



Contrasting size-resolved hygroscopicity of fine particles derived by HTDMA and HR-ToF-AMS measurements between summer and winter in Beijing: the impacts of aerosol aging and local emissions

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Abstract. The effects of aerosols on visibility through scattering and absorption of light and on climate through altering cloud droplet concentration are closely associated with their hygroscopic properties. Here, based on field campaigns in winter and summer in Beijing, we compare the size-resolved hygroscopic parameter (κ_{gf}) of ambient fine particles derived by an HTDMA (hygroscopic tandem differential mobility analyzer) to that (denoted as κ_{chem}) calculated by an HR-ToF-AMS (high-resolution time-of-flight aerosol mass spectrometer) measurements using a simple rule with the hypothesis of uniform internal mixing of aerosol particles. We mainly focus on contrasting the disparity of κ_{gf} and κ_{chem} between summer and winter to reveal the impact of atmospheric processes/emission sources on aerosol hygroscopicity and to evaluate the uncertainty in estimating particle hygroscopicity with the hypothesis. We show that, in summer, the κ_{chem} for 110, 150, and 200 nm particles was on average $\sim 10\%$ – 12% lower than κ_{gf} , with the greatest difference between the values observed around noontime when aerosols experience rapid photochemical aging. In winter, no apparent disparity between κ_{chem} and κ_{gf} is observed for those > 100 nm particles around noontime, but the κ_{chem} is much higher than

κ_{gf} in the late afternoon when ambient aerosols are greatly influenced by local traffic and cooking sources. By comparing with the observation from the other two sites (Xingtai, Hebei and Xinzhou, Shanxi) of north China, we verify that atmospheric photochemical aging of aerosols enhances their hygroscopicity and leads to 10% – 20% underestimation in κ_{chem} if using the uniform internal mixing assumption. The effect is found more significant for these > 100 nm particles observed in remote or clean regions. The lower κ_{chem} likely resulted from multiple impacts of inappropriate application of the density and hygroscopic parameter of organic aerosols in the calculation, as well as influences from chemical interaction between organic and inorganic compounds on the overall hygroscopicity of mixed particles. We also find that local/regional primary emissions, which result in a large number of externally mixed BC (black carbon) and POA (primary organic aerosol) in urban Beijing during traffic rush hour time, cause a 20% – 40% overestimation of the hygroscopic parameter. This is largely due to an inappropriate use of density of the BC particles that is closely associated with its morphology or the degree of its aging. The results show that the calculation can be improved by applying an effective

tive density of fresh BC ($0.25\text{--}0.45\text{ g cm}^{-3}$) in the mixing rule assumption. Our study suggests that it is critical to measure the effective density and morphology of ambient BC, in particular in those regions with influences of rapid secondary conversion/aging processes and local sources, so as to accurately parameterize the effect of BC aging on particle hygroscopicity.

1 Introduction

The effects of aerosols on visibility through scattering and absorption of light and on climate through altering cloud droplet concentration are influenced by their hygroscopic growth. Understanding and reducing the uncertainty in prediction of the aerosol hygroscopic parameter (κ) using chemical composition would improve model predictions of aerosol effects on clouds and climate.

The hygroscopic properties of both the natural and anthropogenic aerosols, in addition to being affected by its chemical composition (Gunthe et al., 2009), are also affected by the particle mixing state and aging (Schill et al., 2015; Peng et al., 2017a). For example, a recent laboratory study showed that the coexisting hygroscopic species have a strong influence on the phase state of particles, thus affecting chemical interactions between inorganic and organic compounds as well as the overall hygroscopicity of mixed particles (C. Peng et al., 2016). The field measurements also demonstrated that the hydrophobic black carbon particles became hygroscopic with atmospheric mixing and aging by organics (i.e., Peng et al., 2017a). In a heavily polluted atmosphere with varied aerosol sources and sinks as well as complex physical and chemical processes, the mixing state and its impact on aerosol hygroscopicity is more complicated. The hygroscopicity of mixed particles and mutual impacts between the components are still poorly understood.

Previous studies have presented the difference between the κ obtained from HTDMA or CCNc (cloud condensation nuclei counter) measurements and that calculated based on the volume mixing ratio of chemical components. Laboratory results from Cruz and Pandis (2000) indicate that κ_{gf} of internally mixed ammonium sulfate and organic matter is higher than κ_{chem} calculated for assumed uniform internal mixing. But C. Peng et al. (2016) found that, for sodium chloride and organic aerosol mixed particles, the measured growth factors by HTDMA were significantly lower than calculations from the mixing rule methods. In some field studies on aged aerosols, the κ was underestimated by the calculation based on a uniform internal mixing assumption, and they thus lead to an underestimation of CCN concentration (Bougiatioti, et al., 2009; Chang et al., 2007; Kuwata et al., 2008; Wang et al., 2010; Ren et al., 2018). However, during primary-emission-dominated periods, the κ value from calculations based on bulk chemical composition was much higher than that mea-

sured by HTDMA measurements (Zhang et al., 2017). The various results from previous studies suggest distinct effects of aerosol mixing state on their hygroscopicity. Overall, (i) to what extent the differences depend on the mixing state and the extent of aging of the particles and (ii) how the different atmospheric processes and what kinds of mixing structure of the particles may result in the disparity between the measured and calculated hygroscopic parameter have not been clearly clarified by the previous studies. A comprehensive investigation on the causes and magnitude of the effect is of great significance to parameterize the effect of atmospheric processes/emissions of aerosols on particle hygroscopicity in models.

In the atmosphere, the κ , which is related to the particle mixing state diversity, varies largely across the size range of ambient fine particles (Rose et al., 2010). However, previous studies just compared the κ calculated from bulk chemical composition to that measured by HTDMA (Zhang et al., 2017). Using size-resolved, not bulk chemical composition measurements in different seasons is expected to provide a more comprehensive understanding and insights into how the aerosol mixing state influences their hygroscopicity, motivating our analysis that employs size-resolved chemical composition measured by a high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS) in this study. The aim of this paper is to study the hygroscopicity and mixing state characteristics of fine particles in the Beijing urban area, as well as to reveal the impact of atmospheric processes/sources and mixing/aging on aerosol hygroscopicity and elucidate the uncertainty in calculating the hygroscopic parameter using simple mixing rule estimates based on size-resolved chemical composition. The experiment and theory in the study are introduced in Sect. 2. The comparison between the hygroscopic parameter obtained from the HTDMA (hygroscopic tandem differential mobility analyzer) and that calculated using size-resolved chemical composition is discussed in Sect. 3. Conclusions from the study are given in Sect. 4.

2 Experiment and theory

2.1 Site and instruments

Two field campaigns are conducted during winter 2016 and summer 2017 of urban Beijing (Fig. 1, BJ: 39.97°N , 116.37°E) for measurements of aerosol physical and chemical properties. The BJ site is located at the Institute of Atmospheric Physics (IAP), Chinese Academy of Sciences, which is between the north Third Ring Road and Fourth Ring Road in northern Beijing. Local traffic and cooking emissions can be important at the site (Sun et al., 2015). The sampling period in the cold season was from 16 November to 10 December 2016, during the domestic heating period in Beijing.

