Variations in $N_{\text{ccn}}$ and $N_{\text{ccn}}$ over marginal seas in China related to marine traffic emissions, new particle formation and aerosol aging

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Abstract. In this study, a cruise campaign was conducted over marginal seas in China to measure the concentrations of condensation nuclei ($N_{\text{cn}}$), cloud condensation nuclei ($N_{\text{ccn}}$) and other pollutants from day of year (DOY) 110 to DOY 135 of 2018. The ship self-emission signals were exhaustively excluded, and the mean values of $N_{\text{ccn}}$ during the cruise campaign were found to slightly increase from $3.2 \pm 1.1 \times 10^3$ cm$^{-3}$ (mean ± standard deviation) at supersaturation (SS) of 0.2% to $3.9 \pm 1.4 \times 10^3$ cm$^{-3}$ at SS of 1.0%, and the mean value for $N_{\text{cn}}$ was $8.1 \pm 4.4 \times 10^3$ cm$^{-3}$. Data analysis showed that marine traffic emissions apparently largely contributed to the increase in $N_{\text{cn}}$ in the daytime, especially in the marine atmospheres over heavily traveled sea zones; however, the fresh sources made no clear contribution to the increase in $N_{\text{ccn}}$. This finding was supported by the quantitative relations between $N_{\text{cn}}$ and $N_{\text{ccn}}$ at SS = 0.2%–1.0% against the mixing ratios of SO$_2$ in the ship self-emission plumes – i.e., a 1 ppb increase in SO$_2$ corresponded to a $1.4 \times 10^3$ cm$^{-3}$ increase in $N_{\text{cn}}$ but only a $30–170$ cm$^{-3}$ increase in $N_{\text{ccn}}$, possibly because of abundant organics in the aerosols. Smooth growth can be observed in the marine-traffic-derived particles, reflecting aerosol aging. The estimated hygroscopicity parameter ($\kappa$) values were generally as high as 0.46–0.55 under the dominant onshore winds, suggesting that inorganic ammonium aerosols likely acted as the major contributor to $N_{\text{cn}}$ largely through aerosol aging processes of decomposing organics. Moreover, the influences of the new transported particles from the continent on the $N_{\text{cn}}$ and $N_{\text{ccn}}$ in the marine atmosphere were investigated.

1 Introduction

Oceans occupy approximately 2/3 of the Earth’s surface, and water evaporation from oceans is a major source of moisture in the atmosphere. Aerosol–cloud interactions in marine atmospheres, ranging from tropical to polar regions, have attracted great attention in the past few decades due to their impact on climate change (Huebert et al., 2003; Yu and Luo, 2009; Quinn and Bates, 2011; Wang et al., 2014; Brooks and Thornton, 2018; Rosenfeld et al., 2019). However, large uncertainties still exist in various marine atmospheres, e.g., the sources of aerosols, the concentrations of bulk cloud condensation nuclei (CCN) and aerosol CCN activation under various supersaturations. (Clarke et al., 2006; Decesari et al., 2011; Quinn and Bates, 2011; Saliba et al., 2019; Rosenfeld et al., 2019). These uncertainties are mainly determined by limited observations in marine atmospheres, although a few additional observations of the number concentrations of aerosols ($N_{\text{cn}}$) and CCN ($N_{\text{ccn}}$) were recently reported in different marine atmospheres, e.g., over the Mediterranean (Bougiatioti et al., 2009), Sea of Japan (Yamashita et al., 2011), Bay of Bengal (Ramana and Devi, 2016), coast of California (Ruehl et al., 2009) and northwest Pacific Ocean (Wang et al., 2019).

In addition to sea-spray aerosols and secondarily formed aerosols from sea-derived gaseous precursors (O’Dowd et al., 1997; Clarke et al., 2006; Quinn and Bates, 2011; Blot et al., 2013; Fossum et al., 2018), marine traffic emits large amounts of aerosols and reactive gases (Chen et al., 2017).
These pollutants may directly or indirectly contribute to CCN to some extent (Langley et al., 2010). In addition, the long-range transport of continental aerosols has been widely reported to act as an important source of CCN in marine atmospheres (Charlson et al., 1987; Huebert et al., 2003; Fu et al., 2017; Royalty et al., 2017; Sato and Suzuki, 2019; Wang et al., 2019). The continent-derived aerosol particles observed in marine atmospheres usually mix with different sources, such as biomass burning, dust and anthropogenic emissions (Feng et al., 2017; Lin et al., 2015; Guo et al., 2014, 2016). An appreciable fraction of organics reportedly exists in marine aerosols and continental aerosols upwind of oceans (O’Dowd et al., 2004; Feng et al., 2012, 2016; Quinn et al., 2015; Song et al., 2018; Ding et al., 2019). However, ammonium sulfate aerosols have been frequently reported to dominantly contribute to CCN-related aerosols in many marine atmospheres and lead to hygroscopicity parameters (κ) larger than 0.5 (Mochida et al., 2010; Cai et al., 2017; Fu et al., 2017; Royalty et al., 2017; Phillips et al., 2018). A question is naturally raised, i.e., where do particulate organics go in the marine aerosols enriched in ammonium sulfate? Anthropogenic emissions in China such as SO2 and NOx have generally increased since the 1980s and recently started to decrease, i.e., SO2 started to decrease in 2006 (Li et al., 2017), whereas NOx started to decrease in 2011 (Li et al., 2017; Liu et al., 2016). Together with the influence of the Asian monsoon, the marginal seas near China are inevitably affected by the outflow of continental aerosols (Guo et al., 2016; Feng et al., 2017). Observations of Ncn and Nccn in marine atmospheres over marginal seas near China help to address the data scarcity, understand the sources and dynamic changes in these parameters, and study their potential climate impacts.

