



# Synergetic effects of $\text{NH}_3$ and $\text{NO}_x$ on the production and optical absorption of secondary organic aerosol formation from toluene photooxidation

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**Abstract.**  $\text{NH}_3$  is the most important alkaline gas in the atmosphere and one of the key species affecting the behaviors of atmospheric aerosols. However, the impact of  $\text{NH}_3$  on secondary organic aerosol (SOA) formation remains poorly understood, especially the dynamic evolution of chemical compositions in the SOA formation process. In this study, a series of chamber experiments were performed to probe the individual and common effects of  $\text{NH}_3$  and  $\text{NO}_x$  on toluene SOA formation through OH photooxidation. The chemical compositions of toluene SOA were characterized using the Aerodyne high-resolution time-of-flight aerosol mass spectrometer (AMS). The SOA yield increased from 28.1 % in the absence of  $\text{NH}_3$  to 34.7 % in the presence of  $\text{NH}_3$  but decreased to 19.5 % in the presence of  $\text{NO}_x$ . However, the highest SOA yield of 42.7 % and the lowest carbon oxidation state ( $\text{OS}_C$ ) occurred in the presence of both  $\text{NH}_3$  and  $\text{NO}_x$ , indicating that the higher-volatility products that formed in the presence of  $\text{NO}_x$  could partition into the particle phase when  $\text{NH}_3$  was added. This resulted in a synergetic effect on SOA formation when  $\text{NH}_3$  and  $\text{NO}_x$  co-existed. The heterogeneous reaction was the main pathway by which  $\text{NH}_3$  participated in SOA formation in the photooxidation process. The synergetic effect of  $\text{NH}_3$  and  $\text{NO}_x$  was also observed in SOA optical absorption. A peak at 280 nm, which is characteristic of organonitrogen imidazole compounds, was observed in the presence of  $\text{NH}_3$ , and its intensity increased when  $\text{NO}_x$  was added into the chamber. This work improves our under-

standing of how the synergistic interactions between  $\text{NH}_3$  and  $\text{NO}_x$  influence SOA formation and offers new insights into mitigating haze pollution.

## 1 Introduction

Secondary organic aerosols (SOAs) are an important component of atmospheric particulate matter (Moise et al., 2015; Liu et al., 2017) and can significantly affect atmospheric visibility, air quality, and human health (Paciga et al., 2014; Yang et al., 2016; Liu et al., 2017). Optical properties of SOA have been directly and indirectly linked to their effects on the climate (Laskin et al., 2015; Xie et al., 2017; Peng et al., 2020). Because of the complexity of their chemical components, oxidation processes, and environmental factors, SOA formation mechanisms are very complex, and the current understanding of SOA formation is incomplete. This limited understanding hampers the ability of models to predict the magnitudes, dynamics, and distributions of atmospheric aerosols from particulate and precursor emissions (Ortiz-Montalvo et al., 2014). In the past decades, although our understanding of SOA formation mechanisms has been constantly improving, there is still a gap between the simulated SOA concentration in large-scale atmospheric models and field observations (Volkamer et al., 2006; Yang et al., 2018).

Ammonia (NH<sub>3</sub>) is the most abundant and ubiquitous alkaline inorganic gas in the atmosphere and one of the critical factors influencing SOA formation (Wang et al., 2016, 2018b; Chen et al., 2019). Some studies have noted that the presence of NH<sub>3</sub> can contribute to the formation of more aerosol mass through photooxidation (Na et al., 2007; Li et al., 2018). Na et al. (2007) observed that aerosol yields in the  $\alpha$ -pinene–ozone oxidation system increased by 8 % when NH<sub>3</sub> was added. Li et al. (2018) reported that the presence of NH<sub>3</sub> in the aromatic hydrocarbon photooxidation system increased aerosol size growth potentials (by 7 %–108 %) and resulted in enhanced SOA formation. Qi et al. (2020) found that the concentration and average diameter of SOA showed an immediate and rapid increase after adding NH<sub>3</sub>. Furthermore, the acid–base reactions between NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> and the carboxyl groups in SOA molecules might enhance SOA formation (Qi et al., 2020; Liu et al., 2015). The condensable ammonium salts formed from the reaction between NH<sub>3</sub> and organic acids reduce the volatility of the organic acids by several orders of magnitude (Paciga et al., 2014), and they act as particle-phase organics that further promote SOA formation (Na et al., 2007; Huang et al., 2012; Chen et al., 2019; Qi et al., 2020; Wu et al., 2020). In addition, carbonyls can undergo nucleophilic attack by NH<sub>3</sub> through the Mailard reactions and form the corresponding iminium intermediates (Noziere et al., 2009; Laskin et al., 2015; Liu et al., 2015). The iminium intermediates can continue to react with carbonyls, which activates further transformations such as the formation of heterocyclic compounds and oligomerization reactions and form condensation (oligomeric) products with more stable secondary imines (Schiff bases) (Laskin et al., 2014). Both Noziere et al. (2009) and Ortiz-Montalvo et al. (2014) reported that NH<sub>3</sub> is an efficient catalyst for reactions with carbonyl compounds to form nitrogen-containing organic aerosols (NOAs). The reaction between carbonyl and NH<sub>3</sub> can significantly decrease the volatility of oxidation products, which further increases the yield of SOA (Lee et al., 2013; Zhang et al., 2015a; Qi et al., 2020). Babar et al. (2017) found that the substantial formation of secondary imines in the presence of NH<sub>3</sub> was responsible for the higher  $\alpha$ -pinene SOA yields. However, not all studies have shown that the presence of NH<sub>3</sub> increases SOA yields. One study observed that NH<sub>3</sub> suppressed SOA formation under certain ozonation conditions (Ma et al., 2018b). Furthermore, the consumption of NH<sub>3</sub> by Criegee intermediates was reported to decrease the secondary ozonide yield and thus affect SOA formation.

Nitrogen oxides (NO<sub>x</sub> = NO + NO<sub>2</sub>), which are mainly emitted from the combustion of fossil fuels, have received significant attention due to their effects on the photooxidation process of volatile organic compounds (VOCs) and SOA formation (Surratt et al., 2006; Ng et al., 2007b; Draper et al., 2015; Berkemeier et al., 2016; Sarrafzadeh et al., 2016; Zhao et al., 2018). A clear increase at first and then a decrease in the SOA yield was found with increasing NO<sub>x</sub> con-

centration from the laboratory experiments with both anthropogenic (trimethylbenzene) and biogenic ( $\beta$ -pinene) VOCs (Sarrafzadeh et al., 2016; Yang et al., 2020). The competitive chemistry of organic peroxy radicals (RO<sub>2</sub>) with hydroperoxyl radicals (HO<sub>2</sub>) and NO was responsible for the variability in SOA formation (Ng et al., 2007a; Xu et al., 2014; Jiang et al., 2020). RO<sub>2</sub> mainly reacts with HO<sub>2</sub> under low-NO<sub>x</sub> conditions to form oxidation products with lower volatility, which may enable it to partition into the particle phase and contribute to the SOA mass (Ng et al., 2007a). While the RO<sub>2</sub> + NO reaction is dominant in high-NO<sub>x</sub> conditions, the increase in volatile products formed through fragmentation was responsible for the decrease in SOA yield with increasing NO<sub>x</sub> (Zhao et al., 2018; Liu et al., 2019a; Xu et al., 2020). In addition, the increase in OH concentration formed through the NO + HO<sub>2</sub> → O<sub>2</sub> + OH reaction at low-NO<sub>x</sub> conditions and a suppressing effect of NO<sub>x</sub> on OH formation under high-NO<sub>x</sub> conditions were partly responsible for the first increasing and then decreasing trend of SOA yield with NO<sub>x</sub> concentration (Sarrafzadeh et al., 2016; Bates et al., 2021).

In the last decade, atmospheric pollutants in China have changed significantly in their concentrations and composition (Wang et al., 2015; Xia et al., 2016) with the emissions of SO<sub>2</sub> and NO<sub>x</sub> decreased by 75 % from 2007–2015 and 10 % from 2011–2015, respectively (de Foy et al., 2016; Vu et al., 2019; Wang et al., 2020). However, owing to the lack of regulation regarding NH<sub>3</sub> emissions, NH<sub>3</sub> emissions increased by ~ 30 % from 2008–2016 over the North China Plain (Liu et al., 2018). As has been pointed out in previous research, the promoting effect of NH<sub>3</sub> on the formation of SOA may counteract the decreases in aerosol formation due to reductions in SO<sub>2</sub> and NO<sub>x</sub> (Zhang et al., 2021a). Indeed, field observation and model simulation have pointed out that the simultaneous control of NH<sub>3</sub> emissions in conjunction with SO<sub>2</sub> emission is more effective in reducing PM<sub>2.5</sub> than the process without NH<sub>3</sub> emissions control, and PM<sub>2.5</sub> concentration can be more effectively reduced if NH<sub>3</sub> emission is decreased as much as that of SO<sub>2</sub> (Erisman and Schaap, 2004). Hence, the mechanism by which NH<sub>3</sub> affects SOA formation has attracted more and more attention (Wang et al., 2018a; Ge et al., 2019; Zhang et al., 2021b). However, previous studies have not paid sufficient attention to the joint impacts of NH<sub>3</sub> and NO<sub>x</sub> on the formation of SOA and its corresponding optical properties. Due to the lack of real-time detection methods for SOA chemical composition, the dynamic characteristics of how NH<sub>3</sub> participates in SOA formation via photooxidation have not been extensively studied.

Toluene is one of the most abundant aromatic VOCs in the urban atmosphere, which is also an important source of brown carbon (Laskin et al., 2010; Ma et al., 2018a). The effects of NH<sub>3</sub> and NO<sub>x</sub> on SOA formation through the toluene photooxidation process were investigated in this study. The chemical composition of toluene SOA was characterized online with an aerosol mass spectrometer, and the characteristics of SOA chemical composition under different conditions

were further explored by applying a positive matrix factorization (PMF) analysis. The optical properties of toluene SOA particles were determined based on a UV–Vis spectrum analysis. Possible mechanisms of the effects of both NH<sub>3</sub> and NO<sub>x</sub> on SOA formation were discussed. The results will help us to better understand SOA formation mechanisms under complex pollution conditions with elevated NH<sub>3</sub> and NO<sub>x</sub> concentrations in an urban atmospheric environment.

## 2 Materials and methods

### 2.1 Photooxidation chamber experiments

All toluene photooxidation experiments were performed in a 4 m<sup>3</sup> chamber. The chamber has been described in detail in our previous study (Liu et al., 2021). Briefly, the chamber was constructed with a 0.08 mm thick Teflon FEP film. Fifty UV-B lamps (TUV36W, Philips) with peak wavelengths of 254 nm were set up around the chamber and used as the light source to drive OH radical formation through hydrogen peroxide (H<sub>2</sub>O<sub>2</sub>) photolysis. Mirror-surfaced stainless steel was used as the interior wall of the enclosure to maximize and homogenize the interior light intensity. All experiments were performed at room temperature (293–298 K), and atmospheric pressure was maintained in the chamber at all times.

