



Supplement of

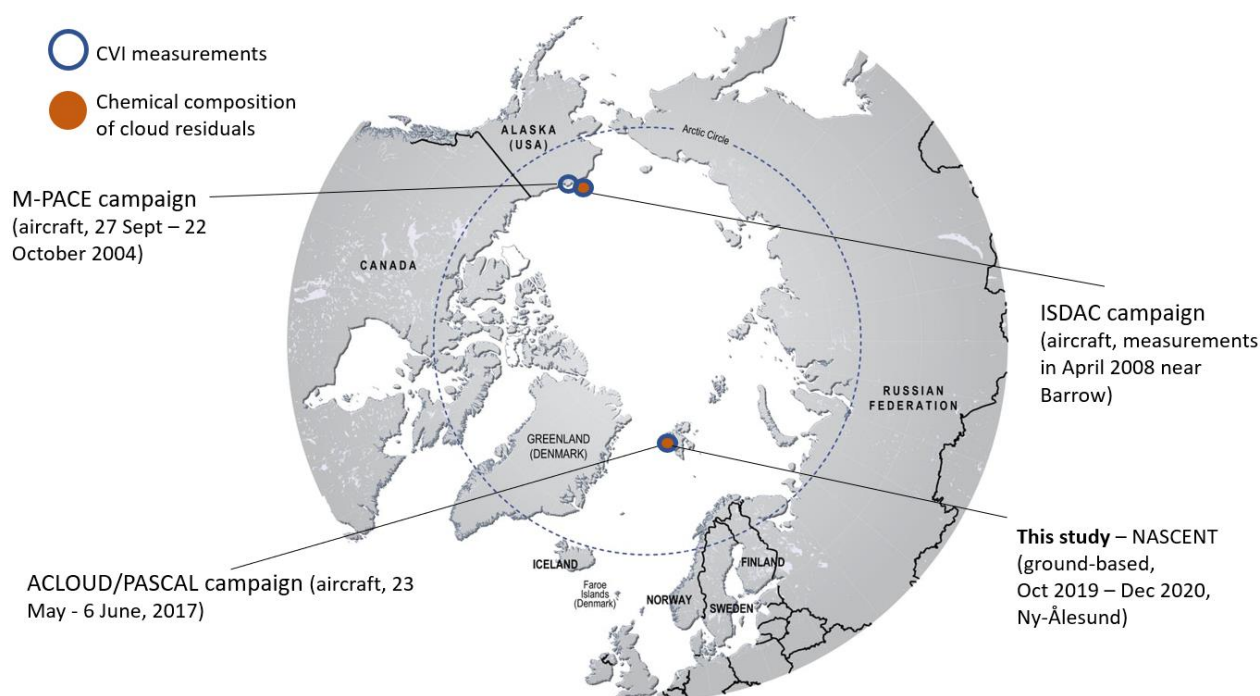
Revealing the chemical characteristics of Arctic low-level cloud residuals – in situ observations from a mountain site

Yvette Gramlich et al.

Correspondence to: Claudia Mohr (claudia.mohr@psi.ch)

The copyright of individual parts of the supplement might differ from the article licence.

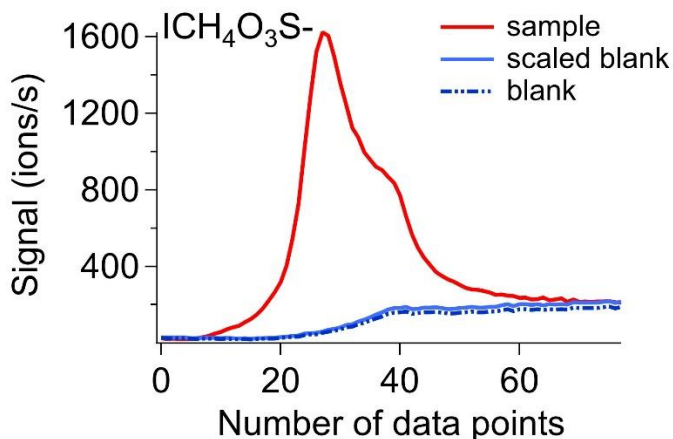
S1 Overview of previous studies on the chemical composition of cloud residuals



5 **Figure S1: Map of the Arctic showing the sampling location of NASCENT and the locations of previous studies on the chemical composition of Arctic cloud residuals.** (Map taken from <https://www.grida.no/resources/8378>.) We note that our cloud residual samples were obtained using a Ground-based Counterflow Virtual Impactor (GCVI), whereas previous cloud residual samples were aircraft based (CVI) (Verlinde et al., 2007; McFarquhar et al., 2011; Wendisch et al., 2019).

S2 Background correction for FIGAERO-CIMS data

- 10 The background correction was done following the approach recommended in Cai et al. (2023). We scaled the blank heating signal to the end of the sample signal and subtracted the integrated scaled blank signal from the integrated sample heating (Fig. S2). Since we did not have a blank for each cloud residual sample, we took the blank that was measured closest (with respect to time) to the sample as the respective background (Table S1).



15

Figure S2: Example of the signal of $\text{CH}_4\text{O}_3\text{S}^-$ (methanesulfonic acid, MSA) measured as iodide cluster in the cloud residual of Jun 26, 2020, showing how the scaled blank signal was used for background determination. The dashed blue line is the original blank signal and the solid blue line the blank signal scaled to the end of the sample heating. The interval between two data points equals a time period of 30 s.

20

Table S1: Cloud residual samples and the respective blanks used for the background correction. The time in the brackets refers to the respective FIGAERO-CIMS sampling time of the blanks. The sampling times for the cloud residual samples can be found in the main text, Table 1.

