

The UV-visible absorption cross-sections of IONO₂

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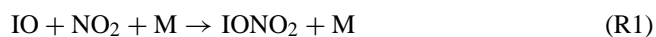
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Abstract. The UV-visible absorption spectrum of gaseous IONO₂ has been measured over the wavelength range 245–415 nm using the technique of laser photolysis with time-resolved UV-visible absorption spectroscopy. IONO₂ was produced in situ in the gas phase by laser flash photolysis of NO₂/CF₃I/N₂ mixtures. Post flash spectra were deconvolved to remove contributions to the observed absorption from other reactant and product species. The resulting spectrum attributed to IONO₂ consists of several overlapping broad absorption bands. Assuming a quantum yield of unity for IONO₂ photolysis, model calculations show that during sunlit hours at noon, 53° N, the first order solar photolysis rate coefficient (*J* value) for IONO₂ is $4.0 \times 10^{-2} \text{ s}^{-1}$.

1 Introduction

Iodine chemistry has been implicated in tropospheric ozone depletion since the 1980s (Chameides and Davis, 1980; Davis et al., 1996; Jenkin et al., 1985). Atmospheric iodine source gases are readily photolysed by sunlight, yielding iodine atoms. Iodine atoms react with ozone forming IO. Any subsequent reactions of IO forming an iodine atom lead to catalytic ozone loss. In competition, IO can be sequestered into reservoir species. One of the main gaseous reservoir species for iodine in the atmosphere is believed to be IONO₂ (Chameides and Davis, 1980; Davis et al., 1996; Jenkin, 1992; Jenkin et al., 1985).

IONO₂ is formed by the reaction of IO with NO₂



The rate of formation of IONO₂ in reaction (R1) is well known (De More et al., 1997) but the amount of IONO₂

residing in the atmosphere and its consequent impact on the atmospheric chemistry depends on its loss processes. Gas phase IONO₂ loss processes include photolysis, thermal decomposition and heterogeneous uptake.

IONO₂ can be photolysed via two possible routes:



However, no studies of the absorption cross-sections, the photolysis quantum yields or the products of IONO₂ photolysis have been reported to date.

Thermal decomposition of IONO₂ could proceed as in reaction (R3).



However, recent observations imply that reaction (R3) is a slow process at ambient temperature (Dillon, 2001; Allan and Plane, 2002).

A recent laboratory study has provided the first experimental evidence that IONO₂ can be lost from the gas phase *via* a heterogeneous process (Holmes et al., 2001). It was shown that IONO₂ is taken up efficiently by dry and frozen salt surfaces to form the di-halogens IBr and ICl, which are then released back into the gas phase.

IONO₂ has been synthesised in the laboratory in solution. However, IONO₂ was found to be unstable and the first attempts to observe its ultraviolet-visible absorption spectrum failed (Schmeisser and Braendle, 1961). The absorption cross-sections for IONO₂ in the UV-visible are essential, in order to evaluate the daytime atmospheric photolysis rate of IONO₂ and the efficiency of the tropospheric ozone loss cycle involving IONO₂.

Broske and Zabel (1998) attempted to measure the gas phase UV-visible absorption spectrum of IONO₂ by photolysis of I₂/NO₂ mixtures in a large reaction chamber equipped

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Table 1. Concentrations used and generated during an experiment

	Before laser photolysis	After laser photolysis/O atom reaction
N ₂ /molecules cm ⁻³	2.5 × 10 ¹⁹	2.5 × 10 ¹⁹
NO ₂ /molecules cm ⁻³	4.0 × 10 ¹⁵	3.9 × 10 ¹⁵
CF ₃ I/molecules cm ⁻³	5.5 × 10 ¹⁶	5.5 × 10 ¹⁶
CF ₃ /molecules cm ⁻³	–	7.5 × 10 ¹³
IO/molecules cm ⁻³	–	7.5 × 10 ¹³
NO/molecules cm ⁻³	–	8.5 × 10 ¹³
NO ₃ /molecules cm ⁻³	–	1.0 × 10 ¹²

for in situ spectroscopy. A diode array detector was used to monitor the UV-visible absorption spectrum. However, because of the low concentrations of IONO₂ produced under their experimental conditions and the presence of strong absorbers such as INO₂, produced by the reaction of iodine atoms from I₂ photolysis with NO₂, they were unable to retrieve the spectrum of IONO₂. Nevertheless, the UV-visible absorption spectrum of gaseous INO₂ was reported (Atkinson et al., 2000).

