



Supplement of

The shifting of secondary inorganic aerosol formation mechanisms during haze aggravation: the decisive role of aerosol liquid water

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Supplement

S1 Data acquisition and analysis methods

S1.1 Data acquisition

On-line ion-chromatograph instrument (MARGA ADI 2080, Metrohm Applikon, Switzerland) was widely used to measure the water-soluble inorganic ions (Na⁺, NH4⁺, Mg²⁺, Ca²⁺, K⁺, Cl⁻, F⁻, SO₄²⁻, NO₃⁻) in PM_{2.5} and corresponding trace gases (SO₂, HNO₂, HNO₃, HCl, NH₃). In detail, 10 ppm H₂O₂ filled wet rotary denuder (WRD) were used to absorb the gaseous pollutants (e.g., HF, HCl, HNO₃, HONO, SO₂, NH₃), while particulate species (e.g., F⁻, Cl⁻, NO₃⁻, SO₄²⁻, NH₄⁺, Na⁺, K⁺, Mg²⁺, Ca²⁺) were collected by steam jet aerosol collector. By synchronizing occupying atmospheric sampling speed of 16.7L/min and WRD rotation speed of \geq 8rpm, the absorption efficiency of whole instrument reached greater than 99.7%, which makes the samples are reliable. Then, both gaseous and particulate samples are simultaneously transferred as liquid (aqueous) samples and analyzed by ion chromatography. An internal calibration standard containing Li⁺ and Br⁻ with 1h-resolution was simultaneously injected with sample to account for any changes in the system. In addition, external standard solutions were also used to ensure perk identification and data quality. The detection limits of the particulate analytes were 0.001, 0.005, 0.004, 0.005, 0.005, 0.009, 0.006, 0.009 µg/m³ for Cl⁻, NO₃⁻, SO₄²⁻, NH4⁺, Na⁺, K⁺, Mg²⁺, Ca²⁺, respectively. Gaseous analytes detection limits were 0.001, 0.005, 0.002, 0.003, 0.005µg/m³ for HCl, HNO₃, HNO₂, SO₂, NH₃, respectively.

S1.2 Calculation of HNO₃-NO₃⁻ partitioning (ε (NO₃^T))

The S curve of $HNO_3-NO_3^-$ partitioning ($\epsilon(NO_3^-)$) has been adopted and discussed explicitly in many works based on the situ observations. To calculate the theoretical $HNO_3-NO_3^-$

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partitioning and estimate the NO_3^- loadings on particulate matter, theoretical calculation method of $HNO_3-NO_3^-$ partitioning (ϵ (NO_3^T)) (Guo et al., 2015; Guo et al., 2016; Guo et al., 2017) was employed in this work. Briefly, this calculation was based on equilibrium between HNO_3 and particle-phase NO_3^- (aqueous nitrate concentration). This equilibrium involves two reversible processes, dissolution of HNO_3 into aqueous phase and dissociation of HNO_3 into H^+ and NO_3^- . These two processes are always reach thermodynamic equilibriums at ambient conditions.

1st
$$HNO_{3(g)} \leftrightarrow HNO_{3(aq)}, H_{HNO_3}$$

2nd $HNO_{3(aq)} \leftrightarrow H^+ + NO_3^-, K_{n1}$

Thus, according to the definition of Henry's law constant and acid dissociation constant, we have

$$H_{HNO_{3}} = \gamma_{HNO_{3}}[HNO_{3}]/p_{HNO_{3}}$$
$$K_{n1} = \frac{\gamma_{NO_{3}}[NO_{3}^{-}]\gamma_{H^{+}}[H^{+}]}{\gamma_{HNO_{3}}[HNO_{3}]}$$

thus, based on calculation by Guo et al. (2016), ϵ (NO₃^T) could be written as :

$$\varepsilon(NO_3^T) = \frac{H_{HNO_3}^* W_i \times 0.987 \times 10^{-14}}{\gamma_{NO_3^-} \gamma_{H^+} 10^{-pH} + H_{HNO_3}^* W_i \times 0.987 \times 10^{-14}}$$

where 0.987 comes from the conversion from 1 atm to 1 bar, Wi is calculated as aerosol liquid water associated with inorganic species in $\mu g/m^3$ (here, aerosol liquid water associated with organics is not considered, due to the extremely low concentrations). R is ideal molar gas constant. T is temperature in K. $H^*_{HNO_3}$ is calculated by the method proposed by Clegg and Brimblecombe (1990).