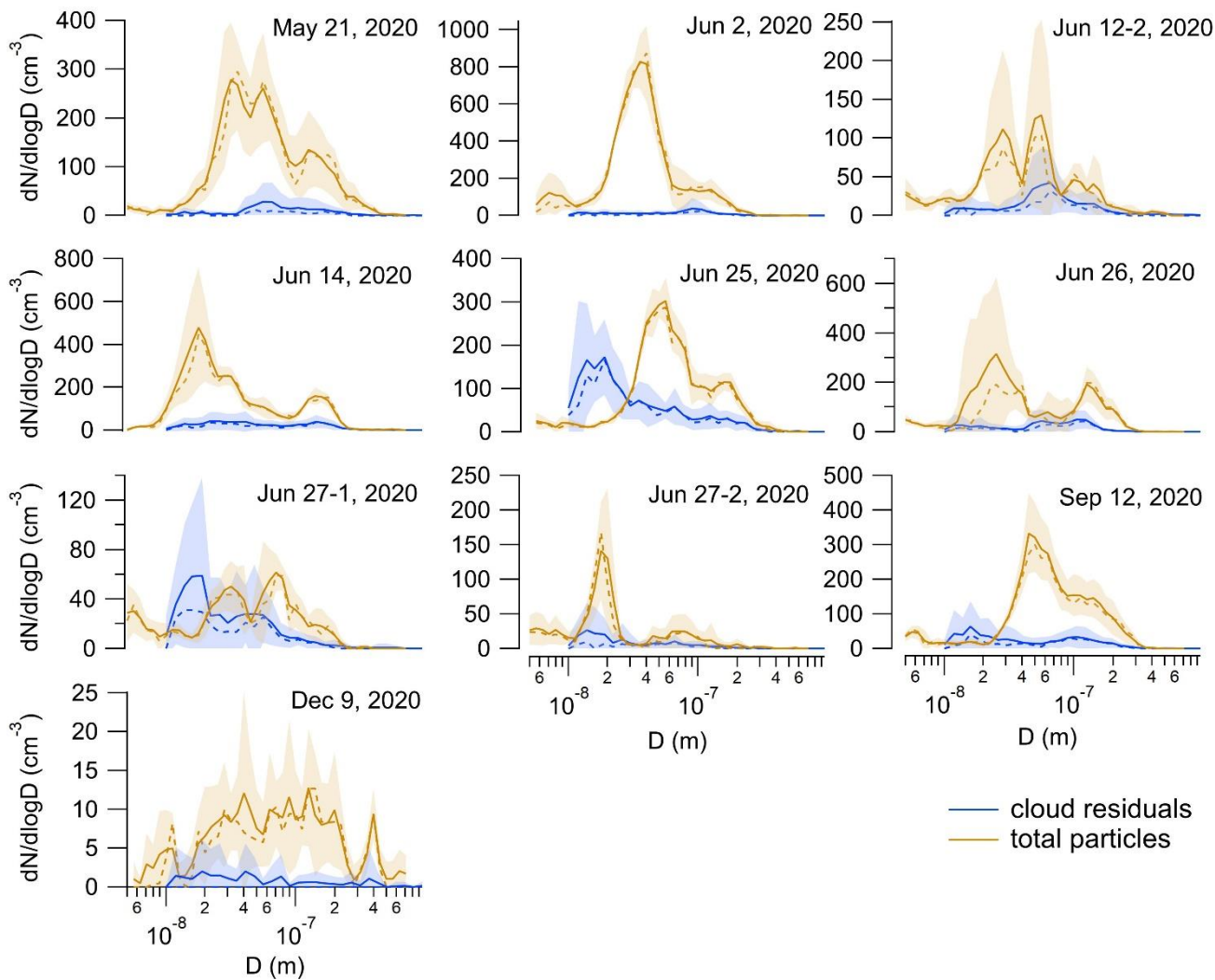
Cloud residual sample	Blank used	Cloud residual sample	Blank used
Dec 25, 2019	Jan 19, 2020 (13:05:51-15:36:14)	Jun 25, 2020	Jun 25, 2020 (13:29:00-15:59:09)
May 18, 2020	May 18, 2020 (11:03:03-13:33:05)	Jun 26, 2020	Jun 26, 2020 (10:29:00-12:59:08)
May 21, 2020	May 21, 2020 (16:17:47-18:47:49)	Jun 27-1, 2020	Jun 27, 2020 (07:28:59-09:59:07)
Jun 2, 2020	Jun 12, 2020 (14:50:11-17:20:19)	Jun 27-2, 2020	Jun 27, 2020 (07:28:59-09:59:07)
Jun 12-1, 2020	Jun 12, 2020 (14:50:11-17:20:19)	Sep 12, 2020	Sep 12, 2020 (23:01:39-Sep 13 01:32:00)
Jun 12-2, 2020	Jun 12, 2020 (14:50:11-17:20:19)	Oct 28, 2020	Sep 12, 2020 (23:01:39-Sep 13 01:32:00)
Jun 14, 2020	Jun 13, 2020 (22:20:25-Jun 14 00:50:03)	Dec 9, 2020	Sep 12, 2020 (23:01:39-Sep 13 01:32:00)

25

S3 Cloud residual size distributions

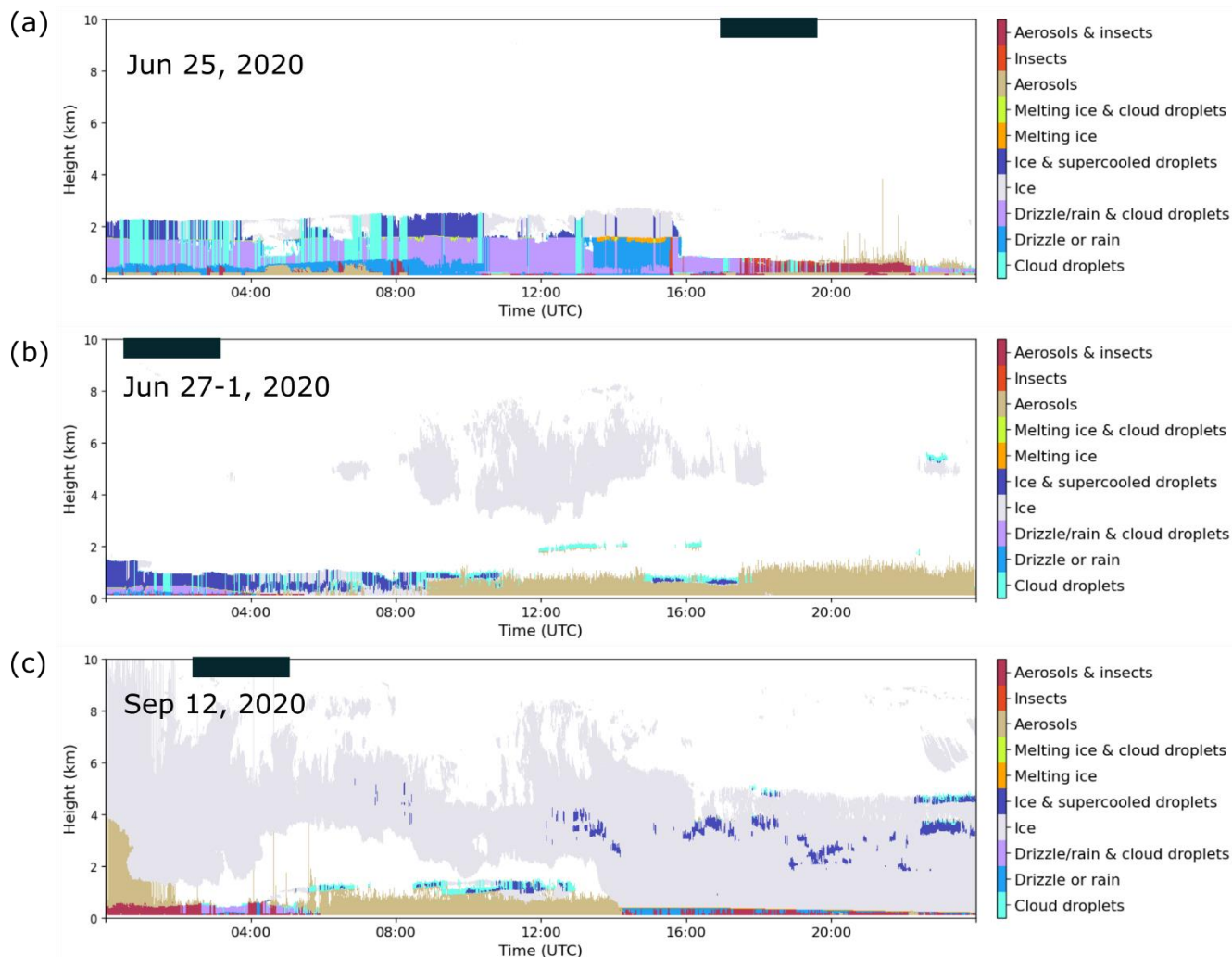
30 The number size distributions of all cloud residuals not shown in the main text are presented in Figure S3.

