



Supplement of

Peroxy radical measurements by ethane – nitric oxide chemical amplification and laser-induced fluorescence during the IRRONIC field campaign in a forest in Indiana

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1 S1. Calibration of NO₂ monitors

1				
2 3	Three cavity attenuated phase-shift spectrometry (CAPS) instruments (Aerodyne Research) were used for measuring NO ₂ (Kebabian et al., 2008;Kebabian et al., 2005). Two were			
4	dedicated for the measurements of NO ₂ as part of the ECHAMP measurement of peroxy radicals.			
5	The third CAPS instrument was used for measuring ambient NO ₂ . The NO ₂ monitors were			
6	calibrated by sampling diluted $NO_2(g)$ from a liquid permeation tube (Kin-Tek). The output of			
7	the permeation tube (held at 40° C) was diluted into 100 sccm of N ₂ and the	nen into variable flow		
8	rates (4000 to 8000 sccm) of either zero air or purified ambient air to make multiple points in the			
9	calibration curve. The purified ambient air was prepared by passing ambient air through a			
10	scrubber filled with sodium permanganate and activated charcoal (Purafil brand SP Blend			
11	Media). The concentrations of NO_2 delivered from the permeation tube were quantified by a			
12	chemiluminescence analyzer (Model 42i Trace Level, Thermo Scientific) where the NO2 was			
13	converted to NO (NO _x mode) in a molybdenum converter held at 325 °C. See section 3 below for			
14	chemiluminescence sensor calibration information. The CAPS NO2 measurements were also			
15	checked by comparing the ECHAMP readings when in "Ox" (background) mode to			
16	measurements of O ₃ by a UV-absorption monitor (2B Tech model 202, accuracy 2%). The two			
17	methods agreed to within 5% (Wood and Charest, 2014).			
18				
19	S2. ECHAMP calibration			
20	The ECHAMP sensor was calibrated using the acetone photolysis method (Wood and Charest,			
21	2014). Photolysis of acetone vapor produces almost equimolar concentrations of methyl peroxy			
22	(CH_3O_2) and peroxyacetyl $(CH_3C(O)OO)$ radicals:			
23				
24	$CH_3C(O)CH_3 + hv (254 nm) + 2O_2 \rightarrow CH_3O_2 + CH_3C(O)OO$	(S1)		
25				
26	Following reaction with excess NO, these RO_2 radicals will produce NO_2 via the following reaction with excess NO, these RO_2 radicals will produce NO_2 via the following reaction with excess NO_2 radicals will produce NO_2 via the following reaction with excess NO_2 radicals will produce NO_2 via the following reaction with excess NO_2 radicals will produce NO_2 via the following reaction with excess NO_2 radicals will produce NO_2 via the following reaction NO_2 via the following r			
27	reactions:			
28		$(\mathbf{C2})$		
29	$CH_3C(0)O + NO - CH_3C(0)O + NO_2$	(S2)		
30 21	$CH_{3}C(0)O + M + O_{2} \neq CH_{3}O_{2} + CO_{2} + M$	(S3)		
22	$CH_{3}O_{2} + NO - CH_{3}O + NO_{2}$ $CH_{2}O_{1} + O_{2} - \Delta HO_{2} + HCHO$	$(\mathbf{S4})$		
22	$CH_{3}O + O_{2} \neq HCHO$ $CH_{2}O + NO + M \rightarrow CH_{2}ONO + M$	(SJa) (S5b)		
33	$HO_{0} + NO \rightarrow OH + NO_{0}$	$(\mathbf{S50})$		
35	102 ± 100 2011 ± 1002	(50)		
36	Ignoring the formation of methyl nitrite (Reaction S5b) each CH_2O_2 radi	cal would produce two		
37	NO ₂ molecules and each $CH_2C(\Omega)OO$ would produce three NO ₂ molecules. The change in NO ₂			
38	observed when the radical source is modulated on and off, effected by diverting the acetone flow			
39	away from the carrier flow that is illuminated by the UV source, is related to the RO ₂ concentration			
40	by the following equation:			

42
$$([CH_3O_2] + [CH_3C(O)OO]) = \Delta NO_2 / (2.44 \times 0.95)$$
 (S7)

the equation would be exactly 2.5. The two factors in the denominator account for these twoprocesses as described in Wood and Charest (2014).