In this study, cruise campaigns were conducted to measure the Nccn, Ncn, particle number size distributions, gaseous pollutants and aerosol composition of water-soluble ionic species over the marginal seas from 20 April 2018 (day of year – DOY 110) to 15 May 2018 (DOY 135), traveling from the East China Sea to the South China Sea and returning to the Yellow Sea. Spatiotemporal variations in the Ncn, Nccn and CCN activities of the aerosol particles were studied. The Kappa values of the aerosol particles from DOY 110 to DOY 118 over the marine environments were calculated and analyzed. Finally, we tried to establish relationship of Ncn and Nccn with the mixing ratios of SO2 in ship self-plumes and ambient marine air. The regression equations are valuable for the estimation of Ncn and Nccn from SO2 when the direct observations of Ncn and Nccn are not available.

2 Experimental design

2.1 Instruments and data sources

A cruise campaign was conducted across marginal seas in China from DOY 110 to DOY 135 of 2018 (Fig. 1a, b). A suite of instruments including a fast mobility particle sizer (FMPS; TSI Model 3091), CCN counter (CCNC; DMT Model 100), condensation particle counter (CPC; TSI Model 3775), gas analyzers, ambient ion monitor-ion chromatography (AIM-IC), etc., were onboard the commercial cargo ship Anqiang 87 for measurements. The FMPS was used to measure the particle number size distributions with mobility diameters from 5.6 to 560 nm in 32 channels at 1 s temporal resolution with an inlet flow of 10 L min−1. The CPC was used to report the Ncn ranging from 4 nm (50 % efficiency) to 3000 nm (Ncn) in 2 s time resolution with an inlet flow of 1.5 L min−1. The Ncn was then used to calibrate the particle number size distributions simultaneously measured by the FMPS on the basis of the procedure proposed by Zimmerman et al. (2015). Due to the severe oceanic conditions and humid weather conditions, the FMPS and CPC were out of service after DOY 118 and DOY 122, respectively. Prior to the campaign, the CCNC was calibrated with ammonium sulfate particles based on the standard procedure detailed by Rose et al. (2008). The calibration curve is shown in Fig. S1 in the Supplement. The total flow rate of CCNC was 0.45 L min−1, with a ratio of sample to sheath at 1/10, and five supersaturations (SS) conditions were selected, including 0.2 %, 0.4 %, 0.6 %, 0.8 % and 1.0 %. More detailed information about the measurement of Nccn can be found in Wang et al. (2019).

During the experiment, ambient particles were first sampled through a conductive tube (TSI, US) and a diffusion dryer filled with silica gel (TSI, US) and then split for analysis by means of different instruments with a splitter. All instruments were placed in an air-conditioned container on the deck of the ship, with an inlet height of approximately 6 m above sea level. Regarding the gas analyzers, the ambient O3 (Model 49i, Thermo Environmental Instruments Inc., USA C-series), SO2 (Model 43i, Thermo Environmental Instruments Inc., USA C-series) and NOx values (Model 42i, Thermo Environmental Instruments Inc., USA C-series) were measured in mixing ratios with a temporal resolution of 1 min. The CCNC and gas analyzers were operated properly throughout the entire campaign. The same was true for the AIM-IC, which was used to measure the water-soluble ionic species in the ambient particles sized smaller than 2.5 μm.

In this study, the Hybrid Single-Particle Lagrangian Integrated Trajectory (HYSPLIT) model from the NOAA Air Resources Laboratory was used to track the particle sources. The input of HYSPLIT, such as wind speed and wind direction, was obtained from the National Center for Environmental Prediction (NCEP) Global Data Assimilation System (GDAS) with a spatial resolution of 0.5°.
The hygroscopicity parameter ($\kappa$) was calculated according to the method proposed by Petters and Kreidenweis (2007).

$$\kappa = \frac{4A^3}{27D_d^3 \ln S_C}, \quad A = \frac{4\sigma_{s/a}M_w}{RT\rho_w},$$

where $D_d$ is the dry diameter, $S_C$ is the supersaturation, $M_w$ is the molecular weight of water, $\sigma_{s/a}$ represents the surface tension over the interface of the solution and air with the value of 0.072 J m$^{-2}$ applied in this study, $R$ is the universal gas constant, $T$ is the ambient temperature and $\rho_w$ is the water density. $D_d$ was not measured directly and was assumed to be equal to the critical diameter for CCN activation ($D_{crit}$). $D_{crit}$ was defined as the particle diameter which is considered as the lower limit of the integral on particle number, together with the upper limit setting as the largest particle diameter, yielding the total integrated particle number concentration equal to the CCN concentration (Hung et al., 2014; Cheung et al., 2020). The FMPS had a low size resolution, particularly at sizes greater than 90 nm, which did not allow accurate calculation of the $Kappa$ values at $SS = 0.2\%$. At $SS = 0.6\%$ and $0.8\%$, the $Kappa$ value was not calculated considering the complication in the explanation of the value, possibly reflecting the combined effects of particle size, mixing state and chemical composition.

