig. 2a and b for 1 month in 2020). However, the prior simulation overestimates the NO₂ TVCDs over urban areas in France compared to the observations. For example, for April 2020, the mean bias between the prior simulation and TROPOMI is about 4.2×10^{14} , 2.0×10^{14} and 3.8×10^{14} molec. cm⁻² over Paris, Lyon and Marseille (Fig. A1), respectively. The inversion brings the simulated TVCDs closer to the TROPOMI–PAL data over urban areas (Fig. 2c): in this case, the mean bias and the mean root mean square over the three cities are reduced by about 30 % and 12%, respectively.

As the CHIMERE prior simulation overestimates the NO₂ TVCDs, the inversion brings the CHIMERE NO₂ columns closer to the TROPOMI–PAL data by reducing NO_x emissions, mainly over dense urban areas (Île-de-France, Lyon–Marseille axis, London area, Benelux, Frankfurt, Po Valley; see Fig. 3b). Over these areas, the relative corrections provided by the inversions to the total (anthropogenic and biogenic) emissions can reach -70% (Fig. 3b).

3.1.2 Estimates of total French NO_x emissions

This section focuses on the results from the NO_x inversions in terms of comparisons between the total posterior NO_x emissions and the prior ones. The focus on total emissions is explained by the fact that the distinction between the signal from the biogenic and anthropogenic emissions in the comparisons between the chemistry-transport model and the satellite NO_2 observations is challenging in the inversion

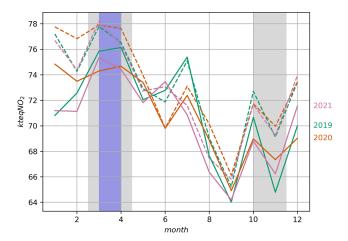


Figure 4. Monthly total NO_x emissions in France as estimated by CAMS-REG/INS and MEGAN (dotted lines) and by the inversions for the years 2019 (in green), 2020 (in orange) and 2021 (in pink), in kteq NO_2 month⁻¹. Gray shaded areas show the French lockdown periods for the year 2020, and the purple shaded area shows the French lockdown period for the year 2021.

framework. Since both types of emissions are solved at the 1 d and model grid cell resolution, the inversion system relies on the respective amplitude and spatial correlations in the prior uncertainties in the biogenic and anthropogenic emissions (in **B**) to discriminate between the corrections to be applied to the prior estimates of the two types of emissions. However, our configuration of **B** yields similar structures of spatial correlations for the two types of emissions. Therefore, our confidence in the split between biogenic and anthropogenic emissions in the posterior emission estimates is low whenever the two types of emissions have comparable levels according to the prior emission estimates.

At the national scale, both the prior and the posterior emissions present a similar seasonal cycle in 2019–2021, with monthly emissions higher than 72 kteq NO₂ during winter and equal to or lower than 66 kteq NO₂ during summer (Fig. 4).

The mean French national budget for the years 2019 to 2021 from the total posterior emission estimates is about 850 kteq NO_2 , which is lower than the total prior emissions estimated from CAMS-REG/INS and MEGAN (average of 875 kteq NO_2 ; Fig. 3, Table 1), with the largest reductions reaching about 8% during fall and winter. This is expected, as public policies led to regular reductions in anthropogenic emissions between 2016 (our version of CAMS-REG/INS) and 2019-2021. The decrease in total NO_x emissions is estimated at -13% from 2016 to 2019 by the French Technical Reference Center for Air Pollution and Climate Change (CITEPA report; we use the estimates for "out-of-scope" natural emissions which are not reported to UNFCCC). According to CITEPA, the reduction is driven by large reductions in emissions from three major sectors between 2016 and 2019:

the energy sector (-28%), the industry sector (-22%) and the transport sector (-15%). However, the decrease found in this work is only about -3% (Table 1). This can be explained by our configuration with neither spatial nor temporal correlation in our **B** matrix, leading to null or very small correction of the prior emissions from the inversions when the coverage of the country by TROPOMI super-observations is very sparse (Zheng et al., 2020). In this case, the posterior emissions remain close to the prior emission estimate and therefore at their 2016 levels (see Sect. 3.2.3).

The 2019–2021 inter-annual variability is smaller with our inversions (Table 1) than in the estimates from CITEPA: the annual budget of the total French posterior NO_x emissions varies by less than 1 % from year to year, while CITEPA estimates a decrease of about 13 % in 2020 compared to 2019 and an increase of about 3 % in 2021 compared to 2020 (Table 1). The similar total annual emissions in 2019 and 2020 nevertheless overlay different sub-annual variations. Higher emissions – partly associated with higher TROPOMI NO₂ tropospheric columns (not shown) - are estimated in January, February, June and November 2020 compared to 2019 (Fig. 4). These increases counterbalance the decrease in NO_x emissions in March/April 2020 (Figs. 4 and 5) which could be associated both with the COVID-19 crisis and with meteorological conditions with a warmer winter in 2020 than in 2019 (see Sect. 3.2.2).