During the times when we sampled the cloud residuals on June 25, June 27-1 and September 12, 2020, there was also drizzle present. This can be seen in data from the condensation particle counter (CPC, model 3772, TSI Inc., USA), as well as in the Cloudnet target classification (Fig. S4). The drizzle droplets can splash when they hit the funnel of the wind tunnel of the
35 GCVI and produce several, smaller droplets. This can then be seen as a spike in the total particle number concentration of the CPC measured at a time resolution of 1 s (N_{tot} 1s). This concentration can be compared to the particle number concentration measured with another CPC (model 3772, TSI Inc., USA) behind the differential mobility analyser (DMA, medium Vienna-type, length 0.28 m, outer radius 0.033 m, inner radius 0.025 m) and integrated over the entire size range (N_{int} , time resolution
40 7 min) by averaging it to the same time resolution (N_{tot} mean). If the size selected by the DMA (D_{scan}) is in the size range of the Aitken mode particles when droplet splashing occurs, the integrated number concentration (N_{int}) will be much higher than N_{tot} mean for the same time interval. Additionally, a large number of Aitken mode particles can be observed in the number size distribution. Therefore, we removed the datapoints for the number size distributions where the median ratio of $N_{\text{int}}/N_{\text{tot}}$ over the entire 2.5 h sampling time was larger than a certain threshold. Fig. S6 shows an example (from September 12, 2020) of how drizzle splashing can be observed in the Differential Mobility Particle Sizer (DMPS) data. The median ratios, the
45 selected threshold and the corresponding number of datapoints that are removed are shown in Table S2.

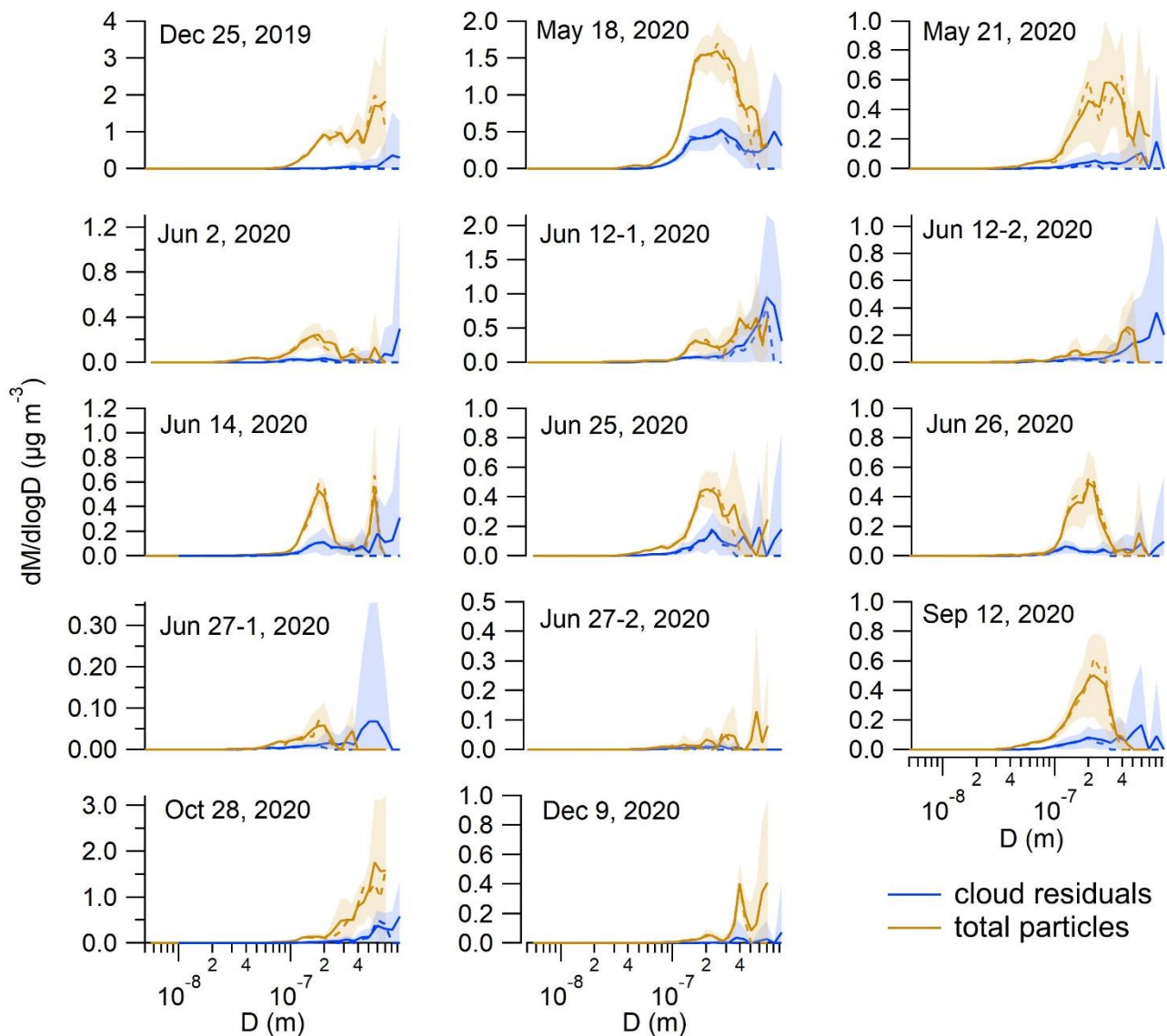


50 **Figure S3: Average number size distributions of the cloud residuals and the total particles during the corresponding 2.5 h FIGAERO-CIMS sampling time of all remaining samples not shown in the main text. The shaded area represents the standard deviation.**

