



# Supplement of

# Temporal variation in $^{129}{\rm I}$ and $^{127}{\rm I}$ in aerosols from Xi'an, China: influence of East Asian monsoon and heavy haze events

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The supplementary information includes five figures and one table.

### S1 Determination of <sup>129</sup>I and <sup>127</sup>I in aerosol samples

#### S1.1 Aerosol sampling

The aerosol samples were collected by a high-volume sampler onto glass fibre filters (200 mm×250 mm, Tianhong Instrument
Ltd., Wuhan, China). The flow rate is 1.5 m<sup>3</sup> min<sup>-1</sup> and the sample duration was 24 h for each filter with air flux of 2100 m<sup>3</sup>. The sampler is installed on the roof of the Xi'an AMS Centre in Xi'an, China (34°13'25"N, 109°0'0"E) with an elevation of 440 m above mean sea level (Fig.1).

#### S1.2. Iodine isotopies analysis

68 aerosol filters, about four filter samples in each month, were selected for measurement of iodine isotopes. Each filter 25 represents one day information. Half of one filter with air flux of about 1000 m<sup>3</sup> was analysed for iodine isotopes, and the other half was reserved for other purpose. Iodine was separated from the aerosol filter using pyrolysis and AgI-AgCl coprecipitation in combination with accelerator mass spectrometry (AMS) for measurement, as described elsewhere (Zhang et al., 2018b). In brief, the aerosol samples were placed into a corundum boat. <sup>125</sup>I in the form of iodide was added for calculation of chemical yield. Iodine in the samples was released as gaseous form at high temperature in the atmosphere of nitrogen and

- 30 oxygen gases in a tube furnace (Hou et al., 2010). The released iodine was trapped into a solution containing 0.5 mol L<sup>-1</sup> NaOH and 0.02 mol L<sup>-1</sup> NaHSO<sub>3</sub>. An aliquot of solution (1.0 mL) was taken for determination of <sup>127</sup>I using ICP-MS (Agilent 8800, USA). Another 1.0 mL solution was taken to a tube and counted for <sup>125</sup>I using a NaI gamma counter (Model FJ2021, Xi'an Nuclear Instrument Factory, China) to calculate the chemical yield of iodine during combustion. After gamma measurement, 0.2 mg <sup>127</sup>I carrier (Woodward Company, USA) was added to the trap solution. For procedure blank samples, 0.2 mg iodine
- 35 and 0.5 mg chloride (as NaCl) was added. 1 mL of 0.5 M NaHSO<sub>3</sub> was used to reduce iodate to iodide. The solution was firstly adjusted to pH < 2 by 3 M HNO<sub>3</sub>, then 1 mL 0.5 M AgNO<sub>3</sub> solution was added to the solution to precipitate iodine as AgI-AgCl coprecipitation. The formed AgI-AgCl precipitate was washed once with 3 M HNO<sub>3</sub> to remove Ag<sub>2</sub>SO<sub>3</sub> and Ag<sub>2</sub>SO<sub>4</sub>, then washed with deionized water once and 5-20% ammonium hydroxide once to remove excessive AgCl, and finally rinsed twice with deionized water. After centrifugation, the AgI-AgCl coprecipitate was ready for AMS measurement. The procedural
- 40 blank was prepared using a blank glass fibre filter with the same procedure as that for samples.

#### S1.3 AMS and ICP-MS determination of <sup>129</sup>I and <sup>127</sup>I

The prepared AgI-AgCl coprecipitates were completely dried at 70°C, then mixed with Nb metal powder (99.9%, 325 mesh, Alfa Aesar, USA) in a mass ratio of 1:5 and pressed into copper target holders. <sup>129</sup>I in the target was measured using a 3MV AMS in the Xi'an AMS Centre (Hou et al., 2010). A voltage of 2.5 MV was applied for measurement of <sup>129</sup>I/<sup>127</sup>I number ratios.

- +5 charge state of iodine ion was selected and extracted from the accelerator by a magnetic analyser. <sup>129</sup>I/<sup>127</sup>I ratios of the iodine carrier are determined to be less than 2×10<sup>-13</sup>. The analytical precision was less than 5% for all the samples. The trapping solution was diluted by a factor of 20-50 with 1% NH<sub>3</sub>·H<sub>2</sub>O, and analysed for <sup>127</sup>I concentration by ICP-MS (Agilent 8800, USA) using the mode of single quadrupole and no dynamic collision-reaction gas. Cs<sup>+</sup> (CsCl) was used as an internal standard in the ICP-MS measurement of iodine. The sensitivity of <sup>127</sup>I is 250 Mcps per 1 mg L<sup>-1</sup> of iodine, and the
- 50 instrumental detection limit is  $0.002 \ \mu g \ L^{-1}$  for <sup>127</sup>I.