When considering the split between biogenic and anthropogenic emissions from the inversions, despite its lack of reliability, the posterior biogenic emission estimates are close to the prior estimates derived from MEGAN (see Figs. C1 and C2): in particular, they do not appear to be stronger in 2020 than in 2019. Therefore, the low decrease (-3%) in the posterior estimate of the national budget of the total NO_x emissions in 2020 can hardly be explained by an increase in the biogenic emissions, which would compensate for the decrease in the anthropogenic emissions, even though the biogenic emissions due to agriculture were overlooked in our inversion setup (see Sect. 2.1).

3.2 Impact of the COVID-19 lockdown in spring 2020

The atmospheric lifetime of NO_2 dictates that the high-spatial-resolution measurements from TROPOMI should readily capture rapid week-to-week changes in near-surface emissions (Levelt et al., 2022) and should therefore make it possible to assess the impact of the COVID-19 lockdown in spring 2020 on French NO_x emissions. Following a usual diagnostic in the literature to assess the change in air pollutant concentrations due to COVID-19 policies, we characterize the impact of the first COVID-19 lockdown in France – occurring from 17 March to 10 May 2020 – in terms of changes from spring 2019 to spring 2020. Our analysis relies on the spatial distribution of the main areas of anthropogenic activities to identify the locations where the total NO_x emissions should be dominated by anthropogenic emissions.

Table 1. Total prior and posterior NO_x emission budgets in kteq NO_2 and their relative differences in % [100 × (posterior – prior)/prior] for different periods in France. Columns "Posterior year_n – year_{n-1}" and "CITEPA year_n – year_{n-1}" show the relative difference between year_n and year_{n-1} posterior fluxes and CITEPA in % [100 × $(F_{year_n} - F_{year_{n-1}})/F_{year_{n-1}}$].

Domain or	Period	CAMS-REG/INS + MEGAN inventories	Posterior in	2019	Posterior in 2020		Posterior in 2021		Posterior 2020–2019	CITEPA 2020–2019	Posterior 2021–2020	CITEPA 2021–2020
region		kteq NO ₂	$kteqNO_2$	%	kteq NO ₂	%	kteq NO ₂	%	%	%	%	%
France	Annual	875	853	-3	852	-3	845	-3	0	-13	-1	+3
	Spring (MAM)	227	224	-1	222	-2	222	-2	-1	-21	0	+12
	Mar	78	76	-3	74	-5	75	-4	-2	-12	+1	+5
	Apr	77	76	-1	75	-4	74	-3	-2	-30	0	+23
	May	73	72	-1	73	-1	72	-1	+2	-22	-2	+10
	Nov	69	65	-6	67	-4	66	-4	+4	-19	-2	+14

Table 2. Changes in TROPOMI–PAL and CHIMERE NO_2 tropospheric columns (in %) and changes in total posterior NO_x emissions (in %), between March/April 2019 and March/April 2020 for the eight French urban areas displayed in Fig. A1.

Urban area	TROPOMI TVCD		CHIMERE prior TVCD		CHIMI posterior		Emissions posterior	
	Mar	Apr	Mar Apr		Mar	Apr	Mar	Apr
France	-11	-28	+7	-4	+2	-12	-2	-1
Bordeaux	-34	-27	-14	+23	-15	+21	-3	-1
Lille	-15	-35	+19	-10	+4	-28	-9	-11
Lyon	-27	-43	+13	+18	+1	+1	-15	-13
Marseille	-20	-27	-2	+46	-9	+19	-2	-7
Nantes	+2	-45	+16	-2	+14	-7	-3	-3
Nice	-28	-37	0	+7	-3	+2	0	-1
Paris	-41	-54	-4	+42	-12	-9	-11	-26
Toulouse	-41	-31	-23	-23	-22	-22	0	+1

3.2.1 General impact on NO₂ TROPOMI TVCDs

The TROPOMI NO₂ TCVDs in March/April 2020 are first compared to 2019 at the national scale (Table 2). The TROPOMI NO₂ TCVDs decrease from 2019 to 2020 by -11 and -28%, respectively, in March/April over France. Similar decreases have been diagnosed in the measurements at surface stations over Europe with reductions in the average concentrations of NO₂ of about -25% for at least 75% of the 1308 European Air Quality e-Reporting database (Air-Base) stations compared with the average of the previous 7 years (2013–2019) for the period 18 March–18 May 2020 (Deroubaix et al., 2021).

The population distribution is heterogeneous in France, with large rural areas (i.e., 75.7% of the country area, according to the French National Institute of Statistics and Economic Studies, 2020). We focus here on eight urban areas which correspond to hotspots of NO_x concentrations and where, as a consequence, we expect a stronger signal due to changes in anthropogenic activities (Fig. A1): Paris, Lyon, Marseille, Lille, Bordeaux, Toulouse, Nice and Nantes.

The strength of the TROPOMI NO₂ signal differs over these eight French cities. Indeed, the absolute changes in NO₂ TVCD in March/April 2020 compared to March/April 2019 are higher for Paris, Lille, Lyon and Nantes (see

Fig. C5) than for Bordeaux, Marseille, Nice and Toulouse (i.e., 1×10^{15} molec. cm⁻² or less). The population density difference between those cities could partly explain such variability (Table A1).