S2 Results and Discussion

S2.1 In-situ high-resolution observations

The average concentrations of PM₁₀, PM_{2.5}, PM_{1.0} were 108.5 \pm 68.3, 84.75 \pm 84.5, 63.18 \pm 96.0 µg/m³, respectively, with a wide range of variability due to several haze events occurred during the studied period (Fig. S2). It was notable that SO₂ was maintained at a relatively low level (18.97 \pm 15.0 µg/m³) since the implementation of pollution controlling policy in recent years, compared with the neighboring city in 2014 (Zhou et al., 2018) and its own SO₂ concentration during 2014 (50 µg/m³). Whereas, concentrations of other reactive precursors were basically maintained at relative high level, such as NO₂ (44.56 \pm 18.6 µg/m³). NH₃ (5.55 \pm 4.07 µg/m³), compared with 2014 (Yearly average NO₂ concentration, 44 µg/m³). Meanwhile, as a typical tracer of combustion, significantly elevated CO emphasized contributions of incomplete combustion on haze intensification. In addition, O₃ was the only significantly declined pollutant with the haze aggravation, due to the weakened solar radiation led by the aggravated haze, which indicating significantly reduces the SIA formation processes through the oxidation between gaseous precursors, OH radicals and stabilized Criegee intermediate (sCI) by

considering the determining role of O_3 in the formation of OH radicals and sCI in the atmosphere (Mauldin Iii et al., 2012; Hua et al., 2008). As for meteorological conditions, low temperature and high humidity were observed during the studied period, with the values were -8.45±4.42 °C and 65.59±14.07 %, respectively.



 $PM_{2.5}$ concentrations presented increasing trends with pollution levels (Fig. S3), and its pH ranged from 0.63 to 6.32 (calculated by ISORROPIA-II), mainly concentrated at 4~5 (Fig. S4). The weak acidic nature of particles was consistent with Tianjin, China (Gao et al., 2020). With the worsening of the haze events, the average concentrations of gaseous precursors, such as SO₂, NO_x, NH₃, were 2.27, 1.33, 1.26 times of those in clean periods, respectively, indicating the increasingly enhanced importance of precursors on SIA generations. Additionally, it was

notable that along with haze aggravation, SIA consecutively sharing the most abundant species in PM_{2.5} and all detected ions, highlighting the determining role of SIA on regional haze aggravation and consistent with previous works (Fig. S3) (Xu et al., 2017; Gao et al., 2020; Huang et al., 2020). Meanwhile, the stabilized conditions, characterized by gradually lower wind speed (from 1.79 ± 1.2 to 1.15 ± 0.47 m/s) and higher RH (from 59.48 ± 15.0 to 71.13 ± 11.9 %) during haze events, further aggravating the regional haze by hindering the diffusion of particulates and accelerating its hygroscopic growth (Tab. S1) (Zhang et al., 2009; Zheng et al., 2015; Zhang et al., 2015).





