



Quantification of carbon monoxide emissions from African cities using TROPOMI

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Abstract. Carbon monoxide (CO) is an air pollutant that plays an important role in atmospheric chemistry and is mostly emitted by forest fires and incomplete combustion in, for example, road transport, residential heating, and industry. As CO is co-emitted with fossil fuel CO₂ combustion emissions, it can be used as a proxy for CO₂. Following the Paris Agreement, there is a need for independent verification of reported activity-based bottom-up CO₂ emissions through atmospheric measurements. CO can be observed daily at a global scale with the Tropospheric Monitoring Instrument (TROPOMI) satellite instrument with daily global coverage at a resolution down to 5.5 × 7 km². To take advantage of this unique TROPOMI dataset, we develop a cross-sectional flux-based emission quantification method that can be applied to quantify emissions from a large number of cities, without relying on computationally expensive inversions. We focus on Africa as a region with quickly growing cities and large uncertainties in current emission estimates. We use a full year of high-resolution Weather Research and Forecasting (WRF) simulations over three cities to evaluate and optimize the performance of our cross-sectional flux emission quantification method and show its reliability down to emission rates of 0.1 Tg CO yr⁻¹. Comparison of the TROPOMI-based emission estimates to the Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa (DACCIWA) and Emissions Database for Global Atmospheric Research (EDGAR) bottom-up inventories shows that CO emission rates in northern Africa are underestimated in EDGAR, suggesting overestimated combustion efficiencies. We see the opposite when comparing TROPOMI to the DACCIWA inventory in South Africa and Côte d’Ivoire, where CO emission factors appear to be overestimated. Over Lagos and Kano (Nigeria) we find that potential errors in the spatial disaggregation of national emissions cause errors in DACCIWA and EDGAR respectively. Finally, we show that our computationally efficient quantification method combined with the daily TROPOMI observations can identify a weekend effect in the road-transport-dominated CO emissions from Cairo and Algiers.

1 Introduction

Carbon monoxide (CO) is an air pollutant that is mostly emitted by anthropogenic sources. It is a product of incomplete combustion in, for example, road transport, residential heating, industry, and forest fires (Zhong et al., 2017). CO is a precursor of ozone, and because it reacts with the hy-

droxyl radical (OH), its presence effectively increases the atmospheric lifetime of methane (Daniel and Solomon, 1998; Jacob, 1999; Wuebbles and Hayhoe, 2002). The concentration of CO is therefore important for climate modeling. Furthermore, as many processes that emit CO also emit carbon dioxide (CO₂), knowledge of CO emission rates can provide additional information about CO₂ emissions (Wu et al., 2022;

Park et al., 2021). The Intergovernmental Panel on Climate Change (IPCC) identified a need for independent verification of the reported greenhouse gas emissions through measurements (IPCC, 2019). From space this is challenging for CO₂, as its long atmospheric residence time results in high background concentrations, making it hard to detect emissions. For this reason, measuring the “short lived” CO can be a useful alternative (Silva et al., 2013; MacDonald et al., 2023). We present a method to quantify CO emission rates over cities in Africa using TROPospheric Monitoring Instrument (TROPOMI) satellite observations.

As transport (23 %) and residential heating (35 %) are key contributors to total anthropogenic CO emissions (Zhong et al., 2017), cities are an important source of CO. In Africa, the contributions of transport and residential heating are estimated at 27 % and 38 % of anthropogenic CO emissions respectively in the DICE-Africa inventory (Marais and Wiedinmyer, 2016) and 17 % and 72 % in the Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa (DACCIIWA) inventory (Keita et al., 2021). The importance of these two sectors is further confirmed by a large number of ground-based measurements specifically aimed at traffic (Diab et al., 2005; Lindén et al., 2008; Zakari et al., 2020; Doumbia et al., 2021) and domestic heating (Havens et al., 2018; Kansime et al., 2022; Saleh et al., 2023) that show CO concentrations in African cities exceeding air quality guidelines by the World Health Organization. Urbanization scenarios predict a growth in both the number of megacities and their populations, leading to larger emission rates and increased health risks. Africa is predicted to have a large urbanization rate in the coming years. Hoornweg and Pope (2017) predict the continent to house 5 of the 10 largest cities by 2100, compared to 1 of 10 today. Africa is also a region for which relatively large uncertainties are present in emission inventories, as only a few are dedicated to the region (Keita et al., 2021). Current emission inventories are based on so-called bottom-up methods, where emissions are estimated by combining activity data (e.g., national fuel consumption statistics) with emission factors and spatially distributing the emission estimates using proxies like population density (Janssens-Maenhout et al., 2019). These bottom-up methods are also used to report country-level greenhouse emission estimates to the United Nations Framework Convention on Climate Change (UNFCCC). However, lack of detailed data results in large uncertainties (Macknick, 2011; Cai et al., 2019; Oda et al., 2019).

Independently of bottom-up methods, emissions can also be estimated by top-down methods, where atmospheric concentrations are measured and used to infer the corresponding emission rates. Multiple studies have investigated urban CO emissions using ground-based measurements (Badarinath et al., 2007; McKain et al., 2012; Bi et al., 2022). Many studies have also shown the capability of satellite measurements for this specific task (Borsdorff et al., 2020; Tian et al., 2022a;

Plant et al., 2022; Wu et al., 2022). For CO, the TROPospheric Monitoring Instrument (TROPOMI) on ESA’s Sentinel 5 precursor satellite is of particular interest (Veefkind et al., 2012). It was launched in 2017 and provides daily global coverage with a resolution of $5.5 \times 7 \text{ km}^2$, which makes it suited to investigate city emissions worldwide.

An advantage of polar-orbiting satellites is their ability to monitor the entire globe. However, most satellite-based studies of CO so far have focused either on regional inversions (Yumimoto et al., 2014; Qu et al., 2022) or on trends in concentrations (Lama et al., 2020; Park et al., 2021; Hedelius et al., 2021), while only a few studies have tried to quantify emissions from individual cities or point sources (Dekker et al., 2017; Borsdorff et al., 2020). These urban emission quantifications use atmospheric inversions, which require computationally expensive high-resolution simulations with chemical transport models (CTMs). Although inversions are able to get relatively accurate emission estimates, they are difficult to apply to a large number of sources. To take full advantage of the TROPOMI data, we adjust the mass balance cross-sectional flux (CSF) method, originally developed for high-resolution point-source quantifications, to be used with TROPOMI data over urban areas. After evaluating the method using atmospheric transport simulations, we use it to estimate emissions from the largest cities in Africa.

2 Data and methods

This section describes the different data products used in the development of the cross-sectional flux method and the simulations that were used to calibrate the model and evaluate its performance using an observing system simulation experiment (OSSE). An OSSE is an experiment where a model or method is applied to synthetic data to evaluate the benefit of using this data and/or method. Which for this work means evaluating whether the CSF can be used to correctly estimate emissions from TROPOMI-like synthetic data. Figure 1 shows the roles of the different data products that are used and further described in Sect. 2.1 to 2.6. In addition, in Sect. 2.6 we show that the CSF method can be successfully applied to simulated data.

2.1 TROPOMI carbon monoxide data product

TROPOMI provides total column carbon monoxide concentrations with daily global coverage at 13:30 local time using the shortwave-infrared band (SWIR) at 2305–2385 nm (Veefkind et al., 2012). From the spectral signal, the CO concentration is inferred using the shortwave-infrared CO retrieval (SICOR) algorithm (Borsdorff et al., 2018). We use 3 years of data (2019–2021) from the operational data product (Landgraf et al., 2018). To assure high quality data, all pixels with a TROPOMI quality flag below 0.7 are removed, leaving data that are cloud-free or only have low altitude clouds. The CO concentration over cloud-free water surfaces

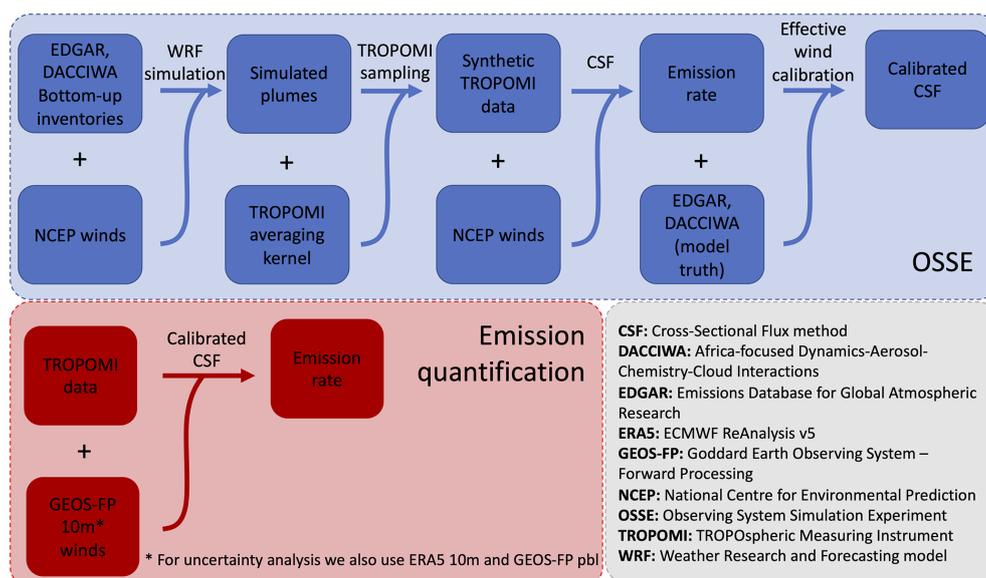


Figure 1. Schematic description of the use of the different data products within the OSSE and the subsequent emission quantification using TROPOMI data. The data products are discussed in Sect. 2.1–2.6. First, the CSF is applied to simulated synthetic plumes in order to determine appropriate values for the various parameters used in our method. Second, an effective wind is calibrated by using the known emission rates of the simulated plumes following the procedure by Varon et al. (2018). The CSF, now calibrated on the synthetic plumes, is subsequently applied to satellite data to estimate emission rates of African cities.

is difficult to retrieve due to the low intensity of reflected light; therefore, we only use observations with a quality flag equal to 0.7 (low altitude clouds) over water. The resulting dataset shows good agreement with ground-based measurements, with a mean difference per station of $2.45 \pm 3.38\%$ to the unscaled Total Carbon Column Observing Network (TC-CON, Wunch et al., 2011) columns and $6.5 \pm 3.54\%$ to the Infrared Working Group of the Network for the Detection of Atmospheric Composition Change (NDACC-IRWG, 2023) measurement stations (Sha et al., 2021).

2.2 EDGAR and DACCIWA bottom-up inventories

We use two different bottom-up inventories to compare the TROPOMI emission estimates: the Emissions Database for Global Atmospheric Research (EDGAR) version 5 (Oreggioni et al., 2021) and the Africa-focused Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa (DACCIWA) inventory (Keita et al., 2021). These inventories are also used in our atmospheric transport simulations (Sect. 2.3) to simulate TROPOMI observations. Both inventories provide yearly gridded emission rates at 0.1° resolution up to 2015. Due to its global scope, the EDGAR inventory relies mostly on international statistics and spatial proxies combined with national data, while its emission factors are based on IPCC methodology for greenhouse gases (Eggleston et al., 2006) and the EMEP/EEA emission inventory guidebook for air pollutants (Nielsen, 2013). The DACCIWA inventory provides emission rates over the

African continent, ranging from -35 to 38° latitude and -25.5 to 63.5° longitude. It uses similar international data but is supplemented by local measurements of emission factors and data from local authorities (Keita et al., 2021). As the DACCIWA inventory characterizes emission from fewer (sub)sectors than EDGAR, we merge different sectors in EDGAR to match those used in the DACCIWA inventory to make them intercomparable. When reporting urban emissions from EDGAR and DACCIWA, we sum emissions over the pixel closest to the city center, its eight neighbors, and all directly attached pixels where the population density exceeds the surroundings by 1.8 standard deviation. Changing the city mask to $0.3 \times 0.3^\circ$ or $0.7 \times 0.7^\circ$ boxes changes the emissions by 10%–20% for 17 out of 29 cities in EDGAR and 16 out of 29 cities in DACCIWA. Although larger deviations up to 50% in densely populated areas like South Africa are observed, the observed patterns discussed in Sect. 3 are valid for these masks as well.

