



# Evaluation of tropical water vapour from CMIP6 global climate models using the ESA CCI Water Vapour climate data records

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**Abstract.** The tropospheric water vapour data record generated within the ESA Climate Change Initiative Water Vapour project (ESA TCWV-COMBI) is used to evaluate the interannual variability of global climate models (CMIP6 framework under AMIP scenarios) and reanalysis (ECMWF ERA5). The study focuses on the tropical belt, with a separation of oceanic and continental situations. The intercomparison is performed according to the probability density function (PDF) of the total column water vapour (TCWV) defined yearly from the daily scale, as well as its evolution with respect to large-scale overturning circulation. The observational diagnostic relies on the decomposition of the tropical atmosphere into percentile of the PDF and into dynamical regimes defined from the atmospheric vertical velocity. Large variations are observed in the patterns among the data records over tropical land, while oceanic situations show more similarities in both interannual variations and percentile extremes. The signatures of El Niño and La Niña events, driven by sea surface temperatures, are obvious over the oceans. Differences also occur over land for both trends (a strong moistening is observed in the ESA TCWV-COMBI data record, which is absent in CMIP6 models and ERA5) and extreme years. The discrepancies are probably associated with the scene selection applied in the data process. Since the results are sensitive to the scene selection applied in the data process, discrepancies are observed among the datasets. Therefore, the normalization process is employed to analyse the time evolution with respect to the mean state. Other sources of differences, linked to the models and their parametrizations, are highlighted.

## 1 Introduction

Water vapour is short-lived yet sufficiently abundant component of the atmosphere that has both direct and indirect impact on weather and environment. It is one of the most important greenhouse gases, and it plays a critical role in the hydrological cycle and climate system (Held and Soden, 2000). It is a radiatively important atmospheric constituent that influences atmospheric energy exchange through interactions with solar and thermal radiations (Raval and Ramanathan, 1989) with strong positive feedbacks (Sherwood et al., 2010). The precipitable water, mainly concentrated in the atmospheric boundary layer, is directly influenced by the surface temperature through robust thermodynamical con-

straints. The concentration of boundary layer water vapour will increase up to  $7\% \text{ } ^\circ\text{C}^{-1}$  globally, confirmed by simulations and observations (Allan et al., 2014). Overall, an increase in atmospheric moisture drives an amplification of feedbacks, yielding changes in evaporation and precipitation patterns at a global scale and amplifying heavy precipitation events (IPCC, 2013; Allan et al., 2020).

Whether it is for weather forecasting, for understanding the evolution of cloud cells or for climate change studies, observing the distribution of atmospheric water vapour at any point in the atmosphere is a central issue. That is why the Global Climate Observing System (GCOS) declared atmospheric water vapour to be one of the Essential Climate Variables (ECVs) (Belward and Dowell, 2016). However, there

are almost 5 orders of magnitude on the water vapour concentration between the surface and the top of the meteorological atmosphere. Evapotranspiration over continental surfaces, atmospheric dynamics or cloud formation reinforces the horizontal and vertical gradients. Unfortunately, the required accuracy at all spatial and temporal scales is difficult to achieve, which leads to the combination of different types of measurements, each with their strengths and weaknesses, depending on the objectives (Wulfmeyer et al., 2015).

Satellite observations have provided global water vapour measurements since the 1970s. Current sensors allow the spatial distribution of water vapour to be observed according to three quantities: the vertical profile of specific humidity ( $q$ , in  $\text{kg kg}^{-1}$ ), the relative humidity of the upper troposphere (UTH, in %) and the total column water vapour (TCWV, in  $\text{kg m}^{-2}$ ). These sensors are deployed either in polar orbit, geostationary orbit or inclined orbit to enhance the temporal sampling of a latitude band. Many programmes have been developed to get long-term observations with high spatial and temporal resolutions. Among them, the European Space Agency (ESA) Climate Change Initiative (CCI) programme has been created to explore the full potential of its Earth observation missions and to generate a climate data record (CDR) associated with each of the 23 ECVs. Among the ESA CCI projects, the CCI Water Vapour project (hereafter CCI\_WV, <https://climate.esa.int/en/projects/water-vapour>, last access: 19 September 2022) started in 2018 with the objective to generate long-term coherent datasets of tropospheric and stratospheric water vapour.

The tropical belt ( $30^{\circ}\text{S}$ – $30^{\circ}\text{N}$ ) is a pivotal region in the Earth's climate. In this part of the globe, regional variations of the hydrological cycle are closely related to the Hadley–Walker overturning cells, which define the tropospheric circulation, and to its long-term changes (such as its observed slowdown and poleward expansion; Ma et al., 2018; Lu et al., 2007).

Studies linking tropical climate evolution and water vapour distribution are done with both models and long-term observations. For instance, the infrared (IR) observations from the High-resolution Infrared Radiation Sounder (HIRS) instrument have been used to investigate the interannual variability of tropical moisture of an atmospheric model (Huang et al., 2005). Following the same approach, Chung et al. (2011) analysed the variability of the simulated UTH from a climate model using both IR and microwave measurements and showed that the wet bias of the model was related to errors in simulating the intensity of large-scale circulation. Globally, improvements have been noticed in the simulation of the tropospheric water vapour distribution and variability between the Coupled Model Intercomparison Project (CMIP5, release in 2014) and the earlier CMIP3 exercise (release in 2010) (Jiang et al., 2012). However, while the models perform best for the boundary layers over the oceans, most likely thanks to the thermodynamic constraint imposed by

the sea surface temperature, strong biases remain in the upper layers of the troposphere, where clouds and atmospheric dynamics have large uncertainties.

A recent intercomparison work of a selection of available long-term datasets has been conducted under the auspices of the Global Energy and Water Exchanges (GEWEX) programme of the World Climate Research Programme (WCRP) (Schröder et al., 2019). This intercomparison not only highlighted the complementarity among the sensors, but also underlined the caveats in the studies of trends and variabilities induced by artificial break points contained in the CDRs such as calibration changes, retrieval algorithms and resolution changes that impact the sampling. Apart from these important points that are inherent to the development of robust time series suitable to investigate climate variability, the evaluation of the distribution and variability of the water vapour with respect to large-scale circulation is still of utmost importance. The last IPCC assessment report (AR6, 2021) contains an entire chapter on the water cycle that highlights the role of the large-scale atmospheric circulation in driving regional changes of atmospheric moisture fluxes and in position and strength of the tropical rain belt (Arias et al., 2021).

The present work follows the previous analysis framework proposed by Bony et al. (2004) and investigates the variability of total column water vapour (TCWV) of a selection of global climate models (GCMs) that have participated in the CMIP6 exercise (Eyring et al., 2016). The ESA CCI\_WV CDRs are studied from the aspect of the large-scale circulation. Section 2 describes the various datasets: the ESA CCI\_WV CDRs, the CMIP6 GCMs and the ERA5 reanalysis. Section 3 is dedicated to the method of evaluation itself, while Sect. 4 discusses the results. Finally, the conclusions are drawn in Sect. 5.

## 2 Data description

We focus on the tropical belt ( $30^{\circ}\text{S}$ – $30^{\circ}\text{N}$ ). Data obtained from ESA CCI\_WV, CMIP6 GCMs, and ERA5 reanalysis are studied at daily scale and over the period 2003–2014. The monthly atmospheric vertical velocity at 500 hPa ( $\omega_{500}$ ) of each data record is used as a proxy of large-scale circulation following Bony et al. (2004). Note that the vertical velocity from ERA5 is also used as the dynamical reference for the CCI\_WV CDRs. The daily water vapour data and monthly mean  $\omega_{500}$  are adopted in our analysis for several reasons. Firstly, our intention is to evaluate the datasets with the highest temporal resolution, as the temporal averaging will mask out the extremes, and the probability density functions (PDFs) would have been smoothened. Secondly, the cloud condition varies significantly over short timescales; therefore, quantification at high temporal resolution is required. Last but not least, the previous study suggests that the  $\omega_{500}$  is sensitive to local dynamics and subject to significant biases at the instantaneous scale (Trenberth et al., 2000). Re-

search shows that the  $\omega 500$  data with shorter timescales are unreliable (Höjgård-Olsen et al., 2020). The monthly vertical motion can represent a mixture of ascending and descending atmospheric conditions. It is worth mentioning that by adopting the monthly mean of  $\omega 500$  in our evaluation, the fluctuations of shorter timescales, where small-scale convection probably dominates, are ignored.