Before each experiment, the chamber was flushed with zero air for at least 18 h, after which the concentration of particles was less than 1 cm<sup>-3</sup>. Zero air was generated by a zero air supply (111-D3N, Thermo Scientific™, USA). The flow rate of zero air was controlled at 20 L min<sup>-1</sup> by a mass flow controller (D088C/ZM, Beijing Sevenstar Electronics Co.) during the process of chamber reaction bag inflating. The relative humidity (RH) of zero air was about 20 %. For each experiment, measured amounts of toluene (Sigma-Aldrich, analytically pure) and H<sub>2</sub>O<sub>2</sub> solution (Sigma-Aldrich, 30 wt % in H<sub>2</sub>O) were injected into a Teflon bulb with microsyringes. Zero air was passed through the injection tube to make sure all the liquids had evaporated to the gas phase and were blown into the chamber. Toluene concentration was measured with a proton-transfer-reaction time-of-flight mass spectrometer (PTR-ToF-MS, Ionicon Analytik, Austria). The evolution of toluene concentration for different experiments was shown in Fig. S1. The OH concentration in the chamber was calculated based on the first-order decay of toluene concentration. There was no obvious difference in OH concentrations in the different NO<sub>x</sub> and NH<sub>3</sub> levels (Fig. S2). The average OH concentration over the entire reaction period was 5.87 × 10<sup>7</sup> molec. cm<sup>-3</sup>. NO<sub>x</sub> (Air Liquid Shanghai, 510 ppm NO<sub>2</sub> in N<sub>2</sub>) and/or NH<sub>3</sub> (Air Liquid Shanghai, 502 ppm NH<sub>3</sub> in N<sub>2</sub>) were introduced directly into the chamber to reach the required concentrations. For experiments with NO<sub>x</sub>, although only NO<sub>2</sub> was introduced into the chamber before photooxidation, NO could be formed through NO<sub>2</sub> photolysis under the UV light irradiation, so NO always co-

existed with NO<sub>2</sub> in the photooxidation system (Zhao et al., 2018). NO<sub>x</sub> concentration was measured online by the NO–NO<sub>2</sub>–NO<sub>x</sub> analyzer (model 42C, Thermo Electron Corporation, USA). Each experiment was performed without seed aerosols. After all the reactants were added, the chamber stood undisturbed for 10 min without turning on the light to ensure that the reactant gases in the chamber were evenly mixed. The photooxidation started when the UV light was turned on.

The experimental conditions for the toluene photooxidation are listed in Table 1. In our work, the OH and toluene concentrations were higher than those of urban conditions. The purpose of the high OH and toluene concentrations is to obtain enough particle production samples for offline collections and accurate measurements. The toluene concentrations remained stable under the different experimental conditions, the variations in toluene-derived SOA mass concentration and yield were only affected by the different NO<sub>2</sub> and/or NH<sub>3</sub> concentrations in this study. Toluene was studied here as the representative of total aromatic VOCs in the urban atmosphere, thus its concentration used in this study was much higher than that in the real atmosphere. However, the concentration ratio of toluene to OH in this study is similar to that under real atmospheric conditions (Prinn et al., 1995; Zou et al., 2015).

### 2.2 Particle concentration measurements

For each experiment, a scanning mobility particle sizer (SMPS) was used to record the particle size distribution and volume concentration of the toluene-derived SOA. The SMPS was composed of a differential mobility analyzer (DMA model 3081, TSI Inc., USA) and a condensation particle counter (CPC model 3776, TSI Inc., USA), which were used for screening particles with specific electrical mobility diameter (from 14.1 to 736.5 nm) and for counting the number of the selected particles, respectively. The sheath gas velocity was 3 L min<sup>-1</sup>, and the sample gas velocity was 0.3 L min<sup>-1</sup>. The size scan was repeated every 5 min. During each scan cycle, the scan time was 240 s, and the particle sizes ranged from 13.6 to 726.5 nm. A density of 1.4 g m<sup>-3</sup>, which was measured by Ng et al. (2007b), was used for the calculation of toluene SOA mass concentration from the particle volume concentration (Ng et al., 2007b).

### 2.3 Chemical characterization

In this study, the toluene SOA chemical compositions were characterized with an online high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS, Aerodyne Research Inc., USA). The sample flow passed through a Nafion dryer, and the RH of the sample gas was reduced to below 20 % before entering the AMS. In the injection port, an aerodynamic lens focused particles with a vacuum aerodynamic diameter below 1 μm into a narrow beam. Parti-

**Table 1.** Summary of experimental conditions in this study.

No.	Tol <sub>0</sub>	ΔTol	NH <sub>3</sub> <sup>a</sup>	NO <sub>2</sub>	RH	SOA mass conc. <sup>b,c</sup>	SOA yield <sup>b</sup>
	(ppb)	(ppb)	(ppb)	(ppb)	(%)	(μg m <sup>-3</sup> )	(%)
Exp. 1	664.1	551.2	–	–	25 ± 1	637 ± 14.6	28.1
Exp. 2	618.7	499.4	~ 200	–	23 ± 1	867 ± 12.7	34.7
Exp. 3	620.9	526.1	~ 200	62	26 ± 1	1020 ± 10.6	42.7
Exp. 4	645.7	539.5	–	63	25 ± 1	452 ± 18.9	19.5

<sup>a</sup> The concentration of NH<sub>3</sub> is estimated by the amount of NH<sub>3</sub> added and the volume of the smog chamber. <sup>b</sup> SOA concentration and yield were calculated after taking into account the wall loss. <sup>c</sup> The reported SOA mass concentrations were the peak values after the wall-loss correction.

cles impacted a flash vaporizer (600 °C) at the rear of the sizing region under high vacuum (~ 10<sup>-7</sup> Torr) and were subsequently ionized by electron impact ionization (70 eV). Then, the positively charged ions entered the ToF section and were separated. V-mode ( $m/\Delta m \approx 2000$ ) was used in the AMS ToF section to achieve the high signal-to-noise ratio. The separated ion fragments were analyzed by a mass spectrometer with scans from 1 to 300  $m/z$ . The composition-dependent collection efficiency (CE) was applied to the data based on the methods established by Middlebrook et al. (2012). For mass concentration calculations, 1.1, 1.2, and 1.4 were applied as the default relative ionization efficiency (RIE) values of nitrate, sulfate, and organic compounds, respectively. The standard AMS data analysis software SQUIRREL 1.63B, coupled with PIKA 1.23B in Igor Pro (WaveMetrics, Inc., Portland, Oregon), which was retrieved from <http://cires1.colorado.edu/jimenez-group/ToFAMSResources/ToFSoftware/> (last access: 26 November 2021), was used for the analysis of elemental ratios and the ion speciated compositions of toluene SOA in the chamber. Note that the elemental ratios (i.e., O/C, H/C, and N/C) and mass-to-carbon ratio (OM/OC) were all calculated using the Aiken-Ambient method for comparability with previous studies (Aiken et al., 2008). In order to further explore the changes in SOA chemical composition, a PMF of the high-resolution mass spectra was performed to determine the different organic aerosol (OA) factors during the toluene photooxidation process. We performed the PMF analysis in the same way as Zhang et al. (2011), the details of which are provided in the Supplement.

## 2.4 Absorption measurements

The changes in absorption spectra and the absorbance of the toluene-derived SOA under different conditions were determined using a UV spectrophotometer (UV-3600, Shimadzu, Japan) with a 1 cm cuvette. The SOA was collected from a 3 m<sup>3</sup> sample gas onto the 46.2 mm PTFE filter (Whatman<sup>TM</sup>, UK). The collected SOA sample was dissolved in 5 mL of methanol (HPLC grade, > 99.8 %) with 30 min of sonication. As reported by Chen and Bond (2010), > 92 % of SOA is ex-

tractable by organic solvents (e.g., methanol), which means that almost all organic matter was extracted in this study. The filter extracts were filtered through 0.2 μm PTFE syringe filters to remove suspended insoluble particles. Before detection of the optical absorbance, a cuvette filled with pure methanol was scanned as a blank to provide a spectral background. The absorption was detected over the range of 200 to 800 nm with a resolution of 0.5 nm<sup>-1</sup>. The light absorption coefficient of the particles at a specific wavelength λ (Abs<sub>λ</sub>, Mm<sup>-1</sup>) was calculated according to Eq. (1):

$$\text{Abs}_\lambda = (A_\lambda - A_{700}) \cdot \frac{V_1}{V_a \cdot L} \cdot \ln(10), \quad (1)$$

where  $A_{700}$  is the background value of light absorption intensity, calculated as the average value of light absorption intensity from 695–705 nm to reduce the limits of error in measurement;  $V_1$  and  $V_a$  are the volumes of methanol with dissolved particles and sampled air, respectively; and  $L$  is the optical path length. Because Abs<sub>λ</sub> was strongly dependent on the amount of SOA, all Abs<sub>λ</sub> results were normalized based on the SOA mass collected on the filter. The normalized result was defined as the mass absorption coefficient (MAC, m<sup>2</sup> g<sup>-1</sup>), calculated using Eq. (2):

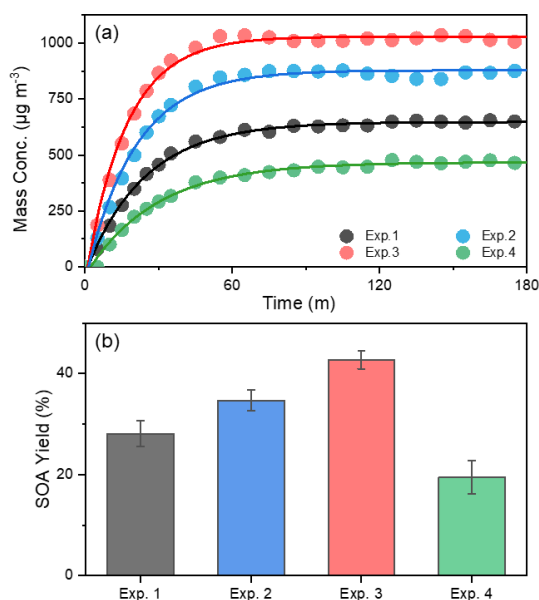
$$\text{MAC}_\lambda = \frac{\text{Abs}_\lambda}{M}, \quad (2)$$

where  $M$  (μg m<sup>-3</sup>) represents the concentration of methanol-soluble organic carbon.

## 3 Result and discussion

### 3.1 SOA formation

In order to investigate the effect of NH<sub>3</sub> and NO<sub>x</sub> on SOA formation from toluene photooxidation, a control test was carried out. SOA yield ( $Y$ ) is defined as  $Y = \Delta M_0 / \Delta \text{HC}$ , where  $\Delta M_0$  is the produced organic aerosol mass concentration (μg m<sup>-3</sup>), and  $\Delta \text{HC}$  is the mass concentration of reacted toluene (μg m<sup>-3</sup>). The evolution of SOA mass concentrations and SOA yield at different conditions during the photooxidation process are shown in Fig. 1. Recent experiments



**Figure 1.** The evolution of mass concentration (a) and yield (b) of toluene-derived SOA in different experiments. All the mass concentrations were wall-loss corrected. The error bars were calculated by the fluctuation of measured SOA concentration after the UV light was turned off at the end of each experiment.