TROPOMI NO₂ tropospheric columns in March/April 2020 are 26 % to 38 % lower on average than in 2019 over the eight urban areas (Table 2). The relative changes from April 2019 to April 2020 range from -54 % for Paris to -27 % in Bordeaux (Table 2). This relative change over Paris is consistent with the decrease of -52 % described by Levelt et al. (2022) and with the decrease of about -56 % estimated by the tropospheric NO₂ columns measured by two UV–visible Système d'Analyse par Observation Zénithale instruments (SAOZ; Pazmiño et al., 2021).

The temporal variability in the changes from spring 2019 to spring 2020 also differs from one urban area to another. With the exception of Bordeaux and Toulouse, the reductions in TVCDs are higher in April 2020 (Fig. 6, Table 2) than in March 2020 (Fig. 5, Table 2). This is consistent with the fact that the French population was confined only from mid-March (i.e., on 17 March versus the whole of April in 2020).

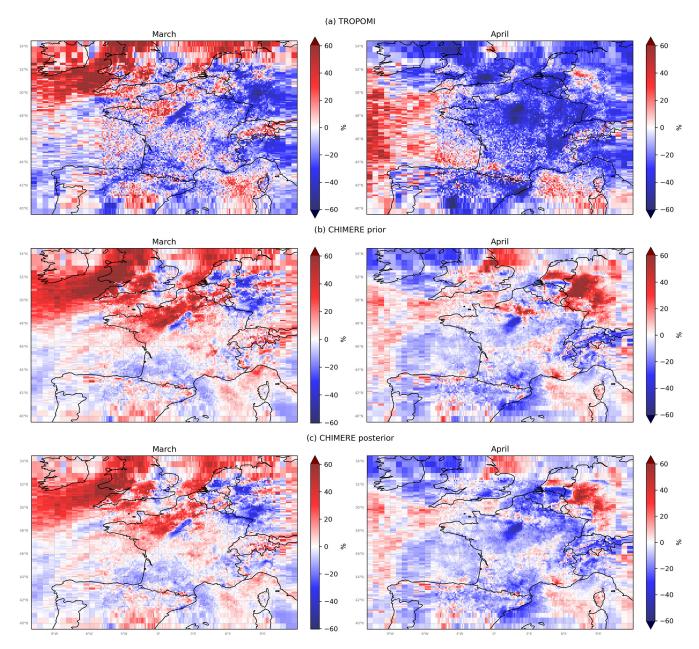


Figure 5. Monthly gridded relative differences between monthly averages of (a) TROPOMI–PAL NO₂ tropospheric columns, (b) CHIMERE prior tropospheric columns and (c) CHIMERE posterior tropospheric columns estimated by the inversions from March/April 2019 to March/April 2020, in % [100 × ($F_{2020} - F_{2019}$)/ F_{2019})].

3.2.2 Various factors contributing to the differences in concentrations from spring 2019 to spring 2020 other than the lockdown

The relative changes in the TROPOMI TVCDs from spring 2019 to spring 2020 are partly due to COVID-19 lockdowns, but they are also driven by the changes in the meteorological and atmospheric chemistry-transport conditions (Menut et al., 2020; Diamond and Wood, 2020; Petetin et al., 2020; Barré et al., 2021; Gaubert et al., 2021; Deroubaix et al., 2021).

The differences from spring 2019 to spring 2020 in the simulated CHIMERE prior NO₂ TVCDs (Fig. 5) are mainly due to changes in the meteorology and to a lesser extent to changes in the model boundary conditions and in the biogenic emissions (Sect. 2.2), as the prior anthropogenic emissions are the same in 2019 and in 2020 (Sect. 2.1).

CHIMERE prior NO₂ TVCDs are 23 % higher in March 2020 than in March 2019 in the northern part of France, except for the plume of Paris. These changes thus have a sign which is opposite to the changes expected from the COVID-

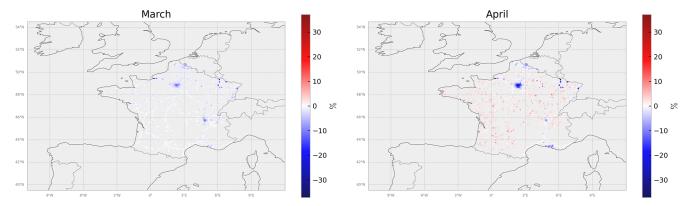


Figure 6. Monthly gridded relative differences between the monthly total (anthropogenic + biogenic) posterior emissions estimated by the inversions from March/April 2019 to March/April 2020, for the months of March/April, using an urban and road land-use proxy (see Fig. C3), in % [$100 \times (F_{2020} - F_{2019})/F_{2019}$].

19 lockdowns. This is in agreement with the analysis of Gaubert et al. (2021), who showed that, when considering only the effect of meteorological variability, the level of NO₂ concentrations would have been high during the 15 March–14 April 2020 period compared to the average NO₂ concentrations in the same period over the 5 previous years (2015–2019) in the northwestern part of France (i.e., the Brittany, Pays de la Loire, Normandy and Hauts-de-France regions; Fig. 5).

CHIMERE prior NO₂ TVCDs are 4 % lower in April 2020 than in April 2019 over almost the entire country (except above Paris and in the Rhône Valley). This is consistent with temperatures above seasonal values (by 3 °C over France, ranking as the third-warmest April on record), and with persistent anticyclonic conditions in April 2020 (MeteoFrance, 2022).

