



Supplement of

Enrichment of calcium in sea spray aerosol: insights from bulk measurements and individual particle analysis during the R/V *Xuelong* cruise in the summertime in Ross Sea, Antarctica

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10 2. Supplementary Figures (Figure S1 to S20)

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12 **Note:** Figures S11-S20 were used when discussing supplementary Text (S3-S5) and were not cited
13 in the main text.

14 **Introduction:** This supplement contains additional information on the comparative analysis.

15 **Text S1. Uncertainty analysis and quality assurance of datasets**

16 **IGAC:** The water-soluble ion mass concentrations were analyzed by using an ion
 17 chromatography (IC) system (Dionex ICS-3000) within an in-situ Gas and Aerosol Compositions
 18 monitoring system (IGAC, Model S-611). Before the sampling, a known concentration of LiBr was
 19 used as an internal standard by adding to the aerosol liquid samples to determine the stability of the
 20 IGAC system. The changes in the internal standard were within 5%. Subsequently, the IC was
 21 calibrated by using six to eight concentrations of stand solutions (0.1-2000 $\mu\text{g L}^{-1}$), depending on
 22 the target species and concentrations, in which the R^2 was above 0.99. The calibration curves for
 23 each ion could be found in our previous studies (Yan et al., 2019; Yan et al., 2020a; Yan et al., 2020b).
 24 The uncertainty of the IC systems was generally less than 5% for all analyzed ionic species. The
 25 detection limits for Na^+ , Cl^- , Ca^{2+} , K^+ , and Mg^{2+} were 0.03, 0.03, 0.019, 0.011, and 0.042 $\mu\text{g L}^{-1}$
 26 (aqueous solution), respectively. During the whole sampling period (hourly temporal resolution),
 27 the detection rate for Na^+ , Cl^- , Ca^{2+} , K^+ , and Mg^{2+} were 98.5% (1178 of 1196), 92.6% (1108 of
 28 1196), 88.2% (1055 of 1196), 98.9% (1183 of 1196), and 98.5% (1178 of 1196), respectively. All
 29 values below the detection limit were omitted before analysis. The undetected rate for both Ca^{2+} and
 30 Na^+ was approximately 12%. **Figure S11** shows the time series of observed ion mass concentrations
 31 and EF_{Ca} . The ion mass concentrations were above the detection limit. Particularly, the mean Na^+
 32 and Ca^{2+} mass concentrations were 364.64 ng m^{-3} (ranging from 6.66 to 4580.10 ng m^{-3}) and 21.20
 33 ng m^{-3} (ranging from 0.27 to 334.40 ng m^{-3}), respectively, which were far above (> 10 times) the
 34 detection limit. Such data indicate that the variations of EF_{Ca} would not suffer the increasing
 35 uncertainties when the ion mass concentrations near the detection limit. Thus, we suggest that the
 36 data on ion mass concentration is reliable and representative.

37 Given that the measured mass concentrations of Ca^{2+} and Na^+ were far above the detection
 38 limit, we only considered the systematic errors ($< 5\%$) of the mass concentrations of Ca^{2+} and Na^+ .
 39 Based on the variations of extreme value, the uncertainty of each EF_{Ca} is calculated as below:

$$40 \quad \text{abs}\left(1 - \frac{0.95 * Ca_{conc}^{2+} / 0.038}{1.05 * Na_{conc}^+}\right) \approx 9.5\% < \text{Unc}_{(\text{EF}_{\text{Ca}})} < \text{abs}\left(1 - \frac{1.05 * Ca_{conc}^{2+} / 0.038}{0.95 * Na_{conc}^+}\right) \approx 10.5\%$$

$$41 \quad \frac{Ca_{conc}^{2+}}{Na_{conc}^+} / 0.038$$

$$42 \quad \frac{Ca_{conc}^{2+}}{Na_{conc}^+} / 0.038$$

41 Where Ca_{conc}^{2+} and Na_{conc}^+ represent the observed ion mass concentration of Ca^{2+} and Na^+ ,
 42 respectively. 0.038 is the ratio of Ca^{2+} to Na^+ in seawater. 0.95 and 1.05 represent the variations of

43 extreme value. The largest uncertainty of EF_{Ca} ($Unc_{(EF_{Ca})}$) would be estimated lower than 11%.

44 **SPAMS:** Even though SPAMS has advantages for investigating the chemical characteristics
45 and evolutionary mechanisms of individual particles, it is still quite challenging for it to provide
46 quantitative evidence, as it is limited by the potential matrix effects, laser inhomogeneities, and
47 transmission efficiencies of the aerodynamic lenses (Qin et al., 2006; Pratt and Prather, 2012).
48 Nevertheless, the results via SPAMS are also reliable because of its broad datasets with high
49 temporal resolution. Therefore, the analyses of particle count, size, and chemical composition (by
50 peak area) can be considered semi-quantitative from a statistical perspective (Healy et al., 2012; Su
51 et al., 2021a; Zhang et al., 2021).

52 **Text S2. Single-particle characteristics of other observed particles**

53 In this study, a total of $\sim 930,000$ particles with mass spectral information and D_{va} were
54 measured using SPAMS and manually clustered into the seven single-particle groups of SS (sea salt
55 aerosol), SS-aged (aged sea salt aerosol), SS-Bio (sea salt with biogenic organic matter), OC-Ca
56 (internally mixed organics with calcium), OC-K (internally mixed organics with potassium), OC
57 (organic-carbon-dominated), and EC (element carbon) (**Table S2**). Based on their different mass
58 spectral characteristics (Prather et al., 2013; Collins et al., 2014; Guasco et al., 2014; Su et al.,
59 2021b), we briefly describe all of them (except for OC-Ca) as follows (**Figure S5**):

60 SS is identified by prominent ion markers that are associated with Na and Cl (e.g., m/z 23 $[Na]^+$,
61 46 $[Na_2]^+$, 81 $[Na_2^{35}Cl]^+$, 83 $[Na_2^{37}Cl]^+$, -35 $[^{35}Cl]^-$, and -37 $[^{37}Cl]^-$), and smaller contributions of
62 other inorganic matter that are known to exist in seawater (i.e., Mg, K, Ca). SS accounted for 16.5%
63 of the particles obtained during the observation cruise, with a peak of $\sim 1. \mu m$ in D_{va} .

64 SS-aged can be regarded as SS with atmospheric aging, with additional characteristic peaks of
65 nitrate (m/z -46 $[NO_2]^-$ and -62 $[NO_3]^-$). SS-aged presented a similar size distribution as SS, with a
66 proportion of 8.1%.

67 SS-Bio is characterized by a large ratio of organic ion signatures of organic nitrogen (m/z -26
68 $[CN]^-$ and -42 $[CNO]^-$), phosphate (m/z -63 $[PO_2]^-$ and -79 $[PO_3]^-$), carbohydrate (m/z -45 $[CHO_2]^-$,
69 -59 $[C_2H_3O_2]^-$, and -73 $[C_3H_5O_2]^-$), siliceous materials (m/z -60 $[SiO_2]^-$), and organic carbon (m/z 27
70 $[C_2H_3]^+$ and 43 $[C_2H_3O_3]^+$), in addition to the aforementioned inorganic salt-related ion signature

71 (e.g., 23 [Na]⁺). SS-Bio accounted for only 3.1% of the particles obtained during the observation
72 cruise, with a peak of ~ 0.9 μm in D_{va} .

73 OC-K is identified by the dominant presence of K (m/z 39) and the aforementioned organic
74 species and exhibited the third largest proportion (13.7%). OC-K was scattered in the D_{va} range of
75 0.2 to 1.0 μm, peaking at 0.6 μm. Similar to OC-Ca, this chemical class may also be associated with
76 biogenic origin, which may originate from intact heterotrophic bacterial cells, fragments of cells,
77 and bacterial exudates (Gaston et al., 2011; Guasco et al., 2014; Sierau et al., 2014).

78 OC exhibits a significant proportion of organic signals of organic carbon and organic nitrogen.
79 This single-particle type may originate from not only heterogeneous nucleation with the oxidation
80 of monoterpene and isoprene but also anthropogenic emissions (e.g., ship emissions). However, we
81 suggest that it is more likely to be related to the biogenic origin of the bacteria and phytoplankton.

82 **Text S3. A comparative analysis of the chemical composition, particle size, counts, and mixing**
83 **state via SPAMS between leg I and leg II**

84 Based on the SPAMS datasets, we found that there was a minimal change in chemical
85 composition (**Figs. S12 and S13**), particle size (**Fig. S14**) and counts (**Fig. S15**), and mixing state
86 (**Figs. S16 and S17**) of the obtained particles (especially SS-Bio, OC-Ca, and OC-K) during leg I
87 (sea ice period) and leg II (period without Sea ice). This suggests that the existence of sea ice may
88 play an insignificant role in the intrinsic physicochemical properties of the obtained particles
89 ranging from 0.2 to 2 μm.

90

91 **Text S4. Potassium (K⁺) and magnesium (Mg²⁺) enrichment in sea spray aerosols induced by**
92 **temperature, wind speed, and sea ice**

93 The enrichment factors of K⁺ (EF_K) and Mg²⁺ (EF_{Mg}) were calculated by using the same
94 equation described in the main text. Generally, the ratios of K⁺ and Mg²⁺ to Na⁺ in seawater are
95 0.0218 and 0.1112, respectively (Hara et al., 2012; Boreddy and Kawamura, 2015; Su et al., 2022).

96 During the observation cruise, EF_K (ranging from 0.20 to 75, with a median and mean of 1.19
97 and 3.61, respectively, n > 1000) and EF_{Mg} (ranging from 0.12 to 26.94, with a median and mean of
98 1.48 and 2.46, respectively, n > 1000) were also observed.

99 **Figure S18** shows the enrichment factors of K⁺ and Mg²⁺ with respect to sodium in bulk
100 aerosols at different ambient temperatures ($\geq -3.5^{\circ}\text{C}$ and $< -3.5^{\circ}\text{C}$), wind speeds ($\geq 7 \text{ m s}^{-1}$ and < 7
101 m s^{-1}), and sea ice fraction (with and without sea ice) during the whole sampling period. The results
102 were very similar to that of calcium (Ca²⁺) enrichment in sea spray aerosols (SSAs), which may
103 present an analogous enrichment mechanism by environmental factors (e.g., ambient temperature
104 and wind speed with sea ice) (**Fig. S19**).