### 2.2 Separating ambient signals of $N_{cn}$ and $N_{ccn}$ from ship self-emissions

The data measured during the cruise campaign were frequently subject to interference from self-emission signals from the ship. The $N_{cn}$ and $N_{ccn}$ over the marginal seas were first distinguished based on the source of the ambient environment or the ship self-emissions. The data measured at 18:00–24:00 LT on DOY 115 are used to illustrate the separation in Fig. 2, and the size distribution of the particle number concentration during DOYs 110–118 is shown in Figs. S2–S10 in the Supplement. At 18:00–21:11 LT (local time), a low $N_{cn}$ of $5.8 \pm 0.4 \times 10^3$ cm$^{-3}$ was observed. The accumulation mode dominated the particle number concentration with a median mobility mode diameter of 105 ± 4 nm (Fig. 2a). Afterwards, the $N_{cn}$ rapidly increased by over 1 order of magnitude (Fig. 2b). The dominant particle number concentration mode changed from accumulation mode to Aitken mode, with the median mobility diameter of the Aitken mode stabilized at 47 ± 4 nm for approximately 90 % of the time. The rapid increase in $N_{cn}$ and the change in the mode size indicated the signal of the emissions of the ship itself. The ship self-emission interference after 21:11 was supported by additional evidence, e.g., a large decrease in the activation ratio (AR), defined as the quotient of $N_{ccn}$ and $N_{cn}$, from $> 0.5$ to $< 0.2$ at $SS = 0.4\%$ (Fig. 2c) due to a large increase in $N_{cn}$ but a much smaller magnitude enhancement of $N_{ccn}$ (Fig. 2b) and a rapid increase in NO$_x$ from < 10 ppb to 192 ± 99 ppb, NO/NO$_2$ from < 0.1 to 0.7 ± 0.3, and SO$_2$ from < 2 ppb to 6.2 ± 2.4 ppb. Large changes were expected because the ship smoke stack was only approximately 10 m away from these detectors. Thus, based upon the features described above, certain criteria were designed in this study to identify ship self-emission signals to separate them from ambient signals, i.e., a net increase in $N_{cn}$ beyond $5 \times 10^4$ cm$^{-3}$ in 5 min, a median mobility mode diameter of approximately 50 nm, NO$_2 > 30$ ppb and NO/NO$_2 > 0.5$.

### 3 Results and discussion

#### 3.1 Spatiotemporal variations in ambient $N_{cn}$ during the cruise period

Figure 3 shows the time series of minutely averaged distributions of $N_{cn}$, $N_{ccn}$ and AR at SSs of 0.4 % and 1.0 % from
with an average value of 8\(\times 10^3\) cm\(^{-3}\) during the first half of the cruise period (Fig. 3a), it was
found that \(N_{\text{cn}}\) spanned a broad range of 0.2–4.5 \(\times 10^3\) cm\(^{-3}\) with an average value of 8.1 \(\pm 4.4\) \(\times 10^3\) cm\(^{-3}\). Specifically, \(N_{\text{cn}}\) was only 6.5 \(\pm 0.8\) \(\times 10^3\) cm\(^{-3}\) at 00:00–06:00 LT on DOY 110 when the ship anchored at the Yangtze River estuary near Shanghai (Fig. 1). The low \(N_{\text{cn}}\) values were comparable to the mean value of \(N_{\text{cn}}\) (5.4 \(\times 10^3\) cm\(^{-3}\)) in the marine-air cases during January–December 2010 in Shanghai reported by Leng et al. (2013). The \(N_{\text{cn}}\) greatly increased to 1.9 \(\pm 0.7\) \(\times 10^4\) cm\(^{-3}\) at 08:00–21:00 LT on DOY 110 when the ship cruised across the Yangtze River estuary. The measured particles in the number concentration were dominantly distributed in the Aitken mode on that day, while the median Aitken mode diameter shifted from 49.5 nm at 00:00–06:00 to 40±5 nm at 08:00–21:00 (Fig. S2). The Yangtze River estuary contains several world-class ports and is heavily traveled by marine traffic in the daytime (Chen et al., 2017). Since the onshore wind dominated on that day (not shown), the increase in \(N_{\text{cn}}\) and the decrease in the median Aitken mode diameter at 08:00–21:00 LT possibly reflected the increased contribution from marine traffic emissions. The marine traffic visibly decreased when the ship left the Yangtze River estuary toward the south. The \(N_{\text{cn}}\) value then significantly decreased, i.e., to 9.5 \(\pm 4.4\) \(\times 10^3\) cm\(^{-3}\) in the marine atmosphere over the sea zone in Zhejiang Province (from 07:00 LT on DOY 111 to 17:00 LT on DOY 114), with \(P < 0.01\). The \(N_{\text{cn}}\) further decreased to 5.8 \(\pm 1.7\) \(\times 10^3\) cm\(^{-3}\) in the marine atmosphere over the sea zone in Fujian Province (from 18:00 LT on DOY 114 to 14:00 LT on DOY 117). All these values were, however, 1–2 orders of magnitude greater than the background values in remote clear marine atmospheres, e.g., < 300 particle cm\(^{-3}\) without the influence of industrial activities in the western Pacific and the tropical Pacific (Ueda et al., 2016) and those reported by Quinn and Bates (2011) and Saliba et al. (2019), indicating overwhelming contributions from non-sea-spray aerosols including marine traffic emissions, long-range continental transport, newly formed particles in marine atmospheres, etc. As reported, the atmospheric particles over marginal seas in China can be further transported to the remote northwest Pacific Ocean (NWPO) in spring under westerly winds – e.g., the \(N_{\text{cn}}\) observed over the NWPO in March–April 2014 was as high as 2.8 \(\pm 1.0\) \(\times 10^3\) cm\(^{-3}\) and approximately half of that over marginal seas in China observed in March 2014 (Wang et al., 2019).

