



# The spatiotemporal relationship between PM<sub>2.5</sub> and aerosol optical depth in China: influencing factors and implications for satellite PM<sub>2.5</sub> estimations using MAIAC aerosol optical depth

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**Abstract.** Satellite aerosol retrievals have been a popular alternative to monitoring the surface-based PM<sub>2.5</sub> concentration due to their extensive spatial and temporal coverage. Satellite-derived PM<sub>2.5</sub> estimations strongly rely on an accurate representation of the relationship between ground-level PM<sub>2.5</sub> and satellite aerosol optical depth (AOD). Due to the limitations of satellite AOD data, most studies have examined the relationship at a coarse resolution (i.e.,  $\geq 10$  km); thus, more effort is still needed to better understand the relationship between “in situ” PM<sub>2.5</sub> and AOD at finer spatial scales. While PM<sub>2.5</sub> and AOD could have obvious temporal variations, few studies have examined the diurnal variation in their relationship. Therefore, considerable uncertainty still exists in satellite-derived PM<sub>2.5</sub> estimations due to these research gaps. Taking advantage of the newly released fine-spatial-resolution satellite AOD data derived from the Multi-Angle Implementation of Atmospheric Correction (MAIAC) algorithm and real-time ground aerosol and PM<sub>2.5</sub> measurements, this study explicitly explored the relationship between PM<sub>2.5</sub> and AOD as well as its plausible impact factors, including meteorological parameters and topography, in mainland China during 2019, at various spatial and temporal scales. The coefficient of variation, the Pearson correlation coefficient and the slope of the linear regression model were used. Spatially, stronger correlations mainly occurred in northern and eastern China, and the linear slope was larger on average in northern inland regions than in other areas. Temporally, the PM<sub>2.5</sub>–AOD correlation peaked at noon and in the afternoon, and reached a maximum in winter. Simultaneously, considering relative humidity (RH) and the planetary boundary layer height (PBLH) in the relationship can improve the correlation, but the effect of RH and the PBLH on the correlation varied spatially and temporally with respect to both strength and direction. In addition, the largest correlation occurred at 400–600 m primarily in basin terrain such as the Sichuan Basin, the Shanxi–Shaanxi basins and the Junggar Basin. MAIAC 1 km AOD can better represent the ground-level fine particulate matter in most domains with exceptions, such as in very high terrain (i.e., Tibetan Plateau) and northern central China (i.e., Qinghai and Gansu). The findings of this study have useful implications for satellite-based PM<sub>2.5</sub> monitoring and will further inform the understanding of the aerosol variation and PM<sub>2.5</sub> pollution status of mainland China.

## 1 Introduction

Aerosols play an important role in regional and global climate change via both their direct and indirect effects (King et al., 1992; Li et al., 2007; Z. Liu et al., 2019, 2020; Liu et al., 2018). Fine particulate matter (PM<sub>2.5</sub>) has attracted a lot of concern from scientists, policymakers and the public due to its negative impact on the environment (Li et al., 2020; Y. Yang et al., 2018, 2020; Yim et al., 2019a) and human health (Schwartz, 1996; Pope III et al., 2002; Gu et al., 2018, 2020; Gu and Yim, 2016; Hou et al., 2019; Shi et al., 2020; Yim et al., 2015). With an extended spatial and temporal coverage, retrieval of the near-surface PM<sub>2.5</sub> concentration from satellite aerosol optical depth (AOD) has become a popular approach to bridging the gap left by the ground-level monitoring network. The use of satellite AOD retrievals facilitates the detection of the large-scale and long-term aerosol loading as well as transboundary transport and also assists in the determination of the population exposure level for epidemiological and health studies (Liu et al., 2017; Zou et al., 2019). The theoretical basis of this technique rests on the strong link between satellite AOD retrievals and ground-level PM<sub>2.5</sub> concentration (Wang and Christopher, 2003). Therefore, it is of great importance to examine the PM<sub>2.5</sub>–AOD relationship, not only with respect to the relationship itself but also regarding the underlying physical understanding involved.

Due to rapid urbanization and industrialization, China has been suffering from serious air pollution (Zhao et al., 2019; Guo et al., 2011; He et al., 2019). Given the heavy haze pollution in China (Luo et al., 2018; Tong et al., 2018b, a; Y. Yang et al., 2019; Yim, 2020; Yim et al., 2019b), observation studies have demonstrated and investigated the PM<sub>2.5</sub>–AOD relationship in China (Wang et al., 2014; Guo et al., 2017, 2009; Kong et al., 2016; Shao et al., 2017; Xin et al., 2016; Q. Yang et al., 2019). For instance, the Moderate Resolution Imaging Spectroradiometer (MODIS) Dark Target (DT) and Deep Blue (DB) 10 km AOD products have been widely used for quantifying the PM<sub>2.5</sub>–AOD relationship at both regional (Kong et al., 2016; Shao et al., 2017) and national scales (Guo et al., 2017; Q. Yang et al., 2019). The relationship found is not always concordant and exhibits large spatiotemporal variability. With the development of satellite remote sensing techniques, a high-spatial-resolution global daily aerosol product – the Multi-Angle Implementation of Atmospheric Correction (MAIAC) 1 km AOD product – has recently been released, providing local-scale aerosol gradients and details missed by previous coarse-resolution (10 km) products (Lyapustin et al., 2018). Because the ground-level PM<sub>2.5</sub> concentration is generally measured at an in situ point, whereas the satellite AOD is reported on a grid (e.g., 1 km × 1 km for MAIAC AOD and 10 km × 10 km for DT AOD), the relationship between a “point” PM<sub>2.5</sub> and a “gridded” AOD at a coarse resolu-

tion might mask their discrepancy or correspondence at a finer spatial resolution. Due to the characterization of fine-resolution aerosol heterogeneity as well as the changes in the retrieval scheme and accuracy (Guo et al., 2017), the use of the high-resolution AOD data probably introduces different/additional uncertainties into the PM<sub>2.5</sub>–AOD association, leading to a different spatiotemporal pattern and further affecting the performance of satellite-derived PM<sub>2.5</sub> predictions. Thus, further exploring the relationship between near-surface PM<sub>2.5</sub> and MAIAC AOD over China seems critically imperative based on AOD products at a much finer resolution.

The PM<sub>2.5</sub>–AOD link has been widely recognized to vary in space and time, which is associated with the defined difference between surface PM<sub>2.5</sub> measurements and satellite AOD stemming from variations in aerosol type, meteorological conditions and topography (Kloog et al., 2014; Lee et al., 2016a; Ma et al., 2014; Wang et al., 2014; Zhang and Li, 2015). Ground-level air quality monitoring equipment contains a dynamic heating system to evaporate water vapor in particles; thus, it measures the dry mass concentration for the PM<sub>2.5</sub> concentration, which is inevitably different from satellite AOD that interprets aerosol extinction without considering the humidity conditions. The vertical structure of aerosols can induce uncertainty between PM<sub>2.5</sub> and AOD, as satellite AOD is an integration of aerosol extinction in the entire atmosphere, whereas the ambient PM<sub>2.5</sub> concentration is only measured at the ground surface. Irrespective of the different conditions, meteorological factors may have a confounding influence, affecting the PM<sub>2.5</sub>–AOD relationship to different extents and even leading to contrary results (i.e., improvement or deterioration). In addition, topography is another major contributor to uncertainty in the PM<sub>2.5</sub>–AOD relationship, but its impact is difficult to quantify due to its complex structure. Therefore, further explicit analyses are required to explore the effect of both terrain and meteorological factors, including hygroscopicity and the vertical structure of aerosols on the PM<sub>2.5</sub>–AOD relationship, especially for mainland China, which has complex terrain and land cover types.

Capturing the diurnal cycle of the surface PM<sub>2.5</sub> concentration has various applications, e.g., establishing the short-term health impact of PM<sub>2.5</sub> as well as aerosol–cloud interactions (Arola et al., 2013; Guo et al., 2016; Kim et al., 2010; Lee et al., 2016b; Zheng et al., 2020). To successfully capture the diurnal variation in the PM<sub>2.5</sub>–AOD relationship and represent the diurnal cycle of PM<sub>2.5</sub>, understanding the diurnal variation involved in the data is crucial. However, this is easily ignored when studying the relationship between surface PM<sub>2.5</sub> and satellite AOD due to the limitations of hourly satellite aerosol observations and the difficulty involved in making hourly PM<sub>2.5</sub>–AOD co-locations. To the best of our

knowledge, few studies have considered the diurnal cycle in the data when investigating the PM<sub>2.5</sub>–AOD relationship.

Therefore, this study aimed at comprehensively investigating the spatiotemporal relationship between ground-level PM<sub>2.5</sub> and satellite AOD in mainland China using the MAIAC 1 km aerosol product with multi-scale analyses in space and time as well as an emphasis on the difference in the spatial variation and temporal trend against previous relationships dependent on coarse-resolution data. We carried out correlation analyses, measured by the Pearson correlation coefficient and the slope of the linear regression, to explore the PM<sub>2.5</sub>–AOD relationship at 1500 monitoring stations, in more than 360 cities and 11 urban agglomerations, by taking the latent impact of the diurnal variation in aerosol particles, relative humidity (RH), the planetary boundary layer height (PBLH) and elevation into account. In addition to monthly variation, seasonality and the annual pattern, real-time ground aerosol observations were also employed in this study to uncover the diurnal variability and its impact on the PM<sub>2.5</sub>–AOD relationship using hourly comparisons during a day when Terra and Aqua data were available. The remainder of this study is organized as follows: descriptions of the data used, including the ground-level PM<sub>2.5</sub> measurements, the MAIAC AOD, RH, the PBLH and elevation in China, are given in Sect. 2. The results of the correlation and the linear slope analyses, with respect to the PM<sub>2.5</sub>–AOD relationship and its impact factors, are presented in Sect. 3 and are discussed in Sect. 4. The major conclusions are summarized in Sect. 5.