## 2.1 ESA CCI\_WV climate data records

Phase 1 of the ESA CCI\_WV project was dedicated to built climate records of both tropospheric and stratospheric water vapour. The project provides daily and monthly water vapour observations on the global scale with spatial resolution of 0.5 and 0.05° for the period of 2002–2017.

The TCWV is commonly defined as the vertically integrated water vapour over the full column with units of kilograms per square metre ( $\text{kg m}^{-2}$ ). Observations from microwave (MW) imagers (namely SSM/I, SSMIS, AMSR-E and TMI) over the ice-free ocean, partly based on a fundamental CDR (Fennig et al., 2020), and near-infrared (NIR) imagers (including MERIS, MODIS-Terra and OLCI) over land, coastal ocean and sea ice have been combined within the ESA CCI\_WV project. Details of the retrieval are discussed in Andersson et al. (2010) and Schröder et al. (2013) for the MW imagers. The algorithms for NIR imagers are discussed in Lindström et al. (2012), Diedrich et al. (2015) and Preusker et al. (2021). The MW and NIR data streams are processed independently and combined afterwards so that the individual TCWV values and their uncertainties remain unchanged. Validation of the dataset against GRUAN and SuomiNet shows that the bias and corrected RMSD (cRMSD) are generally small and within  $\pm 1.5$  and  $2.5 \text{ kg m}^{-2}$ , respectively. The available spatial resolutions of the combined data record (hereafter TCWV-COMBI) are  $0.5^\circ \times 0.5^\circ$  and  $0.05^\circ \times 0.05^\circ$ , where the NIR based data are averaged and the MW-based data are oversampled to produce the results with desired resolution. The daily and monthly mean data are available for the TCWV-COMBI product during July 2002–December 2017. Table 1 summarizes the various original data sources that are used in the TCWV-COMBI CDR.

Here we use the daily  $0.05^\circ \times 0.05^\circ$  TCWV-COMBI dataset. However, as detailed above, the data processing of the TCWV-COMBI is different between land and ocean, which then impacts the processing of the CMIP6 models: over land areas, TCWV is estimated under cloud-free conditions, while over ocean areas, TCWV is estimated until heavy precipitation occurs. Moreover the evaluation period is restricted to before 2014 for consistency with the available period of the CMIP6 experiment. Such a cut in the ESA TCWV-COMBI product excludes the OLCI observations.

## 2.2 CMIP6 models

Seven GCMs participating in CMIP6 are evaluated here, limited by the availability of the required geophysical variables at daily resolution (at least) that is comparable with the CCI\_WV CDRs. However, there was no TCWV field at the daily frequency that was available from the Earth System Grid Federation (ESGF) (node of Institut Pierre Simon Laplace, IPSL). Therefore we recomputed TCWV from the vertical profiles of specific humidity  $q$  (in  $\text{g kg}^{-1}$ ) that were provided at the model vertical resolution. High vertical resolution of specific humidity (more than 19 vertical levels in the troposphere) was then necessary to be certain to capture the full tropospheric water vapour (using the extraction on a selection of pressure levels would bias the computation of TCWV).

The TCWV (in  $\text{kg m}^{-2}$ ) from each model is thus calculated using

$$\text{TCWV} = \int_{\text{surface}}^{\text{top}} q \frac{dp}{g}, \quad (1)$$

where  $g$  is the gravitational acceleration constant, and  $dp$  is the difference between adjacent pressure levels (hPa).

We focus on the AMIP (Atmospheric Model Inter-comparison Project) (Ackerley et al., 2018) scenario with prescribed time-varying sea surface temperature (SST) and sea-ice concentrations from observations, including variations in natural and anthropogenic external forcings (Eyring et al., 2016). The detailed model descriptions are listed in Table 2. In addition to the CMIP6 models, the ensemble mean of the seven models is also included in the following analysis to represent the mean state of the CMIP6 models.

Since the TCWV-COMBI data are cloud-screened over the land area, it is important to carefully analyse the cloud conditions for CMIP6 models before making the quantitative comparison. The cloud screening of the model-to-observation approach must remove the cloudy pixels while maintaining enough data for further evaluation. Since cloud fraction (cf) is the only parameter that is available for the CMIP6 models to be employed as the indicator for cloudiness, a series of threshold tests were conducted by screening the pixels with cf values larger than 5 %, 25 %, 50 % and 75 % at all pressure levels. The distribution of TCWV over tropical land from CanESM5 of the CMIP6 model with different cloud masks, along with data from TCWV-COMBI and ERA5, is shown in Fig. 1. The land area fraction product is adopted as the land–sea mask for the CanESM5 model. Here we defined the area as land where the percentage of the grid cell occupied by land is larger than 50 %. The results indicate that data with cf less than 50 % at all pressure levels show a reasonable spatial coverage comparable to TCWV-COMBI and ERA5. Therefore, this threshold is adopted as the general cloud mask for the CMIP6 models, although this cannot be considered a purely clear sky. It is worth mentioning that

**Table 1.** Summary of the characteristics of CCI water vapour data TCWV-COMBI.

Data source	Spectral domains	Region	Data description	Spatial resolution	Time span	Reference
MERIS	NIR	Land, coastal and sea ice	Daytime, cloud-free	1200 m	2002–2012	Fischer and Bennartz (1997)
MODIS	NIR	Land, coastal and sea ice	Daytime, cloud-free	1000 m	2011–2017	Gao and Kaufman (2003)
OLCI	NIR	Land, coastal and sea ice	Daytime, cloud-free	1200 m	2016–2017	Lindstrot et al. (2012)
HOAPS	MW	Ocean	6-hourly composites, without strong precipitation	0.5°	2002–2017	Lindstrot et al. (2014)

**Table 2.** Main characteristics of the seven CMIP6 models of the study, along with TCWV-COMBI and ERA5. The percentages over land and ocean are computed once the scene selection is applied.

Institution	Model ID	Horizontal resolution	Vertical resolution	Percentage of land data (%)	Percentage of ocean data (%)	Reference
CCCma	CanESM5	$2.81^\circ \times 2.81^\circ$	49 levels (1–1022 hPa)	55.63 %	99.89 %	Swart et al. (2019)
CNRM-CERFACS	CNRM-CM6-1	$1.41^\circ \times 1.41^\circ$	91 levels (0.1–1039 hPa)	62.85 %	99.86 %	Voldoire et al. (2019)
	CNRM-ESM2-1	$1.41^\circ \times 1.41^\circ$	91 levels (0.1–1039 hPa)	62.81 %	99.86 %	S��f��rian et al. (2019)
IPSL	IPSL-CM6A-LR	$1.25^\circ \times 2.50^\circ$	79 levels (0–1028 hPa)	76.10 %	99.79 %	Lurton et al. (2020)
MPI-M	MPI-ESM1-2-HR	$0.94^\circ \times 0.94^\circ$	95 levels (0–1055 hPa)	69.90 %	99.98 %	M��ller et al. (2018)
NCAR	CESM2	$0.94^\circ \times 1.25^\circ$	32 levels (4–993 hPa)	47.14 %	99.97 %	Danabasoglu et al. (2020)
	CESM2-WACCM	$0.94^\circ \times 1.25^\circ$	70 levels (0–993 hPa)	46.14 %	99.97 %	Gettelman et al. (2019)
ESA CCI_WV	TCWV-COMBI	$0.05^\circ \times 0.05^\circ$	–	43.73 %	99.82 %	Ref as in Table 1
ECMWF	ERA5	$0.5^\circ \times 0.5^\circ$	–	52.76 %	97.14 %	Hersbach et al. (2020)

since the results are dependent on the cloud masks, the normalization process is also included in the following analysis to better illustrate the time evolution with respect to the mean state instead of strict biases.