55



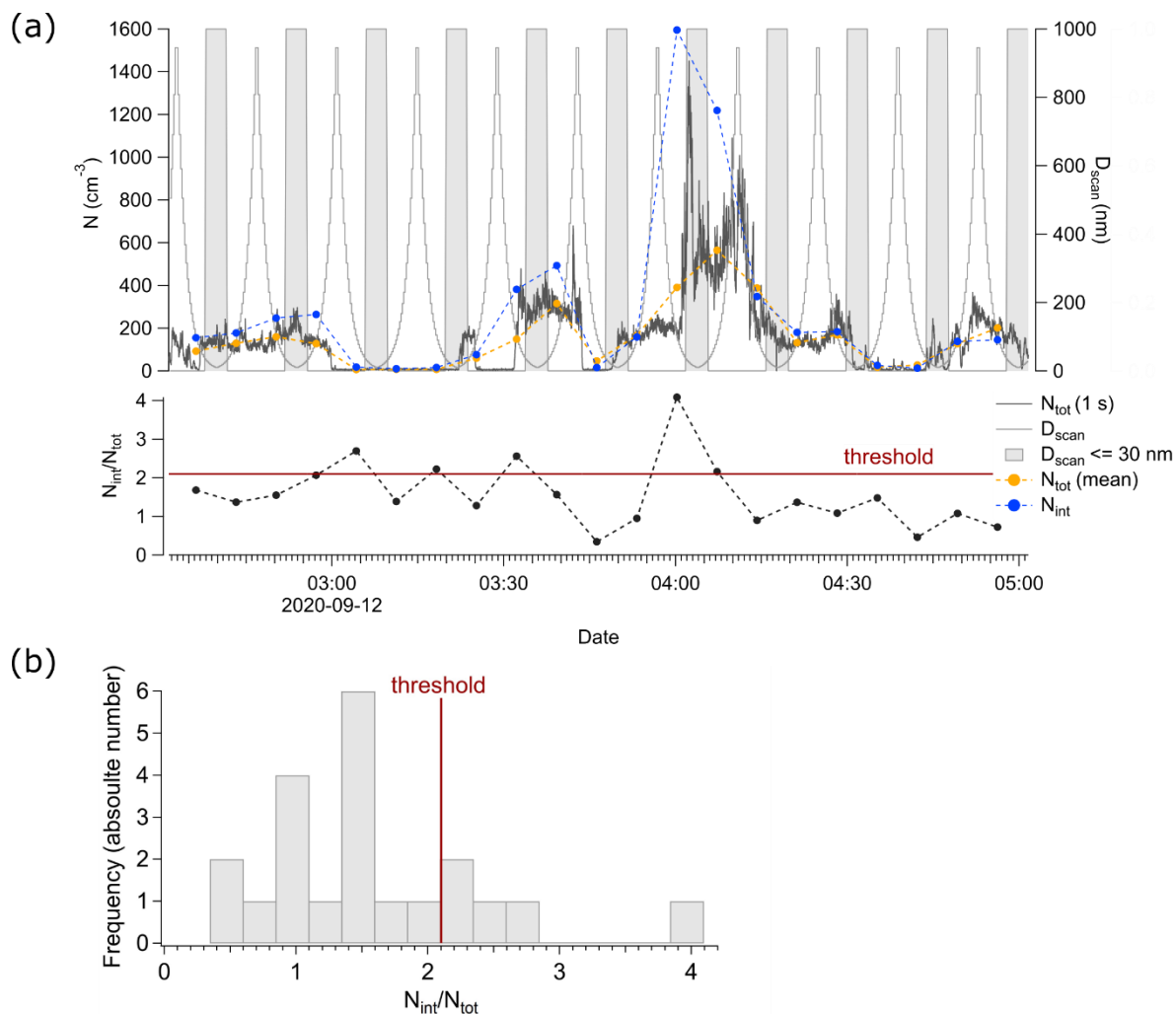
60 **Figure S4: Cloudnet target classification (modified by adding the respective cloud residual sample date, and indicating the approximate sampling times of the cloud with a black bar on top of each subfigure) for (a) June 25, 2020 (Cloudnet (2021), last access: 2022-11-24, 15:35 UTC), (b) June 27-2, 2020 (Cloudnet (2021b), last access: 2022-11-24, 15:28 UTC), (c) September 12, 2020 (Cloudnet (2021a), last access: 2022-11-24, 15:38 UTC), indicating that there was drizzle present during the sampling times of the cloud residuals.**



65

Figure S5: Average mass size distributions of all the cloud residual samples and the total particle population. The shaded area indicates the standard deviation. For the conversion from number to mass a density of 1.3 g cm^{-3} was used, representing secondary organic aerosol (e.g. Alfarrá et al., 2006; Malloy et al., 2009).

70

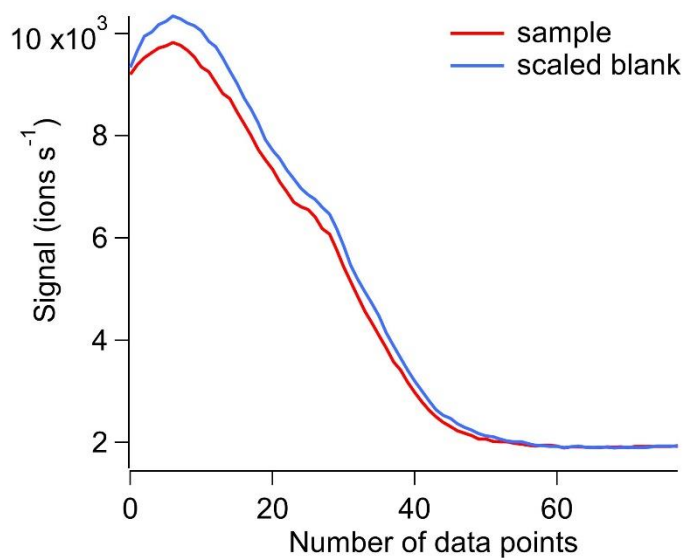


75 **Figure S6:** (a) DMPS data from the cloud residual sample of September 12, 2020, as an example of how the droplet splashing of drizzle droplets can be seen as spikes in the N_{tot} (1 s) time series. $D_{\text{scan}} \leq 30$ indicates the time periods when the size selected by the DMA was in the small size range, selecting particles up to diameters of 30 nm. The lower panel shows the ratio of $N_{\text{int}}/N_{\text{tot}}$ and the location of the threshold. (b) Histogram of the ratio $N_{\text{tot}}/N_{\text{int}}$ indicating the location of the selected threshold for filtering the data. For the ratio $N_{\text{tot}}/N_{\text{int}}$ we took N_{tot} (mean).

Table S2: Cloud case, median ratio of $N_{\text{int}}/N_{\text{tot}}$, selected threshold above which the datapoints were removed, and the corresponding number of datapoints that are removed ($\text{Num}_{\text{removed}}$) from the total number of datapoints ($\text{Num}_{\text{total}}$).