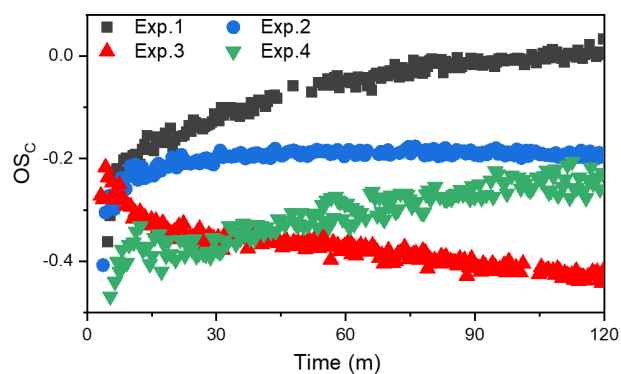
showed that the wall loss of organic vapors to the Teflon walls should not be ignored (Zhang et al., 2014, 2015b) and represented a major challenge in investigating SOA formation with environmental chambers (Krechmer et al., 2020). The formation of SOA in laboratory chambers may be substantially suppressed due to losses of SOA-forming vapors to chamber walls, but this effect on SOA formation has not yet been quantitatively established. However, the particle wall-loss rates were detected at the end of the chamber experiment after the UV lamps were turned off, and the mass concentration was corrected in the same way as by Jiang et al. (2020) and Pathak et al. (2007). After the wall-loss correction, the particle mass concentration was almost constant; the different wall-loss effect caused by gaseous oxidation products formed in the different experiment conditions was considered remedied. There was a noticeable increase in the SOA mass concentration in the presence of NH<sub>3</sub>, which was consistent with previous studies (Qi et al., 2020; Chu et al., 2016). The mass concentration of SOA increased from  $637 \pm 14.6 \mu\text{g m}^{-3}$  without NH<sub>3</sub> to a maximum of  $867 \pm 12.7 \mu\text{g m}^{-3}$  with 200 ppb NH<sub>3</sub>. Previous studies attributed the enhancement of SOA to the formation of NOA from acid–base reactions between NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> and carboxyl groups or Maillard reactions of NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> with carbonyl functional groups (Noziere et al., 2009; Ortiz-Montalvo et al., 2014; Liu et al., 2015; Qi et al., 2020). In contrast, SOA concentrations were lower in the presence of NO<sub>x</sub>, and the maximum mass concentration of toluene SOA was only  $452 \pm 18.9 \mu\text{g m}^{-3}$  with 63 ppb of initial NO<sub>x</sub>. The branching

of RO<sub>2</sub> loss among different pathways has an important influence on the product distribution and SOA formation. The fate of RO<sub>2</sub> mainly depends on the concentrations of NO<sub>x</sub> (Zhao et al., 2018; Liu et al., 2019a; Xu et al., 2020). Several studies have shown that, instead of RO<sub>2</sub> reacting with RO<sub>2</sub> / HO<sub>2</sub>, NO would react with RO<sub>2</sub> to form the RO intermediate and produces more oxidation products with higher volatilities through fragmentation in the presence of NO<sub>x</sub> (Zhao et al., 2018; Liu et al., 2019a; Xu et al., 2020). Highly volatile compounds cannot readily partition into the particle phase, so this substantially suppresses the formation of SOA.

The NO<sub>x</sub> and NH<sub>3</sub> had opposite effects on toluene SOA formation in this study. Interestingly, however, the highest SOA mass concentration ( $1020 \pm 10.6 \mu\text{g m}^{-3}$ ) occurred in the presence of both NO<sub>x</sub> and NH<sub>3</sub>, which was nearly 1.6 times higher than that observed in the absence of NO<sub>x</sub> or NH<sub>3</sub>. Although inorganic aerosol (i.e., NH<sub>4</sub>NO<sub>3</sub>) was formed from the interaction of NH<sub>3</sub> and NO<sub>x</sub> in the chamber, the upper limit of the inorganic matter only accounted for 6.6 % of the total mass of particulate matter (Table S1) in the NH<sub>3</sub> + NO<sub>x</sub> experiment. Therefore, NH<sub>4</sub>NO<sub>3</sub> was not the main cause of the increase in particulate matter; the co-occurrence of NH<sub>3</sub> and NO<sub>x</sub> had a synergistic effect on the SOA formation, because their combined effect on SOA formation was greater than the sum of their separate effects. Qi et al. (2020) found that the promotion of NH<sub>3</sub> on toluene SOA formation was more obviously under high NO<sub>x</sub> concentration; SOA yield increased 3.7 % and 4.6 % for 70 and 160 ppb initial NO<sub>x</sub> concentration, respectively, when 200 ppb NH<sub>3</sub> was added into the chamber. Li et al. (2018) showed that the presence NH<sub>3</sub> can promote the particle size growth of SOA; at the same time, this particle growth rate was higher under low VOC / NO<sub>x</sub> (or high-NO<sub>x</sub>) conditions. All in all, the joint effect of multiple environmental factors on SOA formation is not the simple summation of the influences of various factors on SOA formation. This may at least partly explain why predictions of SOA concentrations in large-scale atmospheric models, which typically describe SOA formation from data derived from chamber experiments, are frequently lower than field observations (Volkamer et al., 2006).

### 3.2 SOA chemical composition

The traditional SOA formation mechanism is based on the chemical compositions obtained through offline detection of the chemical composition of SOA (Jang et al., 2002; Liu et al., 2019a; Xu et al., 2020). SOA is continually evolving in the atmosphere, and the aging process of SOA co-occurs with its formation process, resulting in the transformation of SOA chemical composition continuously proceeding during the photooxidation process. Therefore, the AMS was used for online measurement of the SOA chemical composition and how the chemical composition evolved in the photooxidation process will be discussed in the following section.



**Figure 2.** The OS<sub>C</sub> values for the toluene SOA formed under different NH<sub>3</sub> / NO<sub>x</sub> conditions.

Chemical composition of SOA is very complex. The average carbon oxidation state (OS<sub>C</sub>) has been shown to be an ideal conceptual framework to describe changes in the degree of oxidation undergone by SOA (Kroll et al., 2011) and has been widely applied in field and laboratory studies (Chen et al., 2018; Mandariya et al., 2019). Average OS<sub>C</sub> calculation is shown in the Supplement. Figure 2 shows the changes in the OS<sub>C</sub> of toluene SOA formed in different experiments. Notably, toluene SOA OS<sub>C</sub> values were in the range between  $-0.5$  and  $0$ , which is consistent with that of semi-volatile oxygenated organic aerosols (SV-OOAs). However, the different OS<sub>C</sub> values and the change trends observed for the toluene SOA formed in different conditions (with and without NH<sub>3</sub> / NO<sub>x</sub>) in Fig. 2 indicated that there was a photooxidation mechanism active during SOA formation, which ultimately changed the SOA chemical compositions.

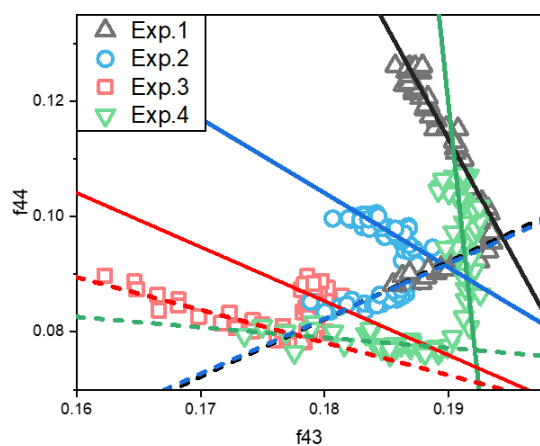
The OS<sub>C</sub> increased over time for all SOAs that were formed in the absence of NH<sub>3</sub>. There are several possible reasons for the increasing trend of OS<sub>C</sub> values. Firstly, a dynamic equilibrium of semi-volatile vapors may have been achieved between the particle phase and gas phase during the earlier toluene oxidation process. The increase in SOA led to a reduction in the concentration of gas-phase semi-volatile organic products. A decreasing concentration of gas-phase semi-volatile organic compound products would suppress their transformation from gas phase to particulate phase. More lower-volatility gas-phase oxidation products with higher OS<sub>C</sub> values would then be shifted to the particle phase, which would be responsible for the continuing increase in SOA and its OS<sub>C</sub>. Secondly, the formed SOA could have further been oxidized by OH radicals through heterogeneous reactions (Kourchev et al., 2015; Liu et al., 2019b). This could be the main reason for the increase in the OS<sub>C</sub> when the SOA concentration was no longer increasing. Finally, as pointed out by Malecha and Nizkorodov (2016), even if there was no OH radical in the chamber, the photodegradation of SOA can produce small oxygenated volatile organic compounds (e.g., acetaldehyde OS<sub>C</sub> =  $-1$ , and acetone OS<sub>C</sub>  $\approx -1.3$ ) under UV light irradiation. The photo-

production of OVOCs (oxygenated VOCs) from SOA had a lower OS<sub>C</sub> value than that of SOA. Although the loss of SOA through photodegradation is small, the OS<sub>C</sub> value of SOA still had increased to a certain extent (Malecha and Nizkorodov, 2016).

The fact that additional photochemical processing results in the dynamic evolution of the OS<sub>C</sub> over time has been demonstrated in both field and laboratory experiments (Jimenez et al., 2009). The atmospheric oxidation of OA tends towards higher OS<sub>C</sub> regardless of the original OA source (Herndon et al., 2008). However, when NH<sub>3</sub> was present, the OS<sub>C</sub> of total SOA was almost unchanged for the whole photooxidation period. Carboxyl and carbonyl are the main oxygen-containing functional groups responsible for the toluene photooxidation production (Ji et al., 2017). An organic ammonium salt with four H atoms can offset an increase in OS<sub>C</sub> value caused by the formation of organic acids or carboxy group with two O atoms through acid–base reactions (Kuwata and Martin, 2012; Liu et al., 2015). In addition, NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> may react with carbonyl functional groups through Maillard reactions, consuming the oxygen atom in the carbonyl group and leading to the formation of species with covalent carbon–nitrogen bonds (Lee et al., 2013; Zhang et al., 2015a; Qi et al., 2020). Xu et al. (2018) showed that imidazole compounds (OS<sub>C</sub>  $\approx -1.3$ ) generated through heterogeneous reaction between NH<sub>3</sub> and carbonyl compounds might contribute to the decrease in the OS<sub>C</sub> of SOA. It is clear that an increase in OS<sub>C</sub> caused by the formation of oxygen-containing functional groups (carboxyl, carbonyl, etc.) would be counteracted through acid–base reactions or Maillard reactions in the presence of NH<sub>3</sub>. After 60 min of UV light irradiation, there was no more SOA formation; however, the OS<sub>C</sub> did decrease slightly in Exp. 2 and 3, illustrating that the NH<sub>3</sub> could continue to react with SOA through heterogeneous processes. Huang et al. (2016) also pointed out that the portion of semi-volatile products with low OS<sub>C</sub> formed at the later stage of photooxidation also contributed to the decreased OS<sub>C</sub>.

The OS<sub>C</sub> of the toluene SOA formed with NO<sub>x</sub> was lower than that formed in the absence of NO<sub>x</sub>, no matter whether NH<sub>3</sub> was present in the chamber or not. This indicated that an increased NO<sub>x</sub> concentration benefits the formation of high volatility oxidation products with lower OS<sub>C</sub> values (Kroll et al., 2011; Jimenez et al., 2009). However, the relationships between OS<sub>C</sub> and SOA mass concentration with and without NH<sub>3</sub> were opposite to each other. Predictably, the SOA formation mechanism in the presence of NO<sub>x</sub> is different from that with NO<sub>x</sub> + NH<sub>3</sub>. In the absence of NH<sub>3</sub>, the RO intermediate, which is easily fragmented to produce relatively high-volatility compounds, was the dominant product of the NO<sub>x</sub> = RO<sub>2</sub> reaction (Zhao et al., 2018; Liu et al., 2019a; Xu et al., 2020). Highly volatile compounds cannot readily partition into the particle phase, which subsequently results in a lower SOA yield in the presence of NO<sub>x</sub> (Yang et al., 2020). Thereby, both OS<sub>C</sub> and the SOA mass concentration were





**Figure 3.** The relationship between total organic signals at 43  $m/z$  ( $f_{43}$ ) vs. 44  $m/z$  ( $f_{44}$ ) from SOA data during the photooxidation process. The  $f_{43}$  vs.  $f_{44}$  plots exhibited inflection points during the photooxidation process. The dashed lines indicate the trends of  $f_{43}$  vs.  $f_{44}$  for the SOA formation stage (before the inflection point), and the solid lines are for the stable stage.