As described in the main text, during the field July 2015 field deployment we produced acetone vapor by flowing air over the headspace of dilute aqueous acetone rather than over pure acetone. Unfortunately this produced variable amounts of blue light-absorbing compounds (possibly glyoxal, methyl glyoxal, or diacetyl) which interfered with the CAPS detection of NO₂.

52 As a result we relied on laboratory calibrations performed in the laboratory rather than in-field

53 calibrations (Fig S1).



54

Fig. S1. Amplification factors obtained for ECHAMP using the acetone photolysis method. Uncertainty bars reflect the 2σ accuracy of 19%.

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59 **S3. Sampling losses in the ECHAMP inlet.**

60 Sampled air flowed through a glass cross that is internally coated with halocarbon wax and into the two FEP/PFA reaction chambers, both of which comprise a 1/4" PFA tee and 1/4" OD, 61 62 0.156" (0.4 cm) ID FEP tubing. The total residence time in the cross was approximately 18 ms. 63 We quantified potential sampling losses in the cross in two ways -1. by quantifying the effective first order wall loss rate constant of HO₂ and isoprene peroxy radicals onto halocarbon wax-coated 64 65 glass of the same dimensions, and 2. by comparing the ECHAMP signal when an HO₂ source was 66 used to overflow the sampling cross and comparing to the signal when the HO₂ source directly 67 overflowed one of the reaction chambers (at the PFA tee).

68 The wall loss rate constant measurements for several types of material will be fully 69 described in a separate manuscript. Briefly, peroxy radicals were produced by illumination of 70 humidified air (8 – 10 LPM) by UV radiation from a mercury lamp: 71

72
$$H_2O + O_2 + UV (185 \text{ nm}) \rightarrow HO_2 + OH$$

(S8)

73

A 50 sccm flow of 0.1% CO was added to convert all OH into HO₂. Similarly, adding 50 sccm of isoprene (40 ppm, balance N_2) to the flow converted all OH into isoprene peroxy radicals,

76 producing a mixture of 50% HO_2 and 50% isoprene peroxy radicals. This source was used to

overflow a quartz tube internally coated with halocarbon wax connected to the sampling cross, and
 the transmitted radicals were quantified by ECHAMP. Four different lengths of tubing were used:

the transmitted radicals were quantified by ECHAMP. Four different lengths of tubing were use $70 = 147 \text{ cm} - 86^{\circ} - 25^{\circ}$ and 0° (i.e., no tube)

79 147 cm, 86", 25", and 0" (i.e., no tube).

80 The loss rate constants increased with RH, and at 60% RH were $1.6 \pm 0.6 \text{ s}^{-1}$ for HO₂ and 81 approximately 0.9 for HO₂/isopreneRO₂, indicating lower losses for isoprene RO₂ than for HO₂. 82 This suggests losses of HO₂ were only 3% during the 18 ms sampling time. Losses of CH₃O₂ 83 radicals were similarly investigated and showed negligible losses (< 1%) onto halocarbon wax 84 and other fluoropolymers for sampling times under 1 second.

Similarly, the second method – comparing the ECHAMP signal when sampling a radical source through the sampling cross or directly into one of the reaction chambers – indicated overall losses of less than 4% for an HO₂ source.

88

89 **S4. Calibration of NO chemiluminescence monitor.**

The Thermo 42i-TL chemiluminescence monitor was calibrated by dilution of gas from a 30 ppm NO standard cylinder with zero grad air using MKS brand mass flow controllers (model 1179A). The flow rates from these flow controllers agreed to within 1% when measured by separately calibrated flow meters (Definer 220, BIOS/Mesa Labs). The humidity dependence of the chemiluminescence sensor was determined by humidification of the diluent zero air.

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96 S5. Baseline measurements for NO, NO₂, and O₃ measurements.

97 Baseline (zero) measurements were executed every 10 minutes for the NO, NO₂, and O₃ 98 measurements by overflowing their common inlet with purified air. This air was prepared by 99 drawing outdoor air sequentially through a PTFE filter, a diaphragm pump, 800 cm³ of KMnO₄(s), 100 600 cm³ of a blend of KMnO₄ and activated charcoal, and finally a second PTFE filter.