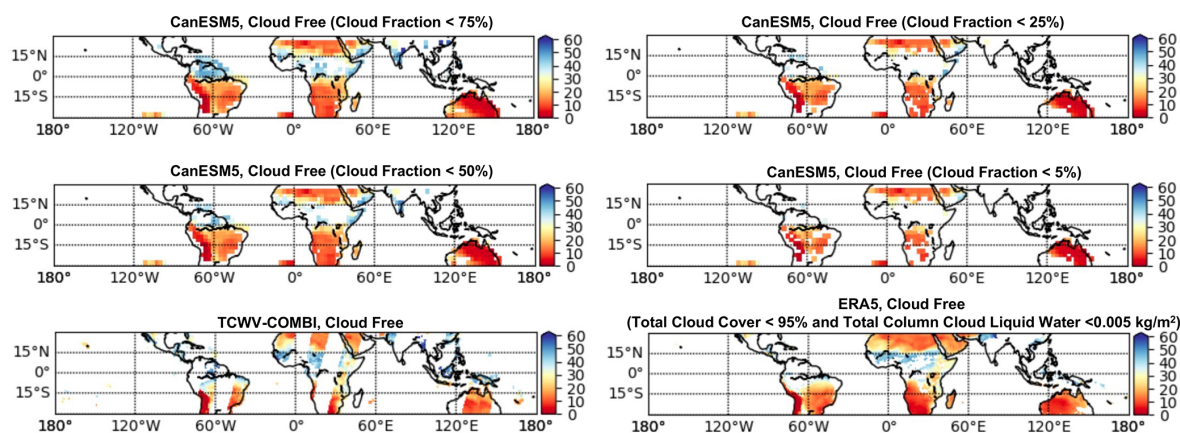
The description of the datasets remaining for the following analysis after the cloud screening over the land and ocean is displayed in Table 2. Over land, the percentage of data that remained for the CMIP6 models after the cloud screening is globally in the range 46.14 %–76.10 %. Over tropical oceans, the percentage of data that remained after removing the pixels under heavy precipitation conditions ranges from 99.79 % to 99.98 %. Although the scene selection is more stringent over land, this indicates that the CMIP6 data used in the fol-

lowing analysis are comparable in terms of size of sample to the data from TCWV-COMBI and ERA5. It is worth mentioning that although the screening thresholds for the models are set to meet the criteria of the TCWV-COMBI product, the number of data retained for the comparison are not exactly the same for all models. Therefore, differences among datasets may be observed in the analysis. This is particularly true over tropical land.

### 2.3 ERA5

The reanalysis data are widely analysed in atmospheric sciences to assess the impact of changes in observation system,





**Figure 1.** Examples of daily mean TCWV over tropical land obtained from CanESM5 with different cloud mask (pixels with cloud fraction larger than 75 %, 50 %, 25 % and 5 % at all pressure levels are considered clouded), along with data from TCWV-COMBI and ERA5 from 1 July 2003.

to scale progress in model simulations and to calculate climatology for forecast-error evaluation (Hersbach et al., 2020). The ECMWF's ERA5 TCWV data are based on the integrated forecasting system (IFS) Cy41r2, with considerably enhanced horizontal resolution of 31 km compared to 80 km for ERA-Interim. Here, the ERA5 TCWV data with hourly frequency are averaged into daily data. To compare the data under the same conditions, the ERA5 land–sea mask is employed for land and ocean separation, and a scene selection is performed and is similar to the process of the CMIP6 data. Hence, data with total cloud cover less than 95 % and total column cloud liquid water less than  $0.005 \text{ kg m}^{-2}$  over land (Sohn and Bennartz, 2008) and data with total precipitation less than  $0.001 \text{ kg m}^{-2} \text{ s}^{-2}$  over ocean are retained.

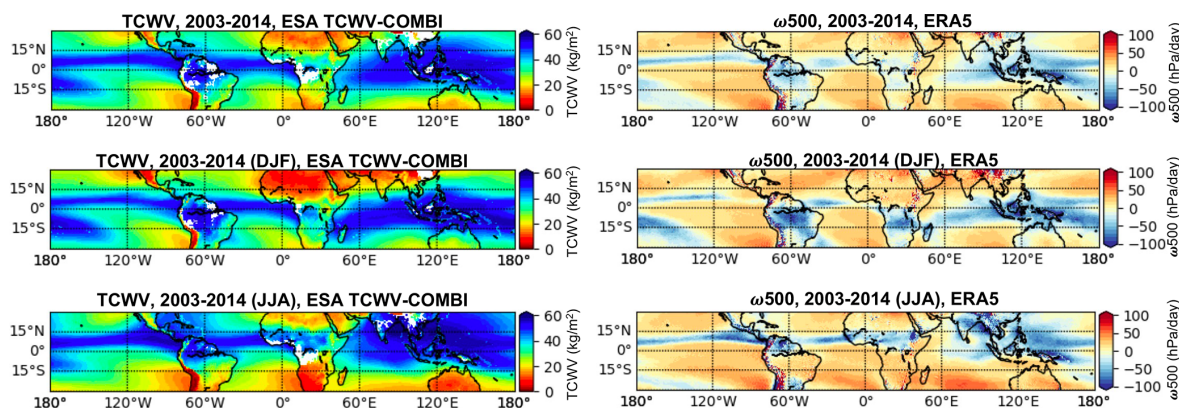
### 3 Methods

The time series of the daily means of the CMIP6, ERA5 and ESA CCI\_WV TCWV-COMBI are analysed with tropical-land and tropical-ocean separation over the common observation period that covers July 2003 to December 2014.

The intercomparisons are conducted according to two approaches.

1. The first approach evaluates the interannual variation of TCWV based on the probability distribution function (PDF) established from the daily records for each year of the period. The percentiles of the TCWV are defined from the yearly distributions, and the data are sorted by intervals of 10 percentiles. Finally, the mean TCWV of each interval is computed and normalized by the corresponding mean TCWV of the whole observation period for this given percentile. This is generalized for every percentile. This approach is meant to highlight the tropical anomalies with respect to the mean and trace back to the inter-annual variability of the tropical atmosphere.

2. The second approach is based on the fact that the water vapour distribution is strongly controlled by the large-scale vertical motion of the atmosphere. Therefore, we can use the mid-tropospheric atmospheric vertical velocity at 500 hPa (noted  $\omega_{500}$  in  $\text{hPa d}^{-1}$ ) as a proxy for the vertical motions in the tropics (Bony et al., 2004). While such framework has been greatly used to study tropical clouds and their distribution (e.g., Konsta et al., 2012; Höjgård-Olsen et al., 2020), this link between vertical motion and TCWV is documented (e.g., Brogniez and Pierrehumbert, 2007) and further illustrated on Fig. 2. Figure 2 presents the TCWV-COMBI averaged over the whole 2003–2014 period as well as the mean winter (December, January and February – DJF) and mean summer (June, July and August – JJA), together with the corresponding  $\omega_{500}$  taken from ERA5 at a monthly scale. As expected, a moist troposphere is associated with large-scale ascending motion ( $\omega_{500} < 0 \text{ hPa d}^{-1}$ ), while a dry troposphere is associated with large-scale subsidence ( $\omega_{500} > 0 \text{ hPa d}^{-1}$ ). The TCWV data are sorted upon  $10 \text{ hPa d}^{-1}$  bins of monthly values of  $\omega_{500}$ . The dynamical decomposition is performed for all TCWV data records for each year of the time period. Moreover, the TCWV data averaged over the whole 2003–2014 period are also sorted into the corresponding  $\omega_{500}$  bins of the period, and this value is considered the reference to normalize the results. This second approach allows us to study the trends in TCWV for a given state of the large-scale dynamics and thus overcome issues associated with variations (such as shifts or expansion) of the atmospheric circulation (Vallis et al., 2015; Mbengue and Schneider, 2017).



**Figure 2.** Maps of the TCWV-COMBI (in  $\text{kg m}^{-2}$ ) during 2003–2014 for the tropical region ( $30^\circ\text{S}$ – $30^\circ\text{N}$ ) for the whole period, winter (December, January and February – DJF) and summer (June, July and August – JJA) and the corresponding maps of ERA5  $\omega_{500}$  (in  $\text{hPa d}^{-1}$ ).