CHIMERE posterior NO₂ TCVDs are about 12 % lower in April 2020 than in April 2019 over almost the entire country (Fig. 5, Table 2); a focus at the city scale shows a decrease of about 9 % over Paris between April 2019 and April 2020, whereas the prior NO₂ TCVDs increase by about 42 % (Table 2).

Finally, the meteorological conditions also have an impact on the availability of the satellite observations. Due to the more favorable meteorological conditions with less cloud coverage leading to an unusual clear-sky period in April 2020 (Gaubert et al., 2021; Deroubaix et al., 2021), there is a higher number of TROPOMI observations in April 2020 than in March 2020 and than in 2019, particularly over the northeastern part of France (Fig. C4), which may allow a better correction of the CHIMERE NO₂ TCVDs for this particular period (Sect. 3.2.3).

3.2.3 Impact on anthropogenic NO_X emissions from spring 2019 to spring 2020

At the French national scale, the total NO_x emission estimates from the inversions present their largest decrease during the first lockdown in March and April 2020 compared to in 2019 (Fig. 4). The emissions between May and September 2020 are similar to those in 2019 (Fig. 4), coinciding with the ease of restrictions. Nevertheless, the decrease in NO_x emissions from March/April 2019 to March/April 2020 of about -2% and -3%, respectively, at the national scale is relatively flatter than the estimations found in the literature. For example, CITEPA estimates reductions of about -12%and -30% (Table 1), respectively, from March/April 2019 to March/April 2020 at the French national scale. Meteorology with a warmer winter can explain part of the changes in emissions between 2019 and 2020 for business sectors such as the residential combustion (Barré et al., 2021; Guevara et al., 2023).

We thus analyze the impact of the COVID-19 policies in terms of differences in retrieved anthropogenic emission estimates from the inversion from spring 2019 to spring 2020. For this, we focus on urban areas, assuming that the emissions in these pixels (see Fig. A1 for the chosen locations) are almost entirely due to anthropogenic activities.

In our inversions, the changes are negative in seven of the eight chosen urban areas (Fig. 6, Table 2), qualitatively consistent with the reduction in the intensity of vehicle traffic (Guevara et al., 2021, 2022). The changes are also negative for urban areas outside France in our domain (Table B2). The impact of the lockdown on anthropogenic NO_x emissions is very different from one urban area to another. These differences between cities cannot be explained by different contributions of industry to NO_x emissions in or around these urban areas, as INS data do not show major differences in terms of sectoral distribution between the main French cities.

The highest reductions are seen in Paris with about -26% in April 2020 compared to April 2019, followed by Lyon

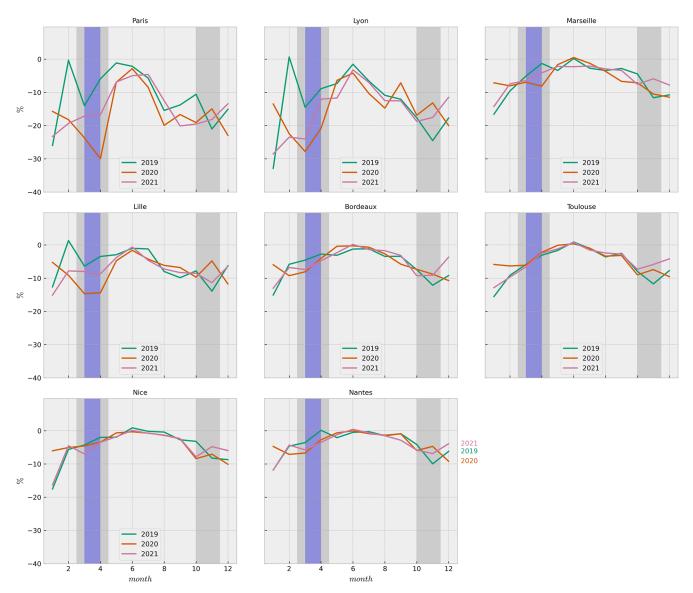


Figure 7. The monthly budget relative differences in anthropogenic NO_x emissions between posterior and prior for the eight French urban areas displayed in Fig. A1, from inversion results for the years 2019 (in green), 2020 (in orange) and 2021 (in pink), in %. Gray shaded areas show the French lockdown periods for the year 2020, and purple shaded areas show the French lockdown period for the year 2021.

(-13%) and Lille (-11%); Table 2). Several urban areas only present a very small drop in emissions in spring 2020 (Table 2, Fig. 7), e.g., the French urban areas Bordeaux, Nice and Nantes. The Toulouse urban area even shows a slight increase (+1%) in April 2020 compared to April 2019 (Table 2, Fig. 7).

These small changes in emissions over Bordeaux, Nice and Toulouse do not seem consistent with the drops in traffic activity estimated by the Centre d'études et d'expertise sur les risques, l'environnement, la mobilité et l'aménagement (CEREMA; CEREMA, 2023). A dedicated study might be necessary to understand the inter-urban variability in detail and is not performed here, as up-to-date local inventories are not available for every city, particularly smaller ones.

The reductions in the anthropogenic NO_x emissions from spring 2019 to spring 2020, while substantial, do not exhibit the same magnitude as the reductions in TROPOMI–PAL NO_2 TVCDs. This is expected because of the non-linearities between anthropogenic NO_x emissions and NO_2 TVCDs but also due to the limitations and assumptions of the inversion itself.