In order to avoid the effects of complicating chemistry involving iodine atoms produced in the IO source reactions, as well as the effects of secondary chemistry, a pulsed laser photolysis system and a photochemical source of IO with no initial iodine atom production (NO₂/CF₃I) was used in this study. We report here, for the first time observation of a gas phase ultraviolet-visible spectrum attributed to IONO₂. The results were used to calculate the atmospheric photolysis rates of IONO₂, and the atmospheric chemistry of IONO₂ is discussed briefly.

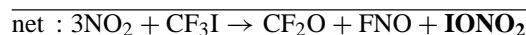
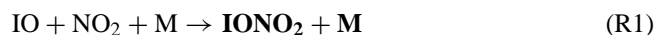
2 Experimental

The IONO₂ absorption spectrum was obtained using the technique of laser flash photolysis with UV absorption spectroscopy, utilising a CCD detection system. The apparatus has been described in detail elsewhere (Rowley et al., 1996), therefore only a brief description is given here.

During an experiment, gases were flowed slowly and continuously through a 98.2 cm long cylindrical quartz cell kept at a total pressure of 760 Torr. Reactions were initiated in the cell using a pulsed excimer laser beam passed longitudinally along the cell. Concentrations of reactant and product gases in the cell were monitored using UV-vis absorption spectroscopy. The analysis light beam from a xenon arc lamp was passed through the cell, counter propagating the laser beam, and focused onto the entrance slit of a 0.25 m focal length astigmatic Czerny-Turner spectrograph fitted with three interchangeable diffraction gratings ruled at 150, 300 and 600 grooves/mm. In order to protect the spectrograph and detector from impinging reflected laser light a narrow

band filter was used in the 350 nm region. Wavelength resolved analysis light from the spectrograph was then imaged across the top 30 rows of a CCD detector, to record spectra. Signal (charge) transfer then enabled recording of sequential spectra, which were acquired before, during and after photolysis.

A NO₂/CF₃I/N₂ chemical system was employed to generate IONO₂ via the following reaction scheme:



The reaction was initiated using 351 nm XeF excimer laser photolysis of NO₂ to form O(³P) (DeMore et al., 1997). N₂ (Messer, 99.996%), 2000 ppm NO₂ in N₂ (Messer) and CF₃I (Fluorochem, 99%) were used as supplied. Gas mixtures were flowed through the reaction cell at a total pressure of 760 Torr and at 298 K via mass flow controllers or ball flow meters (for NO₂ in N₂) and their concentrations (see Table 1 for concentrations used) were calculated from known calibrated flow rates. In a single experiment 1152 spectra were collected on the CCD with a clocking speed of ~ 1 μs per pixel. The actual time resolution was ~ 15 μs resulting from exposure of 30 pixels of the array. Post flash absorbance spectra were calculated relative to preflash spectra using Beer's law ($A = \ln(I_{\text{preflash}}/I_{\text{postflash}})$). Consequently observed absorbances show changes in absorption brought about by the laser flash and by subsequent chemistry over a period of ~ 1 ms. The total absorption spectrum of the gas mixture was recorded using a 150 g/mm and a 300 g/mm grating giving a 130 nm and a 65 nm wavelength coverage, respectively. In the 245–335 nm region the 150 g/mm grating was used with a 50 μm spectrograph entrance slit giving a resolution (fwhm) of 1.67 nm. In the 385–415 nm region the 300 g/mm grating was used with a 10 μm entrance slit and a resolution (fwhm) of 0.55 nm. In both cases time de-

Table 2. Chemical reaction scheme used in the FACSIMILE model (Curtis and Sweetenham, 1987)