2.3 WRF chemical transport model

To test and calibrate our emission quantification approach, we apply our CSF method to simulated TROPOMI data for three urban areas. We use the Weather Research and Forecasting (WRF) chemical transport model version 4.1 (Powers et al., 2017) to simulate column CO mixing ratios over Cairo (Egypt), Bamako (Mali), and Lagos (Nigeria) for 2019, using December 2018 as the spin-up month. These three African cities form a diverse set, with Cairo next to the Nile river,

Bamako at the boundary of the Sahara desert, and Lagos at the coast of the Atlantic Ocean. We simulate CO as an inert tracer and drive the simulations with meteorological fields from the National Center for Environmental Prediction (NCEP, 2000). All simulations have three-layer nested domains, where the outer domain covers $2673 \times 2673 \text{ km}^2$ at a resolution of 27 km; the middle and inner domains cover $891 \times 891 \text{ km}^2$ at 9 km resolution and $315 \times 315 \text{ km}^2$ at 3 km resolution respectively (Fig. 2). Initial and 6-hourly boundary conditions to capture the background CO are taken from the Copernicus Atmosphere Monitoring Service (CAMS) at $0.25^\circ \times 0.25^\circ$ resolution (Inness et al., 2015). The resulting background is scaled to match the mean background observed by TROPOMI over the full year. We use emissions from the global Emissions Database for Global Atmospheric Research (EDGAR) version 5 and the Africa-focused Dynamics–Aerosol–Chemistry–Cloud Interactions in West Africa (DACCIWA) inventory distributed across the vertical model levels according to the sector-specific vertical profiles provided by Bieser et al. (2011). Typical injection heights for CO emissions from transport and the residential sector are 0–20 m, while emissions from industry are typically injected into the atmosphere at 100–200 m (Bieser et al., 2011). City-specific hourly, daily, and monthly temporal profiles for each emission sector are taken from Guevara et al. (2021). To maintain flexibility over model output, the different sectors in the emission inventories (19 for EDGAR and 6 for DACCIWA) are simulated separately. We sample the model output (at the TROPOMI overpass time) to facilitate comparison to the TROPOMI carbon monoxide data as discussed in detail in Sect. 2.6.

While EDGAR and DACCIWA only include the primary production of CO, the concentrations observed by TROPOMI include CO from secondary production as well. CO is produced by oxidation of volatile organic compounds (VOCs), with methane as the main contributor (Rozante et al., 2017). Mixing ratios of non-methane volatile organic compounds (NMVOCs) observed in urban locations are typically of the order of 10^{-3} – 10^{-2} [NMVOC]/[CO] (Von Schneidmesser et al., 2010). Dekker et al. (2019) showed that chemical production of CO by methane and NMVOC over cities only contributes 4% to the total CO signal, justifying the simulation of CO as an inert tracer in our approach. Due to the 10-year atmospheric lifetime of methane, its contribution to CO production will result in a uniform concentration (Park et al., 2013) that is subtracted with the background. NMVOCs have lifetimes of 0.6–10 d (Guo et al., 2007) that are much shorter than the lifetime of CH_4 , but due to their low urban mixing ratios ($\sim 1\%$), their effect on the estimated emission rate is much smaller than the reported uncertainty of the CSF. This is consistent with the observation that the emission estimates of individual transects (that span a timescale of up to ~ 10 h) are stable and do not increase with increasing distance from the city (Sect. 2.4).

2.4 Cross-sectional flux method

The cross-sectional flux method (CSF) has been shown to be an effective way to quantify emission rates of plumes observed by satellites (Varon et al., 2018, 2020; Sadavarte et al., 2021b; Tian et al., 2022b). It is based on the continuity equation, which relates the flux through a closed surface to the associated emission rate:

$$Q = \oint U_{\perp} \Delta\Omega dA, \quad (1)$$

where Q (kg s^{-1}) is the emission rate, U_{\perp} (m s^{-1}) is the wind speed perpendicular to the closed surface, $\Delta\Omega$ (kg m^{-3}) is the enhancement at the closed surface, and dA (m^2) is a surface element. As illustrated in Fig. 3a, the plumes have a distinct direction as they move with the wind, they are very directional, and it suffices to integrate over perpendicular transects that cover the entire plume width (Fig. 3b). Equation (1) can then be rewritten as

$$Q = \int U_{\perp}(x, y) \Delta\Omega(x, y) dy, \quad (2)$$

with x , y coordinates along and perpendicular to the plume respectively as in Varon et al. (2018). Assuming a constant emission rate, transects at different distances downwind of the source should yield the same emission quantification and can be averaged to make the method more robust.

We then optimize the implementation of the CSF on TROPOMI data for city-like sources. Figure 3a shows a CO plume observed by TROPOMI over Cairo on 7 April 2019. As expected, the plume follows the 10 m wind direction given by NASA/GMAO GEOS-FP reanalysis data (Molod et al., 2012). We start by determining the city center for our purposes defined as the location at the center of the urban emissions and therefore best representing the origin of the city's total emission. The location of the center is determined by taking the weighted average position of the pixels in DACCIWA that are part of the city mask introduced in Sect. 2.2. Weights of the pixels are equal to their emission rate. By determining the city location using the emission inventory, we ensure that we are comparing similar regions when we compare our satellite-based emission estimates with the emission inventories in Sect. 3. To make sure the entire plume is downwind, we start the transects of our CSF 0.1° upwind of this city center. As the wind direction is an important source of uncertainty, the downwind direction can not be solely based on the GEOS-FP reanalysis wind data. Instead, following Sadavarte et al. (2021a), we infer the wind direction from the satellite observations by selecting the direction of the highest mean downwind concentration within 90° of the reanalysis wind direction. To do so, we calculate the mean downwind concentration over 180 boxes (0.1° width and 0.4° length) rotated at 1° intervals and pick the direction with the highest downwind concentration. In the absence of a clear plume, this method would create a positive bias, as it would select

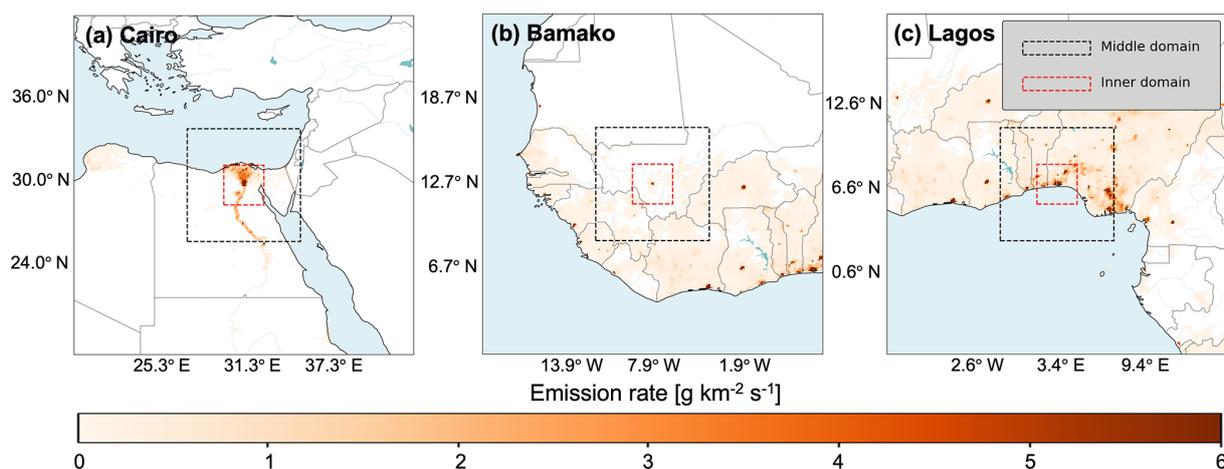


Figure 2. Domain setup of the WRF simulations over Cairo (Egypt), Bamako (Mali), and Lagos (Nigeria). The inner domain (red) spans $315 \times 315 \text{ km}^2$ around the city at 3 km resolution. The middle (black) and outer (full figure) domain cover 893×893 and $2673 \times 2673 \text{ km}^2$ at 9 and 27 km resolution respectively. All panels show the emission rates from the DACCIIWA inventory.

the highest enhancement in the noise. Therefore, we use the reanalysis wind direction if the mean enhancement does not exceed 5 ppb. To calculate CO enhancements, we subtract a background calculated as a mean upwind concentration over a $0.4^\circ \times 0.4^\circ$ square starting 0.3° upwind from the city center (Fig. 3b). If this box contains fewer than five valid pixels, we extend it symmetrically with two arcs of a circle of 10, 20, 45, up to 60° until there are at least five TROPOMI pixels in the background region (Fig. 3c). We use extension in an arc-like fashion rather than increasing the size of the square to be able to get estimate background values for coastal cities. Retrievals over water are only possible if there are clouds present; hence, increasing the size of the background square upwind could result in background pixels far away from the city that are not representative for the local background.

After determining the initial direction of the plume, we need to better capture the shape of the plume to draw transects perpendicular to the plume. The shape of the plume is determined in two steps. First, we select all pixels in a downwind box (0.3° width, 0.8° length) that exceed the mean concentration in the surrounding $3^\circ \times 3^\circ$ area by more than 1.8 standard deviations; these pixels are referred to as the spline mask. We then fit a 2D-spline (0.8° length) through the resulting spline mask. If, due to a lack of signal or missing pixels, the spline mask contains fewer than three pixels, a spline fit is unlikely to capture the true plume shape, and we use a straight line in the (optimized) wind direction instead (Fig. 3c). The transects (0.4° width) are drawn perpendicular to the spline, separated by 0.04° . The transects have a larger width than the box used to determine the spline mask to ensure the transects cover the entire plume width. All pixels overlapping with the transects are used in the emission quantification. We also include days where no clear plume is visible to avoid systematic overestimation of the average emission rate. Using Eq. (2), an emission estimate can be de-

rived for every transect. We stop drawing transects when the emission rate estimates of two consecutive transects are more than 1 standard deviation below the mean estimate of the earlier transects, indicating the end of the plume. Transects with less than 70 % pixel coverage are removed from the estimate, as they will not have a complete integral, resulting in underestimated emissions.

Contrary to studies using high-resolution satellites (Varon et al., 2018, 2020), the plumes observed with TROPOMI cover distances over which there can be significant fluctuations in wind speed and direction. We therefore use the wind speed at each transect instead of a single wind speed for the entire plume. The wind speed at the transects is calculated in two steps. First, a wind speed is calculated for each TROPOMI pixel by interpolation of the reanalysis wind product. Second, the wind speed for each transect is determined by taking the average wind speed of the overlapping TROPOMI pixels, weighted by the length of the overlap. Similar to trends observed in Sadavarte et al. (2021b), the first two transects are found to have roughly 30 % lower emissions than the transects further away, which have a stable mean emission rate. This pattern is consistent across the cities investigated. One reason is that the early plume only captures part of the city's emissions, another explanation is that the associated pixels might see a partial-pixel absorption saturation effect (Pandey et al., 2019). Incorporating the first two transects would result in an average underestimation of emissions by 8 %. We therefore remove the first two transects from the emission estimation.

On the spatial scale relevant to plumes observed by TROPOMI, there can be contamination of the city signal by carbon monoxide produced by open fires (e.g., agricultural fires or wildfires). CO enhancements caused by open fires can result in overestimation of either the background or the downwind urban enhancement, depending on their location.

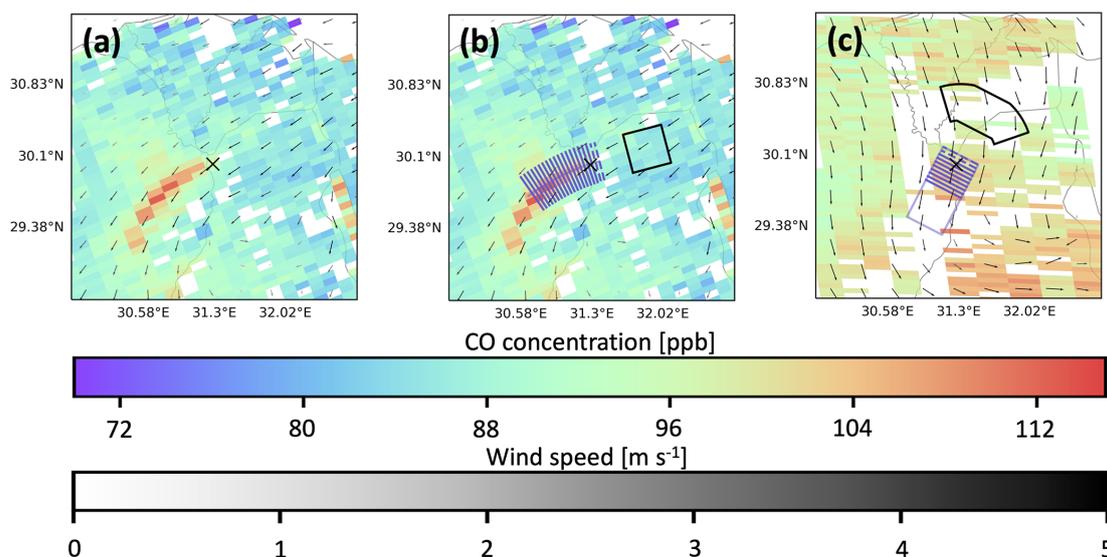


Figure 3. Example of how the cross-sectional flux transects perpendicular to the plume are drawn. **(a)** TROPOMI data over Cairo on 7 April 2019. The city center is shown with a cross and taken from the DACCIWA emission inventory. The arrows show direction and magnitude of GEOS-FP 10 m winds at their native $0.25^\circ \times 0.3125^\circ$ resolution (Molod et al., 2012). **(b)** Pixels downwind of the city that surpass the regional background by more than 1.8σ form a plume mask through which a 2D spline is fitted (grey line). The transects used for quantification (purple lines) are drawn perpendicular to the spline fit starting 0.1° upwind of the city center. The first two transects are dashed to reflect that they are not used for the emission quantification. The background is estimated over the black $0.4^\circ \times 0.4^\circ$ box upwind. **(c)** TROPOMI observation over Cairo on 27 March 2020. Due to a lack of coverage there are insufficient pixels to generate a reliable spline mask. A rectangular box (grey) is therefore used to draw transects instead. The basic background is extended symmetrically, with circle arcs to compensate for a lack of coverage upwind.