Cloud case	Median ratio $N_{\text{int}}/N_{\text{tot}}$	Threshold	$\text{Num}_{\text{removed}}/\text{Num}_{\text{total}}$
Jun 25, 2020	1.5	1.6	9/22
Jun 27-1, 2020	1.2	1.9	4/22
Sep 12, 2020	1.4	2.1	5/21

S4 Signal of formic acid during the heating

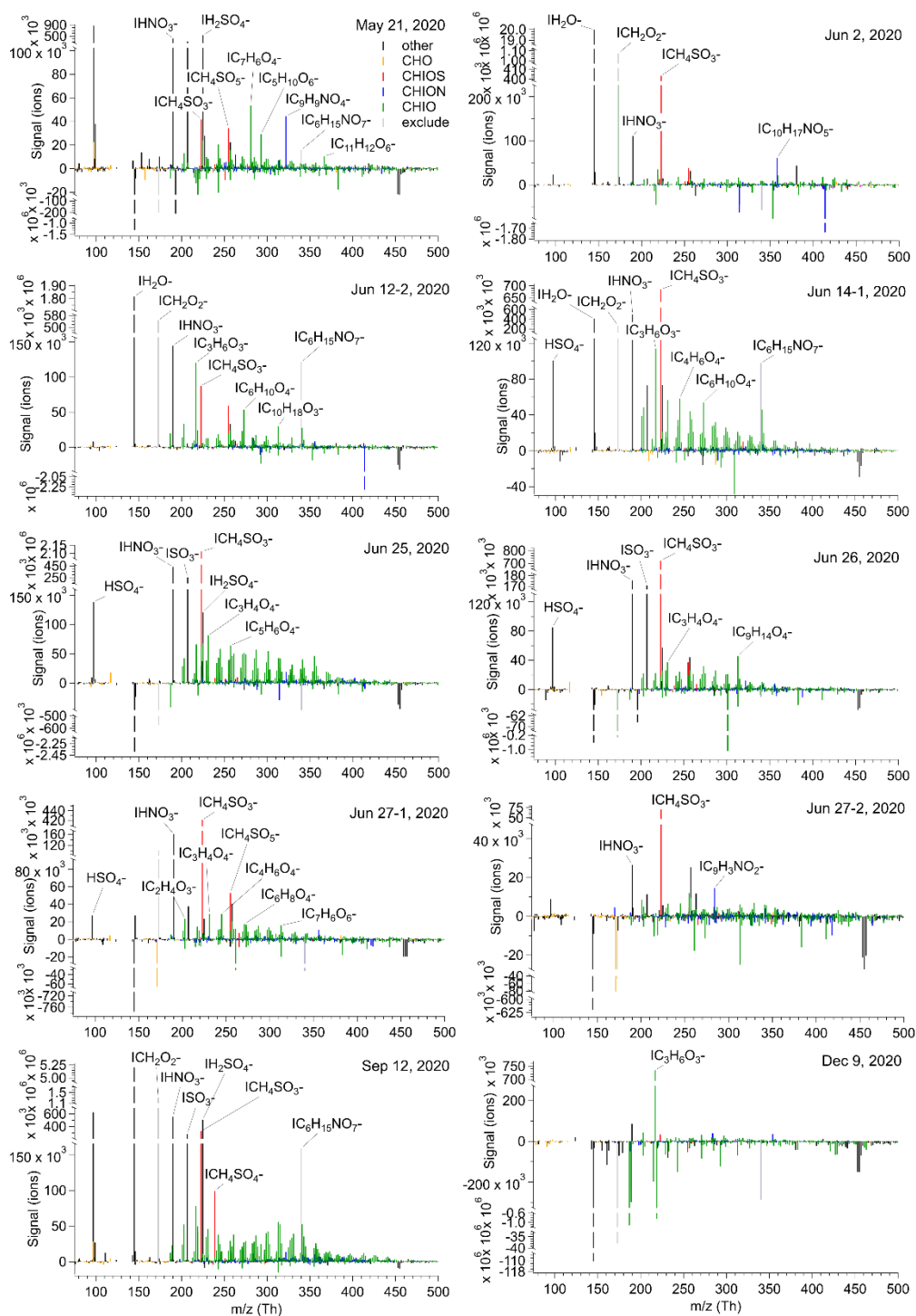


85 **Figure S7: Signal of formic acid (ICH₂O₂⁻) during the heating of the sample and the scaled blank, respectively, as a function of heating time. The interval between two data points equals a time period of 30 s. As an example, the signal here is presented from the cloud residual sample on Jun 25, 2020.**