## 2 Materials and methods

### 2.1 Study area and urban agglomerations

The study area is mainland China (excluding Taiwan, Hong Kong and Macau), covering more than 360 cities and 11 urban agglomerations. The 11 urban agglomerations include five national urban agglomerations, i.e., Beijing–Tianjin–Hebei (BTH), Yangtze River Delta (YRD), Pearl River Delta (PRD), Triangle of Central China (TCC) and Chengdu–Chongqing (CDCQ), and six regional urban agglomerations, i.e., central and southern Liaoning (CSL), Shandong Peninsula (SDP), Harbin–Changchun (HC), Central Plain (CP), Guanzhong Plain (GZP) and northern Tianshan (NTS). According to the “National New-type Urbanization Plan (2014–2020)” ([http://www.gov.cn/zhengce/2014-03/16/content\\_2640075.htm](http://www.gov.cn/zhengce/2014-03/16/content_2640075.htm); last access: 10 November 2021), these agglomerations are densely populated and have relatively well-developed urban systems. The locations of these agglomerations and the associated cities are detailed in Fig. S1 in the Supplement.

### 2.2 Data collection and preprocessing

Hourly ground-level PM<sub>2.5</sub> measurements from ~1500 monitoring stations for 2019 were collected from the official website of the Ministry of Ecology and Environment of China (<https://www.mee.gov.cn>, last access: 10 November 2021). These observed PM<sub>2.5</sub> mass concentrations are measured using the tapered-element oscillating microbalance or beta-attenuation method with quality control (National Ambient Air Quality Standards, GB3095-2012; [http://english.mee.gov.cn/Resources/standards/Air\\_Environment/quality\\_standard1/201605/t20160511\\_337502.shtml](http://english.mee.gov.cn/Resources/standards/Air_Environment/quality_standard1/201605/t20160511_337502.shtml), last access: 10 November 2021). Hourly concentrations < 1 µg m<sup>−3</sup> were discarded, as they are below the monitors’ measurement limit. Daily PM<sub>2.5</sub> values were averaged from hourly concentrations for the period from 00:00 to 23:00 LT (local time), and days with at least 75 % valid hourly observations were kept for statistical analyses.

The MODIS 1 km AOD data archived in Collection 6, derived using the MAIAC algorithm over China in 2019, were used in this study. This new aerosol dataset not only represents aerosol retrievals at a high spatial resolution (Chudnovsky et al., 2013; Lyapustin et al., 2018) but also achieves a comparable accuracy to the previous popular coarse-resolution AOD product (Zhang et al., 2019; N. Liu et al., 2019; Martins et al., 2017), making it possible to detect local aerosol variability. Both Terra (overpass time of 10:30 LT) and Aqua (crossover time of 13:30 LT) MAIAC AOD data were downloaded from the NASA Level-1 and Atmosphere Archive and Distribution System (LAADS) Distributed Active Archive Center (DAAC). To investigate the aerosol diurnal variation, ground-level AOD observations from the Version 3.0 Aerosol Robotic Network (AERONET) dataset with assured quality (Level 2.0) were obtained.

Meteorological variables including the RH, PBLH and terrain were used for impact factor analysis. RH data were obtained from the China Meteorological Data Service Center (<http://data.cma.cn/>, last access: 10 November 2021), and PBLH data were downloaded from the reanalysis dataset provided by the European Centre for Medium-Range Weather Forecasts (ECMWF, <https://www.ecmwf.int/>, last access: 10 November 2021). The Advanced Spaceborne Thermal Emission and Reflection Radiometer (ASTER) global digital elevation model (GDEM) was used to measure the terrain in terms of elevation.

### 2.3 Data co-location and analytical methods

According to previous studies, such as N. Liu et al. (2019), the MAIAC AOD data also confront the missing data problem due to their coarse spatial resolution. To tackle this issue and improve the number of daily co-locations with ground PM<sub>2.5</sub> measurements, an imputation method was conducted to fill the missing AOD values for one satellite when the other was available, and the improved AOD data served as

the daily mean AOD for the co-location with daily PM<sub>2.5</sub> averages. The imputation method was similar to our previous studies (He et al., 2020, 2019): a relationship was constructed between Terra and Aqua retrievals for each observation day, and the obtained coefficients were applied to predict the missing values. In addition, hourly co-location was also accomplished to support the hourly analyses in this study, which directly co-located the hourly PM<sub>2.5</sub> measurements at each observational site with the Aqua/Terra AOD retrievals within the hour of the satellite's overpass. All of the daily and hourly co-locations required that the ground-level monitors spatially fell into the 1 km AOD grid cell. That is, we used the 1 km AOD grid to match the in situ PM<sub>2.5</sub> measurements and averaged those PM<sub>2.5</sub> observations from multiple monitors located within a 1 km grid cell.

The coefficient of variation (CV), defined as the ratio of the standard deviation to the mean, which is used to measure the dispersion of data points around the mean, allows us to compare the degree of difference between ground PM<sub>2.5</sub> and columnar AOD. The relationship between ground PM<sub>2.5</sub> and satellite AOD was quantified using the Pearson correlation coefficient ( $r$ ). The linear regression model was applied using AOD as an independent variable and PM<sub>2.5</sub> a dependent variable. The estimated slopes of the linear models were used for the statistical analysis.

To examine the impact of hygroscopicity and the vertical structure of aerosols on the PM<sub>2.5</sub>–AOD relationship, an RH correction and a vertical correction were applied to obtain respective RH-corrected PM<sub>2.5</sub> and PBLH-corrected AOD for the subsequent correlation analyses. Following previous studies (Kong et al., 2016; Wang et al., 2014), the respective RH-corrected PM<sub>2.5</sub> and PBLH-corrected AOD were calculated as follows:

$$\text{corrected PM}_{2.5} = \text{PM}_{2.5} \left( 1 - \frac{\text{RH}}{100} \right), \quad (1)$$

$$\text{corrected AOD} = \frac{\text{AOD}}{\text{PBLH}}, \quad (2)$$

where corrected PM<sub>2.5</sub> (μg m<sup>-3</sup>) represents PM<sub>2.5</sub> corrected by relative humidity (PM<sub>2.5</sub> is the daily mean PM<sub>2.5</sub> concentration measured by the ground-level air quality network), RH stands for relative humidity (ranging from 0 to 100), and corrected AOD (m<sup>-1</sup>) represents AOD corrected by the PBLH (planetary boundary layer height, in meters).

### 3 Results

#### 3.1 Characteristics of the variation in surface PM<sub>2.5</sub> and satellite AOD

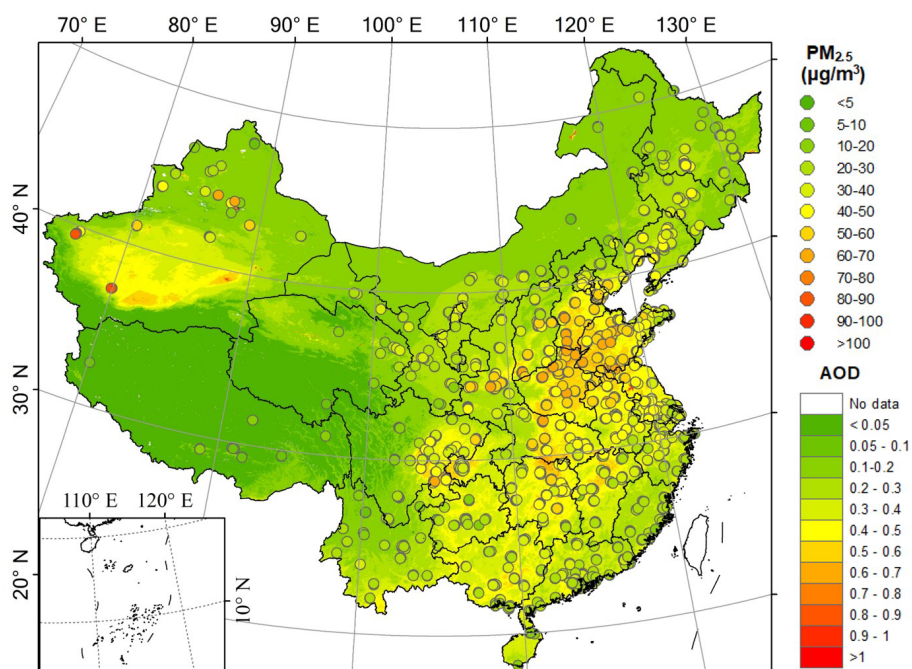
Figure 1 illustrates the annual mean PM<sub>2.5</sub> from in situ data and AOD from all of the satellite-observed grid cells throughout mainland China in 2019. The typical hot spot polluted by fine particles could be identified in the North China

Plain (NCP), in particular in the southeastern Hebei, western Shandong and northern Henan provinces, where the peak PM<sub>2.5</sub> concentration can reach 55 μg m<sup>-3</sup> or even higher. It is noted that southern and western Xinjiang, the margin of the Taklimakan Desert, had a very high annual concentration exceeding 75 μg m<sup>-3</sup>, and this area is highlighted as a PM<sub>2.5</sub> pollution center. In general, the MAIAC AOD shared a similar spatial pattern to PM<sub>2.5</sub>. That is, areas with high AOD values congruously correspond to those with high PM<sub>2.5</sub> concentrations except for the typical mismatch in southwestern China. For example, higher mean AOD values of 0.5–0.8 are found in the Chengdu–Chongqing (CDCQ) agglomeration, which had AOD values as high as those in the NCP, contrary to the relatively low mean PM<sub>2.5</sub> concentration of 38.93 μg m<sup>-3</sup> (mean concentration for the NCP was 49.32 μg m<sup>-3</sup>).