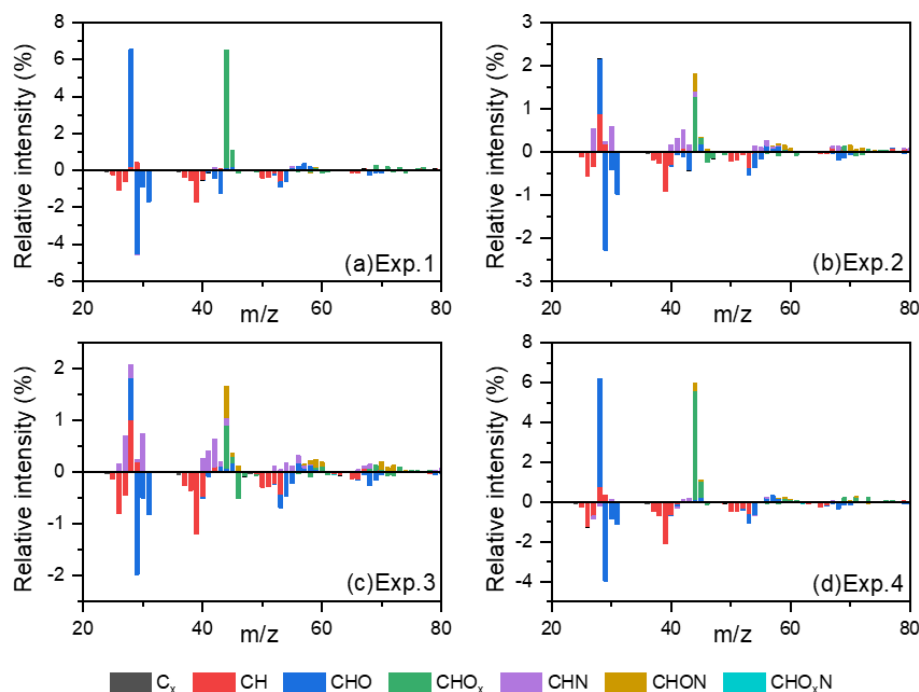
lower when the 60 ppb of NO<sub>x</sub> that was added into the chamber. However, when both NO<sub>x</sub> and NH<sub>3</sub> were present, the toluene-derived SOA had the lowest OS<sub>C</sub> value, but the highest mass concentration. This result suggested that although NO<sub>x</sub> promotes the formation of higher-volatility compounds, these higher-volatility compounds (e.g., glyoxal) can react with NH<sub>3</sub> and partition into the particle phase, which could contribute to the increase in SOA formation. Huffman et al. (2009) observed that aerosol volatility was inversely correlated with the extent of oxidation of OA components. The low value of OS<sub>C</sub> in the presence of NO<sub>x</sub> indicated that NO<sub>x</sub> would promote the formation of the relatively high-volatility compounds. However, the lower OS<sub>C</sub> value in the presence of NH<sub>3</sub> indicated that the high-volatility compounds would promote partitioning into the particle phase when reacting with NH<sub>3</sub>.

Fragments derived from the AMS data have also been widely used to explore the bulk compositions and properties of SOA (Ng et al., 2010, 2017). The  $m/z$  43 ( $f_{43}$ ) frequency was dominated by ion C<sub>2</sub>H<sub>3</sub>O<sup>+</sup>, which is the tracer for organic compounds with alcohol and carbonyl functional groups (Alfarra et al., 2006). Meanwhile, the  $m/z$  44 ( $f_{44}$ ) signal was dominated by CO<sub>2</sub><sup>+</sup> ions, which is the tracer for organic compounds with carboxyl functional groups and an indicator of highly oxygenated organic aerosols (Ng et al., 2010). Here, we use the approach of Ng et al. (2010) by plotting the fractions of the total organic signal at  $m/z$  43 vs.  $m/z$  44 ( $f_{43}$  vs.  $f_{44}$ ). The change of  $f_{43}$  vs.  $f_{44}$ , which has an inflection point during the photooxidation process, is shown in Figs. 3 and S3. In our study, the change before the inflection point was defined as the formation stage, and the linear fit of  $f_{43}$  vs.  $f_{44}$  for the formation stage is shown by the dashed lines. The change in  $f_{43}$  vs.  $f_{44}$  after the inflection

point was defined as the stable stage, and the linear fit of  $f_{43}$  vs.  $f_{44}$  in this stage is shown by the solid lines. The formation and stable stages of the  $f_{43}$  vs.  $f_{44}$  relationship during the experiment are discussed separately here.

In the stable stage, the increase in  $f_{44}$  and decrease in  $f_{43}$  with increasing OH exposure indicated that the carbonyl groups in toluene SOA were oxidized to carboxyl groups by the aging process. For the experiments without NH<sub>3</sub> and NO<sub>x</sub>, the slope ratio of  $f_{43}$  vs.  $f_{44}$  was  $-3.9$ . When there was 60 ppb initial NO<sub>x</sub>, the  $f_{43}$  was almost stable, while the  $f_{44}$  increased with the oxidation process. There was a lower slope ratio of  $f_{43}$  vs.  $f_{44}$ , indicating that organic compounds with more alcohol and carbonyl functional groups had formed in the presence of NO<sub>x</sub>. But for the experiments with 200 ppb initial NH<sub>3</sub>, the slope ratios of  $f_{43}$  vs.  $f_{44}$  were only  $-1.1$  and  $-1.3$  in the presence and absence of NO<sub>x</sub>, respectively. According to the above results, we can see that more carbonyl groups are consumed as carboxyl groups are formed in the presence of NH<sub>3</sub>. The carbonyls can be oxidized to organic acids (Kawamura and Bikkina, 2016), but unreacted carbonyl can be nucleophilically attacked by NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> to form nitrogen-containing heterocyclic compounds, e.g., imidazole (Grace et al., 2019; Lian et al., 2020). Meanwhile, the peak  $f_{44}$  value decreased from 0.13 to 0.10 when NH<sub>3</sub> was added into the chamber. This suggested that the heterogeneous reaction of NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> could promote the consumption of particle-phase carbonyl groups (Xu et al., 2018) and must inhibit the formation of carboxyl groups in the SOA aging process. According to the changing trend of SOA concentration over time, the photooxidation process was divided into formation stage and stable stage. As shown in Fig. 1, the first half hour of photooxidation when SOA concentration increased linearly with time was defined as SOA formation stage. After 60 min of photooxidation, SOA concentration was not changed with reaction time, and it was defined as stable stage. The differences in spectra of toluene SOA in the formation stage and stable stage are shown in Fig. 4. A lower signal intensity variation in CO<sub>2</sub><sup>+</sup> in the presence of NH<sub>3</sub> also illustrated that NH<sub>3</sub> would inhibit heterogeneous reactions that form carboxyl groups.

In the formation stage, the slope ratios of  $f_{43}$  vs.  $f_{44}$  were almost the same for both experiments without NO<sub>x</sub>. It can thus be seen that the presence or absence of NH<sub>3</sub> does not affect the change trend of  $f_{43}$  vs.  $f_{44}$  in the SOA formation stage. Therefore, the gas-phase homogeneous reaction of NH<sub>3</sub> on SOA formation is not important. Clearly, the particle-phase heterogeneous reaction was the main reaction pathway by which NH<sub>3</sub> participated in the photooxidation process and toluene SOA formation. However, negative correlations were observed between  $f_{43}$  and  $f_{44}$  in the presence of NO<sub>x</sub>. Based on this, we concluded that NO<sub>x</sub> not only affects the SOA formation through the particle-phase heterogeneous reactions but also through gas-phase homogeneous reactions.



**Figure 4.** The differential spectra of toluene SOA in the formation and stable stages. Data were taken and analyzed at a high resolution but were summarized to a unit mass resolution for display. Only minimal N-containing fragments could be observed in Exp. 1 without added NH<sub>3</sub> and NO<sub>x</sub>. These N-containing fragments could be attributed to the background NH<sub>3</sub> and NO<sub>x</sub> in the chamber or systematic errors from AMS.

### 3.3 PMF results

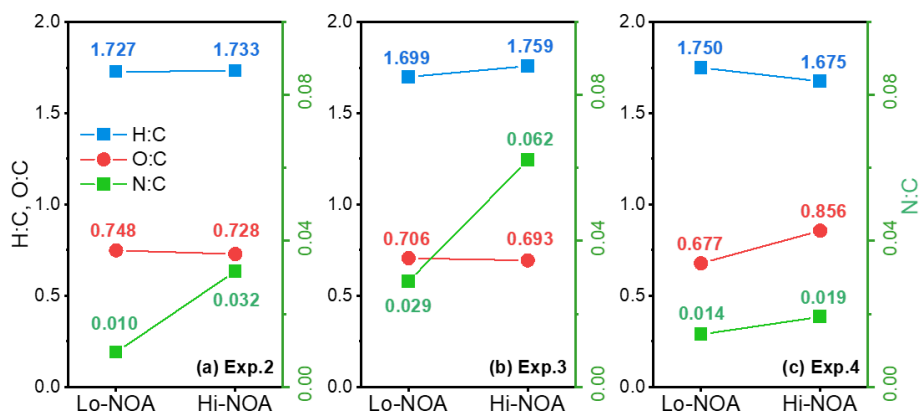
A temporal evaluation of the toluene SOA chemical composition during photooxidation is vital to the analysis of the NOA formation mechanism in the presence of NH<sub>3</sub> and/or NO<sub>x</sub>. Therefore, this study further compared the chemical properties of the SOA generated under different experimental conditions by applying a PMF analysis to the HR-ToF-AMS data (Chen et al., 2019). A summary of the PMF results is presented in Figs. S4–S7. For the toluene OH-photooxidation experiments with NH<sub>3</sub> and/or NO<sub>x</sub> present, two factors were identified from the PMF analysis in the same way as in Chen et al. (2019). The H/C, O/C, and N/C values of these two factors are shown in Fig. 5. The factor with higher N/C values was defined as high-nitrogen OA (Hi-NOA). In contrast, the factor with lower N/C values was defined as low-nitrogen OA (Lo-NOA). Figure 6 exhibits the evolution of Hi-NOA and Lo-NOA during the photooxidation process as resolved from the PMF analysis of different initial NO<sub>x</sub> / NH<sub>3</sub> concentrations. While similar evolutionary trends were observed under different conditions, the relative intensities and the chemical compositions of these two factors in each experiment were not consistent.

For the toluene SOA formed under NH<sub>3</sub> conditions, both Lo-NOA and Hi-NOA had similar O/C values, which were fully oxygenated with an average of  $0.74 \pm 0.04$  (Fig. 5a). These O/C values were comparable to the low-volatility

oxygenated organic aerosols (LV-OOAs) with an O/C value ranging from 0.6 to 1 (Jimenez et al., 2009). The main difference between these two OA sources was the N/C ratio. The N/C ratio of Hi-NOA (N/C = 0.032) was about 3 times higher than that of Lo-NOA (N/C = 0.010) (Fig. 5a). The evolution of these two OA sources during the photooxidation process is shown in Fig. 6a. The components of toluene SOA were mostly Lo-NOA during the initial phase of SOA formation, but Hi-NOA toluene SOA started forming after 1 min and continued to increase. The Lo-NOA reached the maximum mass concentration after 30 min of the photooxidation, and then decreased. Such a declining trend of Lo-NOA at longer reaction times reflected the conversion of Lo-NOA into something else in the particle phase. As the Lo-NOA decreased, the mass concentration of Hi-NOA gradually increased. Thus, the Hi-NOA should be derived from the heterogeneous reaction of Lo-NOA with NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup>. At the same time, it was proved that the formation pathway of Hi-NOA was not through reaction of NH<sub>3</sub> with later-generation gas-phase products in the homogeneous gas phase. With the gradual replacement of Lo-NOA by Hi-NOA, the ratio of [Hi-NOA] / [Lo-NOA] stabilized at 5–6.