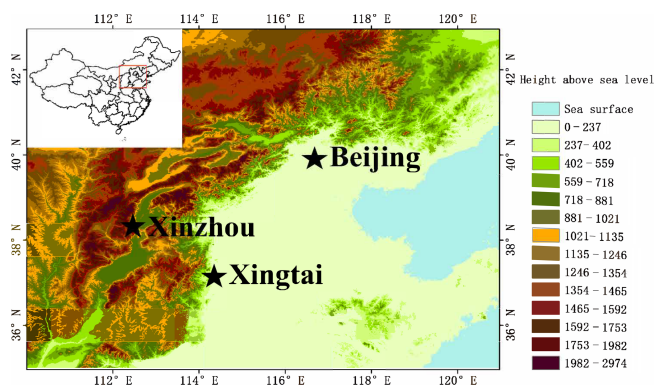


Figure 1. The map location of the sites.

The sampling period in the warm season was from 25 May to 18 June 2017.

The particle number size distribution (PNSD) in the size range from 10 to 550 nm was measured with a scanning mobility particle sizer (SMPS; Wang and Flagan, 1990; Collins et al., 2002), which consists of a long differential mobility analyzer (DMA, model 3081L, TSI Inc.) to classify the particle and a condensation particle counter (CPC, model 3772, TSI Inc.) to detect the size-classified particles. The sampled particles were dried to a relative humidity < 30 % before entering the DMA. The measurement time for each size distribution was 5 min.

The HTDMA system used in this study has been described in detail in previous publications (Tan et al., 2013; Wang et al., 2017; Zhang et al., 2017). Here, only a brief description is given. A Nafion dryer dried the sampled particles to a relative humidity < 20 %, after which the steady-state charge distribution was reached in a bipolar neutralizer. The first differential mobility analyzer (DMA₁, model 3081L, TSI Inc.) selected the quasi-monodisperse particles by applying a fixed voltage. The dry diameters selected in this study were 40, 80, 110, 150, and 200 nm. The quasi-monodisperse particles were humidified to a controlled RH (90 % in this study) using a Nafion humidifier. A second DMA (DMA₂, same model as the DMA₁) coupled with a water-based condensation particle counter (WCPC, model 3787, TSI Inc.) measured the particle number size distributions of the humidified aerosol. RH calibration with ammonium sulfate was carried out regularly during the study.

The hygroscopic growth factor (Gf) is defined as the ratio of the mobility diameter at a given RH to the dry diameter:

$$\text{Gf} = \frac{D(\text{RH})}{D(\text{dry})}. \quad (1)$$

The Gf probability density function (PDF) is retrieved based on the TDMA_{inv} algorithm developed by Gysel et al. (2009). Dry scans in which the RH between the two DMAs was not increased were used to define the width of the transfer function.

Size-resolved nonrefractory submicron aerosol composition was measured with an Aerodyne high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS; Xu et al., 2015). The particle mobility diameter was estimated by dividing the vacuum aerodynamic diameter from the AMS measurements by particle density. Because the uncertainty caused by the fixed density across the size range is negligible (Wang et al., 2016), here, the particle density is assumed to be 1600 kg m⁻³ (Zamora et al., 2019). AMS positive matrix factorization (PMF) with the PMF2.exe (v4.2) method was performed to identify various factors of organic aerosols. Xu et al. (2015) have described the operation and calibration of the HR-ToF-AMS in detail. Black carbon (BC) mass concentration was derived from measurements of light absorption with a seven-wavelength aethalometer (AE33, Magee Scientific Corp.; Zhao et al., 2017).

2.2 Data

The time series of the submicron particle mass concentration PM₁, bulk mass concentrations of the main species in PM₁, mass fraction of the chemical composition of PM₁, and probability density function of growth factor (Gf PDFs) for 40 and 150 nm particles during the campaign are presented in Fig. 2. Quite distinct temporal variability of aerosol chemical and physical properties was observed between winter and summer. The average mass concentration of PM₁ was 55.2 µg m⁻³ in the winter and 16.5 µg m⁻³ in the summer during our study periods. In this study, we define the conditions when the mass concentration in winter period was < 20 and > 80 µg m⁻³ for clean and polluted conditions, respectively. Organic aerosol (OA), consisting of secondary organic aerosol (SOA) and primary organic aerosol (POA), was the major fraction during both the winter and summer sampling periods. POA concentration was higher than that of SOA in the winter, which reflects the influence of primary emissions such as coal combustion OA (COOA) in Beijing (Hu et al., 2016; Sun et al., 2016). In contrast, SOA usually dominated in the summer, which is evidence that secondary aerosol formation played a key role in the source of PM₁. Distinct hydrophobic (with Gf of ~ 1.0) and more hygroscopic (with Gf of ~ 1.5) modes were observed from Gf PDFs of both small and large particles. Sometimes the more hygroscopic mode particles were more concentrated, and at other times the hydrophobic particles were. In general though, the more hygroscopic mode dominated for larger particles (i.e., 150 nm), and the less hygroscopic mode did for the smallest particles (e.g., 40 nm). Occasionally, only the hydrophobic mode was evident for 150 nm particles, which occurred when POA dominated the PM₁. Only the hygroscopic mode was discernable for 40 nm particles during new particle formation (NPF) events that occurred more frequently in summer than winter (Fig. 3).

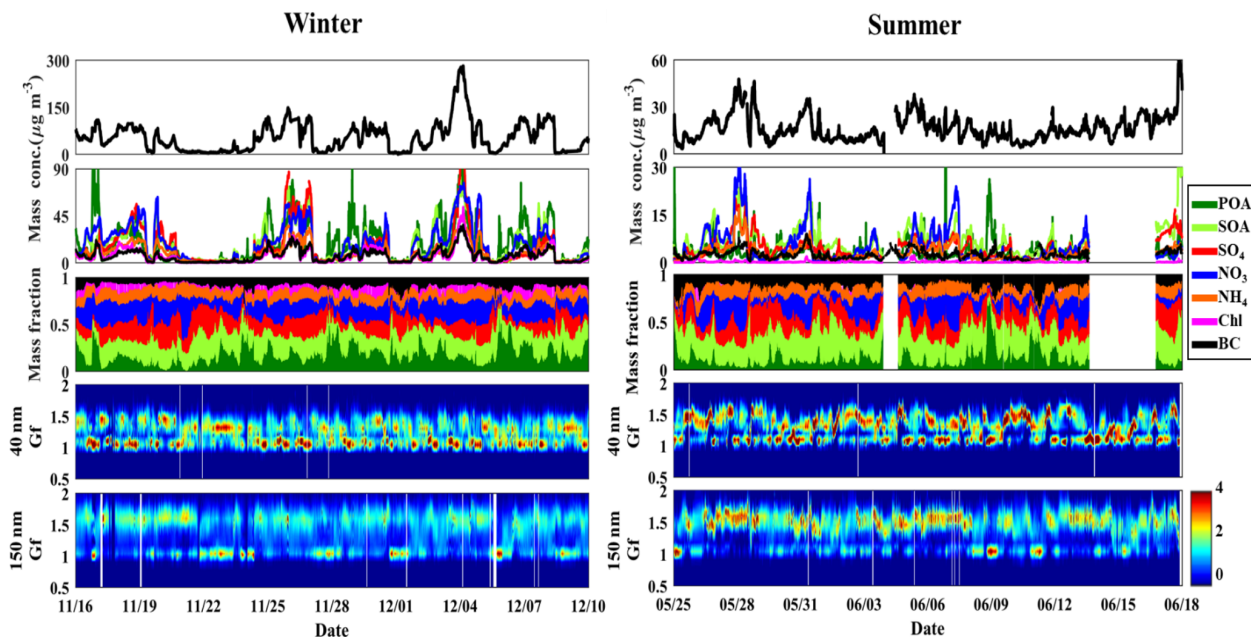


Figure 2. Winter (left) and summer (right) time series of mass concentration of PM₁, bulk mass concentration of the main species in PM₁, mass fraction of the chemical composition of PM₁, and Gf PDFs for 40 and 150 nm particles.