The mean value of \(N_{\text{cn}}\) (8.1 \(\pm 4.4\) \(\times 10^3\)) observed in this study was close to that of 7.6 \(\pm 4.0\) \(\times 10^3\) cm\(^{-3}\) (the number concentrations of particles larger than 10 nm) observed over the eastern part of the Yellow Sea in spring 2017 according by Park et al. (2018). The authors attributed the high number concentrations of particles within nucleation and Aitken modes to the long-range transport of air pollutants over eastern China under the influence of westerly winds. Consistently, larger values of \(N_{\text{cn}}\) were frequently observed in the continental atmospheres upwind of the Yellow Sea, e.g., mean values of 1.8 \(\pm 1.4\) \(\times 10^3\) cm\(^{-3}\) in May 2013 in Qingdao, a coastal city in proximity to the Yellow Sea (Li et al., 2015), 3.18 \(\times 10^4\) cm\(^{-3}\) in February–August 2014 in Beijing.
During which the mean $1.0 \%$. In contrast, after DOY 120, the concentration of NH$_4^+$ on DOY 120) is statistically significantly higher (at significance level of 0.01. There was no significant difference between the $N_{ccn}$ values of the two composite periods at SS values of 0.4 % and 0.6 %. However, the composited $N_{ccn}$ (i.e., only selection of the upper quartile) during the high-NH$_4^+$ period was significantly lower than the composited value during the low-NH$_4^+$ period for $P < 0.01$, e.g., $5.1 \pm 0.5 \times 10^3$ cm$^{-3}$ versus $5.3 \pm 0.7 \times 10^3$ cm$^{-3}$ at SS = 0.8 % and $5.2 \pm 0.5 \times 10^3$ cm$^{-3}$ versus $5.7 \pm 0.7 \times 10^3$ cm$^{-3}$ at SS = 1.0 %. During the low-NH$_4^+$ period, the marine atmospheres over the observational zones may sometimes receive strong continental inputs and/or marine traffic emissions, leading to the larger $N_{ccn}$. The enhanced formation of ammonium salt aerosols during the high-NH$_4^+$ period likely canceled out or even overwhelmed the effects of the continental inputs and/or marine traffic emissions on $N_{ccn}$ at SS = 0.2 %.

In addition, fresh marine traffic emissions likely made a negligible contribution to $N_{ccn}$ in the marine atmosphere because of the large amounts of aged aerosols from various sources therein. For example, the mean values of $N_{ccn}$ were $3.2 \times 10^3$ and $4.5 \times 10^3$ cm$^{-3}$ at SS = 0.4 % and 1.0 % at 08:30–11:30 on DOY 110, respectively. These values were almost the same as the $3.2 \times 10^3$ cm$^{-3}$ at SS = 0.4 % and $3.8 \times 10^3$ cm$^{-3}$ at SS = 1.0 % before 06:00 on that day. The mean values of $N_{ccn}$, however, greatly increased from $6.5 \pm 0.8 \times 10^3$ cm$^{-3}$ before 06:00 to $1.3 \pm 0.3 \times 10^4$ cm$^{-3}$ at 08:30–11:30 when the ship cruised across the Yangtze River estuary (Fig. 3b).

### 3.3 Spatiotemporal variations in CCN activation and Kappa values

The AR values at SSs of 0.4 % and 1.0 % are examined in this section, as shown in Fig. 3c. At SS = 0.4 %, the AR values largely varied from 0.06 to 0.92 with a median value of 0.51. Specifically, the AR values narrowly varied around $0.51 \pm 0.04$ at 00:00–06:00 LT on DOY 110. At 08:00–21:00 LT on that day, when the ship cruised across the Yangtze River estuary, the AR values substantially decreased to $0.26 \pm 0.06$ concurrently with an approximate $200 \%$ increase in the $N_{ccn}$ values, i.e., $N_{ccn}$ values of $6.5 \pm 0.8 \times 10^3$ cm$^{-3}$ at 00:00–06:00 LT and $2.0 \pm 0.7 \times 10^4$ cm$^{-3}$ at 08:00–21:00 LT on DOY 110 (Fig. 3a). The AR values then exhibited an oscillating increase from DOY 111 to DOY 113. A low AR value of $0.12 \pm 0.04$ was suddenly observed at 10:00–18:00 LT on
DOY 114 in the presence of strong new particle signals transported from the upwind continental atmosphere, as discussed later. The AR values, however, reached 0.34±0.04 at 06:00–08:00 LT and 0.39±0.08 at 19:00–24:00 LT on DOY 114, with the new particle signals largely decreased. Even excluding the AR values on DOY 114, a significant difference was still obtained between the AR values of 0.61±0.12 during the high-NH$_4^+$ period and those of 0.55±0.17 during the low-NH$_4^+$ period. The enhanced formation of ammonium salts seemingly increased the CCN activity to some extent. At SS = 1.0 %, the AR values showed large fluctuations with a median value of 0.57±0.17 (Fig. 3c), and the temporal trend was similar to that at SS = 0.4 %.

To minimize the impact from the particle sizes, the Kappa values were further investigated. As reported by Phillips et al. (2018), Kappa values at a high time resolution usually exhibit a broad distribution, reflecting the complexity due to various factors. To reveal the key factors in determining the Kappa values on a large spatiotemporal scale, the daily Kappa values of atmospheric aerosols were estimated on the basis of the daily mean $N_{ccn}$ and the size distributions of the particle number concentration from DOYs 110–118 (Fig. 3c). Please note that for DOY 110, considering the large differences in the particle number concentration between 00:00–06:00 and 08:00–21:00 (Fig. S2), the Kappa values were calculated separately for these two periods. At SS = 0.4 % (green dashed line in Fig. 3c), the estimated Kappa values were as high as 0.66 at 00:00–06:00 LT, while they decreased to 0.37 at 08:00–21:00 LT on DOY 110. The Kappa value varied narrowly from 0.46 to 0.55 on DOYs 111–113, 115 and 117, implying that inorganic aerosols such as completely and incompletely neutralized ammonium salts may make large contributions to the $N_{ccn}$. These values were generally consistent with the reported observations in most marine atmospheres. For example, Cai et al. (2017) reported a Kappa value of approximately 0.5 for particles with sizes of 40–200 nm at a marine site in Okinawa and that sulfate was the dominant component of aerosol particles on 1–9 November 2015, and a similar Kappa value in spring 2008 was reported by Mochida et al. (2010) over this site. Royalty et al. (2017) reported Kappa values for 48, 96 and 144 nm dry particles of 0.57±0.12, 0.51±0.09 and 0.52±0.08 in the subtropical North Pacific Ocean and sulfate-like particles contributing at most 77 %–88 % to the total aerosol number concentration. Over the Atlantic Ocean, Kappa values of approximately 0.54±0.03 were observed for 284 nm particles (Phillips et al., 2018).