90

95

S5 Mass spectra of cloud residuals

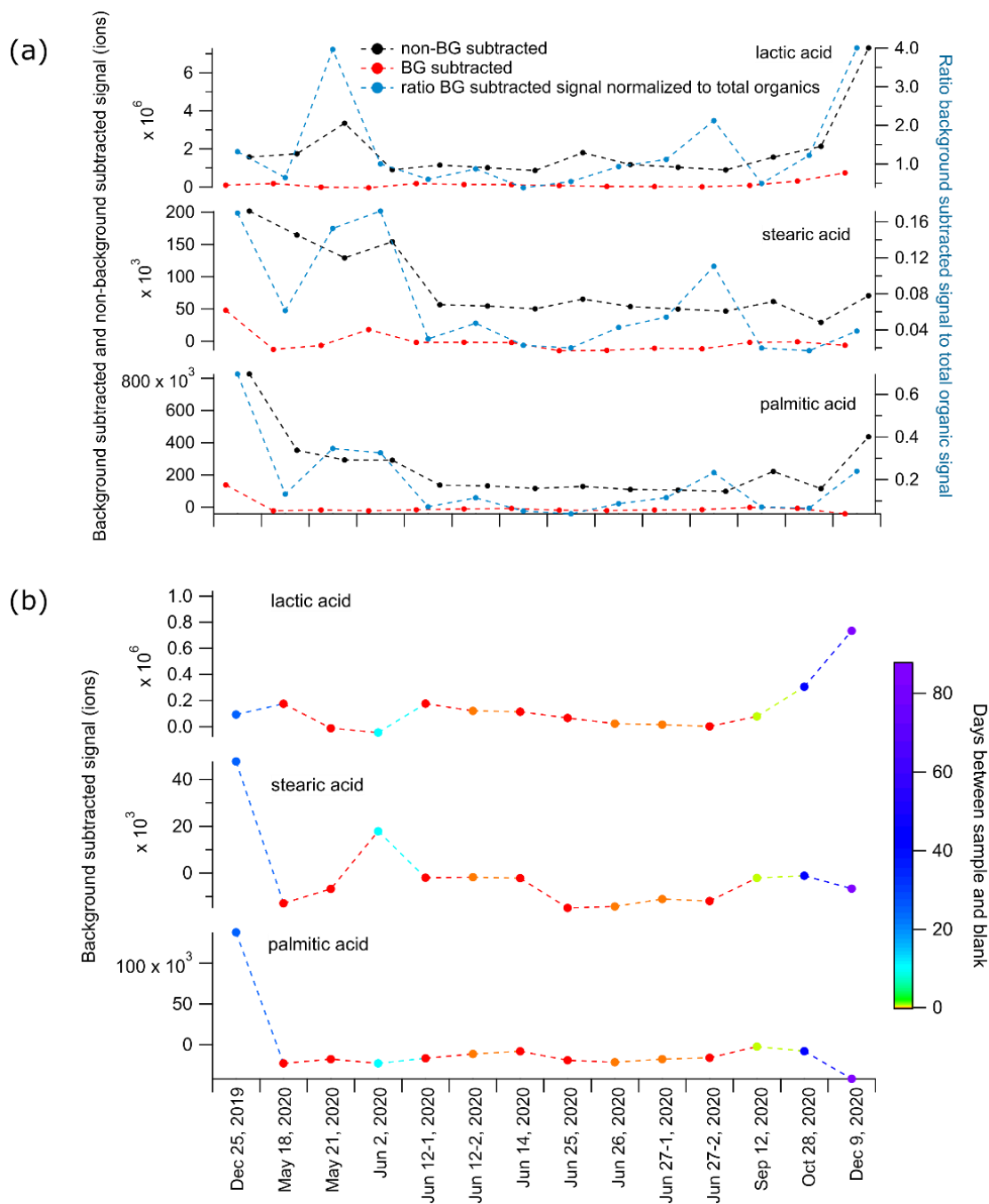


100

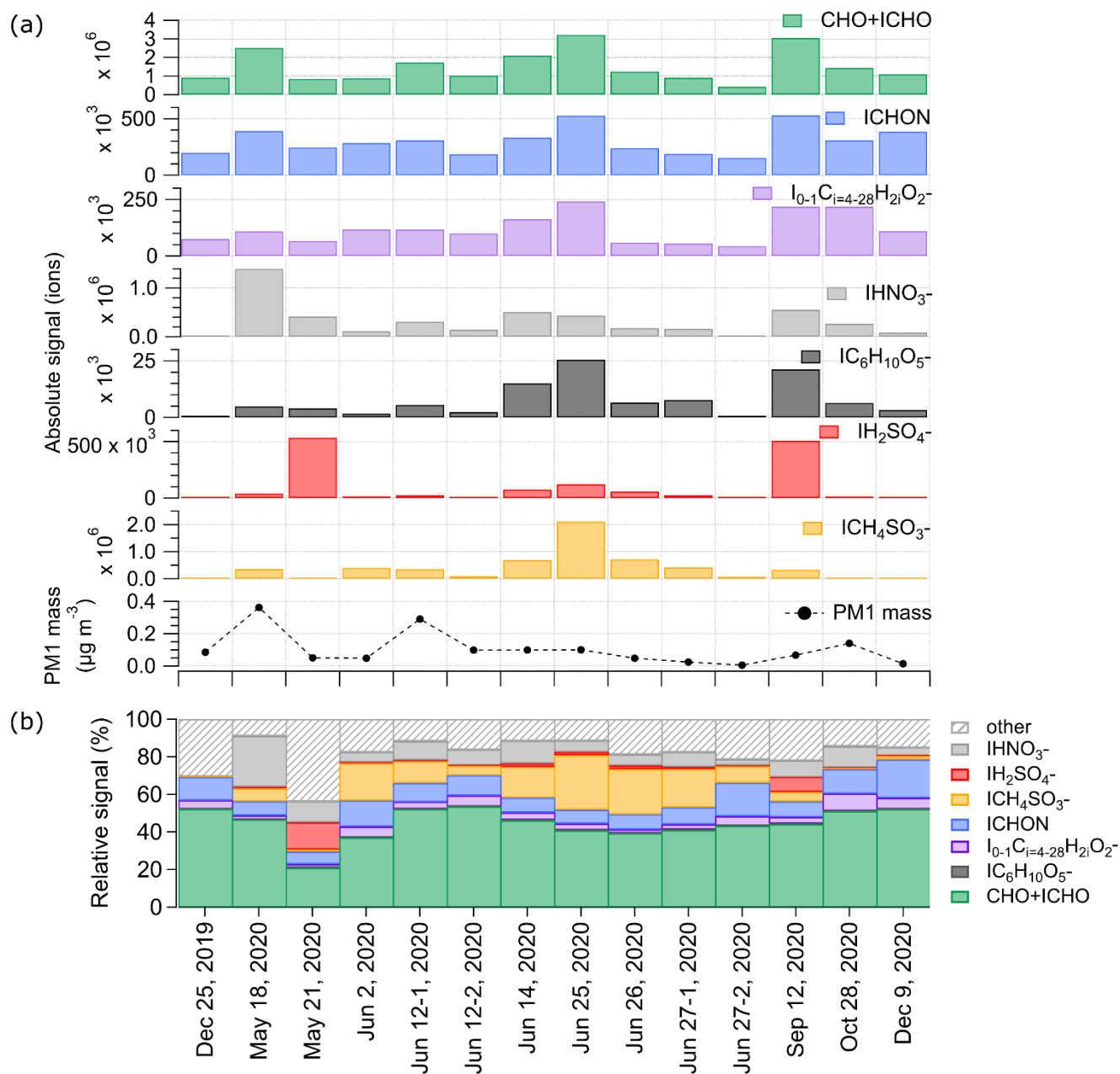
Figure S8: Mass spectra of all cloud residual samples not shown in the main text.

S6 Chemical composition of cloud residuals

105 The chemical formulas referring to lactic acid ($\text{C}_3\text{H}_6\text{O}_3^-$), palmitic acid ($\text{C}_{16}\text{H}_{32}\text{O}_2^-$) and stearic acid ($\text{C}_{18}\text{H}_{36}\text{O}_2^-$) could potentially be related to handling of the GCVI. Based on our data, we were not able to clearly identify if they are only a background signal or if they are actual compounds in the cloud residuals. We observe especially high signals after background subtraction of the compounds in question during the times when the gap in time of sampling and taking a blank was the highest (Dec 25, 2019; Dec 9, 2020, Fig. S9). Fig. S10 shows the chemical composition of the cloud residuals when excluding lactic, palmitic and stearic acid from the absolute and relative signal.



110 **Figure S9: (a) Background (BG) subtracted and non-BG subtracted signals of lactic, stearic and palmitic acid and the ratio of their BG subtracted signal to the total organic signal. (b) Background subtracted signals of lactic, stearic and palmitic acid color coded by the time difference between the sample and the blank.**



115 **Figure S10: (a) Absolute signal of different compound groups (CHO+ICHO, ICHON, $\text{I}_{0-1}\text{C}_{i=4-28}\text{H}_{2i}\text{O}_2^-$ (fatty acids), IHNO_3^- (NA $\text{IC}_6\text{H}_{10}\text{O}_5^-$ (levoglucosan), IH_2SO_4^- (SA), $\text{ICH}_4\text{SO}_3^-$ (MSA)) in the different cloud residual samples, and the respective PM₁ mass. (b) Relative signal of different compound groups in the different cloud residual samples. Note: in the absolute signal view in (a) the CHO+ICHO group contains also the signal of $\text{IC}_6\text{H}_{10}\text{O}_5^-$, and $\text{I}_{0-1}\text{C}_{i=4-28}\text{H}_{2i}\text{O}_2^-$, whereas for the relative signal in (b) the signal from these two groups have been subtracted from CHO+ICHO. This figure is similar to Fig. 5 from the main text, but excluding the three compounds that might be linked to hygiene products (lactic, palmitic and stearic acid).**