101

103 S6. Calculated Ozone Production Rates

104 Net formation of ozone occurs when peroxy radicals oxidize NO to NO₂, followed by 105 photolysis of NO₂ (Seinfeld and Pandis, 2012;Finlayson-Pitts and Pitts Jr, 1999;Haagen-Smit et 106 al., 1954). Therefore, the instantaneous gross O₃ production rate (or more accurately, O_X 107 production rate where $[O_X] \equiv [O_3] + [NO_2]$) can be calculated by the following equation:

108

109 $P(O_3) = k_{XO_2 + NO}[XO_2][NO]$ (S9)

110

111 where k_{XO2+NO} is a weighted rate constant for the reaction of the various peroxy radicals with

112 NO. P(O₃) measurements are useful for assessing the temporal profile of ozone production, help 113 to quantify local production versus transport, and can identify the chemical regime (NOx-limited

114 vs. NOx-saturated) of an air mass. We use a value of 9×10^{-12} cm³ molecule⁻¹ s⁻¹ for the value of

115 k, reflecting a reasonable assumption that isoprene peroxy radicals and HO₂ had large

116 contributions to the total peroxy radical concentration. These two peroxy radicals react with NO

- 117 with rate constants of 9×10^{-12} cm³ molecule⁻¹ s⁻¹ and 8.8×10^{-12} cm³ molecule⁻¹ s⁻¹, respectively
- 118 (Atkinson et al., 2004;Sander et al., 2006). We note that the chemical amplification technique

does not detect the portion of organic peroxy radicals that form organic nitrates (RONO₂) upon reaction with NO; thus no correction for organic nitrate yields are needed in equation 2.

121 P(O₃) values calculated based on 15-min average concentrations of the related chemical 122 species are shown in Fig. S1 along with XO₂ radicals, O₃ and NO during the IRRONIC campaign 123 over the time period of 16 July - 25 July. The missing P(O₃) data on Fig. 6 are due to unavailability 124 of either NO or XO₂ measurements due to calibrations or technical problems with the chemiluminescence instrument. 15-min average P(O₃) values between 9:00 and 21:00 were at most 125 126 9.4 ppb hr⁻¹, with significant inter-day variability. For example $P(O_3)$ exceeded 7.0 ppb/hr for several hours on 18 July but never exceeded 5.0 ppb/hr on 22 or 16 July. Peak P(O₃) values 127 occurred between 9 and 11 am, with average values between 3.3 and 7.8 ppb hr⁻¹. 128



Fig S2. Temporal variations of a) calculated ozone production rate (P(O₃)), b) total peroxy

- radicals (XO₂), c) NO and d) O₃ during the IRRONIC campaign over the time period of 16 July to 25 July. The missing values of $P(O_3)$ are related with the unavailability of either XO₂ or NO
- 134 measurements.

- 135 The observed P(O₃) values at our study site are in general lower than those observed in urban
- 136 areas, which have exceeded 50 ppb h^{-1} in Mexico City and Houston (Cazorla et al.,
- 137 2012;Kleinman et al., 2005;Shirley et al., 2006). The main reason is that both the NO
- 138 concentrations and primary HOx production rates (from $O(^{1}D) + H_{2}O$ and the photolysis of
- 139 HONO and oxygenated VOCs) were significantly lower during the IRRONIC campaign
- 140 compared to those reported in the mentioned urban areas. $P(O_3)$ was highest in the late morning
- 141 (9-11 am) when NO was highest as well. The overall positive correlation between P(O₃) and
- 142 [NO] suggests that ozone production regime was almost always NO_x-limited (see Fig. S2).
- 143
- 144
- 145



- 146 147
- 147
- 148

149 **Fig. S3.** Relationship between $P(O_3)$ and NO during the daytime (09:00 to 21:00) over the time

- 150 period of 13-25 July.
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152 S6. Comparison of Peroxy radical speciation predicted by RACM2, RACM2-LIM1, MCM 153 3.2, and MCM 3.3.1

154

155 The four figures below show the modeled composition of peroxy radicals predicted by the four 156 chemical mechanisms.





59 Fig S4. Peroxy radical concentrations predicted by the four chemical mechanisms

Table S1. Summary of modeled and measured concentrations and ratios between 13:00 and 18:00.