The estimated Kappa values sometimes reached 0.66–0.67 (i.e., on DOY 116), which may be related to unidentified factors. For example, O’Dowd et al. (2014) proposed that some organics derived from sea-spray aerosols may also increase the $N_{ccn}$ to some extent by reducing the surface tension, leading to an increase in the Kappa values. A small fraction of sea-salt aerosols in submicron particles may also increase the Kappa values since its Kappa value was as high as 1.3 (O’Dowd et al., 1997, 2004). A Kappa value of 0.29 was obtained on DOY 118, which is close to the Kappa values widely observed for continental atmospheric aerosols (~0.3) (Andreae and Rosenfeld, 2008; Pöschl et al., 2009; Rose et al., 2010). The estimated Kappa value largely decreased to 0.15 on DOY 114 when new particle formation (NPF) occurred; see Sect. 3.5 for detailed discussion. Moreover, at an SS of 1.0 %, the estimated Kappa value was always smaller than 0.2. The Kappa values of organics were commonly assumed to be 0.1 (Rose et al., 2011; Cai et al., 2017; Singla et al., 2017). In general, the fraction of organics in the nanometer particles increases with decreasing particle size from ~100 to ~50 nm (Rose et al., 2010, 2011; Crippa et al., 2014; Cai et al., 2017). A combination of the two factors likely led the overall Kappa values estimated at SS = 1.0 % to

<table>
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<tr>
<th>Variable</th>
<th>Supersaturation (SS)</th>
<th>Range</th>
<th>Mean ± standard deviation</th>
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<tr>
<td>$N_{cn}$ ($\times 10^3$ cm$^{-3}$)</td>
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<tr>
<td>$N_{ccn}$ ($\times 10^3$ cm$^{-3}$)</td>
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<td></td>
<td>SS = 0.8 %</td>
<td>0.5–11.1</td>
<td>3.8 ± 1.2</td>
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<td></td>
<td>SS = 1.0 %</td>
<td>0.6–12</td>
<td>3.9 ± 1.4</td>
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<tr>
<td>AR</td>
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<tr>
<td></td>
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<td>0.25–9.7</td>
<td>1.7 ± 1.1</td>
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approximately 60% of the total particle number concentration (Fig. 5a), with median Aitken mode diameters of approximately 60% of the total particle number concentrations (Fig. 5b). Concomitantly, the AR values decreased to 0 compared to those in the early morning before 05:30 LT.

NDOY 116 and NDOY 118 were selected. For further investigation of the sources of different modes of particles, the days DOY 112, DOY 116 and DOY 118 were selected.

On DOY 112, the Aitken mode particles accounted for approximately 60% of the total particle number concentration (Fig. 5a), with median Aitken mode diameters of approximately 54 ± 8 nm. Similar to the observations over the Yangtze River estuary, the mean value of $N_{\text{cn}}$ increased by approximately 50% concurrently with a decrease in the median Aitken mode diameters by $\sim$ 9 nm at 05:30–11:40 LT compared to those in the early morning before 05:30 LT (Fig. 5b). Concomitantly, the AR values decreased to 0.31 ± 0.09 at SS of 0.4%, with similar AR decreases at SS of 1.0%. The lowest AR and Kappa values occurred at 06:00–07:00 LT at SSs of both 0.4% and 1.0%. All these results indicated that the increase in Aitken mode particles at 05:30–11:40 LT was likely derived from enhanced marine traffic contributions carried by the onshore wind from the south (Fig. S11). During other times on DOY 112, the onshore wind may also carry marine-traffic-derived particles to the observational sea zones. However, the marine-traffic-derived particles likely aged to some extent, e.g., the median Aitken mode diameters exhibited an oscillating increase from approximately 50 nm at 19:00 to approximately 70 nm at 24:00 LT with a particle growth rate of $\sim$ 4 nm h$^{-1}$. The AR values, however, narrowly varied around 0.47 ± 0.03 at SS = 0.4% and 0.52 ± 0.05 at SS = 1.0% during the particle growth period. The Kappa values at SS = 0.4% gradually decreased from 0.56 at 19:00 to 0.41 at 23:00 LT, reflecting more aged marine-traffic-derived particles growing into CCN sizes.

On DOY 116, the accumulation mode particles dominantly contributed to $N_{\text{cn}}$ rather than Aitken mode particles (Fig. 5d) under the marine air influence from the northeast (Fig. S13). The median accumulation mode diameters narrowly varied around $135 \pm 5$ nm at 01:00–13:00 LT and $102 \pm 5$ nm at 16:20–24:00 LT with a transition period in between (Fig. 5e). The AR and Kappa values, however, showed no statistically significant differences during the two periods at SSs of 0.4% and 1.0%, implying that the size change in the accumulation mode particles had a negligible influence on CCN activation. The hourly variations in the AR and Kappa values may be associated with other factors, e.g., chemical composition and mixing state (Gunthe et al., 2011; Rose et al., 2011).