S2.2 Chemical forms of ammonium species

Although anions could react with metal ions to form non-volatile compounds, the mass fraction of metal ions (Na⁺, K⁺, Mg²⁺, Ca²⁺) to total ions was less than 5.2% and could be neglected in this work. Thus, molar ratios of NH4⁺ vs. anions could be treated as indicators to identify the chemical forms of ammonium salts guided by stoichiometry (Zhou et al., 2018; Wang et al., 2021; Liu et al., 2017). Briefly, NH₃ prefer react with H₂SO₄ to form non-volatile NH₄HSO₄ or (NH₄)₂SO₄ under NH₄⁺-poor conditions, afterwards, semi-volatile NH₄NO₃ and NH₄Cl would subsequently generate by existing sufficient NH₃ through neutralize atmospheric HNO₃ and HCl. Roughly, average molar fractions of NH4⁺ accounting for nearly four times of sulfate molar concentrations throughout the studied periods (Fig. S3), indicating the atmospheric SO4²⁻ were completely neutralized by NH₄⁺ and formed ammonium sulfate and ammonia bisulfate according to the stoichiometry results between NH4⁺ and sulfate (Zhou et al., 2018; Wang et al., 2021; Liu et al., 2017). According to Fig. S5a, significant correlations between n (NH4⁺) and n (SO42-) with correlation coefficient of 0.947 and slope much lower than 0.5 further highlighting the atmospheric SO₄²⁻ completely neutralized by NH₄⁺ to yield ammonium sulfate and ammonia bisulfate. To further deduce the chemical species of the remaining ammonium, the concept of "excess NH_4^+ ", which calculated as $[NH_4^+]-1.5 \times [SO_4^{2-}]$ in molar concentration (Shi et al., 2019), was introduced in this work. The correlations between n (excess NH₄⁺) and $n(NO_3^{-})$ (Fig. S5b) suggested that atmospheric NH₄⁺ not only enough to neutralize both SO₄²⁻ and NO₃⁻ forming two major species of (NH₄)₂SO₄ and NH₄NO₃ during any pollution stage, but also remained on particles as free NH_4^+ according to the gas ratio (GR) >1 (Fig. S6). It indicated NH₄NO₃ formation was limited by the availability of HNO₃ (Paulot et al., 2016). Besides, starting with light pollution periods, the scatters of n (excess NH_4^+) and n (Cl⁺+NO₃⁻) gradually moves from above 1:1 line to below 1:1 line, indicating the coexistence of particle NH4NO3 and NH4Cl and increasingly ammonium ionized on particulate matter from this period (Fig.S5c). Meanwhile, Table S2 showed that the molar concentrations of Cl^{-} and K^{+} significantly elevated from Moderate pollution stage, especially Serious stage, indicating the enhanced emission of biomass burning and coal combustion during these periods. During haze aggravation, the remained-NH₄⁺ ([NH₄⁺]- $1.5 \times$ [SO₄²⁻]-[NO₃⁻]) was observed as 0.045, 0.0865, 0.141, 0.281 and 0.316 µmol/m³ in Clean, Light, Moderate, Heavy and Serious pollution stages, respectively. With the increasing remained-NH4⁺ during haze aggravation, the increasing molar concentrations of Cl⁻ had more opportunities to meet NH₄⁺ and form NH₄Cl, considering the higher molar concentrations of Cl^{-} than (K⁺+Na⁺). Additionally, the remaining ammonium, as well as other detected alkalic cations, may react with undetected CO_3^{2-} , HCO_3^{-} and organic materials (e.g., water-soluble organic acids, sugars, alcohols, etc.) through complexation reaction and form metal coordination compounds (Nozière et al., 2010; Chang-Graham et al., 2011; Cui et al., 2021).









PM2.5 during entire periods. The pentagrams were colored as a function of aerosol liquid water content, and the size of pentagrams corresponding to the PM2.5 mass concentrations, (c) Correlations between Required- NH₄⁺ and ALWC corresponding to PM_{2.5} and excess-NH₄⁺, required NH₄⁺ and excessNH₄⁺ were calculated according to the formula in S.M. text S2







Fig. S11 Correlations between (a) HONO and SOR, (b) NH4⁺/NH3 and HONO response to different pollution stages. To minimize the dependences mixed layer depth during diurnal changes and avoid the fast photolysis of HONO in the daytime, the nighttime hours were selected and corresponding to 20:00-7:00