Reaction	Rate coeff. 295 K/cm ⁻³ molecule ⁻¹ s ⁻¹	Reference
O(³ P) + NO ₂ → NO + O ₂	1.10 × 10 ⁻¹¹	(DeMore et al., 1997)
O(³ P) + NO ₂ → NO ₃	4.20 × 10 ⁻¹²	(DeMore et al., 1997)
O(³ P) + CF ₃ I → CF ₃ + IO	1.10 × 10 ⁻¹¹ (*0.85)	(Addison, et al., 1979)
NO ₂ + CF ₃ → CF ₂ O + FNO	1.53 × 10 ⁻¹¹	(Pagsberg et al., 1998)
NO + CF ₃ → CF ₃ NO	1.89 × 10 ⁻¹¹	(Ley et al., 1995)
CF ₃ + CF ₃ → C ₂ F ₆	8.32 × 10 ⁻¹²	(Rossi and Golden, 1979)
IO + IO → 2I + O ₂	8.60 × 10 ⁻¹¹ (*0.7)	(Bloss et al., 2001)
IO + IO → OIO + I	8.60 × 10 ⁻¹¹ (*0.3)	(Bloss et al., 2001)
O(³ P) + IO → O ₂ + I	1.20 × 10 ⁻¹⁰	(DeMore et al., 1997)
IO + NO ₂ → IONO ₂	9.00 × 10 ⁻¹²	(DeMore et al., 1997)
I + NO ₂ → INO ₂	7.50 × 10 ⁻¹²	(DeMore et al., 1997)
I + NO → INO	4.50 × 10 ⁻¹³	(DeMore et al., 1997)
IO + NO → I + NO ₂	2.20 × 10 ⁻¹¹	(DeMore et al., 1997)
I + INO ₂ → I ₂ + NO ₂	8.32 × 10 ⁻¹¹	(v.d. Bergh and Troe, 1976)
I + NO ₃ → IO + NO ₂	4.50 × 10 ⁻¹⁰	(Chambers et al., 1992)

pendent spectral information from up to three experiments were averaged to increase the signal to noise ratio.

Under the conditions chosen (Table 1) numerical models show that the chemistry of IO and CF₃ dominates, as O(³P) reacts mainly (93%) with CF₃I. In addition, only small amounts of NO are expected (from NO₂ photolysis and the O(³P) + NO₂ reaction (7%)) and I atom production from the reaction IO + NO → I + NO₂ is consequently unimportant. Thus, only small amounts of INO₂ via reaction I + NO₂ + M → INO₂ + M are formed. IONO₂ is expected to be the only major iodine containing product formed under the experimental conditions used, as the reaction of NO₂ with IO dominates over other possible reactions with IO. Thus, the IONO₂ spectrum can in principle be obtained from the experimental spectrum, after subtraction of the absorption due to CF₂O and FNO formed via reaction (R6) and addition of absorbance due to NO₂ and CF₃I consumed in reactions (R4) and (R5), respectively. However, it should be noted that the yield for reaction (R5) is 0.85 (Gilles et al., 1996). The identity of the remaining reaction channels is unknown (Gilles et al., 1996). Other products that could be formed via reaction (R5) such as IF or CF₃IO could also contribute to the overall absorption spectrum obtained. However, since their absorption cross-sections have not been reported to date it is not possible to correct for their contribution to the overall absorption. Other known secondary and side reactions that occur were taken into account by using the numerical integration package FACSIMILE (Curtis and Sweetenham, 1987) to model the post flash chemistry in order to calculate the final product distribution. Details of the chemical reaction scheme used are listed in Table 2. The contribution of secondary products including INO₂ will be discussed in the results and discussion section below. Spectra were not recorded below

245 nm or between 335–385 nm, due to low light levels of the analysis light beam in these regions, the latter due to the attenuation by the laser beam filter in the 350 nm region. The overall error on the IONO₂ cross-sections including the uncertainty due to the unknown reaction channels mentioned above was estimated to be 25%. The chemical mechanism adopted in the FACSIMILE model could not be fully tested, as all the products could not be diagnosed using our detection method and the time resolution of the experiment (~ 15 μs) did not allow the kinetics of product formation to be followed. The estimated error of 25% excludes possible systematic errors that could result from errors in the chemical mechanism used in the FACSIMILE model.