3.2.4 Exploring limitations in our analysis

The discrepancy between the emission changes from 2019 to 2020 between the inversion results and independent estimates from inventories (i.e., CITEPA, CEREMA) can be partly connected to our conservative characterization of the

uncertainties in the prior emission estimates, which ignores potential spatial and temporal correlations in the prior uncertainties in the anthropogenic emissions at the daily and 10 to 50 km resolution. Therefore, the direct information from the satellite observations is not extrapolated by the inversion in space and time via correlations in the **B** matrix; thus the departure of the inversion from the prior emission estimates is highly sensitive to the observation coverage and to its footprint in the emission field.

In this context, the corrections provided by the inversions to the prior emissions can be limited by the fact that the potential of TROPOMI to provide information is hampered by the cloud coverage. When considering the annual to monthly budgets of the emissions over all days (with and without observations), the amplitude of the corrections to the prior estimate of the emissions driven by the satellite observations is artificially decreased by the lack of corrections during days when there are no satellite observations.

We quantify this effect by selecting days when we have at least one super-observation over the pixel of the eight urban areas of interest and extrapolating the retrieved emissions for this subset of days to the whole year, hereafter called "filtered emissions". The reductions in the anthropogenic NO_x filtered emissions from spring 2019 to spring 2020 are almost always higher than for the standard posterior emissions (e.g., -28% and -24% over Paris and over Lyon, respectively, for filtered emissions versus -26% and -13%; Table 3). These results with such a focus on days with the best coverage are in principle much closer to the changes in emissions estimated by CITEPA.

The corrections provided by the inversions to the prior emissions are also highly dependent on the errors associated with the TROPOMI–PAL observations and with the CTM errors in **R**. We illustrate this effect by performing an inversion without model errors in the covariance matrix **R**, giving more weight to the satellite data and regarding the model as perfect. The posterior NO_x emissions retrieved with this error setup at the city scale show a higher reduction from spring 2019 to spring 2020 (Tables 3 and B3), e.g., -31% and -25% over Paris and over Lyon when using all the available observations and -33% and -42% for filtered emissions compared to -26% and -13% for the reference emissions.

Nevertheless, several urban areas still present a very small drop in emissions even when selecting days with observation and regarding the model as perfect to give more weight to the TROPOMI data. This is the case of the French urban areas Bordeaux, Nice and Toulouse with reductions from spring 2019 to spring 2020 lower than 10 % (Table 3).

This may be explained by the strength of the TROPOMI signal over these cities (see Sect. 3.2.1).

Filtered cases are only indicative, as this approach necessitates extrapolating from a notably small sample (days with observations; 16 and 19 d on average for the studied French cities in March and April 2020).

To gain emission representativity, we suggest exploring a fine-tuning of the correlation in **B** (with optionally additional information on NO₂ concentrations like surface stations) to compensate for when TROPOMI data availability is lower. Despite our hypothesis made in Sect. 2.5 on anthropogenic fluxes, this might be necessary in addressing nation-wide NO_x emission monitoring with a high local resolution.

4 Conclusions

We performed a 3-year variational inversion of NO_x emissions from 2019 to 2021 in France at the high resolution of $10\,\mathrm{km} \times 10\,\mathrm{km}$. The TROPOMI–PAL observations were assimilated within the inversion system CIF driving the regional CTM CHIMERE with the MELCHIOR-2 chemical scheme.

The French budget for the years 2019, 2020 and 2021 is about $850 \, \text{kteq} \, \text{NO}_2$. As expected from the implementation of public policies leading to regular reductions in emissions, these national budgets from the inversions are lower than CAMS-REG/INS for 2016, used as the prior in our inversions. In particular, 2020 does not show a clear reduction in emissions compared to 2019: 2020 emissions are higher compared to 2019 in January, February, June and November. These increases dampen the decline in NO_x emissions in March/April 2020 compared to March/April 2019.

We focus on the changes in NO₂ TROPOMI–PAL TVCDs and in anthropogenic NO_x emissions in March and April 2020 compared to 2019 that are due to the lockdowns during the COVID-19 pandemic and to meteorological conditions, including a milder winter (Barré et al., 2021; Guevara et al., 2023). Since inversions mainly detect changes in cities (hotspots), the main changes between March/April 2019 and March/April 2020 are observed at the city scale. However, the impacts of restrictions can vary significantly between different urban areas. Among the eight selected French urban areas, the relative changes between April 2020 and April 2019 in the TROPOMI–PAL observations range from -54%for Paris to -27% in Bordeaux. The highest reductions in anthropogenic NO_x emissions, when we focus on days with observations, are seen in Paris with about -28% in April 2020 compared to April 2019, followed by Lyon (-24%)and Lille (-11%).

Several urban areas, such as Bordeaux, Nice and Toulouse, only show a small decrease in emissions from spring 2019 to spring 2020, even when focusing on the results from the inversion over days with at least one super-observation and even when more weight is given to the TROPOMI data in the inversion process by assuming that the model is perfect. This may be due to weaker TROPOMI signals over these cities, with NO₂ TVCD changes in March/April 2020 compared to March/April 2019 being smaller than those observed in Paris, Lyon and Lille.