For the toluene SOA formed under NO<sub>x</sub> conditions, there was not a large difference between the N/C ratios of Hi-NOA (N/C = 0.019) and Lo-NOA (N/C = 0.014) (Fig. 5c). At the end of the NO<sub>x</sub> experiment, the ratio of [Hi-NOA] / [Lo-NOA] was only 3 : 2 (Fig. 6c). It follows that the contribu-





**Figure 5.** The H/C, O/C, and N/C values of Hi-NOA and Lo-NOA for each experiment. (a) Exp. 2 with 200 ppb NH<sub>3</sub>, (b) Exp. 3 with 200 ppb NH<sub>3</sub> and 62 ppb NO<sub>2</sub>, and (c) Exp. 4 with 63 ppb NO<sub>2</sub>.

tion of the heterogeneous NO<sub>x</sub> reaction to the N/C ratio of toluene SOA was not obvious. Therefore, the formation of NOA in the presence of NO<sub>x</sub> mainly occurred through gas-phase homogeneous reactions, which was consistent with the results in Sect. 3.3.

The changing trend of N/C with time in the presence of NH<sub>3</sub> was different to that with NO<sub>x</sub> present. The evolution trends of the N/C of SOA in different experiments are shown in Fig. 7. In the presence of NH<sub>3</sub>, the N/C value gradually increased throughout the photooxidation process. The increased N/C value in the photooxidation process was attributed to the heterogeneous NH<sub>3</sub> reaction with SOA. But in the presence of NO<sub>x</sub>, the N/C increased rapidly to its maximum value where it was stable for the rest of the reaction. This could mean that the heterogeneous reaction of toluene SOA with NO<sub>x</sub> to form NOA was not as important as the gas-phase homogeneous reaction.

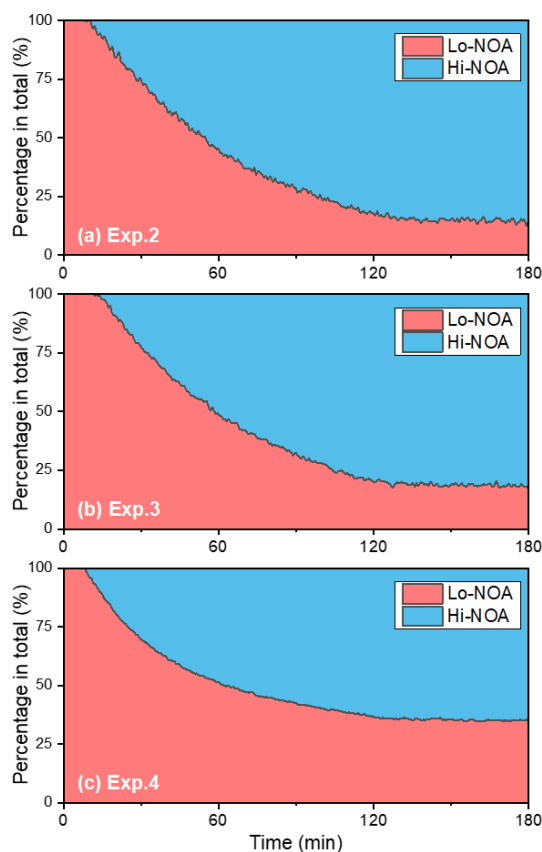
When both NO<sub>x</sub> and NH<sub>3</sub> were added into the chamber, the N/C ratios of Hi-NOA and Lo-NOA were 0.062 and 0.029, respectively (Fig. 5b). The N/C ratio of Hi-NOA, which was comparable to the recently isolated nitrogen-enriched OA value (0.053) observed by Sun et al. (2011), was much higher than that observed in the experiments with only NH<sub>3</sub> or NO<sub>x</sub>. It was even higher than the sum of the N/C ratios from both Exp. 2 with NH<sub>3</sub> and Exp. 4 with NO<sub>x</sub>. In order to calculate the relative contributions of NH<sub>3</sub> and NO<sub>x</sub> to N/C, it was assumed that the effects of NH<sub>3</sub> or NO<sub>x</sub> on the N/C ratio in the Hi-NOA and Lo-NOA factors did not change among different experimental conditions. For Lo-NOA, the contributions of NH<sub>3</sub> and NO<sub>x</sub> to the N/C value were 0.0126 and 0.0164, and their relative intensities were 43 % and 57 %, respectively. For Hi-NOA, the contributions of NH<sub>3</sub> and NO<sub>x</sub> to the N/C values were 0.0404 and 0.0216, and their relative intensities were 65 % and 35 %, respectively. For the experiment with both NH<sub>3</sub> and NO<sub>x</sub>, the contribution of NH<sub>3</sub> to N/C was higher by 26 %, and the contribution of NO<sub>x</sub> to N/C was higher by 17 % compared to the experiments with

the single pollutants. The co-existence of NH<sub>3</sub> and NO<sub>x</sub> further enhanced the N/C value of toluene SOA, indicating that a synergetic interaction between NH<sub>3</sub> and NO<sub>x</sub> further enhanced organic nitrogen formation.

### 3.4 Optical absorption

The optical characteristics of toluene SOA formed from different NH<sub>3</sub> and NO<sub>x</sub> conditions were investigated. The MAC of toluene-derived SOA detected over the range of 230–600 nm is displayed in Fig. 8. Over the entire UV detection range, an increase in light absorption was observed when the toluene SOA formed in the presence of NO<sub>x</sub> or NH<sub>3</sub>.

By looking at Fig. 8 in detail, one can see that the MAC of toluene SOA formed with (red line) and without (black line) NH<sub>3</sub> overlapped at 250 nm, but when the UV wavelength exceeded 250 nm the MAC of the toluene SOA formed in the presence of NH<sub>3</sub> was higher. The red line reflects an obvious characteristic absorption peak at 270–280 nm, which was mainly due to the absorption of the  $n \rightarrow \pi^*$  electronic transitions. The imidazole compounds were formed through the Maillard reactions between NH<sub>3</sub> / NH<sub>4</sub><sup>+</sup> with carbonyl functional groups (Zhang et al., 2015a). The C=N double bonds in the organonitrogen imidazole compounds can act as effective chromophores since both  $\pi \rightarrow \pi^*$  and  $n \rightarrow \pi^*$  transitions are chromatically active (Nguyen et al., 2013). The UV–Vis spectra of imine and pyrrole show broad bands at 270 nm (NIST, 2020), which was consistent with the UV absorption peak of the  $n \rightarrow \pi^*$  band observed here. According to the AMS results, carbonyl was the main functional group of toluene SOA. The emergence of absorption peaks at 270–280 nm demonstrated that some organonitrogen imidazole compounds (e.g., imines and pyrrole) were formed through the heterogeneous reaction of toluene with NH<sub>3</sub>. Meanwhile, the high-molecular-weight nitrogen-containing organic species might have formed through Maillard reactions in the particle phase (Wang et al., 2010). This was also

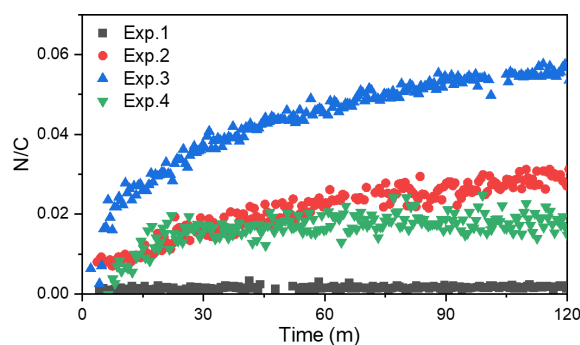


**Figure 6.** The evolution of high-nitrogen OA (Hi-NOA) and low-nitrogen OA (Lo-NOA) during the photooxidation process under different NO<sub>x</sub> / NH<sub>3</sub> concentrations. Hi-NOA and Lo-NOA were not consistent among experiments. (a) Exp. 2 with 200 ppb NH<sub>3</sub>, (b) Exp. 3 with 200 ppb NH<sub>3</sub> and 62 ppb NO<sub>2</sub>, and (c) Exp. 4 with 63 ppb NO<sub>2</sub>.

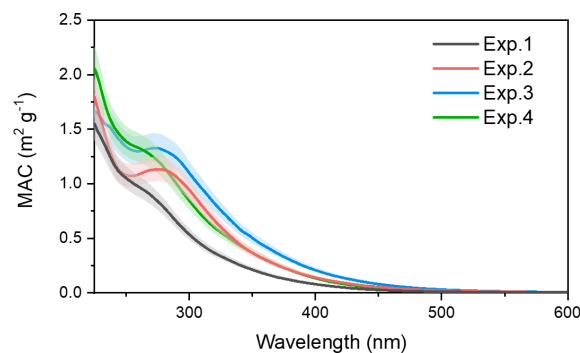
a reason for the increase in SOA mass concentration in the presence of NH<sub>3</sub>.

The green line in Fig. 8 represents the MAC value of toluene-derived SOA in the presence of NO<sub>x</sub>, which was also higher than the black line (control) throughout the UV detection range. When compared with the red line, the green line had no obvious characteristic peak at 280 nm, but it had higher absorbance in the range between 240 and 280 nm. This indicated that both NO<sub>x</sub> and NH<sub>3</sub> increased the absorbance of toluene SOA, while the chromophores generated from the reactions between toluene-derived SOA with NH<sub>3</sub> or NO<sub>x</sub> did not behave in the same way.

The blue line in Fig. 8 represents the absorbance of toluene SOA formed in the presence of both NO<sub>x</sub> and NH<sub>3</sub>. The MAC of toluene SOA formed in the presence of both NO<sub>x</sub> and NH<sub>3</sub> was higher than the toluene SOA formed in the presence of either NH<sub>3</sub> or NO<sub>x</sub>. There might have been a synergetic effect between NO<sub>x</sub> and NH<sub>3</sub> on the absorbance of toluene SOA. Considering that the mass concentration of toluene SOA formed in the presence of both NH<sub>3</sub> and NO<sub>x</sub>



**Figure 7.** The evolution of N/C in different experiments.



**Figure 8.** The MAC over the range of 230–600 nm for the toluene SOA formed under different experiment conditions.

was the highest, as described in Sect. 3.1, the co-existence of NH<sub>3</sub> and NO<sub>x</sub> may also result in the toluene SOA having stronger light absorption and atmospheric radiative forcing. We also noted a higher MAC value at 280 nm, which illustrated that the presence of NO<sub>x</sub> could promote the formation of imines and pyrrole in the photooxidation system of toluene with NH<sub>3</sub>.

#### 4 Conclusions

Here we present the results of a study in which we characterized the mass concentrations, chemical compositions, and optical properties of SOA formed from the photooxidation of toluene under different NH<sub>3</sub> and NO<sub>x</sub> concentrations. When compared with the control experiment, the SOA mass concentration data showed that the formation of toluene-derived SOA was enhanced in the presence of NH<sub>3</sub>, through acid–base reactions between carboxyl groups or Maillard reactions with carbonyls, but inhibited in the presence of NO<sub>x</sub>. Meanwhile, the mass concentration of toluene SOA formed in the presence of both NO<sub>x</sub> and NH<sub>3</sub> was higher than those formed under either NH<sub>3</sub> or NO<sub>x</sub> alone. This result indicated that there was a synergistic interaction between NH<sub>3</sub> and NO<sub>x</sub> that further enhanced toluene-derived SOA formation. At the same time, the lowest OS<sub>C</sub> value was obtained when

both NH<sub>3</sub> and NO<sub>x</sub> were present. We concluded that highly volatile compounds, which were formed from toluene photooxidation in the presence of NO<sub>x</sub>, could react with NH<sub>3</sub> to form products with lower volatilities and promoted the partitioning of these products into the particle phase.