## 4 Results and discussions

This section aims to assess the degree of agreement in the TCWV climatology and interannual variations between the ESA CCI\_WV TCWV-COMBI, CMIP6 models and ERA5 reanalysis data over the tropics ( $30^\circ\text{S}$ – $30^\circ\text{N}$ ). The distribution of the water vapour over tropics and its link to large-scale circulation ( $\omega_{500}$ ) are discussed in detail.

### 4.1 Description of the tropical TCWV: 2003–2014

#### 4.1.1 Time series

Figure 3 shows the time series of the TCWV of the different datasets over land (Fig. 3a; clear skies only) and over ocean (Fig. 3b; all weather without heavy precipitation) for the period July 2003–December 2014. There are differences in the range of daily mean TCWV, but all datasets varied with the seasons. Overall, the different data records agree well with each other despite some outliers observed over tropical land in TCWV-COMBI data that resulted from the differences in available sample size under the clear-sky condition each day. Strong seasonal variations are observed over tropical land with minima reached during JFM and maxima reached during JJA and with a very weak interannual variability. Since the results are dependent on the cloud-screening process, the detailed discussions on the impact of El Niño–Southern Oscillation (ENSO) events are discussed in Sect. 4.2. Over the tropical oceans, the seasonal variations are weaker (the minima still occur during JFM, but the maxima are not always reached in JJA) with a strong interannual variability.

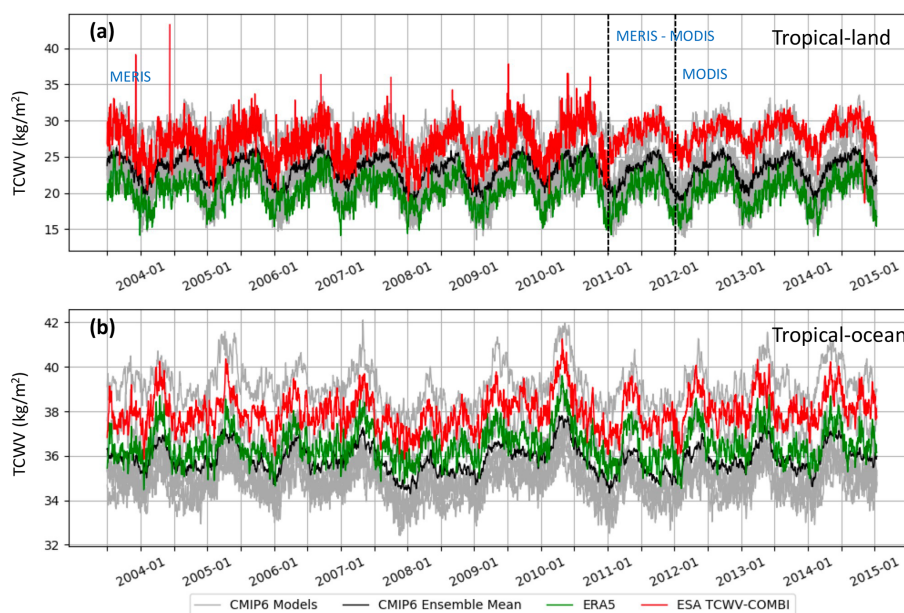
More specifically, the ESA TCWV-COMBI data are moister than ERA5 and most CMIP6 models over both the land and the ocean areas, and this moist bias is even more pronounced over tropical land (Fig. 3a,  $\sim 10\text{ kg m}^{-2}$  over land vs.  $\sim 2\text{ kg m}^{-2}$  over ocean). On the other hand, the daily mean values of water vapour concentration over ocean areas are higher than the values over land areas. This difference can

be explained because the TCWV datasets over land areas are composed of clear-sky-only data, which are likely drier than the nearby cloud area for a given location, thus translating into a dry bias associated with moistening processes by convective clouds (Sohn et al., 2006). Hence, the cloud screening over land makes it difficult to compare the datasets directly. In addition, since the boundary layer is drier in the continental subtropics and the maritime stratocumulus zones are wetter at low levels, the ocean areas should appear moister than the land areas.

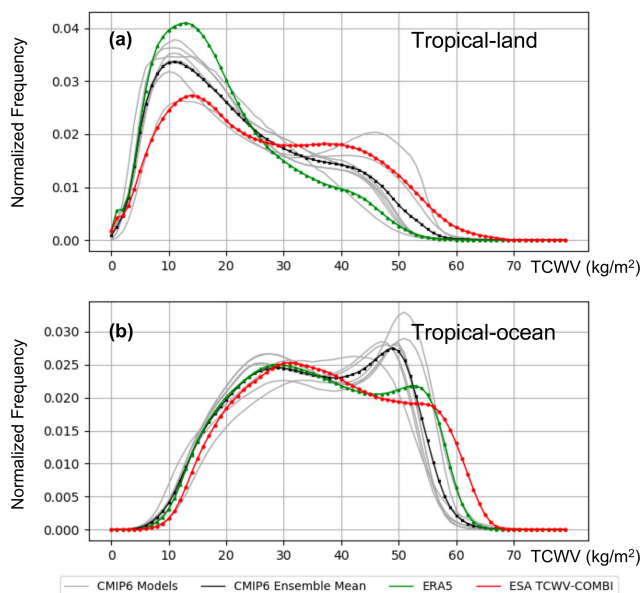
#### 4.1.2 The distribution of tropical TCWV

The normalized PDFs of the daily TCWV obtained from all data records over both land and ocean are displayed in Fig. 4. The bimodal distributions can be explained by the presence of more humid columns in the intertropical convergence zone (ITCZ) and relatively drier ones in subtropical regions. Similar to the time series analysis, the characteristics over land are significantly different to the results over the ocean area because of the data screening. Over land, all datasets reach a first maximum at around  $10$ – $13\text{ kg m}^{-2}$  and present a secondary maximum near  $40$ – $50\text{ kg m}^{-2}$ . Moreover, more lower values are observed in the CMIP6 models and ERA5 than in the ESA TCWV-COMBI dataset, and this cannot be explained by the cloud-screening method alone. Indeed, the results of simulation models and ERA5 are dependent on cloud conditions, consequently leading to differences in the comparison. Although there are changes in variability in TCWV-COMBI data over tropical land with the inclusion of MODIS from 2011, validation of the dataset against GRUAN and SuomiNet shows that the dataset is stable and accurate during the whole observation period.

Over oceans, most of the TCWV data are located around  $20$ – $60\text{ kg m}^{-2}$ . The main peak is around  $30\text{ kg m}^{-2}$ , and a secondary peak appears near  $50\text{ kg m}^{-2}$ . While the mean peak of PDF is nearly identical for TCWV-COMBI, ERA5 and CMIP6, there is a divergence for the secondary peak.



**Figure 3.** Time series of daily mean TCWV in the tropics ( $30^{\circ}$  S– $30^{\circ}$  N) over (a) land areas under clear-sky conditions and (b) ocean areas except for heavy precipitation (see details in Sect. 2). The time series cover the period July 2003–December 2014. The grey lines denote the individual CMIP6 models, while the black line represents their ensemble mean. The green line represents ERA5, and the red line is the ESA CCI<sub>WV</sub> TCWV-COMBI.



**Figure 4.** The normalized PDFs of the TCWV in the tropical area ( $30^{\circ}$  S– $30^{\circ}$  N) over (a) land areas (under clear-sky-only conditions) and over (b) ocean areas (under all-weather conditions except for heavy precipitation). The grey lines denote the individual CMIP6 models, while the black line represents their ensemble mean. The green line represents ERA5, and the red line is the ESA CCI<sub>WV</sub> TCWV-COMBI.