To avoid this, days with considerable CO contributions from open fires have been removed from our estimates. These days were selected based on the fire emission data from the Global Fire Assimilation System fire emission (GFAS) inventory (Kaiser et al., 2012) that is based on satellite measurements of fire radiative power. Days with cumulative fire emissions over 57 Mg h^{-1} (equivalent to 0.5 Tg yr^{-1}) within 1.5° from the city center are removed. Additionally, days with strong burning events closer to the city (23 Mg h^{-1} within a 0.75° radius) are removed as well (Appendix B). Although the change in emission rate by this filtering is limited for most cities, the filter can change estimated emission rates by up to 47 %, as seen in Lusaka (Zambia).

2.5 Uncertainty analysis

To estimate the uncertainty of the estimated emission rates, we compile an ensemble of emission estimates for each city. We generate the ensemble by varying parameters of the quantification method such as the wind database used. For example, we vary parameters such as the number of transects and the distance of the background box. To incorporate the uncertainty on the wind data, we use our method with GEOS-FP 10 m altitude winds, GEOS-FP planetary boundary layer (PBL) averaged winds (Molod et al., 2012) as well as 10 m altitude winds from the ERA5 product, provided by the European Centre for Medium-Range Weather Forecasts

(ECMWF; Hersbach et al., 2020). A complete list of the varied parameters and their ranges can be found in Appendix A. For each city, the spread in the resulting ensemble is reported as uncertainty.

2.6 Calibration and validation

This section describes the application of the CSF to simulated CO column mixing ratios. The simulations are used to determine parameter settings (e.g., spline length and transect width) and to calibrate an effective wind (Varon et al., 2018) for TROPOMI-sized pixels. In addition, the simulations are used to evaluate how well the CSF can quantify emission rates of simulated plumes.

As simulated (and TROPOMI observed) plumes stay within the inner domain, only the inner domain is used to test the performance of the CSF. A set of synthetic TROPOMI observations is created by sampling the simulation output over the TROPOMI footprints, applying its averaging kernel, selecting pixels based on quality value as discussed in Sect. 2.1, and adding Gaussian noise with a standard deviation equal to the reported uncertainty of the respective TROPOMI pixel. The TROPOMI quality value filtering ensures relatively clear sky observations with good surface sensitivity. We also calculate “idealized” pressure weighted columns, which assume a uniform vertical sensitivity (flat averaging kernel), over the TROPOMI footprints without tak-

ing into account whether there is a valid TROPOMI observation as a first check to see whether the CSF can reproduce the emissions used as model input.

We first test the validity of the CSF method using the idealized columns with 10 m winds output by the WRF simulation. The WRF winds are directly responsible for transport within the simulation and can therefore be considered as the true wind fields behind the modeled concentrations. Parameters like the number of transects and distance of the background region are tuned to get optimal quantification estimates on the simulated data, such that the fitted splines capture the observed curvature of the plumes and the background is not affected by the urban emissions. A list of the different parameters and their values can be found in Appendix A. While the true wind field varies with altitude, the CSF method requires just a single (2D) wind field that is representative for the transport of the plume. We use the simulations to calibrate the CSF by introducing an effective wind speed that replaces the wind speed in Eq. (2), following the procedure by Varon et al. (2018). The effective wind speed is the wind that best captures the transport of the plume. It is a parametrization of the true wind speed to account for the effects of turbulence and variation in vertical wind speed and injection height. As the emission rates in the WRF simulations are known, the effective wind can be calculated explicitly for every orbit for each of the simulated cities. Figure 4 shows the relation between the effective wind (U_{eff}) and the WRF 10 m winds U_{10} averaged over the plume; the fitted linear relation is

$$U_{\text{eff}} = a_{10}U_{10} + b_{10}, \quad (3)$$

with $a_{10} = 1.43$ and $b_{10} = -0.92 \text{ m s}^{-1}$ ($R^2 = 0.82$). U_{10} is the wind speed at the time of overpass at 10 m altitude. We determine the effective wind relationship separately for the planetary boundary layer averaged winds, which tend to be higher than the surface winds. The resulting calibration gives $a_{\text{PBL}} = 0.98$ and $b_{\text{PBL}} = -0.20$ ($R^2 = 0.62$). While the absolute value of the PBL winds is closer to the effective wind speeds, using the U_{10} winds captures more of the variability.

After determination of the effective wind on plumes with idealized pressure profiles, we test the performance of the CSF on more realistically sampled plumes, which include the TROPOMI quality filtering and averaging kernel sensitivities as described in Sect. 2.1, to see whether the CSF can correctly quantify emission rates from synthetic observations with quality filtering and non-uniform vertical sensitivity. To test the method's sensitivity, we perform an additional effective wind calibration on these data. The resulting linear fit ($a = 1.42$, $b = -0.86$, $R^2 = 0.27$) yields similar results and shows that the filtering has limited impact on the calibration, while the lower R^2 value reflects the larger variation in estimated emission rates. At the same time, we test the lower limit to which we can trust the resulting emission estimates, as smaller enhancements are more difficult to distinguish from the background. As the modeled output concen-

tration from the WRF simulations without chemistry scales linearly with the magnitude of the input emissions, emissions from the different sectors provided by the bottom-up inventories can be scaled up and down without having to rerun the chemical transport model. This allows us to simulate plumes from cities with different emission rates with limited effect on the simulated background through scaling of the emission sector most concentrated in the considered urban area. We use this to determine the lower limit to which our method can be trusted. Figure 5 shows a comparison between the simulated input and the retrieved emission rates for the three simulated cities. The results suggest that the CSF is able to reproduce input of the WRF simulations when using 1 year of data for cities with emission rates larger than 0.1 Tg yr^{-1} .

To quantify TROPOMI plumes over all major cities in Africa, we will use the NASA/GMAO GEOS-FP wind fields (Molod et al., 2012) rather than the WRF-simulated wind fields, which are available only for the three cities selected for evaluation and calibration. For each TROPOMI pixel, we spatially interpolate the GEOS-FP wind field, which has a $0.25^\circ \times 0.3125^\circ$ spatial resolution and a 1 h time resolution. The NCEP winds that drive the WRF-simulated wind fields have a coarser time resolution of 6 h. Figure 6 shows emission estimates of simulated data using the GEOS-FP wind fields instead of the WRF fields to mimic uncertainties in the wind fields. Individual days are shown as colored dots, while the mean over the full year is shown as a colored line to represent the average estimate. The uncertainty of the average is determined as discussed in Sect. 2.5 and shown as a shaded area. The true emission is shown as a dotted black line and lies within the uncertainty of the estimate for both Cairo and Bamako. For Lagos, the emission rate is underestimated when using the WRF simulations, as the NCEP wind fields that drive the simulations are higher than both the GEOS-FP and ERA5 wind products specifically over Lagos by about 60%. The difference between the wind products might be caused by the fact that Lagos lies in the West African monsoon region, where transport has been shown to be difficult to model (Liu et al., 2014).

3 Results and discussion

After verification of the validity and calibration of the method, we apply it to 29 of the largest cities in Africa. These cities are chosen based on their population or because they are emitting above the CSF's quantification threshold in the DACCIWA inventory. Figure 7 shows the results of our TROPOMI quantification and a comparison with the DACCIWA and EDGAR inventories. These data are also included in Appendix D in Table D1. As discussed in Sect. 2.2, we used different sizes for the city masks applied to the bottom-up inventories to ensure a fair comparison to the satellite-based emission estimates and found the choice of city mask did not impact our conclusions.

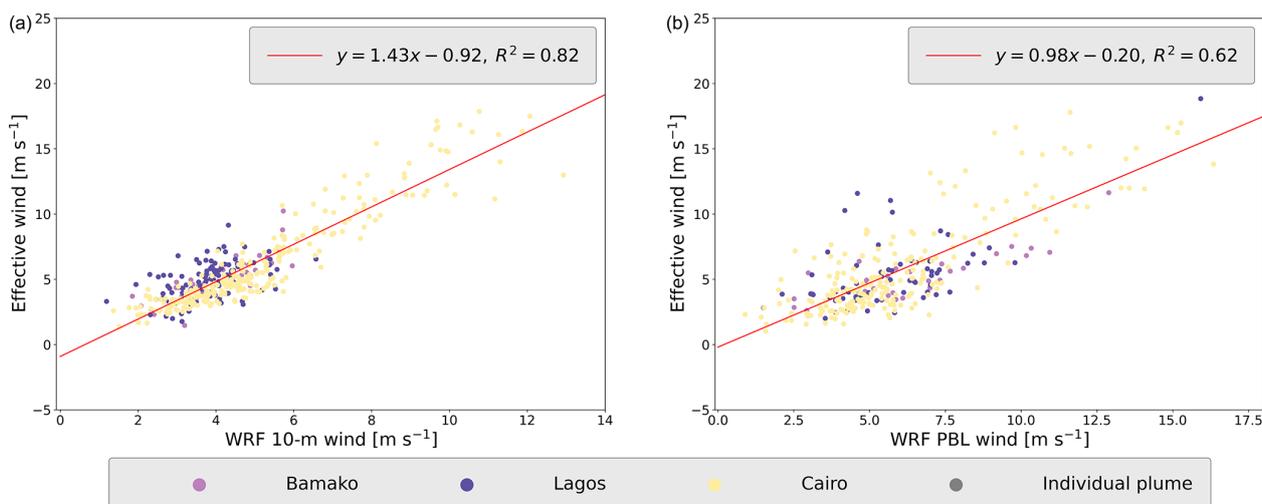


Figure 4. Determination of the relation between the effective wind and both the wind speed at 10 m altitude (a) and the planetary boundary layer averaged winds (b). The effective wind corrects for the effects of turbulence, injection height, and variation in the vertical wind profile. The simulated plumes used in the calibration cover a full year over Cairo, Bamako, and Lagos.

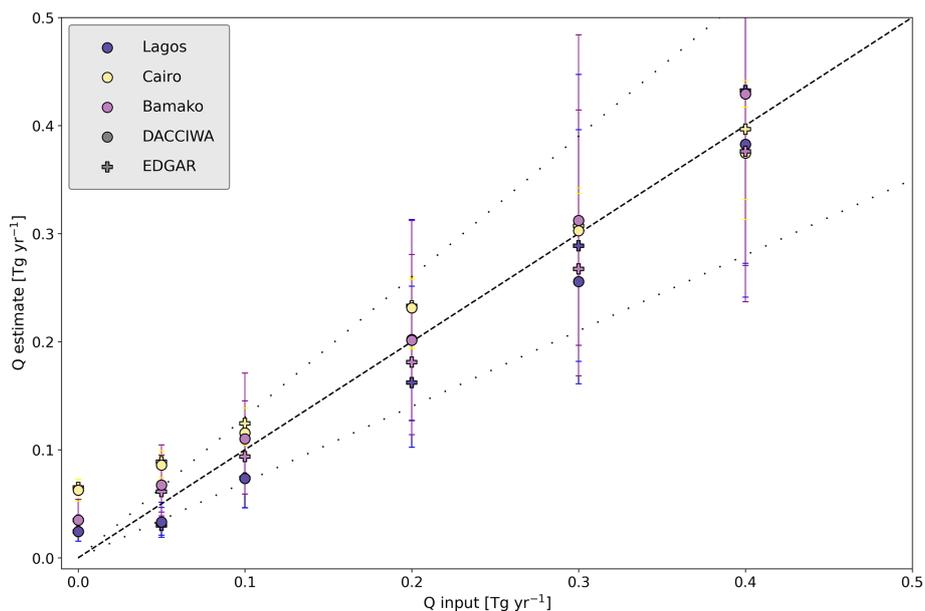


Figure 5. Emission quantification by the CSF on simulated plumes. The plumes are simulated with the WRF model for the year 2019 over Lagos (Nigeria), Cairo (Egypt), and Bamako (Mali). The simulations either used the EDGAR global bottom-up inventory or the DACCIWA inventory. The dotted lines show a 30 % deviation from the (dashed) 1 : 1 line.

On average we find TROPOMI emissions of 0.25 Tg yr^{-1} per city, compared to 0.35 Tg yr^{-1} in DACCIWA and 0.18 Tg yr^{-1} in EDGAR. Except for Abuja (Nigeria) and Khartoum (Sudan), the DACCIWA emission estimates are consistently higher than the EDGAR estimates. Additionally, the two inventories disagree on the sectoral breakdown of the emission estimates, with the domestic sector contributing 59 % to the total emission rate in DACCIWA, while EDGAR attributes 54 % of total emissions to the industry sector. For 10 cities TROPOMI and DACCIWA agree within

the TROPOMI uncertainty, that is the case for 9 cities in EDGAR. For 16 cities, the TROPOMI estimates are closer to DACCIWA than to EDGAR. The largest differences between TROPOMI and DACCIWA are found for Abidjan (627 %) and Lagos (417 %), while estimates for Cairo (2 %) and Antananarivo (9 %) agree best. To explain the differences between TROPOMI and the inventories, we will now focus on some specific areas.