Synergetic effects of NH<sub>3</sub> and NO<sub>x</sub> on the formation of NOA and the optical properties of SOA were also observed in this study. The heterogeneous reaction was responsible for the formation of NOA in the presence of NH<sub>3</sub>. Meanwhile, an absorption peak at 270–280 nm, which is characteristic of imine and pyrrole, was observed. In contrast, the formation of NOA caused by NO<sub>x</sub> alone was mainly due to a gas-phase homogeneous reaction.

In the actual atmosphere, especially in Chinese urban atmosphere, NO<sub>x</sub> and NH<sub>3</sub> abundantly co-exist. Therefore, the findings presented here clearly show that the synergetic effects of NO<sub>x</sub> and NH<sub>3</sub> should not be neglected. In the meantime, our work provides a scientific basis for the consideration of synergistic emission reductions of NH<sub>3</sub> and NO<sub>x</sub> under the compound pollution conditions, which will contribute to reducing the burden of aerosols in the atmosphere. It has to be noted that the concentration of reactants used for the experiments is much higher than that observed in polluted areas. Although the reactant concentrations including NH<sub>3</sub> used in this work are much higher than those in the real urban environment, our results are applicable for the polluted urban atmosphere. In the urban atmosphere aromatic VOCs consist of numerous species, and their total concentration is much higher than a single species such as toluene. On the other hand, carboxylic acids and carbonyls in the urban polluted atmosphere can be produced from aromatics and many other species. Therefore, it is reasonable for our smog chamber experiments to use toluene as a single precursor with a concentration much higher than that in the real atmosphere. Although the mechanisms of SOA formed under high precursor concentrations is expected to be the same as that under low concentrations, the kinetics are probably different. Thus, the effect of NH<sub>3</sub> and NO<sub>x</sub> on the photooxidation of toluene with lower concentrations should be checked in a future study.

**Data availability.** The datasets are available at <https://doi.org/10.6084/m9.figshare.16910953> (last access: 26 November 2021, Liu, 2021.).

**Supplement.** The supplement related to this article is available online at: <https://doi.org/10.5194/acp-21-17759-2021-supplement>.

**Author contributions.** SL and GW designed the experiment. SL, DH, and YW conducted the experiments. SL, DH, YW, and GW performed the data interpretation. SL and GW wrote the paper. YW,

SZ, XL, CW, and WD contributed to the paper with useful scientific discussions or comments.

**Competing interests.** The contact author has declared that neither they nor their co-authors have any competing interests.

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## References

- Alfarra, M. R., Paulsen, D., Gysel, M., Garforth, A. A., Dommen, J., Prévôt, A. S. H., Worsnop, D. R., Baltensperger, U., and Coe, H.: A mass spectrometric study of secondary organic aerosols formed from the photooxidation of anthropogenic and biogenic precursors in a reaction chamber, *Atmos. Chem. Phys.*, 6, 5279–5293, <https://doi.org/10.5194/acp-6-5279-2006>, 2006.
- Babar, Z. B., Park, J.-H., and Lim, H.-J.: Influence of NH<sub>3</sub> on secondary organic aerosols from the ozonolysis and photooxidation of  $\alpha$ -pinene in a flow reactor, *Atmos. Environ.*, 164, 71–84, <https://doi.org/10.1016/j.atmosenv.2017.05.034>, 2017.
- Bates, K., Jacob, D., Li, K., Ivatt, P., Evans, M., Yan, Y., and Lin, J.: Development and evaluation of a new compact mechanism for aromatic oxidation in atmospheric models, *Atmos. Chem. Phys. Discuss.* [preprint], <https://doi.org/10.5194/acp-2021-605>, in review, 2021.
- Berkemeier, T., Ammann, M., Mentel, T. F., Poschl, U., and Shiraiwa, M.: Organic nitrate contribution to new particle formation and growth in secondary organic aerosols from  $\alpha$ -pinene ozonolysis, *Environ. Sci. Technol.*, 50, 6334–6342, <https://doi.org/10.1021/acs.est.6b00961>, 2016.
- Chen, Y. and Bond, T. C.: Light absorption by organic carbon from wood combustion, *Atmos. Chem. Phys.*, 10, 1773–1787, <https://doi.org/10.5194/acp-10-1773-2010>, 2010.
- Chen, C. L., Li, L. J., Tang, P., and Cocker, D. R.: SOA formation from photooxidation of naphthalene and methylnaphthalenes with m-xylene and surrogate mixtures, *Atmos. Environ.*, 180, 256–264, <https://doi.org/10.1016/j.atmosenv.2018.02.051>, 2018.
- Chen, T., Liu, Y., Ma, Q., Chu, B., Zhang, P., Liu, C., Liu, J., and He, H.: Significant source of secondary aerosol: for-

- mation from gasoline evaporative emissions in the presence of SO<sub>2</sub> and NH<sub>3</sub>, *Atmos. Chem. Phys.*, 19, 8063–8081, <https://doi.org/10.5194/acp-19-8063-2019>, 2019.
- Chu, B., Zhang, X., Liu, Y., He, H., Sun, Y., Jiang, J., Li, J., and Hao, J.: Synergetic formation of secondary inorganic and organic aerosol: effect of SO<sub>2</sub> and NH<sub>3</sub> on particle formation and growth, *Atmos. Chem. Phys.*, 16, 14219–14230, <https://doi.org/10.5194/acp-16-14219-2016>, 2016.
- de Foy, B., Lu, Z., and Streets, D. G.: Satellite NO<sub>2</sub> retrievals suggest China has exceeded its NO<sub>x</sub> reduction goals from the twelfth Five-Year Plan, *Sci. Rep.*, 6, 35912, <https://doi.org/10.1038/srep35912>, 2016.
- Draper, D. C., Farmer, D. K., Desyaterik, Y., and Fry, J. L.: A qualitative comparison of secondary organic aerosol yields and composition from ozonolysis of monoterpenes at varying concentrations of NO<sub>2</sub>, *Atmos. Chem. Phys.*, 15, 12267–12281, <https://doi.org/10.5194/acp-15-12267-2015>, 2015.
- Erisman, J. W. and Schaap, M.: The need for ammonia abatement with respect to secondary PM reductions in Europe, *Environ. Pollut.*, 129, 159–163, <https://doi.org/10.1016/j.envpol.2003.08.042>, 2004.
- Ge, S., Wang, G., Zhang, S., Li, D., Xie, Y., Wu, C., Yuan, Q., Chen, J., and Zhang, H.: Abundant NH<sub>3</sub> in China enhances atmospheric HONO production by promoting the heterogeneous reaction of SO<sub>2</sub> with NO<sub>2</sub>, *Environ. Sci. Technol.*, 53, 14339–14347, <https://doi.org/10.1021/acs.est.9b04196>, 2019.
- Grace, D. N., Sharp, J. R., Holappa, R. E., Lugos, E. N., Sebold, M. B., Griffith, D. R., Hendrickson, H. P., and Galloway, M. M.: Heterocyclic product formation in aqueous brown carbon systems, *ACS Earth Space Chem.*, 3, 2472–2481, <https://doi.org/10.1021/acsearthspacechem.9b00235>, 2019.
- Herndon, S. C., Onasch, T. B., Wood, E. C., Kroll, J. H., Canagaratna, M. R., Jayne, J. T., Zavala, M. A., Knighton, W. B., Mazzoleni, C., Dubey, M. K., Ulbrich, I. M., Jimenez, J. L., Seila, R., de Gouw, J. A., de Foy, B., Fast, J., Molina, L. T., Kolb, C. E., and Worsnop, D. R.: Correlation of secondary organic aerosol with odd oxygen in Mexico City, *Geophys. Res. Lett.*, 35, L15804, <https://doi.org/10.1029/2008gl034058>, 2008.
- Huang, D. D., Zhang, X., Dalleska, N. F., Lignell, H., Coggon, M. M., Chan, C. M., Flagan, R. C., Seinfeld, J. H., and Chan, C. K.: A note on the effects of inorganic seed aerosol on the oxidation state of secondary organic aerosol- $\alpha$ -Pinene ozonolysis, *J. Geophys. Res.-Atmos.*, 121, 12476–12483, <https://doi.org/10.1002/2016jd025999>, 2016.
- Huang, Y., Lee, S. C., Ho, K. F., Ho, S. S. H., Cao, N. Y., Cheng, Y., and Gao, Y.: Effect of ammonia on ozone-initiated formation of indoor secondary products with emissions from cleaning products, *Atmos. Environ.*, 59, 224–231, <https://doi.org/10.1016/j.atmosenv.2012.04.059>, 2012.
- Huffman, J. A., Docherty, K. S., Mohr, C., Cubison, M. J., Ulbrich, I. M., Ziemann, P. J., Onasch, T. B., and Jimenez, J. L.: Chemically-resolved volatility measurements of organic aerosol from different sources, *Environ. Sci. Technol.*, 43, 5351–5357, <https://doi.org/10.1021/es803539d>, 2009.
- Jang, M., Czoschke, N. M., Lee, S., and Kamens, R. M.: Heterogeneous atmospheric aerosol production by acid-catalyzed particle-phase reactions, *Science*, 298, 814–817, <https://doi.org/10.1126/science.1075798>, 2002.
- Ji, Y., Zhao, J., Terazono, H., Misawa, K., Levitt, N. P., Li, Y., Lin, Y., Peng, J., Wang, Y., Duan, L., Pan, B., Zhang, F., Feng, X., An, T., Marrero-Ortiz, W., Secrest, J., Zhang, A. L., Shibuya, K., Molina, M. J., and Zhang, R.: Reassessing the atmospheric oxidation mechanism of toluene, *P. Natl. Acad. Sci. USA*, 114, 8169–8174, <https://doi.org/10.1073/pnas.1705463114>, 2017.
- Jiang, X. T., Lv, C., You, B., Liu, Z. Y., Wang, X. F., and Du, L.: Joint impact of atmospheric SO<sub>2</sub> and NH<sub>3</sub> on the formation of nanoparticles from photo-oxidation of a typical biomass burning compound, *Environ. Sci.-Nano*, 7, 2532–2545, <https://doi.org/10.1039/d0en00520g>, 2020.
- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525–1529, <https://doi.org/10.1126/science.1180353>, 2009.
- Kawamura, K. and Bikkina, S.: A review of dicarboxylic acids and related compounds in atmospheric aerosols: Molecular distributions, sources and transformation, *Atmos. Res.*, 170, 140–160, <https://doi.org/10.1016/j.atmosres.2015.11.018>, 2016.
- Kourtchev, I., Doussin, J.-F., Giorio, C., Mahon, B., Wilson, E. M., Maurin, N., Pangu, E., Venables, D. S., Wenger, J. C., and Kalberer, M.: Molecular composition of fresh and aged secondary organic aerosol from a mixture of biogenic volatile compounds: a high-resolution mass spectrometry study, *Atmos. Chem. Phys.*, 15, 5683–5695, <https://doi.org/10.5194/acp-15-5683-2015>, 2015.
- Krechmer, J. E., Day, D. A., and Jimenez, J. L.: Always lost but never forgotten: Gas-phase wall losses are important in all teflon environmental chambers, *Environ. Sci. Technol.*, 54, 12890–12897, <https://doi.org/10.1021/acs.est.0c03381>, 2020.
- Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a metric for describing the chemistry of atmospheric organic aerosol, *Nat. Chem.*, 3, 133–139, <https://doi.org/10.1038/nchem.948>, 2011.
- Kuwata, M. and Martin, S. T.: Phase of atmospheric secondary organic material affects its reactivity, *P. Natl. Acad. Sci. USA*, 109, 17354–17359, <https://doi.org/10.1073/pnas.1209071109>, 2012.
- Laskin, A., Laskin, J., and Nizkorodov, S. A.: Chemistry of atmospheric brown carbon, *Chem. Rev.*, 115, 4335–4382, <https://doi.org/10.1021/cr5006167>, 2015.
- Laskin, J., Laskin, A., Roach, P. J., Slysz, G. W., Anderson, G. A., Nizkorodov, S. A., Bones, D. L., and Nguyen, L. Q.: High-resolution desorption electrospray ionization mass spectrometry