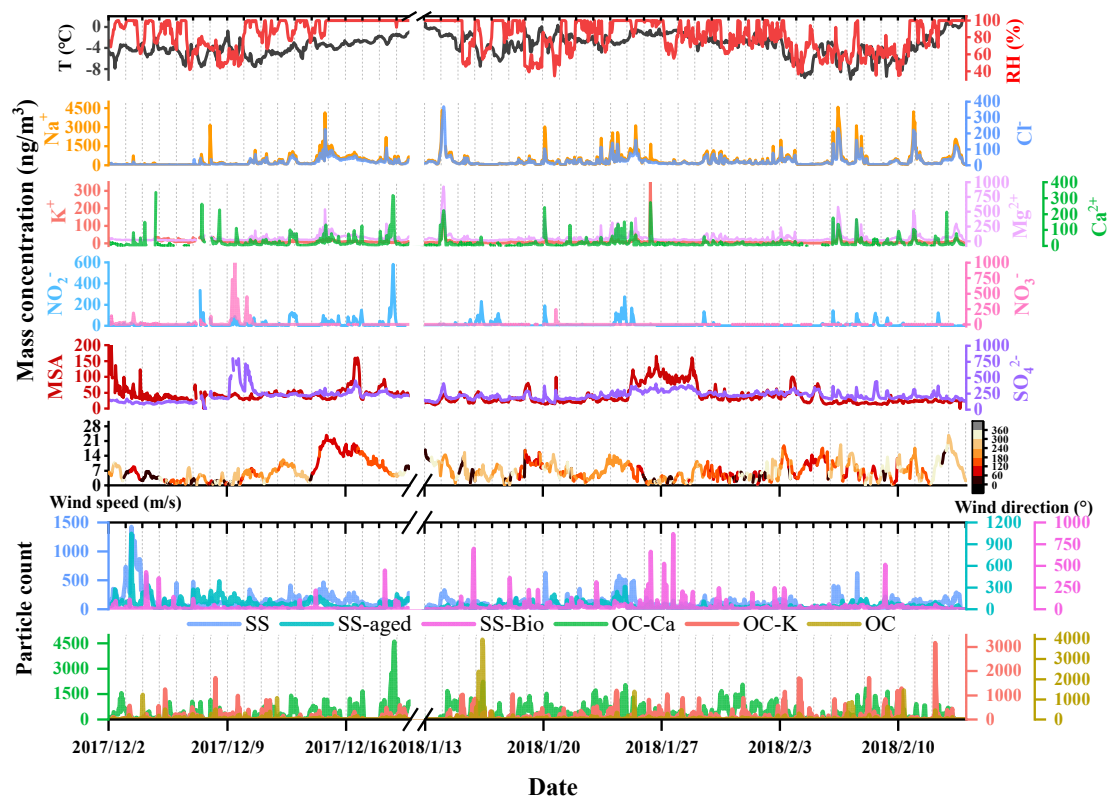
105 We also inferred that the K⁺ and Mg²⁺ enrichments in SSAs were also attributed to organically
106 internally mixed aerosols (e.g., OC-K and SS-Bio). We propose three possible explanations for this:
107 (i) If the released K⁺ and Mg²⁺ are in the form of inorganic salts (e.g., KCl and MgCl₂) within SSAs,
108 they are likely to be associated with sea salt fractionation by precipitations of sylvite ca. -33°C and
109 $10 \text{ H}_2\text{O}\cdot\text{MgCl}_2$ ca. -36°C on the sea ice (Hara et al., 2012). However, the average and minimum
110 ambient temperatures during leg I were only -4.1°C and $\sim -8^{\circ}\text{C}$, respectively. Thus, sea salt
111 fractionation on sea ice is less likely. (ii) We observed a single-particle type of OC-K with abundant
112 K⁺ (m/z 39) and organic ion signatures by using SPAMS during the cruise observations. Other
113 single-particle types involving K⁺ (m/z 39) and Mg²⁺ (m/z 24) also exhibited some characteristic
114 organic peaks (e.g., in SS-Bio). (iii) K⁺ and Mg²⁺ are also greatly capable of stabilizing organic
115 supramolecular structures in the form of organic ligands (i.e., marine microgels) (Gaston et al., 2011;
116 Cochran et al., 2016; Mukherjee et al., 2020). Based on the above discussion, we, therefore, suggest
117 that internally mixed potassium and magnesium organics contribute to corresponding enrichment in
118 SSAs. Analogous to OC-Ca, we also suggest that the marine microgels assembled by K⁺ and Mg²⁺
119 to exopolymer substances (EPSs) may be emitted to the atmosphere by low wind-blown sea ice.

120 **Text S5. Multiple linear and random forest regression**

121 Multiple linear and random forest analyses were applied to describe the relative contribution
122 of possible factors to the variations in EF_{Ca} . In multiple linear models, two of the most common
123 measures of model fit are the residual standard error and proportion of variance explained (R^2), by
124 using least squares fit. In contrast, random forest with nonlinear multiple regression has been widely
125 applied to predict and reproduce the importance of factors, by building multiple decision trees
126 (Lundberg et al., 2020; Zhang et al., 2021; Song et al., 2022).

127 We attempted to estimate the relative importance of the impact factors on EF_{Ca} , including the
128 environmental factors (ambient temperature, wind speed, and chlorophyll-a concentration), particle
129 types (SS-Ca and OC-Ca, by count), and relative fraction of organics in the OC-Ca particles.
130 Multiple linear regression and random forest analysis were applied. As shown in **Fig. S20**, EF_{Ca}
131 could not be well predicted by those possible factors for the sea ice period ($r^2 = 0.19, p < 0.01$) and
132 whole sampling period ($r^2 = 0.15, p < 0.01$), which may be ascribed to other unknown mechanisms
133 and/or OC-Ca with low water solubility. Future studies are needed to explore the mechanisms of
134 calcium enrichment in SSAs.

135 2. Supplementary Figures

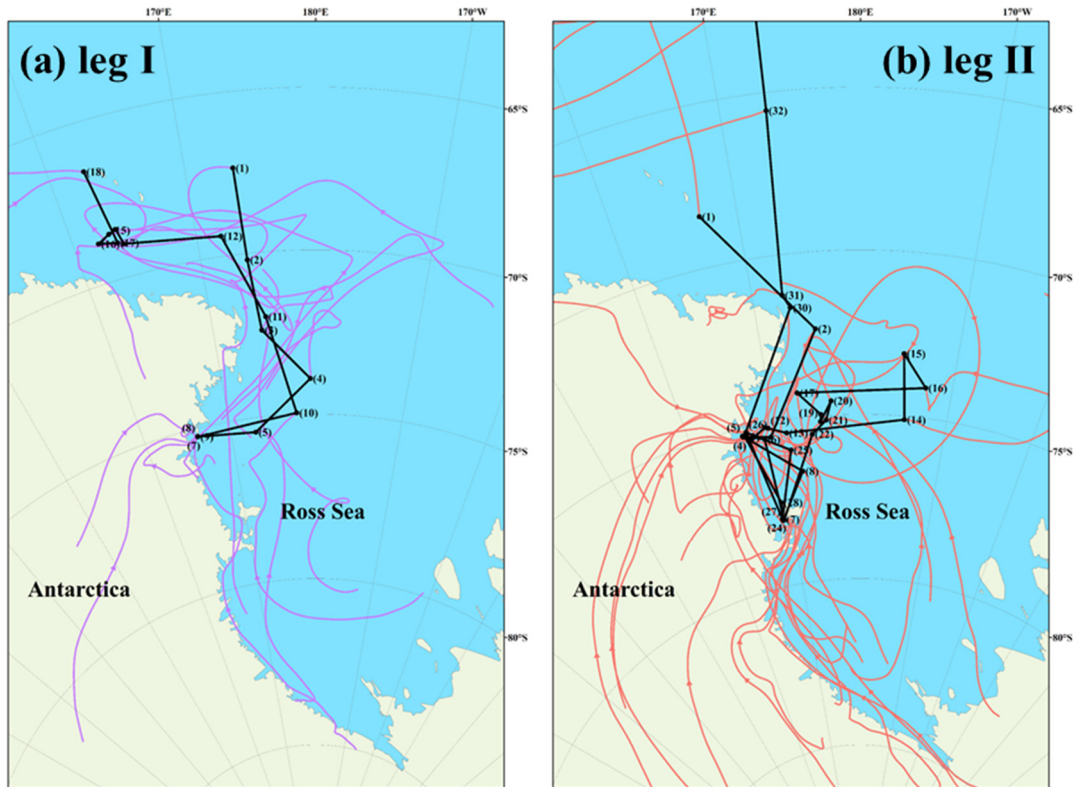


136

137 **Figure S1**

138 Time series (hourly temporal resolution) of meteorological parameters, typical water-soluble ion

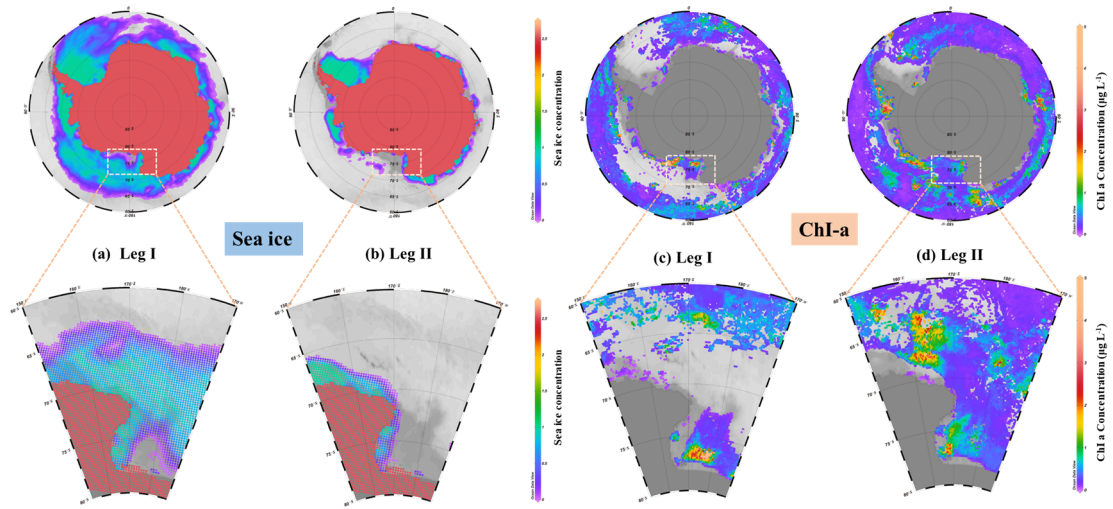
139 mass concentrations, and single-particle counts of individual particle types.



140

141 **Figure S2**

142 HYSPLIT back trajectories (72-hour) arrived at the ship location (50 m above sea level) every day
 143 (0:00 local time) during leg I (purple solid line) and leg II (red solid line). The black solid lines
 144 indicated the cruise track of the R/V *Xuelong*. The marked numbers indicated the ship location at
 145 the local time of 0:00 during the whole observation. For example, “(1)” in the leg I referred to the
 146 ship location at 0:00, on 2 December 2017, and so on.



147

148 **Figure S3**

149 Average sea ice fraction and chlorophyll-a concentration during leg I (a and c) and leg II (b and d)

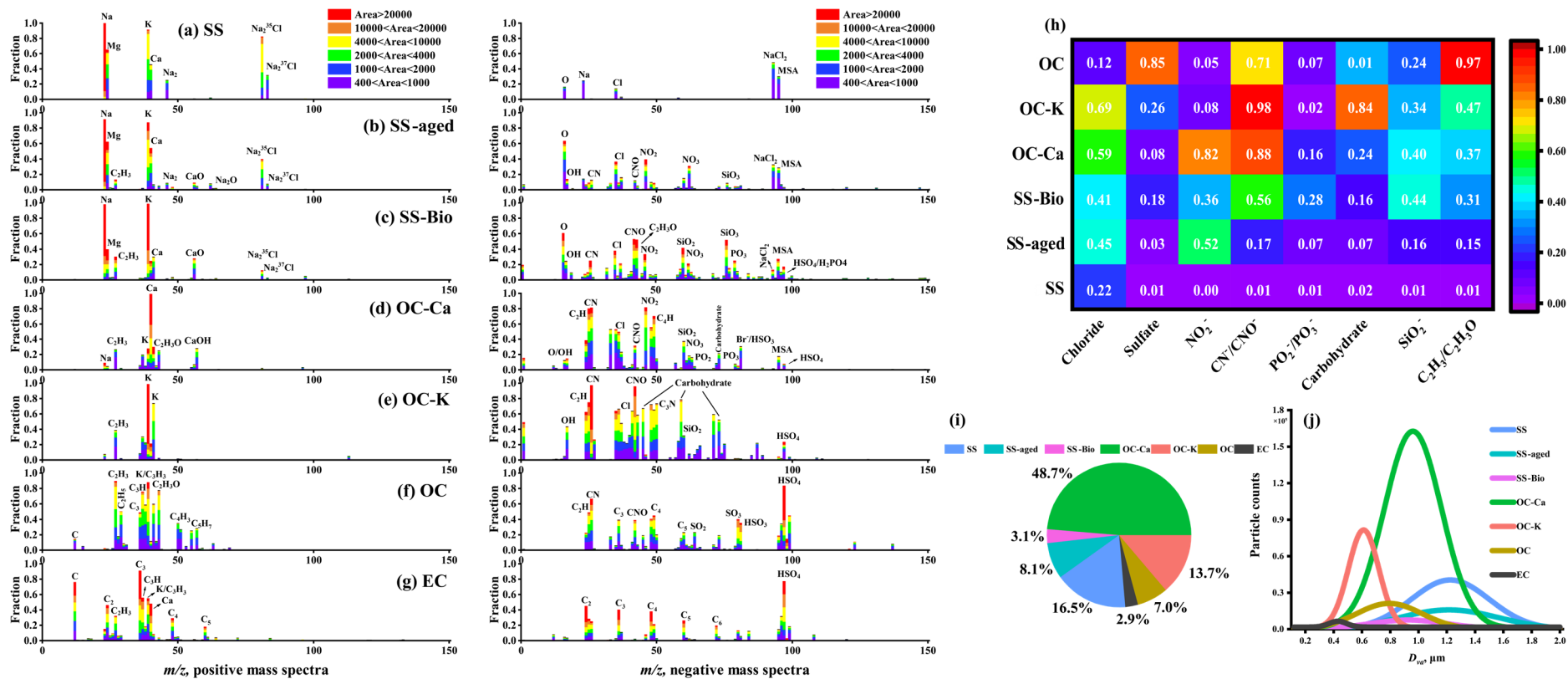
150 from a satellite. This figure was created by using Ocean Data View (Schlitzer, 2002).