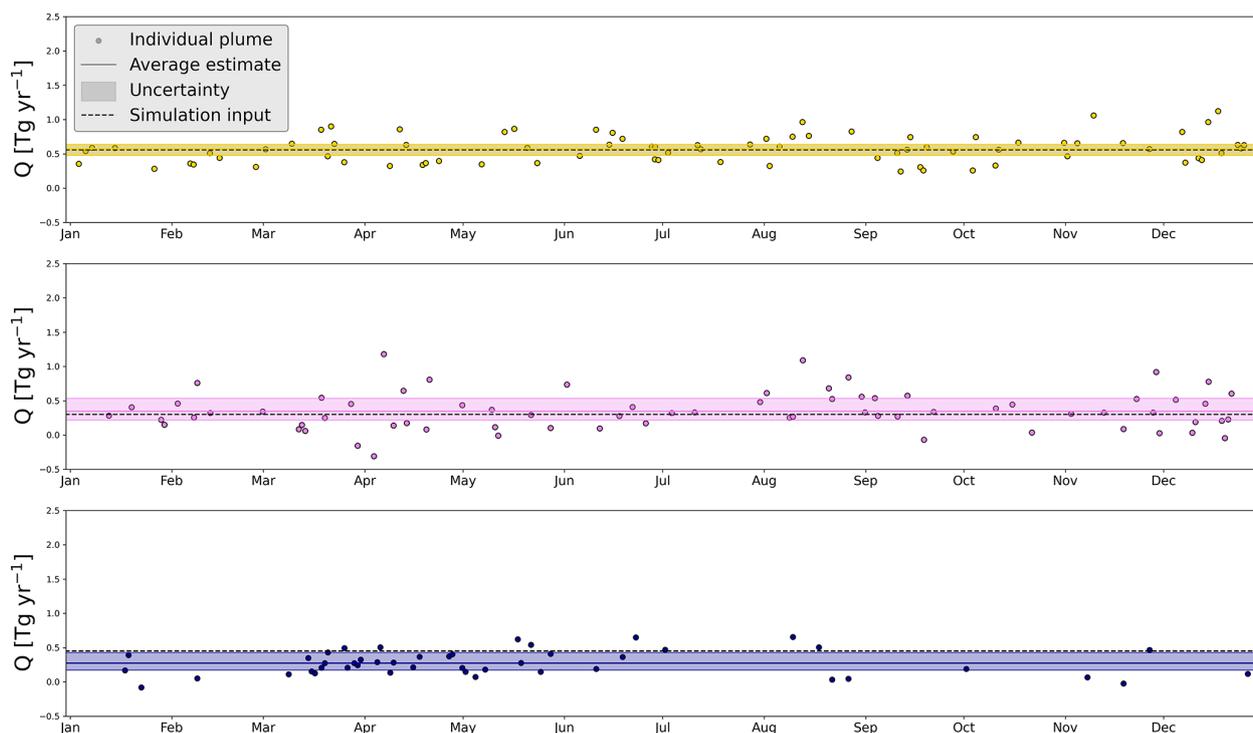


Figure 6. To check the validity of the CSF method for quantification of city emissions, we apply the method to simulated plumes sampled as TROPOMI would see them. The dots show CSF emission estimates for individual days over Cairo, Bamako, and Lagos respectively. The dark-colored line shows the annual CSF mean, with the uncertainty based on the emission ensemble shown by the shaded area. The simulation emission input, dotted black line, lies within the uncertainty of the mean CSF emission estimate for Cairo and Bamako, showing that the CSF can successfully quantify these urban emissions. For Lagos the emissions are underestimated, as the NCEP winds used to drive the simulation are much higher than both the GEOS-FP and ERA5 wind products.

3.1 Northern Africa

Two cities that stand out in Fig. 7 are Algiers and Casablanca. Unlike DACCWA, EDGAR does not include any major emission sources around these cities, even though they both have populations of 4.2 million. EDGAR also appears to largely underestimate the emissions of the two Egyptian cities that were investigated, Cairo and Alexandria, while the DACCWA emissions for these cities agree well with our TROPOMI estimates. As we can not directly obtain the underlying emission factors and activity data that are used in EDGAR and DACCWA, we compare the TROPOMI CO emission rates to the corresponding CO₂ emission rates in EDGAR, as shown in Fig. 8. The CO₂ emission rates are also included in Table D2 in Appendix D. Cairo, Alexandria, Casablanca, and Algiers clearly deviate from the other cities. Their much higher values for CO_{TROPOMI}/CO_{EDGAR} correspond to lower values for CO/CO₂ in EDGAR, indicating that not the activity data but the CO emission factors for these cities are underestimated in EDGAR. This is further confirmed by the higher CO/CO₂ values in DACCWA and the fact that the absolute CO₂ emission rates for these cities agree well between the two inventories. The underestimation in CO may point at an overestimated combustion efficiency

used in the compilation of the EDGAR emissions for this region. Similar observations over Cairo were made by MacDonald et al. (2023) when comparing measured CO and CO₂ concentrations.

3.2 South Africa

In South Africa we find closer agreement between EDGAR and TROPOMI than in northern Africa. However, the emission rates for the four considered cities in DACCWA are on average 2.4 times higher than those based on TROPOMI (Fig. 7). The emission ratios from Fig. 8 show that the South African cities stand out from the other cities in EDGAR, as they have relatively low CO/CO₂ emission ratios, suggesting high average combustion efficiencies. This does not hold for DACCWA, where the South African cities have CO/CO₂ emission ratios comparable to other cities. This indicates that the CO emission factors for South Africa are overestimated in DACCWA, and these cities have higher combustion efficiencies more in line with EDGAR.

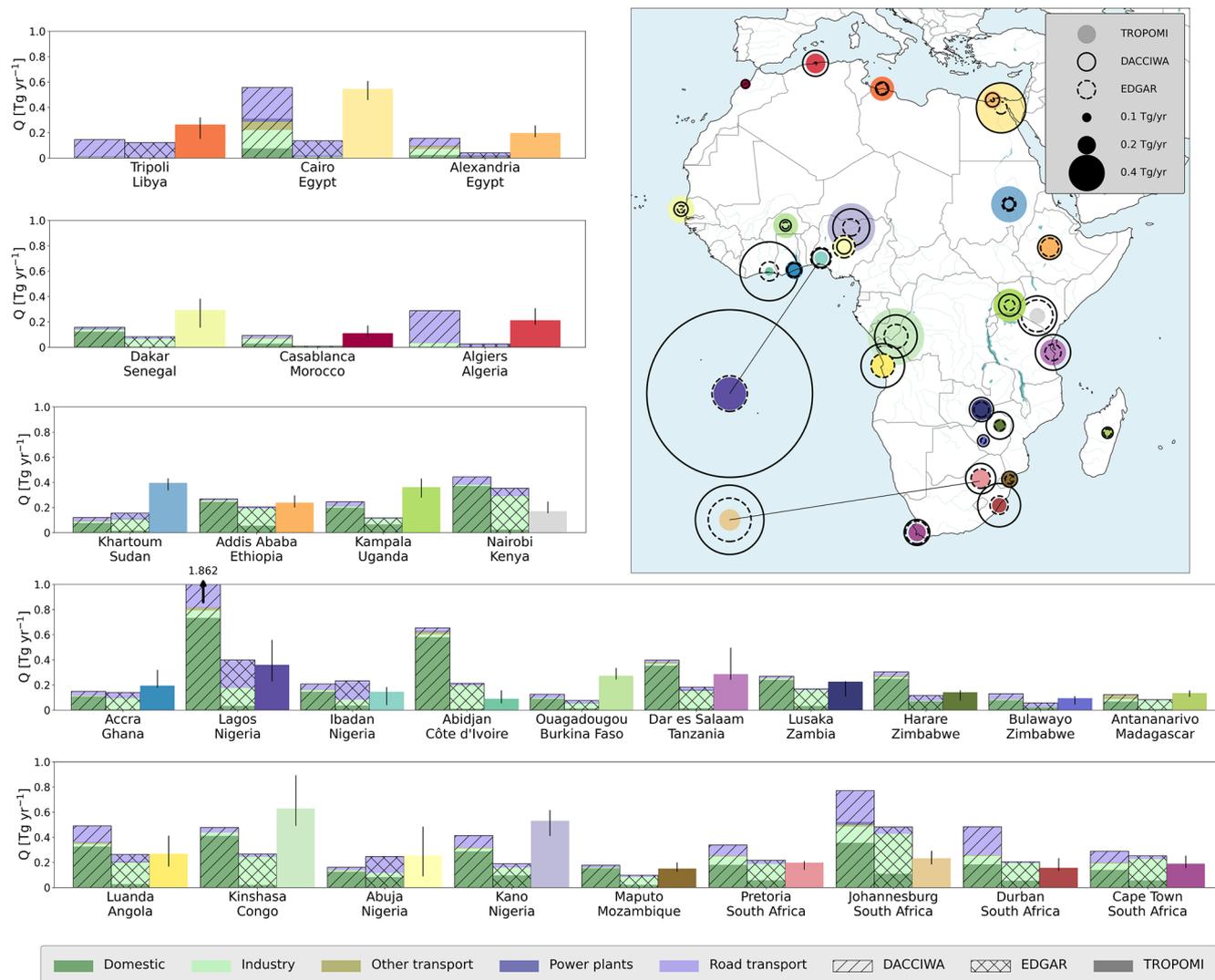


Figure 7. CSF emission quantifications for the largest African cities. Comparison between TROPOMI emission estimates averaged for 2019–2021 (shown as colored circles) and the DACCIWA and EDGAR v5 emission inventories for 2015 shown by the black (dashed) rings. The emission strength is indicated by the size of the circles or rings. The same comparison is made in bar plots, where the first two bars show the emission rates from DACCIWA and EDGAR respectively including the sectoral breakdown. The third bar gives the corresponding TROPOMI estimate, where the uncertainty is given by the range of the ensemble. The cities are ordered by geographical location. The emission estimate for Lagos in DACCIWA extends beyond the figure boundary.

3.3 Nigeria

The four investigated cities in Nigeria show varying results when comparing TROPOMI to the inventories, but the two cities that stand out are Lagos and Kano. In Lagos we estimate emissions of 0.36 (0.23 – 0.56) Tg yr^{-1} that are consistent with EDGAR, but DACCIWA has emissions that are 5.2 times higher, a difference which is much larger than the uncertainty in wind data discussed in Sect. 2.6. For Kano, in contrast to Lagos, we observe an emission rate of 0.53 (0.41 – 0.62) Tg yr^{-1} , which is consistent with DACCIWA but more than twice the EDGAR estimate (0.19 Tg yr^{-1}). The CO/CO_2 ratios of the inventories agree within 50%, but

the differences are caused by the activity data. Figure 9 shows that the CO_2 emission rates in DACCIWA for Lagos, Kano, and Ibadan are respectively 8.1, 3.8, and 3.3 times higher than in EDGAR; these data are also available in Appendix D in Table D2. Comparing Nigeria's national CO_2 budget, there is a 24% difference between the inventories (EDGAR 530 Tg yr^{-1} to DACCIWA 700 Tg yr^{-1}), but the larger regional discrepancies (over 700% for Lagos' CO_2 emissions) suggest differences in spatial allocation as well.

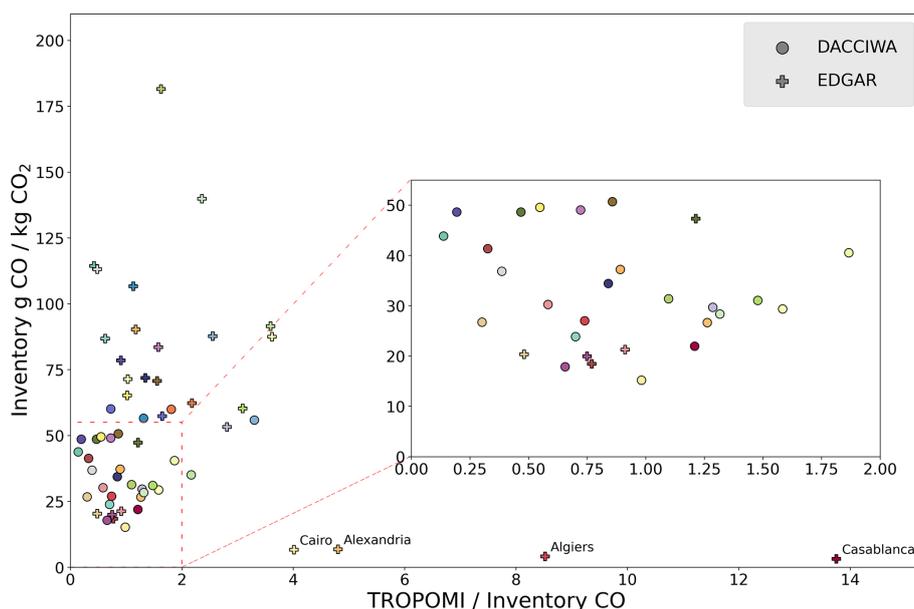


Figure 8. Comparison between inventory combustion completeness and TROPOMI to inventory CO estimates. Each marker represents a single city. The CO₂ values for both inventories include both fossil fuel and biofuel combustion emissions. As power plants hardly emit any CO per kilogram of emitted CO₂ due to their high combustion efficiency, the contributions of this sector are removed from the CO₂ values. Cairo, Alexandria, Algiers, and Casablanca have very low CO emission rates in EDGAR compared to TROPOMI and compared to EDGAR CO₂ emissions, which indicates that EDGAR largely overestimates the combustion efficiency for these cities. The four cities in EDGAR with CO/CO₂ values around 20 are all cities in South Africa, showing lower CO emission rates than the other African cities.