3 Results and discussion

The absorption spectrum recorded in the short wavelength region (245–335 nm) of the spectrum (Fig. 1), i.e. the absorbance calculated relative to pre-laser intensities, represents the change in absorption due to the overall chemical reaction initiated by the laser flash. In the analysis it was assumed that only the 5 principal species (NO₂, CF₃I, FNO, CF₂O and IONO₂) contribute to this absorption. Thus a negative change in absorption results from NO₂ and CF₃I consumption and a positive change in absorption due to IONO₂, CF₂O and FNO production make up the observed spectrum. The post flash spectrum obtained was reproducible and remained constant over the time-scale of the spectral acquisition (the photolysed gas flowed out of the cell over a period of 10–15 s). The total absorption spectrum measured and the absorption spectrum of IONO₂ can be expressed as shown in Eq. (1) and (2), respectively.

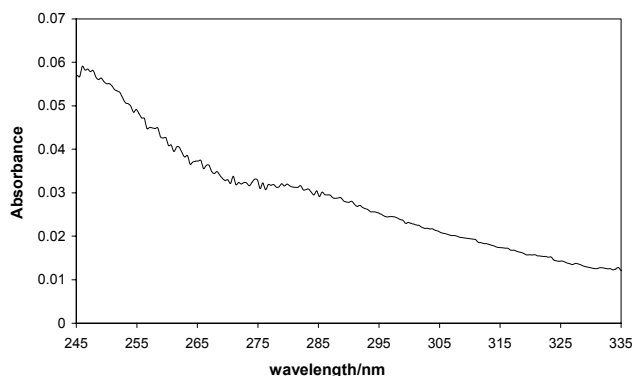


Fig. 1. Post flash absorption spectrum in the low wavelength region (245–335 nm) following flash photolysis of NO₂-CF₃I mixtures.

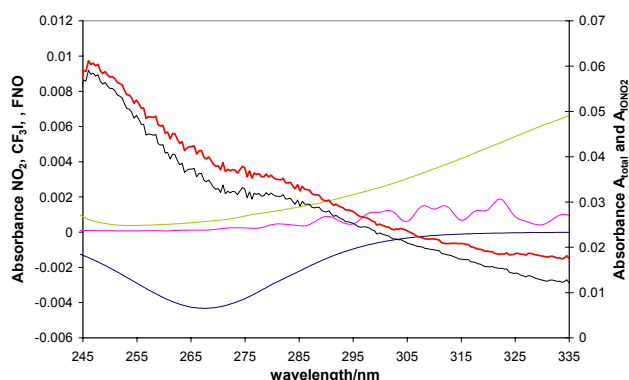


Fig. 2. Spectral stripping procedure in the low wavelength region (245–335 nm). Note changes in scale on the left and right hand axis. Black, pink, green and blue lines show the total absorption spectrum and absorption of FNO, NO₂, and CF₃I, respectively. The red spectrum is attributed to IONO₂ after the stripping procedure. The contribution of CF₂O is not shown as it only absorbs below 245 nm.

$$A_{\text{total}} = A_{\text{IONO}_2} + A_{\text{CF}_2\text{O}} + A_{\text{FNO}} - A_{\text{NO}_2} - A_{\text{CF}_3\text{I}} \quad (1)$$

$$A_{\text{IONO}_2} = A_{\text{total}} - A_{\text{CF}_2\text{O}} - A_{\text{FNO}} + A_{\text{NO}_2} + A_{\text{CF}_3\text{I}} \quad (2)$$

Thus IONO₂ was found by spectral deconvolution. The contribution of each absorber to the total absorption at each wavelength recorded during an experiment is given by:

$$A(\lambda) = \sigma_i(\lambda) \times l \times \Delta c_i \quad (3)$$

where $A(\lambda)$ is the absorbance of a species i at wavelength λ , l is the pathlength of the reaction cell used [cm], Δc = change in concentration of species i [molecules cm⁻³] and $\sigma_i(\lambda)$ = absorption cross-section of species i at that wavelength [molecule cm⁻²].