Table 3. Changes in CHIMERE total posterior NO_x emissions for the standard (in %) and with a simplified observation error setup (in %), **R** matrix without model error; see details in Sect. 3.2.3 from March/April 2019 to March/April 2020 considering different TROPOMI coverage for the eight French urban areas displayed in Fig. A1.

Urban area	Standard posterior emissions		Filtered posterior emissions		Stand poste emiss (R perfec	rior ions	Filtered posterior emissions (R perfect model)	
	Mar	Apr	Mar	Apr	Mar Apr		Mar	Apr
Bordeaux	-3	-1	- 9	-2	-5	-4	-13	-7
Lille	-9	-11	-16	-11	-10	-15	-17	-13
Lyon	-15	-13	-24	-24	-15	-25	-22	-41
Marseille	-2	-7	-5	-9	-2	-10	-5	-13
Nantes	-3	-3	-2	-7	-5	-5	-6	-11
Nice	0	-1	-4	-3	-1	-4	-5	-7
Paris	-11	-26	-15	-28	-13	-31	-15	-32
Toulouse	0	+1	-1	+1	-1	+1	-4	-1

The inversion results show significant decreases in the emissions from the largest cities in France from 2019 to 2020, but this decrease remains lower than that documented in most studies on this topic and by independent inventories (i.e., CITEPA and CEREMA). This can be partly connected to our conservative characterization of the prior emission uncertainties, since we ignore potential spatial and temporal correlations in the prior uncertainties in the anthropogenic emissions, so that the departure from the prior emission estimates in the inversion is highly sensitive to the observation coverage and emission footprints. The corrections provided by the inversions to the prior emissions can indeed be limited by the cloud coverage affecting the TROPOMI observations and by errors in the TROPOMI data and in the CTM. Notably, when TROPOMI observations are unavailable, the correction of prior emissions in the unconstrained pixels through inversions is null, resulting in posterior emissions remaining close to the prior ones. Hence, the aggregation of posterior emissions into monthly or yearly budgets leads to a dampening of the signal provided by TROPOMI. In order to better emphasize the direct information from the satellite observations, some of our analyses of the local urban emissions are focused on days with at least one observation in the targeted pixels, yielding a characterization of the COVID-19 effects which is more consistent with the changes in emissions estimated by CITEPA.

To explore the impact of the setup of error covariance matrices, we considered inversions without error models in the covariance matrix \mathbf{R} , giving more weight to satellite data. In this case, the posterior NO_x emissions at the city scale exhibit their highest reductions between spring 2019 and spring 2020 (e.g., -31% and -25% over Paris and over Lyon, respectively, considering all days within the month). Assimilating the observations without accounting for the model errors may lead to over-fitting and thus project these errors onto the

emission estimates. However, the weight of **R** in our inversions may have to be re-assessed with regard to the relatively conservative option that we use here to assign observation error to the super-observations. A finer assessment would require a good knowledge of the share of retrieval errors between random noise without spatial correlations and more systematic errors with spatial correlations, as well as the typical length scales of such spatial correlations, which are currently challenging to derive (Miyazaki et al., 2012; Boersma et al., 2016; Lambert et al., 2023).

In the absence of alternative information, like new measurements from benchmark cities, reconciling our results with the existing literature would prompt a re-evaluation of the observation errors in our inversion and a potential reassessment of our model error definition. The information contained in TROPOMI TCVDs cannot be fully exploited in our inversion setup to get constraints on diffuse emissions, e.g., in French rural areas.

For improving diagnostics of monthly to annual emissions at national scale for emission hotspots, addressing challenges arising from satellite coverage gaps involves introducing horizontal, temporal and sectoral correlations in the covariance matrix **B** of the uncertainty in the gridded inventories with hourly variations that are used as prior estimates of the emissions by the inversions. Such a characterization would support the extrapolation in space and time of the information obtained locally for some days from the satellite observation. However, obtaining such correlations at high resolution poses a substantial challenge. The usual correlation models based on assumptions of isotropy, homogeneity in space and time, and decrease as a function of distance and time should poorly match the actual derivation and structures of gridded inventories convolved with typical temporal cycles at diurnal to seasonal scales, which is why a conservative configuration was used for the **B** matrix in this study (Super et al., 2020).

The challenge is exacerbated when tackling a period such as 2019-2021, with lockdown measures in response to the COVID-19 crisis highly impacting the emissions and thus the structures of uncertainties in the emission inventories over large spatial scales but limited periods. Exploring corrections to parameters underlying inventories, such as the Fossil Fuel Data Assimilation System (FFDAS) for CO₂, may actually support a cleaner extrapolation. Due to the current lack of knowledge about the statistics of the uncertainties in the gridded inventories used for the inversions, a stepwise approach is probably needed to tackle this general problem, gradually including some temporally varying spatial and temporal correlations in B and, in parallel, increasing efforts to diagnose these uncertainties. If achieved, these additions would assist the inversion system in implementing nationwide emission adjustments.

Furthermore, the incorporation of a hypothetical geostationary satellite (such as Sentinel-4) in conjunction with ground-based monitoring stations could enhance the temporal resolution and enable the capture of daily NO_x cycles while also increasing the sensitivity of the satellite data near the surface.