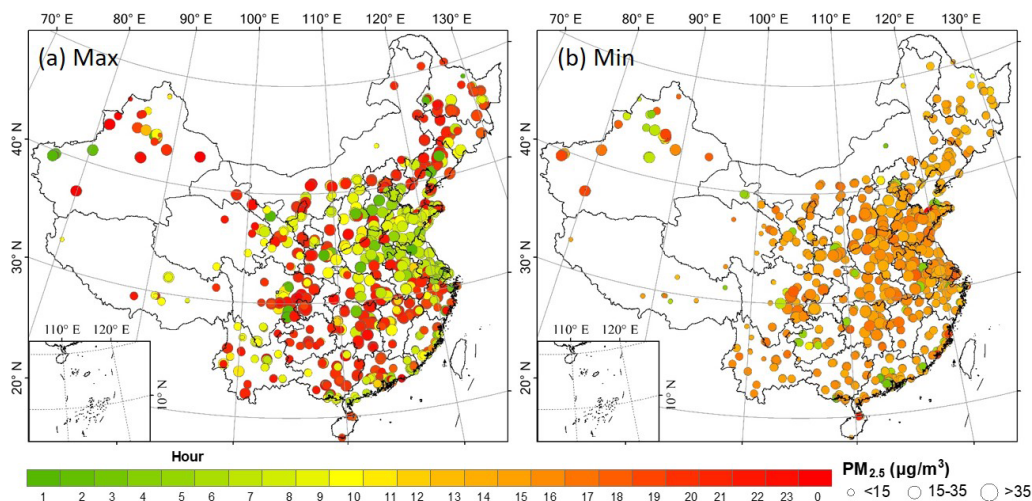
Figure S3 in the Supplement shows the temporal profile of PM<sub>2.5</sub> and AOD for the entire country in 2019. The national PM<sub>2.5</sub> mean was 38.53 μg m<sup>-3</sup>, and the annual mean AOD was 0.22. The monthly changes in the two indicators show a different temporal pattern. The monthly PM<sub>2.5</sub> concentration shows a U-shaped trend, which remained low and stable between June and August, and the monthly concentration in the summer months was almost a third of those in winter months. Compared with winter, summer is generally associated with lower anthropogenic air pollutant emissions (e.g., reduced coal and biomass burning for heating during summer) and favorable meteorological conditions for pollution dispersion (e.g., enhanced convection and more precipitation during summer) (Zhang and Cao, 2015), which may account for the summertime minimum. This result is consistent with the findings reported in previous research (X. Yang et al., 2018). In contrast, the monthly AOD variation shows a rough inverse U-shaped curve with a plateau between March and August. The amplitude of monthly AOD variation was also less than that of PM<sub>2.5</sub>: the maximum AOD level of 0.28 occurred in March, which was 0.1 (~35.42%) higher than the AOD level in the lowest month (November). This further demonstrates the difference between ground-level PM<sub>2.5</sub> and satellite AOD observations in Fig. 1.

The diurnal characteristics of PM<sub>2.5</sub> variations in China in 2019 were investigated. In general, the national maximum PM<sub>2.5</sub> diurnal mean concentration occurred in the morning and evening, whereas the minimum mean values appeared in afternoon (Fig. S4 in the Supplement), which is generally consistent with the diurnal pattern revealed in previous studies (Guo et al., 2017). The two peaks that appeared in the overall diurnal PM<sub>2.5</sub> curves were linked to PM<sub>2.5</sub> spatial variation. Figure 2 shows the spatial distribution of the hours when the maximum and minimum PM<sub>2.5</sub> concentration occurred as well as the corresponding concentration value. Among the 1494 monitoring stations, the largest diurnal PM<sub>2.5</sub> concentration appeared in the evening (19:00–23:00 LT) at 638 sites (~42.70%), 567 sites (~37.95%) displayed a maximum concentration in morning





**Figure 1.** The spatial distribution of annual mean ground-level PM<sub>2.5</sub> and satellite AOD averages in China in 2019. The mean value of the ground-level PM<sub>2.5</sub> concentration was calculated from all measured days, covering almost the whole study period, whereas the mean AOD value was calculated using fewer days because the MAIAC algorithm could not make retrievals over areas with high cloud cover.



**Figure 2.** Spatial distribution of diurnal PM<sub>2.5</sub> averaged over the 2019 period according to the (a) maximum and (b) minimum mean PM<sub>2.5</sub> concentration over 24 h. The solid color and the size of the circle denote the hour with the maximum and minimum diurnal mean concentration and the corresponding concentration value over each monitoring station, respectively.

(07:00–11:00 LT), 50 sites ( $\sim 3.35\%$ ) reached a maximum in the afternoon (12:00–18:00 LT) and the remaining 239 sites ( $\sim 16.00\%$ ) had their maximum after midnight (00:00–06:00 LT). To be more specific, significant spatial variability was found in the highest diurnal PM<sub>2.5</sub> concentration. The morning PM<sub>2.5</sub> peak generally occurred in the NCP, whereas the evening maximum prevailed in southern and northeastern China where an afternoon PM<sub>2.5</sub> maximum also sporadically

appeared. Furthermore, the highest diurnal mean concentration in the morning peak (mean of  $46.60\ \mu\text{g m}^{-3}$ ) was generally greater than that in the evening (mean of  $42.24\ \mu\text{g m}^{-3}$ ). Among the 413 sites in the NCP, there were more than 94 % sites with concentrations exceeding  $35\ \mu\text{g m}^{-3}$ , pinpointing the serious PM<sub>2.5</sub>-pollution risk in the NCP region. We note that there was almost no spatial difference in the minimum

hourly mean concentration, which uniformly appeared between 14:00 and 17:00 LT.

The CV analysis was conducted based on temporally averaged PM<sub>2.5</sub> and AOD values to evaluate the spatial variation of aerosol particles over a given region. As shown in Fig. 3a, the ambient PM<sub>2.5</sub> concentration in 2019 over the CP region was the highest (multiyear mean value of 57.69  $\mu\text{g m}^{-3}$ ) and was associated with moderate spatial variability (CV of 0.79), suggesting that the CP region was the most polluted of the 11 urban agglomerations and that the majority of cities within this area were exposed to very dangerous PM<sub>2.5</sub> levels. In contrast, with the lowest PM<sub>2.5</sub> concentration (27.43  $\mu\text{g m}^{-3}$ ) and the smallest CV value (0.58), almost all PRD cities experienced relatively clean ambient air. The PM<sub>2.5</sub> concentration in the NTS urban agglomeration remained at a moderate level (45.96  $\mu\text{g m}^{-3}$ ), but the CV value was the largest (CV = 1.25), implying a distinct contrast in air quality between cities in the region. Figure 3 also presents a different regional pattern in both mean AOD and CV values.

### 3.2 Spatiotemporal variability in the PM<sub>2.5</sub>–AOD relationship

Figure 4 presents the spatial distribution of the correlation coefficients (Fig. 4a–e) and linear trend (Fig. 4f–j) between the daily mean PM<sub>2.5</sub> concentration and AOD for the entire study period and for the four respective seasons over all monitoring sites. With correlation coefficient ( $r$ ) values ranging from  $-0.5$  to  $-0.18$ , abnormal negative associations between PM<sub>2.5</sub> and AOD were observed in the Tibetan Plateau (TP), where the altitudes are almost higher than 4000 m. In addition to the TP, the correlation coefficient for each monitoring site ranged from 0.15 to 0.84. In general, the stronger correlations ( $r > 0.4$ ) were mostly concentrated in northern and eastern China, whereas  $r$  values for most observational sites dropped below 0.4 in southern areas.

Particularly, the Sichuan Basin (SCB) was highlighted as showing the strongest correlation ( $r > 0.6$ ) in Fig. 4a, whereas very low correlations ( $r < 0.3$ ) were found in northern Xinjiang and northern central China. Figure 4b–e exhibits a distinct seasonal variability in the PM<sub>2.5</sub>–AOD relationship, even though the spatial pattern in each season is generally similar to that shown by all data pairs in Fig. 4a. The overall correlation coefficient in winter was relatively higher, with an overall  $r$  value of 0.61, which demonstrates that the light extinction in the atmosphere in wintertime may be mainly caused by fine particles near the surface, followed by spring (0.57) and then autumn (0.55). The  $r$  value was the lowest in summer (0.47). The distribution of correlations along with various PM<sub>2.5</sub> levels may account for the seasonal difference. As shown in Fig. S5 in the Supplement, correlations between PM<sub>2.5</sub> and AOD were the strongest at extremely high PM<sub>2.5</sub> concentration levels; coincidentally, higher PM<sub>2.5</sub> concentrations typically occurred in winter,

leading to a better PM<sub>2.5</sub>–AOD relationship in this season. Furthermore, the seasonal variation was probably related to the lower RH in winter and the unstable atmospheric conditions in summer.