120

S7 Back trajectories

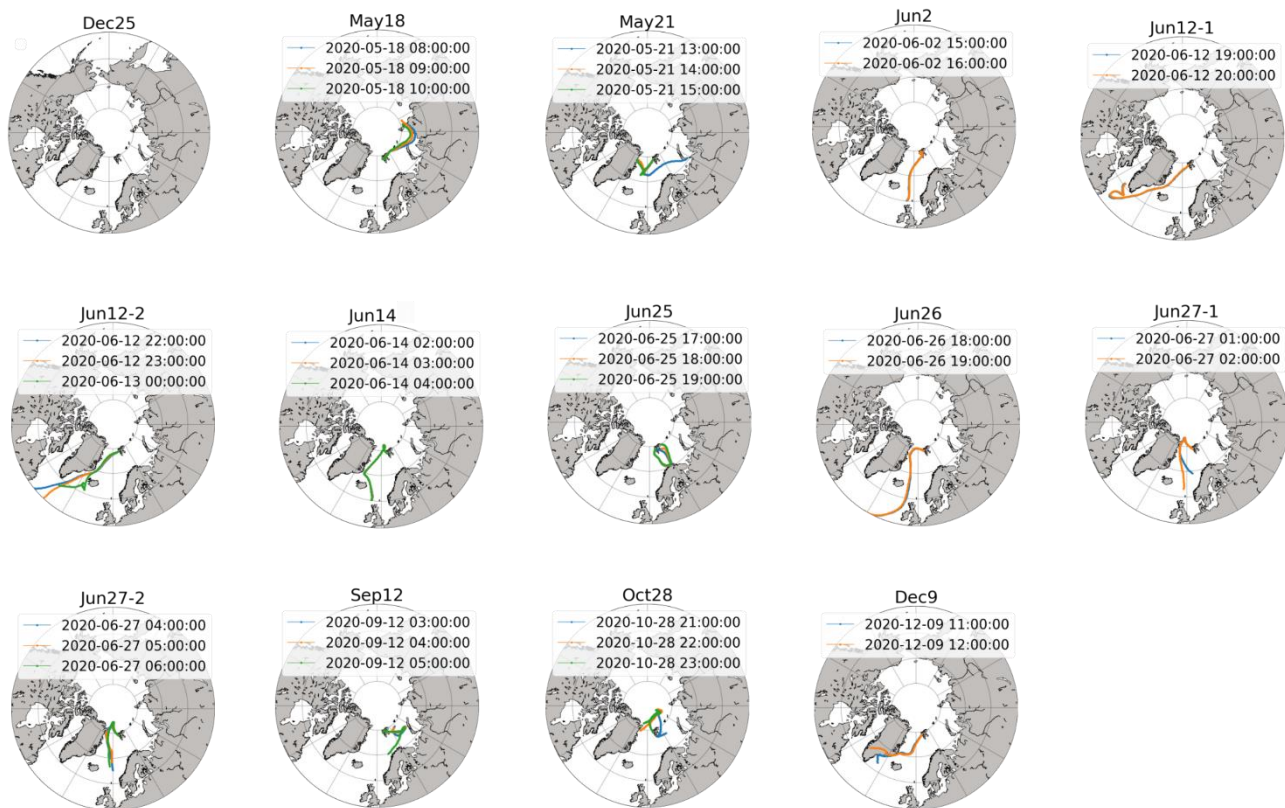
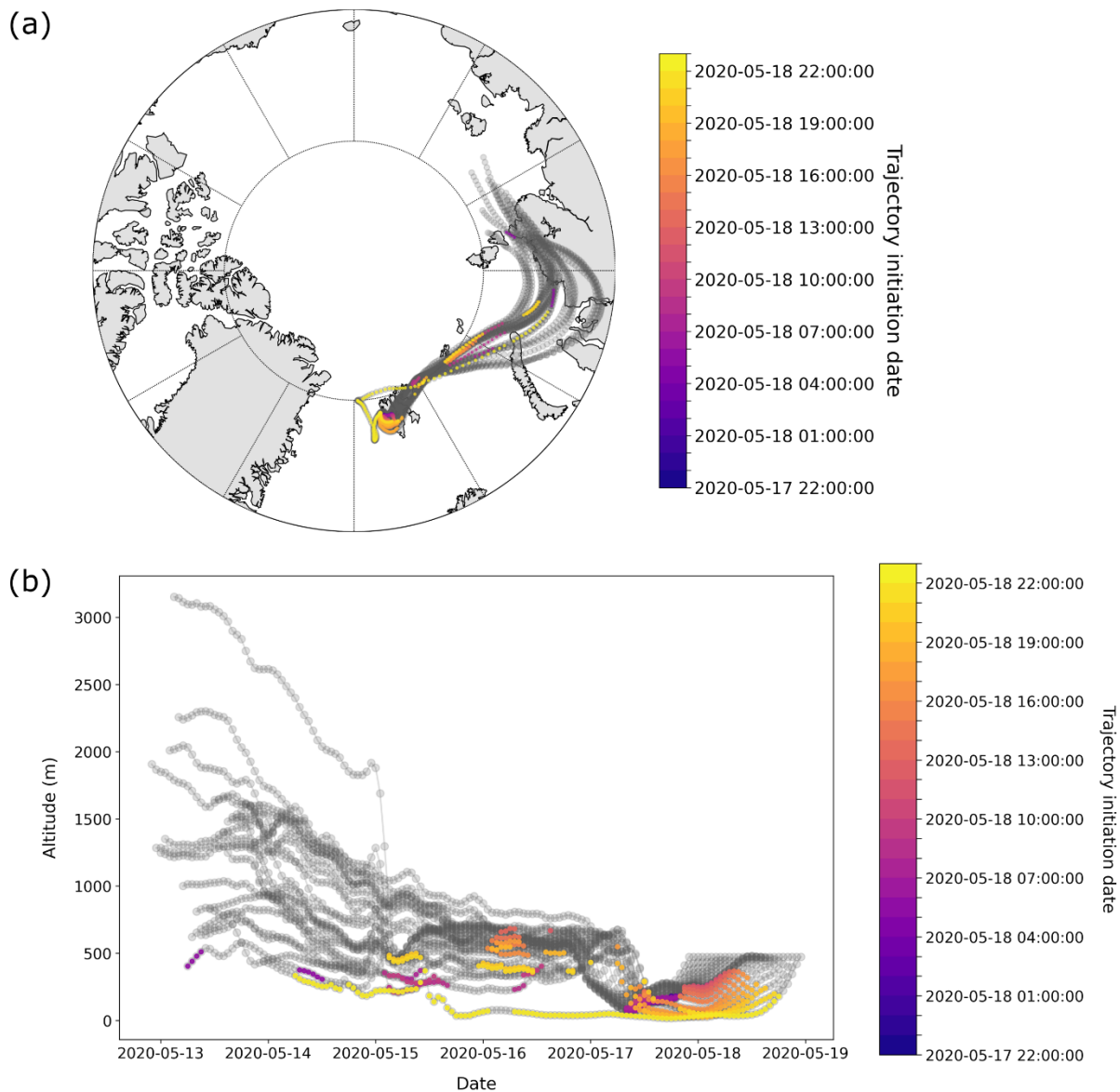


Figure S11: HYSPLIT 5-day back trajectories for the sampling times of all the cloud residual samples.