2.3 Theory and method

2.3.1 Derivation of the hygroscopic parameter, κ , from the growth factor (Gf)

According to κ -Köhler theory (Petters and Kreidenweis, 2007), the hygroscopicity parameter κ can be derived using the growth factor measured by an HTDMA.

$$\kappa = (\text{Gf}^3 - 1) \left(\frac{\exp\left(\frac{A}{D_d \text{Gf}}\right)}{\text{RH}} - 1 \right), \quad (2)$$

$$A = \frac{4\sigma_{s/a}M_w}{RT\rho_w}, \quad (3)$$

where Gf is hygroscopic growth factor measured by HTDMA, D_d is the dry diameter of the particles, RH is the relative humidity in the HTDMA (90 %, in our study), $\sigma_{s/a}$ is the surface tension of the solution/air (assumed here to be the surface tension of pure water, $\sigma_{s/a} = 0.0728 \text{ N m}^{-2}$), M_w is the molecular weight of water, R is the universal gas constant, T is the absolute temperature, and ρ_w is the density of water.

2.3.2 Derivation of the hygroscopic parameter, κ , from chemical composition data

For an assumed internal mixture, κ can also be calculated by a simple mixing rule on the basis of chemical volume fractions (Petters and Kreidenweis, 2007; Gunthe et al., 2009):

$$\kappa_{\text{chem}} = \sum_i \varepsilon_i \kappa_i, \quad (4)$$

where κ_i and ε_i are the hygroscopicity parameter and volume fraction for the individual (dry) component in the mixture, respectively. The AMS provides mass concentrations of organics and of many inorganic ions. The inorganic components mainly consisted of $(\text{NH}_4)_2\text{SO}_4$ and NH_4NO_3 (Zhang et al., 2014). And the values of κ are 0.48 for $(\text{NH}_4)_2\text{SO}_4$ and 0.58 for NH_4NO_3 (Petters and Kreidenweis, 2007). To estimate κ_{org} , we used the following linear function derived by Mei et al. (2013): $\kappa_{\text{org}} = 2.10 \times f_{44} - 0.11$. We derived the volume fraction of each species by dividing mass concentration by its density. The densities are 1.77 g cm^{-3} for $(\text{NH}_4)_2\text{SO}_4$ and 1.72 g cm^{-3} for NH_4NO_3 . The densities of organics are assumed to be 1.2 g cm^{-3} (Turpin and Lim, 2001). The κ and density of BC are assumed to be 0 and 1.7 g cm^{-3} . In the following discussions, κ_{gf} and κ_{chem} denote the values derived from HTDMA measurements and calculated using the Zdanovskii–Stokes–Robinson mixing rule (Stokes and Robinson, 1966), respectively.

In addition, we also compare the results from the field campaigns with those from the other two sites, Xingtai (XT: 37.18° N–114.37° E) and Xinzhou (XZ: 38.24° N–112.43° E), in the North China Plain (Fig. 1). At the XZ site, we use the hygroscopic parameter (defined as κ_{CCNC}) from size-resolved CCN measurements (Zhang et al., 2014, 2016) for comparison. More detailed descriptions of the method to retrieve κ_{CCNC} can be found in Petters and Kreidenweis (2007). Both of the κ_{gf} and κ_{CCNC} are derived based on κ -Köhler theory (Petters and Kreidenweis, 2007). But, different from the κ_{gf} measured by the HTDMA system which is operated at a RH of 90 %, the κ_{CCNC} is derived by mea-

suring aerosol CCN activity under the condition of supersaturations with relative humidity of $> 100\%$. Previous studies from field measurements and laboratory experiments showed that the κ_{CCNc} is generally slightly larger or smaller than κ_{gf} , but they are basically comparable and can well represent an overall aerosol hygroscopicity (e.g., Carrico et al., 2008; Wex et al., 2009; Good et al., 2010; Irwin et al., 2010; Cerully et al., 2011; Wu et al., 2013; Zhang et al., 2017).

3 Results and discussion

3.1 Diurnal variations of ambient-fine-particle physiochemical properties and hygroscopic growth factor

The diurnal variations of the PNSD, mass concentration of PM_{10} , mass concentration and fraction of chemical components in PM_{10} , and Gf PDFs for 40 and 150 nm particles during the campaign are shown in Fig. 3. During the summer an obvious peak value in the PNSD is observed around noon-time due to NPF events that typically started around 10:00 LT (local time). The resulting sharp increase in number concentration of nucleation mode particles was followed by decreased concentration and a rapid growth in diameter of the particles along with increased mass concentration of SOA and sulfate in PM_{10} , indicating strong photochemical and secondary formation processes during daytime in the summer (Peng et al., 2017b; shown by the red box in Fig. 3). In contrast, NPF was not evident during the winter period, which may in part be due to the much higher ($\sim 3\times$) PM_{10} mass concentrations in the winter than in the summer. Note that peak values in number concentration and in mass concentrations of PM_{10} and POA occurred during the early evening (17:00–21:00 LT), indicating the strong impact of local sources from traffic emissions and cooking (shown by the black box in Fig. 3; Peng et al., 2014). In addition, the diurnal cycles of aerosol physical and chemical properties are also influenced by the diurnal changes in the planetary boundary layer (PBL) that lead to the accumulation of particles during nighttime when higher values of both number and mass concentration were observed.

Owing to the continued local and primary emissions near the study site, the Gf PDFs for 40 nm particles generally display a bimodal shape with more and less hygroscopic modes (with Gf of ~ 1.5 and ~ 1.1 respectively) throughout the day in both winter and summer periods, indicating an external mixing state for the 40 nm particles. Note that, during nighttime and early morning in the winter, the more hygroscopic mode dominated and was shifted to higher Gf than during the daytime. This is thought to be due to heterogeneous/aqueous reactions on preexisting primary small particles and/or coagulation/condensation processes that are enhanced at night under lower ambient temperature and higher relative humidity, all of which result in a more hygroscopic and more inter-

nally mixed aerosol (Liu et al., 2011; Massling et al., 2005; Ye et al., 2013; Wu et al., 2016; Wang et al., 2018a). Interestingly, in the summer period, the concentration of the hydrophilic mode increased quickly around noontime and in the early afternoon (12:00–16:00), with a corresponding decrease in the relative concentration of the hydrophobic mode, which likely indicates a transformation of the particles from an externally mixing state to an internally mixing state as a result of the species condensation from the photochemical reaction (Wu et al., 2016; Wang et al., 2017), resulting in an increase in particle hygroscopicity. In addition, it is evident that 40 nm particles after 12:00 were dominated by NPF (Fig. 3). Therefore, the increase of hydrophobic mode particles suggests that a large amount of hydrophilic particles are generated from NPF. For 150 nm particles, the hygroscopic mode in the Gf PDF is more dominant during daytime, in particular during the summer period when the strong solar radiation promotes photochemical aging and growth, thus producing a more internally mixed aerosol. The dominant hydrophobic mode at around 18:00 was observed in both winter and summer and reflects abundant traffic emissions and cooking sources (primarily with POA) during the early evening period.