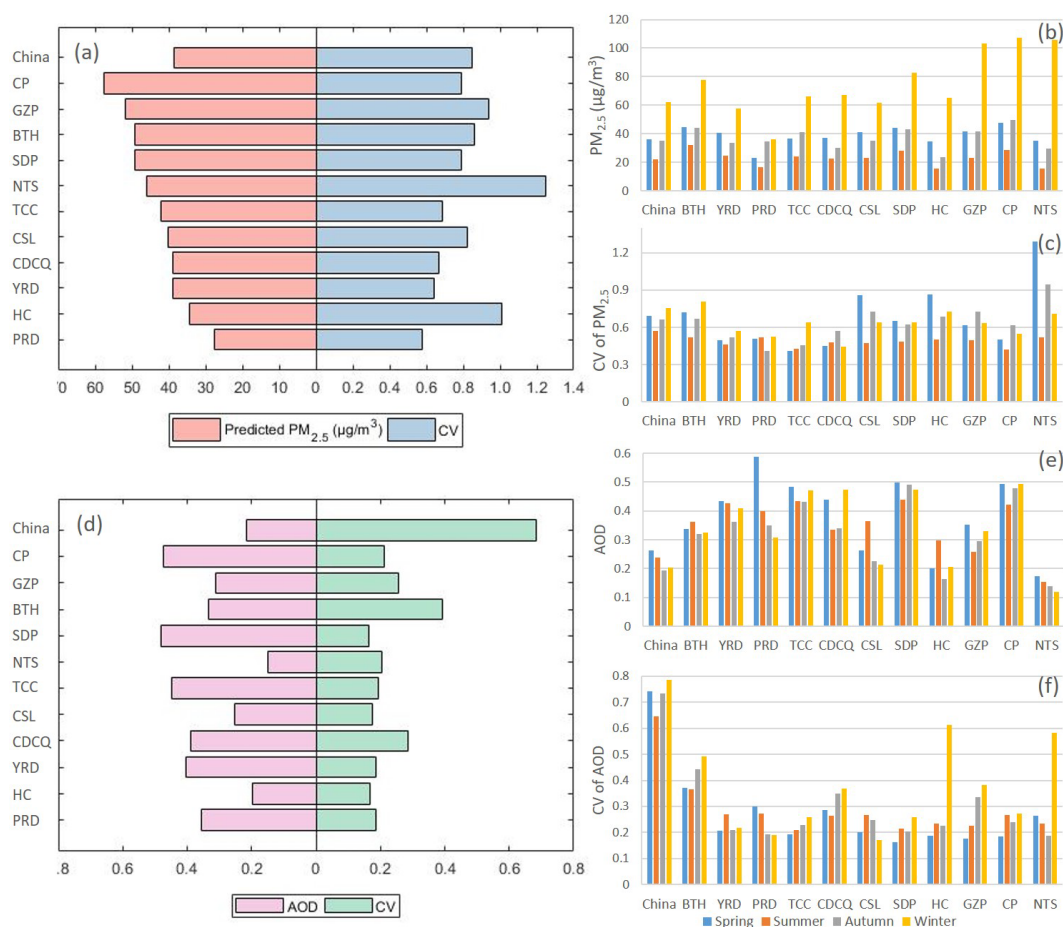
Figure 4f–j shows the spatial distribution of the slopes ( $S$ ) of the linear regressions between daily PM<sub>2.5</sub> and AOD for each monitoring site. Corresponding to the correlation analysis in Fig. 4a–e, the slope values in the TP were also negative; for the other areas over mainland China, the slopes of linear regressions ranged from 9.37 to 158.45 over the entire observation period (Fig. 4f) with significant spatial and seasonal variability (Fig. 4g–j).

To better represent the spatiotemporal features, correlation analyses between ground-level PM<sub>2.5</sub> and satellite AOD were conducted for all of the monitoring stations (Fig. 4) as well as at the urban-agglomeration scale (Fig. 5). In general, the PM<sub>2.5</sub>–AOD relationships at the regional scale mirrored those at the observational-site scale. In terms of the correlation coefficient, higher values were found in urban agglomerations located in northern and eastern China, such as the GZP, BTH and CDCQ agglomerations which had  $r$  values greater than 0.50; the weakest relationship ( $r$  of 0.26) occurred in the NTS agglomeration. As for the slope of the regression function, the maximum slope (81.10) was observed in the GZP region, followed by the CSL (51.48) and BTH (51.37) agglomerations; the lowest value was observed in the PRD urban agglomeration with a slope of 23.00.

Figure S7 in the Supplement demonstrates that the spatial distribution of the PM<sub>2.5</sub>–AOD correlations during the study period using Terra and Aqua AOD matched with PM<sub>2.5</sub> measurements at the respective overpass hour. It is intriguing that correlations for almost all monitoring sites at 14:00 LT were greater than those at 11:00 LT, and the overall correlation at 14:00 LT was 0.54, which was also obviously higher than that of 0.47 at 11:00 LT. To determine whether the difference is caused by aerosol retrieval accuracy from Terra (i.e., at 11:00 LT) and Aqua (i.e., at 14:00 LT), we conducted a comparison analysis using hourly ground PM<sub>2.5</sub> and the spatiotemporally co-located AERONET AOD observations. Figure S8 in the Supplement indicates that the diurnal correlations between ambient PM<sub>2.5</sub> and AERONET AOD were also larger in afternoon.

### 3.3 Influence of meteorology on the PM<sub>2.5</sub>–AOD correlation

To examine the spatiotemporal effect of the RH and PBLH on the daily PM<sub>2.5</sub>–AOD relationship, we conducted a correlation analyses using RH-corrected PM<sub>2.5</sub> data and PBLH-corrected AOD data for monitoring sites (Fig. 6) and for the 11 agglomerations (Table 1). The details of the correction method for PM<sub>2.5</sub> and AOD are provided in Sect. 2.3. In general, after correction by the RH and PBLH (Fig. 6), the mean correlation coefficient ( $r$ ) reached 0.71, and the  $r$  values at  $\sim 16\%$  of sites increased to  $> 0.80$ , which pre-



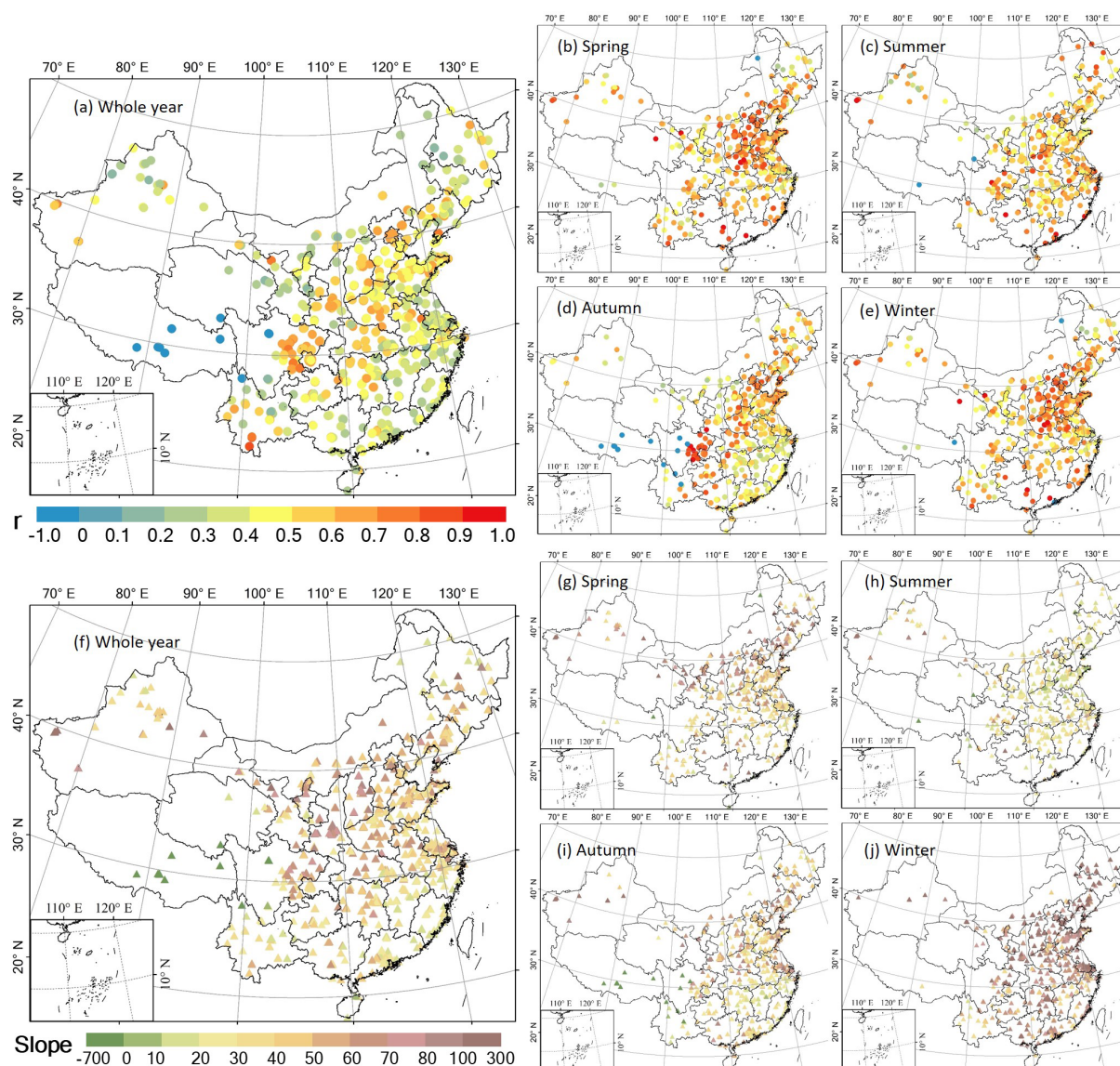
**Figure 3.** Regional mean values and coefficient of variation (CV) values for PM<sub>2.5</sub> and AOD in 2019. Panel (a) shows the mean PM<sub>2.5</sub> and its CV for 11 urban agglomerations for the whole study period, and panels (b) and (c) show the mean PM<sub>2.5</sub> and the CV values for each respective season. Panel (d) shows the mean AOD and its CV in regions for the whole year, and panels (e) and (f) show the mean AOD and the CV values in each respective season.