- for chemical characterization of organic aerosols, *Anal. Chem.*, 82, 2048–2058, <https://doi.org/10.1021/ac902801f>, 2010.
- Laskin, J., Laskin, A., Nizkorodov, S. A., Roach, P., Eckert, P., Gilles, M. K., Wang, B., Lee, H. J., and Hu, Q.: Molecular selectivity of brown carbon chromophores, *Environ. Sci. Technol.*, 48, 12047–12055, <https://doi.org/10.1021/es503432r>, 2014.
- Lee, A. K., Zhao, R., Li, R., Liggio, J., Li, S. M., and Abbatt, J. P.: Formation of light absorbing organo-nitrogen species from evaporation of droplets containing glyoxal and ammonium sulfate, *Environ. Sci. Technol.*, 47, 12819–12826, <https://doi.org/10.1021/es402687w>, 2013.
- Li, K., Chen, L., White, S. J., Yu, H., Wu, X., Gao, X., Azzi, M., and Cen, K.: Smog chamber study of the role of NH<sub>3</sub> in new particle formation from photo-oxidation of aromatic hydrocarbons, *Sci. Total Environ.*, 619, 927–937, <https://doi.org/10.1016/j.scitotenv.2017.11.180>, 2018.
- Lian, X., Zhang, G., Yang, Y., Lin, Q., Fu, Y., Jiang, F., Peng, L., Hu, X., Chen, D., Wang, X., Peng, P. a., Sheng, G., and Bi, X.: Evidence for the formation of imidazole from carbonyls and reduced nitrogen species at the individual particle level in the ambient atmosphere, *Environ. Sci. Tech. Lett.*, 8, 9–15, <https://doi.org/10.1021/acs.estlett.0c00722>, 2021.
- Liu, S.: Synergetic effect of NH<sub>3</sub> and NO<sub>x</sub> on SOA [data set], available at: <https://doi.org/10.6084/m9.figshare.16910953>, last access: 26 November 2021.
- Liu, M., Huang, X., Song, Y., Xu, T., Wang, S., Wu, Z., Hu, M., Zhang, L., Zhang, Q., Pan, Y., Liu, X., and Zhu, T.: Rapid SO<sub>2</sub> emission reductions significantly increase tropospheric ammonia concentrations over the North China Plain, *Atmos. Chem. Phys.*, 18, 17933–17943, <https://doi.org/10.5194/acp-18-17933-2018>, 2018.
- Liu, S., Jia, L., Xu, Y., Tsona, N. T., Ge, S., and Du, L.: Photooxidation of cyclohexene in the presence of SO<sub>2</sub>: SOA yield and chemical composition, *Atmos. Chem. Phys.*, 17, 13329–13343, <https://doi.org/10.5194/acp-17-13329-2017>, 2017.
- Liu, S. J., Jiang, X. T., Tsona, N. T., Lv, C., and Du, L.: Effects of NO<sub>x</sub>, SO<sub>2</sub> and RH on the SOA formation from cyclohexene photooxidation, *Chemosphere*, 216, 794–804, <https://doi.org/10.1016/j.chemosphere.2018.10.180>, 2019a.
- Liu, S. J., Tsona, N. T., Zhang, Q., Jia, L., Xu, Y. F., and Du, L.: Influence of relative humidity on cyclohexene SOA formation from OH photooxidation, *Chemosphere*, 231, 478–486, <https://doi.org/10.1016/j.chemosphere.2019.05.131>, 2019b.
- Liu, S. J., Wang, Y., Wang, G., Zhang, S., Li, D., Du, L., Wu, C., Du, W., and Ge, S.: Enhancing effect of NO<sub>2</sub> on the formation of light-absorbing secondary organic aerosols from toluene photooxidation, *Sci. Total Environ.*, 794, 148714, <https://doi.org/10.1016/j.scitotenv.2021.148714>, 2021.
- Liu, Y., Liggio, J., Staebler, R., and Li, S.-M.: Reactive uptake of ammonia to secondary organic aerosols: kinetics of organonitrogen formation, *Atmos. Chem. Phys.*, 15, 13569–13584, <https://doi.org/10.5194/acp-15-13569-2015>, 2015.
- Ma, P., Zhang, P., Shu, J., Yang, B., and Zhang, H.: Characterization of secondary organic aerosol from photo-oxidation of gasoline exhaust and specific sources of major components, *Environ. Pollut.*, 232, 65–72, <https://doi.org/10.1016/j.envpol.2017.09.018>, 2018a.
- Ma, Q., Lin, X., Yang, C., Long, B., Gai, Y., and Zhang, W.: The influences of ammonia on aerosol formation in the ozonolysis of styrene: Roles of Criegee intermediate reactions, *R. Soc. Open Sci.*, 5, 172171, <https://doi.org/10.1098/rsos.172171>, 2018b.
- Malecha, K. T., and Nizkorodov, S. A.: Photodegradation of secondary organic aerosol particles as a source of small, oxygenated volatile organic compounds, *Environ. Sci. Technol.*, 50, 9990–9997, <https://doi.org/10.1021/acs.est.6b02313>, 2016.
- Mandariya, A. K., Gupta, T., and Tripathi, S. N.: Effect of aqueous-phase processing on the formation and evolution of organic aerosol (OA) under different stages of fog life cycles, *Atmos. Environ.*, 206, 60–71, <https://doi.org/10.1016/j.atmosenv.2019.02.047>, 2019.
- Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of composition-dependent collection efficiencies for the aerodyne aerosol mass spectrometer using field data, *Aerosol Sci. Tech.*, 46, 258–271, <https://doi.org/10.1080/02786826.2011.620041>, 2012.
- Moise, T., Flores, J. M., and Rudich, Y.: Optical properties of secondary organic aerosols and their changes by chemical processes, *Chem. Rev.*, 115, 4400–4439, <https://doi.org/10.1021/cr5005259>, 2015.
- Na, K., Song, C., Switzer, C., and Cocker, D. R.: Effect of ammonia on secondary organic aerosol formation from  $\alpha$ -pinene ozonolysis in dry and humid conditions, *Environ. Sci. Technol.*, 41, 6096–6102, <https://doi.org/10.1021/es061956y>, 2007.
- Ng, N. L., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Kroll, J. H., Kwan, A. J., McCabe, D. C., Wennberg, P. O., Sorooshian, A., Murphy, S. M., Dalleska, N. F., Flagan, R. C., and Seinfeld, J. H.: Effect of NO<sub>x</sub> level on secondary organic aerosol (SOA) formation from the photooxidation of terpenes, *Atmos. Chem. Phys.*, 7, 5159–5174, <https://doi.org/10.5194/acp-7-5159-2007>, 2007a.
- Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.: Secondary organic aerosol formation from m-xylene, toluene, and benzene, *Atmos. Chem. Phys.*, 7, 3909–3922, <https://doi.org/10.5194/acp-7-3909-2007>, 2007b.
- Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H., Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets from Aerosol Mass Spectrometry, *Atmos. Chem. Phys.*, 10, 4625–4641, <https://doi.org/10.5194/acp-10-4625-2010>, 2010.
- Ng, N. L., Brown, S. S., Archibald, A. T., Atlas, E., Cohen, R. C., Crowley, J. N., Day, D. A., Donahue, N. M., Fry, J. L., Fuchs, H., Griffin, R. J., Guzman, M. I., Herrmann, H., Hodzic, A., Iinuma, Y., Jimenez, J. L., Kiendler-Scharr, A., Lee, B. H., Luecken, D. J., Mao, J., McLaren, R., Mutzel, A., Osthoff, H. D., Ouyang, B., Picquet-Varrault, B., Platt, U., Pye, H. O. T., Rudich, Y., Schwantes, R. H., Shiraiwa, M., Stutz, J., Thornton, J. A., Tilgner, A., Williams, B. J., and Zaveri, R. A.: Nitrate radicals and biogenic volatile organic compounds: oxidation, mechanisms, and organic aerosol, *Atmos. Chem. Phys.*, 17, 2103–2162, <https://doi.org/10.5194/acp-17-2103-2017>, 2017.
- Nguyen, T. B., Laskin, A., Laskin, J., and Nizkorodov, S. A.: Brown carbon formation from ketoaldehydes of biogenic monoterpenes, *Faraday Discuss.*, 165, 473–494, <https://doi.org/10.1039/c3fd00036b>, 2013.
- NIST: NIST Chemistry WebBook Standard Reference Database 69, <https://doi.org/10.18434/T4D303>, 2020.