3.2 κ_{gf} dependence on D_p

The size dependence of particle hygroscopicity parameters for the winter and summer periods is presented in Fig. 4. In the winter, the 40 nm particles were the least hygroscopic, and the hygroscopicity of larger particles (> 80 nm) displayed an insignificant dependence on particle size. The size independence for the larger particles is consistent with the observed similarity in mass fractions of inorganic and organic species across the size range as shown in the pie charts in Fig. 4a. A similar dependence of particle hygroscopicity on particle size was also observed in the urban area of Beijing during the wintertime of 2014 (Wang et al., 2018b). In the summer, hygroscopicity increased with increasing particle size, which is expected based on the size-dependent patterns shown in the pie charts, with the mass fraction of POA decreasing with the particles size and the mass fraction of inorganics like sulfate and nitrate increasing with particle size.

3.3 Closure of HTDMA and chemical-composition-derived κ

A closure study was conducted between κ_{chem} and κ_{gf} (Fig. 5) to investigate the uncertainty of the two methods, and especially to further illustrate whether particle hygroscopicity can be well predicted by κ_{chem} calculated by assuming internal mixing. Since a size-resolved BC mass concentration measurement was not available during the campaign, we use the bulk mass fraction of BC particles measured by the AE33 combining with size-resolved BC distribution measured by a single particle soot photometer (SP2) in Beijing

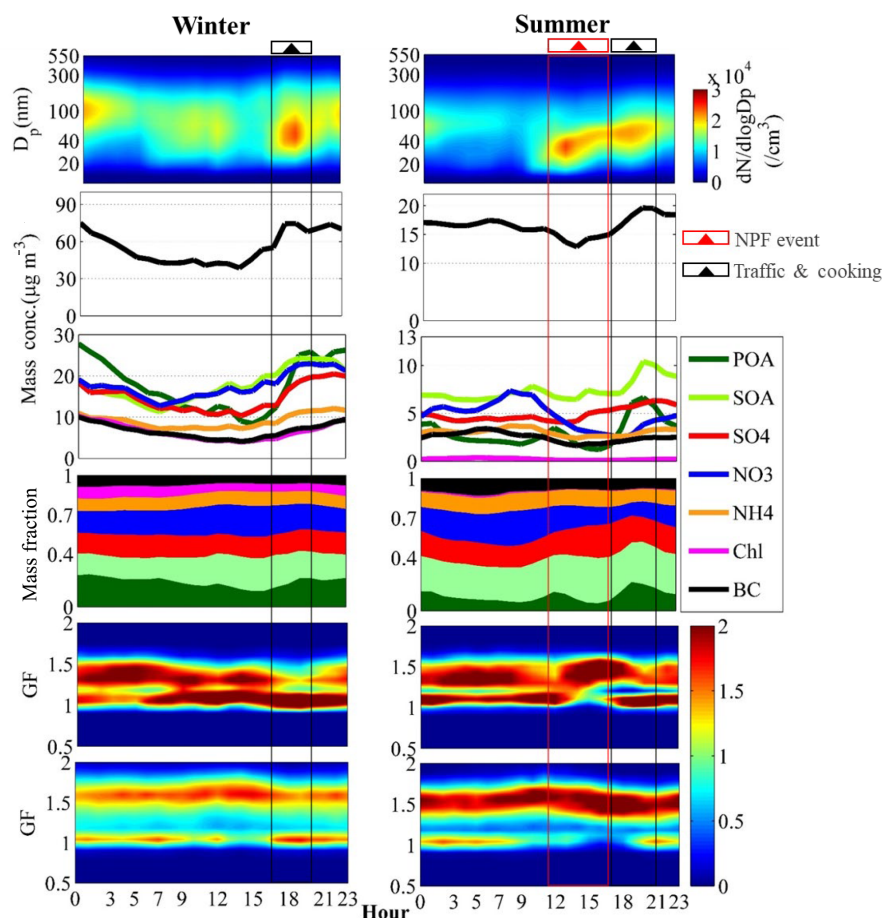


Figure 3. Campaign averaged diurnal variations in particle number size distribution; mass concentration of PM_{10} , bulk mass concentration of main species in PM_{10} , and mass fraction of chemical composition of PM_{10} ; and Gf PDFs for 40 and 150 nm particles in winter (left panels) and summer (right panels) measured in urban Beijing.

(Liu et al., 2019) to estimate κ_{chem} . During the calculation, the BC core diameter measured by SP2 has been converted to the diameter of coated BC particles by multiplying factors of 1.4 and 2.6 under clean (with bulk BC mass concentrations $< 2 \mu\text{g m}^{-3}$) and polluted (with bulk BC mass concentrations $> 2 \mu\text{g m}^{-3}$) conditions respectively (Liu et al., 2019).

Uncertainty in κ is due in part to measurement uncertainty of the HTDMA system and uncertainty resulting from non-ideality effects in the solution droplets, surface tension reduction due to surface-active substances, and the presence of slightly soluble substances that dissolve at a RH higher than that maintained in the HTDMA (e.g., Wex et al., 2009; Good et al., 2010; Irwin et al., 2010; Cerully et al., 2011; Wu et al., 2013). For example, the HTDMA may overestimate the D_p of dry particles for the external mixed BC particles, as BC-containing particles may shrink when humidified, leading to an underestimation of the hygroscopic growth factor. However, our previous study demonstrated that, for this region, estimates using HTDMA data are still better representing the aerosol hygroscopicity than those using the simple mixing

rule based on chemical volume fractions for an assumed internal mixture (Zhang et al., 2017). Therefore, here we focus on discussing and exploring the uncertainty of κ_{chem} by taking κ_{gf} as the reference.