151

152 **Figure S4**

153 An in-situ Gas and Aerosol Composition monitoring system (IGAC) was used to determine the
154 gaseous and aerosol water-soluble ion mass concentrations (red frame, left). A Single Particle
155 Aerosol Mass Spectrometer (SPAMS) was used to measure the chemical compositions, mixing
156 states, and particle size of individual particles (green frame, right).



157

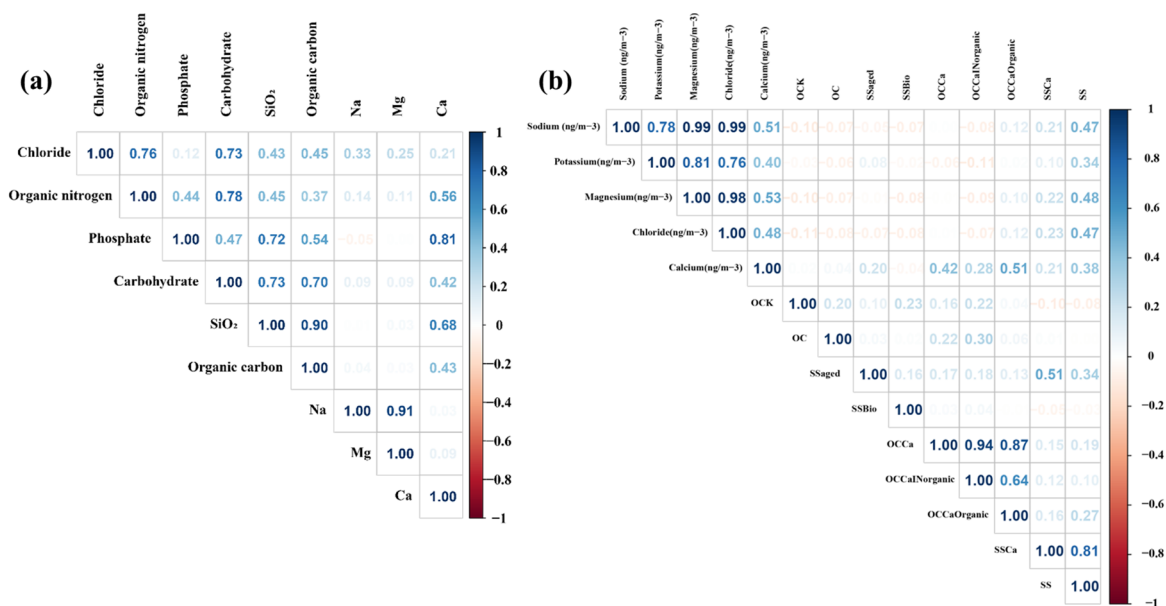
158 **Figure S5**

159 (a) – (g) Average digitalized single-particle mass spectra for seven classes of collected particles. (h) Hot plot of number fractions for major species of obtained single-

160 particle types, including chloride (m/z -35 and -37), sulfate (m/z -97), nitro-containing organic species (m/z -46), organic nitrogen (m/z -26 or -42), phosphate (m/z -63

161 or -79), carbohydrate (m/z -45, -59, or -73), siliceous materials (m/z -60), and organic carbon (m/z 27 or 43). (i) Relative proportion of different single-particle types
162 during the cruise observations. (j) Unscaled size-resolved number distributions of all individual particles.

163 A total of $\sim 580,000$ calcium-containing particles (m/z 40 $[\text{Ca}]^+$) were observed during cruise observations, accounting for $\sim 62\%$ of the total obtained particles.
164 These calcium-containing particles were scattered among all the obtained particle types, with proportions of $\sim 48\%$, $\sim 56\%$, $\sim 25\%$, $\sim 22\%$, $\sim 100\%$, $\sim 12\%$, and $\sim 49\%$ for
165 SS, SS-aged, SS-Bio, OC-K, OC-Ca, OC, and EC, respectively. In particular, the SS-Ca and OC-Ca particle types accounted for $\sim 12\%$ and $\sim 72\%$ of the total calcium-
166 containing particles and $\sim 7\%$ and $\sim 50\%$ of the total obtained particles, respectively.

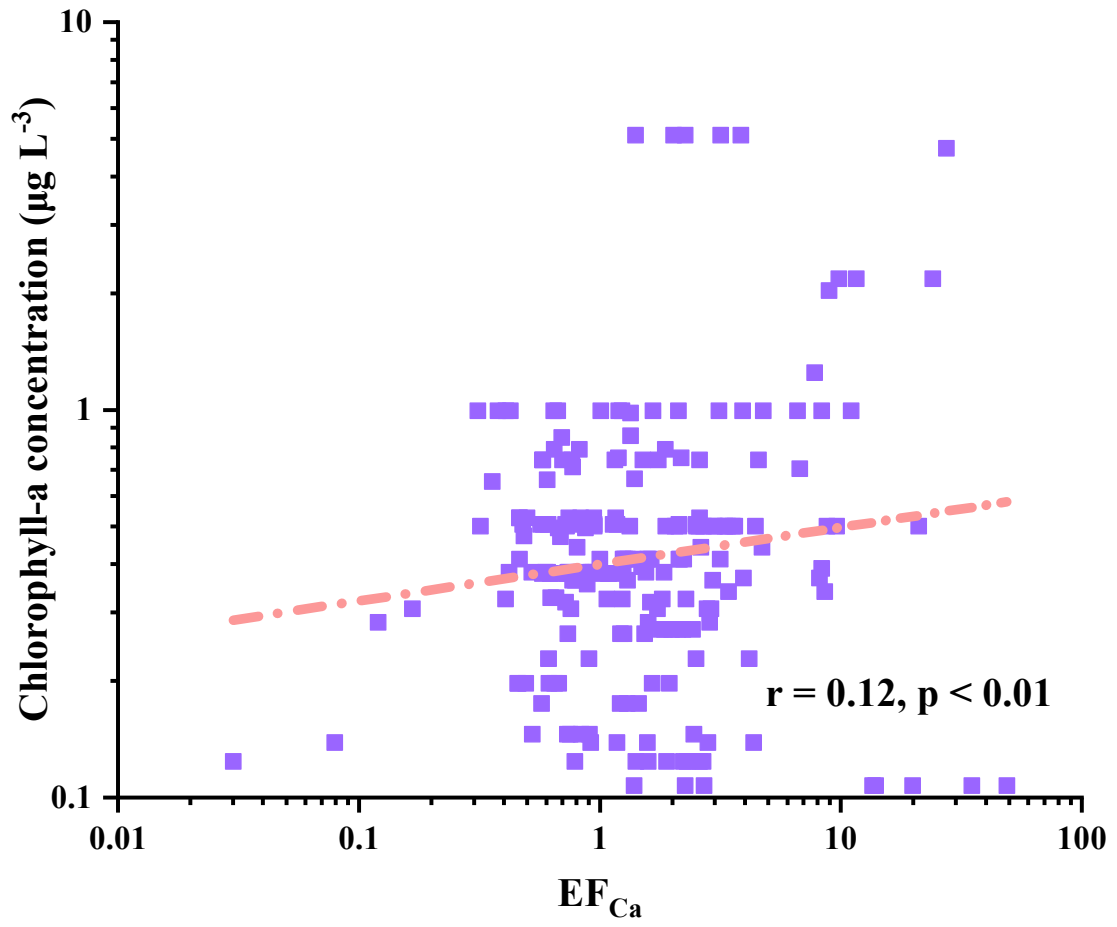


167

168 **Figure S6**

169 (a) Correlation analysis between the single-particle peak area for species in OC-Ca. There were
 170 relatively high correlation coefficients ($r = 0.42$ – 0.81) between the peak area of Ca and organic
 171 species (organic nitrogen, phosphate, carbohydrate, siliceous materials, and organic carbon). (b)
 172 Correlation analysis between the OC-Ca (by count) and mass concentration of Ca²⁺. The first
 173 (sodium) to the fifth (chloride) referred to mass concentration. The sixth (OCK) to the fourteenth
 174 (SS) referred to single-particle types.