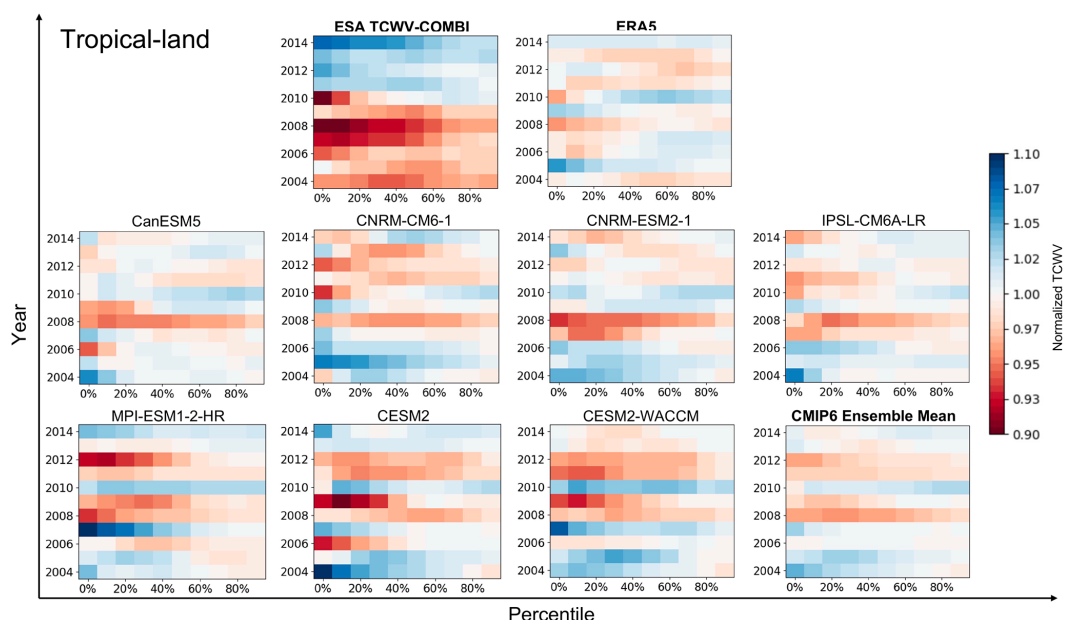
This secondary peak even dominates the PDF of the CMIP6 models, while ERA5 and ESA TCWV-COMBI are still quite similar.

#### 4.1.3 Extremes of the distributions

The data records are then evaluated following the approach (1) described in Sect. 3: the percentiles of the annual distributions of TCWV (at daily resolutions) are sorted into bins of 10 % intervals, and this is done for each year of the period.

The normalized TCWV for land areas and sorted by 10 percentile intervals for each year is displayed in Fig. 5. The bluish colours indicate that the TCWV value of the interval is larger than the reference value, indicating wet anomalies. The reddish colours indicate that the TCWV value is smaller than the reference, indicating dry anomalies. As shown in the figure, the ESA TCWV-COMBI data have quite different characteristics compared to CMIP6 and ERA5 results. A clear moistening trend is observed in the drier percentiles of TCWV-COMBI record. The tipping point seems to be 2011 and thus may be caused by the inhomogeneity of cloud-mask products for different observation instruments when computing the ESA TCWV-COMBI data. Indeed, the data record merged NIR observation from MERIS over 2002–2012, and MODIS observations are included over 2011–2017. Despite individual discrepancies, the CMIP6 ensemble mean is in good agreement with the ERA5 data. Overall, anomalies are observed in the time period for all data records: 2008 appears





**Figure 5.** Normalized percentiles of the TCWV over land areas for every data record. The percentiles are grouped into bins of 10 % intervals. The  $x$  axis represents the percentiles intervals, and the  $y$  axis represents the year. Note that the period starts in 2004 instead of 2003 to focus on full years.

to be a dry year, while 2010 reveals a clear signal of humidification, especially over the high parts of the distributions of TCWV (percentiles  $> 60\%$ ).

A similar comparison is performed over the tropical oceans, and the results are shown in Fig. 6. The TCWV-COMBI results coincided well with CMIP6 models and ERA5 data. Dry anomalies are observed in 2004 for the ESA TCWV-COMBI and ERA5 in the dry end of TCWV. Dry anomalies are also observed in 2008 and 2011 over the highest percentiles ( $> 60\%$ ) for all data records, while 2008 is the driest year of the period. Wet anomalies are observed in 2010 for the highest percentiles in all data records. ERA5 also reveals 2012 as a moister year, in the low range of TCWV, but this anomalous year is not present in the ESA TCWV-COMBI or CMIP6 ensemble mean. The very good agreement among the various datasets is largely due to the fact that the CMIP6 models are evaluated under the AMIP scenario, so with the same prescribed SST for all models and ERA5, and that the relationship between SST and TCWV is largely explained by the Clausius–Clapeyron law (Stephens, 1990). Hence this explains that anomalous years are the results of El Niño–Southern Oscillation (Trenberth et al., 2005): 2008 and 2011 are characterized by a very negative ENSO index, while 2010 is an intermediate year, which starts with a positive ENSO cycle and is followed by a negative one.

## 4.2 TCWV and large-scale circulation

### 4.2.1 General assessment

The interannual variability of TCWV is then analysed from its links with the large-scale atmospheric circulation and follows approach (2) described in Sect. 3. The monthly  $\omega_{500}$  of individual data records is decomposed into intervals of  $10 \text{ hPa d}^{-1}$  in the range of  $-120$  to  $120 \text{ hPa d}^{-1}$ . Figure 7 displays the normalized PDFs of the  $\omega_{500}$  of the CMIP6 models and ERA5. As mentioned earlier, there are no atmospheric circulation data from the ESA TCWV-COMBI data record, so the  $\omega_{500}$  from ERA5 is also employed as the reference for this dataset. Figure 7a and b show that the PDFs of the  $\omega_{500}$  from the CMIP6 ensemble mean agree well with the ERA5 data. Most of the  $\omega_{500}$  reside in around  $10$ – $20 \text{ hPa d}^{-1}$  over both the tropical-land and tropical-ocean area, which characterizes the dominance of the large-scale Hadley subsidence in the subtropical free troposphere and explains the clear-sky radiative cooling of the tropics as discussed in Bony et al. (2004).

The TCWV of each dataset is then sorted into the vertical velocity bins by the corresponding value of  $\omega_{500}$ . Therefore, the variations of the TCWV can be analysed from the perspective of atmospheric vertical motion. As shown in Fig. 7c and d, large-scale downward motion is associated with a dry troposphere, while large-scale ascent is associated with a moister troposphere (except for ERA5 over land, where the most humid regimes are over weak subsidence). The results are in agreement with the maps of Fig. 2. Over-



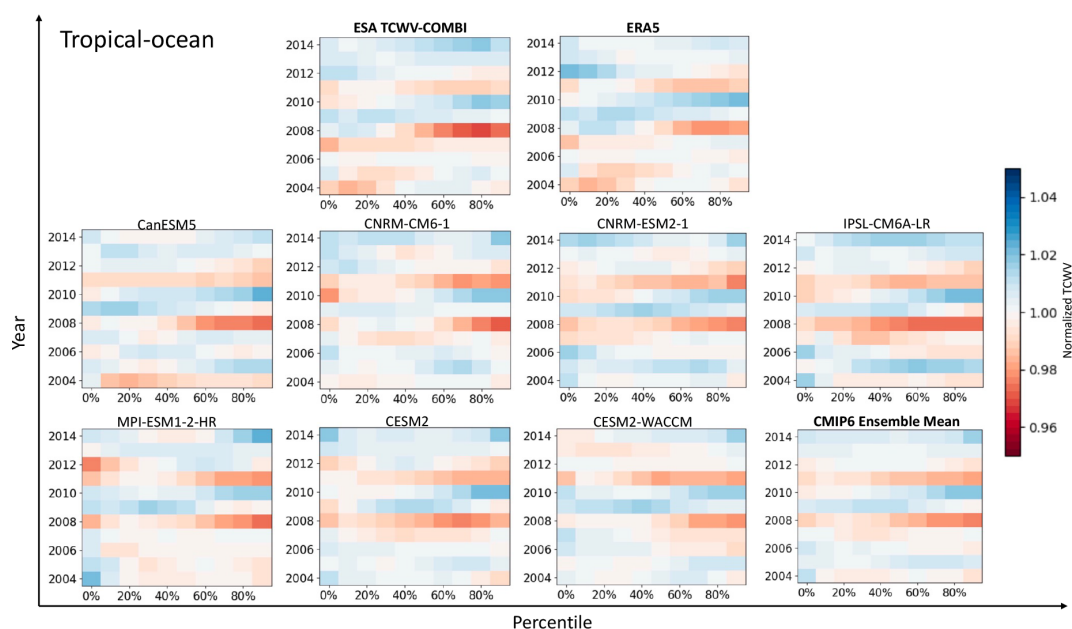


Figure 6. Same as Fig. 5 but for oceans.

