



Supplement of

Influx of African biomass burning aerosol during the Amazonian dry season through layered transatlantic transport of black carbon-rich smoke

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- **Table S1.** List of time sections in which the UPL, CL, LPL MBL and BB were sampled during
- 3 flight AC19.

	Time window	
Offshore profile	16:50:00	19:07:00
UPL	16:21:50	16:22:40
	16:54:22	16:55:02
	17:08:22	17:09:00
	18:02:46	18:03:09
	18:03:47	18:04:18
	18:10:35	18:12:15
	18:34:32	18:35:47
	18:42:51	18:43:39
	18:47:50	18:48:04
	19:58:00	19:58:39
CL	17:09:08	17:09:30
	18:02:04	18:02:39
	18:36:08	18:38:42
	18:40:16	18:42:44
	18:43:43	18:44:24
	18:47:21	18:47:43
LPL	16:52:03	16:53:18
	17:10:00	17:11:37
	17:56:43	17:57:16
	17:57:33	17:58:15
	17:58:31	18:00:08
	18:54:33	18:55:51
MBL	17:23:14	17:27:28
Fresh BB	19:24:03	19:24:37
	19:24:49	19:25:08
	19:25:44	19:27:02
	19:27:11	19:28:16
	19:30:09	19:30:46
	19:32:24	19:34:03
	19:18:24	19:18:45



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Figure S1. Sketch illustrating the initiation of a three-day BT starting at the ATTO site (red

8 circle) mapped to the raster of fire intensities.



- 10 Figure S2. Ensemble of three-day HYSPLIT BTs, starting at every hour at the ATTO site (1000
- 11 m a.s.l.) on 30 September 2014 and corresponding daily fire intensity maps (W m^{-2}) from the
- 12 Global Fire Assimilation System (GFAS).



- 14 **Figure S3.** View from the HALO cockpit, showing the active fire plumes (intersected at ~1 km
- above ground) during flight AC19 at 19:20 UTC on 30 September 2014. See also fire map in Fig.
- 16 1.



Figure S4. Radiosonde sounding at Belem Airport (see Fig 1.) on 30 September 2014 at 12:00 UTC, provided Wyoming University (<u>http://weather.uwyo.edu/upperair/sounding.html</u>, last access on 06 August 2019). The sounding shows similar tropospheric stratification as presented in Fig. 3. The first layer (top around 1000 m) is associated with the boundary layer, the second (top around 3200 m) is related to the shallow clouds top and the third one (around 5000 m) is the large scale inversion.





Figure S5. Particle number size distributions (median and interquartile range) derived from the

26 UHSAS and SP2 (rBC + SC) for the (a) CL, (b) LPL, (c) MPL, (d) UPL and (e) fresh BB plume

27 probed during flight AC19. Panel (f) shows the curve fits of the UHSAS data points.



Figure S6. Scatterplot of the ratios f_{43} (m/z 43 to total organic signal) against f_{44} (m/z 44 29 to total organic signal) expressing the photochemical aging of the organic aerosol 30 measured by the C-ToF-AMS. The blue and green markers correspond to measurements 31 within the UPL and fresh BB, respectively. The signal at m/z 44 relates mostly to CO_2^+ 32 ions and the m/z 43 signal to $C_2H_3O^+$ ions. The triangular region (dashed lines) in the f_{44} 33 vs. f₄₃ space defines the boundaries within which most of the organic aerosol was found in 34 previous studies and can be used as a guide to characterize oxidized organic components: 35 data in the upper left represent more oxidized organics vs. the less oxidized organics in the 36 37 lower right (Ng et al., 2010; Schulz et al., 2018).



Figure S7. Correlation between Δc_{CO} and ΔM_{rBC} within the upper pollution layer (UPL) and fresh biomass burning (BB) plume probed during AC19. A bivariate regression fit was applied to the

42 data set in order to obtain the BC enhancement ration, EnR_{BC} .



43

44 Figure S8. Monthly distribution of AIRS-derived carbon monoxide (400 to 600 hPa) over the Southern Hemisphere. The map shows

45 averages over multiple years (2005-2018).



Figure S9. Hovmöller plot of the daily MODIS AOD (550 nm) distributed over the South
Atlantic region (60 °W to 20 °E) from August to October 2014, averaged over the latitudinal band
of 10° S to 5° N, corresponding to the region of interest (ROI) highlighted in Fig. 7a. Several
events of transatlantic transport of aerosol from Africa towards South America can be easily
identified, with the strongest plume starting approximately on 15 September 2014. A dashed line
for this particular event is also shown in the picture, which arrives at a time close to our
observations on AC19 on 30 September 2014. Westwards of 35 °W, the AOD levels are

- 54 increasingly influenced by the South American continent, which masks the AOD signals of the
- 55 transported African pollution as it approaches the South American continent.



Figure S10. Sensitivity tests showing the DRF-TOA changes due to different assumptions in the
aerosol and surface properties, comparing with the original DRF-TOA estimation. The
simulations show that if the aerosol layer is mostly scattering (SSA = 0.99), a general cooling
(back-scattering by the layer) is observed. On the other hand, if the absorbing ocean is replaced
by a higher surface albedo (0.6), a warming effect by the layer is observed due to the downward
radiation that is scattered back and forth between the aerosol layer and the surface.