	16 Jul	22 Jul	24 Jul
Measured	28.4	38.9	58.6
[XO ₂]			
[HO ₂ *]	26.9	34.5	41.5
[XO ₂]/[HO ₂ *]	1.06	1.13	1.41
MCM32	38.1	44.1	55.2
[XO ₂]			
[HO ₂ *]	29.8	31.4	38.3
[XO ₂]/[HO ₂ *]	1.39	1.41	1.45
MCM331	49.8	47.5	57.2
[XO ₂]			
[HO ₂ *]	35.2	32.8	38.9
[XO ₂]/[HO ₂ *]	1.42	1.46	1.48
RACM2	66.1	56.7	69.4
[XO ₂]			
[HO ₂ *]	50.3	42.4	51.1
[XO ₂]/[HO ₂ *]	1.32	1.34	1.36
RACM2-LIM1	81.3	67.4	79.2
[XO ₂]			
[HO ₂ *]	60.3	49.3	57.5
[XO ₂]/[HO ₂ *]	1.35	1.37	1.38

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167 **References:**

168 Atkinson, R., Baulch, D., Cox, R., Crowley, J., Hampson, R., Hynes, R., Jenkin, M., Rossi, 169 M., and Troe, J.: Evaluated kinetic and photochemical data for atmospheric chemistry: Volume I-170 gas phase reactions of O x, HO x, NO x and SO x species, Atmos Chem Phys, 4, 1461-1738, 2004. 171 Cazorla, M., Brune, W. H., Ren, X., and Lefer, B.: Direct measurement of ozone 172 production rates in Houston in 2009 and comparison with two estimation methods, Atmos. Chem. 173 Phys., 12, 1203-1212, 10.5194/acp-12-1203-2012, 2012. 174 Finlayson-Pitts, B. J., and Pitts Jr, J. N.: Chemistry of the upper and lower atmosphere: 175 theory, experiments, and applications, Academic press, 1999. 176 Haagen-Smit, A., Bradley, C., and Fox, M.: Ozone Formation in Photochemical Oxidation 177 of Organic Substances, Rubber Chemistry and Technology, 27, 192-200, 1954. 178 Kebabian, P. L., Herndon, S. C., and Freedman, A.: Detection of Nitrogen Dioxide by 179 Cavity Attenuated Phase Shift Spectroscopy, Analytical Chemistry, 77, 724-728, 180 10.1021/ac048715y, 2005. 181 Kebabian, P. L., Wood, E. C., Herndon, S. C., and Freedman, A.: A practical alternative to 182 chemiluminescence-based detection of nitrogen dioxide: cavity attenuated phase shift 183 spectroscopy, Environmental science & technology, 42, 6040-6045, 2008. 184 Kleinman, L. I., Daum, P. H., Lee, Y. N., Nunnermacker, L. J., Springston, S. R., 185 Weinstein-Lloyd, J., and Rudolph, J.: A comparative study of ozone production in five U.S. 186 metropolitan areas, Journal of Geophysical Research-Atmospheres, 110. 187 doi:10.1029/2004JD005096, 2005. 188 Sander, S. P., Friedl, R., Golden, D., Kurylo, M., Moortgat, G., Wine, P., Ravishankara, 189 A., Kolb, C., Molina, M., and Finlayson-Pitts, B.: Chemical kinetics and photochemical data for 190 use in atmospheric studies evaluation number 15, 2006. 191 Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics: from air pollution 192 to climate change, John Wiley & Sons, 2012. 193 Shirley, T. R., Brune, W. H., Ren, X., Mao, J., Lesher, R., Cardenas, B., Volkamer, R., 194 Molina, L. T., Molina, M. J., Lamb, B., Velasco, E., Jobson, T., and Alexander, M.: Atmospheric 195 oxidation in the Mexico City Metropolitan Area (MCMA) during April 2003, Atmos. Chem. Phys., 196 6, 2753-2765, 2006. 197 Wood, E. C., and Charest, J.: Chemical Amplification - Cavity Attenuated Phase Shift 198 Spectrometer Measurements of Peroxy Radicals, Anal. Chem., 86, 10266-10273, 2014. 199