



Supplement of

Contrasting effects of secondary organic aerosol formations on organic aerosol hygroscopicity

Ye Kuang et al.

Correspondence to: Shan Huang (shanhuang_eci@jnu.edu.cn) and Min Shao (mshao@pku.edu.cn)

The copyright of individual parts of the supplement might differ from the article licence.

1. Site map



Figure S1. Locations of Heshan (HS) site, and other mega cities (FS:Foshan, GZ:Guangzhou, SZ:Shenzhen) in the Pearl river delta region (map background obtained through python from <u>http://server.arcgisonline.com/ArcGIS/rest/services/ESRI Imagery World 2D/MapServer/</u>, Copyright:© 2013 ESRI, i-cubed, GeoEye).

35
36
37
38
39
40
41

47 2. Details on quantifying uncertainties of using $\kappa_{f(RH),PM_{10}}$ represent κ_{chem,PM_1}

In Eq.3, other than aerosol composition measurements, what's the accuracy of 48 using $\kappa_{f(RH),PM_{10}}$ represent κ_{chem,PM_1} need to be carefully quantified. As discussed 49 in Kuang et al. (2020), the difference between κ_{chem,PM_1} and $\kappa_{f(RH),PM_{10}}$ are both 50 influenced by aerosol PNSD and size-resolved k distribution. To cover as many cases 51 as possible, we have performed a simulation to investigate the relative differences 52 between κ_{chem,PM_1} and $\kappa_{f(RH),PM_{10}}$ under different conditions of aerosol chemical 53 compositions through varying size distributions of ammonium nitrate, ammonium 54 sulfate, organic aerosol, dust and BC, and their mass fractions randomly, and the mass 55 size distributions follow lognormal distributions, that is: 56

57
$$dM/dlogDp = \frac{Mtot*fx}{\sqrt{2\pi}\log(\sigma_{g,x})} \exp\left[-\frac{(\log(D_p) - \log(D_{g,x}))^2}{2\log^2\sigma_{g,x}}\right], \qquad \text{Eq. 1}$$

where x corresponding different aerosol compositions, and fx corresponding to its mass fractions. The parameters used in simulations are listed in Table S1, the number of randomly produced datasets for simulating κ_{chem,PM_1} and $\kappa_{f(RH),PM_{10}}$ is 10000.

- 61
- 62
- 63
- 64
- 65
- 66
- 67
- 68

69 Table S1. used parameters for simulating differences between κ_{chem,PM_1} and $\kappa_{f(RH),PM_{10}}$, m and σ

	fx range (%)	D _g (nm)		σ_{g}		ρ		κ	
		m	σ	m	σ	m	σ	m	σ
Ammonium Nitrate	0-50	450	100	2.2	0.25	1.72		0.56	
Ammonium sulfate	0-50	550	100	2.2	0.25	1.78		0.56	
Organic Aerosol	10-60	450	150	2.2	0.25	1.2	0.05	0.15	0.05
Dust	5-50	2500	500	2.2	0.25	2		0.05	0.015
BC	0-12	250	100	2.2	0.25	1.7		0	

70 corresponds to average and standard deviation respectively.

72 3. Analysis on possible pathways of nitrate increase

Depending on the RH or the RH history that aerosol particles have been experienced, 73 the aerosol particles might be solid (if crystallized) or aqueous phase (dehydration 74 branch from the morning to the afternoon, metastable) after about 16:00. However, 75 when aerosol particles will be deliquescent is also a puzzle if aerosol particle 76 crystallized in the afternoon due to the complex dependence of deliquescence RH on 77 aerosol mixtures of ammonium sulfate, ammonium nitrate and organic aerosols (Kuang 78 et al., 2016), and the peak RH in the morning was usually lower than 80%. To illustrate 79 the possible mechanisms behind the nitrate increase, we assume the following two states 80 of ammonium nitrate: (1) solid state; (2) aqueous phase. 81

For the solid phase case, the NH₃(g) + HNO₃ (g) \leftrightarrow NH₄NO₃ (s) equilibrium theory is applied (reference). If we have the initial concentration of gaseous NH₃, then we can guess the initial HNO₃ (g) based on the K_p(T) and estimate the nitrate concentration increase due to the decrease K_p(T), the K_p(T) changed from 84 ppb² to 24 ppb² from 16:30 to 23:30 for the average case with air temperature decrease from 27.7 to 22.7 °



Figure S2. Simulated variations of nitrate under different NH3 (g) conditions by assuming aerosol particles are solid and using the Kp(T) theory.

⁸⁷ C. Figure S2 shows the variations of nitrate mass concentrations under different ⁸⁸ assumed NH₃ (g) concentration conditions by considering variations of $K_p(T)$. The ⁸⁹ results demonstrate that the temperature induced $K_p(T)$ change is enough for explaining ⁹⁰ observed nitrate increase, and thus a possible mechanism.