In order to obtain the IONO₂ spectrum the contribution of all species to the total absorbance change other than IONO₂

has to be known. The concentrations of all species were calculated by simulating the post flash chemistry, and the main contributions (NO₂, CF₃I, FNO, CF₂O) were determined from the known cross-sections of these species. Other possible contributors were I₂, OIO, I₂O₂, but the yields of these together with their known spectra were judged to be unimportant. Exceptions were the contribution of CF₃IO, for which we have no information about the cross-sections, and the contribution of INO₂, which is discussed in detail below.

The absolute cross-sections of CF₃I, FNO and CF₂O have been reported in the literature (DeMore et al., 1997) and were smoothed to the spectral resolution used for the present experiment (CF₂O only absorbs at wavelength shorter than 245 nm and its contribution to the overall absorption is thus not shown in Fig. 2). The change in concentration of each species after the laser flash is determined by the O atom concentration produced from NO₂ photolysis in the flash and the stoichiometry of subsequent chemistry. In order to determine these concentration changes a kinetic model of the reaction system was constructed in the numerical integration package FACSIMILE (Curtis and Sweetenham, 1987). The model provided the simulated change in concentration of each species during an experiment for all given experimental conditions with the initial post flash O atom concentration being the only variable. In separate experiments NO₂ was photolysed in the absence of CF₃I and the initial O atom concentration was determined from the change in NO₂ concentration following the flash, which was obtained from the measured change in absorption of NO₂ and its cross-section. It was assumed that $[O] = 2 \Delta \text{NO}_2$. Together with the known cross-sections, the contribution of each species to the absorption spectra was then subtracted from the overall post flash spectrum recorded as shown in Fig. 2 for the 245–335 nm region of the spectrum. The same procedure was applied to the data obtained in the longer wavelength region (385–415 nm) as shown in Fig. 3. However, in this wavelength region only NO₂ absorbs significantly and thus the NO₂ spectrum had to be subtracted from the total absorption spectrum. The fact that the structural features of FNO and NO₂ do not appear to a significant extent in the final spectrum provides support for the deconvolution procedure used.

Secondary chemistry generates about 15% of INO₂ compared to IONO₂ in the laser flash. The cross-sections for INO₂ are known (Atkinson et al., 2000) and together with its simulated concentration the contribution of INO₂ to the total absorption spectrum recorded during an experiment was also subtracted. Figure 4 shows the IONO₂ absorption spectrum obtained from subtracting the main components as described above together with the INO₂ absorption spectrum and the final IONO₂ absorption spectrum corrected for the contribution of INO₂ (Atkinson et al., 2000). Subtraction of INO₂ does not significantly change the shape of the IONO₂ absorption spectrum and the change in the magnitude of the cross-sections lies within the estimated uncertainty. The absorption cross-sections of IONO₂ were then obtained assuming

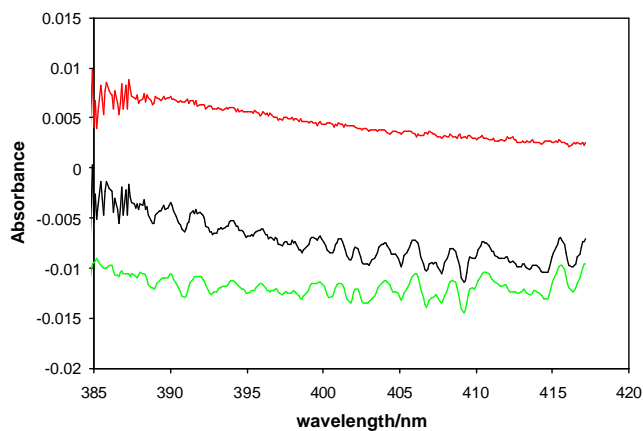


Fig. 3. Spectral stripping procedure in the high wavelength region (385–415 nm). Black, green and red lines show the total absorption spectrum, absorption of NO₂ and IONO₂ after the stripping procedure, respectively.