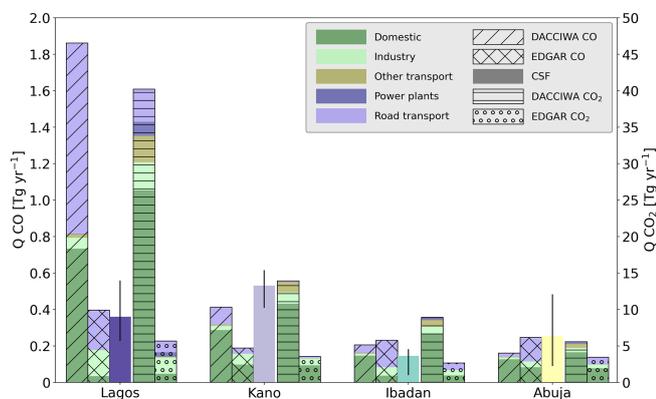


Figure 9. Comparison between the TROPOMI CO emission estimates and EDGAR and DACCIIWA CO and CO₂ for four cities in Nigeria with the same color scheme as Fig. 7. The differences between the two inventories in CO₂ emission rates indicate a different spatial allocation – based on gridded activity data – of the national totals.

3.4 Côte d'Ivoire

Abidjan in Côte d'Ivoire has the largest relative discrepancy between DACCIIWA (0.65 Tg yr⁻¹) and TROPOMI (0.1 (0.05–0.16) Tg yr⁻¹). In DACCIIWA, the domestic sector contributes to 89 % of the city's emissions, and Abidjan is the city with one of the highest CO/CO₂ values of all investigated cities. In EDGAR, the domestic CO/CO₂ ra-

tio for Abidjan is 4 times lower, which would indicate a 4 times lower emission rate. This would bring the DACCIIWA emission rate much closer to the TROPOMI observed emissions, indicating that, similar to South Africa, DACCIIWA may overestimate the city's domestic sector's CO emission factor.

3.5 Libya

Tripoli, the capital of Libya, stands out as its CO emissions in both inventories are almost exclusively 90+ % due to road transport. The TROPOMI estimate for this city of 0.26 (0.15–0.32) Tg yr⁻¹ is 1.7 and 2.2 times higher than DACCIIWA and EDGAR respectively. The difference can be partly explained by considering the domestic and industry sectors. In both emission inventories, the CO/CO₂ ratios for these sectors are 4 to 5 times lower than the mean of the other cities and 2 to 3 times lower than the next lowest city (excluding Egypt, Morocco, and Algeria). This implies that these sectors in Tripoli have high combustion efficiency compared to other cities. However, based on the TROPOMI estimate, both inventories seem to underestimate the emission factor for Tripoli, specifically for the non-road transport sectors.

3.6 Temporal emission patterns

With the 3-year TROPOMI dataset, we can also investigate the temporal variability of emissions. Earlier studies,

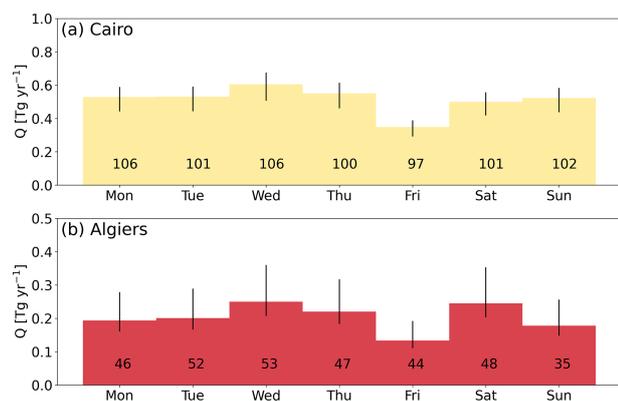


Figure 10. TROPOMI emission estimates over Cairo (a) and Algiers (b) for different days of the week averaged over 2019–2021. The numbers in the bars show the number of days the averages are based on. Both cities show lower emissions on Friday, consistent with road transport being the main contributor to urban emissions and Friday being the standard day off in Islamic countries.

focusing on concentration trends rather than emission estimates, have found that CO concentrations over Cairo are lower on Fridays, which is the day off in the Islamic world (Rey-Pommier et al., 2022). This “weekend effect” has also been observed for nitrogen dioxide (NO₂) and ozone (O₃), which like CO in Cairo are dominated by transport emissions (Beirle et al., 2003; Khoder, 2009; Stavrou et al., 2020). Combined with the fact that both emission inventories agree on road transport as the main contributor to emissions in Cairo, lower CO emissions are indeed expected on Fridays when there is less commuter traffic. Figure 10 indeed shows a 32 % drop in emissions on Fridays over Cairo. A similar reduction in emissions can be seen over Algiers, which can also be attributed to reduced road traffic. Similar significant patterns were not seen for the other cities that tend to have relatively lower contributions from road traffic.

4 Conclusions

We adapted and calibrated the computationally efficient cross-sectional flux (CSF) method to quantify urban carbon monoxide emission rates from major cities in Africa using TROPOMI data. We determined optimal values for the parameters of the CSF by applying the method to a full year of simulated WRF plumes over three distinctly different African cities (Cairo, Lagos, and Bamako), such that the transects drawn best match the shape and curvature of the simulated plumes. These simulations were also used to calibrate the CSF’s effective wind speed relationship for TROPOMI data. By applying the calibrated CSF to the simulated data with known emission rates, we found that we can quantify urban CO emissions down to 0.1 Tg yr⁻¹ within 30 % uncertainty. After calibration, we applied our CSF method to TROPOMI observations of 29 of Africa’s most populated

and/or emitting cities. We focus on Africa as there are relatively few dedicated emission inventories for the continent, and large uncertainties in emission rates are expected.

We compared our TROPOMI-based emission estimates with the global EDGAR emission inventory and the Africa-focused DACCIWA inventory. There are substantial differences between urban CO emissions from both inventories. We did not find vastly better average agreement of either inventory with TROPOMI. DACCIWA is closer to the TROPOMI estimate for 16 out of 29 cities. For 10 cities, the DACCIWA and TROPOMI estimates agree within the uncertainty of the TROPOMI-based estimate, but there are also cities with large significant differences of over 600 %. Compared to EDGAR, we find that nine cities agree within the uncertainty and we similarly find cities with large discrepancies.

We then evaluated our results for different regions. In northern Africa, TROPOMI observes higher emission rates than shown in EDGAR for cities in Egypt, Algeria, and Morocco. The EDGAR CO to CO₂ emission ratios for these four cities are relatively low, implying that the mismatch with TROPOMI may originate from the emission factors, which implies that EDGAR would overestimate the average combustion efficiency in these cities. In South Africa, the TROPOMI estimates agree with EDGAR, but the DACCIWA estimates are high in comparison. EDGAR shows lower CO to CO₂ ratios and hence higher combustion efficiency for South Africa compared to other parts of Africa, implying that those ratios may be too high in DACCIWA. Similarly, DACCIWA appears to underestimate the combustion efficiency in Abidjan (Côte d’Ivoire). For Tripoli (Libya), both inventories estimate lower emission rates than the estimate based on TROPOMI. Specifically the domestic and industry sectors show particularly high combustion efficiencies compared to other cities in both inventories, which can explain part of the discrepancy.

We also found some discrepancies that can be attributed to the activity data used by the inventories. We found a factor ~ 4 lower emissions based on TROPOMI for Lagos (Nigeria) than estimated by DACCIWA. The associated DACCIWA CO₂ emissions are 8 times larger than in EDGAR, which can partly be a difference in activity data but also suggests a mismatch related to the spatial distribution of emissions. For Kano (Nigeria), DACCIWA estimates CO₂ emissions that are 4 times larger than EDGAR. Here, the TROPOMI estimate agrees better with DACCIWA, and the activity rate, which corresponds to CO₂ emission, in EDGAR seems to be an underestimation.

The large TROPOMI data volume enables the identification of temporal emission patterns. Over Cairo and Algiers we find significantly lower emission rates on Fridays – the local rest day – compared to other days of the week. The ability to recognize such patterns builds confidence and shows the strength of TROPOMI’s daily global coverage combined

with a computationally efficient method like the CSF method developed here.

Appendix A: TROPOMI-based CO uncertainty

The CSF method as applied in this work has many parameters that were calibrated on simulated plumes. In order to determine the uncertainty of the estimated emission rate, we have created an ensemble of emission estimates by varying these parameters. The members of the ensemble and the ranges over which they were varied are shown in Table A1. The wind databases are the three wind products as described in Sect. 2.5 and are responsible for a mean uncertainty of -19% and $+12\%$. The standard deviation threshold for spline pixels is the number of standard deviations a pixel has to be above the background concentration in order to be considered part of the plume. Pixels identified as part of the plume are only used for fitting the spline shape. The number of cross sections is the total number of transects drawn perpendicular to the direction of the plume. As described in Sect. 2.4, not all transects are taken into account in the quantification. The number of cross sections is mostly a measure for the line density or the distance between consecutive transects, as the transects are evenly spaced over the full length of the spline. The minimum pixel coverage is the minimum fraction of a line that needs to be covered by pixels and is a balance between retaining enough days with a valid estimate and not underestimating emissions. A lower limit of 50% coverage per line will retain a lot of lines and, thus, more days with an estimate; however, the cross sections will potentially miss large parts of the plume. The distance of the background box is important, as one would like the box to be close to the city to get a good representation of the local background. However, it must not overlap with any urban emissions to have a clean background. As explained in Sect. 2.4, the transects start upwind of the city center to capture the full city plume. We vary the distance between the first transect and the city center for our uncertainty estimate. As a last member of the ensemble, we use the spread in emission estimates of the individual transects. We include the means of the transects with the lowest and highest 50% emission rates in the ensemble.

Table A1. Variables used in the uncertainty analysis and the ranges over which they were varied. The resulting ensemble spreads are reported as uncertainty.

Parameter	Domain	Default
Wind database	GEOS-FP 10m, ERA5 10m, and GEOS-FP PBL	GEOS-FP 10m
Standard deviation threshold for spline pixels	{1.2–2.4}	1.8
Number of transects	{15–25}	20
Minimum pixel coverage per transect	{50%–90%}	70%
Distance of background box	{0.2–0.4°}	0.3°
Varying upwind distance first transect	{0–0.2°}	0.1°
Transects used for estimate	{Lowest 50%–highest 50%}	All

Appendix B: TROPOMI data filtering

Although the CSF has been shown to reproduce simulated emission rates (Fig. 5), it can not be applied to every single overpass of TROPOMI. For example, days with a lot of missing pixels in the TROPOMI data can lead to underestimation of the emission rates. To prevent a positive bias, it is important to not only use days with strong, clearly visible plumes. With the filters chosen in this work, over 400 d are accepted as quantifiable over the 3-year period studied for most non-coastal cities and coastal cities with predominantly inland winds (e.g., Cairo has 713 estimates, Johannesburg 427, and Khartoum 570). With TROPOMI's daily overpasses this means that we estimate emissions on roughly 40 % of all days. Regions with fewer estimates tend to be coastal. For example, we only have estimates for 160 d over Lagos and 113 for Dakar because of limited TROPOMI coverage over water. An additional reason for a small number of valid estimates lies in the occurrence of open fires; for example, 224 orbits (42 %) are removed from our estimate over Lusaka (Zambia) due to fires within 1.5° of the city center and stronger fires within 0.75° .

The filters employed are shown in Table B1.

Table B1. Filtering applied to the data to ensure correct application of the CSF method.

Description		Value
Per pixel	Quality flag TROPOMI (land).	≥ 0.7
	Quality flag TROPOMI (water).	$= 0.7$
Per transect	Misalignment between plume and wind direction.	$< 45^\circ$
	Minimum pixel coverage.	$> 70\%$
Per plume	Downwind coverage in a $0.3^\circ \times 0.8^\circ$ box.	$> 60\%$
	Effective wind speed.	$> 2 \text{ m s}^{-1}$
	Maximum concentration outside the plume.	$< 200 \text{ ppb}$ (1.5° radius)
	Number of transects used for the estimate.	> 3
	Fraction estimate transect 8–20 to transect 3–7. High emission estimates of the far away lines tend to indicate interference of different sources.	$< 2.5 \times$
	Second derivative of spline scaled to 1° pixel size. This represents the dimensionless curvature of the fitted spline.	< 0.05
	Mismatch between the first transect and the starting pixel of the plume.	$< 0.35^\circ$
	Fire emission from Global Fire Assimilation System (GFAS) database, Kaiser et al. (2012).	$< 23 \text{ Mg h}^{-1}$ (0.75° radius), $< 57 \text{ Mg h}^{-1}$ (1.5° radius)

Appendix C: CSF estimates

Table C1. Emission estimates for the studied African cities by applying the CSF method to the TROPOMI CO product (2019–2021), as well as the corresponding emission rates according to the DACCIWA (2015) and EDGAR (2015) inventories. All emission rates are in teragrams per year (Tg yr^{-1}). Population is taken from the Center for International Earth Science Information Network (CIESIN, 2018).