On DOY 118, under the influence of mixtures from marine and coastal areas from the northeast (Fig. S14), the accumulation mode particles generally dominated the contribution to $N_{\text{cn}}$, while the reverse was true on some occasions (Fig. 5g, h). The median accumulation mode diameters exhibited an oscillating increase from approximately 100 to $130$ nm at 00:00–08:00 LT, narrowly varied around $135 \pm 5$ nm at 08:00–13:00 LT and then exhibited an oscillating decrease down to approximately 100 nm at 20:00 LT. The AR values and Kappa values at SS = 0.4%, however, exhibited an inverted bell shape with the lowest values at 0.31 and 0.11 at 13:00. The decreases in the AR values and Kappa may be related to organic condensation on the accumulation mode particles since the median accumulation mode diameters were almost the largest at 13:00. The number concentration of Aitken mode particles was evidently enhanced at 14:00–15:00, but the influence on the AR values and Kappa values at SS = 0.4% was undetectable (Fig. 5i).

### 3.4 Particle number size distributions and CCN activation associated with marine traffic emissions and aerosol aging

The particle number size distributions during DOYs 110–118, shown in Fig. 4, can be generally classified into two categories. Category 1 occurred on DOYs 110–114, when particle number concentrations were mainly distributed in Aitken mode, whereas the accumulation mode was generally undetectable. Category 2 occurred on DOYs 115–118, when the accumulation mode could be clearly identified and generally dominated over the Aitken mode. Hoppel (1986) proposed that cloud-modified aerosols are mainly distributed at 80–150 nm in the remote tropical Atlantic and Pacific oceans. Cloud-modified aerosols are quite common in remote marine atmospheres, likely leading to the dominant accumulation mode particles being observed on DOYs 115–118. Occasionally, the Aitken mode dominated over the accumulation mode, such as on DOY 118. To further investigate the sources of different modes of particles, the days DOY 112, DOY 116 and DOY 118 were selected.

On DOY 112, the Aitken mode particles accounted for approximately 60% of the total particle number concentration (Fig. 5a), with median Aitken mode diameters of approximately 54 ± 8 nm. Similar to the observations over the Yangtze River estuary, the mean value of $N_{\text{cn}}$ increased by approximately 50% concurrently with a decrease in the median Aitken mode diameters by $\sim$ 9 nm at 05:30–11:40 LT compared to those in the early morning before 05:30 LT (Fig. 5b). Concomitantly, the AR values decreased to 0.31 ± 0.09 at SS of 0.4%, with similar AR decreases at SS of 1.0%. The lowest AR and Kappa values occurred at 06:00–07:00 LT at SSs of both 0.4% and 1.0%. All these results indicated that the increase in Aitken mode particles at 05:30–11:40 LT was likely derived from enhanced marine traffic contributions carried by the onshore wind from the south (Fig. S11). During other times on DOY 112, the onshore wind may also carry marine-traffic-derived particles to the observational sea zones. However, the marine-traffic-derived particles likely aged to some extent, e.g., the median Aitken mode diameters exhibited an oscillating increase from approximately 50 nm at 19:00 to approximately 70 nm at 24:00 LT with a particle growth rate of $\sim$ 4 nm h$^{-1}$. The AR values, however, narrowly varied around 0.47 ± 0.03 at SS = 0.4% and 0.52 ± 0.05 at SS = 1.0% during the particle growth period. The Kappa values at SS = 0.4% gradually decreased from 0.56 at 19:00 to 0.41 at 23:00 LT, reflecting more aged marine-traffic-derived particles growing into CCN sizes.

On DOY 116, the accumulation mode particles dominantly contributed to $N_{\text{cn}}$ rather than Aitken mode particles (Fig. 5d) under the marine air influence from the northeast (Fig. S13). The median accumulation mode diameters narrowly varied around $135 \pm 5$ nm at 01:00–13:00 LT and $102 \pm 5$ nm at 16:20–24:00 LT with a transition period in between (Fig. 5e). The AR and Kappa values, however, showed no statistically significant differences during the two periods at SSs of 0.4% and 1.0%, implying that the size change in the accumulation mode particles had a negligible influence on CCN activation. The hourly variations in the AR and Kappa values may be associated with other factors, e.g., chemical composition and mixing state (Gunthe et al., 2011; Rose et al., 2011).

On DOY 118, under the influence of mixtures from marine and coastal areas from the northeast (Fig. S14), the accumulation mode particles generally dominated the contribution to $N_{\text{cn}}$, while the reverse was true on some occasions (Fig. 5g, h). The median accumulation mode diameters exhibited an oscillating increase from approximately 100 to $130$ nm at 00:00–08:00 LT, narrowly varied around $135 \pm 5$ nm at 08:00–13:00 LT and then exhibited an oscillating decrease down to approximately 100 nm at 20:00 LT. The AR values and Kappa values at SS = 0.4%, however, exhibited an inverted bell shape with the lowest values at 0.31 and 0.11 at 13:00. The decreases in the AR values and Kappa may be related to organic condensation on the accumulation mode particles since the median accumulation mode diameters were almost the largest at 13:00. The number concentration of Aitken mode particles was evidently enhanced at 14:00–15:00, but the influence on the AR values and Kappa values at SS = 0.4% was undetectable (Fig. 5i).

### 3.5 The long-range transport of newly grown particles on DOY 114

No hour-long sharp increases were observed in the number concentration of the nucleation mode particles (< 20 nm) during the period from DOY 110 to DOY 118, except on DOY 114 (Fig. 4). According to the conventional definition of NPF events (Kulmala et al., 2004; Dal Maso et al., 2005), the occurrence frequency of NPF events was low in this study. Unlike continental atmospheres where a high occurrence frequency of NPF events has been observed globally in spring (Kulmala et al., 2004; Kerminen et al., 2018), a low occurrence frequency reportedly occurred over the seas.
during the “Meiyu (plum-rain) season” in spring because of frequent rainy, foggy or cloudy weather conditions (Zhu et al., 2019). The lack of NPF events in the marine atmospheres implied that the contributions to $N_{cn}$ and $N_{ccn}$ were mainly from primary emitted aerosols and their aged products.