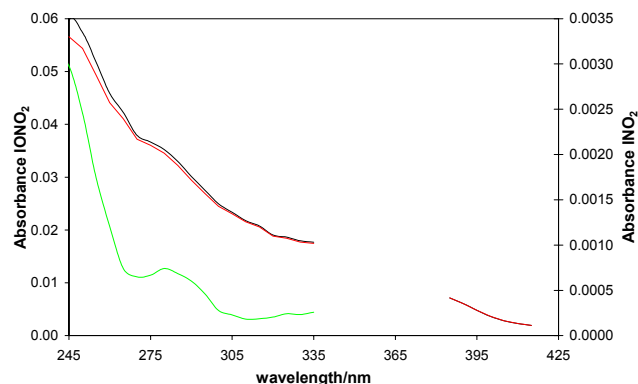


Fig. 4. Comparison of the absorption spectrum of IONO₂ and INO₂ generated during an experiment. Note changes in scale on the left and right hand axis. Black, green and red line show the spectrum originally assigned to IONO₂ obtained from the initial deconvolution process, the absorbance due to INO₂ generated following the laser flash, and the final IONO₂ absorption spectrum after correction for the contribution of INO₂, respectively.

that the remaining absorption was attributed to the simulated amount of IONO₂ generated in the reaction.

The IONO₂ absorption spectrum over the entire wavelength range of 245–415 nm was obtained by fitting the observed spectrum to three Gaussian functions. This enabled interpolation of the IONO₂ absorption in spectral regions where absorbance could not be recorded. The resulting spectrum consists of a broad continuous absorption band as shown in Fig. 5. The IONO₂ cross-sections are listed in Table 3. There is also a possibility of a contribution to the apparent IONO₂ spectrum from formation of light scattering aerosol produced in the IO radical chemistry. A maximum potential contribution to the absorption from aerosol

Table 3. The absorption cross-sections of IONO₂ at 298 K

λ/nm	$10^{18} \sigma_{298}/\text{molecule cm}^{-2}$
245	12.1
250	11.7
255	10.6
260	9.46
265	8.80
270	7.97
275	7.72
280	7.41
285	6.91
290	6.31
295	5.77
300	5.25
305	4.95
310	4.62
315	4.41
320	4.04
325	3.96
330	3.80
335	3.74
340	3.60
345	3.48
350	3.34
355	3.16
360	2.94
365	2.70
370	2.42
375	2.13
380	1.84
385	1.53
390	1.30
395	1.03
400	0.780
405	0.605
410	0.496
415	0.416

was calculated assuming all absorption at $\lambda = 450$ nm was due to aerosol extinction and that the attenuation at lower wavelength followed a $1/\lambda^4$ functional dependence on wavelength (Cox and Coker, 1983). After subtraction of this component the lower curve in Fig. 6 was obtained which corresponds to minimum values for the IONO₂ cross-sections. Examination of time resolved absorption spectra also show no evidence for time delayed aerosol growth for the first two experiments performed in a clean flow tube. After rapid formation of IONO₂ the absorption remained constant over the whole time period of the experiment. Only after multiple experiments in the flow tube did we observe a time delayed increase in overall absorption with time, possibly due to the formation of aerosols, presumably from precursors formed in heterogeneous processes on the walls of the reaction cell.

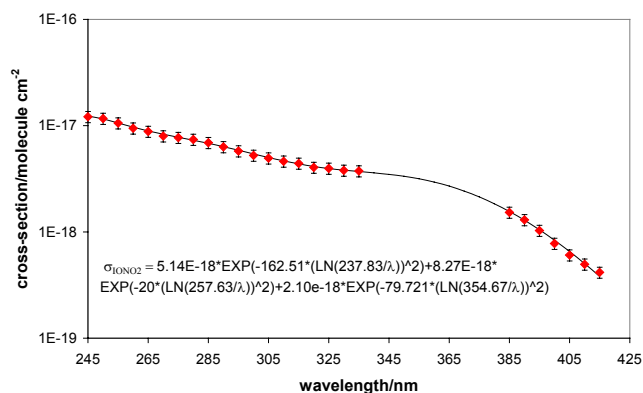


Fig. 5. Absorption cross-sections of IONO₂. Red diamonds: data recorded in the 245–335 nm and in the 385–415 nm wavelength region. Black line: Fit to data.