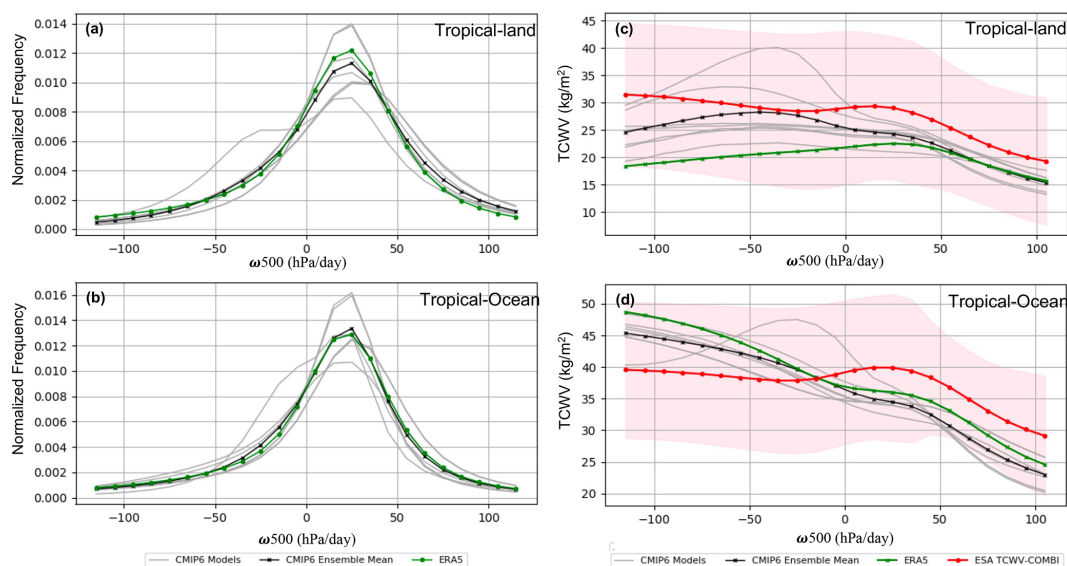


Figure 7. (a, b) Normalized PDFs of  $\omega_{500}$  (in  $\text{hPa d}^{-1}$ ) over land (a) and ocean (b) for CMIP6 models (grey lines), their ensemble mean (black line) and ERA5 (green line). (c, d) Mean TCWV from the CMIP6 models (grey lines), their ensemble mean (black line), ERA5 (green line) and ESA TCWV-COMBI (red line) in different circulation regimes of  $\omega_{500}$  over land (c) and ocean (d) areas. The shaded area in pink represents the  $\sigma$  of each bin in TCWV-COMBI data.

all, the CMIP6 models and ERA5 show differences with the ESA TCWV-COMBI data in the amplitude of the signal and in the gradient of moisture between the ascending and descending regions. The evaporation from the oceans is the primary source of water vapour in the atmosphere; the oceanic boundary layer is humid, even in a weak subsidence regime. Further, as the reference  $\omega_{500}$  for TCWV-COMBI is from ERA5, it is sensible that there are differences observed in the

TCWV-COMBI data comparing to other datasets that are decomposed by the corresponding model products. While the large-scale atmospheric dynamics are consistent among the datasets, the discrepancies in the TCWV reveal difficulties in representing the moistening processes of the tropical atmosphere: lateral mixing (Pierrehumbert and Roca, 1998; Pierrehumbert, 1998), outflows from clouds and too high-/too low precipitation efficiencies of the convective schemes

(Brogniez and Pierrehumbert, 2007). It is worth mentioning that the moistest regime of TCWV from ERA5 over land areas occurred in a weak subsidence regime instead of the strong ascending region. This difference is partly because of the cloud-screening processes. Therefore, the results could not accurately represent the large-scale advection humidification/drying processes.

#### 4.2.2 Trends over lands

This global assessment is further discussed by applying the TCWV- $\omega 500$  approach for every year of each data record to delineate the trends in TCWV. As shown in Fig. 8, all the data records (except for ERA5) agree that the driest troposphere (red) is associated with the most positive  $\omega 500$  bins (meaning the areas of highest downward motion). The moistest troposphere, however, is not always located in the most negative  $\omega 500$  bins (the highest upward motion).

ERA5 and the CMIP6 ensemble mean display the lowest TCWV values ( $< 16 \text{ kg m}^{-2}$ ) that occur all along the 2004–2014 period, while the ESA TCWV-COMBI data record reaches values  $\sim 18 \text{ kg m}^{-2}$ , and this minimum is reached over 2007–2008. There is also a very strong variability amongst the CMIP6 models: IPSL-CM6A-LR is clearly the moistest model, and CanESM5 is the driest. The moist bias of IPSL-CM6A-LR is already documented (Boucher et al., 2020) and is explained by the (too) efficient parametrization scheme of the transport of evaporated air from the surface to the top of the boundary layer. The discrepancy observed from CanESM5 is partly because of its strong effective climate efficiency compared to other CMIP6 models (Virgin et al., 2021). For the CanESM5, the positive low and non-low shortwave cloud feedbacks, as well as subtropical and extratropical free troposphere cloud optical depth, particularly with regards to low clouds across the equatorial Pacific, are the dominant contributors to its increased climate sensitivity (Virgin et al., 2021). ERA5 also displays a very dry troposphere in all dynamical regimes.

To unravel the anomalies of TCWV, the mean values of TCWV of each circulation regime observed during the whole comparison period (2004–2014) are employed as the reference to normalize the results. The normalized TCWV at each circulation interval for the different data records is displayed in Fig. 9. Although different patterns are observed, dry anomalies occurred in 2008 in all of the data records (except for ERA5, Du et al., 2021, where dry anomalies were observed for subsidence and weak upward motion regimes), and wet anomalies occurred in 2010. These extreme years are consistent with the previous findings (Sect. 4.1.3), and this second approach provides another angle of analysis on the assessment.

- The ESA TCWV-COMBI reveals a clear moistening tendency, especially over the subsiding branch of the atmospheric circulation after 2011. One of the major

causes of the turning point is the inclusion of MODIS data from 2011, which would increase the sampling size of the data and in turn affect the tendency.