125

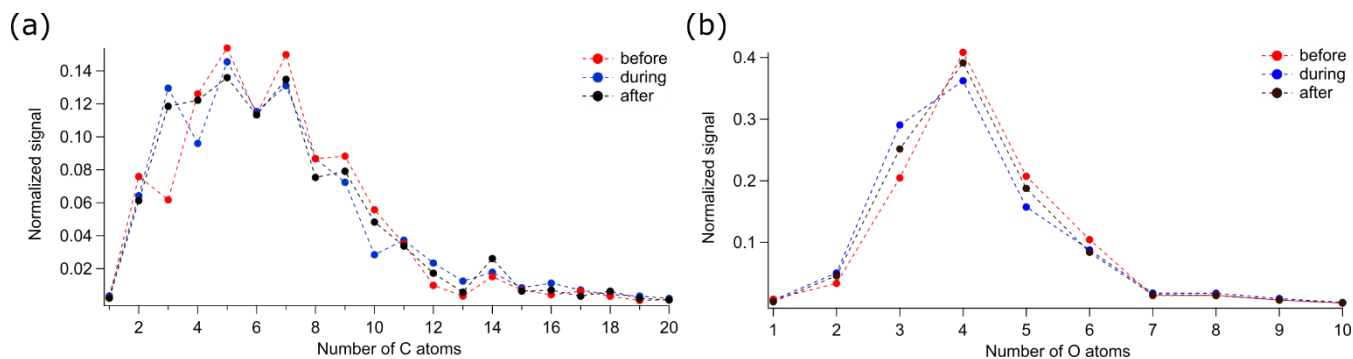
130

S8 Cloud case May 18, 2020



135

Figure S12: 5 days back trajectories of air masses arriving at the Zeppelin Observatory before, during and after the cloud event color coded by time and height with respect to the boundary layer height (BL). (a) Map view, (b) Trajectory height as a function of time. Grey colors indicate times above the BL, and colors indicate times below the BL.



140

Figure S13: Signal contributions of compounds grouped according to different numbers of (a) carbon and (b) and oxygen to the total CHO+CHOI signal for the cloud case May 18, 2020.

145 S References

- Alfarra, M. R., Paulsen, D., Gysel, M., Garforth, A. A., Dommen, J., Prévôt, A. S. H., Worsnop, D. R., Baltensperger, U., and Coe, H.: A mass spectrometric study of secondary organic aerosols formed from the photooxidation of anthropogenic and biogenic precursors in a reaction chamber, *Atmospheric Chem. Phys.*, 6, 5279–5293, <https://doi.org/10.5194/acp-6-5279-2006>, 2006.
- 150 Cai, J., Daellenbach, K. R., Wu, C., Zheng, Y., Zheng, F., Du, W., Haslett, S. L., Chen, Q., Kulmala, M., and Mohr, C.: Characterization of offline analysis of particulate matter with FIGAERO-CIMS, *Atmospheric Meas. Tech.*, 16, 1147–1165, <https://doi.org/10.5194/amt-16-1147-2023>, 2023.
- Cloudnet: Classification data from Ny-Ålesund on 12 September 2020, 2021a.
- Cloudnet: Classification data from Ny-Ålesund on 25 June 2020, 2021b.
- 155 Cloudnet: Classification data from Ny-Ålesund on 27 June 2020, 2021c.
- Malloy, Q. G. J., Nakao, S., Qi, L., Austin, R., Stothers, C., Hagino, H., and Cocker, D. R.: Real-Time Aerosol Density Determination Utilizing a Modified Scanning Mobility Particle Sizer—Aerosol Particle Mass Analyzer System, *Aerosol Sci. Technol.*, 43, 673–678, <https://doi.org/10.1080/02786820902832960>, 2009.
- 160 McFarquhar, G. M., Ghan, S., Verlinde, J., Korolev, A., Strapp, J. W., Schmid, B., Tomlinson, J. M., Wolde, M., Brooks, S. D., Cziczo, D., Dubey, M. K., Fan, J., Flynn, C., Gultepe, I., Hubbe, J., Gilles, M. K., Laskin, A., Lawson, P., Leitch, W. R., Liu, P., Liu, X., Lubin, D., Mazzoleni, C., Macdonald, A.-M., Moffet, R. C., Morrison, H., Ovchinnikov, M., Shupe, M. D., Turner, D. D., Xie, S., Zelenyuk, A., Bae, K., Freer, M., and Glen, A.: Indirect and Semi-direct Aerosol Campaign: The Impact of Arctic Aerosols on Clouds, *Bull. Am. Meteorol. Soc.*, 92, 183–201, <https://doi.org/10.1175/2010BAMS2935.1>, 2011.
- Verlinde, J., Harrington, J. Y., McFarquhar, G. M., Yannuzzi, V. T., Avramov, A., Greenberg, S., Johnson, N., Zhang, G., Poellot, M. R., Mather, J. H., Turner, D. D., Eloranta, E. W., Zak, B. D., Prenni, A. J., Daniel, J. S., Kok, G. L., Tobin, D. C., Holz, R., Sassen, K., Spangenberg, D., Minnis, P., Tooman, T. P., Ivey, M. D., Richardson, S. J., Bahrmann, C. P., Shupe, M., DeMott, P. J., Heymsfield, A. J., and Schofield, R.: The Mixed-Phase Arctic Cloud Experiment, *Bull. Am. Meteorol. Soc.*, 88, 205–222, <https://doi.org/10.1175/BAMS-88-2-205>, 2007.

170 Wendisch, M., Macke, A., Ehrlich, A., Lüpkes, C., Mech, M., Chechin, D., Dethloff, K., Velasco, C. B., Bozem, H., Brückner,
M., Clemen, H.-C., Crewell, S., Donth, T., Dupuy, R., Ebell, K., Egerer, U., Engelmann, R., Engler, C., Eppers, O., Gehrman,
M., Gong, X., Gottschalk, M., Gourbeyre, C., Griesche, H., Hartmann, J., Hartmann, M., Heinold, B., Herber, A., Herrmann,
H., Heygster, G., Hoor, P., Jafariserajehlou, S., Jäkel, E., Järvinen, E., Jourdan, O., Kästner, U., Kecorius, S., Knudsen, E. M.,
175 Köllner, F., Kretzschmar, J., Lelli, L., Leroy, D., Maturilli, M., Mei, L., Mertes, S., Mioche, G., Neuber, R., Nicolaus, M.,
Nomokonova, T., Notholt, J., Palm, M., van Pinxteren, M., Quaas, J., Richter, P., Ruiz-Donoso, E., Schäfer, M., Schmieder,
K., Schnaiter, M., Schneider, J., Schwarzenböck, A., Seifert, P., Shupe, M. D., Siebert, H., Spreen, G., Stapf, J., Stratmann,
F., Vogl, T., Welti, A., Wex, H., Wiedensohler, A., Zanatta, M., and Zeppenfeld, S.: The Arctic Cloud Puzzle: Using
ACLOUD/PASCAL Multiplatform Observations to Unravel the Role of Clouds and Aerosol Particles in Arctic Amplification,
Bull. Am. Meteorol. Soc., 100, 841–871, <https://doi.org/10.1175/BAMS-D-18-0072.1>, 2019.