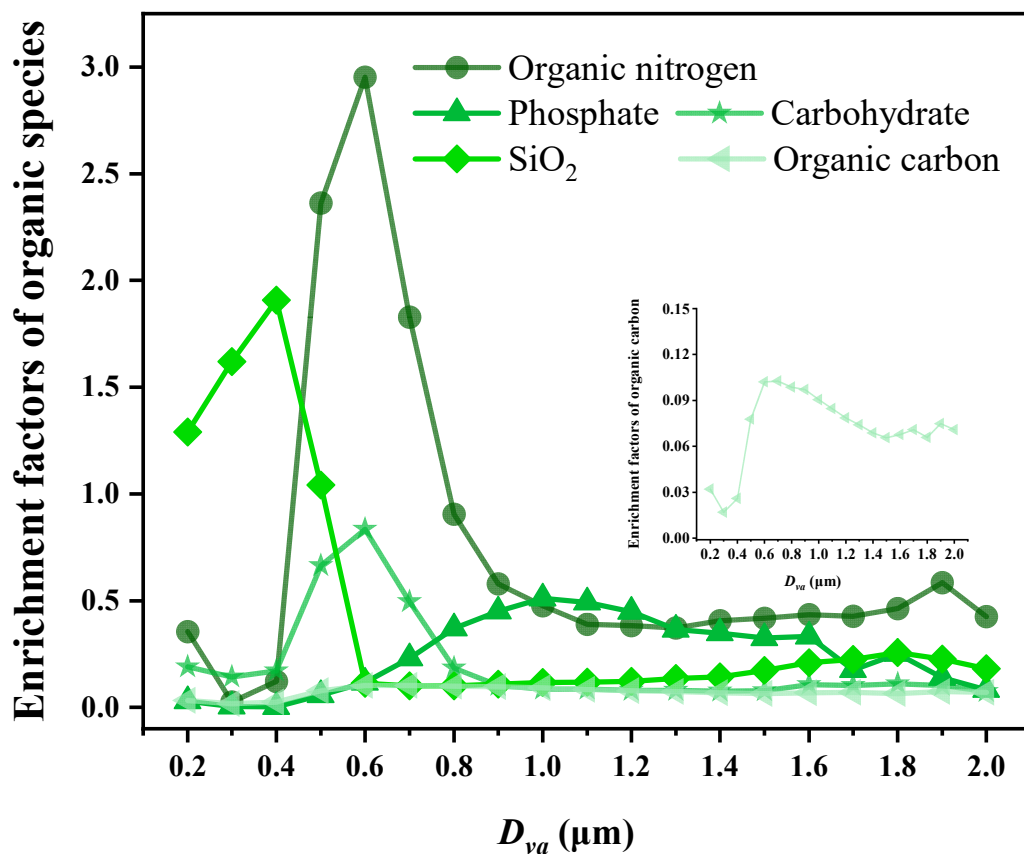
175



176

177 **Figure S7**

178 Correlation analysis between EF_{Ca} and chlorophyll-a concentration during leg II.

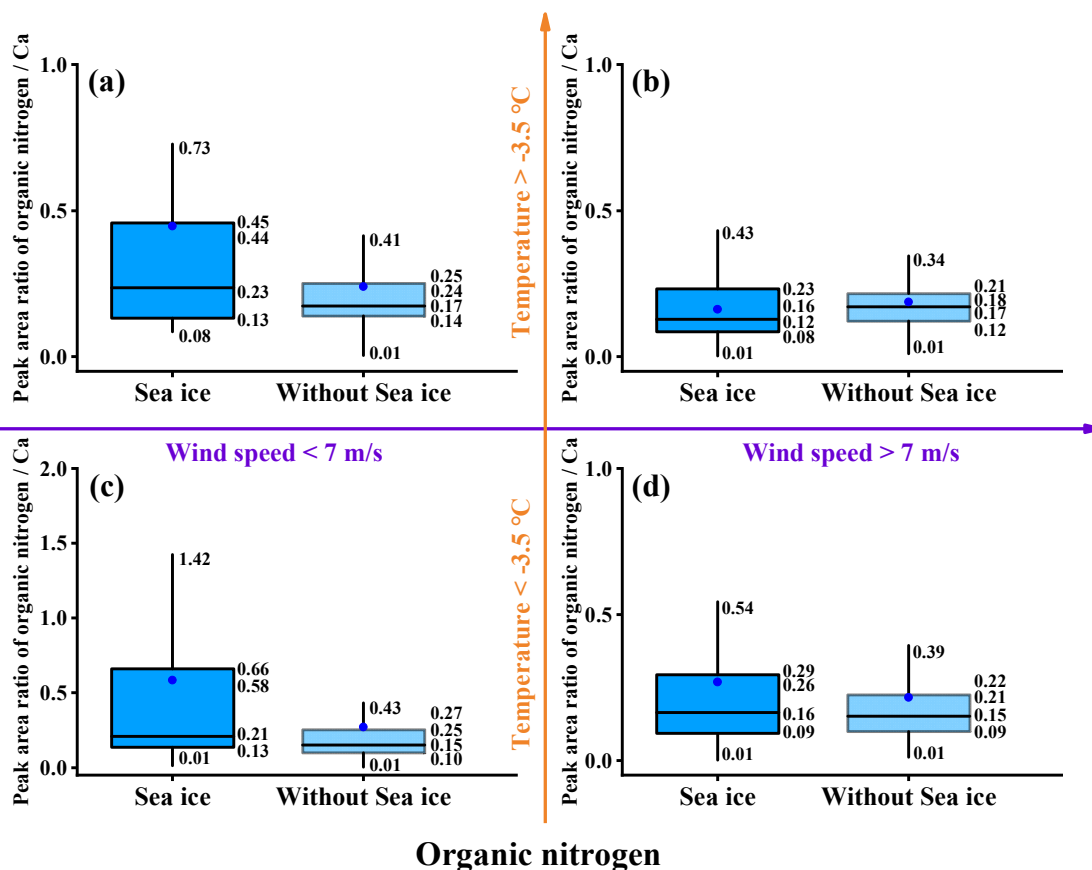


179

180 **Figure S8**

181 Size-dependent single-particle peak area ratio of organic matter to calcium. Based on SPAMS, we
 182 defined the organic matter enrichment factors as the single-particle peak area ratio of organic species
 183 to Ca (m/z 40) in OC-Ca. The single-particle counts of OC-Ca were 420, 000. The organic species
 184 included organic nitrogen (m/z -26 and -42), phosphate (m/z -63 and -79), carbohydrate (m/z -45, -
 185 59, and -73), siliceous materials (SiO_2 , m/z -60), and organic carbon (m/z 27 and 43).

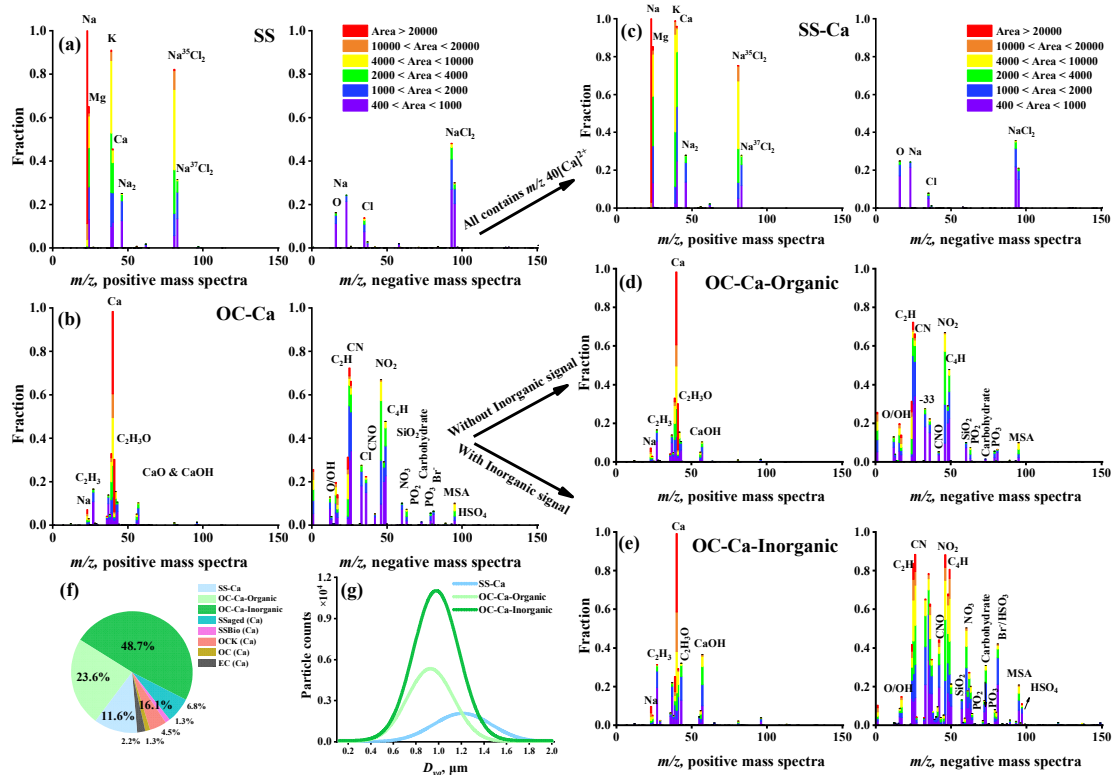
186



187

188 **Figure S9**

189 A box and whisker plot of the peak area ratio of organic nitrogen (m/z -26 and -42) to Ca (m/z 40)
 190 in OC-Ca at different ambient temperatures, wind speeds, and sea ice fraction. As described above,
 191 the organic nitrogen enrichment factor in OC-Ca (EF_{ON}) is defined as the single-particle peak area
 192 ratio of organic nitrogen to calcium. We choose the organic nitrogen within OC-Ca for comparative
 193 analysis because of its large number fraction (0.88). In the box and whisker plot, the marked values
 194 from top to bottom were 90th and 75th percentiles, mean, median, and 25th and 10th percentiles,
 195 respectively.

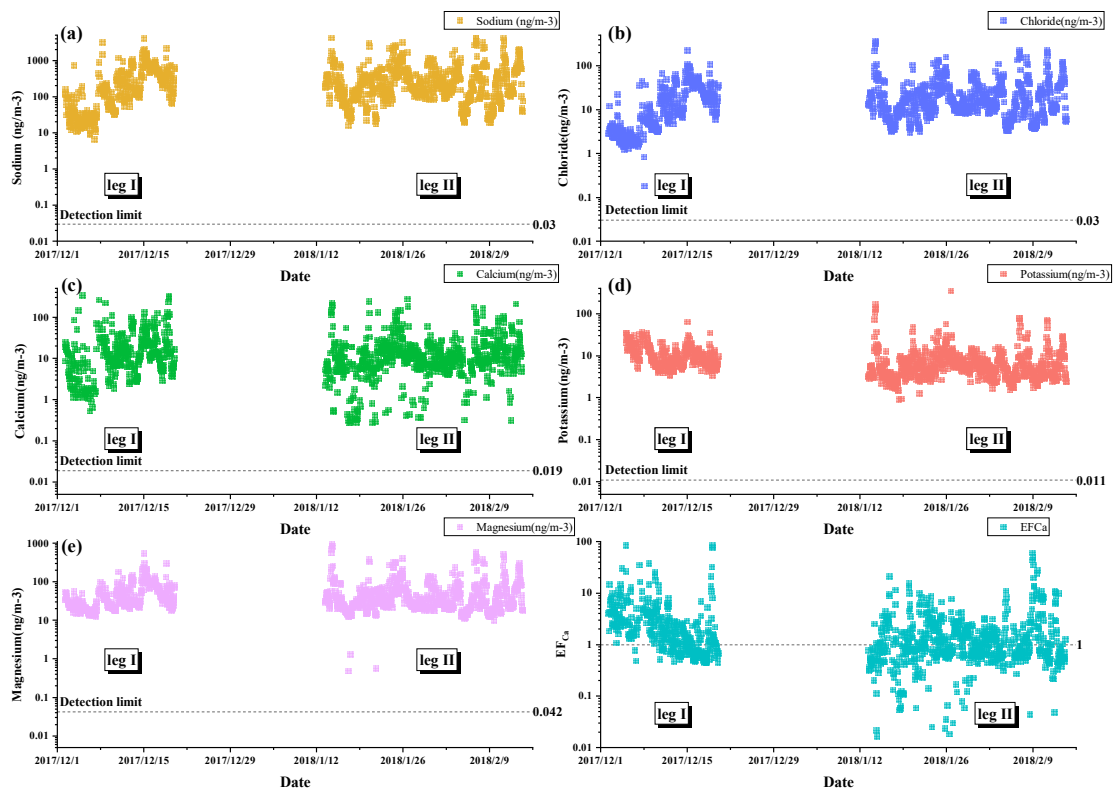


196

197 **Figure S10**

198 (a) and (b) Average digitalized single-particle mass spectra of chemical classes of SS and OC-Ca.
 199 (c) Average digitalized single-particle mass spectra of SS-Ca that are refined by using m/z 40 [Ca]²⁺
 200 upon SS. (d) and (e) Average digitalized single-particle mass spectra of OC-Ca-Organic and OC-
 201 Ca-Inorganic, which are classified by whether inorganic compounds (chloride (m/z -35 and -37),
 202 nitrate (m/z -62), and sulfate (m/z -97) ion signals are present. (f) Relative proportion and (g)
 203 unscaled size-resolved number distributions of single-particle types.