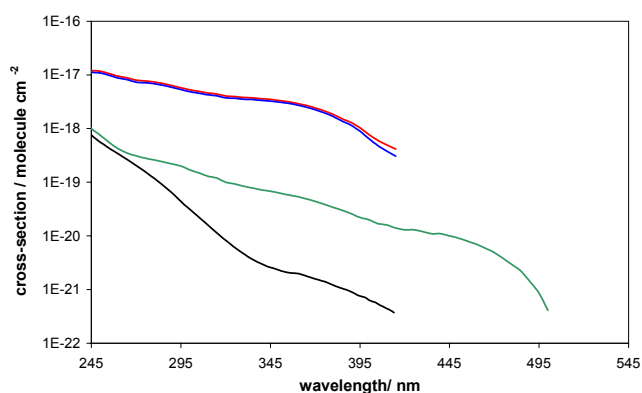


Fig. 7. Comparison of the absorption spectra of ClONO₂ (black), BrONO₂ (green), IONO₂ (red) and (IONO₂-aerosol) contribution (blue).

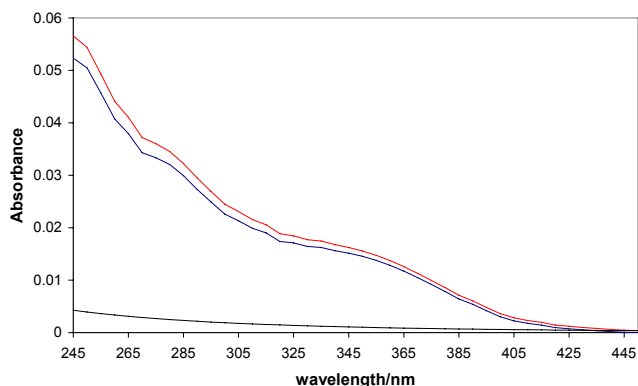


Fig. 6. Comparison of the absorption spectrum of IONO₂ (red), the predicted absorption due to aerosol formation as a function of $1/\lambda^4$ (black) and the contribution due to aerosol absorption subtracted from the original absorption spectrum of IONO₂ (blue).

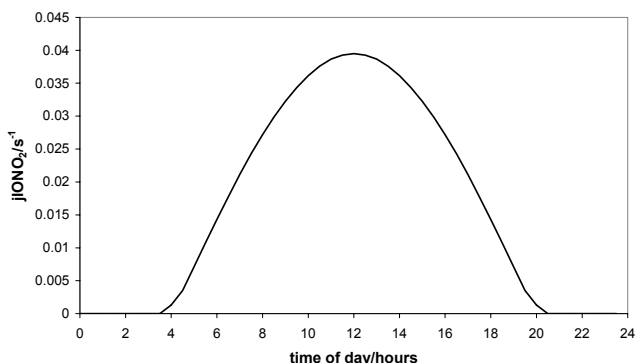


Fig. 8. J -values for IONO₂ as a function of time of day in the marine boundary layer, for midsummer 53° N (Mace Head, Ireland, July 1997).

Figure 7 shows a comparison of the IONO₂ absorption and the (IONO₂-aerosol) absorption with those of ClONO₂ (Burkholder et al., 1994) and BrONO₂ (Burkholder et al., 1995). All spectra show similar broad band absorption features. The absorption due to IONO₂ is red shifted and is significantly more intense at all wavelengths than the chlorine and bromine analogues.