The results show that, although the slopes from linear fitting of κ_{chem} and κ_{gf} are close to 1.0, it shows correlations that are quite poor (typically with correlation coefficients, R^2 , of < 0.3) between κ_{chem} and κ_{gf} of the 80, 110, 150, and 200 nm particles in both winter and summer. The poor correlations reflect large uncertainty in one or both of the calculated parameters that are likely due to the unreasonable assumption of particle mixing state (e.g., Cruz and Pandis, 2000; Svenningsson et al., 2006; Sjogren et al., 2007; Zardini et al., 2008), which varies with their aging and other physiochemical processes in the atmosphere. Note that underestimation of κ_{chem} for the summer occurred mostly in the afternoon (shown in blue dots in Fig. 5). This may be associated with photochemical processes at around noontime. More specific investigations of the particle mixing and aging

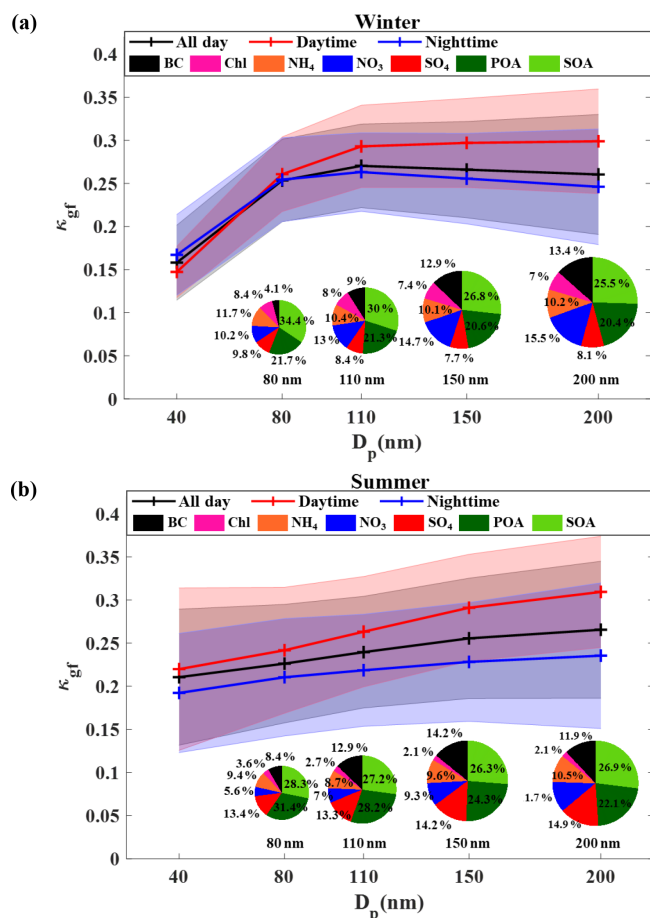


Figure 4. The dependence of κ on D_p at the urban Beijing site during winter (a) and summer (b). The κ values are retrieved from the size-resolved HTDMA measurements. The error bars represent $\pm 1\sigma$. The size-resolved chemical mass fraction at the corresponding D_p is also presented.

impacts on κ_{chem} will be further addressed in the following sections.

3.4 Aerosols aging and source effects indicated by diurnal cycles of κ_{chem} and κ_{gf}

The diurnal cycles of particle hygroscopicity in the summer and winter with the use of the size-resolved chemical composition observations and the ratio of κ_{chem} to κ_{gf} are shown in Fig. 6. In summer, at 09:00–15:00, the disparity between κ_{chem} and κ_{gf} is insignificant for smaller particles (80 and 110 nm), and both of them show a slight decrease from 09:00 or 10:00 to 12:00–13:00 due to the frequent NPF event that usually corresponds to a large fraction of organics (Fig. 3) in urban Beijing. For larger particles (150 and 200 nm), the disparity between κ_{chem} and κ_{gf} around noontime and in the early afternoon is very significant, corresponding to $> 20\%$ underestimation of particle hygroscopicity by κ_{chem} (with a ratio of κ_{chem} to κ_{gf} of ~ 0.8). Similar patterns were also

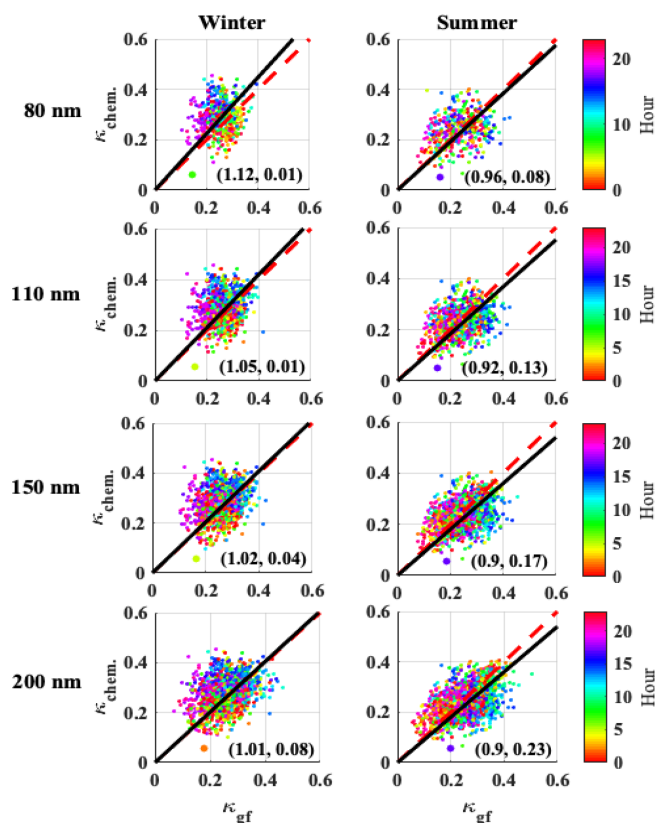


Figure 5. Closure of κ_{chem} calculated from size-resolved chemical composition data and κ_{gf} retrieved from the hygroscopic growth factor by HTDMA measurements in the winter (left panels) and summer (right panels) period. The dots with different colors correspond to the observed time of a day during the campaign as shown by the color bar. In each plot, the red dotted line is the 1 : 1 line, and the black solid line is the fitting line. The numbers in parentheses are slopes of linear fits and correlation coefficients (R^2).

noted by Zhang et al. (2017) but only based on a comparison between κ_{chem} derived from bulk chemical composition and κ_{gf} . Our results based on size-resolved measurements are consistent with that observed by Zhang et al. (2017), which again confirms the effect of the rapid photochemical aging of aerosol particles on their hygroscopicity. In contrast, no significant differences between κ_{chem} and κ_{gf} are observed during nighttime in summer. Note that κ_{chem} is slightly higher than κ_{gf} during early evening traffic rush hour and cooking time, when emissions of primary hydrophobic particles (e.g., BC and POA) are high (Fig. 3), thus resulting in a large percentage of externally mixed particles. Causes of the overestimation of κ_{chem} during the traffic rush hour and cooking time will be discussed in the following paragraph. The particles experience rapid conversion and mixing in urban Beijing due to high precursor gases (Sun et al., 2015; Wu et al., 2016; Ren et al., 2018); thus, the aged particles produced through photochemical processes in the afternoon can mix and interact with the freshly emitted primary particles from traffic and