sented a notable improvement in the PM<sub>2.5</sub>–AOD relationship compared with those based on uncorrected data (mean  $r = 0.44$  and 0.25 % of sites with  $r > 0.80$ ). For individual observational sites, ~96 % of monitoring sites had enhanced PM<sub>2.5</sub>–AOD correlations after both corrections, increasing by 0.01–0.61. The correlation after RH correction slightly increased by an average of ~0.07, and the vertical correction achieved a better improvement in the PM<sub>2.5</sub>–AOD relationship than the former, which contributed to the average  $r$  increase of ~0.18. Specifically, ~22 % of monitoring sites, primarily located in the Guangdong, Guangxi, Yunnan and Heilongjiang provinces, had larger  $r$  values after RH correction than vertical correction. These results indicate that the spatial difference in the influence of the RH and vertical correction on the relationship between PM<sub>2.5</sub> and AOD cannot be ignored due to the significant impact of these parameters on the relationship at most monitoring sites. It should be noted that there was an obvious negative increment after

both corrections, in particular at most monitoring stations in Yunnan Province.

There was also significant seasonal discrepancy with respect the improvement of the PM<sub>2.5</sub>–AOD relationship due to meteorological effects. As shown in the seasonal patterns in Fig. 6, after both corrections were carried out, the correlations had the largest improvement in autumn (the mean  $r$  value increased by 0.21), followed by spring (0.11) and winter (0.10), and summer had the smallest improvement (an increase of 0.09). Specifically, in spring and autumn, both the RH and vertical corrections imposed a positive effect on the PM<sub>2.5</sub>–AOD correlation, although the strength of the effect was different. In spring, the RH correction had a greater contribution to improving the PM<sub>2.5</sub>–AOD correlation than the vertical correction, and the  $r$  value based on RH-corrected data was 0.02 larger at an average level. In autumn, the vertical correction exerted a more important influence than the RH correction, with a mean  $r$  value of 0.68 vs. 0.65. It should be noted that the correlation revised by the PBLH was lower





**Figure 4.** Spatial distributions of the relationship between ambient PM<sub>2.5</sub> and satellite AOD pairs over China in 2019. Panel (a) shows the correlation coefficient for the whole year, and panels (b)–(e) show the correlation coefficients for each respective season. Panels (f)–(j) are based on the slope values of the linear regression model. Please note that only statistically significant results, i.e.,  $p$  level < 0.05, are presented here.

than the original condition in summer, with the mean  $r$  value decreasing by 0.05, especially for monitoring stations in central China, such as the Jiangxi, Shaanxi and Henan provinces, where significant decreases in  $r$  values (changes of < −0.15) occurred.

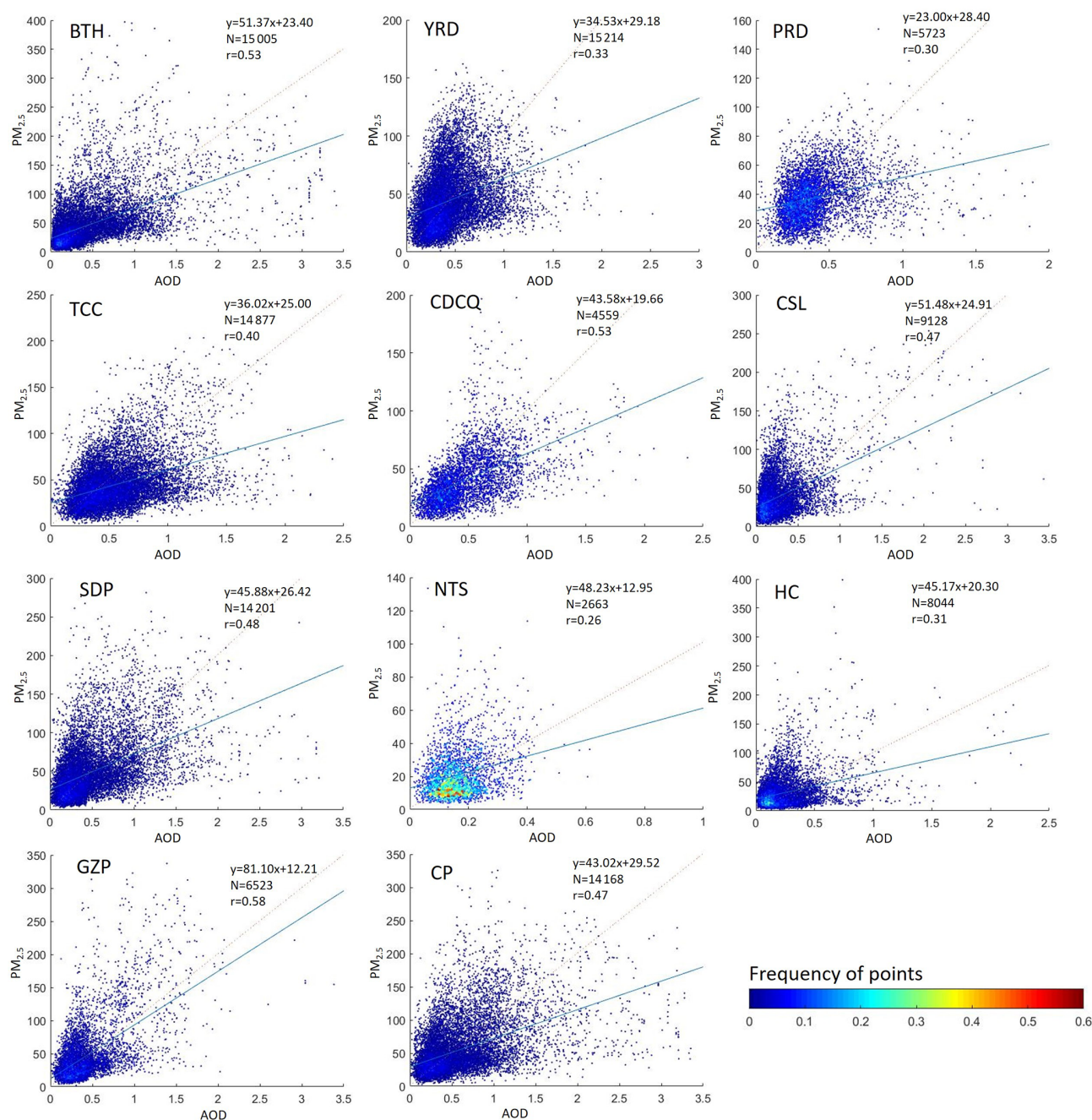
Table 1 illustrates the statistics of the PM<sub>2.5</sub>–AOD relationship based on both corrected and uncorrected data over the 11 urban agglomerations. It follows that there are also obvious spatial and seasonal discrepancies in the influence of humidity and the vertical distribution on the regional PM<sub>2.5</sub>–AOD association. In summer, the correlation coefficient ( $r$ ) of AOD, vertically revised by the PBLH, and ground-based

PM<sub>2.5</sub> showed a significant decrease over the original relationship, except for in the NTS and YRD regions. Despite a generally weak relationship in the NTS region, PBLH-corrected AOD showed a superior relationship with surface PM<sub>2.5</sub> observations compared with the original correlation, not only in summer (0.29 vs. 0.21) but also in spring (0.42 vs. 0.37) and autumn (0.45 vs. 0.23).

### 3.4 Impact of terrain on the PM<sub>2.5</sub>–AOD correlation

The effect of topography on the PM<sub>2.5</sub>–AOD correlation analyses was investigated by separating the co-located samples into 21 elevation bins with an interval of 200 m. As





**Figure 5.** Scatterplots for the paired PM<sub>2.5</sub> concentration (µg m<sup>-3</sup>) and satellite AOD values over the 11 urban agglomerations in China including Beijing–Tianjin–Hebei (BTH), Yangtze River Delta (YRD), Pearl River Delta (PRD), Triangle of Central China (TCC), Chengdu–Chongqing (CDCQ), central and southern Liaoning (CSL), Shandong Peninsula (SDP), Harbin–Changchun (HC), Central Plain (CP), Guanzhong Plain (GZP) and northern Tianshan (NTS). The solid line represents the linear trend.

shown in Fig. 7, a distinct trend and pattern can be identified between terrain and the PM<sub>2.5</sub>–AOD relationship. Corresponding to the abnormal negative relationship over Tibetan areas found in Fig. 4, the correlation coefficients between PM<sub>2.5</sub> and AOD at a higher elevation, namely, greater than ~3200 m, were also less than zero. The correlation peaked at 400–600 m with an  $r$  value of 0.60 and  $p < 0.05$ . These

samples were then overlaid onto the spatial map of topography (Fig. S9 in the Supplement), and it was found that the maximum  $r$  values were primarily located in basin terrain such as the Sichuan Basin, the Shanxi–Shaanxi basins and the Junggar Basin.