Appendix A: Metropolitan masks

The list of all 15 urban areas and their masks is displayed in the following figure. Each mask is made out of $10\,\mathrm{km} \times 10\,\mathrm{km}$ pixels.

Table A1. Cities studied in the paper and their areas in pixels with their population density according to EuroStat (statistics for 2019).

Urban area	Area (number of $10 \text{km} \times 10 \text{km}$ pixels)	Population density (inhab.km ⁻²) EuroStat
Barcelona	1	723
Bordeaux	1	164
Frankfurt	1	633
Lille	1	455
London	6	1435
Lyon	1	580
Marseille	4	289
Milan	3	1591
Nantes	1	209
Nice	1	253
Paris	6	1025
Rotterdam	2	1092
Rhein-Ruhr Süd	121	-
Toulouse	2	221
Turin	1	329

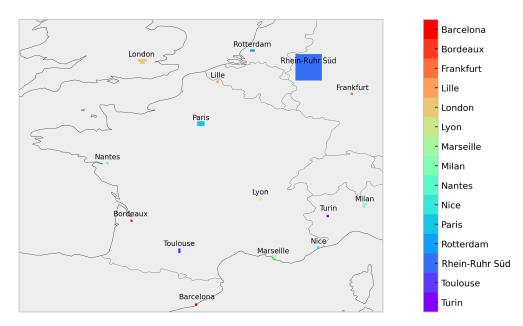


Figure A1. Urban areas studied.

Appendix B: Additional tables

Table B1. Total prior and posterior NO_x emission budgets in kteq NO_2 and their relative differences in % [$100 \times (posterior - prior)/prior$] in Paris for different periods. Columns "Posterior $year_n - year_{n-1}$ " show the relative difference between $year_n$ and $year_{n-1}$ posterior fluxes in % [$100 \times (F_{year_n} - F_{year_{n-1}})/F_{year_{n-1}}$].

Domain or	Period	CAMS-REG/INS + MEGAN inventories	Posterior in	n 2019	Posterior in 2020		Posterior in	n 2021	Posterior 2020–2019	Posterior 2021–2020
region		kteqNO ₂	$kteq NO_2$	%	kteq NO ₂	%	kteq NO ₂	%	%	%
Paris	Annual	27	24	-11	23	-17	23	-15	-6	+2
	Spring (MAM)	7	7	-7	6	-21	6	-14	-15	+9
	Mar	3	2	-14	2	-24	2	-17	-11	+9
	Apr	3	2	-6	2	-30	2	-17	-25	+19
	May	2	2	-1	2	-7	2	-7	-7	+1
	Nov	2	2	-21	2	-15	2	-18	+8	-4

Table B2. Changes in TROPOMI–PAL and CHIMERE NO_2 tropospheric columns (in %) and changes in total posterior NO_x emissions (in %) from March/April 2019 to March/April 2020 for the selection of urban areas displayed in Fig. A1.

Urban area	TROPOMI TVCD		CHIMERE prior TVCD		CHIM posterior		Emissions posterior	
	Mar	Apr	Mar	Apr	Mar	Apr	Mar	Apr
Barcelona	-42	-72	-13	-12	-20	-29	-4	-8
Bordeaux	-34	-27	-14	+23	-15	+21	-3	-1
Frankfurt	-54	+1	-18	+58	-25	+39	-7	-10
Lille	-15	-35	+19	-10	+4	-28	-9	-11
London	-7	-38	+4	+5	-3	-14	-5	-12
Lyon	-27	-43	+13	+18	+1	+1	-15	-13
Marseille	-20	-27	-2	+46	-9	+19	-2	-7
Milan	-30	-21	+28	+2	-8	-13	-13	-2
Nantes	+2	-45	+16	-2	+14	-7	-3	-3
Nice	-28	-37	0	+7	-3	+2	0	-1
Paris	-41	-54	-4	+42	-12	-9	-11	-26
Rotterdam	-35	-17	-15	+39	-19	+12	-5	-6
Rhein-Ruhr Süd	-25	-9	-21	+42	-24	+8	-16	-26
Toulouse	-41	-31	-23	-23	-22	-22	0	+1
Turin	-33	-41	+17	-5	+1	-18	-1	-3

Table B3. Changes in CHIMERE total posterior NO_x emissions for the standard (in %) and with a simplified observation error setup (in %), **R** matrix without model error; see details in Sect. 3.2.3 from March/April 2019 to March/April 2020 considering different TROPOMI coverage for the selection of urban areas displayed in Fig. A1.

Urban area	Standard posterior emissions		post	ered erior sions	Stand poste emiss (R perfec	rior ions	Filtered posterior emissions (R perfect model)	
	Mar	Apr	Mar	Apr	Mar	Apr	Mar	Apr
Barcelona	-4	-8	-7	-16	-6	-13	-10	-23
Bordeaux	-3	-1	-9	-2	-5	-4	-13	-7
Frankfurt	-7	-10	-12	-17	-11	-21	-20	-36
Lille	-9	-11	-16	-11	-10	-15	-17	-13
London	-5	-12	-3	-18	-5	-16	-2	-23
Lyon	-15	-13	-24	-24	-15	-25	-22	-41
Marseille	-2	-7	-5	-9	-2	-10	-5	-13
Milan	-13	-2	-22	-11	-14	-4	-22	-14
Nantes	-3	-3	-2	-7	-5	-5	-6	-11
Nice	0	-1	-4	-3	-1	-4	-5	-7
Paris	-11	-26	-15	-28	-13	-31	-15	-32
Rotterdam	-5	-6	-6	-10	-5	-9	-3	-12
Rhein-Ruhr Süd	-16	-26	-16	-26	-16	-31	-15	-31
Toulouse	0	+1	-1	+1	-1	+1	-4	-1
Turin	-1	-3	-10	-4	+1	-6	-7	-8