During the period of 10:00–18:00 LT on DOY 114, the large increase in the number concentrations of Aitken mode particles (Fig. 6a) likely reflected the long-range transport of newly grown particles from upwind continental atmospheres (Fig. S12). The size distributions of the particle number concentration showed a dominant Aitken mode at 10:00–18:00 LT, when the spatiotemporal variations in $N_{cn}$ and median Aitken mode diameters exhibited bell-shaped patterns (Fig. 6b). The median Aitken mode diameters increased from 26 nm at 10:00 LT to 33 nm at 12:00–13:00 LT and then decreased to 20 nm prior to the signal disappearance, likely reflecting the growth and shrinkage of the Aitken mode particles (Yao et al., 2010; Zhu et al., 2019). The median Aitken mode diameters were evidently smaller than the values, i.e., 40–50 nm for the Aitken mode particles, observed over the Yangtze River estuary on DOY 112 (Fig. 5a). Moreover, the number concentrations of the 20–40 nm particles increased by 5.8 times at 12:00–13:00 LT compared to the mean value at 06:00–09:00 LT, while the total number concentrations of particles greater than 90 nm increased by only 67%. These results implied that the large increases in the number concentrations of Aitken mode particles with a dynamic change in the mode diameter observed at 10:00–18:00 LT were not likely caused by primary emitted and aged particles from marine traffic emissions or other combustion sources. The observations of the gaseous and particulate species during the same period implied that the air masses were well aged and less polluted. For instance, the measured hourly average mixing ratios of $SO_2$ were no larger than 1.2 ppb (Fig. 6c), and the hourly average concentrations of $NH_4^+$ in PM$_{2.5}$ were smaller than 2 µg m$^{-3}$ (Fig. 3b). In addition, the concentrations of K$^+$ were below 0.3 µg m$^{-3}$, suggesting negligible contributions from biomass burning (Fig. 6e).

Before 09:00 LT, a much weaker spike of nucleation mode particles was intermittently observed (Fig. 6a). The weak and intermittent NPF seemed to occur in the marine atmospheres before 09:00 LT when no apparent growth of new particles was observed. Possibly due to transport from the continent (Fig. S12) and an increase in the condensational sink at approximately 10:00 a.m. (Fig. 6a), the weak NPF signal gradually dropped to a negligible level 0.5 h later, concomitant with a large increase in the number concentrations of Aitken mode particles at 10:00–18:00 LT.

$N_{ccn}$ at $SS = 0.4\%$ increased from $1.2 \times 10^3$ cm$^{-3}$ at 06:00–09:00 LT to the peak value of $2.3 \times 10^3$ cm$^{-3}$ at 12:00 LT, with an increase of 92%, and $N_{ccn}$ at $SS = 1.0\%$ increased from $1.6 \times 10^3$ to $4.0 \times 10^3$ cm$^{-3}$, with an increase of 150% (Fig. 6d). The net increase in $N_{ccn}$ at $SS = 0.4\%$ likely reflected the contribution from pre-existing particles since new particles with diameters less than 50 nm were unlikely to be activated as CCN at such low SS (Li et al., 2015; Wu et al., 2016; Ma et al., 2016). The larger net increase in $N_{ccn}$ at $SS = 1.0\%$ may reflect the contributions from pre-existing particles and newly grown particles. The high SS can activate particles as CCN with diameters down to 40 nm (Dusek et al., 2006; Li et al., 2015). The invasion of newly grown particles also led to the large decreases in the AR values from 0.3 to 0.1 at $SS = 0.4\%$ and from 0.4 to 0.2 at $SS = 1.0\%$ (Fig. 6e). After 18:00 LT, the AR values re-
tuned to 0.3–0.4 at SS = 0.4 % and 0.4–0.6 at SS = 1.0 %. The calculated Kappa values were examined (Fig. 6c) and were found to decrease from 0.4 to 0.1–0.2 at SS = 0.4 %. This value returned to 0.3 at 18:00–19:00 LT (FMPS temporarily malfunctioned after 19:20 LT). The Kappa values were below 0.2 at SS = 1.0 % on that day. The decreases in the AR values and Kappa values at the two SS were likely caused by organic vapor condensed on pre-existing particles and new particles (Wu et al., 2016; Zhu et al., 2019).

3.6 Relationship of $N_{cn}$ and $N_{ccn}$ with SO$_2$ in ship self-plumes and ambient air

When ship self-emission signals were detected, the observational values included a combination of contributions from ship self-emissions and ambient concentrations. Although the ambient $N_{cn}$ was negligible in comparison with the $N_{cn}$ derived from the ship self-emissions, this was not the case for $N_{ccn}$ and SO$_2$. Based on the per minute data, the signal was considered to be vessel self-emissions when both $N_{cn}$ was greater than 50 000 cm$^{-3}$ and SO$_2$ was greater than 5 ppb.

The composited data were then used to derive the hourly average $N_{cn}$, $N_{ccn}$ and SO$_2$, which were then subtracted by the ambient hourly mean values during the preceding hour with relatively clean conditions (i.e., concentrations of $N_{cn}$ lower than 10 000 cm$^{-3}$ and SO$_2$ lower than 2.5 ppb). Please note that uncertainties exist in terms of the criteria and separation between the ship self-signals and ambient signals; however, minimal impact is expected in the relationship examined below.