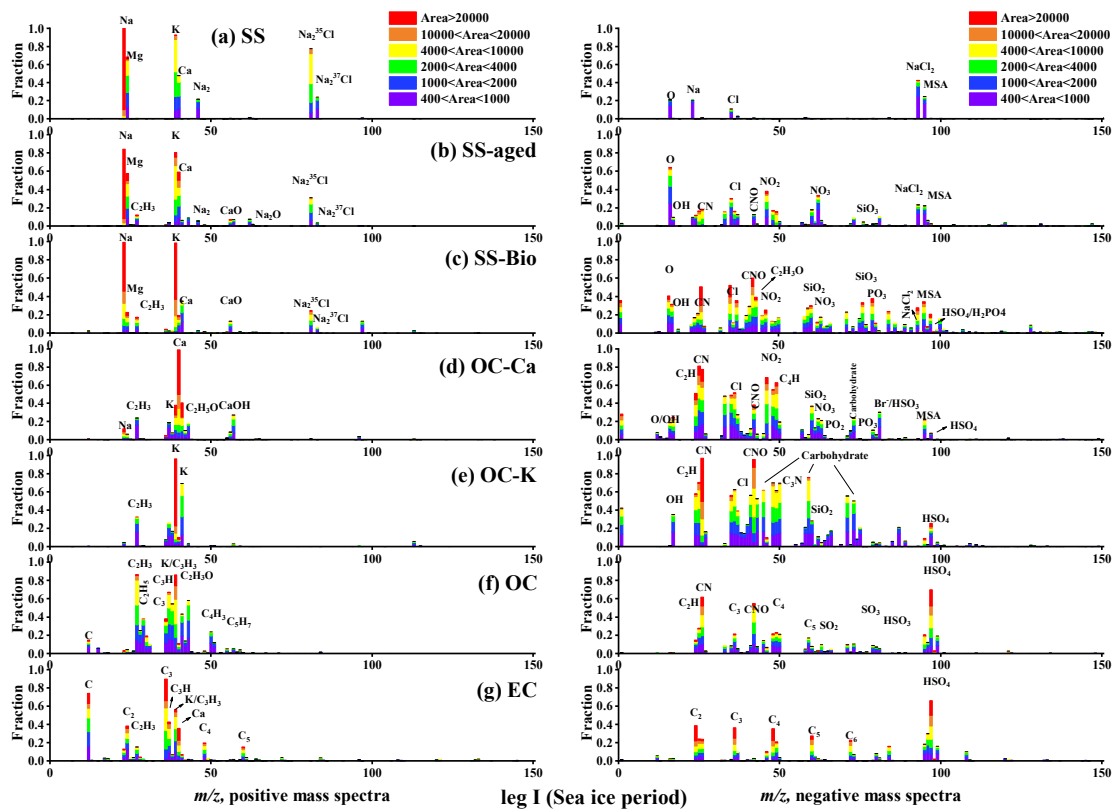
204 **Note:** Figures S11-S20 were used when discussing supplementary Text and were not cited in the
205 main text.



206

207 **Figure S11**

208 The time series of observed ion mass concentrations and EF_{Ca}.

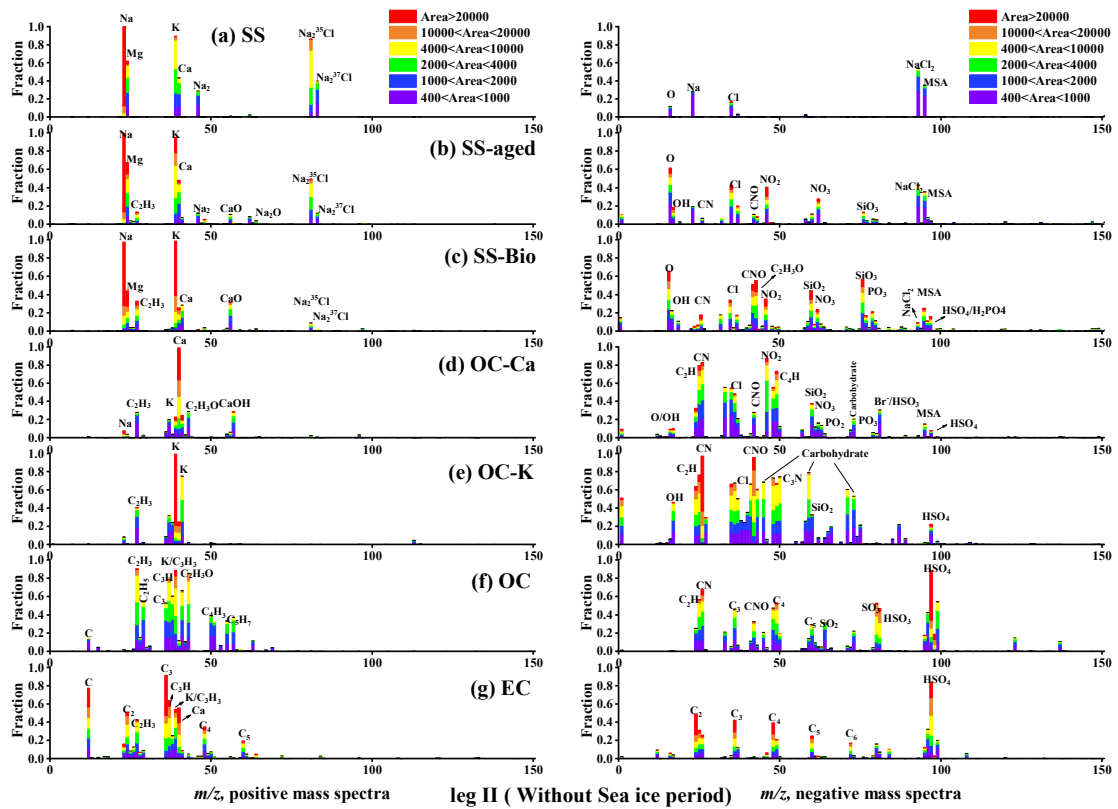


209

210 **Figure S12**

211 Average digitalized single-particle mass spectra for seven classes of collected particles during leg I

212 (sea ice period).

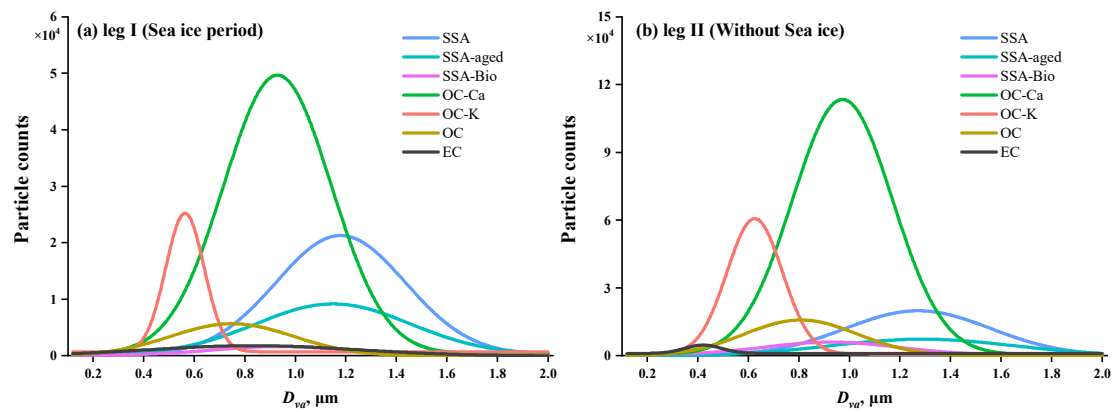


213

214 **Figure S13**

215 Average digitalized single-particle mass spectra for seven classes of collected particles during leg

216 II (period without sea ice).

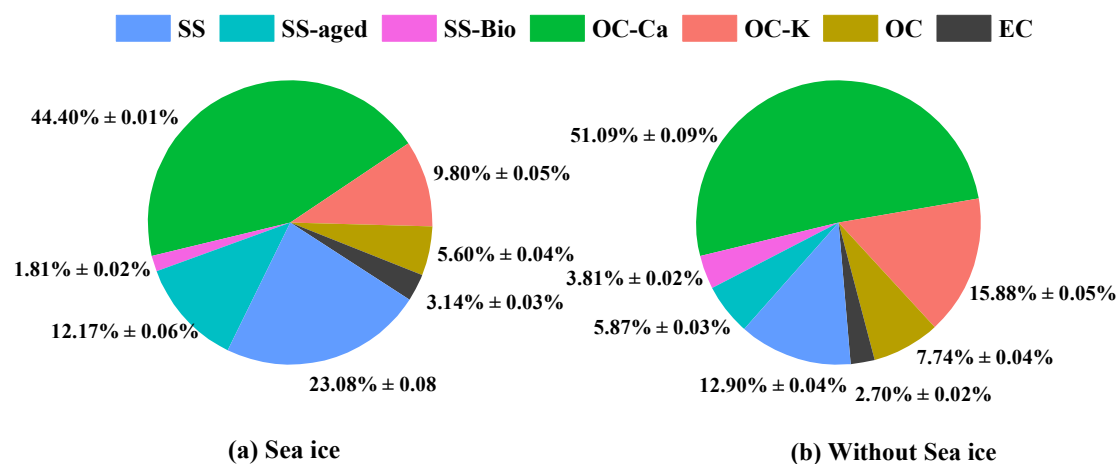


217

218 **Figure S14**

219 Unscaled size-resolved number distributions of all individual particles during leg I (sea ice period)

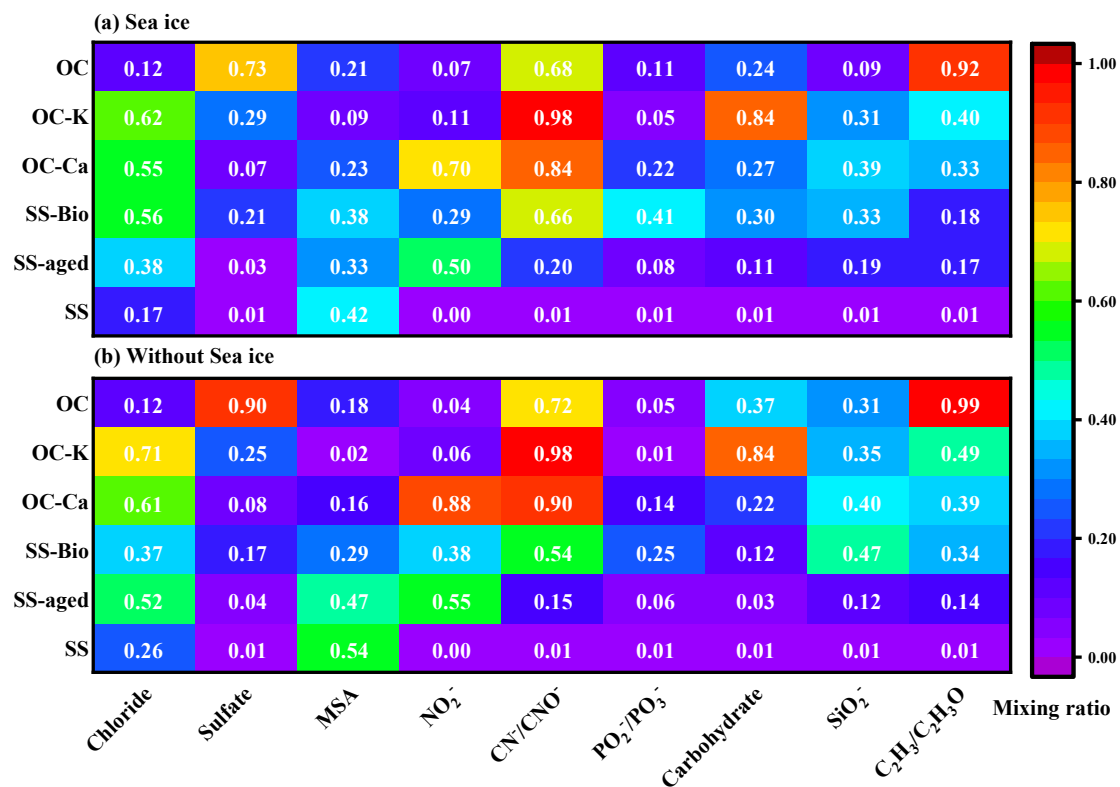
220 and leg II (period without sea ice).