**Table 1.** Correlation coefficients between (RH-corrected) PM<sub>2.5</sub> and (PBLH-corrected) AOD over various urban agglomerations in China during the study period, including Beijing–Tianjin–Hebei (BTH), Yangtze River Delta (YRD), Pearl River Delta (PRD), Triangle of Central China (TCC), Chengdu–Chongqing (CDCQ), central and southern Liaoning (CSL), Shandong Peninsula (SDP), Harbin–Changchun (HC), Central Plain (CP), Guanzhong Plain (GZP) and northern Tianshan (NTS).

|                                | Region | 2019 | Spring | Summer | Autumn | Winter |                    | Region | 2019 | Spring | Summer | Autumn | Winter |
|--------------------------------|--------|------|--------|--------|--------|--------|--------------------|--------|------|--------|--------|--------|--------|
| Original data                  | BTH    | 0.53 | 0.75   | 0.53   | 0.70   | 0.62   | Both corrected     | BTH    | 0.70 | 0.80   | 0.60   | 0.77   | 0.66   |
|                                | YRD    | 0.33 | 0.36   | 0.44   | 0.35   | 0.43   |                    | YRD    | 0.62 | 0.51   | 0.54   | 0.67   | 0.55   |
|                                | PRD    | 0.30 | 0.27   | 0.39   | 0.45   | 0.40   |                    | PRD    | 0.62 | 0.25   | 0.44   | 0.66   | 0.73   |
|                                | TCC    | 0.40 | 0.31   | 0.48   | 0.40   | 0.58   |                    | TCC    | 0.72 | 0.52   | 0.52   | 0.65   | 0.64   |
|                                | CDCQ   | 0.53 | 0.40   | 0.41   | 0.68   | 0.46   |                    | CDCQ   | 0.64 | 0.41   | 0.41   | 0.55   | 0.62   |
|                                | CSL    | 0.47 | 0.73   | 0.40   | 0.55   | 0.49   |                    | CSL    | 0.69 | 0.81   | 0.47   | 0.66   | 0.73   |
|                                | SDP    | 0.48 | 0.67   | 0.42   | 0.56   | 0.62   |                    | SDP    | 0.71 | 0.74   | 0.56   | 0.72   | 0.71   |
|                                | HC     | 0.31 | 0.52   | 0.34   | 0.53   | 0.47   |                    | HC     | 0.59 | 0.68   | 0.47   | 0.71   | 0.61   |
|                                | GZP    | 0.58 | 0.40   | 0.55   | 0.53   | 0.72   |                    | GZP    | 0.83 | 0.58   | 0.69   | 0.76   | 0.83   |
|                                | CP     | 0.47 | 0.67   | 0.54   | 0.55   | 0.61   |                    | CP     | 0.76 | 0.79   | 0.63   | 0.79   | 0.70   |
|                                | NTS    | 0.26 | 0.37   | 0.21   | 0.23   | –      |                    | NTS    | 0.60 | 0.51   | 0.36   | 0.52   | –      |
| RH-corrected PM <sub>2.5</sub> | BTH    | 0.54 | 0.81   | 0.58   | 0.79   | 0.51   | PBLH-corrected AOD | BTH    | 0.68 | 0.75   | 0.43   | 0.71   | 0.69   |
|                                | YRD    | 0.43 | 0.54   | 0.46   | 0.51   | 0.52   |                    | YRD    | 0.54 | 0.31   | 0.46   | 0.56   | 0.49   |
|                                | PRD    | 0.50 | 0.34   | 0.44   | 0.55   | 0.57   |                    | PRD    | 0.36 | 0.16   | 0.34   | 0.50   | 0.52   |
|                                | TCC    | 0.45 | 0.38   | 0.49   | 0.47   | 0.59   |                    | TCC    | 0.69 | 0.40   | 0.35   | 0.58   | 0.65   |
|                                | CDCQ   | 0.50 | 0.41   | 0.38   | 0.52   | 0.58   |                    | CDCQ   | 0.57 | 0.38   | 0.32   | 0.68   | 0.42   |
|                                | CSL    | 0.58 | 0.76   | 0.48   | 0.69   | 0.55   |                    | CSL    | 0.58 | 0.73   | 0.28   | 0.58   | 0.65   |
|                                | SDP    | 0.54 | 0.73   | 0.47   | 0.64   | 0.63   |                    | SDP    | 0.65 | 0.64   | 0.31   | 0.65   | 0.71   |
|                                | HC     | 0.44 | 0.58   | 0.45   | 0.63   | 0.48   |                    | HC     | 0.42 | 0.60   | 0.19   | 0.65   | 0.51   |
|                                | GZP    | 0.62 | 0.51   | 0.60   | 0.58   | 0.72   |                    | GZP    | 0.78 | 0.42   | 0.43   | 0.71   | 0.79   |
|                                | CP     | 0.53 | 0.68   | 0.60   | 0.66   | 0.59   |                    | CP     | 0.71 | 0.73   | 0.42   | 0.71   | 0.68   |
|                                | NTS    | 0.30 | 0.44   | 0.28   | 0.26   | –      |                    | NTS    | 0.49 | 0.42   | 0.29   | 0.45   | –      |

4 Discussion

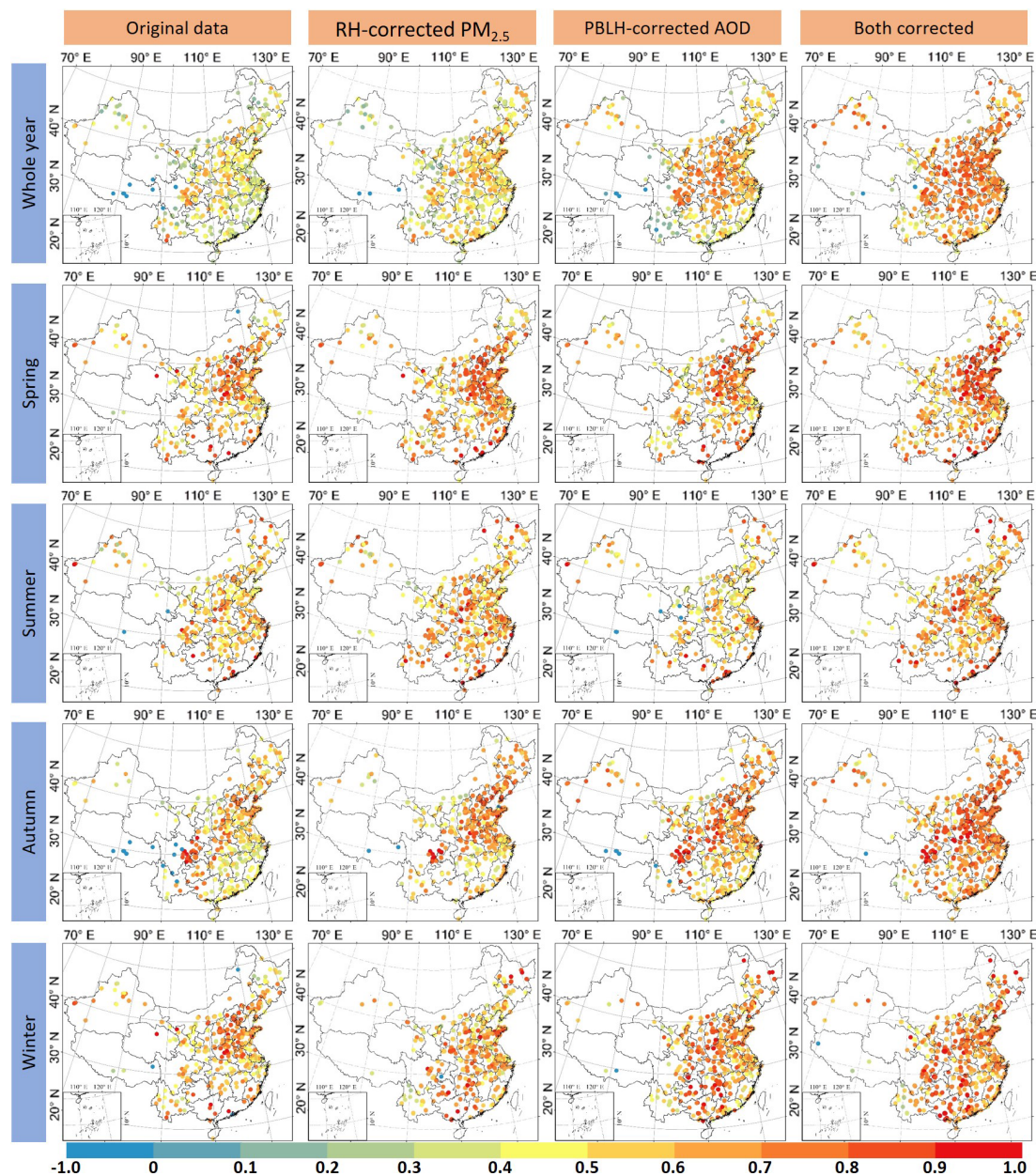
Our results show a mismatch between satellite AOD values and ground-level PM<sub>2.5</sub> concentrations, especially in south-western China. The main cause of the PM<sub>2.5</sub>–AOD mismatch lies in the fundamental discrepancy in the physical concept. AOD is a unitless variable that denotes the total extinction induced by aerosols in the entire atmospheric column, whereas PM<sub>2.5</sub> represents the mass concentration of fine aerosol particles measured at ground level. Another reason was likely the different available observations between the two metrics of aerosol loading: AOD derived from satellite only offered relatively fewer effective retrievals over this area due to cloud contamination (Fig. S2 in the Supplement), whereas surface PM<sub>2.5</sub> observations have almost no such limitations. Consequently, the limited satellite AOD samples cannot fully represent the actual aerosol loading. If the missing AOD values are not filled in, excessive uncertainty will be induced to the subsequent PM<sub>2.5</sub>-related health research. Thus, particular attention should be paid to the mismatch between satellite AOD and ground PM<sub>2.5</sub> caused by the number of observations.