Figure 7a shows the relationship of $N_{cn}$ and $N_{ccn}$ with the mixing ratio of SO$_2$ in the ship self-plumes, prefixed by $\Delta$ for $N_{cn}$, $N_{ccn}$ and SO$_2$ to implicate the removal of ambient signals. A good correlation of 0.66 for $R^2$ ($P < 0.01$) is obtained, and the slope indicates that the increase in $N_{cn}$ by $1.4 \times 10^4$ cm$^{-3}$ for each 1 ppb increase in SO$_2$ resulted from ship emissions (Fig. 7a). High emissions of $N_{cn}$ were generally reported in engine exhausts where high-sulfur-content diesel was used (Yao et al., 2005, 2007). The $N_{ccn}$ at SS of 0.2 % to 1.0 % (Fig. 7b) increased from 30 to 170 cm$^{-3}$ per 1 ppb increase in SO$_2$, showing a statistically significant correlation at the 99th confidence level. The contribution ratios of SO$_2$ to $N_{ccn}$ were 0.002 (SS of 0.2 %), 0.004 (SS of 0.4 %) and 0.012 (SS of 1.0 %) to that of $N_{cn}$, which is in general consistent with a previous study by Ramana and Devi (2016), in which a range of 0.0012–0.57 was observed for CCN at 0.4 % in Bay of Bengal during 13–16 August 2012.

The relationship of hourly averaged $N_{cn}$ and $N_{ccn}$ with SO$_2$ in ambient air were examined and are shown in Fig. 7c, d. The data were segmented into pieces based on SO$_2$ with an interval of 0.2 ppb. A good correlation was ob-
tained between the averaged $N_{cn}$ and SO$_2$ with an $R^2$ of 0.80 ($P < 0.01$), and a 1 ppb increase in SO$_2$ likely increased $N_{cn}$ by $1.6 \times 10^3$ cm$^{-3}$ (Fig. 7c). The increase in $N_{cn}$ with SO$_2$ may reflect the contribution from primary emissions. The intercept was, however, as large as $3.9 \times 10^3$ cm$^{-3}$, likely representing the contribution from well-aged aerosols.

The hourly averaged $N_{ccn}$ at different SSs generally increased with increasing ambient SO$_2$ (Fig. 7d). A good correlation was obtained between the averaged $N_{ccn}$ and SO$_2$, with $R^2 = 0.78$–0.91 ($P < 0.01$). A 1 ppb increase in SO$_2$ likely increased $N_{ccn}$ by $0.6 \times 10^3$ to $0.8 \times 10^3$ cm$^{-3}$ at SSs from 0.2 % to 1.0 %. The increase in $N_{ccn}$ with SO$_2$ may also reflect the contribution from primary emissions. The intercepts of $2.2 \times 10^3$–$2.7 \times 10^3$ cm$^{-3}$ at different SSs were likely contributed by well-aged aerosols. This relationship may be used as an estimation of the $N_{ccn}$ in marine atmospheres over marginal seas in China when no measurements of CCN are available, whereas the ambient SO$_2$ can be estimated from web-based satellite data.

4 Conclusions

The spatiotemporal variations in ambient $N_{cn}$ and $N_{ccn}$ were studied during a cruise campaign on DOYs 110–135 over marginal seas in China. The mean values of $N_{cn}$ (8.1 $\times$ $10^3$ cm$^{-3}$) and $N_{ccn}$ (3.2–3.9 $\times$ $10^3$ cm$^{-3}$) at SSs of 0.2 %–1.0 % were approximately 1 order of magnitude larger than those in remote clear marine atmospheres, indicating overwhelming contributions from non-sea-spray aerosols such as marine traffic emissions, long-range continental transport and others.

The observed ship self-emission signals showed that fresh marine traffic emissions can be important sources of $N_{cn}$ but minor sources of $N_{ccn}$ in the marine atmosphere. The signals showed that a 1 ppb increase in SO$_2$ corresponded to a $1.4 \times 10^4$ cm$^{-3}$ increase in $N_{cn}$ and a 30–170 cm$^{-3}$ increase in $N_{ccn}$ at SS = 0.2 %–1.0 %. Data analysis showed that marine traffic emissions largely increased $N_{cn}$ over heavily traveled sea zones in the daytime.

In ambient marine air, the growth of marine-traffic-derived particles led to a decrease in the estimated bulk Kappa values at 0.4 %, possibly because some of these particles enriched in organics grew into CCN size. However, strong formation of ammonium salts led to aerosol aging and significantly increased $N_{ccn}$ at SS of 0.2 %–1.0 % in comparison with those observed during the period poor in ammonium salt aerosols in PM$_{2.5}$ with $P < 0.01$. The estimated bulk Kappa values from the daily average values varied from 0.46 to 0.55 at SS = 0.4 % in most marine atmospheres, indicating that inorganic ammonium aerosols may dominantly contribute to the $N_{ccn}$ at SS of 0.4 %. The particle number size distributions showed that the high bulk Kappa values could be related to cloud-modified aerosols, which likely led to a large extent of degradation of organics and subsequent loss from the particle phase.

NPF events rarely occurred in the humid ambient marine air. The dominant onshore winds occurred during most of the measurement periods and likely carried primary aerosols and their aged products rather than secondarily formed aerosols to the observational zone. During an occasion when offshore winds blew from the northwest (Fig. S12), new particle signals transported from the continent can be clearly observed. However, the NPF in the marine atmosphere was too weak to be important. The new transported particles from the continent yielded the maximal increases in $N_{ccn}$ of 92 % at SS of 0.4 % and 150 % at SS of 1.0 %. However, consistent with those reported in the literature, the estimated Kappa values largely decreased from 0.4 to 0.1–0.2 at SS of 0.4 % during most of the continent-transporting NPF event because the Kappa value of the organic condensation vapor was as low as 0.1.

Data availability. The data is available upon request to the corresponding authors.

Supplement. The supplement related to this article is available online at: https://doi.org/10.5194/acp-20-9665-2020-supplement.

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Competing interests. The authors declare that they have no conflict of interest.

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