For the aqueous phase case, NH_4NO_3 will be found in the aqueous state, and the corresponding dissociation reaction is then $NH_3(g) + HNO_3(g) \leftrightarrow NH_4^+ + NO_3^-$. For

this case, both temperature decrease and RH increase play roles in the NH₃ and HNO₃ 93 partitioning. We use a thermodynamic model ISORROPIA (forward, metastable) to 94 simulate the possible magnitude of nitrate increase driven by increased RH and reduced 95 temperature for the average case shown in Fig.3 of the manuscript. We also assume 96 different scenarios of $NH_3(g)$ concentrations and for each $NH_3(g)$ concentration scenario, 97 we estimate the $HNO_3(g)$ concentration with which we can reproduce the observed 98 nitrate concentration at the initial time (16:30), to get a total HNO₃ concentration in gas 99 and aerosol phase, then we can estimate the possible increase of nitrate in aerosol phase 100 due to repartitioning and the results are shown in Figure S3. The results show that the 101 more HNO₃ dissolved in aqueous phase can also explain the observed nitrate mass 102 concentration increase. 103

The simulation results shown above demonstrate that repartitioning of HNO₃ (total 104 HNO₃ in aerosol and gas phase remain unchanged) can explain the observed nitrate 105 concentration increase under both assumptions that aerosols are completely aqueous or 106 sold. The actual atmospheric case is more likely that both solid (freshly emitted and 107 crystallized) and aqueous particles (dehydration branch, not crystallized) exist, however, 108 does not affect conclusion that the repartitioning of HNO₃ in gas and aerosol phase can 109 explain the observed nitrate increase. In addition, NO concentration is quite low after 110 16:00 which might be favorable of N_2O_5 formation and the N_2O_5 hydrolysis as a possible 111 nitrate formation pathway cannot be excluded. 112



Figure S3. Simulated variations of nitrate under different NH3 (g) conditions by assuming aerosol particles are aqueous using the ISORROPA thermodynamic model.



Figure S4. (a) Size-resolved κ distributions which are derived from measured size-segregated chemical compositions during HaChi campaign, colors represent corresponding values of average σ_{sp} at 550 nm (Mm^{-1}), black solid line is the average size-resolved κ distribution and error bars are standard deviations. (reprint from (Kuang et al., 2018)); (b) Comparison of simulated $\kappa_{f(RH)}$ of PM10 and κ_{chem} of PM1



Figure S5. Comparison between PM₁ volume concentration derived from SMPS (x-axis) and SP-AMS

- 144 (y-axis) measurements. Dashed red lines represent 20% relative difference lines.





Figure S6. (a) Mass spectral profile in family groups and (b) time series of PMF OA components.Also exhibited are concentration variations of tracer compounds on the right axes.

- 1/7



Figure S7. Comparison between PM_{2.5} mass concentration provided by local authority and total mass
 concentration measured by SP-AMS.



Figure S8. Comparisons of aerosol compositions between SP-AMS (PM1) and external online measurements, including gas aerosol collection system (GAC) with ion chromatography (with PM2.5 impactor), ECOC Analyzer (Sunset, xxx, Inc) and AE33 aethalometer. Note that organic carbon (OC) was calculated by organics from SP-AMS divided ratio of OM and OC provided by AMS HR data, and then compared with OC derived from ECOC analyzer.



Figure S9. Comparisons of ammonium, nitrate, sulfate and chloride between SP-AMS (PM1) andoffline measurements with PM2.5 impactor.



Figure S10. Comparison between κ_{OA} derived with and without random errors



Figure S11. Correlations between κ_{OA} derived with random errors and mass fractions of OA factors in total OA mass

Kuang, Y., Zhao, C. S., Ma, N., Liu, H. J., Bian, Y. X., Tao, J. C., and Hu, M.: Deliquescent phenomena of ambient aerosols on
the North China Plain, Geophysical Research Letters, n/a-n/a, 10.1002/2016GL070273, 2016.

Kuang, Y., Zhao, C. S., Zhao, G., Tao, J. C., Xu, W., Ma, N., and Bian, Y. X.: A novel method for calculating ambient aerosol
liquid water content based on measurements of a humidified nephelometer system, Atmospheric Measurement

241 Techniques, 11, 2967-2982, 10.5194/amt-11-2967-2018, 2018.

Kuang, Y., He, Y., Xu, W., Zhao, P., Cheng, Y., Zhao, G., Tao, J., Ma, N., Su, H., Zhang, Y., Sun, J., Cheng, P., Yang, W., Zhang,
S., Wu, C., Sun, Y., and Zhao, C.: Distinct diurnal variation in organic aerosol hygroscopicity and its relationship with
oxygenated organic aerosol, Atmos. Chem. Phys., 20, 865-880, 10.5194/acp-20-865-2020, 2020.