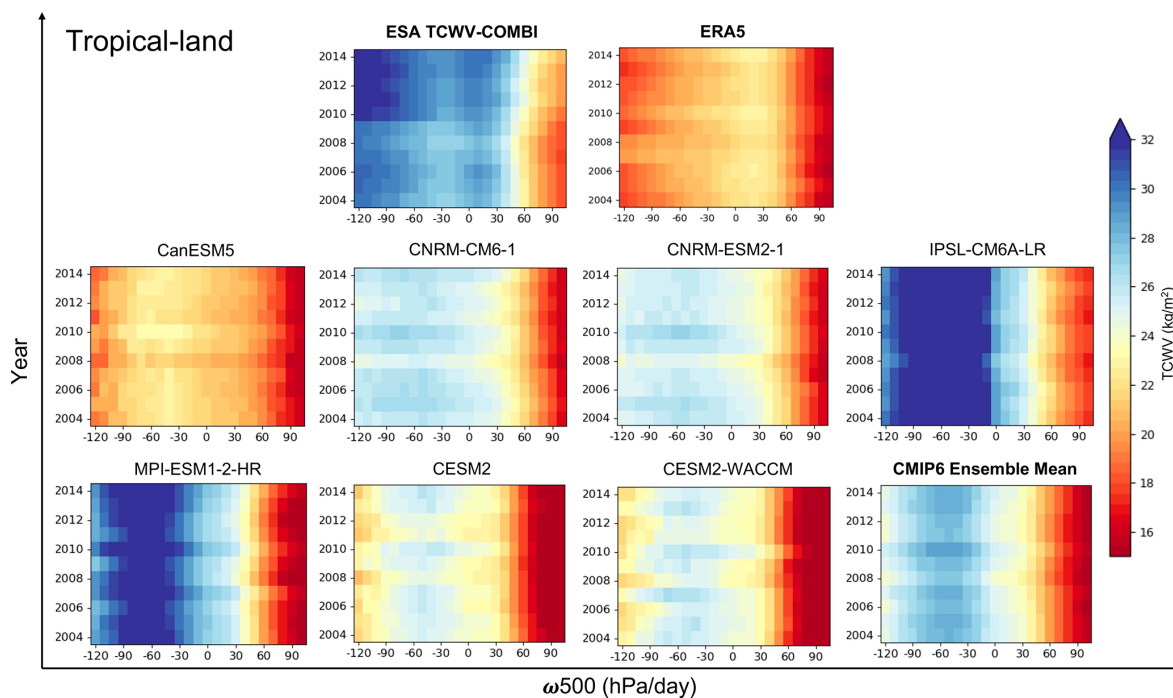
- The TCWV from ERA5 shows only extreme years but no distinct tendency. In the regions of highest upward motions, 2007 and 2008 appear moister than the other years, while 2004 and 2013 appear drier. Once again, this may be due to the scene selection applied over land, but this would not entirely explain the differences with the ESA TCWV-COMBI.
- Finally, the CMIP6 models and their ensemble mean show consistent interannual variabilities: dry anomalies occurred in 2008 in all the data records, and wet anomalies occurred in 2010. The results of the models are in a relative agreement with ERA5 and ESA TCWV-COMBI.

#### 4.2.3 Trends over oceans

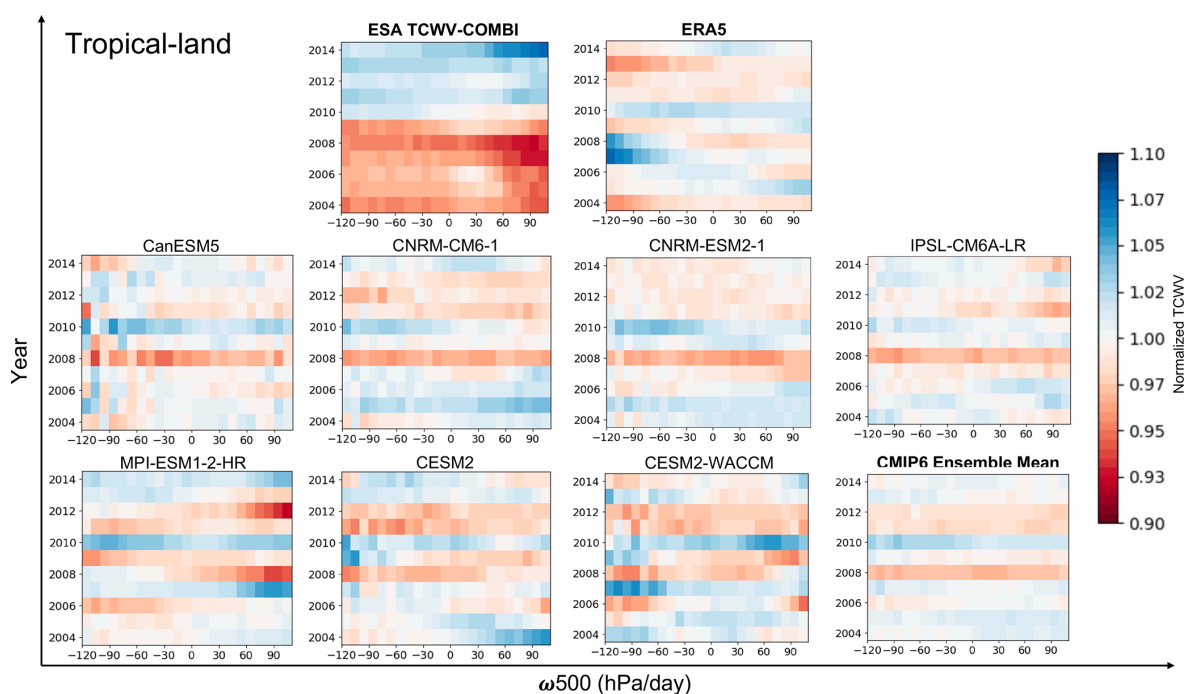
Oceanic situations were also discussed with respect to the large-scale circulation. The results are shown in Fig. 10 (mean TCWV) and Fig. 11 (normalized TCWV). As the data over ocean areas are obtained under all-weather conditions except for heavy precipitation, the impacts from clear-sky biases are greatly reduced. All data records, except for TCWV-COMBI, CanESM5 and IPSL-CM6A-LR, show that the strongest ascending zones correspond to the very humid regions. Different from the results over land, where the CMIP6 models showed strong differences (CanESM5 was the driest, and IPSL-CM6A-LR was the moistest), the amplitudes and gradients of moisture are closer to each other over oceans. Since the transport model in the boundary layer of IPSL-CM6A-LR affects both shallow and deep convection regimes, a compensational bias would be induced; thus a moist bias will be present in the weak upward motion regimes ( $\sim -30 \text{ hPa d}^{-1}$ ) (Boucher et al., 2020). The normalized TCWV with respect to dynamical intervals over ocean areas is shown in Fig. 11. The temporal evolutions of TCWV- $\omega 500$  are consistent with the earlier analysis based on the temporal evolution of the percentiles of TCWV over ocean (Sect. 4.1.3). The extremely dry and moist years (2008 and 2010 respectively) are the same between ESA TCWV-COMBI, ERA5 and the CMIP6 ensemble mean.

## 5 Conclusions

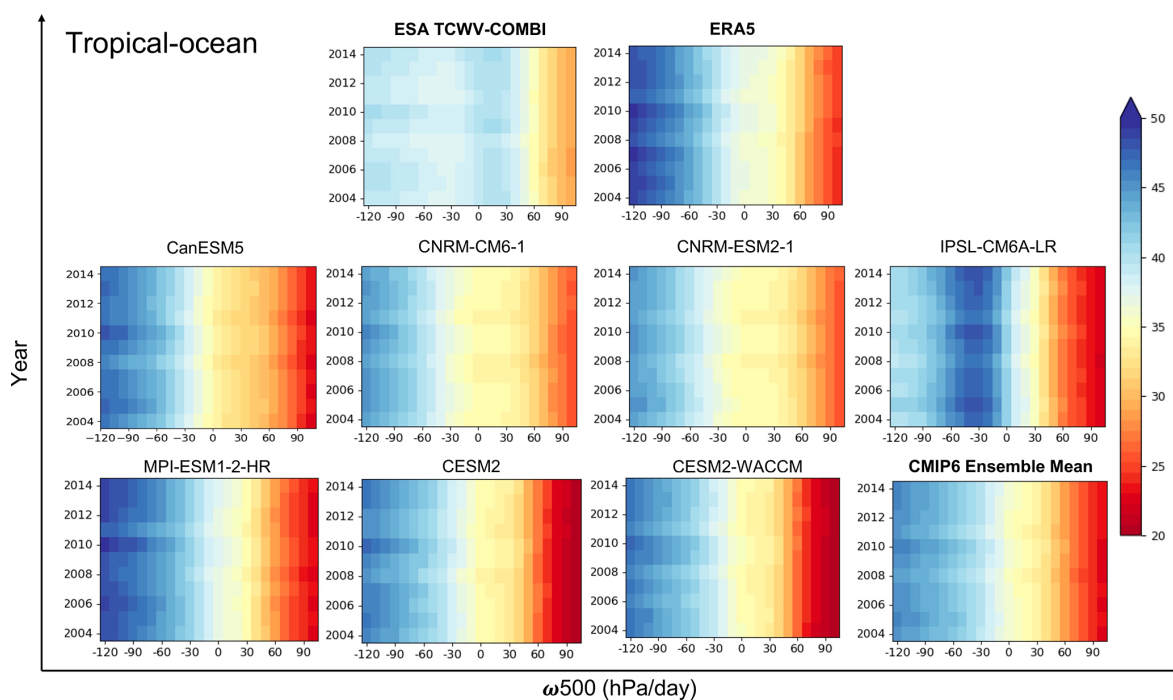
Despite the importance of water vapour in the study of climate variability, our ability to evaluate the water vapour feedback is constrained by its measurements at ranges of scales that are adapted for local, regional and global studies. This deficiency is attributable in part to the fact that it is difficult to quantitatively and accurately measure the distribution of water vapour. To work towards the requirement of GCOS