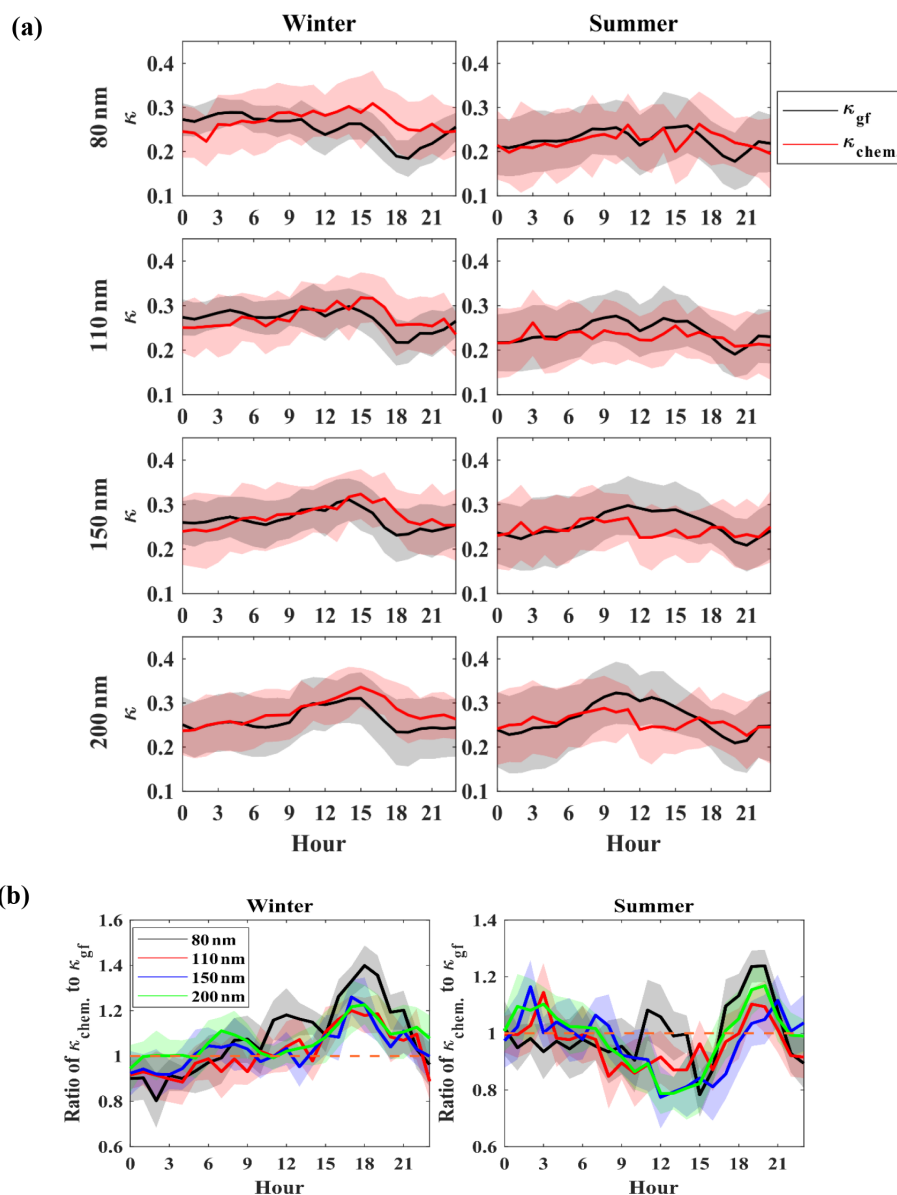


Figure 6. Diurnal variations of (a) κ_{chem} using size-resolved chemical composition data and κ_{gf} in the winter and summer period and (b) the ratio of κ_{chem} to κ_{gf} in the winter and summer period. The shade regions denote the error bars (1σ).

cooking sources (Wu et al., 2008). Therefore, during nighttime (22:00–06:00 LT), the particles are more uniform and internally mixed, which is reflective of the assumption for the calculation of κ_{chem} ; a much better consistency between κ_{chem} and κ_{gf} is hence presented.

In winter, the disparity between κ_{chem} and κ_{gf} is insignificant at 09:00–15:00 due to the weakening effect of photochemical aging. From 15:00 to 21:00 LT, due to the strong vehicle and cooking sources around the site, the particles are dominated by the hydrophobic mode with a large concentration of externally mixed BC and POA particles (Fig. 3); the calculated κ_{chem} is much higher than κ_{gf} , with a maximum ratio of κ_{chem} to κ_{gf} of 1.4; and the greatest disparity

is observed for small particles. The disparity is further enhanced during clean periods when the hydrophobic mode is dominant (Figs. 7, S1 in the Supplement). Note that during the nighttime, κ_{chem} is slight lower than κ_{gf} , with a minimum ratio of κ_{chem} to κ_{gf} of ~ 0.8 for 80 nm particles and ~ 0.9 for 110 and 150 nm particles at 02:00–04:00 LT (Fig. 6b), indicating an underestimation of particle hygroscopicity using composition data. The disparity at nighttime is further increased during heavily polluted events (Fig. S1), when the particles are more internally mixed with only one hygroscopic mode (Fig. 7). We propose the increased underestimation during polluted conditions is likely due to enhanced condensation of secondary hygroscopic compounds (e.g., ni-

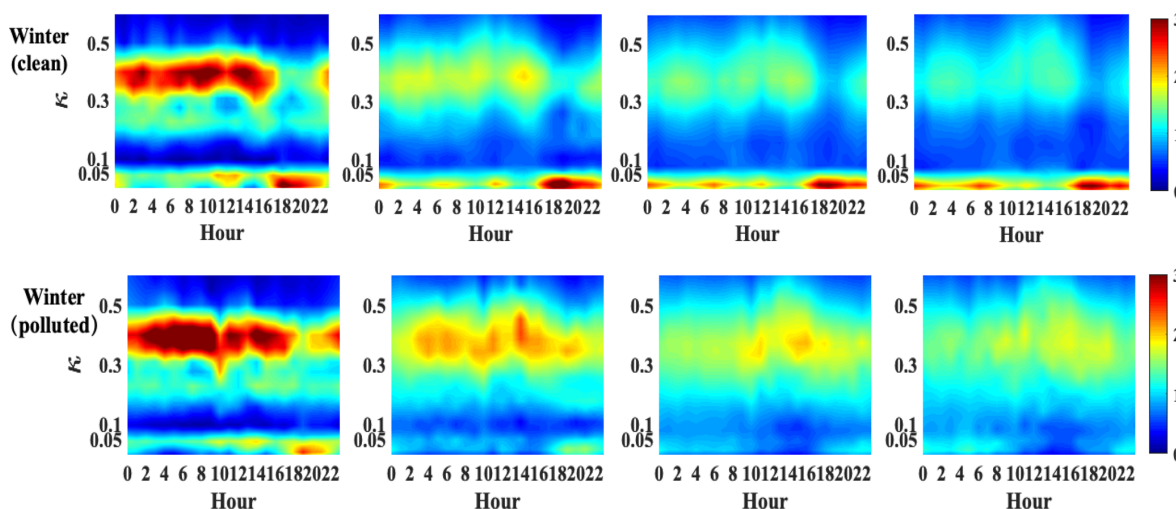


Figure 7. Diurnal cycles of κ_{gf} PDF for 80, 110, 150, and 200 nm particles in clean and polluted events in winter.

trate, sulfate, SOA) on preexisting aerosols at lower temperature and/or hydrophilic SOA formation under higher relative humidity at nighttime (Wu et al., 2008; Wang et al., 2016; An et al., 2019). However, such a condensation effect during nighttime is less significant (indicated by the smaller disparity between κ_{chem} and κ_{gf}) than the aging effect caused by aerosol photochemical processes around noontime (J. F. Peng et al., 2016).