City	Country	Population	TROPOMI estimate	Lower limit	Upper limit	DACCIWA	EDGAR
Algiers	Algeria	4.2 M	0.213	0.177	0.307	0.288	0.025
Luanda	Angola	5.1 M	0.268	0.166	0.413	0.489	0.263
Ouagadougou	Burkina Faso	2.8 M	0.273	0.243	0.335	0.126	0.076
Kinshasa	Congo	7.4 M	0.628	0.49	0.894	0.477	0.266
Abidjan	Côte d'Ivoire	7.1 M	0.09	0.054	0.157	0.654	0.266
Alexandria	Egypt	3.3 M	0.197	0.164	0.257	0.156	0.041
Cairo	Egypt	16.7 M	0.546	0.456	0.608	0.556	0.136
Addis Ababa	Ethiopia	4.4 M	0.239	0.2	0.295	0.268	0.204
Accra	Ghana	3.5 M	0.194	0.176	0.318	0.148	0.172
Nairobi	Kenya	4.8 M	0.17	0.153	0.247	0.441	0.353
Tripoli	Libya	1.2 M	0.264	0.151	0.32	0.146	0.121
Antananarivo	Madagascar	3.0 M	0.135	0.104	0.156	0.123	0.083
Casablanca	Morocco	4.2 M	0.11	0.078	0.171	0.091	0.008
Maputo	Mozambique	2.5 M	0.151	0.125	0.199	0.176	0.097
Abuja	Nigeria	2.6 M	0.255	0.089	0.483	0.161	0.247
Ibadan	Nigeria	2.2 M	0.145	0.039	0.183	0.207	0.232
Kano	Nigeria	5.3 M	0.531	0.409	0.616	0.413	0.189
Lagos	Nigeria	10.9 M	0.36	0.227	0.558	1.862	0.397
Dakar	Senegal	4.2 M	0.293	0.154	0.381	0.157	0.081
Cape Town	South Africa	4.1 M	0.189	0.155	0.252	0.288	0.252
Durban	South Africa	3.3 M	0.157	0.132	0.233	0.482	0.204
Johannesburg	South Africa	9.1 M	0.232	0.184	0.291	0.77	0.482
Pretoria	South Africa	6.4 M	0.197	0.139	0.209	0.338	0.216
Khartoum	Sudan	3.1 M	0.396	0.336	0.432	0.12	0.155
Dar es Salaam	Tanzania	5.4 M	0.286	0.241	0.496	0.396	0.181
Kampala	Uganda	4.3 M	0.362	0.279	0.431	0.245	0.117
Lusaka	Zambia	2.4 M	0.226	0.108	0.23	0.269	0.168
Bulawayo	Zimbabwe	0.8 M	0.094	0.043	0.111	0.13	0.057
Harare	Zimbabwe	2.7 M	0.142	0.076	0.157	0.304	0.117

Appendix D: Inventory emission rates

Table D1. Sectoral breakdown of the CO emission rates for the studied African cities according to the DACCIWA and EDGAR inventory. All emission rates are in gigagrams per year (Gg yr^{-1}).

City	DACCIWA					EDGAR				
	Domestic	Industry	Other	Power	Road	Domestic	Industry	Other	Power	Road
Algiers	0.3	33.7	0.6	0.9	252.2	1.3	2.7	0.2	0.4	20.0
Luanda	326.1	19.4	12.0	3.4	129.0	25.4	174.0	0.3	6.5	56.1
Ouagadougou	90.3	8.0	0.0	0.8	26.4	13.2	39.8	0.1	2.8	20.0
Kinshasa	411.1	25.4	0.0	0.0	40.4	20.1	222.3	0.1	0.1	23.5
Abidjan	579.9	21.4	20.4	7.0	25.5	10.8	183.6	0.3	1.6	15.7
Alexandria	24.3	46.7	20.7	4.4	59.6	3.4	1.8	0.3	15.8	19.9
Cairo	75.8	145.5	64.6	26.1	244.3	3.5	6.4	1.1	9.3	115.9
Addis Ababa	245.3	8.5	1.8	0.0	12.3	55.3	134.8	0.2	0.0	13.9
Accra	108.2	3.2	3.9	0.0	32.4	5.0	90.2	1.0	0.5	43.2
Nairobi	370.8	11.7	0.2	2.3	59.1	23.5	266.9	0.3	2.1	60.0
Tripoli	3.1	4.2	0.0	6.6	131.6	1.9	0.1	0.1	0.0	118.9
Antananarivo	68.8	17.0	30.2	0.1	6.5	7.5	68.3	0.1	0.0	7.3
Casablanca	29.2	37.1	4.6	0.5	20.0	1.5	3.9	0.0	0.2	2.2
Maputo	153.1	5.3	0.1	3.1	14.7	21.0	62.7	0.0	0.3	13.1
Abuja	126.2	10.2	3.6	0.1	21.0	82.6	32.1	0.0	0.0	131.9
Ibadan	145.8	11.7	4.1	0.3	44.7	36.2	46.4	0.0	0.0	149.1
Kano	286.3	23.0	8.1	0.0	95.5	96.9	60.6	0.0	0.0	31.0
Lagos	733.5	59.0	20.8	2.4	1046.3	33.2	143.5	0.1	0.7	219.8
Dakar	123.3	17.4	0.1	6.1	10.0	2.7	62.7	0.2	3.2	12.6
Cape Town	139.9	50.7	5.7	1.4	90.5	56.9	170.1	0.7	0.0	23.8
Durban	183.8	66.6	7.5	0.0	224.4	53.3	139.4	0.3	0.0	10.9
Johannesburg	355.8	128.9	14.5	22.4	247.9	108.8	317.5	0.1	0.0	55.1
Pretoria	180.4	65.4	7.4	0.9	83.4	57.3	129.0	0.1	0.4	29.3
Khartoum	77.2	10.8	3.8	1.6	26.2	16.3	90.0	0.1	5.6	42.6
Dar es Salaam	354.2	15.0	9.6	1.7	15.8	15.1	138.1	0.1	10.4	17.7
Kampala	198.8	10.5	0.1	1.2	34.2	65.6	42.5	0.0	0.0	8.8
Lusaka	239.0	11.7	1.2	0.0	16.7	31.7	126.3	0.0	0.0	9.8
Bulawayo	79.1	4.3	3.5	0.2	42.4	21.5	2.0	0.0	0.2	33.0
Harare	248.2	13.3	10.8	0.2	31.5	69.8	4.5	0.1	0.4	41.7

Table D2. Sectoral breakdown of the CO₂ emission rates for the studied African cities according to the DACCIWA and EDGAR inventory. All emission rates are in teragrams per year (Tg yr⁻¹).

City	DACCIWA					EDGAR				
	Domestic	Industry	Other	Power	Road	Domestic	Industry	Other	Power	Road
Algiers	2.0526	2.3547	0.3026	0.7006	5.9577	2.2141	2.3123	0.1231	0.0446	1.4081
Luanda	4.8853	1.4104	2.2871	0.0	1.2816	1.8331	1.4681	0.1276	1.6398	0.6036
Ouagadougou	2.1571	0.5752	0.701	0.4603	0.1553	0.3469	0.2106	0.1102	0.4868	0.1633
Kinshasa	9.1058	2.4811	4.9228	0.0	0.3205	0.4757	1.1183	0.0897	0.0245	0.2184
Abidjan	5.6168	1.2141	7.6905	4.1632	0.3916	0.3994	1.1479	0.1115	1.1956	0.1948
Alexandria	1.7076	3.1371	0.3293	0.0	0.6755	0.8145	4.1211	0.0639	23.3842	0.9927
Cairo	9.7364	17.8872	3.5146	17.7906	5.4195	3.1328	11.4975	0.7648	13.7938	4.9941
Addis Ababa	5.2452	0.4889	1.0817	0.0	0.3851	1.2733	0.5438	0.2141	0.0005	0.2297
Accra	1.172	0.3122	0.7803	0.6024	0.3533	0.231	0.6034	0.4059	0.3562	0.3729
Nairobi	5.3718	0.9503	4.7633	0.429	0.8817	0.6676	1.6001	0.2878	0.4414	0.5657
Tripoli	0.3477	0.8711	0.3068	7.4683	0.9077	0.252	0.6718	0.022	0.0	0.9969
Antananarivo	1.7812	0.9208	1.0695	0.0798	0.1474	0.205	0.144	0.046	0.0195	0.0622
Casablanca	1.4466	2.1949	0.0158	4.0377	0.4885	0.7635	1.2911	0.0514	1.2078	0.3914
Maputo	2.3379	0.2568	0.6895	0.0	0.1877	0.4888	0.7266	0.0355	0.0795	0.1202
Abuja	4.1275	0.6013	0.5368	0.1024	0.2142	1.9347	0.4962	0.0413	0.0	0.9858
Ibadan	6.678	0.9728	0.8684	0.2404	0.1572	0.944	0.7214	0.0006	0.0006	1.0063
Kano	10.773	1.5693	1.401	0.0311	0.1601	2.3444	0.9346	0.0206	0.0008	0.2502
Lagos	26.3529	3.8388	3.5988	1.9078	4.5136	1.1902	2.2007	0.0889	0.6383	1.5765
Dakar	1.5405	1.01	0.9282	1.9502	0.3954	0.1379	0.4475	0.1115	1.2811	0.2292
Cape Town	3.1332	10.8663	1.2788	1.0896	0.823	1.759	9.0686	0.247	0.0009	1.5695
Durban	2.3207	8.0484	0.543	0.1116	0.7436	1.5599	8.5055	0.0943	0.1335	0.9045
Johannesburg	5.5358	19.1984	2.2775	23.3095	1.7699	3.1791	17.0838	0.1623	4.9948	3.2495
Pretoria	2.2686	7.8676	0.5308	3.0576	0.5046	1.6006	6.6839	0.1254	3.4728	1.7434
Khartoum	0.6315	0.9569	0.3614	0.1673	0.198	0.5049	0.6555	0.101	1.4848	0.5062
Dar es Salaam	5.582	1.0448	1.4222	1.073	0.0259	0.3639	1.5205	0.0908	1.3461	0.1923
Kampala	5.0997	0.8774	1.6523	0.5054	0.2586	1.4997	0.329	0.009	0.0	0.1042
Lusaka	3.8478	1.6338	2.2394	0.0	0.0951	0.7361	1.4979	0.0327	0.0	0.0699
Bulawayo	1.8723	0.1759	0.0062	0.2274	0.1068	0.5132	0.2041	0.0134	0.2151	0.263
Harare	5.522	0.5187	0.0183	0.2416	0.193	1.6508	0.3461	0.059	0.4645	0.4179

Code and data availability. TROPOMI CO data (<https://doi.org/10.5270/S5P-bj3nry0>, Copernicus Sentinel-5P, 2021) are publicly available from the ESA Sentinel-5P data hub at <https://s5phub.copernicus.eu> (last access: 1 June 2023). ERA5 wind data are available at <https://www.ecmwf.int/en/forecasts/dataset/ecmwf-reanalysis-v5> (Hersbach et al., 2022). WRF-Chem code is available at <https://github.com/wrf-model/WRF/releases> (Powers et al., 2023); in this work, version 4.1.5 was used. EDGAR v5 CO data are available at https://edgar.jrc.ec.europa.eu/dataset_ap50 (Crippa et al., 2021). EDGAR v5 CO₂ data are available at https://edgar.jrc.ec.europa.eu/dataset_ghg50 (Crippa et al., 2022). DACCIWA CO and CO₂ data are available at <https://doi.org/10.25326/56> (Keita et al., 2020) at the Emissions of atmospheric Compounds and Compilation of Ancillary Data (ECCAD) system. GPW v4 gridded population density is available at <https://doi.org/10.7927/H49C6VHW> (CIESIN, 2018). Open fire emissions from GFAS are available at <https://atmosphere.copernicus.eu/global-fire-emissions> (Kaiser et al., 2022).

Author contributions. GL, JDM, and IA designed the study. GL performed the TROPOMI analysis with contributions from JDM, HACDvdG, AS, and IA. GL and JDM wrote the paper with contributions from all authors. JDM, HACDvdG, and AS contributed to the comparison between TROPOMI and emission inventories. TB provided the TROPOMI carbon monoxide data and associated support.