We found that mean AOD and CV values had a different regional pattern. This result suggests that the SDP was uniformly polluted by aerosol particles throughout the atmosphere, as the largest mean AOD value (0.48) was observed in this region with the smallest CV value (0.16). In addition,

the regional CV with respect to PM<sub>2.5</sub> generally exhibited larger values than AOD, revealing that the extent of the spatial non-stationarity shown by PM<sub>2.5</sub> was obviously stronger compared with AOD in each individual region. Notably, the national CV of AOD was remarkably larger than the regional values and contrarily higher than the PM<sub>2.5</sub> CV, implying that the overall degree of spatial variability represented by ground-level PM<sub>2.5</sub> measurements was weaker than that of satellite AOD at a national scale. This is probably due to the fact that the monitoring stations are only sparsely located in urban areas, whereas satellite AOD can cover almost the whole study domain. Thus, the PM<sub>2.5</sub> measurements represent the pollution level at a particular location, whereas the satellite AOD interprets the average extinction level for a 1 km “area”. This further corroborates the mismatch between the PM<sub>2.5</sub> and AOD patterns observed in Figs. 1 and S3.

With respect to the spatiotemporal variability in the PM<sub>2.5</sub>–AOD relationship, the TP shows abnormal negative associations between retrieved PM<sub>2.5</sub> and AOD. According to Lyapustin et al. (2018), the MAIAC algorithm did not make reliable aerosol retrievals at very high elevations, and a fixed climatology value should be applied instead; thus, the PM<sub>2.5</sub>–AOD relationship over the TP contrasts with other areas. While we found a large spatial discrepancy in the correlation coefficient for each monitoring site, which was

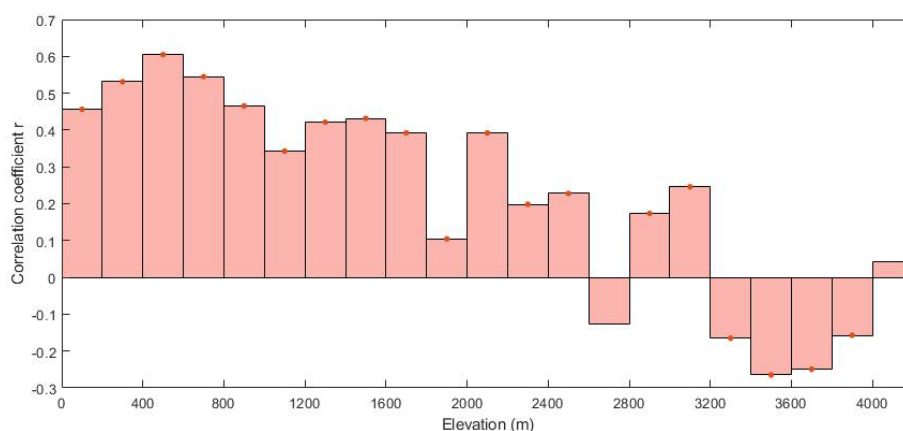




**Figure 6.** The relationship between PM<sub>2.5</sub> and AOD in China for the entire year 2019 and for each respective season based on the original PM<sub>2.5</sub> and AOD observations (left column), only PM<sub>2.5</sub> corrected by RH (middle left column), only AOD corrected by the PBLH (middle right column) and both corrected data (right column). Please note that only stations with statistically significant results (i.e.,  $p$  level < 0.5) are presented in this figure.

consistent with previous studies using MODIS DT aerosol data (Guo et al., 2017; Q. Yang et al., 2019), the relatively low  $r$  values (< 0.4) in southern areas were slightly different from previous studies (Guo et al., 2017). Based on the MODIS DT AOD dataset, Guo et al. (2017) reported that the correlation coefficients between PM<sub>2.5</sub> and AOD in eastern China were apparently larger than those in the western China, as the DT algorithm had a larger uncertainty in the western arid and semiarid areas. Unlike the DT algorithm,

which can only derive reliable AOD values over surfaces with low albedos, MAIAC could provide aerosol retrievals from both bright and dark targets. As shown in Fig. S2, the MAIAC aerosol product can provide more retrievals over the northern arid and semiarid areas than over southern China. Moreover, according to N. Liu et al. (2019), the accuracy of MAIAC AOD over unoccupied land, built-up areas and mixed regions was evidently better than that of DT. Hence, the abovementioned evidence results in relatively lower bi-



**Figure 7.** Correlation coefficients of PM<sub>2.5</sub>–AOD pairs plotted against binned elevations. Note that the red asterisk indicates the statistically significant level at  $p < 0.05$ .

ases for the PM<sub>2.5</sub>–AOD relationship under these conditions and also partly accounts for the difference in the spatial pattern of the PM<sub>2.5</sub>–AOD association between the MAIAC and DT AOD products. In addition to the accuracy and data availability of aerosol retrieval algorithms, the natural condition, including climate and terrain, also had obvious influences on the relationship between PM<sub>2.5</sub> and AOD. For example, compared with southern China, the RH in the north was generally lower, which contributed to the harmony in the PM<sub>2.5</sub>–AOD association in terms of the humidity conditions. In addition, the monsoon could cause aerosols to be unstable in the subtropical climate zone, which may weaken correlations in these areas.

It is obvious that spatial variation in the PM<sub>2.5</sub>–AOD correlation and the slopes of the linear regressions between daily PM<sub>2.5</sub> and AOD for each monitoring site was observed in our results. The largest correlation in the SCB was most likely due to its bowl topography. The surrounding terrain is not favorable for pollutant dispersion and, thus, causes pollutant accumulation in the local environment (Ning et al., 2018, 2019, 2020). This explains the good agreement between ground-level PM<sub>2.5</sub> and total column AOD. For the low correlation in northern central China, e.g., the Qinghai and Gansu provinces, MAIAC AOD retrievals were replaced by a mean value over a mesoscale area of 150 km (Lyapustin et al., 2018), which cannot resolve the 1 km aerosol variability found elsewhere and, thus, introduces additional biases into the correlation between PM<sub>2.5</sub> and AOD, leading to incongruous associations between PM<sub>2.5</sub> and AOD over these areas. Regarding the slopes of the linear regression between daily PM<sub>2.5</sub> and AOD, a larger slope suggests less aerosol extinction. The higher  $S$  values in inland areas and during winter indicate that, under these circumstances, AOD was not very sensitive to PM<sub>2.5</sub> change. This was probably attributed to the spatial discrepancy and seasonal variation under humid conditions. Due to increased hygroscopicity, the size of aerosol particles become larger than dry particles (Silva et al.,

2015), which allows more light to be scattered in the context of higher RH, e.g., over coastal areas and during summer. In addition to humid conditions, the aerosol type was another reason to account for the spatial variability in the slope of the PM<sub>2.5</sub>–AOD relationship. Previous studies have reported that the regression function between the ambient PM<sub>2.5</sub> concentration and AOD corresponds to the typical air masses over a specific region (Kong et al., 2016) and that the aerosol extinction property cannot be as sensitive in areas dominated by coarse-mode aerosols (Kong et al., 2016; Xin et al., 2014). As indicated by the spatial distribution of the mean Ångström exponent (AE or  $\alpha$ ) in Fig. S6 in the Supplement, aerosol particles with a larger size (smaller AE), such as dust, primarily dominated in northern areas over China, whereas fine-mode aerosols prevail in the south, thereby leading to a greater slope of the PM<sub>2.5</sub>–AOD relationship in northern China. It should be noted that the regional pattern of the PM<sub>2.5</sub>–AOD relationship was a little different from the pattern shown by individual monitors, as the regional association was an integration of observational sites within an urban agglomeration rather than a simple averaging of all sites within the region. Moreover, we found that most of the regional correlations based on MAIAC AOD in this study were closed to (and even slightly higher than) those by MODIS DT AOD in previous studies (Guo et al., 2017; Q. Yang et al., 2019). For example, the correlations over the BTH, PRD, SDP and CP regions were 0.55, 0.38, 0.42 and 0.39, respectively, whereas the corresponding correlations in this study were 0.53, 0.30, 0.48 and 0.47, respectively. This was because MAIAC AOD (1 km spatial resolution) had a refined spatial resolution compared with DT AOD (10 km), thereby reducing the variability in aerosol inside a grid cell co-located with the in situ measurements. The improved overall retrieval accuracy of the MAIAC algorithm (Zhang et al., 2019; N. Liu et al., 2019) may be another reason for the bias reduction in the correlation analysis for the PM<sub>2.5</sub>–AOD relationship. These findings suggest that the MAIAC AOD product is a better indicator



of fine particle pollution than the MODIS coarse-resolution aerosol product.