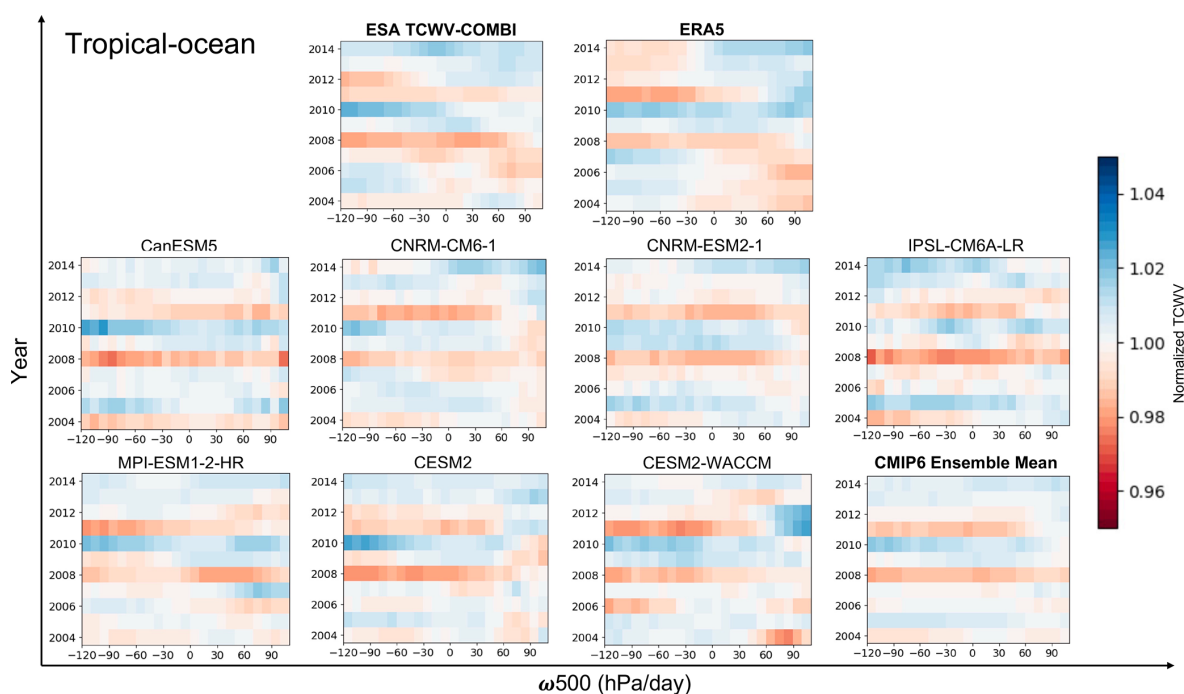
**Figure 8.** Mean of TCWV over tropical-land areas at each dynamical interval ( $\omega 500$ ) in  $10 \text{ hPa d}^{-1}$  computed from each data record.



**Figure 9.** Normalized TCWV with respect to the 2004–2014 mean over tropical-land areas at each dynamical interval ( $\omega 500$ ) in  $10 \text{ hPa d}^{-1}$  computed from each data record.



**Figure 10.** Mean of TCWV over tropical-ocean areas at each dynamical interval ( $\omega 500$ ) in  $10 \text{ hPa d}^{-1}$  computed from each data record.



**Figure 11.** Normalized TCWV with respect to the 2004–2014 mean over tropical-ocean areas at each dynamical interval ( $\omega 500$ ) in  $10 \text{ hPa d}^{-1}$  computed from each data record.



on satellite-based water vapour observation as an ECV, the ESA Climate Change Initiative Water Vapour project (ESA CCI\_WV) tackled this challenge by generating gridded products on stratospheric and tropospheric water vapour from multiple satellite observations suitable for climate and process studies.

We have conducted a comprehensive evaluation of the tropical water vapour (30° N–30° S) of seven GCMs (CMIP6 models, AMIP scenario) and ERA5, using the TCWV-COMBI climate data record developed within the ESA CCI\_WV project as a reference. The study focused over tropical-land and tropical-ocean areas at the daily frequency and over the 2003–2014 period. The variability of TCWV was analysed according to (i) its probability density function (PDF) defined at a yearly scale over the period and to (ii) the large-scale circulation using the atmospheric vertical velocity at 500 hPa ( $\omega$ 500) as a proxy for the tropospheric overturning circulation.

Different patterns of variability are observed among the various datasets, the largest discrepancies being noticed over land areas, while over the oceans, the datasets are closer to each other.

- Over land, the PDFs of the ESA TCWV-COMBI present a clear moistening trend of their driest percentiles, with a tipping point in 2011, probably associated with the addition of the MODIS observation in the climate data record. The projection of the TCWV onto regimes of  $\omega$ 500 shows the same behaviour; dry anomalies occurred in 2008 in all of the data records (except for ERA5, where dry anomalies were observed for subsidence and weak upward motion regimes), and wet anomalies occurred in 2010. Interestingly, the CMIP6 ensemble mean and the ERA5 reanalysis are in good agreement in terms of interannual anomalies, although the ERA5 TCWV is clearly too dry.
- Over ocean, the PDFs of all datasets present the same interannual variability. The extremely dry and moist years, associated with El Niño and La Niña events, are the same. This similarities hold when using the large-scale circulation as an evaluation tool, with the same transition between the dry/subsiding regimes and the moister/ascending regimes.

The results show that the ESA TCWV-COMBI data and ERA5 data vary within the ensemble spread of CMIP6 models, indicating that the mean models could correctly represent the evolution of water vapour with respect to large-scale circulation. The humid area is related to the ascending motion (negative value in  $\omega$ 500), and the dry area is related to the subsiding motion (positive value in  $\omega$ 500) over both tropical-land and tropical-ocean areas. There are discrepancies observed among the data records, probably caused by the lateral mixing, outflows from clouds and the precipitation efficiencies of the convective schemes. It is difficult to track the

reasons for the differences entirely; however, the differences and similarities can be explained by several factors.

1. The use of different satellites with different accuracies and resolutions within the ESA TCWV-COMBI may explain part of the moistening trend observed for this dataset over land.
2. The cloud masks applied to the GCMs and ERA5 and defined to mimic the cloud mask of the observation can also explain the differences.
3. The parametrization of the moisture fluxes at the surface and of convection, as well as the climate efficiency of the GCMs, also contributes to the observed differences.
4. The CMIP6 models under the AMIP scenario (with prescribed sea surface temperatures) and the scene selection that is much more conservative than over land explained the agreement between the ESA TCWV-COMBI, ERA5 and the GCMs over oceans.

It is really necessary to underline the role of the cloud mask in the assessment of water vapour fields in climate models using observations, even though it is easier to compare water vapour than clouds. The water vapour profiles (and sometimes the integrated values) from climate models usually have coarser spatial resolution than satellite observations. The satellite measurements, on the other hand, are often strictly restrained by cloud contamination. This clearly shows the importance of having access to the simulated water vapour (full profiles as well as integrated values) for the clear-sky part of the meshes of the climate models.

**Code availability.** Code is available from the corresponding author upon request (jia.he@latmos.ipsl.fr).

**Data availability.** The CMIP6 data are available from the ESGF system (<https://esgf-node.ipsl.upmc.fr/projects/esgf-ipsl/>, ESGF, 2022), and the CCI Water Vapour data are available from the project website (<https://climate.esa.int/en/projects/water-vapour/>, ESA, 2022).

**Author contributions.** JH carried out the data analysis and prepared all the figures. JH, HB and LP contributed to the interpretation of the results and wrote the paper.

**Competing interests.** The contact author has declared that none of the authors has any competing interests.

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