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References

- Badarinath, K., Kharol, S. K., Chand, T. K., Parvathi, Y. G., Anasuya, T., and Jyothisna, A. N.: Variations in black carbon aerosol, carbon monoxide and ozone over an urban area of Hyderabad, India, during the forest fire season, *Atmos. Res.*, 85, 18–26, 2007.
- Beirle, S., Platt, U., Wenig, M., and Wagner, T.: Weekly cycle of NO₂ by GOME measurements: a signature of anthropogenic sources, *Atmos. Chem. Phys.*, 3, 2225–2232, <https://doi.org/10.5194/acp-3-2225-2003>, 2003.
- Bi, J., Zuidema, C., Clausen, D., Kirwa, K., Young, M. T., Gassett, A. J., Seto, E. Y., Sampson, P. D., Larson, T. V., Szpiro, A. A., Sheppard, L., and Kaufman, J. D.: Within-City Variation in Ambient Carbon Monoxide Concentrations: Leveraging Low-Cost Monitors in a Spatiotemporal Modeling Framework, *Environ. Health Persp.*, 130, 97008, <https://doi.org/10.1289/EHP10889>, 2022.
- Bieser, J., Aulinger, A., Matthias, V., Quante, M., and Van Der Gon, H. D.: Vertical emission profiles for Europe based on plume rise calculations, *Environ. Pollut.*, 159, 2935–2946, 2011.
- Borsdorff, T., Aan de Brugh, J., Hu, H., Aben, I., Hasekamp, O., and Landgraf, J.: Measuring carbon monoxide with TROPOMI: First results and a comparison with ECMWF-IFS analysis data, *Geophys. Res. Lett.*, 45, 2826–2832, 2018.
- Borsdorff, T., García Reynoso, A., Maldonado, G., Mar-Morales, B., Stremme, W., Grutter, M., and Landgraf, J.: Monitoring CO emissions of the metropolis Mexico City using TROPOMI CO observations, *Atmos. Chem. Phys.*, 20, 15761–15774, <https://doi.org/10.5194/acp-20-15761-2020>, 2020.
- Cai, B., Cui, C., Zhang, D., Cao, L., Wu, P., Pang, L., Zhang, J., and Dai, C.: China city-level greenhouse gas emissions inventory in 2015 and uncertainty analysis, *Appl. Energ.*, 253, 113579, <https://doi.org/10.1016/j.apenergy.2019.113579>, 2019.
- CIESIN (Center for International Earth Science Information Network): Gridded Population of the World, Version 4 (GPWv4): Population Density, Revision 11, NASA Socioeconomic Data and Applications Center (SEDAC) [data set], Palisades, New York, <https://doi.org/10.7927/H49C6VHW>, 2018.
- Copernicus Sentinel-5P (processed by ESA): TROPOMI Level 2 Carbon Monoxide total column products, Version 02, European Space Agency [data set], <https://doi.org/10.5270/S5P-bj3nry0>, 2021.
- Crippa, M., Oreggioni, G., Guizzardi, D., Muntean, M., Schaaf, E., Lo Vullo, E., Solazzo, E., Monforti-Ferrario, F., Olivier, J. G. J., and Vignati, E.: Emissions Database for Global Atmospheric Research Air Pollutants, Version 5, European Commission, Joint Research Centre [data set], https://edgar.jrc.ec.europa.eu/dataset_ap50, last access: 24 March 2021.
- Crippa, M., Oreggioni, G., Guizzardi, D., Muntean, M., Schaaf, E., Lo Vullo, E., Solazzo, E., Monforti-Ferrario, F., Olivier, J. G. J., and Vignati, E.: Emissions Database for Global Atmospheric Research Greenhouse Gases, Version 5, European Commission, Joint Research Centre [data set], https://edgar.jrc.ec.europa.eu/dataset_ghg50, last access: 2 November 2022.
- Daniel, J. S. and Solomon, S.: On the climate forcing of carbon monoxide, *J. Geophys. Res.-Atmos.*, 103, 13249–13260, 1998.
- Dekker, I. N., Houweling, S., Aben, I., Röckmann, T., Krol, M., Martínez-Alonso, S., Deeter, M. N., and Worden, H. M.: Quantification of CO emissions from the city of Madrid using MO-PITT satellite retrievals and WRF simulations, *Atmos. Chem. Phys.*, 17, 14675–14694, <https://doi.org/10.5194/acp-17-14675-2017>, 2017.
- Dekker, I. N., Houweling, S., Pandey, S., Krol, M., Röckmann, T., Borsdorff, T., Landgraf, J., and Aben, I.: What caused the extreme CO concentrations during the 2017 high-pollution episode in India?, *Atmos. Chem. Phys.*, 19, 3433–3445, <https://doi.org/10.5194/acp-19-3433-2019>, 2019.
- Diab, R., Foster, S., François, K., Martincigh, B., and Salter, L.: Carbon monoxide levels at a toll plaza near Durban, South Africa, *Environ. Chem. Lett.*, 3, 91–94, 2005.
- Doumbia, M., Kouassi, A. A., Silué, S., Yoboué, V., Liousse, C., Diedhiou, A., Touré, N. E., Keita, S., Assamoi, E.-M., Bamba, A., Zouzoua, M., Dajuma, A., and Kouadio, K.: Road traffic emission inventory in an urban zone of West Africa: case of Yopougon City (Abidjan, Côte d’Ivoire), *Energies*, 14, 1111, <https://doi.org/10.3390/en14041111>, 2021.
- Eggleston, H., Buendia, L., Miwa, K., Ngara, T., and Tanabe, K. (Eds.): 2006 IPCC guidelines for national greenhouse gas inventories, Institute for Global Environmental Strategies (IGES), ISBN 4887880324, 2006.
- Guevara, M., Jorba, O., Tena, C., Denier van der Gon, H., Kuenen, J., Elguindi, N., Darras, S., Granier, C., and Pérez García-Pando, C.: Copernicus Atmosphere Monitoring Service TEMPO profiles (CAM5-TEMPO): global and European emission temporal profile maps for atmospheric chemistry modelling, *Earth Syst. Sci. Data*, 13, 367–404, <https://doi.org/10.5194/essd-13-367-2021>, 2021.
- Guo, H., So, K., Simpson, I., Barletta, B., Meinardi, S., and Blake, D.: C₁–C₈ volatile organic compounds in the atmosphere of Hong Kong: overview of atmospheric processing and source apportionment, *Atmos. Environ.*, 41, 1456–1472, 2007.
- Havens, D., Wang, D., Grigg, J., Gordon, S. B., Balmes, J., and Mortimer, K.: The cooking and pneumonia study (CAPS) in Malawi: a cross-sectional assessment of carbon monoxide exposure and carboxyhemoglobin levels in children under 5 years old, *Int. J. Env. Res. Pub. He.*, 15, 1936, <https://doi.org/10.3390/ijerph15091936>, 2018.
- Hedelius, J. K., Toon, G. C., Buchholz, R. R., Iraci, L. T., Podolske, J. R., Roehl, C. M., Wennberg, P. O., Worden, H. M., and Wunch,

- D.: Regional and urban column CO trends and anomalies as observed by MOPITT over 16 years, *J. Geophys. Res.-Atmos.*, 126, e2020JD033967, <https://doi.org/10.1029/2020JD033967>, 2021.
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren, P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C., Radnoti, G., de Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J.-N.: The ERA5 global reanalysis, *Q. J. Roy. Meteor. Soc.*, 146, 1999–2049, 2020.
- Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J., Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G., Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren, P., Dee, D., Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L., Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C., Radnoti, G., de Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J.-N.: ECMWF atmospheric reanalysis version 5, Copernicus Climate Change Service [data set], <https://www.ecmwf.int/en/forecasts/dataset/ecmwf-reanalysis-v5>, last access: 6 January 2022.
- Hornweg, D. and Pope, K.: Population predictions for the world's largest cities in the 21st century, *Environment and Urbanization*, 29, 195–216, 2017.
- Inness, A., Blechschmidt, A.-M., Bouarar, I., Chabrilat, S., Crepulja, M., Engelen, R. J., Eskes, H., Flemming, J., Gaudel, A., Hendrick, F., Huijnen, V., Jones, L., Kapsomenakis, J., Katragkou, E., Keppens, A., Langerock, B., de Mazière, M., Melas, D., Parrington, M., Peuch, V. H., Razinger, M., Richter, A., Schultz, M. G., Suttie, M., Thouret, V., Vrekoussis, M., Wagner, A., and Zerefos, C.: Data assimilation of satellite-retrieved ozone, carbon monoxide and nitrogen dioxide with ECMWF's Composition-IFS, *Atmos. Chem. Phys.*, 15, 5275–5303, <https://doi.org/10.5194/acp-15-5275-2015>, 2015.
- IPCC: Climate Change and Land: an IPCC special report on climate change, desertification, land degradation, sustainable land management, food security, and greenhouse gas fluxes in terrestrial ecosystems, edited by: Masson-Delmotte, V., Pörtner, H.-O., Skea, J., Buendía, E. C., Zhai, P., Roberts, D., and Shukla, P. R., Intergovernmental Panel on Climate Change (IPCC), 77–131, <https://spiral.imperial.ac.uk/bitstream/10044/1/76618/2/SRCCL-Full-Report-Compiled-191128.pdf> (last access: 1 October 2022), 2019.
- Jacob, D. J.: Introduction to atmospheric chemistry, Princeton University Press, ISBN 9780691001852, 1999.
- Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Muntean, M., Schaaf, E., Dentener, F., Bergamaschi, P., Pagliari, V., Olivier, J. G. J., Peters, J. A. H. W., van Aardenne, J. A., Monni, S., Doering, U., Petrescu, A. M. R., Solazzo, E., and Oreggioni, G. D.: EDGAR v4.3.2 Global Atlas of the three major greenhouse gas emissions for the period 1970–2012, *Earth Syst. Sci. Data*, 11, 959–1002, <https://doi.org/10.5194/essd-11-959-2019>, 2019.
- Kaiser, J. W., Heil, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J., Razinger, M., Schultz, M. G., Suttie, M., and van der Werf, G. R.: Biomass burning emissions estimated with a global fire assimilation system based on observed fire radiative power, *Biogeosciences*, 9, 527–554, <https://doi.org/10.5194/bg-9-527-2012>, 2012.
- Kaiser, J. W., Hell, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J., Razinger, M., Schultz, M. G., Suttie, M., and van der Werf, G. R.: Global Fire Assimilation System fire emission Fire radiative power, Copernicus Atmosphere Monitoring Service [data set], <https://atmosphere.copernicus.eu/global-fire-emissions>, last access: 11 June 2022.
- Kansiime, W. K., Mugambe, R. K., Atusingwize, E., Wafula, S. T., Nsereko, V., Ssekamatte, T., Nalugya, A., Coker, E. S., Ssempebwa, J. C., and Isunju, J. B.: Use of biomass fuels predicts indoor particulate matter and carbon monoxide concentrations; evidence from an informal urban settlement in Fort Portal city, Uganda, *BMC Public Health*, 22, 1723, <https://doi.org/10.1186/s12889-022-14015-w>, 2022.
- Keita, S., Lioussé, C., Assamoi, E.-M., Doumbia, T., N'Datchoh Touré, E., Sylvain, G., Elguindi, N., Granier, C., and Yoboué, V.: African Anthropogenic Emissions Inventory for gases and particles, ECCAD [data set], <https://doi.org/10.25326/56>, 2020.
- Keita, S., Lioussé, C., Assamoi, E.-M., Doumbia, T., N'Datchoh, E. T., Gnamien, S., Elguindi, N., Granier, C., and Yoboué, V.: African anthropogenic emissions inventory for gases and particles from 1990 to 2015, *Earth Syst. Sci. Data*, 13, 3691–3705, <https://doi.org/10.5194/essd-13-3691-2021>, 2021.
- Khoder, M. I.: Diurnal, seasonal and weekdays-weekends variations of ground level ozone concentrations in an urban area in greater Cairo, *Environ. Monit. Assess.*, 149, 349–362, 2009.
- Lama, S., Houweling, S., Boersma, K. F., Eskes, H., Aben, I., Denier van der Gon, H. A. C., Krol, M. C., Dolman, H., Borsdorff, T., and Lorente, A.: Quantifying burning efficiency in megacities using the NO₂/CO ratio from the Tropospheric Monitoring Instrument (TROPOMI), *Atmos. Chem. Phys.*, 20, 10295–10310, <https://doi.org/10.5194/acp-20-10295-2020>, 2020.
- Landgraf, J., de Brugh, J., Scheepmaker, R., Borsdorff, T., Houweling, S., and Hasekamp, O.: Algorithm theoretical baseline document for sentinel-5 precursor: Carbon monoxide total column retrieval, Netherlands Institute for Space Research, the Netherlands, SRON-S5P-LEV2-RP-002, 2018.
- Lindén, J., Thorsson, S., and Eliasson, I.: Carbon Monoxide in Ouagadougou, Burkina Faso – A comparison between urban background, roadside and in-traffic measurements, *Water Air Soil Poll.*, 188, 345–353, 2008.
- Liu, X., Yang, S., Li, Q., Kumar, A., Weaver, S., and Liu, S.: Sub-seasonal forecast skills and biases of global summer monsoons in the NCEP Climate Forecast System version 2, *Clim. Dynam.*, 42, 1487–1508, 2014.
- MacDonald, C. G., Mastrogiacomo, J.-P., Laughner, J. L., Hedelius, J. K., Nassar, R., and Wunch, D.: Estimating enhancement ratios of nitrogen dioxide, carbon monoxide and carbon dioxide using satellite observations, *Atmos. Chem. Phys.*, 23, 3493–3516, <https://doi.org/10.5194/acp-23-3493-2023>, 2023.
- Macknick, J.: Energy and CO₂ emission data uncertainties, *Carbon Manag.*, 2, 189–205, 2011.
- Marais, E. A. and Wiedinmyer, C.: Air quality impact of diffuse and inefficient combustion emissions in Africa (DICE-Africa), *Environ. Sci. Technol.*, 50, 10739–10745, 2016.
- McKain, K., Wofsy, S. C., Nehr Korn, T., Eluszkiewicz, J., Ehleringer, J. R., and Stephens, B. B.: Assessment of ground-