We suppose that the higher/lower κ_{chem} should first be closely associated with temporal changes in actual effective density of BC with the particles aging/diurnal variations of local emissions. It has been demonstrated that rapid aging of BC can occur over a few hours in the polluted urban area (J. F. Peng et al., 2016). The externally mixed BC particles have a fractal structure and chain-like aggregates and have been reported with an effective density of $0.25\text{--}0.45\text{ g cm}^{-3}$ (McMurry et al., 2002), while the BC particles in the κ_{chem} calculation are assumed to be void-free with an effective density of 1.7 g cm^{-3} . This leads to a lower BC volume fraction than it actually is and thus a greater κ_{chem} during the traffic rush hour and cooking time when BC particles are mostly freshly emitted with uncompacted structure. In addition, the significant increase in volume fraction of POA during the late afternoon would result in changes in composition of organic aerosols, and thereby a density much closer to that of POA than the assumed one (1.2 g cm^{-3}) in the calculation should be applied. A sensitivity test has been done to examine the effect of density of BC and organics on the calculated κ_{chem} (Fig. 8). The result shows that the κ_{chem} value can be reduced by 16 %–33 % by decreasing the BC effective density from 1.7 to $0.25\text{--}0.45\text{ g cm}^{-3}$. This basically explains the disparity between κ_{chem} and κ_{gf} during the traffic rush hour when a large amount of BC is freshly emitted. The changes in κ_{chem} are within $\pm 4\%$ by varying the organic density from 1.2 (mixture of SOA and POA) to 1.0 (typically for POA) or

1.4 g cm^{-3} (typically for SOA) (Zamora et al., 2019), showing a much lower impact of variations of organic density on κ_{chem} . In conclusion, the result demonstrated that the disparity between κ_{chem} and κ_{gf} during the late afternoon in winter is largely due to the inappropriate use of the BC particle density that is closely associated with its morphology or the degree of its aging. Our study suggests that, to accurately parameterize the effect of BC aging on particle hygroscopicity, it is critical to measure the effective density and morphology of ambient BC, in particular in those regions with complex influences of rapid secondary conversion/aging processes and local sources.

In that way, the lower κ_{chem} value derived around noontime in summer, when BC aerosols may be more compact through strong photochemical aging, is probably due to application of a lower BC density in the calculation. However, the sensitivity test indicates that, to fill the gap between κ_{chem} and κ_{gf} observed at noontime in summer, the effective density of BC should be extremely high due to decreased sensitivity of κ_{chem} to BC density with its aging. In this case, the density of BC has been assumed to be 1.7 g cm^{-3} , which reflects a very compacted and void-free structure of the BC particles. This currently applied value represents an upper limit for the effective density of ambient BC particles according to previous observations near or in Beijing (Zhang et al., 2015), which suggested the aged BC generally has an effective density of $1.2\text{--}1.4\text{ g cm}^{-3}$. Using these ambient observed values would lead to further underestimation of κ_{chem} . In addition, the photochemical aging can change the overall effective density of organic aerosols by changing their chemical composition. However, the effective density of the photochemical oxidized organic particles (e.g., SOA) does not change much on the timescale of several hours and was observed ranging between 1.2 and 1.3 g cm^{-3} (Bahreini et al., 2005). It can only explain $\sim 4\%$ at most of the underesti-

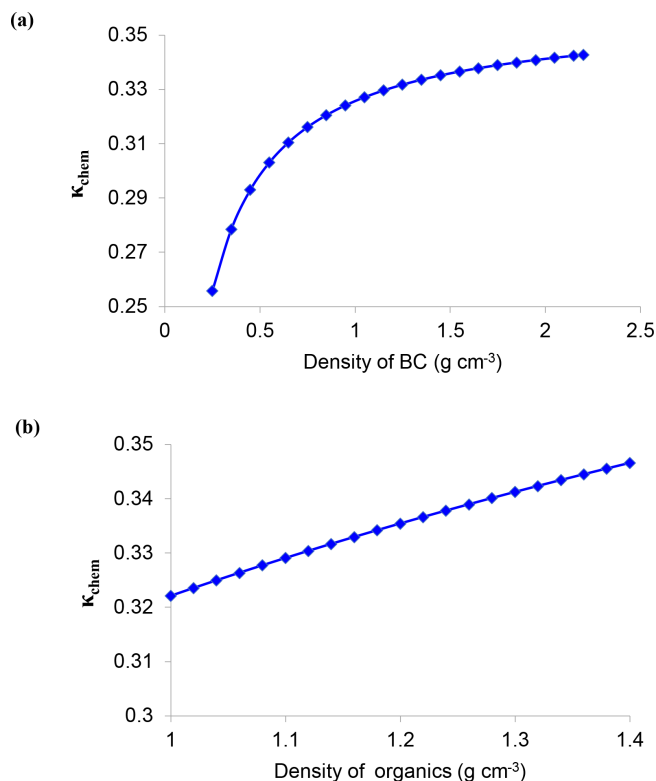


Figure 8. Sensitivity of κ_{chem} to variations in the density of BC (a) and organics (b).

mation in κ_{chem} around noontime in summer by applying a density value of 1.4 g cm^{-3} (typically for SOA). Therefore, the application of higher densities of BC and organics in the calculation cannot fully explain the disparity between κ_{chem} and κ_{gf} during the early afternoon in summer when strong photochemical processes are expected.

The uncertainty in the calculation of κ_{chem} may be also related to the uncertainty caused by the hygroscopic parameter of organics that vary widely over a range of diverse constituents of SOA (Suda et al., 2012). The lower κ_{chem} indicates that the κ of secondary organic aerosols formed through the strong photochemical oxidation processes in summer of urban Beijing is likely underestimated. In this study, the mean κ value of organics derived from the f_{44} parameterized equation is 0.20 ± 0.02 , ranging from 0.17 to 0.23 during 09:00–17:00. While the organic aerosols, especially for particles in accumulated mode, may be more hygrophilic with a much larger κ , i.e., > 0.2 due to large formation of highly oxidized OA. One can easily understand that increasing the κ of organic aerosols from 0.2 to 0.3 can explain about 11 %–13 % of the underestimation of κ_{chem} , but this represents an upper limit of the impact of hygroscopicity of organic aerosols on the calculation. This is because the κ value of 0.3 corresponds to the maximum possible for ambient organic aerosols. Additionally, the f_{44} parameterized equation tends to overestimate the κ according to Fröhlich et al. (2015),

which should yield a larger κ_{chem} . Finally, the coexisting hygroscopic and hydrophobic species may have a strong influence on the phase state of particles, also likely affecting chemical interactions between inorganic and organic compounds as well as the overall hygroscopicity of mixed particles (C. Peng et al., 2016). Overall, the lower κ_{chem} caused by the photochemical aging effect likely resulted from multiple impacts of inappropriate application of the density and hygroscopic parameter of organic aerosols in the calculation, as well as the influences from chemical interaction between organic and inorganic compounds on the overall hygroscopicity of mixed particles. This topic warrants further investigations.