- Noziere, B., Dziedzic, P., and Cordova, A.: Products and kinetics of the liquid-phase reaction of glyoxal catalyzed by ammonium ions (NH<sub>4</sub><sup>+</sup>), *J. Phys. Chem. A*, 113, 231–237, <https://doi.org/10.1021/jp8078293>, 2009.
- Ortiz-Montalvo, D. L., Hakkinen, S. A., Schwier, A. N., Lim, Y. B., McNeill, V. F., and Turpin, B. J.: Ammonium addition (and aerosol pH) has a dramatic impact on the volatility and yield of glyoxal secondary organic aerosol, *Environ. Sci. Technol.*, 48, 255–262, <https://doi.org/10.1021/es4035667>, 2014.
- Paciga, A. L., Riipinen, I., and Pandis, S. N.: Effect of ammonia on the volatility of organic diacids, *Environ. Sci. Technol.*, 48, 13769–13775, <https://doi.org/10.1021/es5037805>, 2014.
- Pathak, R. K., Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Ozonolysis of alpha-pinene at atmospherically relevant concentrations: Temperature dependence of aerosol mass fractions (yields), *J. Geophys. Res.-Atmos.*, 112, D03201, <https://doi.org/10.1029/2006jd007436>, 2007.
- Peng, C., Yang, F., Tian, M., Shi, G., Li, L., Huang, R. J., Yao, X., Luo, B., Zhai, C., and Chen, Y.: Brown carbon aerosol in two megacities in the Sichuan Basin of southwestern China: Light absorption properties and implications, *Sci. Total Environ.*, 719, 137483, <https://doi.org/10.1016/j.scitotenv.2020.137483>, 2020.
- Prinn, R. G., Weiss, R. F., Miller, B. R., Huang, J., Aleya, F. N., Cunnold, D. M., Fraser, P. J., Hartley, D. E., and Simmonds, P. G.: Atmospheric trends and lifetime of CH<sub>3</sub>CCI<sub>3</sub> and global OH concentrations, *Science*, 269, 187–192, <https://doi.org/10.1126/science.269.5221.187>, 1995.
- Qi, X., Zhu, S., Zhu, C., Hu, J., Lou, S., Xu, L., Dong, J., and Cheng, P.: Smog chamber study of the effects of NO<sub>x</sub> and NH<sub>3</sub> on the formation of secondary organic aerosols and optical properties from photo-oxidation of toluene, *Sci. Total Environ.*, 727, 138632, <https://doi.org/10.1016/j.scitotenv.2020.138632>, 2020.
- Sarrafzadeh, M., Wildt, J., Pullinen, I., Springer, M., Kleist, E., Tillmann, R., Schmitt, S. H., Wu, C., Mentel, T. F., Zhao, D., Hastie, D. R., and Kiendler-Scharr, A.: Impact of NO<sub>x</sub> and OH on secondary organic aerosol formation from β-pinene photooxidation, *Atmos. Chem. Phys.*, 16, 11237–11248, <https://doi.org/10.5194/acp-16-11237-2016>, 2016.
- Sun, Y.-L., Zhang, Q., Schwab, J. J., Demerjian, K. L., Chen, W.-N., Bae, M.-S., Hung, H.-M., Hogrefe, O., Frank, B., Rattigan, O. V., and Lin, Y.-C.: Characterization of the sources and processes of organic and inorganic aerosols in New York city with a high-resolution time-of-flight aerosol mass spectrometer, *Atmos. Chem. Phys.*, 11, 1581–1602, <https://doi.org/10.5194/acp-11-1581-2011>, 2011.
- Surratt, J. D., Murphy, S. M., Kroll, J. H., Ng, N. L., Hildebrandt, L., Sorooshian, A., Szmigielski, R., Vermeylen, R., Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Chemical composition of secondary organic aerosol formed from the photooxidation of isoprene, *J. Phys. Chem. A*, 110, 9665–9690, <https://doi.org/10.1021/jp061734m>, 2006.
- Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L. T., Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from anthropogenic air pollution: Rapid and higher than expected, *Geophys. Res. Lett.*, 33, L17811, <https://doi.org/10.1029/2006gl026899>, 2006.
- Vu, T. V., Shi, Z., Cheng, J., Zhang, Q., He, K., Wang, S., and Harrison, R. M.: Assessing the impact of clean air action on air quality trends in Beijing using a machine learning technique, *Atmos. Chem. Phys.*, 19, 11303–11314, <https://doi.org/10.5194/acp-19-11303-2019>, 2019.
- Wang, G., Zhang, F., Peng, J., Duan, L., Ji, Y., Marrero-Ortiz, W., Wang, J., Li, J., Wu, C., Cao, C., Wang, Y., Zheng, J., Seccrest, J., Li, Y., Wang, Y., Li, H., Li, N., and Zhang, R.: Particle acidity and sulfate production during severe haze events in China cannot be reliably inferred by assuming a mixture of inorganic salts, *Atmos. Chem. Phys.*, 18, 10123–10132, <https://doi.org/10.5194/acp-18-10123-2018>, 2018a.
- Wang, G. H., Zhang, R. Y., Gomez, M. E., Yang, L. X., Zamora, M. L., Hu, M., Lin, Y., Peng, J. F., Guo, S., Meng, J. J., Li, J. J., Cheng, C. L., Hu, T. F., Ren, Y. Q., Wang, Y. S., Gao, J., Cao, J. J., An, Z. S., Zhou, W. J., Li, G. H., Wang, J. Y., Tian, P. F., Marrero-Ortiz, W., Seccrest, J., Du, Z. F., Zheng, J., Shang, D. J., Zeng, L. M., Shao, M., Wang, W. G., Huang, Y., Wang, Y., Zhu, Y. J., Li, Y. X., Hu, J. X., Pan, B., Cai, L., Cheng, Y. T., Ji, Y. M., Zhang, F., Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation from London Fog to Chinese haze, *P. Natl. Acad. Sci. USA*, 113, 13630–13635, <https://doi.org/10.1073/pnas.1616540113>, 2016.
- Wang, R. Y., Ye, X. N., Liu, Y. X., Li, H. W., Yang, X., Chen, J. M., Gao, W., and Yin, Z.: Characteristics of atmospheric ammonia and its relationship with vehicle emissions in a megacity in China, *Atmos. Environ.*, 182, 97–104, <https://doi.org/10.1016/j.atmosenv.2018.03.047>, 2018b.
- Wang, S. W., Zhang, Q., Martin, R. V., Philip, S., Liu, F., Li, M., Jiang, X. T., and He, K. B.: Satellite measurements oversee China's sulfur dioxide emission reductions from coal-fired power plants, *Environ. Res. Lett.*, 10, 114015, <https://doi.org/10.1088/1748-9326/10/11/114015>, 2015.
- Wang, X., Gao, S., Yang, X., Chen, H., Chen, J., Zhuang, G., Surratt, J. D., Chan, M. N., and Seinfeld, J. H.: Evidence for high molecular weight nitrogen-containing organic salts in urban aerosols, *Environ. Sci. Technol.*, 44, 4441–4446, <https://doi.org/10.1021/es1001117>, 2010.
- Wang, Y., Chen, Y., Wu, Z., Shang, D., Bian, Y., Du, Z., Schmitt, S. H., Su, R., Gkatzelis, G. I., Schlag, P., Hohaus, T., Voliotis, A., Lu, K., Zeng, L., Zhao, C., Alfarra, M. R., McFiggans, G., Wiedensohler, A., Kiendler-Scharr, A., Zhang, Y., and Hu, M.: Mutual promotion between aerosol particle liquid water and particulate nitrate enhancement leads to severe nitrate-dominated particulate matter pollution and low visibility, *Atmos. Chem. Phys.*, 20, 2161–2175, <https://doi.org/10.5194/acp-20-2161-2020>, 2020.
- Wu, C., Zhang, S., Wang, G., Lv, S., Li, D., Liu, L., Li, J., Liu, S., Du, W., Meng, J., Qiao, L., Zhou, M., Huang, C., and Wang, H.: Efficient heterogeneous formation of ammonium nitrate on the saline mineral particle surface in the atmosphere of east asia during dust storm periods, *Environ. Sci. Technol.*, 54, 15622–15630, <https://doi.org/10.1021/acs.est.0c04544>, 2020.
- Xia, Y., Zhao, Y., and Nielsen, C. P.: Benefits of China's efforts in gaseous pollutant control indicated by the bottom-up emissions and satellite observations 2000–2014, *Atmos. Environ.*, 136, 43–53, <https://doi.org/10.1016/j.atmosenv.2016.04.013>, 2016.
- Xie, M., Chen, X., Hays, M. D., Lewandowski, M., Offenberg, J., Kleindienst, T. E., and Holder, A. L.: Light absorption of secondary organic aerosol: Composition and contribution of nitroaromatic compounds, *Environ. Sci. Technol.*, 51, 11607–11616, <https://doi.org/10.1021/acs.est.7b03263>, 2017.

- Xu, J., Huang, M. Q., Cai, S. Y., Liao, Y. M., Hu, C. J., Zhao, W. X., Gu, X. J., and Zhang, W. J.: Chemical composition and reaction mechanisms for aged p-xylene secondary organic aerosol in the presence of ammonia, *J. Chin. Chem. Soc.*, 65, 578–590, <https://doi.org/10.1002/jccs.201700249>, 2018.
- Xu, L., Kollman, M. S., Song, C., Shilling, J. E., and Ng, N. L.: Effects of NO<sub>x</sub> on the volatility of secondary organic aerosol from isoprene photooxidation, *Environ. Sci. Technol.*, 48, 2253–2262, <https://doi.org/10.1021/es404842g>, 2014.
- Xu, L., Moller, K. H., Crouse, J. D., Kjaergaard, H. G., and Wennberg, P. O.: New insights into the radical chemistry and product distribution in the OH-initiated oxidation of benzene, *Environ. Sci. Technol.*, 54, 13467–13477, <https://doi.org/10.1021/acs.est.0c04780>, 2020.
- Yang, W. Y., Li, J., Wang, M., Sun, Y. L., and Wang, Z. F.: A case study of investigating secondary organic aerosol formation pathways in Beijing using an observation-based SOA box model, *Aerosol Air Qual. Res.*, 18, 1606–1616, <https://doi.org/10.4209/aaqr.2017.10.0415>, 2018.
- Yang, Y., Vance, M., Tou, F. Y., Tiwari, A., Liu, M., and Hochella, M. F.: Nanoparticles in road dust from impervious urban surfaces: Distribution, identification, and environmental implications, *Environ. Sci.-Nano*, 3, 534–544, <https://doi.org/10.1039/c6en00056h>, 2016.
- Yang, Z., Tsona, N. T., Li, J., Wang, S., Xu, L., You, B., and Du, L.: Effects of NO<sub>x</sub> and SO<sub>2</sub> on the secondary organic aerosol formation from the photooxidation of 1,3,5-trimethylbenzene: A new source of organosulfates, *Environ. Pollut.*, 264, 114742, <https://doi.org/10.1016/j.envpol.2020.114742>, 2020.
- Zhang, L., Wang, Y., Feng, C., Liang, S., Liu, Y., Du, H., and Jia, N.: Understanding the industrial NO<sub>x</sub> and SO<sub>2</sub> pollutant emissions in China from sector linkage perspective, *Sci. Total Environ.*, 770, 145242, <https://doi.org/10.1016/j.scitotenv.2021.145242>, 2021a.
- Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.: Understanding atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review, *Anal. Bioanal. Chem.*, 401, 3045–3067, <https://doi.org/10.1007/s00216-011-5355-y>, 2011.
- Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M., and Wang, Y.: Formation of urban fine particulate matter, *Chem. Rev.*, 115, 3803–3855, <https://doi.org/10.1021/acs.chemrev.5b00067>, 2015a.
- Zhang, S., Li, D., Ge, S., Liu, S., Wu, C., Wang, Y., Chen, Y., Lv, S., Wang, F., Meng, J., and Wang, G.: Rapid sulfate formation from synergetic oxidation of SO<sub>2</sub> by O<sub>3</sub> and NO<sub>2</sub> under ammonia-rich conditions: Implications for the explosive growth of atmospheric PM<sub>2.5</sub> during haze events in China, *Sci. Total Environ.*, 772, 144897, <https://doi.org/10.1016/j.scitotenv.2020.144897>, 2021b.
- Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of secondary organic aerosol, *P. Natl. Acad. Sci. USA*, 111, 5802–5807, <https://doi.org/10.1073/pnas.1404727111>, 2014.
- Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C., and Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, *Atmos. Chem. Phys.*, 15, 4197–4214, <https://doi.org/10.5194/acp-15-4197-2015>, 2015b.
- Zhao, D., Schmitt, S. H., Wang, M., Acir, I.-H., Tillmann, R., Tan, Z., Novelli, A., Fuchs, H., Pullinen, I., Wegener, R., Rohrer, F., Wildt, J., Kiendler-Scharr, A., Wahner, A., and Mentel, T. F.: Effects of NO<sub>x</sub> and SO<sub>2</sub> on the secondary organic aerosol formation from photooxidation of  $\alpha$ -pinene and limonene, *Atmos. Chem. Phys.*, 18, 1611–1628, <https://doi.org/10.5194/acp-18-1611-2018>, 2018.
- Zou, Y., Deng, X. J., Zhu, D., Gong, D. C., Wang, H., Li, F., Tan, H. B., Deng, T., Mai, B. R., Liu, X. T., and Wang, B. G.: Characteristics of 1 year of observational data of VOCs, NO<sub>x</sub> and O<sub>3</sub> at a suburban site in Guangzhou, China, *Atmos. Chem. Phys.*, 15, 6625–6636, <https://doi.org/10.5194/acp-15-6625-2015>, 2015.