221

222 **Figure S15**

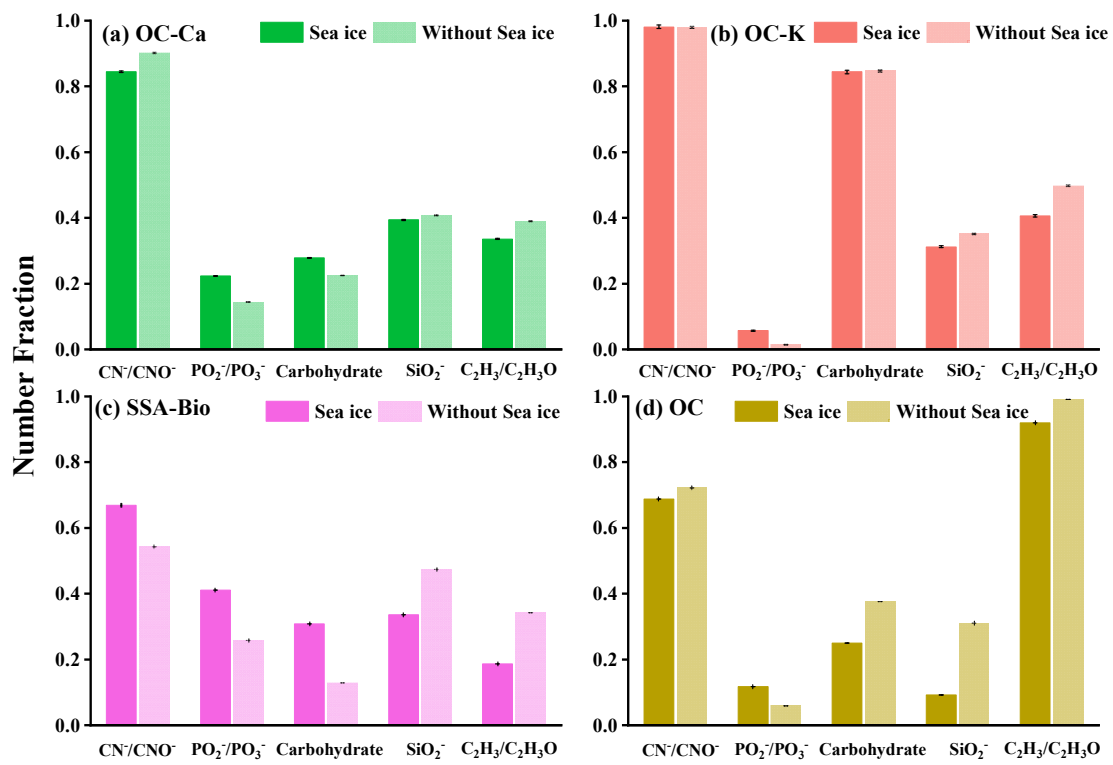
223 The relative proportion of different single-particle types during leg I (sea ice period) and leg II
 224 (period without sea ice). It is still quite a challenge to obtain quantitative measurements using
 225 SPAMS due to the potential inhomogeneities in the transmission efficiencies of the aerodynamic
 226 lenses and desorption/ionization, and the matrix effects of individual particles (Gross et al., 2000;
 227 Qin et al., 2006; Pratt and Prather, 2012). Therefore, it may not be straightforward to use the particle
 228 count in comparison with the absolute mass concentration. We noted that there was little difference
 229 in OC-Ca proportion during the periods of sea ice and without sea ice. The source of OC-Ca for the
 230 two periods may be explained by the low wind-blown sea ice and the land-based ice from Antarctica
 231 (**Fig. S2** and **Table S1**, the influence of air masses from Antarctic land, 40%), respectively. Another
 232 reason for that may be the resuspension of OC-Ca. Also, the bubble bursts within open water and
 233 leads occurred in both periods.



234

235 **Figure S16**

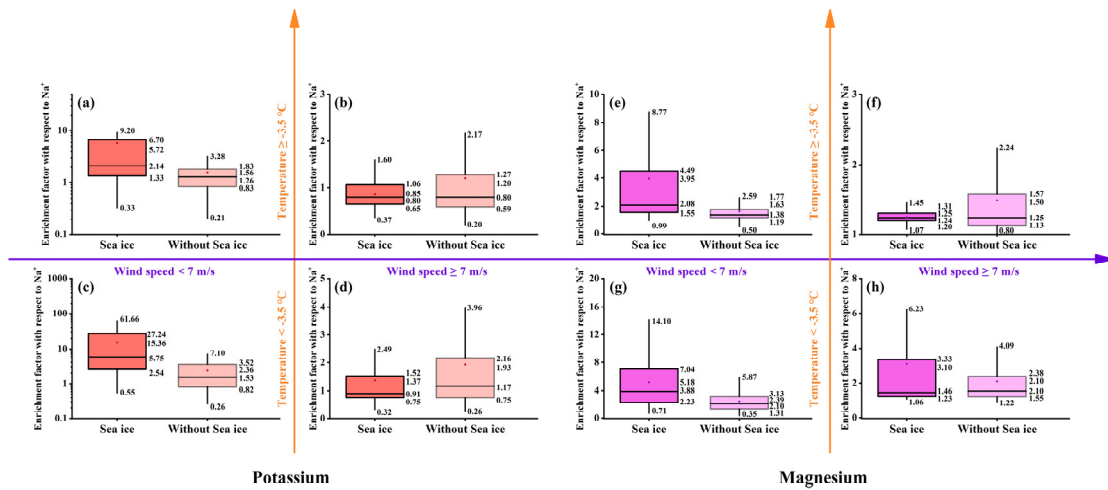
236 The hot plot of number fractions for major species of obtained single-particle types, including
 237 chloride (m/z -35 and -37), sulfate (m/z -97), methanesulfonic acid (MSA, m/z -95), nitro-containing
 238 organic species (m/z -46), organic nitrogen (m/z -26 or -42), phosphate (m/z -63 or -79), carbohydrate
 239 (m/z -45, -59, or -73), siliceous materials (m/z -60), and organic carbon (m/z -27 or 43). (a) leg I (sea
 240 ice period) and leg II (period without sea ice).



241

242 **Figure S17**

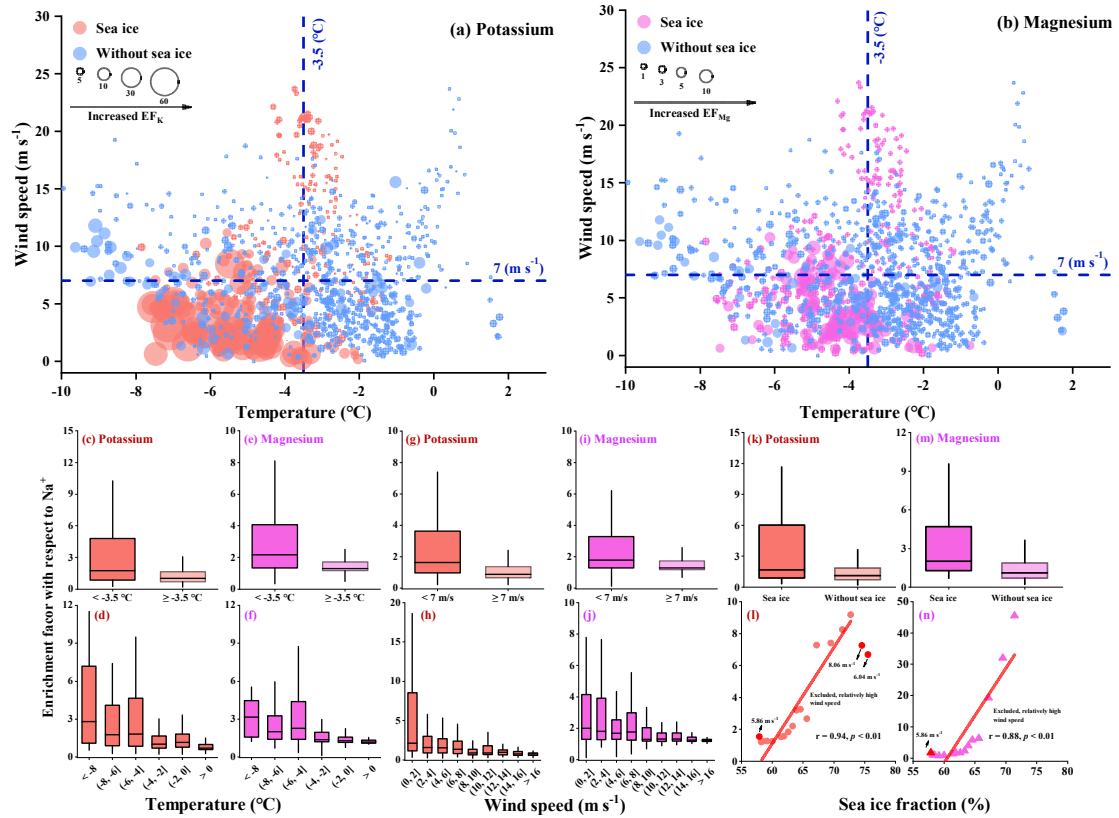
243 Comparison by number fractions for some typical organic chemical components of (a) OC-Ca, (b)
 244 OC-K, and (c) SS-Bio during leg I (sea ice period) and leg II (period without Sea ice). The errors
 245 are calculated assuming Poisson statistics for the obtained particles.



246

247 **Figure S18**

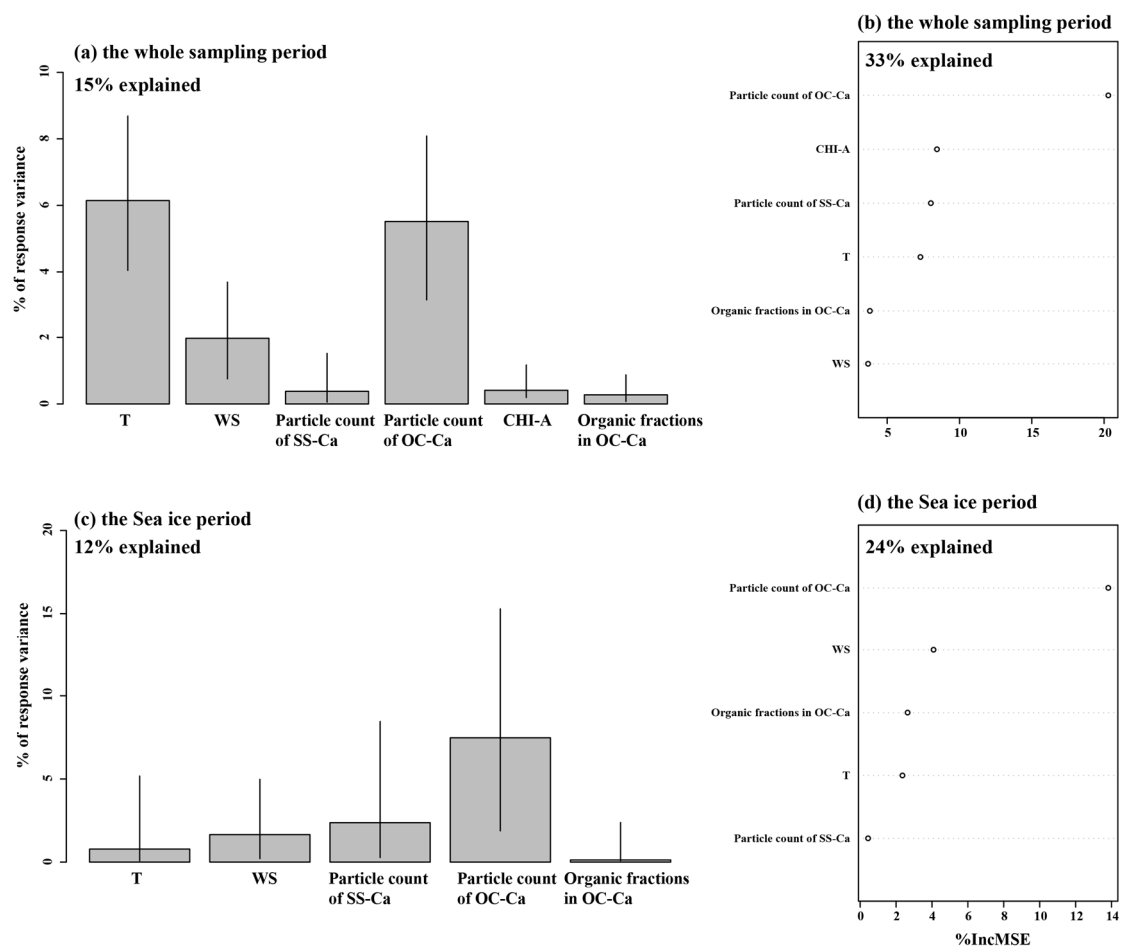
248 Box and whisker plots of hourly Ca^{2+} enrichment factor (EF_{Ca}) with respect to Na^+ at different
 249 ambient temperatures, wind speeds, and sea ice fraction. In the box and whisker plots, the marked
 250 values from top to bottom are the 90th and 75th percentiles, mean, median, and 25th and 10th
 251 percentiles, respectively. In comparison with the bubble charts (Fig. S15), the box and whisker plot
 252 present similar intentions with different forms.