Appendix C: Additional figures

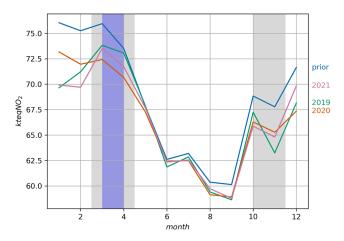


Figure C1. Monthly anthropogenic NO_x emissions in France as estimated by CAMS-REG/INS in 2016 (in blue) and by the inversions for the years 2019 (in green), 2020 (in orange) and 2021 (in pink), in kteq NO_2 month⁻¹. Gray shaded areas show the French lockdown periods for the year 2020, and the purple shaded area shows the French lockdown period for the year 2021.

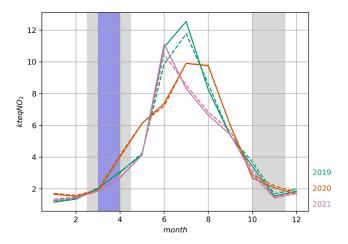


Figure C2. Monthly biogenic NO_x emissions in France as estimated by the MEGAN inventory (dotted lines) and by the inversions for the years 2019 (in green), 2020 (in orange) and 2021 (in pink), in kteq NO_2 month⁻¹. Gray shaded areas show the French lockdown periods for the year 2020, and the purple shaded area shows the French lockdown period for the year 2021.

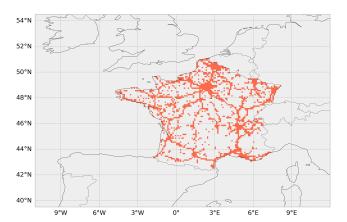


Figure C3. Anthropogenic mask used over France using an urban and road land-use proxy.

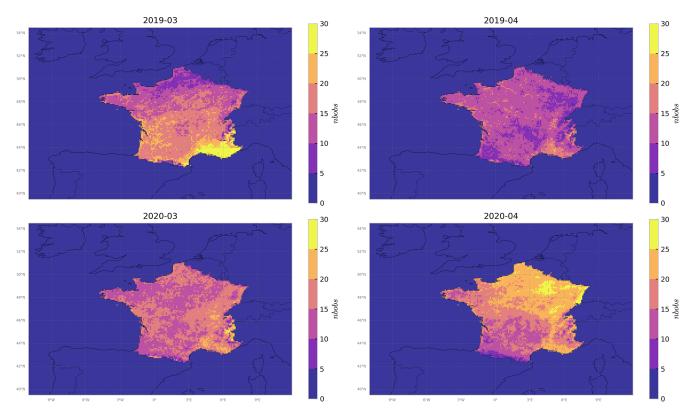


Figure C4. Number of TROPOMI–PAL super-observations in the domain in March/April 2019 and March/April 2020.

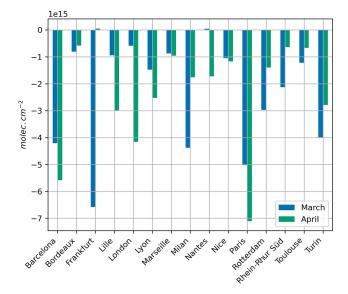


Figure C5. TROPOMI absolute NO_2 TVCD difference between March/April 2019 and March/April 2020 in molec. cm⁻² for the selection of urban areas displayed in Fig. A1.

Code availability. The CHIMERE code (Menut et al., 2013; Mailler et al., 2017) is available at http://www.lmd.polytechnique. fr/chimere/ (CHIMERE, 2024). The CIF code (Berchet et al., 2021) is available at https://doi.org/10.5281/zenodo.6304912 (Berchet et al., 2022).

Data availability. The reprocessed TROPOMI–PAL NO2 dataset (Eskes et al., 2021; S5P-PAL Data Portal, 2024) is available at https://dataspace.copernicus.eu/explore-data/data-collections/sentinel-data/sentinel-5p (Sentinal-5P, 2024). CAMS-REG (Kuenen et al., 2022) is available upon request from TNO (contact: Hugo Denier van der Gon, hugo.deniervandergon@tno.nl). The CITEPA monthly budgets for France (contact: Ariane Druart, ariane.druart@citepa.org) are available at https://www.citepa.org/fr/barometre/ (André et al., 2024). The INS inventory is available at http://emissions-air.developpement-durable.gouv.fr/indexMap.html (INERIS, 2024).

Author contributions. RP, AFC, GB, IP, AB, EP, GD, AC, DS and GS conceptualized the study. GS produced the prior emissions map. HE produced the satellite data. RP and AFC carried out inversions and data analysis. All co-authors contributed to the design of the study and to writing the paper.

Competing interests. The contact author has declared that none of the authors has any competing interests.

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