We also found a significant effect of meteorology on the PM<sub>2.5</sub>–AOD correlation using Terra and Aqua AOD. A previous study evaluated the MAIAC retrieval performance of the Terra and Aqua AOD datasets against ground-level aerosol observations and found that AOD data from the two satellites presented a similar retrieval accuracy (Zhang et al., 2019). Thus, we deduced that the difference in the correlation between the two times (11:00 and 14:00 LT) was most likely due to the diurnal variability in climatology (e.g., the PBLH). The diurnal cycle of the PBLH was more likely one important factor leading to the noon and afternoon maximum. This is because a higher PBLH, which is often observed at noon and in the afternoon (Liu and Liang, 2010; Zhang et al., 2018), would definitely allow air pollutants to be mixed vertically such that the PM<sub>2.5</sub>–AOD relationship should be closer. At nighttime, the relative stable atmosphere may be favorable for the accumulation of aerosol particles near the ground surface; thus, the AOD values may not be able to reflect the ground-level aerosol concentration as well as they do during the daytime. Our results show a significant improvement in the PM<sub>2.5</sub>–AOD correlation after RH and vertical correction, suggesting the importance of applying the correction for the retrieval work. It should be noted that there was an obvious negative increment after both corrections, in particular at most monitoring stations in Yunnan Province. The poorer correlation may be attributed to the vertical correction, as the single correction for humidity can slightly enhance the PM<sub>2.5</sub>–AOD correlation, contributing to an increase of 0.01 in the  $r$  values on average, but the correlations revised by vertical data are remarkably reduced by 0.04–0.42 for individual sites. Due to the East Asian and Indian monsoons, extraneous aerosols such as smoke and dust transported from India, Burma and desert regions in northern China can form elevated aerosol layers in the Yunnan–Guizhou Plateau (Zhu et al., 2017; Liu et al., 2016), leading to the local AOD that is vertically corrected by the PBLH being poorly related to the ground-observed PM<sub>2.5</sub> in this region. The effect of meteorology on the improvement of the PM<sub>2.5</sub>–AOD relationship was also found to vary seasonally. Nevertheless, our results show that a vertical correction in summer may not be useful to improve the PM<sub>2.5</sub>–AOD relationship in central China, such as in the Jiangxi, Shaanxi and Henan provinces. Central China is located in a typical subtropical monsoon climate zone, and the summer monsoon causes the frequent occurrence of unstable weather and increases the dissipation of local aerosols associated with the scour effect of rainfall and atmospheric convection (Gong et al., 2017; Wang et al., 2011), which may account for the weaker PM<sub>2.5</sub>–AOD correlation corrected for the PBLH. The RH correction has a weaker influence in winter, during which the averaged coefficient was found to increase by less than 0.01, which may be related to the dry climate in this season. These findings signify that seasonal variability in the data plays an important role in im-

proving the PM<sub>2.5</sub>–AOD relationship and cannot be ignored, despite the existing differences in the corrected correlations. On the other hand, PBLH-correction can effectively improve the PM<sub>2.5</sub>–AOD relationship in northwestern China. Contrary to the poor effect of vertical revision in summer and spring, seasonal PM<sub>2.5</sub> revised by humidity was more closely linked to satellite AOD than the uncorrected relationship in almost all agglomeration areas, except for summer and autumn in the CDCQ area and winter in the BTH and CP regions. This further corroborates the fact that the regional pattern was a little different from the pattern presented by monitoring stations and that the RH and vertical corrections did not improve the correlation between ground-measured PM<sub>2.5</sub> and satellite AOD with significant scope and should instead be tailored to specific areas.

Besides meteorology, terrain was also found to have impact on the PM<sub>2.5</sub>–AOD correlation, especially in basin areas. This may be because basin terrain tends to form a local atmospheric environment where aerosol particles can mix well in the entire atmosphere. Moreover, the complex topography protects the basins from transboundary transport of air pollution (Ning et al., 2019, 2020). These two reasons contributed to the consistency between near-surface PM<sub>2.5</sub> and column AOD. Figure 7 also shows that the  $r$  values tend to become lower as the elevation rises, which indicates a deterioration in the PM<sub>2.5</sub>–AOD correlation with altitude, especially in eastern China where the majority of observational sites are in operation. Due to the spatially uneven distribution of air quality monitoring stations, there were only 23.68 % pairs of matched samples located at elevations higher than 1000 m; thus, Figs. 7 and S9 mainly represent the effect of terrain on the PM<sub>2.5</sub>–AOD relationship in eastern China.

Finally, it is necessary to discuss the implications for ground PM<sub>2.5</sub> derived from satellite AOD. With the release of the MODIS MAIAC aerosol product, an emerging application of this high-spatial-resolution AOD dataset has become the inference of fine particulate matters in China. This study profiled the spatiotemporal variations in the ground-level PM<sub>2.5</sub> and the 1 km satellite AOD and analyzed the underlying factors that contribute to the current pattern. Generally consistent with previous studies (Q. Yang et al., 2019; Shao et al., 2017; Wang et al., 2014), the PM<sub>2.5</sub>–AOD association in this study presents a significant spatiotemporal variation. The meteorological parameters and topography have a complicated influence on the PM<sub>2.5</sub>–AOD relationship, posing extreme challenges for ground PM<sub>2.5</sub> estimations. For example, the PBLH has an adverse impact on the relationship in Yunnan, whereas humidity may not be useful for improving the association in winter. The aerosol type is another contributor accounting for the uncertainties involved; however, as our results show, the pattern based on the 1 km MAIAC AOD is a little different from previous studies using coarse-resolution aerosol products, such as the MODIS 10 km DT or DB AOD datasets (Guo et al., 2017; Q. Yang et al., 2019), which should be noted for PM<sub>2.5</sub> predictions. For example,

the aerosol retrieval scheme implemented by the MAIAC algorithm at very high altitudes brings about the unnatural negative PM<sub>2.5</sub>–AOD correlation over Tibetan region and, thus, has a direct effect on the PM<sub>2.5</sub> estimation ability. To clarify this issue, we established two separate multiple linear regression models using the same daily samples – one for areas with an elevation higher than 3200 m and one for the remaining area – and obtained significantly different  $R^2$  values of 0.14 vs. 0.41 at  $p < 0.01$ . This result suggests that the PM<sub>2.5</sub> prediction model should treat very high-altitude areas with caution or additional uncertainty within the data may hinder the acquisition of model parameters and further attenuate the prediction accuracy. Therefore, a discreet attitude is required to apply the newly released aerosol product to PM<sub>2.5</sub>, especially for large areas with a complicated terrain and land cover.

## 5 Conclusions

In this study, surface PM<sub>2.5</sub> measurements in 2019 across China were spatially and temporally matched with MAIAC 1 km AOD and combined with humidity, vertical data and topography to perform explicit correlation analyses, which are the cornerstone of retrieving ground-level PM<sub>2.5</sub> from satellite aerosol loading. From a MAIAC high-resolution perspective, we found that the relationship between surface PM<sub>2.5</sub> and column AOD was generally in good agreement with previous results but also showed a slight difference in spatiotemporal variability; thus, caution is required for applications to both aerosol-related and PM<sub>2.5</sub>-related studies. The findings of this study are helpful with respect to better understanding the PM<sub>2.5</sub> pollution status in China as well as its relationship with column AOD, thereby facilitating the reconstruction of high-resolution PM<sub>2.5</sub> from satellite aerosol retrievals.

The relationship between PM<sub>2.5</sub> and AOD varied notably in mainland China, both spatially and temporally, in terms of correlation coefficients and the slopes of the linear regression. Spatially, stronger correlations mainly occurred in northern and eastern China, and the linear slope in northern inland regions was larger on average than those in other areas. Temporally, the PM<sub>2.5</sub>–AOD correlation peaked at noon and in the afternoon and varied greatly between seasons: the maximum correlation occurred in winter ( $r = 0.61$ ). MAIAC 1 km AOD can better represent the ground-level fine particulate matter in most domains with the exception of in very high terrain (i.e., the Tibetan Plateau) and northern central China (i.e., Qinghai and Gansu), where MAIAC does not interpret the actual 1 km aerosol loading, resulting in unnatural or weaker PM<sub>2.5</sub>–AOD relationships for these domains. This merits more attention regarding the application of PM<sub>2.5</sub> estimations.

As far as the confounding impact of meteorology on the relationship between PM<sub>2.5</sub> and AOD was concerned, we found that simultaneously considering the RH and PBLH

in the relationship could improve the correlation, but the effect of the RH and PBLH on the relationship varied spatially and temporally with respect to both strength and direction. The diurnal cycle of the PBLH coincided with that of the PM<sub>2.5</sub>–AOD correlation, and the PBLH may impose an adverse impact on the daily relationship in Yunnan Province and weaken the seasonal relationship in summer in central China, such as in the Jiangxi, Shaanxi and Henan provinces. Compared with the PBLH, the RH exerted a greater contribution to improving the PM<sub>2.5</sub>–AOD correlation in spring and had a weaker influence in winter. In addition, the correlation peaked at 400–600 m primarily in basin terrain such as the Sichuan Basin, the Shanxi–Shaanxi basins and the Junggar Basin. Nevertheless, despite the publicly available high-resolution AOD data during the satellite era, an accurate high-resolution reconstruction from satellite remote sensing remains greatly challenging, not only due to the uncertainties existing in the satellite aerosol retrievals but also due to the uncertainties inherent in the PM<sub>2.5</sub> and AOD measurements, which covary with topography and meteorological conditions. Even though the individual effects of humidity, the vertical structure of aerosols and terrain have been clarified, the synthetic impact of these factors could also be core to regulating the statistical relationship between near-surface PM<sub>2.5</sub> and total column AOD, which requires elucidation in the future. Therefore, more work in this regard is urgently needed. In addition, aerosol properties, such as aerosol type, are plausible factors that could modulate the PM<sub>2.5</sub>–AOD relationship and are worthy of more attention in the future.

**Data availability.** PM<sub>2.5</sub> measurements were sourced from the official website of the Ministry of Ecology and Environment of China: <https://www.mee.gov.cn> (MEEC, 2021). MAIAC aerosol data are available from the NASA LAADS website: <https://ladsweb.modaps.eosdis.nasa.gov/search/> (NASA, 2021a). AERONET data are publicly available from <https://aeronet.gsfc.nasa.gov/> (NASA, 2021b).

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**Author contributions.** QH and SHLY designed the experiment. QH developed the experiment code, and implemented and validated it. The paper was written by QH and revised by SHLY and MW.

**Competing interests.** The authors declare that they have no conflict of interest.

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