253

254 **Figure S19**

255 Bubble charts of hourly (a) K^+ and (b) Mg^{2+} enrichment factors with respect to Na^+ at the ambient
 256 temperatures, wind speeds, and sea ice fraction. The enrichment factors of K^+ and Mg^{2+} with respect
 257 to Na^+ varied as a function of the ambient temperature (c-f), wind speed (g-j), and sea ice fraction
 258 (k-n) during the cruise observation. In the box and whisker plot, the lower, median, and upper lines
 259 of the box denote the 25th, 50th and 75th percentiles, respectively.



260

261 **Figure S20**

262 The relative importance of the variation in EF_{Ca} , as determined by multilinear regression for (a) the
 263 whole sampling period and (c) the sea ice period. The error bars provide 90% confidence intervals
 264 with 100 bootstrap replicates to evaluate the results. The relative importance of the predictors in the
 265 random forest analysis for the EF_{Ca} for (b) the whole sampling period and (d) the sea ice period.
 266 The %IncMSE, which is used as an indicator for the relative contribution to the predicted variable,
 267 refers to the increased mean square error when each independent variable is removed from the
 268 predictors.

no.	Methodology	EF _{Ca}	EF _K	EF _{Mg}	Ref.
1	Laboratory study (fresh and unfiltered seawater)	median = 1.21	n.a.	n.a.	Keene et al. (2007)
2	Field study (Syowa Station, Antarctica)	n.a.	n.a.	1.35-1.55	Hara et al. (2012)
3	Field study (North Atlantic seawater)	~ 100 (45 nm)	~ 5 (45 nm)	~ 2 (45 nm)	Salter et al. (2016)
		~ 10 (80 nm)	~ 1 (80 nm)	~ 1 (80 nm)	
		~ 4 (150 nm)	~ 1 (150 nm)	~ 1 (150 nm)	
		~ 2 (200 nm)	~ 1 (200 nm)	~ 1 (200 nm)	
		~ 1 (500 nm)	~ 1 (500 nm)	~ 1 (500 nm)	
4	Laboratory study (Artificial seawater)	~ 50 (45 nm)	~ 6 (45 nm)	~ 2 (45 nm)	Salter et al. (2016)
		~ 10 (80 nm)	~ 1 (80 nm)	~ 1 (80 nm)	
		~ 4 (150 nm)	~ 1 (150 nm)	~ 1 (150 nm)	
		~ 2.5 (200 nm)	~ 1 (200 nm)	~ 1 (200 nm)	
		~ 2 (500 nm)	~ 1 (500 nm)	~ 1 (500 nm)	
5	Laboratory study (Artificial seawater)	~ 2 (1000 nm)	~ 1 (1000 nm)	~ 1 (1000 nm)	Cochran et al. (2016)
		~ 5 (56 nm)	~ 1.2 (56 nm)	~ 1.1 (56 nm)	
		~ 1.5 (100 nm)	~ 0.3 (100 nm)	~ 1 (100 nm)	
		median = 1.14	median = 0.72	median = 0.83	
6	Field study (Arctic Ocean)	geometric mean = 3.7 (from 1.2 to 39)	1 - 8	0.1 - 1	Mukherjee et al. (2020)
7	Field study (the Ross Sea, Southern Ocean)	0.01 - 85, median = 2.76, mean = 1.18	0.20 - 75, median = 3.61, mean = 1.19	0.12 - 26.94, median = 2.46, mean = 1.48	This study

n.a. refers that the value was unavailable.

Note: Enrichment factors of a specific species X with respect to sodium (EF_x) are defined as the ratio of the mass concentration of a specific species X to the mass concentration of sodium in the particle to the same ratio in bulk seawater.

270 **Table S1**

271 A comparative analysis of enrichment factors of a specific species X (EF_x , in which X indicates
272 Ca^{2+} , K^+ , and Mg^{2+}) with respect to sodium between previous studies and this study.

273

Item	leg1 (02.12.2017- 20.12.2017)	leg2 (13.01.2018- 14.02.2018)	The whole observation
Na (ng m ⁻³)	306.72 ± 421.77	395.84 ± 561.04	364.64 ± 561.04
Ca (ng m ⁻³)	24.54 ± 41.28	19.38 ± 30.80	21.20 ± 34.96
K (ng m ⁻³)	11.39 ± 7.33	8.72 ± 17.65	9.54 ± 15.28
Mg (ng m ⁻³)	50.63 ± 48.67	59.61 ± 88.89	56.59 ± 77.47
Cl (ng m ⁻³)	18.16 ± 22.18	26.57 ± 36.85	23.63 ± 32.74
EF _{Ca}	3.94 ± 8.50	2.11 ± 4.47	2.76 ± 6.27
Positive calcium enrichment event (%)	71.0%	47.7%	56.0%
EF _K	7.93 ± 14.03	1.67 ± 1.69	3.61 ± 8.45
Positive potassium enrichment event (%)	67.9%	54.2%	58.4%
EF _{Mg}	3.74 ± 3.75	1.80 ± 1.05	2.46 ± 2.53
Positive magnesium enrichment event (%)	99.0%	95.2%	96.3%
Temperature (°C)	-4.1 ± 1.4	-3.2 ± 2.2	-3.5 ± 2.0
Wind speed (m s ⁻¹)	7.2 ± 5.5	7.1 ± 4.2	7.1 ± 4.7
Sea ice fraction	64.91 ± 5.57	54.59 ± 0.08	58.38 ± 6.07
ChI-a concentration (µg L ⁻¹)	0.51 ± 0.29	0.44 ± 0.18	0.46 ± 0.23
96-Trajectory coverage (%)			
Sea ice:	92%	30%	52%
Open water:	4%	12%	9%
Antarctic Land:	4%	58%	39%

Note: During the leg I, the sea ice was retreat. During the leg II, almost no sea ice coverage is equivalent to the sea ice fraction below 55.

274

275 **Table S2**

276 Average measured ion mass concentrations, enrichment factors for specific cations, and
277 meteorological parameters for leg I and leg II.

Particle types	Major peaks	Particle count		Total particle count
		leg1 (02.12.2017- 20.12.2017)	leg2 (13.01.2018- 14.02.2018)	
SSA	[Na] ⁺ , [Na ₂] ⁺ , [Na ₂ Cl] ⁺ , [Mg] ²⁺ , [K] ⁺ , [Ca] ²⁺ , and [Cl] ⁻	69982	71930	141912
SSA-aged	Inorganic salt signature and nitrate of [NO ₂] ⁻ and [NO ₃] ⁻	36905	32741	69646
SSA-Bio	Inorganic salt signature and organic matter signals	5489	21276	26765
OC-Ca	[Ca] ²⁺ and organic matter signals	134653	284861	419514
OC-K	[K] ⁺ and organic matter signals	29734	88549	118283
OC	Organic matter signals	16980	88549	105529
EC	Element carbon with m/z ±C _n , n = 1 - 6	9515	15036	24551

Inorganic salt signatures: [Na]⁺, [Na₂]⁺, [Na₂Cl]⁺, [Mg]²⁺, [K]⁺, [Ca]²⁺, and [Cl]⁻.

Organic matter signals: organic nitrogen ([CN]⁻ and [CNO]⁻), phosphate ([PO₂]⁻ and [PO₃]⁻), carbohydrate ([CHO₂]⁻, [C₂H₃O₂]⁻, and [C₃H₅O₂]⁻), siliceous materials ([SiO₂]⁻), and organic carbon ([C₂H₃]⁺ and [C₂H₃O₃]⁺).

278

279 **Table S3**

280 Particle counts and characteristic peaks for the seven single-particle chemical classes were obtained

281 during cruise observation campaigns.

Item	Area 1 (2018.02.08 22:00- 2017.02.10 22:00)	Area 2 (2018.01.22 17:00- 2018.01.24 04:00)	Area 3 (2017.12.02 07:00- 2017.12.04 19:00)	Area 4 (2017.12.05 00:00- 2017.12.05 23:00)	Area 5 (2017.12.18 22:00-2017.12.19 05:00)	leg I	leg II	The whole observation
Duration (h)	48	35	61	24	8	426	769	1195
EF _{Ca}	10.13 ± 13.63	2.96 ± 2.12	5.47 ± 4.64	9.72 ± 18.75	30.98 ± 31.32	3.94 ± 8.50	2.11 ± 4.47	2.76 ± 6.27
EF _K	2.88 ± 2.36	1.49 ± 0.77	n.a.	45.46 ± 14.79	1.22 ± 0.46	7.93 ± 14.03	1.67 ± 1.69	3.61 ± 8.45
EF _{Mg}	2.88 ± 1.54	1.97 ± 0.69	7.89 ± 4.35	8.25 ± 2.90	1.38 ± 0.33	3.74 ± 3.75	1.80 ± 1.05	2.46 ± 2.53
Temperature (°C)	-6.4 ± 1.2	-2.9 ± 0.8	-4.5 ± 0.9	-4.0 ± 0.8	-1.9 ± 2.2	-4.1 ± 1.4	-3.2 ± 2.2	-3.5 ± 2.0
Wind speed (m s ⁻¹)	5.7 ± 3.5	4.7 ± 1.8	6.04 ± 2.2	2.49 ± 1.1	5.1 ± 4.5	7.2 ± 5.5	7.1 ± 4.2	7.1 ± 4.7
Sea ice fraction	54.60 ± 0.02	54.53 ± 0.00	74.28 ± 1.41	71.41	58.06 ± 0.25	64.91 ± 5.57	54.59 ± 0.08	58.38 ± 6.07
Chl-a concentration (µg L ⁻¹)	0.99 ± 1.65	0.10 ± 0.20	Unavailable	Unavailable	Unavailable	0.51 ± 0.29	0.44 ± 0.18	0.46 ± 0.23
96-Trajectory coverage (%) ^a								
Sea ice:	28%	33%	95%	95%	96%	92%	30%	52%
Open water:	15%	8%	5%	2%	0%	4%	12%	9%
Antarctic Land:	57%	59%	2%	3%	4%	4%	58%	39%

Note: (1) Area 1 and 2 are divided during the leg II, whereas the Area 3, 4, and 5 are divided during the leg I. (2) The values of sea ice fraction and chl-a concentration present with daily resolution. Others present with hourly resolution. (3) No sea ice coverage is equivalent to the sea ice fraction below 55.