## S1.4 Calculation of z-score of <sup>129</sup>I concentrations and <sup>129</sup>I/<sup>127</sup>I ratios

Z-scores of <sup>129</sup>I concentrations and <sup>129</sup>I/<sup>127</sup>I number ratios are calculated by subtracting the sample mean from an individual raw score and then dividing the difference by the sample standard deviation (Eq.1). The absolute value of z-score represents the distance between the raw <sup>129</sup>I concentrations and <sup>129</sup>I/<sup>127</sup>I ratios and the sample mean in units of the standard deviation.

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$$z = \frac{x - \bar{x}}{s}$$
 Eq. (1)

Where,  $\bar{x}$  is the mean of the sample; S is the standard deviation of the sample; z is negative when the raw score is below the mean, positive when above.

#### S2. Influence of precipitation and win speed

Precipitation. Wet and dry deposition are vital pathways of iodine removal from the atmosphere. The effect of rainfall on

- 60 iodine concentrations in aerosols is not clear. Xi'an is a warm temperate semihumid continental monsoon climate, with annual precipitation of 522.4-719.5 mm. The annual precipitation was 649.0 mm in 2017, and precipitations in September and October were the most months of 98.6 mm and140 mm, respectively, accounting for 37% of the annual precipitation (Fig.2e) (Xi'an Bureau of Statistics, 2018). Taking the two months for exampling, sixteen aerosols were analysed with eight in rainy days and eight in non-rainy days. <sup>127</sup>I and <sup>129</sup>I concentrations fell in the ranges of 1.88-4.93 ng m<sup>-3</sup> and (1.88-4.93) ×10<sup>5</sup> atoms m<sup>-3</sup> in
- $^{65}$  rainy days, respectively, were comparable to 1.67-8.22 ng m<sup>-3</sup> and (0.44-7.25) ×10<sup>5</sup> atoms m<sup>-3</sup> in non-rainy days. Although the concentration range was narrower in rainy days than non-rainy days, the data in same range suggest that precipitation does not significantly affect the variation of both iodine isotopes in aerosols. Furthermore, this conclusion is also supported by the fact that the frequent rainfall in October did not change iodine concentrations.

Wind speed. Wind speed affect not only the sources of iodine, but also the dispersion rate and retention in local atmospheric system. Controlled by topography, the annually prevailing wind direction is northeasterly wind in Xi'an and daily average wind speed was 1.0-4.1 m s<sup>-1</sup> during the studied periods. <sup>127</sup>I and <sup>129</sup>I varied irregularly with changes of wind speed throughout the year (Fig.2f). These data indicated that small-scale atmospheric circulation limited within a local area unlikely had regular influence on variations of iodine isotopes, which is identical to the observation in Risø, Denmark (Zhang et al., 2016). Large-scale atmospheric circulation, however, might be profound, which will be discussed in the following section.

#### 75 Reference

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90 Fig. S1 Relationship between <sup>127</sup>I and <sup>129</sup>I with a weak correlation (R=0.33, p<0.01) between the two iodine isotopes. This indicates the two iodine isotopes have different sources and their temporal variation patterns were affected by different factors.



Fig. S2 Relationship between iodine isotopes and total suspended particles (TSP) in Xi'an, China (n=68), suggesting significant correlation between <sup>127</sup>I and TSP, and no correlation between <sup>129</sup>I and TSP. The results indicate <sup>127</sup>I was sourced from local input and <sup>129</sup>I was transported to the studied site externally.



Fig. S3 Back trajectories analysis on date of a) 18th April, 2017; b) 18th May, 2017; c) 14th July, 2017; d) 31st August, 2017; e) 6th 100 September, 2017; f) 15th November, 2017; g) 28th December, 2017; h) 17th January, 2018.



Fig. S4 Relations between <sup>127</sup>I and air pollutants including PM10, PM2.5, SO<sub>2</sub>, NO<sub>2</sub>, CO and O<sub>3</sub>, showing significant correlation.



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Fig. S5 850 hPa water vapor transmission flow field on 2 May, 2017 (a), and 21 May, 2017 (b). Data from: https://cmdp.ncccma.net/Monitoring/monsoon.php?ListElem=vt85. The red dot in the figures is the sampling location, Xi'an, China.

Table S1 Mean <sup>129</sup>I concentrations and <sup>129</sup>I/<sup>127</sup>I ratios in three high-level periods (HLP) and two low-level periods (LLP)

No	Туре	Start date	Stop date	$N(^{129}I) / (10^5 \text{ m}^{-3})$		<sup>129</sup> I/ <sup>127</sup> I number ratio / (×10 <sup>-10</sup> )		Managanatara
				Averag e	RSD	Averag e	RSD	inionisoon stage
1	HLP 1	28 Mar, 2017	22 May, 2017	2.37	91%	101	89%	WM and onset of SM
2	LLP 1	23 May, 2017	25 Jul, 2017	0.49	60%	28.5	65%	Active of SM
3	HLP 2	4 Aug, 2017	12 Sep, 2017	1.98	109%	155.	141%	Break of SM
4	LLP 2	21 Sep, 2017	11 Oct, 2017	0.66	44%	40.1	44%	Revival of SM
5	HLP 3	13 Oct, 2017	20 Mar, 2018	2.41	44%	67.9	83%	SM retreat and WM advance then active