4 Atmospheric implications

The IONO₂ absorption cross-sections recorded here were used to calculate the tropospheric solar photolysis rate of IONO₂ using a photochemical box model (McFiggans et al., 2000). We assumed a quantum yield of unity for IONO₂ photolysis throughout the absorption band. The resulting photolysis rate coefficient (J -value) for IONO₂ is shown in Fig. 8 as a function of time of day in the marine boundary layer, for July 53° N at Mace Head, Ireland, where mea-

surements of atmospheric composition had taken place in 1997. The calculated photolytic lifetime of IONO₂ under these conditions is less than one minute. At the same location and time, the surface to volume ratio (S/V) of the fine fraction of the sea-salt aerosol present was determined to be $(6 \pm 1) \times 10^{-7} \text{ cm}^2 \text{ cm}^{-3}$ with an effective particle radius of $2.5 \times 10^{-5} \text{ cm}$ (McFiggans et al., 2000). Using these values, the lifetime of IONO₂ with respect to heterogeneous uptake on sea-salt aerosols, τ_{het} [s], was calculated to be $\sim 10 \text{ h}$ using Eqs. (4) and (5) below, and assuming a reactive uptake coefficient of $\gamma = 0.01$ (Holmes et al., 2001). Equation (4) assumes spherical geometry of the sea-salt particle. It also takes diffusion of IONO₂ to the particle surface into account (Turco et al., 1989).

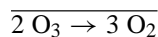
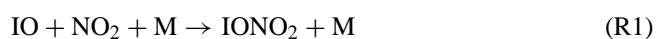
$$k_{\text{het}} = \frac{0.25 \times \gamma \times \omega \times S/V}{(1 + (3 \times \gamma \times r)/(4 \times \lambda))} \quad (4)$$

$$\tau_{\text{het}} = \frac{1}{k_{\text{het}}} \quad (5)$$

where ω = the average molecular speed of IONO₂ at any given temperature [cm s^{-1}], k_{het} = first order rate constant for loss of IONO₂ on the aerosol [s^{-1}], λ = mean free path of a molecule in air at any given temperature and pressure [cm].

During daytime, solar photolysis is thus expected to be the major loss process for IONO₂ and IO_x is not lost from the gas phase. However, the reaction of IONO₂ with sea-salt aerosol may become a significant loss process for IONO₂ during night time. Uptake of IONO₂ is expected to form the di-halogens ICl and IBr (Holmes et al., 2001), which are released into the gas phase during the night. At dawn, IBr and ICl are photolysed. Hence IO_x would not be lost permanently from the gas phase via this reaction as suggested previously (Chatfield and Crutzen, 1990; Davis et al., 1996; Jenkin, 1992), and the uptake of IONO₂ represents a route to liberation of bromide and chloride from sea-salt aerosols, as has been predicted by more recent modelling studies (McFiggans et al., 2000; Vogt et al., 1999). Thermal decomposition of IONO₂ may also contribute to the night time loss of IONO₂. However, recent observations have shown that the lifetime of IONO₂ with respect to thermal decomposition at 298 K is approximately 6 min with a lower limit of 2.4 min (Allan and Plane, 2002).

Depending upon the photolysis product yields, photolysis of IONO₂ could contribute directly to ozone destruction via the following reaction cycle:



However, it is important to note, that this cycle will not operate if the products of IONO₂ photolysis are IO + NO₂. No studies of the photolysis products of IONO₂ have been reported, and such studies are essential to elucidate further the role of atmospheric IONO₂.

5 Conclusions

The UV-visible cross-sections of IONO₂ have been determined over the wavelength ranges 245–335 nm and 385–415 nm. Both data sets were used to interpolate an IONO₂ absorption spectrum over the wavelength range of 245–415 nm. The resulting spectrum consists of several overlapping broad absorption bands. Based on the IONO₂ absorption cross-section, the tropospheric solar photolysis rate coefficient of IONO₂ has been calculated. This indicates that the photochemical lifetime of atmospheric IONO₂ is

less than one minute under most sunlit conditions and consequently that photolysis is the main loss process for gaseous IONO₂. Thus, IONO₂ is thus not expected to act as a stable reservoir of atmospheric iodine during daytime.

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