282

283 Table S4

284 Average enrichment factors for specific cations and metrological parameters over the different areas
 285 mentioned in **Fig. 5** in main text. Although all areas exhibited significant Ca²⁺ enrichment in SSAs,
 286 they may have varied due to synergetic environmental factors rather than a single factor. Ca²⁺
 287 enrichment in SSA was notably observed with low wind speed, underscoring the effect of wind
 288 speed. The back trajectory coverage is labeled as sea ice, open water, and land.

289 For leg I, the major positive Ca²⁺ enrichment events were associated with Areas 3, 4, and 5. In
 290 addition to the lower wind speed, lower temperature, and the presence of sea ice, the air masses
 291 blowing over the large fraction of sea ice and marginal ice zone may play an important role in Ca²⁺
 292 enrichment. For leg II, the major positive Ca²⁺ enrichment events occurred in Areas 1 and 2, which
 293 were mainly associated with lower wind speed and temperature. The air masses were mostly from
 294 the land-based Antarctic ice.

295

296 **References**

- 297 Boreddy, S. K. R. and Kawamura, K.: A 12-year observation of water-soluble ions in TSP aerosols
298 collected at a remote marine location in the western North Pacific: an outflow region of Asian dust,
299 *Atmos Chem Phys*, 15, 6437-6453, <https://doi.org/10.5194/acp-15-6437-2015>, 2015.
- 300 Cochran, R. E., Jayarathne, T., Stone, E. A., and Grassian, V. H.: Selectivity Across the Interface: A Test
301 of Surface Activity in the Composition of Organic-Enriched Aerosols from Bubble Bursting, *J Phys*
302 *Chem Lett*, 7, 1692-1696, <https://doi.org/10.1021/acs.jpcelett.6b00489>, 2016.
- 303 Collins, D. B., Zhao, D. F., Ruppel, M. J., Laskina, O., Grandquist, J. R., Modini, R. L., Stokes, M. D.,
304 Russell, L. M., Bertram, T. H., Grassian, V. H., Deane, G. B., and Prather, K. A.: Direct aerosol
305 chemical composition measurements to evaluate the physicochemical differences between
306 controlled sea spray aerosol generation schemes, *Atmos Meas Tech*, 7, 3667-3683,
307 <https://doi.org/10.5194/amt-7-3667-2014>, 2014.
- 308 Gaston, C. J., Furutani, H., Guazzotti, S. A., Coffee, K. R., Bates, T. S., Quinn, P. K., Aluwihare, L. I.,
309 Mitchell, B. G., and Prather, K. A.: Unique ocean-derived particles serve as a proxy for changes in
310 ocean chemistry, *Journal of Geophysical Research: Atmospheres*,
311 116<https://doi.org/10.1029/2010jd015289>, 2011.
- 312 Gross, D. S., Galli, M. E., Silva, P. J., and Prather, K. A.: Relative sensitivity factors for alkali metal and
313 ammonium cations in single particle aerosol time-of-flight mass spectra, *Anal Chem*, 72, 416-422,
314 <https://doi.org/10.1021/ac990434g>, 2000.
- 315 Guasco, T. L., Cuadra-Rodriguez, L. A., Pedler, B. E., Ault, A. P., Collins, D. B., Zhao, D. F., Kim, M.
316 J., Ruppel, M. J., Wilson, S. C., Pomeroy, R. S., Grassian, V. H., Azam, F., Bertram, T. H., and
317 Prather, K. A.: Transition Metal Associations with Primary Biological Particles in Sea Spray Aerosol
318 Generated in a Wave Channel, *Environ Sci Technol*, 48, 1324-1333,
319 <https://doi.org/10.1021/es403203d>, 2014.
- 320 Hara, K., Osada, K., Yabuki, M., and Yamanouchi, T.: Seasonal variation of fractionated sea-salt particles
321 on the Antarctic coast, *Geophys Res Lett*, 39<https://doi.org/10.1029/2012gl052761>, 2012.
- 322 Healy, R. M., Sciare, J., Poulain, L., Kamili, K., Merkel, M., Muller, T., Wiedensohler, A., Eckhardt, S.,
323 Stohl, A., Sarda-Estevé, R., McGillicuddy, E., O'Connor, I. P., Sodeau, J. R., and Wenger, J. C.:
324 Sources and mixing state of size-resolved elemental carbon particles in a European megacity: Paris,
325 *Atmos Chem Phys*, 12, 1681-1700, <https://doi.org/10.5194/acp-12-1681-2012>, 2012.
- 326 Keene, W. C., Maring, H., Maben, J. R., Kieber, D. J., Pszenny, A. A. P., Dahl, E. E., Izaguirre, M. A.,
327 Davis, A. J., Long, M. S., Zhou, X. L., Smoydzin, L., and Sander, R.: Chemical and physical
328 characteristics of nascent aerosols produced by bursting bubbles at a model air-sea interface, *Journal*
329 *of Geophysical Research: Atmospheres*, 112<https://doi.org/10.1029/2007jd008464>, 2007.
- 330 Lundberg, S. M., Erion, G., Chen, H., DeGrave, A., Prutkin, J. M., Nair, B., Katz, R., Himmelfarb, J.,
331 Bansal, N., and Lee, S. I.: From local explanations to global understanding with explainable AI for
332 trees, *Nat Mach Intell*, 2, 56-67, <https://doi.org/10.1038/s42256-019-0138-9>, 2020.
- 333 Mukherjee, P., Reinfeldt, J. R., and Gao, Y.: Enrichment of calcium in sea spray aerosol in the Arctic
334 summer atmosphere, *Mar Chem*, 227<https://doi.org/10.1016/j.marchem.2020.103898>, 2020.
- 335 Prather, K. A., Bertram, T. H., Grassian, V. H., Deane, G. B., Stokes, M. D., DeMott, P. J., Aluwihare, L.
336 I., Palenik, B. P., Azam, F., Seinfeld, J. H., Moffet, R. C., Molina, M. J., Cappa, C. D., Geiger, F. M.,
337 Roberts, G. C., Russell, L. M., Ault, A. P., Baltrusaitis, J., Collins, D. B., Corrigan, C. E., Cuadra-
338 Rodriguez, L. A., Ebben, C. J., Forestieri, S. D., Guasco, T. L., Hersey, S. P., Kim, M. J., Lambert,

339 W. F., Modini, R. L., Mui, W., Pedler, B. E., Ruppel, M. J., Ryder, O. S., Schoepp, N. G., Sullivan,
340 R. C., and Zhao, D. F.: Bringing the ocean into the laboratory to probe the chemical complexity of
341 sea spray aerosol, *P Natl Acad Sci USA*, 110, 7550-7555, <https://doi.org/10.1073/pnas.1300262110>,
342 2013.

343 Pratt, K. A. and Prather, K. A.: Mass spectrometry of atmospheric aerosolsuRecent developments and
344 applications. Part II: On-line mass spectrometry techniques, *Mass Spectrom Rev*, 31, 17-48,
345 <https://doi.org/10.1002/mas.20330>, 2012.

346 Qin, X. Y., Bhave, P. V., and Prather, K. A.: Comparison of two methods for obtaining quantitative mass
347 concentrations from aerosol time-of-flight mass spectrometry measurements, *Anal Chem*, 78, 6169-
348 6178, <https://doi.org/10.1021/ac060395q>, 2006.

349 Salter, M. E., Hamacher-Barth, E., Leck, C., Werner, J., Johnson, C. M., Riipinen, I., Nilsson, E. D., and
350 Zieger, P.: Calcium enrichment in sea spray aerosol particles, *Geophys Res Lett*, 43, 8277-8285,
351 <https://doi.org/10.1002/2016gl070275>, 2016.

352 Schlitzer, R.: Interactive analysis and visualization of geoscience data with Ocean Data View, *Comput*
353 *Geosci-Uk*, 28, 1211-1218, [https://doi.org/10.1016/S0098-3004\(02\)00040-7](https://doi.org/10.1016/S0098-3004(02)00040-7), 2002.

354 Sierau, B., Chang, R. Y. W., Leck, C., Paatero, J., and Lohmann, U.: Single-particle characterization of
355 the high-Arctic summertime aerosol, *Atmos Chem Phys*, 14, 7409-7430,
356 <https://doi.org/10.5194/acp-14-7409-2014>, 2014.

357 Song, C., Becagli, S., Beddows, D. C. S., Brean, J., Browse, J., Dai, Q., Dall'Osto, M., Ferracci, V.,
358 Harrison, R. M., Harris, N., Li, W., Jones, A. E., Kirchgäßner, A., Kramawijaya, A. G., Kurganskiy,
359 A., Lupi, A., Mazzola, M., Severi, M., Traversi, R., and Shi, Z.: Understanding Sources and Drivers
360 of Size-Resolved Aerosol in the High Arctic Islands of Svalbard Using a Receptor Model Coupled
361 with Machine Learning, *Environ Sci Technol*, 56, 11189-11198,
362 <https://doi.org/10.1021/acs.est.1c07796>, 2022.

363 Su, B., Wang, T., Zhang, G., Liang, Y., Lv, C., Hu, Y., Li, L., Zhou, Z., Wang, X., and Bi, X.: A review
364 of atmospheric aging of sea spray aerosols: Potential factors affecting chloride depletion, *Atmos*
365 *Environ*, 290<https://doi.org/10.1016/j.atmosenv.2022.119365>, 2022.

366 Su, B. J., Zhang, G. H., Zhuo, Z. M., Xie, Q. H., Du, X. B., Fu, Y. Z., Wu, S., Huang, F. G., Bi, X. H., Li,
367 X., Li, L., and Zhou, Z.: Different characteristics of individual particles from light-duty diesel
368 vehicle at the launching and idling state by AAC-SPAMS, *Journal of Hazardous Materials*,
369 418<https://doi.org/10.1016/j.jhazmat.2021.126304>, 2021a.

370 Su, B. J., Zhuo, Z. M., Fu, Y. Z., Sun, W., Chen, Y., Du, X. B., Yang, Y. X., Wu, S., Xie, Q. H., Huang, F.
371 G., Chen, D. H., Li, L., Zhang, G. H., Bi, X. H., and Zhou, Z.: Individual particle investigation on
372 the chloride depletion of inland transported sea spray aerosols during East Asian summer monsoon,
373 *Sci Total Environ*, 765<https://doi.org/10.1016/j.scitotenv.2020.144290>, 2021b.

374 Yan, J., Jung, J., Lin, Q., Zhang, M., Xu, S., and Zhao, S.: Effect of sea ice retreat on marine aerosol
375 emissions in the Southern Ocean, Antarctica, *Sci Total Environ*, 745, 140773,
376 <https://doi.org/10.1016/j.scitotenv.2020.140773>, 2020a.

377 Yan, J., Jung, J., Zhang, M., Xu, S., Lin, Q., Zhao, S., and Chen, L.: Significant Underestimation of
378 Gaseous Methanesulfonic Acid (MSA) over Southern Ocean, *Environ Sci Technol*, 53, 13064-13070,
379 <https://doi.org/10.1021/acs.est.9b05362>, 2019.

380 Yan, J., Jung, J., Zhang, M., Bianchi, F., Tham, Y., Xu, S., Lin, Q., Zhao, S., Li, L., and Chen, L.: Uptake
381 selectivity of methanesulfonic acid (MSA) on fine particles over polynya regions of the Ross Sea,
382 Antarctica, *Atmos Chem Phys*, 20, 3259-3271, <https://doi.org/10.5194/acp-20-3259-2020>, 2020b.

383 Zhang, G., Fu, Y., Peng, X., Sun, W., Shi, Z., Song, W., Hu, W., Chen, D., Lian, X., Li, L., Tang, M.,
384 Wang, X., and Bi, X.: Black Carbon Involved Photochemistry Enhances the Formation of Sulfate in
385 the Ambient Atmosphere: Evidence From In Situ Individual Particle Investigation, Journal of
386 Geophysical Research: Atmospheres, 126<https://doi.org/10.1029/2021jd035226>, 2021.
387