(SO4²⁻) and log(H⁺) corresponding to PM_{2.5} and ALWC













	Whole Campaign	Clean	Light	Moderate	Heavy	Serious
Temperature (°C)	-8.45±4.42	-7.92±5.0	-8.42±4.4	-7.73±4.0	-10.08±3.2	-9.13±3.3
RH (%)	65.59±14.1	59.48±15.0	64.64±11.7	68.33±12.5	77.27±10.3	71.13±11.9
Wind Speed (m/s)	1.43 ± 0.87	$1.79{\pm}1.2$	1.25 ± 0.6	1.07 ± 0.5	1.01±0.5	1.15±0.47
Visibility (km)	37.50±20.4	52.36±18.1	38.12±14.1	30.77±20.3	15.01±10.0	$18.98{\pm}11.7$
Solar Radiation (w/m²)	82.09±135.9	81.31±136.8	83.73±137.7	80.35±144.1	75.7±125.84	73.32±124.7
О з (µg/m ³)	43.41±27.4	45.29±26.5	38.79±22.9	54.23±35.3	38.71±24.3	30.94±14.3
NO2 (μg/m ³)	44.56±18.6	35.72±19.8	45.18±16.4	$47.40{\pm}14.8$	51.99±17.6	57.10±21.1
NO _x (μg/m ³)	65.62±35.2	51.65±32.5	69.91±29.3	63.58±34.3	82.03±41.0	80.93±40.0
SO2 (μg/m ³)	18.97±15.0	14.59±10.2	21.72±11.1	21.52±11.8	15.01±11.1	33.08±43.4
<i>CO (mg/m³)</i>	$1.57{\pm}0.75$	1.06 ± 0.66	$1.58{\pm}0.6$	1.82 ± 0.6	2.18±0.6	2.40±0.8
NH3 (μg/m ³)	5.55±4.07	5.08 ± 3.84	5.24±3.96	6.31±4.43	6.37±3.37	6.42 ± 6.09
$PM_1 (\mu g/m^3)$	63.18±96.0	32.02±23.2	55.27±25.4	86.16±143.9	90.47±27.7	170.7±316.7
PM2.5 (µg/m ³)	84.75±84.5	39.11±31.1	71.21±35.9	112.38±144.6	150.10±46.0	192.14±162.9
PM10 (μg/m ³)	108.5±68.3	63.05±42.8	98.95±45.8	128.77 ± 58.8	177.55±53.2	200.97±114.6
PM1 /PM2.5	$0.84{\pm}0.39$	$0.87{\pm}0.08$	$0.92{\pm}0.06$	$0.94{\pm}0.02$	$0.95 {\pm} 0.02$	0.96±0.01
PM _{2.5} /PM ₁₀	0.71 ± 0.19	0.73±0.15	$0.84{\pm}0.11$	$0.89{\pm}0.06$	$0.93{\pm}0.05$	$0.93{\pm}0.04$
рН	$4.54{\pm}0.51$	4.54±0.59	4.52±0.35	4.63±0.34	4.64±0.22	4.16±1.26
ALWC (µg/m³)	57.49±85.2	13.65±18.8	35.22±32.4	74.41±84.67	154.8±114.9	156.49±153
Notes: all datasets were p	resented as mean± S.D.					

Table S1 Description of meteorological parameters and atmospheric pollutants

	Clean	Light	Moderate	Heavy	Serious
Na ⁺ (µmol/m ³)	$0.007 {\pm} 0.007$	0.009 ± 0.01	$0.009 {\pm} 0.01$	$0.005 {\pm} 0.01$	0.006 ± 0.01
K ⁺ (μmol/m ³)	$0.007 {\pm} 0.007$	0.012 ± 0.01	$0.058 {\pm} 0.23$	0.022 ± 0.01	0.209±0.84
$Mg^{2+}(\mu mol/m^3)$	$0.001 {\pm} 0.002$	0.003 ± 0.01	$0.014 {\pm} 0.04$	$0.005 {\pm} 0.01$	0.050 ± 0.12
$Ca^{2+}(\mu mol/m^3)$	0.006 ± 0.005	0.007 ± 0.01	0.006 ± 0.01	0.004 ± 0.01	0.006 ± 0.01
NH4 ⁺ (µmol/m ³)	0.235 ± 0.21	0.518 ± 0.31	$0.810 {\pm} 0.38$	1.399 ± 0.54	1.572±0.62
NO3 ⁻ (μmol/m ³)	0.094 ± 0.09	0.226 ± 0.12	0.321 ± 0.12	0.461 ± 0.18	0.548±0.32
SO4 ²⁻ (µmol/m ³)	0.064 ± 0.06	0.137 ± 0.10	0.232 ± 0.15	0.438 ± 0.20	0.472 ± 0.34
Ct ⁻ (µmol/m³)	0.051 ± 0.04	0.058 ± 0.04	0.105 ± 0.23	0.093 ± 0.04	0.234±0.74
Remained- NH4 ⁺ (µmol/m ³)	0.045 ± 0.05	$0.08650.087 \pm$	0.141 ± 0.13	0.280 ± 0.121	0.316±0.11

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