- based atmospheric observations for verification of greenhouse gas emissions from an urban region, *P. Natl. Acad. Sci. USA*, 109, 8423–8428, 2012.
- Molod, A., Takacs, L., Suarez, M., Bacmeister, J., Song, I.-S., and Eichmann, A.: The GEOS-5 atmospheric general circulation model: Mean climate and development from MERRA to Fortuna, Tech. rep., NASA technical reports, 11–37, GSFC.TM.01153.2012, 2012.
- NCEP: NCEP FNL Operational Model Global Tropospheric Analyses, continuing from July 1999, Research Data Archive at the National Center for Atmospheric Research, Computational and Information Systems Laboratory [data set], Boulder CO, <https://doi.org/10.5065/D6M043C6>, 2000.
- NDACC-IRWG: <https://www2.acom.ucar.edu/irwg>, last access: 30 March 2023.
- Nielsen, O.-K.: EMEP/EEA air pollutant emission inventory guidebook 2013, Technical guidance to prepare national emission inventories, European Environment Agency, 11–23, ISBN 9789292134037, 2013.
- Oda, T., Bun, R., Kinakh, V., Topylko, P., Halushchak, M., Marland, G., Lauvaux, T., Jonas, M., Maksyutov, S., Nahorski, Z., et al.: Errors and uncertainties in a gridded carbon dioxide emissions inventory, *Mitig. Adapt. Strat. Gl.*, 24, 1007–1050, 2019.
- Oreggioni, G. D., Ferraiolo, F. M., Crippa, M., Muntean, M., Schaaf, E., Guizzardi, D., Solazzo, E., Duerr, M., Perry, M., and Vignati, E.: Climate change in a changing world: Socio-economic and technological transitions, regulatory frameworks and trends on global greenhouse gas emissions from EDGAR v. 5.0, *Global Environ. Chang.*, 70, 102350, <https://doi.org/10.1016/j.gloenvcha.2021.102350>, 2021.
- Pandey, S., Gautam, R., Houweling, S., Van Der Gon, H. D., Sadavarte, P., Borsdorff, T., Hasekamp, O., Landgraf, J., Tol, P., van Kempen, T., Hoogeveen, R., van Hees, R., Hamburg, S. P., Maasackers, J. D., and Aben, I.: Satellite observations reveal extreme methane leakage from a natural gas well blowout, *P. Natl. Acad. Sci. USA*, 116, 26376–26381, 2019.
- Park, H., Jeong, S., Park, H., Labzovskii, L. D., and Bowman, K. W.: An assessment of emission characteristics of Northern Hemisphere cities using spaceborne observations of CO₂, CO, and NO₂, *Remote Sens. Environ.*, 254, 112246, <https://doi.org/10.1016/j.rse.2020.112246>, 2021.
- Park, K., Emmons, L. K., Wang, Z., and Mak, J. E.: Large interannual variations in nonmethane volatile organic compound emissions based on measurements of carbon monoxide, *Geophys. Res. Lett.*, 40, 221–226, 2013.
- Plant, G., Kort, E. A., Murray, L. T., Maasackers, J. D., and Aben, I.: Evaluating urban methane emissions from space using TROPOMI methane and carbon monoxide observations, *Remote Sens. Environ.*, 268, 112756, <https://doi.org/10.1016/j.rse.2021.112756>, 2022.
- Powers, J. G., Klemp, J. B., Skamarock, W. C., Davis, C. A., Dudhia, J., Gill, D. O., Coen, J. L., Gochis, D. J., Ahmadov, R., Peckham, S. E., et al.: The weather research and forecasting model: Overview, system efforts, and future directions, *B. Am. Meteorol. Soc.*, 98, 1717–1737, 2017.
- Powers, J. G., Klemp, J. B., Skamarock, W. C., Davis, C. A., Dudhia, J., Gill, D. O., Coen, J. L., Gochis, D. J., Ahmadov, R., Peckham, S. E., Grell, G. A., Michalakes, J., Trahan, S., Benjamin, S. G., Alexander, C. R., Dimego, G. J., Wang, W., Schwartz, C. S., Romine, G. S., Liu, Z., Snyder, C., Chen, F., Barlage, M. J., Yu, W., and Duda, M. G.: Atmospheric Chemistry Observations and Modeling version 4.1.5, GitHub [code], <https://github.com/wrf-model/WRF/releases>, last access: 1 January 2023.
- Qu, Z., Henze, D. K., Worden, H. M., Jiang, Z., Gaubert, B., Theys, N., and Wang, W.: Sector-based top-down estimates of NO_x, SO₂, and CO emissions in East Asia, *Geophys. Res. Lett.*, 49, e2021GL096009, <https://doi.org/10.1029/2021GL096009>, 2022.
- Rey-Pommier, A., Chevallier, F., Ciais, P., Broquet, G., Christoudias, T., Kushta, J., Hauglustaine, D., and Sciare, J.: Quantifying NO_x emissions in Egypt using TROPOMI observations, *Atmos. Chem. Phys.*, 22, 11505–11527, <https://doi.org/10.5194/acp-22-11505-2022>, 2022.
- Rozante, J. R., Rozante, V., Souza Alvim, D., Ocimar Manzi, A., Barboza Chiquetto, J., Siqueira D’Amelio, M. T., and Moreira, D. S.: Variations of carbon monoxide concentrations in the megacity of São Paulo from 2000 to 2015 in different time scales, *Atmosphere*, 8, 81, <https://doi.org/10.3390/atmos8050081>, 2017.
- Sadavarte, P., Pandey, S., Maasackers, J. D., Lorente, A., Borsdorff, T., Denier van der Gon, H., Houweling, S., and Aben, I.: Methane emissions from superemitting coal mines in Australia quantified using TROPOMI satellite observations, *Environ. Sci. Technol.*, 55, 16573–16580, 2021a.
- Sadavarte, P., Pandey, S., Maasackers, J. D., Lorente, A., Borsdorff, T., Denier van der Gon, H., Houweling, S., and Aben, I.: Methane emissions from superemitting coal mines in Australia quantified using TROPOMI satellite observations, *Environ. Sci. Technol.*, 55, 16573–16580, 2021b.
- Saleh, S., Sambakunsi, H., Makina, D., Chinouya, M., Kumwenda, M., Chirombo, J., Semple, S., Mortimer, K., and Rylance, J.: Personal exposures to fine particulate matter and carbon monoxide in relation to cooking activities in rural Malawi, *Wellcome Open Research*, 7, e251, <https://doi.org/10.12688/wellcomeopenres.18050.2>, 2023.
- Sha, M. K., Langerock, B., Blavier, J.-F. L., Blumenstock, T., Borsdorff, T., Buschmann, M., Dehn, A., De Mazière, M., Deutscher, N. M., Feist, D. G., García, O. E., Griffith, D. W. T., Grutter, M., Hannigan, J. W., Hase, F., Heikkinen, P., Hermans, C., Iraci, L. T., Jeseck, P., Jones, N., Kivi, R., Kumps, N., Landgraf, J., Lorente, A., Mahieu, E., Makarova, M. V., Mellqvist, J., Metzger, J.-M., Morino, I., Nagahama, T., Notholt, J., Ohyama, H., Ortega, I., Palm, M., Petri, C., Pollard, D. F., Rettinger, M., Robinson, J., Roche, S., Roehl, C. M., Röhling, A. N., Rousogonous, C., Schneider, M., Shiomi, K., Smale, D., Stremme, W., Strong, K., Sussmann, R., Té, Y., Uchino, O., Velazco, V. A., Vigouroux, C., Vrekoussis, M., Wang, P., Warneke, T., Wizenberg, T., Wunch, D., Yamanouchi, S., Yang, Y., and Zhou, M.: Validation of methane and carbon monoxide from Sentinel-5 Precursor using TCCON and NDACC-IRWG stations, *Atmos. Meas. Tech.*, 14, 6249–6304, <https://doi.org/10.5194/amt-14-6249-2021>, 2021.
- Silva, S. J., Arellano, A. F., and Worden, H. M.: Toward anthropogenic combustion emission constraints from space-based analysis of urban CO₂/CO sensitivity, *Geophys. Res. Lett.*, 40, 4971–4976, 2013.
- Stavrakou, T., Müller, J.-F., Bauwens, M., Boersma, K., and Van Geffen, J.: Satellite evidence for changes in the NO₂ weekly cycle over large cities, *Sci. Rep.-UK*, 10, 10066, <https://doi.org/10.1038/s41598-020-66891-0>, 2020.

- Tian, Y., Liu, C., Sun, Y., Borsdorff, T., Landgraf, J., Lu, X., Palm, M., and Notholt, J.: Satellite observations reveal a large CO emission discrepancy from industrial point sources over China, *Geophys. Res. Lett.*, 49, e2021GL097312, <https://doi.org/10.1029/2021GL097312>, 2022a.
- Tian, Y., Sun, Y., Borsdorff, T., Liu, C., Liu, T., Zhu, Y., Yin, H., and Landgraf, J.: Quantifying CO emission rates of industrial point sources from Tropospheric Monitoring Instrument observations, *Environ. Res. Lett.*, 17, 014057, <https://doi.org/10.1088/1748-9326/ac3b1a>, 2022b.
- Varon, D. J., Jacob, D. J., McKeever, J., Jervis, D., Durak, B. O. A., Xia, Y., and Huang, Y.: Quantifying methane point sources from fine-scale satellite observations of atmospheric methane plumes, *Atmos. Meas. Tech.*, 11, 5673–5686, <https://doi.org/10.5194/amt-11-5673-2018>, 2018.
- Varon, D. J., Jacob, D. J., Jervis, D., and McKeever, J.: Quantifying time-averaged methane emissions from individual coal mine vents with GHGSat-D satellite observations, *Environ. Sci. Technol.*, 54, 10246–10253, 2020.
- Veefkind, J., Aben, I., McMullan, K., Förster, H., De Vries, J., Otter, G., Claas, J., Eskes, H., De Haan, J., Kleipool, Q., van Weele, M., Hasekamp, O., Hoogeveen, R., Landgraf, J., Snel, R., Tol, P., Ingmann, P., Voors, R., Kruizinga, B., Vink, R., Visser, H., and Levelt, P.F.: TROPOMI on the ESA Sentinel-5 Precursor: A GMES mission for global observations of the atmospheric composition for climate, air quality and ozone layer applications, *Remote Sens. Environ.*, 120, 70–83, 2012.
- Von Schneidemesser, E., Monks, P. S., and Plass-Duelmer, C.: Global comparison of VOC and CO observations in urban areas, *Atmos. Environ.*, 44, 5053–5064, 2010.
- Wu, D., Liu, J., Wennberg, P. O., Palmer, P. I., Nelson, R. R., Kiel, M., and Eldering, A.: Towards sector-based attribution using intra-city variations in satellite-based emission ratios between CO₂ and CO, *Atmos. Chem. Phys.*, 22, 14547–14570, <https://doi.org/10.5194/acp-22-14547-2022>, 2022.
- Wuebbles, D. J. and Hayhoe, K.: Atmospheric methane and global change, *Earth-Sci. Rev.*, 57, 177–210, 2002.
- Wunch, D., Toon, G. C., Blavier, J.-F. L., Washenfelder, R. A., Notholt, J., Connor, B. J., Griffith, D. W., Sherlock, V., and Wennberg, P. O.: The total carbon column observing network, *Philos. T. Roy. Soc. A*, 369, 2087–2112, 2011.
- Yumimoto, K., Uno, I., and Itahashi, S.: Long-term inverse modeling of Chinese CO emission from satellite observations, *Environ. Pollut.*, 195, 308–318, 2014.
- Zakari, M. M., Nguema, F. P., Adamou, A., Esly, E., and Zakari, A.: Air Pollution linked to Road Traffic: Assessment of Carbon Monoxide (CO) Emissions in Zinder City, Niger Republic, *J. Sci. Technol.*, 9, 111–120, 2020.
- Zhong, Q., Huang, Y., Shen, H., Chen, Y., Chen, H., Huang, T., Zeng, E. Y., and Tao, S.: Global estimates of carbon monoxide emissions from 1960 to 2013, *Environ. Sci. Pollut. R.*, 24, 864–873, 2017.