3.5 Observation from other stations

The aging process in the summer period is related to photochemical processing in strong solar radiation conditions. The photochemical reactions produce sulfate and secondary organic aerosol, condensing on the surface of slightly hygroscopic or nonhygroscopic primary aerosols (such as BC) (Zhang et al., 2008). To confirm such a photochemical aging effect on particle hygroscopicity, we further examine the diurnal variations of κ_{chem} and κ_{gf} or κ_{CCNc} (only at the XZ site) based on observations in summer at two other sites in north China (Fig. 1). The XT site is located in the suburbs of XT city, which is about 400 km south of Beijing, with high levels of industrialization and urbanization. Due to industrial emissions and typically weak ventilating winds, concentrations of $\text{PM}_{2.5}$, black carbon, and gaseous precursors are usually high at the site (Fu et al., 2014). Xinzhou is located in the north of Taiyuan and about 360 km southwest of Beijing and is surrounded by mountains on three sides. Local emissions from motor vehicles and industrial activities have relatively little influence on the sampled aerosol (Zhang et al., 2016). Because of its location and elevation, the aerosol at the XZ site is usually aged and transported from other areas. The sampling period was from 22 July to 26 August 2014 and from 17 May to 14 June 2016 at the XZ and XT site respectively.

We find that the case at the XT site is very similar to that observed in BJ (Fig. 9a), with a lower κ_{chem} than κ_{gf} around noon time. But, because of much fewer influences from the local sources at XT compared to that at BJ, such underestimation by κ_{chem} continued until night at XT (Fig. 9b). Interestingly, a noontime lower κ_{chem} was not observed in the diurnal cycles at the XZ site, where κ_{chem} and κ_{CCNc} had similar diurnal patterns (Fig. 9c) with a roughly constant ratio of κ_{chem} to κ_{CCNc} of ~ 0.8 – 0.9 (Fig. 9d). This is probably because the XZ site is usually the recipient of aerosols transported from other areas that are already aged and well-mixed, with minimal impact of further aging (Zhang et al., 2017). Also, the rate of oxidation and condensation may be slow in the relatively remote area where the gas precursors and oxidants are not as high as they are closer to source regions.

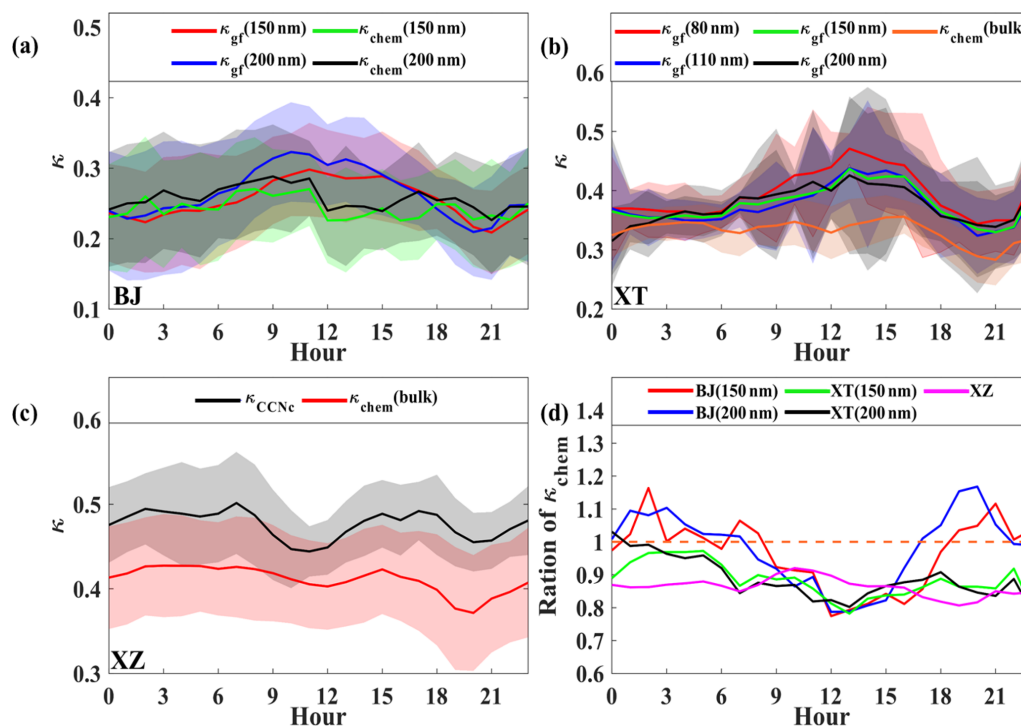


Figure 9. Diurnal variations in (a) κ_{chem} and κ_{gf} for 150 and 200 nm particles at the BJ site; (b) κ_{chem} and κ_{gf} for 40, 80, 110, 150, and 200 nm particles at the XT site; (c) κ_{chem} and mean κ_{CCNe} for particles at the XZ site; and (d) ratio of mean κ_{chem} to κ_{gf} at the three sites.

But at XT, which is located in the heavily polluted area in the North China Plain (Fu et al., 2014), aerosol emissions and processing are more similar to that in urban Beijing. These observations from other sites further confirm the photochemical aging effect that will largely underestimate the particle hygroscopicity using a simple mixing rule based on chemical composition.

4 Conclusions

Using measurements of aerosol composition and hygroscopicity made in Beijing (BJ) during a winter period of 2016 and a summer period of 2017, this paper analyzes the daily variation and seasonal differences of size-resolved aerosol hygroscopicity in urban Beijing. We mainly focus on studying the disparity of κ_{gf} and κ_{chem} between summer and winter to reveal the impact of atmospheric processes and mixing state of the particles on its hygroscopicity. The uncertainty in calculating κ by using chemical composition with a uniform internal mixing hypothesis is elucidated from the diurnal variations of the difference between the calculated values: in summer, a lower κ_{chem} is obtained around noon-time, with a ratio of κ_{chem} to κ_{gf} of about 0.8–0.9 for large particles (i.e., 150 and 200 nm), showing an underestimation of particle hygroscopicity by using a simple mixing rule based on chemical composition. Combining with the observation from XT and XZ, we attribute the underestimation to

the rapid noontime photochemical aging processes in summer, which induces the aging effect that will lead to a lower κ if assuming a uniform mixing of the particles. The lower κ_{chem} likely resulted from multiple impacts of inappropriate application of the density and hygroscopic parameter of organic aerosols in the calculation, as well as the unknown influences from chemical interaction between organic and inorganic compounds on the overall hygroscopicity of mixed particles.

In winter, a larger κ_{chem} than κ_{gf} for > 100 nm particles is derived around noontime and in the early afternoon, with a maximum ratio of κ_{chem} to κ_{gf} of 1.2–1.4 when the particles are dominated by the hydrophobic mode with a large number of externally mixed POA particles from strong vehicle and cooking sources. We attribute this large disparity between κ_{chem} and κ_{gf} to changes in BC morphology that can be indicated by the effective density of BC. The sensitivity test shows that it can well explain the disparity during the traffic rush hour by applying a BC effective density of 0.25–0.45 g cm⁻³. However, we suggest that, to accurately parameterize or account for the effect of BC density on particle hygroscopicity, future investigations need to measure the effective density of ambient BC, in particular in those regions with complex local sources. Our results highlight the impacts of atmospheric processes and sources on aerosol mixing state and hygroscopicity, which should be quantified and considered in models for different atmospheric conditions.

Data availability. All data used in the study are available on <http://www.geodoi.ac.cn/WebEn/doi.aspx?Id=1356> (<https://doi.org/10.3974/geodb.2019.06.11.V1>, Fan et al., 2019) or from the corresponding author upon request (fang.zhang@bnu.edu.cn).

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Competing interests. The authors declare that they have no conflict of interest.

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